



1 **Characterizing water solubility of fresh and aged secondary organic**
2 **aerosol in PM_{2.5} with the stable carbon isotope technique**

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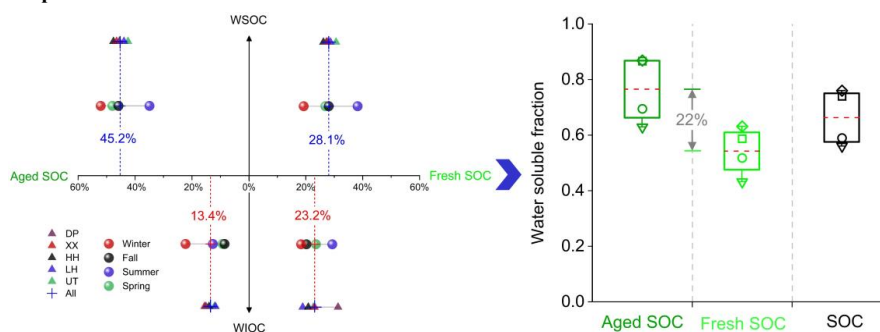
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9 **Abstract:** The investigation of the water-soluble characteristics of secondary organic carbon (SOC) is
10 essential for a more comprehensive understanding of its climate effects. However, due to the limitations
11 of the existing source apportionment methods, the water solubility of different types of SOC remains
12 uncertain. This study analyzed stable carbon isotope and mass spectra signatures of total carbon (TC)
13 and water-soluble organic carbon (WSOC) in ambient PM_{2.5} samples for one year and established stable
14 carbon isotope profiles of fresh and aged SOC. Furthermore, the Bayesian stable isotope mixing (BSIM)
15 model was employed to reveal the water solubility characteristics of fresh and aged SOC in a coastal
16 megacity of China. WSOC was dominated by secondary sources, with fresh and aged SOC contributing
17 28.1 % and 45.2 %, respectively. Water-insoluble organic carbon (WIOC) was dominated by primary
18 sources, to which fresh and aged SOC contributed 23.2 % and 13.4 %. We also found the aging degree
19 of SOC has considerable impacts on its water solubility due to the much higher water-soluble fraction of
20 aged SOC (76.5 %) than fresh SOC (54.2 %). Findings of this study may provide a new perspective for
21 further investigation of the hygroscopicity effects of SOC with different aging degrees on light extinction
22 and climate change.

23 **Keywords:** Fresh SOC; Aged SOC; Water solubility; Stable carbon isotope; BSIM model; Mass
24 spectrometry.

25 **Graphical abstract:**



26



27 **1. Introduction**

28 As a major component of particulate matter ($PM_{2.5}$), secondary organic aerosols (SOA) not only
29 contribute to haze formation but also exert a substantial influence on climate dynamics across various
30 spatial scales, from local to global (Kaul et al., 2011; Shrivastava et al., 2017). The water solubility,
31 considered one of the crucial physical properties of SOA, has been extensively studied recently due to
32 its significant effects on the physicochemical processes in the atmosphere. The water solubility of SOA
33 varied with its aging degrees (Kirillova et al., 2013), while both the water solubility and aging degree of
34 organic aerosols contribute to the hygroscopicity noticeably, which affects the light extinction eventually
35 (Han et al., 2022; Liu et al., 2022). Hence, exploring the water solubility characteristics of SOA with
36 different aging degrees can help elucidate the more detailed extinction mechanism of SOA. In addition,
37 recent studies have also shown that the formation of secondary particulates is one of the main processes
38 determining the amount of CCN in remote oceanic regions (Liu and Matsui 2022). Therefore,
39 investigating the water solubility of SOA with different aging degrees is also meaningful for further
40 exploring its indirect climate effects.

41 Investigating the contributions of SOA with different aging degrees to both organic matter (OM)
42 and water-soluble organic matter (WSOM) is imperative for determining their quantified water solubility.
43 However, due to the constraints of reliable methods, only a limited number of studies have examined the
44 water solubility of SOA using mass spectrometry techniques. Qiu et al. (2019) conducted source
45 apportionment of OM in PM_1 and WSOM in $PM_{2.5}$ based on online and offline AMS-PMF methods
46 respectively (Qiu et al., 2019). This approach faces challenges not only related to the inherent errors of
47 online versus offline methods but also discrepancies in the measured particle sizes of OM and WSOM.
48 Kondo et al. (2007) and Timonen et al. (2013) attempted to apportion water-soluble organic carbon



49 (WSOC) through a multiple linear regression method based on the mass spectral information of OM,
50 which still exhibits large indeterminateness (Timonen et al., 2013; Xiao et al., 2011; Kondo et al., 2007).
51 The carbon isotopic technique offers a promising avenue to overcome the aforementioned limitations,
52 thereby enabling a more in-depth exploration of the water-soluble characteristics of SOA. Carbon isotope
53 techniques have garnered widespread attention and are increasingly employed in source apportionment
54 studies of organic aerosols due to their robust source appointment capabilities. Radioactive carbon
55 isotopes (^{14}C) provide a precise method for quantitatively distinguishing between fossil and non-fossil
56 organic aerosol sources (Fushimi et al., 2011; Zhang et al., 2014). The stable carbon isotope technique
57 (^{13}C), however, can quantitatively assess the contributions of various sources by integrating them into
58 mass balance models (Yao et al., 2022; Widory et al., 2004). The Bayesian mixing model stands out as
59 one of the most widely utilized models (Xiao;Xu and Xiao 2023; Tang et al., 2020). The stable carbon
60 isotope technique can also be combined with other source tracers to further enhance the accuracy of
61 source apportionment of carbonaceous aerosols (Jiang et al., 2022; Plasencia Sánchez et al., 2023;
62 Ceburnis et al., 2011; Lim et al., 2022). However, to our knowledge, no study has employed the carbon
63 isotope technique to estimate the source contribution of both fresh and aged SOA before, owing to the
64 challenging measurement of the carbon isotope profiles for these two sources.

65 Previous studies have predominantly concentrated on assessing the water solubility of SOA at inland
66 urban sites, revealing a strong correlation between SOA water solubility and urban air pollution
67 emissions as well as relative humidity (Wong;Zhou and Abbatt 2015; Pye et al., 2017; Favez et al., 2008;
68 Salma et al., 2007; Weber et al., 2007; Miyazaki et al., 2006). Nevertheless, few researchers have noticed
69 the differences between inland and coastal cities. As dynamic interfaces between urban and marine
70 environments (Donaldson and George 2012), coastal cities exhibit unique characteristics. Shenzhen is a



71 typical representative city for coastal air pollution studies with a coastline spanning 260.5 km and a total
72 sea area of 1145 m². We measured the stable carbon isotope end-members of fresh and aged secondary
73 organic carbon (SOC), which enables us to investigate the source contributions of SOC with different
74 aging degrees to WSOC and their respective water solubility in Shenzhen.

75 The aim of this study is to investigate the water solubility of SOC in PM_{2.5}, emphasizing Shenzhen
76 as a representative mega-coastal city in China. We analyzed stable carbon isotopes and mass spectra
77 signatures of total carbon (TC) and WSOC in ambient PM_{2.5} samples that were collected from five
78 distinct sites in Shenzhen over one year as well as specific emission sources. For the first time, we
79 employed the Bayesian stable isotope mixing (BSIM) model on localized source profiles to quantify the
80 contributions of fresh SOC and aged SOC to WSOC and water-insoluble organic carbon (WIOC). These
81 results would contribute to estimating the water solubility of both fresh and aged SOC, revealing their
82 direct or indirect implications for climate change.

83 **2. Material and methods**

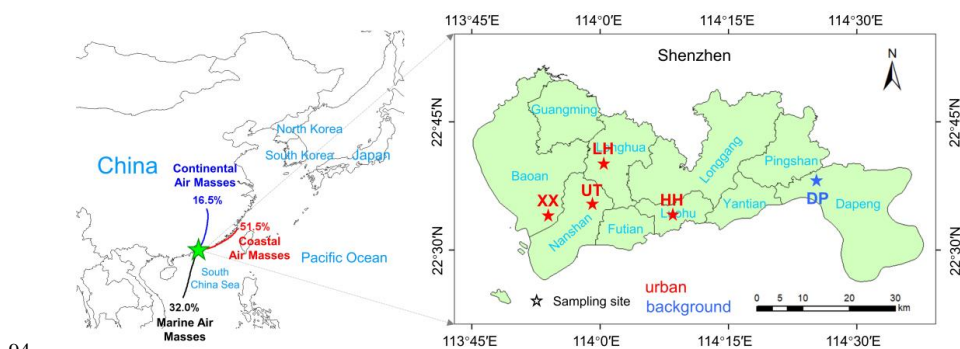
84 **2.1 Ambient PM_{2.5} sampling and chemical analysis**

85 Shenzhen (N22°27' ~ N22°52', E113°46' ~ E114°37'), one megacity of Pearl River Delta, China, is
86 bordered by Daya Bay and Dapeng Bay to the east, the Pearl River Estuary and Lingding Sea to the west,
87 Hong Kong to the south, and Dongguan and Huizhou to the north. As a typical mega-coastal city in China,
88 Shenzhen's air quality is predominantly affected by the continental air mass from northern Guangdong,
89 the eastern coastal air mass, and the southern marine air mass (Fig. 1). For a comprehensive exploration
90 of pollution characteristics in Shenzhen, PM_{2.5} samples were collected from five sites covering the
91 western to eastern regions of the city. The selected sites are Xixiang (XX, urban site), University Town



92 (UT, urban site), Longhua (LH, urban site), Honghu (HH, urban site), and Dapeng (DP, background site)

93 (Fig. 1). Additional details about each sampling site are listed in Table S1.



94

95 **Figure 1.** Spatial distribution of the five sampling sites in Shenzhen for this study.

96 In this study, 24-hour $PM_{2.5}$ sampling was conducted every other day in 2019 at the UT site using a
97 Thermo 2300 atmospheric particulate sampler (Thermo Fisher Scientific Inc., Waltham, Massachusetts,
98 USA), yielding a total of 160 valid samples. For the remaining four sites, a total of 295 valid $PM_{2.5}$
99 samples were collected every other day during typical months of the four seasons in 2019 (March, June,
100 September, and December, Table S2) using a Model TH-16A atmospheric particulate sampler (Tianhong
101 Corp., Wuhan, China). The organic carbon (OC) and elemental carbon (EC) in $PM_{2.5}$ were analyzed using
102 an OC/EC analyzer (2001A, Desert Research Institute, Reno, Nevada, USA) following the IMPROVE
103 A procedure.

104 For WSOC extraction, the $PM_{2.5}$ sample underwent ultrasonication (20 min \times 3 times) in 15 ml
105 ultrapure water (18.2 M Ω ·cm), followed by filtration through a syringe with a 0.45 μ m filter head to
106 eliminate insoluble particles. The extracted $PM_{2.5}$ samples were sequentially analyzed using a long-time-
107 of-flight aerosol mass spectrometer (L-TOF-AMS, Aerodyne, USA) and an ultrasonic nebulizer
108 (U5000AT+, Cetac Technologies Inc., USA) to measure elemental ratios, such as O/C, as well as the



109 mass spectrum signatures of the water-soluble organic fractions, including ion fragments like CO_2^+ ,
110 C_4H_9^+ , and $\text{C}_2\text{H}_4\text{O}_2^+$. The concentration of WSOC was determined using a total organic carbon analyzer
111 (multi N/C 3100, Jena, Germany), and WIOC was calculated as the difference between OC and WSOC.

112 To investigate the stable carbon isotope signatures of carbonaceous aerosols, we built a stable
113 isotope spectrometry system by integrating an OC/EC analyzer with a carbon dioxide isotope
114 spectrometer (QCLAS, Aerodyne). This system reduces the carbon requirement for isotope analysis from
115 5 μgC to 0.5 μgC and improves the accuracy of spectroscopic measurement methods to 0.2‰~0.3‰.
116 The stable carbon isotope values of TC and WSOC in ambient $\text{PM}_{2.5}$ were measured in this study.

117 **2.2 Bayesian stable isotope mixing model**

118 The BSIM model could quantify the contributions of multiple sources to the TC and WSOC based on the
119 principle of mass conservation of stable isotopes, in which the Markov Chain Monte Carlo (MCMC)
120 method was employed. The methodology employed in the BSIM model was detailed in works by Parnell
121 et al. (2013) and Parnell and Inger (2010) (Parnell et al., 2010; Parnell et al., 2013). In brief, the posterior
122 distribution for the Bayesian neural network (BNN) was calculated utilizing the prior distribution and
123 likelihood function based on Bayes theorem. Implementation of the BSIM model in this study utilized
124 the SIMMR package in R software ([https://cran.r-project.org/ web/packages/simmr/index.html](https://cran.r-project.org/web/packages/simmr/index.html)).
125 Gelman diagnostic values, ranging from 1 to 1.01, all met the criteria of the posterior prediction test,
126 indicating robust model performance and reliable results. Additionally, an uncertainty index (UI_{90}) was
127 employed here to further characterize the uncertainty strength of TC and WSOC source apportionments
128 based on their posterior distribution. This index refers to the difference between the proportional
129 contributions of the maximum and minimum values in the rapid increase segment divided by 90 with a
130 90 % cumulative probability ($\text{UI}_{90} = (\text{PC}_{95} - \text{PC}_5) / 90$) (Zaryab et al., 2022; Ji et al., 2017).



131 **2.3 Stable carbon isotope spectrum of PM_{2.5} sources**

132 The BSIM model requires the input of potential sources for carbonaceous aerosols, along with their local
133 source-specific stable carbon isotope values (end-members). The PMF model was employed to identify
134 the TC sources based on PM_{2.5} chemical species concentrations (carbon components, water-soluble
135 inorganic ions, elements, Text S1), and found traffic sources, secondary transformation sources, and
136 biomass combustion sources as the major contributors to carbonaceous aerosols in Shenzhen, which are
137 similar to the previous results in Guangzhou (Huang et al., 2014). Secondary conversion sources could
138 be further subdivided into fresh SOC for the low oxidation state and aged SOC for the high oxidation
139 state (Chen et al., 2019; Presto et al., 2009; Mahrt et al., 2022; Shen et al., 2017). Ultimately, traffic
140 emissions, fresh SOC, aged SOC, and biomass burning (BB) were identified as the four potential sources
141 of TC and WSOC in Shenzhen in this study.

142 Recognizing the regional variability in stable carbon isotope fingerprints of PM_{2.5} sources (Yao et
143 al., 2022), this work obtained representative and locally specific carbon isotope profiles for the four
144 sources in Shenzhen. For the traffic emissions, we measured the stable carbon isotope values of TC and
145 WSOC in PM_{2.5} that were collected from the Mount Tanglang tunnel (dominated by diesel vehicles) and
146 the Jiuweiling tunnel (dominated by petrol vehicles) in Shenzhen. Fresh SOC was simulated through
147 petrol vehicle bench tests, and the oxygen-carbon ratio (O/C) of PM_{2.5} samples ranges from 0.51 to 0.62,
148 indicating a low oxidation state of SOA (Ding et al., 2012). The lowest stable carbon isotope values for
149 TC and WSOC from the simulated sample were chosen as the fresh SOC results. Aged SOC samples
150 were obtained by collecting ambient PM_{2.5} samples at the National Ambient Air Background Monitoring
151 Station (Mount Wuzhi site, Hainan, China), primarily influenced by regional pollution transported by
152 northern continental air masses. These ambient PM_{2.5} samples exhibited a high O/C value of 0.98,



153 suggesting their highly oxidized state (Zhu et al., 2016). Biomass burning emissions were simulated and
 154 analyzed by burning pine wood in the Laboratory of Biomass Burning Simulation at Peking University
 155 Shenzhen Graduate School (He et al., 2010). Additional details about the sampling process are available
 156 in the Supplementary Information (Text S2). Table 1 summarizes the stable carbon isotope values of the
 157 four sources used in this study. Table S3 compares $\delta^{13}\text{C}_{\text{TC}}$ source signatures in this study with global
 158 datasets, indicating that the measurement results fall within the range of global datasets. Previous
 159 research identified $\text{C}_2\text{H}_4\text{O}_2^+$ (m/z 60) as a reliable marker for biomass burning in Shenzhen, with a feature
 160 value of 1.61 ± 0.68 % (Cao et al., 2018). This prior information was also incorporated into the BSIM
 161 model to estimate the biomass burning source.

162 **Table 1.** Stable carbon isotope end-members and f_{60} signatures for TC and WSOC sources.

	Traffic		Fresh SOC		Aged SOC		BB	
	$\delta^{13}\text{C}/\text{‰}$	$f_{60}/\%$	$\delta^{13}\text{C}/\text{‰}$	$f_{60}/\%$	$\delta^{13}\text{C}/\text{‰}$	$f_{60}/\%$	$\delta^{13}\text{C}/\text{‰}$	$f_{60}/\%$
TC	-26.26 ± 0.50	0	-27.31 ± 0.73	0	-25.54 ± 0.28	0	-27.58 ± 0.24	1.61 ± 0.68
WSOC	Traffic		Fresh SOC		Aged SOC		BB	
	$\delta^{13}\text{C}/\text{‰}$	$f_{60}/\%$	$\delta^{13}\text{C}/\text{‰}$	$f_{60}/\%$	$\delta^{13}\text{C}/\text{‰}$	$f_{60}/\%$	$\delta^{13}\text{C}/\text{‰}$	$f_{60}/\%$
	-26.68 ± 0.37	0	-26.18 ± 0.75	0	-24.93 ± 0.39	0	-26.78 ± 0.17	1.61 ± 0.68

163 **2.4 Contributions of SOC to WIOC**

164 Based on the source apportionment results from the BISM model for TC and WSOC, the contributions
 165 of fresh SOC and aged SOC to WIOC were calculated according to the equations (1-2). The uncertainties
 166 (u) in concentrations of Fresh SOC ($_{\text{WIOC}}$) and Aged SOC ($_{\text{WIOC}}$) were assessed using the uncertainty
 167 transfer equations (3-4). Fresh SOC and aged SOC uncertainties in both TC (14.9 %, 30.1 %) and WSOC
 168 (24.1 %, 20.9 %) were determined using the BSIM model. Our findings reveal that the calculated
 169 uncertainties of [Fresh SOC ($_{\text{WIOC}}$)] and [Aged SOC ($_{\text{WIOC}}$)] were 28.3 % and 36.8 %, respectively.

170
$$[\text{Fresh SOC}_{(\text{WIOC})}] = [\text{Fresh SOC}_{(\text{TC})}] - [\text{Fresh SOC}_{(\text{WSOC})}] \quad (1)$$



171
$$[\text{Aged SOC}_{(\text{WIOC})}] = [\text{Aged SOC}_{(\text{TC})}] - [\text{Aged SOC}_{(\text{WIOC})}] \quad (2)$$

172
$$u_{[\text{Fresh SOC}_{(\text{WIOC})}]} = (u_{[\text{Fresh SOC}_{(\text{TC})}]}^2 + u_{[\text{Fresh SOC}_{(\text{WSOC})}]}^2)^{1/2} \quad (3)$$

173
$$u_{[\text{Aged SOC}_{(\text{WIOC})}]} = (u_{[\text{Aged SOC}_{(\text{TC})}]}^2 + u_{[\text{Aged SOC}_{(\text{WSOC})}]}^2)^{1/2} \quad (4)$$

174 **3. Results and discussion**

175 **3.1 Overview of PM_{2.5} and carbonaceous components**

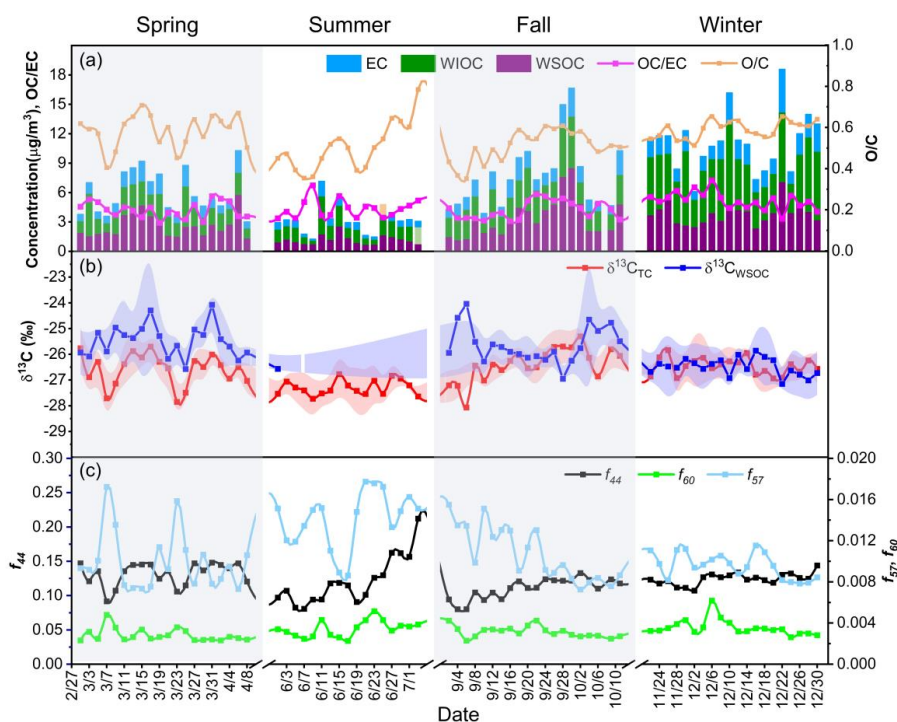
176 The annual mean concentration of PM_{2.5} in Shenzhen was 24.9 µg/m³ in 2019, with TC being the
177 predominant component, exhibiting an annual mean concentration of 7.1 µg/m³ (5.8 and 1.3 µg/m³ for
178 OC and EC, respectively). WSOC accounts for 48 % of OC, presenting an annual mean concentration of
179 2.8 µg/m³. The mean stable carbon isotope values for TC (δ¹³C_{TC}) and WSOC (δ¹³C_{WSOC}) were -26.64 ±
180 0.79 ‰ and -25.80 ± 0.88 ‰, respectively, which is lower than the results of northern cities in China
181 (Wu et al., 2020). This can be attributed to the limited impact of coal combustion (which has high ¹³C
182 values) on PM_{2.5} in Shenzhen (Yao et al., 2022; Vodicka et al., 2022).

183 Seasonal variation revealed that TC, OC, WSOC, and EC exhibited elevated levels in winter and
184 decreased levels in summer (Fig. 2a). This pattern primarily stems from pollution air masses originating
185 from continental regions in the fall and winter, and clean air masses from the southern ocean during the
186 summer months (Fig. S1). The OC to EC ratio, averaging 4.5, was also higher in winter than in summer,
187 consistent with the Oxygen-to-Carbon (O/C) ratio results for WSOC (Fig. 2a), indicating a large
188 influence of aged SOC on carbonaceous aerosols in winter. The stable carbon isotope results support this
189 observation. Figure 2b depicts relatively higher δ¹³C_{TC} and δ¹³C_{WSOC} values in spring (-26.59‰, -
190 25.26‰), fall (-26.38‰, -25.44‰), and winter (-26.46‰, -26.27‰). These higher values are attributed
191 to greater contributions of aged SOC from northern and northeast regional transport processes during



192 these seasons (Fig. S1). In summer, observed low $\delta^{13}\text{C}_{\text{TC}}$ and $\delta^{13}\text{C}_{\text{WSOC}}$ values of -27.29‰ and -26.57‰,
193 respectively, suggest relatively high contributions of fresh SOC to $\text{PM}_{2.5}$. Shenzhen experiences high
194 temperatures in summer, leading to increased gaseous precursor emissions from terrestrial biogenic
195 sources, especially C3 plants. Intense solar radiation and high temperature favor photochemical reactions
196 to generate fresh SOC that depletes ^{13}C in particulate matter during summer (Kirillova et al., 2013).

197 Mass spectra characteristics of CO_2^+ (m/z 44), C_4H_9^+ (m/z 57), and $\text{C}_2\text{H}_4\text{O}_2^+$ (m/z 60) in WSOC were
198 measured to represent oxidized organic aerosol (OOA), hydrocarbon-like organic Aerosol (HOA), and
199 biomass burning organic aerosol (BBOA), respectively. The abundance of these ion fragments, denoted
200 as f_{44} , f_{57} , and f_{60} , is determined by the ratios of signal intensities at m/z 44, m/z 57, and m/z 60 to the sum
201 of signal intensities from all m/z signals in the organic mass spectra. As depicted in Fig. 2c, f_{44} obtained
202 higher values in spring (0.131) and winter (0.125) compared to summer (0.120) and fall (0.112), further
203 indicating an elevated oxidation level of OOA during spring and winter. Considering that f_{60} exceeds
204 0.0030 when biomass burning influences carbonaceous aerosol (Docherty et al., 2008; DeCarlo et al.,
205 2008), the annual average value of f_{60} was 0.0032, suggesting biomass burning was an important source
206 of carbon components in Shenzhen. Winter exhibited higher levels of f_{60} (0.0035) compared to other
207 seasons, suggesting relatively strong impacts of biomass burning on WSOC in winter. Conversely, f_{57}
208 reached its highest level in summer (0.014) and the lowest in winter (0.009), with an annual average
209 value of 0.011, possibly associated with a notable increase in hydrocarbon organic aerosol emissions
210 from traffic and biogenic sources during the summer period.



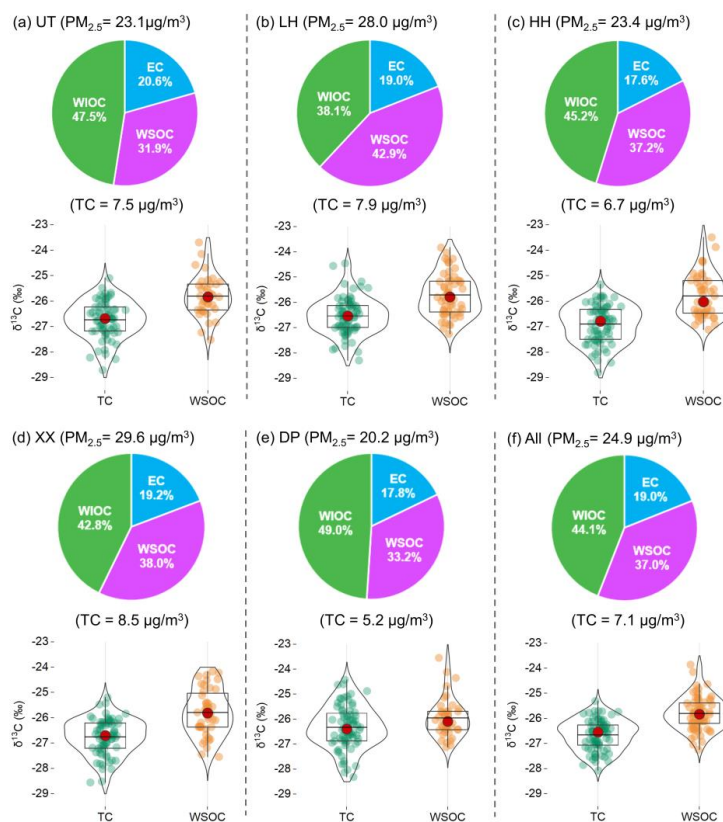
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212 **Figure 2.** Time series of carbonaceous components (a), stable carbon isotope characteristics of TC and
213 WSOC(b), and mass spectra signatures of WSOC in PM_{2.5} (c) from Shenzhen. Each data was averaged
214 from five sampling sites. (Note: Summer samples exhibit elevated analytical errors due to low
215 concentrations, and $\delta^{13}\text{C}_{\text{WSOC}}$ values are computed from combined summer samples).

216 Obvious spatial variations in PM_{2.5} mass concentrations across Shenzhen during 2019 were
217 observed, with XX site registering the highest concentration (29.6 $\mu\text{g}/\text{m}^3$), followed by LH (28.0 $\mu\text{g}/\text{m}^3$),
218 HH (23.4 $\mu\text{g}/\text{m}^3$), UT (23.1 $\mu\text{g}/\text{m}^3$), and DP (20.2 $\mu\text{g}/\text{m}^3$). Figure 3 illustrates that TC made more
219 substantial contributions (28.2 %~32.5 %) to PM_{2.5} at the four urban sites in the central and western
220 regions of Shenzhen compared to the background site (DP, 25.7 %). This suggests that local pollutant
221 emissions significantly influence carbonaceous aerosols in Shenzhen's urban areas. The percentage of
222 WSOC in TC was also higher in urban areas (37.5±3.9 %) compared to the background area (DP, 33.2 %),



223 reaching the highest value at the LH site (42.9 %). However, the percentage of WIOC in TC displayed
224 the opposite trend, suggesting carbonaceous aerosols in urban areas of Shenzhen exhibit higher water
225 solubility than in background areas. Distinct spatial distribution characteristics were also observed in the
226 stable carbon isotopes of TC and WSOC. The background site exhibits higher $\delta^{13}\text{C}_{\text{TC}}$ values (-26.33 %)
227 than the four urban sites (-26.72 ± 0.13 %). This difference may be attributed to the increased contribution
228 of traffic or fresh SOC sources to carbonaceous aerosols at urban sites and the relatively high contribution
229 of aged SOC at the background site. Atmospheric aging processes of organics through photochemical
230 reactions can deplete ^{13}C in aged SOC and enrich ^{13}C in fresh SOC and other related reactants
231 simultaneously (Pavuluri and Kawamura 2017). While the close proximity of the $\delta^{13}\text{C}_{\text{WSOC}}$ values at
232 urban sites (-25.77 ± 0.04 ‰) to the background site (DP, -25.96‰) suggests that the WSOC in different
233 areas of Shenzhen may share a similar origin.



234

235 **Figure 3.** Chemical compositions of TC, δ¹³C_{Tc}, and δ¹³C_{Wsoc} in PM_{2.5} at urban sites (a-d), background

236 site (e), and average result from all five sites (f). The Violin Box-and-Line Plots on the right display

237 spatial variations of δ¹³C_{Tc} and δ¹³C_{Wsoc} at each site, featuring mean values (black lines) and median

238 values (red dots)

239 **3.2 Source apportionment results for TC and WSOC**

240 The BSIM model assessed the contributions of traffic source, fresh SOC, aged SOC, and biomass burning

241 (BB) to TC and WSOC, as shown in Fig. 4. On average, SOC (total of fresh and aged SOC) and traffic

242 emerged as the two major contributors to TC, accounting for 43 % and 40 % respectively, while biomass

243 burning contributed 17 % to TC. The contribution of aged SOC to TC (23 %) is comparable with fresh



244 SOC (20 %). Regarding WSOC, SOC was the dominant source, comprising 45 % of aged SOC and 28 %
245 of fresh SOC, followed by BB (18 %) and Traffic (9 %). The noteworthy contribution of aged SOC to
246 WSOC suggests a comparatively higher water solubility of aged SOC in Shenzhen.

247 To evaluate the BSIM model's performance, we employed the PMF model to apportion the sources
248 of TC and WSOC. The obtained results were subsequently compared with those from the BSIM model,
249 as depicted in Fig. 4a. Seventeen chemical species of PM_{2.5} were applied as the PMF model input to
250 estimate source contributions to TC, encompassing carbon components, soluble inorganic ions, and
251 elements. For the apportionment of WSOC sources, five species including WSOC, WIOC, and three
252 organic mass spectra were applied as the PMF model input. More details about the PMF model and
253 results can be found in the Supplementary Information (Text S1, Fig. S2-S4). PMF identified the traffic
254 as the predominant contributor to TC (55 %), followed by SOC (34 %) and biomass burning (4 %).
255 Concerning WSOC, aged SOC and fresh SOC were the two major sources as well, accounting for 43 %
256 and 27 %, respectively. The traffic contribution to TC apportioned by the PMF model is higher than that
257 of the BSIM model (55 % vs. 40 %), which may be due to the fact that some of the fresh SOC generated
258 by the conversion of primary vehicle emissions was improperly apportioned to the traffic source in the
259 PMF model (Li et al., 2022; Zhao et al., 2014). Previous study also showed that SOA contributes more
260 to carbonaceous aerosols in Shenzhen than the traffic source (Cao et al., 2022). The PMF model results
261 for WSOC were generally consistent with BSIM model results, with deviations primarily attributed to
262 the differences in the principles and uncertainties of the two models.

263 Furthermore, this study examined cumulative frequency distributions to elucidate the inherent
264 uncertainty in source apportionments of TC and WSOC. As shown in Fig. S5a and b, the proportional
265 contributions of BB source to both TC and WSOC were quite stable during the research periods due to

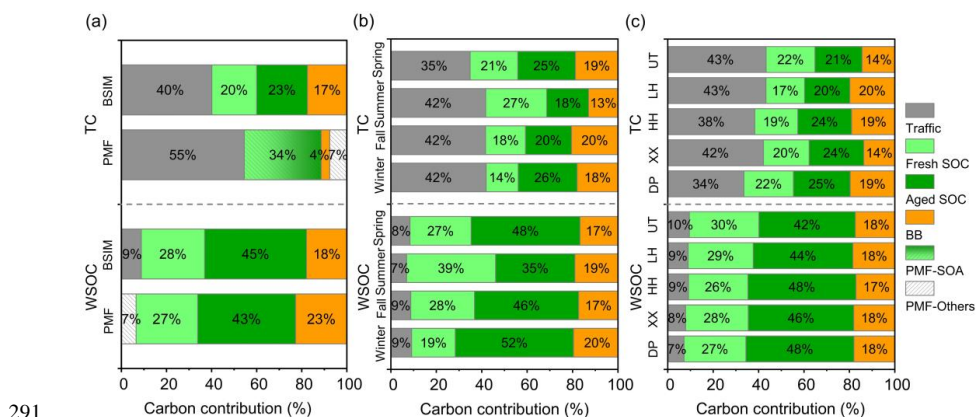


266 its low UI_{90} value (0.02). This may be attributed to the incorporation of mass spectral constraints for the
267 BB source in the BSIM model used in this study. For TC source apportionment results, the largest UI_{90}
268 value (0.46) was observed for the traffic source, indicating that its contribution to TC exhibited relatively
269 high uncertainty. In 90 % probability, its contribution ranged from 19.4 % to 60.9 %. The UI_{90} values for
270 fresh and aged SOC were 0.15 and 0.30, respectively. Regarding WSOC, the calculated UI_{90} value of
271 traffic, fresh SOC, and aged SOC ranged from 0.18 to 0.24. The UI_{90} values obtained through the BSIM
272 model remained within reasonable limits, and were smaller than those calculated in previous related
273 studies (0.23-0.62) (Zaryab et al., 2022; Ji et al., 2017). Consequently, the source contributions of TC
274 and WSOC estimated by the BSIM model in this study were deemed reasonable.

275 For seasonal variations, as shown in Fig. 4b, SOC still was the major source of TC and WSOC
276 during all four seasons, ranging from 38 % ~ 46 % and 71 % ~ 75 % respectively. Significant high
277 contributions of fresh SOC to TC and WSOC occurred in summer (27 %, 39 %), and relatively higher
278 contributions of aged SOC to TC and WSOC were observed in winter (26 %, 52 %). It is because
279 meteorological conditions in winter characterized by inversions and stagnant winds facilitate the
280 accumulation of air pollutants, and Shenzhen is largely influenced by regional pollution transport in
281 winter, favoring the formation of aged SOC (Huang et al., 2018). In contrast, favorable meteorological
282 conditions (e.g. intense and prolonged solar radiation, high temperatures, and relative humidity) in
283 summer enhanced photochemical reactions to generate fresh SOC. In terms of spatial distributions (Fig.
284 4c), the contributions of the traffic source to TC were higher at urban sites (38 % to 43 %) compared to
285 the background site (34 %). This finding aligns with expectations due to increased human activity and
286 vehicle numbers in urban locations. At the DP site, the contributions of SOC to TC were higher than
287 those of other sources (47 %), signifying a predominant influence of regionally transported pollutant



288 emissions on TC at the background site. However, the contributions of SOC and the other two primary
 289 sources at both urban and background sites were all close to each other, indicating the source composition
 290 of WSOC in Shenzhen is less affected by air pollution degree compared to TC.



291
 292 **Figure 4.** (a) Comparison of source apportionment results between BSIM model and PMF model for TC
 293 and WSOC, (b) seasonal and (c) spatial distributions of source apportionment results for TC and WSOC
 294 based on the BSIM model.

295 3.3 Water solubility of fresh SOC and aged SOC

296 The contributions of fresh SOC and aged SOC to WIOC were the differences between the contributions
 297 of those two SOC sources to TC and WSOC from the BSIM model (Sect. 2.4) in this study. As shown in
 298 Fig. 5a, fresh SOC and aged SOC made contributions of 23.2 ± 4.2 % and 13.4 ± 3.8 % to WIOC,
 299 respectively, implying that primary sources are the dominant contributors to WIOC. Further support for
 300 this finding is evident in the strong correlation between WIOC and EC, as depicted in Fig. 5b and c. A
 301 higher WIOC/EC ratio was observed in winter (2.9) than in other seasons, consistent with the highest
 302 contributions of aged SOC to WIOC in winter (22 %). This observation implies that WIOC in winter is
 303 influenced not only by local primary sources but also by the promotion of secondary pollution.

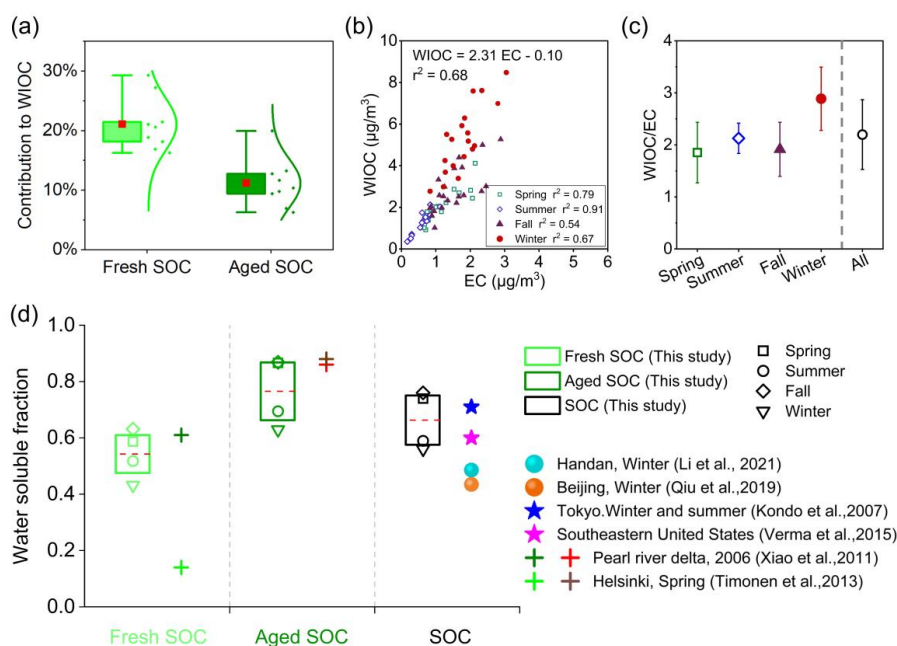


304 To investigate deeply the water solubility characteristics of fresh and aged SOC, we then calculate
305 their water-soluble fraction by comparing their water-soluble portion to the ambient fraction ($[c]_{\text{water-}}$
306 $\text{soluble}/([c]_{\text{water-soluble}} + [c]_{\text{water-insoluble}})$) (Li et al., 2021). As shown in Fig. 5d, the overall water-soluble
307 fraction of SOC in this study was 66.2 % with a range from 58.9 % to 76.0 %. Fresh SOC exhibited a
308 much lower water-solubility of 54.2 %, whereas aged SOC displayed a comparatively higher water-
309 solubility of 76.5 %. The higher water solubility of aged SOC compared to fresh SOC might be due to
310 the positive correlation between aerosol hygroscopicity and oxidation in the sub-saturated state. The
311 water-soluble fraction of SOC in this study was close to that reported in other coastal cities (Tokyo (71 %)
312 and Southeastern United States (60 %)) (Kondo et al., 2007; Verma et al., 2015), while was much higher
313 than that reported in northern Chinese cities (Beijing (42 % ~ 45 %) and Handan (49 %)) (Li et al., 2021;
314 Qiu et al., 2019). In addition, the water-soluble fraction of both fresh SOC and aged SOC, as calculated
315 in this study, was comparable to that reported in Guangzhou (61 % and 86 % for fresh and aged SOC
316 respectively) (Xiao et al., 2011). This could be attributed to Shenzhen's coastal location, which is
317 markedly influenced by regional transport from neighboring urban areas and the eastern seaboard air
318 masses. The high relative humidity facilitates the conversion of aged SOC into WSOC during the
319 pollution transport process. This result is in accordance with previous findings that air masses influenced
320 by anthropogenic emissions could promote the formation of high water-soluble SOA under high relative
321 humidity in urban environments (Miyazaki et al., 2006; Salma et al., 2007; Weber et al., 2007). Given
322 that the aging process of SOA dissolved in water could enhance the cloud condensation nuclei (CCN)
323 activity of the particles (Liu and Matsui 2022), high water-soluble aged SOC in Shenzhen might have
324 significant impacts on the activity of CCN, potentially resulting in more important indirect climate effects.

325 The water-soluble fraction of SOC (especially aged SOC) in Shenzhen exhibits obvious seasonal



326 characteristics, with the highest in fall (76.0 %) and the lowest in winter (56.0 %). This phenomenon is
 327 primary related to the robust atmospheric oxidizing capacity during fall in Shenzhen since the
 328 atmospheric oxidants such as OH and NO₃ radicals play pivotal roles in driving the secondary generation
 329 of WSOC (Wang et al., 2023). Conversely, during winter, the temperature and relative humidity are at
 330 their lowest levels, and the relatively diminished atmospheric oxidizing capacity also constrains the
 331 secondary generation of WSOC.



332
 333 **Figure 5.** (a) Box and whisker plots of fresh and aged SOC contributions to WIOC, the upper and lower
 334 of the box representing the 75th and 25th percentiles, and the red squares featuring mean values, (b)
 335 Scatterplot of WIOC versus EC by season, (c) Seasonal variation of WIOC/EC ratio, (d) Comparison of
 336 the water-soluble fraction of SOC (fresh SOC, aged SOC, SOC) in this study (box and whisker plots)
 337 with those in other related literature (colored markings on the right). The upper and lower of the box
 338 represent the 75th and 25th percentiles and the dashed red lines indicate mean values (Kondo et al., 2007;



339 Li et al., 2021; Qiu et al., 2019; Timonen et al., 2013; Verma et al., 2015; Xiao et al., 2011).

340 **4. Summary and implications**

341 Assessing the impacts of different oxidational SOC on air quality and its water solubility has been
342 challenging, and this work successfully evaluated the water-soluble fraction of fresh and aged SOC
343 employing the BSIM model on one-year observational data for stable carbon isotopes and mass spectra
344 of TC and WSOC in Shenzhen, China. Compared with other methods, e.g. PMF model, EC tracer, and
345 multiple linear regression analyses, the BSIM model successfully calculated the contributions of fresh
346 SOC and aged SOC to WSOC and WIOC, owing to prior and localized information about stable carbon
347 isotopes and mass spectra of PM_{2.5} sources. Therefore, establishing localized carbonaceous aerosol
348 source profiles for stable carbon isotopes becomes crucial for comprehending the relationship between
349 the aging degree and water solubility of SOC.

350 The observed average mass concentration of PM_{2.5} during the sampling period in Shenzhen was
351 24.9 µg/m³, and WSOC accounts for 48 % of OC. The mean stable carbon isotope values for TC ($\delta^{13}\text{C}_{\text{TC}}$)
352 and WSOC ($\delta^{13}\text{C}_{\text{WSOC}}$) were -26.64 ± 0.79 ‰ and -25.80 ± 0.88 ‰, respectively. WSOC was dominated
353 by secondary sources while WIOC was dominated by primary sources. The contribution of fresh SOC
354 and aged SOC to WSOC, WIOC were 28.1 % and 45.2 %, 23.2 % and 13.4 %, respectively. The overall
355 water-soluble fraction of SOC in this study was 66.2 %, with aged SOC constituting 76.5 % and fresh
356 SOC 54.2 %. The water-soluble fraction of aged SOC was 22 % higher than fresh SOC, even though
357 both of them demonstrated remarkable water-soluble characteristics in Shenzhen. This finding highlights
358 the important role of aged SOC in the water uptake process of particulate matter. Considering the strong
359 correlation between the water solubility of SOC and its light extinction effect, further exploration of the



360 extinction effect of SOC with different aging degrees will greatly contribute to a more profound
361 understanding of the extinction mechanism of SOC. Besides, the water solubility of SOC in coastal cities
362 was observed to be higher than that in inland cities, suggesting a more pronounced climate effect of SOC
363 in coastal cities. Therefore, there should be increased emphasis on enhancing the control of SOA
364 precursors in coastal urban areas to better integrate air pollution and climate change management. This
365 is particularly crucial given the observed rise in the proportion of SOA in particulate matter in recent
366 years. Moreover, the results of our study further hinted that the notable water solubility of SOC,
367 particularly aged SOC, may contribute a lot to the formation of CCN above coastal cities, which is also
368 helpful to a better understanding of the cloud microphysical processes and the indirect climate effect of
369 SOC in coastal urban regions.



370 **Data availability.** Datasets are available by contacting the corresponding author, Xing Peng
371 (pengxing@pku.edu.cn)

372

373 **Author contributions.** PX and HX conceptualized the study. WF, CL, TM and FN retrieved and
374 constructed the dataset. WF and PX carried out the statistical analysis. WF prepared the first draft of the
375 manuscript, which was commented on and revised by PX, HL, and HX. All authors reviewed and
376 approved the final version for publication.

377

378 **Competing interests.** The authors declare that they have no conflict of interest.

379

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