

1 **Characterizing water solubility of fresh and aged secondary organic**
2 **aerosol in PM_{2.5} with the stable carbon isotope technique**

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9 **Abstract:** The investigation of the water-soluble characteristics of secondary organic carbon (SOC) is
10 essential for a more comprehensive understanding of its climate effects. However, due to the
11 limitations of the existing source apportionment methods, the water solubility of different types of SOC
12 remains uncertain. This study analyzed stable carbon isotope and mass spectra signatures of total
13 carbon (TC) and water-soluble organic carbon (WSOC) in ambient PM_{2.5} samples for one year and
14 established stable carbon isotope profiles of fresh and aged SOC. Furthermore, the Bayesian stable
15 isotope mixing (BSIM) model was employed to reveal the water solubility characteristics of fresh and
16 aged SOC in a coastal megacity of China. WSOC was dominated by secondary sources, with fresh and
17 aged SOC contributing 28.1 % and 45.2 %, respectively. Water-insoluble organic carbon (WIOC) was
18 dominated by primary sources, to which fresh and aged SOC contributed 23.2 % and 13.4 %. We also
19 found the aging degree of SOC has considerable impacts on its water solubility due to the much higher
20 water-soluble fraction of aged SOC (76.5 %) than fresh SOC (54.2 %). Findings of this study may
21 provide a new perspective for further investigation of the hygroscopicity effects of SOC with different
22 aging degrees on light extinction and climate change.

23 **Keywords:** Fresh SOC; Aged SOC; Water solubility; Stable carbon isotope; BSIM model; Mass
24 spectrometry.

25 **1. Introduction**

26 As a major component of particulate matter (PM_{2.5}), secondary organic aerosols (SOA) not only
27 contribute to haze formation but also exert a substantial influence on climate dynamics across various
28 spatial scales, from local to global (Kaul et al., 2011; Shrivastava et al., 2017). The water solubility,
29 considered one of the crucial physical properties of SOA, has been extensively studied recently due to
30 its significant effects on the physicochemical processes in the atmosphere. The water solubility of SOA
31 varied with its aging degrees (Kirillova et al., 2013), while both the water solubility and aging degree
32 of organic aerosols contribute to the hygroscopicity noticeably, which affects the light extinction
33 eventually (Han et al., 2022; Liu et al., 2022). Hence, exploring the water solubility characteristics of
34 SOA with different aging degrees can help elucidate the more detailed extinction mechanism of SOA.
35 In addition, recent studies have also shown that the formation of secondary particulates is one of the
36 main processes determining the amount of cloud condensation nuclei (CCN) in remote oceanic regions
37 (Liu and Matsui 2022). Therefore, investigating the water solubility of SOA with different aging
38 degrees is also meaningful for further exploring its indirect climate effects.

39 Investigating the contributions of SOA with different aging degrees to both organic matter (OM)
40 and water-soluble organic matter (WSOM) is imperative for determining their quantified water
41 solubility. However, due to the constraints of reliable methods, only a limited number of studies have
42 examined the water solubility of SOA using mass spectrometry techniques. Qiu et al. (2019) conducted
43 source apportionment of OM in PM₁ and WSOM in PM_{2.5} based on online and offline AMS-PMF
44 methods respectively (Qiu et al., 2019). This approach faces challenges not only related to the inherent
45 errors of online versus offline methods but also discrepancies in the measured particle sizes of OM and

46 WSOM. Kondo et al. (2007) and Timonen et al. (2013) attempted to apportion water-soluble organic
47 carbon (WSOC) through a multiple linear regression method based on the mass spectral information of
48 OM, which still exhibits large indeterminateness (Timonen et al., 2013; Xiao et al., 2011; Kondo et al.,
49 2007). The carbon isotopic technique offers a promising avenue to overcome the aforementioned
50 limitations, thereby enabling a more in-depth exploration of the water-soluble characteristics of SOA.
51 Carbon isotope techniques have garnered widespread attention and are increasingly employed in source
52 apportionment studies of organic aerosols due to their robust source appointment capabilities.
53 Radioactive carbon isotopes (^{14}C) provide a precise method for quantitatively distinguishing between
54 fossil and non-fossil organic aerosol sources (Fushimi et al., 2011; Zhang et al., 2014). The stable
55 carbon isotope technique (^{13}C), however, can quantitatively assess the contributions of various sources
56 by integrating them into mass balance models (Yao et al., 2022; Widory et al., 2004). The Bayesian
57 mixing model stands out as one of the most widely utilized models (Xiao;Xu and Xiao 2023; Tang et
58 al., 2020). The stable carbon isotope technique can also be combined with other source tracers to
59 further enhance the accuracy of source apportionment of carbonaceous aerosols (Jiang et al., 2022;
60 Plasencia Sánchez et al., 2023; Ceburnis et al., 2011; Lim et al., 2022). However, to our knowledge, no
61 study has employed the carbon isotope technique to estimate the source contribution of both fresh and
62 aged SOA before, owing to the challenging measurement of the carbon isotope profiles for these two
63 sources.

64 Previous studies have predominantly concentrated on assessing the water solubility of SOA at
65 inland urban sites, revealing a strong correlation between SOA water solubility and urban air pollution
66 emissions as well as relative humidity (Wong;Zhou and Abbatt 2015; Pye et al., 2017; Favez et al.,
67 2008; Salma et al., 2007; Weber et al., 2007; Miyazaki et al., 2006). Nevertheless, few researchers have

68 noticed the differences between inland and coastal cities. As dynamic interfaces between urban and
69 marine environments (Donaldson and George 2012), coastal cities exhibit unique characteristics.
70 Shenzhen is a typical representative city for coastal air pollution studies with a coastline spanning
71 260.5 km and a total sea area of 1145 m². We measured the stable carbon isotope fingerprints of fresh
72 and aged secondary organic carbon (SOC), which enables us to investigate the source contributions of
73 SOC with different aging degrees to WSOC and their respective water solubility in Shenzhen.

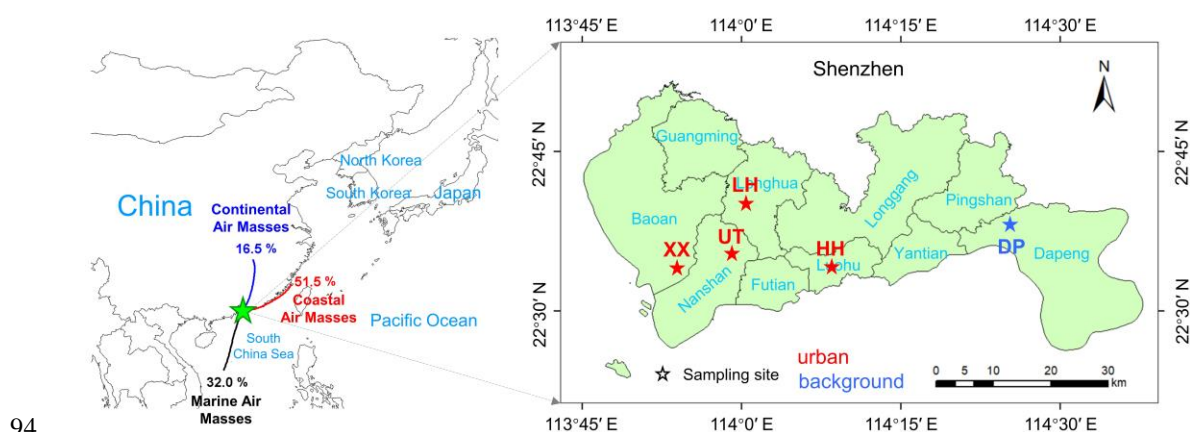
74 The aim of this study is to investigate the water solubility of SOC in PM_{2.5}, emphasizing Shenzhen
75 as a representative mega-coastal city in China. We analyzed stable carbon isotopes and mass spectra
76 signatures of total carbon (TC) and WSOC in ambient PM_{2.5} samples that were collected from five
77 distinct sites in Shenzhen over one year as well as specific emission sources. For the first time, we
78 employed the Bayesian stable isotope mixing (BSIM) model on localized source profiles to quantify
79 the contributions of fresh SOC and aged SOC to WSOC and water-insoluble organic carbon (WIOC).
80 These results would contribute to estimating the water solubility of both fresh and aged SOC, revealing
81 their direct or indirect implications for climate change.

82 **2. Material and methods**

83 **2.1 Ambient PM_{2.5} sampling and chemical analysis**

84 Shenzhen (N22°27' ~ N22°52', E113°46' ~ E114°37'), one megacity of Pearl River Delta, China, is
85 bordered by Daya Bay and Dapeng Bay to the east, the Pearl River Estuary and Lingding Sea to the
86 west, Hong Kong to the south, and Dongguan and Huizhou to the north. As a typical mega-coastal city
87 in China, Shenzhen's air quality is predominantly affected by the continental air mass from northern
88 Guangdong, the eastern coastal air mass, and the southern marine air mass (Fig. 1). For a

89 comprehensive exploration of pollution characteristics in Shenzhen, PM_{2.5} samples were collected from
90 five sites covering the western to eastern regions of the city. The selected sites are Xixiang (XX, urban
91 site), University Town (UT, urban site), Longhua (LH, urban site), Honghu (HH, urban site), and
92 Dapeng (DP, background site) (Fig. 1). Additional details about each sampling site are listed in Table
93 S1.



94
95 **Figure 1.** Spatial distribution of the five sampling sites in Shenzhen for this study.

96 In this study, 24-hour PM_{2.5} sampling was conducted every other day in 2019 at the UT site using
97 a Thermo 2300 atmospheric particulate sampler (Thermo Fisher Scientific Inc., Waltham,
98 Massachusetts, USA), yielding a total of 160 valid samples. For the remaining four sites, a total of 295
99 valid PM_{2.5} samples were collected every other day during typical months of the four seasons in 2019
100 (March, June, September, and December, Table S2) using a Model TH-16A atmospheric particulate
101 sampler (Tianhong Corp., Wuhan, China). The PM_{2.5} samples collected by the quartz filter were used to
102 determine the organic carbon (OC) and elemental carbon (EC) using an OC/EC analyzer (2001A,
103 Desert Research Institute, Reno, Nevada, USA) following the IMPROVE A procedure. In addition, the
104 samples collected by Teflon filter in this study were analyzed for water-soluble ions (mainly SO₄²⁻,
105 NO₃⁻, NH₄⁺, and Cl⁻) within PM_{2.5}, and the mass concentrations of twenty-three metallic elements

106 (primarily Na, Mg, Al, K, Ca, V, Fe, Ni, Zn, Pb, and Cd) within $PM_{2.5}$ were also determined using an
107 inductively coupled plasma mass spectrometer (ICP-MS, Aurora M90; Bruker, Germany). Relevant
108 quality control information is described in the Supplementary Information (Text S1).

109 For WSOC extraction, the $PM_{2.5}$ sample underwent ultrasonication (20 min \times 3 times) in 15 ml
110 ultrapure water (18.2 M Ω ·cm), followed by filtration through a syringe with a 0.45 μ m filter head to
111 eliminate insoluble particles. The extracted $PM_{2.5}$ samples were sequentially analyzed using a long-
112 time-of-flight aerosol mass spectrometer (L-TOF-AMS, Aerodyne, USA) and an ultrasonic nebulizer
113 (U5000AT+, Cetac Technologies Inc., USA) to measure elemental ratios, such as O/C, as well as the
114 mass spectrum signatures of the water-soluble organic fractions, including ion fragments like CO_2^+ ,
115 $C_4H_9^+$, and $C_2H_4O_2^+$. The concentration of WSOC was determined using a total organic carbon analyzer
116 (multi N/C 3100, Jena, Germany), and WIOC was calculated as the difference between OC and WSOC.

117 To investigate the stable carbon isotope signatures of carbonaceous aerosols, we built a stable
118 isotope spectrometry system by integrating an OC/EC analyzer with a carbon dioxide isotope
119 spectrometer (QCLAS, Aerodyne). This system reduces the carbon requirement for isotope analysis
120 from 5 μ gC to 0.5 μ gC and improves the accuracy of spectroscopic measurement methods to
121 0.2‰~0.3‰. The stable carbon isotope values of TC and WSOC in ambient $PM_{2.5}$ were measured in
122 this study.

123 **2.2 Bayesian stable isotope mixing model**

124 The BSIM model could quantify the contributions of multiple sources to the TC and WSOC based on
125 the principle of mass conservation of stable isotopes, in which the Markov Chain Monte Carlo (MCMC)
126 method was employed. The methodology employed in the BSIM model was detailed in works by

127 Parnell et al. (2013) and Parnell and Inger (2010) (Parnell et al., 2010; Parnell et al., 2013). In brief, the
128 posterior distribution for the Bayesian neural network (BNN) was calculated utilizing the prior
129 distribution and likelihood function based on Bayes theorem. Implementation of the BSIM model in
130 this study utilized the SIMMR package in R software ([https://cran.r-project.org/
131 web/packages/simmr/index.html](https://cran.r-project.org/web/packages/simmr/index.html)). Gelman diagnostic values, ranging from 1 to 1.01, all met the
132 criteria of the posterior prediction test, indicating robust model performance and reliable results.
133 Additionally, an uncertainty index (UI₉₀) was employed here to further characterize the uncertainty
134 strength of TC and WSOC source apportionments based on their posterior distribution. This index
135 refers to the difference between the proportional contributions of the maximum and minimum values in
136 the rapid increase segment divided by 90 with a 90 % cumulative probability (UI₉₀ = (PC₉₅-PC₅)/90)
137 (Zaryab et al., 2022; Ji et al., 2017).

138 **2.3 Stable carbon isotope spectrum of PM_{2.5} sources**

139 The BSIM model requires the input of potential sources for carbonaceous aerosols, along with their
140 local source-specific stable carbon isotope values (fingerprints). In this study, we firstly employed the
141 PMF model to identify the potential sources of TC and WSOC (Text S1), with the aim of reducing the
142 uncertainty of the subsequent BSIM model and verifying the reliability of the BSIM results. The PMF
143 results showed that traffic emissions, SOA, and biomass burning are the major contributors to
144 carbonaceous aerosols in Shenzhen, which were similar to the previous results in Guangzhou (Huang et
145 al., 2014). In addition, a literature review indicated that secondary conversion sources could be further
146 subdivided into fresh SOC for the low oxidation state and aged SOC for the high oxidation state (Chen
147 et al., 2019; Presto et al., 2009; Mahrt et al., 2022; Shen et al., 2017). Ultimately, traffic emissions,
148 fresh SOC, aged SOC, and biomass burning (BB) were identified as the four potential sources of TC

149 and WSOC for BSIM model in this study. Since the PMF model lacks the mass spectral information of
150 offline PM_{2.5} samples, it fails to distinguish between fresh SOC and aged SOC in TC, making it
151 challenging to investigate the water solubility characteristics of the SOC based on PMF results. BSIM
152 model simultaneously quantified of fresh and aged SOC separately in both TC and WSOC, thereby
153 enabling an estimation of SOC water solubility. This capability is used for the final analysis in this
154 study.

155 Recognizing the regional variability in stable carbon isotope fingerprints of PM_{2.5} sources (Yao et
156 al., 2022), this work obtained representative and locally specific carbon isotope profiles for the four
157 sources in Shenzhen. The measured profiles of the four sources were used as prior information in the
158 BSIM model for the follow-up analyses. For the traffic emissions, we measured the stable carbon
159 isotope values of TC and WSOC in PM_{2.5} that were collected from the Mount Tanglang tunnel
160 (dominated by diesel vehicles) and the Jiuweiling tunnel (dominated by petrol vehicles) in Shenzhen.
161 Fresh SOC was simulated through petrol vehicle bench tests. The lowest stable carbon isotope values
162 for TC and WSOC from the simulated samples were chosen as the fresh SOC results. The oxygen-
163 carbon ratios (O/C) of fresh SOC samples in this study ranged from 0.51 to 0.62, indicating a low
164 oxidation state (Ding et al., 2012). Aged SOC samples were obtained by collecting ambient PM_{2.5}
165 samples at the National Ambient Air Background Monitoring Station (Mount Wuzhi site, Hainan,
166 China), primarily influenced by regional pollution transported by northern continental air masses.
167 These aged SOC samples exhibited a high O/C value of 0.98, suggesting their highly oxidized state
168 (Zhu et al., 2016). Biomass burning emissions were simulated and analyzed by burning pine wood in
169 the Laboratory of Biomass Burning Simulation at Peking University Shenzhen Graduate School (He et
170 al., 2010). Additional details about the sampling process are available in the Supplementary

171 Information (Text S2). Table 1 summarizes the stable carbon isotope fingerprints of the four sources
 172 and f_{60} signatures used in this study. Table S3 compares $\delta^{13}\text{C}_{\text{TC}}$ source signatures in this study with
 173 global datasets. The stable carbon isotope measurements from the four sources align with the range
 174 observed in global datasets, thus affirming the reliability of the four source fingerprints utilized in this
 175 study. Previous research identified $\text{C}_2\text{H}_4\text{O}_2^+$ (m/z 60) as a reliable marker for biomass burning in
 176 Shenzhen, with a feature value of 1.61 ± 0.68 % (Cao et al., 2018). This prior information was also
 177 incorporated into the BSIM model to estimate the biomass burning source. Although there is some
 178 overlap among the $\delta^{13}\text{C}$ fingerprints of different sources, the Bayesian approach allows for probabilistic
 179 estimation of the contribution of different sources and can also integrate information from multiple
 180 markers and sources to mitigate the effects of overlap. In this study, the PMF model was used to reduce
 181 the uncertainty of interference from unrelated sources, and the chemical tracer marker of biomass
 182 burning source (f_{60}) was also integrated to minimize the effect of this overlap.

183 **Table 1.** Stable carbon isotope fingerprints and f_{60} signatures for TC and WSOC sources.

TC	Traffic		Fresh SOC		Aged SOC		BB	
	$\delta^{13}\text{C}/\text{‰}$	$f_{60}/\%$	$\delta^{13}\text{C}/\text{‰}$	$f_{60}/\%$	$\delta^{13}\text{C}/\text{‰}$	$f_{60}/\%$	$\delta^{13}\text{C}/\text{‰}$	$f_{60}/\%$
	-26.26 ± 0.50	0	-27.31 ± 0.73	0	-25.54 ± 0.28	0	-27.58 ± 0.24	1.61 ± 0.68
WSOC	Traffic		Fresh SOC		Aged SOC		BB	
	$\delta^{13}\text{C}/\text{‰}$	$f_{60}/\%$	$\delta^{13}\text{C}/\text{‰}$	$f_{60}/\%$	$\delta^{13}\text{C}/\text{‰}$	$f_{60}/\%$	$\delta^{13}\text{C}/\text{‰}$	$f_{60}/\%$
	-26.68 ± 0.37	0	-26.18 ± 0.75	0	-24.93 ± 0.39	0	-26.78 ± 0.17	1.61 ± 0.68

184 2.4 Contributions of SOC to WIOC

185 Based on the source apportionment results from the BISM model for TC and WSOC, the contributions
 186 of fresh SOC and aged SOC to WIOC were calculated according to the equations (1-2). The
 187 uncertainties (u) in concentrations of Fresh SOC_(WIOC) and Aged SOC_(WIOC) were assessed using the
 188 uncertainty transfer equations (3-4). Fresh SOC and aged SOC uncertainties in both TC (14.9 %,

189 30.1 %) and WSOC (24.1 %, 20.9 %) were determined using the BSIM model. Our findings reveal that
 190 the calculated uncertainties of [Fresh SOC_(WIOC)] and [Aged SOC_(WIOC)] were 28.3 % and 36.8 %,
 191 respectively.

$$192 \quad [\text{Fresh SOC}_{(WIOC)}] = [\text{Fresh SOC}_{(TC)}] - [\text{Fresh SOC}_{(WSOC)}] \quad (1)$$

$$193 \quad [\text{Aged SOC}_{(WIOC)}] = [\text{Aged SOC}_{(TC)}] - [\text{Aged SOC}_{(WSOC)}] \quad (2)$$

$$194 \quad u_{[\text{Fresh SOC}_{(WIOC)}]} = \left(u_{[\text{Fresh SOC}_{(TC)}]}^2 + u_{[\text{Fresh SOC}_{(WSOC)}]}^2 \right)^{1/2} \quad (3)$$

$$195 \quad u_{[\text{Aged SOC}_{(WIOC)}]} = \left(u_{[\text{Aged SOC}_{(TC)}]}^2 + u_{[\text{Aged SOC}_{(WSOC)}]}^2 \right)^{1/2} \quad (4)$$

196 **3. Results and discussion**

197 **3.1 Overview of PM_{2.5} and carbonaceous components**

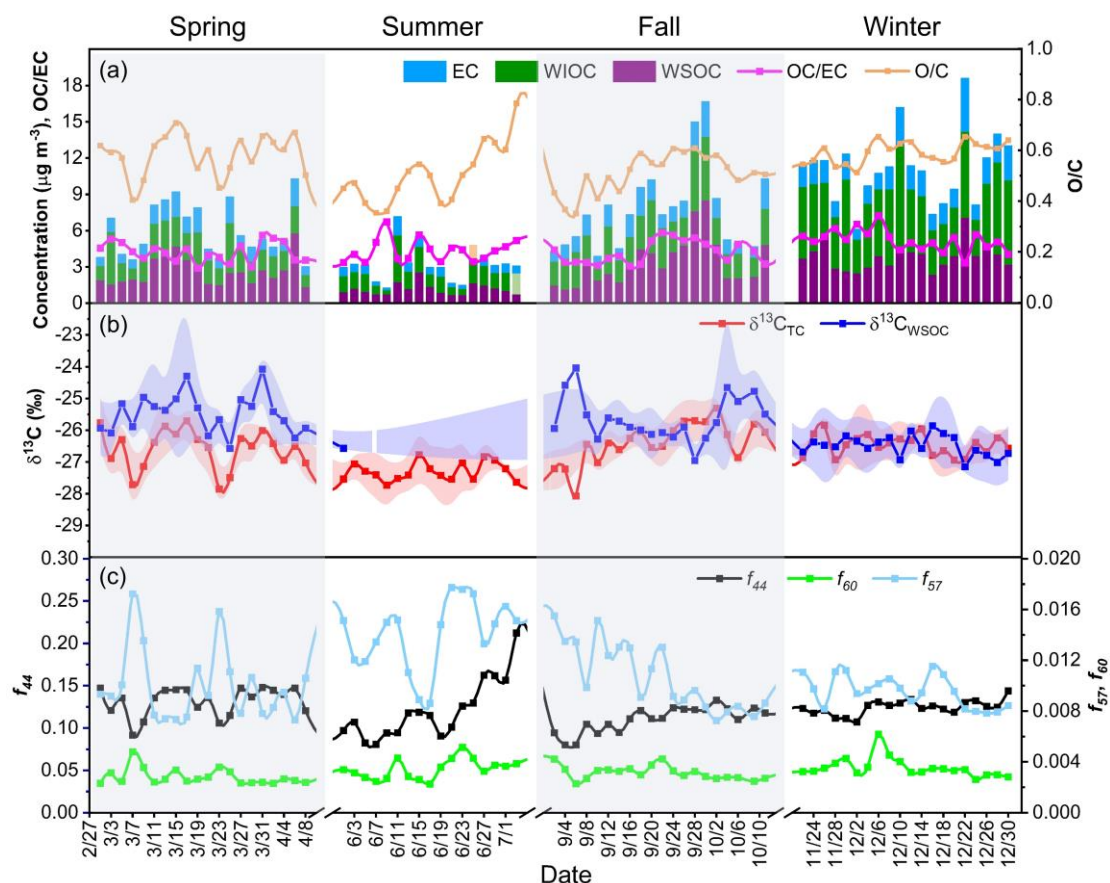
198 The annual mean concentration of PM_{2.5} in Shenzhen was 24.9 μg m⁻³ in 2019, with TC being the
 199 predominant component, exhibiting an annual mean concentration of 7.1 μg m⁻³ (5.8 and 1.3 μg m⁻³ for
 200 OC and EC, respectively). WSOC accounts for 48 % of OC, presenting an annual mean concentration
 201 of 2.8 μg m⁻³. The mean stable carbon isotope values for TC (δ¹³C_{TC}) and WSOC (δ¹³C_{WSOC}) were -
 202 26.64 ± 0.79 ‰ and -25.80 ± 0.88 ‰, respectively, which is lower than the results of northern cities in
 203 China (Wu et al., 2020). This can be attributed to the limited impact of coal combustion (which has
 204 high ¹³C values) on PM_{2.5} in Shenzhen (Yao et al., 2022; Vodicka et al., 2022).

205 Seasonal variation revealed that TC, OC, WSOC, and EC exhibited elevated levels in winter and
 206 decreased levels in summer (Fig. 2a). This pattern primarily stems from pollution air masses
 207 originating from continental regions in the fall and winter, and clean air masses from the southern
 208 ocean during the summer months (Fig. S1). The OC to EC ratio, averaging 4.5, was also higher in

209 winter than in summer, consistent with the Oxygen-to-Carbon (O/C) ratio results for WSOC (Fig. 2a),
210 indicating a large influence of aged SOC on carbonaceous aerosols in winter. The stable carbon isotope
211 results support this observation. Fig. 2b depicts relatively higher $\delta^{13}\text{C}_{\text{TC}}$ and $\delta^{13}\text{C}_{\text{WSOC}}$ values in spring
212 (-26.59‰, -25.26‰), fall (-26.38‰, -25.44‰), and winter (-26.46‰, -26.27‰). These higher values
213 are attributed to greater contributions of aged SOC from northern and northeast regional transport
214 processes during these seasons (Fig. S1). In summer, observed low $\delta^{13}\text{C}_{\text{TC}}$ and $\delta^{13}\text{C}_{\text{WSOC}}$ values of -
215 27.29‰ and -26.57‰, respectively, suggest relatively high contributions of fresh SOC to $\text{PM}_{2.5}$.
216 Shenzhen experiences high temperatures in summer, leading to increased gaseous precursor emissions
217 from terrestrial biogenic sources, especially C3 plants. Intense solar radiation and high temperature
218 favor photochemical reactions to generate fresh SOC that depletes ^{13}C in particulate matter during
219 summer (Kirillova et al., 2013).

220 Mass spectra characteristics of CO_2^+ (m/z 44), C_4H_9^+ (m/z 57), and $\text{C}_2\text{H}_4\text{O}_2^+$ (m/z 60) in WSOC
221 were measured to represent oxidized organic aerosol (OOA), hydrocarbon-like organic Aerosol (HOA),
222 and biomass burning organic aerosol (BBOA), respectively. The abundance of these ion fragments,
223 denoted as f_{44} , f_{57} , and f_{60} , is determined by the ratios of signal intensities at m/z 44, m/z 57, and m/z 60
224 to the sum of signal intensities from all m/z signals in the organic mass spectra. As depicted in Fig. 2c,
225 f_{44} obtained higher values in spring (0.131) and winter (0.125) compared to summer (0.120) and fall
226 (0.112), further indicating an elevated oxidation level of OOA during spring and winter. Considering
227 that f_{60} exceeds 0.0030 when biomass burning influences carbonaceous aerosol (Docherty et al., 2008;
228 DeCarlo et al., 2008), the annual average value of f_{60} was 0.0032, suggesting biomass burning was an
229 important source of carbon components in Shenzhen. Winter exhibited higher levels of f_{60} (0.0035)
230 compared to other seasons, suggesting relatively strong impacts of biomass burning on WSOC in

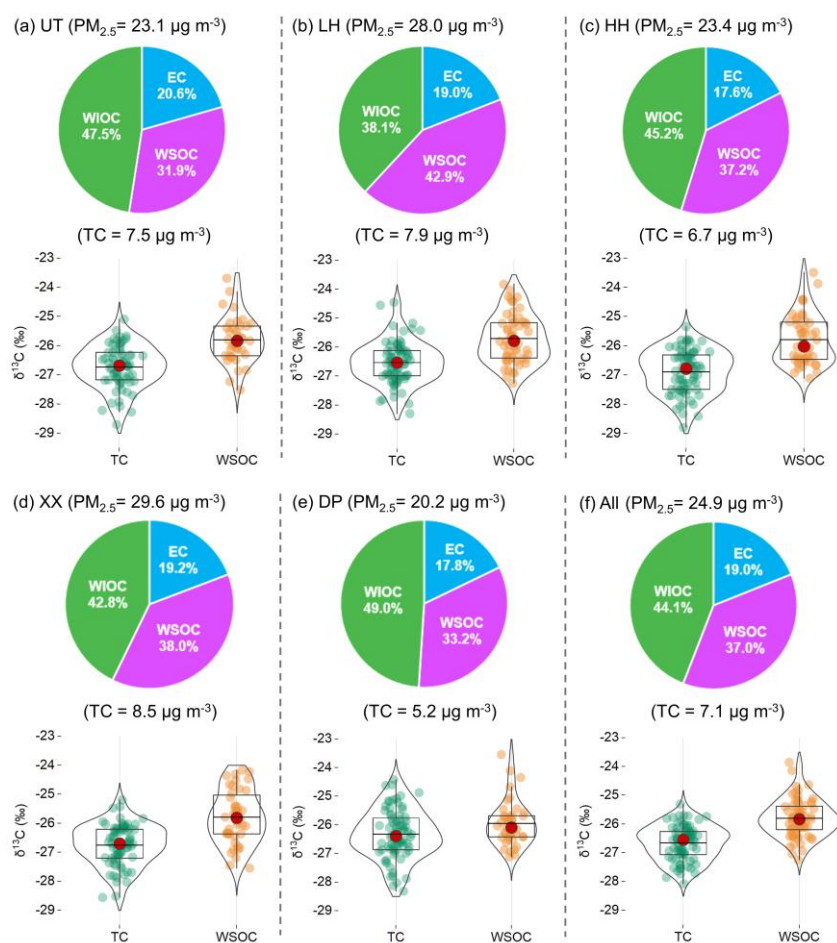
231 winter. Conversely, f_{57} reached its highest level in summer (0.014) and the lowest in winter (0.009),
 232 with an annual average value of 0.011, possibly associated with a notable increase in hydrocarbon
 233 organic aerosol emissions from traffic and biogenic sources during the summer period.



234
 235 **Figure 2.** Time series of carbonaceous components (a), stable carbon isotope characteristics of TC and
 236 WSOC(b), and mass spectra signatures of WSOC in $\text{PM}_{2.5}$ (c) from Shenzhen. Each data was averaged
 237 from five sampling sites. (Note: Summer samples exhibit elevated analytical errors due to low
 238 concentrations, and $\delta^{13}\text{C}_{\text{WSOC}}$ values are computed from combined summer samples).

239 Obvious spatial variations in $\text{PM}_{2.5}$ mass concentrations across Shenzhen during 2019 were
 240 observed, with XX site registering the highest concentration ($29.6 \mu\text{g m}^{-3}$), followed by LH ($28.0 \mu\text{g m}^{-3}$),
 241 HH ($23.4 \mu\text{g m}^{-3}$), UT ($23.1 \mu\text{g m}^{-3}$), and DP ($20.2 \mu\text{g m}^{-3}$). Figure 3 illustrates that TC made more
 242 substantial contributions (28.2 % ~ 32.5 %) to $\text{PM}_{2.5}$ at the four urban sites in the central and western

243 regions of Shenzhen compared to the background site (DP, 25.7 %). This suggests that local pollutant
244 emissions significantly influence carbonaceous aerosols in Shenzhen's urban areas. The percentage of
245 WSOC in TC was also higher in urban areas (37.5 ± 3.9 %) compared to the background area (DP,
246 33.2 %), reaching the highest value at the LH site (42.9 %). However, the percentage of WIOC in TC
247 displayed the opposite trend, suggesting carbonaceous aerosols in urban areas of Shenzhen exhibit
248 higher water solubility than in background areas. Distinct spatial distribution characteristics were also
249 observed in the stable carbon isotopes of TC and WSOC. The background site exhibits higher $\delta^{13}\text{C}_{\text{TC}}$
250 values (-26.33 ‰) than the four urban sites (-26.72 ± 0.13 ‰). This difference may be attributed to the
251 increased contribution of traffic or fresh SOC sources to carbonaceous aerosols at urban sites and the
252 relatively high contribution of aged SOC at the background site. Atmospheric aging processes of
253 organics through photochemical reactions can deplete ^{13}C in aged SOC and enrich ^{13}C in fresh SOC
254 and other related reactants simultaneously (Pavuluri and Kawamura 2017). While the close proximity
255 of the $\delta^{13}\text{C}_{\text{WSOC}}$ values at urban sites (-25.77 ± 0.04 ‰) to the background site (DP, -25.96 ‰) suggests
256 that the WSOC in different areas of Shenzhen may share a similar origin.



257

258 **Figure 3.** Chemical compositions of TC, $\delta^{13}\text{C}_{\text{TC}}$, and $\delta^{13}\text{C}_{\text{WSOC}}$ in $\text{PM}_{2.5}$ at urban sites (a-d), background
 259 site (e), and average result from all five sites (f). The Violin Box-and-Line Plots on the right display
 260 spatial variations of $\delta^{13}\text{C}_{\text{TC}}$ and $\delta^{13}\text{C}_{\text{WSOC}}$ at each site, featuring mean values (black lines) and median
 261 values (red dots).

262 3.2 Source apportionment results for TC and WSOC

263 The BSIM model assessed the contributions of traffic source, fresh SOC, aged SOC, and biomass
 264 burning (BB) to TC and WSOC, as shown in Fig. 4. On average, SOC (total of fresh and aged SOC)
 265 and traffic emerged as the two major contributors to TC, accounting for 43 % and 40 % respectively,
 266 while biomass burning contributed 17 % to TC. The contribution of aged SOC to TC (23 %) is
 267 comparable with fresh SOC (20 %). Regarding WSOC, SOC was the dominant source, comprising 45 %
 268 of aged SOC and 28 % of fresh SOC, followed by BB (18 %) and Traffic (9 %). The noteworthy

269 contribution of aged SOC to WSOC suggests a comparatively higher water solubility of aged SOC in
270 Shenzhen.

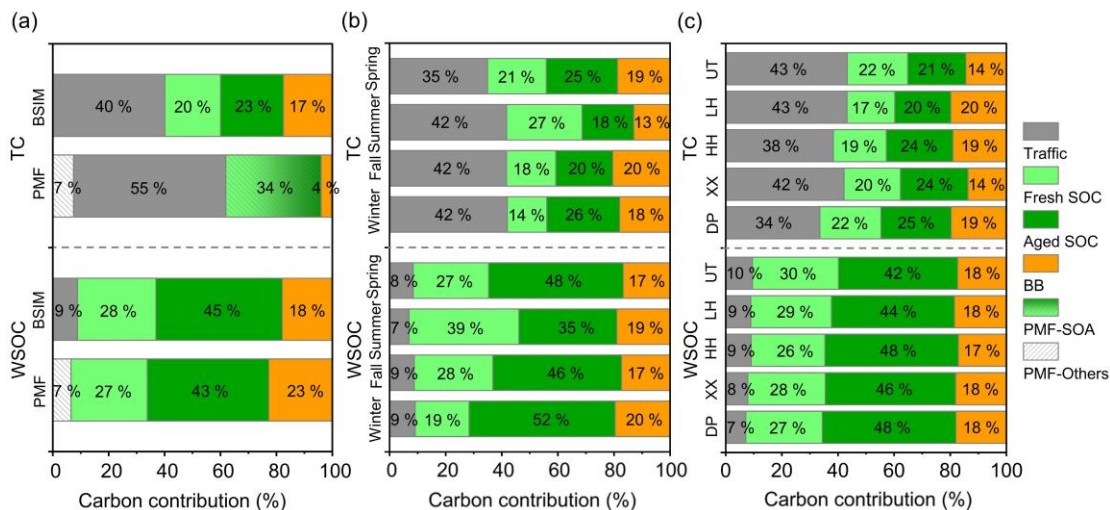
271 To evaluate the BSIM model's performance, we employed the PMF model to apportion the
272 sources of TC and WSOC. The obtained results were subsequently compared with those from the
273 BSIM model, as depicted in Fig. 4a. Seventeen chemical species of PM_{2.5} were applied as the PMF
274 model input to estimate source contributions to TC, encompassing carbon components, soluble
275 inorganic ions, and elements. For the apportionment of WSOC sources, five species including WSOC,
276 WIOC, and three organic mass spectra were applied as the PMF model input. More details about the
277 PMF model and results can be found in the Supplementary Information (Text S1, Fig. S2-S5). PMF
278 identified the traffic as the predominant contributor to TC (55 %), followed by SOC (34 %) and
279 biomass burning (4 %). Concerning WSOC, aged SOC and fresh SOC were the two major sources as
280 well, accounting for 43 % and 27 %, respectively. The traffic contribution to TC apportioned by the
281 PMF model is higher than that of the BSIM model (55 % vs. 40 %), which may be due to the fact that
282 some of the fresh SOC generated by the conversion of primary vehicle emissions was improperly
283 apportioned to the traffic source in the PMF model (Li et al., 2022; Zhao et al., 2014). Previous study
284 also showed that SOA contributes more to carbonaceous aerosols in Shenzhen than the traffic source
285 (Cao et al., 2022). The PMF model results for WSOC were generally consistent with BSIM model
286 results, with deviations primarily attributed to the differences in the principles and uncertainties of the
287 two models.

288 Furthermore, this study examined cumulative frequency distributions to elucidate the inherent
289 uncertainty in source apportionments of TC and WSOC. As shown in Fig. 5a and b, the proportional

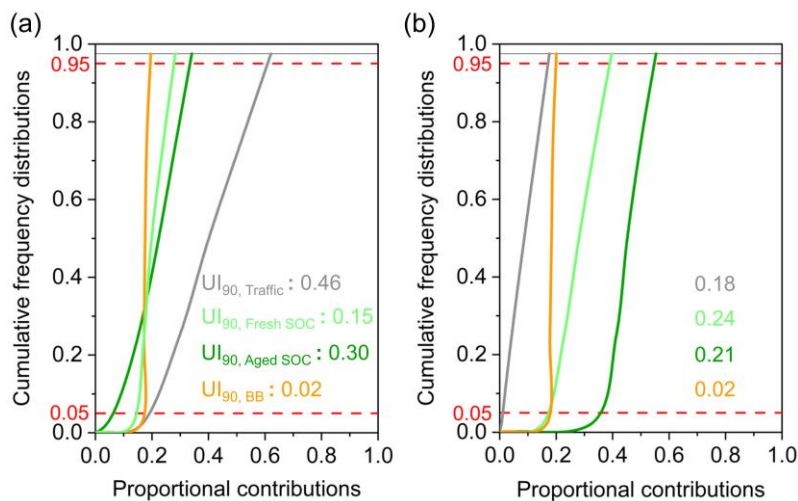
290 contributions of BB source to both TC and WSOC were quite stable during the research periods due to
291 its low UI_{90} value (0.02). This may be attributed to the incorporation of mass spectral constraints for
292 the BB source in the BSIM model used in this study. For TC source apportionment results, the largest
293 UI_{90} value (0.46) was observed for the traffic source, indicating that its contribution to TC exhibited
294 relatively high uncertainty. In 90 % probability, its contribution ranged from 19.4 % to 60.9 %. The
295 UI_{90} values for fresh and aged SOC were 0.15 and 0.30, respectively. Regarding WSOC, the calculated
296 UI_{90} value of traffic, fresh SOC, and aged SOC ranged from 0.18 to 0.24. The UI_{90} values obtained
297 through the BSIM model remained within reasonable limits, and were smaller than those calculated in
298 previous related studies (0.23-0.62) (Zaryab et al., 2022; Ji et al., 2017). Consequently, the source
299 contributions of TC and WSOC estimated by the BSIM model in this study were deemed reasonable.

300 For seasonal variations, as shown in Fig. 4b, SOC still was the major source of TC and WSOC
301 during all four seasons, ranging from 38 % ~ 46 % and 71 % ~ 75 % respectively. Significant high
302 contributions of fresh SOC to TC and WSOC occurred in summer (27 %, 39 %), and relatively higher
303 contributions of aged SOC to TC and WSOC were observed in winter (26 %, 52 %). It is because
304 meteorological conditions in winter characterized by inversions and stagnant winds facilitate the
305 accumulation of air pollutants, and Shenzhen is largely influenced by regional pollution transport in
306 winter, favoring the formation of aged SOC (Huang et al., 2018). In contrast, favorable meteorological
307 conditions (e.g. intense and prolonged solar radiation, high temperatures, and relative humidity) in
308 summer enhanced photochemical reactions to generate fresh SOC. In terms of spatial distributions (Fig.
309 4c), the contributions of the traffic source to TC were higher at urban sites (38 % to 43 %) compared to
310 the background site (34 %). This finding aligns with expectations due to increased human activity and
311 vehicle numbers in urban locations. At the DP site, the contributions of SOC to TC were higher than

312 those of other sources (47 %), signifying a predominant influence of regionally transported pollutant
 313 emissions on TC at the background site. However, the contributions of SOC and the other two primary
 314 sources at both urban and background sites were all close to each other, indicating the source
 315 composition of WSOC in Shenzhen is less affected by air pollution degree compared to TC.



316
 317 **Figure 4.** (a) Comparison of source apportionment results between BSIM model and PMF model for
 318 TC and WSOC, (b) seasonal and (c) spatial distributions of source apportionment results for TC and
 319 WSOC based on the BSIM model.



320
 321 **Figure 5.** Cumulative frequency distributions of the proportional contributions from potential sources
 322 of TC (a) and WSOC (b) based on BSIM model.

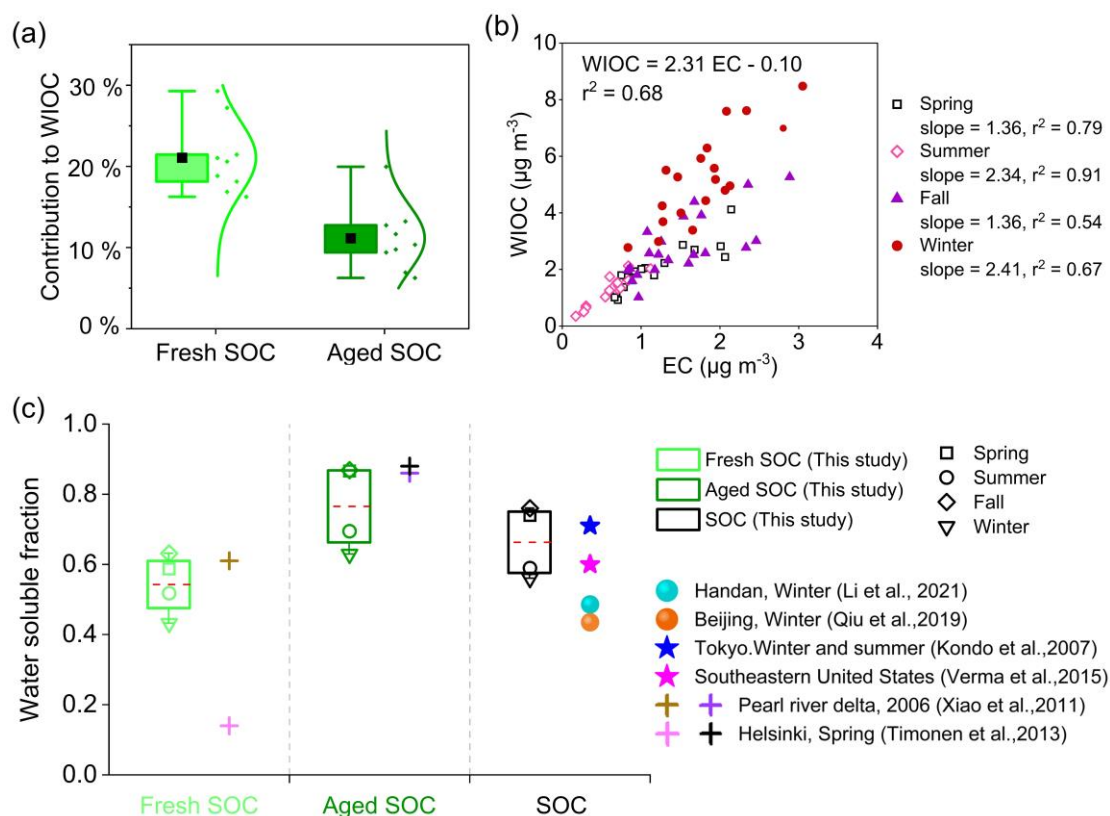
323 3.3 Water solubility of fresh SOC and aged SOC

324 The contributions of fresh SOC and aged SOC to WIOC were the differences between the contributions
325 of those two SOC sources to TC and WSOC from the BSIM model (Sect. 2.4) in this study. As shown
326 in Fig. 6a, fresh SOC and aged SOC made contributions of $23.2\pm 4.2\%$ and $13.4\pm 3.8\%$ to WIOC,
327 respectively, implying that primary sources are the dominant contributors to WIOC. Further support for
328 this finding is evident in the strong correlation between WIOC and EC, as depicted in Fig. 6b. A higher
329 slope was observed in winter (2.4) than in other seasons, consistent with the highest contributions of
330 aged SOC to WIOC in winter (22 %). This observation implies that WIOC in winter is influenced not
331 only by local primary sources but also by the promotion of secondary pollution.

332 To investigate deeply the water solubility characteristics of fresh and aged SOC, we then calculate
333 their water-soluble fraction by comparing their water-soluble portion to the ambient fraction ($[c]_{\text{water-}}$
334 $\text{soluble}/([c]_{\text{water-soluble}} + [c]_{\text{water-insoluble}})$), where $[c]_{\text{water-soluble}}$ and $[c]_{\text{water-insoluble}}$ are the concentrations of fresh
335 SOC or aged SOC in WSOC and WIOC, respectively) (Li et al., 2021). As shown in Fig. 6c, the overall
336 water-soluble fraction of SOC in this study was 66.2 % with a range from 58.9 % to 76.0 %. Fresh
337 SOC exhibited a much lower water-solubility of 54.2 %, whereas aged SOC displayed a comparatively
338 higher water-solubility of 76.5 %. The higher water solubility of aged SOC compared to fresh SOC
339 might be due to the positive correlation between aerosol hygroscopicity and oxidation in the sub-
340 saturated state. The water-soluble fraction of SOC in this study was close to that reported in other
341 coastal cities (Tokyo (71 %) and Southeastern United States (60 %)) (Kondo et al., 2007; Verma et al.,
342 2015), while was much higher than that reported in northern Chinese cities (Beijing (42 % ~ 45 %) and
343 Handan (49 %)) (Li et al., 2021; Qiu et al., 2019). In addition, the water-soluble fraction of both fresh
344 SOC and aged SOC, as calculated in this study, was comparable to that reported in Guangzhou (61 %

345 and 86 % for fresh and aged SOC respectively) (Xiao et al., 2011). This could be attributed to
346 Shenzhen's coastal location, which is markedly influenced by regional transport from neighboring
347 urban areas and the eastern seaboard air masses. The high relative humidity facilitates the conversion of
348 aged SOC into WSOC during the pollution transport process. This result is in accordance with previous
349 findings that air masses influenced by anthropogenic emissions could promote the formation of high
350 water-soluble SOA under high relative humidity in urban environments (Miyazaki et al., 2006; Salma
351 et al., 2007; Weber et al., 2007). Given that the aging process of SOA dissolved in water could enhance
352 the cloud condensation nuclei (CCN) activity of the particles (Liu and Matsui 2022), high water-
353 soluble aged SOC in Shenzhen might have significant impacts on the activity of CCN, potentially
354 resulting in more important indirect climate effects.

355 The water-soluble fraction of SOC (especially aged SOC) in Shenzhen exhibits obvious seasonal
356 characteristics, with the highest in fall (76.0 %) and the lowest in winter (56.0 %). This phenomenon is
357 primarily related to the robust atmospheric oxidizing capacity during fall in Shenzhen since the
358 atmospheric oxidants such as OH and NO₃ radicals play pivotal roles in driving the secondary
359 generation of WSOC (Wang et al., 2023). Conversely, during winter, the temperature and relative
360 humidity are at their lowest levels, and the relatively diminished atmospheric oxidizing capacity also
361 constrains the secondary generation of WSOC.



362

363 **Figure 6.** (a) Left is the box and whisker plots of fresh and aged SOC contributions to WIOC, the
 364 upper and lower of the box representing the 75th and 25th percentiles, and the black squares featuring
 365 mean values. The dots on the right show the contribution of fresh and aged SOC to WIOC across
 366 seasons and sites, the curve demonstrates its normal distribution. (b) Scatterplot of WIOC versus EC by
 367 season, (c) Comparison of the water-soluble fraction of SOC (fresh SOC, aged SOC, SOC) in this
 368 study (box and whisker plots) with those in other related literature (colored markings on the right). The
 369 upper and lower of the box represent the 75th and 25th percentiles and the dashed red lines indicate
 370 mean values.

371 4. Summary and implications

372 Assessing the impacts of different oxidational SOC on air quality and its water solubility has been
 373 challenging, and this work successfully evaluated the water-soluble fraction of fresh and aged SOC
 374 employing the BSIM model on one-year observational data for stable carbon isotopes and mass spectra
 375 of TC and WSOC in Shenzhen, China. Compared with other methods, e.g. PMF model, EC tracer, and

376 multiple linear regression analyses, the BSIM model successfully calculated the contributions of fresh
377 SOC and aged SOC to WSOC and WIOC, owing to prior and localized information about stable carbon
378 isotopes and mass spectra of PM_{2.5} sources. Therefore, establishing localized carbonaceous aerosol
379 source profiles for stable carbon isotopes becomes crucial for comprehending the relationship between
380 the aging degree and water solubility of SOC.

381 The observed average mass concentration of PM_{2.5} during the sampling period in Shenzhen was
382 24.9 μg m⁻³, and WSOC accounts for 48 % of OC. The mean stable carbon isotope values for TC
383 ($\delta^{13}\text{C}_{\text{TC}}$) and WSOC ($\delta^{13}\text{C}_{\text{WSOC}}$) were -26.64 ± 0.79 ‰ and -25.80 ± 0.88 ‰, respectively. WSOC was
384 dominated by secondary sources while WIOC was dominated by primary sources. The contribution of
385 fresh SOC and aged SOC to WSOC, WIOC were 28.1 % and 45.2 %, 23.2 % and 13.4 %, respectively.
386 The overall water-soluble fraction of SOC in this study was 66.2 %, with aged SOC constituting 76.5 %
387 and fresh SOC 54.2 %. The water-soluble fraction of aged SOC was 22 % higher than fresh SOC, even
388 though both of them demonstrated remarkable water-soluble characteristics in Shenzhen. This finding
389 highlights the important role of aged SOC in the water uptake process of particulate matter.
390 Considering the strong correlation between the water solubility of SOC and its light extinction effect,
391 further exploration of the extinction effect of SOC with different aging degrees will greatly contribute
392 to a more profound understanding of the extinction mechanism of SOC. Besides, the water solubility of
393 SOC in coastal cities was observed to be higher than that in inland cities, suggesting a more
394 pronounced climate effect of SOC in coastal cities. Therefore, there should be increased emphasis on
395 enhancing the control of SOA precursors in coastal urban areas to better integrate air pollution and
396 climate change management. This is particularly crucial given the observed rise in the proportion of
397 SOA in particulate matter in recent years. Moreover, the results of our study further hinted that the

398 notable water solubility of SOC, particularly aged SOC, may contribute a lot to the formation of CCN
399 above coastal cities, which is also helpful to a better understanding of the cloud microphysical
400 processes and the indirect climate effect of SOC in coastal urban regions.

401 **Data availability.** Datasets are available by contacting the corresponding author, Xing Peng
402 (pengxing@pku.edu.cn)

403

404 **Author contributions.** PX and HX conceptualized the study. WF, CL, TM and FN retrieved and
405 constructed the dataset. WF and PX carried out the statistical analysis. WF prepared the first draft of
406 the manuscript, which was commented on and revised by PX, HL, and HX. All authors reviewed and
407 approved the final version for publication.

408

409 **Competing interests.** The authors declare that they have no conflict of interest.

410

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