

The Atmospheric Oxidizing Capacity in China: Part 2. Sensitivity to emissions of primary pollutants

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Abstract

45 The Atmospheric Oxidation Capacity (AOC), often referred to as the self-cleansing ability of
the atmosphere, considerably affects the concentrations of photochemical air pollutants.
Despite substantial reductions in anthropogenic emissions of key chemical compounds in
China, the mechanisms that determine the changes in the atmospheric oxidation capacity are
50 still not sufficiently understood. Here, a regional chemical transport model is employed to
quantify the sensitivity of air pollutants and photochemical parameters to specified emission
reductions in China for conditions of January and July 2018 as representative. The model
simulations show that, in winter, a 50% decrease in nitrogen oxides (NO_x) emissions leads to
an 8-10 ppbv (15-20%) increase in surface ozone concentrations across China. In summer, the
ozone concentration decreases by 2-8 ppbv (3-12%) in NO_x-limited areas, while ozone
55 increases by up to 12 ppbv (15%) in volatile organic compounds (VOCs)-limited areas. ~~This
ozone increase is associated with a reduced NO_x-titration effect and higher levels of hydroxyl
(OH) and hydroperoxyl (HO₂) radicals due to reduced loss reacted with nitrogen dioxide (NO₂)
and decreased aerosol uptake. This ozone increase is associated with a reduced NO_x titration
effect and higher levels of hydroperoxyl (HO₂) radical due to decreased aerosol uptake.~~ With
60 an additional 50% reduction in anthropogenic VOCs emission, the predicted ozone
concentration decreases by 5-12 ppbv (6-15%) in the entire geographic area of China, with an
exception in the areas, where the role of BVOCs is crucial to ozone formation. Further, the
adopted reduction in NO_x emission leads to an increase of AOC by 18% in VOC-limited areas.
This specific increase is associated with the combined effect of enhanced radical cycles
65 associated with the photolysis of oxidized VOCs (OVOCs) and the oxidation of alkenes by
hydroxyl (OH) radical and O₃. ~~A large reduction of daytime AOC in summer results from the
reduction in anthropogenic VOCs emission, with a dominant contribution from the reaction of
OH radical with reduced alkenes, followed by the reactions with depleted aromatics and
OVOCs.~~ This study highlights that photolysis of OVOCs and oxidation of alkenes and aromatic
70 in urban areas when NO_x emission is reduced leads to an increase in O₃. To mitigate ozone
risers in urban areas, a joint reduction in the emission of NO_x and specific VOCs species,
including alkenes, aromatics and aromatics and unsaturated photodegradable OVOCs,
including methanol and ~~should~~ethanol, should be implemented.

75 Keywords: ozone pollution, emission reduction, nitrogen chemistry, AOC

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1. Introduction

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To effectively reduce air pollution in China, the government of the country has implemented stringent actions between 2013 and 2020 (Liu et al., 2020; Liu et al., 2023). In the initial phase, from 2013 to 2017, the control of primary pollutants was particularly effective, with a dramatic decrease in the anthropogenic emissions of fine particles (PM_{2.5}), sulfur dioxide (SO₂), and nitrogen oxides (NO_x) (Liu et al., 2020). After that, a sustained reduction in the emission of SO₂, NO_x, and PM_{2.5} was achieved with continuous emission control from 2018 to 2020 (Liu et al., 2023). The implementation of the emission control policies has greatly improved China's air quality. However, a significant increase in the surface ozone (O₃) concentration was observed from 2013 to 2019, with the rising trend slowing down from 2020 to 2021, but rebounding in 2022 (Liu et al., 2023; China Air 2023). Some studies have documented the explanations for the significantly increasing trend in the surface O₃ concentration, including the reduction of NO_x emissions and atmospheric aerosol loading (Li et al., 2019a; Liu et al., 2020). During and after the recent COVID-19 lockdown, ozone pollution has also been reported to happen, which is believed to be favored by the sharp reduction of NO_x and high emissions of VOCs (Li et al., 2021). Looking through these changes over the past decade, we can learn that rapid reductions of emissions may lead to varied ozone chemistry and, thereby, complex changes in ozone concentrations in China.

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The response of ozone to reduced NO_x emissions varies with the local photochemical environment and is different in VOC-limited, NO_x-limited, or transition regimes (Ou et al., 2016; Dai et al., 2023). In VOC-sensitive regimes, the reduction of NO_x tends to increase ozone formation due to the weakening of NO nitration and the competition between NO₂ and VOC for OH radicals (Ou et al., 2016). In NO_x-sensitive regimes, NO_x emission reduction decreases the photolysis of NO₂, leading to less ozone formation (Ou et al., 2016). Several studies using satellite observations (Wang et al., 2021) and regional models (Zhang S. et al., 2023) have shown that the reduction in anthropogenic emissions has generated a change in the geographical distribution of the ozone formation regimes in China. The shift of ozone sensitivity regimes from VOC-sensitive to transition and/or NO_x-sensitive in many metropolitan and suburban regions of East China was also reported by these studies. This shift enables efficient ozone control in NO_x-sensitive areas in response to the continuous decrease in NO_x emissions. In VOC-sensitive and transition areas, NO_x emission reduction fails to effectively mitigate ozone pollution, while a coordinated reduction in anthropogenic VOCs (AVOCs) emissions should effectively limit the ozone formation and should therefore be implemented (Liu et al., 2023; Zhang S. et al., 2023). The source of NO_x in VOC-sensitive areas is mainly from fossil fuel combustion, while AVOCs emissions have a range of sources.

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130 To establish a cost-effective control over AVOCs emission, the assessment of the contribution
of different VOCs species to ozone formation should be accurately estimated for different areas
of China.

135 ~~Aerosol decreases associated with the reduction of primary emissions are expected to
continuously affect the effectiveness of ozone control (Liu et al., 2023). Following the
successful controls on anthropogenic emissions since 2013, a substantial reduction in PM_{2.5}
concentration was observed in China (Zhai et al., 2019). The reduction in the emission of NO_x,
SO₂, and ammonia (NH₃), as the gaseous precursors of secondary inorganic aerosol (SIA, the
140 sum of sulfate (SO₄²⁻), nitrate (NO₃⁻), and ammonium (NH₄⁺)), leads to a decrease in the SIA
concentration (Meng et al., 2022). The decrease in AVOCs emissions will lead to a decrease
in the concentration of secondary organic aerosol (SOA), given the gas-phase photochemical
oxidation of VOCs plays an important role in SOA formation (Yuan et al., 2013; Li et al.,
2022). However, the increasing trends of OH radicals will positively influence SOA formation
145 (Wang W. et al., 2022; Wang W. et al., 2023) and provide an increasing portion of the
secondary formation in aerosol compositions (An et al., 2019). The interaction of aerosol and
O₃ formation has been discussed in many modeling studies (Li et al., 2019; Liu et al., 2020).
However, the influences of aerosol on O₃ production can be varied due to counteracted aerosol
effects and different aerosol concentrations (Tan et al., 2022; Dai et al., 2023). Understanding
150 the changes in aerosol effects on ozone formation when the primary emissions are further
reduced is still necessary to implement an efficient air quality control policy.~~

155 Recent observational studies combined with a source apportionment approach using
observation-based models have highlighted the role of specific VOCs species, including the
alkenes, aromatics, and several OVOCs, in mitigating summertime ozone formation in the
urban areas of China (Shi et al., 2023; Wang W. et al., 2022). The notable contributions of
OVOCs to AOC as well as the formation of SOA in China have been of concern in many studies
(Li et al., 2022; Wang et al., 2023). Since the oxidation of biogenic VOCs (BVOCs) can
160 significantly contribute to the formation of secondary pollutions, the important role of BVOCs
in AOC and the formation of SOA has also been highlighted in vegetated and highly greening
regions in China (Cao et al., 2022; Zhang et al., 2023). However, a comprehensive evaluation
of the influence of different VOCs species on AOC and ozone chemistry in different regions of
China is still needed. Considering the necessity of implementing coordinated actions among
165 large areas to further alleviate air pollution in China, regional chemical transport models are
appropriate tools to assess the quantitative response of various VOCs species and AOC-related
chemical parameters to emission changes.

170 In the companion paper (Part 1; Dai et al., 2023), we used a regional chemical-meteorological
model to quantify the relative contribution of different photochemical processes to the
formation and destruction of near-surface RO_x and O₃ in different chemical environments in
China. In Part 2 of the study, with the evaluated model, we assess the response of the photo-
oxidative species and related parameters to the reduction of primary emissions. This paper is

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175 structured as follows. Section 2 introduces the setups of the model system and describes the
simulations performed for specified reductions in the emissions of primary pollutants. In
180 Section 3, we ~~first analyze the response in the near-surface concentration of ozone precursors
and intermediate to primary emission reductions. We discuss also discuss~~ the changes in the
ozone formation regime, ~~resulting from emission reductions. Further, we derive the associated
changes in ozone, and aerosols to emission reductions. We also derive the associated changes
in the radical and ozone budgets. Further, we analyzed the response in the near surface
concentration of ozone precursors, ozone, and aerosols to emission reductions.~~ Finally, we
describe the sensitivity of ~~the photochemical parameters and~~ the atmospheric oxidative
capacity (AOC) to the reduction in emissions. A summary ~~and implication for policy making~~
185 of our study is provided in Sec. 4.

2. Method

2.1. Model setting

190 We use the WRF-Chem model version 4.1.2 (Skamarock et al., 2019), coupled with the gas-
phase chemistry mechanism MOZART and the aerosol module MOSAIC, to simulate the
meteorological fields as well as the transport, the chemical and physical transformations of
trace gases and aerosols. The months of January and July of 2018 were selected as
195 representative months to conduct the simulations and investigate the changes in secondary
pollution and AOC in response to emission reductions during winter and summer, respectively.
Compared to the standard version of the chemical mechanism, several updates of
heterogeneous uptake over the ambient aerosol were made (Dai et al., 2023). As for SOA
formation in the selected chemical mechanism, the main pathways result from the gas-phase
200 oxidation of VOCs by atmospheric oxidants (OH, O₃, and NO₃) and the heterogeneous
formation of glyoxal SOA over the ambient aerosol. The model domain covers the whole
geographical area of China. Analyses of modeling results at ~~four urban eight sites, including
four urban sites~~ (Beijing, Shanghai, Guangzhou, and Chengdu), ~~two rural sites (Wangdu and
Heshan), and two remote sites (Waliguan and Hok Tsui),~~ were also performed in this study.
205 More detailed information on the model configuration, the model validation, and the sites
selected for our analysis can be found in Part 1 of our paper by Dai et al. (2023).

We used the Multi-resolution Emission Inventory (MEIC v1.3; <http://www.meicmodel.org/>)
to represent anthropogenic emissions in China and the CAMS-GLOB-ANT v4.2 inventory
(<https://eccad.aeris-data.fr/>) provided by the Copernicus Atmosphere Monitoring Service
(CAMS) for anthropogenic emissions in the Asian areas outside China. To explore the
210 sensitivity of secondary pollution and AOC to emission reduction, sensitivity experiments were
designed to separately assess the influence of reduced NO_x, AVOCs, and other emissions. As
shown in Table S1, NO_x emissions include the emissions of NO₂ and NO, and AVOCs
emissions include these of ~~alkenes (C₂H₄, C₃H₆, and BIGENE (alkenes with carbon number ≥
4)), alkanes [(ethane (C₂H₆), propane (C₃H₈), and BIGALK (alkanes with carbon number ≥
4)), alkenes [ethene (C₂H₄), propene (C₃H₆), and BIGENE (alkenes with carbon number ≥ 4)],~~
215 ~~aromatics [(benzene (C₆H₆), toluene (C₆H₅CH₃), and xylene (C₆H₄(CH₃)₂)), alkyne (C₂H₂),~~

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isoprene (C₅H₈), terpenes (C₁₀H₁₆), and OVOCs (~~CH₃CHO~~, methanol (CH₃OH), ethanol (C₂H₅OH), acetaldehyde (CH₃CHO), ~~CH₃OH~~, C₁₀H₁₆, acetone (CH₃COCH₃), methyl acrylate (CH₂CCH₃CHO (MACR), and methyl vinyl ketone (CH₂CHCOCH₃ (MVK)). For other emissions, the emissions of NH₃, SO₂, and carbon monoxide (CO) were considered.

2.2. Design of numerical experiment

To explore the sensitivity of secondary pollutants to emissions changes, five numerical experiments are conducted in both January and July of 2018 (Table 1). In the baseline case, denoted as “BASE”, we adopted emissions as described in Sect. 2.1. The concentrations of the key species modeled in this case have been validated in our companion study (Dai et al., 2023). To quantify the sensitivity of pollutants to the reduction of NO_x and AVOCs emissions, we applied arbitrary reductions in the surface emissions of primary pollutants; In the first two cases, a 50% reduction is applied separately to the baseline NO_x and AVOCs emissions in whole geographical areas of China. These two sensitivity cases are labeled “NO_x” and “AVOCs”, respectively. A third case in which 50% reduction is applied to both NO_x and AVOCs emissions is referred to as “N+A”. ~~The simulation labeled “TOTAL” assumes that all anthropogenic emissions (NO_x, AVOCs, and other primary pollutants including CO, SO₂, and NH₃) are reduced by 50%.~~ The difference between modeled concentrations of pollutants and chemical parameters in the sensitivity cases and the baseline case provides an estimate of the response of secondary pollution and chemistry to different emission reductions.

~~Additionally, a~~ The simulation labeled “TOTAL” assumes that all anthropogenic emissions under consideration (NO_x, AVOCs, and other primary pollutants including CO, SO₂, and NH₃) are reduced by 50%. This case is used to explore the impact of reduction in the emission of CO, as an ozone precursor, and of SO₂ and NH₃, as aerosol precursors, on ozone formation. The spatial distribution of the emission fluxes changes for the different cases is shown in Fig. S1.

3. Model results

3.1. Changes in ozone regimes

~~In order to display the impact of emission reduction on the changes in the ozone regimes, we first show the geographical distribution of NO_x-limited or VOC-limited ozone formation regimes. We adopt the ratio between the production rate of hydrogen peroxide (H₂O₂) and that of nitric acid (HNO₃) [$P(\text{H}_2\text{O}_2)/P(\text{HNO}_3)$] as the indicator to distinguish these regimes. An area is assumed to be VOC-limited or NO_x-limited if $P(\text{H}_2\text{O}_2)/P(\text{HNO}_3) < 0.06$ or $P(\text{H}_2\text{O}_2)/P(\text{HNO}_3) > 0.2$, respectively (Zhang et al., 2009). The regions with ratios between these two limits represent transition situations (Dai et al., 2023).~~

Figure 1 shows how the spatial distribution of ozone regimes varies in response to applied emission reductions in both January and July. Under baseline conditions (BASE case), during January, ozone formation in a large part of China, including the north, east, and

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central areas, as well as the southeastern coastline and PRD regions, is controlled by the availability of VOC (Fig. 1a). In contrast, ozone formation in western China is under the control of NO_x, while a small area in southern China is located in transition areas. Generally, ozone regimes change mainly in the southern part of China when emissions are reduced. With a 50% reduction in NO_x emissions (Fig. 1c), transition or VOC limited regimes in the south and southwest China tends to become NO_x limited for ozone production. With a 50% reduction in AVOCs emissions, some transition areas of southern China are converted to VOC limited areas (Fig. 1e). With the combined reduction in NO_x and AVOCs emissions (Fig. 1g) as well as in all anthropogenic emissions (Fig. 1i), the VOC limited regions evolve towards transition or NO_x limited regions in southern China.

During July, for baseline conditions (BASE case) (Fig. 1b), a large fraction of the Chinese territory corresponds to NO_x limited conditions; VOC limited conditions are found in urban areas including the North China Plain (NCP), the Yangzi River Delta (YRD), the Pearl River Delta (PRD), and the Si Chuan Basin (SCB). The changes in ozone regimes with emission reductions are found mainly in VOC limited areas and surroundings. With the reduction of NO_x emissions, VOC limited areas shrink and become more concentrated in a smaller fraction of metropolitan areas (Fig. 1d). With the reduction of AVOCs emissions, VOC limited areas expand to the surroundings near the metropolitan areas (Fig. 1f). With the combined 50% reduction in the emissions of NO_x and AVOCs (N+A case; Fig. 1h) and of all species (TOTAL case; Fig. 1j), the calculated change pattern of ozone sensitivity is similar, with a smaller VOC limited area relative to the BASE case.

At the specific sites examined here (see Sect. 2.1), emission reduction does not modify ozone sensitivity regimes at the urban and rural sites in January (Fig. S2), remaining in VOCs limited regimes. Except for the remote site Waliguan, at this site, the ozone production is still located in the transition regimes in AVOCs case, while in other cases changing to NO_x limited regime. However, the change in ozone regimes is notable in July (Fig. 2). During this season, in the baseline case, the four urban sites and the rural site of Wangdu are VOC limited; the Heshan site is located in the transition regime and the two remote sites are located in the NO_x limited area. With a 50% reduction in NO_x emissions, only the Guangzhou site remains in a VOC limited region, but also with the value of the $P(\text{H}_2\text{O}_2)/P(\text{HNO}_3)$ ratio increasing from 0.008 to 0.045. The other three urban sites shift from VOC limited regimes to transition areas. With a 50% cut in AVOCs emission, ozone sensitivity at the Hok Tsui site shifts from VOC limited to transition conditions. If we apply a combined 50% reduction to the emissions of NO_x and AVOCs (N+A case) and all emissions (TOTAL case), the three sites of Beijing, Chengdu, and Wangdu shift from a VOC limited to a transition area.

3.21. Changes in precursors and intermediates in ozone formation Changes in the budgets of radicals and ozone

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305 ~~In Part 1 of the present study (Dai et al., 2023), we found that the production of RO_x radicals in China results primarily from the photolysis of O₃, nitrous acid (HONO), and different OVOCs, with a relatively minor role from the ozonolysis of alkenes. Besides, the destruction of RO_x radicals results from the termination reactions between different RO_x radicals and between RO_x radicals and nitric oxide as well as from the heterogeneous uptake of HO₂ on aerosol surfaces. The production rate of odd oxygen (O_x = O₃ + NO₂) is associated with recurrent radical reaction chains involving the oxidation of hydrocarbons in the presence of NO_x. The photochemical destruction of O_x results from several processes, including the photolysis of O₃, followed by the reaction between the electronically excited oxygen atom O(¹D) and water vapor (H₂O). Other O_x loss mechanisms include the reactions of ozone with OH, HO₂, and different alkenes. In the presence of NO_x, additional O_x losses are provided by the titration of O₃ by NO, followed by the conversion of NO₂ to nitric acid (HNO₃).~~

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325 ~~3.2.1. Production and destruction rates of RO_x~~

330 ~~Figure 3 shows the spatial distribution of the changes in the averaged daytime (06:00–19:00 Local Standard Time) production rate of RO_x radicals ($P(\text{RO}_x)$) near the surface resulting from a 50% decrease in NO_x and AVOCs emissions for conditions representative of January and July 2018, respectively.~~

335 ~~In January, with a 50% reduction in NO_x emission, the resulting reduction in $P(\text{RO}_x)$ is of the order of 0.3–0.7 ppbv h⁻¹ (30–33%) in the urban areas, including in the NCP, YRD, and PRD regions (Fig. 3a). The effect of the reduction in AVOCs emissions on the daytime value of $P(\text{RO}_x)$ is less significant than the effect of NO_x emission reduction, with decreases of $P(\text{RO}_x)$ by only 0.1–0.4 ppbv h⁻¹ (10–25%) in the NCP, YRD and SCB (Fig. 3c), but by up to 0.8 ppbv h⁻¹ (35%) in the PRD. More notable effects of AVOCs emission reduction on $P(\text{RO}_x)$ in the PRD are mainly attributed to the reduced contribution from the photolysis of OVOCs (see the results at the Heshan and Hok Tsui sites in Fig. 4). Under the combined reduction in NO_x and AVOCs emissions, the decrease of $P(\text{RO}_x)$ ranges within 0.5–0.9 ppbv h⁻¹ (30–42%), which is overall larger than the sum of the separated effects of NO_x and AVOCs emission reduction (Fig. 3e). With a further decrease in the emissions of other pollutants, the reduction in the daytime $P(\text{RO}_x)$ value, by 0.4–0.8 ppbv h⁻¹ (25–35%; Fig. 3g), is slightly lower than the~~

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350 combined effect of NO_x and AVOCs emissions reduction. The weakened effect on $P(\text{RO}_x)$ is
355 related to the higher concentration of HONO, HCHO, non HCHO OVOCs, and ozone (see
360 Sec. 3.3.1), due to the higher level of oxidants caused by the lower consumption of CO, SO_2 ,
and NH_3 , whose emissions are assumed to be reduced.

In July, the decline of $P(\text{RO}_x)$ due to the reduction in NO_x emissions is larger than in winter
(Fig. 3b), with a maximum decrease of 1.2 ppbv h^{-1} (28%) in urbanized areas. The reduction in
AVOCs emissions leads to a reduction in $P(\text{RO}_x)$ that reaches 0.8 ppbv h^{-1} (18%) in urbanized
areas of East China (Fig. 3d). Under the combined reduction of NO_x and AVOCs emissions,
the reduced value of $P(\text{RO}_x)$ is also larger than when considering their separated effects, which
ranges from 0.5 to 1.2 ppbv h^{-1} (12–28%; Fig. 3f). When the reduction is applied to emissions
of other species, the production rate of RO_x decreases by 0.4 to 1.0 ppbv h^{-1} (9–22% Fig. 3h).

Figure 4 shows the averaged value of daytime $P(\text{RO}_x)$ at eight monitoring sites in all the cases
for January and July, respectively. In January, with the 50% reduced NO_x emissions, the
reduction of $P(\text{RO}_x)$ is predominantly attributed to the reduced contribution from the photolysis
of HONO at urban and rural sites (larger than 70%), which acts as a source of OH. As HONO
mainly emanates from the heterogeneous conversion of NO_2 in polluted East China (Zhang et
al., 2021), the reduced contribution from HONO is associated with the reduced concentrations
of NO_2 . In July, the reduced contribution from the HONO photolysis (due to reduced NO_x
emissions) is also an important factor in explaining the decrease in the $P(\text{RO}_x)$ values (larger
than 70%) in urban areas. When considering the decrease in AVOCs emissions, the reduced
value in daytime $P(\text{RO}_x)$ results to a large extent from the reduction in the photolysis rates of
OVOCs in both seasons.

Figure S3 shows the spatial distribution of the average daytime changes in the destruction rate
of RO_x radicals ($D(\text{RO}_x)$) due to the emission reductions applied in January and July
conditions. The reduction in NO_x emissions has a higher effect on $D(\text{RO}_x)$ than the reduction
in AVOCs emissions. The dominant reason for the decrease in the value of $D(\text{RO}_x)$ is due to
the reduced destruction rate of OH by nitrogen oxides (NO_2 and NO) in January and July (Fig.
S4). A higher contribution from the HO_2 uptake by aerosol, increasing from 3–5% to 5–10%
in January and July, is owing to the higher concentration of HO_2 due to reactions with reduced
NO and less aerosol uptake (See Sec. 3.3.3).

3.2.2. Production and Destruction Rates of O_x

Figure S5 shows the spatial distribution of the changes in the mean daytime production rate of
odd oxygen [$P(\text{O}_x)$] during January and July 2018 resulting from emission changes/reductions.
In January, with the reduction of NO_x emissions, the calculated decrease in $P(\text{O}_x)$ is the largest
in southern and eastern China (by 1.5 – 2.0 ppbv h^{-1} (15–20%); Fig. S5a). Some positive changes
in the O_x production rates are simulated in the metropolitan areas of the YRD and SCB, with
an increase of 0.8 – 1.5 ppbv h^{-1} (7–15%). These positive values are attributed to the increase in
the concentration of HO_2 and RO_2 radicals, as $P(\text{O}_x)$ is strongly contributed by the reaction
rates of NO with the HO_2 and RO_2 radicals. With the reduction in AVOCs emissions, $P(\text{O}_x)$

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395 decreases by up to 4.5 ppbv h^{-1} (42%) in the PRD region (Fig. S5e). When considering
combined reductions of the NO_x and AVOCs emissions, the calculated daytime value of $P(\text{O}_x)$
400 is reduced by 2.3 ppbv h^{-1} (30–60%; Fig. S5e) in the whole of China. When the reduction in all
anthropogenic emissions is considered, the ozone production rate decreases by 3.5 ppbv h^{-1}
(30–70%; Fig. S5g). In different monitoring sites, the highest reduced value of $P(\text{O}_x)$ is up to
4.2 ppbv h^{-1} (40%) at the Guangzhou site, followed by 2.8 ppbv h^{-1} (55%) in the Shanghai site,
in the effect of AVOCs emission reduction only (Fig. S6).

400 The decrease of $P(\text{O}_x)$ for summer conditions is more significant than for winter conditions.
With NO_x emissions reduced by 50% (Fig. S5b), the reduction in the O_x production rate is
within the range of $1.5\text{--}2.0 \text{ ppbv h}^{-1}$ (20–50%). As shown in Fig. S6, the value of $P(\text{O}_x)$ decreases
405 by 50% in Guangzhou, 45% in Shanghai, 30% in Chengdu, and 25% in Beijing. In response to
a 50% reduction in AVOCs emissions, the decrease in the value of $P(\text{O}_x)$ ($0.5\text{--}1.0 \text{ ppbv h}^{-1}$ (10–
28%); Fig. S5d) is smaller than the decrease resulting from the reduction in NO_x emission. We
also found a decrease of $P(\text{O}_x)$ by 28% in Guangzhou, by 10% in Shanghai, by 25% in
410 Chengdu, and by 20% in Beijing (Fig. S6). When the reduction in both NO_x and AVOCs
emissions is applied, the reduced value of $P(\text{O}_x)$ is around $1.6\text{--}2.5 \text{ ppbv h}^{-1}$ (22–55%; Fig. S5f).
With further reduction of other emissions, the values of $P(\text{O}_x)$ decrease by $1.5\text{--}2.2 \text{ ppbv h}^{-1}$ (20–
52%; Fig. S5h).

415 Figure S7 shows the corresponding changes in the destruction rate of odd-oxygen ($D(\text{O}_x)$). The
decrease of $D(\text{O}_x)$ resulting from the 50% reduction of NO_x is in the range of $0.1\text{--}0.5 \text{ ppbv h}^{-1}$
(20–30%) in winter (Fig. S7a), with the highest reduction simulated in the PRD regions. In
summer, the reduction of $D(\text{O}_x)$ reaches $1.0\text{--}2.0 \text{ ppbv h}^{-1}$ (15–25%), with the largest decrease
occurring in the NCP (Fig. S7b). The decrease of $D(\text{O}_x)$ due to the reduction in AVOCs
420 emissions is most prominent in the urbanized areas, with the highest decrease reaching 0.2
 ppbv h^{-1} in winter (25%, Fig. S7c) and 0.8 ppbv h^{-1} (10%) in summer (Fig. S7d).

425 The budgets of radicals and ozone are important indicators for the formation potential of ozone
and other secondary pollutants (Tan et al., 2019). The decrease in the values of the summertime
 $P(\text{RO}_x)$ and $P(\text{O}_x)$ provides some information that can be useful to mitigate secondary
photochemical pollution. The changes in the budgets due to emission reduction suggest that
the most effective way to reduce summertime $P(\text{RO}_x)$ and $P(\text{O}_x)$ is to reduce NO_x emissions.
When applying a combined reduction in the NO_x and AVOCs emissions, the decrease of the
 $P(\text{RO}_x)$ and $P(\text{O}_x)$ values is further enhanced. Thus, reductions in specific AVOCs emissions
are needed to conduct the effective mitigation of air pollution in China.

430 3.3. Changes in the concentrations of ozone and other secondary pollutants

435 Tropospheric ozone is a secondary pollutant, and its formation is largely affected by the levels
of ozone precursors, including NO_x , VOCs, and CO (Wang et al., 2022). The reduction in the
emissions of these primary species leads to changes in the photochemical formation of ozone.
The formation of secondary inorganic aerosols (SIA), including particulate nitrate (NO_3^-),
sulfate (SO_4^{2-}), and ammonium (NH_4^+), is associated with the level of their gas phase

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precursors, such as NO₂, SO₂, and NH₃, and oxidants (Zheng et al., 2015). The oxidizing processes of AVOCs, including benzene, xylene, and toluene, play an important role in the formation of anthropogenic secondary organic aerosol (SOA) (Hu et al., 2017). Owing to high uncertainties in the chemical mechanisms adopted in chemical transport models, only limited studies have assessed the impact of reduced primary emissions on the tropospheric concentration of SOA (Hu et al., 2017; Li et al., 2022). In this section, we quantify the response of ozone precursors, ozone, and secondary aerosols to the reduction in the emissions of primary pollutants.

3.3.1. Precursors and intermediates in ozone formation

Firstly, we describe the changes in the surface concentration of ozone precursors and intermediates, including NO, NO₂, OH radical, HO₂ radical, and specific species of hydrocarbons and OVOCs, in response to different reductions in surface emissions. We show the results for January and July of 2018, as representative for winter and summer conditions. To support the discussion of radical changes induced by emissions reduction, we examine the changes in two parameters. One is the production of RO_x radicals ($P(\text{RO}_x)$), resulting from the photolysis of O₃, nitrous acid (HONO), and of different OVOCs, and from the ozonolysis of alkenes. Another parameter is the destruction of RO_x radicals ($D(\text{RO}_x)$), resulting from the termination reactions between different RO_x radicals and between RO_x radicals and nitric oxide, as well as from the heterogeneous uptake of HO₂ on aerosol surfaces. Detailed estimates of $P(\text{RO}_x)$ and $D(\text{RO}_x)$ can be found in Part 1 of the present study (Dai et al., 2023).

3.1.1. Changes in radicals

Winter condition. Figure 5-1 displays the spatial distribution of the changes in the surface mixing ratios of responses of OH and HO₂ radicals, the ozone precursors and intermediate to resulting from the 50% reduction in the of-NO_x, AVOCs, and combined NO_x and AVOCs emissions for in January of 2018 the daytime. With the reduction in NO_x emissions, the calculated mixing ratio of surface OH radical decreases by 0.05 pptv (40%) in southern China (Fig. 5e), the calculated mixing ratio of surface OH radical is significantly decreased in southern China by up to 40% (0.05 pptv, Fig. 1a), with a relatively lower decrease in the central and western parts of China. The decreases in the NO₂ concentration (Fig. S2a), resulting from the reduced NO_x emissions (Fig. SX), lead to reduced formation of ozone. As a result, the oxidation capacity and levels of oxidants are reduced (see Sect. 3.3). At the same time, an increase in the concentration of the OH radical is found in urban areas, including the North China Plain (NCP), Yangzi River Delta (YRD), Pearl River Delta (PRD), and Si Chuan Basin (SCB) regions, with the maximum increase of 24% in PRD regions. This increase is explained by the reduced destruction of the OH radical by the reaction with NO₂ (Fig. S2b). For the HO₂ radical, a distinct increase in its surface mixing ratio is derived in southern China (by up to 5 pptv (60%), Fig. 1b). This is enhancement is related to the increased OH radical in urban areas, leading to enhanced HO₂ via VOCs oxidation, and with caused by reduced loss of HO₂ via the

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480 aerosol uptake, which is associated with a decrease in the aerosol load (see Sect. 3.2.3), and hence in HO₂ uptake (Song et al., 2021),

485 For the 50% decrease in AVOCs emissions, the mixing ratio of OH is reduced by 4-12% (0.005-0.015 pptv, Fig. 1c) and the mixing ratio of HO₂ radicals is reduced by 20 up to 36% (1-3pptv, Fig. 1d) in the southern part of China. The decreases in these radicals are related to the reduced oxidation of VOCs, due to the reduction in AVOCs emissions. decreased Simultaneously, a slight increase in the mixing ratio of OH radical is derived in the southern part of China. This increase in the OH radical is related to the reduced extinction of radiation associated with the decreased aerosol load due to AVOCs emission reduction.

490 ~~with are calculated with emission reduction~~
When the 50% emission reduction in NO_x is combined with 50% reduction in AVOCs emissions, the distribution of changes in OH radical changes are similar to the pattern induced by the emissions reduction in NO_x alone. However, weakened increase are calculated, as the increase in OH radical with reduced NO_x emissions is largely compensated by the decreased radicals due to the AVOCs emission reduction to radical concentration. -As shown in Fig. 1e, the maximum increase in urban China is also lowered to 12% (from 40%). Besides that, the increase in HO₂ radicals is also largely reduced to 20% (from 60%) by the lowered with AVOCs emissions, with only a mild increase of HO₂ distributed in the southern coast of China (Fig. 1f).

500 With accounting in addition to the reduction of other anthropogenic emissions (NH₃, SO₂, and CO), the mixing ratio of OH radical is positively modified, relative to the results in the combined case (N+A case). As shown in Fig. S5a, the mixing ratio of OH radical is increasing in the PRD and SCB regions (by up to 22%). This increase is due to the reduced consumption of the OH radical by CO (Fig. S4a), due to reduced CO emissions (Fig. S1d). For the HO₂ radicals, the additional emissions reduction in other emissions also contributes to a larger mixing ratio, with a pronounced increase in southern China in January (by up to 18%, Fig. S5b). This increase in HO₂ radical mixing ratio is due to the increase in the oxidation of VOCs by OH radical and the reduced aerosol uptake of HO₂, associated with the decrease in aerosol load, where the ozone sensitivity is controlled by NO_x (Fig. 1c), is due to the decrease in the oxidation capacity of atmospheric (See Sec.3.4.2). At the same time, a relatively strong increase of OH is found in the PRD region (by 0.03 pptv; 24%), due to less consumption by the reduction in NO_x and more production from the photolysis of enhanced OVOCs (Fig. 6g). An increase in the surface mixing ratio of the HO₂ radical is derived in southern China (by up to 5 pptv (60%); Fig. 5g), which is associated with a decrease in the aerosol load (see Sect. 3.3.3), and hence in reduced HO₂ uptake (Song et al., 2021). In January, a decrease of NO (by up to 6 ppbv (40%); Fig. 5a) and NO₂ (by up to 8 ppbv (25%); Fig. 5e) is derived in the metropolitan regions of China, which is consistent with the spatial distribution of reduction in NO_x emissions (Fig. S1a). The calculated mixing ratio of surface OH radical decreases by 0.05 pptv (40%) in southern China (Fig. 5e), where the ozone sensitivity is controlled by NO_x (Fig. 1c), is due to the decrease in the oxidation capacity of atmospheric (See Sec.3.4.2). At the same time, a relatively strong increase of OH is found in the PRD region (by 0.03 pptv; 24%), due to less consumption by the reduction in NO_x and more production from the photolysis of

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enhanced OVOCs (Fig. 6g). An increase in the surface mixing ratio of the HO₂ radical is derived in southern China (by up to 5 pptv (60%); Fig. 5g), which is associated with a decrease in the aerosol load (see Sect. 3.3.3), and hence in reduced HO₂ uptake (Song et al., 2021).

In July, the decrease in the concentration of NO (4 ppbv (50%); Fig. 5b) and NO₂ (6 ppbv (35%); Fig. 5d) due to the reduction in the NO_x emission is smaller than in winter, which is related with the smaller reduction in summertime NO_x emissions (Fig. S1g). The level of summertime decreases in OH radicals (by up to 0.15 pptv (30–40%); Fig. 5f) is larger than the decrease derived in winter and displayed in broader areas. The spatially varied distribution in the changes of OH and HO₂ radicals in two seasons can be explained by the seasonal variations of UV, water vapor, and solar radiation, with high values concentrated in southern China only in winter and evenly distributed in the whole of China in summer (Dai et al., 2023). An increase in the surface mixing ratio of HO₂ radical is derived in the NCP (6–8 pptv; 15–20%) and is due to the decreased uptake by the aerosol surfaces. An increase of alkene is calculated in July by 0.2–0.5 ppbv (10–20%), owing to the reduced consumption by decreased OH radicals.

For the 50% decrease in AVOCs emissions, the calculated surface concentration of alkenes decreased by up to 3.0–4.0 ppbv (30–40%; Fig. 7a) in urban China in January. At the same time, a decrease is calculated in the mixing ratio of OH (by 0.005–0.015 pptv (4–12%); Fig. 7e) and HO₂ radicals (by 3 pptv (36%); Fig. 7e) in the southern part of China. Reasons for the decreases in these radicals are the declined contributions from the VOCs oxidation and OVOCs photolysis. Owing to the reduced consumption of hydrogenated radicals, an increase of NO (by 1.2–2.0 ppbv (7–12%)) is derived in the urban areas of China (Fig. S8a). In July, the calculated decrease of alkenes is estimated to be 1.0–2.0 ppbv (30–50%; Fig. 7b). The reduction in VOCs also leads to a summertime decrease in radicals in urban areas, with the decrease of OH by 0.03–0.05 pptv (8–12%) and HO₂ radicals by 3–5 pptv (6–10%). The relevant summertime enhancement in NO concentration is 0.8–1.0 ppbv (10–15%; Fig. S8b).

When applying the combined 50% reduction in NO_x and AVOCs emissions, the changes in the spatial distribution of radicals are similar to the changes derived when a reduction was applied only to the NO_x emissions alone. Compared with the results in the reduction NO_x emissions alone, a smaller increase of OH radical is provided in urban China in January (by 0.01 pptv (12%); Fig. 8a) and in July (by 0.07 pptv (15–20%); Fig. 8b), which is attributable to the less production of OH radical from the VOCs oxidation due to reduction in AVOCs emission. The decrease of OH radical is also suppressed to 0.02 pptv (16%) in January and 0.1 pptv (10–12%) in July, which is due to the lower destruction rate of OH from the reaction with reduced HO₂ radicals. As the negative effects of declined VOCs oxidation due to AVOCs emissions reduction, the enhanced HO₂ radicals in the NO_x emission reduction case are largely offset. As shown in Fig. 8c and d, the combined case derived less than 1 pptv (12%) of wintertime enhancement in the southern coastal areas of China and 3 pptv (6%) of summertime increase in the urban areas.

When applying a 50% reduction to NO_x, AVOCs, and other anthropogenic emissions, the mixing ratio of OH radical is positively varied compared with the results in the combined case

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(N+A case). A distinct increase of wintertime OH radical is derived in the PRD and SCB regions (by up to 0.03 pptv; Fig. 8e). This increase is owing to the lower consumption of OH radical by the reduced concentration of NH₃, SO₂, and CO, due to their reduced emissions. In July, the increase of OH radical is also enhanced to 0.01 pptv (Fig. 8f) in VOC limited areas. At the same time, the decreases of OH radicals in NO_x limited areas are suppressed in both winter and summer, indicating the increasing atmospheric oxidative capacity in these areas. For HO₂ radicals, the additional emissions reduction also contributed to an increase in its mixing ratio, with a pronounced increase in southern China in January (by 1.5 pptv (18%) Fig. 8g) and in the NCP region in July (by 6 pptv (12%); Fig. 8h). The enhancement in HO₂ radicals is caused by less HO₂ loss via the aerosol uptake, as a decrease is derived in aerosol concentration due to the reduction in the precursor of secondary inorganic aerosols (See Sec. 3.3.3).

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Summer condition. Figure 2 displays the spatial distribution of the changes in OH and HO₂ radicals due to the reduction of NO_x, AVOCs, and combined NO_x and AVOCs emissions for July 2018. Compared with the wintertime variation in radicals, which is limited to southern China, the geographical area covered by these changes is broader during summertime and covers a large fraction of northern China. When applying a reduction in NO_x emission by 50%, a decrease in the concentration of OH radicals is derived in large parts of China, with the maximum decrease reaching 40% (0.15 pptv, Fig. 2a). The concentration of the OH radical increases in the metropolitan areas, including in the YRD and PRD regions. Simultaneously, the surface mixing ratio of HO₂ radical increases by 6 to 8 pptv (15-20%, Fig. 2d) in the North China plain. These changes are affected by the meteorological parameters including the temperature, the water vapor abundance, and the solar radiation intensity, which are involved in the oxidative processes (Dai et al., 2023).

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When AVOCs emissions are reduced by 50%, the concentration of the radicals in urban areas, including in the NCP, YRD, and PRD regions, decreases on the average by 8-12% in the case of OH (0.03-0.05 pptv, Fig. 2b) and by 6-10% (3-5 pptv, Fig. 2e) in the case of HO₂. Simultaneously, a comparable increase in the OH radical is calculated in non-urban areas, due to potential effect by the aerosol light extinction. When applying the combined 50% emissions reduction in AVOCs with NO_x, the distribution pattern of OH radical changes are similar to the distribution derived for the reduction in NO_x emissions alone, but it is also compensated by the offset effect from AVOCs emissions, like in winter conditions. As shown in Fig. 2c, the maximum increase in OH radical is weakened to 20% (from 40%) and the maximum decrease is also reduced to 12% (from 30%). The offset effect of AVOCs emissions reduction is also shown in the enhanced abundance of HO₂ radicals (Fig. 2f), with less than 6% (from 15-20%) increases in the urban areas.

With an additional 50% reduction to other anthropogenic emissions, the mixing ratio of OH radical is increased (Fig. S4c), relative to the results obtained in the combined case, with the highest increase occurring in the NCP region.

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3.1.2 Changes in OVOCs

Oxygenated hydrocarbons (OVOCs) originate from direct surface emissions (primary source), and from the oxidation of primary hydrocarbons (secondary source) in the atmosphere (Wang W. et al., 2022).

Winter condition. Figure 3 shows the spatial distribution of calculated changes in total OVOCs due to the 50% reduction in NO_x, AVOCs and in combined NO_x and AVOCs emissions for January of 2018. With a reduction in NO_x emission, the OVOCs concentration decreases in the non-urban areas in southern China and increases in urbanized China (Fig. 3a), which is consistent with the changes in the mixing ratio OH radical. The highest increase in OVOCs concentration is approximately 10% (2 ppbv), in the urban areas of the YRD and PRD regions; it includes a significant increase in the concentration of formaldehyde (HCHO, Fig. S6a), followed by peroxyacetyl nitrate (PAN, Fig. S6b), and alcohols (CH₃OH+C₂H₅OH, Fig. S6c). At four city sites under consideration, the highest increase in OVOCs is calculated in the site of Shanghai and Guangzhou, with concentrations increasing by about 12% (1.8 ppbv, Fig. 3e) and 8% (1.2 ppbv, Fig. 3f), respectively.

When AVOCs emissions are reduced, OVOCs are decreased in most regions of China (Fig. 3b), with the highest decrease found in the regions of PRD and SCB. At the four city sites under consideration, the decrease in ketones is the most distinct (Fig. 3d-g). The abundance of these species is reduced by almost a factor of two, followed by HCHO and other aldehydes. When combining the emission reduction of AVOCs and NO_x, the decrease in OVOCs concentration resulting from the AVOCs emission reduction is further strengthened in large parts of China (Fig. 3c). However, an increased concentration of 1~2 ppbv is derived for OVOCs at Guangzhou site (Fig. 3f). With additional decreases in other emissions, the OVOCs' concentration is enhanced by 2~4 ppbv in whole China (Fig. S5e), which is consistent with the increased abundance of the OH radical resulting from a reduction in the NH₃, SO₂, and CO emissions. relevant dwith ionfroms OVOCs are both contributed by the primary source, their direct emissions, and the secondary source, the oxidation of hydrocarbons (Wang W. et al., 2022). With the reduction in emissions, the concentrations of OVOCs also change substantially. Figure 6 g and h displays the spatial distribution of the changes in total OVOCs due to the 50% reduction in NO_x emissions in January and July of 2018. During January, there is an increase in the calculated concentration of OVOCs in the NCP, YRD, and PRD. In these areas, the highest wintertime increases in the concentration of OVOCs reach about 0.5-1.0 ppbv (5-10%; Fig. 6g). At specific sites, the highest increase in OVOCs concentrations is calculated at the Shanghai and Guangzhou sites. As shown in Fig. 9, at these two urban sites, the reduction in NO_x emissions leads to an increase in OVOCs concentration by 1.8 ppbv (12%) and 1.2 ppbv (8%), respectively, which is mainly contributed by the increase of HCHO (Fig. 6e) due to the higher secondary formation from the oxidation of VOCs (Li et al., 2021). When AVOCs emissions are reduced, a decrease in OVOCs concentration is simulated in most regions of East China (by up to 8 ppbv (12%); Fig. 7g), especially the SCB up to (X%) and PRD regions by up to (X%), with the (Fig. SX) highest decrease in ketones (by 2-3 ppb (10-

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655 15%); Fig. 9). However, a rebounded OVOCs is derived at Guangzhou site by 1–2 ppbv (Fig. 9); further

660 ~~In summer, with the 50% NO_x emission reduction, the increases of simulated OVOCs concentrations in North China are more pronounced than in the wintertime, with an increase of 0.5–1.5 ppbv (10–15%); Fig. 6h). These OVOCs increases in NO_x limited areas are associated with the enhanced hydrocarbons concentration, including alkenes (Fig. 6d) and isoprene (Fig. S10a) due to less loss via the oxidation by OH radicals. While, in VOC limited areas, the increase of OVOCs can be attributed to the increased mixing ratio of OH radical (Fig. 5f) due to less titration effect. With the 50% AVOCs emission reduction, similarly with the wintertime change, the simulated OVOCs decrease reaches up to about 5 ppbv (20%); Fig. 7h), with the largest contribution calculated in ketones (by 1–3 ppbv (8–15%); Fig. 9). Limited changes (less than 5%) are found in the concentration of OVOCs in the sub-rural sites Heshan, which is relevant to the major contribution of BVOCs, which are not notably influenced by anthropogenic emissions, to the secondary OVOCs formation at this site (Dai et al., 2023).~~

675 ~~Quantification of the impact of emissions reduction on OVOCs abundance provides an important indicator of its impacts on the atmospheric oxidative capacity since the photolysis of OVOCs produces HO_x radicals that drive fast ozone production and accelerate further VOC oxidation (Li et al., 2021). Li et al., (2021) reported that the fast photochemical production of ozone during the Covid-19 lockdown was attributed to the sharp reduction of the NO_x emissions in urban China and the high VOCs emissions, driving ozone production through HCHO photolysis. The study also suggests extending the VOC emission controls year round to avoid the spread of ozone pollution outside the summer season. In our study, with the 50% reduction of NO_x emissions, the increased concentration of OVOCs, including HCHO, is simulated in both winter and summer, which is supportive of Li et al., (2021). The discussion in the following sections explores the changes in other secondary pollutants, including ozone and secondary aerosols.~~

685 the Simultaneously the summertime increases
3.3.2. Ozone

690 Summer conditions. Figure 4 displays the spatial distribution of the changes in total OVOCs due to a 50% reduction in NO_x, AVOCs and in the combined NO_x and AVOCs emissions for July 2018. In summer, Wwith a conducting the 50% reduction in NO_x emissions reduction, the changes in OVOCs concentration do not resemble with changes in the OH radical. A decrease in the OVOCs is derived in southern China. However, in central and northern China, the level of OVOCs generally increases and this increase is not limited to metropolitan areas similar with the impact of changes in OH radical on OVOCs concentration during winter. the increases of simulated OVOCs concentrations in North China are more pronounced than in the wintertime, with an increase of 0.5–1.5 ppbv (10–15%); Fig. 6h). (Fig. 4a). also For specific OVOCs, the changes in HCHO, glyoxal, and PAN (Fig. S7a-c) are consistent with the changes calculated for the OH radical. However, for aldehydes and alcohols (Fig. S8a-b), an increased concentration is derived in the entire geographical area of China, with a consistent distribution

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700 of changes in the level of alkenes (Fig. S8c) and isoprene (Fig. S8d). In summer, the temperature-dependent emissions of biogenic VOCs from vegetations are high. The reduced production rate of OVOCs, resulting from the reduced AVOCs emissions, is compensated to a certain extent by the high natural emissions of BVOCs species.

705 ~~due to the~~ While, in winter, the change of isoprene is ignorable. These OVOCs increases in ~~NO_x limited areas are associated with the enhanced hydrocarbons concentration, including alkenes (Fig. 6d) and isoprene (Fig. S10a) due to less loss via the oxidation by OH radicals. While, in VOC limited areas, the increase of OVOCs can be attributed to the increased mixing ratio of OH radical (Fig. 5f) due to less titration effect. With a the-50% reduction in AVOCs emission reduction, similarly with the wintertime change, the significant the simulated decreases in OVOCs concentrations are significantly reduced decrease is simulated in the NCP and SCB regions, with the maximum decrease by reaches up to about 5 ppbv (20% (Fig. 4b7h), with the largest contribution calculated in ketones (by 1-3 ppbv (8-15%); Fig. 9). Differentsimilarly with what is occurring in large-winter (time decrease in OVOCs concentration reduced by about (-, 50%), wintertime change, the response of OVOCs concentration to AVOCs emissions reduction is milder in summer, with the largest decrease reduced to suppressed to-30%. This is explained by the more effective secondary formation of OVOCs during summertime. Limited changes (less than 5%) are found in the concentration of OVOCs in the sub-rural sites Heshan, which is relevant to the major contribution of BVOCs, which are not notably influenced by anthropogenic emissions, to the secondary OVOCs formation at this site (Dai et al., 2023).~~

715 Quantification of the impact of emissions reduction on OVOCs abundance provides an important indicator of its impacts on the atmospheric oxidative capacity since the photolysis of OVOCs produces HO₂ radicals that drive fast ozone production and accelerate further VOC oxidation (Li et al., 2021). Li et al., (2021) reported that the fast photochemical production of ozone during the Covid-19 lockdown was attributed to the sharp reduction of the NO_x emissions in urban China and the high VOCs emissions, driving ozone production through HCHO photolysis. The study also suggests extending the VOC emission controls year round to avoid the spread of ozone pollution outside the summer season. In our study, with the 50% reduction of NO_x emissions, the increased concentration of OVOCs, including HCHO, is simulated in both winter and summer, which is supportive of Li et al., (2021). The discussion in the following sections explores the changes in other secondary pollutants, including ozone and secondary aerosols.

735 Figure 10 shows the changes in surface daytime ozone concentrations resulting from the 50% emissions reduction for January and July conditions. In winter, the reduction in NO_x emission leads to an increase in the surface ozone concentrations, which is the largest in the YRD and PRD regions (8-10 ppbv (15-20%); Fig. 10a). During wintertime, as shown in Fig. 1a, a large part of China is under a VOC sensitive regime. In this NO_x-reduced case, the weakened titration effect due to the decrease in NO_x concentrations favors ozone formation. In AVOCs emissions reduction case, a reduction in the surface ozone concentration, ranging from 2.0 to

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8.0 ppbv (4–10%; Fig. 10e), is calculated in the southern part of China. This ozone decrease is owing to the decelerated ozone production rate attributed to the decrease in the concentration of hydrocarbons (Fig. 7a) (Jacob et al., 1995) and lower HO_x radicals originating from the photolysis of reduced OVOCs concentration (Fig. 7g).

In the combined emission reduction case, the wintertime ozone changes in a large part of China have primarily followed the ozone changes with NO_x emissions reduction, with an ozone increase of 3.0–7.5 ppbv (4–9%; Fig. 10e) in VOC limited areas. As the distribution of ozone response mainly depends on the sensitive areas, in this combined case, a large part of China, including North China and some urban regions in South China, is located in VOC limited areas (Fig. 1g). Thus, in these areas, the ozone response mainly follows the positive changes in the NO_x reduction case. Simultaneously, a slight ozone decrease (by 2.0–4.5 ppbv; 5–8%) is derived by the model over the southern coast of China. In these areas, the ozone sensitivity is under the control of the NO_x. Therefore, the ozone decrease is relevant to the negative ozone response in AVOCs emissions reduction case. Compared with the combined case, with further emission reduction in other species, including CO, SO₂, and NH₃, a smaller ozone increase (by 3–5 ppbv; 4–6%) is calculated in the southern part of China (Fig. 10g). One of the reasons for the higher ozone formation is the reduced CO concentration, as an ozone precursor, due to the CO emission reduction. Besides, a higher level of HO₂, resulting from less aerosol uptake due to emission reduction of NH₃ and SO₂ (See Sec. 3.3.3), can also lead to higher ozone formation.

In summer, an increase in the surface ozone concentration by up to 12 ppbv (10–20%) is only calculated in the urbanized NCP, YRD, and PRD (Fig. 10b), under the effect of NO_x emissions reduction. These areas are typically located in VOC limited conditions (Fig. 1b); thus, ozone increase can be well explained by reduced ozone titration. At the same time, in NO_x limited areas, the calculated surface ozone concentration is reduced by 2 to 8 ppbv (3–10%). This ozone decrease is associated with less photochemical formation from reduced NO (Fig. 5b) and NO₂ (Fig. 5d) concentrations. With the reduction of AVOCs emissions, the surface ozone concentration decreases in whole areas of China by up to 8.0–12.0 ppbv (8–20%; Fig. 10d). Unlike wintertime ozone decrease, which primarily occurs in southern China, the summertime ozone decline shifts from the southern regions to the northern ones, including the NCP, the PRD, and the YRD regions. This spatial variation is consistent with the distribution of model-derived HO₂ radicals' changes (Fig. 7f), indicating the importance of AVOCs to the formation of HO₂ radicals.

In the combined 50% reduction case of the NO_x and AVOCs emissions, the surface ozone concentration decreases by up to 12 ppbv (15%; Fig. 10f) in NO_x sensitive areas due to the influence of the reduction in the NO_x and AVOCs emissions. In VOC sensitive areas, the increase of ozone associated with the positive impact of the reduction in NO_x emissions is largely compensated by the negative effect of the reduction in AVOCs emissions. An ozone decrease is derived in almost all areas of China and ranges from about 5 to 12 ppbv (6–15%). One exception is found at the Guangzhou site, where ozone slightly increases by 0.5 ppbv (Fig. S11). The important contribution of BVOCs to the formation of ozone in the PRD region has been documented in many studies (Zhang et al., 2023). One reason can be the role of BVOCs

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species, such as isoprene, becomes more important to ozone formation in this area, when NO_x and AVOCs emissions are reduced (see Sec. 3.4.1).

When emission reduction is conducted to all species, the largest ozone decrease is slightly rebounded to 10 ppbv (15%) in NO_x-limited conditions and 11 ppbv (13%) in VOCs limited conditions (Fig. 10h).

This change in two season indicates less sensitive of OVOCs formation to the reduction in AVOCs emissions during summer than during winter. Nevertheless, besides that, a more pronounced distinct decrease in OVOCs is found in North China. For example, at the Beijing site, the decreases in OVOCs concentration increased to reaches 10 ppbv (28%10 ppbv compared to the) from the decrease in winter by 5 ppbv (45%5 ppbv derived in winter), which is consistent with the higher concentration of summertime OVOCs (~34 ppbv) relative to winter (~11 ppbv). When the combined reduction in the emissions reduction in of AVOCs with NO_x are considered, a lower decrease of OVOCs is can be found in the whole geographical area of China (Fig. 4cX), which is relevant to the reduced decrease in the level of the OH radical compared with the results with only AVOCs emission reduction, indicating a more significant impact of OH radical with isoprene when AVOCs emission reduced. The effect on With additional reduction in other emissions, the changes in OVOCs concentration of the reduced emissions in the other species under consideration are mildysmall (Fig. 4SfFig. S).

3.2. Changes in secondary pollutants to emissions reduction

Following the changes in ozone precursors and intermediate, the formation of ozone and secondary aerosol will also be affected by the reduction in emissions. The formation of ozone and secondary aerosol is affected by the changes in the emissions of primary species and in the concentration of ozone precursors and intermediate species. The production rate of odd oxygen ($O_3 = O_2 + NO_2$) is associated with recurrent radical reaction chains involving the oxidation of hydrocarbons in the presence of NO_x. The photochemical destruction of O₃ results from several processes, including the photolysis of O₃, followed by the reaction between the electronically excited oxygen atom O(¹D) and water vapor (H₂O). Other O₃ loss mechanisms include the reactions of ozone with OH, HO₂ and different alkenes. In the presence of NO_x, additional O₃ losses are provided by the titration of O₃ by NO, followed by the conversion of NO₂ to nitric acid (HNO₃).

In order to highlight the regional differences in the response of the existing photochemical regimes to emissions reduction, we first show the changes in the distributions of the areas where the ozone formation is either NO_x- or VOC-limited. In order to display the impact of emission reduction on the changes in the ozone regimes, we first show the geographical distribution of NO_x-limited or VOC limited ozone formation regimes. We distinguish the two regions from adopt the calculated ratio between the production rate of hydrogen peroxide (H₂O₂) and that of nitric acid (HNO₃) [$P(H_2O_2)/P(HNO_3)$] as the indicator to distinguish these regimes. An area is assumed to be VOC-limited or NO_x-limited if the adopted indicator $P(H_2O_2)/P(HNO_3)$ is

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smaller than ≤ 0.06 or if it is larger than $P(\text{H}_2\text{O}_2)/P(\text{HNO}_3) > 0.2$, respectively (Tonnesen and Dennis, 2000) (Zhang et al., 2009). The regions with ratios between these two limits represent transition situations (Dai et al., 2023). The production rate of odd oxygen ($\text{O}_x = \text{O}_3 + \text{NO}_2$) associated with recurrent radical reaction chains involving the oxidation of hydrocarbons in the presence of NO_x and the photochemical destruction of O_x , is also adopted here to support the analysis of the underlying reasons for ozone changes.

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3.2.1 Response of ozone sensitivity regimes to emissions reduction

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In order to display the impact of emission reduction on the changes in the ozone regimes, we first show the geographical distribution of NO_x -limited or VOC-limited ozone formation regimes. We adopt the ratio between the production rate of hydrogen peroxide (H_2O_2) and that of nitric acid (HNO_3) [$P(\text{H}_2\text{O}_2)/P(\text{HNO}_3)$] as the indicator to distinguish these regimes. An area is assumed to be VOC-limited or NO_x -limited if $P(\text{H}_2\text{O}_2)/P(\text{HNO}_3) < 0.06$ or $P(\text{H}_2\text{O}_2)/P(\text{HNO}_3) > 0.2$, respectively (Zhang et al., 2009). The regions with ratios between these two limits represent transition situations (Dai et al., 2023).

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Figure 45 displays the spatial distribution of ozone regimes in response to applied emission reductions in the combined emissions in NO_x and AVOCs emissions in both January and July and July.

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Winter conditions. In January, under baseline conditions (BASE case), during January, ozone formation in a large part of China, including the north, east, and central areas, as well as the southeastern coastline and PRD regions, is controlled by the availability of VOC (Fig. 1a). In contrast, ozone formation in western China is under the control of NO_x , while a small area in southern China is located in transition areas. Generally, ozone regimes change mainly in the southern part of China when emissions are reduced, with a 50% reduction in NO_x emissions (Fig. 5b), transition or VOC-limited regimes in the south and southwest of China (in BASE case, Fig. 5a) tend to be converted into NO_x -limited for ozone production regimes. The change in ozone sensitivity regimes is consistent with (1) the decreased concentration in the simulated HNO_3 (Fig. S9a) due to less NO_2 reacting with OH, and (2) the increased H_2O_2 concentration (Fig. S9e), due to the reduced aerosol HO_2 uptake by aerosols, e-changes in H_2O_2 and HNO_3 . With a 50% reduction in AVOCs emissions, some transition areas of southern China are converted to VOC-limited areas (Fig. 5c). A relevant decrease in H_2O_2 concentration also derived in southern China and is attributed to the decrease in the calculated HO_2 concentration. When considering the combined reduction in NO_x and AVOCs emissions as well as all anthropogenic emissions, the VOC-limited regions of southern China evolve towards transition regions or even NO_x -limited regions (Fig. 5d, Fig. S10a). In these two last cases, the changes in ozone sensitivity regimes are determined by the decrease in the calculated HNO_3 concentrations.

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At the urban sites specific sites examined here (see Sect. 2.1), the emission reduction does not modify the wintertime ozone sensitivity regimes at the urban and rural sites in January (Fig. S11), which remaining in VOCs-limited regimes. Except for the remote site Waliguan, at this

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920 a large part of China, including North China and some urban regions in South China, is located
in VOC limited areas (Fig. 1g). Thus, in these areas, the ozone response mainly follows the
925 positive changes in the NO_x reduction case. Simultaneously, a slight ozone decrease (by 2.0-
4.5 ppbv; 5-8%) is derived over the southern coast of China (by 2.0-4.5 ppbv; 5-8%). In these
areas, the ozone sensitivity is under the control of the NO_x. The ozone decrease is dominant by
the negative ozone response in AVOCs emissions reduction case. Compared with the combined
case, with further emission reduction of other species compared with the combined case, a
higher ozone increase (3-5 ppbv (4-6%); Fig. S5g) relative to the combined case (by 3-5 ppbv;
4-6%) is calculated in the southern part of China, (by 3-5 ppbv (4-6%), concentration of ozone
precursor via aerosols reduction.

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930 *Summer condition.* In July, under the effect of reduction in the NO_x emissions reduction, an
increase in the surface ozone concentration of up to 10 ppbv (17%) is calculated in the
urbanized regions of NCP, YRD, and PRD (Fig. 6d). under the effect of NO_x emissions
935 reduction. These areas are typically located in VOC-limited areas (Fig. 5); thus, ozone increase
is explained by the reduced ozone titration (Fig. S13b). At the same time, in NO_x-limited areas,
the calculated surface ozone concentration is reduced by 2 to 8 ppbv (3-10%), as a result of
less photochemical formation from reduced NO_x, NO (Fig. 5b) and NO₂ (Fig. 5d)
concentrations. With the reduction of AVOCs emissions, the surface ozone concentration
940 decreases by up to 8.0-12.0 ppbv (8-20%; Fig. 6e) in whole areas of China. A spatial shift in
ozone decrease, from the southern regions in winter to the northern ones in summer, also occurs
under this condition, which is relevant with the distribution of model-derived HO₂ radicals'
decrease, as it involved in the reaction with NO and contributed to the decreased ozone
production rate (Fig. S14b).

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945 When combining the 50% reduction in the NO_x and AVOCs emissions, the surface ozone
concentration decreases by up to 12 ppbv (15%; Fig. 6f) in NO_x-sensitive areas. In VOC-
sensitive areas, the increase of ozone associated with the positive impact of the reduction in
NO_x emissions is also surpassed by the negative effect of the reduction in the AVOCs
emissions. In other words, the reduced destruction of ozone associated with the reduced NO
950 level is less important than the decreased production of ozone resulting from the reduced
mixing ratio of HO_x radical. This is explained by the fact that the consumed ozone can be
rapidly compensated by the photochemically ozone formation processes, since the ozone
production rate is accelerated by high temperature and photolysis rate during summertime
(Wang et al., 2022). While the decrease in HO_x radical, due to reduced AVOCs emissions has
a stronger impact on the ozone production, as the radicals are fast consumed in the oxidative
955 processes during summer. One exception can be found at the Guangzhou site, where ozone
slightly increases by 0.5 ppbv (Fig. S15), which is possibly explained by the increasingly
important role of naturally emitted BVOCs species in the oxidation processes when
anthropogenic emissions are reduced. When emission reduction is applied to all species, the
largest ozone decrease is slightly increased to 10 ppbv (15%) in NO_x-limited conditions and
11 ppbv (13%) in VOCs-limited conditions (Fig. S5h) relative to combined cases, due to the
960 increase in the mixing ratio of OH and HO₂ radicals.

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3.2.3. Response of aerosols to emissions reduction and its potential effects on ozone 7T (OT), which is largely contributed

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Figure 7 shows the changes in the average concentrations of secondary aerosol due to emissions reduction in January and July of 2018.

Winter condition. In January, the 50% reduction of NO_x leads to a large decrease of aerosol load (10-18 μg m⁻³ or 12-20%; Fig. 7a) in central and southern China. The aerosol decrease is predominantly results from the decrease in the NO₃⁻ abundance (Fig. S16a) linked to the reduced concentration of NO₂, followed by NO₄⁺ (Fig. S16b). A slight increase in the abundance of secondary organic aerosols (SOA) is derived in the urban areas of NCP, YRD, and PRD regions (1-2 μg m⁻³ or 3-5%; Fig. S16c), which is consistent with the increase in the level of oxidants, including ozone and OH radicals. With a 50% reduction of AVOCs emissions, the changes in aerosol concentration are smaller than with the 50% reduction in NO_x emissions reduction, with a decrease of less than 5 μg m⁻³ (4%; Fig. 7b), which is predominantly resulting from the reduction in SOA (Fig. S17a). With a joint reduction in NO_x and AVOCs (Fig. 7c), the aerosol decrease is larger than the separated effect by emissions decrease, as the increase in SOA by reduced NO_x emissions is compensated by the reduced AVOCs emissions.

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With further reduction in other emissions, the decrease in the concentration of aerosol is deeply enhanced; this is the case for the concentration of NH₄⁺ (Fig. S19a), SO₄²⁻ (Fig. S19b), and NO₃⁻ (Fig. S19c). Conversely, the concentration of their precursors, NH₃ and SO₂, are largely reduced, which can also affect the process of acid replacement and then the level of NO₃⁻. An overestimation in the concentration of NO₂ and PM_{2.5} has been simulated in the baseline condition, which can possibly lead to a higher reduction in aerosol concentration, especially in the concentration of NO₃⁻. This overestimation can also potentially affect the aerosol-related changes in ozone formation.

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Summer condition. In July, the aerosol decrease due to the emissions reduction is much smaller than the decrease in winter. The decreased values range from 1.5 to 5 μg m⁻³ (Fig. 7d), from 2 to 6 μg m⁻³ (Fig. 7e), from 4 to 7 μg m⁻³ (Fig. 7f) and from 8 to 10 μg m⁻³ (Fig. 7g), for a reduction in NO_x, AVOCs, combined NO_x and AVOCs, and TOTAL emissions, respectively. The simulated reductions in aerosols undergo a spatial shift, from the southern part of China in winter to the northern China Plain in summer. This shift is consistent with the calculated changes in oxidants, hydrocarbons, and other gaseous aerosol precursors. The higher decrease in the aerosols loads for the combined case also indicates that the reduction in AVOCs emission will increase the efficiency of the aerosol decrease produced by the reduced NO_x emission.

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In July

The aerosol effect on ozone formation has been discussed in several modeling studies (Li et al., 2019; Liu et al., 2020; Dai et al., 2023). Our results show that the reduction in primary emissions results in a large decrease in aerosol concentrations. The major contribution to aerosol decrease is from the reduction in NO_x emissions, with a strengthened effect on aerosol

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reaches $15\text{--}25\ \mu\text{g}\cdot\text{m}^{-3}$ (20–38%; Fig. 11g), while in July, it reaches $8\text{--}10\ \mu\text{g}\cdot\text{m}^{-3}$ (30–40%; Fig. 11h). This enhancement in the aerosol decrease is caused to a large extent by the decrease in the concentration of NH_4^+ (Fig. S13 g and h) and SO_4^{2-} (Fig. S14 g and h).

Figure 12 shows the changes in the daytime average concentrations of secondary organic aerosol (SOA) due to emissions reduction in January and July. With 50% NO_x emission reduction in winter, a slight increase of SOA ($1.0\text{--}1.5\ \mu\text{g}\cdot\text{m}^{-3}$ (3–5%; Fig. 12a)) is derived in the urban areas of NCP, YRD, and PRD regions, which is consistent with the increase in oxidants, including ozone and OH radicals. When AVOC emissions are reduced, the simulated decrease in SOA concentration is about $2\text{--}5\ \mu\text{g}\cdot\text{m}^{-3}$ (5–11%; Fig. 12c) in the southern part of China. In summer, the decreased level of SOA is comparable to the wintertime decrease. However, the simulated reductions in SOA undergo a spatial shift, from the southern part of China in winter to the northern China Plain. Reasons for the seasonal variations in spatial pattern are the spatial distribution of SOA precursors, including oxidants, hydrocarbons, and OVOCs, and the seasonally changing meteorological parameters, such as temperature and solar radiation.

The aerosol effect on ozone formation has been discussed in several modeling studies (Li et al., 2019; Liu et al., 2020; Dai et al., 2023). Our results show that the SIA concentration is largely decreased in winter due to NO_x emissions reduction, while the declined SOA concentration relies more on the reduction of AVOCs emissions. This aerosol decrease weakens the aerosol extinction effect and therefore enhances the ozone photochemical formation rate (Tan et al., 2022). At the same time, aerosol decreases can result in a decline in aerosol uptake through heterogeneous reactions, resulting in an enhanced concentration of NO_2 and HO_2 radicals. Moreover, the formation of HONO will also decline with NO_x emission reduction (Fig. 6 a and b) through the heterogeneous conversion from NO_2 , photolysis of NO_3^- and direct emission from transport (Dai et al., 2023). Thus, the atmospheric level of the OH radical will also be reduced, as the photolysis of HONO is an important source of HO_x radicals. However, the effect of the increase of NO_2 and the decrease of OH radical associated with the decrease of aerosol, is negligible, due to the offset effect by the decrease in the concentration of NO_2 when NO_x emissions are reduced. The enhancement in HO_2 radicals, contributed by aerosol decrease, is consistent with the weakened titration effect due to reduced NO_2 concentration by NO_x emission reduction and can favor the formation of secondary pollutants, including O_3 and SOA, in the VOC-limited areas. When combined with a decrease in AVOCs emissions, the increase of secondary pollutants, due to reduced NO_x , can be largely offset by the decrease in hydrocarbons and OVOCs. Exceptions can be found in the areas where high enough UV and temperature for the photochemical oxidation processes of VOCs are lacking, such as the northern part of China during wintertime, and where the role of BVOCs is crucial for the formation of ozone and secondary aerosols, including the PRD region during wintertime. A schematic diagram describing the chemical mechanisms involved in the atmospheric ozone chemical in response to the reduction in NO_x emissions is shown in Fig. 14.

3.43. Changes in the photochemical reactivity and the atmospheric oxidative capacity

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The changes in photochemical parameters, including the net production rate of odd oxygen ($O_3 \equiv O_3 + NO_2$), OH reactivity and the atmospheric oxidative capacity (AOC), in response to the adopted reductions in emissions are discussed in this section. The production rate of odd oxygen ($O_3 \equiv O_3 + NO_2$) is associated with recurrent radical reaction chains involving the oxidation of hydrocarbons in the presence of NO_x . The photochemical destruction of O_3 results from several processes, including the photolysis of O_3 , followed by the reaction between the electronically excited oxygen atom $O(^1D)$ and water vapor (H_2O). Other O_3 loss mechanisms include the reactions of ozone with OH, HO_2 , and different alkenes. In the presence of NO_x , additional O_3 losses are provided by the titration of O_3 by NO, followed by the conversion of NO_2 to nitric acid (HNO_3).

The OH reactivity (OH^R) is expressed as the OH loss frequency due to the reactions by VOC (VOC^R) (including CO) and NO_x (NO_x^R). Thus, this parameter can be used to represent the specific role of VOCs, NO_x , as well as CO that determine the photochemical formation of ground level ozone (Tan et al., 2017; Xue et al., 2016; Wang et al., 2022).

Atmospheric Oxidizing Capacity (AOC) is a parameter that characterizes the self-cleansing ability of the atmosphere (Liu et al., 2022). It is derived here as the rate at which CO, methane (CH_4), and non-methane hydrocarbons (NMHCs) are oxidized by atmospheric oxidants, including OH, O_3 , and NO_3 (Xue et al., 2016; Dai et al., 2023). This parameter allows us to characterize the formation process of removal process of secondary species, including O_3 , SIA, and SOA, and therefore, can be used as an indicator to design mitigation control policies for reducing for ozone pollution secondary pollutants. A detailed description of these parameters can be found in Part 1 of the paper (Dai et al., 2023).

3.4.1. OH reactivity

Figure 15 displays the daytime changes in OH reactivity for VOCs (VOC^R) in January and July in response to the specified emission reductions. During wintertime, not surprisingly, the response of VOC^R to the NO_x reduction is considerably smaller than the response to the reduction of AVOCs emissions. Reduction in the NO_x emission slightly reduces the value of VOC^R by $0.2-0.5 s^{-1}$ (<5%) in the YRD and PRD regions (Fig. 15a). The reduction in VOC^R resulting from the 50% reduction in the AVOCs emissions is found to be of the order of $1.5-3 s^{-1}$ (12-25%; Fig. 15c). In summer, the reduction in the NO_x emission increase VOC^R in the southern and northeastern parts of China ($1.5-2 s^{-1}$ (8-15%); Fig. 15b), which is consistent with the increase in the concentration of reactive VOCs species including alkenes (Fig. 6d) and isoprene (Fig. S11a) due to reduced oxidant. The potentially positive role of isoprene in the VOC^R in this case highlights the increasing importance of BVOCs in the summertime atmospheric oxidative capacity as the anthropogenic emissions are reduced. For reduced AVOCs emissions, a decrease in VOC^R of $2-3 s^{-1}$ (10-20%) is derived in the four major metropolitan regions of China, i.e. the NCP, YRD, PRD, and SCB (Fig. 15d).

Figure S15 shows the changes in the value of NO_x^R in response to the 50% reductions in NO_x emissions for the January and July conditions. The reduction in NO_x emissions leads to a decrease in NO_x^R in the range of $4.0-6.0 s^{-1}$ (30-50%) during January (Fig. S15a) and of $3.0-4.0 s^{-1}$ (35-45%) during July (Fig. S15b). The higher decrease in winter is consistent with the

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larger wintertime decrease in NO and NO_2 . When considering the reduction in AVOCs emissions, the changes in NO_x^R are relatively small (less than 3% in both January and July). When combining the emission reductions, the changes in the OH reactivity are substantially larger than individual impacts. For example, the combined reduction leads to a wintertime decrease in NO_x^R up to 8 s^{-1} (65%; Fig. S15e). For the 50% reduction in total anthropogenic emissions, the OH reactivity further decreases in urban areas, which is attributed to the decrease in the concentration of CO due to the relevant emission reduction (Fig. S16d).

For different sites, in winter, the largest decrease is found to occur at the Beijing site (Fig. 16a), with a decrease of 50% in OH reactivity in the combined 50% reductions in VOCs and NO_x emissions. In the combined case, the reduction in the OH reactivity is 45% in Shanghai, 40% in Chengdu, 35% in Wangdu, and 30% in Guangzhou. In summer, the highest decrease also occurs at the Beijing site (by 45%) when a 50% reduction is applied to all emissions. Under these conditions, the OH reactivity is reduced by about 30% in Guangzhou, 25% in Chengdu, and 15% in Shanghai. Besides, the contribution of BVOCs to OH^R becomes considerably larger, with the reduction in NO_x emissions. The most significant increase is calculated at the Heshan sites, with the contribution increasing from 18% to 34%. At the Beijing and Guangzhou sites, the contribution also increases from 3–5% to 10–15%. Small reductions of OH^R (less than 3%) are found at background sites (Waliguan), which suggests that the impact of anthropogenic emissions on the background OH reactivity is limited.

3.4.2. Atmospheric oxidative capacity

The changes in the spatial distribution of daytime (06:00 to 19:00 LST) AOC resulting from the adopted 50% reduction in the emissions of ozone precursors for January and July of 2018 is depicted in Fig. 8.

Winter conditions. In January, the 50% reduction in NO_x emission leads to a decrease in daytime AOC of 10~20% in southern China and an increase of 10~15% in the urban areas, including the PRD, YRD and SCB regions (Fig. 9a). With the 50% reduction in AVOCs emissions, the daytime AOC is reduced in all the major regions of China (Fig. 9b), with the largest decreases (30%) occurred in the southern part of the country. With a combined emission reduction in NO_x and AVOCs emission (Fig. 9c) and with the additional reduction of the other considered emissions (Fig. 9d), the distribution pattern of the changes in daytime AOC are similar with the pattern found in the AVOCs cases but derived with higher levels of decreases in daytime AOC.

Based on the changes in the daytime AOC at four city sites, as shown in Fig. 9, the increase in daytime AOC associated with the reduced NO_x emissions is caused by the enhanced contributions from OH-related reactions, including the reactions of OH with alkenes, followed by OH with OVOCs and aromatics. The reduction in AVOCs emissions leads to a large AOC decrease at four city sites, especially at the Guangzhou site (by 50%), followed by the Shanghai (by 48%) and Beijing sites (by 40%). This decrease in AOC is mainly attributable to the reduced

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1180 contribution from the reactions between OH with alkenes, followed by the reactions of OH
with aromatics and with OVOCs.

1185 During nighttime (20:00 to 05:00 LST), an increase in AOC of up to 50% results from the
reduction in NO_x emissions (Fig. S21a). The increases are contributed by alkenes ozonolysis,
with a relevant enhancement in the concentration of ozone (Fig. 6 a) and alkenes (Fig. S8c).
The largest increase in the alkene ozonolysis is derived at the sites of Shanghai from 31% to
40% (Fig. S22b). These results highlight the importance of ozonolysis associated with the NO_x
emissions reduction.

1190 Summer conditions. During summertime, the decrease in daytime AOC is more pronounced
than in wintertime. With the 50% reduction in NO_x emissions, AOC decreases in large areas
of China (ranging from 10% ~ 20%, Fig. 9e), while, in urban areas, a slight increase of 5% is
also predicted. In the cases of AVOCs, N+A_x, and TOTAL_x, the daytime AOC decreases in the
whole of geographical China (Fig. 9f-h), with more distinct decreases in North China than in
winter conditions. Finally, we show the changes in the spatial distribution of daytime AOC
resulting from the adopted 50% reduction in the emissions for the January and July conditions
in 2018. In January, the reduction in NO_x emission leads to an increase in AOC (by 0.6×10^7
molecules $\text{cm}^{-3} \text{s}^{-1}$ (18%); Fig. 17a) in the metropolitan areas of the YRD, PRD, and SCB and
a decrease in AOC in southern China (by 0.9×10^7 molecules $\text{cm}^{-3} \text{s}^{-1}$; 25%). The increase in
the daytime AOC of the metropolitan areas is consistent with the increase in the concentration
of surface O₃ (Fig. 10a) and SOA (Fig. 12a), which is related to the increase in the
concentration of oxidants, including OH and HO₂ radicals, ozone, and OVOCs.

1205 At the four city sites, during summertime, the reduced NO_x emissions also increase daytime
AOC. However, at the Beijing sites, the daytime value of AOC decreased, due to the shift in
ozone sensitivity regime of the site from VOC-limited to NO_x-limited. During nighttime, the
NO_x emissions reduction also leads to an increase in AOC due to the alkene ozonolysis (Fig.
S21b), with the largest increase derived at Beijing sites (from 10% to 14%, Fig. S22e). With
reduction in AVOCs emissions, a smaller decrease of daytime AOC than the decrease in winter
is derived, especially at Guangzhou site (to 30%). When the emissions of NO_x and AVOCs are
jointly reduced, an increasing role of the reaction between OH and BVOCs in the determination
of AOC is derived at four city sites, with the largest increase (15%) at the Guangzhou sites.
This increase results from the enhanced level of OH radicals and from the concentrations of
biogenic VOCs.

1215 When examining the wintertime impact of NO_x emission reduction at different sites, the largest
daytime AOC increase occurs at the Shanghai site (by 20%) (Fig. 18b), followed by the sites
of Guangzhou (by 18%), Chengdu (by 12%), and Beijing (by 8%). These results are consistent
with the wintertime increase of OH radical (Fig. 5e).

1220 With the reduction in the AVOCs emissions, the daytime AOC is reduced in all the major
regions of China (Fig. 17c). The large decreases occurred in the southern part of China (by 1.0

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1225 $\times 10^7$ molecules $\text{cm}^{-3} \text{s}^{-1}$; 30%), which is consistent with the large decrease of HO_x radicals and ozone due to reduced hydrocarbons and OVOCs.

1230 With the reduction in AVOCs emissions, the largest AOC decrease is found at the Guangzhou site (by 50%), followed by the Shanghai (by 48%) and Beijing sites (by 40%). The decrease in AOC is mainly attributable to the reduced contribution from the reactions between OH with alkenes, followed by OH with aromatics and OVOCs.

1235 With a 50% emission reduction in NO_x , AVOCs, and other primary pollutants (Fig. 17g), the level of daytime AOC is higher than the level with the decrease in NO_x and AVOCs emissions only (Fig. 17e). The higher level of daytime AOC is caused by the enhanced OH and HO_2 radicals (Fig. S9 c and g) resulting from less consumption by the reduced CO, SO_2 , and aerosol uptake.

1240 When examining the wintertime impact of NO_x emission reduction at different sites, the largest daytime AOC increase occurs at the Shanghai site (by 20%) (Fig. 18b), followed by the sites of Guangzhou (by 18%), Chengdu (by 12%), and Beijing (by 8%). These results are consistent with the wintertime increase of OH radical (Fig. 5e). With the reduction in AVOCs emissions, the largest AOC decrease is found at the Guangzhou site (by 50%), followed by the Shanghai (by 48%) and Beijing sites (by 40%). The decrease in AOC is mainly attributable to the reduced contribution from the reactions between OH with alkenes, followed by OH with aromatics and OVOCs.

1245 During summertime, the decrease in daytime AOC, resulting from the reduction in emissions, is more pronounced than in wintertime. With a 50% reduction in NO_x emissions, a decrease in AOC is derived in large areas of China (NO_x limited regimes), with a peak decrease of 1.5×10^7 molecular $\text{cm}^{-3} \text{s}^{-1}$ (20%; Fig. 17b) derived in the NCP region. At the same time, an increase of daytime AOC by 2.5×10^7 molecules $\text{cm}^{-3} \text{s}^{-1}$ (5%) is predicted in urban areas (VOC limited regimes). The spatial pattern of predicted changes in AOC due to reduced NO_2 is consistent with the changes in OH (Fig. 5f), HCHO (Fig. 6f), and O_3 (Fig. 10b), indicating the efficient representation of AOC in the formation of secondary pollutants. At specific sites, the simulated increase (by 5–7%) is derived in the sites of Guangzhou (Fig. 18e), Shanghai, and Chengdu.

1255 The decrease of daytime AOC due to the reduced AVOCs emissions is most pronounced in urbanized areas with values reaching 2×10^7 molecules $\text{cm}^{-3} \text{s}^{-1}$ (25%; Fig. 17d), in the NCP, YRD, and PRD. At four urban sites (Fig. 18), the decrease in AOC ranges from 15% to 30%.

1260 When combining the 50% reduction in AVOCs and NO_x emissions, the decrease in AOC is derived in the whole of geographical China (Fig. 17f). The decreases in different sites range from 10%–20% at the four city sites, and 16% and 20% at the rural sites; a reduction of around 5% is derived at the two remote sites. When adding the decrease in other emissions, the spatial distribution of daytime AOC decrease is not substantially different from the previous cases.

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1270 Interestingly, for nighttime (20:00 to 05:00 LST) AOC, the reduction in NO_x emissions results
in an increase of AOC by up to 2×10^6 molecules cm⁻³ s⁻¹ in both January (50%; Fig. 19a) and
1275 July (10%; Fig. 19b) due to the reaction of alkenes with O₃. The increases are mainly located
in urban China in winter and in southern and northern China in summer, which is associated
with the relevant enhancement in the concentration of ozone (Fig. 10 a and b) and summertime
alkenes (Fig. 6 d). Moreover, the most significant increase from the contribution of alkene
ozonolysis is also derived in city sites. For example, in January, the contribution of ozonolysis
to nighttime AOC increases from 31% to 40% in the Shanghai site (Fig. 20b; Fig. S17b) and
from 10% to 16% in the Chengdu sites (Fig. 20d; Fig. S17d). In summer, the changes are
smaller than in winter, with the increase from 10% to 14% in the Beijing site (Fig. 20a; Fig.
S17a).

1280 Notably, as shown in Fig. 18, among all the emission reduction scenarios, the largest decrease
in daytime AOC at urban sites is provided with a 50% reduction in AVOCs emission. To
mitigate the formation of secondary pollution, the reduction in AVOCs emissions needs to be
implemented. In practice, a reduction of 50% AVOCs emissions is hard to achieve in the short
term, thus efficient control over emissions of key AVOCs species is needed based on their
contributions to AOC. As shown in Fig. S18, at four metropolitan sites (Beijing, Shanghai,
1285 Guangzhou, and Chengdu), the largest contribution to daytime AOC is provided by the reaction
of OH with alkenes (30–44%), carbon monoxide (28–40%) and aromatics (15–25%) in winter.
In summer, a higher contribution from OH with OVOCs (30–40%) is derived at four urban
sites, in comparison to the wintertime conditions (12–22%). The contributions from OH with
alkenes and aromatics are around 15–30% and 10–20%, respectively. Combined with the
consideration of the dominant contribution of alkenes and aromatics to OH reactivity, we
1290 suggest that reducing the emissions of alkenes, aromatics, and unsaturated OVOCs should be
prioritized in the control of AOC and secondary pollutants in urban China. To give more
accurate suggestions for AVOCs emission control, more studies on the contribution of specific
AVOCs emissions to AOC are required.

1295 The aerosol effect on ozone formation has been discussed in several modeling studies (Li et
al., 2019; Liu et al., 2020; Dai et al., 2023). Our results show that the SIA concentration is
largely decreased in winter due to NO_x emissions reduction, while the declined SOA
concentration relies more on the reduction of AVOCs emissions. This aerosol decrease
weakens the aerosol extinction effect and therefore enhances the ozone photochemical
formation rate (Tan et al., 2022). At the same time, aerosol decreases can result in a decline in
1300 aerosol uptake through heterogeneous reactions, resulting in an enhanced concentration of NO₂
and HO₂ radicals. Moreover, the formation of HONO will also decline with NO_x emission
reduction (Fig. 6 a and b) through the heterogeneous conversion from NO₂ photolysis of NO₃⁻
and direct emission from transport (Dai et al., 2023). Thus, the atmospheric level of the OH
radical will also be reduced, as the photolysis of HONO is an important source of HO_x radicals.
1305 However, the effect of the increase of NO₂ and the decrease of OH radical associated with the
decrease of aerosol, is negligible, due to the offset effect by the decrease in the concentration
of NO₂ when NO_x emissions are reduced. The enhancement in HO₂ radicals, contributed by
aerosol decrease, is consistent with the weakened titration effect due to reduced NO₂
concentration by NO_x emission reduction and can favor the formation of secondary pollutants.

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1310 including O₃ and SOA, in the VOC-limited areas. When combined with a decrease in AVOCs
emissions, the increase of secondary pollutants, due to reduced NO_x, can be largely offset by
the decrease in hydrocarbons and OVOCs. Exceptions can be found in the areas where high
1315 enough UV and temperature for the photochemical oxidation processes of VOCs are lacking,
such as the northern part of China during wintertime, and where the role of BVOCs is crucial
for the formation of ozone and secondary aerosols, including the PRD region during
wintertime.

A schematic diagram describing the chemical mechanisms involved in the atmospheric ozone
chemical in response to the reduction in NO_x emissions is shown in Fig. 14.

1320 Notably, as shown in Fig. 18, among all the emission reduction scenarios, the largest decrease
in daytime AOC at urban sites is provided with a 50% reduction in AVOCs emission. To
mitigate the formation of secondary pollution, the reduction in AVOCs emissions needs to be
implemented. In practice, a reduction of 50% AVOCs emissions is hard to achieve in the short
1325 term, thus efficient control over emissions of key AVOCs species is needed based on their
contributions to AOC. As shown in Fig. S18, at four metropolitan sites (Beijing, Shanghai,
Guangzhou, and Chengdu), the largest contribution to daytime AOC is provided by the reaction
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1330 In summer, a higher contribution from OH with OVOCs (30–40%) is derived at four urban
sites, in comparison to the wintertime conditions (12–22%). The contributions from OH with
alkenes and aromatics are around 15–30% and 10–20%, respectively. Combined with the
consideration of the dominant contribution of alkenes and aromatics to OH reactivity, we
suggest that reducing the emissions of alkenes, aromatics, and unsaturated OVOCs should be
1335 prioritized in the control of AOC and secondary pollutants in urban China. To give more
accurate suggestions for AVOCs emission control, more studies on the contribution of specific
AVOCs emissions to AOC are required.

A schematic diagram describing the chemical mechanisms involved in the atmospheric ozone
chemical in response to the reduction in NO_x emissions is shown in Fig. 14.

1340 The distribution pattern of changes in daytime AOC due to emissions reduction is almost
consistent with the relevant changes in the mixing ratio of OH radicals and the concentration
of OVOCs, ozone, and SOA in both winter and summer. These consistent patterns suggest that
the AOC is an appropriate indicator to characterize the changes in secondary pollutants
1345 attributed to emissions reduction. One exception is found when considering the changes in the
ozone concentration resulting from the reduction in the NO_x emission during winter. During
this season, a comparison between the values of AOC with the changes in the ozone
concentration (Fig. 5a) shows that the change in daytime AOC is reflecting primarily the
changes in the net production rate of odd oxygen (Fig. S23); this can be explained by the
1350 important role played by NO₂ in the wintertime formation of ozone.

3. Summary and Policy Implications

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1355 In this study, several model simulations have been conducted to explore the response of radicals, of ozone, and of the atmospheric oxidative processes to a 50% reduction in primary emissions of key pollutants. Our analysis provides insight into the changes in ozone chemistry and in the oxidizing processes to be expected in China in response to future emission reduction.

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1360 In winter, as most areas are VOC-limited (saturated in NO_x) a 50% reduction in NO_x emissions leads to an ozone concentration increase of up to 8-10 ppbv (15-20%) in all geographical regions of China; this increase results from of the reduced titration of ozone by nitric oxide. When combining this NO_x reduction with a 50% reduction in AVOCs emissions, the ozone enhancement found in the rural areas and resulting from the reduced NO_x is considerably reduced. However, in urban areas (VOC-limited), the ozone increase, although weakened, still exists (by 3.0-7.5 ppbv).

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1370 In summer, as most rural areas become NO_x-limited, the geographical regions covered by the ozone increase in response to the 50% emission reduction in NO_x shrink and cover primarily the VOC-limited metropolitan areas. In these urban environments, the ozone increase reaches a maximum of 10 ppbv or 17%. When the NO_x emission reduction is combined with a 50% VOC emission reduction, the increase in ozone almost disappears in all areas of China. This is explained by the significant decrease in ozone production resulting from the reduced level of hydrocarbons. However, in the rural areas where hydrocarbons are dominated by the emissions of biogenic VOCs, the ozone concentration (linked to the photochemical degradation of isoprene) still slightly increases. Our analyses of the model results provide insight into the changes in the atmospheric oxidizing capacity and chemistry in China in response to reductions of anthropogenic emissions.

1380 Paths to mitigation. We conclude this paper by highlighting a few chemical paths that should be considered when designing a mitigation policy for a reduction of ozone in urban areas of China. Figure 10 presents a schematic description of the chemical mechanisms involved in the chemical production of atmospheric ozone and highlights how different reaction paths tend to change the ozone abundance in urban areas in response to a reduction in NO_x and anthropogenic VOC (AVOCs) emissions. This graph shows that a reduction in NO_x emissions tends to increase the ozone concentration by (1) reducing the rate of the NO + O₃ reaction (ozone titration); (2) by increasing the rate of the HO₂ + NO reaction due to an increase in the HO₂ level associated with the reduced uptake of this radical by a lowered aerosol load; (3) by an increase in the atmospheric oxidizing capacity (AOC) through the OH- and ozone-related reactions. The graph also shows that a decrease in AVOCs emissions tends (1) to reduce the level of the HO_x radical and hence the ozone production by the HO₂ + NO reaction; (2) to enhance the level of HO_x due to the reduced aerosol uptake and (3) to reduce the AOC with a negative effect on the ozone concentration. The relative importance of these different chemical mechanisms varies with location and environmental conditions.

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1400 We conclude that in winter when the background concentration is low, the reduction of NO_x emissions tends to increase the ozone concentration, while the reduction in AVOC emissions has the opposite effect. This conclusion applies both in rural and in urban areas. A combined reduction in the emissions of these two primary pollutants tends to decrease the level of ozone in rural areas, but to increase ozone in urban areas. Thus, in urban areas during winter, an effective approach to reduce the surface ozone concentration is through a strong limitation in the emissions of volatile organic compounds.

1405 In summer when the ozone level is generally high, the reduction of NO_x emissions is an effective action to reduce the ozone concentration in rural areas, but this measure is counterproductive in urban areas where ozone is controlled by VOCs. In fact, in urban areas during this season, the mechanisms involved in ozone mitigation are complex. For example, when NO_x emissions are reduced, the atmospheric OH concentration is enhanced because of its reduced destruction by NO₂. Following this increase in the OH concentration, an increase in the level of OVOCs, whose photolysis is an important source of HO_x radicals, also leads to accelerated ozone production and further amplifies the oxidation of VOC. In addition, the increase in AOC, linked to the reaction of OH and ozone with alkenes and the reactions of OH with OVOCs also contributes to an increase in ozone production. Further, the reduction in the aerosol load resulting from a reduction in the emissions of aerosol precursors promotes ozone formation by decreasing the aerosol extinction and by reducing the uptake of HO₂. If combined with a 50% reduction in AVOCs, the increase in OVOCs and AOC, due to reduced NO_x emissions, can be offset. However, the aerosol-related promotion in the level of OH and HO₂ radicals can even be exaggerated, indicating the complexity in mitigating summertime ozone pollution in urban areas.

1425 Table 2 provides quantitative information on the response of ozone at different urban locations for January and July. In short in urban areas, the reduction in the level of surface ozone requires a reduction in the emissions of anthropogenic VOCs. However, in practical terms, a 50% reduction in AVOCs emissions, as assumed in our study, is difficult to implement over a short period of time. With the known contribution of the VOCs-related reactions to the AOC, the reduction in the emissions of alkenes, aromatics, and unsaturated OVOCs, especially the aldehydes and alcohols, should be a priority. The development of efficient mitigation strategies based on the reduction of AVOCs emissions requires, however, more detailed investigations on the reactivity of individual VOCs and on their potential impact on the ozone formation.

1435 With a 50% reduction in NO_x emissions, the production rate of radicals and odd-oxygen decreases substantially in both January and July. These decreases are larger in summer (by 0.5–1.2 ppbv h⁻¹ (12–28%) and 1.5–2.0 ppbv h⁻¹ (20–50%)) than in winter (by 0.3–0.7 ppbv h⁻¹ (30–33%) and 1.5–2.0 (15–20%) ppbv h⁻¹) and can be attributed to the reduction in the concentrations of nitrogen oxides and OH radicals. During wintertime, however, an increase by 0.8–1.5 ppbv h⁻¹ (7–15%) is derived to produce odd-oxygen in the PRD region, which is related to the increasing concentration of the HO₂ radical resulting from the increase in the OVOCs concentration and the reduction in the aerosol load. Moreover, during summertime, the simulated concentration of surface ozone decreases in NO_x-limited areas (up to 2–8 ppbv

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(3–12%) and increases in VOC limited areas (up to 12 ppbv (15%)). The increasing ozone concentration in VOC limited areas is attributed to the reduced titration by NO_x decrease. In the NO_x limited areas, the decrease in the ozone concentration results from the reduced photochemical ozone formation associated with the reduction in the concentration of NO_x with considerate effects from reduced HO_x radicals, due to less formation from the photolysis of HONO and OVOCs.

With the AVOCs emissions lowered by 50%, a lower concentration of alkenes by up to 4.0 ppbv (40%) in January and by 2.0 ppbv (50%) in July is derived in urban China areas. The concentrations of OVOCs are reduced by 2.0–5.0 ppbv (8–20%) in January and by 3.0–5.0 ppbv (12–20%) in July due to the reduced primary emissions and secondary formations from the oxidation of hydrocarbons. Results show that, with a 50% reduction in AVOCs emissions, the RO_x and O_x production rates also reduce; the decreases reach 0.1–0.8 ppbv h⁻¹ (10–35%) and 1.0–4.5 ppbv h⁻¹ (20–42%), respectively, during winter, and 0.3–0.8 ppbv h⁻¹ (7–18%) and 5.0–10.0 ppbv h⁻¹ (10–28%), respectively, during summer. It is attributed to a decrease in the concentration of HO₂ radicals, associated with a decrease in the oxidation of hydrocarbons and photolysis of OVOCs. The simulated ozone concentrations are reduced by 2.0–8.0 ppbv (4–15%) in southern China during winter and by 8.0–12.0 ppbv (9–20%) in urbanized areas during summer.

With the combined 50% reduction in NO_x and AVOCs emissions, the geographic patterns of the simulated ozone changes in January resemble more the patterns associated with NO_x emission reduction than those with AVOCs emission reduction. Generally, the surface ozone concentration increases by 5–12 ppbv (6–15%) in urbanized areas. In summer, the changes in the ozone concentrations in VOC limited areas are affected positively by the reduction in NO_x emissions and negatively by the reduction in AVOCs emissions. The net effect is an ozone reduction in almost all areas of China except the PRD region, where the role of biogenic VOCs emissions is considered to be comparable to or more important than anthropogenic emissions.

For aerosol response to primary emissions reduction, a large decrease (up to 20%) in SIA concentrations is derived with a 50% NO_x emissions reduction in winter, dominantly contributed by decreases in NO₃⁻, followed by NH₄⁺. With a 50% reduction in AVOCs

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1485 emissions, the simulated SIA decrease is limited (less than 5%) during wintertime. However,
when combining the emissions reduction in NO_x with AVOCs, a strengthened effect can be
1490 found in the decrease of each composition of SIA (NO_3^- , NH_4^+ , and SO_4^{2-}), indicating the
necessity of a joint reduction in NO_x and AVOCs emissions in an efficient reduction in the
aerosol concentration in the future. The decrease of SOA is more sensitive to AVOCs
1495 emissions reduction, as the oxidation of several AVOCs is an important source of SOA
formation. A geographical shift can be found in the spatial distribution of SOA decrease, with
large decreases in southern China during wintertime and in the NCP region during
summertime. This shift is consistent with the seasonally varied distribution of HO_x radicals.
With the aerosol decrease, the effect of the associated enhancement of HO_2 will be consistent
1500 with the weakened nitration effect by NO_x emission reduction and lead to an enhancement in
secondary pollutants. When combined with the reduction in AVOCs, the positive effect of
aerosol enhanced HO_2 will still be important to ozone formation in northern China during
wintertime and PRD regions during summertime, due to the relatively small influence of
AVOCs.

1505 Regarding the OH reactivity from VOCs in urban areas, we find that a 50% decrease in the
 NO_x emission leads to a decrease of $0.2\text{--}0.5\text{ s}^{-1}$ (<5%) and $0.6\text{--}3.0\text{ s}^{-1}$ (5–20%) in January and
July, respectively. During summertime, an increase of 3.0 s^{-1} of VOC-related OH reactivity by
up to 3.0 s^{-1} is found in southern and northeastern China. This increase is associated with an
1510 increase in the concentrations of alkenes and biogenic VOCs (e.g. isoprene), due to less
consumption of these reactive VOCs by the reduced OH radical. Regarding the OH reactivity
from NO_x , a slight increase is derived during wintertime in the specific case where AVOCs
emissions are reduced by 50%. This increase is attributed to the increase in the calculated NO_2
concentrations.

1515 For atmospheric oxidative capacity (AOC), a reduction in NO_x emissions during January leads
to a slight decrease ($0.2\text{--}0.4 \times 10^7\text{ molecules cm}^{-3}\text{ s}^{-1}$; 3–6%) in the daytime AOC value of
southern China and a distinct increase ($1.0 \times 10^7\text{ molecules cm}^{-3}\text{ s}^{-1}$; 18%) in the PRD region,
due to the enhanced HO_x radical from the increase alkenes ozonolysis and OVOCs photolysis.
In July, AOC decreases by $0.4\text{--}1.5 \times 10^7\text{ molecules cm}^{-3}\text{ s}^{-1}$ (7–20%) in the NO_x -limited areas
1520 and increases by $2.0 \times 10^7\text{ molecules cm}^{-3}\text{ s}^{-1}$ (25%) in the VOC-limited areas when NO_x
emissions are reduced by 50%. These changes are linked to the corresponding changes in the
levels of OH radicals and ozone. A reduction in the AVOCs emissions leads to a reduction in
the daytime AOC by $1.0 \times 10^7\text{ molecules cm}^{-3}\text{ s}^{-1}$ (13%) in January and $2.0 \times 10^7\text{ molecules cm}^{-3}\text{ s}^{-1}$
(25%) in July. Specifically, the summertime decreases result predominantly from the
weakening of the reaction between OH and alkenes, followed by the reaction of OH with
aromatics and OVOCs.

1525 Our results suggest that a substantial reduction in the NO_x emissions and associated effects in
the aerosols decrease, the photolysis of OVOCs and ozonolysis of alkenes will lead to an
enhancement in surface ozone concentration, OH reactivity from VOCs and AOC parameters
in urbanized areas. A coordinated strategy for the reduction in NO_x and AVOCs emissions is
required to efficiently reduce the ozone levels in metropolitan areas. More detailed

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1530 ~~investigations to characterize the contribution of individual VOCs to summertime ozone~~
~~formation are required to develop efficient mitigation strategies against rising ozone~~
~~concentrations and the related change in the oxidative capacity of the atmosphere in China.~~

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1545 *Code and data availability.* The WRF-Chem model is publicly available at
<https://www2.mmm.ucar.edu/wrf/users/>. The air quality data at surface stations are publicly
available at the website of the Ministry of Ecology and Environment of the People's Republic
of China at <http://english.mee.gov.cn/>.

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1550 *Author contributions.* JD and GB designed the structure of the manuscript, performed the
numerical experiments, analyzed the results, and wrote the manuscript. JD analyzed the data
and established the figures. All co-authors provided comments and reviewed the manuscript.

1555 *Competing interests.* The authors declare that they have no conflict of interest.

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Table 1. Sensitivity experiments

Model Experiments	Description
<i>BASE</i>	Without emission reduction
<i>NO_x</i>	With emission reduction in NO _x ^a by a factor of 2

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AVOCs With emission reduction in anthropogenic VOCs^a by a factor of 2

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N+A With emission reduction in NO_x^a and anthropogenic VOCs^a by a factor of 2

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TOTAL With emissions reduction in all species^a by a factor of 2

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^a Relevant species is shown in Table S1.

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Table 2. Ozone changes due to reduction in emissions (in percentage).

<u>Site type</u>	<u>Location</u>	<u>Sites name</u>	<u>Ozone changes in winter condition (Mean ± SD)</u>			
			<u>NO_x^a</u>	<u>AVOCs^b</u>	<u>N+A^c</u>	<u>TOTAL^d</u>
<u>Urban sites</u>	<u>North</u>	<u>Beijing</u>	<u>35.0 ± 25.2^e</u>	<u>-2.5 ± 1.3</u>	<u>40.0 ± 32.8</u>	<u>30.0 ± 19.5</u>
	<u>East</u>	<u>Shanghai</u>	<u>63.6 ± 35.3</u>	<u>-18.2 ± 13.5</u>	<u>31.8 ± 20.5</u>	<u>32.7 ± 18.8</u>
	<u>South</u>	<u>Guangzhou</u>	<u>31.4 ± 22.6</u>	<u>-17.1 ± 11.2</u>	<u>7.1 ± 3.2</u>	<u>10.0 ± 3.5</u>
	<u>West</u>	<u>Chengdu</u>	<u>31.3 ± 23.8</u>	<u>-9.4 ± 8.5</u>	<u>14.1 ± 8.3</u>	<u>20.3 ± 13.5</u>
<u>non-Urban sites^f</u>	<u>North</u>	<u>Wangdu</u>	<u>45.0 ± 35.0</u>	<u>-2.5 ± 1.8</u>	<u>40.0 ± 28.8</u>	<u>35.0 ± 28.9</u>
	<u>South</u>	<u>Heshan</u>	<u>34.3 ± 28.2</u>	<u>-24.3 ± 19.2</u>	<u>5.7 ± 4.2</u>	<u>7.1 ± 5.8</u>

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Site type	Location	Sites name	Ozone changes in summer condition (Mean ± SD)			
			NO _x	AVOCs	N+A	TOTAL
Urban sites	North	Beijing	6.4 ± 3.8	-21.8 ± 19.2	-5.5 ± 4.2	-7.3 ± 5.0
	East	Shanghai	17.1 ± 12.8	-22.9 ± 20.8	-2.9 ± 2.1	-2.6 ± 1.5
	South	Guangzhou	15.0 ± 13.1	-14.5 ± 13.5	1.3 ± 1.0	1.3 ± 0.9
	West	Chengdu	5.5 ± 4.5	-14.5 ± 10.2	-5.5 ± 2.0	-4.5 ± 1.9
non-Urban sites	North	Wangdu	2.0 ± 1.2	-17.0 ± 15.8	-5.6 ± 4.2	-7.0 ± 5.2
	South	Heshan	-3.3 ± 2.8	-13.3 ± 10.9	-6.7 ± 3.9	-16.7 ± 3.9

a-d. Sensitivity cases with a 50% reduction in NO_x emissions (NO_x), AVOCs emissions (AVOCs), NO_x and AVOCs (N+A), and other species (NO_x, AVOCs, CO, NH₃, SO₂) under consideration (TOTAL).

e. values are displayed in the average ozone changes during daytime (06:00-19:00) in percentage with the standard deviation as the error bar. (ozone changes = (case value -baseline case) *100).

f. two non-Urban sites in North (Wangdu) and South China (Heshan) were selected, detailed information of these two sites can be found in the Part 1 of the paper (Dai et al., 2023).

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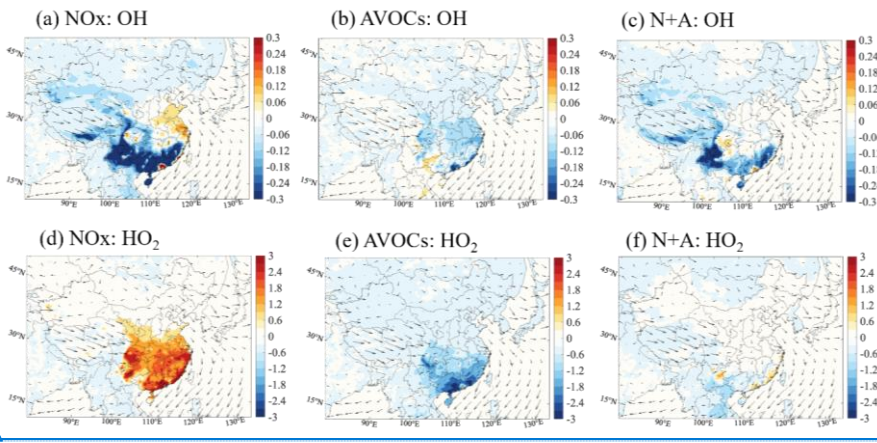
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Figure 1. Changes in the averaged daytime surface mixing ratio of OH radical (a-c, Unit: 0.1 pptv) and HO₂ radical (d-f, Unit: pptv) response to the a 50% reduction in NO_x emissions (a, d; NO_x case), in anthropogenic VOCs (AVOCs) emissions (b, e; AVOCs case) and in NO_x and AVOCs emissions (c, f; N+A case) relative to BASE case for January of 2018. Arrows represent the wind speed and wind direction.

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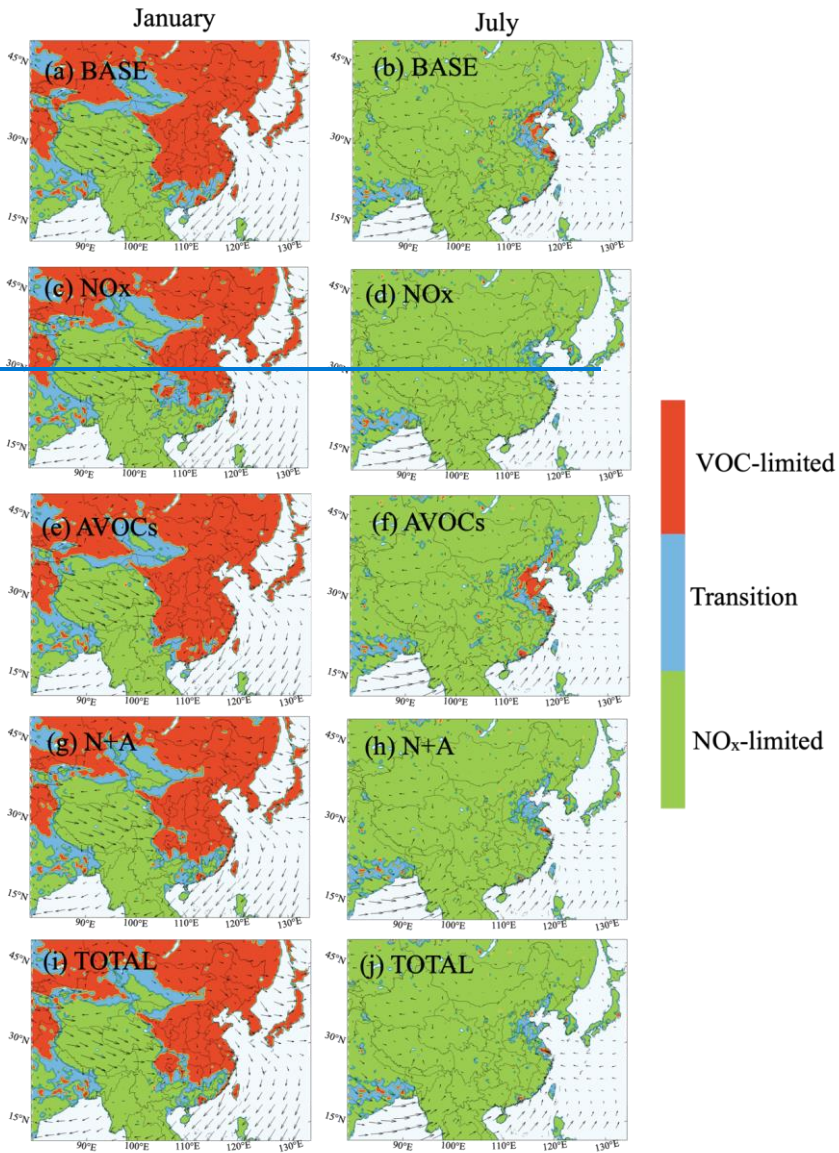
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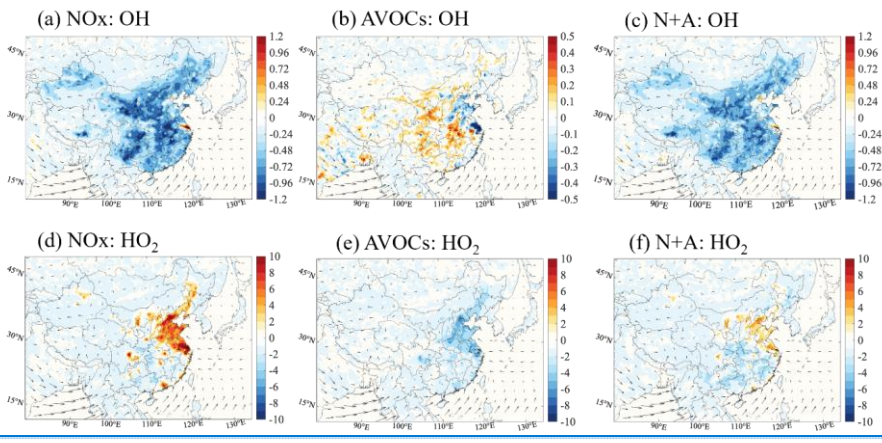


Figure 2. Same as Fig.1, but for July of 2018. Notice the inconsistency in the scale of Figure 2b.

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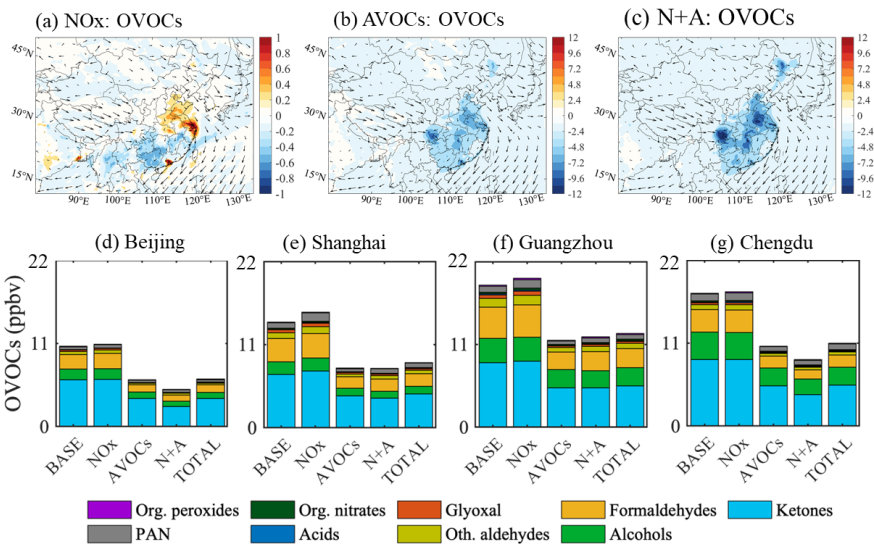


Figure 3. Changes in the surface mixing ratio of total oxidized VOCs (OVOCs) for January of 2018. (a-c) Changes in total OVOCs (Unit: ppbv) response to the reduction in NO_x, AVOCs, and combined NO_x and AVOCs emissions relative to the BASE case. (d-g) Contributions of averaged concentration of OVOC by different species at four city sites (Beijing, Shanghai, Guangzhou, and Chengdu) in China (c-f) in five cases (BASE, NO_x, AVOCs, N+A, and TOTAL cases). Arrows in a-c represent the wind speed and wind direction. Notice the inconsistency in the scale of Figure 3a.

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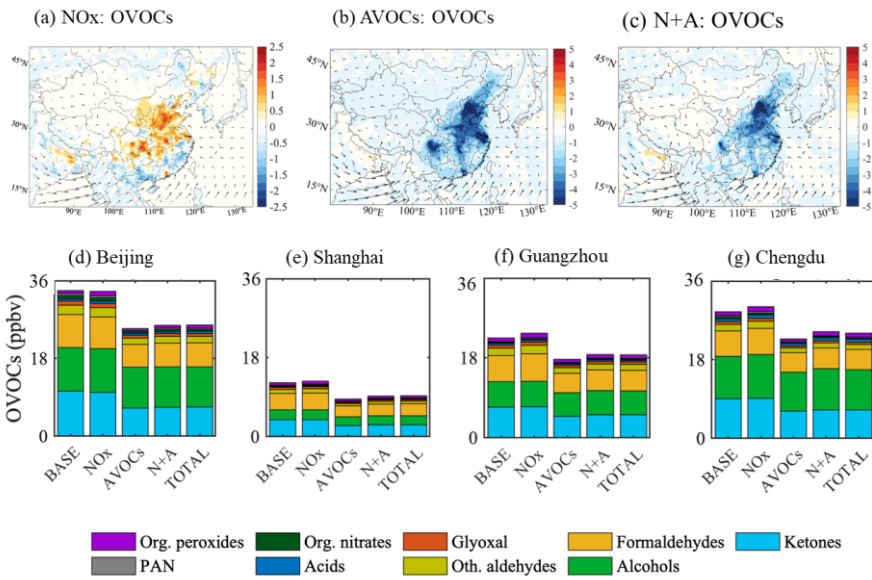


Figure 4. Same as Fig.3 but for July of 2018. Notice the inconsistency in the scale of Figure 4a.

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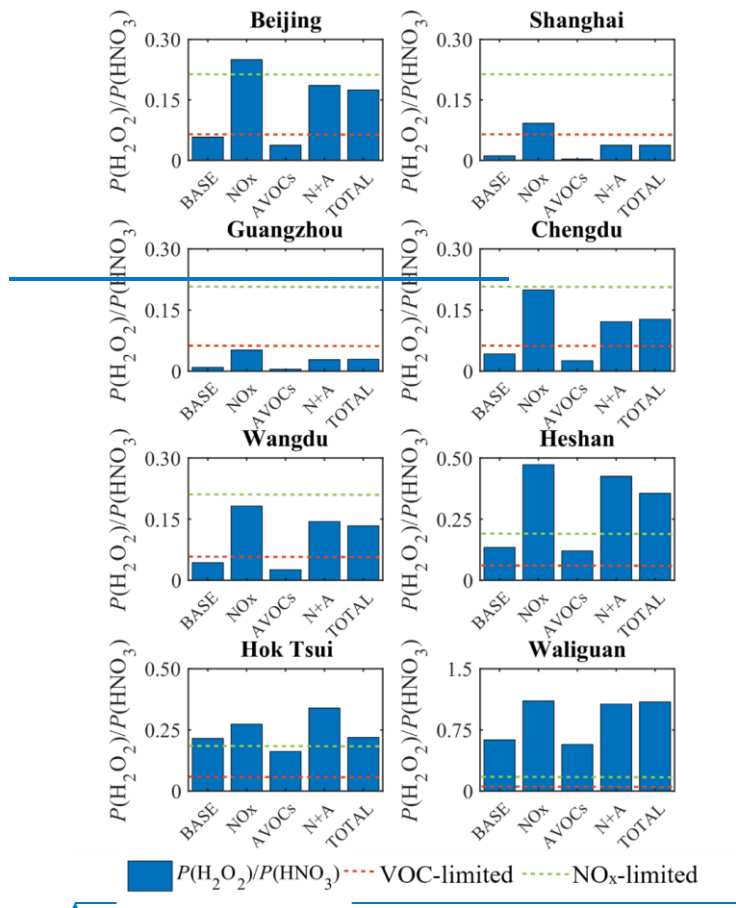
1920

Figure 1. Display of regions in which ozone production is limited by the availability of nitrogen oxides (NO_x -limited, in green), and volatile organic components (VOC limited, in red) from the BASE case (a, b), the NO_x case (c, d), the AVOCs case (e, f), the N+A case (g, h), and the TOTAL case (i, j) in January (a, c, e, g, i) and July (b, d, f, h, j) of 2018. The regions where ozone production is controlled by the availability of both NO_x and VOCs (transition) are shown in blue. The indicator used to define these regions is the ratio between the production rate of hydrogen peroxide (H_2O_2) and nitric acid (HNO_3).

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1930

Figure 2. The daytime (06:00 to 19:00 Local Standard Time (LST)) value of ratio between the production rate of hydrogen peroxide (H_2O_2) and nitric acid (HNO_3) [$P(\text{H}_2\text{O}_2)/P(\text{HNO}_3)$] in different regions of China for July 2018. The value of [$P(\text{H}_2\text{O}_2)/P(\text{HNO}_3)$] below the dotted line in red (0.06), above the dotted line in green (0.2), and in between represents the control by VOCs, NO_x , and in transition.

1935

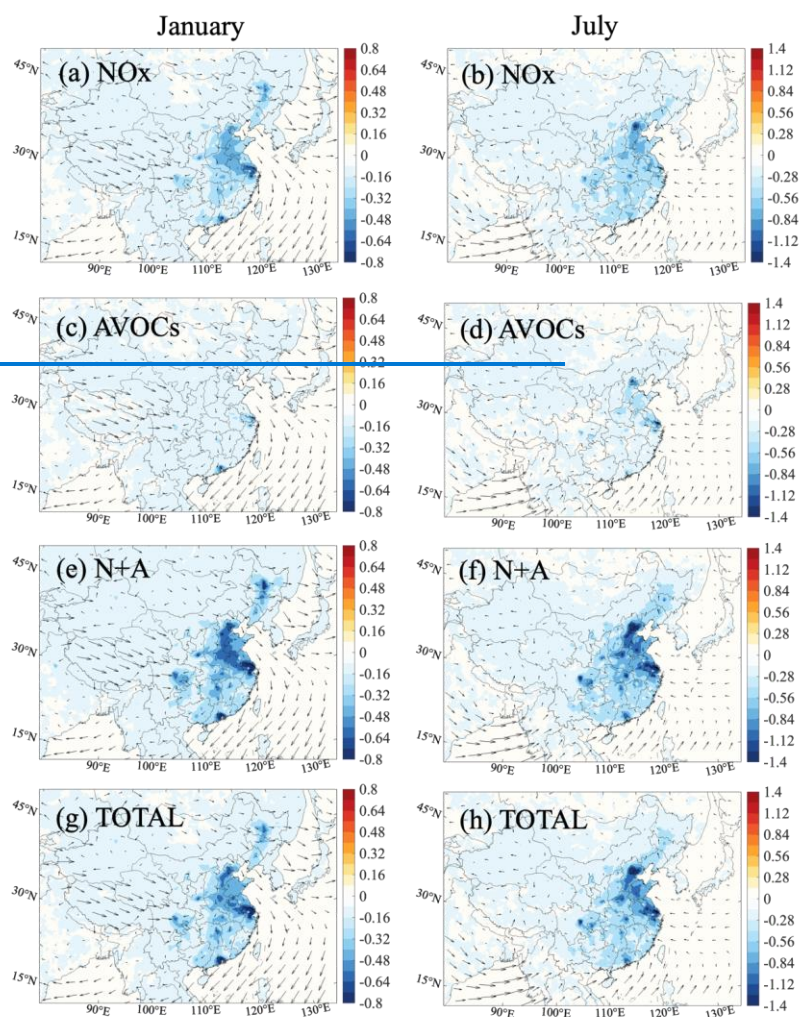


Figure 3. Changes in the averaged daytime surface production rate of RO_x (RO_2+HO_2+OH) [Unit: $ppbv\ h^{-1}$] response to a 50% reduction in NO_x emissions (a, b; NO_x case), in anthropogenic VOCs (AVOCs) emissions (c, d; AVOCs case), in NO_x and AVOCs emissions (e, f; N+A case), and in all anthropogenic emissions (g, h, TOTAL case) relative to BASE case. Results are shown for January (a, c, e, g) and July (b, d, f, h) of 2018. Arrows represent the wind speed and wind direction.

1940

1945

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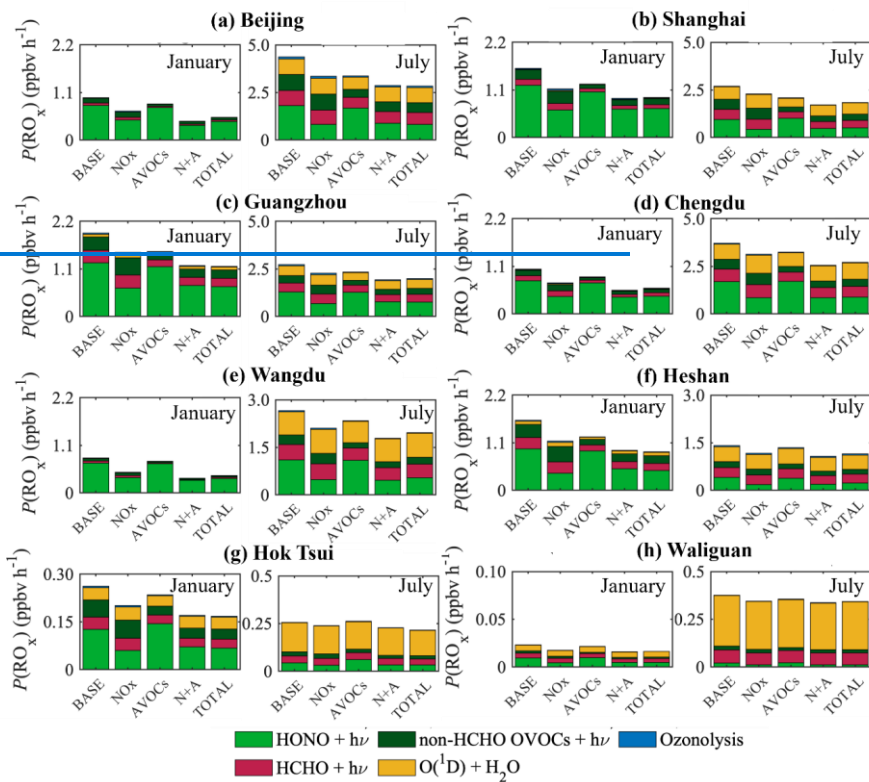


Figure 4. Averaged daytime value of production rate of RO_x ($P(RO_x)$) [Unit: $ppbv\ h^{-1}$] in five different simulated cases (BASE, NO_x, AVOCs, N+A, TOTAL cases) and eight different sites (urban, rural, and remote sites) in January and July of 2018.

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1955

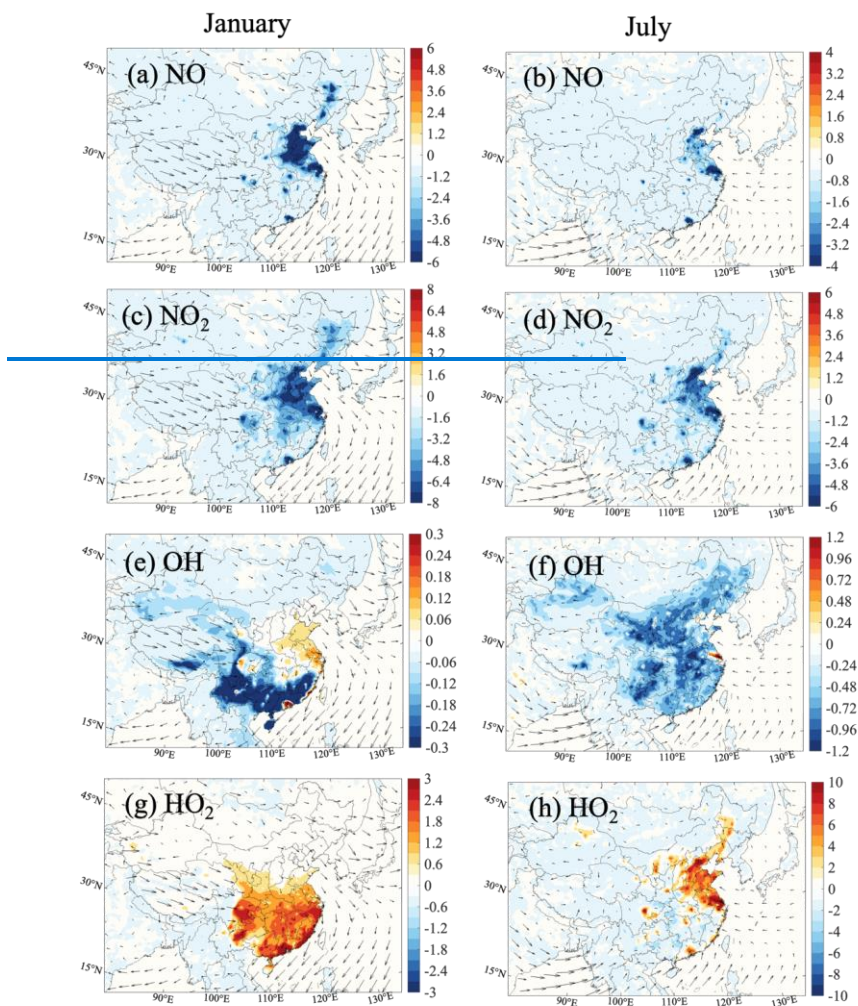


Figure 5. Changes in the surface mixing ratio of NO (a, b) [Unit: ppbv], NO₂ (c, d) [Unit: ppbv], OH radical (e, f) [Unit: 0.1 pptv] and HO₂ radical (g, h) [Unit: pptv] response to the NO_x case relative to BASE case. The results are shown for January (a, c, e, g) and July (b, d, f, h) of 2018. Arrows represent the wind speed and wind direction.

1960

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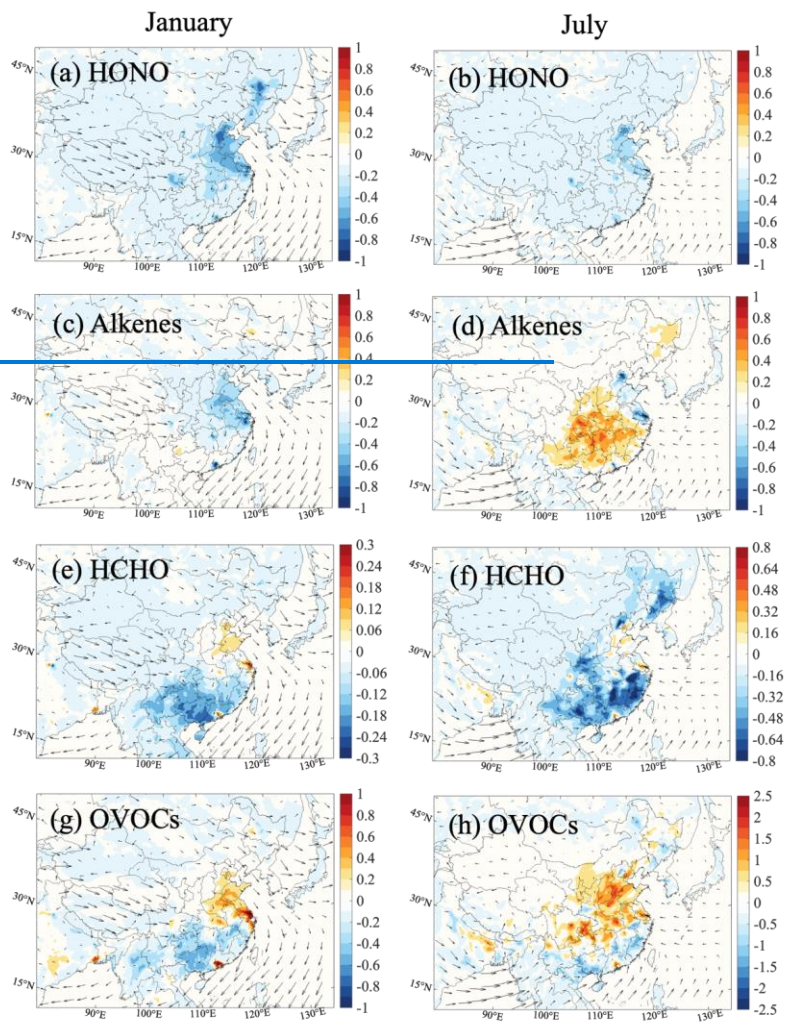


Figure 6. Changes in the surface mixing ratio of HONO (a, b) [Unit: ppbv], alkenes (c, d) [Unit: ppbv], formaldehyde (HCHO; e, f) [Unit: ppbv] and total oxidized VOCs (OVOCs; g, h) [Unit: ppbv] response to the NO_x case relative to the BASE case. The results are shown for January (a, c, e, g) and July (b, d, f, h) of 2018. Arrows represent the wind speed and wind direction.

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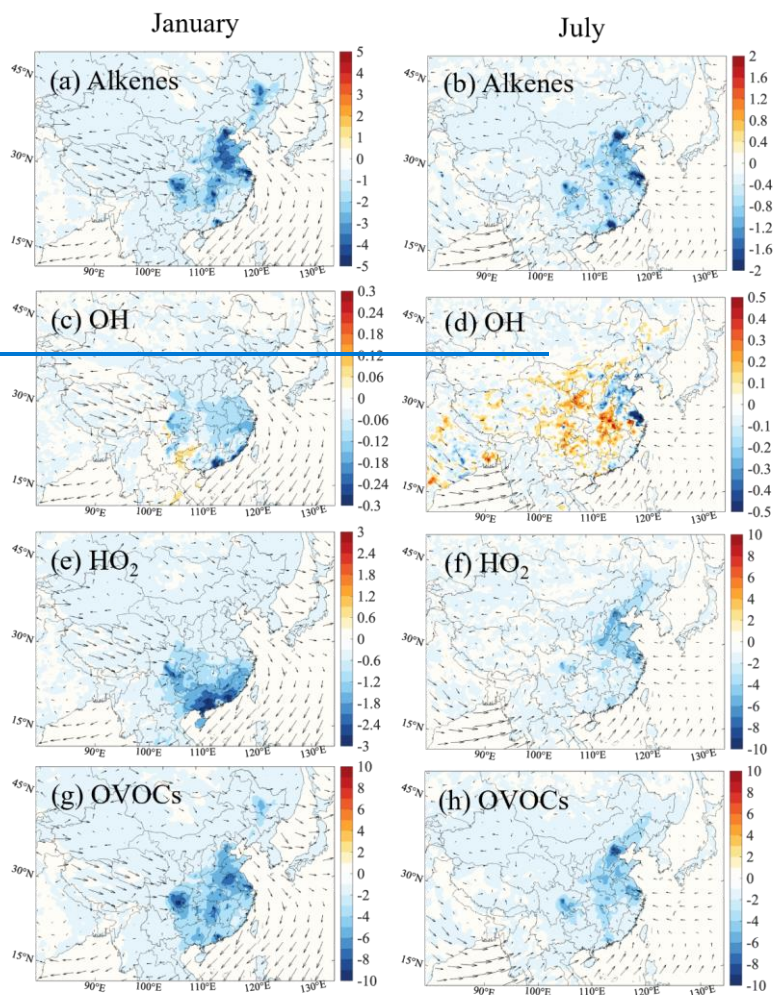


Figure 7. Changes in the surface mixing ratio of alkenes (a, b) [Unit: ppbv], OH radical (c, d) [Unit: 0.1 pptv], HO₂ radical (e, f) [Unit: pptv], and OVOCs (g, h) [Unit: pptv] response to the ratio of 0.5 in AVOCs emissions (AVOCs case) relative to BASE case. Results are shown for January (a, c, e, g) and July (b, d, f, h) of 2018. Arrows represent the wind speed and wind direction.

1980

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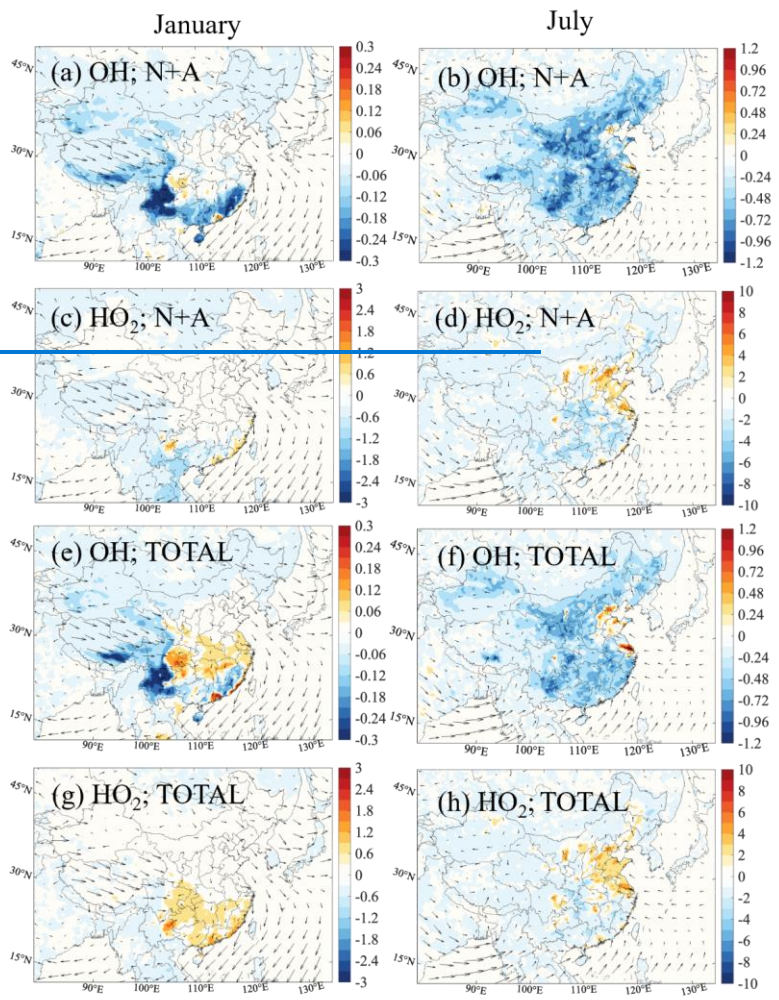


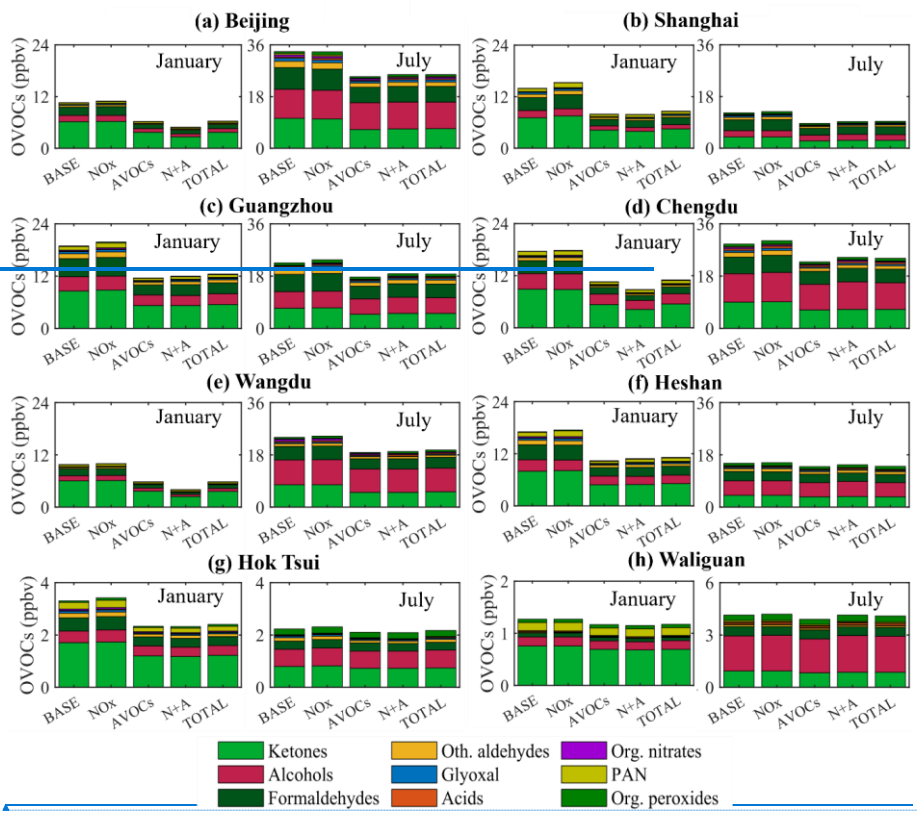
Figure 8. Changes in the surface mixing ratio of OH radical (a, b, e, f) [Unit: 0.1 pptv] and HO₂ radical (c, d, g, h) [Unit: pptv] response to the ratio of 0.5 in NO_x and AVOCs emissions (N+A case) and in all anthropogenic emissions (TOTAL case) relative to BASE case. Results are shown for January (a, c, e, g) and July (b, d, f, h) of 2018. Arrows represent the wind speed and wind direction.

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Figure 9. Averaged mixing ratio of oxidized VOCs (OVOCs; Unit: ppbv) with the contribution from nine types of species in five simulated cases (BASE, NO_x, AVOCs, N+A, TOTAL cases) and at eight sites (urban, rural, and remote sites) in January and July of 2018.

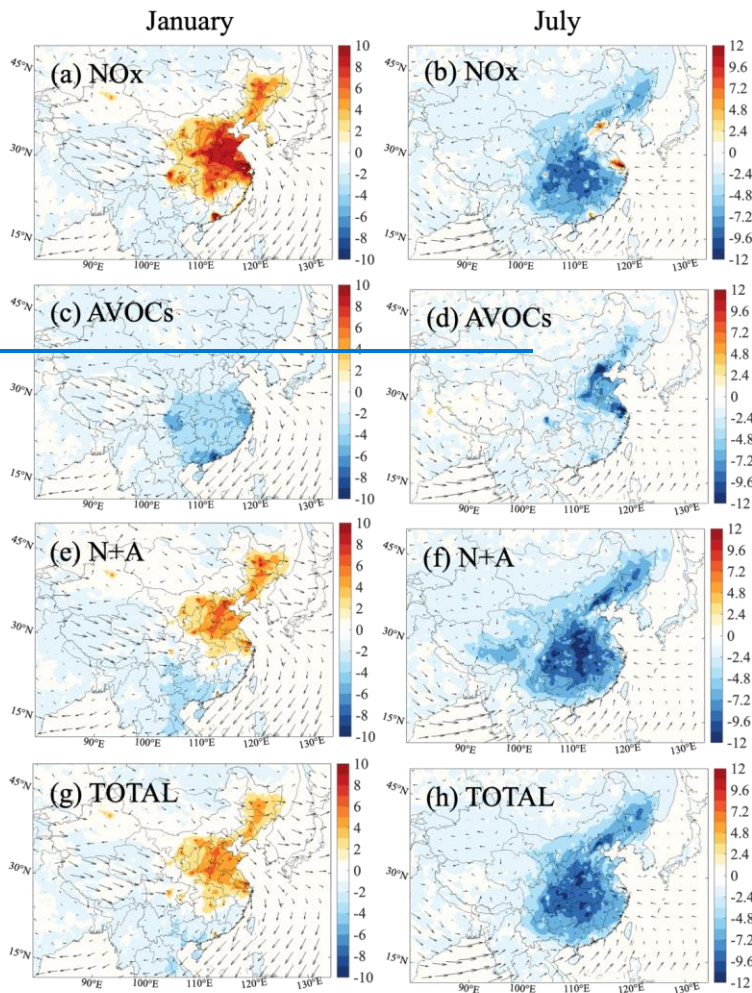
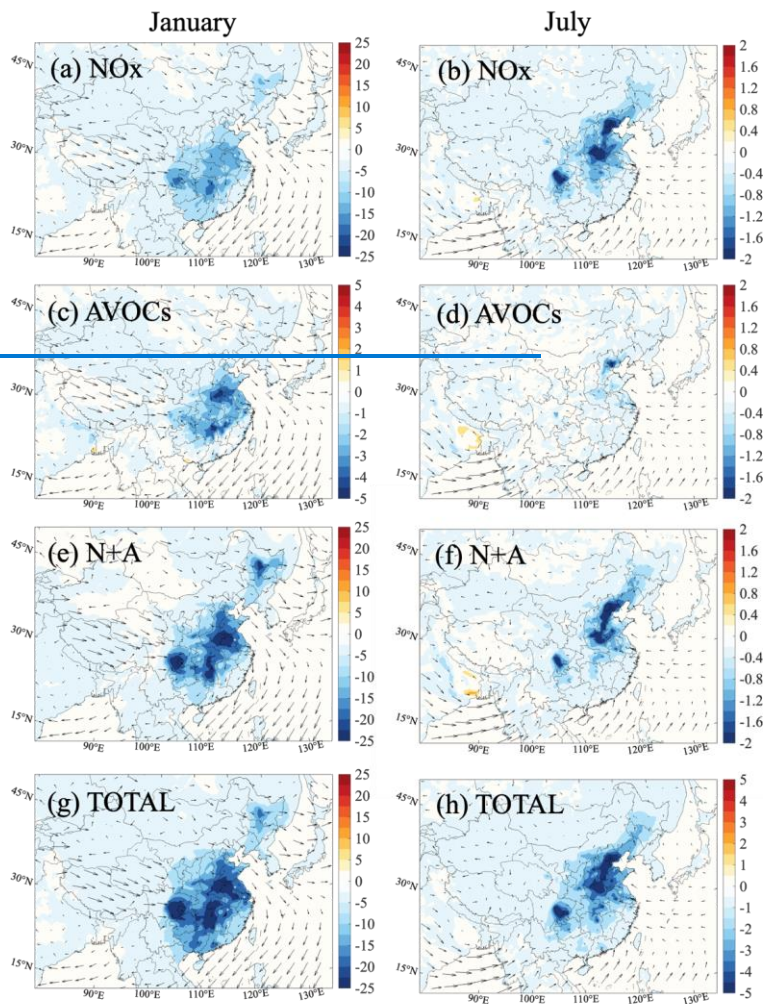


Figure 10. Changes in the averaged daytime surface mixing ratio of ozone [Unit: ppbv] response to the NO_x case (a, b), AVOCs case (c, d), N+A case (e, f), and TOTAL case (g, h) relative to BASE case. The results are shown for January (a, c, e, g) and July (b, d, f, h) of 2018. Arrows represent the wind speed and wind direction.

2010

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2015 Figure 11. Changes in the surface concentration of Secondary Inorganic Aerosol [SIA, Unit: $\mu\text{g m}^{-3}$] response to NO_x case (a, b), AVOCs case (c, d), N+A case (e, f), and TOTAL case (g, h) relative to BASE case. Results are shown for January (a, c, e, g) and July (b, d, f, h) of 2018. SIA represents the sum of nitrate, sulfate, and ammonia. Notice the inconsistency in the scale of Figure 11c and h. Arrows represent the wind speed and wind direction.

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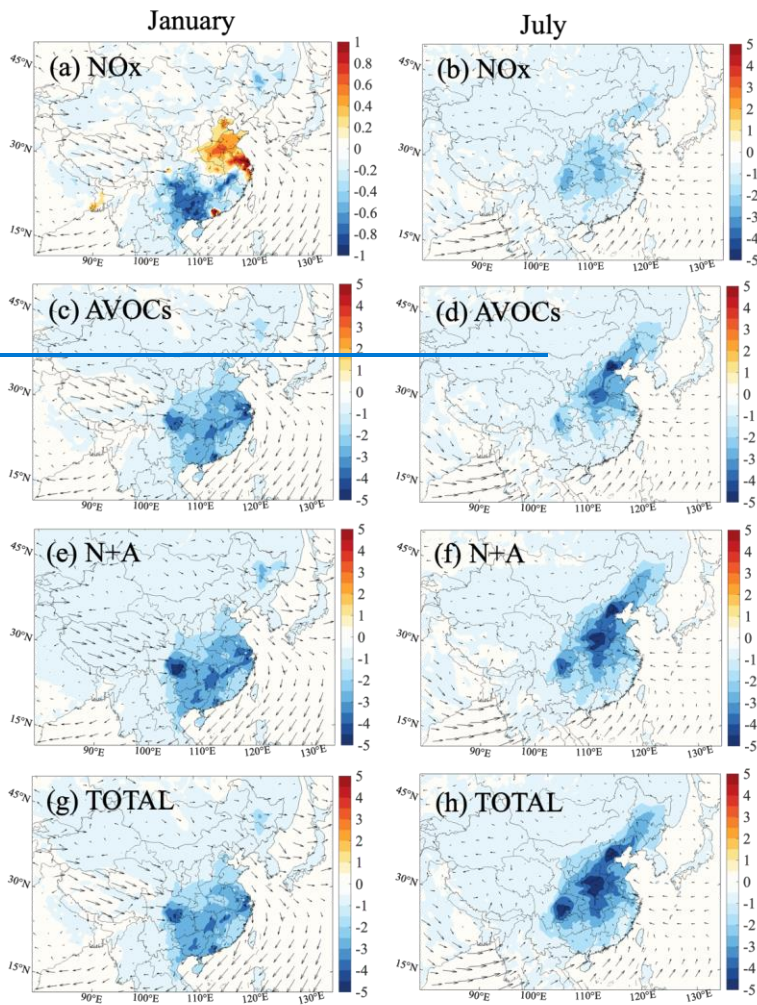


Figure 12. Changes in the surface concentration of Secondary Organic Aerosol [SOA, Unit: $\mu\text{g m}^{-3}$] response to NOx case (a, b), AVOCs case (c, d), N+A case (e, f), and TOTAL case (g, h) relative to BASE case. Results are shown for January (a, c, e, g) and July (b, d, f, h) of 2018. Notice the inconsistency in the scale of Figure 12a. Arrows represent the wind speed and wind direction.

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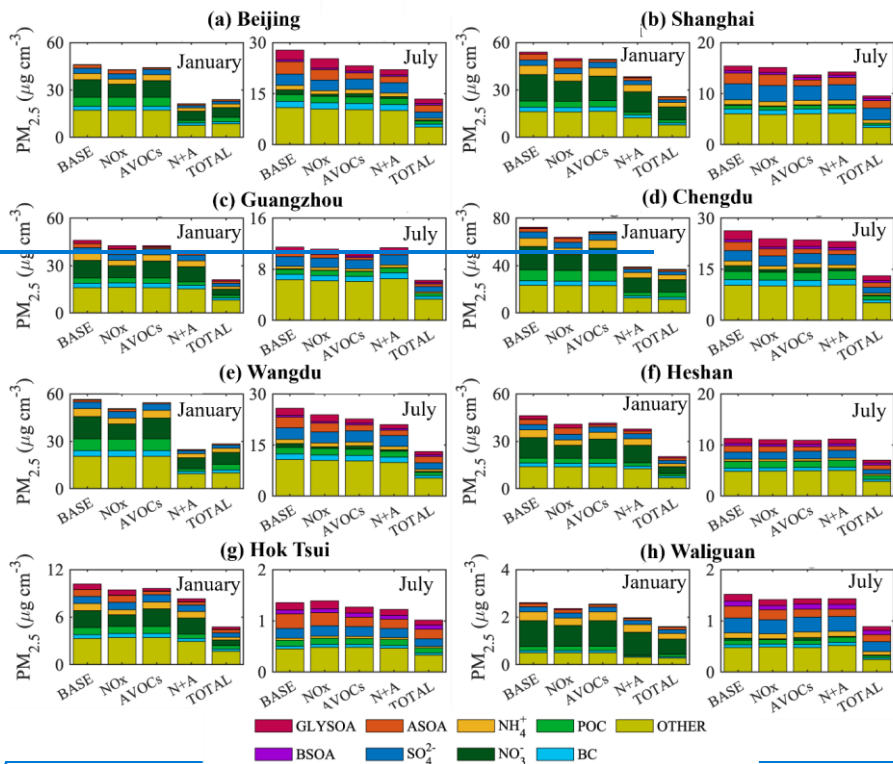


Figure 13. Averaged concentration of $PM_{2.5}$ [Unit: $\mu g m^{-3}$] with the contribution from different chemical compositions in five simulated cases and in eight sites for July 2018. The fine particle is composed of particulate nitrate (NO_3^-), particulate sulfate (SO_4^{2-}), particulate ammonia (NH_4^+), primary organic carbon (POC), black carbon (BC), anthropogenic secondary organic aerosol (ASOA), biogenic secondary organic aerosol (BSOA), secondary organic aerosol from glyoxal (GLYSOA), and other aerosol compositions (OTHER, including sea salt, carbonate, calcium, minerals, and other inorganic mass)

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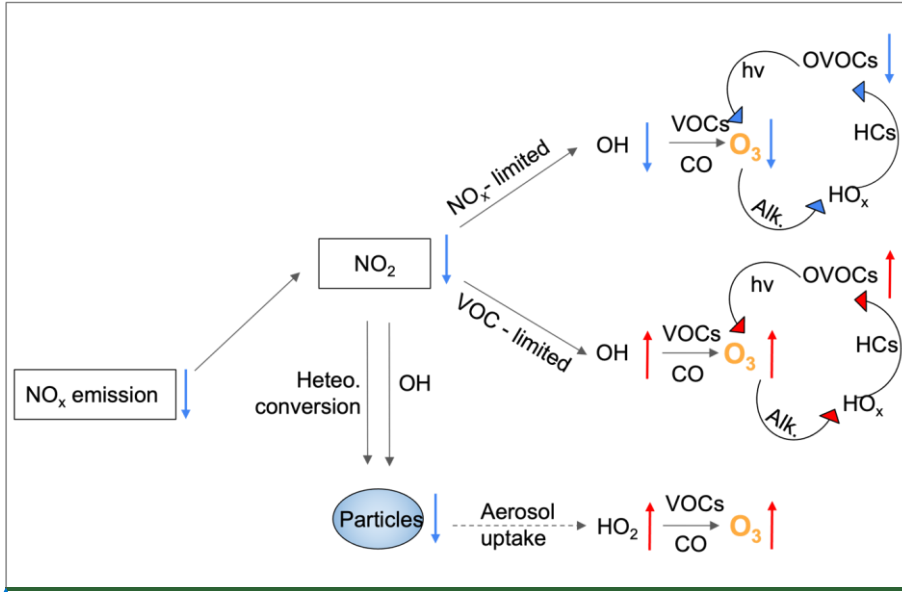


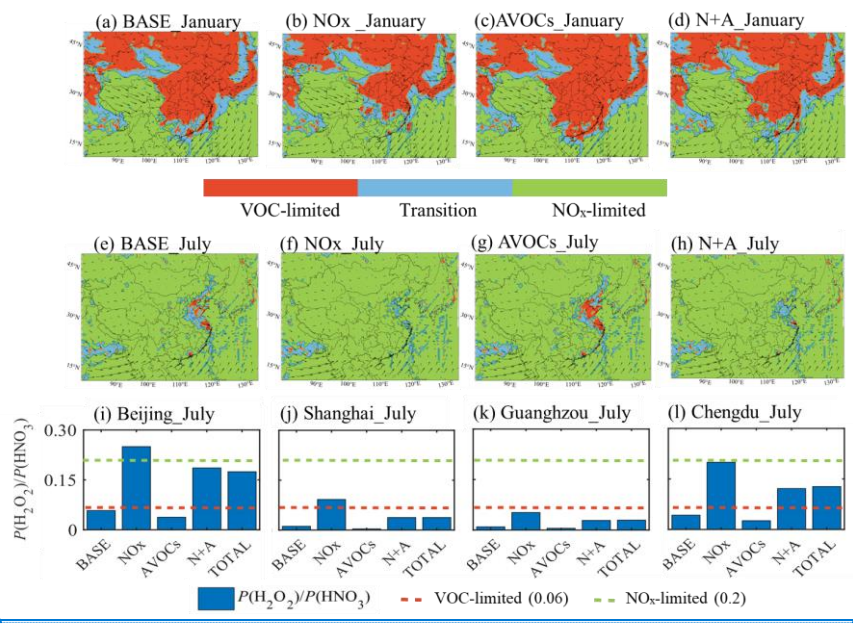
Figure 14. Schematics for the impact of NO_x emission reduction through aerosol effect and photochemical processes on ozone concentration. Arrows represent the changes in the concentration of chemicals associated with the reduction of NO_x emission (decrease trend shown in blue increase trend shown in red). HCs and Alk. are the abbreviations of hydrocarbons and alkenes.

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2055 [Figure 5. Ozone sensitivity regimes and their changes due to emissions reduction. \(a-h\) Display of regions in which ozone production is limited by the availability of nitrogen oxides \(NO_x-limited, in green\), and volatile organic components \(VOC-limited, in red\) from the TOTAL case under the emissions in case of BASE, NO_x, AVOCs, and N+A conditions in January \(a-d\) and July \(e-h\) of 2018. The regions where ozone production is controlled by the availability of both NO_x and VOCs \(transition\) are shown in blue. \(i-l\) The daytime \(06:00 to 19:00 Local Standard Time \(LST\)\) value of ratio between the production rate of hydrogen peroxide \(H₂O₂\) and nitric acid \(HNO₃\) \[\$P\(\text{H}_2\text{O}_2\)/P\(\text{HNO}_3\)\$ \] at four city sites \(Beijing, Shanghai, Chengdu, Guangzhou\) of China for July 2018 in the five emissions cases.](#)

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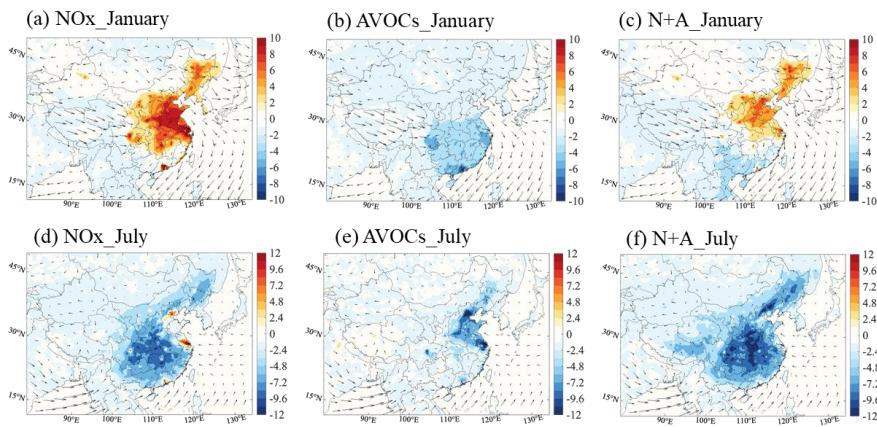
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Figure 6. Changes in the averaged daytime surface mixing ratio ozone response to a 50% reduction in NO_x emissions (NO_x case), in anthropogenic VOCs (AVOCs) emissions (AVOCs case) and in combined NO_x and AVOCs emissions (N+A case) relative to BASE case for January (a-c) and July (d-f) of 2018. Arrows represent the wind speed and wind direction.

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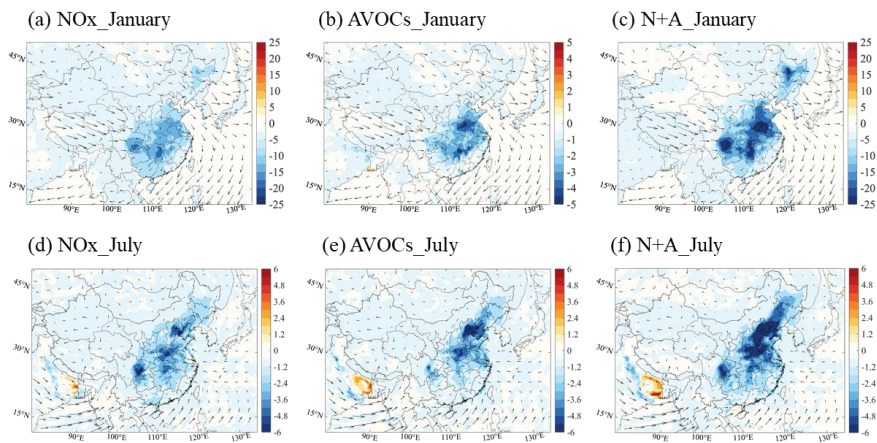


Figure 7. Changes in the surface concentration of fine particulate aerosol (Unit: $\mu\text{g m}^{-3}$) in response to NOx (a, d), AVOCs (b, e) and N+A case (c, f) relative to BASE case for January (a-c) and July (d-f) of 2018. Arrows represent the wind speed and wind direction. Notice the inconsistency in the scale of Figure 7b.

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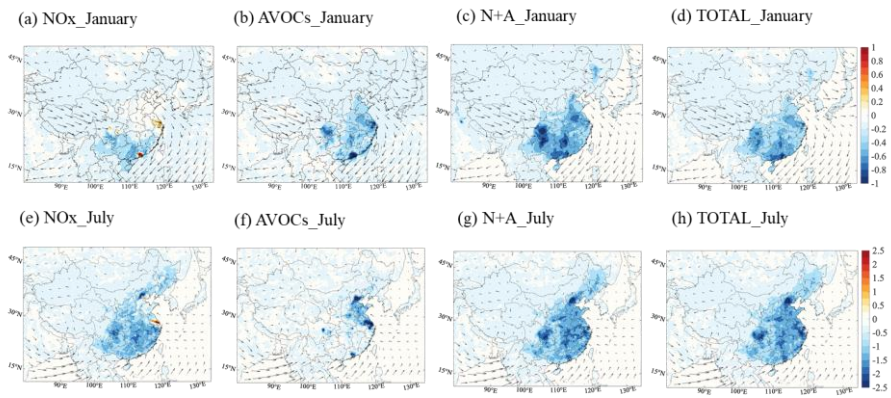
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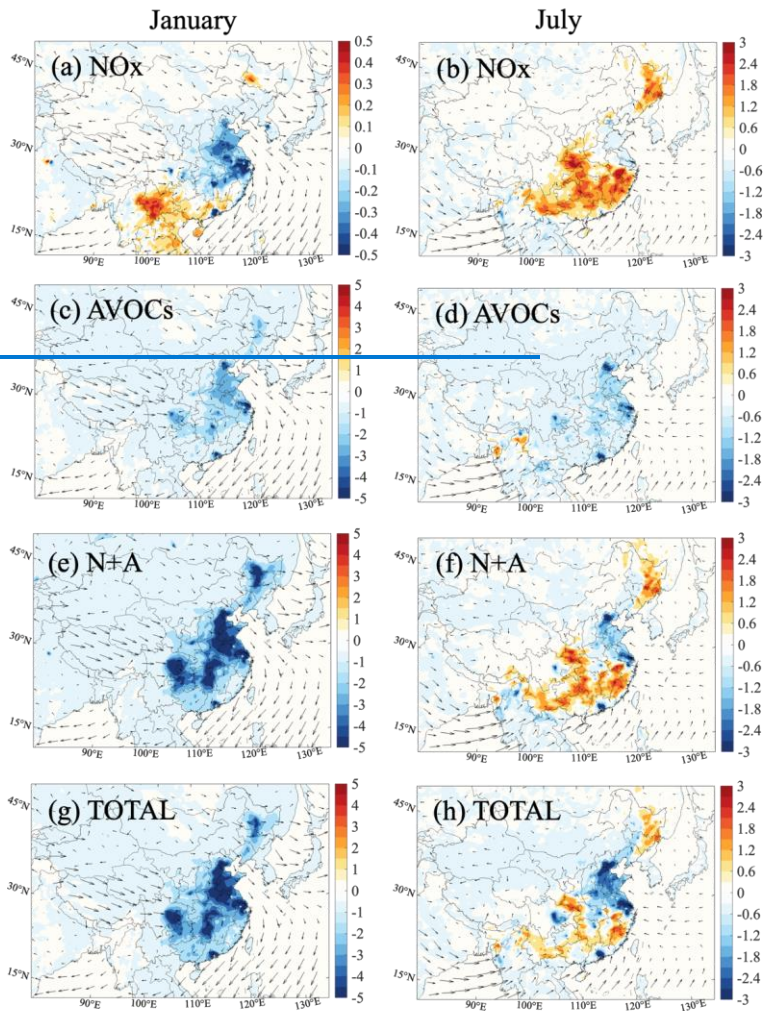
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2115 [Figure 8. Changes in the average daytime atmospheric oxidizing capacity \(AOC, Unit: \$10^7\$ molec. \$\text{cm}^{-3} \text{s}^{-1}\$ \) response to \$\text{NO}_x\$ \(a, e\), AVOCs \(b, f\), N+A \(c, g\), and TOTAL \(d, h\) cases relative to BASE case for January \(a-d\) and July \(e-h\) of 2018.](#)

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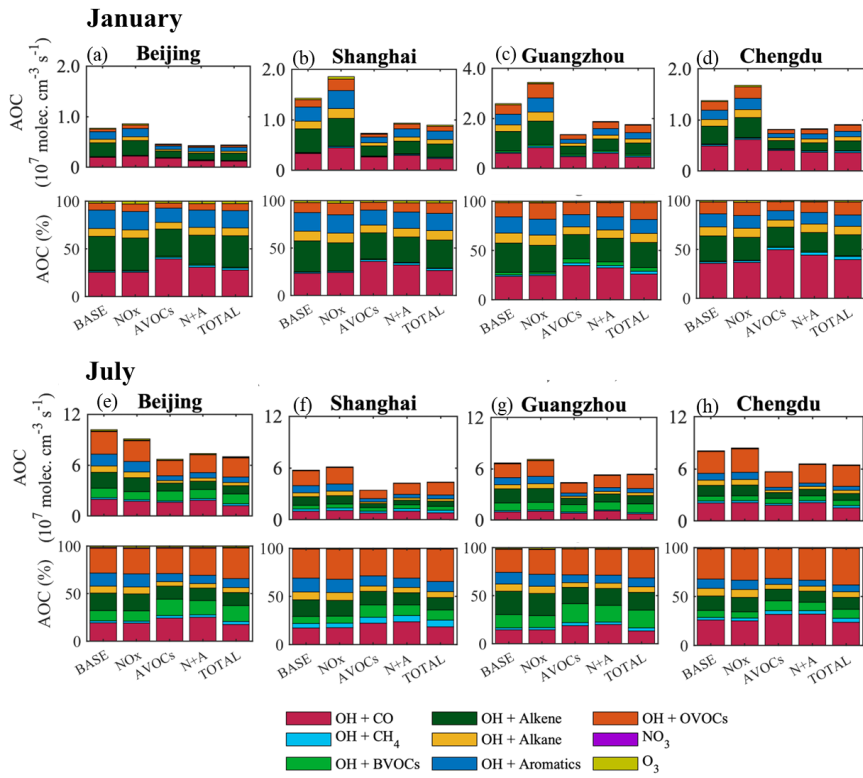


Figure 9. Averaged value and relative terms of daytime (06:00 to 19:00 Local Standard Time) AOC at the sites of Beijing (a, e), Shanghai (b, f), Guangzhou (c, g), and Chengdu (d, h) in five different simulated cases (BASE, NO_x, AVOCs, N+A, TOTAL cases) in January (a-d) and July (e-h) of 2018. Notice the inconsistency in the scale of Figure 9c.

Figure 15. Changes in the average daytime OH reactivity from VOCs and CO [Unit: s⁻¹] response to the NO_x case (a, b), AVOCs case (c, d), N+A case (e, f), and TOTAL case (g, h) relative to BASE case. Results are shown for January (a, c, e, g) and July (b, d, f, h) of 2018. Arrows represent the wind speed and wind direction.

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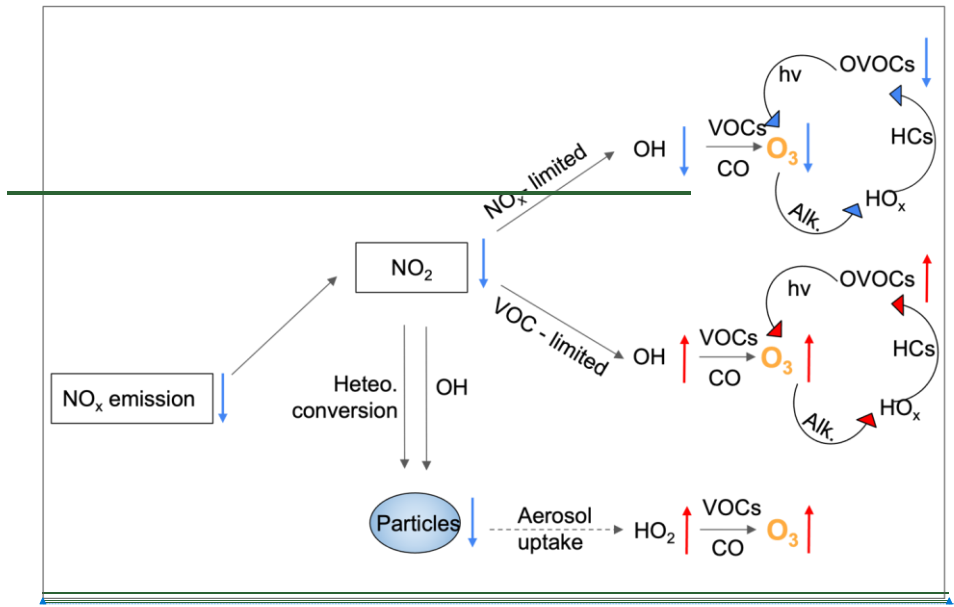
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O₃ changes due to reduction in primary NO_x and AVOCs emissions

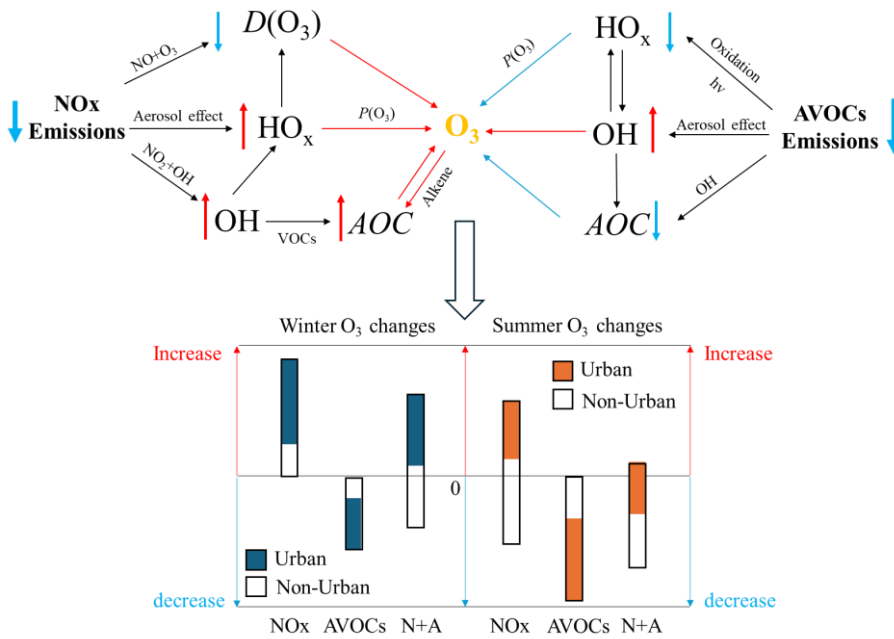


Figure 104. Schematics shows the responses of oxidative processes, associated with ozone formation, to the for the impact of NO_x emission reduction in primary emissions of NO_x and AVOCs in urban areas (VOC-limited) in winter and summer. through aerosol effect and photochemical processes on ozone concentration. Arrows besides the chemicals represent the changes in the concentration of chemicals associated with the reduction of NO_x in emission. (decrease trend shown in blue; increase trend shown in red). Blue and red arrows closing to O₃ represent the positive and negative contributions to the ozone formations. AOC, HCs and P(O₃), and D(O₃) are the abbreviations of the Atmospheric Oxidative Capacity, hydrocarbons and alkenes, and destruction of ozone. Bar figure shows the ranges of ozone changes in whole of China (black bar), in non-urban areas (white part in the bar) in four city sites in three emissions cases (NO_x, AVOCs, and N+A represent the case with emissions reduction in NO_x, Anthropogenic VOCs (AVOCs), and the combined NO_x and AVOCs emissions, respectively) relative to BASE cases in winter and summer.

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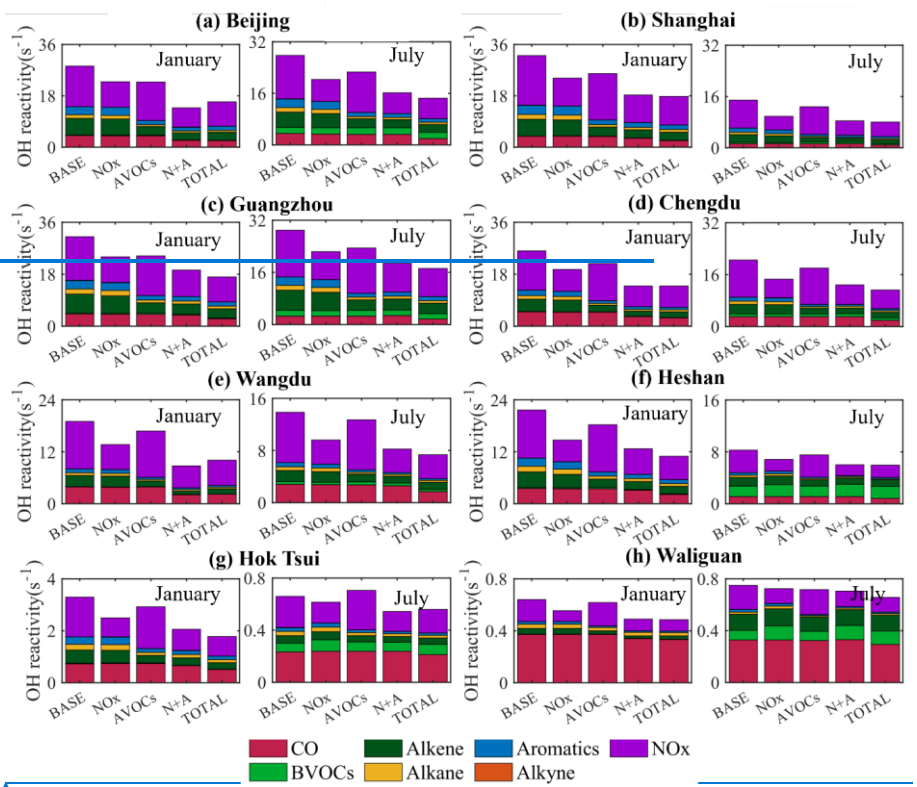
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Figure 16. Averaged value of daytime OH reactivity [Unit: s⁻¹] with the contribution from seven different species in five different simulated cases (BASE, NOx, AVOCs, N+A, TOTAL cases) and in eight different sites in July.

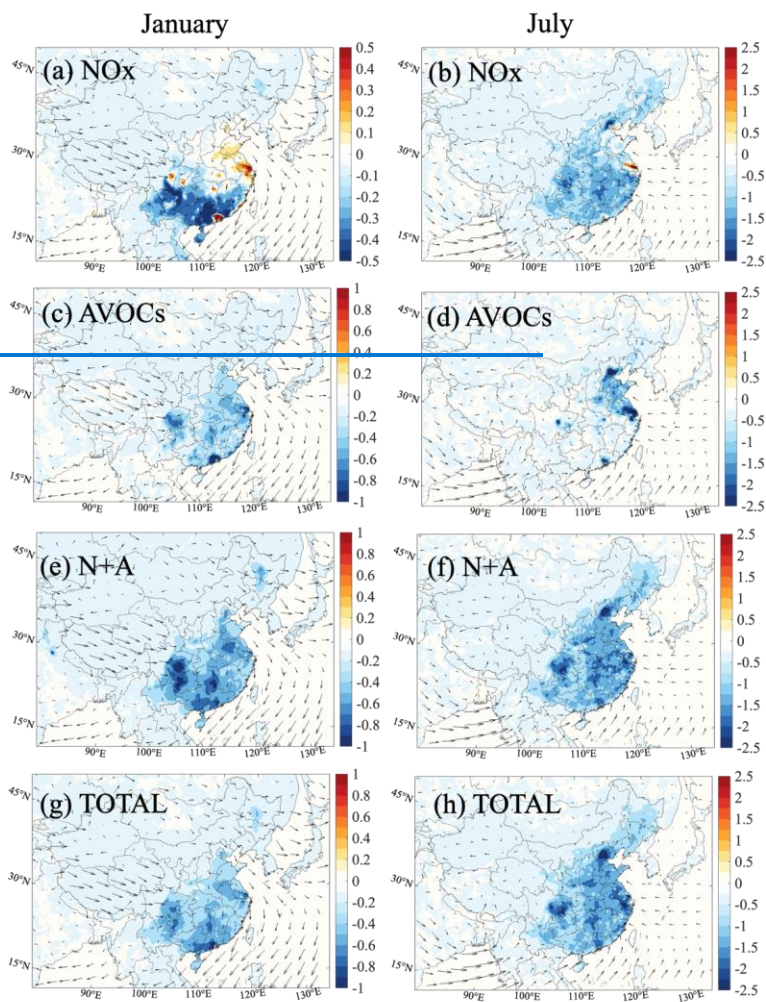


Figure 17. Changes in the daytime averaged atmospheric oxidizing capacity (AOC) [Unit: 10^7 molec. $\text{cm}^{-3} \text{s}^{-1}$] response to the NO_x case (a, b), AVOCs case (c, d), N+A case (e, f), and TOTAL case (g, h) relative to BASE case. Results are shown for January (a, c, e, g) and July (b, d, f, h) of 2018. Arrows represent the wind speed and wind direction.

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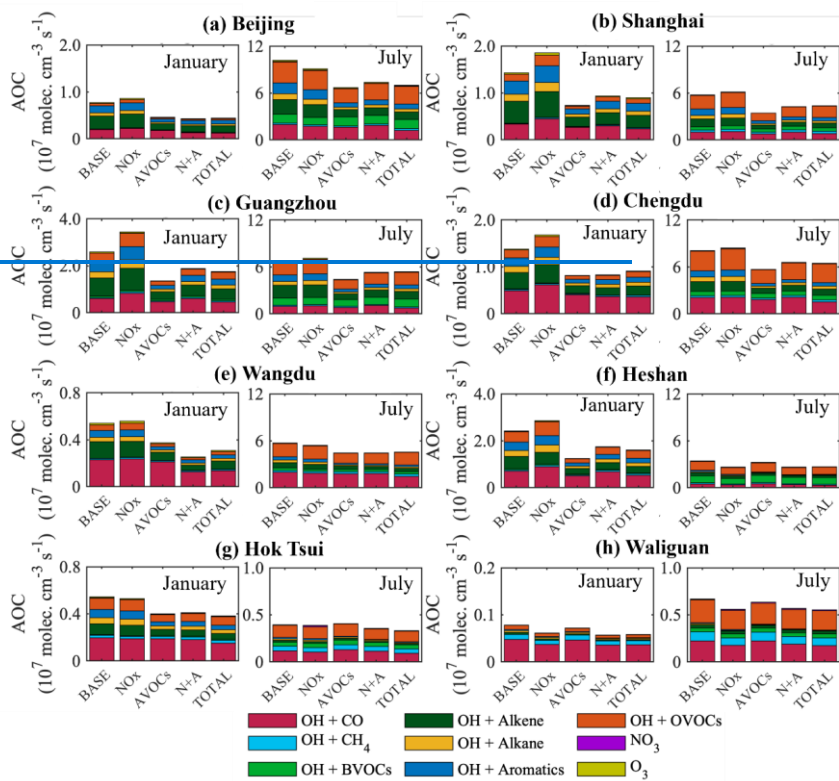
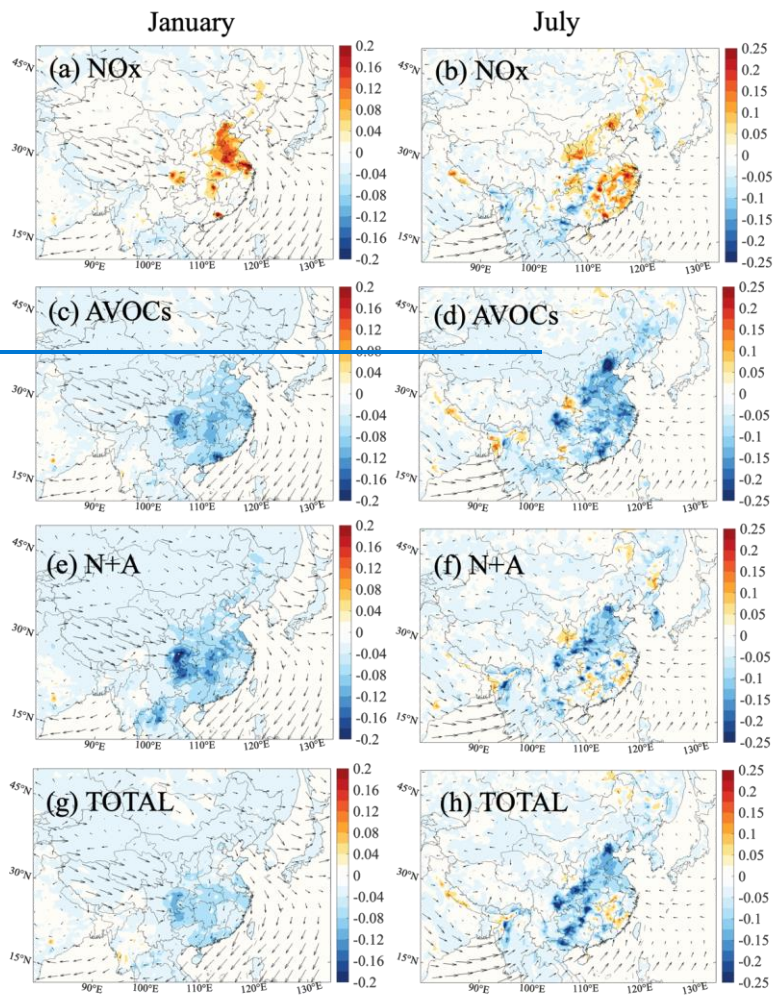


Figure 18. Averaged value of AOC [Unit: $10^7 \text{ molec. cm}^{-3} \text{ s}^{-1}$] during daytime in nine different species in five different simulated cases (BASE, NO_x, AVOCs, N+A, TOTAL cases) and in eight different sites in July.

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2200 Figure 19. Spatial distribution of the averaged nighttime (20:00 to 05:00 LST) atmospheric
 2205 oxidizing capacity (AOC) due to the reactions between ozone and alkenes [Unit: 10^6 molec.
 $\text{cm}^{-3}\text{s}^{-1}$] response to the ratio of 0.5 in NO_x emissions (a, b; NO_x case), in Anthropogenic
 VOCs (AVOCs) emissions (c, d; AVOCs case), in NO_x and AVOCs emissions (e, f; N+A case),
 and all anthropogenic emissions (g, h, TOTAL case) relative to BASE case. Results are shown
 for January (a, c, e, g) and July (b, d, f, h) of 2018. Arrows represent the wind speed and wind
 direction.

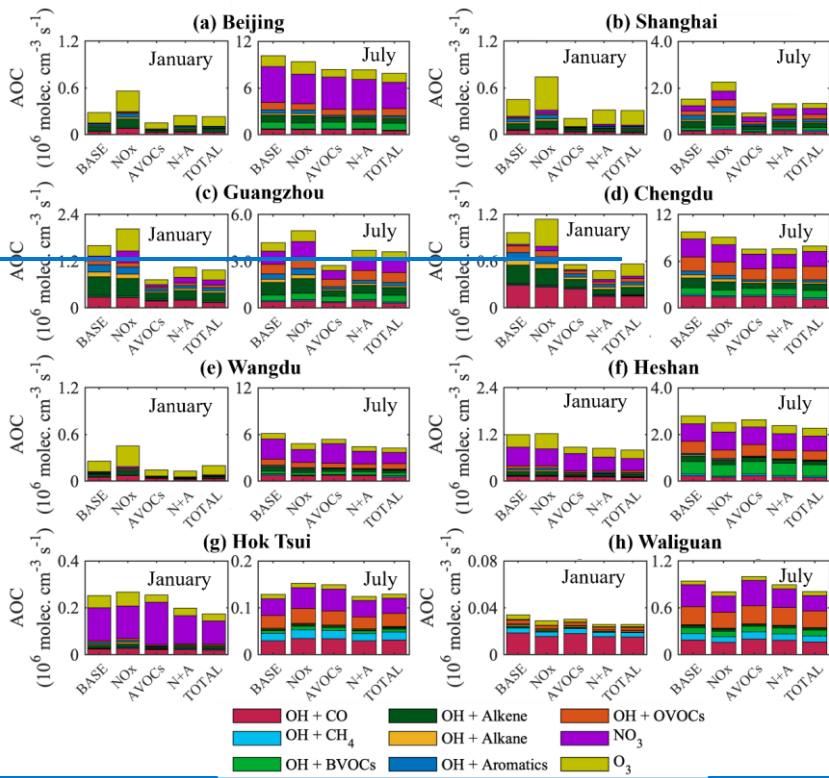


Figure 20. Averaged value of AOC [Unit: $10^6 \text{ molec. cm}^{-3} \text{ s}^{-1}$] during nighttime in nine different species in five different simulated cases (BASE, NO_x, AVOCs, N+A, TOTAL cases) and in eight different sites in July.

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