

1 **Technical note: Ocean Alkalinity Enhancement Pelagic Impact Intercomparison Project**  
2 **(OAEPIIP)**

3

4 Lennart Thomas Bach<sup>1\*</sup>, Aaron Ferderer<sup>1,2</sup>, Julie LaRoche<sup>3</sup>, Kai Georg Schulz<sup>4</sup>

5

6 <sup>1</sup>Institute for Marine and Antarctic Studies, University of Tasmania, TAS 7004, Australia

7 <sup>2</sup>National Collections and Marine Infrastructure, Commonwealth Scientific and Industrial  
8 Research Organisation, Hobart, Tasmania, Australia

9 <sup>3</sup>Department of Biology Dalhousie University, Halifax, NS Canada

10 <sup>4</sup>Faculty of Science and Engineering, Southern Cross University, Lismore, NSW, Australia

11

12 \*corresponding author ([lennart.bach@utas.edu.au](mailto:lennart.bach@utas.edu.au))

13

14 **Abstract**

15

16 Ocean Alkalinity Enhancement (OAE) aims to transfer carbon dioxide (CO<sub>2</sub>) from the  
17 atmosphere to the ocean by increasing the capacity of seawater to store CO<sub>2</sub>. The potential  
18 effects of OAE-induced changes in seawater chemistry on marine biology must be assessed  
19 to understand if OAE, operated at a climate relevant scale, would be environmentally  
20 sustainable. Here, we describe the design of the Ocean Alkalinity Enhancement Pelagic Impact  
21 Intercomparison Project (OAEPIIP) - a standardised OAE microcosm experiment with plankton  
22 communities to be conducted worldwide. OAEPIIP provides funding for participating  
23 laboratories to conduct OAE experiments in their local environments. This paper constitutes  
24 a detailed manual on the standardised methodology that shall be adopted by all OAEPIIP  
25 participants. The individual studies will provide new insights into how plankton communities  
26 respond to OAE. The synthesis of these standardized studies, without publication bias, will  
27 reveal common OAE-responses that occur across geographic and environmental gradients  
28 and are therefore particularly important to determine. The funding available to OAEPIIP and  
29 resulting data will be shared to maximise its value and the accessibility. The globally  
30 coordinated effort has potential to promote scientific consensus about the potential effects  
31 of OAE on diverse plankton communities. Such consensus, through inclusion of the global

32 community, will provide a sounder base to facilitate political decision making whether OAE  
33 should be upscaled or not.

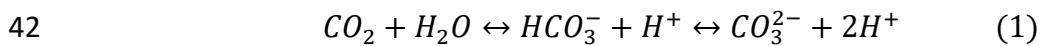
34

35 **1. Rationale for the Ocean Alkalinity Enhancement Pelagic Impact Intercomparison**  
36 **Project (OAEPIIP)**

37

38 Ocean Alkalinity Enhancement (OAE) is an emerging carbon dioxide removal (CDR) approach  
39 (Oschlies et al., 2023, NASEM, 2022) . OAE drives CDR through the introduction of alkaline  
40 substances into seawater which shift the carbonate chemistry equilibrium:

41



43

44 from carbon dioxide (CO<sub>2</sub>) on the left to bicarbonate (HCO<sub>3</sub><sup>-</sup>) and carbonate ions (CO<sub>3</sub><sup>2-</sup>) on  
45 the right. The decline in seawater CO<sub>2</sub> concentration lowers the seawater CO<sub>2</sub> partial pressure  
46 (pCO<sub>2</sub>), thereby enabling an influx of additional atmospheric CO<sub>2</sub>, or alternatively, reducing  
47 the efflux in cases where the surface ocean is a natural source of CO<sub>2</sub> to the atmosphere. The  
48 OAE-induced shift in carbonate chemistry is measurable as an increase in seawater alkalinity  
49 – the name-giving feature of OAE. The viability of OAE to serve as a scalable CDR approach  
50 critically depends on whether it is environmentally safe. Surface ocean habitats are in focus  
51 of the environmental OAE assessment because the surface ocean is where OAE would need  
52 to be implemented to enable CO<sub>2</sub> exchange with the atmosphere (Bach et al., 2019).

53 The environmental OAE assessment is only just starting but seems to be evolving in a  
54 similar way as environmental assessments of other drivers of environmental change (e.g.  
55 ocean acidification) have been set up in the past: Research funding is provided to individual  
56 groups, who will perform individual studies in their local environments, seeking novelty. Each  
57 of these studies will be valuable and exceeding previous research is central to scientific  
58 progress. However, previous research on environmental drivers has also shown that  
59 replication of experiments is perhaps equally important as seeking novelty, since replication  
60 allows us to reveal re-occurring response patterns across various scales and environments  
61 (Benton et al., 2007; Hamm et al., 2022; Stewart et al., 2013). In ocean acidification research  
62 for example, an individual study found that carbon to nitrogen (C/N) stoichiometry of  
63 plankton communities is increased under high CO<sub>2</sub> conditions due to CO<sub>2</sub> fertilization of the

64 phytoplankton community (Riebesell et al., 2007). However, replication of the experiment at  
65 different locations found that zooplankton communities can strongly modify the response, to  
66 the point that the response can be significant in the opposite direction (lower C/N under high  
67 CO<sub>2</sub> (Taucher et al., 2021)). Arguably, the crucial progress in this example was understanding  
68 of the context-dependency of the C/N response to ocean acidification, which was made  
69 possible by replication of a sophisticated experiment across a wide geographical range  
70 (Riebesell et al., 2013). Likewise, the intercomparison of climate models via replicated  
71 numerical experiments (Dingley et al., 2023) has long been recognised as a cornerstone to the  
72 assessment of climate change (Masson-Delmotte et al., 2021), possibly more influential than  
73 the output of individual climate models.

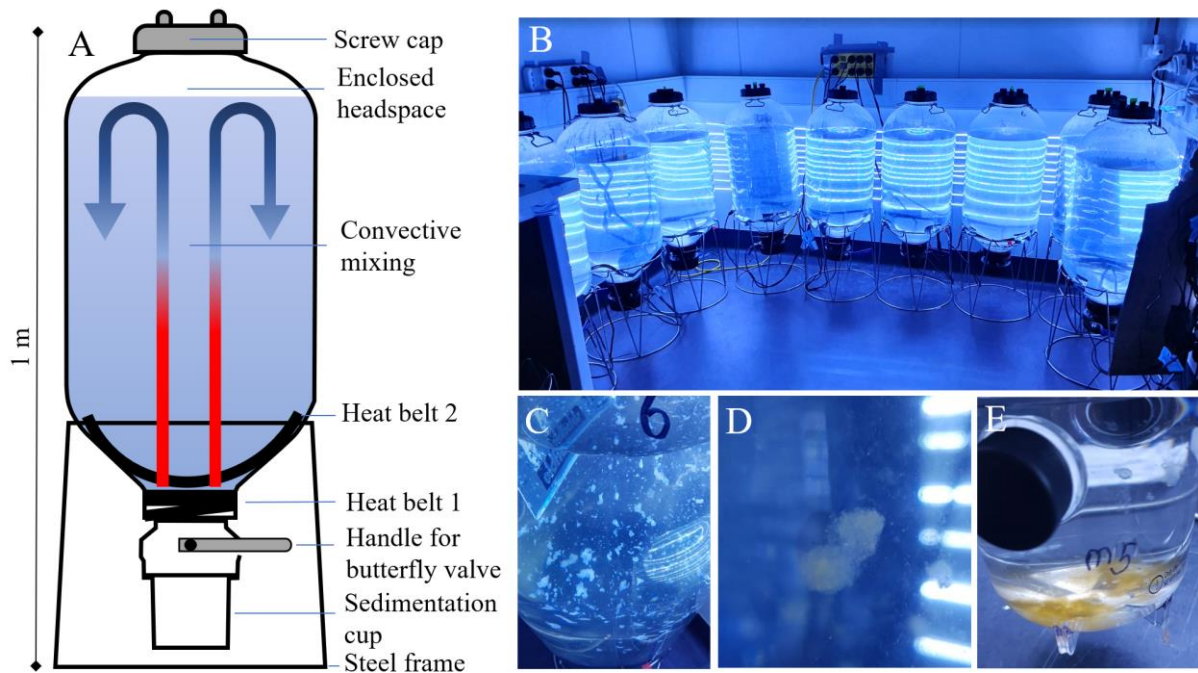
74 The Ocean Alkalinity Enhancement Pelagic Impact Intercomparison Project (OAEPIIP) builds  
75 upon these insights from previous environmental assessments by establishing a platform that  
76 supports replication, while still enabling the pursuit of novelty. In essence, OAEPIIP provides  
77 funding for a cost-efficient and standardised OAE experiment, which can be conducted by  
78 scientists across the globe (section 2). The experiments will use a microcosm setup to study  
79 the response of natural plankton communities to one widely considered OAE implementation  
80 strategy, and they will determine the same set of response variables. Each experiment shall  
81 be published on an individual basis in a special issue of a peer-reviewed scientific journal  
82 under open access with costs largely covered by OAEPIIP (section 3). Individual publication of  
83 OAEPIIP experiments gives room to describe novel observations on how plankton  
84 communities respond to OAE. All datasets will be shared and synthesized in a meta-analysis.  
85 The standardised experimental design facilitates inclusion of individual datasets into the  
86 meta-analysis (Harrison, 2011). Likewise, the collection of all datasets, irrespective of their  
87 outcomes, avoids publication bias, which is a known problem of meta-analyses (Field and  
88 Gillett, 2010). We expect OAEPIIP to promote consensus among scientists concerning the  
89 potential environmental side effects of OAE on plankton communities, with significant  
90 potential for capacity building (section 4). This paper provides a detailed manual for the  
91 OAEPIIP experimental setup and describes its benefits.

92

## 93 **2. Experimental infrastructure, operation, and design**

94

### 95 **2.1. Microcosm setup**



97

98 **Figure 1.** Microcosm setup. (A) Schematic of the microcosm tanks. The 2 heat belts induce  
 99 convective mixing within the tanks. (B) Arrangement of 9 microcosms in a temperature-  
 100 controlled room in front of a light source. Their position should be changed on a daily basis to  
 101 minimize position-dependent differences in light and temperature over the course of the  
 102 study. (C) A picture of a microcosm, just after NaOH addition. The white flakes are brucite  
 103 particles that need to be dissolved after NaOH addition by stirring the seawater within  
 104 microcosms with a plastic spoon. (D) A close-up of a marine snow aggregate, which frequently  
 105 forms after a phytoplankton bloom. (E) Marine snow aggregates collected in the  
 106 sedimentation cup of the microcosm. Sampling these can be interesting, although this is not  
 107 an essential parameter of OAEPiIP (section 2.6.).

108

109 OAEPiIP utilizes the microcosm setup developed by Ferderer et al. (2022), as it is cost-effective  
 110 and relatively easy to set up and operate. The microcosms are 55L Polyethylene terephthalate  
 111 (PET) tanks (FermZilla), which were originally designed for home brewing (Fig. 1), and  
 112 available worldwide (Table S1). It is important that all OAEPiIP participants purchase the same  
 113 type of microcosm incubators so that comparability can be established between individual  
 114 studies (details on the availability of FermZilla tanks are provided in Table S1). The tanks are  
 115 mounted on steel frames and have 120 and 70 mm openings at the top and bottom,  
 116 respectively. The bottom opening is equipped with a butterfly valve and a sedimentation cup,

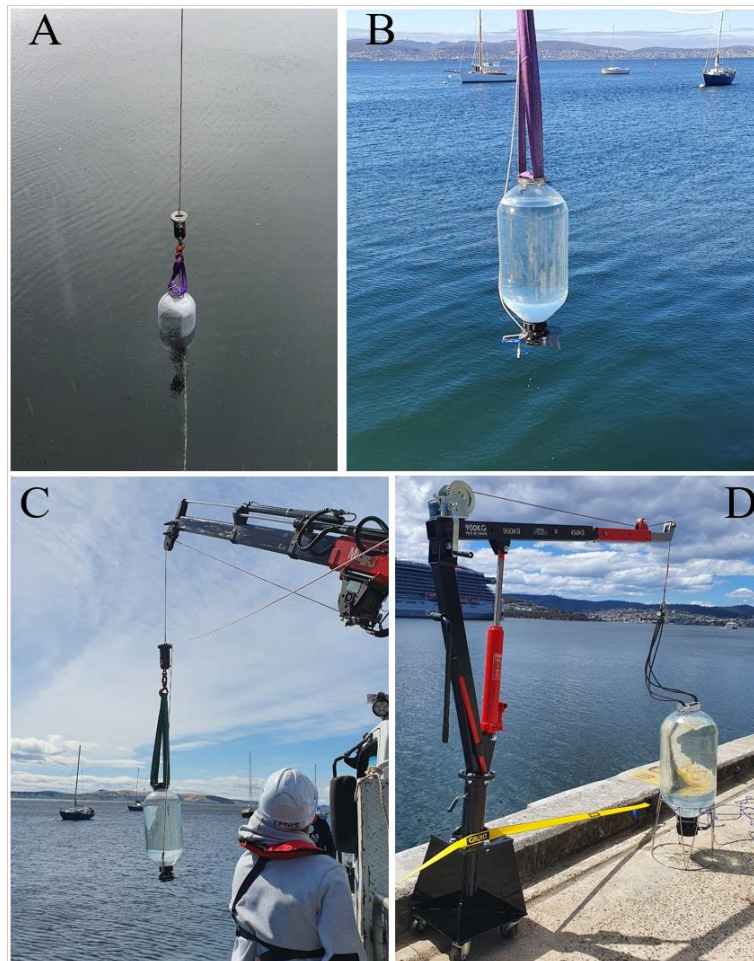
117 used for the collection of settling material. The butterfly valve has a handle so that the  
118 sedimentation cup can be isolated from the water column.

119 The crucial steps for setting up the microcosms, their filling, and their operation are listed in  
120 Table S1, illustrated in Figs. 1 and 2 and Videos S1 (<https://doi.org/10.5446/66751>) and S2  
121 (<https://doi.org/10.5446/66753>), and briefly described here. OAEPIIP experiments occupy  
122 approximately 9 m<sup>2</sup> in a temperature-controlled room with a cooling capacity of roughly 6°C  
123 below the temperature aimed for in the microcosm study (e.g., to 14°C if the desired  
124 experimental water temperature is 20°C). The microcosms need to be thoroughly cleaned in  
125 a two-step procedure before use as detailed in Table S1.

126 Infrastructure needed for filling the microcosms with natural seawater (containing natural  
127 plankton communities) depends on the local environment at an OAEPIIP study site. At our site  
128 in Tasmania, we fill microcosms from a jetty using a small crane or davit (Fig. 2; Video S1  
129 (<https://doi.org/10.5446/66751>)). Natural seawater with plankton communities shall be  
130 collected by opening the top lid and butterfly valve at the bottom and lowering the  
131 microcosms slowly into seawater so that each microcosm is filled from bottom to top. Care  
132 must be taken to not enclose larger debris, nekton or sediments. Once the microcosm is  
133 submersed and only the upper opening is above the sea surface, a rope attached to the handle  
134 of the butterfly valve is pulled so that the bottom opening is closed. The microcosm can now  
135 be lifted back on shore and put back into its metal frame. Another possibility to fill microcosms  
136 is to slowly lower them from a low swimming pontoon or small boat and close the bottom  
137 manually. Filling microcosms by slowly lowering them into seawater is a very gentle way to  
138 collect plankton communities (Video S1 (<https://doi.org/10.5446/66751>)), avoiding the  
139 physical disturbance to plankton imposed by pumping. Based on our experience it takes  
140 roughly 45 minutes to fill 9 microcosms. Longer timescales for the collection (i.e., >>1 hour)  
141 should be avoided to mitigate the risk of changes in seawater communities over the course  
142 of the filling procedure (e.g. through tidal water movement). This potential problem should  
143 also be minimised by filling the microcosms in random order. Furthermore, care should be  
144 taken to not expose the microcosms to excessive sunlight and/or heat after filling.

145 The weight of the enclosed seawater needs to be determined after the filling procedure as  
146 this information is needed for establishing treatments (section 2.5). This could be done using  
147 a balance or (if a balance is not available) volumetrically and determining weight with known  
148 volume, temperature, and salinity. Once the weight has been determined, microcosms need

149 to be transported to the temperature-controlled room where the experiment takes place and  
150 light and temperature control needs to be initiated immediately (see following section).



151  
152 **Figure 2.** Seawater collection for the microcosm experiments. (A) A microcosm slowly lowered  
153 into seawater to gently collect a plankton community. (B and C) A filled microcosm being  
154 pulled back on land. Please note that we mostly used a small crane mounted to a truck or a  
155 davit (as in D) for the seawater collection. However, microcosms filled with seawater only  
156 weigh about 60 kg, so that lighter gear is probably sufficient for collection. A detailed  
157 description of seawater collection is provided in Table S1 and Video S1  
158 (<https://doi.org/10.5446/66751>).

159

## 160 2.2. Mixing, temperature, light, and nutrient conditions in OAEPiIP experiments

161

162 OAEPiIP utilises convection to mix the enclosed microcosm volume and keep plankton in  
163 suspension (Fig. 1). To establish convective mixing, two 30 Watts heat belts (see Table S1 for  
164 where these can be purchased) will be firmly attached to two distinct locations at the bottom

165 of the microcosms (Fig. 1, Video S2 (<https://doi.org/10.5446/66753>)). Based on our  
166 experience, these two 30 W heat belts increase the temperature of the enclosed seawater by  
167 ~6°C relative to the room temperature, so that room temperature needs to be roughly ~6°C  
168 lower than the target temperature in the experiments (please note that testing the  
169 temperature difference will be necessary prior to the experiment as temperature offset may  
170 differ across temperature-controlled rooms). Once heat belts are attached, microcosms  
171 should be placed in front of the light source and heat belts should be plugged in to initiate  
172 the convective mixing.

173 While convection provides gentle and non-invasive mixing, there are several trade-offs  
174 regarding temperature control. Firstly, due to the removal of seawater during sampling, the  
175 total volume within microcosms declines over the course of the experiment. Since the heat  
176 belts cannot be adjusted, there is an increase in heat energy input per liter of enclosed  
177 seawater and thus a gradual warming. To mitigate this issue, the external cooling may need  
178 to be increased over time by lowering the room temperature. In our experience, a reduction  
179 by 1°C for every 5 Liters of seawater sampled from the microcosms is sufficient to keep the  
180 seawater temperature relatively constant over the course of the study. Secondly, small  
181 differences in ventilation at different locations in the temperature-controlled room can lead  
182 to seawater temperature differences of around 2°C between microcosms (Ferderer et al.,  
183 2022; Guo et al., 2023). To mitigate this experimental constraint, the microcosm placement  
184 within the temperature-controlled room must be shuffled daily. Microcosms can easily be  
185 moved when they are being pulled on the steel frame (Fig. 1), but care must be taken to briefly  
186 unplug the heat belts and plug them in again after shuffling their position. Furthermore, fans  
187 can be utilised to remove heat pockets in the room, although care must be taken as the wind  
188 can have a strong cooling effect, resulting in a microcosm that was too warm quickly  
189 becoming too cold. Since temperature is a strong driver of physiological processes, it is highly  
190 advisable to thoroughly test the setup with all microcosms prior to the experiment (the  
191 careful addition of food dye can be used to test advection as explained by Ferderer et al.,  
192 (2022)). The goal should be to have as little variation in temperatures between microcosms  
193 as possible, and the seawater temperature should be as the plankton community would have  
194 experienced it at the location/season it was collected.

195 Like temperature, light conditions set up for the experiment should reproduce the natural  
196 site-specific conditions as much as possible. This applies for the light/dark cycle, the light

197 intensity, and the light spectrum (light spectrum should be between 400 and 750 nm, i.e. cool  
198 white light). Since many OAEPIIP participants may not have access to sophisticated computer-  
199 controlled light sources, we recommend the application of constant light over a fixed  
200 light/dark cycle. In an OAE study in Tasmania, for example, we provided light constantly with  
201  $200 \mu\text{mol photons m}^{-2} \text{ s}^{-1}$  at a 12/12 hours light/dark cycle. These conditions were considered  
202 as representative average level for the surface mixed layer at the location/season  
203 where/when the natural plankton community were sourced in Tasmania. The light/dark cycle  
204 can be achieved by plugging the light source into a timer socket. The microcosms need to be  
205 positioned in such a way that light is very similar inside each microcosm. A light meter shall  
206 be used to determine light intensity inside the microcosms prior to the experiments and at  
207 the end (Table 1). Positioning can be very critical since movement by a few centimetres can  
208 often lead to noticeable changes in measured light intensity that are undetectable by the  
209 human eye. It is therefore important to adjust light conditions before starting the experiment  
210 and marking the spot on the floor where individual microcosms must be placed to ensure  
211 replicable light levels. It is also important to have all the other microcosms at their respective  
212 locations while doing the adjustments as they might shade each other. The daily shuffling of  
213 microcosm positions inside the room, which is essential for the temperature control (see  
214 above), will also help to mitigate systematic bias in light regime between microcosms.  
215 OAEPIIP experiments shall not add organisms, nutrients, or any substances other than  
216 alkalinity/DIC (section 2.4) to the microcosms during or before the experiments.

217

### 218 **2.3. OAE method in focus of OAEPIIP**

219

220 OAE can be implemented with different approaches, applying different alkaline feedstocks  
221 such as solid materials like olivine, calcium/magnesium hydroxides, carbonates, steel slags or  
222 liquid materials like sodium hydroxide (NaOH) dissolved in seawater (Eisaman et al., 2023,  
223 NASEM, 2022). Each alkalinity source has different environmental implications as it is  
224 associated with different environmental perturbations (Bach et al., 2019). A widely  
225 considered approach is electrochemical OAE, where liquid NaOH is the alkalinity source (de  
226 Lannoy et al., 2018). NaOH-based OAE is in focus of OAEPIIP due to the following reasons.  
227 First, electrochemical OAE was recently evaluated to rank among the highest OAE approaches  
228 with regards to their “technological readiness level” (Eisaman et al., 2023), with field trials



229 already underway. Second, liquid NaOH is suitable as an alkalinity source for applications in  
230 pelagic environments as it delivers quasi-instantaneous OAE in seawater. Other methods that  
231 involve more slowly dissolving minerals (e.g. olivine) are considered less suitable for pelagic  
232 applications as they would partially sink into the deep ocean before dissolving (Köhler et al.,  
233 2013; Fakhraee et al., 2023). Third, electrochemical OAE is chemically relatively similar to  
234 other OAE methods such as OAE with magnesium hydroxides or ocean liming based on  
235 calcium hydroxides. Like NaOH, magnesium and calcium hydroxides dissolve relatively quickly  
236 and are comparatively clean sources of alkalinity due to generally less content of bioactive  
237 elements like iron or nickel than for example olivine when derived from carbonates (Bach et  
238 al., 2019; Renforth et al., 2022), or when magnesium hydroxides are produced chemically  
239 (Eisaman et al., 2023). As such, the primary potential of these hydroxides to affect pelagic  
240 communities is by changing seawater carbonate chemistry. Thus, results from NaOH-based  
241 OAE experiments also have potential to inform these other approaches. Fourth, NaOH is  
242 readily available worldwide, which is logistically beneficial for OAEPiIP.

243

#### 244 **2.4. Experimental design**

245

246 NaOH-based OAE reduces seawater  $p\text{CO}_2$  within seconds (Zeebe and Wolf-Gladrow, 2001),  
247 whereas the subsequent equilibration with atmospheric  $\text{CO}_2$  takes months to years (Jones et  
248 al., 2014; Wang et al., 2023, Mu et al., 2023) or potentially even centuries (He and Tyka, 2023).  
249 The carbonate chemistry perturbation is much greater before the equilibration has happened  
250 so that more pronounced effects on communities would be expected shortly after alkalinity  
251 addition (Bach et al., 2019). As such, an argument can be made to study OAE at two different  
252 timepoints when using rapidly dissolving alkalinity sources like NaOH or other hydroxides.  
253 These are the “unequilibrated” timepoint, simulating the fact that  $\text{CO}_2$  influx has not yet  
254 happened right after alkalinity addition, and the “equilibrated” timepoint, assuming the  
255 alkalinity enhanced seawater has already  $\text{CO}_2$ -equilibrated with the atmosphere.  
256 Equilibration could either happen naturally through air-sea  $\text{CO}_2$  influx (Ho et al., 2023), or  
257 even be enforced within a facility so that alkalinity-enhanced seawater equilibrated with the  
258 atmosphere is discharged into the ocean.

259 The nine microcosms available for OAEPiIP experiments will provide triplicate incubations for  
260 controls, unequilibrated and equilibrated treatments. An important aspect for OAEPiIP

261 experiments is that the amount of alkalinity added to the treatments is consistent among all  
262 studies. Modelling studies suggest that gigatonne-scale OAE sustained for 80 years would  
263 increase surface ocean alkalinity by about 100-200  $\mu\text{mol}/\text{kg}$  (Burt et al., 2021; Lenton et al.,  
264 2018). This seemingly modest perturbation is due to dilution by the huge volume of the ocean  
265 (i.e.,  $9.44 \times 10^{17} \text{ m}^3$ ; Sarmiento and Gruber, 2006). However, the perturbation can be more  
266 pronounced at sites where alkalinity is added, before being diluted with unperturbed  
267 seawater with a rate that depends on the location (He and Tyka, 2023; Wang et al., 2023). In  
268 fast dilution regimes, for example, 1 molar NaOH added at 5  $\text{m}^3/\text{s}$  from a large ship would  
269 initially raise pH to 11 but dilute to pH 8.5 within minutes to hours (He and Tyka, 2023). As  
270 such, there is a trade-off for OAE experimentation between the realism and thus the  
271 relevance of a simulated OAE perturbation at the timescales (weeks) proposed here (requiring  
272 rather low alkalinity perturbation) and the detectability biological effects (facilitated when  
273 simulated perturbations are more extreme). For OAEPIIP, we determined an alkalinity  
274 enhancement of 500  $\mu\text{mol}/\text{kg}$  to both the unequilibrated and equilibrated treatments. Our  
275 rationale for the rather high perturbation is that OAEPIIP has a strong focus on capacity  
276 building in OAE research. Setting up clearly distinguishable treatments facilitates data analysis  
277 and interpretation, particularly for those entering the field. We emphasize, however, that a  
278 500  $\mu\text{mol}/\text{kg}$  perturbation over the duration of OAEPIIP studies (i.e., weeks; section 2.7) is on  
279 the higher end for what is plausible for OAE (Wang et al., 2023; He and Tyka, 2023), except  
280 for perhaps in proximity of a continuous NaOH release site. Thus, the relatively extreme  
281 perturbation needs to be taken into account for the eventual interpretation and  
282 communication of OAEPIIP studies – since it is likely that a less extreme perturbations would  
283 also cause less environmental effect.

284

## 285 **2.5. Establishing treatments**

286

287 Alkalinity enhancement shall be performed on day 0 of the experiment, shortly after  
288 microcosms have been positioned in the temperature-controlled room. Before adding  
289 alkalinity, carbonate chemistry samples (i.e., alkalinity and one other carbonate chemistry  
290 parameter; section 2.6) should be collected to constrain carbonate chemistry conditions in all  
291 microcosms before OAE.

292 The three control microcosms will not receive any alkalinity addition and remain untreated.  
293 The three microcosms of the unequilibrated treatment will receive 500  $\mu\text{mol/kg}$  of NaOH. The  
294 simplest way to achieve this is by purchasing and using a 1 molar NaOH solution (ideally in  
295 “analytical quality”) and adding 500  $\mu\text{L}$  per kg of enclosed seawater. For example, if 54.5 kg  
296 of seawater have been enclosed then  $54.5 * 500 = 27250$   $\mu\text{L}$  of 1 molar NaOH solution needs  
297 to be added to the respective microcosm.

298 The equilibrated treatment is slightly more complicated to establish. Here, most of the  
299 alkalinity needs to be added as sodium bicarbonate ( $\text{NaHCO}_3$ ) solution and a smaller amount  
300 as NaOH solution. We provide an R script based on Seacarb (Gattuso et al., 2021) that can be  
301 used to calculate additions of  $\text{NaHCO}_3$  and NaOH (OAEPIIP, 2024). Furthermore, video  
302 tutorials provide detailed instructions on how to use the R script or how to do these  
303 calculations with CO2SYS for MExcel (Pierrot et al., 2021) (Videos S3 and S4;  
304 <https://doi.org/10.5446/66754>, <https://doi.org/10.5446/66752>). Briefly: In a first step, initial  
305 carbonate chemistry conditions need to be calculated for the unperturbed seawater enclosed  
306 in the microcosms. For this calculation one needs to assume a current  $\text{CO}_2$  partial pressure  
307 (e.g., 420  $\mu\text{atm}$ ), the target temperature for the experiment, and a salinity and alkalinity  
308 estimate based on what the experimentalist expects for their region (or ideally has measured  
309 just before collecting the seawater for microcosm experiment). Next, the calculation is  
310 repeated for the same conditions except for alkalinity where 500  $\mu\text{mol/kg}$  is added to the  
311 assumed value (e.g., 2850  $\mu\text{mol/kg}$  when the assumed value of the unperturbed water was  
312 2350  $\mu\text{mol/kg}$ ). The second calculation represents the desired conditions in the equilibrated  
313 treatment after the alkalinity enhancement. The calculated dissolved inorganic carbon (DIC)  
314 concentrations of the initial carbonate system ( $\text{DIC}_{\text{initial}}$ ) need to be subtracted from the  
315 calculated DIC of the calculated treatment ( $\text{DIC}_{\text{equilibrated}}$ ):

316

$$317 \quad \text{NaHCO}_3 \text{ addition} = \text{DIC}_{\text{equilibrated}} - \text{DIC}_{\text{initial}} \quad (2)$$

318

319 where  $\text{NaHCO}_3$  addition is the amount of  $\text{NaHCO}_3$  that needs to be added per kg of enclosed  
320 seawater (in  $\mu\text{mol/kg}$ ). The addition of  $\text{NaHCO}_3$  provides equal amounts of DIC and alkalinity.  
321 However, OAE can only absorb  $\sim 0.85$  mole of DIC per mole of alkalinity added (He and Tyka,  
322 2023; Schulz et al., 2023), so that reaching to +500  $\mu\text{mol/kg}$  requires the addition of slightly

323 more alkalinity without DIC. NaOH is used for this purpose and the exact amount that needs  
324 to be added is calculated as:

325

$$326 \quad \text{NaOH addition} = 500 - \text{NaHCO}_3 \text{ addition} \quad (3)$$

327

328 Where 500 is the targeted alkalinity enhancement in  $\mu\text{mol/kg}$ .  $\text{NaHCO}_3$  and NaOH additions  
329 need to be multiplied with the weight of the enclosed microcosm seawater to calculate how  
330 much  $\text{NaHCO}_3$  and NaOH need to be added per individual microcosm.

331 It is recommended to use 1 molar stock solutions for both  $\text{NaHCO}_3$  and NaOH for treatment  
332 manipulations because in that case required additions in  $\mu\text{mol/microcosm}$  are equivalent to  
333  $\mu\text{L/microcosm}$ . For example, in the equilibrated treatment a typical addition would be 420  
334  $\mu\text{L/kg}$  of  $\text{NaHCO}_3$  and 80  $\mu\text{L/kg}$  of NaOH (i.e., 22.89 mL/microcosm  $\text{NaHCO}_3$  and 4.36  
335 mL/microcosm NaOH when 54.5 kg of seawater were enclosed). One molar  $\text{NaHCO}_3$  stock  
336 solutions can be prepared by dissolving 8.4 g  $\text{NaHCO}_3$  powder (dried at  $60^\circ\text{C}$  overnight; note  
337 that  $\text{NaHCO}_3$  decomposes at higher temperatures) in 100 mL deionised water. One molar  
338 NaOH (ideally in “analytical quality”) should be purchased as such.

339 The addition of NaOH to seawater causes precipitation of brucite ( $\text{Mg}(\text{OH})_2$ ), which appears  
340 as white flakes (Fig. 1C). The brucite flakes bind the alkalinity added via NaOH in particulate  
341 form and need to be re-dissolved so that dissolved alkalinity is increased by the intended 500  
342  $\mu\text{mol/kg}$ . Furthermore, brucite formation can precipitate phosphates (Karl and Tien, 1992).  
343 This must be avoided as the loss of phosphate from the dissolved phase in the treatments  
344 would be a problematic confounding factor. The formation of brucite will be particularly  
345 pronounced in the unequilibrated treatment where all alkalinity is added as NaOH. To dissolve  
346 all brucite, microcosms should be gently stirred with a clean plastic paddle during and after  
347 NaOH additions until all white flakes disappear. Our previous experiments resembling the  
348 OAEPIIP approach (Ferderer et al., 2022) revealed that dissolution of all brucite by gentle  
349 stirring leads to the desired outcome, i.e., alkalinity was increased by 500  $\mu\text{mol/kg}$  and no  
350 phosphate was lost. For consistency, control and equilibrated microcosms should be stirred  
351 as much as the unequilibrated microcosms. If OAEPIIP participants do not have prior practical  
352 experience with seawater carbonate chemistry manipulation, it is advised to test the above-  
353 mentioned procedures (including the measurement of resulting carbonate chemistry  
354 parameter changes such as in TA and DIC) before commencing the main OAEPIIP experiment.

355

356

## 2.6. Essential parameters to be measured in OAEPIIP experiments

357

358 Next to an identical experimental design and setup, the same parameters need to be  
359 measured in individual OAEPIIP experiments to make them comparable (Iglesias-Rodríguez et  
360 al., 2023). A list of “core” parameters with justifications for their choice is provided in Table  
361 1, and additional recommendations on how to sample and process these is provided in Table  
362 S2. The core parameters (Table 1) should provide a relatively comprehensive, yet cost-  
363 efficient insight into processes within the plankton community. Although all core parameters  
364 need to be measured in all participating OAEPIIP studies, there may be unsurmountable  
365 logistical constraints which prohibit a participant from determining a core parameter. Such  
366 cases should be mentioned upon application for OAEPIIP participation so that mitigation  
367 pathways can be explored and that potential participants with less infrastructure capacity still  
368 have the opportunity to participate if possible (see also section 4).

369 If they wish to do so, OAEPIIP participants can also measure additional parameters to  
370 maximise their individual experimental outcomes. However, the following issues should be  
371 considered:

372

373 1) Not more than approximately 1/3 of the microcosm volume should be sampled over  
374 the course of the study (a) to limit the build-up of a headspace and (b) to avoid too  
375 much heat input per Liter of enclosed volume via the heat belts (the room  
376 temperature might need lowering to compensate for reducing volume throughout the  
377 experiment; section 2.2).

378 2) Any type of contamination (particulate or dissolved organic or inorganic) must be kept  
379 at a minimum.

380 3) It is possible to sample mesozooplankton with a customized net (Guo et al., 2023), but  
381 sampling should be restricted to 3 occasions during the experiment (e.g., beginning,  
382 middle, end) to avoid overfishing.

383 4) Aggregation and sedimentation are often observed in these microcosm studies and it  
384 is encouraged to sample settling materials from the sediment trap (Ferderer et al.,  
385 2022). However, care must be taken to not remove significant volumes of seawater.

386

387 **Table 1.** List of core parameters that essentially need to be measured in all individual OAEPIIP  
 388 studies. The “Samplings” column indicates how often all 9 microcosms need to be sampled for  
 389 a specific parameter during the study. “Daily” means that this parameter needs to be  
 390 measured every day, irrespective of the temperature-dependent duration of the study. “b/e”  
 391 means that samples need to be taken at the beginning and the end of the experiment.

Core parameter	Rationale	Samplings
Alkalinity	The treatment-defining parameter of the study.	7*
Second carbonate chemistry parameter (e.g., pH or DIC)	Required to constrain the carbonate system. Also provides insights for net autotrophy/heterotrophy.	Daily*
Salinity	Required to define the marine system under investigation.	b/e
Light	To constrain physical conditions for growth.	b/e
Temperature	To monitor its influence on metabolic rates and assess temperature stability due to convective mixing.	daily
Nutrients (NO <sub>x</sub> <sup>-</sup> , PO <sub>4</sub> <sup>3-</sup> , Si(OH) <sub>4</sub> )	Nitrate+Nitrite (NO <sub>x</sub> <sup>-</sup> ) and phosphate (PO <sub>4</sub> <sup>3-</sup> ) availability largely determines the productivity of the plankton community. Availability of Si(OH) <sub>4</sub> provides insights if productivity will likely be driven by diatoms.	11
Chlorophyll a (chl <sub>a</sub> )	Chl <sub>a</sub> is a widely used proxy for phytoplankton biomass	11
Particulate organic carbon and nitrogen (POC and PON)	POC and PON dynamics are related to the increase and decline of biomass. Their ratio (POC/PON) is an important metric in biogeochemical element cycling.	11
Biogenic silica (BSi)	BSi is a widely used proxy for diatom biomass	11
Flow cytometry (FC)	FC is a cost-efficient tool that reveals shifts in phytoplankton size classes and specific groups with distinguishable fluorescence/scatter characteristics. FC is particularly good for enumeration of small phytoplankton and heterotrophic bacteria.	11
Microscopy	Microscopy is a widely available tool to assess dynamics in phytoplankton and microzooplankton communities. It is complementary to FC as it is better suited for larger phytoplankton/microzooplankton.	7
Nucleic acid sample	Nucleic acid samples (DNA and possibly RNA) will provide a detailed assessment of microbial diversity. Basic requirements for this parameter will be metabarcoding for 16S rRNA genes (variable region of V4-V5). Further analysis for metagenomics and metatranscriptomics will be possible depending on the timing of sample collection but are not essential for the participation.	b/e

392 \*These parameters must be sampled directly before and after establishment of the OAE  
 393 treatments in all 9 microcosms. All other parameters must be sampled for the first time after  
 394 establishment of the treatments.

395

## 396 **2.7. Duration of experiment**

397

398 To the best of our knowledge, there is no general rule for the ideal duration of microcosm  
399 experiments. Experiments that are too short may miss important responses of plankton  
400 communities while long experiments may exacerbate so-called “bottle effects”, non-specific  
401 effects from confinement rather than the experimental perturbation itself (Pernthaler and  
402 Amann, 2005). For example, 3 days of experiment may be too short to observe a  
403 differentiation of plankton species composition between treatments, while a community  
404 contained for 2 months could be dominated by those that best survive in a laboratory  
405 environment. Based on experiments with the OAEPIIP setup in Tasmania we consider 20 days  
406 as a good compromise for an experiment at 15°C. However, metabolic rates increase with  
407 temperature so that experimental duration needs to be adjusted based on respective  
408 locations. Informed by Q10 temperature dependencies (Sherman et al., 2016), we  
409 recommend the following framework: 20 days is the reference duration at 15°C. The duration  
410 (in days) increases/decreases from this reference point using Q10 kinetics:

411

$$412 \quad Duration = \frac{0.5611}{\left(0.5611 \times 1.47^{\frac{T_{exp} - 15}{10}}\right)} \times 20 \quad (4)$$

413

414 where 0.5611 is the reference growth rate at 15°C,  $T_{exp}$  is the anticipated temperature in the  
415 OAEPIIP experiment, and 1.47 is the  $Q_{10}$  factor derived by (Sherman et al., 2016). For example,  
416 an experiment at 25°C should last for 14 days and an experiment at 5°C for 29 days.

417

## 418 **2.8. Sampling operations and logistics**

419

420 All microcosm incubators shall be closed after the filling procedure with the black screw cap  
421 (Fig. 1) and kept closed over the course of the experiment except during the establishment of  
422 treatments (section 2.5) and sampling. The enclosed headspace (Fig. 1) may vary slightly in  
423 between microcosms after the filling procedure (section 2.1) and will increase over the course  
424 of the experiment due to the withdrawal of samples. While an increasing headspace will lead  
425 to some CO<sub>2</sub> exchange between the atmosphere and the enclosed volume, previous studies

426 with the same setting found that this has no effect on the OAE treatments established in the  
427 experiments (Guo et al., 2023; Ferderer et al., 2022). The convective system mixes the water  
428 column so that no manual mixing is needed prior to sampling. A peristaltic pump is  
429 recommended to withdraw the seawater samples from the microcosms.

430 The total number of samplings for specific parameters is listed in Table 1 (for example, POC  
431 and PON need to be sampled 11 times in total). The frequency of sampling needs to be  
432 adjusted based on the temperature-dependent duration of the experiment (section 2.7.).  
433 OAEPIIP experiments at higher temperatures require higher sampling frequency because  
434 metabolic processes are faster. Table 1 lists the minimum number of days a parameter should  
435 be sampled. This number is to guarantee that there will be enough comparable data points  
436 across OAEPIIP experiments. For example, nutrient samples should be taken at least 11 times  
437 in each microcosm during the experiment. For an experiment at 15°C (20 days), this could  
438 mean a sampling on day 0 (directly after establishing treatments) and then days 2, 4, 6,...,20.  
439 However it may also be reasonable to increase frequency during periods of phytoplankton  
440 blooms (e.g., daily) and then reduce the frequency (e.g. every 4 days) when nutrients are  
441 depleted. In general, OAEPIIP experimentalists can best decide on an individual basis what  
442 sampling schedule is most appropriate for their experiment, but the total number of  
443 samplings must be at least as defined in Table 1 for each of the listed parameters.  
444 Sampling for all OAEPIIP experiments should begin two hours after the onset of the light  
445 period on a sampling day. This coordination of initial sampling ensures that the plankton  
446 community is in a similar diurnal growth state. Hence, sampling of all 9 microcosms should  
447 ideally not last longer than 3 hours.

448

## 449 **2.9. Statistical analyses**

450

451 Microcosm data contains complex ecological data which require specific (often complicated)  
452 statistical tools for their analysis. A common issue is the presence of non linear relationships,  
453 which without gross transformation of the variables prevents the fitting of data to linear  
454 models. Furthermore, OAEPIIP microcosms will be sampled several times over an extended  
455 period. This sampling strategy results in temporal-pseudoreplication, where observations are  
456 not independent of each other and therefore violate the assumption of independence  
457 required for simple linear models and Generalised additive models (GAMs) (Zuur et al., 2009;



458 Wood, 2017). The expansion of GAMs to Generalized Additive Mixed Models (GAMMs) allows  
459 for correlations between observations and the modelling of data structures which are nested  
460 as well as for non-linear relationships between the response and explanatory variables.  
461 To facilitate and standardize statistical analyses of individual datasets we provide an R-based  
462 pipeline (OAEPIIP, 2024). This pipeline is tailored towards the evaluation of individual OAEPIIP  
463 data sets using GAMMs. The files contain a workflow which demonstrates the use of GAMMs  
464 and facilitates the seamless integration of individual datasets gathered during OAEPIIP  
465 experiments into the workflow. Theoretical background, knowledge and details on how to fit  
466 such models can be found in the textbooks by Zuur et al. (2009) and Wood (2017).

467

### 468 **3. Logistics and administration**

469

470 Basic instructions and updates on OAEPIIP will be provided on the OAEPIIP website  
471 ([https://appliedbgc.imas.utas.edu.au/ocean-alkalinity-enhancement-pelagic-impact-  
472 intercomparison-project/](https://appliedbgc.imas.utas.edu.au/ocean-alkalinity-enhancement-pelagic-impact-intercomparison-project/)).

473

#### 474 **3.1. Eligibility and funding**

475

476 To join OAEPIIP, participants need to be capable of performing an OAEPIIP study and provide  
477 all data by December 2025. This capacity shall be confirmed on a simple 1 page form (available  
478 on the OAEPIIP website) that potential participants need to fill in and send to the email  
479 provided on the form. Career stage, publication record, or other parameters of a scientist's  
480 curriculum vitae have no relevance for OAEPIIP. As such, application success is determined by  
481 logistical and infrastructure-related aspects, for example whether a participant has access to  
482 a temperature-controlled room and can provide the various data in the given timeframe (but  
483 see also section 4 on suggestions on how to mitigate individual limitations to infrastructure).  
484 Ultimately, participation is restricted by total funding available to OAEPIIP. Should there be  
485 more applications than there is funding, participants will be selected based on two criteria:  
486 First, we will consider the locations of the experiments to obtain the best possible geographic  
487 spread. Second, participants will be selected by chance should there be clusters of  
488 applications in close proximity.

489 OAEPIIP provides a maximum of around 12,000 US\$ per study in materials and funding for  
490 analytical costs and publication fees (the exact amount is slightly variable due to exchange  
491 rates). All materials and standardized components like the microcosms shall be purchased by  
492 the individual participants with ~10,400 US\$ made available to them for the experiments. The  
493 OAEPIIP administration will provide all necessary information for the purchase of  
494 standardized components so that all experiments are conducted in the same type of  
495 incubators. The remaining ~1,600 US\$ will be retained by the OAEPIIP administration and  
496 made available to support the publication of individual OAEPIIP studies(see section 3.2).  
497 The funds for materials and standardised components will be transferred via invoicing. Thus,  
498 participants must have a bank account associated with their affiliation to which funding can  
499 be transferred from Australia. This criterion therefore excludes laboratories in countries  
500 under relevant sanctions from Australia to receive funding, although they are still welcome  
501 to be part of the OAEPIIP community. In practice, participants will send two invoices to the  
502 University of Tasmania, one at the beginning of the experiment to support purchasing of  
503 materials and standardized components and the second one towards the end when the data  
504 is available and has been submitted. OAEPIIP cannot provide funding for salaries. Therefore,  
505 the experiment was designed to be suitable for a Master thesis or a chapter of a PhD thesis.

506

### 507 **3.2. Data management and publication**

508

509 Datasets of individual OAEPIIP studies should be formatted using a standardised template  
510 available on the OAEPIIP homepage (section 3) and submitted to OAEPIIP as soon as they are  
511 available. All data must be uploaded and made available under open access. Participants will  
512 be listed on the OAEPIIP homepage and their individual datasets will be linked to their names  
513 and affiliations as soon as it is made available. OAEPIIP experiments shall be published on an  
514 individual basis in an OAEPIIP special issue (publication fees of up to 1600 US\$ are provided  
515 by OAEPIIP). Individual publication will enable identification of novel observations on how  
516 plankton communities respond to OAE. If participants prefer not to publish their data, they  
517 still need to submit their data to OAEPIIP so that it can be included in the OAEPIIP synthesis.  
518 This is critically important because the synthesis must avoid publication bias.

519 The OAEPIIP synthesis will be prepared once all datasets have been delivered. First and last  
520 authors of individual studies will automatically be co-authors on the synthesis publication(s)  
521 at the end of the project, unless they prefer not to be.

522

#### 523 **4. Capacity building and inclusivity**

524

525 OAEPIIP has potential benefits that go beyond scientific knowledge gain. The community  
526 effort helps to build a network of OAE scientists and provides an incentive and access to those  
527 who have not yet engaged with OAE research. Indeed, growing the OAE research community  
528 is essential to accelerate the OAE assessment and make it more comprehensive. Providing the  
529 same amount of funding, regardless of the location, may increase the attractiveness of  
530 OAEPIIP studies to those that currently have less funding. Participation of scientists worldwide  
531 is what we aim for since the OAE assessment requires the inclusion of the global community.  
532 Indeed, participation in the process of assessing marine CDR methods (such as OAE), rather  
533 than being on the receiving end of information only, has been expressed as an important  
534 aspect by stakeholders from developing countries.

535 We are aware that the infrastructure demands for OAEPIIP (section 2), still put barriers on  
536 participation. To mitigate those barriers, potential participants from more experienced  
537 laboratories can offer to serve as a partner for a less experienced laboratory. Likewise,  
538 potential participants from less experienced laboratories can indicate if they essentially need  
539 support from an experienced laboratory. This information shall be disclosed on the  
540 application form (available on the OAEPIIP website) so that OAEPIIP can establish  
541 partnerships between participants. Partners can support each other through knowledge  
542 exchange but also more practically by analysing samples for each other. For example, if an  
543 interested participant has no capacity to measure alkalinity or flow cytometry samples, it may  
544 partner with another participant to share analytical duties. The distribution of funding for  
545 analytical costs via invoicing allows for such flexibility as it provides an opportunity to easily  
546 re-distribute funding between project participants when this is communicated with the  
547 OAEPIIP administration. For example, when two laboratories partner, they together have  
548 access to twice the funding (~24,000 US\$), which they share among them for the two  
549 experiments they would have to do (the two experiments must be at different locations to  
550 guarantee geographical diversity).

551 Furthermore, potential participants that simply have no chance to measure one (or more)  
552 core parameters due to unsurmountable logistical constraints can still hand in an application,  
553 if they indicate which parameters they are unable to deliver on their application form  
554 (available on the OAEPIIP website). The OAEPIIP administration will then evaluate such  
555 applications on a case-by-case basis and explore if there is a way for participation despite this  
556 limitation. This pathway is set in place specifically for potential participants with less  
557 developed infrastructure and less capacity for collaboration with an experienced (e.g., due to  
558 geographic isolation).

559 Altogether, we hope the cost-efficient design of OAEPIIP, its eligibility criteria that refrain  
560 from classic measures of scientific success, and potential support via an evolving OAEPIIP  
561 community could promote an inclusive assessment of OAE. One primary goal of OAEPIIP is  
562 capacity building to provide more informed decisions concerning OAE that encompass data  
563 from a geographically diverse range of plankton ecosystems.

564

#### 565 **Competing interests**

566

567 The contact author has declared that none of the authors has any competing interests.

568

#### 569 **Acknowledgements**

570

571 We thank the Carbon-to-Sea Initiative for funding OAEPIIP. LTB acknowledges funding from  
572 the Australian Research Council by Future Fellowship (FT200100846).

573

#### 574 **Data availability**

575

576 All code provided for experimental design and statistical analysis can be found here: (OAEPIIP,  
577 2024).

578

#### 579 **References**

580

581 Bach, L. T., Gill, S. J., Rickaby, R. E. M., Gore, S., and Renforth, P.: CO2 Removal With  
582 Enhanced Weathering and Ocean Alkalinity Enhancement: Potential Risks and Co-benefits

583 for Marine Pelagic Ecosystems, *Frontiers in Climate*, 1, 2019.  
584  
585 Benton, T. G., Solan, M., Travis, J. M. J., and Sait, S. M.: Microcosm experiments can inform  
586 global ecological problems, *Trends in Ecology & Evolution*, 22, 516–521,  
587 <https://doi.org/10.1016/j.tree.2007.08.003>, 2007.  
588  
589 Burt, D. J., Fröb, F., and Ilyina, T.: The Sensitivity of the Marine Carbonate System to  
590 Regional Ocean Alkalinity Enhancement, *Frontiers in Climate*, 3, 2021.  
591  
592 Dingley, B., O'Rourke, E., Turner, B., CMIP Panel Members, WGCM Infrastructure Panel  
593 Members, and CMIP7 Task Team Members: CMIP Annual Report 2022-2023,  
594 <https://doi.org/10.5281/zenodo.8101810>, 2023.  
595  
596 Eisaman, M. D., Geilert, S., Renforth, P., Bastianini, L., Campbell, J., Dale, A. W., Foteinis, S.,  
597 Grasse, P., Hawrot, O., Löscher, C. R., Rau, G. H., and Rønning, J.: Assessing the technical  
598 aspects of ocean-alkalinity-enhancement approaches, *State of the Planet*, 2-oae2023, 1–29,  
599 <https://doi.org/10.5194/sp-2-oae2023-3-2023>, 2023.  
600  
601 Fakhraee, M., Li, Z., Planavsky, N. J., and Reinhard, C. T.: A biogeochemical model of mineral-  
602 based ocean alkalinity enhancement: impacts on the biological pump and ocean carbon  
603 uptake, *Environ. Res. Lett.*, 18, 044047, <https://doi.org/10.1088/1748-9326/acc9d4>, 2023.  
604  
605 Ferderer, A., Chase, Z., Kennedy, F., Schulz, K. G., and Bach, L. T.: Assessing the influence of  
606 ocean alkalinity enhancement on a coastal phytoplankton community, *Biogeosciences*, 19,  
607 5375–5399, <https://doi.org/10.5194/bg-19-5375-2022>, 2022.  
608  
609 Field, A. P. and Gillett, R.: How to do a meta-analysis, *British Journal of Mathematical and*  
610 *Statistical Psychology*, 63, 665–694, <https://doi.org/10.1348/000711010X502733>, 2010.  
611  
612 Gattuso, J.-P., Epitalon, J.-M., Lavigne, H., and Orr, J. C.: Seacarb: seawater carbonate  
613 chemistry. R package version 3.3.0., , CRAN, 2021.  
614  
615  
616 Guo, J. A., Strzepek, R. F., Swadling, K. M., Townsend, A. T., and Bach, L. T.: Influence of  
617 ocean alkalinity enhancement with olivine or steel slag on a coastal plankton community in  
618 Tasmania, *Biogeosciences*, 21, 2335–2354, <https://doi.org/10.5194/bg-21-2335-2024>, 2024.  
619  
620 Hamm, T., Barkhau, J., Gabriel, A.-L., Gottschalck, L. L., Greulich, M., Houiller, D., Kawata, U.,  
621 Tump, L. N., Leon, A. S., Vasconcelos, P., Yap, V., Almeida, C., Chase, Z., Hurd, C. L., Lavers, J.  
622 L., Nakaoka, M., Rilov, G., Thiel, M., Wright, J. T., and Lenz, M.: Plastic and natural inorganic  
623 microparticles do not differ in their effects on adult mussels (*Mytilidae*) from different  
624 geographic regions, *Science of The Total Environment*, 811, 151740,  
625 <https://doi.org/10.1016/j.scitotenv.2021.151740>, 2022.  
626  
627 Harrison, F.: Getting started with meta-analysis, *Methods in Ecology and Evolution*, 2, 1–10,  
628 <https://doi.org/10.1111/j.2041-210X.2010.00056.x>, 2011.  
629

630 He, J. and Tyka, M. D.: Limits and CO<sub>2</sub> equilibration of near-coast alkalinity enhancement,  
631 Biogeosciences, 20, 27–43, <https://doi.org/10.5194/bg-20-27-2023>, 2023.

632

633 Ho, D. T., Bopp, L., Palter, J. B., Long, M. C., Boyd, P. W., Neukermans, G., and Bach, L. T.:  
634 Monitoring, reporting, and verification for ocean alkalinity enhancement, State of the  
635 Planet, 2-oae2023, 1–12, <https://doi.org/10.5194/sp-2-oae2023-12-2023>, 2023.

636

637 Iglesias-Rodríguez, M. D., Rickaby, R. E. M., Singh, A., and Gately, J. A.: Laboratory  
638 experiments in ocean alkalinity enhancement research, State of the Planet, 2-oae2023, 1–  
639 18, <https://doi.org/10.5194/sp-2-oae2023-5-2023>, 2023.

640

641 Jones, D. C., Ito, T., Takano, Y., and Hsu, W.-C.: Spatial and seasonal variability of the air-sea  
642 equilibration timescale of carbon dioxide, Global Biogeochemical Cycles, 28, 1163–1178,  
643 <https://doi.org/10.1002/2014GB004813>, 2014.

644

645 Karl, D. M. and Tien, G.: MAGIC: A sensitive and precise method for measuring dissolved  
646 phosphorus in aquatic environments, Limnology and Oceanography, 37, 105–116,  
647 <https://doi.org/10.4319/lo.1992.37.1.0105>, 1992.

648

649 Köhler, P., Abrams, J. F., Völker, C., Hauck, J., and Wolf-Gladrow, D. A.: Geoengineering  
650 impact of open ocean dissolution of olivine on atmospheric CO<sub>2</sub>, surface ocean pH and  
651 marine biology, Environ. Res. Lett., 8, 014009, <https://doi.org/10.1088/1748-9326/8/1/014009>, 2013.

652

653

654 de Lannoy, C.-F., Eisaman, M. D., Jose, A., Karnitz, S. D., DeVaul, R. W., Hannun, K., and  
655 Rivest, J. L. B.: Indirect ocean capture of atmospheric CO<sub>2</sub>: Part I. Prototype of a negative  
656 emissions technology, International Journal of Greenhouse Gas Control, 70, 243–253,  
657 <https://doi.org/10.1016/j.ijggc.2017.10.007>, 2018.

658

659 Lenton, A., Matear, R. J., Keller, D. P., Scott, V., and Vaughan, N. E.: Assessing carbon dioxide  
660 removal through global and regional ocean alkalization under high and low emission  
661 pathways, Earth System Dynamics, 9, 339–357, <https://doi.org/10.5194/esd-9-339-2018>,  
662 2018.

663

664 Masson-Delmotte, V., Zhai, P., Pirani, A., Connors, S. L., Péan, C., Berger, S., Caud, N., Chen,  
665 Y., Goldfarb, L., Gomis, M. I., Huang, M., Leitzell, K., Lonnoy, E., Matthews, J. B. R., Maycock,  
666 T. K., Waterfield, T., Yelekçi, Ö., Yu, R., and Zhou, B. (Eds.): Climate Change 2021: The  
667 Physical Science Basis. Contribution of Working Group I to the Sixth Assessment Report of  
668 the Intergovernmental Panel on Climate Change, Cambridge University Press, Cambridge,  
669 United Kingdom and New York, NY, USA, <https://doi.org/10.1017/9781009157896>, 2021.

670

671 Mu, L., Palter, J. B., and Wang, H.: Considerations for hypothetical carbon dioxide removal  
672 via alkalinity addition in the Amazon River watershed, Biogeosciences, 20, 1963–1977,  
673 <https://doi.org/10.5194/bg-20-1963-2023>, 2023.

674

675 OAEPIIP: Additional resources on carbonate chemistry calculation and statistical evaluations  
676 of OAEPIIP, 2024.

677  
678 Oschlies, A., Bach, L. T., Rickaby, R. E. M., Satterfield, T., Webb, R., and Gattuso, J.-P.:  
679 Climate targets, carbon dioxide removal, and the potential role of ocean alkalinity  
680 enhancement, *State of the Planet*, 2-oae2023, 1–9, [https://doi.org/10.5194/sp-2-oae2023-](https://doi.org/10.5194/sp-2-oae2023-1-2023)  
681 [1-2023](https://doi.org/10.5194/sp-2-oae2023-1-2023), 2023.  
682  
683 Pernthaler, J. and Amann, R.: Fate of Heterotrophic Microbes in Pelagic Habitats: Focus on  
684 Populations, *Microbiology and Molecular Biology Reviews*, 69, 440–461,  
685 <https://doi.org/10.1128/membr.69.3.440-461.2005>, 2005.  
686  
687 Pierrot, D., Epitalon, J.-M., Orr, J. C., Lewis, E., and Wallace, D. W. R.: MS Excel program  
688 developed for CO<sub>2</sub> system calculations – version 3.0, , GitHub repository, 2021.  
689  
690 Riebesell, U., Schulz, K. G., Bellerby, R. G. J., Botros, M., Fritsche, P., Meyerhöfer, M., Neill,  
691 C., Nondal, G., Oschlies, A., Wohlers, J., and Zöllner, E.: Enhanced biological carbon  
692 consumption in a high CO<sub>2</sub> ocean, *Nature*, 450, 545–548,  
693 <https://doi.org/10.1038/nature06267>, 2007.  
694  
695 Riebesell, U., Czerny, J., von Bröckel, K., Boxhammer, T., Büdenbender, J., Deckelnick, M.,  
696 Fischer, M., Hoffmann, D., Krug, S. A., Lentz, U., Ludwig, A., Mücke, R., and Schulz, K. G.:  
697 Technical Note: A mobile sea-going mesocosm system – new opportunities for ocean  
698 change research, *Biogeosciences*, 10, 1835–1847, [https://doi.org/10.5194/bg-10-1835-](https://doi.org/10.5194/bg-10-1835-2013)  
699 [2013](https://doi.org/10.5194/bg-10-1835-2013), 2013.  
700  
701 Schulz, K. G., Bach, L. T., and Dickson, A. G.: Seawater carbonate chemistry considerations  
702 for ocean alkalinity enhancement research: theory, measurements, and calculations, *State*  
703 *of the Planet*, 2-oae2023, 1–14, <https://doi.org/10.5194/sp-2-oae2023-2-2023>, 2023.  
704  
705 Sherman, E., Moore, J. K., Primeau, F., and Tanouye, D.: Temperature influence on  
706 phytoplankton community growth rates, *Global Biogeochemical Cycles*, 30, 550–559,  
707 <https://doi.org/10.1002/2015GB005272>, 2016.  
708  
709 Stewart, R. I. A., Dossena, M., Bohan, D. A., Jeppesen, E., Kordas, R. L., Ledger, M. E.,  
710 Meerhoff, M., Moss, B., Mulder, C., Shurin, J. B., Suttle, B., Thompson, R., Trimmer, M., and  
711 Woodward, G.: Chapter Two - Mesocosm Experiments as a Tool for Ecological Climate-  
712 Change Research, in: *Advances in Ecological Research*, vol. 48, edited by: Woodward, G. and  
713 O’Gorman, E. J., Academic Press, 71–181, [https://doi.org/10.1016/B978-0-12-417199-](https://doi.org/10.1016/B978-0-12-417199-2.00002-1)  
714 [2.00002-1](https://doi.org/10.1016/B978-0-12-417199-2.00002-1), 2013.  
715  
716 Taucher, J., Boxhammer, T., Bach, L. T., Paul, A. J., Schartau, M., Stange, P., and Riebesell, U.:  
717 Changing carbon-to-nitrogen ratios of organic-matter export under ocean acidification, *Nat.*  
718 *Clim. Chang.*, 11, 52–57, <https://doi.org/10.1038/s41558-020-00915-5>, 2021.  
719  
720 Wang, H., Pilcher, D. J., Kearney, K. A., Cross, J. N., Shugart, O. M., Eisaman, M. D., and  
721 Carter, B. R.: Simulated Impact of Ocean Alkalinity Enhancement on Atmospheric CO<sub>2</sub>  
722 Removal in the Bering Sea, *Earth’s Future*, 11, e2022EF002816,  
723 <https://doi.org/10.1029/2022EF002816>, 2023.

724

725 Wood, S. N.: Generalized Additive Models: An Introduction with R, Second Edition, 2nd ed.,  
726 Chapman and Hall/CRC, Boca Raton, 496 pp., <https://doi.org/10.1201/9781315370279>,  
727 2017.

728

729 Zeebe, R. E. and Wolf-Gladrow, D.: CO<sub>2</sub> in Seawater: Equilibrium, Kinetics, Isotopes, Gulf  
730 Professional Publishing, 382 pp., 2001.

731

732 Zuur, A. F., Ieno, E. N., Walker, N., Saveliev, A. A., and Smith, G. M.: Mixed effects models  
733 and extensions in ecology with R, Springer, New York, NY, [https://doi.org/10.1007/978-0-](https://doi.org/10.1007/978-0-387-87458-6)  
734 [387-87458-6](https://doi.org/10.1007/978-0-387-87458-6), 2009.