



Global estimates of reactive nitrogen components during 2000-2100 1 based on the multi-stage model 2 3 Rui Li^{a, b*}, Yining Gao^a, Lijia Zhang^a, Yubing Shen^a, Gehui Wang^a ^a Key Laboratory of Geographic Information Science of the Ministry of Education, School of 4 5 Geographic Sciences, East China Normal University, Shanghai, 200241, PR China 6 ^b Institute of Eco-Chongming (IEC), 20 Cuiniao Road, Chenjia Town, Chongming District, 7 Shanghai, 202162, China 8 * Corresponding author 9 Prof. Li (rli@geo.ecnu.edu.cn) 10 Abstract 11 High contents of reactive nitrogen components aggravate air pollution and could also impact 12 ecosystem structure and function across the terrestrial-aquatic-marine continuum. However, the 13 long-term historical trends and future prediction of reactive nitrogen components at the global scale 14 still remains high uncertainties. In our study, the field observations, satellite products, model output, 15 and many other covariates were integrated into the machine-learning model to capture the global 16 patterns of reactive nitrogen components during 2000-2019. In order to decrease the estimate 17 uncertainties in the future scenarios, the constructed reactive nitrogen component dataset during the 18 historical period was then utilized as the constraint to calibrate the CMIP6 dataset in four scenarios. The results suggested the cross-validation (CV) R^2 values of four species showed satisfied 19 performance ($R^2 > 0.55$). The concentrations of estimated reactive nitrogen components in China 20 experienced persistent increases during 2000-2013, while they suffered from drastic decreases since 21 22 2013 except NH₃. It might be associated with the impact of clean air policy. However, these 23 compounds in Europe and the United States remained relatively stable since 2000. In the future scenarios, SSP3-7.0 (traditional energy scenario) and SSP1-2.6 (carbon neutrality scenario) showed 24 25 the highest and lowest reactive nitrogen component concentrations, respectively. Although the 26 reactive nitrogen concentrations in some heavy-pollution scenarios (SSP3-7.0) also experienced decreases during 2020-2100, SSP1-2.6 and SSP2-4.5 (middle emission scenario) still kept more 27 28 rapid decreasing trends. Our results emphasize the need for carbon-neutrality pathway to reduce 29 global atmospheric N pollution.





30 **1. Introduction**

Along with the development of global urbanization and industrialization, the anthropogenic 31 emissions of reactive nitrogen (e.g., NOx, NH3, HNO3) experienced drastic increases during the past 32 33 decades (Chen et al., 2021; Liu et al., 2020b; McDuffie et al., 2020). The reactive nitrogen released 34 from anthropogenic source could significantly alter the global nitrogen cycle throughout the Earth 35 system (Altieri et al., 2021; Zhang et al., 2020). Reactive nitrogen in the atmosphere dominates the chemical formation of tropospheric O3 and aggravates the particle pollution (Geddes and Martin, 36 37 2017), with implications for global air quality and climate change (He et al., 2022; Von 38 Schneidemesser et al., 2015). Moreover, the ambient reactive nitrogen could be deposited into the 39 land surface and could cause lake eutrophication and soil acidification (Bouwman et al., 2002; Chen et al., 2018). Therefore, it is highly necessary to understand the spatial distributions and temporal 40 41 evolution trends of reactive nitrogen components at the global scale.

42 Despite the global importance, observational constraints on reactive nitrogen in the atmosphere 43 were still scarce in most parts of the world (Liu et al., 2020b). Furthermore, the majority of 44 monitoring sites focused on China, Europe, and the United States (Du et al., 2014; Li et al., 2020; 45 Li et al., 2019a; Li et al., 2016), and these uneven sites only possessed limited spatial 46 representativeness (Shi et al., 2018), which restricted the accurate assessment of global reactive 47 nitrogen pollution. Fortunately, the satellite observations gave us a unprecedent chance to capture the global variations of atmospheric reactive nitrogen. Geddes et al. (2017) used the satellite 48 49 products to calibrate the simulated reactive nitrogen oxides (NO_y) and improved the predictive 50 performance (R = 0.83) compared with the chemical transport model (CTM) output alone (Geddes 51 and Martin, 2017). Afterwards, Liu et al. (2022) also used the similar method to estimate the global 52 wet deposition of reduced nitrogen (NH_4^+) and the R value achieved 0.80(Liu et al., 2021). Although 53 the calibration based on satellite products could improve the predictive accuracy compared with CTM output, the simulated values still largely biased from the ground-level observations. Moreover, 54 55 the method cannot accurately fill the gaps of reactive nitrogen concentrations without satellite coverage. In our previous works, we developed a satellite-based ensemble machine-learning model 56 57 to predict the wet NH₄⁺ deposition across China and the R² value reached 0.76 (R = 0.88) (Li et al., 58 2020). However, this technique was not expanded to the global scale and the high-accuracy and full-

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60	Apart from the historical estimates, the future prediction of reactive nitrogen is also important
61	because these components in the future scenarios could significantly affect the land carbon cycle
62	and greenhouse gas emissions (Chen et al., 2015; Zaehle, 2013). To the best of our knowledge, only
63	two studies focused on global aerosol prediction in the future scenarios. Chen et al. (2023) predicted
64	the global $PM_{2.5}$ levels and associated mortalities in 2100 under different climate scenarios and
65	found that SSP3-7.0 scenario was linked with the highest $PM_{2.5}$ exposure. Li et al., (2022) also
66	simulated the global $\mathrm{NO}_3^{\text{-}}$ and $\mathrm{NH}_4^{\text{+}}$ levels in the future four scenarios and demonstrated that both
67	of these components showed marked decreases in most cases except SSP5-8.5 scenario. However,
68	this study predicted the future reactive nitrogen based on historical CTM output alone, which lacks
69	of observation constraints. The result might increase the uncertainty of assessment.
70	In our study, we developed a multi-stage model to estimate the concentrations of four reactive
71	nitrogen species (NO ₃ ⁻ , HNO ₃ , NH ₃ , and NH ₄ ⁺) during 2000-2019 because these species were most
72	important reactive nitrogen components for human health and ecological ecosystem and also
73	showed abundant ground-level observations. Then, the species over the 2020-2100 period under the
74	SSP1-2.6, SSP2-4.5, SSP3-7.0, and SSP5-8.5 scenarios were also corrected based on the historical
75	estimates. Finally, the long-term dataset of reactive nitrogen during 2000-2100 was constructed. Our
76	results were beneficial to assess the impacts of reactive nitrogen components on air pollution and
77	climate change in the future.

coverage global ambient reactive nitrogen dataset was still lack.

78 2. Material and methods

79 2.1 Reactive nitrogen observations

80 Most of reactive nitrogen observations focused on East Asia, Europe, and the United States. The 81 monthly reactive nitrogen components monitoring data during 2010-2015 in China were downloaded from nationwide nitrogen deposition monitoring network (NNDMN) including 32 sites, 82 and these sites could be classified into three types mixed with urban, rural, and background sites 83 84 (Xu et al., 2019) (Table S1). The concentrations of reactive nitrogen components were determined 85 using the active DEnuder for Long-Term Atmospheric sampling system (DELTA). The detailed 86 sampling and analysis procedures have been described by Xu et al. (2019). The dataset of reactive 87 nitrogen components in other countries of East Asia during 2000-2019 were download from the





- Acid Deposition Monitoring Network in East Asia (EANET). The European Monitoring and Evaluation Programme (EMEP) provides records of long-term reactive nitrogen components in 86 sites of most countries across West Europe. Monthly reactive nitrogen components dataset in 84 locations across the United States could be obtained from the Clean Air Status and Trends Network (CASTNET) (Figure S1).
- 93 2.2 Data preparation

The GEOS-Chem (v13.4.0) model driven by MERRA2 meteorological parameters was applied 94 to simulate the historical reactive nitrogen components (daily) during 2000-2019(Feng et al., 2021). 95 96 The GEOS-Chem model was composed of detailed ozone-NOx-VOC-PM-halogen tropospheric chemistry. The grid version of the model with a horizontal resolution of $2^{\circ} \times 2.5^{\circ}$ was utilized. Wet 97 98 deposition contained many processes including sub-grid scavenging in convective updrafts, in-99 cloud rainout, and below-cloud washout (Liu et al., 2001). Dry deposition was estimated based on 100 a resistance-in-series model (Wesely, 2007). The anthropogenic emission inventory in 2000-2019 101 was downloaded from the website of Community Emissions Data System (CEDS) (Hoesly et al., 102 2018). Then, the daily reactive nitrogen components were averaged to the monthly scale.

103 The IASI instrument aboard on the polar sun-synchronous MetOp platform traverses the 104 equation twice each day (9:30 a.m. and 9:30 p.m. local solar time) (Whitburn et al., 2016a). The 105 measurements in the daytime usually shows the better accuracy than those at night due to the high 106 sensitivity to ambient NH_3 (Van Damme et al., 2017; Whitburn et al., 2016a; Whitburn et al., 2016b). 107 In our study, we used the IASI NH_3 columns in morning during 2008-2019 to estimate the NH_3 and 108 NH_4^+ concentrations globally. Besides, the NH_3 column dataset with a cloud shield higher than 25% 109 and relative error above 100% were eliminated.

The tropospheric vertical column density (VCD) of NO₂ retrieved from OMI aboard on Aura satellite crosses the earth once a day (Kim et al., 2016). OMI-derived tropospheric NO₂ column densities during 2005-2019 was applied to develop the model. The tropospheric NO₂ column density data with cloud radiance fraction > 0.5, terrain reflectivity > 30%, and solar zenith angles > 85° were screened (Cooper et al., 2022). Additionally, the NO₂ columns from GOME (1995-2003), SCIAMACHY (2002-2011) and GOME-2 (2007-) were also collected to simulate the NO₃⁻ and HNO₃ levels. The similar overpass time of these three instruments (from about 09:30 to 10:30 LT,





117	local time) facilitates the simultaneous use to capture consistent long-term coverage. However, the
118	dataset cannot cover the NO_2 columns since 2017. To overcome the inconsistency of these satellite
119	products, we applied the linear regression technique to construct the relationship between $OMI-NO_2$
120	columns and GOME/SCIAMACHY NO_2 columns. The results suggested these satellite products
121	showed good relationship ($R^2 > 0.6$). At last, the long-term (2000-2019) NO ₂ columns at the global
122	scale were constructed.
123	The monthly meteorological parameters derived from ERA-5 comprise of 2 m dewpoint
124	temperature (D_{2m}), 2 m temperature (T_{2m}), surface pressure (S_p), and total precipitation (T_p), 10 m
125	U wind component (U10), and 10 m V wind component (V10). The population density data during
126	2000-2020 around the world were downloaded from
127	https://hub.worldpop.org/geodata/listing?id=64. The elevation data was extracted from ETOPO at a
128	spatial resolution of 1' (Amante and Eakins, 2009) (https://rda.ucar.edu/datasets/ds759.4/). In

addition, the land use data including cropland, forest, grassland, shrubland, tundra, barren land, and snow/ice were obtained from Liu et al. (2020a). Besides, the CMIP6 dataset in four scenarios were also applied to predict the reactive nitrogen concentrations during 2020-2100. The dataset includes 2-m air temperatures, wind speed at 850 and 500 hPa, total cloud cover, precipitation, relative humidity, and short-wave radiation. The modelled meteorological parameters derived from 16 earth system models were incorporated into the machine-learning model. The detailed models are summarized in Table S2.

136 2.3 Model development

137 A three-stage model was established to capture the full-coverage reactive nitrogen dataset at the global scale (Figure 1). In the first stage, the ground-level reactive nitrogen species, satellite 138 139 products (e.g., OMI-NO2 and IASI-NH3 columns), meteorological factors, land use types, population, and simulated reactive nitrogen components derived from GEOS-Chem model were 140 collected as the independent variables to estimate the gridded reactive nitrogen species at the 141 period/grid with satellite product based on LightGBM algorithm. In the second stage, the 142 143 meteorological parameters, GEOS-Chem output, land use data, and population were applied to fill the gaps without satellite retrievals. Then, the simulated results based on these models were fused 144 145 to obtain the full-coverage reactive nitrogen components and the ground-level observations were





- 146 further used to calibrate the full-coverage dataset and the final reactive nitrogen components at the global scale were simulated. In the last stage, the reactive nitrogen components and meteorological 147 parameters in four scenarios (SSP1-2.6, SSP2-4.5, SSP3-7.0, and SSP5-8.5) during 2020-2100 were 148 149 collected from CMIP6 dataset including 16 earth system models (Table S2). Then, the data in the 150 future scenarios were integrated into the ensemble model including XGBoost, LightGBM, and convolutional neural networks (CNN) to further calibrate the modeling results based on historical 151 152 dataset (2000-2019) derived from previous two-stage model. The detailed equations of multiple 153 machine-learning models are summarized as follows:
- 154 (1) XGBoost model

155
$$F^{(t)} = \sum_{i=1}^{n} [l(y_i, y^{\Lambda^{(t-1)}}) + \partial_{y^{(t-1)}} l(y_i, y^{\Lambda^{(t-1)}}) f_t(x_i) + \frac{1}{2} \partial_{y^{(t-1)}}^2 l(y_i, y^{\Lambda^{(t-1)}}) f_t^2(x_i)] + \Omega(f_t) \quad (1)$$

156 where $F^{(t)}$ is the cost function at the t-th period; ∂ is the derivative of the function; $\partial^2_{y^{(t-1)}}$

represents the second derivative of the function; *l* denotes the differentiable convex loss function that reveals the difference of the predicted value $\begin{pmatrix} x \\ y \end{pmatrix}$ of the i-th instance at the t-th period and the target value (y_i); f_i(x) represents the increment; $\Omega(f_t)$ reflects the regularizer.

160 (2) LightGBM model

161
$$\hat{f} = \arg\min_{f} E_{y,X} Q(y, f(x)) \quad (2)$$

162
$$f_T(X) = \sum_{t=1}^T f_t(X) \quad (3)$$

163 where Q(y,f(x)) reflects the a specific loss function; $\sum_{t=1}^{T} f_t(X)$ denotes the regression trees.

164 (3) CNN model

The reactive nitrogen species and meteorological parameters in the future scenarios were applied to
CNN model based on the historical (2000-2019) reactive nitrogen species derived from stage 1-2
model.

168
$$x \xrightarrow{f:U-Net} y$$
 (4)

169 where x (x1, x2,...., xn) represents the reactive nitrogen species and meteorological parameters





- 170 derived from CMIP6 dataset; y (y1, y2,, yn) denotes the historical (2000-2019) reactive nitrogen 171 species. All of the convolution layers showed the same kernel size of 3×3 and used rectified linear unit 172 173 (ReLU) as the activation function. Max-pooling layers were employed for adjusting the size of 174 images to capture better bottleneck information. After each block, the image size could be halved 175 by using the max pooling layer with kernel size of 2×2, but the number of channels will be doubled. 176 In our study, the learning rate was set as 0.1 to achieve the best performance. All of the independent variables collected from multiple sources were resampled to 0.25° grids 177 178 using appropriate algorithms. For example, both of the population density and land use type in each grid were calculated using spatial clipping toolbox. Later on, all of these variables were combined 179 to develop the model. During the development of multi-stage model, it was highly imperative to 180 181 remove some redundant explanatory variables and then determine the optimal variable group. The 182 redundant variables means that the overall predictive accuracy could degrade after the removal of 183 these variables. 184 3. Results and discussion 185 3.1 The modelling performance of historical reactive nitrogen estimates 186 The multi-stage model was applied to capture the spatiotemporal variations of reactive nitrogen 187 concentrations during 2000-2100. In our study, we employed XGBoost model to construct the full-188 coverage reactive nitrogen dataset during 2000-2020. The cross-validation (CV) R² values of the 189 model for NO₃, HNO₃, NH₃, and NH₄⁺ estimates reached 0.67, 0.62, 0.58, and 0.60, respectively (Figure 2). RMSE of NO₃, HNO₃, NH₃, and NH₄⁺ were 0.55, 0.23, 2.32, and 1.71 μ g N m⁻³, 190 respectively. MAE of NO3, HNO3, NH3, and NH4+ reached 0.19, 0.13, 1.23, and 0.59 µg N m-3. The 191 192 $CV R^2$ values of NO_3^- , HNO_3 , and NH_4^+ estimates were significantly higher than Jia et al. (2016) (0.22, 0.41, and 0.49), while the CV R² value of NO₃⁻ estimate in our study was comparable to 193 Geddes et al. (2017) (0.68) (Geddes and Martin, 2017). The CV R² value of NH₃ estimates were 194 also close to the results obtained by Liu et al. (2019) (0.45-0.71) (Liu et al., 2019). Overall, the 195 predictive performances historical reactive nitrogen was satisfied. Although the CV R² values in our 196 197 study were not significantly higher than those in some previous studies, our study developed the
- 198 full-coverage (gap-free) ambient reactive nitrogen dataset, which was superior to some previous





199	studies. Based on the constructed full-coverage reactive nitrogen dataset, we also developed the
200	ensemble model to calibrate the CMIP6 dataset in the future scenarios. The CV R^2 values of the
201	model for NO $_3$ ⁻ , HNO $_3$, NH $_3$, and NH $_4$ ⁺ estimates in the future scenarios reached 0.62, 0.67, 0.56,
202	and 0.60, respectively (Figure S2). RMSE of NO_3^- , HNO ₃ , NH ₃ , and NH ₄ ⁺ were 0.58, 0.26, 2.12,
203	and 1.91 μg N m $^{-3}$, respectively. MAE of NO3 $^{-}$, HNO3, NH3, and NH4 $^{+}$ reached 0.22, 0.22, 1.04, and
204	0.65 μg N m 3 . Overall, the ensemble model for these species in the future scenarios still showed
205	satisfied performance, and thus the result could be treated to be robust.
206	3.2 The spatial patterns of nitrogen reactive components
207	The global annual mean concentrations of NO_3^- , HNO_3 , NH_3 , and NH_4^+ during 2000-2019
208	ranged from 0.03 to 9.08, 0.03 to 1.73, 0.21 to 13.9, and 0.08 to 17.1 $\mu gNm^{\text{-}3}$ with the mean values
209	of 0.43 \pm 0.24 (standard deviation over grids), 0.28 \pm 0.13, 1.79 \pm 0.85, and 0.65 \pm 0.36 μg N m $^{-3}$
210	(Figure S3), respectively. China, West Europe, and the United States obtained widespread attention
211	due to the developed economy and dense anthropogenic activity.
212	In China, the overall mean ambient NO $_3^-$, HNO $_3$, NH $_3$, and NH $_4^+$ concentrations reached
213	$1.05\pm0.62,0.35\pm0.19,4.05\pm1.84,and2.38\pm1.26~\mu g$ N m 3 , ranging from 0.07-9.08, 0.06-1.73,
214	0.84-11.6, and 0.18-13.1 μg N m^-3. At the regional scale, the annual mean NO_3^-, HNO_3, NH_3, and
215	$\rm NH_{4^+}$ concentrations followed the order of North China Plain (NCP) (4.38, 1.12, 7.22, and 7.69 μg
216	N m^-3) $>$ Sichuan Basin (2.40 \pm 1.01, 0.52 \pm 0.28, 4.92 \pm 1.71, and 6.02 \pm 1.82 μg N m^-3) (Figure
217	3). NCP displayed the higher NO_3^- and HNO_3 concentrations owing to dense human activities and
218	strong industry foundation (Qi et al., 2023; Wen et al., 2018), which could emit a large amount of
219	NO_{x} to the atmosphere. In both of Yangtze River Delta (YRD) and Pearl River Delta (PRD), the
220	combustion of fossil fuels and traffic emissions might be the major source of $\ensuremath{\mathrm{NO}}_x$ emission, which
221	aggravated nitrate events via gas-particle conversion processes (Huang et al., 2017; Li et al., 2017).
222	For Sichuan Basin, the poor topographical or meteorological conditions were major factors
223	responsible for the severe nitrate pollution (Zhang et al., 2019). It was not surprising that high
224	ambient $\rm NH_3$ concentrations focused on NCP and Sichuan Basin because most of Chinese croplands
225	are distributed on these regions (Karra et al., 2021; Potapov et al., 2022), which was the major source
226	of NH_3 emissions with frequent N fertilizer applications (Ma et al., 2022). Besides, N manure was
227	another major source of $\rm NH_3$ emissions in China, and the percentage of N manure to $\rm NH_3$ emissions





228	exceeds 50% (Kang et al., 2016). The spatial pattern of NH_{4^+} level was in good agreement with the
229	$\rm NH_3$ concentration because $\rm NH_4^+$ was often generated from the reaction of $\rm NH_3$ with SO_2 and NO_2
230	(Ehrnsperger and Klemm, 2021).
231	In Europe, the ambient NO_3^- , HNO_3 , NH_3 , and NH_4^+ concentrations ranged from 0.13 to 2.84,
232	0.06 to 0.92, 0.35 to 7.81, and 0.22 to 3.77 μg N m 3 , respectively. The annual mean $NO_3^-,$ HNO3,
233	NH3, and NH4^+ levels reached 0.57 \pm 0.28, 0.25 \pm 0.11, 1.58 \pm 0.68, and 0.89 \pm 0.42 μg N m^3,
234	respectively (Figure 3). High concentrations of reactive nitrogen components focused on the
235	northern part of Italy, central and southern part of Germany, North France, Poland, and the western
236	part of Russia, which was in good agreement with the spatial pattern of NO_{x} and NH_{3} emissions
237	(Luo et al., 2022; Qu et al., 2020). Emissions Database for Global Atmospheric Research (EDGAR)
238	suggested that N fertilization and N manure accounted for 43% and 53% of total NH_3 emissions in
239	western Europe (Liu et al., 2019), respectively. Furthermore, Liu et al. (2019) confirmed that a good
240	relationship between ambient $\rm NH_3$ level and N fertilization plus N manure (r = 0.62) was observed
241	in Europe. Cooper et al. (2017) employed the inversion model to estimate NO_x emission in Europe
242	and also found that high NO_{x} emission was also mainly distributed on North France, Germany, the
243	northern part of Italy, and Russia, which partly explained the higher concentrations of reactive
244	nitrogen components in these regions.
245	In the United States, the ambient NO3 ⁻ , HNO3, NH3, and NH4 ⁺ concentrations reached 0.28 \pm
0.40	

246 0.12, 0.19 ± 0.08 , 2.12 ± 0.66 , and $0.49 \pm 0.25 \ \mu g \ N \ m^{-3}$, with the range of 0.03-2.35, 0.03-1.31, 247 0.26-9.96, and 0.10-6.09 $\mu g \ N \ m^{-3}$ (Figure 3), respectively. The hotspots of NO₃⁻, HNO₃, and NH₄⁺ 248 levels focused on the eastern part of the United States, while the higher NH₃ concentration focused 249 on Central Great Plains and some regions in California such as San Joaquin Valley (6.15 $\mu g \ N \ m^{-3}$). 250 Both of bottom-up and top-down NO_x and NH₃ emissions suggested that the spatial distributions of 251 reactive nitrogen components were strongly dependent on the precursor emissions (McDuffie et al., 2020; Qu et al., 2020).

253 Besides, some other regions such as the northern part of India also experienced severe N 254 pollution in the atmosphere. Meanwhile, some countries in South America such Brazil and 255 Argentina also suffered from serious HNO₃ and NH₃ pollution. The higher ambient NH₃ 256 concentration focused on the northern part of India might be contributed by two major reasons. First





257	of all, the intensive agricultural activities and high air temperature might be responsible for the
258	higher NH_3 level (Cui, 2023; Wang et al., 2020). Moreover, the relatively low sulfur dioxide (SO ₂)
259	and nitrogen oxides (NO _x) emissions coupled with high air temperature restricted the gas-to-particle
260	conversion of $\rm NH_3$ (Wang et al., 2020). The severe $\rm HNO_3$ and $\rm NH_3$ pollution in Brazil and Argentina
261	might be also linked with the dense agricultural activities (Huneeus et al., 2017).
262	3.3 The seasonal variations of reactive nitrogen components
263	The ambient NO3-, HNO3, NH3, and NH4+ concentrations exhibited significant seasonal
264	variations (Figure S4-8). NO $_3^-$, HNO $_3$, and NH $_4^+$ displayed the highest and lowest values in winter
265	(December-February) and summer (June-August), respectively. On the one hand, the anthropogenic
266	$\mathrm{NO}_{\boldsymbol{x}}$ emission for domestic heating might be higher in winter compared with other seasons (Lin et
267	al., 2011). On the other hand, the stagnant meteorological conditions limited the pollutant diffusion
268	(Li et al., 2019b; Liu et al., 2020c). Meanwhile, the higher relative humidity in winter facilitated the
269	formation of $\rm NH_4NO_3$ (Huang et al., 2016; Xu et al., 2012). However, both of ambient $\rm NO_3^-$ and
270	$\mathrm{NH_{4}^{\scriptscriptstyle +}}$ concentrations showed the lower concentrations in summer, which might be attributable to
271	the decomposition of $\mathrm{NH_4NO_3}$ under the condition of high air temperature. In contrast to the
272	secondary inorganic nitrogen, the ambient NH_3 level showed the highest concentration in summer
273	(1.71 \pm 0.45 μg N m^-3). China, Europe, and the United States suffered from similar NH_3 peaks in
274	summer (4.20 \pm 1.85, 1.77 \pm 0.65, and 2.21 \pm 1.04 μg N m^3). There are two reasons accounting for
275	the fact. At first, mineral N fertilizer or manure application was mainly performed in spring and
276	early summer (Paulot et al., 2014). Many field observations have obtained similar $N\mathrm{H}_3$ peak in
277	summer (He et al., 2021; Pan et al., 2018). Moreover, summer often showed the higher air
278	temperature, which promotes the volatilization of ammonium and limits the gas-to-particle of
279	gaseous NH ₃ (Liu et al., 2019).
280	3.4 The historical trends of reactive nitrogen components during 2000-2019

The long-term trends of ambient NO_3^- , HNO_3 , NH_3 , and NH_4^+ concentrations are shown in Figure 4 and Figure S9-12. The NO_3^- concentration in China displayed rapid increase (9.7%/yr) during 2000-2007, and then it kept the moderate increase (4.2%/yr) during 2007-2013. However, the NO_3^- concentration in China experienced the drastic decrease (-2.6%/yr) since 2013. The ambient HNO_3 and NH_4^+ concentrations showed similar trends during this period. Due to the impact





of Clean Air Action, the concentrations of gaseous precursors (e.g., SO2 and NOx) suffered from 286 287 substantial decreases, which could be transformed into nitrate and ammonium via heterogeneous 288 reactions (Huang et al., 2019). However, the decreasing rates of NO3- were still much lower than 289 those of gaseous precursors (Li et al., 2023). On the one hand, it might be associated with the 290 increased O₃ level and enhanced atmospheric oxidation capacity (AOC), which led to an increase 291 in the photochemical reaction rate of the secondary components (Wang et al., 2019). On the other 292 hand, strong SO2 emission control under the Clean Air Action allowed more gaseous NH3 to form 293 nitrate. The ambient NH₃ level remained relatively stable status during 2000-2013, while it 294 experienced rapid increases after 2013. The result was in good agreement with Liu et al. (2019). In fact, the ambient NH3 level in North China Plain still experienced dramatic increase (> 0.2 µg N m-295 296 ³/yr) during 2000-2013 because enhanced agricultural activities. Zhang et al. (2017) have 297 demonstrated that the livestock manure and fertilizer application generally accounted for 43.1% and 298 36.4% of the agricultural NH₃ emission, respectively. Since 2013, the NH₃ concentration in the 299 entire China suffered from rapid increase, which might be associated with the drastic decrease of 300 sulfate. It was well known that NH₃ could react with HNO₃ and gaseous H₂SO₄ to generate ammonia 301 sulfate and ammonia nitrate (Wang et al., 2022; Wang et al., 2019). Substantial decreases of acidic 302 gases (e.g., SO₂) lead to the reduction of NH₃ conversion to ammonia salts in the atmosphere (Chen 303 et al., 2019), which result in excess NH₃ remaining in the gaseous phase.

304 Compared with China, the long-term trends of reactive nitrogen components Europe and the 305 United States were relatively stable. In the Europe, the concentrations of NO₃, HNO₃, and NH₃ 306 exhibited increases during 2000-2007 (0.7%/yr, 2.3%/yr, and 2.1%/yr), while they experienced 307 slight decreases after 2007. The NH4⁺ level displayed continuous decrease since 2000. The result 308 was closely linked with the trends of NOx and NH₃ emissions derived from satellite retrieval 309 (Cooper et al., 2017; Luo et al., 2022). In the United States, both of NO₃ and NH₄⁺ showed persistent decreases during 2000-2019. Zhang et al. (2018) have confirmed that NO_x emission in the eastern 310 US has experienced persistent decrease since 1990, which facilitated the decreases of NO3- and 311 NH4⁺ levels. However, the ambient HNO3 and NH3 concentrations displayed slight increases during 312 313 2000-2007 (2.1%/yr), and then remained relatively stable since 2007. Liu et al. (2019) also found similar characteristic of ambient NH3 trend in the United States. In fact, the NH3 concentrations in 314





- 315 the Middle Plain and eastern US still showed increases due to the lack of NH₃ emission control
- 316 policies as well as the decline in acidic gases (Warner et al., 2017).
- 317 3.5 Projection of future ambient reactive nitrogen components

318 For the future reactive nitrogen component estimates, the ensemble model was applied to 319 predict the reactive nitrogen component concentrations under the SSP1-2.6, SSP2-4.5, SSP3-7.0, 320 and SSP5-8.5 scenarios. The whole period (2021-2100) could be uniformly classified into four 321 periods including 2021-2040 (2030s), 2041-2060 (2050s), 2061-2080 (2070s), and 2081-2100 322 (2090s), respectively. In SSP1-2.6, the projected global mean NO₃ concentrations decreased from 323 $0.40 \pm 0.12 \ \mu g \ N \ m^{-3}$ during 2030s to $0.38 \pm 0.11 \ \mu g \ N \ m^{-3}$ during 2090s (Figure 5a-d and Figure S13). SSP2-4.5 scenario represents the middle range of plausible future pathways (Nazarenko et al., 324 325 2022), and the predicted global average NO₃ concentrations in this scenario decreased from 0.40 \pm $0.14 \ \mu g \ N \ m^{-3}$ during 2030s to $0.37 \pm 0.12 \ \mu g \ N \ m^{-3}$ during 2090s. SSP3-7.0 and SSP5-8.5 denote 326 327 the less investment in the environment and heavily relies on traditional energy for rapid economic 328 development, respectively. The ambient NO₃ levels in these scenarios showed more remarkable 329 decreases from 2030s to 2090s. For instance, the NO₃ concentrations in China decreased from 1.07 330 \pm 0.36 and 1.01 \pm 0.30 μg N m 3 to 0.84 \pm 0.46 and 0.60 \pm 0.34 μg N m 3 , respectively. The NO $_3$ 331 levels in India also experienced rapid increases from 2.21 ± 1.04 to $3.30 \pm 1.58 \ \mu g \ N \ m^{-3}$ in SSP3-332 7.0 scenario, while they suffered from marked decreases from 2.33 ± 1.12 to $1.60 \pm 0.84 \,\mu g \,\mathrm{N \,m^{-3}}$ 333 in SSP5-8.5 scenario.

334 The projected global average HNO₃ concentrations remained relatively stable from 0.26 ± 0.13 and 0.27 \pm 0.14 μg N m 3 during 2030s to 0.25 \pm 0.12 and 0.26 \pm 0.13 μg N m 3 during 2090s in 335 336 SSP1-2.6 and SSP2-4.5 scenarios, respectively (Figure 6a-d and Figure S14). However, some 337 developing countries such as China and India experienced drastic HNO₃ changes during these 338 scenarios. The mean HNO₃ concentrations in China and India decreased from 0.29 ± 0.17 and 0.43 \pm 0.25 µg N m⁻³ to 0.26 \pm 0.15 and 0.35 \pm 0.22 µg N m⁻³ in SSP1-2.6 scenario, respectively. In SSP2-339 4.5 scenario, the average HNO₃ concentrations in China and India decreased from 0.31 ± 0.20 and 340 $0.55 \pm 0.29 \ \mu g \ N \ m^{-3}$ in 2030s to 0.27 ± 0.18 and $0.45 \pm 0.25 \ \mu g \ N \ m^{-3}$ in 2090s, respectively. 341 Compared with SSP1-2.6 and SSP2-4.5, the heavy-pollution scenarios (e.g., SSP3-7.0 and SSP5-342 343 8.5) showed the higher HNO3 concentrations. The HNO3 concentrations in China decreased from





344	0.34 ± 0.21 and $0.31\pm0.20~\mu g$ N m 3 in 2030s to 0.31 ± 0.20 and $0.28\pm0.14~\mu g$ N m 3 in 2090s for
345	SSP3-7.0 and SSP5-8.5 scenarios, respectively. The HNO3 concentrations in India also experienced
346	rapid increases from 0.54 \pm 0.29 to 0.70 \pm 0.38 μg N m^{-3} in SSP3-7.0 scenario. It was assumed that
347	the government gave less investment in environment improvement and the anthropogenic emission
348	did not show marked decrease under the condition of SSP3-7.0 scenario. However, the ambient
349	HNO_3 levels suffered from marked decreases from 0.63 ± 0.28 to $0.51\pm0.22~\mu g$ N m 3 in SSP5-8.5
350	scenario. Although the SSP5-8.5 scenario is heavily dependent on the fossil fuel production (Chen
351	et al., 2020), the anthropogenic emission still displayed gradual decrease during the 80-year period.
352	The higher ambient NH3 concentrations also focused on China and India (Figure 7 and Figure
353	S15). In SSP1-2.6, the ambient NH ₃ concentrations in China and India decreased from 2.92 and
354	5.59 μg N m $^{\text{-3}}$ during 2030s to 1.77 and 2.73 μg N m $^{\text{-3}}$ during 2090s, respectively. In SSP2-4.5, the
355	ambient mean NH_3 concentrations in China and India decreased from 3.43 and 7.57 $\mu gNm^{\text{-3}}$ during
356	2030s to 2.06 and 5.33 μg N m 3 during 2090s, respectively. Compared with SSP1-2.6 and SSP2-
357	4.5, the ambient NH ₃ concentrations in heavy-pollution scenarios did not show marked decreases
358	from 2020-2100, which might be associated with ineffective NH_3 emission control. The temporal
359	variations of ambient $\mathrm{NH_{4}^{+}}$ levels in the future scenarios show similar trends with $\mathrm{NH_{3}}.$ The
360	atmospheric NH4+ levels in China decreased from 1.53 (SSP1-2.6), 1.95 (SSP2-4.5), 2.31 (SSP3-
361	7.0), and 1.87 μg N m $^{-3}$ (SSP5-8.5) during 2030s to 0.60, 0.84, 1.78, and 0.97 μg N m $^{-3}$ during 2090s,
362	respectively (Figure 8 and Figure S16). Meanwhile, the ambient NH_{4^+} levels in India changed from
363	3.68 (SSP1-2.6), 5.31 (SSP2-4.5), 5.49 (SSP3-7.0), and 6.07 μg N m 3 (SSP5-8.5) during 2030s to
364	1.38, 3.49, 6.89, and 2.96 $\mu g \ N \ m^{\text{-3}}$ during 2090s, respectively.
365	3.6 Implications and limitations

Global trends of four reactive nitrogen components during 2000-2100 emphasizes the urgent mitigation measures (carbon neutrality pathway) to reduce precursor emissions in order to decrease the concentrations and depositions of reactive nitrogen components especially in China and India. Furthermore, our result could give valuable insights into the impact of reactive nitrogen components on human health and ecological environment. However, this study still shows some limitations. First of all, the observation networks mainly focus on China, Europe, and the United States, and thus the simulations in many other regions might show large uncertainties. Secondly, the CMIP6 future





- 373 climate scenario data also exhibits large uncertainties, which could impact the reliability of this
- 374 study. Lastly, our predictions were performed on the basis of the premise that the world was steadily
- 375 developing, and cannot predict the impacts of uncontrollable factors (e.g., COVID-19, Russia-
- 376 Ukraine War).
- 377 Competing interests
- 378 The contact author has declared that none of the authors has any competing interests.
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- 381 Data availability
- 382 The CMIP6 dataset used in this publication is available in <u>https://esgf.nci.org.au/projects/cmip6-</u>
- 383 <u>nci/</u>.
- 384 Author contributions
- 385 LR and WGH designed the study; LR developed the model; GYN, ZLJ, and SYB analyzed the
- 386 observation and model data. LR wrote this manuscript.
- 387





- 388 Figure 1 The workflow of global full-coverage reactive nitrogen estimates during 2000-2100.
- 389 GC_{output} denotes the GEOS-Chem output.







391Figure 2 The predictive performances of four reactive nitrogen components including NO_3^- (a),392HNO3 (b), NH3 (c), and NH4⁺ (d). The model was constructed with 90% original data and the393remained data was applied to validate the model. The black solid line denotes the best-fitting curve394for all of the points, while the black dashed line represents the diagonal, which means the same395observed and simulated values. The color scale denotes the sample size.







Figure 3 The spatiotemporal variations of NO₃⁻, HNO₃, NH₃, and NH₄⁺ concentrations in East Asia
(a-d), Europe (e-h), and North America (i-l) (Unit: μg N m⁻³).







- 400 Figure 4 The long-term variations of NO_3^- , HNO_3 , NH_3 , and NH_4^+ concentrations in China (pink),
- 401 Europe (green), and the United States (cyan) (Unit: μ g N m⁻³).







Figure 5 Spatial variations of projected global ambient concentrations of reactive nitrogen
components under different climate change scenarios (Unit: μg N m⁻³). Panels (a-d) represent the
annual mean concentrations of ambient NO₃⁻ under SSP1-2.6, SSP2-4.5, SSP3-7.0, and SSP5-8.5
during 2021-2100, respectively.







- 409 Figure 6 Spatial variations of projected global ambient concentrations of reactive nitrogen components under different climate change scenarios (Unit: µg N m⁻³). Panels (a-d) represent the 410
- 411 annual mean concentrations of ambient HNO3 under SSP1-2.6, SSP2-4.5, SSP3-7.0, and SSP5-8.5
- 412 during 2021-2100, respectively.







Figure 7 Spatial variations of projected global ambient concentrations of reactive nitrogen
 components under different climate change scenarios (Unit: μg N m⁻³). Panels (a-d) represent the

annual mean concentrations of ambient NH₃ under SSP1-2.6, SSP2-4.5, SSP3-7.0, and SSP5-8.5
during 2021-2100, respectively.







- 421 Figure 8 Spatial variations of projected global ambient concentrations of reactive nitrogen
 422 components under different climate change scenarios (Unit: μg N m⁻³). Panels (a-d) represent the
- annual mean concentrations of ambient NH4⁺ under SSP1-2.6, SSP2-4.5, SSP3-7.0, and SSP5-8.5
 during 2021-2100, respectively.







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