



## 30 1. Introduction

31 Along with the development of global urbanization and industrialization, the anthropogenic  
32 emissions of reactive nitrogen (e.g.,  $\text{NO}_x$ ,  $\text{NH}_3$ ) experienced drastic increases during the past  
33 decades, and caused the higher concentrations of  $\text{NO}_2$ ,  $\text{NH}_3$ , and many secondary components such  
34 as  $\text{NO}_3^-$  ( $\text{NO}_3\text{-N}$ ),  $\text{NH}_4^+$  ( $\text{NH}_4\text{-N}$ ), and  $\text{HNO}_3$  (Chen et al., 2021; Liu et al., 2020b; McDuffie et al.,  
35 2020). The reactive nitrogen released from anthropogenic source could significantly alter the global  
36 nitrogen cycle throughout the Earth system (Altieri et al., 2021; Zhang et al., 2020). Reactive  
37 nitrogen in the atmosphere dominates the chemical formation of tropospheric  $\text{O}_3$  and aggravates the  
38 particle pollution (Geddes and Martin, 2017), with implications for global air quality and climate  
39 change (He et al., 2022; Von Schneidmesser et al., 2015). Moreover, the ambient reactive nitrogen  
40 could be deposited into the land surface and could cause lake eutrophication and soil acidification  
41 (Bouwman et al., 2002; Chen et al., 2018). Therefore, it is highly necessary to understand the spatial  
42 distributions and temporal evolution trends of reactive nitrogen components at the global scale.

43 Despite the global importance, observational constraints on reactive nitrogen in the atmosphere  
44 were still scarce in most parts of the world (Liu et al., 2020b). Furthermore, the majority of  
45 monitoring sites focused on China, Europe, and the United States (Du et al., 2014; Li et al., 2020;  
46 Li et al., 2019a; Li et al., 2016), and these uneven sites only possessed limited spatial  
47 representativeness (Shi et al., 2018), which restricted the accurate assessment of global reactive  
48 nitrogen pollution. Fortunately, the satellite observations gave us a unprecedented chance to capture  
49 the global variations of atmospheric reactive nitrogen. Geddes et al. (2017) used the satellite  
50 products to calibrate the simulated reactive nitrogen oxides ( $\text{NO}_y$ ) and improved the predictive  
51 performance ( $R = 0.83$ ) compared with the chemical transport model (CTM) output alone (Geddes  
52 and Martin, 2017). Afterwards, Liu et al. (2022) also used the similar method to estimate the global  
53 wet deposition of reduced nitrogen ( $\text{NH}_4^+$ ) and the R value achieved 0.80 (Liu et al., 2021). Although  
54 the calibration based on satellite products could improve the predictive accuracy compared with  
55 CTM output, the simulated values still largely biased from the ground-level observations. Moreover,  
56 the method cannot accurately fill the gaps of reactive nitrogen concentrations without satellite  
57 coverage. In our previous works, we developed a satellite-based ensemble machine-learning model  
58 to predict the wet  $\text{NH}_4^+$  deposition across China and the  $R^2$  value reached 0.76 ( $R = 0.88$ ) (Li et al.,

59 2020). However, this technique was not expanded to the global scale and the high-accuracy and full-  
60 coverage global ambient reactive nitrogen dataset was still lack.

61 Apart from the historical estimates, the future prediction of reactive nitrogen is also important  
62 because these components in the future scenarios could significantly affect the land carbon cycle  
63 and greenhouse gas emissions, both of which could aggravate the global climate change and affect  
64 the earth system safety (Chen et al., 2015; Zaehle, 2013). To the best of our knowledge, only two  
65 studies focused on global aerosol prediction in the future scenarios. Chen et al. (2023) predicted the  
66 global PM<sub>2.5</sub> levels and associated mortalities in 2100 under different climate scenarios and found  
67 that SSP3-7.0 scenario was linked with the highest PM<sub>2.5</sub> exposure. Li et al., (2022) also simulated  
68 the global NO<sub>3</sub><sup>-</sup> (NO<sub>3</sub>-N) and NH<sub>4</sub><sup>+</sup> (NH<sub>4</sub>-N) levels in the future four scenarios and demonstrated  
69 that both of these components showed marked decreases in most cases except SSP5-8.5 scenario.  
70 However, this study predicted the future reactive nitrogen based on historical CTM output alone,  
71 which lacks of observation constraints. The result might increase the uncertainty of assessment.

72 In our study, we developed a multi-stage model to estimate the concentrations of four reactive  
73 nitrogen species (NO<sub>3</sub><sup>-</sup> (NO<sub>3</sub>-N), HNO<sub>3</sub>, NH<sub>3</sub>, and NH<sub>4</sub><sup>+</sup> (NH<sub>4</sub>-N)) during 2000-2019 because these  
74 species were most important reactive nitrogen components for human health and ecological  
75 ecosystem and also showed abundant ground-level observations. Then, the species over the 2020-  
76 2100 period under the SSP1-2.6, SSP2-4.5, SSP3-7.0, and SSP5-8.5 scenarios were also corrected  
77 based on the historical estimates. Finally, the long-term dataset of reactive nitrogen during 2000-  
78 2100 was constructed. Our results were beneficial to assess the impacts of reactive nitrogen  
79 components on air pollution and climate change in the future.

## 80 **2. Material and methods**

### 81 2.1 Reactive nitrogen observations

82 Most of reactive nitrogen observations focused on East Asia, Europe, and the United States. The  
83 monthly reactive nitrogen components monitoring data during 2010-2015 in China were  
84 downloaded from nationwide nitrogen deposition monitoring network (NNDMN) including 32 sites,  
85 and these sites could be classified into three types mixed with urban, rural, and background sites  
86 (Xu et al., 2019) (Table S1). The concentrations of reactive nitrogen components were determined  
87 using the active DENuder for Long-Term Atmospheric sampling system (DELTA). The detailed

88 sampling and analysis procedures have been described by Xu et al. (2019). The dataset of reactive  
89 nitrogen components in other countries of East Asia during 2000-2019 were download from the  
90 Acid Deposition Monitoring Network in East Asia (EANET), which includes 41 sites. The European  
91 Monitoring and Evaluation Programme (EMEP) provides records of long-term reactive nitrogen  
92 components in 86 sites of most countries across West Europe. Monthly reactive nitrogen  
93 components dataset in 84 locations across the United States could be obtained from the Clean Air  
94 Status and Trends Network (CASTNET) (Figure S1).

## 95 2.2 Data preparation

96 The GEOS-Chem (v13.4.0) model driven by MERRA2 meteorological parameters was applied  
97 to simulate the historical reactive nitrogen components (daily) during 2000-2019 (Feng et al., 2021).  
98 The GEOS-Chem model was composed of detailed ozone-NO<sub>x</sub>-VOC-PM-halogen tropospheric  
99 chemistry. The grid version of the model with a horizontal resolution of 2° × 2.5° was utilized. Wet  
100 deposition contained many processes including sub-grid scavenging in convective updrafts, in-  
101 cloud rainout, and below-cloud washout (Liu et al., 2001). Dry deposition was estimated based on  
102 a resistance-in-series model (Wesely, 2007). The estimates of aerosol optical properties account for  
103 the hygroscopic growth (Drury et al., 2010). Vertical mixing in the boundary layer follows a non-  
104 local scheme implemented by Lin and McElroy (2010), and convection employs the relaxed  
105 Arakawa-Schubert scheme. The anthropogenic emission inventory in 2000-2019 was downloaded  
106 from the website of Community Emissions Data System (CEDS) (Hoesly et al., 2018). CEDS  
107 emission inventory includes eight sectors such as agriculture, energy, industry, residential, shipping,  
108 solvents, transportation, and waste incineration. Then, the daily reactive nitrogen components were  
109 averaged to the monthly scale.

110 The IASI instrument aboard on the polar sun-synchronous MetOp platform traverses the  
111 equator twice each day (9:30 a.m. and 9:30 p.m. local solar time) (Whitburn et al., 2016a). The  
112 measurements in the daytime usually shows the better accuracy than those at night due to the high  
113 sensitivity to ambient NH<sub>3</sub> (Van Damme et al., 2017; Whitburn et al., 2016a; Whitburn et al., 2016b).  
114 In our study, we used the IASI NH<sub>3</sub> columns in morning during 2008-2019 to estimate the NH<sub>3</sub> and  
115 NH<sub>4</sub><sup>+</sup> concentrations globally. Besides, the NH<sub>3</sub> column dataset with a cloud shield higher than 25%  
116 and relative error above 100% were eliminated.

117 The tropospheric vertical column density (VCD) of NO<sub>2</sub> retrieved from OMI aboard on Aura  
118 satellite crosses the earth once a day (Kim et al., 2016). OMI-derived tropospheric NO<sub>2</sub> column  
119 densities during 2005-2019 was applied to develop the model. The tropospheric NO<sub>2</sub> column density  
120 data with cloud radiance fraction > 0.5, terrain reflectivity > 30%, and solar zenith angles > 85°  
121 were screened (Cooper et al., 2022). Additionally, the NO<sub>2</sub> columns from GOME (1995-2003),  
122 SCIAMACHY (2002-2011) and GOME-2 (2007-) were also collected to simulate the NO<sub>3</sub><sup>-</sup> (NO<sub>3</sub>-  
123 N) and HNO<sub>3</sub> levels. The similar overpass time of these three instruments (from about 09:30 to  
124 10:30 LT, local time) facilitates the simultaneous use to capture consistent long-term coverage.  
125 However, the dataset cannot cover the NO<sub>2</sub> columns since 2017. To overcome the inconsistency of  
126 these satellite products, we applied the linear regression technique to construct the relationship  
127 between OMI-NO<sub>2</sub> columns and GOME/SCIAMACHY NO<sub>2</sub> columns. The results suggested these  
128 satellite products showed good relationship ( $R^2 > 0.6$ ). At last, the long-term (2000-2019) NO<sub>2</sub>  
129 columns at the global scale were constructed.

130 The monthly meteorological parameters derived from ERA-5 comprise of 2 m dewpoint  
131 temperature ( $D_{2m}$ ), 2 m temperature ( $T_{2m}$ ), surface pressure ( $S_p$ ), and total precipitation ( $T_p$ ), 10 m  
132 U wind component ( $U_{10}$ ), and 10 m V wind component ( $V_{10}$ ). The population density data during  
133 2000-2020 around the world were downloaded from  
134 <https://hub.worldpop.org/geodata/listing?id=64>. The elevation data was extracted from ETOPO at a  
135 spatial resolution of 1' (Amante and Eakins, 2009) (<https://rda.ucar.edu/datasets/ds759.4/>). In  
136 addition, the land use types including cropland, forest, grassland, shrubland, tundra, barren land,  
137 and snow/ice were obtained from Liu et al. (2020a). Besides, the CMIP6 dataset in four scenarios  
138 were also applied to predict the reactive nitrogen concentrations during 2020-2100. The dataset  
139 includes 2-m air temperatures, wind speed at 850 and 500 hPa, total cloud cover, precipitation,  
140 relative humidity, and short-wave radiation. The modelled meteorological parameters derived from  
141 16 earth system models were incorporated into the machine-learning model. The detailed models  
142 are summarized in Table S2.

### 143 2.3 Model development

144 A three-stage model was established to capture the full-coverage reactive nitrogen dataset at  
145 the global scale (Figure 1). In the first stage, the ground-level reactive nitrogen species, satellite

146 products (e.g., OMI-NO<sub>2</sub> and IASI-NH<sub>3</sub> columns), meteorological parameters, land use types,  
 147 population, and simulated reactive nitrogen components derived from GEOS-Chem model were  
 148 collected as the independent variables to estimate the gridded reactive nitrogen species at the  
 149 period/grid with satellite product based on XGBoost algorithm. In the second stage, the  
 150 meteorological parameters, GEOS-Chem output, land use types, and population were applied to fill  
 151 the gaps without satellite retrievals. Then, the simulated results based on these models were fused  
 152 to obtain the full-coverage reactive nitrogen components and the ground-level observations were  
 153 further used to calibrate the full-coverage dataset and the final reactive nitrogen components at the  
 154 global scale were simulated. In the last stage, the reactive nitrogen components and meteorological  
 155 parameters in four scenarios (SSP1-2.6, SSP2-4.5, SSP3-7.0, and SSP5-8.5) during 2020-2100 were  
 156 collected from CMIP6 dataset including 16 earth system models (Table S2). Then, the data in the  
 157 future scenarios were integrated into the ensemble model including XGBoost, LightGBM, and  
 158 convolutional neural networks (CNN) to further calibrate the modeling results based on historical  
 159 dataset (2000-2019) derived from previous two-stage model. The detailed equations of multiple  
 160 machine-learning models are summarized as follows:

161 (1) XGBoost model

$$162 \quad F^{(t)} = \sum_{i=1}^n [l(y_i, y^{\Lambda(t-1)}) + \partial_{y^{\Lambda(t-1)}} l(y_i, y^{\Lambda(t-1)}) f_t(x_i) + \frac{1}{2} \partial_{y^{\Lambda(t-1)}}^2 l(y_i, y^{\Lambda(t-1)}) f_t^2(x_i)] + \Omega(f_t) \quad (1)$$

163 where  $F^{(t)}$  is the cost function at the  $t$ -th period;  $\partial$  is the derivative of the function;  $\partial_{y^{\Lambda(t-1)}}^2$   
 164 represents the second derivative of the function;  $l$  denotes the differentiable convex loss function  
 165 that reveals the difference of the predicted value ( $y^{\Lambda}$ ) of the  $i$ -th instance at the  $t$ -th period and the  
 166 target value ( $y_i$ );  $f_t(x)$  represents the increment;  $\Omega(f_t)$  reflects the regularizer. Maximum tree  
 167 depth and learning rate reached 15 and 0.1, respectively.

168 (2) LightGBM model

$$169 \quad f^{\Lambda} = \arg \min_f E_{y,x} Q(y, f(x)) \quad (2)$$

$$170 \quad f_T(X) = \sum_{t=1}^T f_t(X) \quad (3)$$

171 where  $Q(y, f(x))$  reflects the a specific loss function;  $\sum_{t=1}^T f_t(X)$  denotes the regression trees.

172 Maximum tree depth, learning rate, and feature fraction reached 25, 0.2, and 0.7, respectively.

173 (3) CNN model

174 The reactive nitrogen species and meteorological parameters in the future scenarios were applied to  
175 CNN model based on the historical (2000-2019) reactive nitrogen species derived from stage 1-2  
176 model.

$$177 \quad x \xrightarrow{f:U-Net} y \quad (4)$$

178 where  $x (x_1, x_2, \dots, x_n)$  represents the reactive nitrogen species and meteorological parameters  
179 derived from CMIP6 dataset;  $y (y_1, y_2, \dots, y_n)$  denotes the historical (2000-2019) reactive nitrogen  
180 species.

181 All of the convolution layers showed the same kernel size of  $3 \times 3$  and used rectified linear unit  
182 (ReLU) as the activation function. Max-pooling layers were employed for adjusting the size of  
183 images to capture better bottleneck information. After each block, the image size could be halved  
184 by using the max pooling layer with kernel size of  $2 \times 2$ , but the number of channels will be doubled.  
185 In our study, the learning rate was set as 0.1 to achieve the best performance.

186 All of the independent variables collected from multiple sources were resampled to  $0.25^\circ$  grids  
187 using Kriging interpolation. For example, both of the population density and land use types in each  
188 grid were calculated using spatial clipping toolbox. Later on, all of these variables were combined  
189 to develop the model. During the development of multi-stage model, it was highly imperative to  
190 remove some redundant explanatory variables and then determine the optimal variable group. The  
191 redundant variables means that the overall predictive accuracy could degrade after the removal of  
192 these variables.

### 193 **3. Results and discussion**

#### 194 3.1 The modelling performance of historical reactive nitrogen estimates

195 The multi-stage model was applied to capture the spatiotemporal variations of reactive nitrogen  
196 concentrations during 2000-2100. In our study, we employed XGBoost model to construct the full-  
197 coverage reactive nitrogen dataset during 2000-2020. The cross-validation (CV)  $R^2$  values of the  
198 model for  $\text{NO}_3^-$  ( $\text{NO}_3\text{-N}$ ),  $\text{HNO}_3$ ,  $\text{NH}_3$ , and  $\text{NH}_4^+$  ( $\text{NH}_4\text{-N}$ ) estimates reached 0.67, 0.62, 0.58, and

199 0.60, respectively (Figure 2). RMSE of  $\text{NO}_3^-$  ( $\text{NO}_3\text{-N}$ ),  $\text{HNO}_3$ ,  $\text{NH}_3$ , and  $\text{NH}_4^+$  ( $\text{NH}_4\text{-N}$ ) were 0.55,  
200 0.23, 2.32, and  $1.71 \mu\text{g N m}^{-3}$ , respectively. MAE of  $\text{NO}_3^-$  ( $\text{NO}_3\text{-N}$ ),  $\text{HNO}_3$ ,  $\text{NH}_3$ , and  $\text{NH}_4^+$  ( $\text{NH}_4\text{-N}$ )  
201  $\text{N}$ ) reached 0.19, 0.13, 1.23, and  $0.59 \mu\text{g N m}^{-3}$ . The CV  $R^2$  values of  $\text{NO}_3^-$  ( $\text{NO}_3\text{-N}$ ),  $\text{HNO}_3$ , and  
202  $\text{NH}_4^+$  ( $\text{NH}_4\text{-N}$ ) estimates were significantly higher than Jia et al. (2016) (0.22, 0.41, and 0.49), while  
203 the CV  $R^2$  value of  $\text{NO}_3^-$  estimate in our study was comparable to Geddes et al. (2017) (0.68)  
204 (Geddes and Martin, 2017). The CV  $R^2$  value of  $\text{NH}_3$  estimates were also close to the results  
205 obtained by Liu et al. (2019) (0.45-0.71) (Liu et al., 2019). Overall, the predictive performances  
206 historical reactive nitrogen was satisfied. Although the CV  $R^2$  values in our study were not  
207 significantly higher than those in some previous studies, our study developed the full-coverage (gap-  
208 free) ambient reactive nitrogen dataset, which was superior to some previous studies. Based on the  
209 constructed full-coverage reactive nitrogen dataset, we also developed the ensemble model to  
210 calibrate the CMIP6 dataset in the future scenarios. The CV  $R^2$  values of the model for  $\text{NO}_3^-$  ( $\text{NO}_3\text{-N}$ -  
211  $\text{N}$ ),  $\text{HNO}_3$ ,  $\text{NH}_3$ , and  $\text{NH}_4^+$  ( $\text{NH}_4\text{-N}$ ) estimates in the future scenarios reached 0.62, 0.67, 0.56, and  
212 0.60, respectively (Figure S2). RMSE of  $\text{NO}_3^-$  ( $\text{NO}_3\text{-N}$ ),  $\text{HNO}_3$ ,  $\text{NH}_3$ , and  $\text{NH}_4^+$  ( $\text{NH}_4\text{-N}$ ) were 0.58,  
213 0.26, 2.12, and  $1.91 \mu\text{g N m}^{-3}$ , respectively. MAE of  $\text{NO}_3^-$  ( $\text{NO}_3\text{-N}$ ),  $\text{HNO}_3$ ,  $\text{NH}_3$ , and  $\text{NH}_4^+$  ( $\text{NH}_4\text{-N}$ )  
214  $\text{N}$ ) reached 0.22, 0.22, 1.04, and  $0.65 \mu\text{g N m}^{-3}$ . Overall, the ensemble model for these species in the  
215 future scenarios still showed satisfied performance, and thus the result could be treated to be robust.

### 216 3.2 The spatial patterns of nitrogen reactive components

217 The global annual mean concentrations of  $\text{NO}_3^-$ ,  $\text{HNO}_3$ ,  $\text{NH}_3$ , and  $\text{NH}_4^+$  during 2000-2019  
218 ranged from 0.03 to 9.08, 0.03 to 1.73, 0.21 to 13.9, and 0.08 to  $17.1 \mu\text{g N m}^{-3}$  with the mean values  
219 of  $0.43 \pm 0.24$  (standard deviation over grids),  $0.28 \pm 0.13$ ,  $1.79 \pm 0.85$ , and  $0.65 \pm 0.36 \mu\text{g N m}^{-3}$   
220 (Figure S3), respectively. East Asia especially China, West Europe, and the United States obtained  
221 widespread attention due to the developed economy and dense anthropogenic activity.

222 In China, the overall mean ambient  $\text{NO}_3^-$  ( $\text{NO}_3\text{-N}$ ),  $\text{HNO}_3$ ,  $\text{NH}_3$ , and  $\text{NH}_4^+$  ( $\text{NH}_4\text{-N}$ )  
223 concentrations reached  $1.05 \pm 0.62$ ,  $0.35 \pm 0.19$ ,  $4.05 \pm 1.84$ , and  $2.38 \pm 1.26 \mu\text{g N m}^{-3}$ , ranging  
224 from 0.07-9.08, 0.06-1.73, 0.84-11.6, and 0.18-13.1  $\mu\text{g N m}^{-3}$ . At the regional scale, the annual mean  
225  $\text{NO}_3^-$ ,  $\text{HNO}_3$ ,  $\text{NH}_3$ , and  $\text{NH}_4^+$  concentrations followed the order of North China Plain (NCP) ( $4.38$ ,  
226  $1.12$ ,  $7.22$ , and  $7.69 \mu\text{g N m}^{-3}$ ) > Sichuan Basin ( $2.40 \pm 1.01$ ,  $0.52 \pm 0.28$ ,  $4.92 \pm 1.71$ , and  $6.02 \pm$   
227  $1.82 \mu\text{g N m}^{-3}$ ) (Figure 3). NCP displayed the higher  $\text{NO}_3^-$  and  $\text{HNO}_3$  concentrations owing to dense



228 human activities and strong industry foundation (Qi et al., 2023; Wen et al., 2018), which could emit  
229 a large amount of NO<sub>x</sub> to the atmosphere. In both of Yangtze River Delta (YRD) and Pearl River  
230 Delta (PRD), the combustion of fossil fuels and traffic emissions might be the major source of NO<sub>x</sub>  
231 emission, which aggravated nitrate events via gas-particle conversion processes (Huang et al., 2017;  
232 Li et al., 2017). For Sichuan Basin, the poor topographical or meteorological conditions were major  
233 factors responsible for the severe nitrate pollution (Zhang et al., 2019). It was not surprising that  
234 high ambient NH<sub>3</sub> concentrations focused on NCP and Sichuan Basin because many croplands (dry  
235 land) are distributed on these regions (Karra et al., 2021; Potapov et al., 2022), which was the major  
236 source of NH<sub>3</sub> emissions with frequent N fertilizer applications (Ma et al., 2022). Besides, N manure  
237 was another major source of NH<sub>3</sub> emissions in China, and the percentage of N manure to NH<sub>3</sub>  
238 emissions exceeds 50% (Kang et al., 2016). The spatial pattern of NH<sub>4</sub><sup>+</sup> level was in good agreement  
239 with the NH<sub>3</sub> concentration because NH<sub>4</sub><sup>+</sup> was often generated from the reaction of NH<sub>3</sub> with SO<sub>2</sub>  
240 and NO<sub>2</sub> (Ehrnsperger and Klemm, 2021). Apart from China, many other countries such as South  
241 Korea and Japan also showed the higher ambient reactive nitrogen concentrations. As shown in  
242 Figure 3, the higher reactive N concentrations occurred on the western coasts of South Korea than  
243 on the eastern coasts. The higher reactive N concentrations in Japan mainly focused on the urban  
244 areas around Tokyo, which might be linked with the dense anthropogenic emission in this region  
245 (Li et al., 2024). In Southeast Asia, Indonesia (NO<sub>3</sub><sup>-</sup> (NO<sub>3</sub>-N), HNO<sub>3</sub>, NH<sub>3</sub>, and NH<sub>4</sub><sup>+</sup> (NH<sub>4</sub>-N): 0.18,  
246 0.47, 5.72, and 0.44 μg N m<sup>-3</sup>) suffered from the most serious reactive N pollution compared with  
247 other surrounding countries.

248 In Europe, the ambient NO<sub>3</sub><sup>-</sup> (NO<sub>3</sub>-N), HNO<sub>3</sub>, NH<sub>3</sub>, and NH<sub>4</sub><sup>+</sup> (NH<sub>4</sub>-N) concentrations ranged  
249 from 0.13 to 2.84, 0.06 to 0.92, 0.35 to 7.81, and 0.22 to 3.77 μg N m<sup>-3</sup>, respectively. The annual  
250 mean NO<sub>3</sub><sup>-</sup> (NO<sub>3</sub>-N), HNO<sub>3</sub>, NH<sub>3</sub>, and NH<sub>4</sub><sup>+</sup> (NH<sub>4</sub>-N) levels reached 0.57 ± 0.28, 0.25 ± 0.11, 1.58  
251 ± 0.68, and 0.89 ± 0.42 μg N m<sup>-3</sup>, respectively (Figure 4). High concentrations of reactive nitrogen  
252 components focused on the northern part of Italy, central and southern part of Germany, North  
253 France, Poland, and the western part of Russia, which was in good agreement with the spatial pattern  
254 of NO<sub>x</sub> and NH<sub>3</sub> emissions (Luo et al., 2022; Qu et al., 2020). Emissions Database for Global  
255 Atmospheric Research (EDGAR) suggested that N fertilization and N manure accounted for 43%  
256 and 53% of total NH<sub>3</sub> emissions in western Europe (Liu et al., 2019), respectively. Furthermore, Liu

257 et al. (2019) confirmed that a good relationship between ambient NH<sub>3</sub> level and N fertilization plus  
258 N manure ( $r = 0.62$ ) was observed in Europe. Cooper et al. (2017) employed the inversion model to  
259 estimate NO<sub>x</sub> emission in Europe and also found that high NO<sub>x</sub> emission was also mainly distributed  
260 on North France, Germany, the northern part of Italy, and Russia, which partly explained the higher  
261 concentrations of reactive nitrogen components in these regions.

262 In the United States, the ambient NO<sub>3</sub><sup>-</sup> (NO<sub>3</sub>-N), HNO<sub>3</sub>, NH<sub>3</sub>, and NH<sub>4</sub><sup>+</sup> (NH<sub>4</sub>-N)  
263 concentrations reached  $0.28 \pm 0.12$ ,  $0.19 \pm 0.08$ ,  $2.12 \pm 0.66$ , and  $0.49 \pm 0.25$   $\mu\text{g N m}^{-3}$ , with the  
264 range of 0.03-2.35, 0.03-1.31, 0.26-9.96, and 0.10-6.09  $\mu\text{g N m}^{-3}$  (Figure 5), respectively. The  
265 hotspots of NO<sub>3</sub><sup>-</sup> (NO<sub>3</sub>-N), HNO<sub>3</sub>, and NH<sub>4</sub><sup>+</sup> (NH<sub>4</sub>-N) levels focused on the eastern part of the United  
266 States, while the higher NH<sub>3</sub> concentration focused on Central Great Plains and some regions in  
267 California such as San Joaquin Valley ( $6.15 \mu\text{g N m}^{-3}$ ). Both of bottom-up and top-down NO<sub>x</sub> and  
268 NH<sub>3</sub> emissions suggested that the spatial distributions of reactive nitrogen components were  
269 strongly dependent on the precursor emissions (McDuffie et al., 2020; Qu et al., 2020).

270 Besides, some other regions such as India (1.4, 0.5, 6.6, and 4.4  $\mu\text{g N m}^{-3}$ ) especially the  
271 northern part of India (3.1, 0.8, 12.6, and 8.4  $\mu\text{g N m}^{-3}$ ) also experienced severe reactive N pollution  
272 in the atmosphere. Meanwhile, some countries in South America such Brazil and Argentina and in  
273 Africa such as West Africa Coast (Nigeria, Ivory Coast, Ghana, Togo, and Benin) (HNO<sub>3</sub> and NH<sub>3</sub>:  
274 0.3 and 5.0  $\mu\text{g N m}^{-3}$ ) and Democratic Congo (0.4 and 1.6  $\mu\text{g N m}^{-3}$ ) also suffered from serious  
275 HNO<sub>3</sub> (Brazil and Argentina: 0.3 and 0.2  $\mu\text{g N m}^{-3}$ ) and NH<sub>3</sub> (3.6 and 2.8  $\mu\text{g N m}^{-3}$ ) pollution. The  
276 higher ambient NH<sub>3</sub> concentration focused on the northern part of India might be contributed by  
277 two major reasons. First of all, the intensive agricultural activities and high air temperature might  
278 be responsible for the higher NH<sub>3</sub> level (Cui, 2023; Wang et al., 2020). Moreover, the relatively low  
279 sulfur dioxide (SO<sub>2</sub>) and nitrogen oxides (NO<sub>x</sub>) emissions coupled with high air temperature  
280 restricted the gas-to-particle conversion of NH<sub>3</sub> (Wang et al., 2020). The severe HNO<sub>3</sub> and NH<sub>3</sub>  
281 pollution in Brazil, Argentina, and West Africa Coast might be also linked with the dense agricultural  
282 activities (Huneus et al., 2017).

### 283 3.3 The seasonal variations of reactive nitrogen components

284 The ambient NO<sub>3</sub><sup>-</sup> (NO<sub>3</sub>-N), HNO<sub>3</sub>, NH<sub>3</sub>, and NH<sub>4</sub><sup>+</sup> (NH<sub>4</sub>-N) concentrations exhibited  
285 significant seasonal variations (Figure S4-8). NO<sub>3</sub><sup>-</sup>, HNO<sub>3</sub>, and NH<sub>4</sub><sup>+</sup> displayed the highest and

286 lowest values in winter (December-February) and summer (June-August), respectively. On the one  
287 hand, the anthropogenic NO<sub>x</sub> emission for domestic heating might be higher in winter compared  
288 with other seasons (Lin et al., 2011). On the other hand, the stagnant meteorological conditions  
289 limited the pollutant diffusion (Li et al., 2019b; Liu et al., 2020c). Meanwhile, the higher relative  
290 humidity in winter facilitated the formation of NH<sub>4</sub>NO<sub>3</sub> (Huang et al., 2016; Xu et al., 2012).  
291 However, both of ambient NO<sub>3</sub><sup>-</sup> and NH<sub>4</sub><sup>+</sup> concentrations showed the lower concentrations in  
292 summer, which might be attributable to the decomposition of NH<sub>4</sub>NO<sub>3</sub> under the condition of high  
293 air temperature. In contrast to the secondary inorganic nitrogen, the ambient NH<sub>3</sub> level showed the  
294 highest concentration in summer ( $1.71 \pm 0.45 \mu\text{g N m}^{-3}$ ). China, Europe, and the United States  
295 suffered from similar NH<sub>3</sub> peaks in summer ( $4.20 \pm 1.85$ ,  $1.77 \pm 0.65$ , and  $2.21 \pm 1.04 \mu\text{g N m}^{-3}$ ).  
296 There are two reasons accounting for the fact. At first, mineral N fertilizer or manure application  
297 was mainly performed in spring and early summer (Paulot et al., 2014). Many field observations  
298 have obtained similar NH<sub>3</sub> peak in summer (He et al., 2021; Pan et al., 2018). Moreover, summer  
299 often showed the higher air temperature, which promotes the volatilization of ammonium and limits  
300 the gas-to-particle of gaseous NH<sub>3</sub> (Liu et al., 2019).

#### 301 3.4 The historical trends of reactive nitrogen components during 2000-2019

302 The long-term trends of ambient NO<sub>3</sub><sup>-</sup> (NO<sub>3</sub>-N), HNO<sub>3</sub>, NH<sub>3</sub>, and NH<sub>4</sub><sup>+</sup> (NH<sub>4</sub>-N) concentrations  
303 are shown in Figure 6 and Figure S9-12. The NO<sub>3</sub><sup>-</sup> concentration in China displayed rapid increase  
304 (9.7%/yr) during 2000-2007, and then it kept the moderate increase (4.2%/yr) during 2007-2013.  
305 However, the NO<sub>3</sub><sup>-</sup> (NO<sub>3</sub>-N) concentration in China experienced the drastic decrease (-2.6%/yr)  
306 since 2013. The ambient HNO<sub>3</sub> and NH<sub>4</sub><sup>+</sup> (NH<sub>4</sub>-N) concentrations showed similar trends during this  
307 period. Due to the impact of Clean Air Action, the concentrations of gaseous precursors (e.g., SO<sub>2</sub>  
308 and NO<sub>x</sub>) suffered from substantial decreases, which could be transformed into nitrate and  
309 ammonium via heterogeneous reactions (Huang et al., 2019). However, the decreasing rates of NO<sub>3</sub><sup>-</sup>  
310 were still much lower than those of gaseous precursors (Li et al., 2023). On the one hand, it might  
311 be associated with the increased O<sub>3</sub> level and enhanced atmospheric oxidation capacity (AOC),  
312 which led to an increase in the photochemical reaction rate of the secondary components (Wang et  
313 al., 2019). On the other hand, strong SO<sub>2</sub> emission control under the Clean Air Action allowed more  
314 gaseous NH<sub>3</sub> to form nitrate. The ambient NH<sub>3</sub> level remained relatively stable status during 2000-

315 2013, while it experienced rapid increases after 2013. The result was in good agreement with Liu et  
316 al. (2019). In fact, the ambient  $\text{NH}_3$  level in North China Plain still experienced dramatic increase  
317 ( $> 0.2 \mu\text{g N m}^{-3}/\text{yr}$ ) during 2000-2013 because enhanced agricultural activities. Zhang et al. (2017)  
318 have demonstrated that the livestock manure and fertilizer application generally accounted for 43%  
319 and 36% of the agricultural  $\text{NH}_3$  emission, respectively. Since 2013, the  $\text{NH}_3$  concentration in the  
320 entire China suffered from rapid increase, which might be associated with the drastic decrease of  
321 sulfate. It was well known that  $\text{NH}_3$  could react with  $\text{HNO}_3$  and gaseous  $\text{H}_2\text{SO}_4$  to generate ammonia  
322 sulfate and ammonia nitrate (Wang et al., 2022; Wang et al., 2019). Substantial decreases of acidic  
323 gases (e.g.,  $\text{SO}_2$ ) lead to the reduction of  $\text{NH}_3$  conversion to ammonia salts in the atmosphere (Chen  
324 et al., 2019), which result in excess  $\text{NH}_3$  remaining in the gaseous phase. Different from China, the  
325 reactive N concentrations in some other Asia and Africa countries especially India ( $\text{NO}_3^-$ ,  $\text{HNO}_3$ ,  
326  $\text{NH}_3$ , and  $\text{NH}_4^+$ : 54%, 46%, 11%, and 94%), South Korea (76%, 42%, 40%, and 9%), Indonesia  
327 (21%, 5%, 14%, and 41%), Democratic Congo (9%, 16%, 145%, 41%), and West Africa Coast (4%,  
328 11%, 37%, and 106%) still exhibited stable increases during 2000-2019. The results indicated that  
329 no strong reactive N emission control measures were implemented in these countries, which should  
330 be further exerted imperatively.

331 Compared with China, the long-term trends of reactive nitrogen components Europe and the  
332 United States were relatively stable. In the Europe, the concentrations of  $\text{NO}_3^-$  ( $\text{NO}_3\text{-N}$ ),  $\text{HNO}_3$ , and  
333  $\text{NH}_3$  exhibited increases during 2000-2007 (0.7%/yr, 2.3%/yr, and 2.1%/yr), while they experienced  
334 slight decreases after 2007. The  $\text{NH}_4^+$  ( $\text{NH}_4\text{-N}$ ) level displayed continuous decrease since 2000. The  
335 result was closely linked with the trends of  $\text{NO}_x$  and  $\text{NH}_3$  emissions derived from satellite retrieval  
336 (Cooper et al., 2017; Luo et al., 2022). In the United States, both of  $\text{NO}_3^-$  and  $\text{NH}_4^+$  showed persistent  
337 decreases during 2000-2019. Zhang et al. (2018) have confirmed that  $\text{NO}_x$  emission in the eastern  
338 US has experienced persistent decrease since 1990, which facilitated the decreases of  $\text{NO}_3^-$  and  
339  $\text{NH}_4^+$  levels. However, the ambient  $\text{HNO}_3$  and  $\text{NH}_3$  concentrations displayed slight increases during  
340 2000-2007 (2.1%/yr), and then remained relatively stable since 2007. Liu et al. (2019) also found  
341 similar characteristic of ambient  $\text{NH}_3$  trend in the United States. In fact, the  $\text{NH}_3$  concentrations in  
342 the Middle Plain and eastern US still showed increases due to the lack of  $\text{NH}_3$  emission control  
343 policies as well as the decline in acidic gases (Warner et al., 2017). The reactive N concentrations

344 in some countries in South America such as Brazil (9%, 0%, 13%, and 34%) and Argentina (10%,  
345 12%, 18%, and 7%) also remained relatively stable because local anthropogenic emission of reactive  
346 N did not show dramatic increases in the past two decades (McDuffie et al., 2020).

### 347 3.5 Projection of future ambient reactive nitrogen components

348 For the future reactive nitrogen component estimates, the ensemble model was applied to  
349 predict the reactive nitrogen component concentrations under the SSP1-2.6, SSP2-4.5, SSP3-7.0,  
350 and SSP5-8.5 scenarios. SSP1-2.6 represents the low emission pathways. In SSP1-2.6, the projected  
351 average  $\text{NO}_3^-$  concentrations in most countries experienced rapid decreases from 2020 to 2100  
352 (Figure 7 and Table 1). The mean concentrations of  $\text{NO}_3^-$  in China, India, Europe, and the United  
353 States decreased from  $1.16 \pm 0.35$ ,  $1.23 \pm 0.42$ ,  $0.41 \pm 0.14$ , and  $0.27 \pm 0.09 \mu\text{g N m}^{-3}$  to  $0.33 \pm 0.10$ ,  
354  $0.65 \pm 0.21$ ,  $0.10 \pm 0.03$ , and  $0.06 \pm 0.02 \mu\text{g N m}^{-3}$  during 2020-2100 in SSP1-2.6 scenario,  
355 respectively. Besides, the  $\text{NO}_3^-$  concentrations in many other countries of Africa and South America  
356 such as Brazil (-127%) and Democratic Congo (-162%) also suffered from drastic decreases in this  
357 scenario. SSP2-4.5 scenario represents the middle range of plausible future pathways (Nazarenko  
358 et al., 2022). In this scenario, the predicted average  $\text{NO}_3^-$  concentrations in China, India, Europe,  
359 and the United States decreased from  $1.19 \pm 0.40$ ,  $1.43 \pm 0.35$ ,  $0.44 \pm 0.13$ , and  $0.24 \pm 0.08 \mu\text{g N}$   
360  $\text{m}^{-3}$  to  $0.41 \pm 0.14$ ,  $0.95 \pm 0.32$ ,  $0.24 \pm 0.08$ , and  $0.05 \pm 0.02 \mu\text{g N m}^{-3}$  during 2020-2100, respectively.  
361 SSP3-7.0 and SSP5-8.5 denote the less investment in the environment and heavily relies on  
362 traditional energy for rapid economic development, respectively. The ambient  $\text{NO}_3^-$  in these  
363 scenarios generally showed the higher concentrations compared with other scenarios. For instance,  
364 the  $\text{NO}_3^-$  concentrations in China reduced from  $1.25 \pm 0.40$  (SSP3-7.0) and  $1.21 \pm 0.39$  (SSP5-8.5)  
365 to  $0.75 \pm 0.25$  (SSP3-7.0) and  $0.58 \pm 0.18$  (SSP5-8.5)  $\mu\text{g N m}^{-3}$  during 2020-2100 (Figure 8),  
366 respectively. The higher  $\text{NO}_3^-$  concentrations in SSP3-7.0 and SSP5-8.5 might be associated with  
367 the higher anthropogenic  $\text{NO}_x$  emission. Compared with SSP1-2.6 and SSP2-4.5, the  $\text{NO}_3^-$   
368 concentrations in some countries during SSP3-7.0 and SSP5-8.5 scenarios displayed slight increases  
369 from 2020 to 2040. For instance, the ambient  $\text{NO}_3^-$  concentrations in Indonesia increased by 12%  
370 (SSP3-7.0) and 5% (SSP5-8.5), respectively.

371 The temporal variations of ambient  $\text{HNO}_3$  were similar to those of  $\text{NO}_3^-$  concentrations. The  
372 mean concentrations of  $\text{HNO}_3$  in China, India, Europe, and the United States decreased from 0.25

373  $\pm 0.09$ ,  $0.50 \pm 0.16$ ,  $0.18 \pm 0.06$ , and  $0.08 \pm 0.03$   $\mu\text{g N m}^{-3}$  to  $0.05 \pm 0.01$ ,  $0.24 \pm 0.08$ ,  $0.05 \pm 0.02$ ,  
374 and  $0.03 \pm 0.01$   $\mu\text{g N m}^{-3}$  during 2020-2100 in SSP1-2.6 scenario (Figure S13-S14 and Table S3),  
375 respectively. However, the decreasing ratios of ambient  $\text{HNO}_3$  levels especially in some developing  
376 countries such as Democratic Congo (-13%) and West Africa Coast (-47%) were much less than  
377 those of ambient  $\text{NO}_3^-$  levels. For the SSP3-7.0 and SSP5-8.5 scenarios, the  $\text{HNO}_3$  levels in some  
378 developing countries such as Democratic Congo (18%), West Africa Coast (16%), Indonesia (13%)  
379 even experienced moderate increases. It was assumed that the government gave less investment in  
380 environment improvement and the anthropogenic emission did not show marked decrease under the  
381 condition of SSP3-7.0 scenario (Chen et al., 2023; Chen et al., 2020).

382 As shown in Figure S15-S18 and Table S4-S5, the higher ambient  $\text{NH}_3$  and  $\text{NH}_4^+$   
383 concentrations also focused on India and North China. In SSP1-2.6, the ambient  $\text{NH}_3$  ( $\text{NH}_4^+$ )  
384 concentrations in China, India, Europe, and the United States decreased from  $3.51 \pm 1.12$  ( $2.00 \pm$   
385  $0.62$ ),  $6.30 \pm 2.12$  ( $4.26 \pm 1.42$ ),  $1.54 \pm 0.51$  ( $0.75 \pm 0.24$ ), and  $1.79 \pm 0.59$  ( $0.53 \pm 0.17$ )  $\mu\text{g N m}^{-3}$   
386 to  $1.75 \pm 0.58$  ( $0.58 \pm 0.19$ ),  $2.57 \pm 0.85$  ( $1.25 \pm 0.41$ ),  $1.15 \pm 0.36$  ( $0.50 \pm 0.16$ ), and  $1.58 \pm 0.52$   
387 ( $0.45 \pm 0.15$ )  $\mu\text{g N m}^{-3}$  during 2020-2100. Compared with SSP1-2.6, the ambient  $\text{NH}_3$  and  $\text{NH}_4^+$   
388 concentrations in heavy-pollution scenarios (SSP3-7.0, and SSP5-8.5) scenarios did not show  
389 marked decreases from 2020-2100. Some developing countries such as Argentina (9%), Democratic  
390 Congo (25%), and West Africa Coast (24%) even suffered from persistent increases of ambient  $\text{NH}_3$   
391 and  $\text{NH}_4^+$  levels. It might be associated with ineffective control of  $\text{NH}_3$  emission compared with  
392  $\text{NO}_x$  emission.

### 393 3.6 Conclusions and limitations

394 The ground-level ambient reactive nitrogen observations, satellite retrievals, GEOS-Chem  
395 model output, and many other geographical covariates were integrated into the multi-stage model  
396 to reveal the global patterns of ambient reactive nitrogen components during 2000-2019. Then, these  
397 high-resolution reactive nitrogen dataset during the historical period was then utilized as the  
398 constraint to calibrate the CMIP6 dataset in four scenarios during 2020-2100. The results indicated  
399 the cross-validation (CV)  $R^2$  values of four reactive nitrogen species showed satisfied performance  
400 ( $R^2 > 0.55$ ). At the spatial scale, four reactive nitrogen components exhibited the higher  
401 concentrations in China and India. For the temporal variations, the concentrations of estimated

402 ambient reactive nitrogen components in China experienced persistent increases during 2000-2013,  
403 while they suffered from drastic decreases since 2013 except NH<sub>3</sub>, which might be linked with the  
404 impact of clean air policy. However, the concentrations of these species in Europe and the United  
405 States remained relatively stable since 2000. In the future scenarios, SSP3-7.0 (traditional energy  
406 scenario) and SSP1-2.6 (carbon neutrality scenario) displayed the highest and lowest reactive  
407 nitrogen component concentrations, respectively.

408 Global trends of four reactive nitrogen components during 2000-2100 emphasizes the urgent  
409 mitigation measures (carbon neutrality pathway) to reduce precursor emissions in order to decrease  
410 the concentrations and depositions of reactive nitrogen components especially in China and India.  
411 Furthermore, our result could give valuable insights into the impact of reactive nitrogen components  
412 on human health and ecological environment. However, this study still shows some limitations. First  
413 of all, the observation networks mainly focus on China, Europe, and the United States, and thus the  
414 simulations in many other regions might show large uncertainties. Secondly, both of the GEOS-  
415 Chem output and CMIP6 future climate scenario data also exhibits large uncertainties, which could  
416 impact the reliability of this study. Lastly, our predictions were performed on the basis of the premise  
417 that the world was steadily developing, and cannot predict the impacts of uncontrollable factors  
418 (e.g., COVID-19, Russia-Ukraine War).

#### 419 **Competing interests**

420 The contact author has declared that none of the authors has any competing interests.

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#### 423 **Data availability**

424 The CMIP6 dataset used in this publication is available in <https://esgf.nci.org.au/projects/cmip6-nci/>.

#### 426 **Author contributions**

427 LR and WGH designed the study; LR developed the model; GYN, ZLJ, and SYB analyzed the  
428 observation and model data. LR wrote this manuscript.

429

430 **References**

- 431 Altieri, K.E., Fawcett, S.E., Hastings, M.G. (2021) Reactive nitrogen cycling in the atmosphere and  
432 ocean. *Annual Review of Earth and Planetary Sciences* 49, 523-550.
- 433 Amante, C., Eakins, B.W. (2009) ETOPO1 arc-minute global relief model: procedures, data sources and  
434 analysis.
- 435 Bouwman, A., Van Vuuren, D., Derwent, R., Posch, M. (2002) A global analysis of acidification and  
436 eutrophication of terrestrial ecosystems. *Water, Air, and Soil Pollution* 141, 349-382.
- 437 Chen, H., Li, D., Gurnesa, G.A., Yu, G., Li, L., Zhang, W., Fang, H., Mo, J. (2015) Effects of nitrogen  
438 deposition on carbon cycle in terrestrial ecosystems of China: A meta-analysis. *Environmental  
439 pollution* 206, 352-360.
- 440 Chen, W., Lu, X., Yuan, D., Chen, Y., Li, Z., Huang, Y., Fung, T., Sun, H., Fung, J.C. (2023) Global PM2.  
441 5 Prediction and Associated Mortality to 2100 under Different Climate Change Scenarios.  
442 *Environmental science & technology* 57, 10039-10052.
- 443 Chen, X., Wang, Y.-h., Ye, C., Zhou, W., Cai, Z.-c., Yang, H., Han, X. (2018) Atmospheric nitrogen  
444 deposition associated with the eutrophication of Taihu Lake. *Journal of Chemistry* 2018.
- 445 Chen, Y., Liu, A., Cheng, X. (2020) Quantifying economic impacts of climate change under nine future  
446 emission scenarios within CMIP6. *Science of the Total Environment* 703, 134950.
- 447 Chen, Y., Shen, H., Russell, A.G. (2019) Current and future responses of aerosol pH and composition in  
448 the US to declining SO<sub>2</sub> emissions and increasing NH<sub>3</sub> emissions. *Environmental science &  
449 technology* 53, 9646-9655.
- 450 Chen, Y., Zhang, L., Henze, D.K., Zhao, Y., Lu, X., Winiwarter, W., Guo, Y., Liu, X., Wen, Z., Pan, Y.  
451 (2021) Interannual variation of reactive nitrogen emissions and their impacts on PM2. 5 air pollution  
452 in China during 2005–2015. *Environmental Research Letters* 16, 125004.
- 453 Cooper, M., Martin, R.V., Padmanabhan, A., Henze, D.K. (2017) Comparing mass balance and adjoint  
454 methods for inverse modeling of nitrogen dioxide columns for global nitrogen oxide emissions.  
455 *Journal of Geophysical Research: Atmospheres* 122, 4718-4734.
- 456 Cooper, M.J., Martin, R.V., Hammer, M.S., Levelt, P.F., Veefkind, P., Lamsal, L.N., Krotkov, N.A., Brook,  
457 J.R., McLinden, C.A. (2022) Global fine-scale changes in ambient NO<sub>2</sub> during COVID-19 lockdowns.  
458 *Nature* 601, 380-387.
- 459 Cui, L. (2023) Impact of COVID-19 restrictions on the concentration and source apportionment of  
460 atmospheric ammonia (NH<sub>3</sub>) across India. *Science of the Total Environment* 881, 163443.
- 461 Drury, E., Jacob, D.J., Spurr, R.J.D., Wang, J., Shinozuka, Y., Anderson, B.E., Clarke, A.D., Dibb, J.,  
462 McNaughton, C. Weber, D., Synthesis of satellite (MODIS), aircraft (ICARTT), and surface  
463 (IMPROVE, EPA-AQS, AERONET) aerosol observations over eastern North America to improve  
464 MODIS aerosol retrievals and constrain surface aerosol concentrations and sources *J. Geophys. Res.  
465 Atmos.*, 115 (2010), p. D14204.
- 466 Du, E., de Vries, W., Galloway, J.N., Hu, X., Fang, J. (2014) Changes in wet nitrogen deposition in the  
467 United States between 1985 and 2012. *Environmental Research Letters* 9, 095004.
- 468 Ehrnsperger, L., Klemm, O. (2021) Source apportionment of urban ammonia and its contribution to  
469 secondary particle formation in a Mid-size European City. *Aerosol and Air Quality Research* 21,  
470 200404.
- 471 Feng, X., Lin, H., Fu, T.-M., Sulprizio, M.P., Zhuang, J., Jacob, D.J., Tian, H., Ma, Y., Zhang, L., Wang,  
472 X. (2021) WRF-GC (v2. 0): online two-way coupling of WRF (v3. 9.1. 1) and GEOS-Chem (v12. 7.2)  
473 for modeling regional atmospheric chemistry–meteorology interactions. *Geoscientific Model*



474 Development 14, 3741-3768.

475 Geddes, J.A., Martin, R.V. (2017) Global deposition of total reactive nitrogen oxides from 1996 to 2014  
476 constrained with satellite observations of NO<sub>2</sub> columns. *Atmos. Chem. Phys.* 17, 10071-10091.

477 He, Y., Pan, Y., Gu, M., Sun, Q., Zhang, Q., Zhang, R., Wang, Y. (2021) Changes of ammonia  
478 concentrations in wintertime on the North China Plain from 2018 to 2020. *Atmospheric Research* 253,  
479 105490.

480 He, Z., Liu, P., Zhao, X., He, X., Liu, J., Mu, Y. (2022) Responses of surface O<sub>3</sub> and PM<sub>2.5</sub> trends to  
481 changes of anthropogenic emissions in summer over Beijing during 2014–2019: A study based on  
482 multiple linear regression and WRF-Chem. *Science of the Total Environment* 807, 150792.

483 Hoesly, R.M., Smith, S.J., Feng, L., Klimont, Z., Janssens-Maenhout, G., Pitkanen, T., Seibert, J.J., Vu,  
484 L., Andres, R.J., Bolt, R.M. (2018) Historical (1750–2014) anthropogenic emissions of reactive gases  
485 and aerosols from the Community Emissions Data System (CEDS). *Geoscientific Model Development*  
486 11, 369-408.

487 Huang, L., An, J., Koo, B., Yarwood, G., Yan, R., Wang, Y., Huang, C., Li, L. (2019) Sulfate formation  
488 during heavy winter haze events and the potential contribution from heterogeneous SO<sub>2</sub>+NO<sub>2</sub>  
489 reactions in the Yangtze River Delta region, China. *Atmospheric Chemistry and Physics* 19, 14311-  
490 14328.

491 Huang, T., Chen, J., Zhao, W., Cheng, J., Cheng, S. (2016) Seasonal variations and correlation analysis  
492 of water-soluble inorganic ions in PM<sub>2.5</sub> in Wuhan, 2013. *Atmosphere* 7, 49.

493 Huang, X., Liu, Z., Liu, J., Hu, B., Wen, T., Tang, G., Zhang, J., Wu, F., Ji, D., Wang, L. (2017) Chemical  
494 characterization and source identification of PM<sub>2.5</sub> at multiple sites in the Beijing–Tianjin–Hebei  
495 region, China. *Atmospheric Chemistry and Physics* 17, 12941-12962.

496 Huneus, N., Granier, C., Dawidowski, L., van Der Gon, H.D., Alonso, M., Castesana, P., Diaz, M., Frost,  
497 G.J., Gallardo, L., Gomez, D., (2017) Anthropogenic emissions in South America for air quality and  
498 climate modelling, 2017 International Emission Inventory Conference “Applying Science and  
499 Streamlining Processes to Improve Inventories”.

500 Kang, Y., Liu, M., Song, Y., Huang, X., Yao, H., Cai, X., Zhang, H., Kang, L., Liu, X., Yan, X. (2016)  
501 High-resolution ammonia emissions inventories in China from 1980 to 2012. *Atmospheric Chemistry  
502 and Physics* 16, 2043-2058.

503 Karra, K., Kontgis, C., Statman-Weil, Z., Mazzariello, J.C., Mathis, M., Brumby, S.P., (2021) Global  
504 land use/land cover with Sentinel 2 and deep learning, 2021 IEEE international geoscience and remote  
505 sensing symposium IGARSS. IEEE, pp. 4704-4707.

506 Kim, H.C., Lee, P., Judd, L., Pan, L., Lefer, B. (2016) OMI NO<sub>2</sub> column densities over North American  
507 urban cities: the effect of satellite footprint resolution. *Geoscientific Model Development* 9, 1111-1123.

508 Li, H., Yang, Y., Wang, H., Wang, P., Yue, X., Liao, H. (2022) Projected aerosol changes driven by  
509 emissions and climate change using a machine learning method. *Environmental science & technology*  
510 56, 3884-3893.

511 Li, H., Zhang, Q., Zhang, Q., Chen, C., Wang, L., Wei, Z., Zhou, S., Parworth, C., Zheng, B., Canonaco,  
512 F. (2017) Wintertime aerosol chemistry and haze evolution in an extremely polluted city of the North  
513 China Plain: significant contribution from coal and biomass combustion. *Atmospheric Chemistry and  
514 Physics* 17, 4751-4768. Li, M.; Kurokawa, J.; Zhang, Q.; Woo, J.-H.; Morikawa, T.; Chatani, S.; Lu,  
515 Z.; Song, Y.; Geng, G.; Hu, H., MIXv2: a long-term mosaic emission inventory for Asia (2010-2017).  
516 *Atmospheric Chemistry and Physics* 2024, 24, (7), 3925-3952.

517 Li, R., Cui, L., Fu, H., Zhao, Y., Zhou, W., Chen, J. (2020) Satellite-Based Estimates of Wet Ammonium

518 (NH<sub>4</sub>-N) Deposition Fluxes Across China during 2011–2016 Using a Space–Time Ensemble Model.  
519 Environmental science & technology 54, 13419-13428.

520 Li, R., Cui, L., Zhao, Y., Zhang, Z., Sun, T., Li, J., Zhou, W., Meng, Y., Huang, K., Fu, H. (2019a) Wet  
521 deposition of inorganic ions in 320 cities across China: spatio-temporal variation, source  
522 apportionment, and dominant factors. Atmospheric Chemistry and Physics 19, 11043-11070.

523 Li, R., Gao, Y., Xu, J., Cui, L., Wang, G. (2023) Impact of Clean Air Policy on Criteria Air Pollutants  
524 and Health Risks Across China During 2013-2021. Journal of Geophysical Research: Atmospheres  
525 128, e2023JD038939.

526 Li, R., Wang, Z., Cui, L., Fu, H., Zhang, L., Kong, L., Chen, W., Chen, J. (2019b) Air pollution  
527 characteristics in China during 2015-2016: Spatiotemporal variations and key meteorological factors.  
528 Science of the Total Environment 648, 902-915.

529 Li, Y., Schichtel, B.A., Walker, J.T., Schwede, D.B., Chen, X., Lehmann, C.M., Puchalski, M.A., Gay,  
530 D.A., Collett Jr, J.L. (2016) Increasing importance of deposition of reduced nitrogen in the United  
531 States. Proceedings of the National Academy of Sciences 113, 5874-5879.

532 Lin, J.T., McElroy, M.B., (2010) Impacts of boundary layer mixing on pollutant vertical profiles in the  
533 lower troposphere: implications to satellite remote sensing Atmos. Environ., 44, pp. 1726-1739.

534 Lin, W., Xu, X., Ge, B., Liu, X. (2011) Gaseous pollutants in Beijing urban area during the heating period  
535 2007-2008: variability, sources, meteorological, and chemical impacts. Atmospheric Chemistry and  
536 Physics 11, 8157-8170.

537 Liu, H., Gong, P., Wang, J., Clinton, N., Bai, Y., Liang, S. (2020a) Annual dynamics of global land cover  
538 and its long-term changes from 1982 to 2015. Earth System Science Data 12, 1217-1243.

539 Liu, H., Jacob, D.J., Bey, I., Yantosca, R.M. (2001) Constraints from 210Pb and 7Be on wet deposition  
540 and transport in a global three-dimensional chemical tracer model driven by assimilated  
541 meteorological fields. Journal of Geophysical Research: Atmospheres 106, 12109-12128.

542 Liu, L., Yang, Y., Xi, R., Zhang, X., Xu, W., Liu, X., Li, Y., Liu, P., Wang, Z. (2021) Global Wet-Reduced  
543 Nitrogen Deposition Derived From Combining Satellite Measurements With Output From a Chemistry  
544 Transport Model. Journal of Geophysical Research: Atmospheres 126, e2020JD033977.

545 Liu, L., Zhang, X., Wong, A.Y., Xu, W., Liu, X., Li, Y., Mi, H., Lu, X., Zhao, L., Wang, Z. (2019)  
546 Estimating global surface ammonia concentrations inferred from satellite retrievals. Atmospheric  
547 Chemistry and Physics 19, 12051-12066.

548 Liu, L., Zhang, X., Xu, W., Liu, X., Lu, X., Wei, J., Li, Y., Yang, Y., Wang, Z., Wong, A.Y. (2020b)  
549 Reviewing global estimates of surface reactive nitrogen concentration and deposition using satellite  
550 retrievals. Atmospheric Chemistry and Physics 20, 8641-8658.

551 Liu, Y., Zhou, Y., Lu, J. (2020c) Exploring the relationship between air pollution and meteorological  
552 conditions in China under environmental governance. Scientific reports 10, 14518.

553 Luo, Z., Zhang, Y., Chen, W., Van Damme, M., Coheur, P.-F., Clarisse, L. (2022) Estimating global  
554 ammonia (NH<sub>3</sub>) emissions based on IASI observations from 2008 to 2018. Atmospheric Chemistry  
555 and Physics 22, 10375-10388.

556 Ma, R., Yu, K., Xiao, S., Liu, S., Ciais, P., Zou, J. (2022) Data-driven estimates of fertilizer-induced soil  
557 NH<sub>3</sub>, NO and N<sub>2</sub>O emissions from croplands in China and their climate change impacts. Global  
558 Change Biology 28, 1008-1022.

559 McDuffie, E.E., Smith, S.J., O'Rourke, P., Tibrewal, K., Venkataraman, C., Marais, E.A., Zheng, B.,  
560 Crippa, M., Brauer, M., Martin, R.V. (2020) A global anthropogenic emission inventory of atmospheric  
561 pollutants from sector-and fuel-specific sources (1970–2017): an application of the Community

562 Emissions Data System (CEDS). *Earth System Science Data* 12, 3413-3442.

563 Nazarenko, L.S., Tausnev, N., Russell, G.L., Rind, D., Miller, R.L., Schmidt, G.A., Bauer, S.E., Kelley,  
564 M., Ruedy, R., Ackerman, A.S. (2022) Future climate change under SSP emission scenarios with  
565 GISS-E2. 1. *Journal of Advances in Modeling Earth Systems* 14, e2021MS002871.

566 Pan, Y., Tian, S., Zhao, Y., Zhang, L., Zhu, X., Gao, J., Huang, W., Zhou, Y., Song, Y., Zhang, Q. (2018)  
567 Identifying ammonia hotspots in China using a national observation network. *Environmental science  
568 & technology* 52, 3926-3934.

569 Paulot, F., Jacob, D.J., Pinder, R., Bash, J., Travis, K., Henze, D. (2014) Ammonia emissions in the  
570 United States, European Union, and China derived by high-resolution inversion of ammonium wet  
571 deposition data: Interpretation with a new agricultural emissions inventory (MASAGE\_NH3). *Journal  
572 of Geophysical Research: Atmospheres* 119, 4343-4364.

573 Potapov, P., Hansen, M.C., Pickens, A., Hernandez-Serna, A., Tyukavina, A., Turubanova, S., Zalles, V.,  
574 Li, X., Khan, A., Stolle, F. (2022) The global 2000-2020 land cover and land use change dataset  
575 derived from the Landsat archive: first results. *Frontiers in Remote Sensing* 3, 856903.

576 Qi, L., Zheng, H., Ding, D., Wang, S. (2023) Responses of sulfate and nitrate to anthropogenic emission  
577 changes in eastern China-in perspective of long-term variations. *Science of the Total Environment* 855,  
578 158875.

579 Qu, Z., Henze, D.K., Cooper, O.R., Neu, J.L. (2020) Impacts of global NO<sub>x</sub> inversions on NO<sub>2</sub> and  
580 ozone simulations. *Atmospheric Chemistry and Physics* 20, 13109-13130.

581 Shi, X., Zhao, C., Jiang, J.H., Wang, C., Yang, X., Yung, Y.L. (2018) Spatial representativeness of PM<sub>2.5</sub>  
582 concentrations obtained using observations from network stations. *Journal of Geophysical Research:  
583 Atmospheres* 123, 3145-3158.

584 Van Damme, M., Whitburn, S., Clarisse, L., Clerbaux, C., Hurtmans, D., Coheur, P.-F. (2017) Version 2  
585 of the IASI NH<sub>3</sub> neural network retrieval algorithm: near-real-time and reanalysed datasets.  
586 *Atmospheric Measurement Techniques* 10, 4905-4914.

587 Von Schneidemesser, E., Monks, P.S., Allan, J.D., Bruhwiler, L., Forster, P., Fowler, D., Lauer, A.,  
588 Morgan, W.T., Paasonen, P., Righi, M. (2015) Chemistry and the linkages between air quality and  
589 climate change. *Chemical reviews* 115, 3856-3897.

590 Wang, M., Xiao, M., Bertozzi, B., Marie, G., Rörup, B., Schulze, B., Bardakov, R., He, X.-C., Shen, J.,  
591 Scholz, W. (2022) Synergistic HNO<sub>3</sub>-H<sub>2</sub>SO<sub>4</sub>-NH<sub>3</sub> upper tropospheric particle formation. *Nature* 605,  
592 483-489.

593 Wang, T., Song, Y., Xu, Z., Liu, M., Xu, T., Liao, W., Yin, L., Cai, X., Kang, L., Zhang, H., Zhu, T. (2020)  
594 Why is the Indo-Gangetic Plain the region with the largest NH<sub>3</sub> column in the globe during pre-  
595 monsoon and monsoon seasons? *Atmos. Chem. Phys.* 20, 8727-8736.

596 Wang, Y., Li, W., Gao, W., Liu, Z., Tian, S., Shen, R., Ji, D., Wang, S., Wang, L., Tang, G. (2019) Trends  
597 in particulate matter and its chemical compositions in China from 2013–2017. *Science China Earth  
598 Sciences* 62, 1857-1871.

599 Warner, J., Dickerson, R., Wei, Z., Strow, L.L., Wang, Y., Liang, Q. (2017) Increased atmospheric  
600 ammonia over the world's major agricultural areas detected from space. *Geophysical Research Letters*  
601 44, 2875-2884.

602 Wen, L., Xue, L., Wang, X., Xu, C., Chen, T., Yang, L., Wang, T., Zhang, Q., Wang, W. (2018)  
603 Summertime fine particulate nitrate pollution in the North China Plain: increasing trends, formation  
604 mechanisms and implications for control policy. *Atmospheric Chemistry and Physics* 18, 11261-11275.

605 Wesely, M. (2007) Parameterization of surface resistances to gaseous dry deposition in regional-scale

606 numerical models. *Atmospheric Environment* 41, 52-63.

607 Whitburn, S., Van Damme, M., Clarisse, L., Bauduin, S., Heald, C., Hadji-Lazaro, J., Hurtmans, D.,  
608 Zondlo, M.A., Clerbaux, C., Coheur, P.F. (2016a) A flexible and robust neural network IASI-NH<sub>3</sub>  
609 retrieval algorithm. *Journal of Geophysical Research: Atmospheres* 121, 6581-6599.

610 Whitburn, S., Van Damme, M., Clarisse, L., Turquety, S., Clerbaux, C., Coheur, P.F. (2016b) Doubling  
611 of annual ammonia emissions from the peat fires in Indonesia during the 2015 El Niño. *Geophysical*  
612 *Research Letters* 43, 11,007-011,014.

613 Xu, L., Chen, X., Chen, J., Zhang, F., He, C., Zhao, J., Yin, L. (2012) Seasonal variations and chemical  
614 compositions of PM<sub>2.5</sub> aerosol in the urban area of Fuzhou, China. *Atmospheric Research* 104, 264-  
615 272.

616 Xu, W., Zhang, L., Liu, X. (2019) A database of atmospheric nitrogen concentration and deposition from  
617 the nationwide monitoring network in China. *Scientific data* 6, 1-6.

618 Zaehle, S. (2013) Terrestrial nitrogen-carbon cycle interactions at the global scale. *Philosophical*  
619 *Transactions of the Royal Society B: Biological Sciences* 368, 20130125.

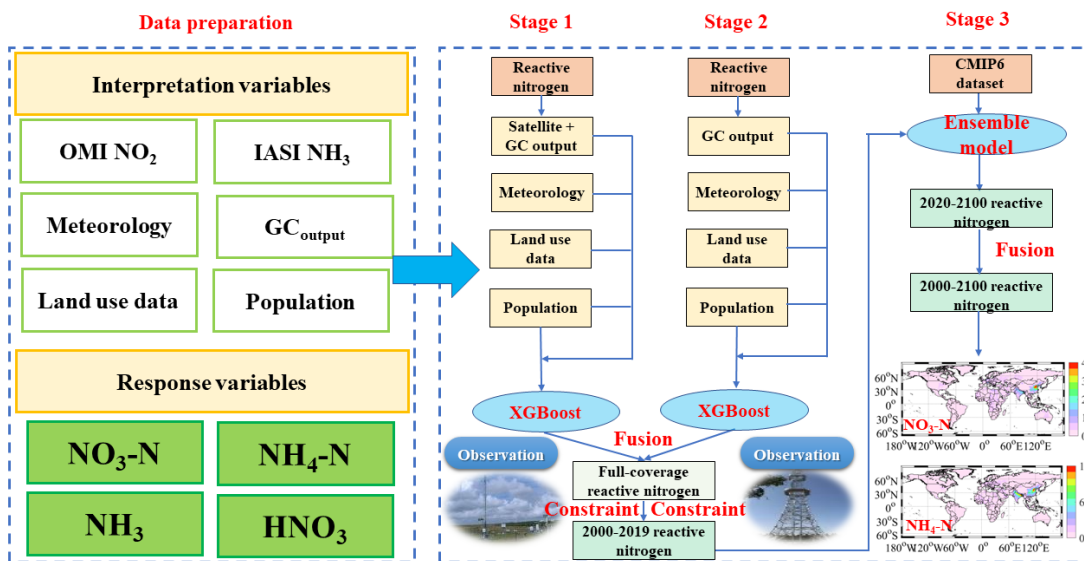
620 Zhang, L., Guo, X., Zhao, T., Gong, S., Xu, X., Li, Y., Luo, L., Gui, K., Wang, H., Zheng, Y. (2019) A  
621 modelling study of the terrain effects on haze pollution in the Sichuan Basin. *Atmospheric*  
622 *Environment* 196, 77-85.

623 Zhang, X., Ward, B.B., Sigman, D.M. (2020) Global nitrogen cycle: critical enzymes, organisms, and  
624 processes for nitrogen budgets and dynamics. *Chemical reviews* 120, 5308-5351.

625 Zhang, X., Wu, Y., Liu, X., Reis, S., Jin, J., Dragosits, U., Van Damme, M., Clarisse, L., Whitburn, S.,  
626 Coheur, P.-F. (2017) Ammonia emissions may be substantially underestimated in China.  
627 *Environmental science & technology* 51, 12089-12096.

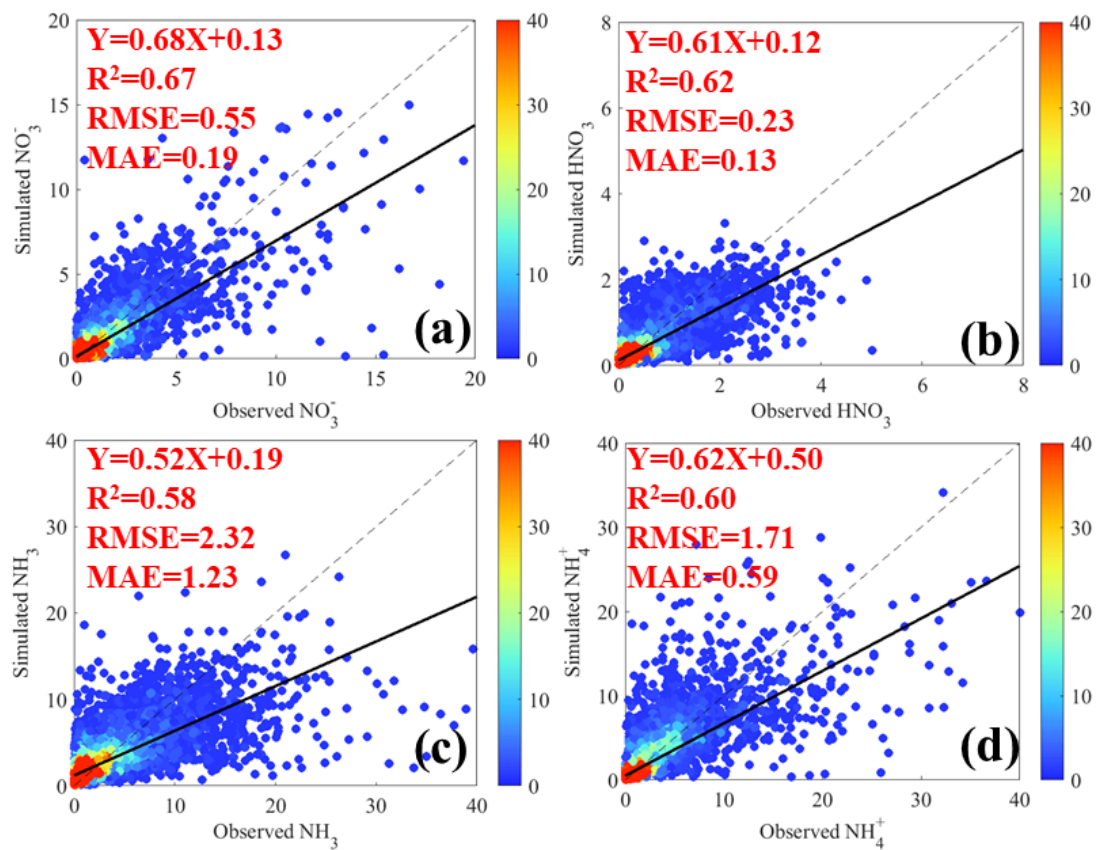
628 Zhang, Y., Mathur, R., Bash, J.O., Hogrefe, C., Xing, J., Roselle, S.J. (2018) Long-term trends in total  
629 inorganic nitrogen and sulfur deposition in the US from 1990 to 2010. *Atmospheric Chemistry and*  
630 *Physics* 18, 9091-9106.

631 **Figure 1** The workflow of global full-coverage reactive nitrogen estimates during 2000-2100.  
 632 GC<sub>output</sub> denotes the GEOS-Chem output.



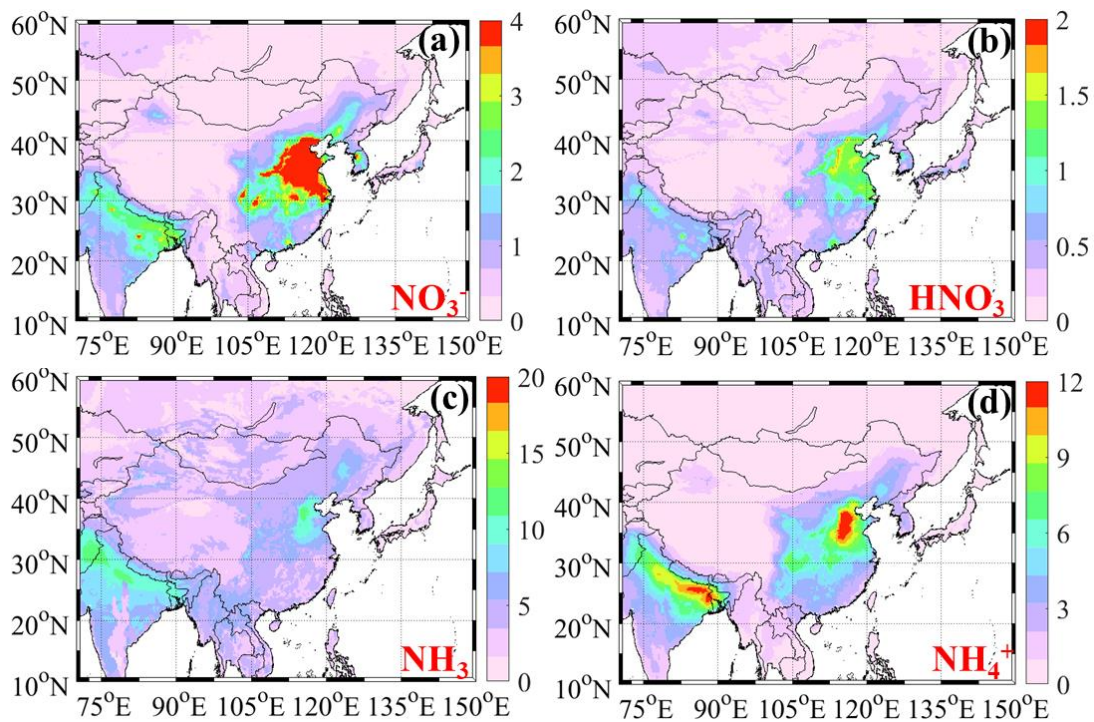
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634 **Figure 2** The predictive performances of four reactive nitrogen components including  $\text{NO}_3^-$  (a),  
635  $\text{HNO}_3$  (b),  $\text{NH}_3$  (c), and  $\text{NH}_4^+$  (d). The model was constructed with 90% original data and the  
636 remained data was applied to validate the model. The black solid line denotes the best-fitting curve  
637 for all of the points, while the black dashed line represents the diagonal, which means the same  
638 observed and simulated values. The color scale denotes the sample size.



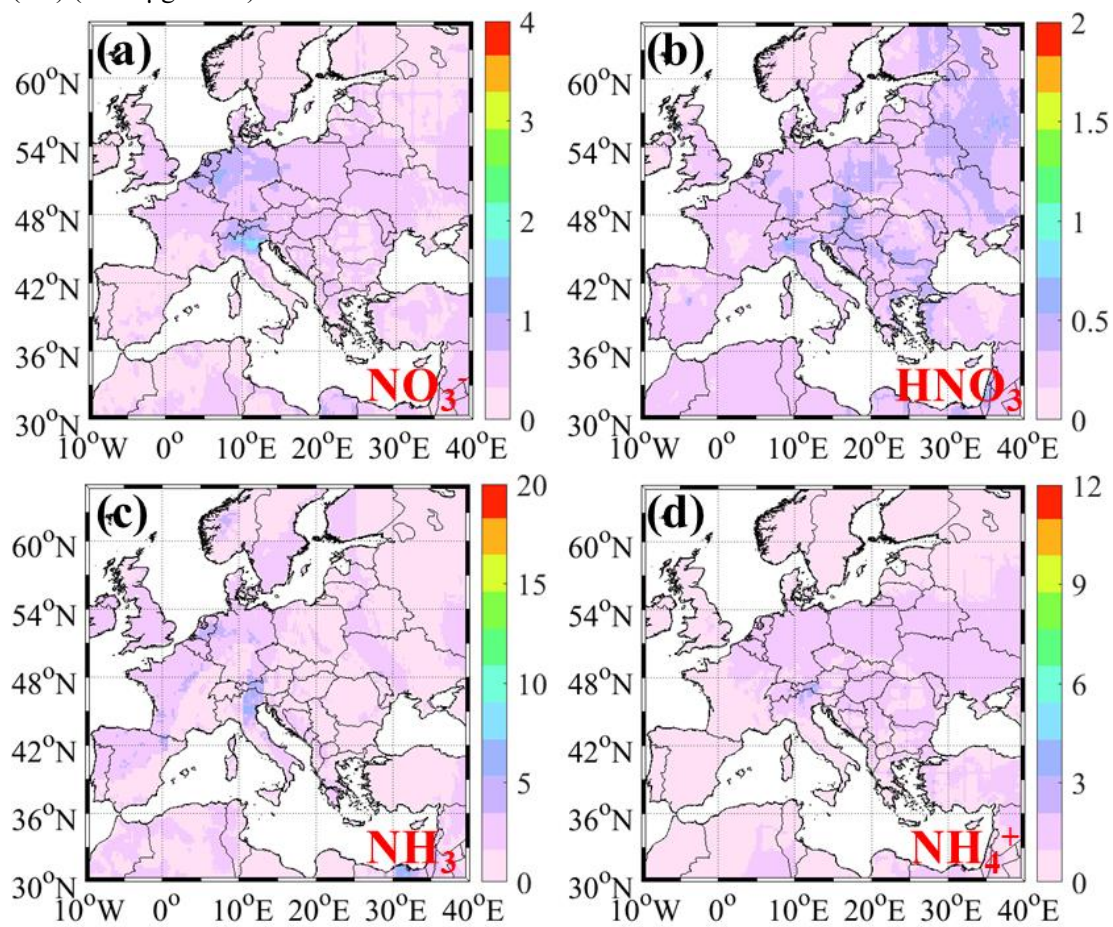
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640 **Figure 3** The spatiotemporal variations of  $\text{NO}_3^-$ ,  $\text{HNO}_3$ ,  $\text{NH}_3$ , and  $\text{NH}_4^+$  concentrations in East Asia  
641 (a-d) (Unit:  $\mu\text{g N m}^{-3}$ ).



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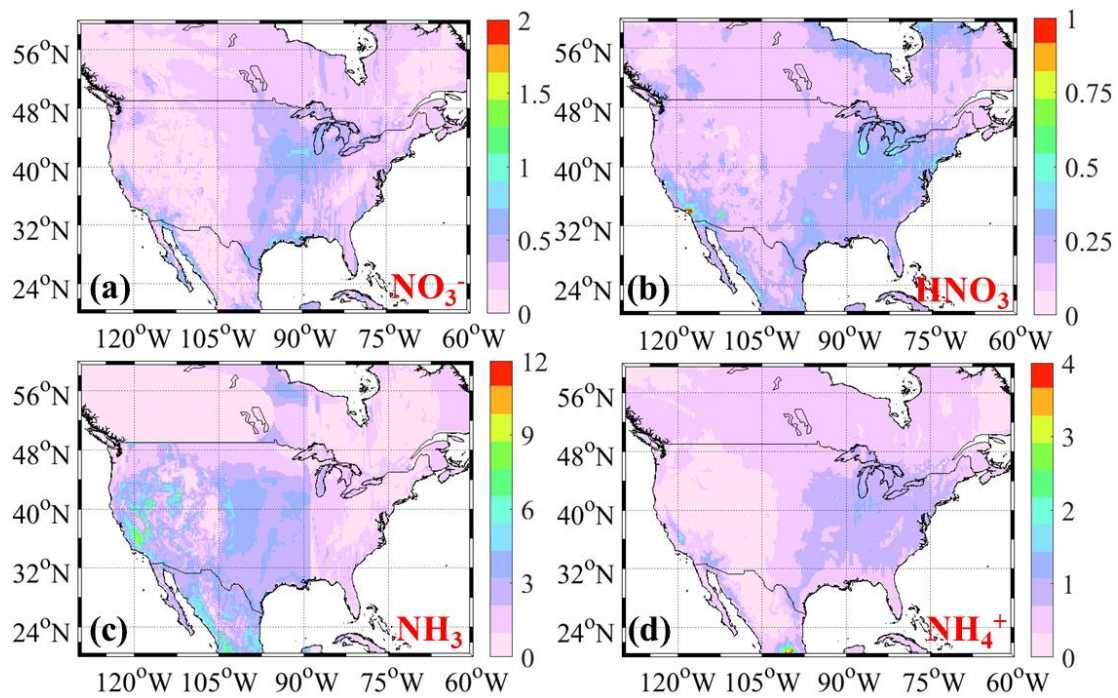
644 **Figure 4** The spatiotemporal variations of  $\text{NO}_3^-$ ,  $\text{HNO}_3$ ,  $\text{NH}_3$ , and  $\text{NH}_4^+$  concentrations in Europe  
645 (a-d) (Unit:  $\mu\text{g N m}^{-3}$ ).



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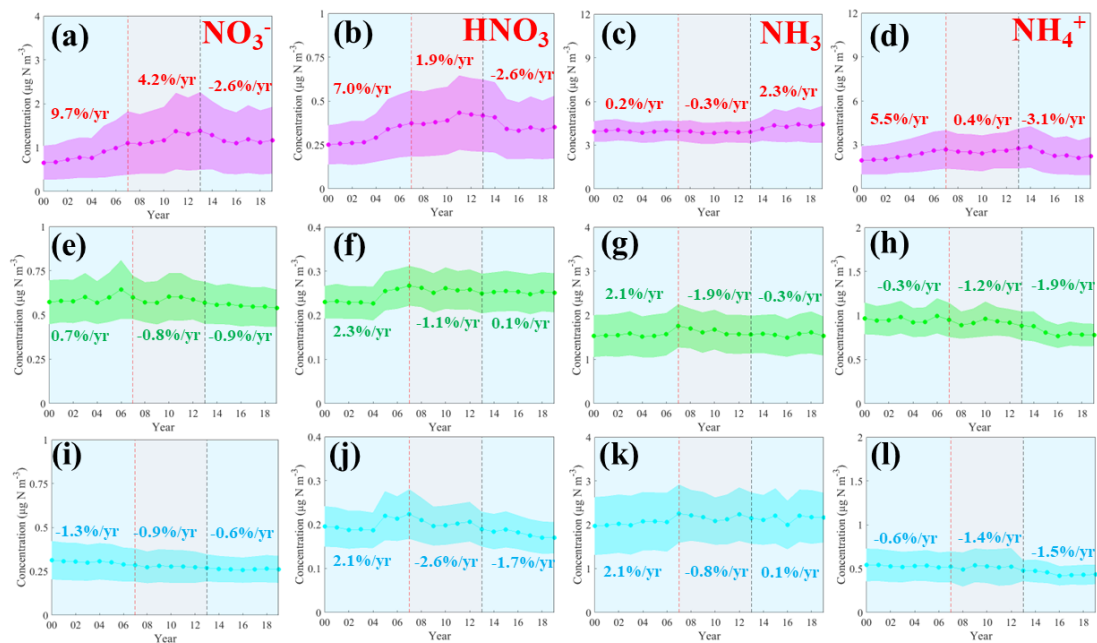


648 **Figure 5** The spatiotemporal variations of  $\text{NO}_3^-$ ,  $\text{HNO}_3$ ,  $\text{NH}_3$ , and  $\text{NH}_4^+$  concentrations in North  
649 America (a-d) (Unit:  $\mu\text{g N m}^{-3}$ ).



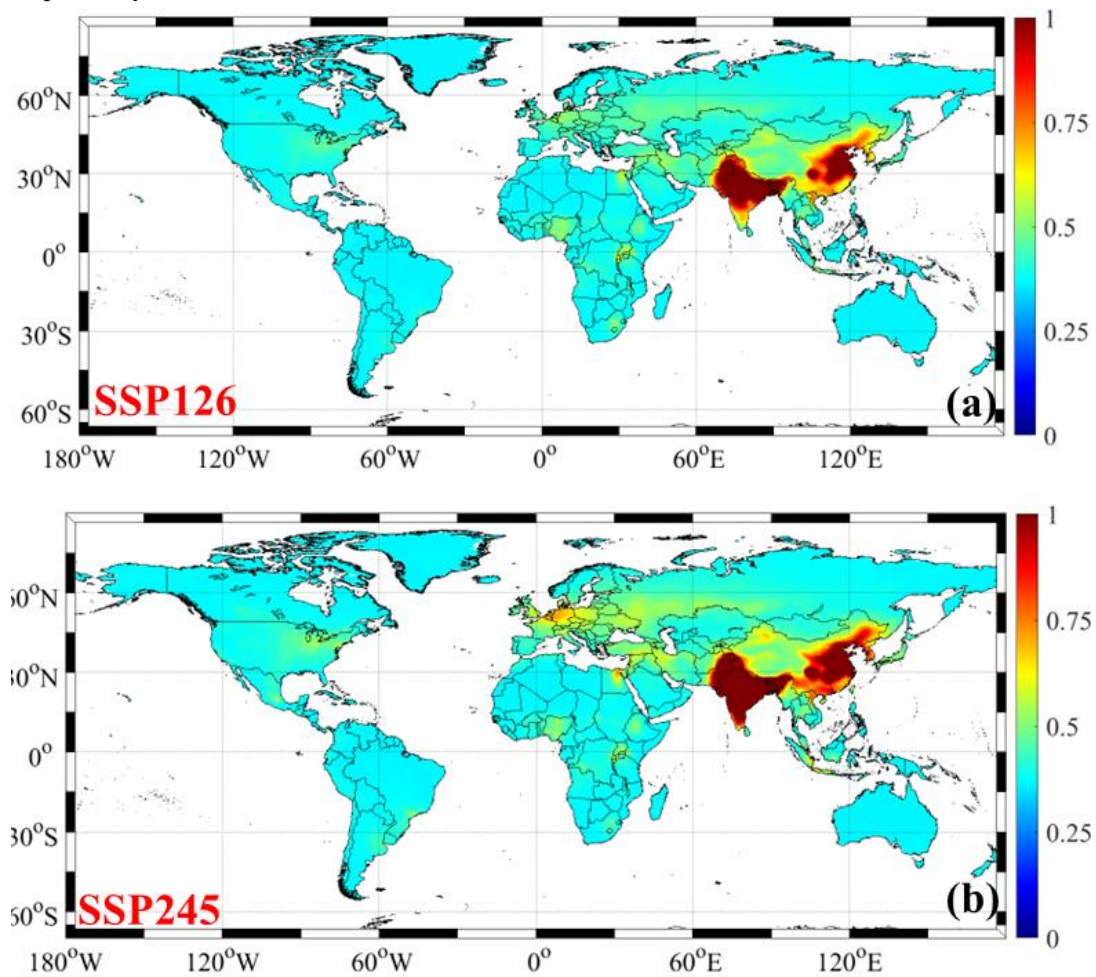
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651 **Figure 6** The long-term variations of  $\text{NO}_3^-$ ,  $\text{HNO}_3$ ,  $\text{NH}_3$ , and  $\text{NH}_4^+$  concentrations in China (pink),  
 652 Europe (green), and the United States (cyan) (Unit:  $\mu\text{g N m}^{-3}$ ).



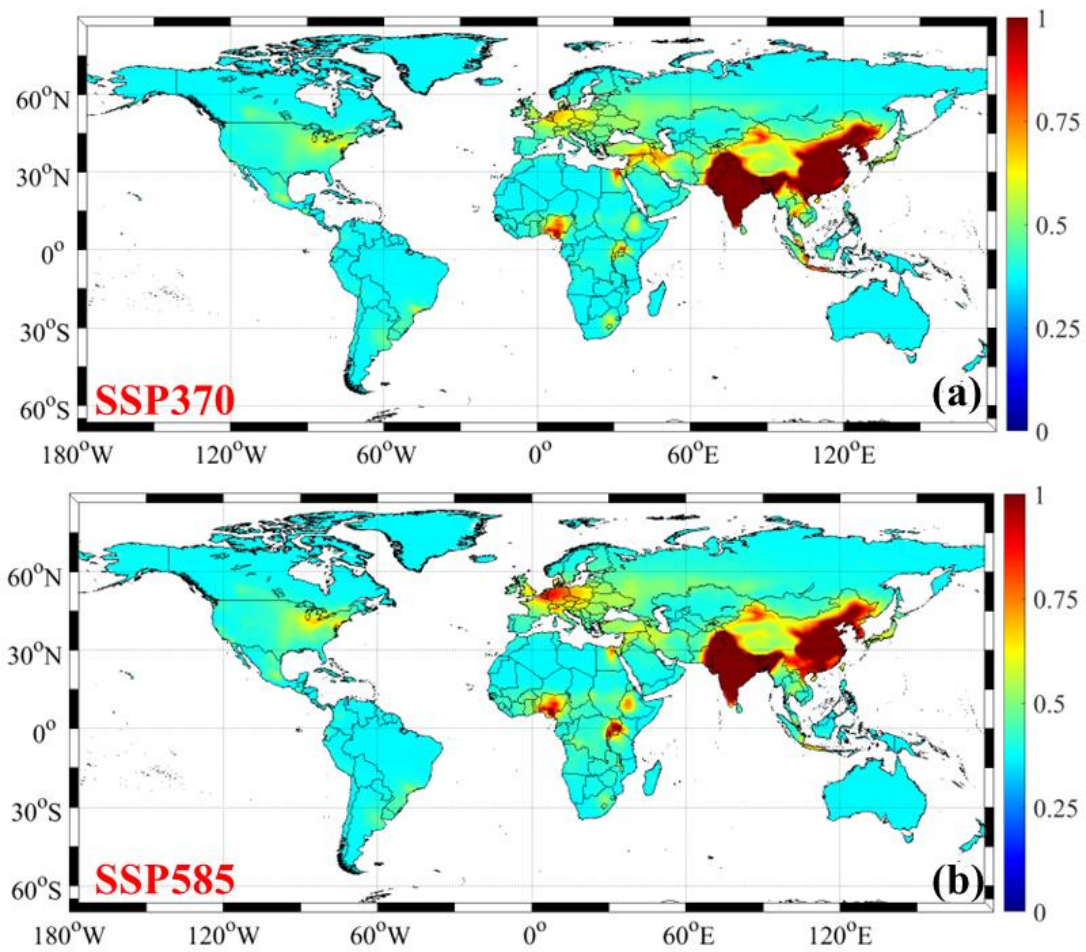
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654 **Figure 7** Spatial variations of projected global ambient concentrations of reactive nitrogen  
655 components under different climate change scenarios (Unit:  $\mu\text{g N m}^{-3}$ ). Panels (a-b) represent the  
656 annual mean concentrations of ambient  $\text{NO}_3^-$  under SSP1-2.6, SSP2-4.5 during 2021-2100,  
657 respectively.



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660 **Figure 8** Spatial variations of projected global ambient concentrations of reactive nitrogen  
661 components under different climate change scenarios (Unit:  $\mu\text{g N m}^{-3}$ ). Panels (a-b) represent the  
662 annual mean concentrations of ambient  $\text{NO}_3^-$  under SSP3-7.0, and SSP5-8.5 during 2021-2100,  
663 respectively.



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**Table 1** The temporal variations of ambient NO<sub>3</sub><sup>-</sup> (NO<sub>3</sub>-N) concentrations (average concentrations) in many countries during 2000-2100.

Scenario	NO <sub>3</sub> <sup>-</sup>	China	India	Europe	United States	Brazil	Argentina	Democratic Congo	West Africa Coast	Indonesia	South Korea
Historical	2000	0.66	0.95	0.57	0.31	0.29	0.21	0.56	0.30	0.31	1.06
	2005	0.91	1.26	0.60	0.30	0.31	0.22	0.56	0.30	0.34	1.38
	2010	1.17	1.53	0.60	0.28	0.33	0.23	0.58	0.31	0.32	1.43
	2013	1.39	1.63	0.57	0.27	0.30	0.24	0.58	0.31	0.33	1.57
	2015	1.16	1.34	0.56	0.26	0.32	0.22	0.58	0.32	0.45	1.88
	2019	1.18	1.46	0.54	0.26	0.32	0.23	0.61	0.32	0.37	1.87
SSP1-2.6	2020	1.16	1.23	0.41	0.27	0.25	0.18	0.55	0.28	0.34	1.81
	2040	0.82	1.12	0.28	0.10	0.21	0.14	0.62	0.32	0.28	1.46
	2060	0.69	1.01	0.15	0.05	0.18	0.13	0.43	0.25	0.24	0.69
	2080	0.41	0.89	0.12	0.05	0.17	0.12	0.32	0.17	0.20	0.39
	2100	0.33	0.65	0.10	0.06	0.11	0.08	0.21	0.11	0.16	0.23
SSP2-4.5	2020	1.19	1.43	0.44	0.24	0.26	0.19	0.52	0.29	0.37	1.85
	2040	1.09	1.35	0.43	0.16	0.22	0.16	0.57	0.31	0.35	1.80
	2060	0.89	1.22	0.35	0.11	0.20	0.15	0.51	0.27	0.32	1.25
	2080	0.63	1.06	0.29	0.07	0.17	0.12	0.42	0.23	0.23	0.68
	2100	0.41	0.95	0.24	0.05	0.14	0.10	0.36	0.19	0.20	0.38
SSP3-7.0	2020	1.25	1.59	0.53	0.33	0.31	0.22	0.64	0.34	0.42	1.95
	2040	1.36	1.50	0.47	0.24	0.26	0.19	0.61	0.33	0.47	1.89
	2060	1.18	1.35	0.42	0.19	0.22	0.16	0.56	0.30	0.41	1.56
	2080	0.96	1.15	0.36	0.16	0.18	0.13	0.54	0.29	0.35	1.35
	2100	0.75	1.08	0.33	0.12	0.15	0.11	0.51	0.27	0.30	1.24
SSP5-8.5	2020	1.21	1.50	0.53	0.28	0.28	0.20	0.57	0.31	0.37	1.91
	2040	1.28	1.42	0.49	0.23	0.25	0.18	0.60	0.32	0.39	1.85
	2060	1.05	1.30	0.47	0.24	0.20	0.15	0.55	0.29	0.35	1.44
	2080	0.86	1.10	0.44	0.25	0.15	0.11	0.50	0.27	0.29	1.26
	2100	0.58	1.02	0.31	0.20	0.13	0.09	0.46	0.25	0.25	1.05

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