- 1 Using observed urban NO_x sinks to constrain VOC reactivity and the ozone and radical
- 2 budget in the Seoul Metropolitan Area
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29 For ACP

Abstract

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Ozone (O₃) is an important secondary pollutant that impacts air quality and human health. Eastern Asia has high regional O₃ background due to the numerous sources and increasing and rapid industrial growth, which impacts the Seoul Metropolitan Area (SMA). However, SMA has also been experiencing increasing O₃ driven by decreasing NO_x emissions, highlighting the role of local, in-situ O₃ production on SMA. Here, comprehensive gas-phase measurements collected on the NASA DC-8 during the NIER/NASA Korea United States-Air Quality (KORUS-AQ) study are used to constrain the instantaneous O₃ production rate over the SMA. The observed NO_x oxidized products support the importance of non-measured peroxy nitrates (PNs) in the O₃ chemistry in SMA, as they accounted for ~49% of the total PNs. Using the total measured PNs (ΣPNs) and alkyl and multifunctional nitrates (ΣANs) , unmeasured volatile organic compound (VOC) reactivity (R(VOC)) is constrained and found to range from $1.4 - 2.1 \text{ s}^{-1}$. Combining the observationally constrained R(VOC) with the other measurements on the DC-8, the instantaneous net O₃ production rate, which is as high as ~10 ppbv hr⁻¹, along with the important sinks of O₃ and radical chemistry, are constrained. This analysis shows that ΣPNs play an important role in both the sinks of O_3 and radical chemistry. Since ΣPNs are assumed to be in steady-state, the results here highlight the role ΣPNs play in urban environments in altering net O_3 production, but ΣPNs can potentially lead to increased net O₃ production downwind due to their short lifetime (~1 hr). The results provide guidance for future measurements to identify the missing R(VOCs) and Σ PNs production.

Short Summary

- Ozone (O₃) is a pollutant formed from the reactions of gases emitted from various sources. In urban areas, the density of human activities can increase the O₃ formation rate (P(O₃)); thus, impact
- air quality and health. Observations collected over Seoul, South Korea, are used to constrain P(O₃).
- A high local P(O₃) was found; however, local P(O₃) was partly reduced due to compounds typically
- 55 ignored. These observations also provide constraints for unmeasured compounds that will impact
- 56 $P(O_3)$.

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1. Introduction

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Representing global and urban tropospheric ozone (O₃) in chemical transport models (CTMs) is still challenging due to uncertainty in physical and chemical processes that control the O₃ budget (Archibald et al., 2020). One area of uncertainty is underestimated urban volatile organic compounds (VOCs) emissions (von Schneidemesser et al., 2023), which arise from a large number of sources, including some that are very hard to quantify (e.g., cooking and chemical products) (e.g., McDonald et al., 2018; Simpson et al., 2020). Intensive research is also ongoing as to why O₃ is increasing in recent years in urban areas, even with reductions in combustion emissions (Colombi et al., 2023; e.g., Lyu et al., 2017). This O₃ impacts the large populations in urban areas with harmful health effects, including premature mortality (e.g., Cohen et al., 2017). Tropospheric O_3 production is driven by the catalytic cycling of nitrogen oxides (NO_x = NO + NO₂) fueled by the photoxidation of VOCs, both of which can come from anthropogenic emissions. The chemistry producing O_3 is described in R1 – R6 in Table 1. During daylight hours, VOCs are oxidized by OH (or undergo photolysis) to form an organic peroxy radical (RO₂) in R1a (R1b). If the RO₂ then proceeds through R2a, at least two O₃ molecules are produced. The first O_3 molecule is formed by the photolysis of NO_2 and the reaction of $O(^3P)$ with oxygen (R3 – R4). The second O₃ molecule is formed through the reaction of the alkoxy radical (RO) with oxygen to form the hydroperoxyl radical (HO₂) (R5), which goes on to react with NO to produce

 NO_2 (R6) and the subsequent reactions described above (R3 – R4). However, some fraction of the

time, depending on the number of carbons and functional group (e.g., Espada and Shepson, 2005;

Perring et al., 2013; Yeh and Ziemann, 2014), alkyl or multifunctional nitrates (ANs \equiv RONO₂)

are formed (R2b). The fraction of reactions to form ANs is described by the branching ratio, α .

Reaction R2b has been shown to impact O₃ production, depending on the types of VOC emitted,

by reducing the fraction of NO_2 that photolyzes to form O_3 in source regions (R3 – R4) (Farmer et al., 2011). As α is a function of the individual VOC's carbon backbone and functional group (e.g., Perring et al., 2013), any uncertainty related to primary VOC emissions and secondary chemistry will directly impact the ability to describe urban O_3 production.

One important subclass of VOCs are aldehydes (RCHO), which can either be directly emitted or produced via photooxidation of VOCs (de Gouw et al., 2018; Mellouki et al., 2015; Wang et al., 2022; Yuan et al., 2012). The photooxidation of the aldehyde (R7) in the presence of NO_x can either form acyl peroxy nitrates (R8, PNs = R(O)O₂NO₂) or an organic peroxy radical (RO₂) (R9). The competition between R8 to form PNs versus R9 to form RO₂ depends on the NO-to-NO₂ ratio (Nihill et al., 2021). Further, R8 is in thermodynamic equilibrium due to the weak bond strength between the acyl peroxy radical (R(O)O₂) and NO₂. Thus, formation of PNs pose only a temporary loss of NO₂. Finally, it has been observed that aldehydes with longer carbon backbones (e.g., C8s and C9s) from various anthropogenic activities, such as cooking (Coggon et al., 2024; Rao et al., 2010), may have mixing ratios as high as aldehydes typically quantified in field experiments (acetaldehyde and propaldehyde). However, there is larger uncertainty associated with these higher aldehydes in their fate to produce both PNs and ANs (e.g., Hurst Bowman et al., 2003). Missing both these emissions and subsequent chemistry would impact estimates of urban O₃ chemistry.

The fraction of RO_2 forming ANs in R2b and the fraction of $R(O)O_2$ forming PNs in R8 alter the instantaneous O_3 production ($P(O_3)$) by removing NO_2 and/or the radical species. This is further shown in Figure S1, where an analytical equation to describe R1 – R6 (Farmer et al., 2011), is used to explore how changes in the VOC reactivity (R(VOC)), radical production ($P(HO_x)$), and ANs production and branching ratio, α (R2b), impact the instantaneous $P(O_3)$ (see Sect. S1 for the

analytical equation and description). Any changes in $P(HO_x)$, R(VOC), and/or α will impact both the instantaneous $P(O_3)$ as well as the NO_x mixing ratio corresponding to the maximum $P(O_3)$. As these parameters are generally interconnected, investigating all three is important to understand the sources and control of instantaneous $P(O_3)$. Further, R7 - R9 are not included in this traditional description of the analytical equation, as it is assumed PNs are in steady-state (Farmer et al., 2011). Thus, if PNs are not in steady-state, their role in altering $P(O_3)$ may be underestimated.

Increasing surface O₃ is a concern throughout East Asia, including South Korea (Colombi et al., 2023; Gaudel et al., 2018; Kim et al., 2021; Yeo and Kim, 2021). The emissions associated with industry and other anthropogenic activities and the associated photochemistry have impacted regional air quality, leading to high O₃ backgrounds that can impact a country's ability to achieve reduced O₃ exposure for new air quality standards (e.g., Colombi et al., 2023). However, local emissions and photochemistry still play an important role. For example, during the Korea-United States Air Quality (KORUS-AQ) campaign, it was observed between morning and afternoon in the Seoul Metropolitan Area (SMA), O₃ increased by ~20 parts per billion by volume (ppbv) over a background concentration of over 75 ppbv (Crawford et al., 2021). Thus, an understanding of the variables highlighted in Figure S1 are necessary to control both local and regional P(O₃).

One tool typically used to understand the role of regional O₃ and transported O₃ on local O₃ and impacts of local emission controls on O₃ are CTMs. As shown in Park et al. (2021), for the SMA, CTMs typically underestimate the observed O₃ and formaldehyde. While the low O₃ could be partially related to underestimated transport (e.g., Seo et al., 2018) or resolution of the CTM (e.g., Jo et al., 2023; Park et al., 2021), the low bias also observed for modeled formaldehyde indicates overall (a) too little VOCs and thus too low R(VOC) (Brune et al., 2022; H. Kim et al., 2022), (b) missing photochemical products from missing VOCs, including oxygenated VOCs

(OVOCs) that contribute to $P(HO_x)$ (Brune et al., 2022; H. Kim et al., 2022; Lee et al., 2022; Wang et al., 2022), and (c) likely missing PNs and ANs from the underestimated VOCs related to the underestimated R(VOC) (Lee et al., 2022; Park et al., 2021). Missing (a) – (c) will bias the instantaneous $P(O_3)$ (Figure S1), impacting the ability to investigate what policies should be implemented to reduce O_3 .

To better understand what controls the instantaneous $P(O_3)$ over SMA, observations collected on the NASA DC-8 during KORUS-AQ are used to constrain the three variables highlighted in Figure S1—R(VOC), HO_x production and loss, and ANs and PNs production. Observational constraints on these three parameters provide a means to investigate the instantaneous $P(O_3)$ over SMA and the major classes of contributors to O_3 and HO_x production and loss. These results are discussed and placed into the context of improving our knowledge about O_3 production in an urban environment.

2. Methods and Data Description

2.1 KORUS-AQ and DC-8 Descriptions

The KORUS-AQ campaign was a multi-national project that was conducted in May – June, 2016, led by South Korea's National Institute of Environmental Research (NIER) and United States National Aeronautics and Space Administration (NASA). The project was conducted in South Korea and the surrounding seas with numerous airborne platforms, research vessels, and ground sites (Crawford et al., 2021). The study here focuses on the observations collected on the NASA DC-8.

The instrument payload, flights, and observations have been described in other studies (Brune et al., 2022; Crawford et al., 2021; Lee et al., 2022; Schroeder et al., 2020). Briefly, the

DC-8 was stationed at Osan Air Force Base, Pyeongtaek, South Korea, which is approximately 60 km south of Seoul. A total of 20 research flights were conducted with the DC-8. Part of each research flights included a stereo-route in the SMA in the morning (~09:00 local time), midday (~12:00 local time), and afternoon (~15:00 local time), which included a missed approach over Seoul Air Base (< 15 km from Seoul city center) and a fly-over of the Olympic Park and Taehwa Forest Research sites (Figure 1). A total of 55 descents over Olympic Park and 53 spirals over Taehwa Forest Research site were conducted (Crawford et al., 2021). Only observations from the DC-8 after 11:00 local time are used here to ensure that the boundary layer has grown and stabilized and to minimize any influence from residual layer mixing into the boundary layer and/or titration of O₃ by NO (R10). We analyze data collected below 2 km and between 127.10 – 127.67°E and 37.22 – 37.69°N to focus on the boundary layer in the SMA without influence from industrial emissions along the western South Korean coast (Crawford et al., 2021).

During KORUS-AQ, four different meteorological periods, as described by Peterson et al. (2019), impacted the region. These periods included a Dynamic period from 1 – 16 May, where there were a series of frontal passages; a Stagnant period from 17 – 22 May, where it was dry, clear, and stagnant; Transport/Haze period from 25 – 31 May, where long-range transport and hazy conditions with high humidity and cloud cover prevailed; and, a Blocking period from 1 – 7 June, where a high pressure ridge is located to an area north of lower pressure, which can preclude significant changes in synoptic meteorology and results in occasional stagnant conditions/minimal pollution transport (Peterson et al., 2019). However, as discussed in Sect. 3.2, conditions did not impact the general trends and chemistry and thus the whole campaign has been analyzed together.

The observations used for the analysis are shown in Table 2, along with the associated references. The 1-min merged data from the DC-8 is used here (KORUS-AQ Science Team, 2023).

For data missing due to frequency of measurements (e.g., VOCs from WAS), data was filled in a similar approach as Schroeder et al. (2020), in that VOCs with missing data were filled by the linear relationship of that VOC with VOCs measured more frequently. This step was necessary for the observations used in the diel steady-state calculations described in Sect. 2.2. Note, the TD-LIF NO₂ (see Table 2) was used throughout this study and discussed in Sect. S2 and Figure S2 – S4 as it generally agreed better with steady-state calculated NO₂-to-NO ratios and steady-state NO₂ than the chemiluminescence NO₂. Further, though PBzN was measured by GT-CIMS (Table 2), it is not compared with calculated PBzN from F0AM (Sect. 2.2) as it may be underestimated due to possible inlet losses, as discussed in Zheng et al. (2011).

2.2 F0AM Box Model Diel Steady-State Calculations for Missing Reactivity and Peroxynitrate Budget Analysis

We use the F0AM box model (Wolfe et al., 2016) with chemistry from the MCMv3.3.1 (Jenkin et al., 2015) to simulate production of PNs and formaldehyde using 1-min merged data, as described in Sect. 2.1. As in Schroeder et al. (2020), we simulate each 1-min merged aircraft observation through the full diurnal solar cycle (i.e., diel steady-state) until the diurnal cycle for each unconstrained species reaches convergence within 1%. These unconstrained species, such as formaldehyde, NO₂, and OH, are then evaluated to ensure consistency between F0AM model and aircraft observations. We constrain concentrations of NO, O₃, H₂O₂, HNO₃, CO, CH₄, H₂, PAN, PPN, and all measured or estimated VOCs given in Table 2 and Table S1 to calculate HO₂, all organic peroxy and acyl peroxy radicals, and unmeasured PNs. To calculate the PAN and PNs budget, we allow the model to freely calculate NO₂, formaldehyde, acetaldehyde, and all PNs, including PAN and PPN, for when calculating the budget of PNs. We use a dilution constant of 12

hours, according to Brune et al. (2022). Model evaluation is discussed in Sect. 3.4. The contribution of individual VOCs to PAN was calculated by reducing precursor VOCs by 20% and multiplying the resulting impact on the peroxy acetyl radical (CH₃C(O)O₂) by 5. Other acyl peroxy nitrates (higher PNs) are lumped into categories based on their primary precursor species from Table S2, species currently typically measured (e.g., PPN) or contributes a large fraction of the total higher PNs budget (greater than >2%; e.g., PHAN and MPAN).

Note, the reason PAN and PPN were constrained were due to uncertainties in the thermal lifetime, temperature history, and dilution rate used in F0AM, which had larger impacts on the CH₃C(O)O₂ and PAN than on other unconstrained compounds (e.g., OH and formaldehyde and not shown; Brune et al. (2022)).Part of this larger impact is due to CH₃C(O)O₂ being one of the most abundant radicals and one of the final radical products in the oxidation of numerous compounds (e.g., Jenkin et al., 2015). We do not expect these uncertainties to impact the higher PNs as (a) there are less precursors to form them compared to PAN and (b) they are expected to have higher thermal stability compared to PAN due to longer carbon backbone (Kabir et al., 2014).

2.3 Calculation of Instantaneous Ozone and HO_x Production and Loss

An experimental budget for the production and loss of O_x ($O_x = O_3 + NO_2$) and HO_x ($HO_x = OH + HO_2 + RO_2^- + R(O)O_2^-$) is described here. NO_2 and O_3 are combined to reduce any potential impact from titration via O_3 reaction with NO (R10). The budget analysis includes field-measured quantities (mixing ratios and photolysis rates, Table 2), results from F0AM (Sect. 2.2), estimated missing R(VOC) (Sect. 3.2) and published kinetic rate constants (see Table 1 for references). The rate of production or destruction is calculated with the following equations (Eq. 1 – 7) below. Note, these equations differ from Schroeder et al. (2020) in that (a) ANs and PNs chemistry are explicitly

included and (b) the reaction of O₃ with alkenes is excluded as this reaction contributed a minor

220 loss to O_3 (< 1%).

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$$P_{O_x} = \sum_{i} (1 - \alpha_{ieff}) k_{RO_{2,i} + NO} [RO_{2,i}] [NO] + k_{HO_2 + NO} [HO_2] [NO]$$
 (1)

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$$L_{O_x} = k_{NO_2 + OH}[NO_2][OH] + k_{O_3 + OH}[O_3][OH] + f \times j_{O_1}[O_3] +$$

223
$$k_{HO_2+O_3}[HO_2][O_3] + net(PNs)$$
 (2)

net(PNs)=
$$\beta k_{R(O)O_2+NO_2}[R(O)O_2][NO_2] - (1-\beta)k_{decomposition}[PNs]$$
 (3)

225
$$\beta = \frac{k_{RC(O)O_2 + NO_2}[NO_2]}{k_{RC(O)O_2 + NO_2}[NO_2] + k_{RC(O)O_2 + NO}[NO]}$$
(4)

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$$P(HO_x) = 2f \times j_{O^1D}[O_3] + 2j_{H_2O_2}[H_2O_2] + 2j_{CH_2O \to H + HCO}[CH_2O] + 2j_{CHOCHO}[CHOCHO] +$$

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$$2j_{CH_3OOH}[CH_3OOH] + 2j_{CH_3CHO}[CH_3CHO] + 2j_{CH_3C(O)CH_3}[CH_3C(O)CH_3] +$$

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$$2j_{CH_3CH_2C(O)CH_3}[CH_3CH_2C(O)CH_2]$$
 (5)

229
$$L(HO_x) = k_{NO_2+OH}[NO_2][OH] + \sum_{i} \alpha_{eff} k_{RO_2,i} + NO[RO_{2,i}][NO] +$$

$$2k_{HO_2+HO_2}[HO_2][HO_2] + 2k_{RO_2+RO_2}[RO_2][RO_2] + 2k_{HO_2+RO_2}[HO_2][RO_2] + net(PNs) \qquad (6)$$

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$$[RO_{2}^{\cdot}] = \frac{\sum_{i} k_{OH+VOC,i} [VOC_{i}] [OH]}{(1-\alpha_{eff}) k_{RO_{2}+NO} [NO] + k_{RO_{2}+HO_{2}} [HO_{2}]}$$
 (7)

Here, k is the rate constant for compound, i, with the associated compound listed, α_{eff} is the effective branching ratio for R2a and R2b for the observations (Sect. 3.2), f is the fraction that O^1D that reacts with water to form OH versus reacting with a third body molecule to form O^3P , β is the fraction the R(O)O₂ that reacts with NO₂ versus NO, and j is the measured photolysis frequency (Table 2). In Eq. 5, only directly values directly measured on the DC-8 during KORUS-AQ are included. As discussed in Wang et al. (2022) and Sect. 4.3,, this is most likely an underestimations of P(HO_x). Note, R(O)O₂ is not included in Eq. 7 as (a) it is assumed the initial production of R(O)O₂ is captured with the reaction of OH with VOC and (b) R(O)O₂ accounts for a small

fraction of the total RO_2 (< 10%). Not including $R(O)O_2$ in Eq. 7 may lead to a small underestimation of total RO_2 . Finally, HO_2 calculated from F0AM, rather than aircraft measurements (Crawford et al., 2021), is used in the equations to determine the O_x and HO_x budget (see Sect. S3, Figure S6).

3. Observational constraints on NO_x organic oxidation chemistry

In the Sect. 3.1, the detailed observations from the DC-8 during KORUS-AQ provided measurements that allow us to test our understanding of NO_x oxidation into total NO_z (NO_z = higher NO_x oxides, including ΣPNs , ΣANs , HNO_3 and particulate nitrate, pNO_3), which is needed for the remainder of the analysis. Sect. 3.2 to 3.4 will focus on the organic NO_z chemistry. This is due to the chemistry and dynamics impacting the total inorganic nitrate chemistry that has been discussed recently (Jordan et al., 2020; Travis et al., 2022).

3.1 NO_x and its oxidation products

The average NO_x mixing ratios observed by the NASA DC-8 in the SMA below 2 km after 11:00 local time is shown in Figure 1. As NO_x is mainly emitted from anthropogenic activities, such as combustion emissions, in an urban environment, the largest NO_x mixing ratios are observed between Olympic Park and the missed approach, as this area included downtown SMA. The missed approach included low level sampling at a military airport, which may have contributed to the NO_x mixing ratios along with the activities throughout the SMA. As the DC-8 flies from the missed approach toward Taehwa Research Site, the NO_x mixing ratios decreases. The combination of reduced emissions, chemical reactions, and dilution and mixing reduces the

 NO_x mixing ratios away from the city. An understanding of these processes is important for urban $P(O_x)$.

On the DC-8, there were multiple measurements of various speciated and total family contribution towards NO_z (Table 2). The comparison of the speciated and measured NO_z is investigated in Figure 2 for observations over SMA. When only speciated PNs (GT) and ANs (CIT + WAS) and gas-phase nitrate (HNO₃) are compared to the NO_z (NO_y (NCAR) – (NO (NCAR) + NO₂ (TD-LIF)), only 46% of the NO_z can be explained. This is not completely unexpected, as multiple studies have indicated that the speciated ANs measurements are typically lower than the total ANs measurements (Fisher et al., 2016; Perring et al., 2010). Further, pNO₃ has been found to be important for total nitrate budget in the SMA (e.g., Travis et al., 2022). Chemiluminescence measurements of gas-phase NO_y have been found to efficiently measure pNO₃, depending on the sensitivity to pNO₃ enhancements or exclusions (Bourgeois et al., 2022); thus, it is expected that missing ANs and pNO₃ are necessary to close the NO_z budget. Adding the measured pNO₃ to the speciated PNs (GT) and ANs (CIT + WAS) and gas-phase nitric acid, 81% of NO_z can be explained. This barely overlaps the combined uncertainty of the measurements (~26%). Total PNs and ANs, measured by TD-LIF, are needed to close of the total NO_z budget.

The breakdown of the NO_z budget over the SMA as the airmasses photochemically ages (decreasing NO_x contribution to total NO_y) is shown in Figure 2b. During KORUS-AQ, ~56% of NO_z was inorganic (gas- and particle-phase nitrate), ranging from 52% to 62%; the remaining NO_z was organic (PNs and ANs). Approximately 74% of the total ANs were not speciated (range 73% to 76%). Speciated PNs species, such as PAN (peroxy acetyl nitrate), account for a mean 51% of the total PNs (range 47 to 59%), much lower than typically observed in prior studies (e.g., Wooldridge et al., 2010). In these prior studies, the speciated PN species (typically PAN + PPN

(peroxy propionyl nitrate)) accounted for 90 – 100% of the ΣPNs, except for some select cases attributed to poor inlet design (Wooldridge et al., 2010). PAN accounted for the majority of the speciated PNs, with the remaining speciated PNs (PPN + PBzN (peroxy benzoyl nitrate) + APAN (peroxy acryloyl nitrate)) accounting for ~1%. However, during KORUS-AQ, Lee et al. (2022) observed that PAN contributed only 60% of calculated total PNs in industrial plumes near the SMA. Thus, the VOC emissions in and near SMA potentially lead to PNs typically not directly measured; this is explored more in Sect. 3.4

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As NO_x decreases from ~30 ppby to 4 ppby, the contribution of organic NO_z increases (Figure 2b). At about 4 ppbv, the contribution of organic NO_z starts to decrease. Further, the contribution of the different organic NO_z species changes. For example, from ~30 ppbv to 4 ppbv, the un-speciated ΣPNs contributes the majority of the organic NO_z budget (~39%). Below ~4 ppbv, the contribution of un-speciated ΣPNs decreases and the PAN contribution increases. The change in contribution of PNs is due to changes in the PN precursors (e.g., combination short-lived precursors oxidizing to $CH_3C(O)O_2$ and thermal decomposition of the higher PNs (higher PNs = $\Sigma PNs - PAN$)). On the other hand, the contribution of un-speciated ΣANs remains relatively constant with NO_x (~6% of total NO_z). However, the type of ANs is most likely changing with NO_x due to the lifetime of the ANs precursors and/or the lifetime of ANs. Less is known about the lifetime of ANs derived from anthropogenically emitted VOCs compared to those from biogenic VOCs (González-Sánchez et al., 2023; Picquet-Varrault et al., 2020; Zare et al., 2018). On average unknown ANs and PNs account for ~24% of the observed NO_z on average. The differences in the binned mean value for each species is greater than the uncertainty associated with its measurements (maximum uncertainty 30%) and greater than the standard error of the mean, indicating that all the percent differences shown here are real.

3.2 Meteorological impact on NO_x oxidation

As discussed in Sect. 2.1 and various prior studies, four different meteorological conditions impacted the observations during KORUS-AQ (Peterson et al., 2019). The impact of the meteorological conditions on NO_x oxidation was investigated by plotting two metrics of NO_x oxidation— O_x versus Σ ANs and Σ PNs versus formaldehyde (Figure 3). The implications of both plots are further discussed in Sect. 3.3 and 3.4, respectively. Briefly, O_x versus Σ ANs and Σ PNs versus formaldehyde are competitive products from the reaction of RO_2 or $R(O)O_2$ with NO_x (R2a versus R2b or R8 versus R9). The different meteorological periods corresponded to differences in temperatures and amount of photolysis due to cloud cover (Peterson et al., 2019). Thus, these different periods may impact gas-phase chemistry and/or VOC emissions. However, as demonstrated in Figure 3, there are minimal systematic differences in the trends observed for the two NO_x oxidation products as there is no systematic shift in the trends or scatter observed in Figure 3. This suggests that the data does not have to be separated by meteorological conditions.

3.3 Production of ANs to constrain R(VOC)

Observations of un-speciated ANs and PNs imply missing VOCs that impact O_3 chemistry. The relationship of ANs to O_x can provide a method to investigate this source. This relationship provides an estimate of the effective branching ratio, α , for the observed VOC mix (Perring et al., 2013 and references therein). The value of this relationship stems from the reactions discussed above (R1 – R6) in that upon the oxidation of VOCs, some fraction of the time, RO_2 reacts with NO to form an AN molecule and the remainder of the time the reaction goes to form O_3 . This is expressed with the following equations:

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$$P_{\Sigma ANS} = \sum \alpha_i k_{OH+VOCi} [OH] [VOC_i]$$
 (8)

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$$P(O_x) = \sum_i \gamma_i (1 - \alpha_i) k_{OH + VOC_i} [OH] [VOC_i]$$
 (9)

Here, α is the effective branching ratio in the reaction of RO₂ with NO to form ANs versus RO (R2), k is the OH rate constant with VOC, i, and γ is the number of O₃ molecules formed per oxidation of VOC, i. The reactivity weighted γ is calculated for the observed and F0AM calculated species with Eq. 10, where γ for each compound is taken from MCM (Jenkin et al., 2015) and accounts for potential for difference number of O₃ molecules produced per channel per oxidation (e.g., xylene produces two O₃ molecules 60% of the time and one O₃ molecule 40% of the time). All the terms were defined for Eq. 8 – 9.

$$\gamma_{eff} = \frac{\sum_{i} \gamma_{i} k_{OH+VOC_{i}} [OH][VOC_{i}]}{\sum_{i} k_{OH+VOC_{i}} [OH][VOC_{i}]}$$
(10)

The reactivity weighted γ is found to be, on average, 1.53, which is lower than the value of 2 typically assumed in prior studies (e.g., Perring et al., 2013). This lower reactivity weighted γ is due to the role of CO ($\gamma = 1$) and CH₂O ($\gamma = 1$) to the total reactivity. After the boundary layer height has stabilized (e.g., after 11:00 am LT used here) and is near enough (e.g., less than 1 day aging) to the VOC source to ignore deposition and entrainment, Eq. 8 and 9 can be combined to approximate the change in O_x per molecule Σ AN formed:

$$\frac{\Delta O_x}{\Delta \Sigma A N s} \approx \frac{P_{O_x}}{P \Sigma A N s} \approx \frac{1.53(1-\alpha)}{\alpha} \tag{11}$$

For this equation to be valid, α needs to be relatively small ($\alpha << 1$), which is true for VOCs, as maximum α for the conditions of KORUS-AQ is expected to be 0.35 (Orlando and Tyndall, 2012; Perring et al., 2013; Yeh and Ziemann, 2014). Note, though Eq. 11 can be used at short photochemical ages due to minimal impact from physical loss processes, chemical loss processes may impact the assumptions in Eq. 11 and are discussed in more detail below.

Over the SMA during KORUS-AQ, the slope between O_x and Σ ANs was observed to be 40.5 ± 1.8 (Figure 3a), with an $R^2=0.60$. Using Eq. 11, this translates to an effective branching ratio (α_{eff}), of 0.036. For other urban locations around the world, this slope has ranged from 13 – 47 (Farmer et al., 2011; Kenagy et al., 2020; Perring et al., 2010; Rosen et al., 2004), leading to an effective α between 0.04 and 0.15, assuming a γ of 2 instead of the calculated γ used here. Thus, the α_{eff} observed over SMA during KORUS-AQ is similar to other urban locations (Houston = 0.05 (Rosen et al., 2004) and South Korea = 0.05 (Kenagy et al., 2021)) but much lower than observed for Mexico City = 0.07 – 0.12 (Farmer et al., 2011; Perring et al., 2010) and Denver = 0.16 (Kenagy et al., 2020). This suggests that VOCs with low α dominate the total R(VOC) and production of ANs in SMA. The VOCs in SMA that dominate R(VOCs), including OVOCs, alkenes, and aromatics (Schroeder et al., 2020; Simpson et al., 2020), generally have lower α (Orlando and Tyndall, 2012; Perring et al., 2013 and references therein).

We use the observed VOCs (Table 2) and estimated secondary products from F0AM to calculate α_{eff} from this mixture to compare to the calculated α_{eff} of 0.036 derived from the slope of O_x versus ΣANs in Figure 3a, as shown in Figure 4. To derive α_{eff} , Eq. 12 was used, where all the terms are the same as Eq. 8 – 9.

$$\alpha_{eff} = \frac{\sum_{i} \alpha_{i} k_{OH+VOC_{i}} [OH] [VOC_{i}]}{\sum_{i} \gamma_{i} k_{OH+VOC_{i}} [OH] [VOC_{i}]}$$
(12)

The R(VOC) calculated from the observed VOCs and from the intermediates produced by the F0AM model, described in Sect. 2.2, are shown in Figure 4a, and the reactivity weighted α for the observations is shown in Figure 4b. As has been observed in other urban environments (e.g., Hansen et al., 2021; Whalley et al., 2016; Whalley et al., 2021; Yang et al., 2022), measured OVOCs contribute the most to the calculated R(VOC) for all NO_x mixing ratios (32 – 48%). The unmeasured OVOCs (F0AM species) contributed 17 – 28% of the calculated reactivity. The F0AM

species reactivity ranged from $0.45 - 1.78 \text{ s}^{-1}$, which is a similar increase in total OH reactivity observed by Brune et al. (2022) over South Korea. At higher NO_x mixing ratios, primary, more reactive VOCs (e.g., alkanes, alkenes, aromatics) contribute an important fraction (> 25%) of the R(VOC). As there are interferences in the total OH reactivity measurement at high NO_x (Brune et al., 2022), we are unable to determine the extent to which the observed and modeled reactivity captures total OH reactivity in the SMA above a NO_x value of approximately 4 ppbv. At lower NO_x mixing ratios, ~33% of the R(VOC) is missing (calculated R(VOC), including F0AM species, ~3.0 s⁻¹ and measured R(VOC) from Penn State—see Table 2—is 4.5 s⁻¹).

Numerous other urban studies have observed unmeasured OH reactivity, which is assumed to be unmeasured R(VOC), as the inorganic OH reactivity is typically well covered by measurements. Here, we are defining unmeasured R(VOC) as the reactivity not represented by measurements on the DC-8 or by F0AM predicted species and reactivity. This unmeasured R(VOC) has ranged from $\sim 3~s^{-1}$ to $\sim 10~s^{-1}$ (e.g., Brune et al., 2022; Hansen et al., 2021; Kim et al., 2016; Ma et al., 2022; Tan et al., 2019; Whalley et al., 2016; Whalley et al., 2021). Over the SMA, the difference between measured and calculated R(VOC) was $\sim 1.5~s^{-1}$ at low NO_x and unknown at high NO_x mixing ratios. The lower difference may be related to the comparison occurring for observations at low NO_x, when the very reactive material has either reacted into compounds measured on the DC-8 (e.g., formaldehyde, acetaldehyde, etc.), diluted to low enough concentrations to be negligible for R(VOC), or undergone deposition or partitioning to the particle-phase.

At higher NO_x mixing ratios, which is more representative of fresh emissions, these more reactive compounds typically not measured are expected to lead to a higher difference between the calculated and observed R(VOC). Prior studies with more comprehensive measurements found

these more reactive compounds and their secondary products contributed an important fraction towards the R(VOC) (e.g., Whalley et al., 2016). Thus, to determine if these unmeasured VOCs potentially contribute to the R(VOC), and thus $P(O_x)$, in SMA, another means to constrain their contributions is necessary. One potential means to constrain the total R(VOC) is by using the observed Σ ANs and O_x and assuming the observations are from the instantaneous production of both species (e.g., the assumption used for Figure 3a).

To estimate the unmeasured R(VOC), Eq. 13 is used without cancelling out terms and expanded into the measured and unmeasured R(VOC) and α :

$$\frac{\Delta O_x}{\Delta \sum ANs} = \frac{\gamma RVOC_m[OH] + \gamma RVOC_u[OH] - \gamma \alpha_m RVOC_m[OH] - \gamma \alpha_u RVOC_u[OH]}{\alpha_m RVOC_m[OH] + \alpha_u RVOC_u[OH]}$$
(13)

Here, $\frac{\Delta O_x}{\Delta \Sigma ANs}$ is the slope from Figure 3a, γ is the number of O₃ molecules formed per oxidation of VOC, which is 1.53 for this study, R(VOC) is the VOC reactivity, which is its OH oxidation rate constant and its concentration (k×[VOC]) in units s⁻¹, α is the branching ratio for R2 (Table 1), and m and u correspond to "measured" (measured VOCs on DC-8 along with secondary species predicted by F0AM) and "unmeasured" (unmeasured VOCs that are not represented by DC-8 observations and not predicted by F0AM) RVOC and α . The rate constants for the measured VOCs are listed in Table 1, the reactivity for F0AM is taken directly from F0AM, and α is either from MCM (Jenkin et al., 2015) or Perring et al. (2013) for observations or assumed to be 0.05 for F0AM secondary products. The equation is rearranged and solved for $RVOC_u$, using different values of α_u (e.g., 0.00 – 0.30, values typical α).

As discussed in Sect. S3 in the Supp. Information, there are numerous assumptions and potential sources of uncertainty in the simplified version of Eq. 13. A thorough analysis and discussion of these assumptions are discussed in Sect. S3. The potentially most important assumption is that chemical loss is negligible in solving Eq. 13. However, due to the expected

relatively short lifetime of \sum ANs, the chemical loss of both O_x and ANs nearly cancel each other, leading to similar results in considering or neglecting these loss terms in Eq. 13. Further, as \sum ANs chemical loss has uncertainty, especially for ANs produced from anthropogenic VOC oxidation, the use of Eq. 13 reduces some of these uncertainties in comparison to Eq. S9. Thus, for the remainer of the paper, the values calculated from Eq. 13 will be used. Another limitation in this study is assuming a constant α and γ across all NO_x mixing ratios to estimate the unmeasured R(VOC). At higher NO_x mixing ratios, where the VOC mixing ratios would be highest due to being closer to emissions, it would be expected that both α and γ would change. However, the direction that these values would change is uncertain as both α and γ depend on the structure of the VOC, which is currently unknown.

For the range of missing α assumed, an α = 0.10 for the unmeasured R(VOC) provides the best agreement with the observed R(VOC) ("From PSU" is the Penn State OH Reactivity with inorganic reactivity subtracted out) for all observations where NO_x < 4 ppbv. Further, it is found that α ranging from 0.075 – 0.125 encompasses the associated uncertainty with the observed R(VOC) (\pm 0.64 s⁻¹ (Brune et al., 2019)). This leads to an average unmeasured R(VOC) of 1.7 $^{+1.1}_{-0.4}$. The associated total missing R(VOC) for the assumed α of 0.10 ranges from 1.5 to 2.8 s⁻¹ (Figure 4a upper panel). Assuming typical rate constants for emitted VOCs, assuming it is comparable to semi- and intermediate-VOCs, and their associated secondary products (~1 – 4×10⁻¹¹ cm³ molec. S⁻¹ (Ma et al., 2017; Zhao et al., 2014)), the total missing reactivity would be equivalent to ~1 – 8 ppbv. Zhao et al. (2014) observed ~12 µg m⁻³ of semi- and intermediate-VOCs

near Los Angeles, CA, during the CalNex study. Depending on the molecular weight assumed,

this translates to ~ 1 to 2 ppbv. Nault et al. (2018) found that $\sim 5-8$ ppbv of VOCs were needed to

explain the observed secondary organic aerosol production over the SMA, depending on the

molecular weight assumed for the VOC. Further, Kenagy et al. (2021) also found that known chemistry could only account for ~33% of the observed ANs and missing sources of lower volatility VOCs to produce anthropogenically-derived ANs were necessary. Finally, Whalley et al. (2016) found that addition of unassigned VOCs and their associated oxidation products led to a reactivity of ~1.6 s⁻¹, leading to ~1 – 6 ppbv missing R(VOC). Thus, the reactivity and equivalent mixing ratios estimated here appear plausible and warrant future measurements to understand this unmeasured reactivity sources.

One important aspect of this unmeasured R(VOC) is that it should not be considered one or a couple of VOCs emitted and contributing 1-8 ppbv of VOC in the atmosphere. Instead, it will be the emitted VOCs and its oxidation products summed together to form the 1-8 ppbv of unmeasured VOCs in the atmosphere.

One possible group of missing VOCs are long-chain aldehydes from cooking and vegetative emissions, including nonanal (Hurst Bowman et al., 2003; Rao et al., 2010; Sai et al., 2012; Schauer et al., 2002). Kim et al. (2018) observed cooking organic aerosols at a ground site in SMA, indicating that there should be associated gas-phase emissions from cooking. Higher carbon aldehydes (or cycloalkanes) have been recently suggested to be a potential interference compound with isoprene measurements on a PTR-MS (Coggon et al., 2024; Wargocki et al., 2023). Comparisons of isoprene measured by the PTR-MS and WAS during KORUS-AQ (Figure S7) shows at increasing NO_x mixing ratios (closer to emission sources), the difference between the PTR-MS and WAS isoprene mixing ratios increases. This suggests that there are potential unmeasured OVOCs and/or other C₅H₈ alkenes at high NO_x ratios that cannot be easily determined by the difference between the PTR-MS and WAS. Continuing to use nonanal as a surrogate for this unmeasured OVOC, nonanal has a rate constant consistent with the values used above for the

missing R(VOC) (3.6×10^{-11} cm³ molec.⁻¹ s⁻¹ (Hurst Bowman et al., 2003)). Further, nonanal has an estimated high α of ~0.2 (Hurst Bowman et al., 2003). As typical nonanal mixing ratios have been observed or estimated to be < 500 pptv, this suggests that nonanal or similar OVOCs may contribute to some of the missing reactivity (< 0.45 s⁻¹). Finally, nonanal and other long-chain aldehydes may be an important higher PNs precursor (see Sect. 3.4 for more discussion about unspeciated higher PNs).

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OVOC emissions from multiple sources, including solvent evaporation and other nontransportation emissions, are generally considered to be an important fraction of R(VOC) for urban emissions but may not be measurable by PTR or GC, such as glycols (de Gouw et al., 2018; Gkatzelis et al., 2021; McDonald et al., 2018; Ma et al., 2022; Simpson et al., 2020; Wang et al., 2022; Yang et al., 2022). However, the α for OVOC is potentially smaller than alkanes, though it is highly unconstrained (Orlando and Tyndall, 2012). Note, higher OVOCs have been understudied and thus may have higher α (e.g., nonanal). Thus, if the missing reactivity is mainly OVOCs and it is assumed their α is low, compounds with $\alpha > 0.15$ will be needed for the budget closure shown here. Likely compounds with high α include alkanes, cycloalkenes/alkenes, and aromatics, though the latter is also highly uncertain. Alkanes have typically been a small source for the R(VOC) in urban environments (e.g., McDonald et al., 2018; Simpson et al., 2020; Whalley et al., 2016). Though aromatics contribute a significant fraction of R(VOC) in different Asian urban environments (Brune et al., 2022; Schroeder et al., 2020; Simpson et al., 2020; Whalley et al., 2021), the majority of the aromatic R(VOC) is considered to be measured by WAS over SMA during KORUS-AQ (e.g., measured aromatics account for ~81% of aromatic reactivity in McDonald et al. (2018) and 98% of aromatic reactivity in Whalley et al. (2016), where both studies had more complete VOC measurements). Finally, the cycloalkenes/alkenes originate from

numerous anthropogenic sources (e.g., McDonald et al., 2018; Simpson et al., 2020). One subclass of cycloalkenes includes monoterpenes. Similar to the comparison of isoprene between PTR-MS and WAS, the difference in monoterpenes between these two measurements increases with increasing NO_x (Figure S8). As the interfering compound(s) measured by the PTR-MS and whether they are oxygenated or not is not known, only the WAS monoterpenes are used in this analysis of calculating R(VOC). Assuming the limonene rate constant, the difference between the PTR-MS and WAS monoterpenes raises the terpene reactivity by $0.05 - 0.30 \, \text{s}^{-1}$. Though this does not include any associated photochemical products from the oxidation of monoterpenes and can improve the closure, it does not explain the total missing reactivity $(1.4 - 2.1 \, \text{s}^{-1})$. Thus, the missing R(VOC) is most likely a combination of OVOCs and cycloalkenes/alkenes.

3.4 Sources of PNs over SMA

As shown in Figure 2, ΣPNs account for a larger fraction of the total NO_z budget than ΣANs . ΣPNs are known to be a temporary sink of NO_x and radicals $(R(O)O_2^-)$ due to their short thermal lifetime (~1 hr). Thus, the NO_x emitted in SMA is being transported regionally, impacting the $P(O_x)$.

In Figure 3b, ΣPNs shows some correlation with formaldehyde. Both are secondary products from the photooxidation of VOCs and have short lifetimes, leading to the correlation. However, above 4 ppbv formaldehyde, the correlation shifts as ΣPNs increases more rapidly than formaldehyde. As shown in Figure S9, this change in the relationship between ΣPNs versus formaldehyde is due to changes in the competition in the reaction of the acyl peroxy radical $(R(O)O_2)$ between NO_2 and NO. At low NO-to- NO_2 ratios, R8 is more favorable, as the $R(O)O_2$ is more likely to react with NO_2 compared with NO, leading to more efficient production over

formaldehyde. As the NO-to-NO₂ ratios increase (NO becomes comparable to NO₂, leading to more equal probability in R(O)O₂ reacting to NO and NO₂, leading to production of alkoxy radicals that can form formaldehyde), R9 becomes more dominant, leading to less production of PNs.

To further explore the sources of both PAN and the higher Σ PNs, the F0AM model (Wolfe et al., 2016) was used to predict both unmeasured Σ PNs, constrained by the observed VOCs precursors, PAN, and PPN (Table 2) and budget of all Σ PNs, where PAN and PPN were not constrained. F0AM shows minimal bias in the predicted formaldehyde, NO₂, and OH (Figure S10) when PAN and PPN were constrained. As discussed in Sect. 3.3, though, there is missing R(VOC) of $1.7^{+1.1}_{-0.4}$ s⁻¹. A sensitivity analysis in adding this missing reactivity to F0AM on predicted OH and formaldehyde was conducted (Sect. S4 and Figure S11 – S12). Both OH and formaldehyde are found to be buffered with the addition of this low amount of R(VOC). Thus, though there is good agreement in these intermediate products between observation and F0AM, this analysis for the sources of PAN and higher Σ PNs is expected to be a lower limit. This missing R(VOC) is further observed in the F0AM-predicted higher PNs (Σ PNs without PAN, or Σ PNs-PAN for short) versus formaldehyde, as a general underestimation in the total higher PNs compared to observations is observed (Figure 5a). (Schroeder et al., 2020)

The classes of compounds producing higher PNs in F0AM are shown in Figure 5b. The classes of compounds were selected from the parent VOC which was oxidized into the higher PN (Table S2). Individual PNs with high contributions and/or are typically measured (PPN, PBzN, and MPAN (methacryloyl peroxy nitrate)) or are a large fraction of PNs but have yet to be measured in ambient conditions (PHAN) are shown without any connection to the parent VOC. Further, both PHAN and PPN have numerous precursors while many of the other higher PNs

modeled by F0AM are generally associated with one precursor. At high NO_x mixing ratios, the more reactive VOCs (aromatics, terpenes) contribute a large fraction of the higher PNs (>35% for $NO_x > 4$ ppbv). As the air moves away from SMA (lower NO_x mixing ratios), contributions of higher PNs from longer-lived compounds (e.g., alkanes) and later generation oxidation products start dominating.

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An interesting trend is observed for PPN and PHAN. Both peroxy acyl radicals for PPN and PHAN (C₂H₅C(O)O₂ and CH₂(OH)C(O)O₂, respectively) are products from photooxidation of many VOCs, including aromatics, alkanes, and methyl ethyl ketone (MEK). However, the fractional contribution of PPN to higher PNs remains constant with decreasing NOx while the fractional contribution of PHAN increases with decreasing NO_x (Figure 5b). This stems from the sources of C₂H₅C(O)O₂ versus CH₂(OH)C(O)O₂. The MCM mechanism, which is used for F0AM, produces C₂H₅C(O)O₂ from the photooxidation from both short- and long-lived species (isoprene, C8-aromatics, toluene, ethanol, MEK, propane, and C4-alkanes) while CH₂(OH)C(O)O₂ is produced from the photooxidation of isoprene and ethene. For CH₂(OH)C(O)O₂, the production is through minor channels in the photooxidation of isoprene (~3% yield directly from isoprene and ~20% as a secondary product (Galloway et al., 2011)). Ethene is relatively long-lived, with a lifetime ~ 7 hrs (OH = 5×10^6 molec. cm⁻³) leading to the delay in the production of PHAN. Note, PHAN formation in MCM/F0AM may be overestimated, as Butkovskaya et al. (2006) found that the radical formed from the photooxidation of glycolaldehyde decomposes to form formaldehyde and CO₂, potentially competing with the pathway to form PHAN. Other studies also found that PNs were not observed by photooxidation of glycoaldehyde (Magneron et al., 2005).

The results here in general indicate more speciated measurements of higher PNs are needed. However, as highlighted in Figure 5, improved detection or measurements of PBzN, PHAN, and MPAN would allow for furthering our knowledge in PNs chemistry in urban environments and their role in controlling O_x production.

A qualitative investigation of the precursors of PAN predicted by F0AM are shown in Figure 5c. This provides a basis for further investigation of the sources over the SMA region for PAN as (a) F0AM over-predicts PAN, as noted above, (b) ethanol is currently estimated, similar to Schroeder et al. (2020), and (c) R(VOC) in F0AM is low due to missing precursors. Like the higher PNs, highly reactive R(VOC) contributes a large portion of the PAN budget at high NO_x. The short-lived compounds contribute ~80% of PAN over SMA at the highest NO_x mixing ratios. At lower NO_x mixing ratios, moving away from SMA, longer-lived compounds, such as ethanol, contribute the most towards PAN production (~70%).

One of the interesting contributions not typically observed for PAN is MEK, which also contributes to PPN and PHAN. In prior studies, MEK mixing ratios were typically 0.5 to 2.0 ppbv (Bon et al., 2011; de Gouw et al., 2018; Liu et al., 2015). Over the SMA, 1.5 ppbv of MEK was observed on average with values as high as 8.3 ppbv. Due to the long lifetime of MEK (\sim 30 hrs for the average photolysis rate measured and OH = 5×10^6 molec. cm⁻³), the high mixing ratios of MEK are most likely due to direct emissions (e.g., de Gouw et al., 2005; Liu et al., 2015). Thus, there are potentially large sources of MEK in SMA that need to be considered in properly representing PAN chemistry.

Another potentially important compound for PAN production is ethanol. However, this compound was not measured during KORUS-AQ; instead, it was estimated based on previous ground-based observations, similar to Schroeder et al. (2020). Ethanol is considered to mainly

come from both vehicle emissions (e.g., Millet et al., 2012) and non-transportation emissions, including cleaning agents and solvents (e.g., McDonald et al., 2018). As ethanol use is predicted to increase in the future (e.g., de Gouw et al., 2012) and cleaning agents and other volatile chemical products appear to scale with population (Gkatzelis et al., 2021), ethanol and MEK may continue contributing towards the PAN budget in the SMA in the future.

As a note, two other compounds potentially important for PAN production that were not measured on the DC-8 during KORUS-AQ include methylglyoxal and biacetyl (LaFranchi et al., 2009). In a forested environment that was partially impacted by urban outflow, these two components contributed on average 25% of the PAN budget (LaFranchi et al., 2009). In urban environments, methylglyoxal is believed to mainly originate from aromatic oxidation (Ling et al., 2020); whereas, biacetyl is believed to come from anthropogenic emissions (Xu et al., 2023). Further, as discussed in Sect. 4.3, these two compounds may potentially be important missing HO_x sources, as well. Thus, measurements of these two compounds along with ethanol is necessary to better understand PAN chemistry.

4. Observational constraints of the HO_x and O_x budget over SMA

As highlighted in Figure S1, the three factors impacting instantaneous $P(O_x)$ are R(VOC), $P(HO_x)$, and NO_x loss processes. In Sect. 3, the NO_x loss processes were investigated and provided a constraint for R(VOC) to improve the investigation of $P(O_x)$. With R(VOC) constrained, the RO_2 concentration can be estimated, providing a means to calculate the net $P(O_x)$ and to investigate the major reactions leading to O_x loss and total HO_x ($OH + HO_2 + RO_2$ + $R(O)O_2$) loss. With the latter, this allows for an investigation of the major $P(HO_x)$ reactions, assuming $L(HO_x)$ equals $P(HO_x)$ (see Eq. 1 – 7 in Sect. 2.3).

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4.1 Net O_x production and sources of O_x loss

Using the total R(VOC) from Sect. 3.3 (Figure 4a), the net $P(O_x)$ (Eq. 1 – 2) over SMA during KORUS-AQ has been determined (Figure 6a). The net P(O_x) peaked at 9.3 ppbv hr⁻¹ at ~8 ppbv NO_x. If only the measured and estimated R(VOC) from F0AM secondary products is used to calculate net $P(O_x)$, the value decreases to 7.6 ppbv hr⁻¹, but at the same NO_x mixing ratio. This value is similar to values observed in other urban locations around the world ($\sim 2-20$ ppbv hr⁻¹), showing that many urban areas are still impacted by high P(O_x) values (Brune et al., 2022; Griffith et al., 2016; Ma et al., 2022; Ren et al., 2013; Schroeder et al., 2020; Whalley et al., 2016, 2018). The NO_x distribution over SMA (Figure 1) shows a large area (~127.53°E to 127.18°E, or \sim 39 km), which corresponds to the NO_x mixing ratio that results in maximum P(O_x), as shown in Figure 6a. Thus, a large portion of the SMA will have high instantaneous $P(O_x)$ of ~9 ppbv hr⁻¹. As the median wind speed over SMA during KORUS-AQ was ~5 m s⁻¹, an air parcel would remain at the highest $P(O_x)$ for ~2 hrs, leading to ~18 ppbv O_3 being produced (not including dilution). This agrees with the ~20 ppbv increase in O₃ observed over the Taehwa Research Forest supersite between midday and afternoon overpasses by the DC-8 during KORUS-AQ (Crawford et al., 2021). Thus, though there is a substantial O₃ background observed over SMA (Colombi et al., 2023; Crawford et al., 2021), a large contribution of the O₃ is due to photochemical production. The major reactions leading to O_x loss $(L(O_x))$ are shown in Figure 6b. The two major reactions that lead to O_x loss are net R8 (light and dark red), or the net production of PNs (which includes losses), and R11, reaction of NO₂ with OH (blue) (see Table 1). Note, as discussed in Sect. 2.2, for the budget analysis conducted here, PAN and PPN were constrained to observations. At high NO_x (near emissions, ~30 ppbv), R11 (NO₂ + OH) dominates the L(O_x) budget (> 60%),

with net R8 (net PAN, dark red, and higher PNs, light red) contributing ~25%, and R12 – 14 accounting for the remaining 15% of O_x loss. As NO_x mixing ratios decrease (moving away from emissions), the net R8 reaction, producing both PAN and higher PNs, starts contributing to larger total $L(O_x)$, ranging from 30 – 40%. Furthermore, the net R8 reaction contribution towards $L(O_x)$ remains relatively constants with NO_x mixing ratios as the contribution from R11 (OH + NO_2) decreases. At NO_x mixing ratios < 3 ppbv is when non- NO_x reactions (R12 – 14) contribute greater than 30% of the $L(O_x)$ budget. Thus, proper representation of PAN and higher PNs, both in precursors and speciation, is important in properly understanding the O_x budget in SMA.

4.2 HO_{x} loss over the SMA

Similar to $L(O_x)$, the major reactions leading to $L(HO_x)$ over the SMA during KORUS-AQ were the reactions of NO_x with HO_x , specifically NO_2 with OH (R11) and net PAN (dark red) and higher PNs (light red) production (R8) (Figure 6c). Reaction R11 is most important for NO_x mixing ratios greater than 15 ppbv (50 – 65%). Between 5 and 15 ppbv, R11 is comparable to the net PN production (R8), where R11 comprises 35 – 50% of $L(HO_x)$ while net R8 (sum of higher Σ PNs and PAN) comprises 30 – 40% of $L(HO_x)$. At lower NO_x mixing ratios, R11 is always smaller for $L(HO_x)$ than net R8, where R11 is about a factor of 2 lower than net R8. Production of Σ ANs played a minor role due to the low α_{eff} .

The self-reaction of HO_x species (R15 – R16) contributes minimally to $L(HO_x)$ (less than 10%) for NO_x mixing ratios greater than 8 ppbv. At lower NO_x mixing ratios, R16 starts dominating $L(HO_x)$ budget, increasing from 8% at 8 ppbv to 50% of $L(HO_x)$ at NO_x mixing ratios less than 2 ppbv. Reaction R15 remains relatively small for the $L(HO_x)$ budget, only reaching 7% of the $L(HO_x)$ budget at NO_x mixing ratios less than 2 ppbv.

4.3 Sources of HO_x over SMA

The analysis conducted leads to the ability to constrain HO_x losses over the SMA during KORUS-AQ. This is important as not all typical HO_x sources were measured on the DC-8 during the project (e.g., nitrous acid, or HONO), and HO_x production rates directly impacts $P(O_x)$ (Figure S1). Prior studies (e.g., Griffith et al., 2016; Tan et al., 2019; Whalley et al., 2018) have demonstrated that in urban environments, sources of HO_x include photolysis of O_3 and subsequent reaction with water vapor, formaldehyde photolysis, and HONO photolysis. Furthermore, recent studies have highlighted the potential importance of typically non-measured OVOCs in their contribution to $P(HO_x)$ and subsequent $P(O_x)$ in an urban environment (Wang et al., 2022). To constrain the $P(HO_x)$ over SMA during KORUS-AQ, the $P(HO_x)$ was assumed to be equal to the observationally constrained $L(HO_x)$. Then, $P(HO_x)$ was calculated for the measurements on the DC-8, including photolysis of O_3 , formaldehyde, H_2O_2 , and other measured OVOCs (Table 2).

Comparing the calculated $P(HO_x)$ and $L(HO_x)$, ~1.5 ppbv hr⁻¹ $P(HO_x)$ (range 1.3 – 1.8 ppbv hr⁻¹) is not accounted for, leading to ~45% of the necessary $L(HO_x)$ to maintain steady-state (Figure 7). For the calculated $P(HO_x)$ budget, O_3 and formaldehyde photolysis contributed ~50% and 40% of the budget, respectively, with the remainder coming from photolysis of H_2O_2 and other measured OVOCs. Accounting for the unobserved $P(HO_x)$, O_3 and formaldehyde photolysis contributed ~25% and ~20%, respectively.

Potential missing sources of $P(HO_x)$ are briefly speculated here. First, one potential source is the photolysis of methylglyoxal. Using the F0AM predicted methylglyoxal, as it was not measured on the DC-8, methylglyoxal would contribute ~0.24 ppbv hr⁻¹ $P(HO_x)$, or ~16% of the unobserved $P(HO_x)$. Another OVOC not measured on the DC-8 and expected to originate from

anthropogenic emissions and not from chemistry is 2,3-butanedione, or biacetyl (de Gouw et al., 2018; Grosjean et al., 2002; Schauer et al., 2002; Xu et al., 2023; Zhou et al., 2020). Prior studies observed 20 – 400 pptv of biacetyl (de Gouw et al., 2018; Xu et al., 2023), correspond to 0.04 – 0.74 ppbv hr⁻¹, or 3 – 49% of the unobserved P(HO_x). Thus, between these two OVOCs, 19 – 66% of the unobserved P(HO_x) could be explained. Other unmeasured OVOCs could potentially contribute to the observed P(HO_x) (e.g., Wang et al., 2022); however, there is less constraints both on the speciation and photolysis rates for these OVOCs (e.g., Mellouki et al., 2015). Finally, HONO could contribute to this observed P(HO_x). Up to 700 pptv of HONO was observed in SMA during KORUS-AQ (Gil et al., 2021), though, this would quickly photolyze to the altitudes the DC-8 flew over SMA (Tuite et al., 2021). Even at 50 – 100 pptv HONO, photolysis of HONO would lead to 0.2 – 0.4 ppbv hr⁻¹ P(HO_x), or 13 – 27% of the unobserved P(HO_x). Thus, between methylglyoxal, biacetyl, and HONO, between 32 – 92% of the unobserved P(HO_x) could be accounted for. This analysis highlights the importance of measuring these HO_x sources to better understand and constrain O_x chemistry in SMA and other urban environments.

A comparison of HO_x sources from F0AM is shown in Figure S14. As it has more complete OVOCs than the observations, the contributions are different than shown in Figure 7. Both observations and F0AM agree that photolysis of O₃ and subsequent reaction with water (R12) and photolysis of formaldehyde are the two largest sources of HO_x. F0AM also shows that methylglyoxal is an important source of HO_x, which is not shown in Figure 7 as methylglyoxal was not measured. However, the total F0AM P(HO_x) was ~2.4 ppbv hr⁻¹, which was lower than the observationally constrained value. This further supports either potential unmeasured OVOCs coming from both emissions and chemistry and/or uncertainty in the photolysis rate constants for these OVOCs (e.g., Wang et al., 2022).

One note about this analysis is that particulate matter collected onto the downwelling CAFS optics during KORUS-AQ (see Sect. S5, Table S3, and Figure S13). Corrections of up to 20% were determined, and the associated uncertainties were also increased by 20% due to the corrections. Thus, the exact amount of unmeasured $P(HO_x)$ is potentially smaller than discussed.

5. Conclusions and Implications

In the Seoul Metropolitan Area (SMA), the ozone (O_3) mixing ratio often exceeds current standards and is increasing. Many processes can impact the O_3 mixing ratios and exceedances. Here, the processes that impact instantaneous O_3 production ($P(O_x)$, where O_x is $O_3 + NO_2$ to account for possible O_3 titration) were investigated for observations collected on the NASA DC-8 during the 2016 NIER/NASA Korea United-States Air Quality (KORUS-AQ) study. The observations indicate missing oxidized NO_x products (NO_z) that include both the short-lived peroxy nitrates (ΣPNs) and alkyl and multi-functional nitrates (ΣANs). ΣPNs contributed the most for the organic NO_z species. Only $\sim 50\%$ of the ΣPNs were speciated over SMA, which is atypical as prior studies typically show closure between the speciated and total PN measurements.

The un-speciated ΣPNs and ΣANs were used to constrain the missing volatile organic compound (VOC) reactivity (R(VOC)), as R(VOC) is important in constraining the instantaneous P(O₃). The missing R(VOC) was found to be 1.4 to 2.1 s⁻¹. The F0AM box model further supports the role of unmeasured ΣPNs as an important temporary NO_x and radical sink over SMA. F0AM predicts ~50% of the higher ΣPNs (higher $\Sigma PNs = \Sigma PNs - PAN$), indicating missing R(VOCs) may explain the other 50%. Constraints from both the ΣPNs and ΣANs suggest that this missing R(VOC) would include oxygenated VOCs (OVOCs), including aldehydes such as octanal and nonanal from cooking, and alkenes from anthropogenic emissions.

With the constraints on the R(VOC), the net instantaneous $P(O_x)$ was determined for SMA. It was found to peak at ~10 ppbv hr⁻¹ at ~8 ppbv NO_x. A large fraction of the SMA area was, on average, at this mixing ratio of NO_x, indicating high local $P(O_x)$. This supports the increase of ~20 ppbv of O₃ observed in a downwind site (Taehwa Research Forest supersite) from midday to afternoon during KORUS-AQ.

With the comprehensive measurements on-board the DC-8, the F0AM model results, and the observationally constrained R(VOC), a budget analysis on the sinks of O_3 (L(O_x)) and HO_x (L(HO_x), where HO_x = OH + HO₂ + RO₂ + R(O)O₂) was performed. Due to the high R(VOC), type of VOC, and the NO₂-to-NO ratio, net Σ PNs production is surprisingly a large and important sink of O_x and HO_x over SMA (~25 – 40% and 15 – 40% for L(O_x) and L(HO_x), respectively), with production of HNO₃ and radical self-reactions accounting for the other L(O_x) and L(HO_x) losses. Net Σ PNs production as an important L(O_x) and L(HO_x) term is significant, as Σ PNs is a temporary reservoir of both NO₂ and R(O)O₂ but has not traditionally been included in these calculations. Downwind locations separated from the local NO_x and VOC emissions of the SMA will experience increased P(O_x) due to the release of NO₂ and R(O)O₂. With the constraint of L(HO_x), P(HO_x) was investigated, assuming steady-state, and unmeasured HONO plus unmeasured OVOCs were found to be necessary to explain the missing HO_x sources. Both sources of HO_x are either missing or highly uncertain in chemical transport models.

Though the high regional background and foreign sources of O_3 and its precursors elevate the O_3 levels in SMA and potentially already causes the SMA to be in exceedance for O_3 concentrations, this study highlights the importance local, in-situ $P(O_x)$ to the SMA area, which can further exacerbate the O_3 concentrations for SMA and the surrounding region. The results support the observations of increasing O_3 with decreasing NO_x that has been observed for SMA in

prior studies. Further, the study highlights the important role of unmeasured VOCs and OVOCs and the necessity to understand their sources and role in NO_x and O_3 chemistry. Further, the study demonstrates the interplay of direct emissions or secondary production of PN precursors and its role in net $P(O_x)$. Attempts at specifically reducing the sources of PN may adversely impact net $P(O_x)$, as lower net PN chemistry may increase O_3 due to more NO_2 being available.

Competing Interests

At least one of the (co-)authors is a member of the editorial board of Atmospheric Chemistry and

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Data Availability

770 Version **R6** 1-min merged data used in this analysis available at 771 DOI:10.5067/Suborbital/KORUSAQ/DATA01. The F0AM setup file, input file, and output files are all available at https://doi.org/10.5281/zenodo.10723227. 772

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Author Contribution

- BAN, KRT, and JHC designed the experiment and wrote the paper. BAN and KRT analyzed the
- data. KRT ran the F0AM model and KRT and BAN analyzed the model output. BAN, DRB, PCJ,
- 777 RCC, JPD, GSD, SRH, LGH, JLJ, K-EK, YL, IJS, KU, and AW collected and QA/QC the data
- during KORUS-AQ. All authors contributed to the writing and editing of the paper.

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1251 Tables

Table 1. Reactions described in text along with associated rate constants and references for those rate constants.

	Reaction	Reaction Rate	Reference
R1a	$VOC+OH \xrightarrow{O_2} RO_2^{\cdot}$	Varies	Atkinson (2003); Atkinson and Arey(2003); Atkinson et al. (2006); Bohn and Zetzsch (2012); Sprengnether et al. (2009)
R1b	$VOC+hv \xrightarrow{O_2} RO_2$	Varies/Measured	Shetter & Müller (1999)
R2a	$RO_2^{\cdot}+NO \rightarrow (1-\alpha) RO^{\cdot}+(1-\alpha) NO_2$	2.7×10 ⁻¹¹ ×exp(390/T)	Burkholder et al. (2020)
R2b	$RO_2^{\cdot}+NO \rightarrow \alpha RONO_2$	2.7×10 ⁻¹¹ ×exp(390/T)	Burkholder et al. (2020)
R3	$NO_2 + h\nu \rightarrow NO + O(^3P)$	Measured on DC-8	Shetter & Müller (1999)
R4	$O(^3P) + O_2 \rightarrow O_3$	$3.2 \times 10^{-11} \times \exp(67/T)$	Saunders et al. (2003)
R5	$RO'+O_2 \rightarrow R(O)+HO_2$	Assumed Instantaneous	
R6	$HO_2+NO \rightarrow OH+NO_2$	$3.45 \times 10^{-12} \times \exp(270/T)$	Saunders et al. (2003)
R7	RCHO+OH $\stackrel{\text{O}_2}{\rightarrow}$ R(O)O ₂	Varies	Atkinson (2003); Atkinson and Arey (2003); Atkinson et al. (2006)
R8 ^a	$R(O)O_2^{\cdot}+NO_2 \leftrightarrow R(O)O_2NO_2$	F: 8.69×10 ⁻¹² cm ³ molec. ⁻¹ s ⁻¹ R: 4.30×10 ⁻⁴ s ⁻¹	Burkholder et al. (2020)
R9	$R(O)O_2+NO \rightarrow RO_2+NO_2$	8.1×10 ⁻¹² ×exp(270/T)	Burkholder et al. (2020)
R10	$O_3 + NO \rightarrow O_2 + NO_2$	2.07×10 ⁻¹² ×(-1400/T)	Burkholder et al. (2020)
R11 ^b	$OH+ NO_2 \rightarrow HNO_3$	1.24×10 ⁻¹¹ cm ³ molec. ⁻¹ s ⁻¹	Burkholder et al. (2020)
R12	O_3 +hv $\xrightarrow{H_2O}$ 2O(1 D)	hv measured on DC-8; $2.14 \times 10^{-10} \text{ cm}^3 \text{ molec.}^{-1} \text{ s}^{-1}$	Shetter & Müller (1999); Saunders et al. (2003)
R13	$O_3+OH \rightarrow HO_2+O_2$	$1.7 \times 10^{-12} \times \exp(-940/T)$	Saunders et al. (2003)
R14	$O_3 + HO_2 \rightarrow OH + 2O_2$	$1.0 \times 10^{-14} \times \exp(-490/T)$	Burkholder et al. (2020)
R15 ^b	$HO_2+HO_2 \xrightarrow{H_2O} H_2O_2$	5.06×10 ⁻¹² cm ³ molec. ⁻¹ s ⁻¹	Saunders et al. (2003)
R16	$HO_2 + RO_2 \rightarrow Products$	$2.91 \times 10^{-13} \times \exp(1300/T)$	Saunders et al. (2003)
R17	$HO_2+OH \rightarrow Products$	4.80×10 ⁻¹¹ ×exp(250/T)	Burkholder et al. (2020)

R18 ^b	OH+NO → HONO	7.40×10 ⁻¹² cm ³ molec. ⁻¹ s ⁻¹	Burkholder et al. (2020)
R19	$HO_2+R(0)O_2 \rightarrow Products$	$4.30 \times 10^{-13} \times \exp(1040/T)$	Burkholder et al. (2020)

^aOnly showing forward (F) and reverse (R) rate constant at 298 K and 1013 hPa and being a termolecular reaction.

^bTermolecular reaction; only showing rate at 298 K and 1013 hPa

Table 2. List of instruments, compounds measured, accuracy/precision, and associated references used in this study.

Instrument	Species	References
University of California, Irvine, Whole Air Sampler (WAS)	Ethane, Ethene, Ethyne, Propane, Propene, i-Butane, n-Butane, 1-Butene, i-Butene, trans-2-Butene, cis-2-Butene i-Pentane, n-Pentane, 1,3-Butadiene, Isoprene, n-Hexane, n-Heptane, n-Octane, n-Nonane, n-Decane, 2,3-Dimethylbutane, 2-Methylpentane, 3-Methylpentane, Cyclopentane, Methylcyclopentane, Cyclohexane, Methylcyclohexane, Benzene, Toluene, m+p-Xylene, o-Xylene, Ethylbenzene, Styrene, i-Propylbenzene, n-Propylbenzene, 3-Ethyltoluene, 4-Ethyltoluene, 2-Ethyltoluene, 1,3,5-Trimethylbenzene, 1,2,4-Trimethylbenzene, 1,2,3-Trimethylbenzene, α-Pinene, β-Pinene, Methyl nitrate, Ethyl nitrate, i-Propyl nitrate, n-Propyl nitrate, 2-Butyl nitrate, 3-Pentyl nitrate, 2-Pentyl nitrate, 3-Methyl-2-Butyl nitrate	Simpson et al. (2020)
The Pennsylvania State University Airborne Tropospheric Hydrogen Oxides Sensor (ATHOS)	OH, HO ₂ , OH Reactivity	Faloona et al. (2004), Mao et al. (2009), Brune et al. (2019)
University of California, Berkeley, Thermal Dissociation-Laser Induced Fluorescence (TD-LIF)	NO_2 , ΣPNs , ΣANs	Thornton et al. (2000), Day et al. (2002), Wooldridge et al. (2010)
NASA Langley Diode Laser Hygrometer (DLH)	H_2O	Diskin et al. (2002)
NASA Langley Diode Laser Spectrometer Measurements (DACOM)	CO, CH ₄	Sachse et al. (1987)
University of Colorado, Boulder, Compact Atmospheric Multi-species Spectrometer (CAMS)	CH ₂ O, C ₂ H ₆	Richter et al. (2015), Fried et al. (2020)
Gwangju Institute of Science and Technology Korean Airborne Cavity Enhances Spectrometer (K-ACES)	СНОСНО	Min et al. (2016), D. Kim et al. (2022)

NCAR CCD Actinic Flux Spectroradiometers (CAFS)	j-values	Shetter & Müller (1999)
Georgia Institute of Technology Chemical Ionization Mass Spectrometer (GT)	SO ₂ , PAN, PPN, APAN, PBzN	Kim et al. (2007), Lee et al. (2022)
University of Colorado, Boulder, High-Resolution Time-of-Flight Aerosol Mass Spectrometer	pNO_3	Nault et al. (2018), Day et al. (2022)
NCAR 4-Channel Chemiluminescence Instrument (NCAR)	NO, NO ₂ , O ₃ , NO _y	Weinheimer et al. (1994)
California Institute of Technology Chemical Ionization Mass Spectrometer (CIT)	Butene Hydroxynitrates, Butadiene Hydroxnitrates, Ethene Hydroxynitrates, Ethanal Nitrate, Isoprene Hydroxynitrates, Propene Hydroxynitrates, Propanal Nitrate, CH ₃ OOH, Peroxyacetic Acid, HNO ₃ , Hydroxyacetone, H ₂ O ₂	Crounse et al. (2006), Teng et al. (2015)
University of Oslo Proton Transfer Reaction Time-of- Flight Mass Spectrometer (PTR-MS)	Methanol, Acetaldehyde, Acetone+Propanal, Isoprene, MVK+MACR+ISOPOOH, Benzene, Toluene, C8-alkylbenzenes, Monoterpenes, MEK	Müller et al. (2014)
NSRC Meteorological and Geographical Data	Latitude, Longitude, Altitude, Temperature, Pressure	Crawford et al. (2021)

1260 Figures

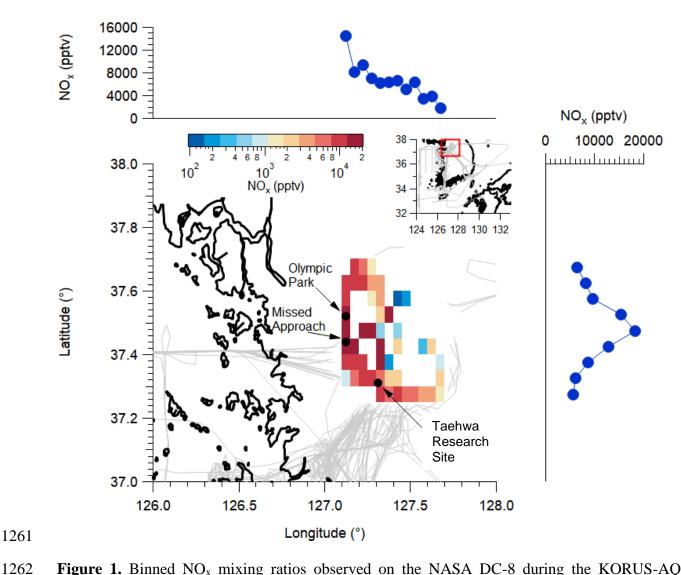


Figure 1. Binned NO_x mixing ratios observed on the NASA DC-8 during the KORUS-AQ campaign. Note, the color bar scale is logarithmic. The binning is along the flight paths of the NASA DC-8 for any observations collected below 2.0 km and after 11:00 local time. The rest of the NASA DC-8 flight paths not included in the analysis are shown in grey. Three key areas from KORUS-AQ are highlighted—the Olympic Park ground site, the airfield where the NASA DC-8 conducted routine missed approaches, and the Taehwa Research ground site. The histograms above and to the left are the distribution of NO_x mixing ratios longitudinally and latitudinally, respectively.

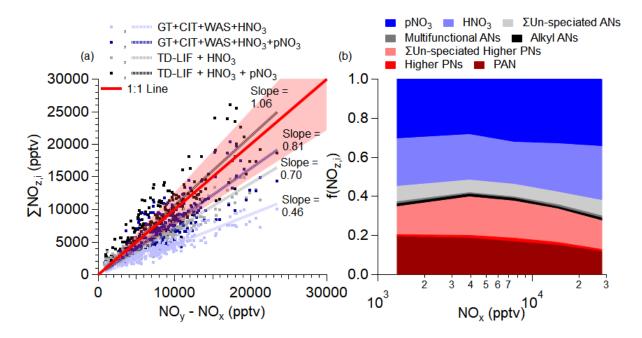


Figure 2. (a) Scatter plot of the summation of individual NO_z (NO_z is higher oxide NO_x products) measured by GT, CIT, WAS, TD-LIF, and AMS versus NO_z measured by difference between NO_y and NO_x (see Table 2 for compounds measured by each instrument). NO_x is NO measured by NCAR and NO₂ measured by LIF. The observations are for when the DC-8 was over the SMA. (b) Average contribution of measured speciated NO_z over the SMA during KORUS-AQ versus NO_x. Higher PNs is PPN + APAN + PBZN. ΣUn-speciated PNs is total peroxnitrates from TD-LIF minus total measurement from GT. Alkyl RONO₂ is the total small alkyl nitrate measurements from WAS. Multifunctional RONO₂ is the total measurements from CIT. ΣUn-speciated ANs is the total alkyl nitrates from TD-LIF minus total RONO₂ from CIT and WAS.

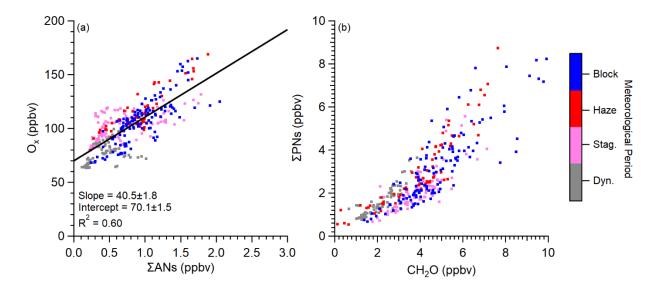
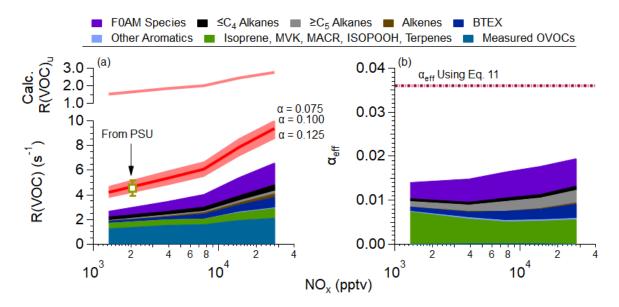


Figure 3. Scatter plot of (a) O_x versus ΣANs and (b) ΣPNs versus formaldehyde (CH₂O) over SMA (see Figure 1 for area studied). Data is colored by meteorological periods discussed in Peterson et al. (2019). Data plotted here is after 11:00 am LT to minimize impact of growing boundary layer and nocturnal residual layer mixing. The curvature in (b) is further explored in Figure S9. Eq. 8, 9, and 11 is used to convert the slope in (a) into α_{eff} . The units of the slope are ppbv ppbv⁻¹.



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Figure 4. (a) Upper panel is the binned calculated (calc.) unmeasured VOC reactivity (R(VOC)₁₁). Note, unmeasured is for any species not measured on DC-8 or constrained by F0AM and is calculated using Eq. 13. Lower panel is binned VOC reactivity versus NO_x observed over SMA during KORUS-AQ (see Figure 1 for the area studied). The measured observed R(VOC), labeled as "From PSU", where PSU is Pennsylvania State University, is the VOC reactivity calculated from the measured total OH reactivity with inorganic OH reactivity removed. As discussed in Brune et al. (2022), the OH reactivity has interferences at high NO_x mixing ratios. The error bar is the uncertainty in the OH reactivity measurement (Brune et al., 2022). The red line represents the calculated unmeasured R(VOC), using Eq. 11, with an assumed $\alpha = 0.10$. The shaded area represents different calculated unmeasured R(VOC), assuming different α for the unmeasured R(VOC) (see Eq. 11). (b) The calculated effective α from observations versus NO_x. The dashed purple line is the effective α estimated from Eq. 11, using the slope from Figure 3a. For both (a) and (b), the colored stacked data is the calculated VOC reactivity (a) and weighted effective α (b). The values from (b) are calculated using Eq. 11. Finally, for both (a) and (b), F0AM species is the reactivity for compounds not measured on the DC-8 predicted by F0AM with an estimated $\alpha =$ 0.05. The associated uncertainty in using different α for the F0AM predicted reactivity is explored in Figure S5.

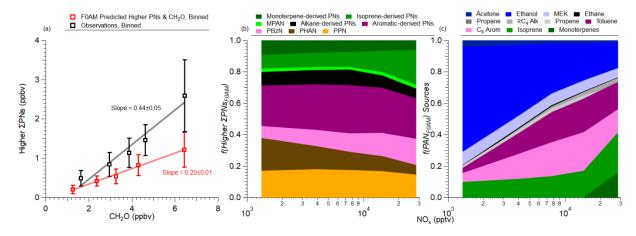


Figure 5. (a) Scatter plot of binned higher ΣPNs calculated using F0AM (red) or binned higher ΣPNs from observations (black) versus formaldehyde (CH₂O). Slopes shown are ODR fits to the binned data. PPN and PAN were constrained by observations for F0AM while all the other higher PNs were not constrained (b) Fractional contribution of the higher PNs predicted from F0AM versus NO_x. (c) Fractional contribution of different precursors to PAN, predicted by F0AM versus NO_x. For both (b) and (c), Alk is all alkanes, Arom is all aromatics, and $\geq C_4$ Alk is all alkanes with 4 or more carbons. See Figure S10 for comparison of F0AM.

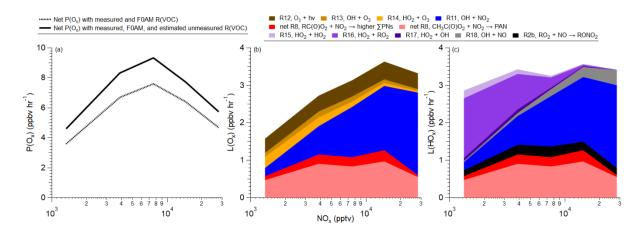


Figure 6. (a) Net O_x ($O_3 + NO_2$) production (see Eq. 1 and 2) predicted for SMA using measured and F0AM R(VOC) (dashed) or total R(VOC) (solid), from Figure 4a, versus NO_x . (b) Contribution of different reactions to the total O_x loss versus NO_x . (c) Contribution of different reactions to total HO_x ($HO_x = OH + HO_2 + RO_2 + R(O)O_2$) loss versus HO_x . The predicted HO_x comes from the total VOC reactivity calculated in Figure 4a assuming steady-state (Eq. 7), and HO_2 the acyl peroxy radicals are from F0AM results. Note for both (b) and (c), net HO_x RC(O)O₂ + HO_x and HO_x are described in Eq. 3. Radical reactions contributing < 1% to the HO_x or HO_x are not included. Also note that F0AM HO_x , HO_x CH₃C(O)O₂, HO_x , and F0AM secondary products are used here along with observations.

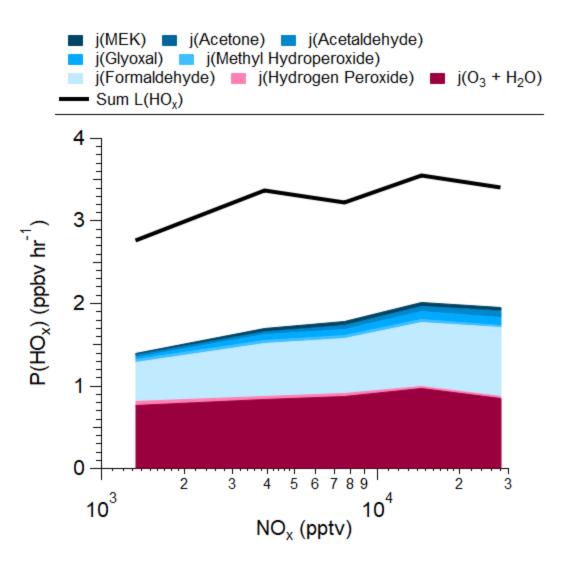


Figure 7. Calculated HO_x production from observations (colored stack) compared with the calculated HO_x loss from Figure 6c over the SMA during KORUS-AQ.