#### 1 Using observed urban NO<sub>x</sub> sinks to constrain VOC reactivity and the ozone and radical

#### 2 budget in the Seoul Metropolitan Area

- 3 Benjamin A. Nault<sup>1,2,\*</sup>, Katherine R. Travis<sup>3</sup>, James H. Crawford<sup>3</sup>, Donald R. Blake<sup>4</sup>, Pedro
- 4 Campuzano-Jost<sup>5</sup>, Ronald C. Cohen<sup>6</sup>, Joshua P. DiGangi<sup>3</sup>, Glenn S. Diskin<sup>3</sup>, Samuel R. Hall<sup>7</sup>, L.
- 5 Gregory Huey<sup>8</sup>, Jose L. Jimenez<sup>5</sup>, Kyung-Eun Kim<sup>9</sup>, Young Ro Lee<sup>8,a</sup>, Isobel J. Simpson<sup>4</sup>, Kirk
- 6 Ullmann<sup>7</sup>, Armin Wisthaler<sup>10,11</sup>
- 7 <sup>1</sup>CACC, Aerodyne Research, Inc., Billerica, MA, USA
- <sup>2</sup>Department of Environmental Health and Engineering, Johns Hopkins University, Baltimore,
   MD, USA
- 10 <sup>3</sup>NASA Langley Research Center, Hampton, VA, USA
- <sup>4</sup>Department of Chemistry, University of California, Irvine, CA, USA
- 12 <sup>5</sup>CIRES and Department of Chemistry, University of Colorado, Boulder, CO, USA
- 13 <sup>6</sup>Department of Chemistry, University of California, Berkeley, CA, USA
- 14 <sup>7</sup>Atmospheric Chemistry Observations & Modeling Laboratory, NCAR, Boulder, CO, USA
- 15 <sup>8</sup>School of Earth & Atmospheric Sciences, Georgia Institute of Technology, Atlanta, GA, USA
- <sup>9</sup>School of Environmental Sciences and Environmental Engineering, Gwangju Institute of Science
   and Technology, Gwangju, South Korea
- <sup>10</sup>University of Oslo, Oslo, Norway
- 19 <sup>11</sup>University of Innsbruck, Innsbruck, Austria
- 20
- 21 22

<sup>a</sup>Now at Division of Geological and Planetary Sciences, California Institute of Technology, Pasadena, CA, USA

- 23 24 25
- 26 \*Corresponding author:
- 27 Email: bnault@aerodyne.com, bnault1@jh.edu
- 28 29 For ACP

#### 30 Abstract

31 Ozone (O<sub>3</sub>) is an important secondary pollutant that impacts air quality and human health. Eastern 32 Asia has high regional O<sub>3</sub> background due to the numerous sources and increasing and rapid industrial growth, which impacts the Seoul Metropolitan Area (SMA). However, SMA has also 33 been experiencing increasing O3 driven by decreasing NOx emissions, highlighting the role of 34 35 local, in-situ O<sub>3</sub> production on SMA. Here, comprehensive gas-phase measurements collected on 36 the NASA DC-8 during the NIER/NASA Korea United States-Air Quality (KORUS-AQ) study 37 are used to constrain the instantaneous O3 production rate over the SMA. The observed NOx 38 oxidized products support the importance of non-measured peroxy nitrates (PNs) in the O<sub>3</sub> chemistry in SMA, as they accounted for ~49% of the total PNs. Using the total measured PNs 39 40 (ΣPNs) and alkyl and multifunctional nitrates (ΣANs), unmeasured volatile organic compound (VOC) reactivity (R(VOC)) is constrained and found to range from 1.4 - 2.1 s<sup>-1</sup>. Combining the 41 42 observationally constrained R(VOC) with the other measurements on the DC-8, the instantaneous 43 net O3 production rate, which is as high as ~10 ppbv hr-1, along with the important sinks of O3 and radical chemistry, are constrained. This analysis shows that  $\Sigma PNs$  play an important role in both 44 45 the sinks of  $O_3$  and radical chemistry. Since  $\Sigma PNs$  are assumed to be in steady-state, the results 46 here highlight the role ΣPNs play in urban environments in reducing altering net O<sub>3</sub> production, 47 but  $\Sigma$ PNs can potentially lead to increased net O<sub>3</sub> production downwind due to their short lifetime 48 (~1 hr). The results provide guidance for future measurements to identify the missing R(VOCs) 49 and  $\Sigma PNs$  production.

#### 50 Short Summary

51	Ozone $(O_3)$ is a pollutant formed from the reactions of gases emitted from various sources. In
52	urban areas, the density of human activities can increase the $O_3$ formation rate (P(O_3)); thus, impact
53	air quality and health. Observations collected over Seoul, South Korea, are used to constrain P(O <sub>3</sub> ).
54	A high local P(O <sub>3</sub> ) was found; however, local P(O <sub>3</sub> ) was partly reduced due to compounds typically
55	ignored. These observations also provide constraints for unmeasured compounds that will impact
56	P(O <sub>3</sub> ).

#### 57 1. Introduction

58 Representing global and urban tropospheric ozone (O<sub>3</sub>) in chemical transport models 59 (CTMs) is still challenging due to uncertainty in physical and chemical processes that control the O3 budget (Archibald et al., 2020). One area of uncertainty is underestimated urban volatile organic 60 61 compounds (VOCs) emissions (von Schneidemesser et al., 2023), which arise from a large 62 number of sources, including some that are very hard to quantify (e.g., cooking and chemical 63 products) (e.g., McDonald et al., 2018; Simpson et al., 2020). Intensive research is also ongoing 64 as to why O<sub>3</sub> is increasing in recent years in urban areas, even with reductions in combustion 65 emissions (Colombi et al., 2023; e.g., Lyu et al., 2017). This  $O_3$  impacts the large populations in 66 urban areas with harmful health effects, including premature mortality (e.g., Cohen et al., 2017). Tropospheric  $O_3$  production is driven by the catalytic cycling of nitrogen oxides (NO<sub>x</sub> = 67 68  $NO + NO_2$ ) fueled by the photoxidation of VOCs, both of which can come from anthropogenic 69 emissions. The chemistry producing  $O_3$  is described in R1 – R6 in Table 1. During daylight hours, 70 VOCs are oxidized by OH (or undergo photolysis) to form an organic peroxy radical (RO<sub>2</sub>) in 71 R1a (R1b). If the RO<sub>2</sub> then proceeds through R2a, at least two  $O_3$  molecules are produced. The 72 first O<sub>3</sub> molecule is formed by the photolysis of NO<sub>2</sub> and the reaction of O(<sup>3</sup>P) with oxygen (R3 – 73 R4). The second  $O_3$  molecule is formed through the reaction of the alkoxy radical (RO) with 74 oxygen to form the hydroperoxyl radical (HO<sub>2</sub>) (R5), which goes on to react with NO to produce 75  $NO_2$  (R6) and the subsequent reactions described above (R3 – R4). However, some fraction of the 76 time, depending on the number of carbons and functional group (e.g., Espada and Shepson, 2005; 77 Perring et al., 2013; Yeh and Ziemann, 2014), alkyl or multifunctional nitrates (ANs = RONO<sub>2</sub>) 78 are formed (R2b). The fraction of reactions to form ANs is described by the branching ratio,  $\alpha$ . 79 Reaction R2b has been shown to impact  $O_3$  production, depending on the types of VOC emitted,

by reducing the fraction of NO<sub>2</sub> that photolyzes to form O<sub>3</sub> in source regions (R3 – R4) (Farmer et
al., 2011). As α is a function of the individual VOC's carbon backbone and functional group (e.g.,
Perring et al., 2013), any uncertainty related to primary VOC emissions and secondary chemistry
will directly impact the ability to describe urban O<sub>3</sub> production.

One important subclass of VOCs are aldehydes (RCHO), which can either be directly 84 emitted or produced via photooxidation of VOCs (de Gouw et al., 2018; Mellouki et al., 2015; 85 86 Wang et al., 2022; Yuan et al., 2012). The photooxidation of the aldehyde (R7) in the presence of 87  $NO_x$  can either form acyl peroxy nitrates (R8, PNs = R(O)O\_2NO\_2) or an organic peroxy radical 88  $(RO_2)$  (R9). The competition between R8 to form PNs versus R9 to form  $RO_2$  depends on the 89 NO-to-NO<sub>2</sub> ratio (Nihill et al., 2021). Further, R8 is in thermodynamic equilibrium due to the weak bond strength between the acyl peroxy radical (R(O)O2<sup>-</sup>) and NO2. Thus, formation of PNs pose 90 91 only a temporary loss of NO<sub>2</sub>. Finally, it has been observed that aldehydes with longer carbon 92 backbones (e.g., C8s and C9s) from various anthropogenic activities, such as cooking (Coggon et 93 al., 2024; Rao et al., 2010), may have mixing ratios as high as aldehydes typically quantified in 94 field experiments (acetaldehyde and propaldehyde). However, there is larger uncertainty associated with these higher aldehydes in their fate to produce both PNs and ANs (e.g., Hurst 95 96 Bowman et al., 2003). Missing both these emissions and subsequent chemistry would impact 97 estimates of urban O3 chemistry.

The fraction of  $RO_2^{-1}$  forming ANs in R2b and the fraction of  $R(O)O_2^{-1}$  forming PNs in R8 alter the instantaneous  $O_3$  production (P( $O_3$ )) by removing NO<sub>2</sub> and/or the radical species. This is further shown in Figure S1, where an analytical equation to describe R1 – R6 (Farmer et al., 2011), is used to explore how changes in the VOC reactivity (R(VOC)), radical production (P(HO<sub>x</sub>)), and ANs production and branching ratio,  $\alpha$  (R2b), impact the instantaneous P(O<sub>3</sub>) (see Sect. S1 for the

103	analytical equation and description). Any changes in $P(HO_x)$ , $R(VOC)$ , and/or $\alpha$ will impact both
104	the instantaneous $P(O_3)$ as well as the $NO_x$ mixing ratio corresponding to the maximum $P(O_3)$ . As
105	these parameters are generally interconnected, investigating all three is important to understand
106	the sources and control of instantaneous $P(O_3)$ . Further, $R7 - R9$ are not included in this traditional
107	description of the analytical equation, as it is assumed PNs are in steady-state (Farmer et al., 2011).
108	Thus, if PNs are not in steady-state, their role in altering P(O <sub>3</sub> ) may be underestimated.

109 Increasing surface O3 is a concern throughout East Asia, including South Korea (Colombi 110 et al., 2023; Gaudel et al., 2018; Kim et al., 2021; Yeo and Kim, 2021). The emissions associated 111 with industry and other anthropogenic activities and the associated photochemistry have impacted 112 regional air quality, leading to high  $O_3$  backgrounds that can impact a country's ability to achieve 113 reduced O<sub>3</sub> exposure for new air quality standards (e.g., Colombi et al., 2023). However, local 114 emissions and photochemistry still play an important role. For example, during the Korea-United 115 States Air Quality (KORUS-AQ) campaign, it was observed between morning and afternoon in 116 the Seoul Metropolitan Area (SMA), O<sub>3</sub> increased by ~20 parts per billion by volume (ppbv) over a background concentration of over 75 ppbv (Crawford et al., 2021). Thus, an understanding of 117 the variables highlighted in Figure S1 are necessary to control both local and regional P(O<sub>3</sub>). 118

One tool typically used to understand the role of regional  $O_3$  and transported  $O_3$  on local O<sub>3</sub> and impacts of local emission controls on  $O_3$  are CTMs. As shown in Park et al. (2021), for the SMA, CTMs typically underestimate the observed  $O_3$  and formaldehyde. While the low  $O_3$  could be partially related to underestimated transport (e.g., Seo et al., 2018) or resolution of the CTM (e.g., Jo et al., 2023; Park et al., 2021), the low bias also observed for modeled formaldehyde indicates overall (a) too little VOCs and thus too low R(VOC) (Brune et al., 2022; H. Kim et al., 2022), (b) missing photochemical products from missing VOCs, including oxygenated VOCs 126 (OVOCs) that contribute to  $P(HO_x)$  (Brune et al., 2022; H. Kim et al., 2022; Lee et al., 2022; Wang 127 et al., 2022), and (c) likely missing PNs and ANs from the underestimated VOCs related to the 128 underestimated R(VOC) (Lee et al., 2022; Park et al., 2021). Missing (a) – (c) will bias the 129 instantaneous P(O<sub>3</sub>) (Figure S1), impacting the ability to investigate what policies should be 130 implemented to reduce O<sub>3</sub>.

To better understand what controls the instantaneous  $P(O_3)$  over SMA, observations collected on the NASA DC-8 during KORUS-AQ are used to constrain the three variables highlighted in Figure S1—R(VOC), HO<sub>x</sub> production and loss, and ANs and PNs production. Observational constraints on these three parameters provide a means to investigate the instantaneous P(O<sub>3</sub>) over SMA and the major classes of contributors to O<sub>3</sub> and HO<sub>x</sub> production and loss. These results are discussed and placed into the context of improving our knowledge about O<sub>3</sub> production in an urban environment.

138

#### 139 2. Methods and Data Description

#### 140 2.1 KORUS-AQ and DC-8 Descriptions

141 The KORUS-AQ campaign was a multi-national project that was conducted in May – June, 142 2016, led by South Korea's National Institute of Environmental Research (NIER) and United 143 States National Aeronautics and Space Administration (NASA). The project was conducted in 144 South Korea and the surrounding seas with numerous airborne platforms, research vessels, and 145 ground sites (Crawford et al., 2021). The study here focuses on the observations collected on the 146 NASA DC-8.

147 The instrument payload, flights, and observations have been described in other studies
148 (Brune et al., 2022; Crawford et al., 2021; Lee et al., 2022; Schroeder et al., 2020). Briefly, the

149	DC-8 was stationed at Osan Air Force Base, Pyeongtaek, South Korea, which is approximately 60
150	km south of Seoul. A total of 20 research flights were conducted with the DC-8. Part of each
151	research flights included a stereo-route in the SMA in the morning (~09:00 local time), midday
152	(~12:00 local time), and afternoon (~15:00 local time), which included a missed approach over
153	Seoul Air Base (< 15 km from Seoul city center) and a fly-over of the Olympic Park and Taehwa
154	Forest Research sites (Figure 1). A total of 55 descents over Olympic Park and 53 spirals over
155	Taehwa Forest Research site were conducted (Crawford et al., 2021). Only observations from the
156	DC-8 after 11:00 local time are used here to ensure that the boundary layer has grown and
157	stabilized and to minimize any influence from residual layer mixing into the boundary layer and/or
158	titration of O <sub>3</sub> by NO (R10). We analyze data collected below 2 km and between $127.10 - 127.67^{\circ}E$
159	and $37.22 - 37.69^{\circ}N$ to focus on the boundary layer in the SMA without influence from industrial
160	emissions along the western South Korean coast (Crawford et al., 2021).

161 During KORUS-AQ, four different meteorological periods, as described by Peterson et al. 162 (2019), impacted the region. These periods included a Dynamic period from 1 - 16 May, where 163 there were a series of frontal passages; a Stagnant period from 17 - 22 May, where it was dry, 164 clear, and stagnant; Transport/Haze period from 25-31 May, where long-range transport and hazy conditions with high humidity and cloud cover prevailed; and, a Blocking period from 1 - 7 June, 165 166 where blocking conditions minimized transporta high pressure ridge is located to an area north of 167 lower pressure, which can preclude significant changes in synoptic meteorology and results in 168 occasional stagnant conditions/minimal pollution transport (Peterson et al., 2019). However, as 169 discussed in Sect. 3.2, conditions did not impact the general trends and chemistry and thus the 170 whole campaign has been analyzed together.

171	The observations used for the analysis are shown in Table 2, along with the associated
172	references. The 1-min merged data from the DC-8 is used here (KORUS-AQ Science Team, 2023).
173	For data missing due to frequency of measurements (e.g., VOCs from WAS), data was filled in a
174	similar approach as Schroeder et al. (2020), in that VOCs with missing data were filled by the
175	linear relationship of that VOC with VOCs measured more frequently. This step was necessary for
176	the observations used in the diel steady-state calculations described in Sect. 2.2. Note, the TD-LIF
177	NO <sub>2</sub> (see Table 2) was used throughout this study and discussed in Sect. S2 and Figure S2 – S $\underline{43}$
178	as it generally agreed better with steady-state calculated NO <sub>2</sub> -to-NO ratios and steady-state $NO_2$
179	than the chemiluminescence NO2. Further, though PBzN was measured by GT-CIMS (Table 2), it
180	is not compared with calculated PBzN from F0AM (Sect. 2.2) as it may be underestimated due to
181	possible inlet losses, as discussed in Zheng et al. (2011).
182	
182 183	2.2 F0AM Box Model Diel Steady-State Calculations for Missing Reactivity and
182 183 184	2.2 F0AM Box Model Diel Steady-State Calculations for Missing Reactivity and Peroxynitrate Budget Analysis
182 183 184 185	<ul><li>2.2 F0AM Box Model Diel Steady-State Calculations for Missing Reactivity and</li><li>Peroxynitrate Budget Analysis</li><li>We use the F0AM box model (Wolfe et al., 2016) with chemistry from the MCMv3.3.1</li></ul>
182 183 184 185 186	2.2 F0AM Box Model Diel Steady-State Calculations for Missing Reactivity and Peroxynitrate Budget Analysis We use the F0AM box model (Wolfe et al., 2016) with chemistry from the MCMv3.3.1 (Jenkin et al., 2015) to simulate production of PNs and formaldehyde using 1-min merged data, as
182 183 184 185 186 187	2.2 F0AM Box Model Diel Steady-State Calculations for Missing Reactivity and Peroxynitrate Budget Analysis We use the F0AM box model (Wolfe et al., 2016) with chemistry from the MCMv3.3.1 (Jenkin et al., 2015) to simulate production of PNs and formaldehyde using 1-min merged data, as described in Sect. 2.1. As in Schroeder et al. (2020), we simulate each <u>1-min merged_aircraft</u>
182 183 184 185 186 187 188	2.2 F0AM Box Model Diel Steady-State Calculations for Missing Reactivity and Peroxynitrate Budget Analysis We use the F0AM box model (Wolfe et al., 2016) with chemistry from the MCMv3.3.1 (Jenkin et al., 2015) to simulate production of PNs and formaldehyde using 1-min merged data, as described in Sect. 2.1. As in Schroeder et al. (2020), we simulate each <u>1-min merged aircraft</u> observation <u>in-through the full diurnal solar</u> cycle <u>mode-(i.e., diel steady-state)</u> until the diurnal
182 183 184 185 186 187 188 189	2.2 F0AM Box Model Diel Steady-State Calculations for Missing Reactivity and Peroxynitrate Budget Analysis We use the F0AM box model (Wolfe et al., 2016) with chemistry from the MCMv3.3.1 (Jenkin et al., 2015) to simulate production of PNs and formaldehyde using 1-min merged data, as described in Sect. 2.1. As in Schroeder et al. (2020), we simulate each 1-min merged_aircraft observation in-through the full_diurnal_solar cycle mode_(i.e., diel_steady-state)_until the diurnal cycle for each_unconstrained species reaches convergence within 1%. These unconstrained species,
182 183 184 185 186 187 188 189 190	2.2 F0AM Box Model Diel Steady-State Calculations for Missing Reactivity and Peroxynitrate Budget Analysis We use the F0AM box model (Wolfe et al., 2016) with chemistry from the MCMv3.3.1 (Jenkin et al., 2015) to simulate production of PNs and formaldehyde using 1-min merged data, as described in Sect. 2.1. As in Schroeder et al. (2020), we simulate each 1-min merged_aircraft observation in-through the full_diurnal_solar cycle mode_(i.e., diel_steady-state)_until the diurnal cycle for each_unconstrained species reaches convergence within 1%. These unconstrained species, such as formaldehyde, NO <sub>2</sub> , and OH, are then evaluated to ensure consistency between F0AM
182 183 184 185 186 187 188 189 190 191	2.2 F0AM Box Model Diel Steady-State Calculations for Missing Reactivity and Peroxynitrate Budget Analysis We use the F0AM box model (Wolfe et al., 2016) with chemistry from the MCMv3.3.1 (Jenkin et al., 2015) to simulate production of PNs and formaldehyde using 1-min merged data, as described in Sect. 2.1. As in Schroeder et al. (2020), we simulate each 1-min merged_aircraft observation in-through the full_diurnal_solar cycle mode_(i.e., diel_steady-state)_until the diurnal cycle for each_unconstrained species reaches convergence within 1%. These unconstrained species, such as formaldehyde, NO <sub>2</sub> , and OH, are then evaluated to ensure consistency between F0AM model and aircraft observations. We constrain concentrations of NO, O <sub>3</sub> , H <sub>2</sub> O <sub>2</sub> , HNO <sub>3</sub> , CO, CH <sub>4</sub> ,
182         183         184         185         186         187         188         189         190         191         192	2.2 F0AM Box Model Diel Steady-State Calculations for Missing Reactivity and Peroxynitrate Budget Analysis We use the F0AM box model (Wolfe et al., 2016) with chemistry from the MCMv3.3.1 (Jenkin et al., 2015) to simulate production of PNs and formaldehyde using 1-min merged data, as described in Sect. 2.1. As in Schroeder et al. (2020), we simulate each 1-min merged_aircraft observation in-through the full_diurnal_solar cycle mode-(i.e., diel steady-state) until the diurnal cycle for each unconstrained species reaches convergence within 1%. These unconstrained species, such as formaldehyde, NO <sub>2</sub> , and OH, are then evaluated to ensure consistency between F0AM model and aircraft observations. We constrain concentrations of NO, O <sub>3</sub> , H <sub>2</sub> O <sub>2</sub> , HNO <sub>3</sub> , CO, CH <sub>4</sub> , H <sub>2</sub> , PAN, PPN, and all measured or estimated VOCs given in Table 2 and Table S1_to calculate

194	<u>PNs budget, <math>\frac{W}{W}</math></u> allow the model to freely calculate NO <sub>2</sub> , formaldehyde, <u>acetaldehyde</u> , and all
195	PNs, including PAN and PPN, for when calculating the budget of PNs. However, for the acyl
196	peroxy radical mixing ratios to calculate Ox and HOx budget (Sect. 2.3), PAN and PPN were
197	constrained by observations. We use a dilution constant of 12 hours, according to Brune et al.
198	(2022). Model evaluation is discussed in Sect. 3.4. The contribution of individual VOCs to PAN
199	was calculated by reducing precursor VOCs by 20% and multiplying the resulting impact on the
200	peroxy acetyl radical (CH <sub>3</sub> C(O)O <sub>2</sub> ) by 5. Other acyl peroxy nitrates (higher PNs) are lumped into
201	categories based on their primary precursor species from Table S2, species currently typically
202	measured (e.g., PPN) or contributes a large fraction of the total higher PNs budget (greater than
203	>2%; e.g., PHAN and MPAN).
204	Note, the reason PAN and PPN were constrained were due to uncertainties in the thermal
205	lifetime, temperature history, and dilution rate used in FOAM, which had larger impacts on the
206	CH <sub>3</sub> C(O)O <sub>2</sub> and PAN than on other unconstrained compounds (e.g., OH and formaldehyde and
207	not shown; Brune et al. (2022)).Part of this larger impact is due to CH <sub>3</sub> C(O)O <sub>2</sub> being one of the
208	most abundant radicals and one of the final radical products in the oxidation of numerous
209	compounds (e.g., Jenkin et al., 2015). We do not expect these uncertainties to impact the higher
210	PNs as (a) there are less precursors to form them compared to PAN and (b) they are expected to
211	have higher thermal stability compared to PAN due to longer carbon backbone (Kabir et al., 2014).
212	
213	
214	2.3 Calculation of Instantaneous Ozone and HO <sub>x</sub> Production and Loss

# 215 An experimental budget for the production and loss of $O_x$ ( $O_x = O_3 + NO_2$ ) and $HO_x$ ( $HO_x$

210	The experimental subjection are production and loss of $O_X (O_X = O_3 + 1(O_2))$ and $HO_X (HO_X)$
216	= $OH + HO_2 + RO_2^{-} + R(O)O_2^{-})$ is described here. NO <sub>2</sub> and O <sub>3</sub> are combined to reduce any potential
217	impact from titration via O3 reaction with NO (R10). The budget analysis includes field-measured

**Formatted:** Indent: First line: 0.5"

219	missing R(VOC) (Sect. 3.2) and published kinetic rate constants (see Table 1 for reference	ces). The
220	rate of production or destruction is calculated with the following equations (Eq. $1 - 7$ ) below	ow. Note,
221	these equations differ from Schroeder et al. (2020) in that (a) ANs and PNs chemistry are o	explicitly
222	included and (b) the reaction of O <sub>3</sub> with alkenes is excluded as this reaction contributed	l a minor
223	loss to $O_3$ (< 1%).	
224	$P_{O_x} = \sum_{i} (1 - \alpha_{ieff}) k_{RO_{2,i} + NO} [RO_{2,i}^{\cdot}] [NO] + k_{HO_2 + NO} [HO_2] [NO]$	(1)
225	$L_{O_x} = k_{NO_2+OH}[NO_2][OH] + k_{O_3+OH}[O_3][OH] + f \times j_{O^1D}[O_3] +$	
226	$k_{HO_2+O_3}[HO_2][O_3] + net(PNs)$	(2)
227	$net(PNs) = \beta k_{R(O)O_2 + NO_2}[R(O)O_2][NO_2] - (1-\beta)k_{decomposition}[PNs]$	(3)
228	$\beta = \frac{k_{RC(O)O_2 + NO_2}[NO_2]}{k_{RC(O)O_2 + NO_2}[NO_2] + k_{RC(O)O_2 + NO}[NO]}$	(4)
229	$P(HO_x) = 2f \times j_{O^1D}[O_3] + 2j_{H_2O_2}[H_2O_2] + 2j_{CH_2O \to H+HC0}[CH_2O] + 2j_{CHOCHO}[CHO]$	CHO] +
230	$2j_{CH_3OOH}[CH_3OOH] + 2j_{CH_3CHO}[CH_3CHO] + 2j_{CH_3C(O)CH_3}[CH_3C(O)CH_3] + 2j_{CH_3C(O)CH_3C(O)CH_3}[CH_3C(O)CH_3] + 2j_{CH_3C(O)CH_3C(O)CH_3C(O)CH_3}[CH_3C(O)C(O)CH_3C(O)C(O)CH_3C(O)$	
231	$2j_{CH_3CH_2C(O)CH_3}[CH_3CH_2C(O)CH_2]$	(5)
232	$L(HO_{x}) = k_{NO_{2}+OH}[NO_{2}][OH] + \sum_{i} \alpha_{eff} k_{RO_{2},i} + NO[RO_{2},i][NO] + NO[RO_{2},i][NO$	
233	$2k_{HO_2+HO_2}[HO_2][HO_2] + 2k_{RO_2+RO_2}[RO_2][RO_2] + 2k_{HO_2+RO_2}[HO_2][RO_2] + net(PNs)$	(6)
234	$[\mathrm{RO}_{2}^{\cdot}] = \frac{\sum_{i  k_{\mathrm{OH+VOC},i}} [\mathrm{VOC}_{i}] [\mathrm{OH}]}{(1 - \alpha_{\mathrm{eff}}) k_{\mathrm{RO}_{2} + \mathrm{NO}} [\mathrm{NO}] + k_{\mathrm{RO}_{2} + \mathrm{HO}_{2}} [\mathrm{HO}_{2}]}$	(7)
235	Here, k is the rate constant for compound, i, with the associated compound listed, o	a <sub>eff</sub> is the
236	effective branching ratio for R2a and R2b for the observations (Sect. 3.2), f is the fraction	that O <sup>1</sup> D
237	that reacts with water to form OH versus reacting with a third body molecule to form $O^{3}$	P, $\beta$ is the

quantities (mixing ratios and photolysis rates, Table 2), results from FOAM (Sect. 2.2), estimated

218

238

the

fraction the R(O)O2<sup>-</sup> that reacts with NO2 versus NO, and j is the measured photolysis frequency

### 11

239 (Table 2). In Eq. 5, only directly values directly measured on the DC-8 during KORUS-AO are 240 included. As discussed in Wang et al. (2022) and Sect. 4.3,, this is most likely an underestimations 241 of  $P(HO_x)$ . Note,  $R(O)O_2$  is not included in Eq. 7 as (a) it is assumed the initial production of 242  $R(O)O_2$  is captured with the reaction of OH with VOC and (b)  $R(O)O_2$  accounts for a small 243 fraction of the total RO<sub>2</sub> (< 10%). Not including R(O)O<sub>2</sub> in Eq. 7 may lead to a small 244 underestimation of total RO2. Finally, HO2 calculated from F0AM, rather than aircraft 245 <u>measurements</u> (Crawford et al., 2021), is used in the equations to determine the  $O_x$  and  $HO_x$  budget 246 (see Sect. S3, Figure S6).

247

#### 248 3. Observational constraints on NO<sub>x</sub> organic oxidation chemistry

In the Sect. 3.1, the detailed observations from the DC-8 during KORUS-AQ provided measurements that allow us to test our understanding of NO<sub>x</sub> oxidation into total NO<sub>z</sub> (NO<sub>z</sub> = higher NO<sub>x</sub> oxides, including  $\Sigma$ PNs,  $\Sigma$ ANs, HNO<sub>3</sub> and particulate nitrate, pNO<sub>3</sub>), which is needed for the remainder of the analysis. Sect. 3.2 to 3.4 will focus on the organic NO<sub>z</sub> chemistry. This is due to the chemistry and dynamics impacting the total inorganic nitrate chemistry that has been discussed recently (Jordan et al., 2020; Travis et al., 2022).

255

#### 256 3.1 NO<sub>x</sub> and its oxidation products

The average  $NO_x$  mixing ratios observed by the NASA DC-8 in the SMA below 2 km after 11:00 local time is shown in Figure 1. As  $NO_x$  is mainly emitted from anthropogenic activities, such as combustion emissions, in an urban environment, the largest  $NO_x$  mixing ratios are observed between Olympic Park and the missed approach, as this area included downtown SMA. The missed approach included low level sampling at a military airport, which may have 267 On the DC-8, there were multiple measurements of various speciated and total family 268 contribution towards NOz (Table 2). The comparison of the speciated and measured NOz is 269 investigated in Figure 2 for observations over SMA. When only speciated PNs (GT) and ANs (CIT 270 + WAS) and gas-phase nitrate (HNO<sub>3</sub>) are compared to the NO<sub>z</sub> (NO<sub>y</sub> (NCAR) – (NO (NCAR) + 271  $NO_2$  (TD-LIF)), only 46% of the  $NO_z$  can be explained. This is not completely unexpected, as 272 multiple studies have indicated that the speciated ANs measurements are typically lower than the 273 total ANs measurements (Fisher et al., 2016; Perring et al., 2010). Further, pNO<sub>3</sub> has been found 274 to be important for total nitrate budget in the SMA (e.g., Travis et al., 2022). Chemiluminescence 275 measurements of gas-phase NO<sub>y</sub> have been found to efficiently measure pNO<sub>3</sub>, depending on the 276 sensitivity to  $pNO_3$  enhancements or exclusions (Bourgeois et al., 2022); thus, it is expected that 277 missing ANs and pNO<sub>3</sub> are necessary to close the NO<sub>z</sub> budget. Adding the measured pNO<sub>3</sub> to the 278 speciated PNs (GT) and ANs (CIT + WAS) and gas-phase nitric acid, 81% of NOz can be 279 explained. This barely overlaps the combined uncertainty of the measurements (~26%). Total PNs 280 and ANs, measured by TD-LIF, are needed to close of the total NOz budget.

The breakdown of the NO<sub>z</sub> budget over the SMA as the airmasses photochemically ages (decreasing NO<sub>x</sub> contribution to total NO<sub>y</sub>) is shown in Figure 2b. During KORUS-AQ, ~56% of NO<sub>z</sub> was inorganic (gas- and particle-phase nitrate), ranging from 52% to 62%; the remaining NO<sub>z</sub> was organic (PNs and ANs). Approximately 74% of the total ANs were not speciated (range 73% 285 to 76%). Speciated PNs species, such as PAN (peroxy acetyl nitrate), account for a mean 51% of the total PNs (range 47 to 59%), much lower than typically observed in prior studies (e.g., 286 287 Wooldridge et al., 2010). In these prior studies, the speciated PN species (typically PAN + PPN (peroxy propionyl nitrate)) accounted for 90 – 100% of the  $\Sigma$ PNs, except for some select cases 288 289 attributed to poor inlet design (Wooldridge et al., 2010). PAN accounted for the majority of the 290 speciated PNs, with the remaining speciated PNs (PPN + PBzN (peroxy benzoyl nitrate) + APAN 291 (peroxy acryloyl nitrate)) accounting for ~1%. However, during KORUS-AQ, Lee et al. (2022) 292 observed that PAN contributed only 60% of calculated total PNs in industrial plumes near the 293 SMA. Thus, the VOC emissions in and near SMA potentially lead to PNs typically not directly 294 measured; this is explored more in Sect. 3.4

295 As NO<sub>x</sub> decreases from  $\sim$ 30 ppbv to 4 ppbv, the contribution of organic NO<sub>z</sub> increases 296 (Figure 2b). At about 4 ppbv, the contribution of organic  $NO_z$  starts to decrease. Further, the 297 contribution of the different organic NOz species changes. For example, from ~30 ppbv to 4 ppbv, 298 the un-speciated  $\Sigma$ PNs contributes the majority of the organic NO<sub>z</sub> budget (~39%). Below ~4 ppbv, 299 the contribution of un-speciated  $\Sigma$ PNs decreases and the PAN contribution increases. The change 300 in contribution of PNs is due to changes in the PN precursors (e.g., combination short-lived 301 precursors oxidizing to  $CH_3C(O)O_2^{-1}$  and thermal decomposition of the higher PNs (higher PNs = 302  $\Sigma PNs - PAN$ )). On the other hand, the contribution of un-speciated  $\Sigma ANs$  remains relatively 303 constant with NO<sub>x</sub> (~6% of total NO<sub>z</sub>). However, the type of ANs is most likely changing with 304 NO<sub>x</sub> due to the lifetime of the ANs precursors and/or the lifetime of ANs. Less is known about the 305 lifetime of ANs derived from anthropogenically emitted VOCs compared to those from biogenic 306 VOCs (González-Sánchez et al., 2023; Picquet-Varrault et al., 2020; Zare et al., 2018). On average 307 unknown ANs and PNs account for ~24% of the observed NOz on average. The differences in the

14

308	binned mean value for each species is greater than the uncertainty associated with its measurements
309	(maximum uncertainty 30%) and greater than the standard error of the mean, indicating that all the
310	percent differences shown here are real.

311

### 312 **3.2 Meteorological impact on NOx oxidation**

313 As discussed in Sect. 2.1 and various prior studies, four different meteorological conditions 314 impacted the observations during KORUS-AQ (Peterson et al., 2019). The impact of the 315 meteorological conditions on NOx oxidation was investigated by plotting two metrics of NOx 316 oxidation— $O_x$  versus  $\Sigma ANs$  and  $\Sigma PNs$  versus formaldehyde (Figure 3). The implications of both 317 plots are further discussed in Sect. 3.3 and 3.4, respectively. Briefly,  $O_x$  versus  $\Sigma ANs$  and  $\Sigma PNs$ 318 versus formaldehyde are competitive products from the reaction of  $RO_2$  or  $R(O)O_2$  with  $NO_x$ 319 (R2a versus R2b or R8 versus R9). The different meteorological periods corresponded to 320 differences in temperatures and amount of photolysis due to cloud cover (Peterson et al., 2019). 321 Thus, these different periods may impact gas-phase chemistry and/or VOC emissions. However, 322 as demonstrated in Figure 3, there are minimal systematic differences in the trends observed for 323 the two NO<sub>x</sub> oxidation products as there is no systematic shift in the trends or scatter observed in 324 Figure 3. This suggests that the data does not have to be separated by meteorological conditions. 325

#### 326 **3.3 Production of ANs to constrain R(VOC)**

327 Observations of un-speciated ANs and PNs imply missing VOCs that impact  $O_3$  chemistry. 328 The relationship of ANs to  $O_x$  can provide a method to investigate this source. This relationship 329 provides an estimate of the effective branching ratio,  $\alpha$ , for the observed VOC mix (Perring et al., 330 2013 and references therein). The value of this relationship stems from the reactions discussed

331	above $(R1 - R6)$ in that upon the oxidation of VOCs, some fraction of the time, $RO_2$ reacts with	
332	NO to form an AN molecule and the remainder of the time the reaction goes to form $O_3$ . This is	
333	expressed with the following equations:	
334	$P_{\Sigma ANS} = \sum \alpha_i  k_{OH+VOCi} [OH] [VOC_i] \tag{8}$	
335	$P(O_{x}) = \sum_{i} \gamma_{i} (1 - \alpha_{i}) k_{OH+VOC_{i}}[OH][VOC_{i}] $ (9)	
336	Here, $\alpha$ is the effective branching ratio in the reaction of RO <sub>2</sub> <sup>-</sup> with NO to form ANs versus RO <sup>-</sup>	
337	(R2), k is the OH rate constant with VOC, i, and $\gamma$ is the number of O <sub>3</sub> molecules formed per	
338	oxidation of VOC, i. The <u>reactivity weighted</u> $\gamma$ is calculated for the observed and <u>FOAM</u> calculated	
339	compounds species from F0AM using the values from MCM with Eq. 10, where $\gamma$ for each	
340	compound is taken from MCM (Jenkin et al., 2015) and accounts for potential for difference	
341	number of O <sub>3</sub> molecules produced per channel per oxidation (e.g., xylene produces two O <sub>3</sub>	
342	molecules 60% of the time and one O <sub>3</sub> molecule 40% of the time). All the terms were defined for	
343	<u>Eq. 8 – 9.</u>	
344	$\gamma_{eff} = \frac{\sum_{i} \gamma_i k_{OH+VOC_i} [OH] [VOC_i]}{\sum_{i} k_{OH+VOC_i} [OH] [VOC_i]} $ (10)	/
345	<u>The reactivity weighted <math>\gamma_{\tau}</math> is found to be, on average, 1.53, which is lower than the value of 2</u>	

Formatted: Font: Times New Roman, Not Italic Formatted: Font: Times New Roman, Not Italic

The reactivity weighted  $\gamma_7$  is found to be, on average, 1.53, which is lower than the value of 2 typically assumed in prior studies (e.g., Perring et al., 2013). This lower reactivity weighted  $\gamma$  is due to the role of CO ( $\gamma = 1$ ) and CH<sub>2</sub>O ( $\gamma = 1$ ) to the total reactivity. After the boundary layer height has stabilized (e.g., after 11:00 am LT used here) and is near enough (e.g., less than 1 day aging) to the VOC source to ignore deposition and entrainment, Eq. 8 and 9 can be combined to approximate the change in O<sub>x</sub> per molecule  $\Sigma$ AN formed:

$$\frac{\Delta O_x}{\Delta \Sigma A N_S} \approx \frac{P_{O_x}}{P \Sigma A N_S} \approx \frac{1.53(1-\alpha)}{\alpha} \tag{110}$$

For this equation to be valid,  $\alpha$  needs to be relatively small ( $\alpha \ll 1$ ), which is true for VOCs, as maximum  $\alpha$  for the conditions of KORUS-AQ is expected to be 0.35 (Orlando and Tyndall, 2012;

16

Perring et al., 2013; Yeh and Ziemann, 2014). Note, though Eq. 110 can be used at short
photochemical ages due to minimal impact from physical loss processes, chemical loss processes
may impact the assumptions in Eq. 110 and are discussed in more detail below.

357 Over the SMA during KORUS-AQ, the slope between  $O_x$  and  $\Sigma ANs$  was observed to be 40.5±1.8 (Figure 3a), with an  $R^2 = 0.60$ . Using Eq. 110, this translates to an effective branching 358 ratio ( $\alpha_{eff}$ ), of 0.036. For other urban locations around the world, this slope has ranged from 13 – 359 360 47 (Farmer et al., 2011; Kenagy et al., 2020; Perring et al., 2010; Rosen et al., 2004), leading to an 361 effective  $\alpha$  between 0.04 and 0.15, assuming a  $\gamma$  of 2 instead of the calculated  $\gamma$  used here. Thus, 362 the  $\alpha_{eff}$  observed over SMA during KORUS-AQ is similar to other urban locations (Houston = 363 0.05 (Rosen et al., 2004) and South Korea = 0.05 (Kenagy et al., 2021)) but much lower than observed for Mexico City = 0.07 - 0.12 (Farmer et al., 2011; Perring et al., 2010) and Denver = 364 365 0.16 (Kenagy et al., 2020). This suggests that VOCs with low  $\alpha$  dominate the total R(VOC) and 366 production of ANs in SMA. The VOCs in SMA that dominate R(VOCs), including OVOCs, 367 alkenes, and aromatics (Schroeder et al., 2020; Simpson et al., 2020), generally have lower  $\alpha$ 368 (Orlando and Tyndall, 2012; Perring et al., 2013 and references therein).

We use the observed VOCs (Table 2) and estimated secondary products from F0AM to calculate  $\alpha_{eff}$  from this mixture to compare to the calculated  $\alpha_{eff}$  of 0.036 derived from the slope of O<sub>x</sub> versus ΣANs in Figure 3a, as shown in Figure 4. To derive  $\alpha_{eff}$ , Eq. 12 was used, where all the

 $\frac{1}{2} \frac{1}{2} \frac{1}$ 

373

 $\underline{\alpha}_{eff} = \frac{\sum_{i} \alpha_{i} k_{OH+VOC_{i}} [OH][VOC_{i}]}{\sum_{i} \gamma_{i} k_{OH+VOC_{i}} [OH][VOC_{i}]}$ 

(12)

The R(VOC) calculated from the observed VOCs and from the intermediates produced by the FOAM model, described in Sect. 2.2, are shown in Figure 4a, and the reactivity weighted  $\alpha$  for the observations is shown in Figure 4b. As has been observed in other urban environments (e.g., Formatted: Font: Times New Roman, Not Italic Formatted: Font: Times New Roman, Not Italic

377	Hansen et al., 2021; Whalley et al., 2016; Whalley et al., 2021; Yang et al., 2022), measured
378	OVOCs contribute the most to the calculated R(VOC) for all NO <sub>x</sub> mixing ratios $(32 - 48\%)$ . The
379	unmeasured OVOCs (F0AM species) contributed $17-28\%$ of the calculated reactivity. The F0AM
380	species reactivity ranged from $0.45 - 1.78 \text{ s}^{-1}$ , which is a similar increase in total OH reactivity
381	observed by Brune et al. (2022) over South Korea. At higher NO <sub>x</sub> mixing ratios, primary, more
382	reactive VOCs (e.g., alkanes, alkenes, aromatics) contribute an important fraction (> 25%) of the
383	R(VOC). As there are interferences in the total OH reactivity measurement at high NO <sub>x</sub> (Brune et
384	al., 2022), we are unable to determine the extent to which the observed and modeled reactivity
385	captures total OH reactivity in the SMA above a $NO_x$ value of approximately 4 ppbv. At lower
386	$NO_x$ mixing ratios, ~33% of the R(VOC) is missing (calculated R(VOC), including F0AM species,
387	~3.0 s <sup>-1</sup> and measured R(VOC) from Penn State—see Table 2—is 4.5 s <sup>-1</sup> ).

388 Numerous other urban studies have observed unmeasured OH reactivity, which is assumed 389 to be unmeasured R(VOC), as the inorganic OH reactivity is typically well covered by 390 measurements. Here, we are defining unmeasured R(VOC) as the reactivity not represented by 391 measurements on the DC-8 or by FOAM predicted species and reactivity. This unmeasured R(VOC) has ranged from  $\sim 3 \text{ s}^{-1}$  to  $\sim 10 \text{ s}^{-1}$  (e.g., Brune et al., 2022; Hansen et al., 2021; Kim et al., 392 393 2016; Ma et al., 2022; Tan et al., 2019; Whalley et al., 2016; Whalley et al., 2021). Over the SMA, the difference between measured and calculated R(VOC) was ~1.5 s<sup>-1</sup> at low NO<sub>x</sub> and unknown at 394 395 high NO<sub>x</sub> mixing ratios. The lower difference may be related to the comparison occurring for 396 observations at low NO<sub>x</sub>, when the very reactive material has either reacted into compounds 397 measured on the DC-8 (e.g., formaldehyde, acetaldehyde, etc.), diluted to low enough 398 concentrations to be negligible for R(VOC), or undergone deposition or partitioning to the particle-399 phase.

400	At higher $NO_x$ mixing ratios, which is more representative of fresh emissions, these more
401	reactive compounds typically not measured are expected to lead to a higher difference between the
402	calculated and observed R(VOC). Prior studies with more comprehensive measurements found
403	these more reactive compounds and their secondary products contributed an important fraction
404	towards the R(VOC) (e.g., Whalley et al., 2016). Thus, to determine if these unmeasured VOCs
405	potentially contribute to the $R(VOC)$ , and thus $P(O_x)$ , in SMA, another means to constrain their
406	contributions is necessary. One potential means to constrain the total R(VOC) is by using the
407	observed $\Sigma ANs$ and $O_x$ and assuming the observations are from the instantaneous production of
408	both species (e.g., the assumption used for Figure 3a).

409 To estimate the unmeasured R(VOC), Eq. 130 is used without cancelling out terms and 410 expanded into the measured and unmeasured R(VOC) and  $\alpha$ :

411 
$$\frac{\Delta O_x}{\Delta \sum ANs} = \frac{\gamma RVOC_m[OH] + \gamma RVOC_u[OH] - \gamma \alpha_m RVOC_m[OH] - \gamma \alpha_u RVOC_u[OH]}{\alpha_m RVOC_m[OH] + \alpha_u RVOC_u[OH]}$$
(13+)

Here,  $\frac{\Delta O_x}{\Delta \sum ANS}$  is the slope from Figure 3a,  $\gamma$  is the number of O<sub>3</sub> molecules formed per oxidation of 412 VOC, which is 1.53 for this study, R(VOC) is the VOC reactivity, which is its OH oxidation rate 413 constant and its concentration (k×[VOC]) in units s<sup>-1</sup>,  $\alpha$  is the branching ratio for R2 (Table 1), and 414 415 m and u correspond to "measured" (measured VOCs on DC-8 along with secondary species 416 predicted by F0AM) and "unmeasured" (unmeasured VOCs that are not represented by DC-8 417 observations and not predicted by FOAM) RVOC and a. The rate constants for the measured VOCs 418 are listed in Table 1, the reactivity for F0AM is taken directly from F0AM, and  $\alpha$  is either from 419 MCM (Jenkin et al., 2015) or Perring et al. (2013) for observations or assumed to be 0.05 for 420 FOAM secondary products. The equation is rearranged and solved for RVOC<sub>u</sub>, using different 421 values of  $\alpha_u$  (e.g., 0.00 – 0.30, values typical  $\alpha$ ).

422	As discussed in Sect. S3 in the Supp. Information, there are numerous assumptions and
423	potential sources of uncertainty in the simplified version of Eq. 1 $\underline{3}$ +. A thorough analysis and
424	discussion of these assumptions are discussed in Sect. S3. The potentially most important
425	assumption is that chemical loss is negligible in solving Eq. 1 $\underline{34}$ . However, due to the expected
426	relatively short lifetime of $\sum$ ANs, the chemical loss of both $O_x$ and ANs nearly cancel each other,
427	leading to similar results in considering or neglecting these loss terms in Eq. 134. Further, as $\sum ANs$
428	chemical loss has uncertainty, especially for ANs produced from anthropogenic VOC oxidation,
429	the use of Eq. $134$ reduces some of these uncertainties in comparison to Eq. S9. Thus, for the
430	remainer of the paper, the values calculated from Eq. 1 <u>3</u> 4 will be used. <u>Another limitation in this</u>
431	study is assuming a constant $\alpha$ and $\gamma$ across all $NO_x$ mixing ratios to estimate the unmeasured
432	<u>R(VOC)</u> . At higher $NO_x$ mixing ratios, where the VOC mixing ratios would be highest due to
433	being closer to emissions, it would be expected that both $\alpha$ and $\gamma$ would change. However, the
434	direction that these values would change is uncertain as both $\alpha$ and $\gamma$ depend on the structure of
435	the VOC, which is currently unknown.
436	For the range of missing $\alpha$ assumed, an $\alpha = 0.10$ for the unmeasured R(VOC) provides the
437	best agreement with the observed R(VOC) ("From PSU" is the Penn State OH Reactivity with
438	inorganic reactivity subtracted out) for all observations where $NO_x < 4$ ppbv. Further, it is found
439	that $\alpha$ ranging from 0.075 $-$ 0.125 encompasses the associated uncertainty with the observed
440	R(VOC) (±0.64 s <sup>-1</sup> (Brune et al., 2019)). This leads to an average unmeasured R(VOC) of $1.7^{+1.1}_{-0.4}$ .
441	The associated total missing R(VOC) for the assumed $\alpha$ of 0.10 ranges from 1.54 to 2.84
442	s <sup>-1</sup> (Figure 4a upper panel). Assuming typical rate constants for emitted VOCs, assuming it is
443	comparable to semi- and intermediate-VOCs, and their associated secondary products (~ $1-4 \times 10^{-1}$
444	<sup>11</sup> cm <sup>3</sup> molec. <sup>-1</sup> s <sup>-1</sup> (Ma et al., 2017; Zhao et al., 2014)), the total missing reactivity would be

equivalent to  $\sim 1-8$  ppbv. Zhao et al. (2014) observed  $\sim 12 \ \mu g \ m^{-3}$  of semi- and intermediate-VOCs 445 near Los Angeles, CA, during the CalNex study. Depending on the molecular weight assumed, 446 447 this translates to  $\sim 1$  to 2 ppbv. Nault et al. (2018) found that  $\sim 5 - 8$  ppbv of VOCs were needed to explain the observed secondary organic aerosol production over the SMA, depending on the 448 449 molecular weight assumed for the VOC. Further, Kenagy et al. (2021) also found that known 450 chemistry could only account for ~33% of the observed ANs and missing sources of lower 451 volatility VOCs to produce anthropogenically-derived ANs were necessary. Finally, Whalley et 452 al. (2016) found that addition of unassigned VOCs and their associated oxidation products led to a reactivity of ~1.6 s<sup>-1</sup>, leading to ~1 – 6 ppbv missing R(VOC). Thus, the reactivity and equivalent 453 454 mixing ratios estimated here appear plausible and warrant future measurements to understand this 455 unmeasured reactivity sources.

One important aspect of this unmeasured R(VOC) is that it should not be considered one or a couple of VOCs emitted and contributing 1 - 8 ppbv of VOC in the atmosphere. Instead, it will be the emitted VOCs and its oxidation products summed together to form the 1 - 8 ppbv of unmeasured VOCs in the atmosphere.

460 One possible group of missing VOCs is-are long-chain aldehydes from cooking and 461 vegetative emissions, including nonanal, which is associated with cooking emissions (Hurst 462 Bowman et al., 2003; Rao et al., 2010; Sai et al., 2012; Schauer et al., 2002) and vegetative 463 emissions. Kim et al. (2018) observed cooking organic aerosols at a ground site in SMA, indicating 464 that there should be associated gas-phase emissions from cooking. Higher carbon aldehydes (or 465 cycloalkanes)Nonanal has have been recently been suggested to be a potential interference 466 compound with isoprene measurements on a PTR-MS (Coggon et al., 2024; Wargocki et al., 2023). 467 Comparisons of isoprene measured by the PTR-MS and WAS during KORUS-AQ (Figure S75)

21

468 shows at increasing  $NO_x$  mixing ratios (closer to emission sources), the difference between the 469 PTR-MS and WAS isoprene mixing ratios increases. This suggests that there are potential 470 unmeasured OVOCs and/or other C<sub>5</sub>H<sub>8</sub> alkenes at high NO<sub>x</sub> ratios that cannot be easily determined 471 by the difference between the PTR-MS and WAS. Continuing to use nonanal as a surrogate for 472 this unmeasured OVOC, nonanal has a rate constant consistent with the values used above for the missing R(VOC) (3.6×10<sup>-11</sup> cm<sup>3</sup> molec.<sup>-1</sup> s<sup>-1</sup> (Hurst Bowman et al., 2003)). Further, nonanal has 473 474 an estimated high  $\alpha$  of ~0.2 (Hurst Bowman et al., 2003). As typical nonanal mixing ratios have 475 been observed or estimated to be < 500 pptv, this suggests that nonanal or similar OVOCs may 476 contribute to some of the missing reactivity ( $< 0.45 \text{ s}^{-1}$ ). Finally, nonanal and other long-chain 477 aldehydes may be an important higher PNs precursor (see Sect. 3.4 for more discussion about un-478 speciated higher PNs).

479 OVOC emissions from multiple sources, including solvent evaporation and other non-480 transportation emissions, are generally considered to be an important fraction of R(VOC) for urban 481 emissions but may not be measurable by PTR or GC, such as glycols (de Gouw et al., 2018; 482 Gkatzelis et al., 2021; McDonald et al., 2018; Ma et al., 2022; Simpson et al., 2020; Wang et al., 483 2022; Yang et al., 2022). However, the  $\alpha$  for OVOC is potentially smaller than alkanes, though it 484 is highly unconstrained (Orlando and Tyndall, 2012). Note, higher OVOCs have been understudied 485 and thus may have higher  $\alpha$  (e.g., nonanal). Thus, if the missing reactivity is mainly OVOCs and 486 it is assumed their  $\alpha$  is low, compounds with  $\alpha > 0.15$  will be needed for the budget closure shown 487 here. Likely compounds with high  $\alpha$  include alkanes, cycloalkenes/alkenes, and aromatics, though 488 the latter is also highly uncertain. Alkanes have typically been a small source for the R(VOC) in 489 urban environments (e.g., McDonald et al., 2018; Simpson et al., 2020; Whalley et al., 2016). 490 Though aromatics contribute a significant fraction of R(VOC) in different Asian urban

491	environments (Brune et al., 2022; Schroeder et al., 2020; Simpson et al., 2020; Whalley et al.,
492	2021), the majority of the aromatic $R(\mbox{VOC})$ is considered to be measured by WAS over SMA
493	during KORUS-AQ (e.g., measured aromatics account for ~81% of aromatic reactivity in
494	McDonald et al. (2018) and 98% of aromatic reactivity in Whalley et al. (2016), where both studies
495	had more complete VOC measurements). Finally, the cycloalkenes/alkenes originate from
496	numerous anthropogenic sources (e.g., McDonald et al., 2018; Simpson et al., 2020). One subclass
497	of cycloalkenes includes monoterpenes. Similar to the comparison of isoprene between PTR-MS
498	and WAS, the difference in monoterpenes between these two measurements increases with
499	increasing NOx (Figure S§6). As the interfering compound(s) measured by the PTR-MS and
500	whether they are oxygenated or not is not known, only the WAS monoterpenes are used in this
501	analysis of calculating R(VOC). Assuming the limonene rate constant, the difference between the
502	PTR-MS and WAS monoterpenes raises the terpene reactivity by $0.05 - 0.30 \text{ s}^{-1}$ . Though this does
503	not include any associated photochemical products from the oxidation of monoterpenes and can
504	improve the closure, it does not explain the total missing reactivity $(1.4 - 2.1 \text{ s}^{-1})$ . Thus, the missing
505	R(VOC) is most likely a combination of OVOCs and cycloalkenes/alkenes.

506

#### 507 3.4 Sources of PNs over SMA

508 As shown in Figure 2,  $\Sigma$ PNs account for a larger fraction of the total NO<sub>z</sub> budget than 509  $\Sigma$ ANs.  $\Sigma$ PNs are known to be a temporary sink of NO<sub>x</sub> and radicals (R(O)O<sub>2</sub><sup>-</sup>) due to their short 510 thermal lifetime (~1 hr). Thus, the NO<sub>x</sub> emitted in SMA is being transported regionally, impacting 511 the P(O<sub>x</sub>).

512 In Figure 3b, ΣPNs shows some correlation with formaldehyde. Both are secondary
513 products from the photooxidation of VOCs and have short lifetimes, leading to the correlation.

514	However, above 4 ppbv formaldehyde, the correlation shifts as $\Sigma$ PNs increases more rapidly than
515	formaldehyde. As shown in Figure S27, this change in the relationship between $\Sigma PNs$ versus
516	formaldehyde is due to changes in the competition in the reaction of the acyl peroxy radical
517	$(R(O)O_2^{-})$ between NO <sub>2</sub> and NO. At low NO-to-NO <sub>2</sub> ratios, R8 is more favorable, as the $R(O)O_2^{-}$
518	is more likely to react with NO <sub>2</sub> compared with NO, leading to more efficient production over
519	formaldehyde. As the NO-to-NO2 ratios increase (NO becomes comparable to NO2, leading to
520	more equal probability in R(O)O2- reacting to NO and NO2, leading to production of alkoxy
521	radicals that can form formaldehyde), R9 becomes more dominant, leading to less production of
522	PNsleading to more efficient production of PNs over formaldehyde. As NO to NO2 ratios
523	increase (NO becomes comparable to NO2), R9 becomes more dominant, leading to less
524	production of PNs.
525	To further explore the sources of both PAN and the higher $\Sigma$ PNs, the F0AM model (Wolfe
526	et al., 2016) was used to predict both unmeasured <b>SPNs</b> , constrained by the observed VOCs
527	precursors, PAN, and PPN (Table 2) and budget of all <b>SPNs</b> , where PAN and PPN were not
528	constrained. FOAM shows minimal bias in the predicted formaldehyde, NO <sub>2</sub> , and OH (Figure

S108) when PAN and PPN were constrained. As discussed in Sect. 3.3, though, there is missing 529 R(VOC) of  $1.7^{+1.1}_{-0.4}$ s<sup>-1</sup>. A sensitivity analysis in adding this missing reactivity to F0AM on 530 531 predicted OH and formaldehyde was conducted (Sect. S4 and Figure  $S_{119}^{19} - S_{129}^{20}$ ). Both OH and 532 formaldehyde are found to be buffered with the addition of this low amount of R(VOC). Thus, 533 though there is good agreement in these intermediate products between observation and FOAM, 534 this analysis for the sources of PAN and higher  $\Sigma$ PNs is expected to be a lower limit. This missing R(VOC) is further observed in the F0AM-predicted higher PNs (<u>SPNs\_without PAN, or SPNs-</u> 535 PAN for short) versus formaldehyde, as a general underestimation in the total higher PNs 536

537	compared to observations is observed (Figure 5a). PAN was excluded as FOAM overestimated the
538	mixing ratios of PAN by approximately a factor of 2 (Figure S8e). Note, F0AM also overpredicted
539	the PPN mixing ratios, but to a lesser extent than PAN (~50%; Figure S8f). The differences in
540	predicted versus observed PNs may be associated with assumed background, dilution, and/or
541	temperature used to reach steady-state (Schroeder et al., 2020). Thus, the results from FOAM will
542	provide qualitative insight into sources and chemistry that should be investigated to better
543	understand PN-chemistry in SMA.

544 The classes of compounds producing higher PNs in F0AM are shown in Figure 5b. The 545 classes of compounds were selected from the parent VOC which was oxidized into the higher PN 546 (Table S2). Individual PNs with high contributions and/or are typically measured (PPN, PBzN, 547 and MPAN (methacryloyl peroxy nitrate)) or are a large fraction of PNs but have yet to be 548 measured in ambient conditions (PHAN) are shown without any connection to the parent VOC. 549 Further, both PHAN and PPN have numerous precursors while many of the other higher PNs 550 modeled by F0AM are generally associated with one precursor. At high NO<sub>x</sub> mixing ratios, the 551 more reactive VOCs (aromatics, terpenes) contribute a large fraction of the higher PNs (>35% for  $NO_x > 4$  ppby). As the air moves away from SMA (lower NO<sub>x</sub> mixing ratios), contributions of 552 553 higher PNs from longer-lived compounds (e.g., alkanes) and later generation oxidation products 554 start dominating.

An interesting trend is observed for PPN and PHAN. Both peroxy acyl radicals for PPN and PHAN ( $C_2H_5C(O)O_2$  and  $CH_2(OH)C(O)O_2$ , respectively) are products from photooxidation of many VOCs, including aromatics, alkanes, and methyl ethyl ketone (MEK). However, the fractional contribution of PPN to higher PNs remains constant with decreasing NO<sub>x</sub> while the fractional contribution of PHAN increases with decreasing NO<sub>x</sub> (Figure 5b). This stems from the 560 sources of  $C_2H_5C(O)O_2$  versus  $CH_2(OH)C(O)O_2$ . The MCM mechanism, which is used for 561 F0AM, produces  $C_2H_5C(O)O_2$  from the photooxidation from both short- and long-lived species (isoprene, C8-aromatics, toluene, ethanol, MEK, propane, and C4-alkanes) while 562 563  $CH_2(OH)C(O)O_2$  is produced from the photooxidation of isoprene and ethene. For 564  $CH_2(OH)C(O)O_2$ , the production is through minor channels in the photooxidation of isoprene (~3% yield directly from isoprene and ~20% as a secondary product (Galloway et al., 2011)). 565 Ethene is relatively long-lived, with a lifetime  $\sim$ 7 hrs (OH = 5×10<sup>6</sup> molec. cm<sup>-3</sup>) leading to the 566 567 delay in the production of PHAN. Note, PHAN formation in MCM/F0AM may be overestimated, 568 as Butkovskaya et al. (2006) found that the radical formed from the photooxidation of 569 glycolaldehyde decomposes to form formaldehyde and CO<sub>2</sub>, potentially competing with the pathway to form PHAN. Other studies also found that PNs were not observed by photooxidation 570 571 of glycoaldehyde (Magneron et al., 2005).

572 The results here in general indicate more speciated measurements of higher PNs are 573 needed. However, as highlighted in Figure 5, improved detection-of or measurements of PBzN, 574 PHAN, and MPAN would allow for furthering our knowledge in PNs chemistry in urban 575 environments and their role in controlling O<sub>x</sub> production.

A qualitative investigation of the precursors of PAN predicted by F0AM are shown in Figure 5c. This provides a basis for further investigation of the sources over the SMA region for PAN as (a) F0AM over-predicts PAN, as noted above, (b) ethanol is currently estimated, similar to Schroeder et al. (2020), and (c) R(VOC) in F0AM is low due to missing precursors. Like the higher PNs, highly reactive R(VOC) contributes a large portion of the PAN budget at high NO<sub>x</sub>. The short-lived compounds contribute ~80% of PAN over SMA at the highest NO<sub>x</sub> mixing ratios. 582 At lower NO<sub>x</sub> mixing ratios, moving away from SMA, longer-lived compounds, such as ethanol,
583 contribute the most towards PAN production (~70%).

584 One of the interesting contributions not typically observed for PAN is MEK, which also contributes to PPN and PHAN. In prior studies, MEK mixing ratios were typically 0.5 to 2.0 ppbv 585 586 (Bon et al., 2011; de Gouw et al., 2018; Liu et al., 2015). Over the SMA, 1.5 ppbv of MEK was observed on average with values as high as 8.3 ppby. Due to the long lifetime of MEK (~30 hrs 587 588 for the average photolysis rate measured and  $OH = 5 \times 10^6$  molec. cm<sup>-3</sup>), the high mixing ratios of 589 MEK are most likely due to direct emissions (e.g., de Gouw et al., 2005; Liu et al., 2015). Thus, 590 there are potentially large sources of MEK in SMA that need to be considered in properly 591 representing PAN chemistry.

592 Another potentially important compound for PAN production is ethanol. However, this 593 compound was not measured during KORUS-AQ; instead, it was estimated based on previous 594 ground-based observations, similar to Schroeder et al. (2020). Ethanol is considered to mainly 595 come from both vehicle emissions (e.g., Millet et al., 2012) and potentially-non-transportation emissions, including cleaning agents and solvents (e.g., McDonald et al., 2018). As ethanol use is 596 597 predicted to increase in the future (e.g., de Gouw et al., 2012) and cleaning agents and other volatile 598 chemical products appear to scale with population (Gkatzelis et al., 2021), ethanol and MEK may 599 continue contributing towards the PAN budget in the SMA in the future.

As a note, two other compounds potentially important for PAN production that were not measured on the DC-8 during KORUS-AQ include methylglyoxal and biacetyl (LaFranchi et al., 2009). In a forested environment that was partially impacted by urban outflow, these two components contributed on average 25% of the PAN budget (LaFranchi et al., 2009). In urban environments, methylglyoxal is believed to mainly originate from aromatic oxidation (Ling et al., 2020); whereas, biacetyl is believed to come from anthropogenic emissions (Xu et al., 2023).
Further, as discussed in Sect. 4.3, these two compounds may potentially be important missing HO<sub>x</sub>
sources, as well. Thus, measurements of these two compounds along with ethanol is necessary to
better understand PAN chemistry.

609

#### 610 4. Observational constraints of the HO<sub>x</sub> and O<sub>x</sub> budget over SMA

As highlighted in Figure S1, the three factors impacting instantaneous  $P(O_x)$  are R(VOC), P(HO<sub>x</sub>), and NO<sub>x</sub> loss processes. In Sect. 3, the NO<sub>x</sub> loss processes were investigated and provided a constraint for R(VOC) to improve the investigation of  $P(O_x)$ . With R(VOC) constrained, the RO<sub>2</sub><sup>-</sup> concentration can be estimated, providing a means to calculate the net  $P(O_x)$  and to investigate the major reactions leading to O<sub>x</sub> loss and total HO<sub>x</sub> (OH + HO<sub>2</sub> + RO<sub>2</sub><sup>-</sup> +  $R(O)O_2^{-}$ ) loss. With the latter, this allows for an investigation of the major  $P(HO_x)$  reactions, assuming L(HO<sub>x</sub>) equals  $P(HO_x)$  (see Eq. 1 – 7 in Sect. 2.3).

618

#### 619 4.1 Net O<sub>x</sub> production and sources of O<sub>x</sub> loss

620 Using the total R(VOC) from Sect. 3.3 (Figure 4a), the net  $P(O_x)$  (Eq. 1 – 2) over SMA 621 during KORUS-AQ has been determined (Figure 6a). The net  $P(O_x)$  peaked at <u>9.340.3</u> ppbv hr<sup>-1</sup> 622 at ~8 ppbv NOx. If only the measured and estimated R(VOC) from F0AM secondary products is 623 used to calculate net  $P(O_x)$ , the value decreases to 7.68.8 ppbv hr<sup>-1</sup>, but at the same NO<sub>x</sub> mixing ratio. This value is similar to values observed in other urban locations around the world ( $\sim 2 - 20$ 624 625 ppbv  $hr^{-1}$ ), showing that many urban areas are still impacted by high P(O<sub>x</sub>) values (Brune et al., 626 2022; Griffith et al., 2016; Ma et al., 2022; Ren et al., 2013; Schroeder et al., 2020; Whalley et al., 627 2016, 2018).

628 The NO<sub>x</sub> distribution over SMA (Figure 1) shows a large area ( $\sim$ 127.53°E to 127.18°E, or 629 ~39 km), which corresponds to the is near the NO<sub>x</sub> mixing ratio that with results in the maximum 630  $P(O_x)$ , as shown in (Figure 6a). Thus, a large portion of the SMA will have high instantaneous  $P(O_x)$  of  $\sim 190$  ppbv hr<sup>-1</sup>. As the median wind speed over SMA during KORUS-AQ was  $\sim 5 \text{ m s}^{-1}$ , 631 632 an air parcel would remain at the highest P(O<sub>x</sub>) for ~2 hrs, leading to ~ $\frac{1820}{100}$  ppbv O<sub>3</sub> being 633 produced (not including dilution). This agrees with the  $\sim 20$  ppbv increase in O<sub>3</sub> observed over the 634 Taehwa Research Forest supersite between midday and afternoon overpasses by the DC-8 during 635 KORUS-AQ (Crawford et al., 2021). Thus, though there is a substantial O3 background observed 636 over SMA (Colombi et al., 2023; Crawford et al., 2021), a large contribution of the O<sub>3</sub> is due to 637 photochemical production.

638 The major reactions leading to  $O_x$  loss (L( $O_x$ )) are shown in Figure 6b. The two major 639 reactions that lead to O<sub>x</sub> loss are net R8 (light and dark red), or the net production of PNs (which 640 includes losses), and R11, reaction of NO2 with OH (blue) (see Table 1). Note, as discussed in 641 Sect. 2.2, for the budget analysis conducted here, PAN and PPN were constrained to observations. 642 At high NO<sub>x</sub> (near emissions,  $\sim$ 30 ppbv), R11 (NO<sub>2</sub> + OH) dominates the L(O<sub>x</sub>) budget (> 60%), with net R8 (net PAN, dark red, and higher PNs, light red) contributing  $\sim 25\%$ , and R12 – 14 643 644 accounting for the remaining 15% of Ox loss. As NOx mixing ratios decrease (moving away from 645 emissions), the net R8 reaction, producing both PAN and higher PNs, starts contributing to larger 646 total L(O<sub>x</sub>), ranging from 30 - 40%. Furthermore, the net R8 reaction contribution towards L(O<sub>x</sub>) 647 remains relatively constants with  $NO_x$  mixing ratios as the contribution from R11 (OH +  $NO_2$ ) 648 decreases. At NO<sub>x</sub> mixing ratios < 3 ppbv is when non-NO<sub>x</sub> reactions (R12 – 14) contribute greater 649 than 30% of the  $L(O_x)$  budget. Thus, proper representation of PAN and higher PNs, both in 650 precursors and speciation, is important in properly understanding the O<sub>x</sub> budget in SMA.

651

#### 652 **4.2 HO**<sub>x</sub> loss over the SMA

653 Similar to L(Ox), the major reactions leading to L(HOx) over the SMA during KORUS-AQ 654 were the reactions of NOx with HOx, specifically NO2 with OH (R11) and net PAN (dark red) and 655 higher PNs (light red) production (R8) (Figure 6c). Reaction R11 is most important for NOx mixing ratios greater than 15 ppbv (50 - 65%). Between 5 and 15 ppbv, R11 is comparable to the net PN 656 657 production (R8), where R11 comprises 35 - 50% of L(HO<sub>x</sub>) while net R8 (sum of higher  $\Sigma$ PNs 658 and PAN) comprises 30 - 40% of L(HOx). At lower NOx mixing ratios, R11 is always smaller for 659  $L(HO_x)$  than net R8, where R11 is about a factor of 2 lower than net R8. Production of  $\Sigma ANs$ 660 played a minor role due to the low  $\alpha_{eff}$ .

The self-reaction of HO<sub>x</sub> species (R15 – R16) contributes minimally to  $L(HO_x)$  (less than 10%) for NO<sub>x</sub> mixing ratios greater than 8 ppbv. At lower NO<sub>x</sub> mixing ratios, R16 starts dominating  $L(HO_x)$  budget, increasing from 8% at 8 ppbv to 50% of  $L(HO_x)$  at NO<sub>x</sub> mixing ratios less than 2 ppbv. Reaction R15 remains relatively small for the  $L(HO_x)$  budget, only reaching 7% of the  $L(HO_x)$  budget at NO<sub>x</sub> mixing ratios less than 2 ppbv.

666

#### 667 4.3 Sources of HOx over SMA

The analysis conducted leads to the ability to constrain HO<sub>x</sub> losses over the SMA during KORUS-AQ. This is important as not all typical HO<sub>x</sub> sources were measured on the DC-8 during the project (e.g., nitrous acid, or HONO), and HO<sub>x</sub> production rates directly impacts  $P(O_x)$  (Figure S1). Prior studies (e.g., Griffith et al., 2016; Tan et al., 2019; Whalley et al., 2018) have demonstrated that in urban environments, sources of HO<sub>x</sub> include photolysis of O<sub>3</sub> and subsequent reaction with water vapor, formaldehyde photolysis, and HONO photolysis. Furthermore, recent 674 studies have highlighted the potential importance of typically non-measured OVOCs in their contribution to  $P(HO_x)$  and subsequent  $P(O_x)$  in an urban environment (Wang et al., 2022). To 675 676 constrain the  $P(HO_x)$  over SMA during KORUS-AQ, the  $P(HO_x)$  was assumed to be equal to the 677 observationally constrained  $L(HO_x)$ . Then,  $P(HO_x)$  was calculated for the measurements on the 678 DC-8, including photolysis of O<sub>3</sub>, formaldehyde, H<sub>2</sub>O<sub>2</sub>, and other measured OVOCs (Table 2). Comparing the calculated  $P(HO_x)$  and  $L(HO_x)$ , ~1.5 ppbv hr<sup>-1</sup>  $P(HO_x)$  (range 1.3 – 1.8 ppbv 679 680 hr<sup>-1</sup>) is not accounted for, leading to ~45% of the necessary L(HO<sub>x</sub>) to maintain steady-state 681 (Figure 7). For the calculated  $P(HO_x)$  budget,  $O_3$  and formaldehyde photolysis contributed ~50%

and 40% of the budget, respectively, with the remainder coming from photolysis of  $H_2O_2$  and other measured OVOCs. Accounting for the unobserved P(HO<sub>x</sub>), O<sub>3</sub> and formaldehyde photolysis contributed ~25% and ~20%, respectively.

685 Potential missing sources of P(HO<sub>x</sub>) are briefly speculated here. First, one potential source 686 is the photolysis of methylglyoxal. Using the FOAM predicted methylglyoxal, as it was not 687 measured on the DC-8, methylglyoxal would contribute ~0.24 ppbv hr<sup>-1</sup> P(HO<sub>x</sub>), or ~16% of the 688 unobserved P(HO<sub>x</sub>). Another OVOC not measured on the DC-8 and expected to originate from 689 anthropogenic emissions and not from chemistry is 2,3-butanedione, or biacetyl (de Gouw et al., 690 2018; Grosjean et al., 2002; Schauer et al., 2002; Xu et al., 2023; Zhou et al., 2020). Prior studies 691 observed 20 - 400 pptv of biacetyl (de Gouw et al., 2018; Xu et al., 2023), correspond to 0.04 -692 0.74 ppbv hr<sup>-1</sup>, or 3-49% of the unobserved P(HO<sub>x</sub>). Thus, between these two OVOCs, 19-66%693 of the unobserved  $P(HO_x)$  could be explained. Other unmeasured OVOCs could potentially contribute to the observed P(HO<sub>x</sub>) (e.g., Wang et al., 2022); however, there is less constraints both 694 695 on the speciation and photolysis rates for these OVOCs (e.g., Mellouki et al., 2015). Finally, 696 HONO could contribute to this observed P(HOx). Up to 700 pptv of HONO was observed in SMA

697	during KORUS-AQ (Gil et al., 2021), though, this would quickly photolyze to the altitudes the
698	DC-8 flew over SMA (Tuite et al., 2021). Even at 50 – 100 pptv HONO, photolysis of HONO
699	would lead to $0.2 - 0.4$ ppbv hr <sup>-1</sup> P(HO <sub>x</sub> ), or $13 - 27\%$ of the unobserved P(HO <sub>x</sub> ). Thus, between
700	methylglyoxal, biacetyl, and HONO, between $32 - 92\%$ of the unobserved P(HO <sub>x</sub> ) could be
701	accounted for. This analysis highlights the importance of measuring these HO <sub>x</sub> sources to better
702	2 understand and constrain O <sub>x</sub> chemistry in SMA and other urban environments.
703	A comparison of HO <sub>x</sub> sources from F0AM is shown in Figure S14. As it has more complete
704	OVOCs than the observations, the contributions are different than shown in Figure 7. Both

observations and FOAM agree that photolysis of O3 and subsequent reaction with water (R12) and

photolysis of formaldehyde are the two largest sources of HOx. FOAM also shows that

707 methylglyoxal is an important source of HO<sub>x</sub>, which is not shown in Figure 7 as methylglyoxalwas not measured. However, the total FOAM P(HO<sub>x</sub>) was ~2.4 ppbv hr<sup>-1</sup>, which was lower than 708 709 the observationally constrained value. This further supports either potential unmeasured OVOCs 710 coming from both emissions and chemistry and/or uncertainty in the photolysis rate constants for 711 these OVOCs (e.g., Wang et al., 2022). 712 One note about this analysis is that particulate matter collected onto the downwelling CAFS 713 optics during KORUS-AQ (see Sect. S5, Table S3, and Figure S134). Corrections of up to 20% 714 were determined, and the associated uncertainties were also increased by 20% due to the 715 corrections. Thus, the exact amount of unmeasured  $P(HO_x)$  is potentially smaller than discussed. 716

#### 717 5. Conclusions and Implications

705

706

In the Seoul Metropolitan Area (SMA), the ozone (O<sub>3</sub>) mixing ratio often exceeds current
 standards and is increasing. Many processes can impact the O<sub>3</sub> mixing ratios and exceedances.

Here, the processes that impact instantaneous  $O_3$  production (P(O<sub>x</sub>), where  $O_x$  is  $O_3 + NO_2$  to account for possible  $O_3$  titration) were investigated for observations collected on the NASA DC-8 during the 2016 NIER/NASA Korea United-States Air Quality (KORUS-AQ) study. The observations indicate missing oxidized NO<sub>x</sub> products (NO<sub>z</sub>) that include both the short-lived peroxy nitrates ( $\Sigma$ PNs) and alkyl and multi-functional nitrates ( $\Sigma$ ANs).  $\Sigma$ PNs contributed the most for the organic NO<sub>z</sub> species. Only ~50% of the  $\Sigma$ PNs were speciated over SMA, which is atypical as prior studies typically show closure between the speciated and total PN measurements.

727 The un-speciated  $\Sigma PNs$  and  $\Sigma ANs$  were used to constrain the missing volatile organic 728 compound (VOC) reactivity (R(VOC)), as R(VOC) is important in constraining the instantaneous  $P(O_3)$ . The missing R(VOC) was found to be 1.4 to 2.1 s<sup>-1</sup>. The FOAM box model further supports 729 730 the role of unmeasured  $\Sigma$ PNs as an important temporary NO<sub>x</sub> and radical sink over SMA. F0AM 731 predicts ~50% of the higher  $\Sigma PNs$  (higher  $\Sigma PNs = \Sigma PNs - PAN$ ), indicating missing R(VOCs) 732 may explain the other 50%. Constraints from both the  $\Sigma PNs$  and  $\Sigma ANs$  suggest that this missing 733 R(VOC) would include oxygenated VOCs (OVOCs), including aldehydes such as octanal and 734 nonanal from cooking, and alkenes from anthropogenic emissions.

With the constraints on the R(VOC), the net instantaneous P( $O_x$ ) was determined for SMA. It was found to peak at ~10 ppbv hr<sup>-1</sup> at ~8 ppbv NO<sub>x</sub>. A large fraction of the SMA area was, on average, at this mixing ratio of NO<sub>x</sub>, indicating high local P(O<sub>x</sub>). This supports the increase of ~20 ppbv of O<sub>3</sub> observed in a downwind site (Taehwa Research Forest supersite) from midday to afternoon during KORUS-AQ.

740 With the comprehensive measurements on-board the DC-8, the F0AM model results, and 741 the observationally constrained R(VOC), a budget analysis on the sinks of O<sub>3</sub> (L(O<sub>x</sub>)) and HO<sub>x</sub> 742 (L(HO<sub>x</sub>), where HO<sub>x</sub> = OH + HO<sub>2</sub> + RO<sub>2</sub><sup>-</sup> + R(O)O<sub>2</sub><sup>-</sup>) was performed. Due to the high R(VOC),

33

743	type of VOC, and the NO <sub>2</sub> -to-NO ratio, net $\Sigma$ PNs production is surprisingly a large and important
744	sink of $O_x$ and $HO_x$ over SMA (~25 – 40% and 15 – 40% for $L(O_x)$ and $L(HO_x)$ , respectively),
745	with production of $HNO_3$ and radical self-reactions accounting for the other $L(O_x)$ and $L(HO_x)$
746	losses. Net $\Sigma$ PNs production as an important $L(O_x)$ and $L(HO_x)$ term is significant, as $\Sigma$ PNs is a
747	temporary reservoir of both NO <sub>2</sub> and $R(O)O_2$ but has not traditionally been included in these
748	calculations. Downwind locations separated from the local $\ensuremath{\text{NO}}_x$ and $\ensuremath{\text{VOC}}$ emissions of the SMA
749	will experience increased $P(O_x)$ due to the release of NO <sub>2</sub> and $R(O)O_2$ . With the constraint of
750	$L(HO_x)$ , $P(HO_x)$ was investigated, assuming steady-state, and unmeasured HONO plus
751	unmeasured OVOCs were found to be necessary to explain the missing $HO_x$ sources. Both sources
752	of $HO_x$ are either missing or highly uncertain in chemical transport models.

753 Though the high regional background and foreign sources of O<sub>3</sub> and its precursors elevate 754 the  $O_3$  levels in SMA and potentially already causes the SMA to be in exceedance for  $O_3$ concentrations, this study highlights the importance local, in-situ P(Ox) to the SMA area, which 755 756 can further exacerbate the O<sub>3</sub> concentrations for SMA and the surrounding region. The results 757 support the observations of increasing  $O_3$  with decreasing  $NO_x$  that has been observed for SMA in 758 prior studies. Further, the study highlights the important role of unmeasured VOCs and OVOCs and the necessity to understand their sources and role in NOx and O3 chemistry. Further, the study 759 760 demonstrates the interplay of direct emissions or secondary production of PN precursors and its 761 role in net P(Ox). Attempts at specifically reducing the sources of PN may adversely impact net 762 P(O<sub>x</sub>), as lower net PN chemistry may increase O<sub>3</sub> due to more NO<sub>2</sub> being available.

#### 763 Competing Interests

At least one of the (co-)authors is a member of the editorial board of Atmospheric Chemistry andPhysics.

## 766

#### 767 Acknowledgements

768 The authors acknowledge Michelle Kim, Alex Teng, John Crounse, and Paul O. Wennberg for 769 their measurements with CIT-CIMS (HNO3, multifunctional alkyl nitrates, and OVOCs), William 770 H. Brune for his measurements with ATHOS (OH, OH reactivity), Alan Fried for his 771 measurements with CAMS (CH2O and C2H6), Paul Romer-Present for his contribution to 772 collecting data with TD-LIF, Sally Pusede for her contributions to collecting data with DACOM 773 and DLH, and Andrew J. Weinheimer for his measurements of NO, O3, and NOy. The PTR-MS 774 instrument team (P. Eichler, L. Kaser, T. Mikoviny, M. Müller) are acknowledged for their 775 support.

776

#### 777 Funding

BAN and KRT acknowledge NASA grant 80NSSC22K0283. LGH and YL acknowledge NASA
grant NNX15AT90G for the PAN measurements. SRH and KU were supported by NASA grant
NNX15AT99G for photolysis measurements. AW acknowledges support by the Austrian Federal
Ministry for Transport, Innovation, and Technology (bmvit-FFG-ASA) for the PTR-MAS
measurements. PCJ and JLJ were supported by NASA 80NSSC21K1451 and 80NSSC23K0828.

783

784 Data Availability

785 Version R6 1-min merged data in this analysis available used at DOI:10.5067/Suborbital/KORUSAQ/DATA01. The F0AM setup file, input file, and output files 786 787 are all available at https://doi.org/10.5281/zenodo.10723227.

788

#### 789 Author Contribution

- 790 BAN, KRT, and JHC designed the experiment and wrote the paper. BAN and KRT analyzed the
- 791 data. KRT ran the F0AM model and KRT and BAN analyzed the model output. BAN, DRB, PCJ,
- 792 RCC, JPD, GSD, SRH, LGH, JLJ, K-EK, YL, IJS, KU, and AW collected and QA/QC the data
- 793 during KORUS-AQ. All authors contributed to the writing and editing of the paper.

#### 794 References

- 795 Archibald, A. T., Neu, J. L., Elshorbany, Y. F., Cooper, O. R., Young, P. J., Akiyoshi, H., Cox, R.
- 796 A., Coyle, M., Derwent, R. G., Deushi, M., Finco, A., Frost, G. J., Galbally, I. E., Gerosa, G.,
- 797 Granier, C., Griffiths, P. T., Hossaini, R., Hu, L., Jöckel, P., Josse, B., Lin, M. Y., Mertens, M.,
- 798 Morgenstern, O., Naja, M., Naik, V., Oltmans, S., Plummer, D. A., Revell, L. E., Saiz-Lopez, A.,
- 799 Saxena, P., Shin, Y. M., Shahid, I., Shallcross, D., Tilmes, S., Trickl, T., Wallington, T. J., Wang,
- 800 T., Worden, H. M., and Zeng, G.: Tropospheric ozone assessment report: A critical review of
- k1 changes in the tropospheric ozone burden and budget from 1850 to 2100,
  k1tps://doi.org/10.1525/elementa.2020.034, 2020.
- 803 Atkinson, R.: Kinetics of the gas-phase reactions of OH radicals with alkanes and cycloalkanes,
- 804 Atmos Chem Phys, 3, 2233–2307, https://doi.org/10.5194/acp-3-2233-2003, 2003.
- Atkinson, R. and Arey, J.: Atmospheric Degradation of Volatile Organic Compounds, Chem Rev,
  103, 4605–4638, https://doi.org/10.1021/CR0206420, 2003.
- 807 Atkinson, R., Baulch, D. L., Cox, R. A., Crowley, J. N., Hampson, R. F., Hynes, R. G., Jenkin, M.
- 808 E., Rossi, M. J., and Troe, J.: Evaluated kinetic and photochemical data for atmospheric chemistry:
- Volume II gas phase reactions of organic species, Atmos Chem Phys, 6, 3625–4055,
  https://doi.org/10.5194/acp-6-3625-2006, 2006.
- 811 Bohn, B. and Zetzsch, C.: Kinetics and mechanism of the reaction of OH with the
- trimethylbenzenes experimental evidence for the formation of adduct isomers, Physical
  Chemistry Chemical Physics, 14, 13933, https://doi.org/10.1039/c2cp42434g, 2012.
- 814 Bon, D. M., Ulbrich, I. M., De Gouw, J. a., Warneke, C., Kuster, W. C., Alexander, M. L., Baker,
- 815 a., Beyersdorf, a. J., Blake, D., Fall, R., Jimenez, J. L., Herndon, S. C., Huey, L. G., Knighton, W.
- 816 B., Ortega, J., Springston, S., and Vargas, O.: Measurements of volatile organic compounds at a
- 817 suburban ground site (T1) in Mexico City during the MILAGRO 2006 campaign: measurement

- 818 comparison, emission ratios, and source attribution, Atmos Chem Phys, 11, 2399–2421,
- 819 https://doi.org/10.5194/acp-11-2399-2011, 2011.
- 820 Bourgeois, I., Peischl, J., Neuman, J. A., Brown, S. S., Allen, H. M., Campuzano-jost, P., Coggon,
- 821 M. M., Digangi, J. P., Diskin, G. S., Gilman, J. B., Gkatzelis, G. I., Guo, H., Halliday, H. A.,
- 822 Hanisco, T. F., Holmes, C. D., Nault, B. A., Nowak, J. B., Pagonis, D., Rickly, P. S., Robinson,
- 823 M. A., Veres, P. R., Warneke, C., Wennberg, P. O., Washenfelder, R. A., and Wiggins, E. B.:
- 824 Comparison of airborne measurements of NO, NO<sub>2</sub>, HONO, NO<sub>y</sub>, and CO during FIREX-AQ,
- 825 Atmos Meas Tech, 15, 4901–4930, https://doi.org/10.5194/amt-15-4901-2022, 2022.
- 826 Brune, W. H., Miller, D. O., Thames, A. B., Allen, H. M., Apel, E. C., Blake, D. R., and Bui, T.
- 827 P.: Exploring Oxidation in the Remote Free Troposphere: Insights From Atmospheric
- Tomography (ATom), Journal of Geophysical Research : Atmospheres, 125, c2019JD031685,
  https://doi.org/10.1029/2019JD031685, 2019.
- 830 Brune, W. H., Miller, D. O., Thames, A. B., Brosius, A. L., Barletta, B., Blake, D. R., Blake, N.
- 831 J., Chen, G., Choi, Y., Crawford, J. H., Digangi, J. P., Diskin, G., Fried, A., Hall, S. R., Hanisco,
- 832 T. F., Huey, G. L., Hughes, S. C., Kim, M., Meinardi, S., Montzka, D. D., Pusede, S. E., Schroeder,
- 833 J. R., Teng, A., Tanner, D. J., Ullmann, K., Walega, J., Weinheimer, A., Wisthaler, A., and
- 834 Wennberg, P. O.: Observations of atmospheric oxidation and ozone production in South Korea,
- 835 Atmos Environ, 269, 118854, https://doi.org/10.1016/j.atmosenv.2021.118854, 2022.
- 836 Burkholder, J. B., Sander, S. P., Abbatt, J. P. D., Barker, J. R., Cappa, C. D., Crounse, J. D., Dibble,
- 837 T. S., Huie, R. E., Kolb, C. E., Kurylo, M. J., Orkin, V. L., Percival, C. J., Wilmouth, D. M., and
- 838 Wine, P. H.: Chemical Kinetics and Photochemical Data for Use in Atmospheric Studies,
- 839 Evaluation No. 19, Pasadena, CA, USA, 2020.

- 840 Butkovskaya, N. I., Pouvesle, N., Kukui, A., and Le Bras, G.: Mechanism of the OH-Initiated
- 841 Oxidation of Glycolaldehyde over the Temperature Range 233–296 K, J Phys Chem A, 110,
- 842 13492–13499, https://doi.org/10.1021/jp064993k, 2006.
- 843 Coggon, M. M., Stockwell, C. E., Claflin, M. S., Pfannerstill, E. Y., Lu, X., Gilman, J. B.,
- 844 Marcantonio, J., Cao, C., Bates, K., Gkatzelis, G. I., Lamplugh, A., Katz, E. F., Arata, C., Apel, E.
- 845 C., Hornbrook, R. S., Piel, F., Majluf, F., Blake, D. R., Wisthaler, A., Canagaratna, M., Lerner, B.
- 846 M., Goldstein, A. H., Mak, J. E., and Warneke, C.: Identifying and correcting interferences to
- 847 PTR-ToF- MS measurements of isoprene and other urban volatile organic compounds, Atmos
- 848 Meas Tech, 17, 801–825, https://doi.org/10.5194/amt-17-801-2024, 2024.
- 849 Cohen, A. J., Brauer, M., Burnett, R., Anderson, H. R., Frostad, J., Estep, K., Balakrishnan, K.,
- 850 Brunekreef, B., Dandona, L., Dandona, R., Feigin, V., Freedman, G., Hubbell, B., Jobling, A.,
- 851 Kan, H., Knibbs, L., Liu, Y., Martin, R., Morawska, L., Pope, C. A., Shin, H., Straif, K., Shaddick,
- 852 G., Thomas, M., van Dingenen, R., van Donkelaar, A., Vos, T., Murray, C. J. L., and Forouzanfar,
- 853 M. H.: Estimates and 25-year trends of the global burden of disease attributable to ambient air
- pollution: an analysis of data from the Global Burden of Diseases Study 2015, The Lancet, 389,
- 855 1907–1918, https://doi.org/10.1016/S0140-6736(17)30505-6, 2017.
- 856 Colombi, N. K., Jacob, D. J., Yang, L. H., Zhai, S., Shah, V., Grange, S. K., Yantosca, R. M., Kim,
- 857 S., and Liao, H.: Why is ozone in South Korea and the Seoul metropolitan area so high and
- 858 increasing?, Atmos Chem Phys, 23, 4031–4044, https://doi.org/10.5194/acp-23-4031-2023, 2023.
- 859 Crawford, J. H., Ahn, J. Y., Al-Saadi, J., Chang, L., Emmons, L. K., Kim, J., Lee, G., Park, J. H.,
- 860 Park, R. J., Woo, J. H., Song, C. K., Hong, J. H., Hong, Y. D., Lefer, B. L., Lee, M., Lee, T., Kim,
- 861 S., Min, K. E., Yum, S. S., Shin, H. J., Kim, Y. W., Choi, J. S., Park, J. S., Szykman, J. J., Long,
- 862 R. W., Jordan, C. E., Simpson, I. J., Fried, A., Dibb, J. E., Cho, S. Y., and Kim, Y. P.: The Korea-

- 863 United States air quality (KORUS-AQ) field study, Elementa, 9, 1–27,
  864 https://doi.org/10.1525/elementa.2020.00163, 2021.
- Crounse, J., McKinney, K. A., Kwan, A. J., and Wennberg, P. O.: Measurement of gas-phase
  hydroperoxides by chemical ionization mass spectrometry, Anal Chem, 78, 6726–6732,
  https://doi.org/doi:10.1021/ac0604235, 2006.
- Day, D. A., Wooldridge, P. J., Dillon, M. B., Thornton, J. A., and Cohen, R. C.: A thermal
  dissociation laser-induced fluorescence instrument for in situ detection of NO<sub>2</sub>, peroxy nitrates,
  alkyl nitrates, and HNO<sub>3</sub>, Journal of Geophysical Research-Atmospheres, 107, 4046,
  https://doi.org/10.1029/2001JD000779, 2002.
- 872 Day, D. A., Campuzano-Jost, P., Nault, B. A., Palm, B. B., Hu, W., Guo, H., Wooldridge, P. J.,
- 873 Cohen, R. C., Docherty, K. S., Huffman, J. A., De Sá, S. S., Martin, S. T., and Jimenez, J. L.: A
- 874 systematic re-evaluation of methods for quantification of bulk particle-phase organic nitrates using
- 875 real-Time aerosol mass spectrometry, Atmos Meas Tech, 15, 459–483,
  876 https://doi.org/10.5194/amt-15-459-2022, 2022.
- 877 Diskin, G. S., Podolske, J. R., Sachse, G. W., and Slate, T. A.: Open-path airborne tunable diode
- 878 laser hygrometer, in: Diode Lasers and Applications in Atmospheric Sensing, edited by: Fried, A.,
- 879 Proceedings of the Society of Photo-Optical Instrumentation Engineers (SPIE), 4817, 196–204,
- 880 https://doi.org/doi:10.1117/12.453736, 2002.
- 881 Espada, C. and Shepson, P. B.: The production of organic nitrates from atmospheric oxidation of
- 882 ethers and glycol ethers, Int J Chem Kinet, 37, 686–699, https://doi.org/10.1002/kin.20121, 2005.
- 883 Faloona, I. C., Tan, D., Lesher, R. L., Hazen, N. L., Frame, C. L., Simpas, J. B., Harder, H.,
- 884 Martinez, M., Di Carlo, P., Ren, X., and Brune, W. H.: A Laser-induced Fluorescence Instrument

- for Detecting Tropospheric OH and HO<sub>2</sub>: Characteristics and Calibration, J Atmos Chem, 47,
  139–167, https://doi.org/10.1023/B:JOCH.0000021036.53185.0e, 2004.
- 887 Farmer, D. K., Perring, A. E., Wooldridge, P. J., Blake, D. R., Baker, A., Meinardi, S., Huey, L. G., Tanner, D., Vargas, O., and Cohen, R. C.: Impact of organic nitrates on urban ozone 888 889 production, Atmos Chem Phys, 11, 4085–4094, https://doi.org/10.5194/acp-11-4085-2011, 2011. 890 Fisher, J. A., Jacob, D. J., Travis, K. R., Kim, P. S., Marais, E. A., Miller, C. C., Yu, K., Zhu, L., 891 Yantosca, R. M., Sulprizio, M. P., Mao, J., Wennberg, P. O., Crounse, J. D., Teng, A. P., Nguyen, 892 T. B., Clair, J. M. S., Cohen, R. C., Romer, P., Nault, B. A., Wooldridge, P. J., Jimenez, J. L., 893 Campuzano-Jost, P., Day, D. A., Hu, W., Shepson, P. B., Xiong, F., Blake, D. R., Goldstein, A. 894 H., Misztal, P. K., Hanisco, T. F., Wolfe, G. M., Ryerson, T. B., Wisthaler, A., and Mikoviny, T.: 895 Organic nitrate chemistry and its implications for nitrogen budgets in an isoprene- and 896 monoterpene-rich atmosphere: Constraints from aircraft (SEAC<sup>4</sup>RS) and ground-based (SOAS) 897 observations in the Southeast US, Atmos Chem Phys, 16, https://doi.org/10.5194/acp-16-5969-898 2016, 2016. 899 Fried, A., Walega, J., Weibring, P., Richter, D., Simpson, I. J., Blake, D. R., Blake, N. J., Meinardi,
- 900 S., Barletta, B., Hughes, S. C., Crawford, J. H., Diskin, G., Barrick, J., Hair, J., Fenn, M.,
- 901 Wisthaler, A., Mikoviny, T., Woo, J., Park, M., Kim, J., Min, K., Jeong, S., Wennberg, P. O., Kim,
- 902 M. J., Crounse, J. D., Teng, A. P., Bennett, R., Yang-martin, M., Shook, M. A., Huey, G., Tanner,
- 903 D., Knote, C., and Kim, J.: Airborne formaldehyde and volatile organic compound measurements
- 904 over the Daesan petrochemical complex on Korea's northwest coast during the Korea-United
- 905 States Air Quality study: Estimation of emission fluxes and effects on air quality, Elementa:
- 906 Science of the Anthropocene, 8, 1, https://doi.org/10.1525/elementa.2020.121, 2020.

- 907 Galloway, M. M., Huisman, A. J., Yee, L. D., Chan, A. W. H., Loza, C. L., Seinfeld, J. H., and
- 908 Keutsch, F. N.: Yields of oxidized volatile organic compounds during the OH radical initiated
- 909 oxidation of isoprene, methyl vinyl ketone, and methacrolein under high-NO<sub>x</sub> conditions, Atmos
- 910 Chem Phys, 11, https://doi.org/10.5194/acp-11-10779-2011, 2011.
- 911 Gaudel, A., Cooper, O. R., Ancellet, G., Barret, B., Boynard, A., Burrows, J. P., Clerbaux, C.,
- 912 Coheur, P. F., Cuesta, J., Cuevas, E., Doniki, S., Dufour, G., Ebojie, F., Foret, G., Garcia, O.,
- 913 Granados-Muñoz, M. J., Hannigan, J. W., Hase, F., Hassler, B., Huang, G., Hurtmans, D., Jaffe,
- 914 D., Jones, N., Kalabokas, P., Kerridge, B., Kulawik, S., Latter, B., Leblanc, T., Le Flochmoën, E.,
- 915 Lin, W., Liu, J., Liu, X., Mahieu, E., McClure-Begley, A., Neu, J. L., Osman, M., Palm, M.,
- 916 Petetin, H., Petropavlovskikh, I., Querel, R., Rahpoe, N., Rozanov, A., Schultz, M. G., Schwab,
- 917 J., Siddans, R., Smale, D., Steinbacher, M., Tanimoto, H., Tarasick, D. W., Thouret, V.,
- 918 Thompson, A. M., Trickl, T., Weatherhead, E., Wespes, C., Worden, H. M., Vigouroux, C., Xu,
- 919 X., Zeng, G., and Ziemke, J.: Tropospheric Ozone Assessment Report: Present-day distribution
- 920 and trends of tropospheric ozone relevant to climate and global atmospheric chemistry model
- 921 evaluation, Elementa, 6, https://doi.org/10.1525/elementa.291, 2018.
- 922 Gil, J., Kim, J., Lee, M., Lee, G., Ahn, J., Soo, D., Jung, J., Cho, S., Whitehill, A., Szykman, J.,
- and Lee, J.: Characteristics of HONO and its impact on  $O_3$  formation in the Seoul Metropolitan
- Area during the Korea-US Air Quality study, Atmos Environ, 247, 118182,
  https://doi.org/10.1016/j.atmosenv.2020.118182, 2021.
- 926 Gkatzelis, G. I., Coggon, M. M., Mcdonald, B. C., Peischl, J., Aikin, K. C., Gilman, J. B., Trainer,
- 927 M., and Warneke, C.: Identifying Volatile Chemical Product Tracer Compounds in U.S. Cities,
- 928 Environmental Science & Technology, 55, 188–199, https://doi.org/10.1021/acs.est.0c05467, 2021.

- 929 González-Sánchez, J. M., Brun, N., Wu, J., Ravier, S., and Clément, J.: On the importance of
- 930 multiphase photolysis of organic nitrates on their global atmospheric removal, Atmos Chem Phys,
- 931 23, 5851–5866, https://doi.org/10.5194/acp-23-5851-2023, 2023.
- 932 de Gouw, J. A., Middlebrook, A. M., Warneke, C., Goldan, P. D., Kuster, W. C., Roberts, J. M.,
- 933 Fehsenfeld, F. C., Worsnop, D. R., Canagaratna, M. R., Pszenny, A. A. P., Keene, W. C.,
- 934 Marchewka, M. L., Bertman, S. B., and Bates, T. S.: Budget of organic carbon in a polluted
- 935 atmosphere: Results from the New England Air Quality Study in 2002, Journal of Geophysical
- 936 Research: Atmospheres, 110, D16305, https://doi.org/10.1029/2004JD005623, 2005.
- 937 de Gouw, J. A., Gilman, J. B., Borbon, A., Warneke, C., Kuster, W. C., Goldan, P. D., Holloway,
- 938 J. S., Peischl, J., Ryerson, T. B., Parrish, D. D., Gentner, D. R., Goldstein, A. H., and Harley, R.
- A.: Increasing atmospheric burden of ethanol in the United States, Geophys Res Lett, 39, L15803,
  https://doi.org/10.1029/2012GL052109, 2012.
- de Gouw, J. A., Gilman, J. B., Kim, S.-W., Alvarez, S. L., Dusanter, S., Graus, M., Griffith, S. M.,
- 942 Isaacman-VanWertz, G., Kuster, W. C., Lefer, B. L., Lerner, B. M., McDonald, B. C.,
- 943 Rappenglück, B., Roberts, J. M., Stevens, P. S., Stutz, J., Thalman, R., Veres, P. R., Volkamer, R.,
- 944 Warneke, C., Washenfelder, R. A., and Young, C. J.: Chemistry of Volatile Organic Compounds
- 945 in the Los Angeles Basin: Formation of Oxygenated Compounds and Determination of Emission
- 946 Ratios, Journal of Geophysical Research: Atmospheres, 123, 2298–2319,
- 947 https://doi.org/10.1002/2017JD027976, 2018.
- 948 Griffith, S. M., Hansen, R. F., Dusanter, S., Michoud, V., Gilman, J. B., Kuster, W. C., Veres, P.
- 949 R., Graus, M., de Gouw, J. A., Roberts, J., Young, C., Washenfelder, R., Brown, S. S., Thalman,
- 950 R., Waxman, E., Volkamer, R., Tsai, C., Stutz, J., Flynn, J. H., Grossberg, N., Lefer, B., Alvarez,
- 951 S. L., Rappenglueck, B., Mielke, L. H., Osthoff, H. D., and Stevens, P. S.: Measurements of

- 952 hydroxyl and hydroperoxy radicals during CalNex-LA: Model comparisons and radical budgets,
- 953 Journal of Geophysical Research: Atmospheres, 121, 4211–4232,
  954 https://doi.org/10.1002/2015JD024358, 2016.
- Grosjean, D., Grosjean, E., and Gertler, A. W.: On-Road Emissions of Carbonyls from Light-Duty
  and Heavy-Duty Vehicles, Environmental Science & Technology, 35, 45–53,
- 957 https://doi.org/10.1021/es001326a, 2002.
- 958 Hansen, R. F., Griffith, S. M., Dusanter, S., Gilman, J. B., Graus, M., Kuster, W. C., Veres, P. R.,
- 959 de Gouw, J. A., Warneke, C., Washenfelder, R. A., Young, C. J., Brown, S. S., Alvarez, S. L.,
- 960 Flynn, J. H., Grossberg, N. E., Lefer, B., Rappenglueck, B., and Stevens, P. S.: Measurements of
- 961 Total OH Reactivity During CalNex-LA, Journal of Geophysical Research: Atmospheres, 126,
- 962 e2020JD032988, https://doi.org/10.1029/2020JD032988, 2021.
- Hurst Bowman, J., Barket, D. J., and Shepson, P. B.: Atmospheric chemistry of nonanal,
  Environmental Science & Technology, 37, 2218–2225, 2003.
- 965 Jenkin, M. E., Young, J. C., and Rickard, A. R.: The MCM v3.3.1 degradation scheme for isoprene,
- 966 Atmos Chem Phys, 15, 11433–11459, https://doi.org/10.5194/acp-15-11433-2015, 2015.
- 967 Jo, D. S., Emmons, L. K., Callaghan, P., Tilmes, S., and Woo, J.: Comparison of Urban Air Quality
- 968 Simulations During the KORUS-AQ Campaign With Regionally Refined Versus Global Uniform
- 969 Grids in the Multi-Scale Infrastructure for Chemistry and Aerosols (MUSICA) Version 0, J Adv
- 970 Model Earth Syst, 15, e2022MS003458, https://doi.org/10.1029/2022MS003458, 2023.
- 971 Jordan, C. E., Crawford, J. H., Beyersdorf, A. J., Eck, T. F., Halliday, H. S., Nault, B. A., Chang,
- 972 L. S., Park, J. S., Park, R., Lee, G., Kim, H., Ahn, J. Y., Cho, S., Shin, H. J., Lee, J. H., Jung, J.,
- 973 Kim, D. S., Lee, M., Lee, T., Whitehill, A., Szykman, J., Schueneman, M. K., Campuzano-Jost,
- 974 P., Jimenez, J. L., DiGangi, J. P., Diskin, G. S., Anderson, B. E., Moore, R. H., Ziemba, L. D.,

- 975 Fenn, M. A., Hair, J. W., Kuehn, R. E., Holz, R. E., Chen, G., Travis, K., Shook, M., Peterson, D.
- 976 A., Lamb, K. D., and Schwarz, J. P.: Investigation of factors controlling PM2.5 variability across
- 977 the South Korean Peninsula during KORUS-AQ, Elementa, 8,
  978 https://doi.org/10.1525/elementa.424, 2020.
- Kabir, M., Jagiella, S., and Zabel, F.: Thermal Stability of n-Acyl Peroxynitrates, Int J Chem Kinet,
  46, 462–469, https://doi.org/10.1002/kin.20862, 2014.
- 981 Kenagy, H. S., Sparks, T. L., Ryerson, T. B., Blake, D. R., and Cohen, R. C.: Evidence of
- 982 Nighttime Production of Organic Nitrates During SEAC<sup>4</sup>RS, FRAPPE, and KORUS-AQ,
- 983 Geophys Res Lett, 47, e2020GL087860, https://doi.org/10.1029/2020GL087860, 2020.
- 984 Kenagy, H. S., Romer Present, P. S., Wooldridge, P. J., Nault, B. A., Campuzano-Jost, P., Day, D.
- 985 A., Jimenez, J. L., Zare, A., Pye, H. O. T., Yu, J., Song, C. H., Blake, D. R., Woo, J. H., Kim, Y.,
- 986 and Cohen, R. C.: Contribution of Organic Nitrates to Organic Aerosol over South Korea during
- 87 KORUS-AQ, Environ Sci Technol, 55, 16326–16338, https://doi.org/10.1021/acs.est.1c05521,
  88 2021.
- 989 Kim, D., Cho, C., Jeong, S., Lee, S., Nault, B. A., Campuzano-jost, P., Day, D. A., Schroder, J.
- 990 C., Jimenez, J. L., Volkamer, R., Pusede, S. E., Hall, S. R., Ullmann, K., Huey, L. G., Tanner, D.
- 991 J., and Dibb, J.: Field observational constraints on the controllers in glyoxal (CHOCHO) reactive
- 992 uptake to aerosol, Atmos Chem Phys, 22, 805–821, https://doi.org/10.5194/acp-22-805-2022,
  993 2022a.
- Kim, H., Zhang, Q., and Heo, J.: Influence of Intense secondary aerosol formation and long-range
  transport on aerosol chemistry and properties in the Seoul Metropolitan Area during spring time :
  Results from KORUS-AQ, Atmos. Chem. Phys., 18, 7149–7168, https://doi.org/10.51944/acp2017-947, 2018.

- 998 Kim, H., Park, R. J., Kim, S., Brune, W. H., Diskin, G. S., Fried, A., Hall, S. R., Weinheimer, A.
- 999 J., Wennberg, P., Wisthaler, A., Blake, D. R., and Ullmann, K.: Observed versus simulated OH
- 1000 reactivity during KORUS-AQ campaign: Implications for emission inventory and chemical
- 1001 environment in East Asia, Elementa, 10, 1–26, 2022b.
- 1002 Kim, J., Lee, J., Han, J., Choi, J., Kim, D.-G., Park, J., and Lee, G.: Long-term Assessment of
- 1003 Ozone Nonattainment Changes in South Korea Compared to US, and EU Ozone Guidelines, Asian
- 1004 Journal of Atmospheric Environment, 15, 20–32, https://doi.org/10.5572/ajae.2021.098, 2021.
- 1005 Kim, S., Huey, L. G., Stickel, R. E., Tanner, D. J., Crawford, J. H., Olson, J. R., Chen, G., Brune,
- 1006 W. H., Ren, X., Lesher, R., Wooldridge, P. J., Bertram, T. H., Perring, A., Cohen, R. C., Lefer, B.
- 1007 L., Shetter, R. E., Avery, M., Diskin, G., and Sokolik, I.: Measurement of HO2NO2 in the free
- troposphere during the Intercontinental Chemical Transport Experiment–North America 2004,
  Journal of Geophysical Research: Atmospheres, 112, D12S01,
  https://doi.org/10.1029/2006JD007676, 2007.
- 1011 Kim, S., Sanchez, D., Wang, M., Seco, R., Jeong, D., Hughes, S., Barletta, B., Blake, D. R., Jung,
- 1012 J., Kim, D., Lee, G., Lee, M., Ahn, J., Lee, S.-D., Cho, G., Sung, M.-Y., Lee, Y.-H., Kim, D. B.,
- 1013 Kim, Y., Woo, J.-H., Jo, D., Park, R., Park, J.-H., Hong, Y.-D., Hong, J.-H., Zhang, D. Y., Liu, J.
- 1014 J., Li, B. J., Davis, D. L., Bell, M. L., Fletcher, T., Haagen-Smit, A. J., Blacet, F. E., Edinger, J.
- 1015 G., Yum, S. S., Roberts, G., Kim, J. H., Song, K. Y., Kim, D. Y., Lim, Y. J., Armendariz, A., Son,
- 1016 Y. S., Kim, J. C., Kim, S., Kim, S. Y., Lee, M., Shim, H., Wolfe, G. M., Guenther, A. B., He, A.,
- 1017 Hong, Y., Han, J., Kim, S., Lee, M., Kim, S., Choi, S., Seok, S., Kim, S., Kim, S. Y., Jiang, X. Y.,
- 1018 Lee, M., Turnipseed, A., Guenther, A., Kim, J. C., Lee, S. J., Kim, S., Lee, K. Y., Kwak, K. H.,
- 1019 Ryu, Y. H., Lee, S. H., Baik, J. J., Ryu, Y. H., Baik, J. J., Kwak, K. H., Kim, S., Moon, N., Bao,
- 1020 H., Shrestha, K. L., Kondo, A., Kaga, A., Inoue, Y., Ran, L., Zhao, C. S., Xu, W. Y., Lu, X. Q.,

- 1021 Han, M., Lin, W. L., Yan, P., Xu, X. B., Deng, Z. Z., Ma, N., Liu, P. F., Yu, J., Liang, W. D.,
- 1022 Chen, L. L., Geng, F., Tie, X., Guenther, A., Li, G., et al.: OH reactivity in urban and suburban
- 1023 regions in Seoul, South Korea an East Asian megacity in a rapid transition, Faraday Discuss.,
- 1024 189, 231–251, https://doi.org/10.1039/C5FD00230C, 2016.
- 1025 KORUS-AQ Science Team: KORUS-AQ Data, [Dataset], NASA Langley Research Center.,
  1026 https://doi.org/10.5067/Suborbital/KORUSAQ/DATA01, 2023.
- 1027 LaFranchi, B. W., Wolfe, G. M., Thornton, J. a., Harrold, S. a., Browne, E. C., Min, K. E.,
- 1028 Wooldridge, P. J., Gilman, J. B., Kuster, W. C., Goldan, P. D., DeGouw, J. a., McKay, M.,
- 1029 Goldstein, a. H., Ren, X. R., Mao, J. Q., Cohen, R. C., de Gouw, J. a., Welsh-Bon, D., Chen, Z.,
- 1030 and Brune, W. H.: Closing the peroxy acetyl (PA) radical budget: Observations of acyl peroxy
- 1031 nitrates (PAN, PPN and MPAN) during BEARPEX 2009, Abstracts of Papers of the American
- 1032 Chemical Society, 9, 289, https://doi.org/10.5194/acp-9-7623-2009, 2009.
- 1033 Lee, Y. R., Huey, L. G., Tanner, D. J., Takeuchi, M., Qu, H., Liu, X., Ng, N. L., Crawford, J. H.,
- 1034 Fried, A., Richter, D., Simpson, I. J., Blake, D. R., Blake, N. J., Meinardi, S., Kim, S., Diskin, G.
- 1035 S., Digangi, J. P., Choi, Y., Pusede, S. E., Wennberg, P. O., Kim, M. J., Crounse, J. D., Teng, A.
- 1036 P., Cohen, R. C., Romer, P. S., Brune, W., Wisthaler, A., Mikoviny, T., Jimenez, J. L.,
- 1037 Campuzano-Jost, P., Nault, B. A., Weinheimer, A., Hall, S. R., and Ullmann, K.: An investigation
- 1038 of petrochemical emissions during KORUS-AQ: Ozone production, reactive nitrogen evolution,
- and aerosol production, Elementa, 10, 1–24, https://doi.org/10.1525/elementa.2022.00079, 2022.
- 1040 Ling, Z., Xie, Q., Shao, M., Wang, Z., Wang, T., Guo, H., and Wang, X.: Formation and sink of
- 1041 glyoxal and methylglyoxal in a polluted subtropical environment: Observation-based
- 1042 photochemical analysis and impact evaluation, Atmos Chem Phys, 20, 11451-11467,
- 1043 https://doi.org/10.5194/acp-20-11451-2020, 2020.

- 1044 Liu, Y., Yuan, B., Li, X., Shao, M., Lu, S., Li, Y., Chang, C., Wang, Z., Hu, W., Huang, X., He,
- 1045 L., Zeng, L., Hu, M., and Zhu, T.: Impact of pollution controls in Beijing on atmospheric
- 1046 oxygenated volatile organic compounds ( OVOCs ) during the 2008 Olympic Games : observation
- and modeling implications, Atmos Chem Phys, 15, 3045–3062, https://doi.org/10.5194/acp-153045-2015, 2015.
- 1049 Lyu, X. P., Zeng, L. W., Guo, H., Simpson, I. J., Ling, Z. H., Wang, Y., Murray, F., Louie, P. K.
- K., Saunders, S. M., Lam, S. H. M., and Blake, D. R.: Evaluation of the effectiveness of air
  pollution control measures in Hong Kong, Environmental Pollution, 220, 87–94,
  https://doi.org/10.1016/j.envpol.2016.09.025, 2017.
- 1053 Ma, P. K., Zhao, Y., Robinson, A. L., Worton, D. R., Goldstein, A. H., Ortega, A. M., Jimenez, J.-
- L., Zotter, P., Prévôt, A. S. H., Szidat, S., and Hayes, P. L.: Evaluating the impact of new
  observational constraints on P-S/IVOC emissions, multi-generation oxidation, and chamber wall
  losses on SOA modeling for Los Angeles, CA, Atmos Chem Phys, 17, 9237–9259,
  https://doi.org/10.5194/acp-17-9237-2017, 2017.
- 1058 Ma, X., Tan, Z., Lu, K., Yang, X., Chen, X., Wang, H., Chen, S., Fang, X., Li, S., Li, X., Liu, J.,
- 1059 Liu, Y., Lou, S., Qiu, W., and Wang, H.: OH and HO 2 radical chemistry at a suburban site during
- 1060 the EXPLORE-YRD campaign in 2018, Atmos Chem Phys, 22, 7005–7028,
  1061 https://doi.org/10.5194/acp-22-7005-2022, 2022.
- 1062 Magneron, I., Mellouki, A., Le Bras, G., Moortgat, G. K., Horowitz, A., and Wirtz, K.: Photolysis
- and OH-Initiated Oxidation of Glycolaldehyde under Atmospheric Conditions, J Phys Chem A,
- 1064 109, 4552–4561, https://doi.org/10.1021/jp044346y, 2005.
- 1065 Mao, J., Ren, X., Brune, W. H., Olson, J. R., Crawford, J. H., Fried, a., Huey, L. G., Cohen, R. C.,
- 1066 Heikes, B., Singh, H. B., Blake, D. R., Sachse, G. W., Diskin, G. S., Hall, S. R., and Shetter, R.

- 1067 E.: Airborne measurement of OH reactivity during INTEX-B, Atmos Chem Phys, 9, 163–173,
- 1068 https://doi.org/10.5194/acp-9-163-2009, 2009.
- 1069 McDonald, B. C., de Gouw, J. A., Gilman, J. B., Jathar, S. H., Akherati, A., Cappa, C. D., Jimenez,
- 1070 J. L., Lee-Taylor, J., Hayes, P. L., McKeen, S. A., Cui, Y. Y., Kim, S.-W., Gentner, D. R.,
- 1071 Isaacman-VanWertz, G., Goldstein, A. H., Harley, R. A., Frost, G. J., Roberts, J. M., Ryerson, T.
- 1072 B., and Trainer, M.: Volatile chemical products emerging as largest petrochemical source of urban
- 1073 organic emissions, Science, 359, 760–764, https://doi.org/10.1126/science.aaq0524, 2018.
- Mellouki, A., Wallington, T. J., and Chen, J.: Atmospheric chemistry of oxygenated volatile
  organic compounds: Impacts on air quality and climate, Chem Rev, 115, 3984–4014,
  https://doi.org/10.1021/cr500549n, 2015.
- 1077 Millet, D. B., Apel, E., Henze, D. K., Hill, J., Marshall, J. D., Singh, H. B., and Tessum, C. W.:
- 1078 Natural and Anthropogenic Ethanol Sources in North America and Potential Atmospheric Impacts
  1079 of Ethanol Fuel Use, Environmental Science & Technology, 46, 8484–8492, 2012.
- 1080 Min, K.-E., Washenfelder, R. a., Dubé, W. P., Langford, a. O., Edwards, P. M., Zarzana, K. J.,
- 1081 Stutz, J., Lu, K., Rohrer, F., Zhang, Y., and Brown, S. S.: A broadband cavity enhanced absorption
- 1082 spectrometer for aircraft measurements of glyoxal, methylglyoxal, nitrous acid, nitrogen dioxide,
- 1083 and water vapor, Atmos Meas Tech, 9, 423–440, https://doi.org/10.5194/amt-9-423-2016, 2016.
- 1084 Müller, M., Mikoviny, T., Feil, S., Haidacher, S., Hanel, G., Hartungen, E., Jordan, A., Märk, L.,
- 1085 Mutschlechner, P., Schottkowsky, R., Sulzer, P., Crawford, J. H., and Wisthaler, A.: A compact
- 1086 PTR-ToF-MS instrument for airborne measurements of volatile organic compounds at high
- 1087 spatiotemporal resolution, Atmos Meas Tech, 7, 3763–3772, https://doi.org/10.5194/amt-7-3763-
- 1088 2014, 2014.

- 1089 Nault, B. A., Campuzano-Jost, P., Day, D. A., Schroder, J. C., Anderson, B., Beyersdorf, A. J.,
- 1090 Blake, D. R., Brune, W. H., Choi, Y., Corr, C. A., de Gouw, J. A., Dibb, J., DiGangi, J. P., Diskin,
- 1091 G. S., Fried, A., Huey, L. G., Kim, M. J., Knote, C. J., Lamb, K. D., Lee, T., Park, T., Pusede, S.
- 1092 E., Scheuer, E., Thornhill, K. L., Woo, J.-H., and Jimenez, J. L.: Secondary organic aerosol
- 1093 production from local emissions dominates the organic aerosol budget over Seoul, South Korea,
- 1094 during KORUS-AQ, Atmos Chem Phys, 18, 17769–17800, https://doi.org/10.5194/acp-18-17769-
- 1095 2018, 2018.
- 1096 Nihill, K. J., Ye, Q., Majluf, F., Krechmer, J. E., Canagaratna, M. R., and Kroll, J. H.: Influence
- 1097 of the NO/NO<sub>2</sub> Ratio on Oxidation Product Distributions under High-NO Conditions, Environ Sci
- 1098 Technol, 55, 6594–6601, https://doi.org/10.1021/acs.est.0c07621, 2021.
- Orlando, J. J. and Tyndall, G. S.: Laboratory studies of organic peroxy radical chemistry: an
  overview with emphasis on recent issues of atmospheric significance, Chem Soc Rev, 41, 6294–
  6317, https://doi.org/10.1039/c2cs35166h, 2012.
- 1102 Park, R. J., Oak, Y. J., Emmons, L. K., Kim, C. H., Pfister, G. G., Carmichael, G. R., Saide, P. E.,
- 1103 Cho, S. Y., Kim, S., Woo, J. H., Crawford, J. H., Gaubert, B., Lee, H. J., Park, S. Y., Jo, Y. J.,
- 1104 Gao, M., Tang, B., Stanier, C. O., Shin, S. S., Park, H. Y., Bae, C., and Kim, E.: Multi-model
- 1105 intercomparisons of air quality simulations for the KORUS-AQ campaign, Elementa, 9, 1–29,
- 1106 https://doi.org/10.1525/elementa.2021.00139, 2021.
- 1107 Perring, A. E., Bertram, T. H., Farmer, D. K., Wooldridge, P. J., Dibb, J., Blake, N. J., Blake, D.
- 1108 R., Singh, H. B., Fuelberg, H., Diskin, G., Sachse, G., and Cohen, R. C.: The production and
- 1109 persistence of  $\Sigma RONO_2$  in the Mexico City plume, Atmos Chem Phys, 10, 7215–7229,
- 1110 https://doi.org/10.5194/acp-10-7215-2010, 2010.

- 1111 Perring, A. E., Pusede, S. E., and Cohen, R. C.: An observational perspective on the atmospheric
- 1112 impacts of alkyl and multifunctional nitrates on ozone and secondary organic aerosol., Chem Rev,
- 1113 113, 5848–70, https://doi.org/10.1021/cr300520x, 2013.
- 1114 Peterson, D. A., Hyer, E. J., Han, S. O., Crawford, J. H., Park, R. J., Holz, R., Kuehn, R. E.,
- 1115 Eloranta, E., Knote, C., Jordan, C. E., and Lefer, B. L.: Meteorology influencing springtime air
- 1116 quality, pollution transport, and visibility in Korea, Elementa, 7,
  1117 https://doi.org/10.1525/elementa.395, 2019.
- 1118 Picquet-Varrault, B., Suarez-Bertoa, R., Duncianu, M., Cazaunau, M., Pangui, E., David, M.,
- 1119 Doussin, J., Cnrs, U. M. R., Créteil, U. P., Paris, U. De, Pierre, I., and Laplace, S.: Photolysis and
- 1120 oxidation by OH radicals of two carbonyl nitrates: 4-nitrooxy-2-butanone and 5-nitrooxy-2-
- 1121 pentanone, Atmos Chem Phys, 20, 487–498, https://doi.org/10.5194/acp-20-487-2020, 2020.
- 1122 Rao, H., Fullana, A., Sidhu, S., and Carbonell-barrachina, Á. A.: Emissions of volatile aldehydes
- 1123 from heated cooking oils, Food Chem, 120, 59–65,
- 1124 https://doi.org/10.1016/j.foodchem.2009.09.070, 2010.
- 1125 Ren, X., Duin, D. Van, Cazorla, M., Chen, S., Mao, J., Zhang, L., Brune, W. H., Flynn, J. H.,
- 1126 Grossberg, N., Lefer, B. L., Rappenglück, B., Wong, K. W., Tsai, C., Stutz, J., Dibb, J. E., Jobson,
- 1127 B. T., Luke, W. T., and Kelley, P.: Atmospheric oxidation chemistry and ozone production :
- 1128 Results from SHARP 2009 in Houston, Texas, Journal of Geophysical Research: Atmospheres,
- 1129 118, 5770–5780, https://doi.org/10.1002/jgrd.50342, 2013.
- 1130 Richter, D., Weibring, P., Walega, J. G., Fried, A., Spuler, S. M., and Taubman, M. S.: Compact
- highly sensitive multi-species airborne mid-IR spectrometer, Applied Physics B, 119, 119–131,
- 1132 https://doi.org/10.1007/s00340-015-6038-8, 2015.

- 1133 Rosen, R. S., Wood, E. C., Wooldridge, P. J., Thornton, J. A., Day, D. A., Kuster, W., Williams,
- 1134 E. J., Jobson, B. T., and Cohen, R. C.: Observations of total alkyl nitrates during Texas Air Quality
- 1135 Study 2000 : Implications for O<sub>3</sub> and alkyl nitrate photochemistry, J Geophys Res, 109, D07303,
- 1136 https://doi.org/10.1029/2003JD004227, 2004.
- 1137 Sachse, G. W., Hill, G. F., Wade, L. O., and Perry, M. G.: Fast-Response, High-Precision Carbon
- 1138 Monoxide Sensor using a Tunable Diode Laser Absorption Technique, Journal of Geophysical
- 1139 Research: Atmospheres, 92, 2071–2081, https://doi.org/doi:10.1029/JD092iD02p02071, 1987.
- 1140 Sai, S., Ho, H., Yu, J. Z., Chu, K. W., Yeung, L. L., Sai, S., Ho, H., Yu, J. Z., Chu, K. W., and
- 1141 Yeung, L. L.: Carbonyl Emissions from Commercial Cooking Sources in Hong Kong Carbonyl
- 1142 Emissions from Commercial Cooking Sources in Hong Kong, J Air Waste Manage Assoc, 56,
- 1143 1091–1098, https://doi.org/10.1080/10473289.2006.10464532, 2012.
- 1144 Saunders, S. M., Jenkin, M. E., Derwent, R. G., and Pilling, M. J.: Protocol for the development
- 1145 of the Master Chemical Mechanism, MCM v3 (Part A): tropospheric degradation of non-aromatic
- 1146 volatile organic compounds, Atmos Chem Phys, 3, 161–180, https://doi.org/10.5194/acp-3-1611147 2003, 2003.
- 1148 Schauer, J. J., Kleeman, M. J., Cass, G. R., and Simoneit, B. R. T.: Measurement of Emissions
- 1149 from Air Organic Compounds from Cooking with Seed Oils, Environ Sci Technol, 36, 567–575,
- 1150 https://doi.org/10.1021/es002053m, 2002.
- 1151 von Schneidemesser, E., McDonald, B. C., Denier van der Gon, H., Crippa, M., Guizzardi, D.,
- 1152 Borbon, A., Dominutti, P., Huang, G., Jansens-Maenhout, G., Li, M., Ou-Yang, C. F., Tisinai, S.,
- 1153 and Wang, J. L.: Comparing Urban Anthropogenic NMVOC Measurements With Representation
- 1154 in Emission Inventories—A Global Perspective, Journal of Geophysical Research: Atmospheres,
- 1155 128, https://doi.org/10.1029/2022JD037906, 2023.

- 1156 Schroeder, J. R., Crawford, J. H., Ahn, J. Y., Chang, L., Fried, A., Walega, J., Weinheimer, A.,
- 1157 Montzka, D. D., Hall, S. R., Ullmann, K., Wisthaler, A., Mikoviny, T., Chen, G., Blake, D. R.,
- 1158 Blake, N. J., Hughes, S. C., Meinardi, S., Diskin, G., Digangi, J. P., Choi, Y., Pusede, S. E., Huey,
- 1159 G. L., Tanner, D. J., Kim, M., and Wennberg, P.: Observation-based modeling of ozone chemistry
- 1160 in the Seoul metropolitan area during the Korea-United States Air Quality Study (KORUS-AQ),
- 1161 Elementa, 8, https://doi.org/10.1525/elementa.400, 2020.
- 1162 Seo, J., Park, D. R., Kim, J. Y., Youn, D., Lim, Y. Bin, and Kim, Y.: Effects of meteorology and
- 1163 emissions on urban air quality : a quantitative statistical approach to long-term records (1999 –
- 1164 2016) in Seoul, South Korea, Atmos Chem Phys, 18, 16121–16137, 2018.
- 1165 Shetter, R. E. and Müller, M.: Photolysis frequency measurements using actinic flux
- 1166spectroradiometry during the PEM-Tropics mission: Instrumentation description and some results,1167JournalofGeophysicalResearch-Atmospheres,104,5647–5661,
- 1168 https://doi.org/10.1029/98JD01381, 1999.
- 1169 Simpson, I. J., Blake, D. R., Blake, N. J., Meinardi, S., Barletta, B., Hughes, S. C., Fleming, L. T.,
- 1170 Crawford, J. H., Diskin, G. S., Emmons, L. K., Fried, A., Guo, H., Peterson, D. A., Wisthaler, A.,
- 1171 Woo, J., Barré, J., Gaubert, B., Kim, J., Kim, M. J., Kim, Y., Knote, C., Mikoviny, T., Sally, E.,
- 1172 Schroeder, J. R., Wang, Y., Wennberg, P. O., and Zeng, L.: Characterization , sources and
- 1173 reactivity of volatile organic compounds ( VOCs ) in Seoul and surrounding regions during
- 1174 KORUS-AQ, Elementa: Science of the Anthropocene, 8, 37,
- 1175 https://doi.org/10.1525/elementa.434, 2020.
- 1176 Sprengnether, M. M., Demerjian, K. L., Dransfield, T. J., Clarke, J. S., Anderson, J. G., and
- 1177 Donahue, N. M.: Rate Constants of Nine C6-C9 Alkanes with OH from 230 to 379 K: Chemical
- 1178 Tracers for [OH], J Phys Chem A, 113, 5030–5038, https://doi.org/10.1021/jp810412m, 2009.

- 1179 Tan, Z., Lu, K., Hofzumahaus, A., Fuchs, H., Bohn, B., Holland, F., Liu, Y., Rohrer, F., Shao, M.,
- 1180 Sun, K., Wu, Y., Zeng, L., Zhang, Y., Zou, Q., Kiendler-Scharr, A., Wahner, A., and Zhang, Y.:
- 1181 Experimental budgets of OH, HO<sub>2</sub>, and RO<sub>2</sub> radicals and implications for ozone formation in the
- Pearl River Delta in China 2014, Atmos Chem Phys, 19, 7129–7150, https://doi.org/10.5194/acp19-7129-2019, 2019.
- 1184 Teng, A. P., Crounse, J. D., Lee, L., St Clair, J. M., Cohen, R. C., and Wennberg, P. O.: Hydroxy
- 1185 nitrate production in the OH-initiated oxidation of alkenes, Atmos Chem Phys, 15, 4297–4316,
- 1186 https://doi.org/10.5194/acp-15-4297-2015, 2015.
- Thornton, J. A., Wooldridge, P. J., and Cohen, R. C.: Atmospheric NO<sub>2</sub>: In-situ laser-induced
  fluorescence detection at parts per trillion mixing ratios, Anal Chem, 72, 528–539,
  https://doi.org/doi:10.1021/ac9908905, 2000.
- 1190 Travis, K. R., Crawford, J. H., Chen, G., Jordan, C. E., Nault, B. A., Kim, H., Jimenez, J. L.,
- 1191 Campuzano-Jost, P., Dibb, J. E., Woo, J. H., Kim, Y., Zhai, S., Wang, X., McDuffie, E. E., Luo,
- 1192 G., Yu, F., Kim, S., Simpson, I. J., Blake, D. R., Chang, L., and Kim, M. J.: Limitations in
- 1193 representation of physical processes prevent successful simulation of PM2.5 during KORUS-AQ,
- 1194 Atmos Chem Phys, 22, 7933–7958, https://doi.org/10.5194/acp-22-7933-2022, 2022.
- 1195 Tuite, K., Thomas, J. L., Veres, P. R., and Roberts, J. M.: Quantifying Nitrous Acid Formation
- 1196 Mechanisms Using Measured Vertical Profiles During the CalNex 2010 Campaign and 1D
- Column Modeling, Journal of Geophysical Research: Atmospheres, 126, e2021JD034689,
  https://doi.org/10.1029/2021JD034689, 2021.
- 1199 Wang, W., Yuan, B., Peng, Y., Su, H., Cheng, Y., Yang, S., Wu, C., Qi, J., Bao, F., Huangfu, Y.,
- 1200 Wang, C., Ye, C., Wang, Z., Wang, B., Wang, X., Song, W., Hu, W., Cheng, P., Zhu, M., Zheng,
- 1201 J., and Shao, M.: Direct observations indicate photodegradable oxygenated volatile organic

- 1202 compounds (OVOCs) as larger contributors to radicals and ozone production in the atmosphere,
- 1203 Atmos Chem Phys, 22, 4117–4128, https://doi.org/10.5194/acp-22-4117-2022, 2022.
- 1204 Wargocki, P., Weschler, J., and Williams, J.: Assessment of aldehyde contributions to PTR-MS
- m/z 69.07 in indoor air measurements, Environmental Science: Atmospheres, Advance Ar,
  https://doi.org/10.1039/d3ea00055a, 2023.
- 1207 Weinheimer, A. J., Walega, J. G., Ridley, B. A., Gary, B. L., Blake, D. R., Blake, N. J., Rowland,
- 1208 F. S., Sachse, G. W., Anderson, B. E., and Collins, J. E.: Meridional distributions of NOx, NOy,
- 1209 and other species in the lower stratosphere and upper troposphere during AASE II, Geophys Res
- 1210 Lett, 21, 2583–2586, https://doi.org/10.1029/94GL01897, 1994.
- 1211 Whalley, L. K., Stone, D., Bandy, B., Dunmore, R., Hamilton, J. F., Hopkins, J., Lee, J. D., Lewis,
- 1212 A. C., and Heard, D. E.: Atmospheric OH reactivity in central London: observations, model
- predictions and estimates of in situ ozone production, Atmos Chem Phys, 16, 2109–2122,
  https://doi.org/10.5194/acp-16-2109-2016, 2016.
- 1215 Whalley, L. K., Stone, D., Dunmore, R., Hamilton, J., Hopkins, J. R., Lee, J. D., Lewis, A. C.,
- 1216 Williams, P., Kleffmann, J., Laufs, S., and Woodward-massey, R.: Understanding in situ ozone
- 1217 production in the summertime through radical observations and modelling studies during the Clean
- 1218 air for London project ( ClearfLo ), Atmos Chem Phys, 18, 2547-2571,
- 1219 https://doi.org/10.5194/acp-18-2547-2018, 2018.
- 1220 Whalley, L. K., Slater, E. J., Woodward-massey, R., Ye, C., Lee, J. D., Squires, F., Mehra, A.,
- 1221 Worrall, S. D., Bacak, A., Bannan, T. J., Coe, H., and Percival, C. J.: Evaluating the sensitivity of
- 1222 radical chemistry and ozone formation to ambient VOCs and NOx in Beijing, Atmos Chem Phys,
- 1223 21, 2125–2147, https://doi.org/10.5194/acp-21-2125-2021, 2021.

- Wolfe, G. M., Marvin, M. R., Roberts, S. J., Travis, K. R., and Liao, J.: The Framework for 0-D
  Atmospheric Modeling (F0AM) v3.1, Geosci Model Dev, 9, 3309–3319,
- 1226 https://doi.org/10.5194/gmd-9-3309-2016, 2016.
- 1227 Wooldridge, P. J., Perring, A. E., Bertram, T. H., Flocke, F. M., Roberts, J. M., Singh, H. B., Huey,
- 1228 L. G., Thornton, J. A., Wolfe, G. M., Murphy, J. G., Fry, J. L., Rollins, A. W., LaFranchi, B. W.,
- 1229 and Cohen, R. C.: Total Peroxy Nitrates (*SPNs*) in the atmosphere: the Thermal Dissociation-Laser
- 1230 Induced Fluorescence (TD-LIF) technique and comparisons to speciated PAN measurements,
- 1231 Atmos Meas Tech, 3, 593–607, https://doi.org/DOI 10.5194/amt-3-593-2010, 2010.
- 1232 Xu, Y., Feng, X., Chen, Y., Zheng, P., Hui, L., Chen, Y., Yu, J. Z., and Wang, Z.: Development
- 1233 of an enhanced method for atmospheric carbonyls and characterizing their roles in photochemistry
- 1234 in subtropical Hong Kong, Science of The Total Environment, 896, 165135,
  1235 https://doi.org/10.1016/j.scitotenv.2023.165135, 2023.
- 1236 Yang, G., Huo, J., Wang, L., Wang, Y., Wu, S., Yao, L., Fu, Q., and Wang, L.: Total OH Reactivity
- 1237 Measurements in a Suburban Site of Shanghai J, Journal of Geophysical Research: Atmospheres,
- 1238 127, 1–20, https://doi.org/10.1029/2021JD035981, 2022.
- 1239 Yeh, G. K. and Ziemann, P. J.: Alkyl nitrate formation from reactions of C8-C14 n-alkanes with
- OH radicals in the presence of NOx: Measured yields with essential corrections for gas-wall partitioning, Journal of Physical Chemistry A, 118, 8147–8157,
- 1242 https://doi.org/10.1021/jp500631v, 2014.
- 1243 Yeo, M. J. and Kim, Y. P.: Long-term trends of surface ozone in Korea, J Clean Prod, 294, 125352,
- 1244 https://doi.org/10.1016/j.jclepro.2020.125352, 2021.
- 1245 Yuan, B., Shao, M., de Gouw, J., Parrish, D. D., Lu, S., Wang, M., Zeng, L., Zhang, Q., Song, Y.,
- 1246 Zhang, J., and Hu, M.: Volatile organic compounds (VOCs) in urban air: How chemistry affects

- the interpretation of positive matrix factorization (PMF) analysis, Journal of Geophysical
  Research: Atmospheres, 117, n/a-n/a, https://doi.org/10.1029/2012JD018236, 2012.
- 1249 Zare, A., Romer, P. S., Nguyen, T., Keutsch, F. N., Skog, K., and Cohen, R. C.: A comprehensive
- 1250 organic nitrate chemistry: Insights into the lifetime of atmospheric organic nitrates, Atmos Chem
- 1251 Phys, 18, 15419–15436, https://doi.org/10.5194/acp-18-15419-2018, 2018.
- 1252 Zhao, Y., Hennigan, C. J., May, A. A., Tkacik, D. S., De Gouw, J. A., Gilman, J. B., Kuster, W.
- 1253 C., Borbon, A., and Robinson, A. L.: Intermediate-volatility organic compounds: A large source
- 1254 of secondary organic aerosol, Environ Sci Technol, 48, 13743–13750,
  1255 https://doi.org/10.1021/es5035188, 2014.
- 1256 Zheng, W., Flocke, F. M., Tyndall, G. S., Swanson, A., Orlando, J. J., Roberts, J. M., Huey, L. G.,
- 1257 and Tanner, D. J.: Characterization of a thermal decomposition chemical ionization mass
- 1258 spectrometer for the measurement of peroxy acyl nitrates (PANs) in the atmosphere, Atmos Chem
- 1259 Phys, 11, 6529–6547, https://doi.org/10.5194/acp-11-6529-2011, 2011.
- 1260 Zhou, Z., Tan, Q., Deng, Y., Song, D., Wu, K., Zhou, X., Huang, F., Zeng, W., and Lu, C.:
- Compilation of emission inventory and source profile database for volatile organic compounds : 1261 for 1262 A case study Sichuan, China, Atmos Pollut Res, 11, 105 - 116, 1263 https://doi.org/10.1016/j.apr.2019.09.020, 2020.
- 1264
- 1265

### **Tables**

1267 Table 1. Reactions described in text along with associated rate constants and references for those1268 rate constants.

	Reaction	Reaction Rate	Reference
R1a	VOC+OH $\xrightarrow{O_2} \operatorname{RO}_2^{\cdot}$	Varies	Atkinson (2003); Atkinson and Arey(2003); Atkinson et al. (2006); Bohn and Zetzsch (2012); Sprengnether et al. (2009)
R1b	$\text{VOC}+hv \xrightarrow{O_2} \text{RO}_2^{\cdot}$	Varies/Measured	Shetter & Müller (1999)
R2a	$\text{RO}_2^+\text{NO} \rightarrow (1-\alpha) \text{ RO}^+(1-\alpha) \text{ NO}_2$	2.7×10 <sup>-11</sup> ×exp(390/T)	Burkholder et al. (2020)
R2b	$RO_2^{\cdot}+NO \rightarrow \alpha RONO_2$	2.7×10 <sup>-11</sup> ×exp(390/T)	Burkholder et al. (2020)
R3	$NO_2 + hv \rightarrow NO + O(^3P)$	Measured on DC-8	Shetter & Müller (1999)
R4	$O(^{3}P)+O_{2} \rightarrow O_{3}$	3.2×10 <sup>-11</sup> ×exp(67/T)	Saunders et al. (2003)
R5	$RO'+O_2 \rightarrow R(O)+HO_2$	Assumed Instantaneous	
R6	$HO_2 + NO \rightarrow OH + NO_2$	3.45×10 <sup>-12</sup> ×exp(270/T)	Saunders et al. (2003)
R7	RCHO+OH $\xrightarrow{O_2} R(O)O_2$	Varies	Atkinson (2003); Atkinson and Arey (2003); Atkinson et al. (2006)
R8ª	$R(O)O_2^{\cdot}+NO_2 \leftrightarrow R(O)O_2NO_2$	F: 8.69×10 <sup>-12</sup> cm <sup>3</sup> molec. <sup>-1</sup> s <sup>-1</sup> R: 4.30×10 <sup>-4</sup> s <sup>-1</sup>	Burkholder et al. (2020)
R9	$R(O)O_2 + NO \rightarrow RO_2 + NO_2$	8.1×10 <sup>-12</sup> ×exp(270/T)	Burkholder et al. (2020)
R10	$O_3 + NO \rightarrow O_2 + NO_2$	2.07×10 <sup>-12</sup> ×(-1400/T)	Burkholder et al. (2020)
R11 <sup>b</sup>	$OH+ NO_2 \rightarrow HNO_3$	1.24×10 <sup>-11</sup> cm <sup>3</sup> molec. <sup>-1</sup> s <sup>-1</sup>	Burkholder et al. (2020)
R12	$O_3$ +hv $\xrightarrow{H_2O} 2O(^1D)$	hv measured on DC-8; 2.14×10 <sup>-10</sup> cm <sup>3</sup> molec. <sup>-1</sup> $s^{-1}$	Shetter & Müller (1999); Saunders et al. (2003)
R13	$O_3 + OH \rightarrow HO_2 + O_2$	1.7×10 <sup>-12</sup> ×exp(-940/T)	Saunders et al. (2003)
R14	$O_3 + HO_2 \rightarrow OH + 2O_2$	$1.0 \times 10^{-14} \times exp(-490/T)$	Burkholder et al. (2020)
R15 <sup>b</sup>	$HO_2 + HO_2 \xrightarrow{H_2O} H_2O_2$	$5.06 \times 10^{-12} \text{ cm}^3 \text{ molec.}^{-1} \text{ s}^{-1}$	Saunders et al. (2003)
R16	$HO_2 + RO_2 \rightarrow Products$	2.91×10 <sup>-13</sup> ×exp(1300/T)	Saunders et al. (2003)
R17	$HO_2$ +OH $\rightarrow$ Products	4.80×10 <sup>-11</sup> ×exp(250/T)	Burkholder et al. (2020)

R18 <sup>b</sup>	$OH+NO \rightarrow HONO$	$7.40 \times 10^{-12} \text{ cm}^3 \text{ molec.}^{-1} \text{ s}^{-1}$	Burkholder et al. (2020)	
R19	$HO_2 + R(0)O_2 \rightarrow Products$	4.30×10 <sup>-13</sup> ×exp(1040/T)	Burkholder et al. (2020)	

<sup>a</sup>Only showing forward (F) and reverse (R) rate constant at 298 K and 1013 hPa and being a termolecular reaction. <sup>b</sup>Termolecular reaction; only showing rate at 298 K and 1013 hPa 1270 1271

### 1273 **Table 2.** List of instruments, compounds measured, accuracy/precision, and associated references used in this study.

Instrument	Species	References
University of California, Irvine, Whole Air Sampler (WAS)	Ethane, Ethene, Ethyne, Propane, Propene, i- Butane, n-Butane, 1-Butene, i-Butene, trans-2- Butene, cis-2-Butene i-Pentane, n-Pentane, 1,3- Butadiene, Isoprene, n-Hexane, n-Heptane, n- Octane, n-Nonane, n-Decane, 2,3- Dimethylbutane, 2-Methylpentane, 3- Methylpentane, Cyclopentane, Methylcyclopentane, Cyclohexane, Methylcyclohexane, Benzene, Toluene, m+p- Xylene, o-Xylene, Ethylbenzene, Styrene, i- Propylbenzene, n-Propylbenzene, 3- Ethyltoluene, 4-Ethyltoluene, 2-Ethyltoluene, 1,3,5-Trimethylbenzene, 1,2,4- Trimethylbenzene, 1,2,3-Trimethylbenzene, $\alpha$ - Pinene, $\beta$ -Pinene, Methyl nitrate, Ethyl nitrate, i-Propyl nitrate, n-Propyl nitrate, 2-Butyl nitrate, 3-Pentyl nitrate, 2-Pentyl nitrate, 3- Methyl-2-Butyl nitrate	Simpson et al. (2020)
The Pennsylvania State University Airborne Tropospheric Hydrogen Oxides Sensor (ATHOS)	OH, HO <sub>2</sub> , OH Reactivity	Faloona et al. (2004), Mao et al. (2009), Brune et al. (2019)
University of California, Berkeley, Thermal Dissociation-Laser Induced Fluorescence (TD-LIF)	NO2, ΣPNs, ΣANs	Thornton et al. (2000), Day et al. (2002), Wooldridge et al. (2010)
NASA Langley Diode Laser Hygrometer (DLH)	H <sub>2</sub> O	Diskin et al. (2002)
NASA Langley Diode Laser Spectrometer Measurements (DACOM)	$CO, CH_4$	Sachse et al. (1987)
University of Colorado, Boulder, Compact Atmospheric Multi-species Spectrometer (CAMS)	CH <sub>2</sub> O, C <sub>2</sub> H <sub>6</sub>	Richter et al. (2015), Fried et al. (2020)
Gwangju Institute of Science and Technology Korean Airborne Cavity Enhances Spectrometer (K-ACES)	СНОСНО	Min et al. (2016), D. Kim et al. (2022)

NCAR CCD Actinic Flux Spectroradiometers (CAFS)	j-values	Shetter & Müller (1999)
Georgia Institute of Technology Chemical Ionization Mass Spectrometer (GT)	SO2, PAN, PPN, APAN, PBzN	Kim et al. (2007), Lee et al. (2022)
University of Colorado, Boulder, High-Resolution Time-of-Flight Aerosol Mass Spectrometer	pNO <sub>3</sub>	Nault et al. (2018), Day et al. (2022)
NCAR 4-Channel Chemiluminescence Instrument (NCAR)	NO, NO <sub>2</sub> , O <sub>3</sub> , NO <sub>y</sub>	Weinheimer et al. (1994)
California Institute of Technology Chemical Ionization Mass Spectrometer (CIT)	Butene Hydroxynitrates, Butadiene Hydroxnitrates, Ethene Hydroxynitrates, Ethanal Nitrate, Isoprene Hydroxynitrates, Propene Hydroxynitrates, Propanal Nitrate, CH <sub>3</sub> OOH, Peroxyacetic Acid, HNO <sub>3</sub> , Hydroxyacetone, H <sub>2</sub> O <sub>2</sub>	Crounse et al. (2006), Teng et al. (2015)
University of Oslo Proton Transfer Reaction Time-of- Flight Mass Spectrometer (PTR-MS)	Methanol, Acetaldehyde, Acetone+Propanal, Isoprene, MVK+MACR+ISOPOOH, Benzene, Toluene, C8-alkylbenzenes, Monoterpenes, MEK	Müller et al. (2014)
NSRC Meteorological and Geographical Data	Latitude, Longitude, Altitude, Temperature, Pressure	Crawford et al. (2021)





1277 Figure 1. Binned NO<sub>x</sub> mixing ratios observed on the NASA DC-8 during the KORUS-AQ campaign. Note, the color bar scale is logarithmic. The binning is along the flight paths of the 1278 1279 NASA DC-8 for any observations collected below 2.0 km and after 11:00 local time. The rest of the NASA DC-8 flight paths not included in the analysis are shown in grey. Three key areas from 1280 1281 KORUS-AQ are highlighted—the Olympic Park ground site, the airfield where the NASA DC-8 conducted routine missed approaches, and the Taehwa Research ground site. The histograms 1282 1283 above and to the left are the distribution of NOx mixing ratios longitudinally and latitudinally, 1284 respectively.



1286 Figure 2. (a) Scatter plot of the summation of individual  $NO_z$  ( $NO_z$  is higher oxide  $NO_x$  products) 1287 measured by GT, CIT, WAS, TD-LIF, and AMS versus NO<sub>z</sub> measured by difference between NO<sub>y</sub> 1288 and  $NO_x$  (see Table 2 for compounds measured by each instrument).  $NO_x$  is NO measured by 1289 NCAR and NO<sub>2</sub> measured by LIF. The observations are for when the DC-8 was over the SMA. 1290 (b) Average contribution of measured speciated NOz over the SMA during KORUS-AQ versus 1291 NO<sub>x</sub>. Higher PNs is PPN + APAN + PBZN. ΣUn-speciated PNs is total peroxnitrates from TD-1292 LIF minus total measurement from GT. Alkyl RONO2 is the total small alkyl nitrate measurements from WAS. Multifunctional RONO<sub>2</sub> is the total measurements from CIT. 2Un-speciated ANs is 1293 1294 the total alkyl nitrates from TD-LIF minus total RONO<sub>2</sub> from CIT and WAS.



1296Figure 3. Scatter plot of (a)  $O_x$  versus ΣANs and (be) ΣPNs versus formaldehyde (CH<sub>2</sub>O) over1297SMA (see Figure 1 for area studied). Data is colored by meteorological periods discussed in1298Peterson et al. (2019). Data plotted here is after 11:00 am LT to minimize impact of growing1299boundary layer and nocturnal residual layer mixing. The curvature in (be) is further explored in1300Figure S<u>9</u>7. Eq. 8, 9, and 11 is used to convert the slope in (a) into  $\alpha_{eff}$ . The units of the slope are1301ppbv ppbv<sup>-1</sup>.



1304 Figure 4. (a) Upper panel is the binned calculated (calc.) unmeasured VOC reactivity (R(VOC)<sub>u</sub>). 1305 Note, unmeasured is for any species not measured on DC-8 or constrained by F0AM and is 1306 calculated using Eq. 13. Lower panel is bBinned VOC reactivity versus NOx observed over SMA 1307 during KORUS-AQ (see Figure 1 for the area studied). The measured observed R(VOC), labeled 1308 as "From PSU", where PSU is Pennsylvania State University, is the VOC reactivity calculated 1309 from the measured total OH reactivity with inorganic OH reactivity removed. As discussed in 1310 Brune et al. (2022), the OH reactivity has interferences at high NO<sub>x</sub> mixing ratios. The error bar is 1311 the uncertainty in the OH reactivity measurement (Brune et al., 2022). The red line represents the 1312 calculated unmeasured R(VOC), using Eq. 11, with an assumed  $\alpha = 0.10$ . The shaded area 1313 represents different calculated unmeasured R(VOC), assuming different  $\alpha$  for the unmeasured 1314 R(VOC) (see Eq. 11). (b) The calculated effective  $\alpha$  from observations versus  $NO_x$ . The dashed 1315 purple line is the effective  $\alpha$  estimated from Eq. 110, using the slope from Figure 3a. For both (a)

- 1316 and (b), the colored stacked data is the calculated VOC reactivity (a) and weighted effective  $\alpha$  (b).
- 1317 The values from (b) are calculated using Eq. 11. Finally, for both (a) and (b), FOAM species is the reactivity for compounds not measured on the DC-8 predicted by FOAM with an estimated  $\alpha =$
- 1318
- 1319 0.05. The associated uncertainty in using different  $\alpha$  for the F0AM predicted reactivity is explored
- 1320 in Figure S<u>5</u>4.



1322

1323 Figure 5. (a) Scatter plot of binned higher  $\Sigma$ PNs calculated using F0AM (red) or binned higher 1324 ΣPNs from observations (black) versus formaldehyde (CH<sub>2</sub>O). Slopes shown are ODR fits to the 1325 binned data. PPN and PAN were constrained by observations for F0AM while all the other higher 1326 PNs were not constrained (b) Fractional contribution of the higher PNs predicted from F0AM 1327 versus NO<sub>x</sub>. (c) Fractional contribution of different precursors to PAN, predicted by F0AM versus 1328 NO<sub>x</sub>. For both (b) and (c), Alk is all alkanes, Arom is all aromatics, and  $\geq C_4$  Alk is all alkanes with 1329 4 or more carbons. See Figure S108 for comparison of F0AM.



1332 Figure 6. (a) Net  $O_x$  ( $O_3 + NO_2$ ) production (see Eq. 1 and 2) predicted for SMA using measured 1333 and F0AM R(VOC) (dashed) or total R(VOC) (solid), from Figure 4a, versus NOx. (b) Contribution of different reactions to the total O<sub>x</sub> loss versus NO<sub>x</sub>. (c) Contribution of different 1334 1335 reactions to total HO<sub>x</sub> (HO<sub>x</sub> = OH + HO<sub>2</sub> + RO<sub>2</sub> + R(O)O<sub>2</sub>) loss versus NO<sub>x</sub>. The predicted RO<sub>2</sub> comes from the total VOC reactivity calculated in Figure 4a assuming steady-state (Eq. 7), and 1336 1337 HO<sub>2</sub> the acyl peroxy radicals are from F0AM results. Note for both (b) and (c), net RC(O)O<sub>2</sub> + 1338  $NO_2$  and net  $CH_3C(O)O_2 + NO_2$  are described in Eq. 3. Radical reactions contributing < 1% to the 1339 L(O<sub>x</sub>) or L(HO<sub>x</sub>) are not included. Also note that F0AM HO<sub>2</sub>, CH<sub>3</sub>C(O)O<sub>2</sub>, R(O)O<sub>2</sub>, and F0AM 1340 secondary products are used here along with observations.

Formatted: Subscript



### 

1342Figure 7. Calculated HOx production from observations (colored stack) compared with the1343calculated HOx loss from Figure 6c over the SMA during KORUS-AQ.