

*Supplement of*

## **Aerosol Size Distribution Properties Associated with Cold-Air Outbreaks in the Norwegian Arctic**

Abigail S. Williams<sup>1</sup>, Jeramy L. Dedrick<sup>1</sup>, Lynn M. Russell<sup>1</sup>, Florian Tornow<sup>2,3</sup>, Ann M. Fridlind<sup>3</sup>, and Israel Silber<sup>4,\*</sup>, Benjamin Swanson<sup>5</sup>, Paul J. DeMott<sup>5</sup>, Paul Zieger<sup>6,7</sup>, and Radovan Krejci<sup>6,7</sup>

<sup>1</sup>Scripps Institution of Oceanography, University of California San Diego, La Jolla, CA, USA

<sup>2</sup>Center for Climate Systems Research, Columbia University, New York, NY, USA

<sup>3</sup>NASA Goddard Institute for Space Studies, New York, NY, USA

<sup>4</sup>Department of Meteorology and Atmospheric Science, Pennsylvania State University, University Park, PA, USA

\*now at Atmospheric, Climate, and Earth Sciences Division, Pacific Northwest National Laboratory, Richland, WA, USA

<sup>5</sup>Department of Atmospheric Science, Colorado State University, Fort Collins, CO, USA

<sup>6</sup>Department of Environmental Science, Stockholm University, Stockholm, Sweden

<sup>7</sup>Bolin Centre for Climate Research, Stockholm University, Stockholm, Sweden

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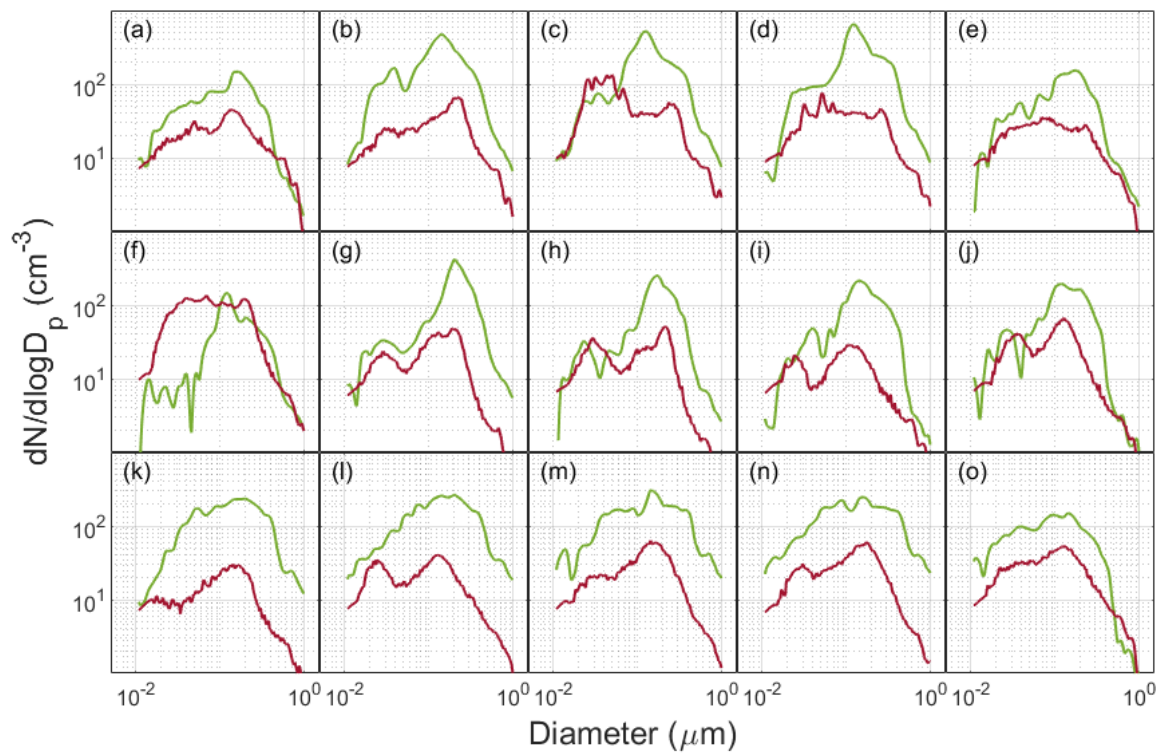
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**Text S1.**

A very weak negative correlation between  $D_{\text{HM}}$  and  $N_{\text{HM}}$  ( $r = -0.15$ ,  $p < 0.05$ ) is observed during CAO events at Andenes. If the supersaturation increase were aerosol-limited for non-precipitating low-cloud conditions, then a positive correlation of  $D_{\text{HM}}$  with  $N_{\text{HM}}$  is expected (Dedrick et al., 2024). If the supersaturation is not limited by the number of particles that activate, then  $D_{\text{HM}}$  may be limited by updraft velocity (Chen et al., 2016). No correlation is observed between  $D_{\text{HM}}$  and the updraft velocity retrievals available from Bear Island (COMBLE supplemental site) for time periods corresponding to when air mass trajectories arriving at Andenes passed near Bear Island ( $r = -0.1$ ,  $p > 0.05$ ). The absence of either correlation may result from scavenging of activated particles during precipitation.

**Figure S1.** Measured aerosol size distributions at Zeppelin Observatory (green; upwind) and Andenes (red; downwind) for 15 CAO trajectories (Sect. 4). For subplots (a)-(o), refer to Table S1 for corresponding dates and times.



**Table S1.** The dates and times that air masses passing by Zeppelin Observatory arrived at Andenes for each of the 15 CAO trajectories (Sect. 4) shown in Fig. S1.

Figure S1 subplot	Time ( <i>mm/dd/yy HH:MM in UTC</i> ) of passage by Zeppelin Observatory	Time ( <i>mm/dd/yy HH:MM in UTC</i> ) of arrival at Andenes
(a)	12/1/19 12:00	12/2/19 15:00
(b)	1/3/20 09:00	1/4/20 15:00
(c)	1/3/20 18:00	1/4/20 18:00
(d)	1/3/20 21:00	1/4/20 21:00
(e)	1/21/20 11:00	1/22/20 18:00
(f)	2/4/20 13:00	2/5/20 18:00
(g)	2/17/20 19:00	2/19/20 00:00
(h)	2/17/20 23:00	2/19/20 03:00
(i)	3/10/20 22:00	3/12/20 09:00
(j)	3/11/20 02:00	3/12/20 12:00
(k)	3/12/20 04:00	3/13/20 06:00
(l)	3/12/20 22:00	3/13/20 18:00
(m)	3/13/20 0:00	3/13/20 21:00
(n)	3/13/20 01:00	3/14/20 00:00
(o)	3/24/20 04:00	3/25/20 15:00

**Table S2.** Summary of reported CCN concentrations and the CCN/CN ratios measured at high latitudes (above 60°N) at supersaturations (SS) near 0.4 %. Values reported are mean except where noted.

Reference	CCN (cm <sup>-3</sup> )	CCN/CN	SS (%)	Location
Moore et al., 2011	550 <sup>a</sup>	~ 0.90	0.42	Arctic boundary layer
Jung et al., 2018	45-81 <sup>b</sup>	~ 0.3-0.85	0.4	Zeppelin Observatory
Paramonov et al., 2015	31-149	~ 0.08-1.8	0.3	Pallas (Northern Finland)
	50-176	~ 0.1-0.38	0.5	Pallas (Northern Finland)
Dall'Osto et al., 2017	69-117	n/a	0.4	Canadian Arctic
Herenz et al., 2018	139	n/a	0.3	Canadian Arctic
	164	n/a	0.5	Canadian Arctic
Zabori et al., 2015	~75-230 <sup>b</sup>	~ 0.1-0.65	0.4	Zeppelin Observatory
Martin et al., 2011	35	n/a	0.41	High Arctic
Latham et al., 2013	247 <sup>b</sup>	0.52	0.55	Arctic

<sup>a</sup>95th percentile

<sup>b</sup>Median

## References

- Chen, J., Liu, Y., Zhang, M., and Peng, Y.: New understanding and quantification of the regime dependence of aerosol-cloud interaction for studying aerosol indirect effects, *Geophysical Research Letters*, 43, 1780–1787, <https://doi.org/10.1002/2016GL067683>, 2016.
- Dedrick, J. L., Russell, L. M., Sedlacek, A. J., III, Kuang, C., Zawadowicz, M. A., & Lubin, D.: Aerosol-correlated cloud activation for clean conditions in the tropical Atlantic boundary layer during LASIC, *Geophysical Research Letters*, 51, e2023GL105798, <https://doi.org/10.1029/2023GL105798>, 2024.