Response to Reviewer #2’s Comments:

Yuxin Zhao et al. (Author)

We are deeply grateful to Reviewer #2 for the suggested revisions to our manuscript. The issues raised highlighted some shortcomings in our manuscript, which we have subsequently addressed. Our conclusions regarding the different impacts of aerosols on cold and warm cloud bases was formed after analyzing observational data and considering mechanisms proposed in previous studies. Following the reviewer's advice, we revised the description in "Conclusions and Discussions" to clarify the potential mechanisms and acknowledge the uncertainty in the causal relationship between results and mechanisms. Additionally, in response to the reviewer's suggestion to average the instantaneous observations of cloud radiative effect throughout the daytime, we processed the data using the recommended method and included the results in our response (the response to Comment #15). In our understanding, this method is reliable to analyze daily mean cloud radiative effects from a climatological perspective. The method is based on the premise that clouds remain unchanged throughout the daytime, with only solar radiation exhibiting periodic variations. However, although our study is based on over a decade of satellite observations, the total sample size is only a few hundred. Hence, our results must be considered cautiously in the sense of climatology. Consequently, we revised the vague expression “daytime” and “nighttime” to “~1:30 p.m.” and “~1:30 a.m.” separately, emphasizing the "instantaneity". And the discussion of the limitations of daily-twice measurements is added. Please see our point-to-point reply to comments. All revisions were shown in revised manuscript by using track changes.

General Comments

The focus of this study was to capture deep convective system properties over the Tibetan Plateau using satellite remote sensing observations. This research is deemed novel as not much is understood of the cloud structure and radiative effects of full deep
convective systems over this region. Furthermore, it is challenging to decouple the relative influences of meteorological conditions and aerosol concentrations on the deep convective cloud structure and precipitation. Therefore, the analysis also investigated how dynamical properties such as vertical wind shear, convective instability, and vertical velocity influence cloud and precipitation development under differing aerosol loading environments. The results were supported by explanations of potential mechanisms that were theorized in previous work. However, the exact mechanisms, such as how the aerosol invigoration effect differs between warm-base and cold-base clouds, or how aerosol invigoration and aerosol radiative effects individually influence entrainment suppression, were not themselves tested or observed in this study. While it seems appropriate to speculate the potential mechanisms, the conclusions heavily relied on such mechanisms to explain the results. Therefore, I suggest that it would serve the paper better if the concluding remarks pointed to such mechanisms without stating that these mechanisms are the reason for the results.

Response: We appreciate the constructive suggestions provided. Indeed, our conclusions remain speculative, drawing on mechanisms proposed in previous studies. We add a description of these limitations in the “conclusions and discussion” Section. According to the reviewer's comments, we revise the descriptions about the possible mechanisms of the different effects of aerosols on warm-base and cold-base clouds in the conclusion. We are grateful for the reviewer's suggestion, which are very helpful in improving the rigor and readability of this paper.

Specific Comments

1. L41: What did Luo et al. (2011) find?

Response: Luo et al. (2011) found that deep convection core over the TP is shallower, less frequent, and embedded in smaller-size convection systems. And the cloud tops of deep convection cores are more densely packed. We have
incorporated the more detailed conclusions mentioned above into the citation of Luo et al. (2011) in the revised manuscript.

**See the line: 42-43.**

2. *L70-76:* These sentences are the start of a different point, so I think they would best be served in a different paragraph that discusses how spaceborne measurements—and particularly the ones that you are using—have been used for aerosol-cloud-precipitation measurements.

**Response:** Thank you for your suggestion. We have moved this part to a new paragraph following this section to describe the application of satellite measurements in cloud, aerosol and precipitation.

**See the lines: 82-90.**

3. *L106-107:* Measurements are two times a day at 1:30 am/pm LST, not for the full day, between 2006-2011.

**Response:** We appreciate the reviewer pointing out the imprecise description in our manuscript. In the revised manuscript, it has been corrected to: "the 2B-CLDCLASS-LIDAR data is only available at ~1:30 a.m./p.m. local time from 2006 to 2011 and solely ~1:30 p.m. from 2012 to 2019."

**See the lines: 114-115.**

4. *L111-112:* “It is important to note that CloudSat will no longer operate during nighttime due to the battery anomaly (Witkowski et al., 2018)” seems out of place. Is this in reference to the DO-Op switch made in 2012?

**Response:** Yes, we are indeed referring to the switch to DO-Op. We have revised the inappropriate description pointed out by the reviewer to the following sentence:
"It is important to note that all CloudSat data since the switch to DO-Op in late 2011 have been daytime-only (at ~1:30 p.m.)."

See the lines: 113 and 119-120.

5. *L115*: Is daytime SW CRE normalized to account for variability in solar insolation throughout the day? If not, that could drastically impact the results — CRE would be much more enhanced at 1:30 pm compared to the rest of the day.

Response: In this study, the SW CRE has not been normalized and only represents the instantaneous radiative effects observed at 1:30 p.m. As described at the beginning of our response, we think that the normalization method, which assumes clouds remain unchanged during daytime, is more applicable to climatological studies with large sample sizes. However, this method seems unsuitable for our study, which involves only a few hundred cloud clusters. In the revised manuscript, we have added the description of the limitations of instantaneous radiative effect observations. Additionally, we also attempted to normalize the SW CRE following the reviewer's suggestion and presented the results in our response to comment #15, though we did not include it in the main text.

See the line: 123-126.

6. *L177*: So anvil can be precipitating or non-precipitating?

Response: Generally, the thicker parts of the anvil near the core can be precipitating, while the thinner cirrus parts at the edges are non-precipitating. For example, as shown in Figure 1, several cumulus profiles near DCC are observed. These profiles may exist drizzle or snow. However, most of the anvil profiles are non-precipitating. As results, the precipitation from anvils are not included in the results in our study. The precipitation in our study is the precipitation from DCC.
7. *Figure 1:* Just to clarify, is this example not included in your analysis of CRE since you are ignoring systems with multiple cloud-layer profiles (i.e., systems with underlying low-level cloud)?

**Response:** The example shown in Figure 1 is included in the CRE analysis. Although we excluded multilayer cloud profiles in DCS to eliminate their influence on the DCS radiative effect, we did not exclude the entire DCS. In other words, the CRE of the DCS shown in Figure 1 is calculated as the average CRE of its single-layer cloud profiles. We represent the average cloud radiative forcing of the entire DCS using the average results of these single-layer cloud profiles. This approach is because DCS, as a cloud system with a large horizontal extent, rarely consists entirely of single-layer clouds. Therefore, in calculating the CRE of DCS that includes multilayer cloud profiles, we exclude only the multilayer cloud profiles and use the results of the single-layer cloud profiles from that DCS. In the revised manuscript, we have improved the description of the method used for excluding multilayer cloud profiles: “In other words, in calculating the CRE of DCS that includes multi-layer cloud profiles, we average the CRE of the single-layer cloud profiles in this DCS and ignore the multi-layer cloud profiles (the results of this DCS are not excluded).”

**See the lines:** 128-131.

8. *L208-209:* What is your motivation for selecting meteorological factors the hour before the DCS was detected by CloudSat/CALIPSO? Do you want to select them before convection is initiated, or before the DCS advects into that region? Or to match up with the aerosol information?

**Response:** The influence of meteorological factors on the development of deep convective clouds (DCS) exhibits a lag. In other words, the state of the DCS at the time of observation is more closely related to the meteorological factors from before that moment. Therefore, we analyzed the meteorological field from one hour
prior to the DCS observation. And the one-hour movement distance of DCS relative to the ERA5 grid size does not have a significant impact (details shown in the response to Comment #9). This matching method is commonly used in studies analyzing the impact of thermal and dynamic factors on clouds and precipitation (e.g., Sun et al., 2023). The aerosol information is matched to the DCS in time to account for the effects of wet scavenging, hence the choice of aerosol information from before the precipitation event. Consequently, the selected aerosol and meteorological factor data may not be from the same time. However, when precipitation does not occur, aerosol changes are relatively minor. Consequently, the impact of this temporal matching on the results is smaller compared to wet scavenging.

9. **L209-210**: When you say that the convective system movement under advection is ignored, what are you suggesting? Are environments fairly homogeneous such that you do not need to consider the meteorological conditions in the region that convection moves into?

**Response**: I'm sorry for any confusion our expression may have caused the reader. “the convective system movement under advection is ignored” means that the one-hour movement distance of DCS relative to the ERA5 grid size does not have a significant impact. Due to the lag effect of meteorological factors on clouds, we used meteorological data from the hour before the DCS was detected by CloudSat/CALIPSO. For spatial matching, we selected meteorological data spatially proximate to the DCS/DCC. Considering the grid size of the meteorological data is 0.25°*0.25° (longitude*latitude), the DCS is unlikely to move out of its located grid during this hour, or it may move to an adjacent grid, but not far enough to experience significant differences in meteorological conditions. Referencing the study by Sherwood and Wahrlich (1999) on tropical ocean convective clouds, they found that the movement speed of convective cloud systems correlates highest with 700 hPa wind speed. We also examined the 700
hPa wind speeds corresponding to our TO region samples. Only 11.7% of the samples had 700 hPa horizontal wind speeds exceeding 10 m/s (samples that moved more than the diagonal length of a 0.25° grid in one hour). Previous studies analyzing the influence of meteorological fields on clouds or precipitation have also commonly used the method of spatial proximity with a temporal lead for matching. For example, in Sun et al. (2023), the CAPE and wind shear used are before precipitation start time in the analysis of the impact of meteorological factors on precipitation. In the revised manuscript, we have modified the expression.

**See the lines:** 235-239.

10. **Section 3.1:** Is this section considering both daytime and nighttime DCSs?

**Response:** Yes, in the statistics presented in Section 3.1, we have included both daytime and nighttime DCS without making any distinction between them. In the revised manuscript, the differences in daytime and nighttime DCSs are discussed in reference to the Comment #12.

11. **L230:** Can one cloud cluster or DCS contain multiple DCCs in your analysis?

**Response:** Yes, in our study, there are samples where DCC profiles within a DCS are discontinuous, or in other words, multiple DCCs exist. Additionally, we examined some samples where the DCC profiles were not adjacent. We found that in some cases, there were only a few interrupted profiles (the cloud type of these profiles are not deep convection) between DCCs, and these interrupted profiles exhibited fairly thick clouds. This could be due to precipitation events or other factors preventing these profiles from meeting the thresholds defined for deep convection in 2B-CLDCLASS-LIDAR. We do not count the DCCs within the same DCS separately but consider them together. For example, if a DCS contains 50 DCC profiles, even if they are not contiguous, such as 20 DCC profiles followed by a gap and then another 30 DCC profiles, the width of the DCC in this DCS is
considered as 50*1.1 km (the horizontal spacing of the profiles). When matching meteorological factors at the locations of DCCs in Section 3.3, the average of the meteorological factor data at the locations of all DCC profiles within a DCS will be matched to that DCS, even if the DCC profiles are interrupted.

12. L232: It should also be discussed what the differences in convection between land and ocean are, particularly with respect to the diurnal cycle. Since land and ocean have different diurnal cycles of convection and precipitation, and that CloudSat-CALIPSO are confined to only twice-daily measurements that do not capture the full diurnal cycle, how might this influence the results that you are seeing? For example, how might the diurnal cycle be influencing the differences in the frequency and structure of convection between TO and TP?

Response: Thank you very much for the reviewer's helpful comments. The different diurnal cycles of convection between land and ocean is indeed a noteworthy topic. Due to the CloudSat switch to DO-Op in the late 2011 and only daytime (~1:30 p.m.) data is available after 2011. We recalculate the spatial statistics of daytime (~1:30 p.m.) and nighttime (~1:30 a.m.) DCSs based on 2B-CLDCLASS-LIDAR from 2006 to 2011. The results are as follows:

<table>
<thead>
<tr>
<th>Region</th>
<th>Sample number</th>
<th>Width of DCSs (km) / SD</th>
<th>Width of DCCs (km) / SD</th>
<th>Width of anvil (km) / SD</th>
</tr>
</thead>
<tbody>
<tr>
<td>TO</td>
<td>285</td>
<td>612.4/564.6</td>
<td>54.0/62.6</td>
<td>558.4/542.3</td>
</tr>
<tr>
<td>TP (total)</td>
<td>111</td>
<td>201.4/192.4</td>
<td>21.3/15.5</td>
<td>180.1/186.3</td>
</tr>
<tr>
<td>TP (NW)</td>
<td>10</td>
<td>198.2/133.2</td>
<td>14.6/5.8</td>
<td>183.6/132.1</td>
</tr>
<tr>
<td>TP (NE)</td>
<td>18</td>
<td>136.2/148.0</td>
<td>21.6/18.0</td>
<td>114.6/134.3</td>
</tr>
<tr>
<td>TP (SW)</td>
<td>35</td>
<td>225.3/220.4</td>
<td>22.2/15.4</td>
<td>203.1/215.7</td>
</tr>
<tr>
<td>TP (SE)</td>
<td>48</td>
<td>209.1/195.2</td>
<td>22.0/16.0</td>
<td>187.1/188.7</td>
</tr>
</tbody>
</table>

<table>
<thead>
<tr>
<th>Region</th>
<th>DCCs /DCSs (%)</th>
<th>Thickness of DCCs (km) / SD</th>
<th>DCCs/DCSs penetrating tropopause (%)</th>
<th>Mean precipitation of DCCs (mm hr⁻¹)</th>
</tr>
</thead>
<tbody>
<tr>
<td>TO</td>
<td>14.5</td>
<td>14.4/1.3</td>
<td>31.6/45.3</td>
<td>3.4</td>
</tr>
<tr>
<td>TP (total)</td>
<td>18.3</td>
<td>9.7/1.3</td>
<td>14.4/26.1</td>
<td>0.9</td>
</tr>
<tr>
<td>TP (NW)</td>
<td>14.0</td>
<td>9.1/1.3</td>
<td>20.0/20.0</td>
<td>0.2</td>
</tr>
<tr>
<td>TP (NE)</td>
<td>22.5</td>
<td>9.9/1.3</td>
<td>22.2/27.8</td>
<td>1.1</td>
</tr>
</tbody>
</table>
Table R1: The spatial statistics of daytime (~1:30 p.m.) DCSs during 2006-2011 in different subregions. The definition of different parts of TP are as follows: TP(NW) (33.5°N–37°N, 78°E–90.5°E); TP(NE) (33.5°N–37°N, 90.5°E–103°E); TP(SW) (30°N–33.5°N, 78°E–90.5°E); TP(SE) (30°N–33.5°N, 90.5°E–103°E). SD is an abbreviation for standard deviation.

<table>
<thead>
<tr>
<th>Region</th>
<th>Sample number</th>
<th>Width of DCSs (km)</th>
<th>Width of DCCs (km)</th>
<th>Width of anvil (km)</th>
</tr>
</thead>
<tbody>
<tr>
<td></td>
<td></td>
<td>(km) / SD</td>
<td>(km) / SD</td>
<td>(km) / SD</td>
</tr>
<tr>
<td>TO</td>
<td>357</td>
<td>725.4/592.6</td>
<td>72.1/73.8</td>
<td>653.3/574.5</td>
</tr>
<tr>
<td>TP (total)</td>
<td>10</td>
<td>974.3/856.0</td>
<td>37.0/30.4</td>
<td>937.3/865.8</td>
</tr>
<tr>
<td>TP (NW)</td>
<td>0</td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>TP (NE)</td>
<td>0</td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>TP (SW)</td>
<td>3</td>
<td>775.1/719.8</td>
<td>45.5/8.3</td>
<td>729.7/723.3</td>
</tr>
<tr>
<td>TP (SE)</td>
<td>7</td>
<td>1059.6/947.7</td>
<td>33.3/36.3</td>
<td>1056.3/958.8</td>
</tr>
</tbody>
</table>

Table R2: The spatial statistics of nighttime (~1:30 a.m.) DCSs during 2006-2011 in different subregions. The definition of different parts of TP are as follows: TP(NW) (33.5°N–37°N, 78°E–90.5°E); TP(NE) (33.5°N–37°N, 90.5°E–103°E); TP(SW) (30°N–33.5°N, 78°E–90.5°E); TP(SE) (30°N–33.5°N, 90.5°E–103°E). SD is an abbreviation for standard deviation.

In terms of frequency, there are more DCSs occur at ~1:30 a.m. over the TO, whereas DCSs over the TP primarily occurs during daytime (~1:30 p.m.). The difference in DCS daytime and nighttime frequency is quite pronounced between the TP and the TO. However, due to the limited number of nighttime DCSs over the TP (only 10), the representativeness of their width, thickness, and precipitation statistics is very limited. The results show that, compared to daytime, nighttime DCSs on the TP are wider, thicker, and have more precipitation. On the contrary,
these parameters are greater during nighttime than during daytime in the TO. As an aside, the total sample size of the results in Table R1 and R2 is slightly larger than the sample size of results in Section 3.2 (TP-116; TO-623, shown in L115 in the manuscript). For several DCS samples, the corresponding 2B-FLXHR-lidar data files could not be available.)

Since the only twice-daily measurements CloudSat and CALIPSO cannot capture the full diurnal cycle, the different diurnal cycles of DCS in the TP and TO regions are bound to affect the results to some extent. Previous studies found that the occurrence frequency of deep convection reach the daily maximum at around 16:00-18:00 LT (Devasthale and Fueglistaler, 2010; Xu and Zipser, 2011). Shown in the results based on ISCCP brightness temperature in Kottayil et al. (2021) (Fig. R1), the diurnal peak time of deep convection occurrence frequency over Indian ocean occurs between ~5:00 and ~12:00 LT. And they found that the diurnal amplitude is smaller at sea than over land. Although the subjects of the referenced studies do not exactly match the DCS in our study, the diurnal cycle results of deep convection can provide some degree of reference. Overall, the diurnal cycle of convection has a greater impact on the twice-daily measurements CloudSat and CALIPSO results in the TP region due to the larger diurnal amplitude over land. The analysis of difference in DCS daytime and nighttime frequency is added in “Results (Section 3.1)” and the limitations of twice-daily measurements are added in “Conclusions and discussion”.

See the lines: 266-277 and 659-666.
13. **L248-250**: Why does deep convection over the TP contribute so much stratospheric pollutants?

**Response**: The tropics are far from the main sources of biomass and biofuel burning pollutants. And the pollutants have a strong sink from contact with the ocean. The results based on satellites observations show that much of the air in the tropical upper troposphere is relatively depleted in hydrogen cyanide. For these reasons, although the transport of air from the troposphere to the stratosphere occurs primarily in the tropics, associated with the ascending branch of the Brewer-Dobson circulation, the tropical upwelling cannot be the main source. The Tibetan Plateau is the heat source of Asian summer monsoon, and the strong convergence near the surface in summer is conducive to the transport of pollutants. And the Asian summer monsoon circulation contains a strong anticyclonic vortex in the upper troposphere and lower stratosphere. As evidenced by satellite observations, a mean upward circulation on the eastern side of the anticyclone extends the transport into the lower stratosphere. The cross-tropopause high mixing ratio of hydrogen cyanide near 30°N also show the strong transport of pollution over the
Tibetan Plateau. The above evidence and analysis are from Randel et al. (2010). To clarify, we revised the expression.

**See the lines:** 290-293.

14. *Table 2*: are these exclusively single-cell DCSs? Also, I would suggest switching the order of the TP (total) and TO so that the TP regions are consecutive.

**Response:** In our study, we did not distinguish between single-cell and multi-cell DCSs. The DCSs we used include both cases. If discontinuous DCC profiles appear within the same DCS, they will be integrated for analysis together. For the calculation method of DCC statistics in multi-cell DCSs, please refer to our response to Comment #11. Currently, our differentiation of components (DCC and anvils) within DCS is based on whether they are identified as deep convection in 2B-CLDCLASS-LIDAR, which is a relatively simple criterion. We have not conducted detailed research into whether DCS are single-cell. This may require future study to utilize additional auxiliary data and conduct more in-depth investigations. We appreciate the reviewer's suggestion, and we have swapped the order of TP (total) and TO in Table 2 in the revised manuscript.

15. *L302*: How are you getting these LW and SW CRE values? The SW CRE is from daytime-only at the 1:30 pm overpass, correct? And are you averaging over the daytime and nighttime overpasses for the LW CRE? Since the SW CRE is calculated from radiative fluxes at 1:30 pm local solar time, these measurements are more enhanced than at other times of day due to the near peak in solar insolation at this time. To capture cloud radiative effects that are more representative of what they would be throughout the day, you would need to multiply the SW radiative fluxes by the diurnally averaged insolation for that day and location (L'Ecuyer et al., 2019) and then recalculate the net CRE.

**Response:** The LW CRE value is from averaging both 1:30 p.m. and 1:30 a.m. overpass, and SW CRE value is from only 1:30 p.m. overpass. Thank you very
much for the reviewers' suggestions. Following the reviewers' comment and the method from L'Ecuyer et al. (2019), we recalculate the average shortwave fluxes as follows:

\[ F_{\text{daily}} = F_{\text{inst}} \times \frac{Q_{\text{daily}}}{Q_{\text{inst}}} \] (1)

where \( F_{\text{daily}} \) is the daily mean shortwave fluxes, \( F_{\text{inst}} \) is the instantaneous fluxes from 2B-FLXHR-LIDAR, \( Q_{\text{daily}} \) is the daily mean solar insolation, \( Q_{\text{inst}} \) is the incoming solar of each profile from 2B-FLXHR-LIDAR. We recalculate all the results in Section 3.2. The new results are as follows:

<table>
<thead>
<tr>
<th>Region</th>
<th>TOA NCRE (W m(^{-2}))</th>
<th>BOA NCRE (W m(^{-2}))</th>
<th>ATM NCRE (W m(^{-2}))</th>
</tr>
</thead>
<tbody>
<tr>
<td></td>
<td>DCC</td>
<td>DCS</td>
<td>anvil</td>
</tr>
<tr>
<td>TP</td>
<td>-51.8</td>
<td>-41.4</td>
<td>-33.2</td>
</tr>
<tr>
<td>TO</td>
<td>-80.8</td>
<td>-13.6</td>
<td>4.7</td>
</tr>
</tbody>
</table>

Table R1: The mean net cloud radiative effects (NCRE) of DCSs, DCCs and anvil in different regions. The estimated SW fluxes are normalized to the diurnally averaged insolation.

<table>
<thead>
<tr>
<th>Region</th>
<th>TOA CRE (W m(^{-2}))</th>
<th>BOA CRE (W m(^{-2}))</th>
</tr>
</thead>
<tbody>
<tr>
<td></td>
<td>SWCRE</td>
<td>LWCRE</td>
</tr>
<tr>
<td></td>
<td>DCC</td>
<td>DCS</td>
</tr>
<tr>
<td>TP</td>
<td>-205.3</td>
<td>-153.5</td>
</tr>
<tr>
<td>TO</td>
<td>-244.1</td>
<td>-123.7</td>
</tr>
</tbody>
</table>

Table R2. The SW CRE (W m\(^{-2}\)) and LW CRE (W m\(^{-2}\)) at TOA and BOA of DCSs, DCCs and anvil in different regions. The estimated SW fluxes are normalized to the diurnally averaged insolation.
Figure R2: The proportion of each bin in the total sample of DCC, DCS and anvil in the TP (blue), TO (red) of the net CRE at the top of the atmosphere (TOA) (a, d, g), at the bottom of the atmosphere (BOA) (b, e, h) and in the atmosphere (ATM) (c, f, i). The dash lines show the peak of DCC. The estimated SW fluxes are normalized to the diurnally averaged insolation.
Figure R3: The same to Figure R2, but the results of the SW CRE.

Figure R4: The same as Figure R2, but the results of the LW CRE.
Figure R5: The mean vertical profiles of longwave (a, d), shortwave (b, e), and net (c, f) cloud radiative heating rates (CRH) (K d$^{-1}$) in the TP and TO. The red lines represent the results of DCCs and the blue lines represent the results of DCSs. The shadow range represents the standard deviation. The estimated SW fluxes are normalized to the diurnally averaged insolation.

Figure R6. The mean vertical profiles of cloudy (all-sky) heating rates (K d$^{-1}$) of DCCs over the TP, and TO. The shadow range represents the standard deviation. The estimated SW fluxes are normalized to the diurnally averaged insolation.
This method roughly reduces the instantaneous observed SW flux to 40% of its original value. The relative magnitudes of the CRE of DCS, DCC, and anvils between the two regions remain unchanged. However, the absolute values of both the net CRE and the SW CRE have overall decreased. Due to the reduction in the absolute value of SW CRH after normalizing the SW fluxes, the net CRH no longer shows the vertical distribution similar to that of SW CRH but instead resembles that of LW CRH.

However, as mentioned in comment #12, twice-daily measurements cannot capture the full diurnal cycle of convective clouds. This method of normalizing the SW flux assumes that cloud cover remains constant throughout the day, as detected at 13:30 LT. In reality, the diurnal cycle of deep convection is significant. And individual DCSs have different life cycles. Therefore, we chose to present the instantaneous observation results of cloud radiative effects of DCSs in the main text. Additionally, we described the limitations of the twice-daily instantaneous
observations and the summary of the above results in the “Conclusions and discussion” section.

**See the lines:** 659-666.

16. *L315:* The sentence starting with "In particular" is the start of a new discussion on LW BOA CRE so I would suggest starting a new paragraph.

**Response:** Thank you very much for the reviewer’s suggestion. It is revised according to the suggestion.

**See the line:** 364.

17. *L388:* What does "the high value of CRH, the difference between all-sky and clear-sky heating rates" mean?

**Response:** We mean that the CRH is calculated as the all-sky heating rates minus clear-sky heating rates. Therefore, we can separately analyze the heating rates under all-sky and clear-sky conditions to elucidate the reasons for the high values of CRH. The description in our manuscript may cause misunderstanding for readers. In the revised manuscript, we have changed the sentence to: “CRH is calculated as the all-sky heating rates minus clear-sky heating rates. Therefore, the reasons for the high value of CRH can be analysed through the results of heating rates under all-sky and clear-sky conditions.”

**See the lines:** 442-444.

18. *L461:* What is the sample size of each bin for the different ACs in Figure 6?

**Response:** When analyzing the role of aerosols, we select the top 30% and bottom 30% of samples based on aerosol concentration. Each group is further divided into four bins. For each ACs group, the TP region has 17 samples per bin, while the TO region has 55 samples per bin. Limited sample sizes are unavoidable when studying cloud clusters rather than cloud profiles. Therefore, as noted in L461, the limited
sample size could potentially impact the results. Nevertheless, it is encouraging that despite the sample size of each bin is small, some results still demonstrate obvious correlations.

19. L604-605: Your reasoning would for the enhanced SW CRE of anvils over the Tibetan Plateau could be improved if you verify whether the cloud tops are more densely packed. Perhaps take a look at cloud optical depth, which can be found in the 2B-FLXHR-LIDAR data set.

Response: Thank you very much for the valuable suggestions from the reviewers. Cloud optical thickness indeed serves as a robust parameter for describing cloud density. Following the reviewer's suggestion, we have added the analysis of cloud optical thickness based on the parameters provided by 2B-FLXHR-LIDAR. The results are as follows:

The results indicate that the cloud optical depth of anvils over the TP is larger, though the thickness is thinner. This suggests that anvils over the TP are more densely packed. Furthermore, this finding further supports our discovery of stronger shortwave radiative cooling of the anvils over the TP. These results have
been included in the supplement and corresponding descriptions have been added to the revised manuscript.

See the lines: 388-391.

Technical Comments

1.  \textit{L25-26: “convective activity” instead of “convection activities”}

   \textbf{Response:} It is corrected in the revised manuscript.

   \textbf{See the line:} 26-27.

2.  \textit{L32-32: Modify this sentence to “Although deep convective cloud is less frequent compared to other cloud types, it has a more complicated vertical structure and larger extent, thus exerting a great influence on radiation and precipitation over the Tibetan Plateau region.”}

   \textbf{Response:} Thank you for the suggestion. It is corrected in the revised manuscript.

   \textbf{See the lines:} 33-35.

3.  \textit{L41: Remove “Such as”}

   \textbf{Response:} Revisions have been made according to the reviewer’s suggestion.

   \textbf{See the line:} 42.

4.  \textit{L51-52: “A complete deep convection system (DCS) should include both the deep convective core (DCC) and the anvils” is out of place and can be removed.}

   \textbf{Response:} This sentence is removed.

5.  \textit{L110 & 151: L'Ecuyer}
**Response:** Sorry, this is a citation formatting error. It is corrected in the revised manuscript.

**See the lines:** 119 and 171.

6. **L233:** "Since the TO"

   **Response:** It is revised according to the reviewer’s suggestion.

   **See the line:** 263.

7. **L239:** I think you mean DCC, not DCS

   **Response:** Yes, the thickness of the DCC is used to characterize the thickness of the entire DCS in our study. To enhance readability, we revised the expression to “the DCCs over the TP are thinner”.

   **See the line:** 279.

8. **L250-251:** In the tropical region? Please rephrase.

   **Response:** It is revised to “In the tropics”.

   **See the line:** 293.

9. **Fig S2-S3:** Are these mean values?

   **Response:** Yes, the results in Fig. S2 and Fig. S3 are mean values. This explanation is added in the figure captions.

10. **L363:** By "larger", do you mean less negative?

    **Response:** Yes, as the CRE is negative, “larger” means less negative. To improve readability, it is revised to “less negative”.

    **See the line:** 415.
11. L366-367: Are you comparing the anvil structure and radiative effects to the structure and radiative effects of the DCCs? Please clarify.

Response: Yes, we mean that compared with DCCs, the anvil clouds are thinner and higher. For this reason, the absolute value of SW CRE and LW CRE of DCSs, which is the average of the results of DCCs and anvils, is smaller than the results of DCCs. To enhance clarity, we have revised the sentence as follows: “Due to anvils being thinner and higher relative to DCCs, anvils reflect less solar radiation and emit less LW radiation. This results in the weaker CRE of DCSs compared to DCCs.”.

See the lines: 418-419.

12. L383: Do you mean to say "below" instead of "within"?

Response: Yes, we mean the LW CRH is positive “below” 10 km. It is corrected in the revised manuscript.

See the line: 436.

13. L385-386: This is an incomplete sentence; please modify.

Response: It is revised to “This peak of LW CRH near the surface is more distinct in the TP compared with the TO, due to the colder boundary layer emitting less upwelling LW radiation in the TP (Haynes et al., 2013).”

See the lines: 438-440.

14. L388: Is "height bins with only several valid data are not shown" a typo? What is the value of "several valid data"?

Response: We only show the results of height bins with more than 50 samples. The results below 4.56 km are not shown. The value is added in the revised manuscript.
15. L398: Remove comma after LW CRH

Response: It is corrected in the revised manuscript.

See the line: 452.

16. L463-465: This sentence is incomplete.

Response: Thank you for the reviewer's suggestion. The logical relationship between these two sentences is indeed somewhat confusing. The sentence "Although... reduced convection" is not closely related to the preceding or following text. To avoid confusing the readers, we have deleted and revised this part accordingly.

17. L555: Remove comma after While

Response: It is corrected in the revised manuscript.

See the line: 582.

Reference


