

1 **Non biogenic source is an important but**  
2 **overlooked contributor to aerosol isoprene-**  
3 **derived organosulfates during winter in**  
4 **northern China**

5

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21 **Abstract:** Previous measurement-model comparisons of atmospheric isoprene levels  
22 showed a significant unidentified source of isoprene in some northern Chinese cities  
23 during winter. Here, spatial variability in winter aerosol organosulfate (OS) formation  
24 in typical southern (Guangzhou and Kunming) and northern (Xi'an and Taiyuan)  
25 cities, China, was investigated to reveal the influence of potential non biogenic  
26 contributor on aerosol OS pollution levels. Monoterpene-derived OSs were  
27 significantly higher in southern cities than in northern cities, which was attributed to  
28 temperature dependent emission of monoterpenes (i.e., higher temperatures in  
29 southern cities drove more monoterpene emissions). However, isoprene-derived OSs  
30 ( $OS_i$ ) showed the opposite trend, with significantly higher levels in northern cities.  
31 Principal component analysis combined with field simulation combustion experiments  
32 suggested that biomass burning rather than gasoline, diesel, and coal combustion  
33 contributed significantly to the abundance of  $OS_i$  in northern cities. The comparison  
34 of anthropogenic OS molecular characteristics between particles released from  
35 various combustion sources and ambient aerosol particles suggested that stronger  
36 biomass and fossil fuel combustion activities in northern cities promoted the  
37 formation of considerable anthropogenic OSs. Overall, this study provides direct  
38 molecular evidence for the first time that non biogenic sources can significantly  
39 contribute to the formation of  $OS_i$  in China during winter.

40

41 **Keywords:** Aerosol organosulfates, Biogenic precursors, Anthropogenic precursors,  
42 Spatial variation, Influencing factors, Biomass burning

43 **1. Introduction**

44 Organosulfates (OSs) with a sulfate ester functional group typically contribute 3–  
45 30% of the organic aerosol mass in atmospheric fine particles (PM<sub>2.5</sub>) (Luk'Ac's et al.  
46 2009). Moreover, OSs have been estimated to account for up to 12% of the total sulfur  
47 mass in fine particles, playing significant roles in the global biogeochemical cycling  
48 of sulfur (Luk'Ac's et al. 2009). In particular, OSs can impact the properties of  
49 aerosols, such as hygroscopicity, acidity, viscosity, and morphology, which are closely  
50 associated with the organic aerosol formation and urban air quality (Riva et al. 2019;  
51 Fleming et al. 2019). Thus, aerosol OSs have attracted significant attention over the  
52 years. However, the mechanisms and key factors impacting the formation and  
53 abundance of aerosol OSs in the real world remain considerable uncertainty, despite  
54 the important insights gained from laboratory simulation experiments (Wang et al.  
55 2021; Yang et al. 2023; Wang et al. 2020).

56 Previous field studies have indicated that acidity (Duporté et al. 2019), sulfate  
57 (Aoki et al. 2020), aerosol liquid water (Duporté et al. 2016), and oxidants (e.g.,  
58 ozone) (Wang et al. 2021) represent critical factors controlling the formation of OSs  
59 via heterogeneous and liquid phase processes (Brüggemann et al. 2020b). Precursor  
60 emission intensities (e.g., isoprene, monoterpenes, polycyclic aromatic hydrocarbons,  
61 and alkanes) also play an important role in impacting abundance of biogenic and  
62 anthropogenic OSs in ambient aerosols (Wang et al. 2022; Bryant et al. 2021; Yang et  
63 al. 2024). Furthermore, previous studies have identified a large number of CHOS  
64 compounds in smoke particles (e.g., pine branches, corn straw, rice straw, and coal)

65 (Song et al. 2019; Song et al. 2018; Tang et al. 2020). However, limited studies have  
66 focused on the contribution of different smoke particles to urban aerosol OSs. This  
67 may be an overlooked source of OSs. In general, few field studies have conducted a  
68 comprehensive investigation into the relationship between biogenic and  
69 anthropogenic impacting factors and regional differences in aerosol OS pollution.  
70 This complicates our understanding of how aerosol OS pollution is formed and what  
71 limits it in a complex polluted atmosphere across different cities in China.

72 The considerable variations in climatic conditions and air pollution levels in the  
73 northern and southern regions of China during winter (Ding et al. 2014; Ding et al.  
74 2016b) provide a distinctive opportunity to examine the complex influences of  
75 precursors, humidity, acidity, atmospheric oxidants, and anthropogenic pollution on  
76 the formation and abundance of aerosol OSs in the real world (Yang et al. 2024; Yang  
77 et al. 2023; Wang et al. 2021; Hettiyadura et al. 2019). In this study, we conducted the  
78 simultaneous observations of OSs and other chemical components in  $PM_{2.5}$  collected  
79 from typical southern (Guangzhou and Kunming) and northern (Xi'an and Taiyuan)  
80 cities in China during winter. Moreover, we also attempted to identify OSs in smoke  
81 particles emitted from combustion of different materials (i.e., rice straw, pine branch,  
82 diesel, gasoline, and coal). The principal aims of this study are 1) to investigate the  
83 spatial differences in aerosol OS pollution in northern and southern China during  
84 winter and 2) to elucidate the key factors that contribute to the spatial variability of  
85 OS pollution, with a focus on the OSs derived from smoke particles.

86

87 **2. Materials and Methods**

88 **2.1. Site description and sample collection**

89 The research sites are located in four urban areas in China, including Xi'an (XA)  
90 Taiyuan (TY), Guangzhou (GZ), and Kunming (KM) (**Figure S1a**). XA and TY are  
91 typical northern cities with cold winters (average temperature below 2 °C during the  
92 study period; **Table S1**). Thus, burning coal and biomass for heating is prevalent in  
93 these two cities during winter (Zhou et al. 2017; Ma et al. 2017), which significantly  
94 deteriorated the local air quality (**Figure S1b**). GZ and KM represent typical southern  
95 cities, with an average air temperature of over 10 °C during the winter sampling  
96 period (**Table S1**). Clearly, the distinctive climatic conditions in the northern and  
97 southern cities during winter may lead to significant spatial differences in the level of  
98 air pollution and the emission intensity of biogenic volatile organic compounds  
99 (VOCs) (Ding et al. 2014; Xu et al. 2024b).

100 From 10 December 2017 to 8 January 2018, sampling was performed  
101 simultaneously in four cities. Filters contained PM<sub>2.5</sub> were collected at regular two- to  
102 three-day intervals, with the collection duration being 24 hours, using a high-volume  
103 air sampler (Series 2031, Laoing, China) at a flow rate of  $\sim$ 1.05 m<sup>3</sup> min<sup>-1</sup> (Xu et al.  
104 2024a). A blank filter was sampled at each of the study sites. A total of 48 ambient  
105 samples were collected and stored at a temperature of  $-30^{\circ}\text{C}$ . Meteorological data,  
106 including wind speed, relative humidity (RH), and temperature, were obtained from  
107 nearby environmental stations. Concurrently, the concentrations of various pollutants,  
108 such as O<sub>3</sub>, NO<sub>2</sub>, and SO<sub>2</sub>, were also recorded.

109 **2.2. Smoke particle collection**

110 The controlled burning experiments conducted in the field were designed to  
111 simulate the emissions of “real world” burning cases in China (**Figure S2**), with the  
112 methodology being improved according to the previous reports (He et al. 2010; Wang  
113 et al. 2017). Rice straw and pine branch are typical materials for biomass burning in  
114 China (Zhou et al. 2017). In addition, the combustion of coal, gasoline, and diesel was  
115 representative of fossil fuel combustion (Yu et al. 2020). Accordingly, the smoke  
116 particles emitted from rice straw, pine branch, coal combustion, gasoline vehicle  
117 exhausts, and diesel vehicle exhausts were separately collected using self-made  
118 devices (**Figure S2**).

119 Briefly, the smoke from the combustion of rice straw, pine branch, and coal was  
120 sampled through a combustion furnace pumped with ambient air (particulate matter is  
121 removed) (**Figure S2a**). It should be noted that introducing ambient air with removed  
122 particulate matter into the combustion furnace is to minimize the pollution of ambient  
123 particulate matter to the smoke particle samples. Each combustion experiment for  
124 straw, pine branch, and coal lasted for 30–40 min. Regarding the smoke particles  
125 emitted from gasoline vehicle exhausts and diesel vehicle exhausts, they were  
126 collected for 3 hours by directly connecting to the car exhaust pipe (**Figure S2b**). All  
127 smoke particle samples are collected onto prebaked quartz fiber filters via a high-  
128 volume air sampler (Series 2031, Laoing, China). Four repeated experiments were  
129 conducted for each combustion material, one of which was collected as a blank  
130 sample. All smoke particle samples were stored at  $-30^{\circ}\text{C}$ .

131 **2.3. Chemical analysis and predictions of aerosol acidity and water concentration**

132 The extraction, measurement procedures, and identification of OSs were  
133 described in detail in our recent publications (Yang et al. 2024). Briefly, the filter  
134 sample was extracted using methanol, then filtered through a 0.22  $\mu\text{m}$  PTFE syringe  
135 filter and concentrated by a gentle stream of nitrogen gas. Subsequently, the  
136 concentrated sample with adding ultrapure water (300  $\mu\text{L}$ ) was thoroughly mixed  
137 using a mixer. The mixture was centrifuged to obtain the supernatant for analysis of  
138 UPLC-MS/MS system (Waters, USA) (Wang et al. 2021). The reverse-phase liquid  
139 chromatography (RPLC) method was performed on an Acquity UPLC HSS T3  
140 column (2.1mm  $\times$  100 mm, 1.8  $\mu\text{m}$  particle size; Waters, USA) in this study. Although  
141 our method is quite effective in retaining and separating low molecular weight (MW)  
142 OSs, as demonstrated in our recent publication (Yang et al. 2024), we also  
143 acknowledge that the developed hydrophilic interaction liquid chromatography  
144 method may provide another solution for the measurement of low-MW OSs (Cui et al.  
145 2018; Hettiyadura et al. 2015).

146 In addition, it has been indicated in previous studies (Brüggemann et al. 2020a;  
147 Kristensen et al. 2016) that the levels of OSs can be affected by the sampling  
148 procedure, especially when  $\text{SO}_2$  removal procedures are not employed. On the  
149 assumption that  $\text{SO}_2$  reacts with organics on filters to form OSs, similar processes  
150 must also occur on ambient particles prior to sampling. Moreover, there is currently  
151 no study evaluating the relative efficiency of OS generation in filters and ambient  
152 particles. Consequently, the possible consequences of sampling without denuding  $\text{SO}_2$

153 for the quantification of OSs were not taken into account in our studies (Yang et al.  
154 2024; Yang et al. 2023). In total, 212 OSs were identified. However, only 111 OS  
155 species were quantified using surrogate standards in this study (**Table S2** and **S3**). The  
156 study divided the several principal OS groups as follows: monoterpene-derived OSs  
157 (OS<sub>m</sub>), isoprene-derived OSs (OS<sub>i</sub>), C<sub>2</sub>–C<sub>3</sub> OSs (i.e., OSs with two or three carbon  
158 atoms), and anthropogenic OSs (i.e, aliphatic and aromatic OSs) (Yang et al. 2023).  
159 The terms "OS<sub>m</sub>" and "OS<sub>i</sub>" refer to organosulfates generated from monoterpene and  
160 isoprene, respectively. These compounds were generally classified as biogenic OSs  
161 due to their natural origin (Wang et al. 2021; Wang et al. 2018). The specific  
162 classification and quantification methods were detailed in our recent publications  
163 (Yang et al. 2023; Yang et al. 2024) and **Supporting Information**.

164 An additional portion of each filter was extracted using ultrapure water for  
165 determining the inorganic ions (Huang et al. 2023). The concentrations of SO<sub>4</sub><sup>2-</sup>, Ca<sup>2+</sup>,  
166 NO<sub>3</sub><sup>-</sup>, Na<sup>+</sup>, K<sup>+</sup>, Mg<sup>2+</sup>, Cl<sup>-</sup>, and NH<sub>4</sub><sup>+</sup> were analyzed using ICS5000+ ion  
167 chromatography (Thermo, USA) (Yang et al. 2024; Lin et al. 2023). The mass  
168 concentration of aerosol liquid water (ALW) and pH value were calculated by a  
169 thermodynamic model (ISORROPIA-II) in the forward mode and thermodynamically  
170 metastable state, which was detailed in our previous studies (Liu et al. 2023; Xu et al.  
171 2022; Xu et al. 2023b; Xu et al. 2020). The influence of OSs on ALW and pH was not  
172 taken into account in the present study due to their negligible contribution to the  
173 prediction outcomes, as indicated by Riva et al. (2019) and Yang et al. (2024).

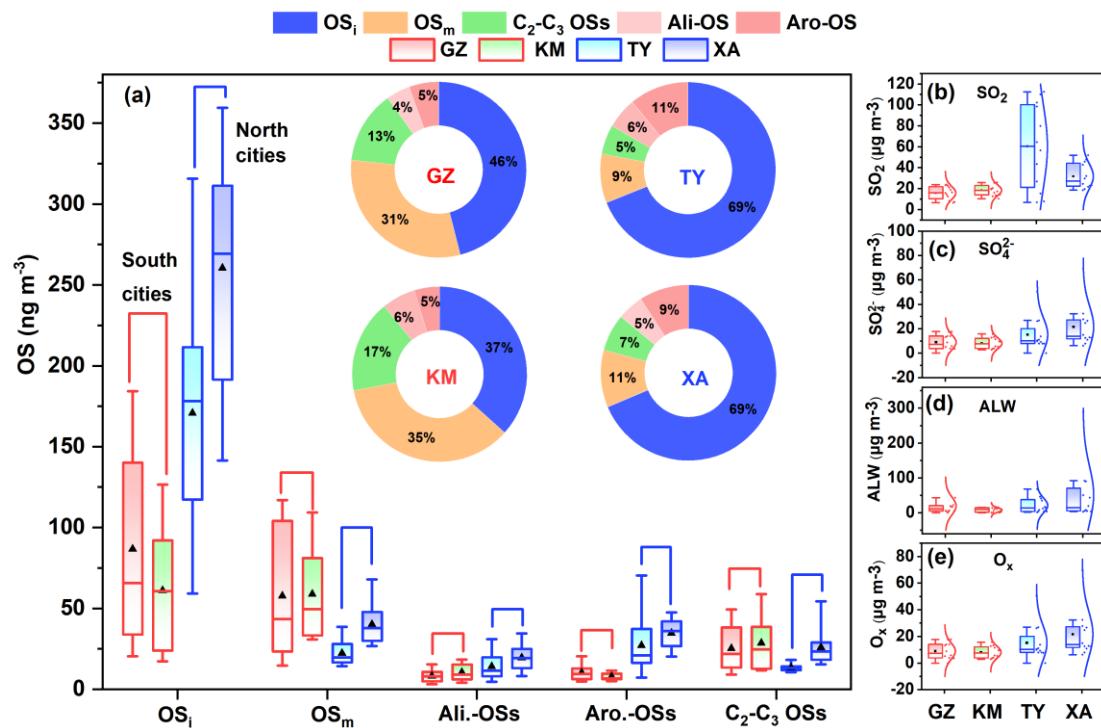
174

175 **3. Results and Discussion**

176 **3.1. Spatial variations in concentrations and compositions of different OSs**

177 **Figure 1a** shows the spatial distributions in mass concentrations and mass  
178 fractions of OS<sub>i</sub>, OS<sub>m</sub>, aliphatic OSs, aromatic OSs, and C<sub>2</sub>–C<sub>3</sub> OSs in PM<sub>2.5</sub> collected  
179 in southern (KM and GZ) and northern (TY and XA) cities during winter. On average,  
180 OS<sub>i</sub> was the dominant OS subgroup, which accounted for 37% – 46% and 68% – 69%  
181 of the total OS mass in southern and northern cities, respectively. The predominance  
182 of OS<sub>i</sub> in aerosol OSs was also reported by previous studies in cities in northern (e.g.,  
183 Beijing and Tianjin) (Wang et al. 2018; Ding et al. 2022) and southern (e.g.,  
184 Guangzhou and Shanghai) (Wang et al. 2022; Wang et al. 2021) China, as well as in  
185 coastal (the Yellow Sea and Bohai Sea) (Wang et al. 2023) and European (Sweden)  
186 (Kanellopoulos et al. 2022) and American regions (Chen et al. 2021; Hettiyadura et al.  
187 2017; Hettiyadura et al. 2019) (**Table S4**). Moreover, the concentrations of OS<sub>i</sub> were  
188 significantly lower in southern cities ( $61 \pm 38 \text{ ng m}^{-3}$  –  $87 \pm 60 \text{ ng m}^{-3}$ ) than in  
189 northern cities ( $171 \pm 69 \text{ ng m}^{-3}$  –  $260 \pm 71 \text{ ng m}^{-3}$ ) (**Table S1**), showing a  
190 concentration range overlapped with previous observations (**Table S4**). From southern  
191 to northern cities, the mass concentrations and mass fractions of OS<sub>m</sub> tended to  
192 decrease, which was opposite to the spatial variation pattern of OS<sub>i</sub> (**Figure 1a**). Both  
193 OS<sub>i</sub> and OS<sub>m</sub> are generally considered as typical biogenic OSs (Hettiyadura et al.  
194 2019; Wang et al. 2018), the abundances of which were tightly associated with  
195 biogenic VOC emissions when acidity, sulfate, atmospheric oxidation capacity, and  
196 ALW are not limiting factors (Bryant et al. 2021; Wang et al. 2022; Yang et al. 2024).

197 Thus, these dissimilarities in the spatial variations of  $OS_i$  and  $OS_m$  can be attributed to  
 198 large differences in the intensity of biogenic VOC emissions (Wang et al. 2022)  
 199 and/or the key factors that constrain OS formation between the northern and southern  
 200 regions of China (**Table S1**).



201 **Figure 1** Box and whisker plots showing the variations in the concentration of  
 202 different OS groups in PM<sub>2.5</sub> collected in southern (GZ and KM) and northern (TY  
 203 and XA) cities of China during winter. Each box encompasses the 25th–75th  
 204 percentiles. Whiskers are the minimum and maximum values. The triangles and solid  
 205 lines inside boxes indicate the mean and median. The spatial variation in average  
 206 abundance of anthropogenic OSs (i.e., OS<sub>a</sub>, including aliphatic and aromatic  
 207 OSs) was also shown in panel (a). The other figures show the spatial variations in (b) SO<sub>2</sub>, (c) SO<sub>4</sub><sup>2-</sup>, (d) ALW, and (e) O<sub>3</sub> levels.

209

210 The abundance of anthropogenic OSs (i.e., OS<sub>a</sub>, including aliphatic and aromatic

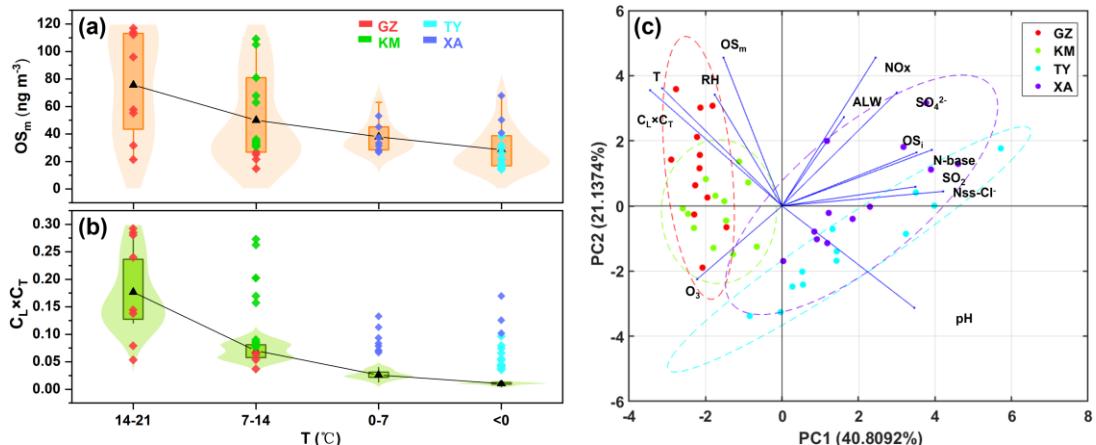
211 OSs, **Sect. S1**) in southern cities was lower than that of  $OS_m$ , which was opposite to  
212 the case in the northern cities showing higher anthropogenic OS abundance (**Figure**  
213 **1a** and **Table S1**). Moreover, we found that the spatial variation patterns of  $OS_i$  and  
214  $OS_a$  were similar to those of  $SO_2$ ,  $SO_4^{2-}$ , ALW, and  $O_x$  (**Figures 1b–e**), as indicated by  
215 significant ( $P < 0.05$ ) correlations of  $OS_i$  and  $OS_a$  with those factors (**Figure S3**).  
216 However,  $OS_m$  and  $C_2–C_3$  OSs showed the opposite spatial variation pattern to  $SO_2$ ,  
217  $SO_4^{2-}$ , ALW, and  $O_x$  (**Figure 1**). If both  $OS_i$  and  $OS_m$  are assumed to be formed mainly  
218 from the oxidation of biologically emitted VOCs, the higher  $SO_2$ ,  $SO_4^{2-}$ , ALW, and  $O_x$   
219 levels could theoretically lead to higher  $OS_m$  in northern cities, just as these factors  
220 leaded to higher  $OS_i$  abundance in northern cities (**Figure 1** and **Table S1**).  
221 Accordingly, the above differentiated spatial variation patterns among different OS  
222 subgroups likely indicated that other sources of isoprene contributed to the formation  
223 of  $OS_i$  in northern cities. Further given the significant ( $P < 0.05$ ) correlations between  
224  $OS_i$  and  $OS_a$ , non biogenic isoprene emissions may play an important role in the  
225 formation of aerosol  $OS_i$  in northern cities. This will be further demonstrated in the  
226 following discussion.

227

228 **3.2. Key factors affecting spatial differences in monoterpane-derived OS**  
229 **abundance**

230 **Figure 2a** shows the distribution of  $OS_m$  concentration as a function of air  
231 temperature. We found that the  $OS_m$  concentration tended to increase with the increase  
232 of air temperature. Specifically, the air temperature in the southern cities was mainly

233 in the range of 7–14°C during the sampling period, corresponding to higher aerosol  
234 OS<sub>m</sub> abundance. In contrast, the low temperature (< 7°C) in the northern cities  
235 corresponded to a significant decrease in OS<sub>m</sub> abundance. This finding was similar to  
236 the previously observed decrease in aerosol OS<sub>m</sub> compounds with decreasing  
237 temperature during winter in Guangzhou (Bryant et al. 2021). Furthermore, the  
238 indicator (C<sub>L</sub> × C<sub>T</sub>) of biogenic VOC emission rate (Ding et al. 2016a; Guenther et al.  
239 1993) was also higher in southern cities than in northern cities (**Figure 2b**), which  
240 implied higher monoterpene emissions in southern cities. It has been suggested that  
241 the emission rates of biogenic VOCs (e.g., monoterpene and isoprene) can be driven  
242 by increased air temperature and lighting (Ding et al. 2016a; Ding et al. 2016b). A  
243 previous study also found that the concentrations of atmospheric monoterpenes during  
244 the winter season were higher in warmer southern Chinese cities than in colder  
245 northern Chinese cities (Ding et al. 2016b; Li et al. 2020). In particular, GZ and KM,  
246 which encompass extensive areas of coniferous and broad-leaved forests, have been  
247 identified as hotspots for monoterpene and isoprene emissions (Li and Xie 2014).  
248 Considering the lower levels of key factors affecting OS formation observed in  
249 southern cities (**Figures 1b–e** and **Table S1**), it can be inferred that the significant  
250 spatial differences in OS<sub>m</sub> abundances were largely attributed to temperature  
251 dependent emission of monoterpenes.



252

253 **Figure 2** Distribution of (a)  $OS_m$  and (b)  $C_L \times C_T$  data in different temperature ranges  
 254 during winter. The triangles inside boxes indicate the mean. Principal component  
 255 analysis result (c) deciphering the relationship among  $OS_i$ ,  $OS_m$ , and key factors  
 256 influencing OS formation.

257

258 To further determine the key factors affecting the spatial differences of  $OS_m$ ,  
 259 principal component analysis was conducted (**Figure 2c**). It can be easily determined  
 260 that the abundance of aerosol  $OS_m$  was closely related to changes in air temperature  
 261 and  $C_L \times C_T$  value. This further explained the changes in  $OS_m$  data in the southern  
 262 cities. In contrast, the abundance of aerosol  $OS_i$  in the northern cities was more  
 263 influenced by anthropogenic factors, as indicated by combustion source tracers such  
 264 as nitrogen-containing bases (N-bases) and non-sea-salt  $Cl^-$  (nss- $Cl^-$ ) (Wang et al.  
 265 2017; Jiang et al. 2023) (**Figure 2c**). Thus, principal component analysis can perfectly  
 266 distinguish the main factors causing changes in  $OS_m$  and  $OS_i$  abundances between the  
 267 northern and southern cities. In general, the above results confirm that the spatial  
 268 variation of  $OS_m$  was predominantly controlled by temperature-related monoterpane

269 emissions. However, this cannot account for the observed spatial variation of OS<sub>i</sub>  
270 (**Figure 2c** and **Figure S4**). Interestingly, the spatial distribution patterns of OS<sub>m</sub> and  
271 OS<sub>i</sub> in northern and southern China exhibited consistency during summer, closely  
272 resembling the spatial distribution of biogenic VOC emission intensities (Wang et al.  
273 2022). Thus, this case together with our observations during winter further imply that  
274 non biogenic sources of isoprene were important contributors to the formation of OS<sub>i</sub>  
275 in northern China during winter.

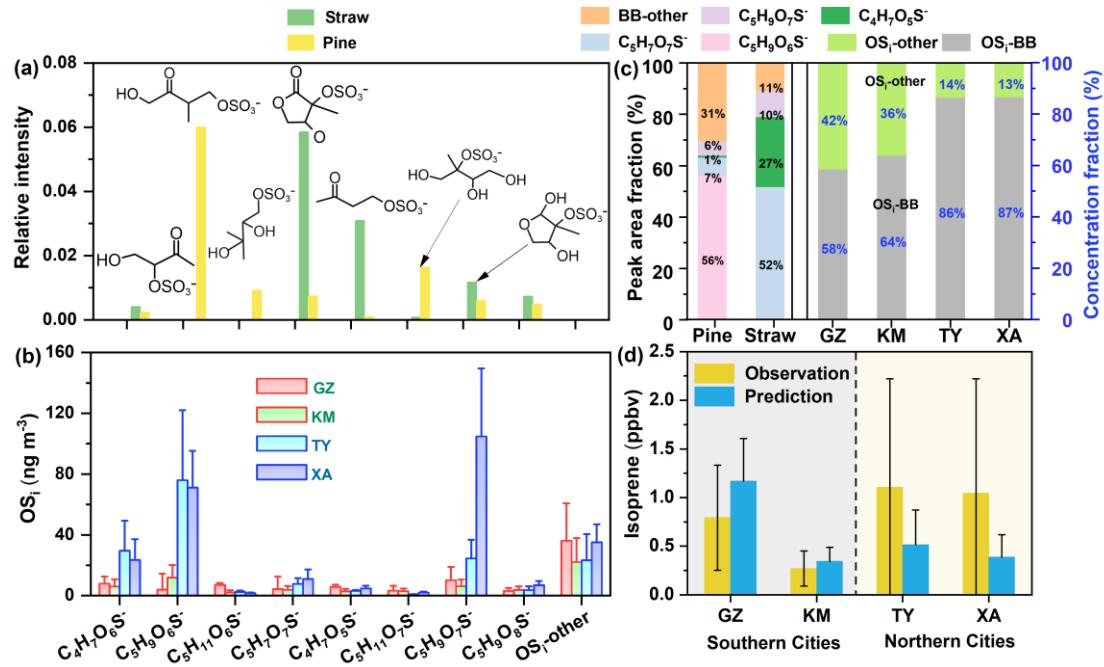
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277 **3.3. Significant contribution of biomass burning to isoprene-derived OSs in**  
278 **Northern China**

279 The previous principal component analysis has suggested that the abundance of  
280 OS<sub>i</sub> in northern cities was closely related to the levels of combustion source tracers  
281 (e.g., N-base compounds and nss-Cl<sup>-</sup>). N-base compounds are CHN species that  
282 contain exclusively C, H, and N atoms, and have been demonstrated to exhibit high  
283 sensitivity as molecular indicators in identifying biomass burning (Wang et al. 2017).

284 To further confirm the potential contribution of combustion release to aerosol OS<sub>i</sub>,  
285 OSs in smoke particles emitted from rice straw, pine branch, and coal combustion, as  
286 well as from gasoline vehicle exhausts, and diesel vehicle exhausts (**Figure S2**), were  
287 investigated. A total of 8 distinct OS<sub>i</sub> were identified in both the smoke particles  
288 emitted from biomass burning (rice straw and pine branch) and the ambient aerosol  
289 particles, including C<sub>4</sub>H<sub>7</sub>O<sub>6</sub>S<sup>-</sup>, C<sub>5</sub>H<sub>9</sub>O<sub>6</sub>S<sup>-</sup>, C<sub>5</sub>H<sub>11</sub>O<sub>6</sub>S<sup>-</sup>, C<sub>5</sub>H<sub>7</sub>O<sub>7</sub>S<sup>-</sup>, C<sub>4</sub>H<sub>7</sub>O<sub>5</sub>S<sup>-</sup>,  
290 C<sub>5</sub>H<sub>11</sub>O<sub>7</sub>S<sup>-</sup>, C<sub>5</sub>H<sub>9</sub>O<sub>7</sub>S<sup>-</sup>, and C<sub>5</sub>H<sub>9</sub>O<sub>8</sub>S<sup>-</sup>. Moreover, the peak intensities of these 8 OS<sub>i</sub>

291 in smoke particles emitted from fossil fuel combustion (gasoline and diesel vehicle  
292 exhausts and coal) were close to those in the blank sample. A previous investigation  
293 into CHOS compounds in smoke particles emitted from residential coal combustion  
294 and biomass burning also failed to identify OS<sub>i</sub> species (Song et al. 2019; Song et al.  
295 2018), which further supported the reliability of the combustion experiment  
296 conducted in this study. C<sub>5</sub>H<sub>9</sub>O<sub>6</sub>S<sup>-</sup> was dominant OS<sub>i</sub> species in pine-derived smoke  
297 particles (**Figure 3a,c**). We found that the average concentration of C<sub>5</sub>H<sub>9</sub>O<sub>6</sub>S<sup>-</sup> in  
298 ambient aerosol samples was much higher in northern cities than in southern cities  
299 (**Figure 3b**). A reasonable explanation for this is that pine branches are commonly  
300 used as solid fuel for heating and cooking in northern suburbs and rural areas (Zhou et  
301 al. 2017). C<sub>5</sub>H<sub>7</sub>O<sub>7</sub>S<sup>-</sup> and C<sub>4</sub>H<sub>7</sub>O<sub>5</sub>S<sup>-</sup> dominated OS<sub>i</sub> species in straw-derived smoke  
302 particles (**Figure 3a,c**). However, these two types of OS<sub>i</sub> have relatively low  
303 abundance in ambient aerosol samples in both northern and southern cities. This may  
304 be attributed to the fact that straw burning was mainly concentrated in autumn rather  
305 than winter in China (Zhou et al. 2017; Yang et al. 2015). On average, the biomass  
306 burning-related OS<sub>i</sub> accounted for 58% – 64% and 86% – 87% of the total OS<sub>i</sub>  
307 concentration in southern and northern cities, respectively (**Figure 3c**). Although  
308 these biomass burning-related OS<sub>i</sub> can also be formed through atmospheric  
309 transformation of biogenic isoprene, the higher proportion of these OS<sub>i</sub> in northern  
310 cities together with previous principal component analysis results still support our  
311 previous consideration that non biogenic OS<sub>i</sub> may be an important contributor to  
312 aerosol OS<sub>i</sub> in northern cities.



313

314 **Figure 3** Relative signal intensity of (a) identified major OS<sub>i</sub> species in different types  
 315 of smoke particle samples. Spatial variation in the concentration of several specific  
 316 OS<sub>i</sub> (identified in smoke particles) in (b) ambient PM<sub>2.5</sub> samples. Peak area and  
 317 concentration fraction of (c) OS<sub>i</sub> species identified in both ambient PM<sub>2.5</sub> samples  
 318 collected in different cities and smoke particles. Comparison of (d) isoprene mixing  
 319 ratios obtained from observation and modeling in different cities (Zhang et al. 2020).

320

321 Previous laboratory studies have suggested that these identified OS<sub>i</sub> species in  
 322 biomass burning-derived smoke particles are typically formed through heterogeneous  
 323 and multiphase reactions involving isoprene, its oxidation intermediates, and sulfate  
 324 or sulfur dioxide (Surratt et al. 2008; Surratt et al. 2007; Darer et al. 2011).  
 325 Specifically, C<sub>5</sub>H<sub>9</sub>O<sub>6</sub>S<sup>-</sup>, as a sulfate ester of C<sub>5</sub>-alkene triols, was formed mainly  
 326 through the uptake of gas-phase isoprene oxidation products onto acidified sulfate  
 327 aerosol (Surratt et al. 2007). The formation of C<sub>5</sub>H<sub>7</sub>O<sub>7</sub>S<sup>-</sup> and C<sub>5</sub>H<sub>9</sub>O<sub>7</sub>S<sup>-</sup> begins with

328 the gas-phase oxidation of isoprene (Surratt et al. 2008).  $\text{C}_4\text{H}_7\text{O}_6\text{S}^-$  can be generated  
329 both from isoprene photooxidation and sulfate radical reaction with methacrolein  
330 (MACR) or methyl vinyl ketone (MVK) (Schindelka et al. 2013; Wach et al. 2019;  
331 Nozière et al. 2010).  $\text{C}_5\text{H}_{11}\text{O}_7\text{S}^-$  was produced by reactive uptake of isoprene-derived  
332 epoxide (IEPOX) on sulfate under low-NO<sub>x</sub> conditions. Since our combustion  
333 experiments have excluded the direct contribution of ambient aerosol particles to OS<sub>i</sub>  
334 in smoke particles, it can be expected that these detected OS<sub>i</sub> compounds were mainly  
335 generated within smoke plumes through the isoprene oxidation pathway mentioned  
336 above. It has been demonstrated that directly emitted organic aerosols or VOCs can  
337 undergo a chemical reaction within smoke plumes, forming secondary organic  
338 compounds within a matter of hours (Wang et al. 2017; Song et al. 2018; Mason et al.  
339 2001). A field study conducted by Zhu et al. (2016) at a rural site (Yucheng) in the  
340 North China Plain (NCP) region has observed that the concentration of ambient  
341 isoprene during the period of straw combustion was approximately twice as high as  
342 that observed during periods of non combustion. In addition, Li et al. (2018) found  
343 that isoprene-derived epoxides increased significantly during field open burning of  
344 straw. Generally, despite the fact that a few of the mechanisms by which OSs are  
345 formed have been verified through field studies, the formation of CHOS and CHONS  
346 compounds has been observed to occur in the biomass burning plume (Zhang et al.  
347 2024; Song et al. 2018; Tang et al. 2020). Thus, these previous case studies further  
348 support our consideration that OS<sub>i</sub> compounds formed in biomass burning-derived  
349 smoke particles in this study can be attributed to increasing isoprene emission caused

350 by field biomass burning (Zhu et al. 2016) and favorable aqueous secondary organic  
351 aerosols (SOA) formation during the aging process of the biomass burning plume  
352 (Gilardoni et al. 2016).

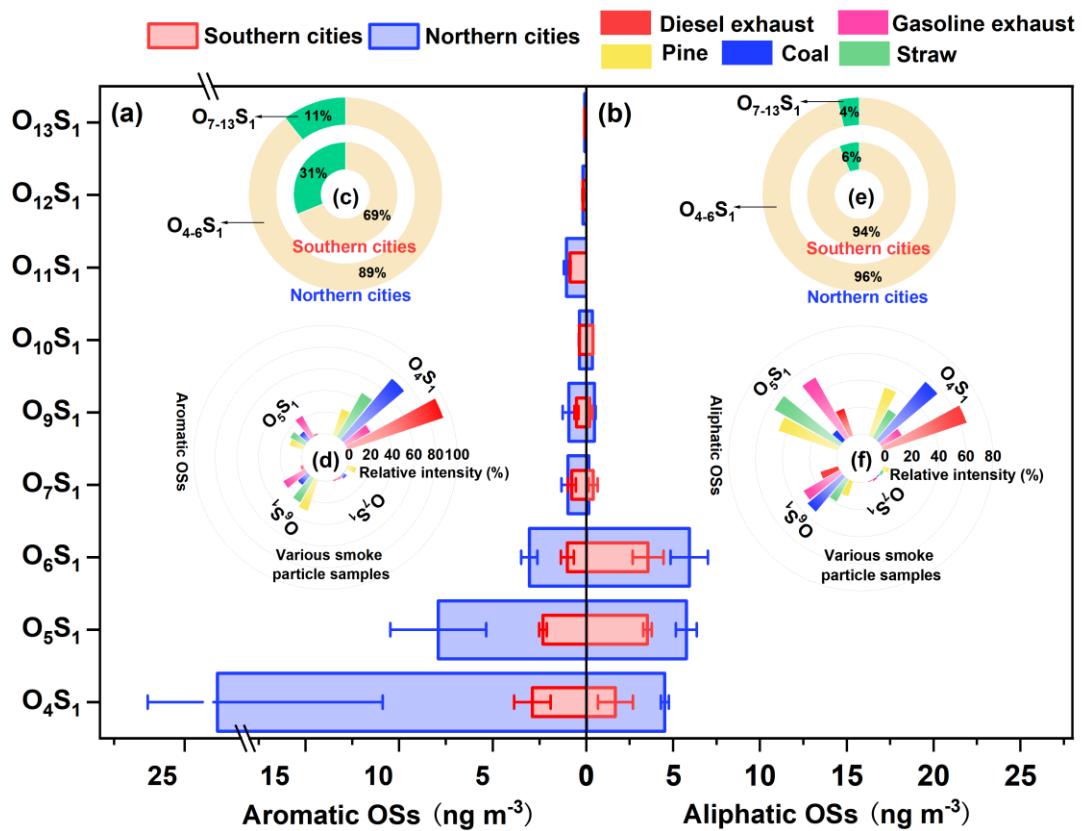
353 **Figure 3d** presents a comparison between the isoprene mixing ratios derived  
354 from model simulations (plant functional type related model) and those observed in  
355 the field in different Chinese cities during winter (December and January) (Zhang et  
356 al. 2020). Overall, the levels of isoprene observed in northern cities during winter  
357 were higher than those in southern cities. In addition, the predicted values in southern  
358 cities were slightly higher than the observed values, which may be attributed to the  
359 lag in model prediction results caused by the rapid urbanization rates in these southern  
360 cities (Zhang et al. 2020). However, the observed values in these two northern cities  
361 were 53% to 63% higher than the predicted values, on average. Clearly, this plant  
362 functional type related isoprene prediction model cannot explain the large amount of  
363 “missing” isoprene sources in northern cities. Thus, the observed spatial differences in  
364 OS<sub>i</sub> (**Figure 1**) and field combustion experiments (**Figure 3**) can suggest that these  
365 “missing” isoprene sources were mainly derived from biomass burning, significantly  
366 contributing to the production of aerosol OS<sub>i</sub> in northern cities. This can also be  
367 supported by previous principal component analysis involving combustion source  
368 tracers and OS<sub>i</sub> compounds (**Figure 2c**).

369

370 **3.4. Formation of anthropogenic OSs mainly driven by fossil fuel and biomass**  
371 **combustion**

372 **Figures 4a,b** show the average concentration distribution of anthropogenic OSs  
373 classified based on the number of O atoms in their molecules in southern (GZ and  
374 KM) and northern (TY and XA) cities. The O<sub>4</sub>S<sub>1</sub> subgroup was the most abundant  
375 aromatic OSs in both southern and northern cities, among which C<sub>9</sub>H<sub>9</sub>O<sub>4</sub>S<sup>-</sup>, phenyl  
376 sulfate (C<sub>6</sub>H<sub>5</sub>O<sub>4</sub>S<sup>-</sup>), and benzyl sulfate (C<sub>7</sub>H<sub>7</sub>O<sub>4</sub>S<sup>-</sup>) were dominant species (**Table S3**).  
377 C<sub>7</sub>H<sub>7</sub>O<sub>4</sub>S<sup>-</sup> and C<sub>6</sub>H<sub>5</sub>O<sub>4</sub>S<sup>-</sup> have been suggested to be formed mainly through the  
378 photooxidation of 2-methylnaphthalene and naphthalene (Riva et al. 2015), or  
379 alternatively, by the sulfate radical reaction with aromatic compounds, including  
380 toluene and benzoic acid, in an aqueous phase environment (Riva et al. 2015). The  
381 formation mechanism of C<sub>9</sub>H<sub>9</sub>O<sub>4</sub>S<sup>-</sup> is rarely reported. However, C<sub>9</sub>H<sub>9</sub>O<sub>4</sub>S<sup>-</sup>, C<sub>6</sub>H<sub>5</sub>O<sub>4</sub>S<sup>-</sup>,  
382 and C<sub>7</sub>H<sub>7</sub>O<sub>4</sub>S<sup>-</sup> were also detected in both fossil fuel combustion-derived smoke  
383 particles and biomass burning-derived smoke particles (**Figure S5** and **Table S5**),  
384 indicating that the aromatic VOCs produced by fuel combustion are closely related to  
385 the formation of these aromatic OSs. Overall, aerosol aromatic OS compounds in both  
386 southern and northern cities were mainly distributed between four and six O atoms  
387 (**Figure 4c**), which was similar to the distribution of aromatic OSs identified in  
388 various smoke particles emitted from different combustion sources (**Figure 4d**).  
389 However, the average abundances of aromatic O<sub>4-6</sub>S<sub>1</sub> compounds in northern cities  
390 were 3–6 times higher than those in southern cities. The above results suggest that  
391 aromatic OSs originated from fossil fuel and biomass combustion activities are  
392 important contributors to urban aerosol anthropogenic OSs in winter in China,  
393 especially in northern cities. We found that the correlations between aromatic OSs and

394 anthropogenic indicators ( $\text{SO}_2$ ,  $\text{SO}_4^{2-}$ , N-base, and nss- $\text{Cl}^-$ ) were stronger in northern  
 395 cities than in southern cities (**Figure S6**), and that the release of polycyclic aromatic  
 396 hydrocarbons from fossil fuel combustion was also higher in northern cities (**Figure**  
 397 **S7**). This further indicates that higher aerosol aromatic OSs in northern cities was  
 398 mainly attributed to stronger combustion activities in those cities.



399  
 400 **Figure 4** Concentration distribution of different (a) aromatic and (b) aliphatic OS  
 401 subgroups (classification based on oxygen atoms) in southern and northern cities.  
 402 Ring charts (c,e) show the percentage contributions of  $\text{O}_{4-6}\text{S}_1$  and  $\text{O}_{7-13}\text{S}_1$  subgroups.  
 403 Radial bar charts (d,f) illustrate the relative signal intensity of different OS subgroups  
 404 in different smoke particle samples.

405

406 Aliphatic OSs were also predominantly distributed between  $\text{O}_4\text{S}_1$  and  $\text{O}_6\text{S}_1$

407 subgroups in both southern and northern cities (**Figures 4b,e**), which was similar to  
408 the case found in both fossil fuel combustion-derived smoke particles and biomass  
409 burning-derived smoke particles (**Figure 4f**). It has been suggested that the long-chain  
410 alkanes derived from traffic emissions can largely contribute to the formation of  
411 CHOS compounds with aliphatic carbon chains (Tao et al. 2014). In addition, Tang et  
412 al. (2020) analyzed the molecular compositions of smoke particles from open biomass  
413 burning, household coal combustion and vehicle emissions and suggested that the  
414 aliphatic CHOS compounds can be derived from both vehicle emissions and coal and  
415 biomass combustion. In this study, aliphatic OSs showed a significant ( $P < 0.05$ )  
416 positive correlation with nss-Cl<sup>-</sup>, SO<sub>2</sub>, NO<sub>x</sub>, and N-base compounds in both southern  
417 and northern cities (**Figure S6 and S8**), indicating aerosol aliphatic OSs were affected  
418 by a combination of biomass burning and vehicle emissions in those cities during  
419 winter. Thus, the significantly higher level of aliphatic O<sub>4-6</sub>S<sub>1</sub> species in northern  
420 cities indicated that the formation of aliphatic OSs in northern cities was more driven  
421 by pollutants released from the combustion of fossil fuels and biomass compared to  
422 southern cities. This consideration is highly consistent with the fact that the  
423 concentrations of air pollutants (e.g., SO<sub>2</sub> and NO<sub>2</sub>) in northern cities with a large  
424 demand for heating during winter are usually higher than those in warmer southern  
425 cities (**Table S1 and Figure S1b**) (Yu et al. 2020; Ding et al. 2017; Ma et al. 2017;  
426 Zhou et al. 2017).

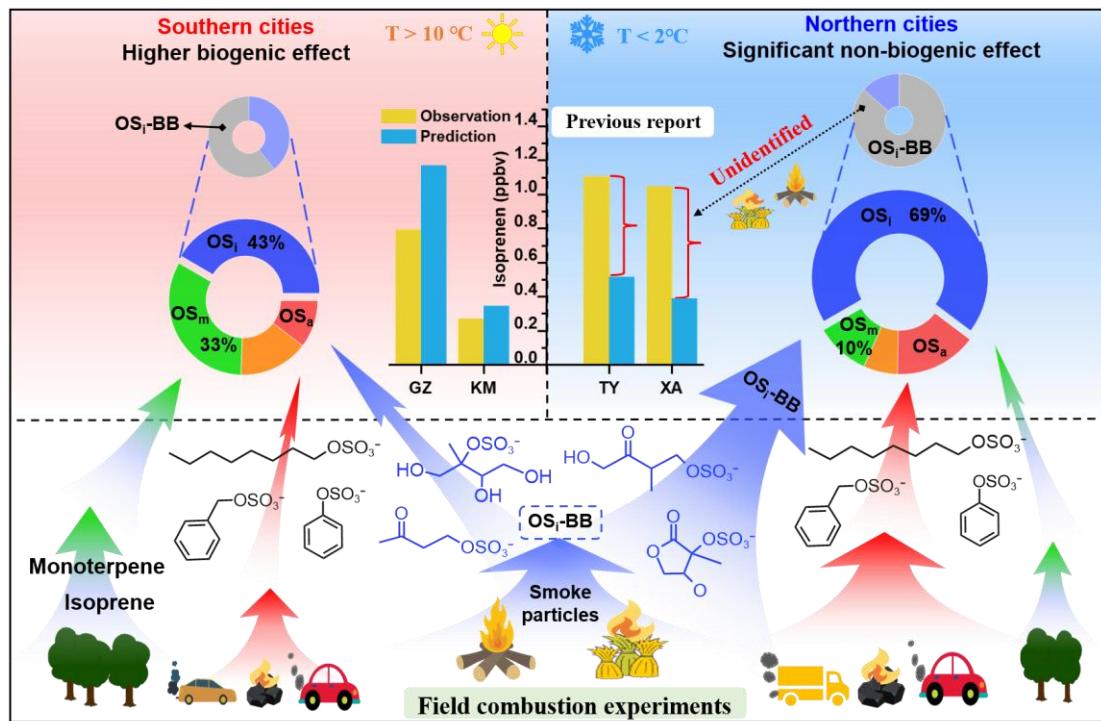
427

428 **4. Conclusion and atmospheric implications**

429 It has been previously suggested that isoprene can also be released into the  
430 atmosphere as a result of open burning of agricultural residues and forest fires  
431 (Andreae 2019; Simpson et al. 2011). A field study conducted by (Wang et al. 2019)  
432 in Beijing during winter inferred that the prevalence of OS<sub>i</sub> compounds in total  
433 aerosol OSs may be partially attributable to biomass burning emissions, although  
434 there was a paucity of compelling evidence to support this hypothesis. This work  
435 combines strongly contrasting observational studies (northern Chinese Cities vs  
436 southern Chinese Cities) with in situ combustion modelling experiments to provide  
437 the first direct evidence that biomass burning emission, rather than fossil fuel  
438 combustion emission, is a significant contributor to aerosol OS<sub>i</sub> in northern cities  
439 (**Figure 5**). In Chinese cities, particularly those in the northern region, biomass  
440 materials are extensively utilized for domestic heating and cooking purposes during  
441 the winter season (Zhou et al. 2017). Clearly, the isoprene emissions from biomass  
442 combustion sources would result in higher isoprene mixing ratios than those  
443 simulated by the model (Zhang et al. 2020) that only considers natural isoprene  
444 emissions. Thus, isoprene prediction models applied to Chinese winters in the future  
445 should also take into account the various biomass combustion source releases. Given  
446 the potential for both biomass burning and biogenic isoprene to contribute to OS<sub>i</sub>  
447 formation, separating their respective contributions remains challenging. Furthermore,  
448 biogenic OSs are important SOA constituents and have been frequently serve as  
449 important tracers for biogenic SOA (Ding et al. 2014; Ding et al. 2016a). The overall  
450 results suggest that some OS<sub>i</sub> species may not be suitable as biogenic SOA markers,

451 especially in areas with intensive biomass burning activities, such as northern Chinese  
452 cities during winter.

453



454

455 **Figure 5** Conceptual picture showing the characteristics and main contributors of OSs  
456 in northern and southern China during winter. It is noteworthy that OS<sub>i</sub>-BB can  
457 originate not only from biomass combustion, but also from the secondary formation  
458 of isoprene emitted from biogenic sources.

459

460 We found that different fossil fuel combustion emissions (e.g., vehicle emissions  
461 and coal combustion emissions) and biomass burning emissions can contribute to  
462 aerosol anthropogenic OSs. However, current studies have not been able to accurately  
463 distinguish between the contributions of various material combustion to different  
464 types of anthropogenic OSs. Future research is necessary to develop more

465 comprehensive models to further explore the effects of various combustion sources on  
466 the generation and reduction of urban aerosol OS pollution. Of particular importance  
467 is that although the production of various OSs was directly observed through our  
468 simulated combustion experiments, it is not clear whether the chemical mechanisms  
469 involved are similar to those derived from the laboratory simulations. This is because  
470 the combustion process is accompanied by the effects of high temperatures. In  
471 general, although our results provide direct evidence for the release of OSs from  
472 combustion of various combustion sources, further mechanistic studies and  
473 environmental impact assessment are still urgently needed. This may be important for  
474 effective control of urban wintertime organic aerosol pollution in China.

475

#### 476 **Data availability**

477 The data presented in this work are available upon request from the corresponding  
478 authors.

479

#### 480 **Competing interests**

481 The authors declare no conflicts of interest relevant to this study.

482

#### 483 **Author contributions**

484 YX designed the study. TY, YJM, HWX, and HX performed field measurements and  
485 sample collection; TY and YJM performed chemical analysis; YX and TY performed  
486 data analysis; YX and TY wrote the original manuscript; and YX, TY, YCW, and

487 HYX reviewed and edited the manuscript.

488

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492

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