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**Deep-Learning-derived Planetary Boundary Layer Height from
Conventional Meteorological Measurements**

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21 **Abstract.** The planetary boundary layer (PBL) height (PBLH) is an important
22 parameter for various meteorological and climate studies. This study presents a multi-
23 structure deep neural network (DNN) model, which can estimate PBLH by integrating
24 the morning temperature profiles and surface meteorological observations. The DNN
25 model is developed by leveraging a rich dataset of PBLH derived from long-standing
26 radiosonde records and augmented with high-resolution micro-pulse lidar and Doppler
27 lidar observations. We assess the performance of the DNN with an ensemble of ten
28 members, each featuring distinct hidden layer structures, which collectively yield a
29 robust 27-year PBLH dataset over the Southern Great Plains from 1994 to 2020. The
30 influence of various meteorological factors on PBLH is rigorously analyzed through
31 the importance test. Moreover, the DNN model's accuracy is evaluated against
32 radiosonde observations and juxtaposed with conventional remote sensing
33 methodologies, including Doppler lidar, ceilometer, Raman lidar, and Micro-pulse
34 lidar. The DNN model exhibits reliable performance across diverse conditions and
35 demonstrates lower biases relative to remote sensing methods. In addition, the DNN
36 model, originally trained over a plain region, demonstrates remarkable adaptability
37 when applied to the heterogeneous terrains and climates encountered during the
38 GoAmazon (Tropical Rainforest) and CACTI (Middle Latitude Mountain) campaigns.
39 These findings demonstrate the effectiveness of deep learning models in estimating
40 PBLH, enhancing our understanding of boundary layer processes with implications for
41 enhancing the representation of PBL in weather forecasting and climate modeling.

42

43 **1 Introduction**

44 The Planetary Boundary Layer (PBL) is the atmosphere's lowest part, where the
45 Earth's surface directly influences meteorological variables, impacting the climate
46 system (Garratt, 1994; Kaimal and Finnigan, 1994). The PBL height (PBLH) is a
47 meteorological factor that strongly influences surface-atmosphere exchanges of heat,
48 moisture, and energy (Stull, 1988; Caughey, 1984; Holtslag and Nieuwstadt, 1986;
49 Mahrt, 1999; Helbig et al., 2021; Guo et al., 2024; Beamesderfer et al., 2022). In
50 addition, PBLH is a crucial variable for monitoring and simulating surface pollutant
51 behaviors since it determines the volume available for near-surface pollutant dispersion
52 (Li et al., 2017; Su et al., 2024a; Tucker et al., 2009; Wang et al. 2020). Due to its
53 impact on cloud evolution and the development of convective systems, PBLH is also a
54 key parameter in numerical weather forecasts and climate projections (Deardorff, 1970;
55 Kaimal et al. 1976; Menut et al., 1999; Park et al., 2001; Emanuel, 1994; Guo et al.,
56 2017, 2019; Lilly, 1968; Matsui et al., 2004).

57 Radiosonde (SONDE) remains the standard method for estimating PBLH, yet it is
58 hampered by limitations in temporal frequency, restricting its ability to capture the
59 whole diurnal cycle of PBL development (Stull, 1988; Seidel et al. 2010; Guo et al.
60 2021; Liu and Liang, 2010). To overcome these challenges, there has been an increasing
61 dependence on remote sensing techniques, especially lidar systems. These techniques
62 capture atmospheric vertical information (e.g., aerosols, temperature, humidity, and
63 wind) at high temporal and vertical resolutions, leading to remote sensing-based
64 retrievals of PBLH (Menut et al., 1999; Kotthaus et al., 2023; Sawyer and Li, 2013;

65 Wang et al., 2023). The remote sensing systems, including Doppler lidar (Barlow et al.
66 2011), ceilometer (Zhang et al. 2022), Raman lidar (Summa et al. 2013), and Micro-
67 pulse lidar (Melfi et al., 1985), utilize laser-based technology to track PBLH diurnal
68 evolutions, helping us understand the PBL evolutions (Cohn and Angevine, 2000;
69 Davis et al., 2000). In addition, wind profilers can estimate the PBLH using algorithms
70 that analyze the signal-to-noise ratio from wind profiler data (Molod et al. 2015; Solanki
71 et al. 2022; Liu et al. 2019; Salmun et al. 2023; Bianco and Wilczak 2002; Bianco et al.
72 2008; Tao et al. 2021).

73 However, the advancement in remote sensing for the estimation of PBLH
74 challenges is still posing in bridging the results obtained by different remote sensing
75 instruments with those obtained from the SONDE measurements (Zhang et al. 2022;
76 Chu et al., 2019). Specifically, interpreting aerosol, turbulence, and moisture profiles
77 derived from remote sensing techniques to determine PBLH bears inherent limitations
78 due to the unstable signal-to-noise ratio (Kotthaus et al., 2023; Krishnamurthy et al.,
79 2021). This issue is compounded by the different measurement methodologies and
80 definitions employed by various remote sensing tools, leading to uncertainties when
81 comparing their PBLH estimates to the retrievals derived from SONDE measurements
82 (Zhang et al. 2022; Sawyer and Li, 2013).

83 As machine learning (ML) has shown potential in atmospheric science (McGovern
84 et al., 2017; Gagne et al., 2019; Su et al. 2020a; Vassallo et al., 2020; Cadeddu et al.,
85 2009; Molero et al. 2022), this technique presents a promising tool for refining the
86 estimation of PBLH to resolve the inherent complexity and variability of PBL. For

87 example, several studies use ML to identify PBLH using thermodynamic profiles
88 Atmospheric Emitted Radiance Interferometer (AERI) or using backscatter profiles
89 from lidar, highlighting the ML's superiority over conventional techniques under
90 different scenarios (Sleeman et al. 2020; Rieutord et al. 2021; Liu et al. 2022; Ye et al.
91 2021). For example, Li et al. (2023) applied an ML algorithm for retrieving PBLH
92 under complex atmospheric conditions with account of the vertical distribution of
93 aerosols. Krishnamurthy et al. (2021) incorporated a random forest model, along with
94 machine learning, to use Doppler lidar data for the extraction of PBLH with better
95 results compared to the results retrieved by traditional methods.

96 While existing ML methodologies have made great progress in estimating PBLH,
97 these studies mainly focus on refining retrievals from remote sensing data, particularly
98 lidar-based technologies. Thus, there is an inherent limitation to the applicability due to
99 a reliance on specific remote sensing instruments. To address this issue, we aim to
100 leverage and integrate the comprehensive field observations (i.e., radiosonde and
101 remote sensing techniques) to develop a deep learning model for direct PBLH
102 estimation from conventional meteorological data. This strategy circumvents the
103 limitations of relying on particular remote sensing technologies. Furthermore, our
104 model employs an advanced deep neural network (DNN) approach (Sze et al. 2017;
105 Schmidhuber, 2015; Nielsen, 2015; Pang et al. 2020), diverging from traditional ML
106 methods like random forest. This deep learning model utilizes ensemble techniques,
107 constructing arrays of various structures and using their average for the final estimation.
108 This approach method provides particular advantages in the context of complex and

109 nonlinear processes (Ganaie et al. 2022; Mohammed and Kora. 2023). Ensemble DNN
110 with multi-structure designs shows very strong flexibility and robustness, so it
111 relatively performs better and has high stability across a wide range of conditions (Xue
112 et al. 2020; Dong et al. 2020). This facilitates the adaptability of DNN as a tool for
113 PBLH estimation, which can be utilized under different scenarios and locations.

114 By focusing on the interaction between surface meteorology and the PBL, this study
115 introduces a DNN-based method to estimate the daytime evolution of PBLH from
116 morning temperature profiles and surface meteorology. We evaluate the model's
117 performance using extensive datasets over the Southern Great Plains (SGP) for a period
118 spanning 27 years (1994-2020) and includes comparisons with PBLH estimations
119 obtained from Doppler lidar, ceilometer, Raman lidar, and micro-pulse lidar.
120 Furthermore, we explore the generalizability of the model to different geographic
121 regions and climates, as tested during the field campaigns, e.g., Green Ocean Amazon
122 (GoAmazon) and Cloud, Aerosol, and Complex Terrain Interactions (CACTI).

123

124 **2 Data and instruments**

125 **2.1 ARM Sites**

126 The Atmospheric Radiation Measurement (ARM) program, funded by the U.S.
127 Department of Energy, has been employed at the Southern Great Plains (SGP) site in
128 Oklahoma (36.607°N, 97.488°W), situated 314 meters above mean sea level. This study
129 use comprehensive field observations at the SGP site during 1994 to 2020. In addition
130 to the SGP site, this study utilizes data from the ARM GoAmazon (3.213°S, 60.598°W)

131 and ARM CACTI (32.126°S, 64.728°W) field campaigns to carry out independent
132 tests for the deep learning model. Specifically, the GoAmazon campaign is located in the
133 Amazon tropical forests and provides rich field observations data during 2014-2015
134 (Martin et al. 2016). Meanwhile, the CACTI central site, at an elevation of 1141 meters
135 within the Sierras de Córdoba Mountain range in north-central Argentina, offers the
136 observations during the 2018-2019 period (Varble et al. 2021). Utilizing these
137 comprehensive ARM datasets, our study includes thermodynamic profiles derived from
138 radiosondes, data from the Active Remote Sensing of Clouds (ARSCL, Clothiaux et al.
139 2000, 2001; Kollias et al. 2020), in-situ surface flux measurements, and standard
140 meteorological observations at the surface, as documented by Cook (2018) and Xie et
141 al. (2010).

142 SONDE measurements at the ARM sites launch routinely several times a day and
143 provide detailed information into the thermodynamic conditions of the atmosphere. The
144 technical details of the ARM SONDE data are documented in Holdridge et al. (2011).
145 Moreover, we use the surface meteorological parameters at the standard meteorological
146 station. In-situ measurements at 2 meters above ground level provide data on
147 temperature, relative humidity, and vapor pressure. Moreover, this study obtains the
148 surface sensible and latent heat fluxes from the surface instruments (Wesely et al.,
149 1995). In SGP, we use the best-estimate surface fluxes in the Bulk Aerodynamic Energy
150 Balance Bowen Ratio (BAEBBR) product, which is derived from the measurements by
151 Energy Balance Bowen Ratio (EBBR). Due to the availability, we utilize the surface

152 fluxes from Quality Controlled Eddy CORrelation (QCECOR) datasets from CACTI
153 and GoAmazon sites (Tang et al. 2019).

154

155 **2.2 Existing PBLH datasets over the ARM sites**

156 For analyzing PBLH, we have utilized a variety of datasets to get a full picture of
157 PBLH derived from different instruments. These datasets are developed by using
158 different methodologies and instruments and jointly offer a detailed information of
159 PBLH under various meteorological conditions. Among these datasets, SONDE- and
160 ceilometer-derived PBLH are available for all three sites, other datasets are only
161 available over the SGP. The technique details for these datasets can be found in the
162 corresponding publications or technique reports.

163 *(1) SONDE-derived PBLH by Liu and Liang (2010):*

164 PBLHs are retrieved using a method developed by Liu and Liang (2010), based on
165 potential temperature gradients from SONDE. We focus on daytime data during 05:00–
166 18:00 Local Time (LT), with a resampled vertical resolution of 5-hPa. The SONDE
167 dataset is available at DOI: <https://doi.org/10.5439/1595321>.

168 *(2) Doppler Lidar-derived PBLH by Sivaraman and Zhang (2021):*

169 Doppler lidar PBLH estimates are derived using a vertical velocity variance method
170 during 2010-2019 (Tucker et al., 2009; Lareau et al., 2018; Sivaraman and Zhang 2021).
171 The dataset is available at DOI: <https://doi.org/10.5439/1726254>.

172 *(3) Combined MPL and SONDE PBLH by Su et al. (2020b):*

173 We utilize a PBLH dataset that merges lidar and SONDE measurements during
174 1998-2023, ensuring vertical coherence and temporal continuity (Su et al. 2020b). An
175 additional method for handling cloudy conditions is detailed in Su et al. (2022). The
176 dataset is available at DOI: <https://doi.org/10.5439/2007149>.

177 *(4) Ceilometer-derived PBLH by Zhang et al. (2022):*

178 The Vaisala CL31 ceilometer, with a 7.7 km vertical range, provides detailed
179 backscatter profiles used for PBLH estimation via gradient methods during 2011-2023
180 (Zhang et al. 2022). Enhanced algorithms ensure robust estimations under all weather
181 conditions. The dataset is available at DOI: <https://doi.org/10.5439/1095593>.

182 *(5) MPL-derived PBLH by Sawyer and Li (2013):*

183 Micropulse lidar (MPL) is utilized for its high temporal resolution to retrieve PBLH
184 during 2009-2020. MPL-derived PBLH, validated against SONDE and AERI data,
185 improves understanding of boundary-layer processes (Sawyer and Li. 2013). The
186 dataset is available at DOI: <https://doi.org/10.5439/1637942>.

187 *(6) Combined Raman Lidar and AERI PBLH by Ferrare (2012):*

188 PBLH is calculated using merged potential temperature profiles from Raman lidar
189 and AERI, with criteria established for the SGP site. PBL heights are computed hourly
190 for 2009-2011. The dataset is available at DOI: <https://doi.org/10.5439/1169501>.

191 In the datasets, (1-3) serve as the foundation for training. Concurrently, considering
192 radiosonde as the benchmark standard, we utilized dataset (1) for validating PBLH
193 retrievals obtained from various sources. Meanwhile, datasets (4-6) are used for the
194 intercomparisons between PBLH derived from DNN and remote sensing techniques.

195

196 **3 Deep Learning Model to Estimate PBLH**

197 **3.1 The Multi-Structure Deep Learning Model**

198 Our deep learning model for estimating PBLH leverages the robustness of ensemble
199 learning using a multi-structure DNN (Sze et al. 2017; Schmidhuber, 2015; Nielsen,
200 2015; Pang et al. 2020). This model used the TensorFlow Package, developed by
201 Google (Abadi et al., 2016; <https://www.tensorflow.org/>). By employing an array of
202 varied network architectures, we capitalize on the unique strengths of each structure to
203 synthesize a more accurate and reliable estimation of PBLH. Figure 1 outlines the
204 DNN's comprehensive design, beginning with the input layer that ingests a suite of
205 morning meteorological features. The DNN model derives the PBLH from surface
206 meteorological parameters. We also incorporate boundary layer heights derived from
207 sensible heat and parcel methods (BLH_{Parcel} and BLH_{SH}) as inputs. Specifically,
208 BLH_{Parcel} is calculated based on the morning profile of potential temperature
209 (Holzworth. 1964), while BLH_{SH} is determined using the surface temperature
210 combined with surface sensible heat, following the methodologies (Stull, 1988; Su et
211 al. 2023). We first present a preliminary run for the model to obtain the importance of
212 each input feature. Then, these inputs undergo a filtration process based on their
213 importance (Date and Kikuchi, 2018; Altmann et al. 2010), ensuring that only the
214 impactful data guide the model (detailed in Section 3.3). Subsequently, the filtered
215 inputs traverse through an ensemble of ten structures with distinct hidden layers. Each
216 structure here represents an ensemble member and contributes to the prediction of

217 PBLH in its unique way (Ganaie et al. 2022). The ensemble employs a three-layer base
218 structure [52, 28, 16] for neural networks, from which ten unique configurations are
219 derived by applying random perturbations to the default settings of the base structure.
220 These different structures for ensembles 1-10 are presented in Table 1.

221 At the final stage, the model use the PBLH esimations from different ensembles to
222 get a mean value as the final PBLH retrieval. This process allows the model to leverage
223 the different results of all structures and enhance the generalizability of results. In the
224 DNN model, neuron biases in the output and hidden layers are important for the
225 network's architecture (Battaglia et al. 2018). These biases serve as fine-tuning
226 parameters to adjust the activation thresholds of neurons in different layers and further
227 refine the model's predictive capabilities. Neuron biases are initialized with small
228 random values at the start of the training process and then iteratively adjusted according
229 to the network weights during the training. Normalization is a preprocessing technique
230 that often leads to improvements in model training by scaling the input features and
231 target values to a standard range (Raju et al. 2020). The normalization process was
232 applied to each input data to ensure that they have a zero mean and a standard deviation
233 of one, as well as the target data. This standardization scales the different input data to
234 a similar range, and thus, contributes a more stable and efficient training process.

235 The hidden layers of the DNN model incorporate L2 regularization to curtail
236 overfitting, while batch normalization aids in stabilizing learning. Moreover, a dropout
237 rate of 0.2 helps the model to generalize better by reducing reliance on any specific
238 neurons during training. We choose the Adam optimizer and mean squared error as the

239 loss function, which aligns with one of the best practices for regression models (Zhang.
240 2018). The mean absolute error is selected as a metric to evaluate the model's accuracy
241 during the training. We incorporate the early stopping and learning rate reduction
242 callbacks in the model's training for regularization and fine-tuning (Liu et al. 2019).
243 Such measures ensure optimal performance by terminating training at the right juncture
244 and avoid the overfitting in the final results.

245

246 **3.2 Training the DNN Model**

247 The training of the DNN model was conducted using a PBLH dataset enriched by
248 SONDE and lidar measurements during 1994 to 2016 over the SGP. Table 2 presents
249 the distribution of dataset samples under different local time, which were important for
250 both the training and validation processes of the DNN model. The primary dataset (i.e.,
251 PBLH derived from SONDE measurements) is listed in the first column and are
252 available routinely for 5, 11, and 17 LT. The training dataset was augmented with the
253 combined MPL-SONDE PBLH dataset (Su et al. 2020b) and the Doppler Lidar-derived
254 PBLH (Sivaraman and Zhang, 2021) to address the gaps where SONDE measurements
255 were not available. In instances where radiosonde data are unavailable, the lidar datasets
256 are used for training, contingent upon their agreement with radiosonde measurements
257 within a margin of 0.2 km over a 3-hour window. Specifically, out of the total
258 comparisons during the study period, 40.2% of the lidar measurements do not agree
259 within the 0.2 km threshold with the SONDE results. The cases with relatively larger
260 inconsistencies stem from various factors, including instrumental errors, rainy

261 conditions, stable PBL conditions, differing definitions, and lidar signal attenuation, as
262 discussed in previous studies (Su et al., 2020b; Kotthaus et al., 2023). These cases were
263 excluded from the DNN model training to maintain the quality of the process.

264 For the purpose of training the DNN model, 70% of the hourly data from both
265 SONDE and the lidar combined dataset were randomly selected. The remaining 30%
266 dataset, comprises the portion of SONDE measurements set aside for validation
267 purposes, including a separate subset from the years 2017 to 2020 to test the model's
268 predictive capabilities on independent data. This training and validation scheme ensures
269 that the DNN model is not only well-trained but also thoroughly evaluated, reinforcing
270 its reliability in accurately estimating PBLH. As morning SONDE data constitute the
271 primary input and boundary conditions for the model, the validation of PBLH retrievals
272 is consequently confined to the 08:00 to 18:00 LT.

273

274 **3.3 Feature Importance Score**

275 In the DNN model, we quantified the significance of each input parameter using the
276 permutation importance technique, which is a widely-used method for the deep learning
277 (Date and Kikuchi, 2018; Altmann et al. 2010; Breiman, L., 2001). Initially, we carry
278 out a test run to determine a baseline performance by calculating the mean absolute
279 error (MAE) on the validation set. Then, each feature within this set was then
280 individually shuffled, severing its correlation with the target PBLH, and the MAE was
281 recalculated. Compared to the baseline performance, the increase in MAE from this
282 shuffled state indicates the feature's predictive value: the greater the increase, the more

283 significant the feature. We repeat this shuffling and evaluation for 15 times, each with
284 a unique random seed to ensure statistical robustness. Furthermore, we calculated the
285 average MAE increase across these iterations as the importance score. These scores are
286 expressed as percentages, with each feature's importance score normalized to sum to
287 100%. Each score quantitatively represents how much the shuffling of a feature
288 increases the MAE, indicating the relative significance of that feature in the model's
289 predictive accuracy and facilitating a straightforward comparison of the influence of
290 each feature within the model. Therefore, we derived a composite importance metric
291 for feature groups to represent their significance as the cumulative sum of related inputs.

292 Figure 2 presents the importance scores to demonstrate the relative influence of
293 different feature groups on the model's performance. Prominently, features such as the
294 BLH_{Parcel} , morning potential temperature profiles (θ profile), and surface relative
295 humidity are identified as most important three features, with their substantial impact
296 on the accuracy of PBLH estimation being highlighted. BLH_{Parcel} is defined as the
297 height where the morning potential temperature first exceeds the current surface
298 potential temperature by more than 1.5 K (Holzworth, 1964; Chu et al., 2019). Among
299 these features, BLH_{Parcel} captures the response of the PBL to surface heating, which
300 can drastically affect local convection and thus serves as one of the key parameters in
301 the DNN model. Incorporating this parameter and its association with PBL
302 development better simulates diurnal variations of PBLH in the DNN model.
303 Meanwhile, the morning θ profile represents the vertical stratification of
304 thermodynamics and is essential for understanding stability and mixing processes

305 within the PBL. Thus, θ profile serves as the initial boundary condition for the PBLH
306 estimation with a significant importance score. Surface relative humidity also emerges
307 as a key influencer, affecting the model's performance significantly. Humidity levels
308 influence the condensation and evaporation processes within the PBL, which are
309 important in determining its vertical extent layer and structure. Fair-weather and dry
310 conditions are typically associated with a more turbulent and higher PBL. Conversely,
311 high surface humidity often contributes to the formation of boundary layer clouds,
312 which introduces complex interactions with PBL thermodynamics.

313 In this analysis, each feature, such as θ profile, comprises several different inputs,
314 and the relative importance scores presented in Figure 2 are calculated as the cumulative
315 sum of these inputs. Complementing this, Table 3 offers an exhaustive breakdown of
316 importance scores for all considered input features within the deep learning model. In
317 refining the model, features contributing a negligible or negative effect on performance
318 (i.e., importance scores less than zero) are excluded. As a result, this selection criterion
319 has led to the inclusion of 58 out of the original 64 features. This process ensures we
320 only use inputs with a proven positive influence in the DNN model.

321

322 **4 Evaluation of Deep Learning Model**

323 **4.1 Comparative analysis of biases among different datasets**

324 A critical component of evaluating our deep learning model's efficacy is analyzing
325 the biases of individual ensemble members and their collective output. Figure 3 offers
326 a visual assessment of the mean absolute error (MAE), root mean square error (RMSE),

327 and correlation coefficient (R) for each ensemble member, alongside a comparison with
328 the ensemble mean (average of all individual ensemble members). The plotted data
329 points reveal the variation in performance across different model architectures, while
330 the ensemble mean, represented by the horizontal dashed lines, indicates the collective
331 accuracy of the ensemble approach. The structures of different hidden layer
332 configurations are listed in the Table 1.

333 This methodological consolidation results in a more reliable and accurate PBLH
334 estimation, leveraging the strengths and mitigating the weaknesses of individual
335 models. By integrating multiple neural network configurations, we revealed that an
336 ensemble prediction that consistently outperforms the individual models. This strategy
337 can improve the MAE by up to 4.4%, rendering the model less dependent on any
338 specific structural configuration.

339 An in-depth comparative analysis of biases among various PBLH estimation
340 methods is essential for validating the reliability and accuracy of the DNN developed
341 in this study. Figure 4 illustrates the MAE trends for several methods over a multi-year
342 span, with the SONDE-derived PBLH serving as the benchmark for ground truth. The
343 analysis reveals the performance of different methodologies: the DNN approach,
344 doppler lidar, ceilometer, MPL, and Raman lidar. Significantly, the DNN model,
345 depicted in black, maintains a consistent MAE trend throughout the trained period
346 (1994-2016) as well as the subsequent untrained period (2017-2020), demonstrating
347 robust predictive stability. In contrast, the remote sensing-based methods show a
348 reduction in bias from 2010 to 2022, possibly due to the improvement of remote sensing

349 data quality. The discrepancy in PBLH estimates between the DNN and SONDE
350 remains consistently lower than those observed with conventional remote sensing
351 techniques.

352 Figure 5 provides a detailed evaluation of the DNN model in comparison to
353 ceilometer and doppler lidar-derived PBLH, as these two methods have demonstrated
354 the high quality with more than nine years of datasets. Figure 5a-b contrast the PBLH
355 predictions from the DNN model for both the trained period (1994-2016) and untrained
356 periods (2017-2020), respectively, showcasing strong correlations and low MAEs,
357 indicative of the model's robust training and generalization capabilities. Figure 5c-d
358 further this examination with ceilometer and Doppler lidar comparisons, respectively.
359 Overall, Doppler lidar exhibits a closer alignment with SONDE-derived PBLH than the
360 ceilometer. However, the MAE from Doppler lidar-based estimates is still
361 approximately 48% higher than those derived from the DNN model. The correlation
362 coefficient for the DNN-derived PBLH estimates has seen a substantial improvement,
363 rising from the 0.5-0.6 range typically observed with remote sensing-based PBLH
364 methods to exceed 0.8 when compared to SONDE-derived PBLH measurements. This
365 comparative analysis not only confirms the DNN model's accuracy but also offers
366 insights into the relative performance of various contemporary PBLH estimation
367 methodologies.

368

369 **4.2 Performances of PBLH retrievals under different conditions**

370 The performance of PBLH retrievals under varying atmospheric conditions is a
371 crucial aspect of model evaluation. In Figure 6, the seasonal diurnal cycles of PBLH
372 estimated by different methods are presented, offering information into the diurnal and
373 seasonal evolution of PBL. As PBLH demonstrates notable variations for different
374 seasons and local time with large differences between summer and winter, the DNN
375 and Doppler lidar estimates show good agreement and closely track the variations
376 observed in SONDE data. Meanwhile, the ceilometer presents an underestimation of
377 PBLH, especially for the summer afternoon, indicating the potential bias of ceilometer
378 derived PBLH under a convective environment.

379 Figure 7 illustrates the diurnal variation in the model's performance by comparing
380 the correlation coefficient, RMSE, and MAE against SONDE-derived PBLH as the
381 reference. The bar graphs for each local time hour offer a comparison of the RMSE and
382 MAE, as well as the correlation, showcasing the model's precision and consistency
383 relative to remote sensing methods (i.e., ceilometer and Doppler lidar). The ceilometer-
384 derived PBLH exhibits the greatest variations during different hours, particularly
385 around noon, suggesting a time-dependent bias in its measurements. Conversely, both
386 the DNN and Doppler lidar-derived PBLH demonstrate stable performance in term of
387 MAE and RMSE throughout the day. Regarding the correlation, remote sensing
388 methods like ceilometer and Doppler lidar exhibit a lower correlation with SONDE-
389 derived PBLH, especially in the early hours (8-9 LT) with a value of 0.1-0.3, indicating
390 potential limitations in their reliability during these times. On the other hand, the DNN

391 model shows a relatively good correlation with SONDE retrievals (above 0.6 under
392 different hours). This comparison shows the efficacy of DNN in tracking the diurnal
393 cycle of PBLH.

394 Continuing our assessment of the DNN model, we analyze the DNN model's
395 monthly performance in estimating PBLH, as shown in Figure 8. The analysis compares
396 MAE, RMSE, and correlation coefficients for each month to assess the model's
397 precision and dependability. The summer months (June-July-August) exhibit higher
398 biases, with MAE values for the DNN, ceilometer, and Doppler lidar at 0.3 km, 0.56
399 km, and 0.45 km, respectively. In contrast, the winter months (December-January-
400 February) show reduced biases, with MAE values of 0.2 km for the DNN, 0.27 km for
401 the ceilometer, and 0.24 km for the Doppler lidar. Specifically, the DNN model shows
402 a much lower bias during the summer season. Compared to the remote sensing-based
403 retrievals, the DNN-derived PBLH shows a much better agreement with SONDE-
404 derived PBLH, increasing from 0.3-0.6 to approximately 0.8 in term of correlation
405 coefficients.

406 Figure 9 presents the biases of PBLH retrievals under clear-sky and low cloud
407 conditions. We calculated biases as the absolute deviation from the mean PBLH for
408 each condition, focusing particularly on the differences between low cloud (maximum
409 cloud fraction between 0-4 km exceeding 1%) and clear-sky (total cloud fraction below
410 1%) scenarios. The threshold of 1% for cloud fraction is also used to identify cloud base
411 height (CBH) in the European Centre for Medium-Range Weather Forecasts' fifth-
412 generation global reanalysis (ERA-5, Hersbach et al., 2023). The violin plots in this

413 figure illustrate the data distribution of biases for each method to demonstrate their
414 variability. For the DNN model and ceilometer, the relative biases between clear and
415 cloudy conditions are comparable and the difference is less than than 1%. This suggests
416 a consistent performance across these atmospheric states. However, the Doppler lidar
417 exhibits a larger disparity, showing a 5.5% bias under cloudy conditions compared to
418 clear skies. Moreover, the spread of biases (shaded areas and error bars) is notably wider
419 for both the ceilometer and Doppler lidar. This indicates large variability in their
420 performance. For all three methods, the mean biases are notably higher than the median
421 values. Such differences indicate that the mean values are notably influenced by outliers
422 under both clear-sky and cloudy conditions.

423 The evolution of the PBLH under shallow cumulus conditions offers insights into
424 the interactions between clouds, PBL, and land surface (Zhang and Klein, 2010, 2013).
425 Figure 10 demonstrates the variations of PBLH measurements from different methods
426 during conditions typical of shallow cumulus clouds. Shallow cumulus clouds were
427 identified following Su et al. (2024b). Specifically, these coupled clouds form post-
428 sunrise; and the sky must not be overcast, characterized by a cloud fraction less than
429 90%. This selection criterion ensures that the observed cloud formations are primarily
430 driven by surface heating and local convection. The DNN model closely matches the
431 SONDE-derived PBLH and the CBH from ARSCL. This alignment underscores the
432 physical validity of the DNN approach, confirming its capability to replicate traditional
433 measurement techniques to a good extend of accuracy. Meanwhile, Doppler lidar-
434 derived PBLH retrievals also show high consistency with SONDE measurements,

435 whereas ceilometer-derived PBLH generally underestimates values under shallow
436 cumulus conditions.

437 Figure 10 also demonstrates the general relationship between the development of
438 shallow cumulus clouds and the PBL, which are driven by local convection and
439 turbulence. The formation of these cumulus clouds is linked to rising thermals and an
440 increase in surface heat fluxes, essential for driving vertical mixing within the sub-cloud
441 layer. This relationship is evidenced by the increased occurrence of cumulus clouds
442 along with an increase in DNN-derived PBLH from morning to late afternoon.
443 Specifically, during periods with a high frequency of shallow cumulus, the DNN-
444 derived PBLH often surpasses the CBH. This indicates that rising air parcels extend
445 beyond the condensation level, facilitating the formation and development of coupled
446 cumulus clouds.

447 In this context, these analyses confirm the physical consistency of DNN-derived
448 PBLH with traditional measurement techniques and highlight its physically reasonable
449 variations during cloudy conditions. The results presented in this section illustrate the
450 effectiveness of the DNN model in capturing the PBLH variations across different local
451 times, seasons, and cloudy conditions. Compared to the traditional remote sensing
452 methods, the DNN model exhibits relatively good accuracy in aligning with SONDE-
453 derived PBLH, indicating its capability and stable performance under different
454 scenarios.

455

456 **4.3 Testing the DNN Model's Adaptability**

457 The DNN model relies on the incorporation of morning temperature profiles as
458 inputs, such as detailed in Table 3. This dependency prompts the question of how to
459 proceed the DNN model in the absence of SONDE data at specific locations. As a
460 solution, we suggest employing morning temperature profiles from the ERA-5
461 (Hersbach et al., 2020) dataset when radiosonde data is not available to maintain the
462 model's operational integrity for the conditions without SONDE data. As one of the
463 most advanced reanalysis data, the ERA-5 is generated by the Integrated Forecasting
464 System coupled with a data assimilation system, and offer the meteorological data at a
465 spatial resolution of 0.25° - 0.25° .

466 Figure 11 assess the performance of DNN produced by multi-sources field
467 observations in estimating the PBLH by using morning temperature profiles from ERA-
468 5 (5 LT) and observed surface meteorological data. The temperature profiles in ERA-5
469 have a vertical resolution of 25-hPa in the lower atmosphere and are interpolated into
470 different levels described in Table 3. By utilizing ERA-5 morning profiles, the model
471 demonstrates similar performance to those results achieved with radiosonde inputs, as
472 evidenced by comparing Figure 11a and Figure 5. Moreover, this alternative approach
473 also shows enhanced accuracy over the native PBLH model outputs from ERA-5,
474 increasing the correlation coefficient from 0.74 to 0.86 and reducing the MAE from 0.3
475 km to 0.25 km. In addition, it is important to acknowledge that the PBLH represented
476 in ERA-5 is indicative of a grid-average value, approximately 25 km in scale, and
477 therefore inherently differs from site-specific data.

478 These findings highlight the alternative DNN model's robustness, offering a reliable
479 substitute for radiosonde data by leveraging reanalysis data with similar performance.
480 This demonstrates the DNN model's adaptability and potential as a practical tool for
481 PBLH estimation across various meteorological sites, especially in regions or periods
482 where radiosonde data may be lacking.

483 We further test the adaptability and generalizability of the DNN model, by applying
484 across different climatic and geographic regions. To this end, we extended our model
485 evaluation to include SONDE and surface meteorological data from the GoAmazon
486 (Tropical Rainforest) and CACTI (Middle Latitude Mountain) field campaigns.
487 Seasonality is accounted for as an input variable in the DNN model, with months in the
488 Southern Hemisphere adjusted to reflect their Northern Hemisphere seasonal
489 counterparts (e.g., July inputs are treated as January). The normalization process
490 (Section 3.1) was reapplied for the CACTI campaign data to adjust for notable pressure
491 level variations, ensuring input standardization with zero mean and unit variance.

492 Figure 12 presents the model's performance, in comparison to SONDE observations
493 for both GoAmazon and CACTI campaigns. The DNN model demonstrates
494 commendable adaptability, maintaining a strong correlation (0.86-0.88) with SONDE
495 measurements (Figure 12a-b). Further comparison is provided, which assess the
496 performance of ceilometer derived PBLH against SONDE for the same campaigns.
497 When assessing the performance of the ceilometer-derived PBLH against SONDE for
498 the same campaigns, the DNN model exhibited both stronger correlations and smaller
499 biases, as shown in Figure 12b-d.

500 Nevertheless, the analysis highlighted the presence of systematic biases, with
501 relatively larger MAE at the GoAmazon and CACTI sites compared to the SGP site.
502 Figure 13 underscores this by presenting a comparative analysis of PBLH means and
503 standard deviations across the three ARM sites. The early morning measurements
504 during 05-07 LT are excluded. The results, derived from SONDE, the DNN model,
505 ceilometer, and Doppler lidar data, reveal average differences in PBLH means relative
506 to SONDE. These differences suggest an overestimation (+15%) and underestimation
507 (-23%) by the DNN model for the GoAmazon and CACTI sites, respectively, compared
508 to the more consistent PBLH values at the SGP site.

509 The evident systematic deviations when applying the SGP-trained DNN model to
510 the diverse environments of GoAmazon and CACTI underscore the challenges in
511 generalizing the model to regions with significantly different meteorological
512 backgrounds. These findings point to the potential of DNN models for PBLH estimation
513 while also highlighting the necessity for region-specific model adjustments.

514

515 **5 Summary**

516 This study has developed a Multi-Structure DNN model for estimating PBLH using
517 conventional meteorological data. The DNN model is developed by leveraging a long-
518 term dataset of PBLH derived from radiosonde data and augmented with high-
519 resolution MPL and Doppler lidar observations. This model produced an PBLH dataset
520 over the SGP with robust accuracy, consistently yielding lower bias values across
521 various conditions and datasets. Utilizing conventional meteorological data, this

522 method generates a 27-year dataset over the SGP, encompassing periods with limited
523 remote sensing data availability. In situations where morning radiosonde data is
524 unavailable, ERA-5 data can be effectively employed to initiate the model, offering a
525 practical alternative.

526 An important aspect of this research involved comparing DNN models with diverse
527 remote sensing instruments. Although these instruments offer high temporal and
528 vertical resolution, discrepancies in PBLH estimation remain. Our DNN model,
529 leveraging a broad range of input features refined by their importance, constructs a
530 representation of PBL evolutions, frequently demonstrating a closer agreement with
531 SONDE-derived PBLH. In the absence of remote sensing data, the DNN model can
532 produce high-quality PBLH results from the conventional meteorology data.

533 The study has shown the DNN model's ability to synthesize complex patterns from
534 meteorological data, reflecting the versatility of machine learning in simulating the
535 boundary layer processes. Its application to varied geographic terrains and climates
536 during the GoAmazon and CACTI campaigns has further validated its adaptability,
537 demonstrating a high correlation between DNN-derived PBLH and SONDE-derived
538 PBLH. Nonetheless, systematic biases in regions outside the SGP highlight the
539 influence of regional factors in PBLH estimation and suggest the need for region-
540 specific refinements to the model.

541 In summary, this research introduces a machine learning framework for PBLH
542 estimation that is able to generate high-quality PBLH using meteorological data,
543 independent of remote sensing instruments. This methodology, alongside the datasets

544 derived from the deep learning model, is beneficial in advancing our understanding of
545 PBL daytime development including thermodynamics and dynamics. It also has
546 implications for improved representation of the PBL processes in weather forecasting
547 and climate models, particularly by offering the potential to diagnose PBL in models
548 through the integration of modeled meteorological data as input. Future efforts will be
549 directed towards refining this model to ensure its wide applicability over a global scale.
550 These developments aim to effectively tackle the challenges of systematic biases and
551 regional variability in PBLH estimation.

552

553 **Data Availability.** ARM radiosonde data, surface fluxes, and cloud masks are available
554 at <https://doi.org/10.5439/1333748> (ARM User Facility, 1994). The datasets of
555 planetary boundary layer height used in this study can be downloaded from
556 https://adc.arm.gov/discovery/#/results/instrument_class_code::pblht (last access: 7
557 January 2024; ARM User Facility, 2024). Climate Data Store offers the ERA-5
558 reanalysis data (<https://doi.org/10.24381/cds.adbb2d47>, Hersbach et al., 2023). The
559 DNN-derived PBLH datasets over the SGP, CACTI, and GoAmazon are available at
560 Zenodo (<https://zenodo.org/records/10633811>, Su, 2024) and will be uploaded to ARM
561 data archive as a product with detailed information upon acceptance.

562

563 **Author contributions.** TS conceptualized this study and carried out the analysis. TS
564 and YZ interpreted the data and wrote the manuscript. YZ supervised the project.

565

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568

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581

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906 **Table list:**

907 **Table 1.** This table lists the varying structures of hidden layers used by each ensemble
908 member for PBLH estimation. Each configuration is expressed as an array, with the
909 number of elements indicating the number of layers and each value specifying the
910 number of neurons activated in the corresponding layer. For instance, a structure
911 denoted as [52, 28, 16] comprises three hidden layers containing 52, 28, and 16 neurons,
912 respectively.

913

Ensemble Member	Different Structures in Hidden Layer	Ensemble Member	Different Structures in Hidden Layer
Member 1	[52, 28, 16]	Member 6	[57, 44, 19]
Member 2	[61, 43, 20]	Member 7	[55, 43, 19]
Member 3	[59, 45, 19]	Member 8	[57, 43, 15]
Member 4	[60, 45, 23]	Member 9	[59, 41, 20, 10]
Member 5	[57, 45, 23]	Member 10	[57, 43, 18, 9]

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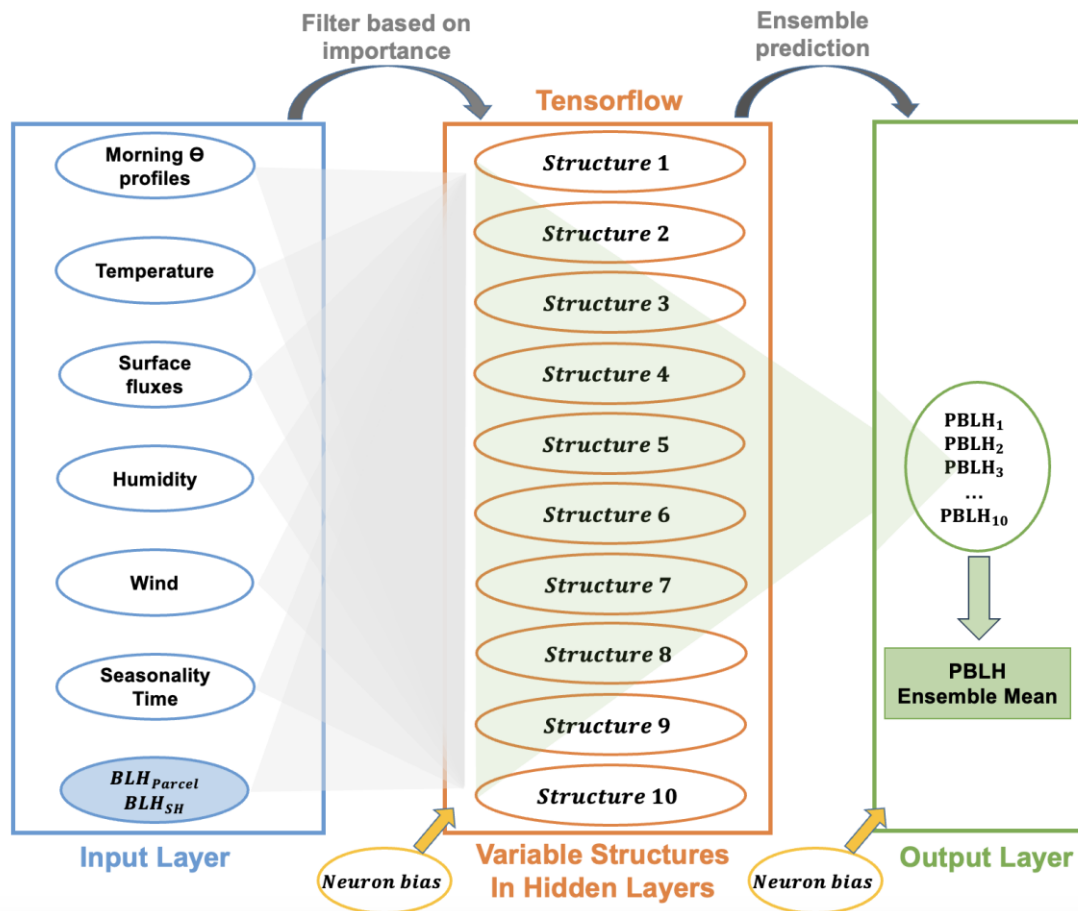
927 **Table 2.** Distribution of Dataset Samples for deep learning neural network (DNN)
 928 Training and Validation. This table details the sample data in different local time used
 929 for the development and validation of DNN to estimate planetary boundary layer height
 930 (PBLH). The first column lists the available PBLH derived from radiosonde (SONDE,
 931 Liu and Liang, 2010) during various local hours from 1994 to 2016. The second column
 932 supplements the dataset with a combined MPL and SONDE approach (Su et al. 2020b)
 933 and Doppler Lidar-derived PBLH (Sivaraman and Zhang, 2021) used in the absence of
 934 SONDE measurements. Seventy percent of the combined dataset from the first and
 935 second columns was randomly selected for the model's training. The third column
 936 provides the number of SONDE measurements available for validation purposes. Since
 937 morning SONDE serves as the input and boundary condition.

Local Time (h)	SONDE	Supplement Lidar Dataset	SONDE for Validation
5	7163	0	0
6	22	1181	0
7	3	1186	0
8	1225	2541	453
9	16	2629	8
10	9	2732	3
11	6513	13	3307
12	26	2797	9
13	14	2694	47
14	2131	2334	728
15	28	2555	9
16	3	2730	1
17	6503	2	3348

938 **Table 3.** The relative importance scores (%) of each input feature used in the deep
939 learning model to estimate the planetary boundary layer height. The features include
940 local time, month, relative humidity, U and V wind components, surface pressure,
941 precipitation, temperature, lifting condensation level (LCL), boundary layer height
942 derived from sensible heat and parcel methods (Sensible Heat BLH and Parcel Method
943 BLH), sensible and latent heat, and profiles of potential temperature (θ) at different
944 heights. The importance scores are expressed as percentages, indicating each feature's
945 relative contribution to the model's predictive accuracy, normalized to sum to 100%.

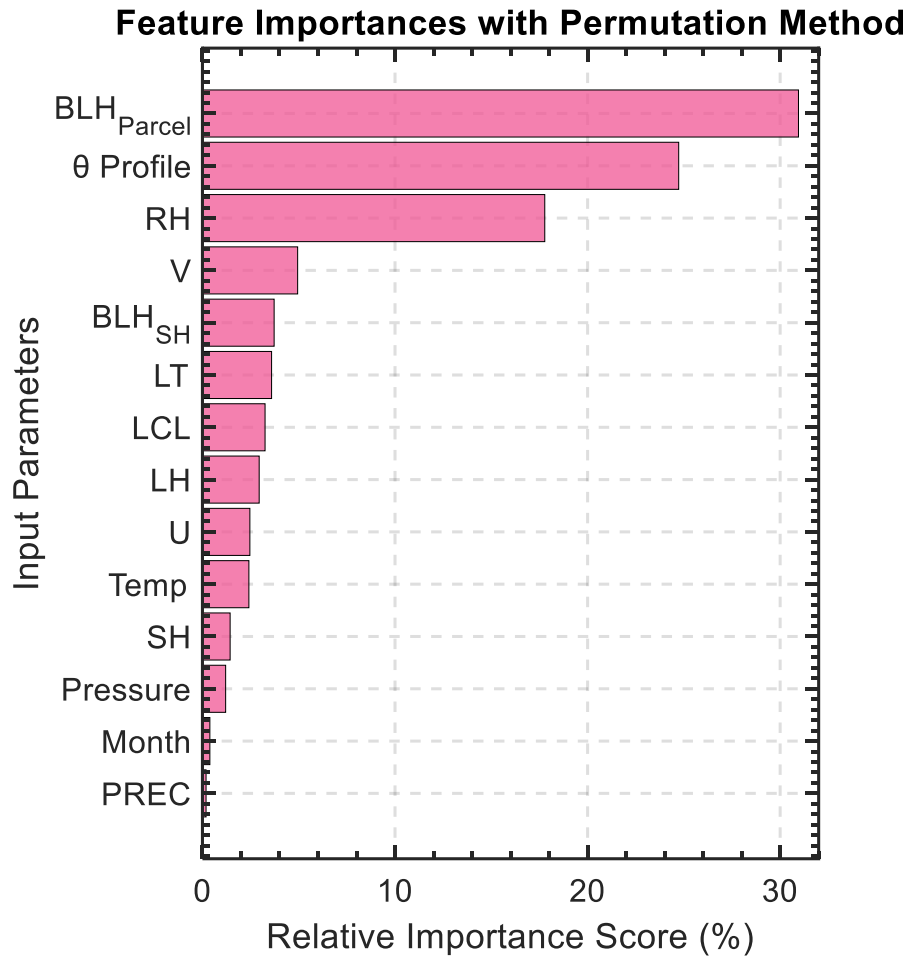
Feature	Importance (%)	Feature	Importance (%)
Local Time	0.385238096	θ 0.45km	0.589744268
Month	3.589829217	θ 0.5km	0.537731259
RH (i-1)	1.525447612	θ 0.55km	0.534610382
RH (i)	16.25123402	θ 0.6km	0.552997086
U Wind (i-1)	0.385834048	θ 0.65km	0.431060615
U Wind (i)	2.076794013	θ 0.7km	0.342764903
V Wind (i-1)	2.537910928	θ 0.75km	0.310147803
V Wind (i)	2.405275378	θ 0.8km	0.380120894
Surface Pressure (i-1)	0.187890954	θ 0.85km	0.468503984
Surface Pressure (i)	1.016443163	θ 0.9km	0.413498983
Rain Rate (i-1)	0.077638613	θ 0.95km	0.263411835
Rain Rate (i)	0.10979265	θ 1km	0.132168034
Temperature (i-1)	1.028603672	θ 1.1km	0.163035362
Temperature (i)	1.382663171	θ 1.2km	0.042643843
LCL (i-1)	0.330188472	θ 1.3km	-0.020619871
LCL (i)	2.92117154	θ 1.4km	-0.117425464
Sensible Heat BLH (i-1)	1.071904572	θ 1.5km	-0.020003889
Sensible Heat BLH (i)	2.650567178	θ 1.6km	0.10811159
Parcel Method BLH (i-1)	8.796298485	θ 1.7km	0.211953821
Parcel Method BLH (i)	22.15513884	θ 1.8km	0.092761568
Sensible Heat (i-1)	1.09273529	θ 1.9km	0.134436502
Sensible Heat (i)	0.344360459	θ 2km	0.109195516
Latent Heat (i-1)	1.240177933	θ 2.2km	-0.10805866
Latent Heat (i)	1.705848738	θ 2.4km	-0.217483536
θ 0.05km	13.55861389	θ 2.6km	-0.178324068
θ 0.1km	1.19646809	θ 2.8km	0.08071272
θ 0.15km	0.025100917	θ 3km	0.249503653
θ 0.2km	0.193888217	θ 3.2km	0.143137953
θ 0.25km	0.445161715	θ 3.4km	0.19819078
θ 0.3km	0.572192811	θ 3.6km	0.158828504
θ 0.35km	0.751498918	θ 3.8km	0.185359544
θ 0.4km	0.768690105	θ 4km	1.046682377

Deep Neural Networks for estimating boundary layer height



947

948 **Figure 1.** Schematic of the multi-structure deep neural networks (DNN) used for
 949 estimating the planetary boundary layer height (PBLH). Input features, including
 950 morning potential temperature profiles, temperature, wind, humidity, surface fluxes,
 951 seasonality, and time, are filtered based on importance and fed into the network. The
 952 system comprises ten distinct hidden layer structures, each processing the inputs to
 953 model PBLH. The outputs from these structures are then synthesized to determine the
 954 final PBLH value, leveraging the diverse representations of atmospheric dynamics
 955 captured by each neural network configuration. Neuron biases are applied at the output
 956 and hidden layers to fine-tune the model's performance.



957

958 **Figure 2.** Feature importance with permutation method in the deep learning model.

959 This table presents the importance scores of each input feature used in the deep learning

960 model to estimate the PBLH. The features include local time (LT), month, relative

961 humidity (RH), surface U and V wind components, pressure at the surface (Pressure),

962 precipitation (PREC), surface temperature (Temp), sensible and latent heat (SH and

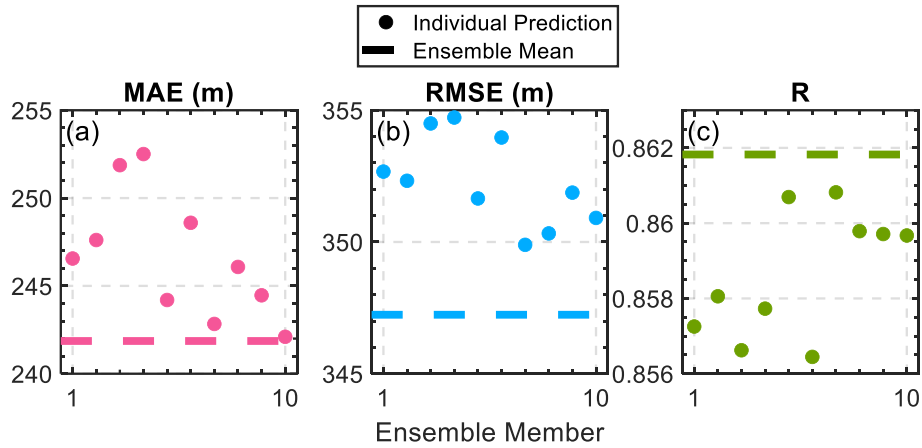
963 LH), surface-derived lifting condensation level (LCL), boundary layer height derived

964 from sensible heat and parcel methods (BLH_{Parcel} and BLH_{SH}), and morning profiles

965 of potential temperature (θ Profile). The importance scores are presented as

966 percentages, representing each feature's relative contribution to the model's predictive

967 accuracy, normalized to sum to 100%.

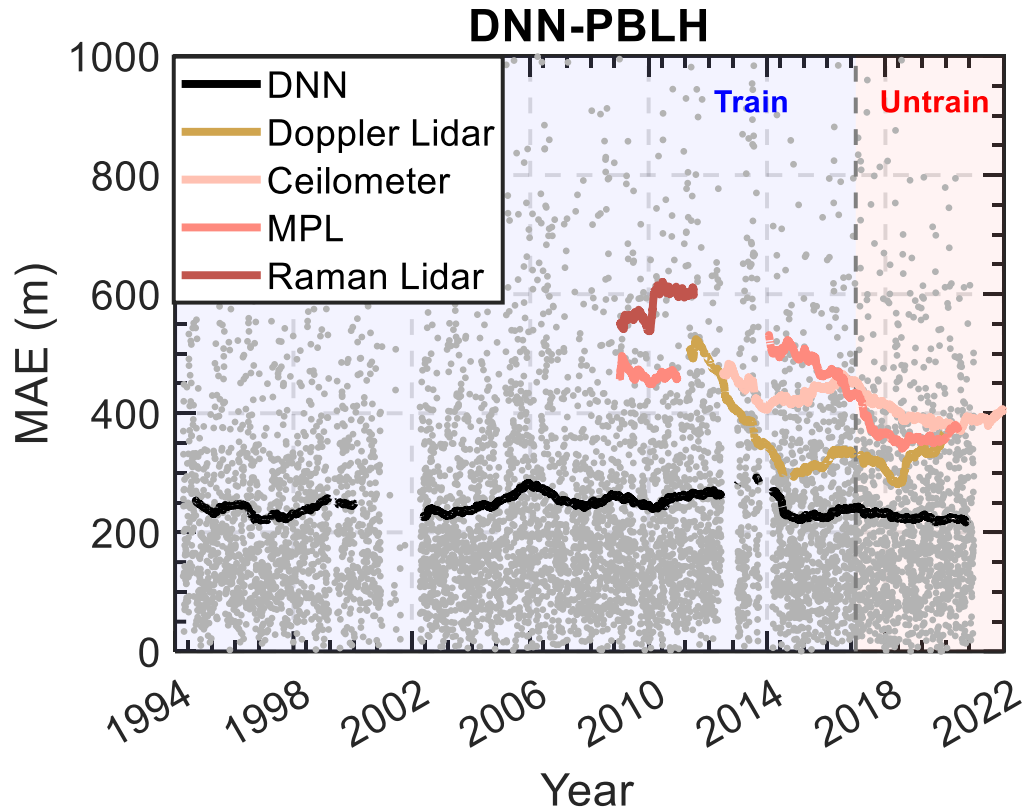


968

969 **Figure 3:** Performance metrics of individual ensemble members and the ensemble
 970 mean in estimating planetary boundary layer height (PBLH). Panel (a) displays the
 971 mean absolute error (MAE), panel (b) the root mean square error (RMSE), and panel
 972 (c) the correlation coefficient (R) for each of the ten ensemble members (represented
 973 by dots) and the ensemble mean (indicated by the horizontal dash line). The ensemble
 974 approach demonstrates improved accuracy and reliability in PBLH estimation as
 975 evidenced by the aggregation of individual model predictions into a robust ensemble
 976 mean.

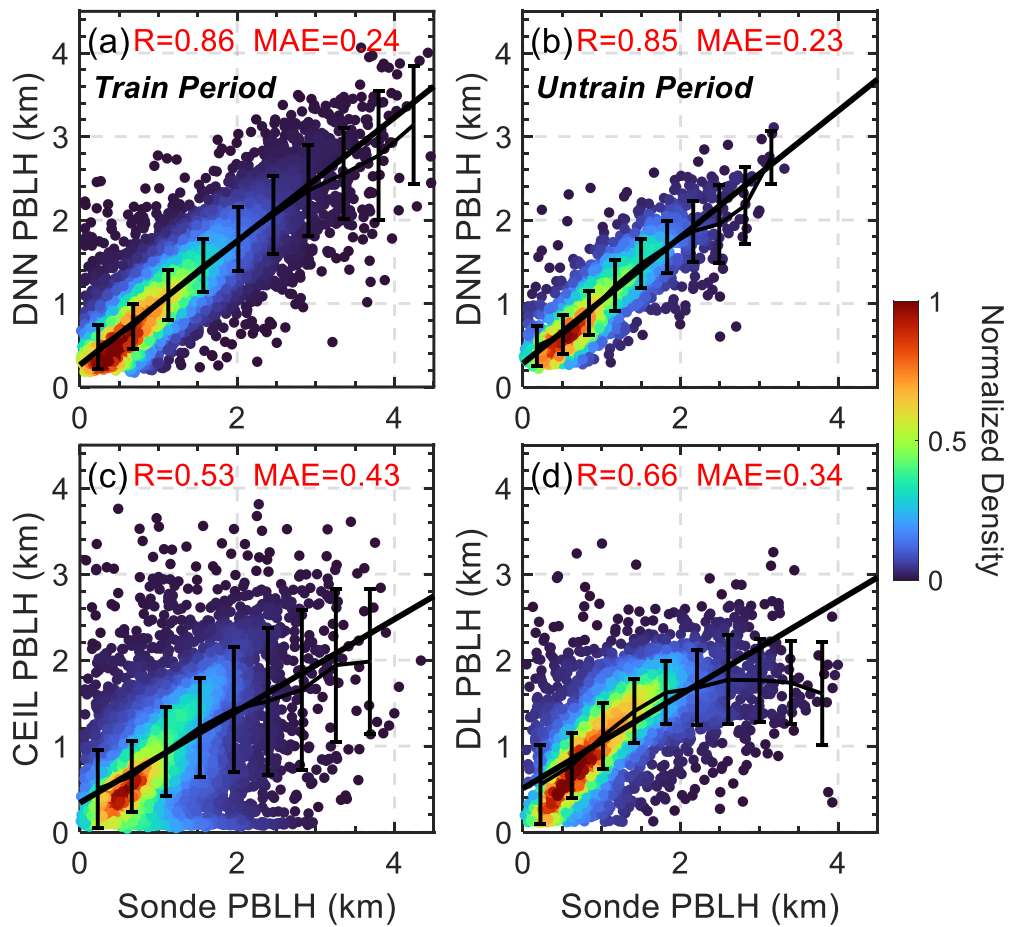
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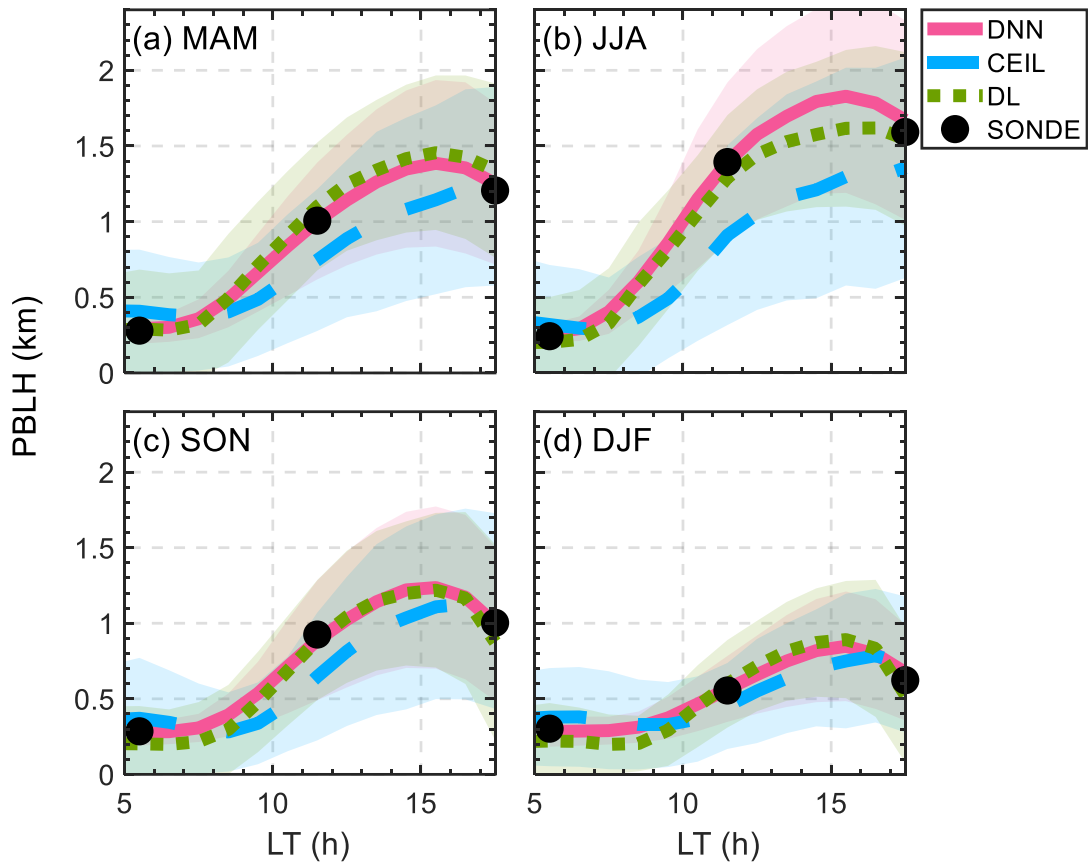
979

980 **Figure 4:** Comparative analysis of the mean absolute error (MAE) in PBLH estimation
 981 using different methodologies. PBLH derived from SONDE is considered as the ground
 982 truth. The DNN approach is shown in black, doppler lidar (Sivaraman and Zhang. 2021)
 983 in yellow, ceilometer (Zhang et al. 2022) in pink, micro-pulse lidar (MPL, Sawyer and
 984 Li. 2013) in light red, and Raman lidar (Ferrare. 2012) in dark red. DNN model is
 985 trained during 1994-2016. Individual MAE values for DNN are represented by gray
 986 dots, while the solid lines denote the smoothed MAE for each method with a 2-year
 987 smooth window.



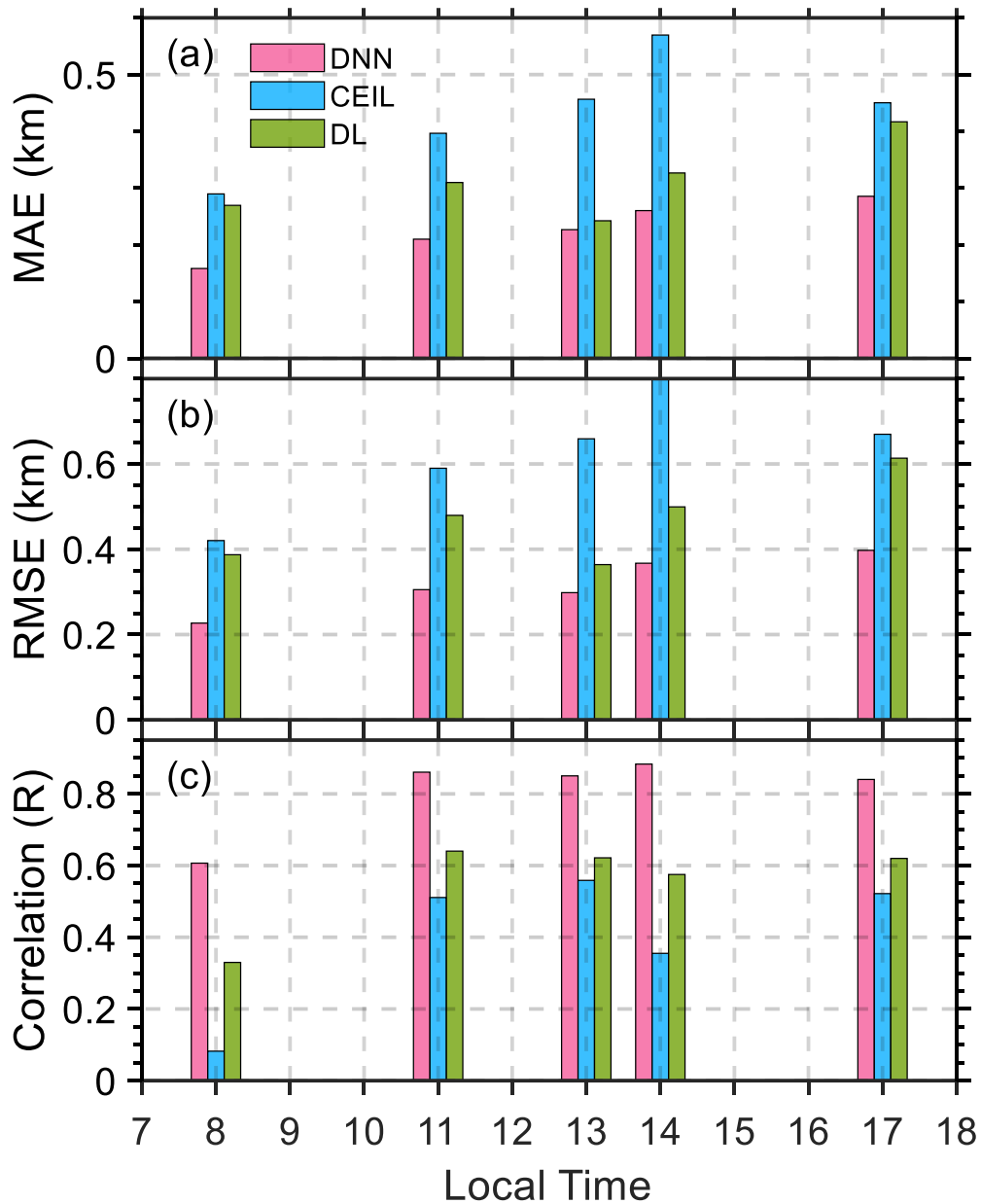
988

989 **Figure 5:** Scatter plots comparing observed radiosonde (SONDE) PBLH with estimates
 990 from the machine learning model and lidar observations. Panels (a) and (b) show the
 991 PBLH estimated by the deep neural network (DNN) during the trained period (1994-
 992 2016) and the untrained period (2017-2020), respectively, with corresponding
 993 correlation coefficients (R) and mean absolute errors (MAE). Panels (c) and (d) display
 994 comparisons of Sonde PBLH with ceilometer (CEIL) and doppler lidar (DL) derived
 995 PBLH, respectively. The color gradient indicates the normalized density of data points,
 996 while the solid black line represents the line of best fit and error bars indicates the mean
 997 and standard deviations for each bin.



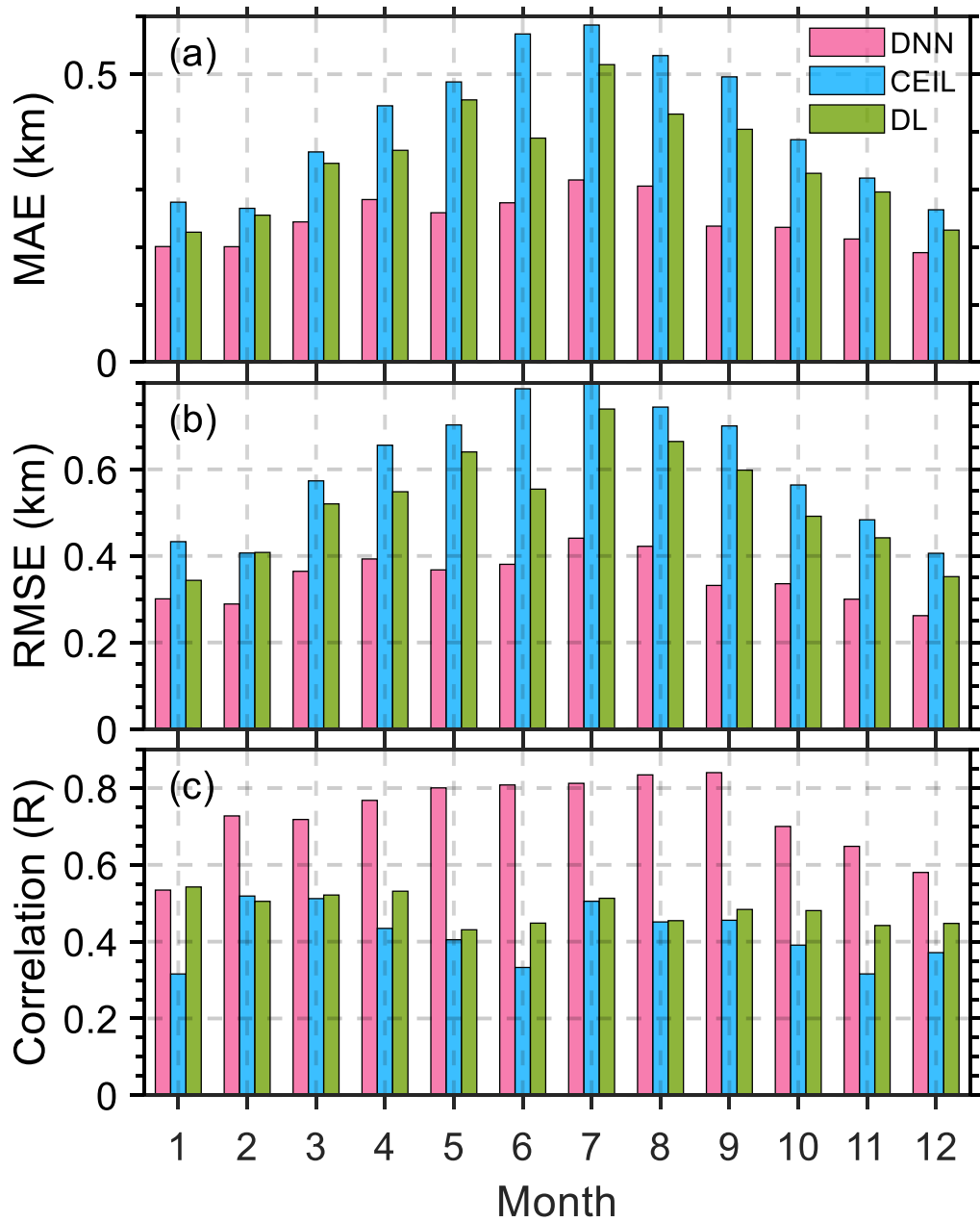
998

999 **Figure 6:** Seasonal-averaged daytime evolution of planetary boundary layer height
 1000 (PBLH) derived from various methods. The panels represent the mean PBLH values
 1001 throughout the day for different seasons: (a) March-April-May (MAM), (b) June-July-
 1002 August (JJA), (c) September-October-November (SON), and (d) December-January-
 1003 February (DJF). The PBLH values estimated by the deep neural network (DNN) are
 1004 shown in red, ceilometer (CEIL) estimates in blue, Doppler lidar (DL) in green, and
 1005 observed radiosonde (SONDE) data in black. Shaded areas around the lines indicate the
 1006 standard deviations within each method.



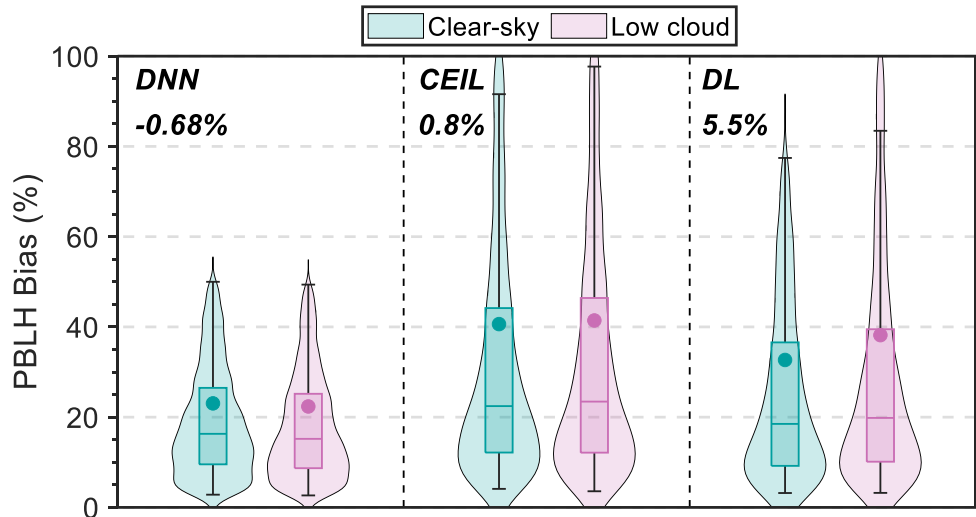
1007

1008 **Figure 7:** Diurnal variations in the performance metrics for estimating PBLH using
 1009 different datasets. (a) Shows the correlation coefficient (R), (b) represents the root mean
 1010 square error (RMSE), and (c) depicts the mean absolute error (MAE) at various local
 1011 times throughout the day. The deep learning neural network (DNN) estimates are in
 1012 blue, ceilometer (CEIL) derived estimates are in pink, and doppler lidar (DL) estimates
 1013 are in green. Note that these biases metrics are calculated using SONDE PBLH as the
 1014 standard. The availability of SONDE data for different hours is detailed in Table 2.



1015

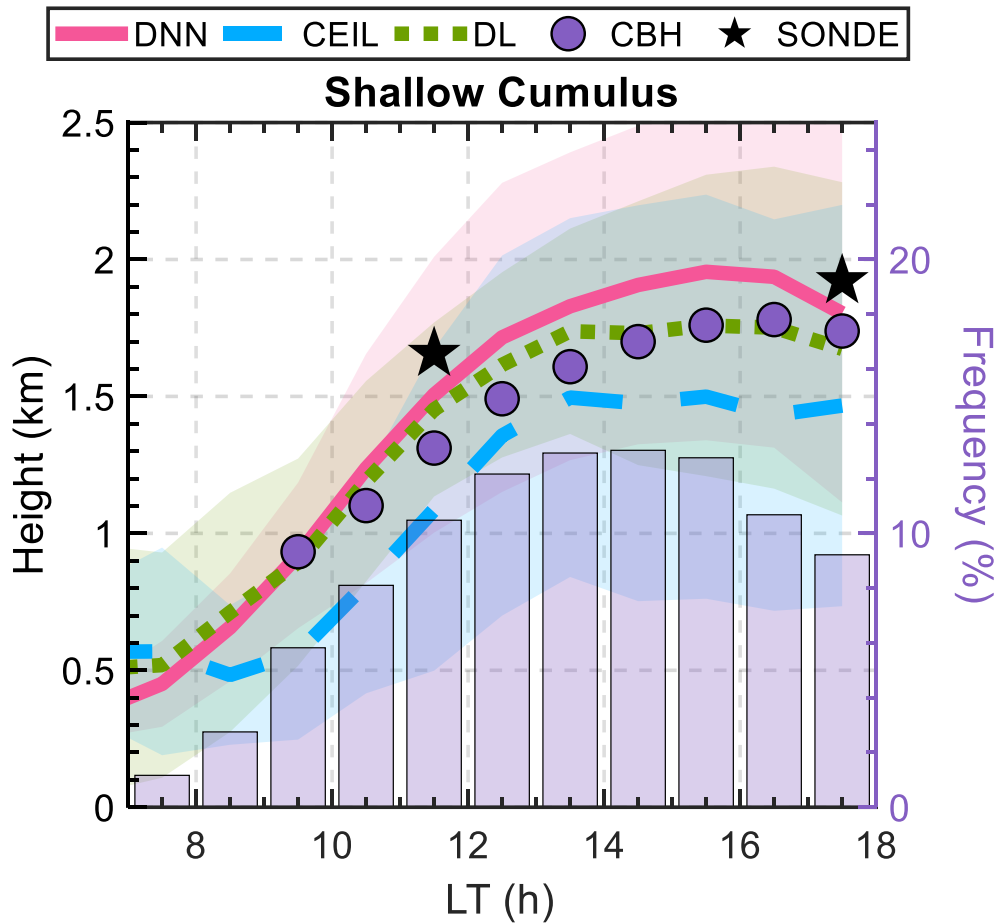
1016 **Figure 8:** Similar to Figure 7, but for MAE, RMSE, and R for different month.



1017

1018 **Figure 9:** Comparative analysis of PBLH estimation bias under clear-sky and low cloud
 1019 conditions for various methods. Bias percentages are computed as the absolute bias
 1020 normalized by the mean PBLH for each condition, with the number above each method
 1021 indicating the difference in bias between low cloud and clear-sky scenarios. The
 1022 boxplots detail the 10th, 25th, 50th, 75th, and 90th percentiles, while shaded areas in
 1023 violin plots illustrate the distribution of dataset biases. The dots indicate the mean value
 1024 for each condition.

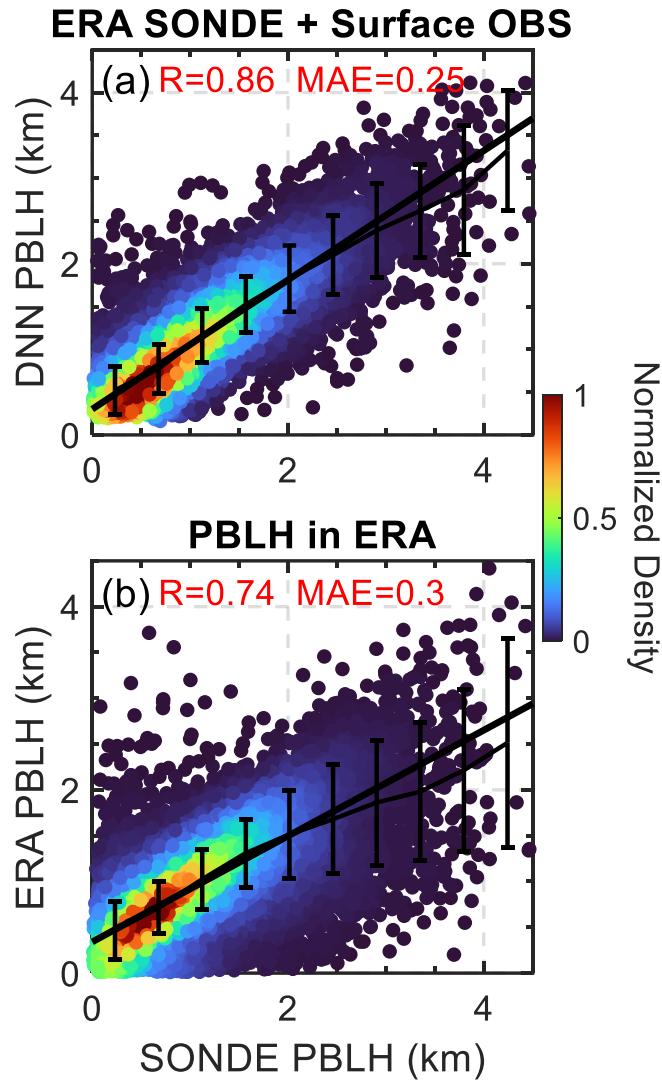
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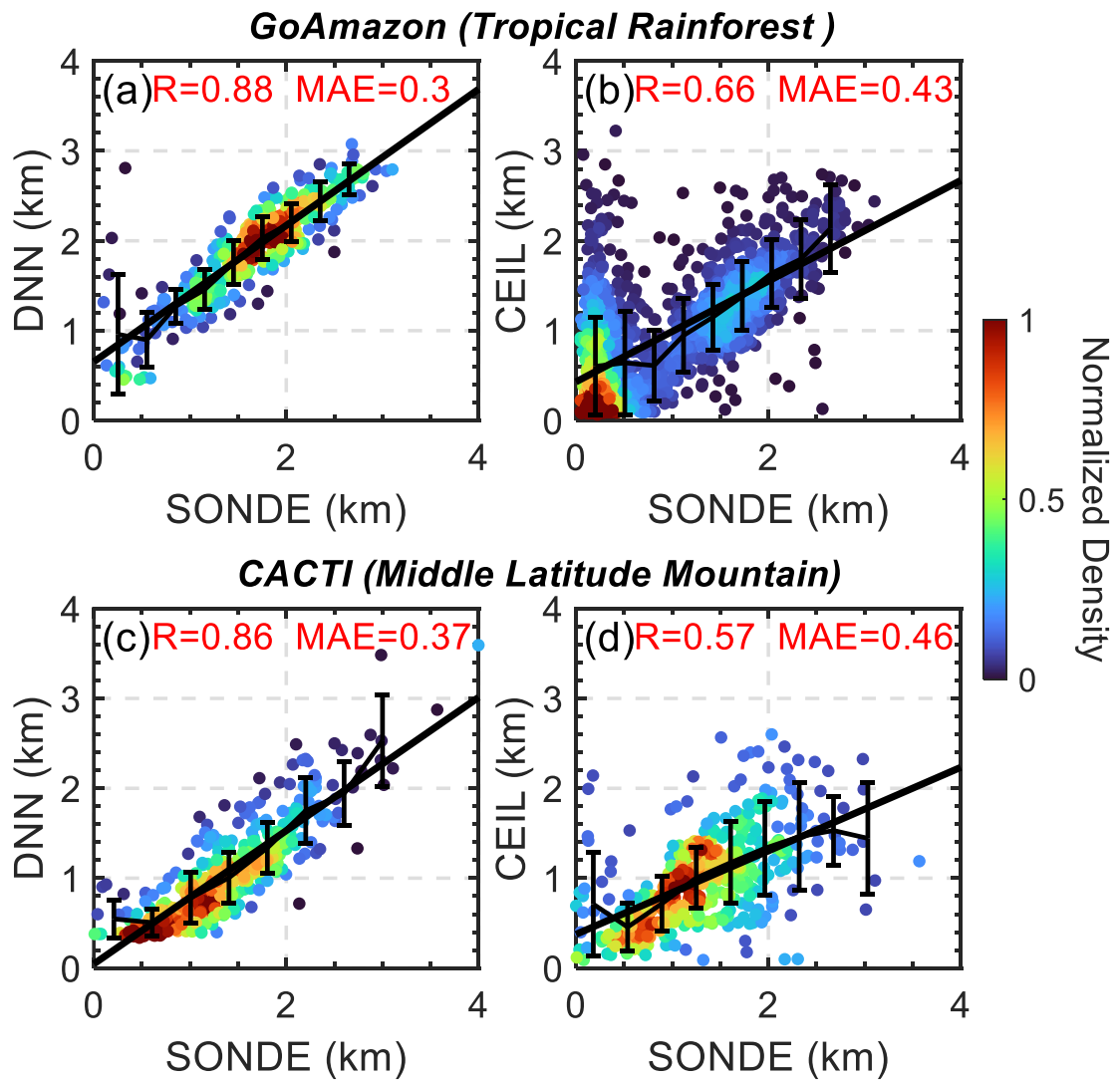
1027 **Figure 10:** Daytime evolution of planetary boundary layer height (PBLH) derived from
 1028 various methods under the shallow cumulus condition. PBLH values estimated by the
 1029 deep neural network (DNN) are shown in red, ceilometer (CEIL) estimates in blue,
 1030 Doppler lidar (DL) in green. Observed radiosonde (SONDE) data are represented by
 1031 black stars. Purple bars show the relative frequency of shallow cumulus occurrences
 1032 throughout the day, while purple dots mark the corresponding cloud-base heights
 1033 (CBH). Shaded areas around each line reflect the standard deviations for each method.

1034



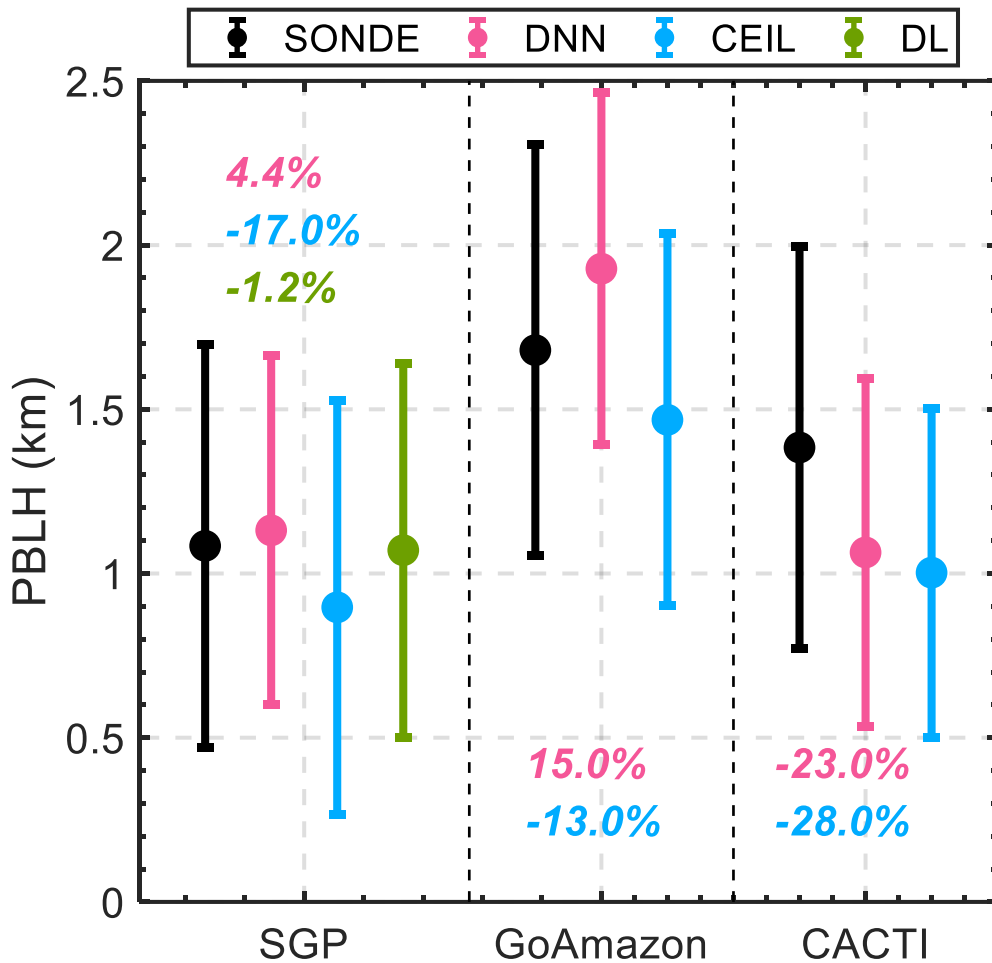
1035

1036 **Figure 11:** Scatter plots comparing SONDE PBLH with estimates from the DNN and
 1037 ERA-5. (a) The comparison between observed SONDE PBLH and estimates from the
 1038 DNN model, which utilizes morning temperature profiles (5 LT) from ERA-5 (ERA
 1039 Profile) and observed surface meteorological data (surface OBS) as inputs. (b) The
 1040 correlation comparison observed SONDE PBLH and PBLH model outputs from the
 1041 ERA-5 datasets. The color gradient in both panels represents the normalized density of
 1042 data points, while the solid black line indicates the linear regression, and the error bars
 1043 denote the mean and standard deviations for each bin.



1044

1045 **Figure 12:** Validation of the DNN trained over the SGP for the GoAmazon (Tropical
 1046 Rainforest) and CACTI (Middle Latitude Mountain) field campaigns. Panels (a) and (c)
 1047 illustrate the correlation (R) and mean absolute error (MAE) between DNN predictions
 1048 and SONDE observations for GoAmazon and CACTI, respectively. Panels (b) and (d)
 1049 show the performance of ceilometer (CEIL) derived PBLH compared to SONDE for
 1050 the same campaigns. The color gradient indicates the normalized density of data points,
 1051 while the solid black line represents the line of best fit and error bars indicates the mean
 1052 and standard deviations for each bin.



1053

1054 **Figure 13:** Comparative PBLH mean (dots) and standard deviations (error bars) across
 1055 ARM sites (SGP, GoAmazon, and CACTI). The datasets are derived from radiosonde
 1056 (SONDE, in black), the DNN model (in pink), ceilometer (CEIL, in blue), and Doppler
 1057 lidar (DL, in green), respectively. Noted the DL-derived PBLH is only available at the
 1058 SGP. The percentages in various colors denote the differences in PBLH means derived
 1059 from the DNN, CEIL, and DL methods relative to SONDE observations. To mitigate
 1060 sampling bias, these mean values and standard deviations are computed exclusively for
 1061 intervals where all instruments have concurrently available data.