

We thank the reviewer for carefully reading our manuscript and helpful comments. Below, the reviewer’s comments are marked in blue and our answers to the comments are written in black.

The manuscript “Retrieval of stratospheric aerosol extinction coefficients from OMPS-LP measurements” by Rozanov et al. presents a new retrieval algorithm to obtain vertical profiles of the aerosol extinction coefficient. The main claim of the paper is that by avoiding altitude normalization, the algorithm becomes almost completely independent of the “prior aerosol extinction profile.” However, in my view, the authors did not provide sufficient evidence to prove this point. Particularly, in Section 5, the authors wrote that uncertainties about the aerosol concentration at the normalization altitude would lead to a strong sensitivity to the a priori extinction profile across the entire vertical range. Figure 3 shows results for two algorithms (V1.0.9 and V2.1) and concludes that by removing altitude normalization in V2.1, the retrieved profiles become almost insensitive to a priori. However, I am afraid that the authors are comparing apples and oranges here. We (readers) do not know if the two algorithms use the same L1 data or different data because the differences in the magnitude of retrieved aerosol extinction coefficients are quite large between the two algorithms, as shown in Figure 5. The authors do not describe all the algorithmic differences between the two algorithms to convince the reader that the changes they see in Figure 3 are caused by the normalization at higher altitude.

Actually, the investigation was done in exactly the same way as suggested by the reviewer. The version we denoted in this part of the manuscript as V1.0.9 was actually V2.1 with the normalization switched to the upper tangent height (same as in V1.0.9). We thought using the notation V1.0.9 makes the things more easy for the reader, this seems, however, to cause confusion. We apologize for this. In the revised manuscript we changed the notations and revised the corresponding text.

The authors listed three main reasons why altitude normalization can negatively affect aerosol retrievals: larger stray light at the normalization altitude, uncertainties about the aerosol amount at the normalization altitude (that comes from a priori), and scene reflectivity (albedo). I agree with all three points; however, I don’t understand how any of these factors can lead to a strong dependence on the a priori throughout the entire vertical range. To prove the claim, the authors perturbed a priori profiles by increasing them by a factor of 2 and 3 and ran retrievals using the two models. First of all, if the authors believe that it’s the uncertainties in aerosol concentration at the normalization altitude that affect aerosol retrievals below, then they should arbitrarily increase aerosol at the normalization altitude rather than the entire profile.

We prefer to keep the plot with scaling the entire a priori profile in the main text of the manuscript for the reason that this makes our point easier to understand (changing the prior profiles only above the reference tangent height does not change them within the plot range). To address the reviewer’s comment

we added a plot to the supplement showing the changes in the resulting profiles when we change the a priori profile only at and above the reference tangent height and only below the reference tangent height, respectively.

The retrieval sensitivity to a priori can also be estimated using AKs (see Rodgers, 2000). The AKs for V2.1 are shown in Figure 2, but V1.0.9 AKs had not been demonstrated to readers. Can you please plot them as well? Can you estimate sensitivity to a priori using the equation (Rodgers et al., 2000) and check if it's consistent with what you observe from direct perturbations of the a priori?

The averaging kernels for V1.0.9 look very similar to those from V2.1. The reason for that is that both retrievals are almost independent of the a priori information in the retrieval range, i.e. below the reference tangent height. The fact that V1.0.9 retrieval is strongly sensitive to the a priori profile at and above the reference tangent height, i.e. outside the retrieval range, is not reflected in the averaging kernels, as it is not a dependence on the a priori information in the sense of Rodgers (2000). Following the terminology of Rodgers (2000), it is rather a parameter error. For this reason, the formalism for the estimation of the influence of the a priori information described by Rodgers (2000) is not useful in this case.

Secondly, the authors claim that scene albedo R derived at 40 km depends on aerosol, which is true, but I am not sure how that can increase the sensitivity to the a priori aerosol. Since the background aerosol amount is negligibly small at 40.5 km, its contribution to R is quite small compared to the pure Rayleigh atmosphere. I agree that the change in R will affect aerosol retrievals, but I am not sure how that can increase the sensitivity to the a priori aerosol. The authors should provide evidence to support this claim.

We would like to point out that the reference tangent height in V1.0.9 was at 37.5 km rather than at 40.5 km. At least at this altitude the aerosol signal is not negligibly small. If this was the case, there would be no dependence on the a priori profile used at these altitudes (see Fig. S1 in the new supplement). As the assumed aerosol around 37.5 km has a non negligible contribution to the measured signal it is also non-negligible when retrieving the effective albedo. The smaller our assumed aerosol concentration at 37.5 km, the larger the retrieved effective albedo becomes. This is because it compensates for the signal missing due to the underestimation of the aerosols. A bias in the retrieved albedo produces, in turn, an additional retrieval error. We agree with the reviewer that this effect is small for very clean upper stratospheric conditions but this is certainly not the case if a strongly elevated aerosol plume is observed as e.g. for the Hunga-Tonga eruption (see Fig. 4 of the paper).

I believe the method proposed in this paper can be useful for aerosol retrievals in perturbed conditions like those after the Hunga eruption, when the aerosol at 40 km was significantly different from the climatology. However, in

my view, the paper needs to be substantially revised, and the authors have to provide more supporting evidence for their main claim.

As we discussed above, we consider that the concerns of the reviewer were a result of our not explaining adequately the dependence on the a priori profile shown in Fig. 3 of the manuscript. Figure 3 shows the difference between V2.1 running similarly to V 1.0.9 and V2.1. This difference is determined by the aerosol concentration at and above the reference tangent height. We hope that our clarification about these differences between the retrievals and new Fig. S1 in the supplement sufficiently address the issue raised by the reviewer.

Major Comments: Title: The title of the paper is too vague and does not reflect the content of the paper. As the authors pointed out, there are multiple groups and multiple aerosol retrieval algorithms that use OMPS LP measurements to derive aerosol extinction. The title should be changed to reflect the paper’s content.

We added “sun-normalized” to the title to make it reflect more clearly the objectives and ambition of our manuscript.

Abstract, line 4: There is a statement in the paper, which is repeated multiple times, saying that the novelty of the presented algorithm is that “the method employs the normalization of the limb radiances to the solar irradiance in contrast to the normalization by a limb measurement at an upper tangent height, which is used by most of the other published limb-scatter retrievals.” A search in the literature reveals that, for example, NASA’s retrieval algorithm (Loughman et al., 2018; Chen et al., 2018; Taha et al., 2021) uses sun-normalized radiances to derive aerosol extinction. Indeed, the NASA algorithm requires the altitude normalization at higher altitudes, but it is incorrect to state that nobody uses sun-normalization.

We disagree with the reviewer. Loughman et al. (2018) use the altitude-normalized radiances (ANRs) for their retrieval, which is the radiance at a tangent height of interest divided by the radiance at a selected normalization tangent height (see the beginning of their section 3.4). Chen et al. (2018) write at the beginning of their Sect. 2 “ the radiances are normalized (i.e., divided by their value at the normalization altitude, 40.5 km) in all cases”, Taha et al. (2021) write in their section 2.2.2: “To reduce the stray light effect on the retrieval at longer wavelengths, h_n was lowered to 38.5 km in V2.0 (from the 40.5 km value used in previous versions).” In summary, all these manuscripts are reporting the results of algorithms which use the reference tangent height normalization. They certainly have used the sun-normalized radiances before the additional normalization as it is provided in Level-1 OMPS-LP product, this has, however, no meaning with respect to the discussed topic as the solar normalization cancels out when dividing by the reference tangent height. We reworded the sentence in the abstract to clarify that the solar irradiance normalization means the absence of the normalization to the reference tangent height.

Abstract: A large fraction of the paper is dedicated to comparisons with other instruments (SAGE III and OSIRIS). The statement in the abstract declares that differences are mostly within 25%, but such agreement is only seen in a relatively narrow vertical range, and outside that range, the differences are much larger. In my view, the authors should clearly identify in the abstract the vertical and latitudinal ranges where the agreement is within the desired 25%.

We included the vertical range for both comparisons and the latitudinal limitation for OSIRIS in the abstract

Page 2, lines 24-26: The authors stated that substantial ozone losses were observed after the 2020 Australian fires and the Hunga eruption and provided references. In my view, the words “significant losses” exaggerate the losses described in the cited studies. Instead of using the words “significant ozone losses,” the authors should quote numbers from the cited publications.

We added the numbers from the papers as recommended by the reviewer.

Page 3, line 57: There is an extensive list of publications that estimate the SO₂ amount injected by the Hunga eruption. It would be better to quote numbers rather than say “a significant amount.”

We quoted numbers in accordance with the reviewer’s recommendation and added additional references to support these numbers.

Page 4, Section 3, line 115: Are you solving Equation 1 with respect to the initial guess or a priori profile? Is the first guess in your terminology the same as a priori?

Yes, initial guess profile is the a priori profile. This is clarified in the revised manuscript.

Page 4, line 116: By removing the altitude normalization, you need to accurately know the surface albedo. Can you reduce the number of iterations by retrieving reflectivity R₀ at, say, 40 or 45 km first and use this as the initial guess for R?

In the standard retrieval, the effective surface albedo is quite close to the final value already after 2-3 iterations. Thus, there is no advantage in doing a pre-retrieval, especially taking into account that the effective albedo retrieved from measurements at upper tangent heights might be wrong because of a correlation with the stratospheric aerosol signal or stray light contamination.

Page 5, Section 3, lines 26-28: It is not the normalization to solar radiances that makes retrievals more sensitive to upwelling radiances. It is the absence of the altitude normalization.

Changed to “As the normalization by the solar irradiance instead of the reference tangent height makes the retrievals more sensitive to the surface reflectance ...” to avoid a confusion.

Page 6, lines 170-173: The described convergence criteria are questionable and definitely are not optimal. The range between 15 and 28 km might be reasonable for the background aerosol conditions. However, for the case with a dense aerosol cloud like after the Hunga eruption, the line-of-sight optical depth becomes incredibly high. This means that the measurements at lower tangent points are not sensitive to changes at those altitudes, and the signal rather comes from upper levels that lie closer to the instrument. Under those conditions, instead of focusing on improving retrievals in places where the measurements are the most sensitive (based on K), the algorithm is pushed to retrieve hard in places with no sensitivity.

We agree, the convergence criteria we use might be sub-optimal for some situations but they ensure a reasonable convergence speed and a satisfactory quality of the results for the majority of the runs. In our opinion the only drawback of using suboptimal convergence criteria is a high number of iterations. No overall quality drawback is expected from this as no displacement of the retrieval focus at each particular iteration occurs.

Page 6, line 173: I am not sure that the algorithm with 100 iterations can be used in the operational environment. Can you plot a histogram showing the number of iterations under background conditions and under perturbed conditions (like volcanic eruptions or wildfires)?

Our algorithm has been developed as a research exercise and certainly needs an optimization to be used for a near real time processing. This optimization is however relatively straight forward. For now, we can easily process the entire data set within a few weeks using high performance computing facilities. The requested histograms are now provided in the supplement.

Page 10, lines 247-250: The measurement noise can be quite different between LP and SCIAMACHY. I would not extrapolate conclusions derived from the analysis of SCIAMACHY spectra to OMPS LP.

We agree with the reviewer's comment and added the following comment to the manuscript text: "Although, results from SCIAMACHY cannot be directly transferred to other instruments, a degradation of the measurement quality with an increasing tangent height is rather common for limb-scatter observations."

Page 12, Figure 5: What do the horizontal green lines at 9 km represent? Does it mean that positive differences for v1.09 switch to negative below 9 km? If I interpret the error bars correctly, the standard deviation for differences is larger than +/-100% at lower altitudes (depending on latitude zone). How meaningful are the comparisons with a standard deviation greater than 100%?

Yes, these lines mean a switch between positive and negative differences. The reason for this is that V1.0.9 retrieves only down to 12.5 km and switches to a priori below. From the mathematical inversion theory point of view, signals with amplitudes below the noise level can still be detected. However, the obtained results should be handled with care.

Page 12, line 275: Have you described the collocation criteria for comparisons with SAGE III?

Indeed, we forgot to describe collocation criteria for the comparison with SAGE III. They are discussed at the beginning of Sect. 6 of the revised manuscript.

Page 14, Figure 7: Are you calculating zonal means from all available measurements for each instrument independently? If so, then the difference in the temporal and spatial sampling (particularly with SAGE III) can produce biases that are not accounted for.

Yes, monthly zonal means are calculated from each instrument independently. We agree this can introduce biases. The usage of colocated measurements is, however, disadvantageous for a time series comparison as the sampling of the measurements is strongly reduced. As the end users are rather interested in monthly mean climatologies [1, 2], we think our comparison is appropriate. Our internal tests show that general conclusions remain the same when collocated measurements are used in the comparison.

Page 14, lines 294-295 and Figure 7: OMPS LP aerosol retrievals at 869 nm were converted to 750 nm to compare with OSIRIS. Since this study validates OMPS LP retrievals, it would be better to run comparisons at the “native” wavelength and rather convert OSIRIS extinction to 869 nm.

As OMPS-LP data described in this paper are used to create a climatological data record of stratospheric aerosol extinction coefficients at 750 nm [2], we prefer to keep the comparison at this wavelength.

Page 15, lines 330-333: Why do you use 2022 data when you stated above that OSIRIS data quality degraded in 2022? Or is it a typo?

This is indeed a typo. We changed the range to 2012 – 2021.

Page 16, line 335: Both OSIRIS and LP have relatively dense sampling, so I am not sure what you meant here. Can you please elaborate on this?

Both instruments cannot measure in the darkness which leads to gaps at high latitudes in winter. As a result of the difference in the times of the equator crossing for OSIRIS and OMPS-LP orbits, the gaps are larger for OSIRIS.

Page 16, line 336: I agree that the cloud correction is one of the many factors that contribute to reduced quality of aerosol data and larger differences in the troposphere, but I will not say that it is the only one.

We think there is a consensus in the limb-scatter community that clouds is a major issue for the aerosol data comparisons in the troposphere. Certainly it is not the only reason for differences but this reason is enough to spoil the comparison if clouds are treated in different ways.

Multi-panel figures: Please add labels (a, b, c, etc.) on all figures that have multiple panels.

The labels have been added and the manuscript text has been adjusted to reference the panels in accordance with their labels.

Figure 10, legend: What do you mean by “Relative mean differences”? Do you calculate zonal means first and then calculate the difference between the two monthly zonal means? Then it should be “Relative differences.” Otherwise, clarify that in the text.

We agree, these are relative differences.

Section 7: There have been many publications in the last two years that describe the transport of volcanic aerosol after the Hunga eruption, which are not acknowledged here. Is there any reason for that? How do the conclusions of this study agree with previously published results?

We added a paragraph at the end of Sect. 7 comparing our results to those from other studies.

Minor comments: Page 2, line 46: it’s not clear from the context what “this range” refer to. It might be better to say “. . . to the aerosol at the normalization altitude”.

Corrected as suggested by the reviewer.

Page 2, line 46: It doesn’t sound right when you state that the knowledge is the major source of uncertainty. Perhaps, “the lack of knowledge” or “incomplete knowledge”.

Changed to “a lack of the knowledge”

Page 8, lines 213: should “for example”.

“example” is meant here as an adjective which belongs to “measurements”, the construction “an example measurement” is correct.

Page 8, line 214: should be “with the tangent point ground coordinates”

Changed

Page 8, line 216: should be “every third AK”

Changed

Page 11, line 258: the word “tangent” is used twice.

Corrected

References

- [1] Kovilakam, M., Thomason, L. W., Ernest, N., Rieger, L., Bourassa, A., and Millán, L.: The Global Space-based Stratospheric Aerosol Climatology (version 2.0): 1979 –2018, *Earth Syst. Sci. Data*, 12, 2607–2634, <https://doi.org/10.5194/essd-12-2607-2020>, 2020.
- [2] Sofieva, V. F., Rozanov, A., Szlag, M., Burrows, J. P., Retscher, C., Damadeo, R., Degenstein, D., Rieger, L. A., and Bourassa, A.: A Climate Data Record of Stratospheric Aerosols, *Earth Syst. Sci. Data Discuss.* [preprint], <https://doi.org/10.5194/essd-2023-538>, in review, 2024.