



Ionospheric Chaos in Solar quiet Current due to Sudden Stratospheric Warming Events Across Europe-Africa Sector

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Abstract

This study examines the ionospheric chaos in the solar quiet current, across Europe and Africa sectors during 2009 and 2021 Sudden Stratospheric Warming (SSW). The SSW was categorized into precondition, ascending, peak, descending, after and no-SSW phases based on the rising stratospheric temperature. Thirteen magnetometer stations, located within the geographical longitude of 26° to 40° across Europe and Africa sectors were considered. The magnetometer data obtained during the periods of SSW were used to derived the ionospheric solar quiet current time series. This solar quiet current time series was transformed into a complex network representation using the Horizontal Visibility Graph (HVG) approach and Fuzzy Entropy was employed on the



transformed solar quiet current time series to quantify the presence ionospheric chaos during the periods of SSW. The results revealed that the latitudinal distribution of entropy depicts high entropy values indicating the presence of ionospheric chaos in most of the stations situated within the European sector compared to stations in the African sector. A consistent low entropy values unveiling the presence of orderliness behavior was found to be prominent in the Africa sector. This dominance of orderliness behavior in the Africa sector during SSW means that the influence of SSW on the regional ionosphere of this sector is minimal. However, the pronounced features of ionospheric chaos found in the European sector reveal evidence of significant effects of SSW on the regional ionosphere in this sector. Finally, we found that after the peak phase of SSW, the ionospheric chaos is more pronounced.

Keywords: Sudden Stratospheric Warming (SSW), Solar quiet current Sq(H), Ionospheric chaos, Fuzzy Entropy, Horizontal Visibility Graph (HVG)

1. Introduction.

Natural systems, including the troposphere, stratosphere, and ionosphere, are significantly affected by regional climate changes, particularly temperature increases (Li et al., 2023; Fortin, 2017). As the temperature rises, there are a wide range of effects on weather conditions. For instance, less snowpack in mountain ranges and polar areas is experienced as a result of climate change and the snow melts faster (Williams et al., 2024; Fortin, 2017). The increasing number of extreme weather events is caused by climate change. Climate change does also impact the frequent occurrence of sudden stratospheric warming (SSW), with the consequence of communication disruptions, economic disruptions, transportation disruption and energy demand fluctuations (Yasyukevich et al., 2022; Li et al., 2023; Wright et al., 2021; Baldwin et al., 2021). The emergence of SSW is described by an atmospheric phenomenon where the stratospheric temperature increases rapidly



52 within a couple of days (>30 to $40k$). SSW is one of the usual meteorological events, where the
53 stratospheric temperature increases rapidly in the winter polar region due to the rapid growth of
54 quasi-stationary planetary waves (Baldwin et al., 2021; Butler et al., 2015). These quasi-stationary
55 planetary waves interact nonlinearly with the atmospheric tides, leading to the generation of both
56 migrating and non-migrating tidal components (Wang et al., 2021, 2023; Antokhina et al., 2023;
57 Ern et al., 2021; Liu et al., 2010; Chau et al., 2012). During SSW, the planetary waves and tidal
58 interaction can generate electromotive force that drive the electric currents and move the electrical
59 conducting air in the ionosphere across the earth's magnetic field (Siddiqui et al., 2018; Yamazaki
60 et al., 2020; Yamazaki, 2014; Wagner et al., 1980; Richmond, 1989; Yamazaki and Maute, 2017;
61 Yamazaki and Richmond, 2013). It has been reported that during SSW, the coupling processes at
62 lower atmosphere can propagate forcing that can reshape the plasma density variability of the
63 ionosphere. These reshape in ionospheric plasma possesses some dynamical characteristics that
64 are of nonlinear nature. The main mechanism responsible for the connections of this influences
65 includes planetary waves, atmospheric tides, and gravity waves (Goncharenko et al., 2021; Chau
66 et al., 2012; Goncharenko et al., 2010, 2012; Yiğit et al., 2024; Zhou et al., 2023; Gordiyenko et
67 al., 2024; Klimenko et al., 2018, 2019; Gómez-Escolar et al., 2014).

68 In the upper atmosphere, high energy injections into the ionosphere due to the solar wind-
69 magnetosphere coupling processes during geospace disturbances such as geomagnetic storms,
70 HILDCAAs and others are considered to be the major factors influencing the dynamics of the
71 ionosphere (Zou et al., 2014; Hajra et al., 2015; Tsurutani et al., 2020; Raeder et al., 2016;
72 Oludehinwa et al., 2023; Lin et al., 2022). However, (Goncharenko et al., 2010) and others have
73 demonstrated that the reshaping of electron density variability during SSWs is similar in magnitude
74 to that of moderate geomagnetic storms. Therefore, the occurrence of SSW can infer a



75 phenomenon where the ionospheric plasma density is reshaped and destabilized, thereby posing
76 an operational hazard on radio communication, navigation, and imaging system.

77 This calls for the need to examine the state of the ionosphere during SSW by investigating the
78 dynamics of the ionospheric current system during quiet periods as SSW emerges. Notably, the
79 dynamics of the ionospheric current changes with time. They exhibit complex fluctuations in form
80 of nonlinear behavior due to their continuous interactions in response to lower atmospheric
81 circulation and the energy deposition of solar wind-magnetosphere coupling processes. The
82 ionospheric currents during these meteorological and space weather events are propelling
83 phenomena that can unveil the connections between the lower atmosphere and upper atmosphere
84 (Pedatella and Forbes, 2010; Liu et al., 2010; Eswaraiah et al., 2017). According to the dynamo
85 theory, the neutral wind (U) in the dynamo region transports conducting plasma across the Earth's
86 magnetic fields (B), generating electric fields and currents that flow in the E-region of the
87 ionosphere (Wagner et al., 1980; Yamazaki and Richmond, 2013; Yamazaki et al., 2012b). This
88 ionospheric current system during geomagnetically quiet periods is referred to as Solar quiet Sq(H)
89 currents.

90 The ionospheric response to SSW introduces spatial and temporal variability in the dynamics of
91 the ionospheric current system and exhibits complex fluctuation signatures in its underlying
92 dynamics, which require further investigation. In addition, the research question of the contribution
93 of SSW formation to the regional ionosphere across the European-African sector needs special
94 attention. This observation forms the basis of this study to quantitatively examine the dynamics of
95 the ionospheric solar quiet current Sq(H) before its orderly and disorderly (chaotic) behavior
96 during the anomalous stratospheric temperature changes over the Europe-Africa sector. The
97 present study seeks to unravel the ionospheric chaotic dynamics in the solar quiet current Sq(H)



98 due to the emerging influence of the SSW over this sector. Also, to reveal the common features of
99 the ionospheric dynamics due to the occurrence of SSWs during solar minimum years.

100 Several studies have already investigated the solar quiet current during SSW (Yamazaki, 2014;
101 Bolaji et al., 2016a; Charles Owolabi and Babatunde Rabi and Kayode Oluyo, 2015; Bolaji et al.,
102 2015; Yamazaki and Maute, 2017; Yamazaki et al., 2011). (Fejer et al., 2011) studied the evidence
103 of possible equatorial vertical plasma drifts associated with SSW. They found a strongly enhanced
104 lunar semi-diurnal vertical plasma drift amplitude during early morning solar flux warming.
105 (Maute et al., 2014) also found evidence of ionospheric vertical drift changes during SSW and
106 attribute these drift changes to the interaction of specific tides and planetary waves. (Yamazaki et
107 al., 2012a) examined the ionospheric current system during SSW of 2002 and 2003 considering
108 the Northern and Southern Hemisphere across the east Asia region. They reported an additional
109 current system that would be superposed on the normal Sq current system and further suggest that
110 abnormally large lunar tidal winds played the main role in producing the additional current and
111 counter electrojet (CEJ). (Yamazaki et al., 2012b) further looked at the solar quiet current during
112 SSW. They noticed a significant decrease and increase in the intensity of the solar quiet current in
113 the Northern and Southern Hemispheres, respectively, with a decrease in the longitudinal
114 separation between the Northern and Southern eddies. (Yamazaki, 2014) examined solar and lunar
115 ionospheric tidal drives, which are thought to cause ionospheric electrodynamics effects during
116 SSW, by estimating the average solar and lunar ionospheric current systems for SSW and non-
117 SSW periods. He found that the lunar current intensity is enhanced during SSWs by approximately
118 75% while the solar current intensity is much smaller by approximately 10%. (Bolaji et al., 2016b)
119 reported the solar quiet current response in the Africa sector. They categorized the rising
120 stratospheric temperature during SSW into six phases (precondition, ascending, peak, descending,



121 after and no) and noticed a counter electrojet (CEJ) after the polar stratospheric temperature
122 reached its maximum value (SSW peak). In addition, a decrease in the Sq(H) magnitude was
123 observed at geomagnetic latitudes between 21.13°N (Fayum) in Egypt and 39.51°S (Durban) in
124 South Africa. (Siddiqui et al., 2018) investigated the evidence for an Equatorial Electrojet (EEJ)
125 during a major SSW event. They found a significant increase in the amplitude of the EEJ
126 semidiurnal lunar tides during all major SSW events investigated. The mechanism responsible for
127 the anomalies in the tropical lower thermosphere-ionosphere due to SSW 2009 was demonstrated
128 by (Klimenko et al., 2019). The authors claim that perturbations in the ionospheric conductivity
129 also make a significant contribution to the formation of the electric field response to SSW. They
130 also show that the phase change of the semidiurnal migrating solar tide (SW2) in the neutral wind
131 caused by the 2009 SSW at the altitude of the dynamo electric field generation has a crucial
132 importance for the SW2 phase change in the zonal electric field.

133 Notably, one of the atmospheric activities unveiling the coupling features between the lower and
134 upper atmospheric circulation is the dynamics of the ionospheric current system (Yamazaki et al.,
135 2012b; Bolaji et al., 2016b; Yamazaki and Richmond, 2013). The ionospheric current system is a
136 key factor that can reveal the dynamical characteristics of the ionosphere during SSW.
137 Interestingly, among all the aforementioned studies on ionospheric currents during SSW events, it
138 is noteworthy that these currents have not been examined from the aspect of chaos theory. Since
139 then, it has been reported that the lower and middle atmospheric circulation (i.e. troposphere-
140 stratosphere coupling) nonlinearly interact during SSW, modulating the dynamics of the
141 ionosphere. Therefore, this nonlinearity (out-of-equilibrium) features is the prevailing condition
142 for the existence of chaotic behavior in the ionosphere. Important characteristics of chaotic
143 behavior are the existence of a strange attractor and that a small change in the ionosphere due to



SSW can amplify the dynamics, leading to instability and divergence from its initial state (Takens, 1981). Additionally, the ionosphere is considered a dynamical system due to its continuous response to atmospheric forcing. Bifurcation behaviour within the ionospheric system is eminent (Hysell, 2020; Lapenta, 2020; Materassi, 2020). In some occasions, the ionospheric dynamics tend to exhibit either orderliness or disorderliness (chaotic) behavior in its underlying dynamics due to influences of atmospheric forces from lower and upper atmospheric circulations. Therefore, it is crucial to examine the dynamics of the ionospheric current system during quiet periods as the stratospheric temperature rises during SSW. Implementing the concept of nonlinear dynamics, informed by information theory and graph theory, presents an innovative approach to reveal the contributions of SSW to the ionospheric region across the Europe and Africa sectors. Therefore, this study focuses on the ionospheric chaos (order-disorder behavior) in the ionospheric current system during quiet periods when the stratospheric temperature increases anomalously.

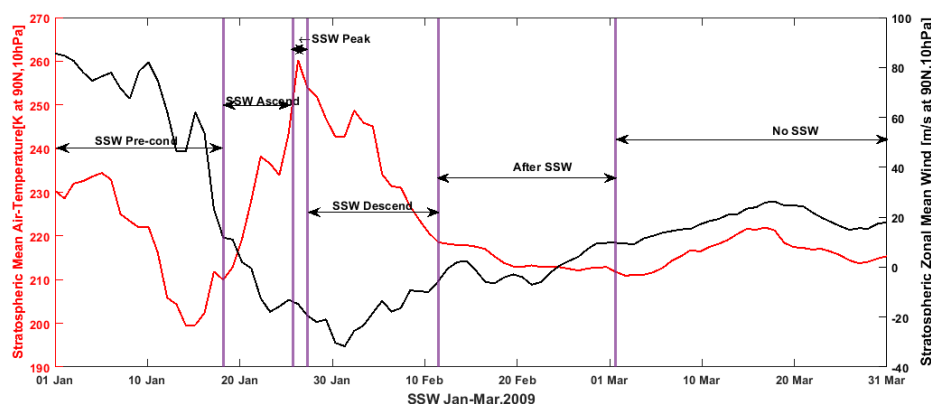
In this context, a theoretically robust method known as the Horizontal Visibility graph (HVG), derived from graph theory, is employed to preprocess the solar quiet current time series, while the ionospheric chaos during SSW is revealed through Fuzzy Entropy (FuzzyEn). HVG and FuzzyEn have been demonstrated to be useful tools in inter-disciplinary applications and when analyzing challenging data (Conejero et al., 2024; O’Pella, 2019; Gonçalves et al., 2016; Simons et al., 2018; Velichko et al., 2023, 2022; Velichko and Heidari, 2021).

1.1 Characterization of 2009 and 2021 Major SSW Events

The 2009 SSW event occurred from January to March. Its rising stratospheric temperature and the corresponding stratospheric zonal mean wind was categorized into six phases during these periods: SSW Pre-Condition Phase, SSW Ascending Phases, SSW Peak Phase, SSW Descending Phase,



166 After SSW Phase, and No SSW phase (Figure (1)). The red color represents the stratospheric mean
 167 air temperature (k) at 10hPa. While the black color is the stratospheric zonal mean wind at 10hPa.



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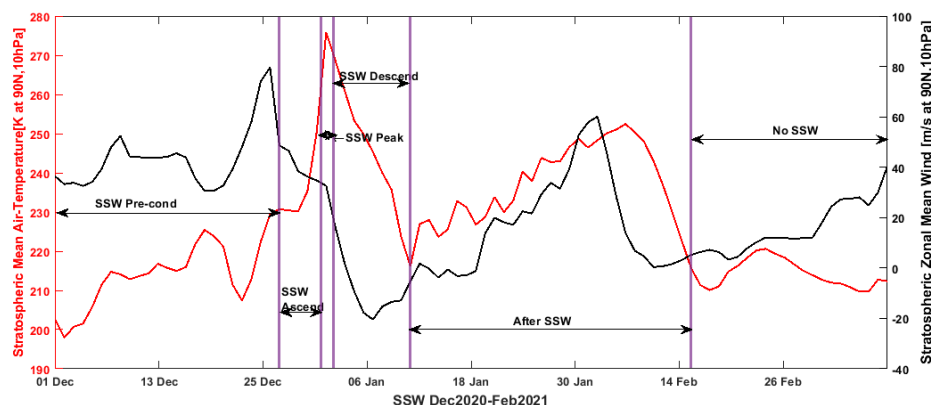
169 Figure 1: The stratospheric zonal mean air temperature and zonal mean wind occurring January-
 170 March, 2009 showing the SSW Precondition Phase, SSW Ascending Phase, SSW peak Phase,
 171 SSW Descending Phase, After SSW Phase and no SSW phase.

172 The SSW precondition phase represents the start of the increase in the stratospheric temperature
 173 (1-16 January) shown in Figure 1, the SSW ascending phase signifies when the stratospheric
 174 temperature increases (17-21 January), the SSW peak phase is when the stratospheric temperature
 175 reaches its maximum (22-24 January), the SSW Descending phase indicates the begin of the
 176 stratospheric temperature decline (25 January-12 February), the After SSW phase represents when
 177 the stratospheric temperature begins to recover to its normal state (13 February-2 March), and the
 178 No SSW phase represents when the stratospheric temperature finally recovers to its normal state
 179 (3-31 March). The SSW event was categorized in accordance to the work of (Bolaji et al., 2016b).

180 The stratospheric temperature and its corresponding zonal mean wind for the 2021 SSW event is
 181 shown in Figure 2. The SSW occurred from December 2020 to February 2021 and was also



182 categorized into six phases. The SSW precondition phase spans between 1 and 27 December, 2020,
 183 the SSW ascending phase is from 28 December 2020 to 2 January 2021, while the SSW peak phase
 184 ranges from 3-5 January 2021, the SSW descending phase spans within 6-14 January 2021, and
 185 the After-SSW phase is from 15 January-16 February 2021. Finally, the No SSW phase emerges
 186 from 17-28 February 2021.



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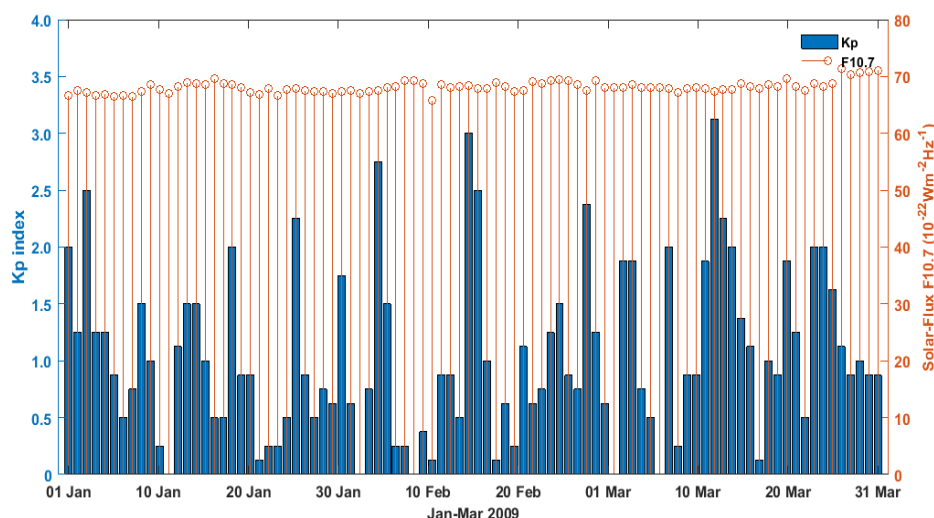
188 Figure 2: The stratospheric zonal mean air temperature and zonal mean wind from December 2020
 189 until February, 2021 revealing the SSW Precondition Phase, SSW Ascending Phase, SSW peak
 190 Phase, SSW Descending Phase, After SSW Phase and no SSW phase.

191 1.2 Global Scale of geomagnetic activities during the 2009 and 2021 SSW Events

192 The year 2009 was the beginning of a solar minimum of the solar cycle 24 while the year 2021
 193 was a solar minimum at the end of the solar cycle 24. Generally, solar minimum years are periods,
 194 where the geomagnetic disturbances mostly record quiet periods. This indicates that the years 2009



195 and 2021 of solar cycle 24 are periods where the geomagnetic disturbances were minimal.

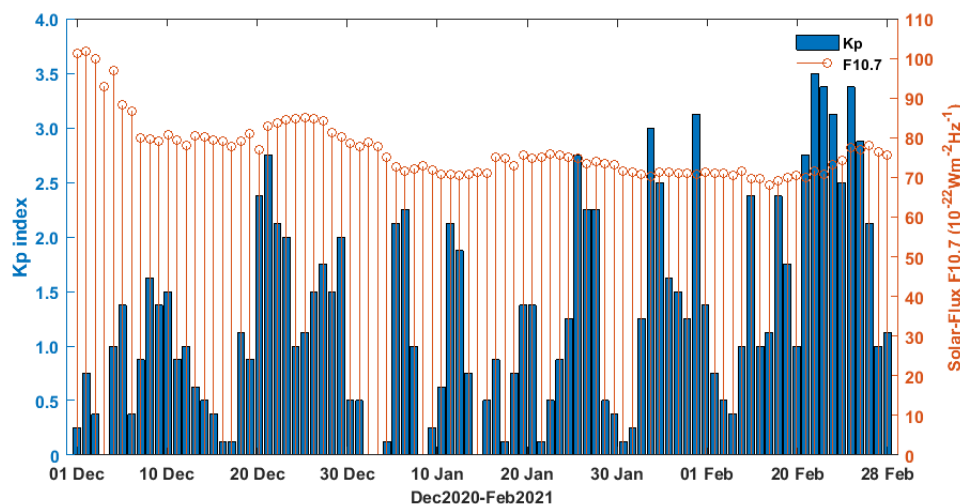


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197 Figure 3: The planetary index K_p (blue bars) and Solar flux $F_{10.7}$ (red lines) from January to March,
 198 2009.

199 The planetary index (K_p) and the solar flux activity during the months of 2009 and 2021 SSW
 200 event are shown in Figures (3 & 4) respectively. During the 2009 SSW event (January-March), the
 201 planetary index (K_p) recorded values of $K_p \leq 2^+$, while the solar flux index ($F_{10.7}$) was around
 202 $F_{10.7} \sim 72$ (Figure (3)). Notably, the planetary index and solar flux activity recorded low values
 203 indicating that the geomagnetic activities were quiet.

204



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206 Figure 4: The planetary index: K_p (blue bar) and Solar flux $F_{10.7}$ (red lines) from December 2020-
 207 February, 2021.

208 Also, the planetary index during the occurrence of the 2021 SSW (December 2020-February 2021)
 209 was within $K_p \leq 3^+$ and the solar flux activity was within $F_{10.7} \sim 100$ (Figure 4). This further
 210 indicates that the geomagnetic activities during the periods of 2009 and 2021 SSW occurrences
 211 were quiet. This observed characteristic testifies the uniqueness of the selected SSW events in
 212 2009 and 2021. Examining the ionospheric chaos in solar quiet currents during periods of minimal
 213 solar and geomagnetic activity will unveil the contributing influence of SSW on ionospheric
 214 dynamics across the Europe and Africa sector.

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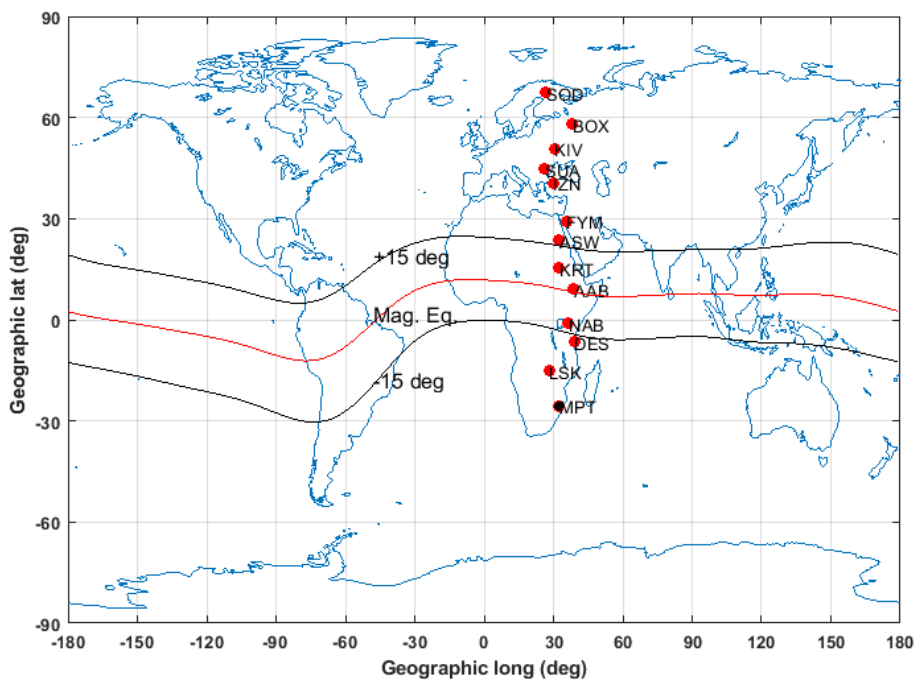
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219 **2. Data Acquisition and Method of Analysis**

220 Ground-based magnetometer data acquired from the Magnetic Data Acquisition System
221 (MAGDAS) at the International Research Centre for Space and Planetary Environment Science
222 (*i – SPES*), Fukuoka, Japan (<http://magdas2.serc.kyushu-u.ac.jp/>) was used in this study. The
223 geographical location of the studied areas is shown in Figure 5.



224 Figure 5: The geographical location of the magnetometer observatories stations investigated across
225 the Europe and Africa sector.
226 The acquired magnetic data from MAGDAS covers 8 magnetometer stations situated in the Africa
227 sector distributed within the geographical latitudes of both Northern and Southern Hemisphere
228 shown in Table 1.
229



230 **TABLE 1: STATIONS INVESTGATED ACROSS EUROPE AND AFRICA SECTOR**

S/ No	Stations	Country	Geographical latitude	Geographical Longitude	Geomagnetic Latitude	Geomagnetic Longitude	Local Time (LT)	Magnetometer Network
1	Sodankyla (SOD)	Finland	67.37	26.63	63.70	107.68	UTC+3	INTERMAGNET
2	Borok (BOX)	Russia	58.07	38.23	53.92	113.32	UTC+3	INTERMAGNET
3	Kiev Dymmer (KIV)	Ukraine	50.70	30.30	46.32	104.39	UTC+3	INTERMAGNET
4	Surlari (SUA)	Romania	44.68	26.25	39.52	99.52	UTC+3	INTERMAGNET
5	Iznik (IZN)	Turkey	40.50	29.72	34.74	102.13	UTC+3	INTERMAGNET
6	Fayum (FYM)	Egypt	29.18	35.50	21.78	106.00	UTC+2	MAGDAS
7	Aswan (ASW)	Egypt	23.50	32.51	14.56	103.89	UTC+2	MAGDAS
8	Khartoum (KRT)	Sudan	15.33	32.32	5.69	103.80	UTC+2	MAGDAS
9	Addis Ababa (AAB)	Ethiopia	9.01	38.74	0.14	110.46	UTC+3	MAGDAS
10	Nairobi (NAB)	Kenya	-1.10	36.48	-10.58	108.18	UTC+3	MAGDAS
11	Dar es Salaam (DES)	Tanzania	-6.80	39.28	-16.26	110.72	UTC+3	MAGDAS
12	Lusaka (LSK)	Zambia	-15.23	28.19	-26.06	98.31	UTC+2	MAGDAS
13	Maputo (MPT)	Mozambi que	-25.50	32.36	-35.92	99.56	UTC+2	MAGDAS

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232 Also, magnetometer observations across Europe were acquired from the International Real-time
 233 Magnetic Observatory Network (INTERMAGNET) (available online at www.intermagnet.org).
 234 The collection consists of 5 magnetometer observatory stations across Europe, see Table 1. The
 235 magnetometer data were acquired during the periods of the 2009 and 2021 SSW occurrences.

236 The planetary (K_p) index during 2009 and 2021 SSW periods was downloaded from GFZ Indices
 237 of Global Geomagnetic Activity (<https://www.gfz-potsdam.de/Kp-index/>) while the Solar flux
 238 activity ($F_{10.7}$) at 2009 and 2021 SSW periods were collected online from the archive of the Space
 239 Physics Data Facility, NASA (<https://omniweb.gsfc.nasa.gov/form/dx1.html>).

240 The daily mean values of zonal mean air temperature and wind during the periods of 2009 and
 241 2021 SSW were acquired from the National Oceanic and Atmospheric Administration (NOAA)



(<https://psl.noaa.gov/data/getpage/>). The ionospheric solar quiet currents, Sq(H), was derived using the H-component of the magnetic field data with daily resolution (with time unit in minutes).

2.1 Ionospheric Solar Quiet Current Sq(H) as Observational Time Series

The dynamics of the Sq(H) current system changes in its daily variation in sequence with time. As a result of its changes in dynamical behavior, the Sq(H) current can be regarded as an observational time series data. To derive the daily dynamics of the Sq(H) current, magnetic field data from various magnetometer stations across Europe and Africa, within the geographical longitude range of 26° to 40°, were archived at one-minute intervals. The H-component (magnetic northward) was collected in minutes from all the stations under investigation and the values of the geomagnetic storm index in minutes (SYM-H) were subtracted from the H-component to minimize the disturbance field arising from magnetospheric currents (Yamazaki et al., 2012b; Bolaji et al., 2016b; Yamazaki et al., 2012a). The derivation of the Sq(H) current is mathematically expressed as:

$$\Delta H_{ground} = H_{local} - (SYM H \cos(L)) \quad (1)$$

where $SYM H \cos(L)$ is the calculated disturbance field arising from the magnetospheric current effect and L is the geomagnetic latitude of the stations under investigation.

To estimate the solar quiet current, the average value between 24:00 and 1:00 local time (LT) in minutes for a particular day was calculated.

$$BLV = \frac{\Delta H_{01} + \Delta H_{24}}{2} \quad (2)$$

The notation ΔH_{01} and ΔH_{24} are the 60 minutes values of H component at 1:00 and 24:00 LT while BLV is the Baseline.



263 The residual value after subtracting the baseline value from the H-component gives the solar quiet
264 current time series considered in minutes.

$$265 \quad S_q(H)_t = H_t - BLV \quad (3)$$

266 $S_q(H)_t$ is the solar quiet current considered in minutes. The analysis of the Sq(H) was deduced for
267 day-to-day activities of the 2009 SSW (January-March) and 2021 SSW (December 2020-February)
268 periods for all stations under investigation.

269 **2.2 Detrending of Solar quiet Sq(H) Current Time Series by Horizontal Visibility Graph** 270 **(HVG).**

271 HVG is a simplified visibility graph that transforms a time series into a graph structure maintaining
272 the inherent characteristics of the transformed time series (Conejero et al., 2024; Gonçalves et al.,
273 2016; O’Pella, 2019; Luque et al., 2009; Zou et al., 2019). It considers the time series in a two-
274 dimensional plane and determines which data are mutually visible. Each value in the time series
275 data is represented by a vertical bar. Two data points share visibility if a line can be drawn from
276 one bar to the other without intersecting any other vertical bars between them. Transforming the
277 signal via HVG, we obtain a graph where nodes represent points in the signals, and the connection
278 between them are based on horizontal visibility. This transformation allows for a shift from a time
279 series to a graph structure while preserving topological information about the signal’s structure,
280 making it useful for analyzing the dynamic properties of time series data.

281 Let us consider a time series of N -data points be represented as:

$$282 \quad [x_i, i = 1, 2, \dots, N] \quad (4)$$



283 Two nodes i and j in the graph are connected, if it is possible to trace a horizontal line in the series
 284 linking x_i and x_j not intersecting intermediate data height, fulfilling:

$$285 \quad x_i, x_j > x_n \text{ for all } i < n < j \quad (5)$$

286 In this study, the time series of the solar quiet current. Sq(H) derived during the periods of the 2009
 287 and 2021 SSW was subjected to the HVG method which transform the series into a complex
 288 network. The calculations were performed using the ts2vg Python module, specifically utilizing
 289 the HorizontalVG class, which represents one of the types of visibility graphs, namely the
 290 'Horizontal Visibility Graph'. The input to this process is a time series, which is transformed into
 291 a network where each point in the series becomes a node, and edges are formed based on the
 292 visibility criteria between points. By applying this method, we calculate the degree (number of
 293 connections) for each point in the time series, capturing the number of other points it can 'see' in
 294 the horizontal visibility graph. The output is a list of degrees for each point, reflecting the local
 295 connectivity structure within the time series, which helps reveal patterns, dependencies, and
 296 variability within the data.

297 **2.3 Fuzzy Entropy (FuzzyEn)**

298 Fuzzy entropy (FuzzyEn) is a powerful and popular nonlinear tool used to assess the dynamical
 299 characteristics of time series data (Ishikawa and Mieno, 1979; Li et al., 2017; Azami et al., 2019;
 300 Chen et al., 2007). It provides a quantitative measure of a signal's complexity and chaos. High
 301 entropy indicates a more chaotic structure, whereas low entropy suggests a more regular or
 302 periodic nature. FuzzyEn was developed to overcome the shortcoming of approximate entropy
 303 (ApEn) and sample entropy (SampEn) (Azami et al., 2019; Li et al., 2017; Dass et al., 2019).



FuzzyEn uses exponential functions with Fuzzy boundaries. It is expressed mathematically as follows. Given a time series X_i , Eq. (4), we embed it using a given embedding dimension (m). Then, a new m -dimensional vector (X_m) is formed as

$$X_m(i) = [X_i, X_{i+1}, \dots, X_N] - x0_i \quad (6)$$

These vectors represent m consecutive x -values, starting with the i -th points, and with the baseline $x0_i = \frac{1}{m} \sum_{j=0}^{m-1} x_{i+j}$ removed. Then, the distance between vectors $X_m(i)$ and $X_m(j)$, $d_{i,j,m}$ can be defined as the maximum absolute difference between their scalar components. Given n and r , the degree of similarity ($D_{i,j,m}$) of the vectors $X_m(i)$ and $X_m(j)$ is calculated using the fuzzy function.

$$D_{i,j,m} = \mu(d_{i,j,m}, r) = \exp\left(\frac{-(d_{i,j,m})^n}{r}\right) \quad (7)$$

Where n and r are the FuzzyEn power and threshold respectively. The function ϕ_m is defined as

$$\phi_m(n, r) = \frac{1}{N-m} \sum_{i=1}^{N-m} \left(\frac{1}{N-m-1} \sum_{\substack{j=1 \\ j \neq i}}^{N-m} D_{i,j,m} \right) \quad (8)$$

Repeating the same procedure from equation (7) and equation (8) for the vector $X_{m+1}(j)$, i.e., for dimension $(m + 1)$, the function ϕ_{m+1} is obtained. Therefore, FuzzyEn can be estimated as

$$\text{FuzzyEn}(m, n, r, N) = \ln \phi_m(n, r) - \ln \phi_{m+1}(n, r) \quad (9)$$

In this study, we applied the FuzzyEn to the solar quiet current, Sq(H), to investigate ionospheric chaos during SSW across the Europe and Africa sector. The computational parameters used for FuzzyEn analysis included an embedding dimension ($m = 1$) and a tolerance threshold defined as $r_1 = 0.2 \times \text{std}$, where std represents the standard deviation of the time series X . Additionally, the



323 argument exponent (pre-division) $r_2=3$ was applied, along with a time delay of $\tau=1$. The window
324 size used for the analysis was $s=200$.

325 The calculations were performed using Python with the EntropyHub library [EntropyHub. An
326 Open-Source Toolkit for Entropic Time Series Analysis. Available online:
327 <https://www.entropyhub.xyz/> (accessed on 27 February 2024).] (version 0.2), which provides a
328 reliable and standardized method for calculating FuzzyEn, ensuring that the results can be
329 compared across different studies. EntropyHub integrates many established entropy methods into
330 a single package, available for Python, MatLab, and Julia users. By utilizing this library, we
331 ensured the consistency and reproducibility of our entropy calculations. The FuzzyEn was
332 computed using the solar quiet current time series for the SSW periods of 2009 (January-March)
333 and 2021 (December 2020-February 2021) after applying the Horizontal Visibility Graph (HVG)
334 transformation.

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344 3. Results

345 A sample of solar-quiet current $Sq(H)$ on 31st of March 2009 at Addis Ababa, Ethiopia is shown
 346 in Figure 6 (a-d). The panel (a) represent the time series of the solar quiet current derived in
 347 minutes, while panel (b) is the detrended time series of solar quiet current transformed through
 348 Horizontal Visibility Graph (HVG). The panel (c) is the FuzzyEn depicting the changes in entropy
 349 of the solar quiet current without HVG transformation. The result in panel (d) depicts the changes
 350 in entropy of solar quiet current transformed through HVG. The solar quiet current gradually
 351 enhances in magnitude during the peak noon periods in its daily variation, while at pre-noon and
 352 post-noon periods, a gradual increment and a decrease in the magnitude of the solar quiet current
 353 was observed. The results of Fuzzy Entropy values after HVG transformation reveals a gradual
 354 decrease in entropy at noon-periods. Notably, during pre-noon and post-noon periods, the result
 355 revealed an increment in entropy changes. These distinct features of entropy changes obtained in
 356 Fuzzy Entropy after HVG transformation of the solar quiet time series was not obvious in the
 357 results of Fuzzy Entropy obtained without HVG transformation method. The HVG transformation
 358 series unveils the transient changes in entropy for ionospheric current system during SSW. This
 359 indicates that the HVG transformation method captures the dynamical characteristics of the
 360 ionospheric current enabling, the unveiling of chaotic behavior obtained through changes in Fuzzy
 361 Entropy during the periods of 2009 (January-March) and (December 2020-February 2021) SSW
 362 events.

363 The day-to-day latitudinal distribution of entropy across Europe and Africa sector in January 2009
 364 is display in Figure 7. The contour map depicts the entropy changes in color representation. The
 365 yellow color unveils ranges of Fuzzy Entropy values between 1.2 and 1.4 indicating high entropy.
 366 This high entropy indicates the presence of ionospheric chaos (disorderliness) in the ionospheric



current system. The light blue color ranging at approximately 1.2~0.8 reveals a declining Fuzzy Entropy value which indicates a declining to orderliness behavior in the ionospheric current system. While the deem blue color ranging at 0.8~0.6 depicts low entropy values. The low entropy indicates the presence of orderliness behavior in the ionospheric dynamics. We noticed that in the day-to-day latitudinal distribution of entropy across the stations in the Europe and Africa sector. Some days in January 2009 depicts the presence of ionospheric chaos at most of the stations shown in Figure 5. Such that high entropy values were recorded on 1st, 4th, 8th, 10th, 14th, 18th-25th and 28th-31st of January at stations IZN, SUA, KIV, BOX and SOD (see Tab. 1). On 25th of January, a higher entropy values, unveiling chaotic behavior in the ionospheric dynamics, was observed in all the investigated stations in the European and African sector. Notably, low entropy values, signifying orderliness (periodic) behavior in the dynamics of ionosphere, was observed on 3rd, 5th-7th, 9th, 11th-16th, and 26th of January. The lowest entropy values were found on 3rd of January, indicating that the state of the ionosphere exhibits a highly orderliness behavior in its underlying dynamics across the Europe-Africa sector during this period.

We display the contour plots of the day-to-day latitudinal distribution of entropy on February 2009 in Figure 8. The changes in entropy unveils high values of Fuzzy Entropy indicating the presence of ionospheric chaos in most of the station across the Europe and Africa sector. This trend of high entropy values was noticed on 1st-2nd, 9th-10th, 15th, 17th-18th of February. This high entropy observed in the aforementioned days signifies that the dynamics of the ionosphere is chaotic. In addition, 1st and 17th February exhibited even more chaotic behavior across the stations. This observation further strengthening the evident influence of SSW on the ionospheric dynamics on these dates. Notably, low entropy implying the presence of orderliness behavior in the dynamics of the ionosphere was obvious on 5th-8th, 11th-12th, 14th, 16th, 19th-22nd, and 25th-27th of February



390 2009. However, 5th and 14th of February exhibited the lowest entropy values, revealing that the
391 state of the ionosphere is at a minimal perturbation state.

392 The day-to-day latitudinal distribution of entropy across the European and African sector on March
393 2009 is shown in Figure 9. Most of the Fuzzy entropy values are associated with low entropy. For
394 instance, 1st-6th, 8th-18th, 20th-23rd, 26th-28th and 31st of March depicts low changes in entropy
395 in most of the stations. High entropy values associated with the presence of ionospheric chaos was
396 found on 7th, 24th, and 29th-30th of March.

397 The Fuzzy Entropy of the day-to-day latitudinal distribution across European sector on December
398 2020 is presented in Figure 10. We observed that the changes in entropy from 1st to 6th of December
399 at the SOD, BOK and KIV stations reveals a consistently low entropy values, indicating that the
400 state of the ionosphere at these stations exhibits a consistent orderliness behavior during these
401 periods. However, at SUA and IZN, consistent high entropy values indicating the presence of
402 ionospheric chaos from 1st to 5th of December was observed. Notably, on 6th of December, all the
403 stations in the European sector (SOD, BOK, KIV, SUA, and IZN) depicts low entropy signifying
404 that the dynamics of ionosphere is less perturbed by the influence of SSW on December 2020.
405 However, a high entropy values indicating chaotic behavior in the ionospheric dynamics was
406 observed at SOD, BOK, and KIV stations on 7th and 8th of December. Interestingly, SUA and IZN
407 were consistently associated with a low entropy distribution from 8th to 22nd of December. This
408 observed consistent low entropy at SUA and IZN signifies that the ionosphere in this region is less
409 perturbed by the influence of SSW. Furthermore, the possible emergence of ionospheric
410 disturbances at SUA and IZN during those observed days is extremely minimal. A consistent
411 decline from disorderliness (chaotic) to orderliness (periodic) was observed at SOD, BOK, and



412 KIV from 25th to 31st of December, while SUA and IZN exhibited some features of ionospheric
 413 chaos on 30th and 31st of December.

414 For SSW January 2021, the day-to-day latitudinal distribution of entropy across Europe sector is
 415 displayed in Figure 11. The latitudinal distribution depicts low entropy across the day-to-day
 416 observation in January. For instance, 1st-6th, 8th-11th, 13th-16th, 19th-24th, and 26th-28th of
 417 January unveils low entropy values across the stations investigated in the Europe sector implying
 418 that the ionosphere at this region exhibits orderliness (periodic) behavior in its underlying
 419 dynamics. Notably, high entropy, indicating the presence of ionospheric chaos (disorderliness),
 420 was observed at SUA and IZN on 7th, 10th, and 18th of January.

421 In Figure 12 displaying the day-to-day latitudinal distribution of entropy in February 2021 across
 422 Europe sector. The changes in entropy reveals high entropy, suggesting the presence of ionospheric
 423 chaos on 9th, 12th, 16th-18th, 21st, and from 24th to 27th of February at SOD, BOX, SUA and IZN.
 424 In contrast, KIV consistently exhibited low values of entropy from 1st to 8th of February.

425 **3.1 Ionospheric Chaos During Phases of 2009 SSW**

426 Display in Figure 13 is the latitudinal distribution of entropy across Europe and Africa sector
 427 during the phases of 2009 SSW. The phases of SSW are categorized into six namely: precondition
 428 phase, ascending phase, peak phase, descending phase, after SSW phase and no SSW phase. The
 429 entropy values during the precondition phase of SSW depicts a high entropy across the stations in
 430 the Europe sector signifying that the precondition phase of 2009 SSW were associated with the
 431 emergence of ionospheric chaos. The observed features of high entropy values unveiling
 432 pronounced ionospheric chaos at SUA, KIV, BOX and SOD were obvious. This observation
 433 further reveals the emergence of ionospheric disturbances in SUA, KIV, BOX and SOD due to the



434 influence of SSW. Interestingly, the periods from January to March 2009, characterized by SSW,
435 were strongly associated with geomagnetic quiet periods (see, Figure 2a). In the African sector,
436 SSW preconditioning phase is characterized by a declining in entropy values, indicating that the
437 state of the ionosphere across Africa sector is transiting to orderliness behavior in its underlying
438 dynamics at the preconditioning phase of 2009 SSW. A low entropy was seen at AAB, NAB and
439 DES during the preconditioning phase signifying that the underlying dynamics of the ionosphere
440 at AAB, NAB, and DES are exhibiting orderliness as the preconditioning phase of SSW emerges.
441 During the ascending phase of SSW, we observed also a low entropy values at stations in Africa
442 up to IZN. However, high values of entropy implying the presence of ionospheric chaos
443 (disorderliness) were found at KIV, BOX, and SOD. The high entropy observed at these stations
444 during the ascending phase indicates that the influence of SSW on the state of the ionosphere at
445 KIV, BOX, and SOD is present. This suggests that SSW events significantly affect ionospheric
446 dynamics, leading to increased variability and changes in electron density patterns across Europe
447 sector during the ascending phase of 2009 SSW. In the African sector, during the ascending phase
448 of SSW, we observed that, with the exception of the ASW and KRT stations, high entropy values
449 were recorded around 16-24hr (LT). All other stations investigated in Africa depict low entropy
450 values. This further signifies that the ionospheric dynamics in most of the stations investigated in
451 the Africa sector during the ascending phase exhibits orderliness (periodic) behavior in its
452 underlying dynamics.

453 During the peak phase of SSW 2009, evidence of ionospheric chaos was observed at SOD, BOX
454 and SUA. The changes in entropy during this period depicts high values, indicating significant
455 disturbances in the ionosphere. A declining and steady low entropy change was observed from
456 SUA to DES. In contrast, KRT, AAB and NAB exhibited very low entropy values around 4-10



457 hours (LT) during the peak phase of SSW. This observed low entropy changes at KRT, AAB and
458 NAB signifies that the contributing influence of SSW to ionospheric dynamics at KRT, AAB and
459 NAB is not pronounced at the peak phase of 2009 SSW compared to the stations situated in the
460 European sector. At the descending phase of SSW, high entropy values associated with the
461 presence of ionospheric chaos was evident at SOD, BOX, KIV, SUA, IZN, FYM and ASW.
462 However, AAB and NAB depicts low entropy values during the descending phase unveiling that
463 ionospheric dynamics at AAB and NAB are not influenced by SSW during this period.

464 The after-SSW phase of 2009 also exhibits a distribution of high entropy values across the
465 investigated stations from the Europe to the Africa sector. Notably, at 8-16hr (LT), a declining
466 entropy was noticed at KRT while low entropy values were seen at AAB, NAB, DES, LSK and
467 MPT. At 20-24hr (LT), ionospheric chaos associating with high entropy was evident at LSK, DES,
468 AAB, KRT, ASW and FYM. During the no-SSW phase, the changes in entropy were observed to
469 decline at SOD, BOX, KIV and IZN. In contrast, at SUA high entropy values indicative for
470 ionospheric chaos were noted. A consistent low entropy change was found in ASW, KRT, AAB,
471 NAB, DES, LSK and MPT.

472 **3.2 Ionospheric chaos during phases of 2021 SSW**

473 Shown in Figure 14 is the 2021 SSW latitudinal distribution of entropy across Europe sector at
474 precondition, ascending, peak, descending, After-SSW and no-SSW phases. The Precondition
475 phase at SOD, BOX, and KIV depicts high values of entropy signifying that ionospheric chaos is
476 dominant during this period, while a declining and low entropy change was observed at IZN and
477 SUA. The ascending, peak, descending and after SSW phases exhibit a consistent low entropy
478 value in their latitudinal distribution, suggesting orderliness behavior in the ionosphere across
479 SOD, BOX and KIV. In contrast, at SUA a consistent increase in entropy was evident. This



480 consistent increase in entropy at SUA during all the phases of SSW implies that the ionospheric
481 dynamics is consistently exhibiting chaotic behavior due to the influence of SSW.

482 **4. Discussion of Results**

483 The chaotic behavior of the ionosphere is prevailing due to the dynamical characteristics of the
484 atmospheric layers. Evidence of chaos in the ionosphere, magnetosphere due to space weather
485 effects have been investigated using various chaotic quantifiers such as entropy, Lyapunov
486 exponent, correlation dimension and recurrence plot (Unnikrishnan, 2008; Pavlos et al., 1992;
487 Pavlos, 2012; Donner et al., 2019; Ogunsua et al., 2014; Balasis et al., 2023). However, the
488 ionospheric chaos with regards to regional information has not been ascertained, especially during
489 extreme meteorological event such as SSW. To ascertain the degree of SSW influence on the
490 dynamics of the regional ionosphere within 26° to 40° Eastern geographical longitude. This present
491 study unveils the ionospheric chaos response to solar quiet current, Sq(H) dynamics across Europe
492 and Africa sector during 2009 and 2021 SSW event. The ionospheric dynamics is a complex
493 system, driven by a set of atmospheric elements that continually interacts with the geospace
494 environment. They are sensitive to any small change in the atmospheric drivers that can amplify
495 and lead the ionospheric plasma to instability (disorderliness) or stability (orderliness) from its
496 initial state. This sensitivity defines the potential for ionospheric chaos (Marwan et al., 2007;
497 Conejero et al., 2024; Velichko and Heidari, 2021; Boriskov et al., 2022). The chaotic behavior is
498 strongly associated with high entropy while the orderliness behavior is associated with declining
499 and low entropy. These features are characteristic of a dynamical system, and the ionosphere serves
500 as a prime example of such a system (Materassi, 2020; Radicella and Nava, 2020; Hokkanen, 2000;
501 Ogunsua et al., 2014).



502 Previous studies by (Goncharenko et al., 2010; Fejer et al., 2011; Yamazaki et al., 2012c, a;
503 Yamazaki, 2014; Yamazaki et al., 2016; Gómez-Escolar et al., 2014; Baldwin et al., 2021;
504 Goncharenko et al., 2021; Hocke et al., 2024; Fuller-Rowell et al., 2016; Gupta et al., 2021; Ma
505 and Hocke, 2024) and other researchers have established the potential reshaping of ionospheric
506 plasma associated with the emergence of SSW. Our present findings unveils the evidence of
507 ionospheric chaos been more pronounced at SUA, KIV, BOX and SOD compared to the stations
508 situated in the African sector. A consistent dynamics of orderliness behavior in the ionosphere was
509 noticed at AAB, NAB and DES. Furthermore, we noticed that, as the stratospheric temperature
510 rises, a clear orderliness behavior becomes evident in the African sector. For instance, our results
511 unveil that low entropy begins to converge around AAB during the preconditioning phase of SSW.
512 While at the ascending phase, the low entropy spread from AAB across MPT and FYM, the region
513 of both Southern and Northern Hemisphere of Africa sector. Then during the SSW peak, low
514 entropy diminishes and converges at KRT, AAB and NAB. From the SSW descending phase till
515 No-SSW phase, we noticed that the features of low entropy begin to expand and spread across all
516 the stations in the Africa sector. This low entropy indicative for orderliness behavior is found to be
517 more pronounced during the no-SSW phase. Remarkably, the observation of low entropy values
518 across all the stations situated in the Africa sector reveals that the influence of SSW to ionospheric
519 dynamics in the Africa sector is minimal. Furthermore, the evidence of low entropy, an indicative
520 of orderliness implies that the chaotic behavior at the regional ionosphere of Africa sector is
521 extremely low facilitating steady dynamics of ionospheric current system during the SSW.
522 However, ionospheric chaos was evident in most of the station situated in the Europe sector. In
523 that, at the precondition phase of SSW, the ionospheric chaos spans from FYM till SOD while at
524 the ascending phase, the chaotic behavior was only observable at KIV, BOX and SOD. Notably,



525 during the peak phase, the ionospheric chaos enhanced at SUA, KIV, BOX and SOD. This
526 observed enhancement in ionospheric chaos spans further during the descending phases of SSW.
527 At the after SSW phase, the ionospheric chaos extends to some station in the Southern Hemisphere
528 of the Africa sector around 0-6hr (LT) in FYM, ASW, KRT, and AAB. During the no-SSW phase,
529 the stations situated in the Europe sector exhibited a decline of ionospheric chaos at SOD, BOX,
530 KIV and IZN. These findings of ionospheric chaos in most of the stations situated in the Europe
531 sector unveils the evidence that the 2009 SSW indeed significantly influenced the ionospheric
532 dynamics in this region. Interestingly, during the 2009 SSW periods, geomagnetic activities were
533 characterized by a planetary index of $K_p \leq 2^+$ and a solar flux ($F_{10.7}$) of approximately 72 (see
534 Figure 3). This record indicates that geomagnetic disturbances were extremely low during the 2009
535 SSW. Therefore, the finding that ionospheric chaos was dominant in most stations located in the
536 European sector, compared to those in the African sector, provides additional evidence of the
537 SSW's contributing influence on the regional ionosphere in Europe sector. In the 2021 SSW, the
538 entropy distribution at KIV, BOX, and SOD unveils significant enhancement of ionospheric chaos
539 during the preconditioning phase of SSW while orderliness behavior was noticed at SUA and IZN.
540 The orderliness behavior of the ionospheric dynamics was evident during the ascending, peak and
541 descending phase of SSW at SOD, BOX, and KIV. Then, at the after SSW and no-SSW phases,
542 the presence of ionospheric chaos begins to emanate at SOD, BOX and KIV. Evidence of
543 ionospheric chaos was prominent at SUA during the ascending, peak, descending, after, and no-
544 SSW phases. This observed feature of ionospheric chaos at SUA unveils the evidence of the SSW
545 influence on the ionospheric dynamics in some regions of the European sector during the 2021
546 SSW.



547 The observation of prevailing ionospheric chaos due to SSW across European sector suggest that
548 for a relatively good predictability of SSW effects on the ionosphere. The effort of understanding
549 the regional ionospheric chaos characteristics will enhance ionospheric forecasting skill. Because
550 some of our findings of change in entropy associated with regional ionosphere are noticed to be
551 less sensitive to SSW influences propagating from the lower atmosphere (See Figures 13 and 14).
552 This observation also highlights orderliness behavior where the ionospheric disturbances due to
553 SSW influences are relatively suppress. It is worthy to note that emergence of orderliness behavior
554 in the ionospheric dynamics due to SSW does not completely erase chaotic behavior. Rather, it's
555 an evidence of suppression of chaos due to SSW event. We suspect that the regions where
556 orderliness behavior are pronounced due to SSW influences could suggest that the coupling
557 processes between the atmospheric drivers are in strong synchrony leading to synchronization
558 characteristics in ionospheric dynamics. For instance, in the Africa sector, stations within the
559 equatorial region i.e. AAB consistently exhibits orderliness behavior such that low entropy values
560 were consistently observed during the phases of SSW. This observation will be extensively
561 research in our future investigation with primary focus on equatorial region.

562 The evidence of ionospheric chaos could result to the formation of small/large disturbances in the
563 European sector owing to the impact of disruption on communication and navigation signals
564 during SSW. This observation of ionospheric chaos may be attributed to the enhancement of the
565 solar and lunar migrating semidiurnal tides during SSWs, which influences the generation of
566 electric fields through the E-region dynamo mechanism (Goncharenko et al., 2021; Pedatella et
567 al., 2014; Siddiqui et al., 2018)

568

569



570 **5.0 Conclusion**

571 This study aimed to unveil the contributing influence of Sudden Stratospheric Warming (SSW) on
572 the regional ionosphere across Europe and Africa by examining ionospheric chaos in the solar
573 quiet current during the SSW events of 2009 and 2021. These SSW events occurred during the
574 solar minimum years of solar cycle 24. The SSW was categorized according to the rising
575 stratospheric temperature into six phases namely precondition, ascending, peak, descending, after,
576 and no-SSW phases. The study covers 13 magnetometer stations across the Europe and Africa
577 sector located within the geographical longitude of 26° to 40° East. Magnetometer data obtained
578 during the periods of SSW were used to derive time series of the ionospheric solar quiet current,
579 Sq(H). These solar quiet current time series were transformed into a network representation
580 through the Horizontal Visibility Graph (HVG) approach and analysed by Fuzzy Entropy to
581 quantify the presence of ionospheric chaos during the periods of SSW. Low entropy is associated
582 with orderliness behavior. We found that the latitudinal distribution of entropy depicts high entropy
583 indicating the presence of ionospheric chaos in most of the stations situated within the Europe
584 sector compared to stations in the Africa sector. A consistent low entropy distribution unveiling
585 the presence of orderliness behavior was found to be prominent in the Africa sector. This prevailing
586 evidence of orderliness behavior in the Africa sector during SSW signifies that the contribution
587 influence of SSW to the regional ionosphere of Africa sector is minimal. However, the pronounced
588 features of ionospheric chaos found in the Europe sector unveil the evidence of significant effects
589 of SSW on the regional ionosphere in Europe. Finally, we found that after the peak phase of SSW,
590 the ionospheric chaos is more pronounced.

591

592



593 **Code availability**

594 The code is a collection of routines in MATLAB (MathWorks) and is available upon request to the
595 corresponding author.

596 **Data availability**

597 The magnetometer data are publicly available and provided by Magnetic Data Acquisition System
598 (MAGDAS) at the International Research Centre for Space and Planetary Environment Science,
599 Fukuoka, Japan (<http://magdas2.serc.kyushu-u.ac.jp/>). The magnetometer can also be access at
600 the International Real-time Magnetic Observatory Network (INTERMAGNET) (available online
601 at www.intermagnet.org). While the stratospheric temperature can be access from National
602 Oceanic and Atmospheric Administration (NOAA) (<https://psl.noaa.gov/data/getpage/>). The
603 planetary index are provided and access GFZ Indices of Global Geomagnetic Activity
604 (<https://www.gfz-potsdam.de/Kp-index/>). The data of disturbance storm time in minute resolution
605 (SYM-H) is available at the World Data Centre for Geomagnetism, Kyoto, Japan:
606 <https://wdc.kugi.kyoto-u.ac.jp/>, while the solar flux index are archived at the National Aeronautics
607 and Space Administration (NASA), Space Physics Facility: [https://omniweb.gsfc.](https://omniweb.gsfc.nasa.gov/form/dx1.html)
608 [nasa.gov/form/dx1.html](https://omniweb.gsfc.nasa.gov/form/dx1.html)

609 **Author contributions**

610 **IAO** developed the idea behind the problem being solved, supervise the project, analyzed the data;
611 developed the codes; and draft the manuscript; **VA** supervise the project, developed the codes,
612 analyzed the data and contribute to the drafting of the manuscript; **MN** supervise the project and
613 contribute to the drafting of the manuscript; **OOI, BOS, and OBO** contributes to the interpretation
614 and Discussion of results; **NAN and OOT** read and made useful comments to the manuscript.



615 **Competing interests**

616 Some authors are members of the editorial board of nonlinear processes in geophysics journal, and
617 the authors declare they have no conflict of interest.

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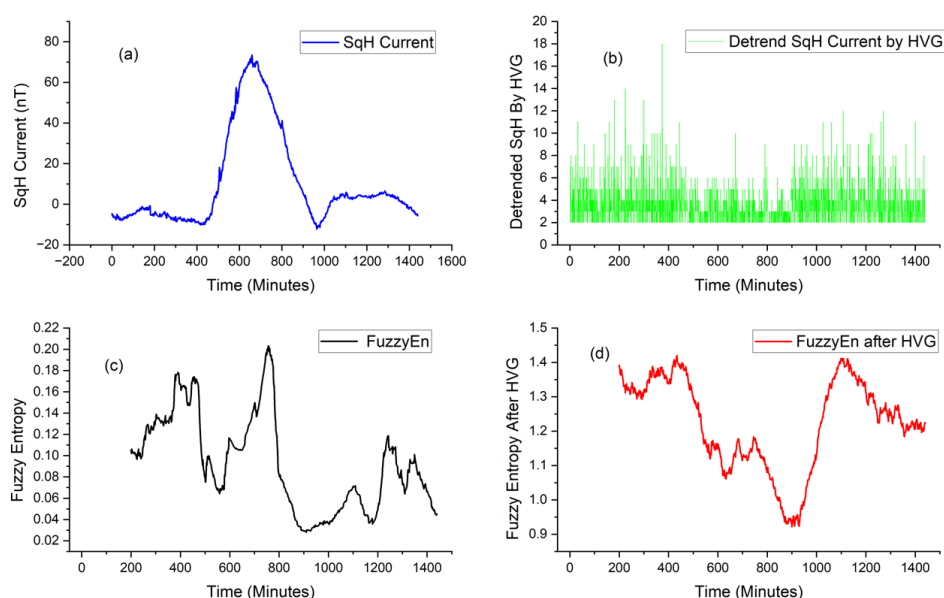
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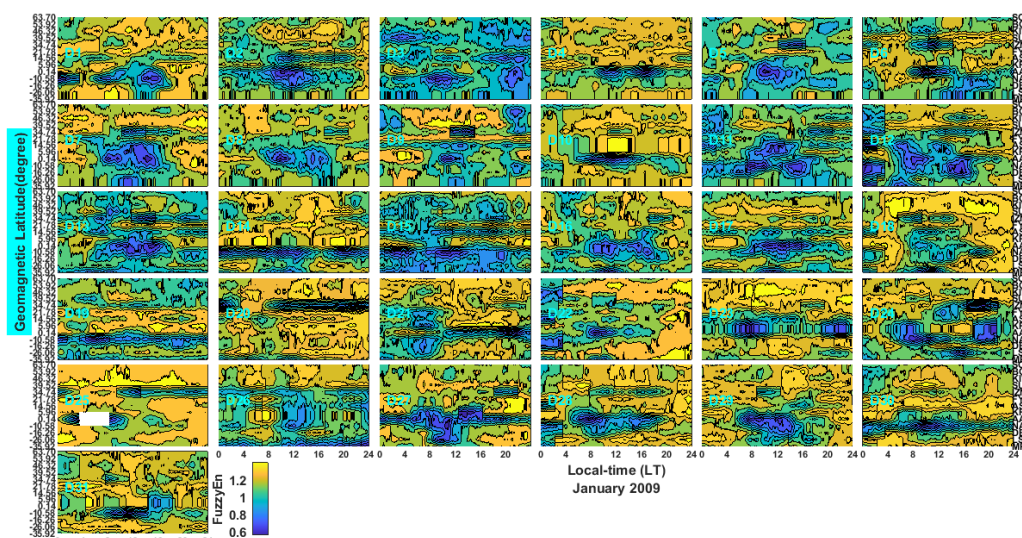
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893 **Figure 6:** The sample of solar quiet current Sq(H) on 31st of March 2009 at Addis Ababa, Ethiopia:

894 (a) The time series of solar quiet current, Sq(H) derived in minutes, (b) The detrended time series
 895 of solar quiet current transformed through Horizontal Visibility Graph (HVG), (c) The changes in
 896 Fuzzy Entropy of solar quiet current without HVG transformation, (d) The changes in Fuzzy
 897 Entropy of solar quiet current with HVG transformation.

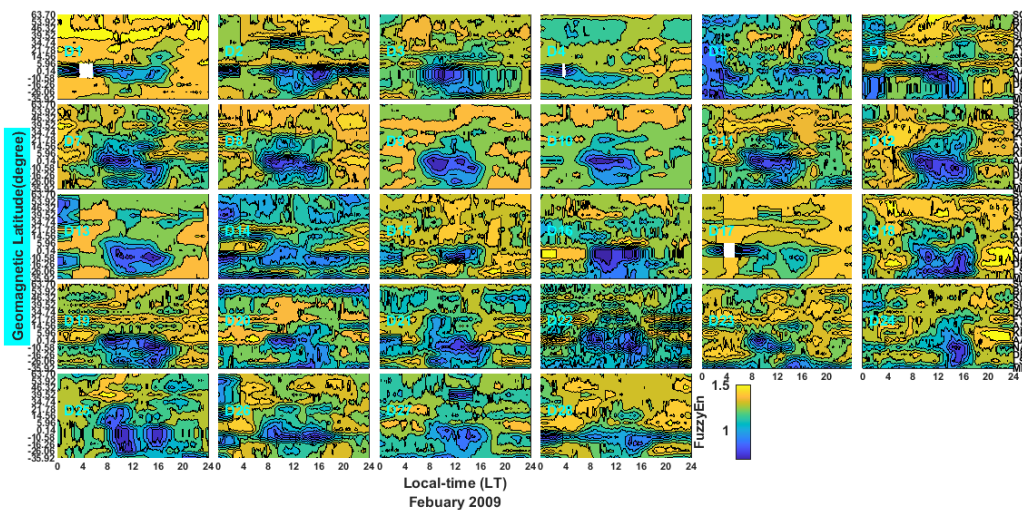
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901 **Figure 7:** The Day-to-Day latitudinal Distribution of Fuzzy Entropy across Europe-Africa sector
 902 on January 2009



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904 **Figure 8:** The Day-to-Day latitudinal Distribution of Fuzzy Entropy across Europe-Africa sector
 905 on February 2009.

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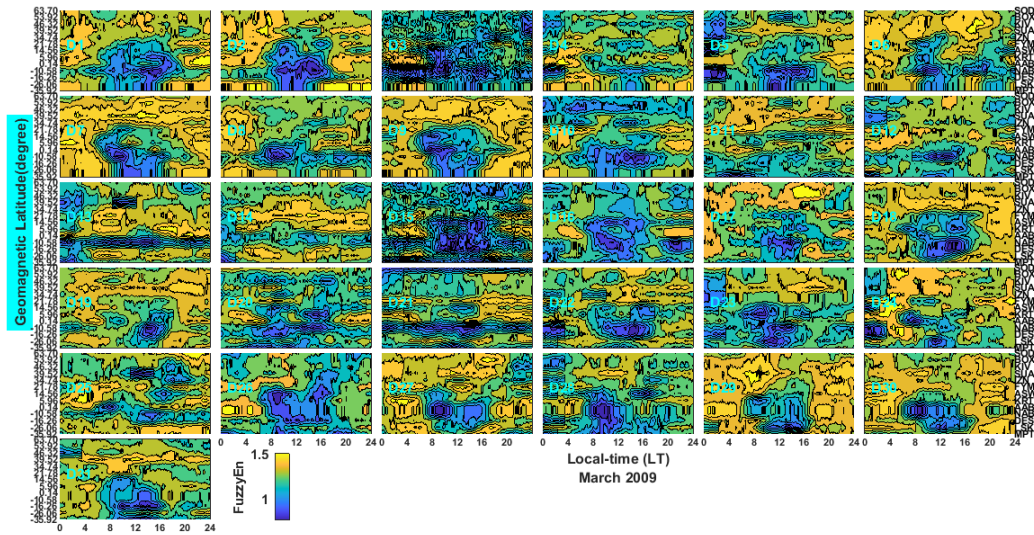


Figure 9: The Day-to-Day latitudinal Distribution of Fuzzy Entropy across Europe-Africa sector on March 2009

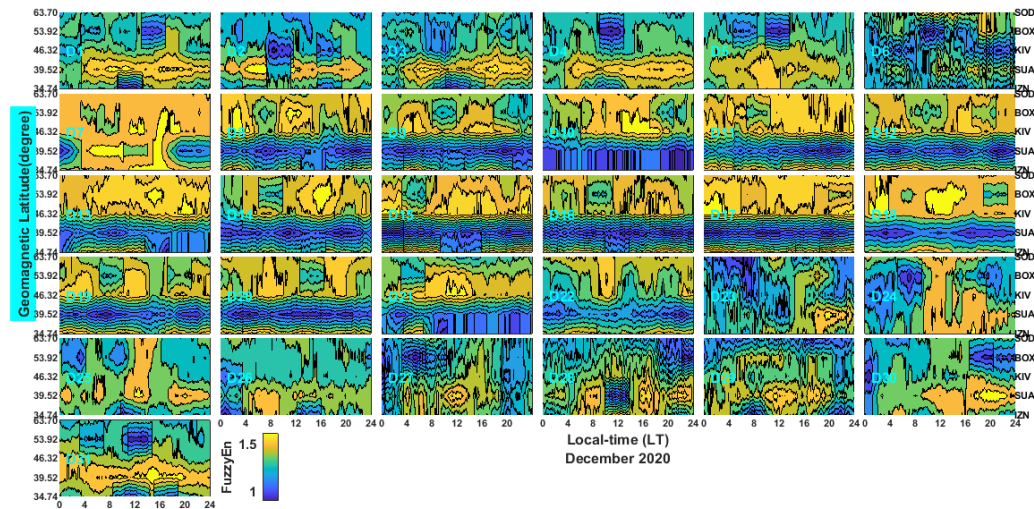
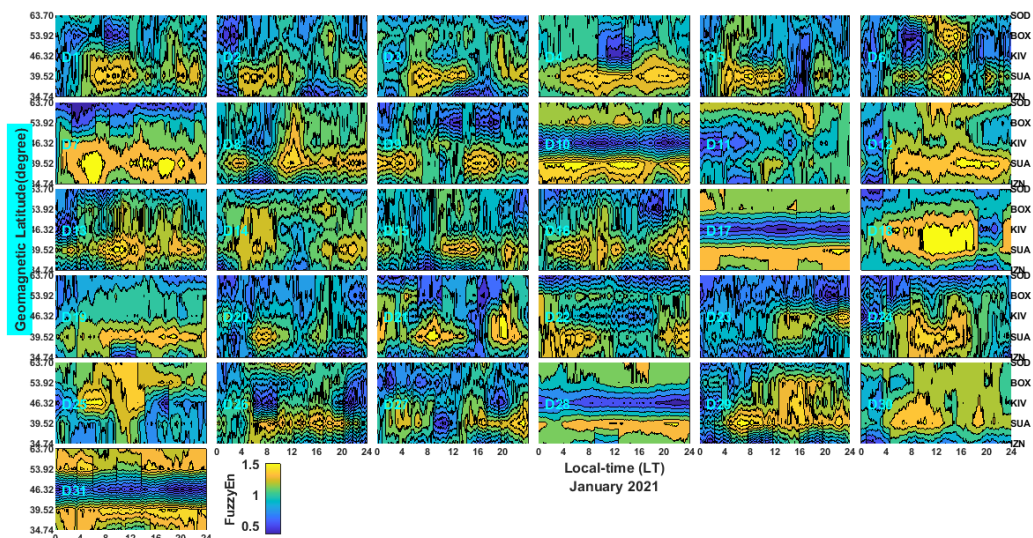
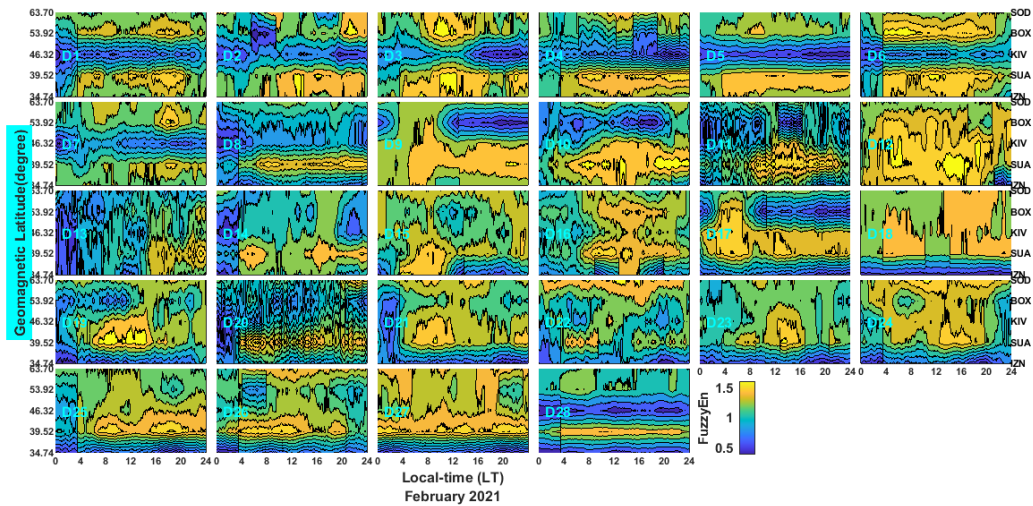


Figure 10: The Day-to-Day latitudinal Distribution of Fuzzy Entropy across Europe sector on December 2020



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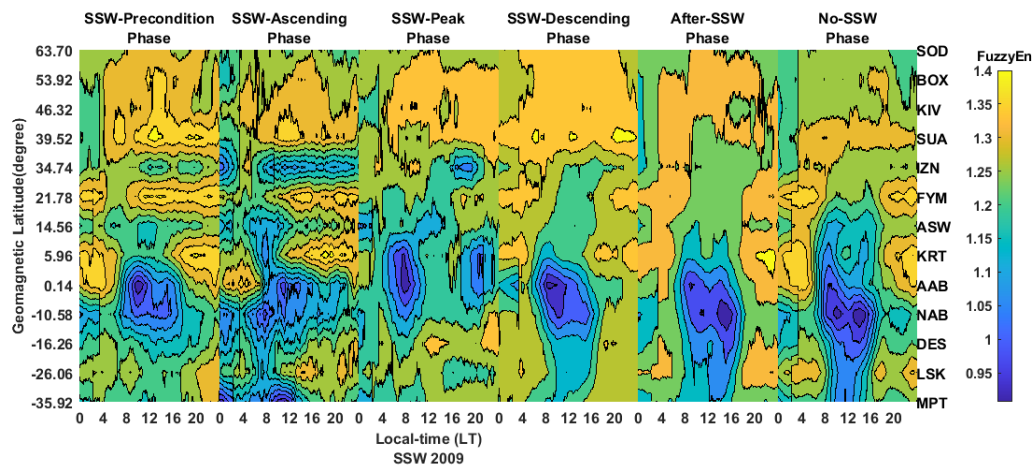
915 **Figure 11:** The Day-to-Day latitudinal Distribution of Fuzzy Entropy across Europe sector on
916 January 2021.



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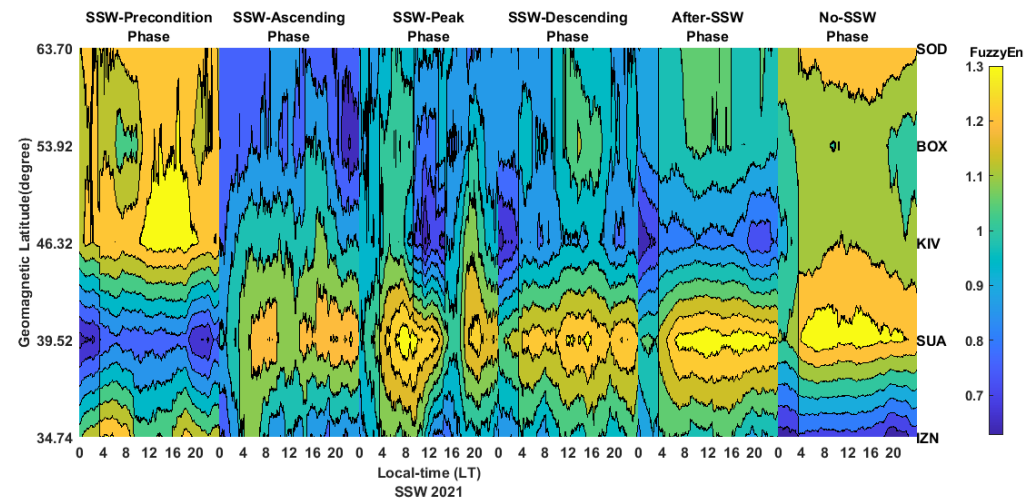
918 **Figure 12:** The Day-to-Day latitudinal Distribution of Fuzzy Entropy across Europe sector on
919 February 2021.

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922 **Figure 13:** The latitudinal distribution of Fuzzy Entropy across Europe-Africa sector during the
923 phases of 2009 SSW.



924

925 **Figure 14:** The latitudinal distribution of Fuzzy Entropy across Europe sector during the phases
926 of 2021 SSW.