~~Pelagic~~ ~~c~~Coccolithophore abundance and production and their impacts on particulate inorganic carbon cycling in the western North Pacific

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Abstract. Coccolithophores are globally~~, a type of single-celled phytoplankton that is~~ abundant single-celled and shelled phytoplankton that play an important role in the marine carbon cycle due to their contribution to the carbonate pump~~in global oceans, are closely associated with the carbonate pump and thus play a crucial role in the marine carbon cycle~~. However, the current distribution of coccolithophore species and their dependence on environmental conditions are poorly known, hindering our ability to predict the response of the marine carbonate pump to changing climates. Here we investigated coccolithophore abundance~~s~~, species composition~~s~~, coccolithophore calcium carbonate ($CaCO_3$ as calcite) and particulate inorganic carbon (PIC) concentrations in the upper water column of the western North Pacific Ocean, along a meridional transect spanning the oligotrophic subtropical gyre and the nutrient-richer Kuroshio-Oyashio transition region. Our ~~results~~ samples and data revealed that *Umbellosphaera tenuis* was the numerically dominant coccolithophore species in the ~~former~~subtropical gyre, while *Emiliania huxleyi* and *Syracosphaera* spp. dominated in the transition region~~latter~~. The coccolithophore community composition showed significant depth- and latitude-dependent variations. Calcite from coccolithophores accounted for an average of 79 ± 27 % of the $CaCO_3$ standing stock in Niskin bottle samples in the euphotic zone, with a higher contribution observed in the subtropical gyre (91 ± 30 %) compared to the Kuroshio-Oyashio transition region (70 ± 24 %). This pattern was further supported by size-fractionated PIC concentrations of in situ pump samples, which showed a greater contribution of small PIC to total PIC in the subtropical gyre (76 ± 11 %) than in the transition region (67 ± 13 %). During the sampling period, coccolithophore $CaCO_3$ production rate ranged from 0.8 to 2.1 mmol m^{-2} d^{-1}, averaging 1.5 ± 0.7 mmol m^{-2} d^{-1} in the subtropical gyre and 1.2 ± 0.4 mmol m^{-2} d^{-1} in the transition region.~~Coccolithophore calcite contributed a major fraction of the PIC standing stocks above a depth of 150 m, among which *E. huxleyi* was the most important producer while less abundant and larger species also played a role. The coccolithophore $CaCO_3$ production rate in the subtropical gyre (0.62 mol m^{-2} yr^{-1}) was ~5 fold higher than that in the Kuroshio-Oyashio transition region (0.14 mol m^{-2} yr^{-1}), indicating that inorganic carbon metabolism driven by coccolithophores is relatively strong in oligotrophic ocean waters. Using a box model including coccolithophore $CaCO_3$ production and metabolic calcite saturation state, we demonstrated that $CaCO_3$ dissolution associated with organic carbon metabolism can generate excess alkalinity in the oversaturated upper water column of the western North Pacific Ocean.~~ Results of our study highlight the critical role of coccolithophores in the pelagic $CaCO_3$ ~~production and dissolution~~cycle, particularly in oligotrophic ocean waters.~~; knowledge of these processes is important to assess PIC cycling~~

~~and carbonate pump efficiency in the pelagic ocean.~~

1 Introduction

Calcium carbonate ($CaCO_3$) production and dissolution comprise~~are two major processes associated with~~ $CaCO_3$ cycling in the ocean, and are a key component of the global oceanic carbon cycle (Broecker and Peng, 1982) ~~via the so-called~~through the carbonate pump (Volk and Hoffert, 1985). Production of biogenic $CaCO_3$ by calcifying plankton in the euphotic zone elevates the partial pressure of carbon dioxide (CO_2) in seawater (e.g., Feely et al., 2002), while ballasting of sinking particles can promote the transport of carbon from the surface to deep sea and marine sediments (e.g., Armstrong et al., 2001; Klaas and Archer, 2002). Dissolution of $CaCO_3$ in the water column acts as a buffer to facilitate ocean sequestration of atmospheric CO_2 and ~~has the effect of reducing~~reduces the rate of ocean acidification (Feely et al., 2004; Barrett et al., 2014). Over the last decade, ocean acidification, a global reduction in seawater pH caused by the uptake of anthropogenic CO_2, has emerged as a significant feedback mechanism, ~~This acidification feedback mechanism makes~~ making it harder for calcifying organisms to produce their skeletons, and thus adversely affects marine ecosystems (Feely et al., 2004; Ma et al., 2023)~~(Feely et al., 2004; Steiner et al., 2018)~~. Therefore, quantification of marine $CaCO_3$ production and dissolution is of vital importance in determining the response of marine ecosystems to changes in the partial pressure of CO_2.

Marine $CaCO_3$ occurs in the form of calcite, aragonite and high-magnesium calcite. Coccolithophores are a key, single-celled phytoplankton taxonomic group, responsible for a large percentage (30–60 %) of modern oceanic $CaCO_3$ production and 10–20 % of marine primary production on a global scale (Poulton et al., 2006, 2013). Coccolithophore calcite accounts for a major fraction (24–80 %) of the $CaCO_3$ exported to the deep sea and sediments (Broerse et al., 2000; Young and Ziveri, 2000; Rigual Hernández et al., 2020). Field observations~~Data assessment~~ along a northeast Pacific transect from Hawaii to Alaska suggested that coccolithophore calcite comprises 90 % of the total $CaCO_3$ production in the euphotic zone, while pteropods and foraminifera only play a minor role (Ziveri et al., 2023). However, large uncertainties remain in estimates of the production rate of $CaCO_3$ in the upper ocean, as well as the contributions of different plankton groups, which are still unclear and vary across regions (Balch et al., 2007; Berelson et al., 2007; Smith and Mackenzie, 2016; Ziveri et al., 2023). Based on a global compilation of $CaCO_3$ production using in situ ^{14}C incubations, Daniels et al. (2018) found that calcification rate ranged from ~~10~~ ≤0.1 to ~~600~~ 6 ~~mg~~ mmol m^{-2} d^{-1} in the euphotic zone. A recent estimate of $CaCO_3$ biomass from three main pelagic calcifying plankton groups also suggested large variation in $CaCO_3$ production in the eastern North Pacific Ocean,

ranging from ~~110 to 729 mg~~1.1 to 7.3 mmol m^{-2} d^{-1} (Ziveri et al., 2023). ~~The generally low coverage of observations and the considerable spatiotemporal variation shown by available data, along with the current scarcity of studies, most likely result in deviations among regional estimates.~~

~~$CaCO_3$ dissolution is generally assumed to mainly occur below the saturation horizon. However, this assumption has been challenged by an increasing number of studies which suggest considerable dissolution in the oversaturated upper water column where $CaCO_3$ saturation state (Ω) is greater than 1 (Sabine et al., 2002; Chung et al., 2003; Berelson et al., 2007; Barrett et al., 2014). Shallow-water $CaCO_3$ dissolution is supported by excess alkalinity/calcium (Feely et al., 2002, 2004; Cao and Dai, 2011; Subhas et al., 2022), decreased sinking fluxes of particulate inorganic carbon (PIC) measured using the ^{234}Th method, and sediment trap data form the upper ocean (Dong et al., 2019; Roca-Martí et al., 2021). One possible mechanism for $CaCO_3$ dissolution in calcite-oversaturated shallow waters is the dissolution of more soluble forms of $CaCO_3$ including aragonitic pteropods and high-Mg fish calcites (Honjo et al., 2008; Wilson et al., 2009; Dong et al., 2019; Folkerts et al., 2024; Oehlert et al., 2024). An alternative hypothesis is that microbial oxidation of organic matter produces an acidic microenvironment conducive for carbonate dissolution (Bishop et al., 1980; Milliman et al., 1999). The digestive system of grazing zooplankton may also contain acidic conditions facilitating $CaCO_3$ dissolution (Pond et al., 1995; White et al., 2018).~~

The North Pacific Ocean is a vital region for modulating the carbon cycle, as it accounts for ~25 % of the global ocean sink for atmospheric CO_2 (Takahashi et al., 2009). In the eastern North Pacific Ocean, $CaCO_3$ production, export, and dissolution have been studied along a transect from Hawaii to Alaska (Dong et al., 2019, 2022; Naviaux et al., 2019; Subhas et al., 2022; Ziveri et al., 2023)~~.~~. Ziveri et al. (2023) found ~~which revealed~~ that depth-integrated $CaCO_3$ production in the nutrient-rich subpolar gyre is twice as high as that in the nutrient-poor subtropical gyre. This contrast, however, is smaller than the sixfold to sevenfold difference based on satellite estimates of surface particulate inorganic carbon (PIC), indicating the importance of coccolithophore $CaCO_3$ production over a deeper euphotic zone~~in deep waters~~ and the limitation of satellite products as highlighted by Neukermans et al. (2023). ~~These investigations also suggested that widespread shallow-water $CaCO_3$ dissolution is driven by metabolic activity along the Hawaii-to-Alaska transect.~~

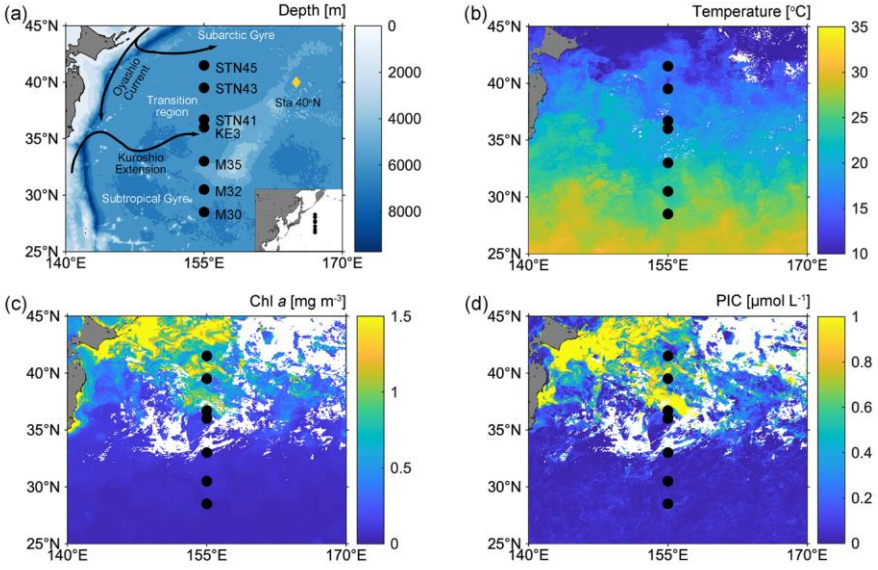
Here, we determined the abundance~~s~~ and species composition~~s~~ of coccolithophore~~s~~, as well as the concentrations of coccolithophore calcite and PIC based on both Niskin bottle and in situ pump sampling in the upper water column of the

western North Pacific Ocean. Additionally, we conducted measurements of environmental conditions such as nutrient and carbonate chemistry parameters. The aims of this research were to answer the following questions: (1) What is the distribution of coccolithophore abundances and species compositions across the oligotrophic-nutrient replete environmental gradient? (2) What is the contribution of coccolithophores to $CaCO_3$ production in the euphotic zone? ~~(3) Does shallow-water $CaCO_3$ dissolution occur in the western North Pacific Ocean, and what is the significance of metabolic activities in driving dissolution in oversaturated ambient conditions?~~

2 Methods

2.1 Sample collection

Sampling was conducted onboard R/V *Tan Kah Kee* during cruise NORC2022-306 from 09 June to 25 July 2022. The cruise trajectory crossed from the oligotrophic North Pacific Subtropical Gyre (NPSG) to the relatively nutrient-rich Kuroshio-Oyashio transition region along the 155°E meridian (Fig. 1a; Table S1). Seven sampling stations can be divided into those located in the NPSG region, including stations M30, M32 and M35, characterized by high sea-surface temperature (SST) and low surface chlorophyll *a* (Chl *a*) and PIC concentrations, and those located in the Kuroshio-Oyashio transition region, including stations KE3, STN41, STN43 and STN45, featuring lower SST, but higher Chl *a* and PIC concentrations (Fig. 1b–d).


(a)
Depth [m]
Subarctic Gyre
Oyashio Current
Transition region
Kuroshio Extension
Subtropical Gyre
STN45
STN43
Sta 40°N
STN41
KE3
M35
M32
M30
45°N
40°N
35°N
30°N
25°N
140°E
155°E
170°E
0
2000
4000
6000
8000
(b)
Temperature [°C]
35
30
25
20
15
10
(c)
Chl a [mg m-3]
1.5
1
0.5
0
(d)
PIC [µmol L-1]
1
0.8
0.6
0.4
0.2
0

104

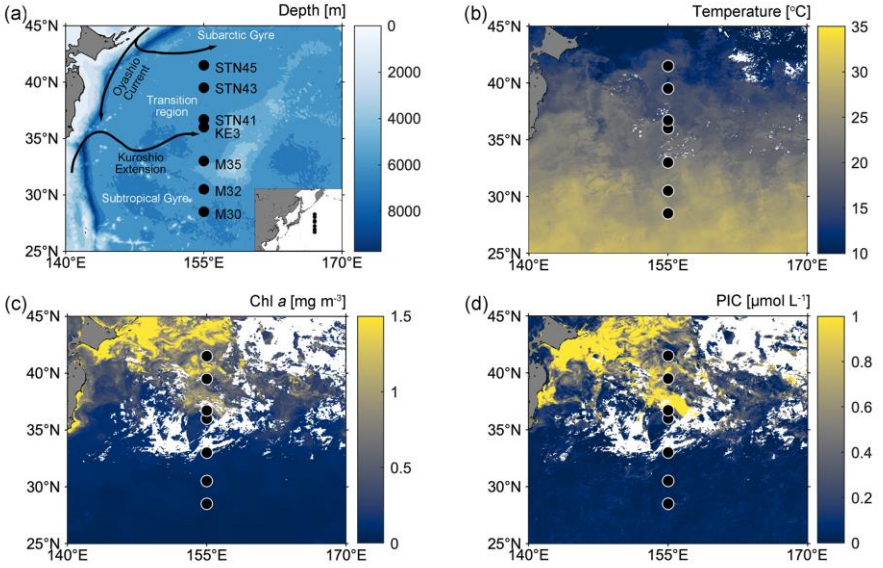


Fig 1. (a) Map of the western North Pacific Ocean showing sampling stations (black filled circles) and major surface currents (solid black lines)~~. The yellow diamond indicates station 40° N at which particulate inorganic carbon (PIC) sinking flux was investigated by Honda et al., (2002)~~; (b–d) satellite-based temperature, chlorophyll *a* (Chl *a*) and particulate inorganic carbon (PIC) concentrations in surface water from 1st to 30th June 2022~~in June 2022~~ (data from the Moderate Resolution Imaging Spectroradiometer (MODIS)-Aqua satellite; https://oceancolor.gsfc.nasa.gov/l3/).

Water samples were collected within the water column above 300 m depth using Niskin bottles on a rosette system equipped with SBE-911 conductivity-temperature-depth (CTD) sensors (Sea-Bird Electronics, Inc., Bellevue, WA, USA). For PIC analyses, 24 L of seawater were collected using acid-cleaned fluorinated bottles and filtered through two quartz microfiber (QMA) filters (1.0 µm pore size, 25 mm diameter). For coccolithophore analyses, 2–4 L of seawater were collected and gently filtered through polycarbonate membranes (0.8 µm pore size, 25 mm diameter), using a vacuum pump at <20 mm Hg pressure.

Membrane ~~F~~filters were oven-dried at 60°C and stored in plastic petri dishes.

Size-fractionated particles were collected using McLane Research in situ pumps. Filter holders were loaded with a 51 µm Sefar polyester mesh prefilter followed by paired Whatman QMA filters. Hereafter, we refer to the two particle size fractions as large (> 51 µm) and small (1–51 µm) size fractions ~~(LSF, > 51 µm and SSF, 1–51 µm, respectively)~~. A 1/4 subsample of the 51 µm polyester mesh prefilter and two circles of 23 mm diameter subsample of the QMA filter were analyzed for ~~LSF~~ large and ~~SSF~~ small PIC concentrations, respectively, and the sum of the two fractions yielded the total PIC ~~(PIC_{total})~~ concentration.

2.2 Sample analyses

PIC concentrations were determined by measuring the amount of CO_2 released after acid treatment of the filters using a Thermo Delta V Plus isotope ratio mass spectrometer (IRMS, Thermo Fisher, USA) coupled with a Thermo Gasbench II system at the Center for Isotope Geochemistry and Geochronology of the Laoshan Laboratory (Li et al., 2021). International reference materials of calcite NBS-18 and IAEA-603 were measured for calibration. The ~~PIC~~ analytical precision of PIC determination was ~~better than~~ <10 % (one standard deviation, 1SD).

~~A portion of the filters was~~Filters were cut and mounted with a carbon sticky tab on a stub and gold-coated prior to analysis using a Quanta 650 FEG field-emission scanning electron microscope (SEM). The coccosphere cell or detached coccolith concentrations (CC, cells or coccoliths L^{-1}) were estimated as follows:

$$CC = (F * C)/(V * S) \quad (1)$$

where F is the effective filtration area (336.9 mm^2), C is the total number of coccosphere cells or detached coccoliths, V is the filtered seawater volume, and S is the total area of fields of view (mm^2). This cell counting strategy gives a detection limit of at least 1.87 cells mL^{-1} (Bollmann et al., 2002). The coccolithophore abundance in four samples, that were collected at 10 m and 200 m at station M30, 200 m at station KE3 and 200 m at station STN45, fell below the detection limit. Despite potential inaccuracies, these values are still meaningful as they indicate an exceptionally low coccolithophore presence. Coccolithophore species identification followed Young et al. (2003) and the Nannotax3 website (http://ina.tmsoc.org/Nannotax3/). Aggregates formed by clusters of multiple coccolithophores were quantified in terms of abundance but were excluded from the coccolithophore calcite calculations, mainly due to the difficulty in accurately determining the number of individual coccoliths within the aggregates.~~Individual coccolithophore calcite content was calculated by multiplying the number of coccoliths per~~

~~cell by the average coccolith calcite mass of a given species (Young and Ziveri, 2000; Yang and Wei, 2003; Boeckel and Baumann, 2008; Beuvier et al., 2019; Jin et al., 2022). All the biometry work based on SEM images was conducted using imageJ free software (imagej.nih.gov/ij/) and Coccobiom2-SEM measuring macro (Young, 2015).~~ Individual coccolithophore calcite content was calculated by multiplying the number of coccoliths per cell by the average coccolith calcite mass of a given species. The average coccolith mass was estimated based on the coccolith size (usually using coccolith length) and a factor related to coccolith cross-sectional shape (Young and Ziveri, 2000):

$$m\ (pg\ CaCO_3) = 2.7 * K_s * l^3 \quad (2)$$

where l is the coccolith size (μm), K_s is a species-specific shape constant, and 2.7 is the calcite density ($CaCO_3$; pg μm^{-3}). The specific coccolith distal shield length or process height used in the calculation was measured from SEM images. Measurements were conducted using ImageJ free software (imagej.nih.gov/ij/) and Coccobiom2-SEM measuring macro (Young, 2015). The K_s values used were from Young and Ziveri (2000) and Jin et al. (2016). The number of coccoliths per coccosphere was obtained from Yang and Wei (2003) and Boeckel and Baumann (2008). The calculation of coccolith PIC is detailed in Table S2 in the Supplement. Sheward et al. (2024) have extensively discussed the potential errors of the morphometric-based calcite estimation method, suggesting that an additional uncertainty of 5–40 % may arise from slight variations in K_s and size between coccoliths on the same coccosphere, as well as errors in coccolith number estimation. Additionally, it is important to note that further uncertainties can be introduced by counting inaccuracies, particularly in cases where clumps or overlapping coccoliths are present. Despite these possible errors and limitations, our data and results offer robust and comparable insights into coccolithophore calcite dynamics.

Chl *a* concentrations were measured after being extracted with 90 % acetone for 14 h at −20°C using a Trilogy Laboratory Fluorometer with non-acidification module (Turner Designs, USA) (Welschmeyer, 1994). Nutrient samples were collected in acid-washed Nalgene high-density polyethylene bottles and determined onboard the vessel using a Four-channel Continuous-Flow Technicon AA3 Autoanalyzer (Bran+Luebbe GmbH). The detection limits were 0.1 μmol L^{-1}, 0.08 μmol L^{-1}, and 0.16 μmol L^{-1} for dissolved inorganic nitrogen (DIN, nitrate plus nitrite), soluble reactive phosphate (SRP), and dissolved silicate (DSi), respectively. The analytical precisions (derived from repeat measurements of aged deep seawater) were 0.44% for DIN, 0.91% for SRP, and 0.28% for DSi (n = 82). Analysis of reference standard LOT.CM (KANSO TECHNOS CO., LTD.)

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produced concentrations of 33.72 ± 0.13 µmol L^{-1} for DIN, 2.460 ± 0.025 µmol L^{-1} for SRP, and 102.2 ± 0.3 µmol L^{-1} for DSi (n = 20), which agree well with consensus values (http://www.kanso.co.jp/eng/pdf/certificate_cb.pdf). For measurements of DIN and SRP concentrations in surface samples below the detection limit of the AA3 Autoanalyzer, duplicate samples were collected and frozen separately at −20°C until analysis. Nanomolar DIN concentrations were determined using a continuous-flow analysis system combined with a liquid waveguide capillary flow cell as described by Zhang (2000). The detection limit was 5.2 nmol L^{-1} and the analytical precision was 7.5% (derived from repeat measurements of aged deep seawater with 1000-fold dilution, 36.2 ± 2.7 nmol L^{-1}, n = 57). Nanomolar SRP concentrations were measured using an automated analyzer including a syringe pump and multiposition selection valve combined with a solid-phase extraction cartridge (Deng et al., 2020). The detection limit was 2.5 nmol L^{-1} and the analytical precision was 5% (derived from repeat measurements of aged deep seawater with 1000-fold dilution, 26.0 ± 1.2 nmol L^{-1}, n = 56). Seawater ammonium (NH_4^+) concentrations were measured onboard using solid-phase extraction combined with fluorescence determination with a detection limit of 3.6 nmol L^{-1} (Zhu et al., 2013, 2018).

Samples for analysis of dissolved inorganic carbon (DIC) and total alkalinity (TA) were collected in 250 mL PYREX® borosilicate glass bottles, and poisoned with 250 µL of a $HgCl_2$-saturated solution upon sample collection. DIC was measured using an infrared CO_2 detector (Apollo ASC-3) , with a precision of ± 2 µmol L^{-1} (Cai et al., 2004). TA was determined on 25 mL samples using an open-cell setting based on the Gran titration technique (Cai et al., 2010) with a Kloehn digital syringe pump. The analytical precision was ± 2 µmol L^{-1}. Both DIC and TA concentrations were calibrated against certified reference materials provided by Andrew G. Dickson (the Scripps Institution of Oceanography, University of California, San Diego, USA).

2.3 Estimation of $CaCO_3$ production rate

The euphotic zone bottom at each station was defined as the depth where surface photosynthetically active radiation (PAR) reaches 0.1 % (Table S1). $CaCO_3$ production rates ~~above 150 m~~in the euphotic zone were determined by dividing measurements of the living $CaCO_3$ standing stock (which only included whole coccosphere cells and excluded loose coccoliths) by the coccolithophore turnover time, which is 0.7–10 days with a growth rate ranging from 0.1 to 1.5 cell divisions day^{-1} (Krumhardt et al., 2017; Ziveri et al., 2023). The coccolithophore turnover time was derived from both laboratory and field

estimates, as well as simulations from a generalized coccolithophore model, which has also been applied to the eastern North Pacific Ocean (Krumhardt et al., 2017; Ziveri et al., 2023). We are aware that different coccolithophore species exhibit widely varying growth rates and cell growth phase differs. Smaller cells produce fewer coccoliths during the exponential growth phase characterized by rapid division, whereas larger cells generate more coccoliths during the early stationary phase when cell division slows down (Raven and Crawfurd, 2012; Krumhardt et al., 2017). We also acknowledge that estimating coccolithophore calcite and production rates using an average coccolith calcite value introduces uncertainties, as this approach does not fully account for the complexity of coccolith dynamics, including rapid cycling and reabsorption (Johns et al., 2023). Despite these possible errors and uncertainties, our estimations generally comparable with those of prior work (e.g., Daniels et al., 2018), remain a reliable basis for assessing coccolithophore calcification. Uncertainty in the $CaCO_3$ standing stock estimates, which were obtained by vertically integrating PIC concentrations in the euphotic zone ~~above a depth of 150 m~~, was typically ±10 % (1SD).

A Monte Carlo-based probabilistic approach was used to determine the $CaCO_3$ production rate and the uncertainties associated with the turnover time using the R package violoto ~~a flat probability distribution in MATLAB visualized by a violin plot (Hoffmann, 2015)~~. To obtain an annual $CaCO_3$ production based on our field observations, we used the ratio of satellite-derived PIC for July 2022 to annual climatology PIC (data from the NASA Goddard Space Flight Center's Ocean Ecology Laboratory) to calibrate for potential seasonal variability (Ziveri et al., 2023).

2.4 Influence of environmental conditions on coccolithophores

The redundancy analysis (RDA) is a widely used multivariate analytical method to identify relationships among individual variables in different categories. Prior to the RDA, statistical differences in environmental variables were evaluated using an analysis of variance (one-way ANOVA), while collinearity between environmental variables was accounted for by calculating variance inflation factors (VIF). Forward selection of variables was subsequently carried out until all VIF scores were <10, in order to only including variables that are not significantly correlated. These criteria reduced the number of environmental variables used in the RDA. Monte Carlo permutation tests, based on 1000 randomizations, were performed to identify the most significant and independent effect on variation in the coccolithophore community composition. The overall significance of the explanatory variables after forward selection was evaluated through ANOVA (α <0.05) and coefficient of determination (r^2),

and adjusted r^2 were calculated to assess the power of a selected RDA model using the vegan package (Oksanen, 2010).~~The redundancy analysis (RDA) function in the vegan package in R, in combination with Monte Carlo permutation, were performed to assess the relative importance of environmental variables in explaining the overall variation in the composition of the coccolithophore community (Oksanen et al., 2007).~~ The contribution of each environmental variable to community variation was determined by hierarchical partitioning in canonical analysis via the 'dbRDA' function in the "rdacca.hp" package in R (Lai et al., 2022).~~We further conducted random forest analyses to identify the main predictors of coccolithophore abundance with the "randomForest" package in R (Liaw and Wiener, 2002).~~

~~2.5 Evaluation of shallow-water $CaCO_3$ dissolution~~

~~To estimate the effect of microenvironment undersaturation driven by microbial oxidation of organic matter, we assumed that aerobic metabolic activity of marine particles consumes all ambient oxygen and alters ambient dissolved inorganic carbon (DIC) and total alkalinity (TA). This process decreases microenvironment Ω and thus causes $CaCO_3$ dissolution. We calculated Ω due to in situ metabolism (defined as Ω_{met}) using the concurrently altered DIC, TA and soluble reactive phosphate (SRP) concentrations following Subhas et al., (2022):~~

~~$$DIC_{met} = DIC + [O_2] * R_{CO} \quad (2a)$$~~

~~$$TA_{met} = TA - [O_2] * R_{NO} \quad (2b)$$~~

~~$$SRP_{met} = SRP + [O_2] * R_{PO} \quad (2c)$$~~

~~where R_{CO}, R_{NO}, and R_{PO} denote the Redfield ratio of carbon to oxygen (0.688), nitrate to oxygen (0.0941), and phosphate to oxygen (0.0059), respectively (Anderson and Sarmiento, 1994).~~

~~A one-dimensional (1-D) model was used to diagnose the sinking flux and $CaCO_3$ dissolution rate in the upper 1000 m of the water column by coccolithophore $CaCO_3$ production, coccolith dissolution kinetics and Ω_{met} (Dong et al., 2019; Subhas et al., 2022). The model assumes that dissolution of $CaCO_3$ particles occurs when the water column $\Omega_{met} < 1$. The magnitude of dissolution thus depends on the initial sinking flux, Ω_{met} and particle residence time as follows:~~

~~$$Flux_{zt} = Flux_{zt-1} * \left(1 - R_{diss} * \frac{z_t - z_{t-1}}{w}\right) \quad (3)$$~~

~~where $Flux_{zt}$ and $Flux_{zt-1}$ (mmol m^{-2} d^{-1}) denote the $CaCO_3$ sinking flux at a given depth Z_t and its overlying depth Z_{t-1},~~

~~respectively, R_{diss} (g g^{-1} d^{-1}) denotes $CaCO_3$ dissolution rate which is a function of the Ω_{mct} at depth Z_{i-1} (Subhas et al., 2018; Subhas et al., 2022) and w (m d^{-1}) denotes the particle sinking rate.~~

~~TA regeneration rate (R_{TA}, μmol kg^{-1} yr^{-1}) at depth Z_i resulting from shallow-water $CaCO_3$ dissolution was calculated based on the following equation:~~

~~$$R_{TA\_Zi} = \frac{2 * Flux_{Zi} * R_{diss}}{\rho * w} \quad (4)$$~~

~~where ρ denotes the density of seawater (1029 kg m^{-3}).~~

3 Results

3.1 Hydrography ~~and nutricline~~

~~Temperature and salinity were high in surface water and gradually decreased with increasing depth above 300 m.~~ Hydrochemical variables ~~clearly~~ exhibited a south to north trend. Temperature and salinity were highest at the surface of station M30, due to strong net evaporation in the subtropical gyre (Fig. 2a and b). There was a northward decrease in temperature and salinity due to the influence of upwelling in the subarctic gyre. ~~The surface mixed layer depth, defined as the depth where potential density increases by 0.03 kg m^{-3} compared to that at the sea surface, varied around 11–25 m.~~

In contrast to temperature and salinity and as expected, the distribution of ~~dissolved inorganic nitrogen (~~DIN~~, nitrate plus nitrite)~~, SRP and ~~dissolved silicate (~~DSi~~)~~ showed a generally northward increasing pattern (Figs. 2c–d and S1a). Surface DIN concentrations were ~~below the detection limit~~on average 0.006 μmol L^{-1} in the NPSG region and ~~averaged~~ 0.02 μmol L^{-1} in the Kuroshio-Oyashio transition region. ~~The top of the nutricline, defined as the depth where DIN concentrations reach 0.1 μmol L^{-1}, ranged from 110 m at station M30 to 20 m at station STN45.~~ The ~~ammonium (~~NH_4^+~~)~~ concentration above 100 m at station STN45 was notably higher than that at other stations (Fig. S1b). ~~Similar to the nutricline distribution, t~~The deep chlorophyll maximum (DCM) depth gradually shoaled northward from 110 m at station M30 in the NPSG region to 33 m at station STN45 in the Kuroshio-Oyashio transition region (Fig. 2e).

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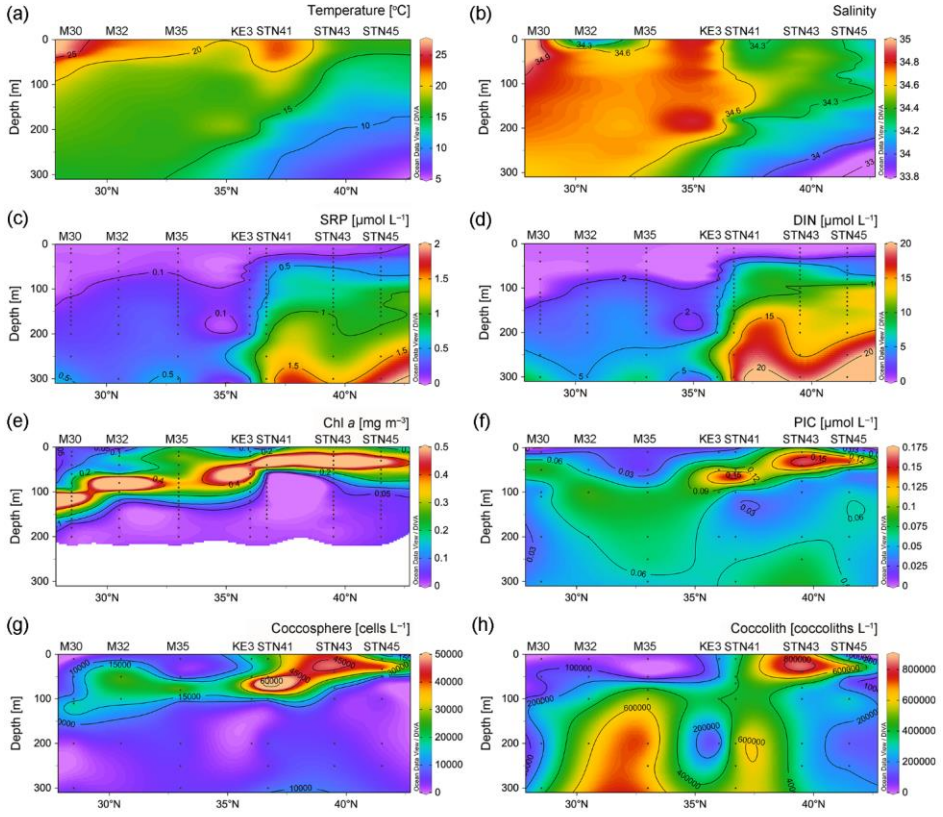

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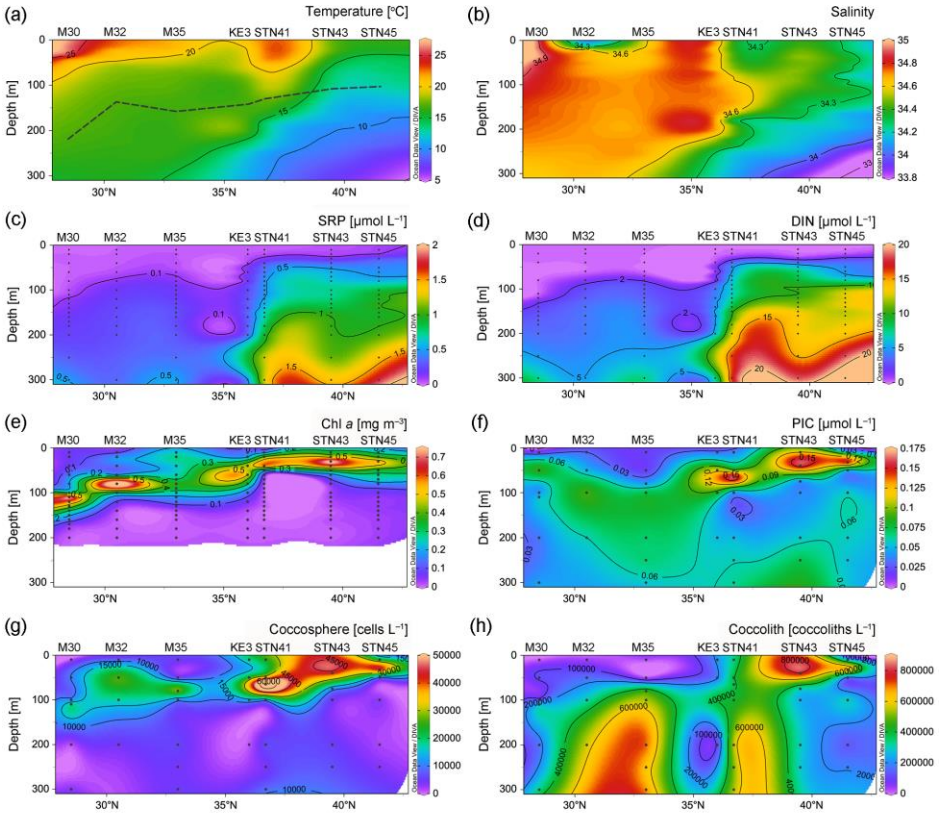

Fig. 2. Vertical depth distributions of (a) temperature, (b) salinity and concentrations of (c) soluble reactive phosphate (SRP), (d) dissolved inorganic nitrogen (DIN, nitrate plus nitrite), (e) Chlorophyll *a* (Chl *a*), (f) particulate inorganic carbon (PIC), (g) coccosphere cell and (h) detached coccoliths in the upper 300 m of the water column in the study area. In (a), the black dashed line indicates the bottom of the euphotic zone.

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3.2 Vertical distribution of PIC and coccolithophore concentrations

PIC concentrations along the 155°E transect ranged from 0.02 to 0.17 μmol L^{-1}, with an average of 0.06 ± 0.04 μmol L^{-1} in the upper 300 m of the water column (Fig. 2f). Generally, PIC concentrations were lower at the surface and increased with increasing depth to attain a maximum in the DCM layer, and decreased with depth thereafter. In the DCM layer, PIC concentrations ranged from 0.06 μmol L^{-1} at 110 m of station M30 in the subtropical gyre to 0.16 μmol L^{-1} at 33 m of station STN45 in the Kuroshio-Oyashio transition region. The vertical distribution pattern of bottle-derived PIC and coccosphere cell concentrations overall followed that of Chl *a*, showing a northward shoaling of the subsurface maximum.

Concentrations of coccosphere cells ranged from ca. 970 to 75,000 cells L^{-1} (Fig. 2g). Along the transect, a subsurface maximum was evidenced around the DCM layer with an average of 42,000 cells L^{-1}, followed by a steep decrease below 100 m. The highest coccosphere cell concentration was observed at 65 m of station STN41, corresponding to the highest PIC concentration. The average coccosphere cell concentration was notably lower in the NPSG region (9,800 cells L^{-1}) than in the transition region (18,000 cells L^{-1}). The detached coccolith concentration averaged 340,000 coccoliths L^{-1}, with a range of 11,000 to 800,000 coccoliths L^{-1} (Fig. 2h). The highest concentration was observed around 10–40 m of station STN43. High coccolith concentrations were also observed below 100 m at stations M32, M35 and STN41.

Size-fractionated PIC concentrations from in situ pumps varied from 0.01 to 0.09 μmol L^{-1} in the ~~SSF~~ small size fraction and from 0.01 to 0.06 μmol L^{-1} in the large size fraction~~LSF~~. ~~PIC_{total}~~ Total PIC concentrations averaged 0.07 ± 0.02 μmol L^{-1}, and were comparable to bottle-derived PIC concentrations (Fig. 3). Roughly 70 % of the PIC was contributed by the small size fraction~~SSF~~ at each sampling station. Generally, large size fraction PIC~~LSF PIC~~ concentrations increased northward from stations M30–M35 to stations KE3–STN45 and accounted for 22 % and 36 % of total PIC~~PIC_{total}~~ concentrations in the NPSG region and the Kuroshio-Oyashio transition region, respectively. The maximum concentration of large size-fractionated PIC ~~LSF PIC~~ (0.06 μmol L^{-1}) was observed at 26 m of station STN45 (Fig. 3g).

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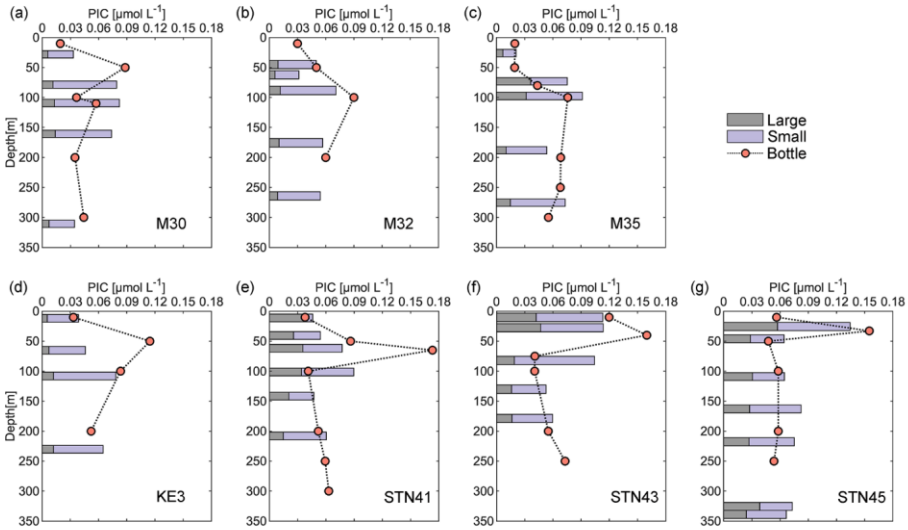
(a)
PIC [μmol L-1]
0 0.03 0.06 0.09 0.12 0.15 0.18
Depth[m]
M30
(b)
M32
(c)
M35
Large
Small
Bottle
(d)
KE3
(e)
STN41
(f)
STN43
(g)
STN45

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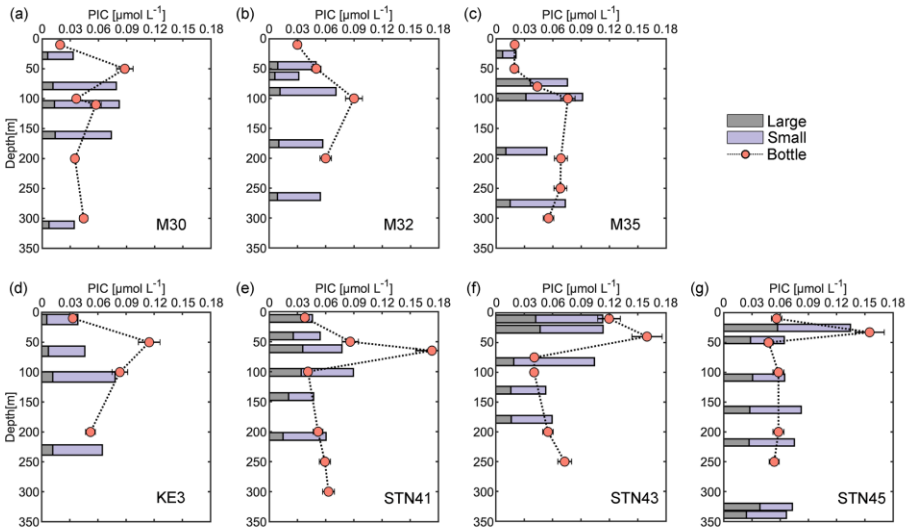
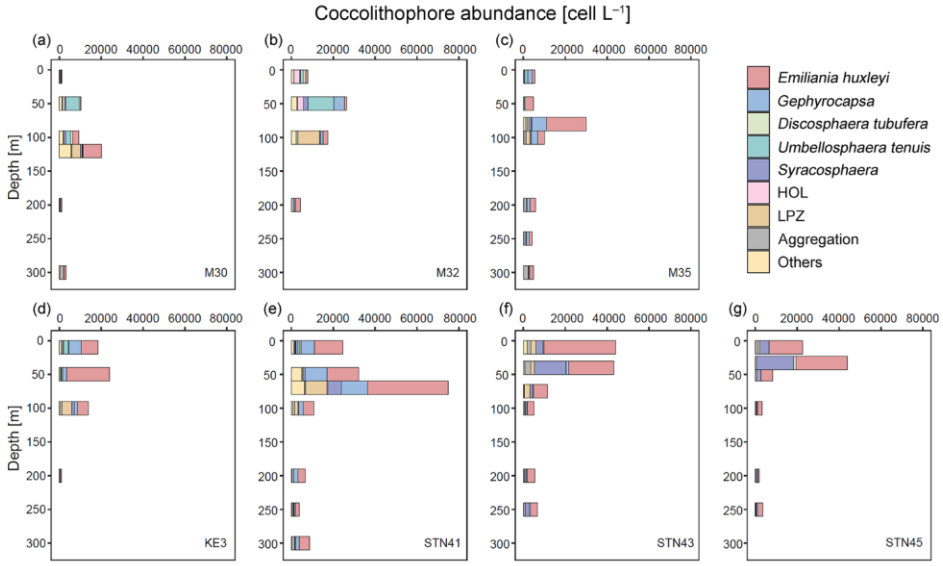


Fig. 3. Vertical depth distributions of particulate inorganic carbon (PIC) concentrations derived from sampling using both Niskin bottles and in situ pumps (small size fraction of 1–51 μm and large size fraction of > 51 μm) in the upper 350 m of the water column at sampling stations in the study area.

3.3 Characteristics of the coccolithophore assemblage

Coccolithophore populations were predominantly represented by *Emiliania huxleyi*, *Gephyrocapsa ericsonii*, *Gephyrocapsa oceanica*, *Umbellosphaera tenuis*, *Syracosphaera* spp., holo-coccolithophores (HOL), *Algirosphaera robusta,* and *Florisphaera profunda* (each comprising > 1 % of total coccosphere abundance; Fig. 4). In surface water, coccolithophore cells were dominated by *Dicosphaera tubifera*, *U. tenuis* and HOL at stations M30 and M32 (Fig. 4a and b) and by *G. ericsonii* at stations M35, KE3 and STN41 (Fig. 4c, d and e), while high abundance of *E. huxleyi* and *Syracosphaera* spp. was clearly observed at stations STN43 and STN45 (Fig. 4f and g). It is noteworthy that *E. huxleyi* contributed the largest fraction (50 %) to the total coccolithophore ~~assemblage~~ cells and was also found to be the dominant species in the DCM layer. *U. tenuis* was

mainly observed in subtropical gyre waters, with peak abundance at 50 m and lower abundance at the surface and in the DCM ~~waters~~ (Fig. 4a and b). Lower euphotic zone (LPZ, defined as the region of the water column that receives 10–1% of surface PAR) coccolithophore species (including *A. robusta* and *F. profunda*) were commonly found in the subsurface population below 50 m, accounting for 7 % of the entire coccolithophore community (Jin et al., 2016; Poulton et al., 2017). Overall, coccolithophores were scarce in the NPSG region and dominated by *U. tenuis*, whereas their abundance notably increased in the Kuroshio-Oyashio transition region where it was dominated by *E. huxleyi*, *Gephyrocapsa* and *Syracosphaera* spp..



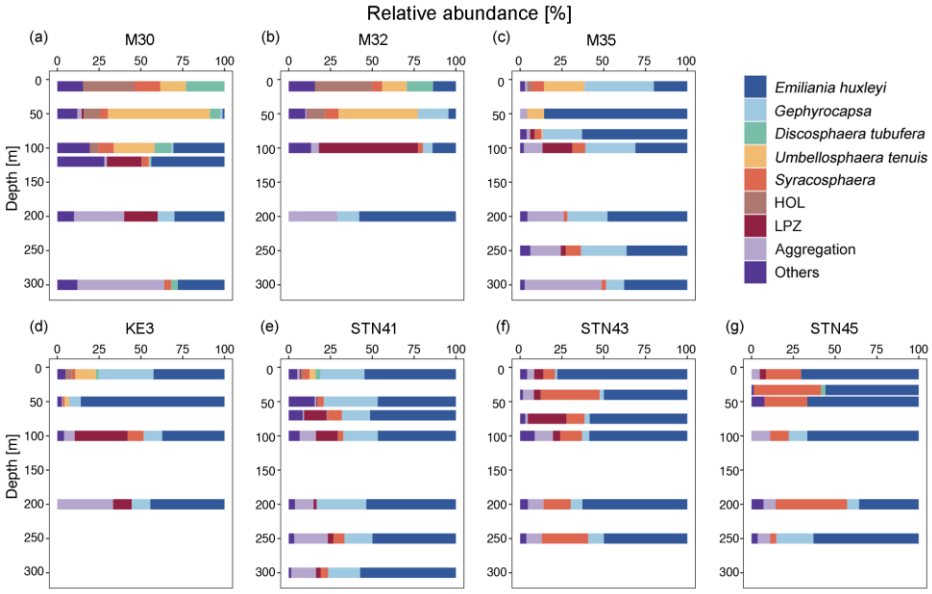
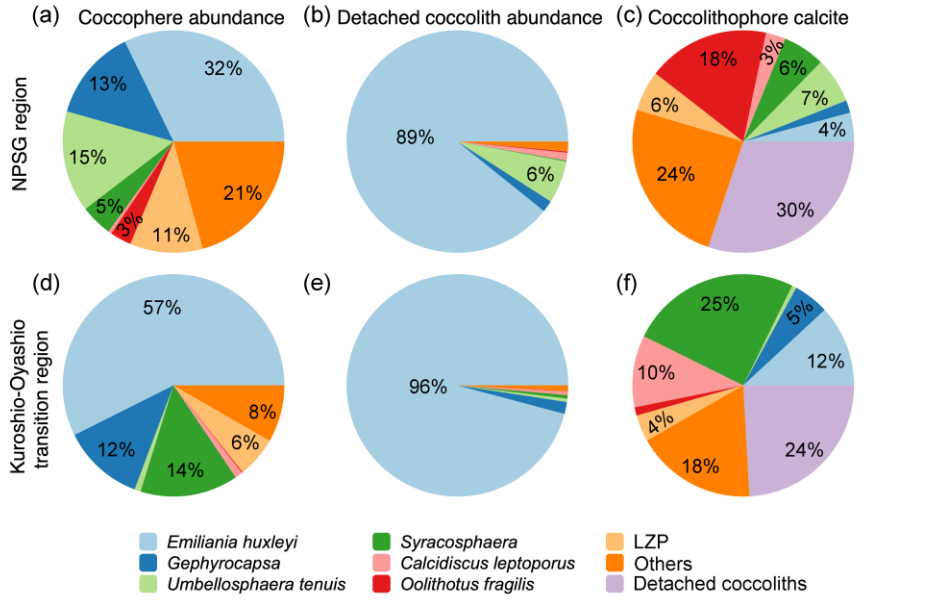


Fig. 4. Relative ~~Abundance~~ abundance of different coccolithophore groups in the upper 300 m of the water column ~~at sampling stations in the study area~~. Lower euphotic zone (LPZ) species include *Florisphaera profunda* and *Algirosphaera robusta*; HOL indicates holo-coccolithophores.

The estimated coccolithophore calcite concentrations ranged from <0.01~~0.00~~ to 0.23 µmol L^{-1}, averaging 0.05 ± 0.04 µmol L^{-1} above 300 m along the 155°E transect. The coccospheres of *E. huxleyi* accounted for ~~50~~ 32 % and 57 % of the total coccolithophore ~~assemblage~~ cells but represented only ~~9~~ 4 % and 12 % of the coccolithophore calcite concentration in the NPSG region and the Kuroshio-Oyashio transition region, respectively (Fig. 5~~a and c~~). In the NPSG region, *U. tenuis* accounted for 15 % of the total coccolithophore cells and 7 % of the coccolithophore calcite concentration, both notably higher than in the transition region, where its contribution was <1 % for both measures. ~~The less abundant (<1 %) species *Calcidiscus leptoporus* and *Oolithotus fragilis* accounted for 7.5 % and 7.7 % of the coccolithophore calcite concentration, respectively.~~

Syracosphaera spp. was the largest contributor in the Kuroshio-Oyashio transition region, accounting for ~~17.7~~25 % of the coccolithophore calcite concentration (Fig. 5~~e~~f). The less abundant (<3 %) species *Calcidiscus leptoporus* and *Oolithotus fragilis* accounted for 21 % and 12 % of the coccolithophore calcite concentration in the NPSG region and the Kuroshio-Oyashio transition region, respectively. Additionally, ~~*E. huxleyi*~~ detached coccoliths ~~, comprising 94 % of the total detached coccoliths, made up~~contributed to ~~16 %~~30 % and 24 % of the total coccolithophore calcite concentration in the two regions, respectively (Fig. ~~5b~~ 5c and ~~e~~f).



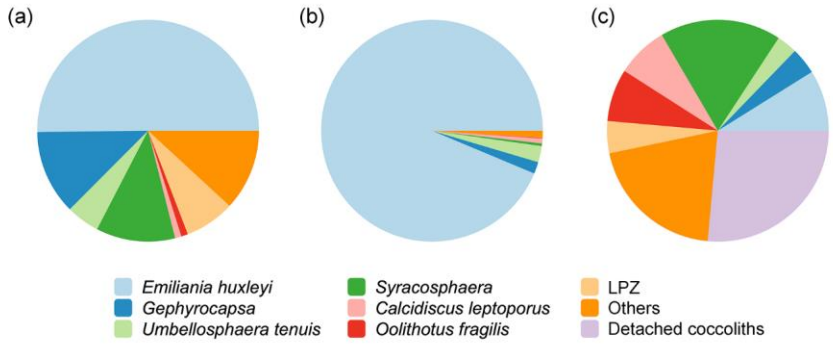


Fig. 5. Contribution of different coccolithophore groups to coccosphere cell abundance, detached coccolith abundance, and coccolithophore calcite concentrations in the upper 300 m of the water column in (a–c) the North Pacific Subtropical Gyre (NPSG, stations M30, M32 and M35) and (d–f) the Kuroshio-Oyashio transition region (stations KE3, STN41, STN43 and STN45). ~~**Fig. 5.** Contribution of different coccolithophore groups to (a) coccosphere cell abundance, (b) detached coccolith abundance and (c) coccolithophore calcite concentration in the upper 300 m of the water column.~~ Lower euphotic zone (LPZ) species include *Florisphaera profunda* and *Algirosphaera robusta*.

3.4 $CaCO_3$ standing stock and production

The standing stocks of $CaCO_3$ in the euphotic zone were determined using data from Niskin bottles, coccolithophore calcite, and size-fractionated samples (Fig. 6a). ~~The standing stock of $CaCO_3$ was determined by considering the first shallow sampling depth to a consistent depth of 150 m (Figs. 6a and S2a).~~ $CaCO_3$ standing stock derived from Niskin bottle-sampling ranged from ~~660~~ 7.0 to ~~1,200~~11.1 mmol m^{-2} ~~mg m^{-2}~~, and was slightly lower in the oligotrophic NPSG region (~~790~~ 8.7 ± 1.7 mmol m^{-2} ~~mg m^{-2}~~) than in the relatively nutrient-high Kuroshio-Oyashio transition region (9.2 ± 1.7 mmol m^{-2} ~~1,100 mg m^{-2}~~). Based on the estimated coccolithophore calcite concentrations, $CaCO_3$ standing stocks ranged from ~~370 to 1,000 mg m^{-2}~~ 4.0 to 11.3 mmol m^{-2} and peaked at station ~~STN41~~M30 due to its deepest euphotic zone (Fig. 2a and 6a). Calcite from coccolithophores comprised on average ~~76 %~~79 ± 27 % of the $CaCO_3$ standing stock from Niskin bottle samples, and the contribution was

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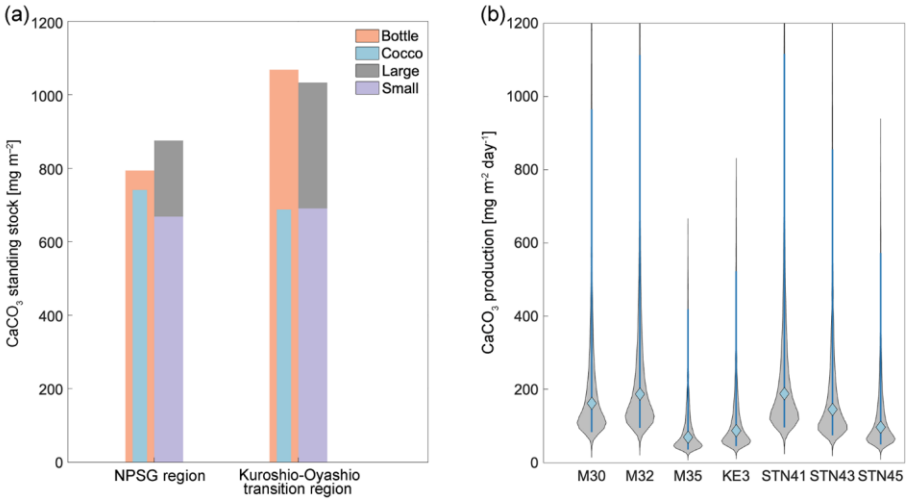
higher in the NPSG region (91 ± 30 %) than in the Kuroshio-Oyashio transition region (~~65 %~~70 ± 24 %)~~ ~~; Fig. 6b)~~(~~Fig. 6a~~),

demonstrating the vital role of coccolithophores in $CaCO_3$ production, particularly in oligotrophic ocean waters.



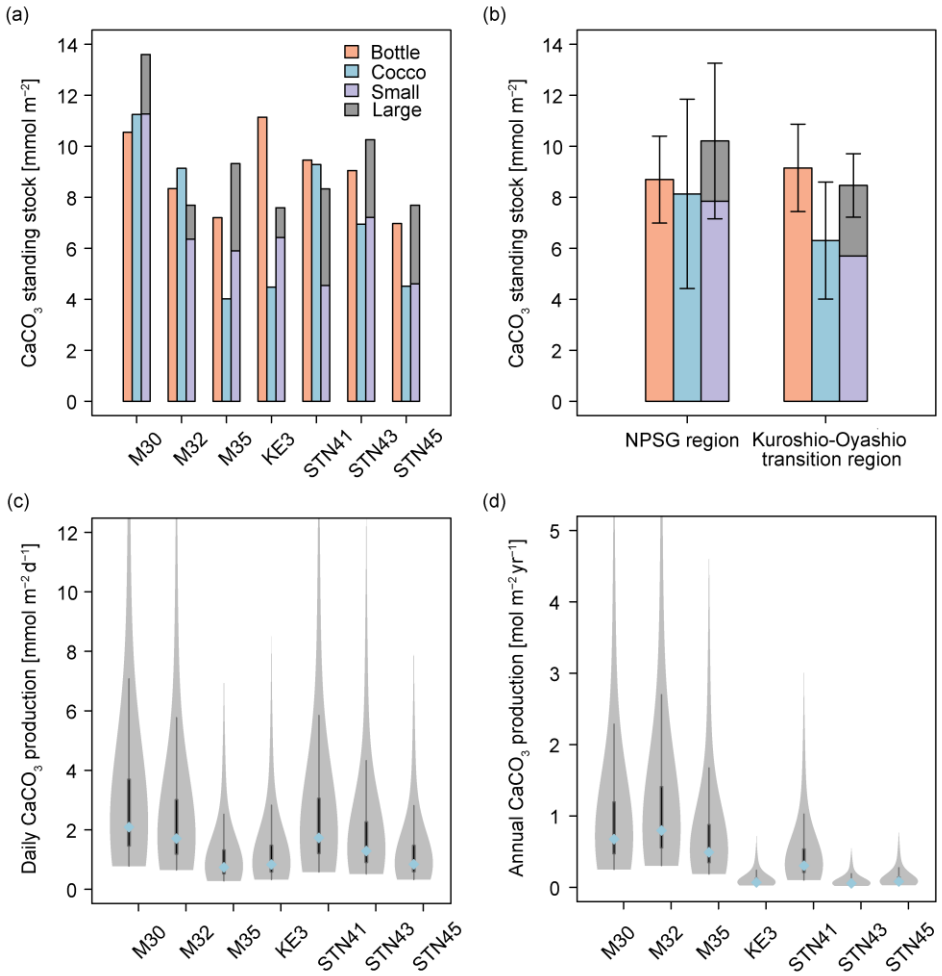


Fig. 6. Calcium carbonate ($CaCO_3$) standing stock in the euphotic zone estimated from Niskin bottle particulate inorganic carbon (PIC), total calcite (Cocco) and size-fractionated (large and small fractions indicate > 51 and 1–51 µm, respectively) PIC concentrations (a) at each sampling station and (b) in the North Pacific Subtropical Gyre (NPSG) and Kuroshio-Oyashio

transition regions; (c) $CaCO_3$ production by coccolithophores in the euphotic zone at indicated sampling stations in June-July 2022; (d) annual $CaCO_3$ production corrected for seasonal bias using satellite-derived PIC concentrations. In (c) and (d), the blue diamond marks the median value, while the shaded area displays the probability density of the estimates. The grey lines denote the 25% and 75% quartiles.

~~**Fig. 6.** (a) Calcium carbonate ($CaCO_3$) standing stock in the upper 150 m water column estimated from Niskin bottle particulate inorganic carbon (PIC), total calcite (Cocco) and size-fractionated (large and small fractions indicate > 51 and 1–51 μm, respectively) PIC concentrations in the North Pacific Subtropical Gyre (NPSG) and Kuroshio-Oyashio transition regions; (b) $CaCO_3$ production by coccolithophores in the upper 150 m water column at indicated sampling stations in June-July 2022.~~

Total $CaCO_3$ standing stock derived from in situ pump samples ranged from 7.6 to 13.6 mmol m^{-2}~~820 to 1,300 mg m^{-2}~~, averaging 10.2 ± 3.1 mmol m^{-2}~~880 mg m^{-2}~~ in the subtropical gyre and 8.5 ± 1.2 mmol m^{-2}~~1,030 mg m^{-2}~~ in the transition region~~, consistent with results from Niskin bottle samples~~. The $CaCO_3$ standing stock of the ~~SSF~~ small PIC ranged from 4.5 to 11.3 mmol m^{-2}~~514 to 904 mg m^{-2}~~ and accounted for 71 ± 12 %~~71 %~~ of the total standing stock in the entire research domain (Fig. 6a).

Given that coccolithophores have a turnover time of 0.7–10 days (Krumhardt et al., 2017; Ziveri et al., 2023), $CaCO_3$ production rate in the ~~upper 150 m of the water column~~euphotic zone ranged from ~~70 to 190 mg m^{-2} d^{-1}~~0.8 to 2.1 mmol m^{-2} d^{-1} during the sampling period (Fig. ~~6b~~6c). Generally, the coccolithophore $CaCO_3$ production was comparable in the subtropical gyre and the Kuroshio-Oyashio transition region, averaging 1.5 ± 0.7~~140 ± 62~~ and 1.2 ± 0.4 mmol m^{-2} d^{-1}~~130 ± 46 mg m^{-2} d^{-1}~~, respectively. Coccolithophore $CaCO_3$ production in the euphotic zone was maximal at station ~~STN41~~ M30~~where it reached 190 mg m^{-2} d^{-1}, corresponding to the maximum coccosphere cell concentration (Fig. 2e).~~ and ~~The~~ the lowest coccolithophore $CaCO_3$ production ~~of 70 mg m^{-2} d^{-1}~~was observed at station M35 ~~in the NPSG region~~.

4 Discussion

4.1 Contribution of coccolithophore calcite to PIC

In this study, bottle- and pump-derived PIC concentrations generally agreed with each other (Fig. 3), and both were on the

same order of magnitude as suspended PIC concentrations measured in the Atlantic, Indian and Pacific Oceans (Beaufort et al., 2008; Barrett et al., 2014; Lam et al., 2015, 2018; Maranón et al., 2016). Coccolithophore calcite concentrations showed a significant positive correlation with PIC concentrations ($r^2 = 0.52$, $p < 0.01$, n = 40; Fig. 7a), highlighting the major contribution of coccospheres and detached coccoliths (68 %) to total $CaCO_3$ in the upper 300 m of the water column. This is consistent with findings from the eastern North Pacific Ocean where coccolithophores dominate $CaCO_3$ production (Ziveri et al., 2023). It is noteworthy that detached coccolith concentrations of *E. huxleyi, U. tenuis* and *Syracosphaera* spp. showed a significant positive relationship with their coccosphere cell concentrations (Fig. 7b–d), indicating that those detached particles were likely shed by cells as part of the dynamic calcification process, during which coccoliths are continuously produced and released (Johns et al., 2023). However, other potential sources and processes, such as advection, cell disintegration from viral lysis and grazing, fecal pellets, or the dissolution associated with microbial respiration could also contribute to the observed detached coccolith concentrations (Subhas et al., 2022; Vincent et al., 2023; Dean et al., 2024). Coccolith production and shedding vary among species. Fast-growing species like *E. huxleyi* produce and shed coccoliths rapidly during exponential growth phases, whereas other species exhibit different patterns, which are influenced by their distinct physiological and ecological characteristics (Johns et al., 2023).

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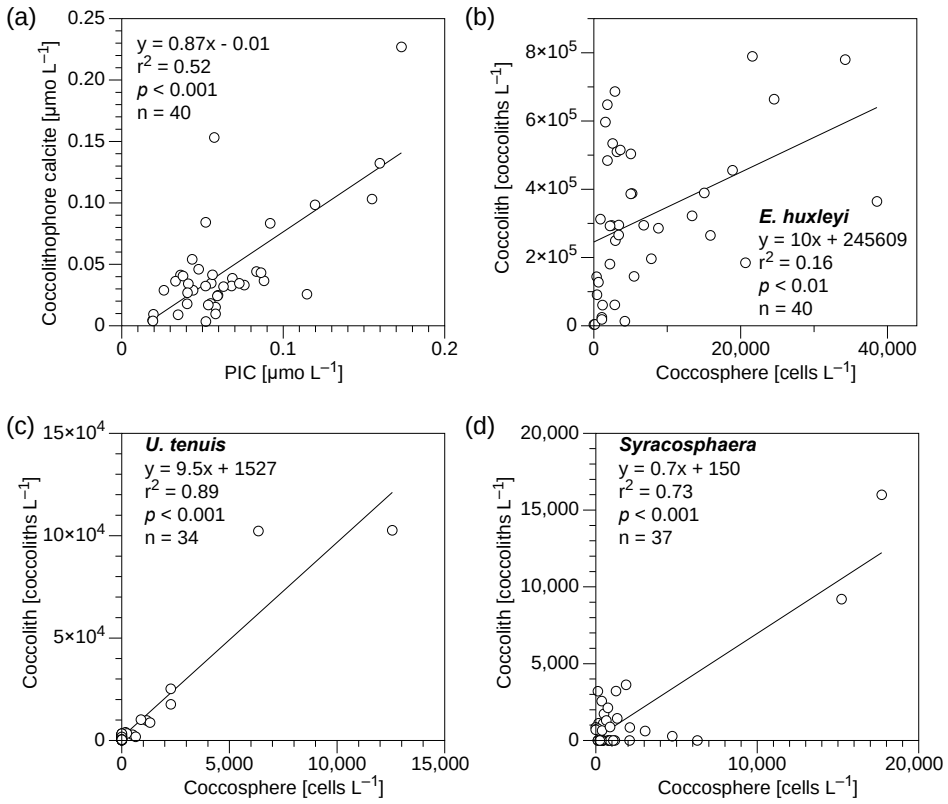

Fig. 7. Relationship of (a) coccolithophore calcite (coccospheres and detached coccoliths) vs particulate inorganic carbon (PIC) concentrations and (b–d) detached coccolith vs coccosphere cell concentrations for (b) *Emiliania huxleyi*, (c) *Umbellosphaera tenuis* and (d) *Syracosphaera* spp. in the upper 300 m water column in the study area. Equations describing the fitted straight lines are also shown.

The less abundant (<3 %) species such as *C. leptoporus* and *O. fragilis* also made a large contribution to calcite concentrations, accounting for 21 % and 12 % of the coccolithophore calcite concentration in the NPSG region and the

Kuroshio-Oyashio transition region, respectively (Fig. 5). It has been reported that despite the relatively low numeric abundance (<2 %), some larger species of the coccolithophore community such as *C. leptoporus*, *Helicosphaera carteri* and *Coccolithus pelagicus* may account for most of the coccolithophore $CaCO_3$ flux to the deep ocean (Rigual Hernández et al., 2020). Some rare coccolithophore species with high coccolith and coccosphere cell concentrations have also been identified as important contributors to both upper-ocean calcite production (Daniels et al., 2016) and deep-sea calcite fluxes (Ziveri et al., 2007). Thus, larger and less abundant coccolithophore species can play an important role in $CaCO_3$ production and export.

Higher $CaCO_3$ standing stock in the euphotic zone of the Kuroshio-Oyashio transition region (Fig. 6a) is consistent with satellite observations suggesting that higher surface PIC concentrations occur at high latitudes (Balch et al., 2005; Berelson et al., 2007). In the present study, however, the relative contribution of coccolithophores to the $CaCO_3$ standing stock was higher in the NPSG region (~91 %) than in the Kuroshio-Oyashio transition region (~70 %) (Fig. 6a). To date, most studies estimated $CaCO_3$ standing stocks using satellite-derived data, which might be challenging to use in subtropical gyres where the DCM depth usually lies below 100 m (Cornec et al., 2021). In these oligotrophic oceans with low productivity, a subsurface PIC maximum can develop within the euphotic zone, and the highly variable subsurface PIC concentrations are poorly reflected by satellites, potentially limiting the ability to fully capture coccolithophore contributions.

In oligotrophic ocean gyres, subsurface $CaCO_3$ production could still occur even if surface PIC is low (Balch et al., 2018). Along our studied transect, maximum coccolithophore abundances increased about twofold from the subtropical gyre to the transition region (Fig. 2g), while a much smaller difference was found in the integrated coccolithophore $CaCO_3$ between the two regions (Fig. 6a). This suggests that subsurface coccolithophore $CaCO_3$ contributed substantially to the total upper water column PIC concentration in the NPSG region. Coccolithophore groups were diverse in the subtropical gyre, including some rare but relatively large and heavily calcified species that contribute significantly to $CaCO_3$ production. In the Southern Ocean, coccolithophores contribution to the annual $CaCO_3$ export is highest in waters with low algal biomass accumulations (Rigual Hernández et al., 2020). Given that low surface PIC regions (<0.1 mmol m^{-3}) occupy about 87 % of the global ocean surface (Ziveri et al., 2023), our data highlight the notable contribution of these regions to global coccolithophore $CaCO_3$ production.

Size-fractionated PIC concentrations showed a smaller contribution of coccolithophores to the $CaCO_3$ standing stock in the Kuroshio-Oyashio transition region (67 ± 13 %) than in the NPSG region (76 ± 11 %) (Fig. 6b). This pattern is consistent with

that observed in the eastern North Pacific Ocean (Fig. 8), which suggests that the contribution of small PIC to $CaCO_3$ standing stock is lower in the subpolar gyre (65 %) than in the subtropical gyre (84 %). In other words, the contribution of large size fraction PIC (e.g., zooplanktonic foraminifera, pteropods and heteropods) to $CaCO_3$ standing stock is higher in the subpolar gyre (35 %) than in the subtropical gyre (16 %) of the eastern North Pacific Ocean (Ziveri et al., 2023). Betzer et al. (1984) reported that foraminifera calcite is more abundant in northern regions (north of 42°N) of the western North Pacific. At Ocean Station Papa in the northeast Pacific (50°N, 145°W), model results showed that foraminifera calcite accounts for only 18–30 % of the total $CaCO_3$ production, whereas coccolithophores are the main producer, contributing to 59–77 % of the total $CaCO_3$ production (Fabry, 1989). These findings support our results and suggest that the relatively high contribution of large size fraction PIC in the northern region of the western North Pacific is likely attributed to foraminifera.

In the Atlantic Ocean, coccolithophore calcite fluxes and species richness are higher in subtropical than in temperate waters, which is ascribed to the reduced competition with diatoms in the former (Broerse et al., 2000). Note that a clear latitudinal gradient of diatom biomass was observed along 160°E in the North Pacific Ocean, consistent with findings from phytoplankton pigment analysis and ocean-color satellite observations (Hirata et al., 2011; Sugie and Suzuki, 2017). The distribution of planktic foraminifera in the North Pacific has been linked to phytoplankton productivity and food availability, with higher abundance in the transitional region compared to the subtropical region (Taylor et al., 2018). Based on these findings, we suggest that differences in ecosystem structure among sites modulate the relative contribution of various calcifiers to pelagic PIC production. The higher abundance of non-calcareous phytoplankton (e.g., diatoms) in the transition zone could also reduce coccolithophore biomass via resource competition (Quere et al., 2005; Sinha et al., 2010) and stimulate the growth of foraminifera (Schiebel et al., 2017), resulting in the observed decreased contribution of small coccolithophores to total $CaCO_3$ production. Sediment trap data from the North Pacific also support this pattern, indicating lower fluxes of planktonic foraminifera, organic matter, and biogenic opal in the subtropical region but elevated fluxes in the transitional and subarctic regions (Eguchi et al., 2003).

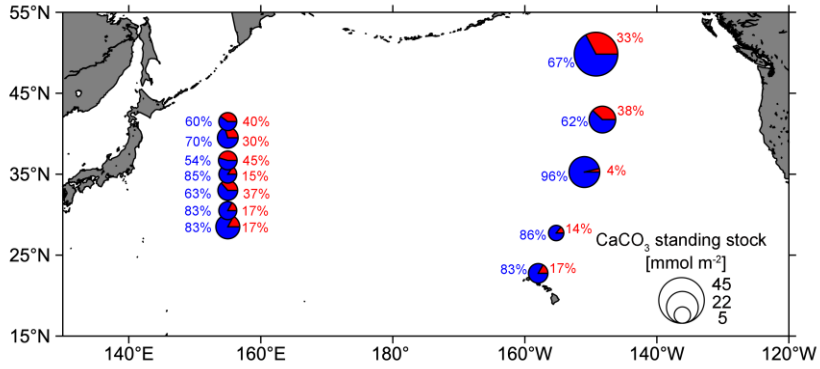

Fig. 8. Pie charts showing the composition of the total calcium carbonate ($CaCO_3$) standing stock in the euphotic zone of the western (this study) and eastern North Pacific Ocean (data from the CDisK-IV cruise; Ziveri et al., 2023). Red represents the standing stock of large size-fractionated (> 51 µm) $CaCO_3$ from this study, and planktonic foraminifera, pteropods and heteropods from the CDisK-IV cruise. Blue represents the standing stock of small size-fractionated (1–51 µm) $CaCO_3$ from this study and coccolithophores from the CDisK-IV cruise.

~~4.1~~4.2 Coccolithophore responses to environmental factors

~~Coccolithophores are an important component of phytoplankton biomass and fill a variety of ecological niches in global oceans. They inhabit different marine environments from oligotrophic to eutrophic, warm to cold, euphotic to aphotic, and stratified to mixed (Balch, 2018). It is of critical importance to evaluate the response of coccolithophore species to different environmental conditions to better understand changes in coccolithophore species diversity, community composition, and their role in the oceanic carbon cycle.~~

Although biogeographical zones of coccolithophores in the North and Central Pacific were identified a couple of decades ago, few studies have investigated coccolithophore distributions in the North Pacific over the recent two decades (Okada and Honjo, 1973; Hagino et al., 2005). In the western North Pacific Ocean, higher diversity and less abundant coccolithophore assemblages were observed in the oligotrophic subtropical gyres, whereas the Kuroshio-Oyashio transition region tended to exhibit a lower

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diversity corresponding to higher PIC and coccolithophore concentrations (Figs. 2 and ~~S3~~S2). This finding is consistent with results from the Atlantic Ocean, and a result of the different survival strategies of various coccolithophore species (Poulton et al., 2017; Balch et al., 2019). Coccolithophores are nutrient stress tolerant and have low~~er~~ iron cell quotas, ~~and are~~ thus generally abundant in the open ocean (Gregg and Casey, 2007; Brun et al., 2015). Prior studies have shown that coccolithophores, particularly *E. huxleyi*, can grow more effectively under low iron conditions than other phytoplankton such as diatoms (Hartnett et al., 2012; Balch, 2018). However, ~~When~~ when nutrients and light are plentiful, the heavy coccoliths of this group of phytoplankters pose a selective disadvantage over diatoms and chlorophytes (Gregg and Casey, 2007). The dominance of coccolithophores in the Great Calcite Belt is primarily driven by their adaptation to low iron levels, which, together with low surface DSi concentrations, limit diatom growth (Balch et al., 2016). The majority of coccolithophore species are K-selected, characterized by relatively slow-growth, large cell size and are more competitive in low-nutrient and well-stratified regions (Brand, 1994), whereas only few r-selected species ~~(~~, such as the fast-growing and small-sized *E. huxleyi* ~~)~~ ~~mainly survive~~thrive in relatively dynamic and nutrient-rich regions (Charalampopoulou, 2011; Brun et al., 2015; O'~~b~~Brien et al., 2016). In the present study, the most abundant and widely distributed coccolithophore species was *E. huxleyi*, which showed increasing abundance northward along the study transect (Fig. 4). This is consistent with prior observations demonstrating that *E. huxleyi* is the most abundant coccolithophore species in the subarctic, subantarctic and bordering transitional regions (Saavedra-Pellitero et al., 2014).

According to the RDA results, environmental variables accounted for 47.6 % of the total variation in coccolithophore community composition (Fig. ~~7a~~9a). The first two RDA axes suggested that there were significant spatial differences in the coccolithophore community across depths and regions (Fig. S43). In the tropical and subtropical Atlantic Ocean, coccolithophore communities exhibit greater variability vertically within the water column than horizontally, at spatial scales of hundreds to thousands of kilometers (Poulton et al., 2017). Moreover, distinct species distributions are identified based on the depth zones (upper euphotic, lower euphotic, and subeuphotic zones), which reflect the lifestyle of the species (Poulton et al., 2017; Balch, 2018). In the NPSG region, our results also reveal a distinct vertical distribution pattern (Fig. 4), which may have been driven by factors such as light availability, temperature, and nutrient levels. These environmental variables likely contribute to the physiological diversity of coccolithophores. A shift in dominant species occurred from *U. tenuis* and *E. huxleyi*

in the NPSG region to *Syracosphaera* spp. and *E. huxleyi* in the transition region (Fig. 5). This is consistent with the prior observations of Balch et al. (2019). Correspondingly, hierarchical partitioning analysis showed that depth and latitude had a significant effect on coccolithophore community variation ($p < 0.05$). Other environmental factors, such as temperature, salinity, Chl *a* and TA ~~were~~ also ~~important influencing~~ influenced the coccolithophore community (Fig. ~~7b~~9b).

Based on Spearman's correlation analysis, coccolithophore abundance showed a significant positive relationship with temperature, $\Omega_{calcite}$ and pH, and a significant negative relationship with depth, DIC and macro-nutrient concentrations, especially for *D. tubifera*, *U. tenuis* and HOL that are more sensitive to environmental factors (Fig. ~~7e~~9c). The positive correlation with temperature is consistent with field observations and model simulations pointing to a general trend of increasing coccolithophore abundance in the context of global warming (Rivero-Calle et al., 2015; Rousseaux and Gregg, 2015). More abundant species like *E. huxleyi* and *Syracosphaera* spp., however, only showed a highly positive correlation with depth, latitude and Chl *a* concentration, suggesting that these species are more adaptable to varying environmental conditions (Schlüter et al., 2014). ~~Using random forest analysis, we also determined that the best predictors of coccolithophore abundance were Chl *a* concentration, depth and $\Omega_{calcite}$ ($p < 0.01$; Fig. 7d).~~ In the Atlantic Ocean, *E. huxleyi* has been observed to exhibit an increasing relative abundance with increasing latitude (Balch et al., 2019; Holligan et al., 2010; Poulton et al., 2017). Unlike many other species, *E. huxleyi* has a widespread distribution attributed to its ability to adapt to diverse environments through both phenotypic plasticity and genetic selection (Lohbeck et al., 2012; Rickaby et al., 2016b; Taylor et al., 2017). Our results indicate that less abundant species, such as *C. leptoporus* and *O. fragilis*, also contributed to coccolithophore calcite concentrations (Fig. 5). Their calcification is species-specific, predominantly driven by inherent biological traits, including cell shapes, coccolith types, and architectural variations, which are conservative features of coccolithophore biology (Rickaby et al., 2016a). However, the weak correlation of *C. leptoporus* and *O. fragilis* with environmental factors might be due to their low abundance. Overall, our study highlights the significant influence of depths and latitude on coccolithophore community composition, emphasizing the complex interplay between biotic and abiotic factors.

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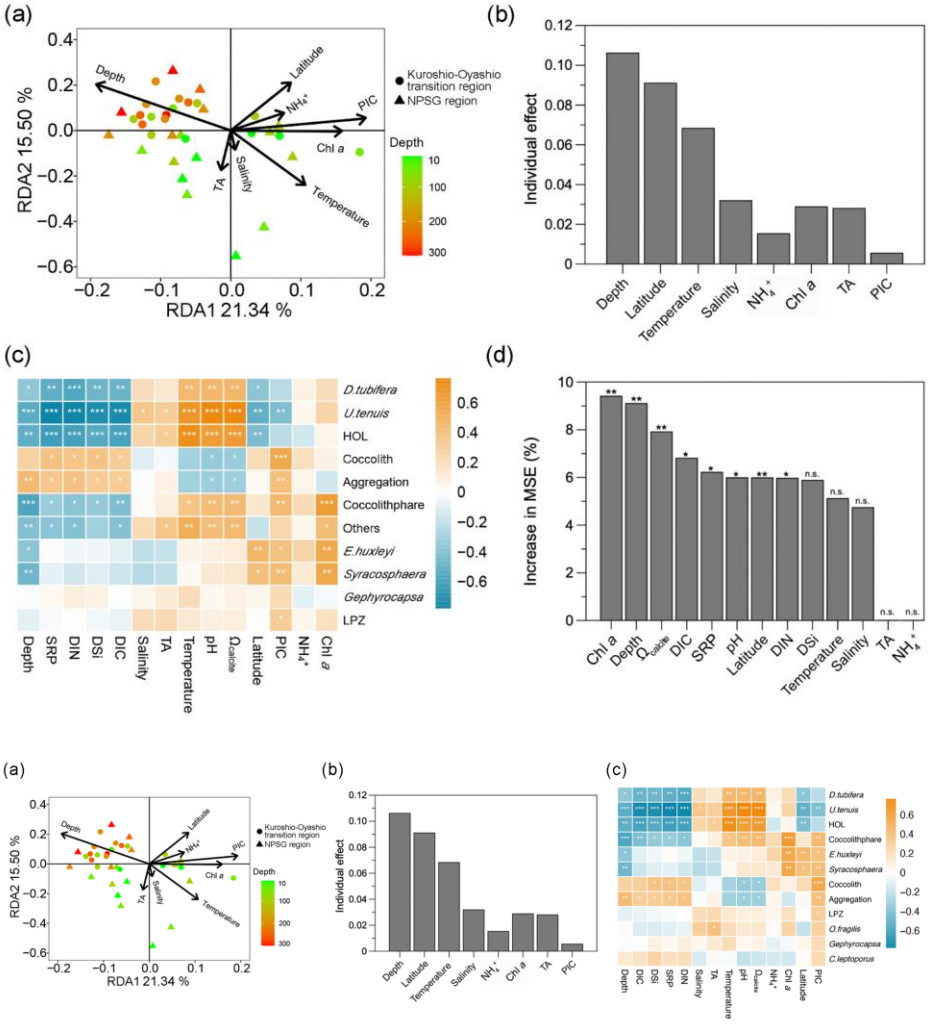

Fig. ~~79~~. (a) Redundancy analysis (RDA) diagram illustrating the relationship between the coccolithophore community and

environmental factors; (b) independent contribution of each environmental factor to coccolithophore community variation using hierarchical partitioning-based canonical analysis; (c) correlations between coccolithophore groups and environmental factors with color gradients denoting the significance of the Spearman's correlation coefficient r. Asterisks represent the statistical significance (***$p < 0.001$, **$p < 0.01$, *$p < 0.05$)~~; (d) random forest mean predictor importance, i.e., the percentage of increase in the mean variance error (MSE) of environmental factors on coccolithophore abundance (**$p < 0.01$, *$p < 0.05$, n.s., non-significant)~~. Chl *a*: chlorophyll *a*, DIC: dissolved inorganic carbon, TA: total alkalinity, $\Omega_{calcite}$: saturation state with respect to calcite, PIC: particulate inorganic carbon, DIN: dissolved inorganic nitrogen (nitrate plus nitrite), NH_4^+: ammonium, SRP: soluble reactive phosphate, DSi: dissolved silicate, HOL: holo-coccolithophores and LPZ: lower euphotic zone species *Florisphaera profunda* and *Algirosphaera robusta*.

~~**4.2 Contribution of coccolithophore calcite to PIC**~~

~~In this study, bottle- and pump-derived PIC concentrations generally agreed with each other (Fig. 3), and both were comparable to and of the same order of magnitude as the suspended PIC concentrations detected in the Atlantic, Indian and Pacific Oceans (Beaufort et al., 2008; Barrett et al., 2014; Lam et al., 2015, 2018; Maranón et al., 2016). Coccolithophore calcite concentrations showed a significant positive correlation with PIC concentrations ($r^2 = 0.75$, $p < 0.01$, n = 40; Fig. 8a), highlighting the major contribution of coccospheres and detached coccoliths (68 %) to the total $CaCO_3$ in the upper 300 m of the water column. This is consistent with findings from the eastern North Pacific Ocean where coccolithophores dominate $CaCO_3$ production (Ziveri et al., 2023). It is noteworthy that detached coccolith concentrations of *E. huxleyi*, *U. tenuis* and *Syracosphaera* spp. showed a significant positive relationship with their coccosphere cell concentrations (Fig. 8b–d), indicating that those detached particles were likely to have originated from living cells.~~

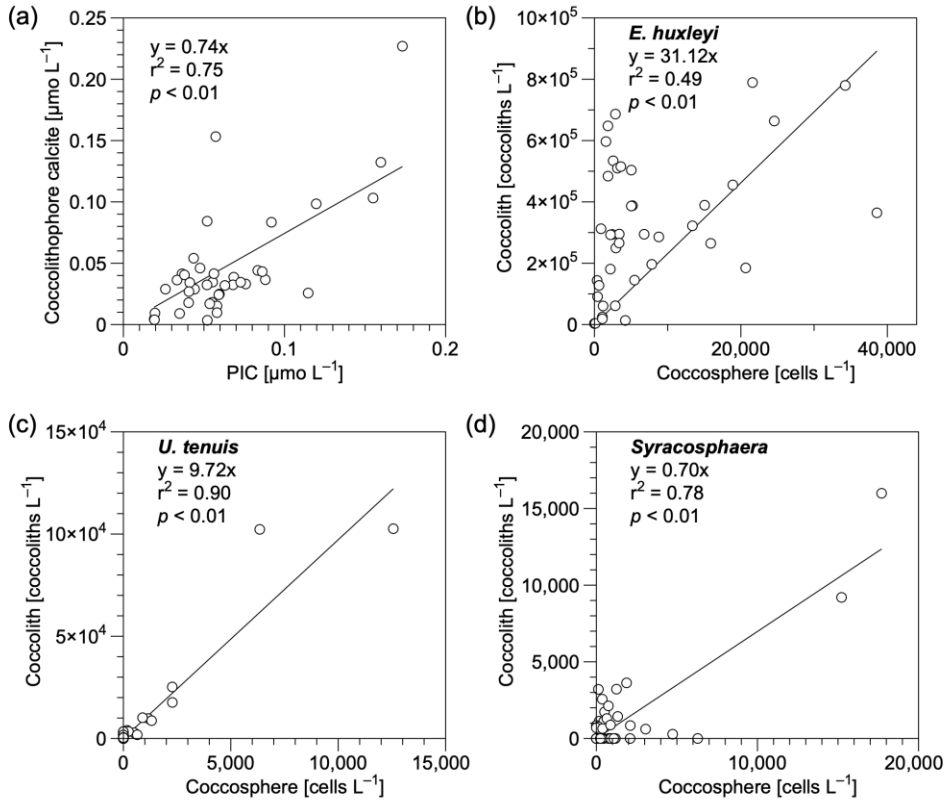

Fig. 8. Relationship of (a) coccolithophore calcite vs particulate inorganic carbon (PIC) concentrations and (b–d) detached coccolith vs coccosphere cell concentrations for (b) *Emiliania huxleyi*, (c) *Umbellosphaera tenuis* and (d) *Syracosphaera* spp. in the upper 300 m water column at the study site. Equations describing the fitted straight lines are also shown.

Abundant coccolithophore groups, including *E. huxleyi*, *Syracosphaera* spp., LPZ species and aggregation, showed a significant positive relationship with PIC concentration (Fig. 7e), but less abundant species like *C. leptoporus* and *O. fragilis* also made a large contribution to calcite concentrations. It has been reported that despite the relatively low abundance (< 2 %)

of the coccolithophore community, some larger species such as *C. leptoporus*, *Helicosphaera carteri* and *Coccolithus pelagicus* could account for most of the coccolithophore $CaCO_3$ flux to the deep ocean (Rigual Hernández et al., 2020). Some rare coccolithophore species with high coccolith and coccosphere cell concentrations have also been identified as important contributors to both upper-ocean calcite production (Daniels et al., 2016) and deep-sea calcite fluxes (Ziveri et al., 2007). Thus, although *E. huxleyi* is one of the most abundant species in the ocean, larger coccolithophore species can also play an important role in $CaCO_3$ export.

Higher $CaCO_3$ standing stock in the upper 150 m of the Kuroshio-Oyashio transition region (Fig. 6a) is consistent with satellite observations suggesting that higher surface PIC concentrations occur at high latitudes (Balch et al., 2005; Berelson et al., 2007). In the present study, however, the relative contribution of coccolithophores to the $CaCO_3$ standing stock was higher in the NPSG region (~91 %) than in the Kuroshio-Oyashio transition region (~65 %) (Fig. 6a). To date, most studies estimated $CaCO_3$ standing stock using satellite-derived data, which might be challenging to use in subtropical gyres where the DCM depth usually lies below 100 m. Because coccolithophore $CaCO_3$ is largely produced in the lower layer of the euphotic zone and is thus difficult to detect by satellites, coccolithophore contributions could be underestimated.

In oligotrophic ocean gyres, subsurface $CaCO_3$ production could still occur even if surface PIC is low (Balch et al., 2018). Along our studied transect, maximum coccolithophore abundances increased about twofold from the subtropical gyre to the transition region (Fig. 2e), while a much smaller difference was found in the integrated coccolithophore $CaCO_3$ between the two regions (Fig. 6a). This suggests that subsurface coccolithophore $CaCO_3$ contributed substantially to the total upper water column PIC flux in the NPSG. Coccolithophore groups were diverse in the subtropical gyre, where the environmental conditions favor slow-growing, large and heavy species, which account for a large fraction of $CaCO_3$ production. In the Southern Ocean, coccolithophores contribute to the highest annual $CaCO_3$ export in waters with low algal biomass accumulations (Rigual Hernández et al., 2020). Given that low surface PIC regions (< 0.1 mmol m^{-3}) occupy ~87 % of the global ocean surface (Ziveri et al., 2023), our data suggest that coccolithophore $CaCO_3$ production in subsurface waters constitutes an important part of global pelagic $CaCO_3$ production.

Size-fractionated PIC concentrations showed a smaller contribution of coccolithophores to the $CaCO_3$ standing stock in the Kuroshio-Oyashio transition region (67 %) than in the NPSG region (76 %) (Fig. 6a), consistent with observations from the

eastern North Pacific Ocean (Fig. 9). The contribution of LSF PIC (e.g., zooplanktonic foraminifera, pteropods and heteropods) to $CaCO_3$ standing stock is higher in the subpolar gyre (35 %) than in the subtropical gyre (16 %) of the eastern North Pacific Ocean (Ziveri et al., 2023). Betzer et al., (1984) reported that foraminifera calcite is more abundant in northern regions (north of 42°N) of the western North Pacific. At Ocean Station Papa in the northeast Pacific (50°N, 145°W), model results showed that foraminifera calcite accounts for only 18–30 % of the total $CaCO_3$ production, whereas coccolithophores are the major producer, contributing up to 59–77 % of the total $CaCO_3$ production (Fabry, 1989). In the Atlantic Ocean, coccolithophore calcite fluxes and species richness are higher in subtropical than in temperate waters, which is ascribed to the reduced competition with diatoms in the former (Broerse et al., 2000). Based on these findings, we suggest that differences in ecosystem structure among sites modulate the relative contribution of various calcifiers to pelagic PIC production. The higher abundance of non-calcareous phytoplankton (e.g., diatoms) in the transition zone could also reduce coccolithophore biomass via resource competition (Quere et al., 2005; Sinha et al., 2010) and stimulate the growth of foraminifera (Schiebel et al., 2017), resulting in the observed decreased contribution of small coccolithophores to total $CaCO_3$ production.

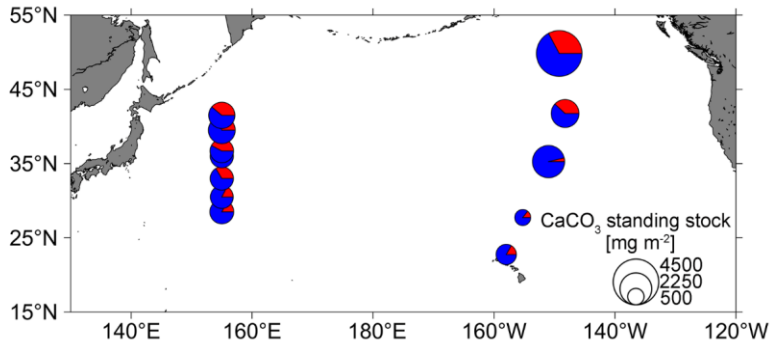

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Fig. 9. Pie charts showing the composition of the total calcium carbonate ($CaCO_3$) standing stock in the upper 150 m of the water column in the western (this study) and eastern North Pacific Ocean (data from the CDisK-IV cruise; Ziveri et al., 2023). Red represents the standing stock of large size-fractionated (> 51 µm) $CaCO_3$ from this study, and planktonic foraminifera,

~~pteropods and heteropods from the CDisK-IV cruise. Blue represents the standing stock of small size-fractionated (1–51 μm) $CaCO_3$ from this study and coccolithophores from the CDisK-IV cruise.~~

4.3 $CaCO_3$ production compared with the eastern North Pacific

While ^{14}C incubations can provide a direct and precise measurement of in situ calcification rates, the calculation method we used offers a practical approach to convert concentration data into production estimates using turnover time (Graziano et al., 2000; Ziveri et al., 2023). This approach has limitations, particularly due to uncertainties in the estimation of coccolithophore calcite, which relies on cell counts and a morphometric-based calcite estimation method, with potential errors reaching up to 50% (Young and Ziveri, 2000; Sheward et al., 2024). The calculation of production rates introduces further uncertainty, as it depends on the coccolithophore calcite standing stock and a broad range of turnover time estimates. Despite these challenges, this method produces reasonable results that are comparable to field observations and thus helps fill a critical data gap in the study region.

Our results indicate that the coccolithophore $CaCO_3$ production ranged from 0.8 to 2.1 mmol m^{-2} d^{-1} during the sampling period, align with globally reported in situ calcification rates and are consistent with observations from the North Atlantic subtropical region (Poulton et al., 2006; Daniels et al., 2018). Although station M30 is located in the oligotrophic NPSG region, it exhibits the highest coccolithophore $CaCO_3$ production in the euphotic zone of the study area (Fig. 6c). This is primarily because of the deepest euphotic zone at this site, reaching up to 219 m, and the relatively high coccolithophore species diversity. While the coccolithophore abundance at station M30 was lower than at other stations, the less abundant but larger species play an important role in contributing to the $CaCO_3$ production at this site.

Using a seasonal-correction method (Ziveri et al., 2023), the average coccolithophore $CaCO_3$ production ~~above 150 m~~in the euphotic zone was estimated to be ~~0.35~~0.4 ± 0.3 mol m^{-2} yr^{-1} for the entire research domain~~, which agrees well with the global estimate of 0.4 mol m^{-2} yr^{-1} (Balch et al., 2007) and model result of 0.3 mol m^{-2} yr^{-1} in the North Pacific (Hopkins and Balch, 2018)~~. In particular, this production was ~~0.62~~0.66 ± 0.2 mol m^{-2} yr^{-1} in the subtropical gyre and ~~0.14~~0.13 ± 0.1 mol m^{-2} yr^{-1} in the Kuroshio-Oyashio transition region (Fig. ~~S2b~~6d). However, the latter is much lower than the recent estimate of 0.9–1.0 mol m^{-2} yr^{-1} by Ziveri et al. (2023) based on data from the transition zone and subpolar gyre in the eastern North Pacific Ocean

using the same seasonal-correction method.

Several factors may lead to the above discrepancy. First, $CaCO_3$ production rate on the present study was estimated based only on coccolithophores, whereas estimates by Ziveri et al. (2023) also included the contribution from planktonic foraminifera, pteropods and heteropods. Second, in the CDisK-IV cruise to the eastern North Pacific Ocean, coccolithophore calcite concentrations were significantly higher than suspended seawater PIC concentrations collected by in situ pumps in the transition zone and subpolar gyre (Fig. ~~S5~~S4; Dong et al., 2019, 2022). Calculations based on these apparently inconsistent data may result in an overestimation of actual $CaCO_3$ production. Third, high spatial and seasonal variations in PIC production might occur between the two oceanic environments~~, in particular the high dynamics~~. Particularly, the complex environmental gradients and variability in the ~~Kuroshio-Oyashio~~ transition regions between the subtropical and subpolar gyres may have ~~essentially altered~~skewed the coccolithophore community and associated $CaCO_3$ production.

Overall, our results suggest that the calibration of satellite-derived PIC should be unreliable. There was a significant positive relationship between surface coccolithophore calcite concentrations and satellite-derived PIC concentrations ($r^2 = 0.84$; $p < 0.01$; Fig S5a), which implies the latter can reflect the distribution tendency of the former but not the true values, because satellite-derived PIC in high latitude areas is likely overestimated. Over the entire euphotic zone, our results indicate no correlation between satellite-derived PIC concentrations and actual PIC production, a finding that is also highlighted by Ziveri et al. (2023), in which the linear correlation is primarily driven by the highest data value (Fig. S5b). More in situ calcification rates determined by ^{14}C incubations, as well as direct measurements of coccolithophore turnover time, are required to reduce uncertainties in the estimation of PIC production and the assessment of the oceanic $CaCO_3$ budget.

~~**4.4 Shallow-water $CaCO_3$ dissolution in the western North Pacific**~~

~~We used a box model to calculate the magnitude of metabolic $CaCO_3$ dissolution in shallow waters of the western North Pacific Ocean (Eqs. 3 and 4; Dong et al., 2019, 2022). Focusing on $CaCO_3$ produced by coccolithophores and assuming that production occurs in the euphotic zone, we applied the model to stations STN43 and STN45 because the calculated metabolic calcite saturation horizon varied from 100 m to 150 m in the Kuroshio-Oyashio transition region. The results suggest that $CaCO_3$ might start to dissolve in setting marine particles after sinking out of the euphotic zone at the two stations, despite apparently oversaturated ambient conditions (Fig. S6). Three different particle sinking rates (1, 10 and 100 m d^{-1}) were used in our model~~

calculations (Fig. 10a). To examine the possible influence by lateral transport around the Kuroshio Extension, our data in the upper 1200 m were compared with those obtained during the CDisK-IV cruise (Fig. S5d–f). There was no significant difference in PIC distribution patterns between western and eastern basins across 27°N–42°N in the North Pacific Ocean (one-way ANOVA, $p > 0.05$), with the latter having relatively limited water parcel transport in the horizontal direction. We thus contend that in our first order estimation, the effect of lateral transport on PIC distributions could be considered negligible (Dong et al., 2019; Subhas et al., 2022).

Assuming that all coccolithophore production is exported out of the euphotic zone, our model results show that the PIC sinking flux using the 10 m d^{-1} sinking rate agrees well with data obtained from sediment traps deployed in July 1999 at a 40°N station (Honda et al., 2002) near our research domain (Fig. 10a). At this rate, the R_{TA} driven by Ω_{met} showed a vertical distribution pattern similar to TA^*–chlorofluorocarbon (CFC) age-based R_{TA} in the entire North Pacific Ocean (Feely et al., 2002), with both displaying a maximum at a density of 26.58 kg m^{-3} corresponding to a water depth of 300 m (Fig. 10b). In addition, our R_{TA} varied within a magnitude comparable to Alk^*–transit time distribution (TTD) and ^{14}C age-based R_{TA} values. Instead, those driven by ambient Ω were essentially equal to zero with densities < 27.0 kg m^{-3} and became considerable with further increasing densities below 600 m. Therefore, our model results indicate that shallow-water $CaCO_3$ dissolution indeed occurs in the western North Pacific Ocean mainly as a result of metabolic acidification in the particulate microenvironment.

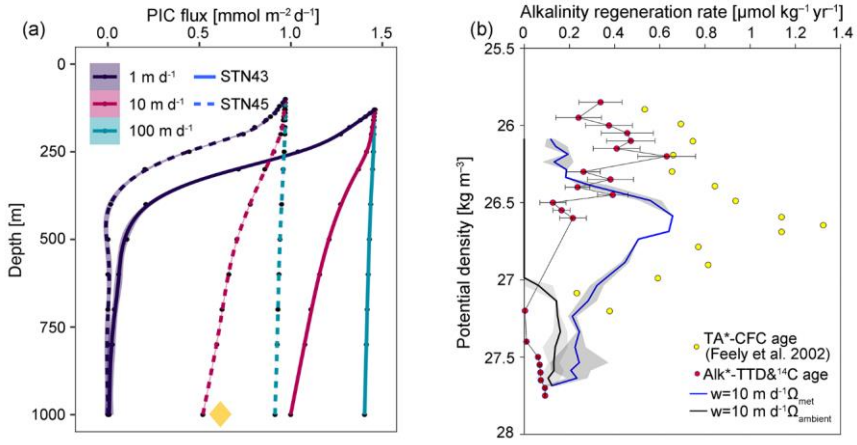

Fig. 10. (a) Particulate inorganic carbon (PIC) fluxes above 1,000 m at stations STN43 and STN45 generated by a box model using three different sinking rates (1, 10 and 100 m d^{-1}); (b) vertical distribution of the alkalinity regeneration rate (R_{TA}) at stations STN43 and STN45 generated by a 1-D model driven by metabolism-altered calcite saturation state (Ω_{met}) and ambient calcite saturation state ($\Omega_{ambient}$) at a sinking rate of 10 m d^{-1}. Lines indicate R_{TA} binned and averaged at 0.05 kg m^{-3} potential density intervals. In (a), the yellow diamond marks the PIC flux at 1,000 m at station 40°N, 165°E in the western North Pacific Ocean (Honda et al., 2002; Fig. 1a). In (b), TA^*-chlorofluorocarbon (CFC) age represents the excess alkalinity-based R_{TA} from (Feely et al., 2002) and Alk^*-TTD&^{14}C age represents the excess alkalinity-based R_{TA} using transit-time distribution (TTD) ages and ^{14}C ages combined with Alk^* (Key et al., 2004; Gebbie and Huybers, 2012; Carter et al., 2014; Jeansson et al., 2021; Sulpis et al., 2021). Details of the estimation method are provided in the Supplementary Materials.

This finding is consistent with that obtained in the eastern North Pacific Ocean (Subhas et al., 2022), which suggests widely occurring shallow-water dissolution throughout the entire North Pacific Ocean associated with organic carbon respiration. The maximum R_{TA} value of 0.66 μmol kg^{-1} yr^{-1} at a density of 26.58 kg m^{-3} in our study aligns exactly with that of 0.6 μmol kg^{-1} yr^{-1} at a density of 26.54 kg m^{-3} in the eastern North Pacific Ocean (Subhas et al., 2022). It is noteworthy that the R_{TA}

~~discrepancy between this and previous studies might be ascribed to various dissolution mechanisms (Fig. 10b), since our model only accounted for coccolithophore calcite excluding other calcifying plankton groups such as those producing aragonite and high-Mg calcite. Jansen and Wolf-Gladrow (2001) also suggested that dissolution within zooplankton guts may account for 25 % of the shallow-water dissolution signal.~~

5 Conclusions

We have demonstrated that coccolithophore abundance~~s~~ and species composition~~s~~ ~~had~~ had distinct geographic and vertical distribution patterns, with *U. tenuis* ~~dominated~~ dominating in the NPSG region while *E. huxleyi* and *Syracosphaera* spp. in the Kuroshio-Oyashio transition region. The environmental variables that best described varying coccolithophore communities were depth and latitude. Calcite derived from coccolithophores contributed ~~on~~ 79 ± 27 %~~average ~ 76 %~~ of the PIC standing stocks ~~above 150 m~~in the euphotic zone, with a relatively greater contribution in the subtropical gyre than in the transition region. Less abundant (<3 %) species such as *C. leptoporus* and *O. fragilis* also made a large contribution of 21 % and 12 % to the coccolithophore calcite concentration in the NPSG region and the Kuroshio-Oyashio transition region, respectively. ~~Our results suggests that coccolithophore $CaCO_3$ production was ~ 5-fold higher in the former than in the latter, which highlights the importance of coccolithophores in oligotrophic environments.~~ During the sampling period, coccolithophore $CaCO_3$ production ranged from 0.8 to 2.1 mmol m^{-2} d^{-1} in the entire research domain, averaging 1.5 ± 0.7 mmol m^{-2} d^{-1} in the subtropical gyre and 1.2 ± 0.4 mmol m^{-2} d^{-1} in the Kuroshio-Oyashio transition region.

~~This study also inferred that extensive $CaCO_3$ dissolution occurs above the ambient calcite saturation horizon, and is primarily driven by the metabolic activity associated with organic carbon respiration.~~ Given the important role of $CaCO_3$ ~~production and dissolution~~dynamics in the marine alkalinity and carbon cycles, ~~additional studies are required that target~~ coccolithophore production at different scales from seasonal to annual and from regional to global needs further examination.~~, as well as processes leading to $CaCO_3$ dissolution in the apparently oversaturated upper ocean.~~

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Data availability. Data for temperature, salinity, coccolithophore cell and coccolith abundances, coccolithophore calcite ~~and ,~~ PIC ~~concentrations~~and nutrients concentrations can be downloaded from the Science Data Bank (https://www.scidb.cn/en/s/i6bMFn~~https://www.scidb.cn/en~~). Satellite-based temperature, Chl *a* and PIC concentration data were obtained from the MODIS-Aqua satellite (https://oceancolor.gsfc.nasa.gov/l3/).

Supplement link.

Author Contributions. YH, ZC, and MD conceived and designed the study. YH, ZS, DF, and JC contributed to data acquisition and analysis. YH, ZS, ZC, and MD wrote the first draft of the manuscript. YH, ZS, ZC, JY, and MD discussed results and edited the paper. All authors read and approved the final version of the manuscript.

Competing interests. The authors declare that they have no conflict of interests.

Disclaimer.

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