



# An extension of the logistic function to account for nonstationary drought losses

Tongtiegang Zhao<sup>1</sup>, Zecong Chen<sup>1</sup>, Yongyong Zhang<sup>2</sup>

5 <sup>1</sup> Southern Marine Science and Engineering Guangdong Laboratory (Zhuhai), School of Civil Engineering, Sun Yat-Sen University, Guangzhou 510275, China

<sup>2</sup> Key Laboratory of Water Cycle and Related Land Surface Processes, Institute of Geographic Sciences and Natural Resources Research, Chinese Academy of Sciences, Beijing 100101, China

*Corresponding author:* Tongtiegang Zhao (zhaottg@mail.sysu.edu.cn)

10

## Highlights:

1. The drought-affected population in mainland China exhibits significant correlation not only with standard precipitation index, but also with time;
2. The nonstationarity of drought losses is effectively characterized by incorporating time into the parameters of the classic  
15 logistic function;
3. The three nonstationary intensity loss functions built upon the logistic function are demonstrated to be a useful tool for drought impact assessment.



20 **Abstract:** While the intensity loss function is fundamental to drought impact assessment, the relationship between drought  
loss and intensity can be nonstationary, i.e., changing as time progresses, owing to socio-economic developments. This paper  
builds three novel intensity loss functions upon the classic logistic function to account for nonstationary drought losses.  
Specifically, the time is explicitly formulated as an explanatory variable and respectively incorporated into the magnitude,  
25 shape and location parameters of the logistic function to derive three nonstationary intensity loss functions. To examine the  
effectiveness, a case study is devised for the drought-affected population by province in mainland China during the period  
from 2006 to 2023. The results highlight the existence of nonstationarity in that the drought-affected population exhibits  
significant correlation not only with standard precipitation index but also with year. The three nonstationary intensity loss  
functions are shown to outperform the classic logistic function and also the linear regression. They present effective  
30 characterizations of observed drought loss in different ways: 1) the nonstationary function with the flexible magnitude  
parameter fits the data by adjusting the maximum drought loss by year; 2) the nonstationary function with the flexible shape  
parameter works by modifying the growth rate of drought loss with intensity; and 3) the nonstationary function with the  
flexible shape parameter acts by shifting the response curves along the axis by year. In general, the nonstationary function  
with the flexible magnitude parameter is shown to be the most promising in terms of high coefficient of determination, low  
Bayesian information criterion and explicit physical meaning. Taken together, the nonstationary intensity loss functions  
35 developed in this paper can serve as an effective tool for drought management.

**Short summary:** The classic logistic function characterizes the stationary relationship between drought loss and intensity.  
This paper incorporates the time into the three parameters of logistic function and derives three nonstationary intensity loss  
functions. The functions are tested through a case study of drought-affected population by province in mainland China  
40 during the period from 2006 to 2023. Overall, the nonstationary intensity loss functions are shown to be a useful tool for  
drought management.



## 1 Introduction

Droughts are one of the most destructive natural hazards (Baez-Villanueva et al., 2024; Van Dijk et al., 2013; Zhang et al., 2022). In general, there exist meteorological, hydrological, agricultural and socio-economic droughts (Mishra and Singh, 2010). Originating from precipitation deficits and high atmospheric evaporative demands, droughts propagate through hydrological processes and eventually impair human beings and natural ecosystems (Gao et al., 2024a; Liu et al., 2024; Zhao et al., 2024a). From 2001 to 2009, the “Millennium Drought” in Southeast Australia amplified median rainfall reduction by up to 4 times in streamflow and reduced irrigated rice and cotton production respectively by 99% and 84% (Van Dijk et al., 2013). The 2012 summertime drought arrived at the Central Great Plains in North America without early warning and caused more than US\$30 billion of economic losses (Hoerling et al., 2014; Yuan et al., 2023). The 2021/22 drought event made 76.2% of the Euro-Mediterranean region under mild drought, 61.4% under moderate drought and 39.4% under severe drought (Garrido-Perez et al., 2024). Under climate change, droughts are expected to not only increase worldwide (Dai, 2011) but also intensify more rapidly (Yuan et al., 2023).

Socio-economic losses are an integral part of droughts in environment management (AghaKouchak et al., 2021; Hoerling et al., 2014; Van Dijk et al., 2013). Although there exist extensive studies on hydroclimatic processes relating to droughts (Entekhabi, 2023; Mishra and Singh, 2010; Wang et al., 2023b; Yang et al., 2024; Zhang et al., 2021), far less attention is paid to socio-economic impacts of droughts (AghaKouchak et al., 2021; Apurv and Cai, 2021; Su et al., 2018). One possible cause is that in situ observations, satellite remote sensing and earth system models generate a vast amount of hydroclimatic data (Hersbach et al., 2020; Pradhan et al., 2022; Zhang et al., 2024, 2021; Zhao et al., 2024b). Plenty of spatial-temporal data facilitate drought investigations at catchment, regional, continental and global scales and in pentad, monthly, seasonal and annual timesteps (Gao et al., 2024b; Ma et al., 2022; Wang et al., 2023a). On the other hand, there are limited data on drought-related socio-economic losses (AghaKouchak et al., 2021). Usually, drought losses have to be collected from statistical yearbooks issued by local and central governments and from survey reports provided by international organizations and commercial services (Chen et al., 2015; Hou et al., 2019).

The intensity loss function, which is also called exposure-response function, plays a critical part in drought impact assessment (AghaKouchak et al., 2021; Qiu et al., 2023; West et al., 2019). On the one hand, the classic logistic function is effective in characterizing the growth of socio-economic loss with drought intensity (Chen et al., 2015; Hou et al., 2019; Todisco et al., 2013). On the other hand, the relationship between socio-economic loss and drought intensity can be nonstationary, i.e., temporally changing, considering that economic growth can increase the exposure to droughts and that infrastructure developments can decrease the vulnerability to droughts (Apurv and Cai, 2021; Haile et al., 2020; Long et al., 2020). In this paper, we build three non-stationary intensity loss functions upon the classic logistic function that represents a stationary intensity loss function. As will be illustrated in the methods and results, the proposed functions tend to capture the non-stationary characteristics of drought-affected population in mainland China. The effects of drought intensity and time on population in different provinces are effectively characterized.



## 2 Methods

### 2.1 Intensity loss function

Drought indices are essential for drought impact assessment (Montanari et al., 2023; Todisco et al., 2013; West et al., 2019).  
80 Among the popular indices are the standardized precipitation index (SPI), the Palmer drought severity index (PDSI), the  
standardized precipitation evapotranspiration index (SPEI) and the standardized runoff index (SRI) (AghaKouchak et al.,  
2021; Apurv and Cai, 2021; Zhao et al., 2024b). The intensity is derived from drought indices (Hao et al., 2017; Mishra and  
Singh, 2010; Su et al., 2018). Since 0 is both the mean and median values of the standard normal distribution, the extent to  
which drought indices falling below 0 indicates the degree of dryness. Thresholds can be employed to identify drought  
85 events (Wang et al., 2023b). For example,  $(-0.99, 0]$  is near normal,  $(-1.49, -1.00]$  is moderately dry,  $(-1.99, -1.50]$  is  
severely dry and  $(-\infty, -2.00]$  is extremely dry. Therefore, drought events can be defined by the combinations of multiple  
indices, e.g., by  $SPI \leq -1.0$ ,  $PDSI \leq -2.0$  and  $SPEI \leq -1.0$  (Su et al., 2018).

Denoting the drought intensity as  $I$ , the intensity loss function is formulated as:

$$L = f(I) \quad (1)$$

in which  $L$  is the drought loss corresponding to the intensity  $I$ . Empirically, there are four important characteristics of  $f(I)$ : 1)  
90 there is minimal loss when there is minimal  $I$ ; 2) there is maximal loss when there is maximal  $I$ ; 3)  $f(I)$  is a monotonically  
increasing function, i.e., drought loss increases with drought intensity; and 4) drought loss initially grows approximately  
exponentially with  $I$ , slows to linear as saturation begins and finally stops at maturity.

The above four characteristics can mathematically be formulated by the renowned logistic function (Chen et al., 2015;  
Jonkman et al., 2008; Kucharavy and De Guio, 2011):

$$L(I) = \frac{A}{1 + e^{-k(I-c)}} \quad (2)$$

95 in which there are three parameters: 1) the magnitude parameter  $A$  representing the maximum drought loss; 2) the shape  
parameter  $k$  controlling the growth rate of  $L$  with  $I$ ; and 3) the location parameter  $c$  indicating the point at which the  
saturation begins.

As to drought indices that represent the intensity, they can be derived from the target hydroclimatic variable's cumulative  
distribution function (CDF) and the inverse CDF of the standard normal distribution (Hao et al., 2017; Mishra and Singh,  
100 2010; Montanari et al., 2023; Zhang et al., 2024; Zhao et al., 2024b). For example, the SPI is calculated as:

$$SPI_t = CDF_{N(0,1)}^{-1} \left( CDF_p(p_t) \right) \quad (3)$$

in which  $SPI_t$  in period  $t$ , which follows the standard normal distribution, is derived from precipitation amount  $p_t$  in period  $t$ .  
There are two steps. Firstly,  $p_t$  is converted into a standard uniform variate between 0 and 1 by its CDF, i.e.,  $CDF_p(\cdot)$ .



Secondly, the standard uniform variate is converted into the standard normal variate  $SPI$ , by the inverse CDF of  $N(0,1^2)$ , i.e.,  $CDF_{N(0,1)}^{-1}(\cdot)$ .

105

## 2.2 Formulation of the logistic function

There is an inverse relationship between drought intensity and drought indices like SPI, PDSI, SPEI and SRI. It is because the extent of dryness is generally characterized by how negative drought indices are (Haile et al., 2020; Liu et al., 2024; Zhang et al., 2021). That is, the more intensive dryness, the more negative drought indices. Taking the SPI as an indicator of drought intensity, the logistic function is modified by removing the negative sign in front of  $k$ :

110

$$L(SPI) = \frac{A}{1 + e^{k(SPI-c)}} \quad (4)$$

The ranges of the three parameters can be predetermined in accordance with the physical meanings of the parameters. First of all,

$$A > 0 \quad (5)$$

which means that drought loss is always above zero. Secondly,

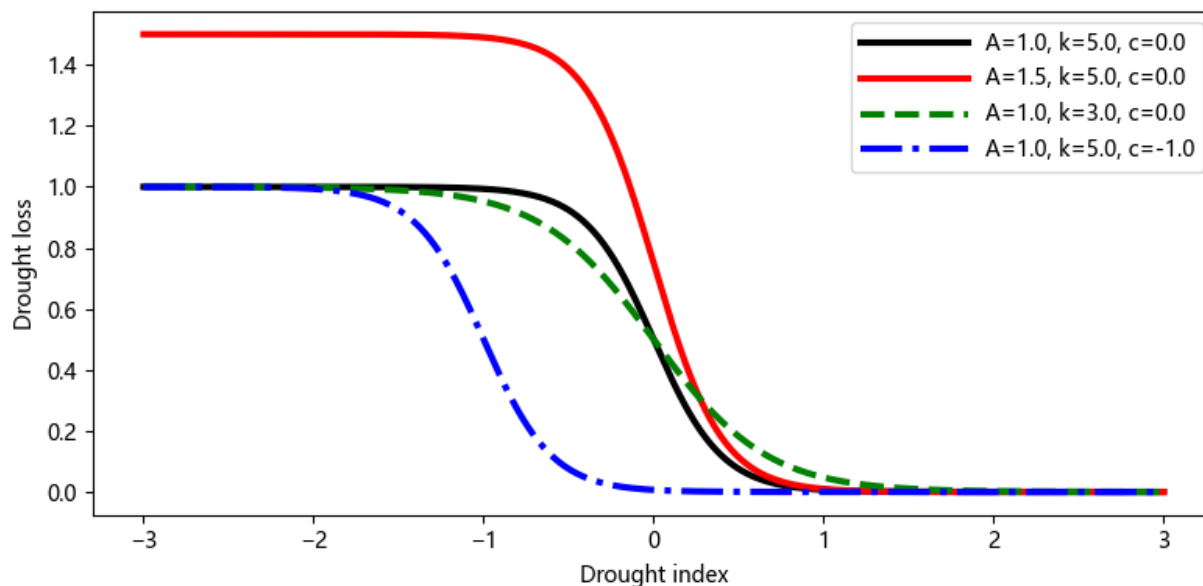
$$k > 0 \quad (6)$$

which means that as  $SPI$  increases from  $-\infty$  to  $+\infty$ , the denominator in Eq. (4) increases and leads to the reduction of drought loss. Eventually, the increasing denominator makes drought loss approach zero when SPI is large enough. On the other hand, it is noted that the loss would turn to increase with SPI when  $k$  is negative. Thirdly,

115

$$-\infty < c < +\infty \quad (7)$$

which means that the value of  $c$  depends on the case under investigation and can change freely.



120 **Figure 1. An illustrative example of the logistic function under four sets of parameters.**

An illustrative example of the logistic function in Eq. (4) is presented in Figure 1. The result under the basic parameter set of ( $A=1.0$ ,  $k=5.0$ ,  $c=0.0$ ) is marked in black. There are three one-factor-at-a-time experiments (Chen and Zhao, 2020). Firstly, the value of  $A$  is increased to 1.5. As is shown by the red line, the maximum drought loss evidently increases but the shape of the line stays the same. Secondly, the value of  $k$  is reduced to 3.0. As is shown by the green line, the shape of the line becomes flatter but the maximum loss remains the same. Thirdly, the value of  $c$  is decreased to  $-1.0$ . As is shown by the blue line, the plot is shifted to the left as a whole while both the maximum loss and shape do not change.

### 2.3 Stationary and non-stationary formulations

130 There are socio-economic factors contributing to temporal changes, i.e., nonstationarity, of the intensity loss function (AghaKouchak et al., 2021; Chiang et al., 2021; Long et al., 2020). Firstly, the maximum drought loss can increase with time owing to increases of population and accumulations of wealth. Secondly, the loss under a given level of drought intensity may decrease with time considering engineering measures, such as constructions of water storage reservoirs and inter-basin water diversion projects. Thirdly, the growth rate of drought loss with drought intensity can be attenuated by non-

135 engineering drought management measures such as hydroclimatic forecasting and forecast-informed reservoir operation. In order to account for temporal changes, the time  $t$  is explicitly taken as an explanatory variable and then respectively incorporated into the parameters  $A$ ,  $k$  and  $c$  (Cheng et al., 2014; Xiong et al., 2015):



$$\begin{cases} A_t = A_0 + A_1 \times t \\ k_t = k_0 + k_1 \times t \\ c_t = c_0 + c_1 \times t \end{cases} \quad (8)$$

in which  $A_0$ ,  $k_0$  and  $c_0$  are intercept terms while  $A_1$ ,  $k_1$  and  $c_1$  are trend terms. Without the trend terms, there is a stationary logistic function  $L_{A_0k_0c_0}(\cdot)$ :

$$L_{A_0k_0c_0}(SPI_t) = \frac{A_0}{1 + e^{k_0(SPI_t - c_0)}} \quad (9)$$

140 Incorporating  $A_t$  into Eq. (9), the logistic function  $L_{A_1k_0c_0}(\cdot)$  with a nonstationary magnitude parameter is derived:

$$L_{A_1k_0c_0}(SPI_t) = \frac{A_0 + A_1 \times t}{1 + e^{k_0(SPI_t - c_0)}} \quad (10)$$

Incorporating  $k_t$  into Eq. (9), the logistic function  $L_{A_0k_1c_0}(\cdot)$  with a nonstationary shape parameter is derived:

$$L_{A_0k_1c_0}(SPI_t) = \frac{A_0}{1 + e^{(k_0 + k_1 \times t) \times (SPI_t - c_0)}} \quad (11)$$

Incorporating  $c_t$  into Eq. (9), the logistic function  $L_{A_0k_0c_1}(\cdot)$  with a nonstationary location parameter is derived:

$$L_{A_0k_0c_1}(SPI_t) = \frac{A_0}{1 + e^{k_0(SPI_t - (c_0 + c_1 \times t))}} \quad (12)$$

In Eqs. (9) to (12), there are subscripts “Ax”, “kx” and “cx” for the intensity loss function. As to “x”, 0 and 1 respectively indicate the non-use and use of time  $t$ . Accordingly, there is one stationary function  $L_{A_0k_0c_0}(\cdot)$  and three non-stationary  
 145 functions  $L_{A_1k_0c_0}(\cdot)$ ,  $L_{A_0k_1c_0}(\cdot)$  and  $L_{A_0k_0c_1}(\cdot)$ . The fitting of these functions is considered as a nonlinear least-squares problem by searching for the set of parameters that minimize the sum of squares of residuals. It is performed by the `curve_fit` function under the `scipy` optimization toolbox in Python (Virtanen et al., 2020).

### 3 Case study

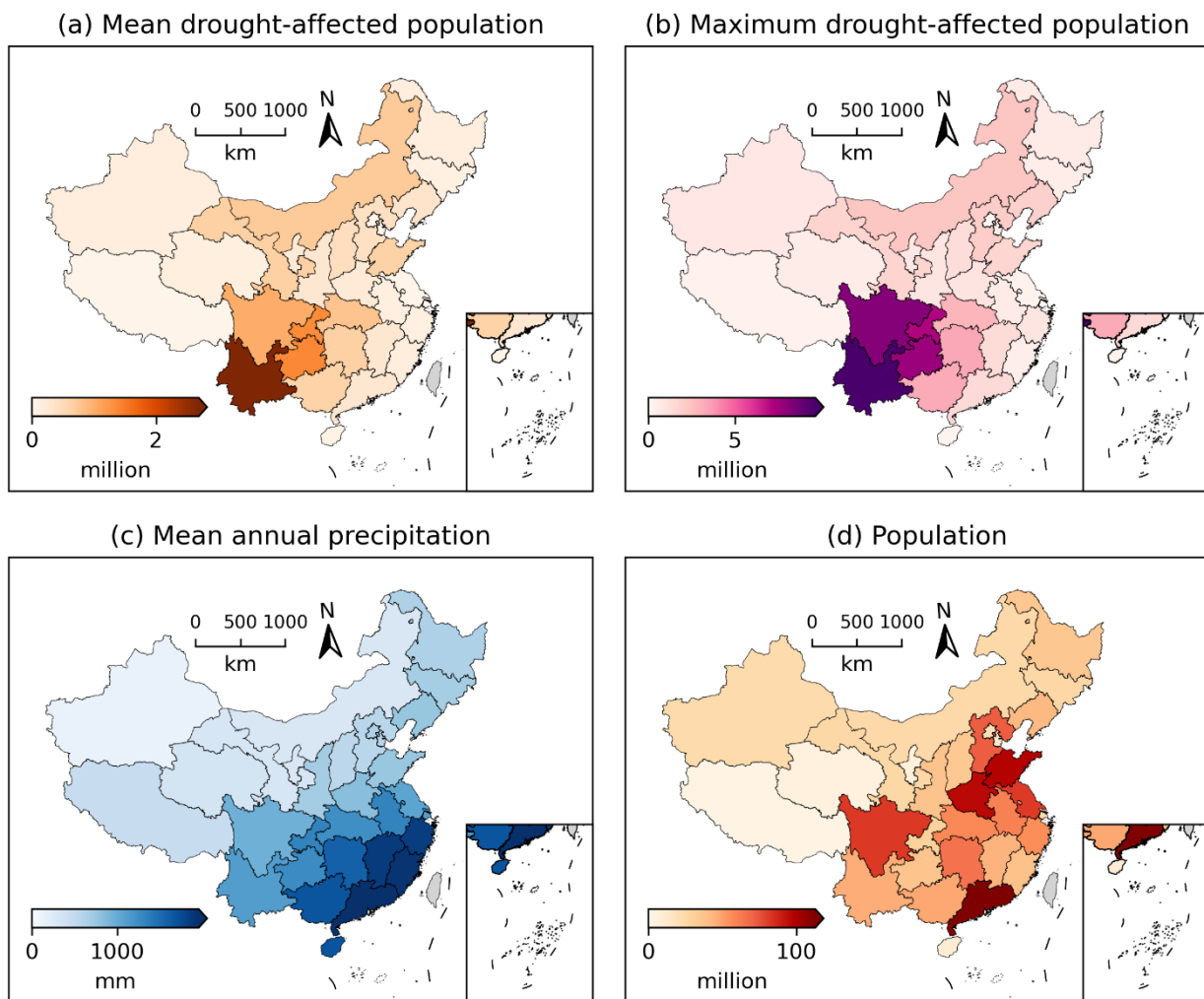
#### 150 3.1 Data description

To test the stationary and nonstationary intensity loss functions, the drought loss data is sourced from the Ministry of Water Resources (MWR) of China. It is noted that the MWR has since 2006 published by year “Bulletin of Flood and Drought Disaster in China”. The name of the bulletin was changed to “China Flood and Drought Disaster Prevention Bulletin” in  
 155 presented in the bulletin major events of droughts and floods across the 31 provinces in mainland China. As to droughts and floods in each province, the bulletin provides by year the quantitative socio-economic losses, contingency plans and retrospective analysis of prevention and control measures.

The attention is paid to the drought-affected population. In Figure 2 are the multi-annual mean drought-affected population, maximum annual drought-affected population, mean annual precipitation and total population. From Figures 2a and 2b, it



160 can be observed that provinces in Southwest China, including Yunnan, Guizhou and Sichuan Provinces, tend to have the  
largest drought-affected population. Particularly in 2010, 8.82 million people in Yunnan Province and 5.44 million people in  
Guizhou Province were struck by a record-breaking drought event induced by the persistently positive Madden-Julian  
Oscillation (Lü et al., 2012). On the other hand, it can be seen from Figures 2c and 2d that there is neither low precipitation  
nor large population in Southwest China. In general, the large drought-affected population in Yunnan and Sichuan Provinces  
165 is attributed to the Karst landscape, which features small storage capacity, high infiltration rate and fast groundwater flow  
(Wan et al., 2016).



170 **Figure 2. Spatial plots of (a) mean annual drought-affected population, (b) maximum annual drought-affected population, (c) mean annual precipitation and (d) population by province in mainland China.**





The precipitation data used for the calculation of SPI is obtained from the Climate Hazards Group InfraRed Precipitation with Station data (CHIRPS) (Funk et al., 2015). The intersection operation is performed to use provincial polygons to extract spatially averaged precipitation from the raster CHIRPS precipitation field. To better characterize the climatological distribution of precipitation, time series of annual precipitation are extracted by province for the period from 1981 to 2023. The 43 years' annual precipitation is firstly converted into CDF by the Weibull's plotting position (Ye et al., 2018) and then converted into SPI by the inverse CDF of  $N(0,1^2)$ . Then, the SPI in the years from 2006 to 2023 is used in the fitting of the logistic functions.

### 180 3.2 Model evaluation

The coefficient of determination, i.e.,  $R^2$ , is evaluated for the stationary logistic function (Eq. 9) and the three types of nonstationary logistic functions (Eqs. 10, 11 and 12). That is, the sum of squares of residuals for the estimations provided by the functions is compared to the baseline sum of squares of residuals for the mean value. As a result,  $R^2$  represents the ratio of total variation of the drought loss that is explained:

$$R^2 = 1 - \frac{\sum_t (L_t - \hat{L}_t)^2}{\sum_t (L_t - \bar{L})^2} \quad (13)$$

185 in which  $L_t$  is the drought loss in year  $t$ ,  $\hat{L}_t$  is the loss estimated by the function under investigation and  $\bar{L}$  is the mean value of all  $L_t$ .

The number of parameters plays a critical part in statistical modelling. That is, more parameters facilitate more flexible fitting of observed data but in the meantime, are more prone to overfitting (Neath and Cavanaugh, 2012). There are respectively 3 and 4 parameters in the stationary and nonstationary logistic functions:

$$\begin{cases} n_{A0k0c0} = 3 \\ n_{A1k0c0} = 4 \\ n_{A0k1c1} = 4 \\ n_{A0k0c1} = 4 \end{cases} \quad (14)$$

190 The Bayesian information criterion (BIC) takes into account both the sum of squares of residuals and the number of parameters (Neath and Cavanaugh, 2012):

$$BIC = T \times \ln \left( \frac{\sum_t (L_t - \hat{L}_t)^2}{T} \right) + n \times \ln(T) \quad (15)$$

in which  $\ln(\cdot)$  is the natural logarithmic function,  $T$  is the number of observations and  $n$  is the number of parameters. BIC is negatively oriented, meaning that a lower value indicates a better fit. As a result, both larger sum of squares of residuals and more parameters are penalized by the BIC.

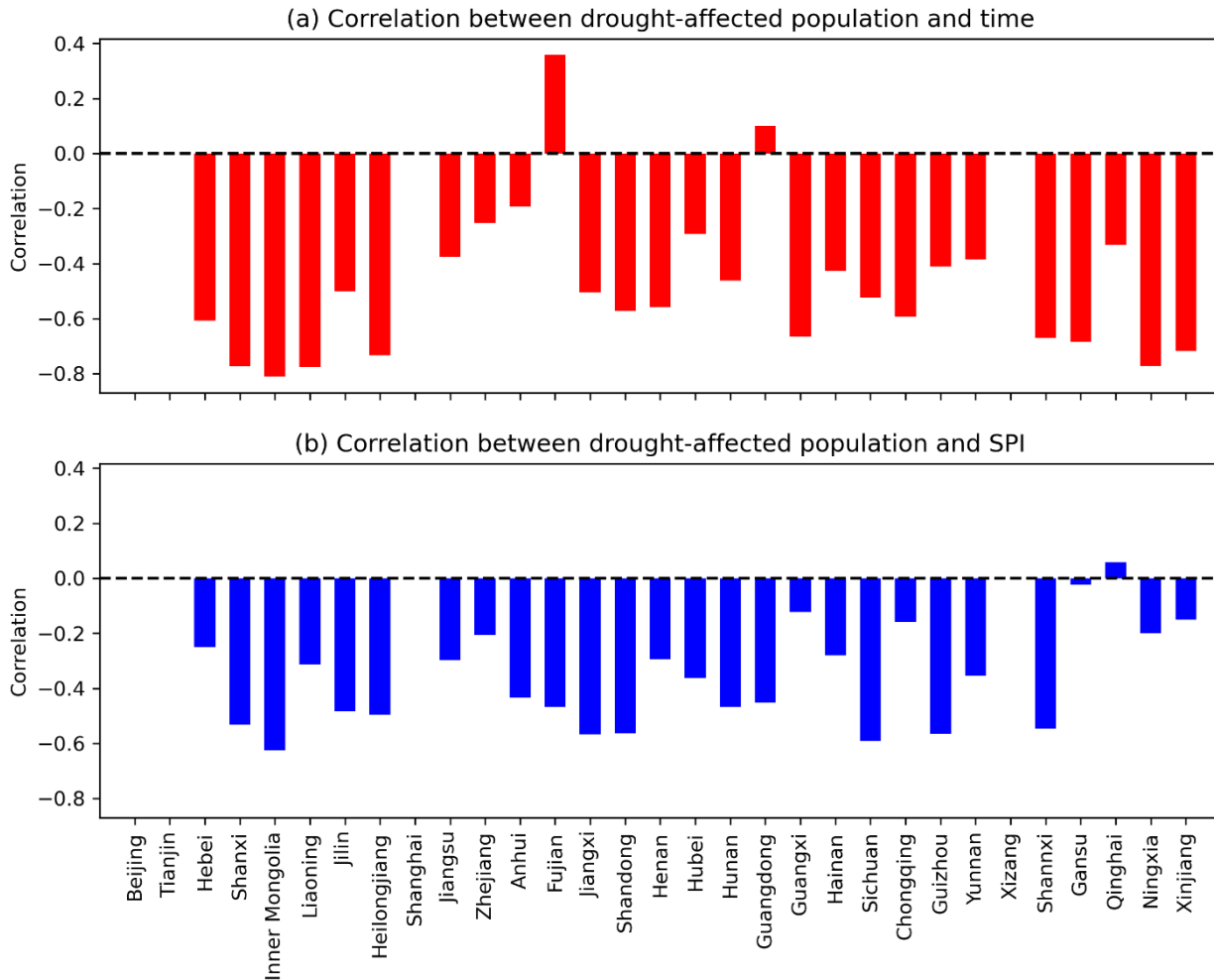
195



## 4 Results

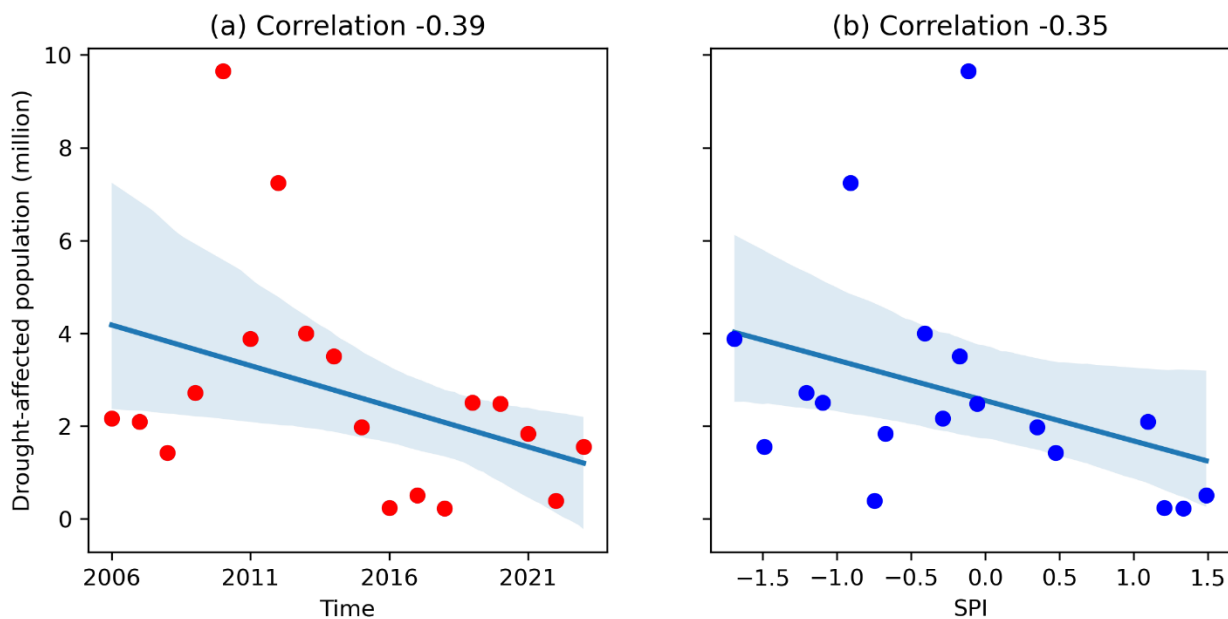
### 4.1 Correlation analysis

The Pearson's correlation coefficient between drought-affected population and time as well as SPI are illustrated by bar plots in Figure 3. There are in total 31 provincial administrative regions in mainland China. Beijing, Tianjin, Shanghai and Xizang are not considered since they are free from drought-affected population in most years. This outcome is mainly due to ample water supply by infrastructure developments (Long et al., 2020; Sun et al., 2021). For the other 27 provincial administrative regions, it can be observed that the correlation coefficient between drought-affected population and SPI is negative except for Guangdong and Fujian Provinces. The implication is that the drought-affected population mostly exhibits a decreasing trend as time progresses and sometimes shows an increasing trend. In the meantime, the correlation coefficient between drought-affected population and SPI is in general significantly negative. This result suggests that drought-affected population tends to decrease as the amount of precipitation increases. Overall, the correlation coefficients in Figure 3 point out that it is reasonable to use both time and SPI as explanatory variables of drought-affected population.

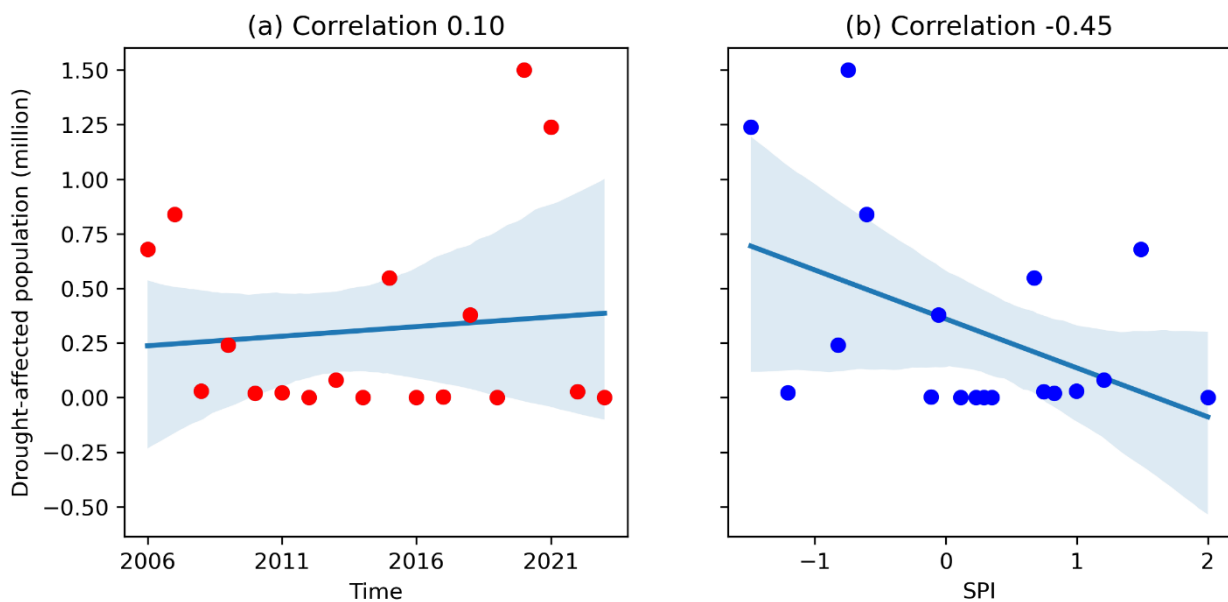


210 **Figure 3. Correlation coefficient between drought-affected population and (a) time as well as (b) SPI by province.**

The drought-affected population is plotted against time and SPI for Yunnan Province in Figure 4 and for Guangdong Province in Figure 5. The scatter plots on the left-hand side of the two figures imply the complexity of drought impact assessment. That is, owing to socio-economic developments, the drought-affected population can decrease or increase as time progresses (Apurv and Cai, 2021; Haile et al., 2020; Long et al., 2020). In the meantime, the scatter plots on the right-hand side suggest that the increase of precipitation amounts effectively reduces the population subject to droughts (AghaKouchak et al., 2021; Qiu et al., 2023; West et al., 2019).



220 **Figure 4.** Scatter plots of drought-affected population against (a) time and (b) SPI in Yunnan Province.



**Figure 5.** As for Figure 4, but for Guangdong Province.

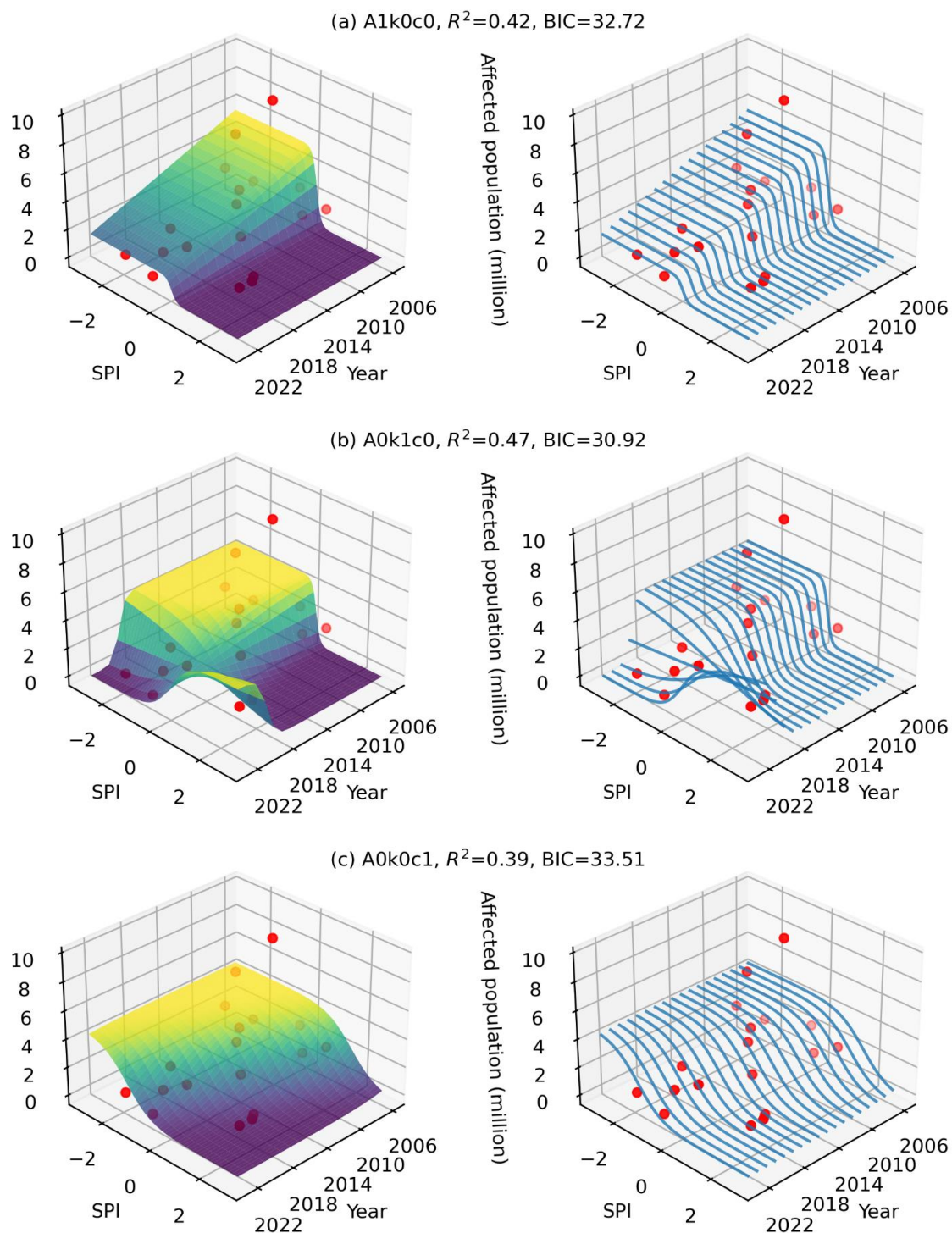


## 225 4.2 Decreasing drought-affected population

The three nonstationary logistic functions are one-by-one fitted by relating the drought-affected population to time and SPI for Yunnan Province. The results are visualized by the surface and wireframe plots in Figure 6. The 3d scatter points suggest that the drought-affected population tends to decrease with SPI and that it exhibits a decreasing trend as time progresses from 2006 to 2023. The three nonstationary logistic functions are shown to be effective in generating 3d response surfaces to  
230 characterize the dependency relationships, with reasonable  $R^2$  and BIC values. Meanwhile, they perform differently in capturing the observed data:

- 1) The flexible magnitude parameter in A1k0c0 tends to fit the observed data by reducing the maximum drought loss by year, as shown in Figure 4a. As can be seen from the wireframe plot, the maximum drought loss evidently reduces from 2006 to 2023 while the shape and location of the curves remain the same.
- 235 2) The flexible shape parameter in A0k1c0 fits the observed data by changing the response surface, as shown in Figure 4b. Although it exhibits the highest  $R^2$  and the lowest BIC, the fitted drought-affected population is shown to counterintuitively increase with SPI in 2021, 2022 and 2023. That is, more people could be subject to drought as precipitation increases in these three years. This wrong outcome is owing to the flexibility of the shape parameter. Specifically, the value of the shape parameter can be forced by the trend term to turn from positive to negative as time  
240 progresses. When the shape parameter is negative, the estimated drought impact would increase with the precipitation amount.
- 3) The flexible location parameter in A0k0c1 tends to fit the observation data by shifting the response curves by year, as shown in Figure 4c. Due to that the maximum drought-affected population is fixed from 2006 to 2023, it is observed that the maximum affected population in 2010 is not effectively captured.

245



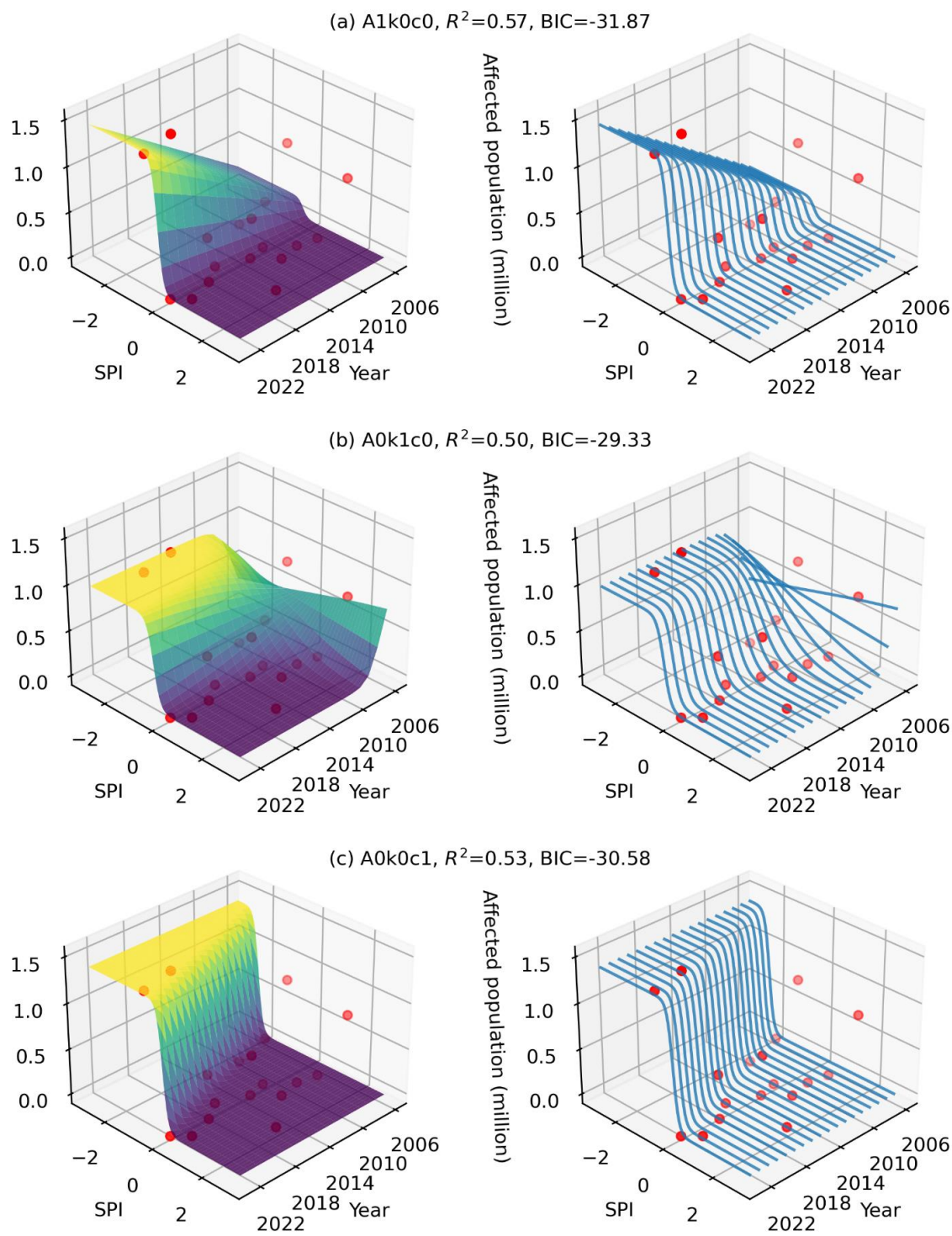
**Figure 6.** Surface plots (left) and wireframe plots (right) for the nonstationary logistic functions (a) A1k0c0, (b) A0k1c0 and (c) A0k0c1 relating the drought-affected population to time and SPI for Yunnan Province.



### 250 4.3 Increasing drought-affected population

Given the importance of the Guangdong-Hong Kong-Macao Greater Bay Area (Shao et al., 2020), the dependency of drought-affected population on time and SPI in Guangdong Province is illustrated by the surface and wireframe plots in Figure 7. Since the population of Guangdong is concentrated on the Pearl River Delta, recent years have witnessed serious water scarcity due to upstream reservoir impoundments and estuary saltwater intrusion (Weng et al., 2024). From Figure 7, it can be observed that the three non-stationary logistic functions characterize the increasing drought-affected population in different ways:

- 1) The flexible magnitude parameter in  $A1k0c0$  tends to fit the increase by enlarging the maximum drought loss by year. As shown in Figure 7a, it tends to capture the maximum drought-affected population of 1.50 million in 2020 and the second maximum drought-affected population of 1.24 million in 2021.
- 260 2) The flexible shape parameter in  $A0k1c0$  is observed to fit the observation data by changing the shape of response surface by year. As shown in Figure 7b, although the affected population in 2020 and 2021 is to some extent characterized, drought-affected population is seen to surprisingly increase with SPI in 2006. These results highlight the role that the shape parameter plays in determining the growth (reduction) rate.
- 265 3) The flexible location parameter in  $A0k0c1$  is shown to fit the observation data by fixing the maximum drought loss but shifting the response curves by year. As shown in Figure 7c, it tends to characterize the maximum and second maximum drought-affected population in recent years but does not seem to be as effective in characterizing drought-affected population in early years.



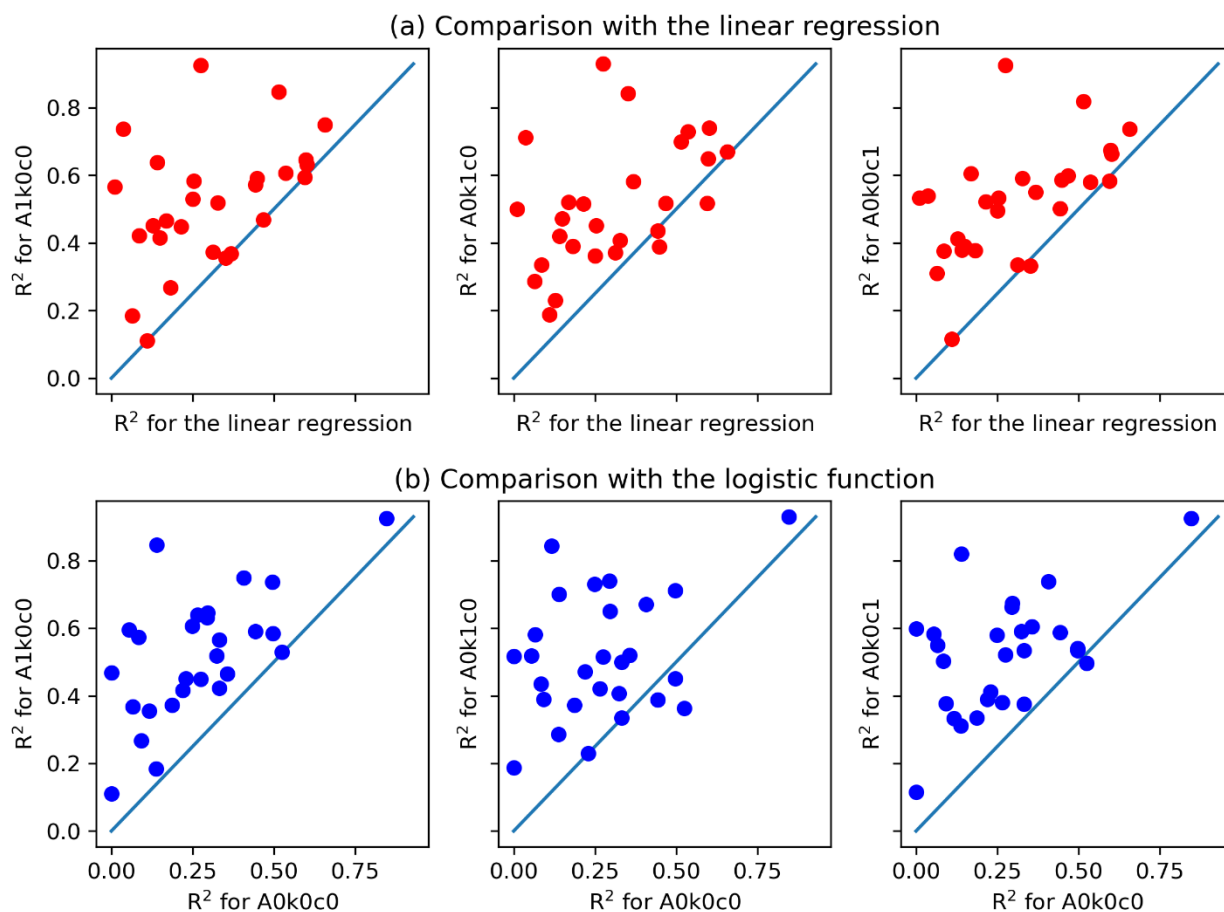
270 Figure 7. As for Figure 6, but for Guangdong Province.





#### 4.4 Goodness-of-fit

The stationary and nonstationary logistic functions are set up to account for the drought-affected population based on the explanatory variables of time and SPI for 27 provincial administrative regions other than Beijing, Tianjin, Shanghai and Xizang. The  $R^2$  for the three nonstationary logistic functions are plotted against that of linear regression based on time (Figure 8a) and also against that of the stationary logistic function (Figure 8b). The three scatter plots are generally above the 1:1 line. This result indicates that the consideration of time  $t$  evidently enhances the proportion of total variation explained by the non-stationary logistic functions. It is noted that the mean  $R^2$  is respectively 0.307 for linear regression and 0.269 for the stationary logistic function A0k0c0. By contrast, the mean  $R^2$  is respectively 0.512, 0.506 and 0.509 for A1k0c0, A0k1c0 and A0k0c1. Overall, the nonstationary logistic function A1k0c0 is of the highest  $R^2$ . This result highlights that the incorporation of time into the magnitude parameter can effectively deal with the non-stationary drought-affected population.



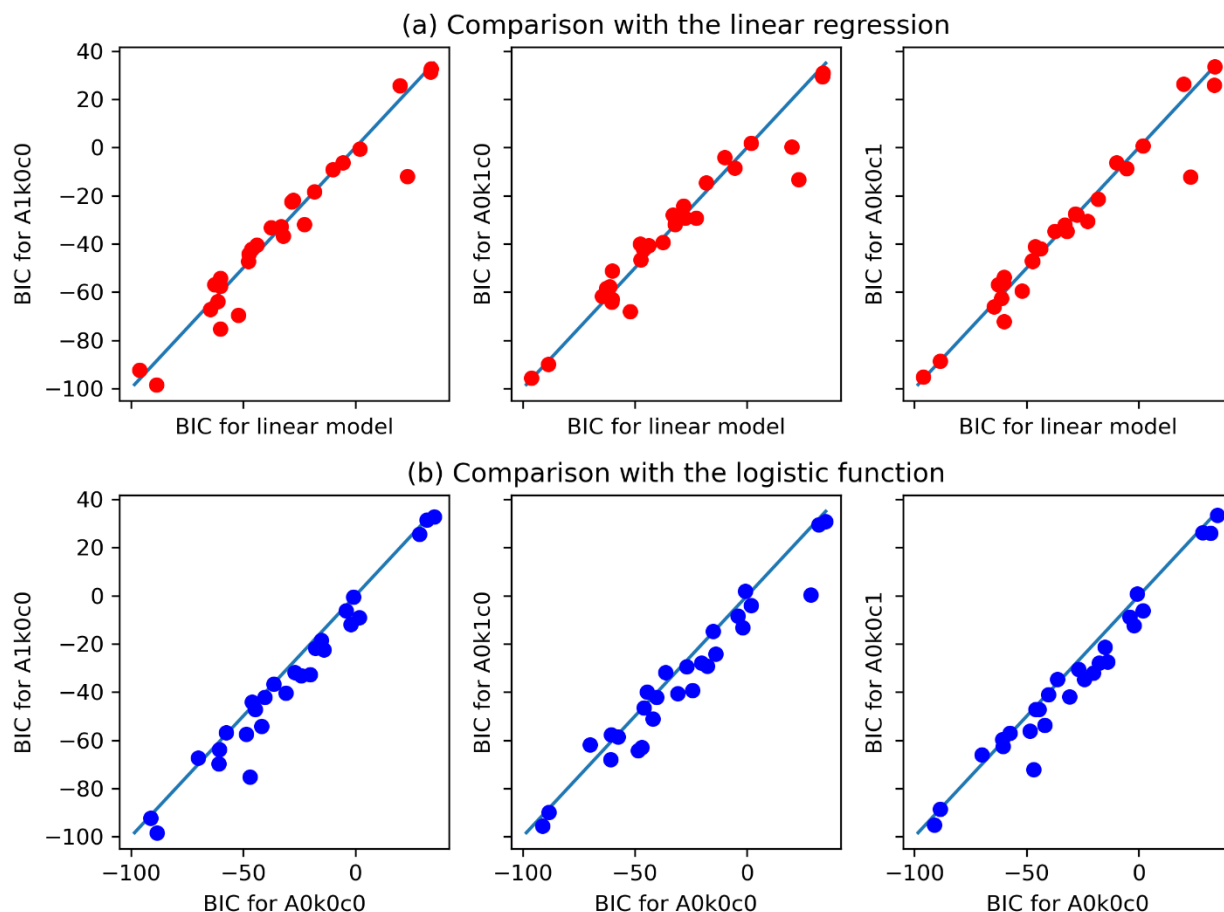


285 **Figure 8. Scatter plots of the  $R^2$  for the three nonstationary logistic functions against the  $R^2$  for (a) the linear regression and (b) the stationary logistic function A0k0c0.**

Furthermore, the BIC of the three nonstationary logistic functions is plotted against the BIC of the linear regression in Figure 9a and against that of the stationary logistic function in Figure 9b. Since the higher  $R^2$  of the nonstationary logistic functions in Figure 8 is at the cost of an additional parameter (Neath and Cavanaugh, 2012), the BIC takes into account not only the number of parameters but also the mean squared error. It can be observed that the scatter plots in Figure 5 are largely below the 1:1 line. Considering that the BIC is a negatively oriented metric, this result suggests that there is a low risk of overfitting and that the information hidden in the significant correlation (Figure 3) is deemed to be effectively exploited by the three non-stationary logistic functions. It is noted that the mean BIC is respectively  $-33.105$  for linear regression and  $-29.365$  for the stationary logistic function A0k0c0. By contrast, the mean BIC is respectively  $-34.980$ ,  $-34.772$  and  $-34.740$  for A1k0c0, A0k1c0 and A0k0c1. As the nonstationary logistic function A1k0c0 is of the lowest BIC, it is highlighted that the incorporation of time into the magnitude parameter of the logistic function is effective in accounting for the non-stationarity of drought losses.

290

295



300 **Figure 9.** As for Figure 8, but for the BIC.

## 5 Discussion

The nonstationary intensity loss functions developed in this paper complement existing studies on hydroclimatic processes of droughts (Garrido-Perez et al., 2024; Haile et al., 2020; Todisco et al., 2013). Focusing on drought indices such as SPI, PDSI, 305 SPEI and SRI, previous studies have presented in-depth investigations about past changes and future projections of meteorological, hydrological, agricultural and socio-economic droughts (Apurv and Cai, 2021; Hao and Singh, 2015; Mishra and Singh, 2010). One remarkable feature of the proposed intensity loss function is the explicit estimation of drought loss under different combinations of SPI and time. In future studies, the relationship between drought loss and other drought indices can readily be investigated at local and regional scales. Given that the logistic function is already an established 310 growth model in biosciences (Tsoularis and Wallace, 2002), it is expected that the proposed functions can be used to



characterize the growth of drought loss with drought conditions characterized by different drought indices. More case studies are in demand to test the usefulness.

The frequency, duration and intensity are three important characteristics of drought (Baez-Villanueva et al., 2024; Entekhabi, 2023; Liu et al., 2024; Mishra and Singh, 2010; Yang et al., 2024). Given a threshold for the identification of drought events, 315 the frequency is generally defined as the number of drought events in a certain period (one year for example), the duration as the timespan of a drought event and the intensity as the cumulative sum of the drought index (AghaKouchak et al., 2021; Chiang et al., 2021). Given that the SPI is derived for annual precipitation in this paper, the SPI values are expected to reflect the conditions of drought frequency, duration and intensity across different years. It is noted that the use of annual precipitation is mainly due to the fact that the drought-affected population by province is available at the annual timescale. It 320 is possible that drought losses are available on an event scale. In that case, event-based analysis becomes feasible. That is, both drought loss and intensity can be quantified for each drought event; and then the effectiveness of the logistic function can be tested.

## 6 Conclusions

325 This paper has presented three novel nonstationary intensity loss functions for drought impact assessment. On the one hand, the classic logistic function that has three parameters, i.e., magnitude, shape and location, presents a stationary formulation of the growth of drought losses with drought conditions. On the other hand, the incorporations of time respectively into the magnitude, shape and location parameters facilitate three nonstationary logistic functions. A case study is presented for the drought-affected population by province in China during the period from 2006 to 2023. The results highlight that despite the 330 fact that drought-affected population can either decrease or increase with time, the joint use of both time and SPI as explainable variables lead to effective characterization of drought-affected population. In comparison with the stationary logistic function, the effectiveness of the nonstationary logistic functions is indicated not only by higher  $R^2$ , which indicates reasonable proportion of total explained variation, but also by lower BIC, which suggests low risk of overfitting. Among the three nonstationary logistic functions, the function with nonstationary magnitude parameter generally outperform the other 335 two in terms of higher  $R^2$ , lower BIC and clearer physical meanings. Overall, the nonstationary intensity loss functions developed in this paper can serve as a useful tool for drought management.

## Acknowledgments

This research is supported by the National Natural Science Foundation of China (2023YFF0804900 and 52379033) and the 340 Guangdong Provincial Department of Science and Technology (2019ZT08G090).



### **CRediT authorship contribution statement**

Tongtiegang Zhao: Writing – original draft, Visualization, Software, Methodology, Conceptualization. Zecong Chen: Validation, Resources, Data curation. Yongyong Zhang: Investigation, Formal analysis.

345

### **Declaration of competing interest**

The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.

### **350 Data Availability Statement**

The drought-affected population is available from the Ministry of Water Resources of China (<http://www.mwr.gov.cn/sj/tjgb/zgshzhgb/>). The CHIRPS precipitation data is available from the Climate Hazards Center at the University of California, Santa Barbara (<https://www.chc.ucsb.edu/data/chirps>)

### **355 References**

AghaKouchak, A., Mirchi, A., Madani, K., Di Baldassarre, G., Nazemi, A., Alborzi, A., Anjileli, H., Azarderakhsh, M., Chiang, F., Hassanzadeh, E., Huning, L. S., Mallakpour, I., Martinez, A., Mazdiyasn, O., Moftakhari, H., Norouzi, H., Sadegh, M., Sadeqi, D., Van Loon, A. F., and Wanders, N.: Anthropogenic Drought: Definition, Challenges, and Opportunities, *Reviews of Geophysics*, 59, e2019RG000683, <https://doi.org/10.1029/2019RG000683>, 2021.

360 Apurv, T. and Cai, X.: Regional Drought Risk in the Contiguous United States, *Geophysical Research Letters*, 48, e2020GL092200, <https://doi.org/10.1029/2020GL092200>, 2021.

Baez-Villanueva, O. M., Zambrano-Bigiarini, M., Miralles, D. G., Beck, H. E., Siegmund, J. F., Alvarez-Garretón, C., Verbist, K., Garreaud, R., Boisier, J. P., and Galleguillos, M.: On the timescale of drought indices for monitoring streamflow drought considering catchment hydrological regimes, *Hydrol. Earth Syst. Sci.*, 28, 1415–1439, <https://doi.org/10.5194/hess-28-1415-2024>, 2024.

365

Chen, H. and Zhao, T.: Modeling power loss during blackouts in China using non-stationary generalized extreme value distribution, *Energy*, 195, 117044, <https://doi.org/10.1016/j.energy.2020.117044>, 2020.



- Chen, M., Ma, J., Hu, Y., Zhou, F., Li, J., and Yan, L.: Is the S-shaped curve a general law? An application to evaluate the damage resulting from water-induced disasters, *Nat Hazards*, 78, 497–515, <https://doi.org/10.1007/s11069-015-1723-9>, 2015.
- 370 Cheng, L., AghaKouchak, A., Gilleland, E., and Katz, R. W.: Non-stationary extreme value analysis in a changing climate, *Climatic Change*, 127, 353–369, <https://doi.org/10.1007/s10584-014-1254-5>, 2014.
- Chiang, F., Mazdiyasi, O., and AghaKouchak, A.: Evidence of anthropogenic impacts on global drought frequency, duration, and intensity, *Nat Commun*, 12, 2754, <https://doi.org/10.1038/s41467-021-22314-w>, 2021.
- Dai, A.: Drought under global warming: a review, *WIREs Climate Change*, 2, 45–65, <https://doi.org/10.1002/wcc.81>, 2011.
- 375 Entekhabi, D.: Propagation in the Drought Cascade: Observational Analysis Over the Continental US, *Water Resources Research*, 59, e2022WR032608, <https://doi.org/10.1029/2022WR032608>, 2023.
- Funk, C., Peterson, P., Landsfeld, M., Pedreros, D., Verdin, J., Shukla, S., Husak, G., Rowland, J., Harrison, L., Hoell, A., and Michaelsen, J.: The climate hazards infrared precipitation with stations—a new environmental record for monitoring extremes, *Scientific Data*, 2, 150066, <https://doi.org/10.1038/sdata.2015.66>, 2015.
- 380 Gao, H., Hrachowitz, M., Wang-Erlandsson, L., Fenicia, F., Xi, Q., Xia, J., Shao, W., Sun, G., and Savenije, H. H. G.: Root zone in the Earth system, *Hydrology and Earth System Sciences*, 28, 4477–4499, <https://doi.org/10.5194/hess-28-4477-2024>, 2024a.
- Gao, Y., Zhao, T., Tu, T., Tian, Y., Zhang, Y., Liu, Z., Zheng, Y., Chen, X., and Wang, H.: Spatiotemporal links between meteorological and agricultural droughts impacted by tropical cyclones in China, *Science of The Total Environment*, 912, 169119, <https://doi.org/10.1016/j.scitotenv.2023.169119>, 2024b.
- 385 Garrido-Perez, J. M., Vicente-Serrano, S. M., Barriopedro, D., García-Herrera, R., Trigo, R., and Beguería, S.: Examining the outstanding Euro-Mediterranean drought of 2021–2022 and its historical context, *Journal of Hydrology*, 630, 130653, <https://doi.org/10.1016/j.jhydrol.2024.130653>, 2024.
- Haile, G. G., Tang, Q., Li, W., Liu, X., and Zhang, X.: Drought: Progress in broadening its understanding, *WIREs Water*, 7, e1407, <https://doi.org/10.1002/wat2.1407>, 2020.
- 390 Hao, Z. and Singh, V. P.: Drought characterization from a multivariate perspective: A review, *Journal of Hydrology*, 527, 668–678, <https://doi.org/10.1016/j.jhydrol.2015.05.031>, 2015.
- Hao, Z., Yuan, X., Xia, Y., Hao, F., and Singh, V. P.: An Overview of Drought Monitoring and Prediction Systems at Regional and Global Scales, *Bulletin of the American Meteorological Society*, 98, 1879–1896, <https://doi.org/10.1175/BAMS-D-15-00149.1>, 2017.
- 395 Hersbach, H., Bell, B., Berrisford, P., Hirahara, S., Horányi, A., Muñoz-Sabater, J., Nicolas, J., Peubey, C., Radu, R., Schepers, D., Simmons, A., Soci, C., Abdalla, S., Abellan, X., Balsamo, G., Bechtold, P., Biavati, G., Bidlot, J., Bonavita, M., De Chiara, G., Dahlgren, P., Dee, D., Diamantakis, M., Dragani, R., Flemming, J., Forbes, R., Fuentes, M., Geer, A., Haimberger, L., Healy, S., Hogan, R. J., Hólm, E., Janisková, M., Keeley, S., Laloyaux, P., Lopez, P., Lupu, C., Radnoti, G., De Rosnay, P., Rozum, I., Vamborg, F., Villaume, S., and Thépaut, J.: The ERA5 global reanalysis, *Quart J Royal Meteorol Soc*, 146, 1999–2049, <https://doi.org/10.1002/qj.3803>, 2020.
- 400 Hoerling, M., Eischeid, J., Kumar, A., Leung, R., Mariotti, A., Mo, K., Schubert, S., and Seager, R.: Causes and Predictability of the 2012 Great Plains Drought, *Bulletin of the American Meteorological Society*, 95, 269–282, <https://doi.org/10.1175/BAMS-D-13-00055.1>, 2014.



- 405 Hou, W., Chen, Z.-Q., Zuo, D.-D., and Feng, G.: Drought loss assessment model for southwest China based on a hyperbolic tangent function, *International Journal of Disaster Risk Reduction*, 33, 477–484, <https://doi.org/10.1016/j.ijdrr.2018.01.017>, 2019.
- Jonkman, S. N., Vrijling, J. K., and Vrouwenvelder, A. C. W. M.: Methods for the estimation of loss of life due to floods: a literature review and a proposal for a new method, *Nat Hazards*, 46, 353–389, <https://doi.org/10.1007/s11069-008-9227-5>,  
410 2008.
- Kucharavy, D. and De Guio, R.: Application of S-shaped curves, *Procedia Engineering*, 9, 559–572, <https://doi.org/10.1016/j.proeng.2011.03.142>, 2011.
- Liu, R., Yin, J., Slater, L., Kang, S., Yang, Y., Liu, P., Guo, J., Gu, X., Zhang, X., and Volchak, A.: Machine-learning-constrained projection of bivariate hydrological drought magnitudes and socioeconomic risks over China, *Hydrol. Earth Syst. Sci.*, 28, 3305–3326, <https://doi.org/10.5194/hess-28-3305-2024>, 2024.  
415
- Long, D., Yang, W., Scanlon, B. R., Zhao, J., Liu, D., Burek, P., Pan, Y., You, L., and Wada, Y.: South-to-North Water Diversion stabilizing Beijing’s groundwater levels, *Nat Commun*, 11, 3665, <https://doi.org/10.1038/s41467-020-17428-6>, 2020.
- Lü, J., Ju, J., Ren, J., and Gan, W.: The influence of the Madden-Julian Oscillation activity anomalies on Yunnan’s extreme drought of 2009–2010, *Sci. China Earth Sci.*, 55, 98–112, <https://doi.org/10.1007/s11430-011-4348-1>, 2012.  
420
- Ma, M., Qu, Y., Lyu, J., Zhang, X., Su, Z., Gao, H., Yang, X., Chen, X., Jiang, T., Zhang, J., Shen, M., and Wang, Z.: The 2022 extreme drought in the Yangtze River Basin: Characteristics, causes and response strategies, *River*, 1, 162–171, <https://doi.org/10.1002/rvr.2.23>, 2022.
- Mishra, A. K. and Singh, V. P.: A review of drought concepts, *Journal of Hydrology*, 391, 202–216, <https://doi.org/10.1016/j.jhydrol.2010.07.012>, 2010.  
425
- Montanari, A., Nguyen, H., Rubinetti, S., Ceola, S., Galelli, S., Rubino, A., and Zanchettin, D.: Why the 2022 Po River drought is the worst in the past two centuries, *Sci. Adv.*, 9, eadg8304, <https://doi.org/10.1126/sciadv.adg8304>, 2023.
- Neath, A. A. and Cavanaugh, J. E.: The Bayesian information criterion: background, derivation, and applications, *WIREs Computational Stats*, 4, 199–203, <https://doi.org/10.1002/wics.199>, 2012.
- 430 Pradhan, R. K., Markonis, Y., Vargas Godoy, M. R., Villalba-Pradas, A., Andreadis, K. M., Nikolopoulos, E. I., Papalexiou, S. M., Rahim, A., Tapiador, F. J., and Hanel, M.: Review of GPM IMERG performance: A global perspective, *Remote Sensing of Environment*, 268, 112754, <https://doi.org/10.1016/j.rse.2021.112754>, 2022.
- Qiu, M., Ratledge, N., Azevedo, I. M. L., Diffenbaugh, N. S., and Burke, M.: Drought impacts on the electricity system, emissions, and air quality in the western United States, *Proc. Natl. Acad. Sci. U.S.A.*, 120, e2300395120, <https://doi.org/10.1073/pnas.2300395120>, 2023.  
435
- Shao, Q., Liu, X., and Zhao, W.: An alternative method for analyzing dimensional interactions of urban carrying capacity: Case study of Guangdong-Hong Kong-Macao Greater Bay Area, *Journal of Environmental Management*, 273, 111064, <https://doi.org/10.1016/j.jenvman.2020.111064>, 2020.
- 440 Su, B., Huang, J., Fischer, T., Wang, Y., Kundzewicz, Z. W., Zhai, J., Sun, H., Wang, A., Zeng, X., Wang, G., Tao, H., Gemmer, M., Li, X., and Jiang, T.: Drought losses in China might double between the 1.5 °C and 2.0 °C warming, *Proc. Natl. Acad. Sci. U.S.A.*, 115, 10600–10605, <https://doi.org/10.1073/pnas.1802129115>, 2018.



- Sun, S., Zhou, X., Liu, H., Jiang, Y., Zhou, H., Zhang, C., and Fu, G.: Unraveling the effect of inter-basin water transfer on reducing water scarcity and its inequality in China, *Water Research*, 194, 116931, <https://doi.org/10.1016/j.watres.2021.116931>, 2021.
- 445 Todisco, F., Mannocchi, F., and Vergni, L.: Severity–duration–frequency curves in the mitigation of drought impact: an agricultural case study, *Nat Hazards*, 65, 1863–1881, <https://doi.org/10.1007/s11069-012-0446-4>, 2013.
- Tsoularis, A. and Wallace, J.: Analysis of logistic growth models, *Mathematical Biosciences*, 179, 21–55, [https://doi.org/10.1016/S0025-5564\(02\)00096-2](https://doi.org/10.1016/S0025-5564(02)00096-2), 2002.
- 450 Van Dijk, A. I. J. M., Beck, H. E., Crosbie, R. S., De Jeu, R. A. M., Liu, Y. Y., Podger, G. M., Timbal, B., and Viney, N. R.: The Millennium Drought in southeast Australia (2001–2009): Natural and human causes and implications for water resources, ecosystems, economy, and society, *Water Resources Research*, 49, 1040–1057, <https://doi.org/10.1002/wrcr.20123>, 2013.
- 455 Virtanen, P., Gommers, R., Oliphant, T. E., Haberland, M., Reddy, T., Cournapeau, D., Burovski, E., Peterson, P., Weckesser, W., Bright, J., Van Der Walt, S. J., Brett, M., Wilson, J., Millman, K. J., Mayorov, N., Nelson, A. R. J., Jones, E., Kern, R., Larson, E., Carey, C. J., Polat, İ., Feng, Y., Moore, E. W., VanderPlas, J., Laxalde, D., Perktold, J., Cimrman, R., Henriksen, I., Quintero, E. A., Harris, C. R., Archibald, A. M., Ribeiro, A. H., Pedregosa, F., Van Mulbregt, P., SciPy 1.0 Contributors, Vijaykumar, A., Bardelli, A. P., Rothberg, A., Hilboll, A., Kloeckner, A., Scopatz, A., Lee, A., Rokem, A., Woods, C. N., Fulton, C., Masson, C., Häggström, C., Fitzgerald, C., Nicholson, D. A., Hagen, D. R., Pasechnik, D. V., Olivetti, E., Martin, E., Wieser, E., Silva, F., Lenders, F., Wilhelm, F., Young, G., Price, G. A., Ingold, G.-L., Allen, G. E., 460 Lee, G. R., Audren, H., Probst, I., Dietrich, J. P., Silterra, J., Webber, J. T., Slavič, J., Nothman, J., Buchner, J., Kulick, J., Schönberger, J. L., De Miranda Cardoso, J. V., Reimer, J., Harrington, J., Rodríguez, J. L. C., Nunez-Iglesias, J., Kuczynski, J., Tritz, K., Thoma, M., Newville, M., Kümmerer, M., Bolingbroke, M., Tartre, M., Pak, M., Smith, N. J., Nowaczyk, N., Shebanov, N., Pavlyk, O., Brodtkorb, P. A., Lee, P., McGibbon, R. T., Feldbauer, R., Lewis, S., Tygier, S., Sievert, S., Vigna, S., Peterson, S., More, S., et al.: SciPy 1.0: fundamental algorithms for scientific computing in Python, *Nat Methods*, 465 17, 261–272, <https://doi.org/10.1038/s41592-019-0686-2>, 2020.
- Wan, L., Zhou, J., Guo, H., Cui, M., and Liu, Y.: Trend of water resource amount, drought frequency, and agricultural exposure to water stresses in the karst regions of South China, *Natural Hazards*, 80, 23–42, <https://doi.org/10.1007/s11069-015-1954-9>, 2016.
- 470 Wang, F., Lai, H., Li, Y., Feng, K., Tian, Q., Zhang, Z., Di, D., and Yang, H.: Terrestrial ecological drought dynamics and its response to atmospheric circulation factors in the North China Plain, *Atmospheric Research*, 294, 106944, <https://doi.org/10.1016/j.atmosres.2023.106944>, 2023a.
- Wang, Y., Wang, S., Chen, Y., Wang, F., Liu, Y., and Zhao, W.: Anthropogenic drought in the Yellow River basin: Multifaceted and weakening connections between meteorological and hydrological droughts, *Journal of Hydrology*, 619, 129273, <https://doi.org/10.1016/j.jhydrol.2023.129273>, 2023b.
- 475 Weng, P., Tian, Y., Zhou, H., Zheng, Y., and Jiang, Y.: Saltwater intrusion early warning in Pearl river Delta based on the temporal clustering method, *Journal of Environmental Management*, 349, 119443, <https://doi.org/10.1016/j.jenvman.2023.119443>, 2024.
- 480 West, H., Quinn, N., and Horswell, M.: Remote sensing for drought monitoring & impact assessment: Progress, past challenges and future opportunities, *Remote Sensing of Environment*, 232, 111291, <https://doi.org/10.1016/j.rse.2019.111291>, 2019.





- Xiong, L., Du, T., Xu, C.-Y., Guo, S., Jiang, C., and Gippel, C. J.: Non-Stationary Annual Maximum Flood Frequency Analysis Using the Norming Constants Method to Consider Non-Stationarity in the Annual Daily Flow Series, *Water Resour Manage*, 29, 3615–3633, <https://doi.org/10.1007/s11269-015-1019-6>, 2015.
- 485 Yang, X., Wu, F., Yuan, S., Ren, L., Sheffield, J., Fang, X., Jiang, S., and Liu, Y.: Quantifying the Impact of Human Activities on Hydrological Drought and Drought Propagation in China Using the PCR-GLOBWB v2.0 Model, *Water Resources Research*, 60, e2023WR035443, <https://doi.org/10.1029/2023WR035443>, 2024.
- Ye, L., Hanson, L. S., Ding, P., Wang, D., and Vogel, R. M.: The probability distribution of daily precipitation at the point and catchment scales in the United States, *Hydrol. Earth Syst. Sci.*, 22, 6519–6531, <https://doi.org/10.5194/hess-22-6519-2018>, 2018.
- 490 Yuan, X., Wang, Y., Ji, P., Wu, P., Sheffield, J., and Otkin, J. A.: A global transition to flash droughts under climate change, *Science*, 380, 187–191, <https://doi.org/10.1126/science.abn6301>, 2023.
- Zhang, L., Yuan, F., and He, X.: Probabilistic Assessment of Global Drought Recovery and Its Response to Precipitation Changes, *Geophysical Research Letters*, 51, e2023GL106067, <https://doi.org/10.1029/2023GL106067>, 2024.
- 495 Zhang, X., Hao, Z., Singh, V. P., Zhang, Y., Feng, S., Xu, Y., and Hao, F.: Drought propagation under global warming: Characteristics, approaches, processes, and controlling factors, *Science of The Total Environment*, 838, 156021, <https://doi.org/10.1016/j.scitotenv.2022.156021>, 2022.
- Zhang, Y., Keenan, T. F., and Zhou, S.: Exacerbated drought impacts on global ecosystems due to structural overshoot, *Nat Ecol Evol*, 5, 1490–1498, <https://doi.org/10.1038/s41559-021-01551-8>, 2021.
- 500 Zhao, T., Chen, Z., Tian, Y., Zhang, B., Li, Y., and Chen, X.: A decomposition approach to evaluating the local performance of global streamflow reanalysis, *Hydrology and Earth System Sciences*, 28, 3597–3611, <https://doi.org/10.5194/hess-28-3597-2024>, 2024a.
- Zhao, T., Li, X., Li, Y., Zhang, B., and Zhang, Y.: Concurrent droughts across Major River Basins of the World modulated by El Niño–Southern Oscillation, *Journal of Hydrology*, 644, 132112, <https://doi.org/10.1016/j.jhydrol.2024.132112>, 2024b.