Acidification, warming, and nutrient management are projected to cause reductions in shell and tissue weights of oysters in a coastal plain estuary

Catherine R. Czajka[1], Marjorie A.M. Friedrichs[1], Emily B. Rivest[1], Pierre St-Laurent[1], Mark J. Brush[1], and Fei Da[2]

[1]Virginia Institute of Marine Science, William & Mary, Gloucester Point, VA, 23062, USA
[2]Department of Geosciences, Princeton University, Princeton, NJ, 08544, USA

Correspondence to: Catherine R. Czajka (czajkacatherine@gmail.com)

Abstract. Coastal acidification, warming, and nutrient management actions all alter water quality conditions that marine species experience, with potential impacts to their physiological processes. Decreases in calcite saturation state (Ω_{Ca}) and food availability, combined with warming water temperatures, pose a threat to calcifying organisms; however, the magnitude of future changes in estuarine systems is challenging to predict and is not well known. This study aims to determine how and where oysters will be affected by future acidification, warming, and nutrient reductions, and the relative effects of these stressors. To address these goals, an oyster bioenergetics model for Eastern oysters (*Crassostrea virginica*) was embedded in a 3-D coupled hydrodynamic-biogeochemistry model implemented for two tributaries in the lower Chesapeake Bay. Model simulations were forced with projected future conditions (mid-21st century atmospheric CO_2 and atmospheric temperature under representative concentration pathway (RCP) 8.5, as well as managed nutrient reductions) and compared with a realistic present-day reference run. Together, all three stressors are projected to reduce Ω_{Ca} and growth of oyster shell and tissue. Increased atmospheric CO_2 and temperature are both projected to cause widespread reductions in Ω_{Ca}. The resulting reductions in oyster shell and tissue growth will be most severe along the tributary shoals. Future warming during peak oyster growing seasons is projected to have the strongest negative influence on tissue and shell growth, due to summer water temperatures reducing filtration rates, enhancing shell dissolution and oyster respiration rates, and increasing organic matter remineralization rates, thus reducing food availability. Nutrient reductions will exacerbate deficits in oyster food availability, contributing to further reductions in growth. Quantifying the effects of these stressors provides insight on the areas in the lower bay where oysters will be most vulnerable to mid 21st-century conditions.

Short summary. Under future acidification, warming, and nutrient management, substantial reductions in shell and tissue weights of Eastern oysters are projected for the Chesapeake Bay. Lower oyster growth rates will be largely driven by reduced calcium carbonate saturation states and reduced food availability. Oyster aquaculture practices in the region will likely be affected, with site selection becoming increasingly important as impacts will be highly spatially variable.

1 Introduction

Anthropogenic climate change and its associated impacts on water quality may threaten marine organisms and economic systems reliant on them. Oceanic uptake of increasing anthropogenic atmospheric carbon dioxide (CO_2) causes a decrease in seawater pH and saturation states of calcium carbonate (e.g., Ω_{Ca}), a phenomenon known as ocean acidification (Caldeira and Wickett, 2003; Doney et al., 2009). Globally, the ocean has absorbed about 30% of anthropogenic atmospheric CO_2 since pre-industrial times (Gruber et al., 2019), and open-ocean surface pH is anticipated to decrease by 0.3 units on average relative to the 2010s by the end of the century under 'business-as-usual' conditions (Riahi et al., 2011; IPCC, 2019). The percent volume of ocean water undersaturated with calcite ($\Omega_{Ca} < 1$), a condition associated with elevated energetic costs of shell-building for calcifying organisms, is predicted to expand to 91% by 2100 from 76% in the 1990s (Caldeira and Wicket, 2005; Gattuso et al., 2015). Concurrently, the global ocean has absorbed approximately 93% of the atmospheric heat produced by anthropogenic activity, leading to a global sea surface temperature increase of 0.7°C since 1900 (Jewett and Romanou, 2017). Ocean warming is expected to continue, with global averages increasing by 2.7°C by 2100 and greater increases expected in shallow coastal regions (Jewett and Romanou, 2017).

Since estuaries have lower and more variable pH than the open-ocean, the effects of increased CO_2 and warming on estuarine water quality and biota are often amplified. In coastal and estuarine systems, acidification may be exacerbated by local-level processes, such as the input of acidic freshwater and nutrient runoff from precipitation, a process termed coastal acidification (Salisbury et al., 2008; Wallace et al., 2014; Carstensen and Duarte, 2019). Coastal acidification may accelerate as warming of coastal waters increases rates of biogeochemical processes; increased respiration rates may drive larger diel variations in pH, dissolved oxygen, and associated water quality (Du et al., 2018; Tian et al., 2021). Freshwater has relatively low total alkalinity (TA), or buffering capacity, so estuarine areas with greater relative freshwater influence cannot resist changes to pH as easily as more saline or open-ocean waters (Hasler et al., 2018; Pacella et al., 2024). Eutrophication, the increased rate of organic matter input to a system (Nixon, 1995), may also drive large variations in local pH and overall water quality. Elevated nutrient inputs cause pH to increase in surface waters due to higher primary productivity, which will reduce surface acidification; however, pH will decrease in deeper bottom waters as the additional organic matter sinks and is remineralized (Cai et al., 2021). Management actions to reduce eutrophication and improve water quality in bottom waters have been successful but may also enhance acidification in shallow surface waters by lowering primary productivity (Borges and Gypens, 2010; Da et al., 2021). The overall effect of future changes in nutrient inputs on coastal biogeochemistry is thus unclear, particularly in the presence of warming and acidification.

Characterizing spatiotemporal patterns of acidification in estuarine waters is important, as negative impacts of acidification on the biology of marine organisms may be substantial. Acidification disrupts the formation of calcium carbonate ($CaCO_3$) during shell-building, i.e., biocalcification, which is a vital process for growth and survival of many aquatic invertebrate species (e.g., Orr et al., 2005; Gazeau et al., 2007; Dong et al., 2023). Under acidified conditions, water concentrations of CO_2 and H^+ increase, and concentrations of carbonate ions ($[CO_3^{2-}]$) decrease. A low ambient $[CO_3^{2-}]$ inhibits

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105 calcifying organisms from forming $CaCO_3$ for their shells, as more energy is required to precipitate CO_3^{2-} from acidified waters

106 (e.g., Guinotte and Fabry, 2008; Lutier et al., 2022; Matoo et al., 2020; Mederios and Souza, 2023). Low Ω_{Ca} may also lead to

107 net dissolution of $CaCO_3$, leading to weaker shells and greater juvenile susceptibility to predation (e.g., Waldbusser et al.,

108 2011; Amaral et al., 2012; Barclay et al., 2020). Acidification may further reduce shell growth through adverse physiological

109 effects that limit energy availability for calcification. Because acidification is often more extreme in estuaries, coastal bivalve

110 species experience stronger effects of climate change than organisms living in open-ocean environments (Poach et al., 2019;

111 Melzner et al., 2020; Cai et al., 2021). Oysters in particular are of high management and research interest due to their

112 commercial value in the United States (USDA, 2023). Prior experiments have revealed negative effects of acidification,

113 warming, and nutrient reductions on oyster biocalcification and growth (Beniash et al., 2010; Waldbusser et al., 2011; Gobler

114 and Talmage, 2014), but it is yet to be determined how the impacts of these stressors on oyster shell and tissue growth will

115 vary spatially in highly dynamic systems.

116 　　The Chesapeake Bay is an excellent study system for examining the interacting influences of acidification, warming,

117 and nutrient reductions (hereafter referred to collectively as "future stressors") on estuarine biogeochemistry and the organisms

118 living there. The bay exhibits high temporal and spatial variability in pH due to seasonal phytoplankton blooms, eutrophication,

119 and acidic freshwater input (Da et al., 2021; St-Laurent et al., 2020; Kemp et al., 2005; Cai et al., 2021). From the mid-1980s

120 to mid-2010s, surface waters in the upper bay experienced pH increases between +0.2 and +0.4 pH units in early spring and

121 fall due to increased riverine TA from reduced acid mine drainage and lowered nitrate inputs, while surface waters in the

122 nitrogen-limited middle bay decreased up to -0.24 pH units during late spring and summer as a result of decreased primary

123 production (Da et al., 2021). Over the same time period, the bay warmed by $0.24 \pm 0.15°C$ per decade (Hinson et al., 2022),

124 more than double the average rate of warming for the upper 75m of the global ocean (Rhein et al., 2013). Warming has also

125 led to more severe hypoxia (Irby et al., 2018; Ni et al., 2020; Frankel et al., 2022; Hinson et al., 2023; Hinson et al., 2024). In

126 2010, the Environmental Protection Agency mandated a Total Maximum Daily Load (TMDL) of pollutants from point and

127 non-point sources to be achieved by 2025 (EPA, 2010). As nutrient reductions negatively affect pH in surface waters of the

128 bay (Shen et al., 2020; Da et al., 2021), achieving the TMDLs may actually worsen acidification in shallow and near-shore

129 regions. Much of the research effort devoted to characterizing present-day carbonate chemistry and its historical trends has

130 focused on the mainstem and upper Chesapeake Bay (Cai et al., 2017; Shen et al., 2020, Su et al., 2020), and less is known

131 about these conditions throughout the tributaries of the lower bay (Shadwick et al., 2019).

132 　　The combined effects of future stressors will impact calcifying organisms in the lower Chesapeake Bay as well as the

133 economic systems reliant upon them. The Eastern oyster *Crassostrea virginica* (Gmelin, 1791) is a foundation species native

134 to the bay (Dayton, 1972). Eastern oyster aquaculture in this region has grown rapidly in the past few decades, with Virginia

135 becoming the third most productive oyster fishery in 2018 (Hudson, 2019), largely a result of the development of disease-

136 resistant oyster strains (Frank-Lawale et al., 2014). Negative impacts of acidification on aquaculture practices in other parts of

137 the world (Barton et al., 2015) have already stirred concern over the vulnerability of oysters in the Chesapeake Bay. For

138 example, in the Pacific Northwest, major larval mortality occurred at a shellfish hatchery following an upwelling event that

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142 lowered pH and Ω of aragonite, which had cascading impacts on the oyster industry all along the West Coast (Barton et al.,

143 2015). While most effects of acidification on aquaculture have been observed in oyster larvae in hatcheries, fewer studies have

144 examined acidification's influence on adult oysters when deployed in the field. Water quality conditions within oyster farms

145 can be highly spatially variable, so the impacts of acidification may vary with growing conditions (Saavedra et al., 2024;

146 Simpson et al., 2024). To support the future of the oyster aquaculture industry in Chesapeake Bay, it is critical to identify

147 which areas in the bay will be most vulnerable to acidification at mid-century and how each driver of change contributes to

148 acidification and its impacts on growth.

149 This study addresses the following primary research question: How and where will carbonate chemistry and Eastern

150 oyster growth in the lower Chesapeake Bay change in the future and which future stressors will drive these changes? A three-

151 dimensional hydrodynamic-biogeochemical model is coupled with an oyster bioenergetics model and is applied to two major

152 Virginia tributaries of the Chesapeake Bay. The model provides detailed information on present-day environmental conditions,

153 and when combined with climate projections from Earth System Models, allows for simulations of the independent and

154 interacting influences of future environmental change on carbonate chemistry and Eastern oysters. This study provides insight

155 into which areas are most vulnerable to mid 21st-century acidification and how acidification, warming, and nutrient loading

156 may each impact oyster growth in isolation as well as via simultaneous co-stressors.

157 **2 Methods**

158 **2.1 Model description**

159 **2.1.1 Hydrodynamic model**

160 This study uses the three-dimensional hydrodynamic Regional Ocean Modeling System (ROMS; Shchepetkin and

161 McWilliams, 2005), implemented similarly to St-Laurent and Friedrichs (2024) but on a higher resolution grid focused on two

162 of the lower Virginia Chesapeake Bay tributaries (Fig. 1). The model domain (Da et al., 2024) includes the York and

163 Rappahannock Rivers, as well as a portion of the mainstem shoal north of the Rappahannock. The model grid consists of

164 620x740 horizontal grid cells with a horizontal resolution of 120 m, allowing for greater resolution of coastlines than many

165 other Chesapeake Bay model grids (Irby et al., 2016). The hydrodynamic model includes 20 terrain-following vertical levels

166 and two primary state variables: practical salinity and potential temperature. A wetting and drying scheme has been

167 implemented to represent water levels and currents in coastal grid cells (Warner et al., 2013; St-Laurent and Friedrichs, 2024).

168 **2.1.2 Carbon and biogeochemistry model**

169 The Estuarine-Carbon-Biogeochemistry model (ECB) embedded in ROMS and used in this study has previously been

170 implemented in the Chesapeake Bay (Feng et al., 2015; St-Laurent et al., 2020; Frankel et al., 2022; Hinson et al., 2023) as

171 well as in the lower Virginia tributaries (Da et al., 2024). ECB simulates full carbon and nitrogen cycles of the lower trophic

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levels, represented by the following state variables: nitrate, ammonium, phytoplankton and zooplankton nitrogen, small and large detrital nitrogen and carbon, semi-labile and refractory dissolved organic nitrogen, DIC, TA, and dissolved oxygen (O_2). Phytoplankton and zooplankton carbon are calculated from fixed C:N ratios (Redfield, 1934; Hopkinson et al., 1998). C:N ratios for dissolved organic matter in the estuary are allowed to freely evolve with time. Biogeochemical processes include primary production, aggregation, sinking, basal metabolism, exudation, sloppy feeding, excretion, metabolism, nitrification/denitrification, remineralization, grazing, and mortality. Biogeochemical sources and sinks are included in the bottom vertical level (e.g., burial, resuspension, nitrification/denitrification, remineralization, sediment O_2 and CO_2 exchange). Light attenuation throughout the water column is based on the diffuse attenuation coefficient (Kd), which is parameterized as a function of surface total suspended solids (TSS; including inorganic and organic components) and salinity as a proxy for colored dissolved organic matter (Feng et al., 2015; Turner et al., 2021). TSS is calculated within the model as the sum of the four inorganic suspended sediment size classes and particulate organic matter. Particulate organic carbon (POC) is calculated as the sum of phytoplankton carbon, zooplankton carbon, and small and large detrital carbon. The sediment transport module within ECB accounts for sediment-water exchange processes, such as deposition and resuspension of inorganic sediment and particulate organic matter, and is comprised of two vertical seabed layers that simulate four suspended sediment size classes (Turner et al., 2021).

The inorganic carbon module within ECB has been fine-tuned in this implementation of the model, allowing for improved performance in carbonate system simulations (Da et al., 2024). The model grid includes tidal wetlands along the York River based on estimated wetland areas (Mitchell et al., 2017), which further contribute to TA fluxes through sulfate reduction in sediments (Raymond et al., 2000; Najjar et al., 2020). Ω_{Ca} is calculated from DIC, TA, temperature, and salinity using CO2SYS (van Heuven et al., 2011) using the equilibrium constants of Cai and Wang (1998) as they are suitable for both fresh and estuarine waters (Dinauer and Mucci, 2017; Herrmann et al., 2020). Although submerged aquatic vegetation is a possible source of $CaCO_3$ (Mazarrasa et al., 2015; Su et al., 2020), $CaCO_3$ precipitation and dissolution are not simulated in ECB due to both insufficient observations and low submerged aquatic vegetation presence throughout the model domain (Orth et al., 1998; Moore et al., 2009).

Several additional updates have been made in this implementation of ROMS-ECB to better represent oxygen and primary production dynamics in the lower Virginia tributaries. The maximum phytoplankton growth rate has been increased to 2.15 d^{-1}, and the growth rate is limited in the fresh portion of the tributaries using a Michaelis-Menten function of salinity and a half-saturation of 1.5 (Da et al., 2024). The sediment bed climatology from Moriarty et al. (2021) has been adjusted to better represent the sand class distributions published in Nichols (1991) and observations taken by the USGS (Reid et al., 2005). Specifically, the changes include a greater percentage of small clay-rich flocs throughout the main stem of the York River as well as more sand and large silt-rich flocs in the Rappahannock River. Previously, the sediment module assumed the same critical shear stress for large silt-rich flocs, small clay-rich flocs, and unaggregated mud; here, the critical shear stress of smaller particles is lower than larger particles, meaning smaller particles resuspend more easily. The updated critical shear stress coefficient for erosion and deposition is 0.5 Pa for large silt-rich flocs and 0.4 Pa for both small clay-rich flocs and

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212 unaggregated mud, which represent a small portion of the sediment bed. The ballasting formulation of Turner et al. (2021) has
213 also been added to simulate the increase in particle sinking rates due to the aggregation of particles in turbid waters.

2.1.3 Oyster bioenergetics model

215   As part of this study, the oyster bioenergetics model EcoOyster (Brush and Kellogg, 2018; Kellogg et al., 2018) has
216 been one-way coupled to ROMS-ECB in the deepest (bottom) level (see Supplementary Tables S1-S4 for EcoOyster
217 equations). The one-way coupling means that environmental conditions from ROMS-ECB influence oyster growth in
218 EcoOyster, but the oysters do not influence water quality conditions in ROMS-ECB. This allows the focus of our analysis to
219 be on the effect of future climate change on oyster growth, rather than the effect of oyster growth on water quality, which has
220 been previously studied in the Chesapeake Bay (e.g., Gawde et al., 2024). By focusing on the deepest vertical level of the
221 model, the assumption is that oysters are growing on the bottom, and not inside floating cages. This is representative of
222 conditions similar to on-bottom or bottom cage aquaculture methods that are common in Virginia. The coupled model, referred
223 to hereafter as ROMS-ECBO, simulates daily somatic tissue dry weight, gonadal tissue dry weight, shell dry weight, and shell
224 height of diploid and triploid oysters as a function of filtration, respiration, egestion, allocation to reproduction, calcification,
225 and dissolution (Brush and Kellogg, 2018; Kellogg et al., 2018; Rivest et al., 2023). For the purpose of this study, only diploid
226 oysters were included, as model equations were developed from a study on diploid oysters, and triploid allometric equations
227 are not as well constrained. Tissue growth rates depend on individual weight together with temperature, salinity, O_2, TSS, and
228 POC from ROMS-ECB. Chl$a$ is required for the filtration function and is calculated from ROMS-ECB phytoplankton carbon
229 and Kd, in combination with seasonal carbon:chl$a$ ratios that are computed using equations from Cerco and Noel (2004). The
230 calcification function includes a threshold value of $\Omega_{Ca} = 0.93$, determined through laboratory experiments with Eastern oysters
231 (Rivest et al., 2023).

232   The *EcoOyster* equations were developed from a meta-analysis of existing oyster bioenergetics models and laboratory
233 experiments with diploid oysters (Brush and Kellogg, 2018; Kellogg et al., 2018; Rivest et al., 2023). Allometric relationships
234 between shell dry weight, tissue dry weight, and shell height used for initial conditions were derived from observational data
235 in the Chesapeake Bay (VOSARA, 2024). Total dry tissue weight is calculated as the sum of somatic tissue weight and gonadal
236 weight. Reproduction is simulated through gonadal weight, a function of growth of gonadal tissue, resorption of gonadal tissue,
237 and spawning (Hofmann et al., 1994). Somatic tissue weight is a function of assimilation, respiration, growth of gonadal tissue,
238 and resorption of gonadal tissue. Assimilation is a function of filtration and POC. Filtration is a function of a maximum
239 filtration rate based on tissue weight, limited by sub-optimal temperature, salinity, TSS, O_2, and chl$a$ (Cerco and Noel, 2005;
240 Fulford et al., 2007; Ehrich and Harris, 2015). The optimal temperature for oyster filtration (T_{opt}) is set to 27 °C (Jordan, 1987).
241 Filtration is also multiplied by p, a tunable factor representing the proportion of computed filtration actually performed by
242 oysters, which accounts for processes excluded from the model such as time spent filtering and is constrained by published
243 growth rates. Respiration is a function of tissue weight, temperature, and assimilation. While filtration has a temperature
244 limitation, respiration increases exponentially with temperature (Fig. S1). Tissue growth functions are not affected by

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249 carbonate chemistry variables, as experimental studies have found that neither filtration (Lemasson et al., 2018) nor respiration
250 (Beniash et al., 2010; Matoo et al., 2013) of oysters are affected by pH changes; however, weight-specific net calcification is
251 a function of Ω_{Ca} and temperature (Rivest et al., 2023). Shell growth is a function of both total tissue weight and net
252 calcification.

2.2 Present day reference simulation

254 A realistic reference simulation was generated to represent 2017 conditions. The year 2017 was chosen for
255 atmospheric, terrestrial, and open-ocean boundary conditions as this represents a relatively typical hydrological year.
256 Atmospheric forcings (air temperature, long- and short-wave radiation, precipitation, winds, dewpoint temperature, and air
257 pressure) are obtained from the ERA5 atmospheric reanalysis (Copernicus Climate Change Service, 2017; Hersbach et al.,
258 2020). Surface atmospheric variables are available at 3-hourly intervals with a 0.25° resolution and are interpolated to a 0.2°
259 grid. Terrestrial inputs of freshwater, nitrogen, carbon, and sediment are derived from the Phase 6 CBP Watershed Model
260 (CBPWM; Bhatt et al., 2023) and USGS data. Daily estimates of freshwater discharge, water temperature, and loadings of
261 nitrate, ammonium, organic nitrogen, and four classes of sediment from the CBPWM were concatenated to 74 locations
262 throughout the model domain. To compute carbon loadings, constant carbon-to-nitrogen ratios are used, specifically 10:1 for
263 dissolved organic matter (Hopkinson et al., 1998) and 6.625:1 for particulate organic matter (Redfield, 1934). Riverine TA
264 concentrations are computed as in Da et al. (2024), using monthly-varying linear relationships between historical USGS
265 observations of discharge and USGS TA estimates determined using the Weighted Regression on Time, Discharge, and Season
266 (WRTDS; Hirsch et al., 2010) approach. Riverine DIC is calculated from daily riverine TA and daily DIC:TA ratios, linearly
267 interpolated from the monthly climatology of USGS WRTDS DIC:TA in each tributary. Open boundary conditions are derived
268 from a recent 600 m resolution whole-bay implementation of ROMS (St-Laurent and Friedrichs, 2024). Initial conditions for
269 the six-month spin-up were derived from Da et al., (2024).

270 Since spring-spawned oysters are typically deployed in late spring through summer on oyster farms, the reference run
271 was started on July 1st and spanned one full year, ending June 30th of the following year. Oyster sizes were initialized based
272 on shell height approximations of a typical spring-spawned oyster at deployment in July (i.e., a few months old). Starting dry
273 tissue weight was assumed to be 0.001 g for all oysters, back-calculated from the approximate height of an oyster at the time
274 of deployment. Starting shell dry weights and heights were calculated from allometric relationships to be 0.144 g and 11.6
275 mm, respectively.

2.3 Comparison of reference simulation to in situ observations

277 *In situ* water quality monitoring observations are available since 1984 throughout the Chesapeake Bay. Specifically,
278 the Chesapeake Bay Program's Water Quality Monitoring Program (CBP WQMP) conducts cruises in the Bay and its
279 tributaries. On average, stations are sampled once monthly, with the exception of June through August in the mainstem, when
280 sampling occurs twice. In this study, measurements of water temperature, salinity, O_2, pH (NBS scale), TSS, and POC are

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284 used from 16 CBP stations throughout the model domain, with depths ranging from 5 to 16 m (Fig. 1a; CBP, 2024). For all

285 variables except TSS and POC, measurements are taken *in situ* using a YSI or Hydrolab® sonde roughly every one to two

286 meters of the water column. TSS and POC are obtained from bottle samples at the surface, bottom, and at deeper stations, two

287 additional depths above and below the pycnocline. TSS is determined by filtering a known volume of water through a pre-

288 weighted filter and then re-weighing the filter after filtration and drying. POC is determined through combustion of a filter

289 using an elemental analyzer (Olsen, 2012).

290 Model skill was evaluated by comparing results from the reference simulation to the CBP WQMP observations

291 described above. Hourly outputs from the four closest grid cells to each CBP station were spatially interpolated to obtain results

292 at each respective station. Multiple variables in ECB at the bottom level of the model, including temperature, salinity, O_2, pH,

293 TSS, and POC, were compared with observations from the same station and time, within the bottom 10% of the water column

294 (Table 1). Model bias and root-mean squared difference (RMSD) were computed for all aforementioned variables. Seasonal

295 skill was also evaluated by comparing the 2017 reference run to CBP decadal averages. Decadal means were used for these

296 comparisons, as once-monthly or once-seasonally sampling dates in 2017 bias outputs toward conditions on the time of the

297 month when the measurements were taken in 2017, and the purpose of the comparison was to examine how the model

298 reproduces average seasonal variability.

2.4 Future simulations

300 In addition to the reference run, this study generated five future simulations (Table 1) to investigate the change in

301 carbonate chemistry conditions and oyster growth resulting from three drivers of future change in the bay: increased

302 atmospheric CO_2 (*AtmCO₂*), atmospheric warming (*Temp*), and reduced nutrient loading (*TMDL*). Model forcings were

303 modified for each simulation to represent mid-century conditions. A *Combined Future* simulation was run including forcings

304 of all future stressors, in addition to three sensitivity simulations to isolate the impacts of each stressor on oyster growth.

305 Atmospheric CO_2 concentration for the *AtmCO₂* and *Combined Future* simulations was set to 655 ppm, derived from the

306 Coupled Model Intercomparison Project Phase 5 report RCP8.5 (business-as-usual) scenario projected for 50 years in the

307 future relative to the reference run (Meinshausen et al., 2011). While it is reasonable to assume that warming temperatures will

308 always accompany increases in atmospheric CO_2, separate experiments for each stressor were conducted. Isolating the

309 physiological effects of each stressor helps us understand the mechanisms effecting the change in oyster growth and

310 furthermore indicates the degree to which management strategies for offsetting acidification can reduce the effects of climate

311 change on oysters. Future atmospheric temperature for the *Temp* and *Combined Future* simulations was obtained from the

312 IPSL-CM5B-LR Earth System Model (Dufresne et al., 2013), statistically downscaled with the Multivariate Adaptive

313 Constructed Analogs method (Abatzoglou and Brown, 2012). IPSL-CM5B-LR was selected as in Hinson et al. (2024), since

314 it was deemed the most representative downscaled ESM of the 20 available (Hinson et al., 2023). As in Hinson et al. (2024),

315 the delta method was used to calculate the daily average change in atmospheric temperatures between present-day and future

316 conditions. To calculate this change, two 30-year climatologies, centered on 2000 and 2050 respectively, were computed and

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daily averaged 50-year differences between the two climatologies (Fig. 2) were added to the atmospheric temperatures used in the reference run. Future watershed inputs for the *TMDL* and *Combined Future* simulations included a climatology of nitrate, ammonium, dissolved organic matter, and particulate organic matter concentrations, derived from a Phase 6 CBPWM 1991-2000 run using reduced nutrient concentrations assuming the TMDLs had been successfully achieved (approximately a 20-25% mean reduction in total nutrient loading; Bhatt et al., 2023). Freshwater discharge in this run was set to be identical to the reference run, to isolate the effects of lowered nutrient concentrations on water chemistry and oyster growth. Since future climate change is expected to impact terrestrial inputs much less than future management actions (Hinson et al., 2023), the direct impact of climate change on the watershed is not considered in this analysis. A fifth simulation (*AtmCO₂ + Temp*) was run to compare the influences of local management actions to the combined drivers of climate change, which includes both future atmospheric CO_2 concentration and atmospheric temperature. Preliminary investigations revealed a minimal impact of sea level rise on Ω_{Ca} in the bay; therefore, it was not included in the simulated climate change variables.

To generate open boundary conditions for each future simulation, a full bay model (St-Laurent and Friedrichs, 2024) was run with the same atmospheric and river forcings as in this 120-meter model implementation. As in the reference run, all future simulations were spun up for six months (January 1 – June 30) before beginning on July 1, but represent 50 years in the future from the reference simulation (i.e., July 1, 2067). Initial conditions for all spin-ups are identical to the reference simulation. Analysis confirmed the effects of initial conditions are negligible by July 1. To examine results most relevant to oysters, model output was extracted at locations that support oyster production, defined as all grid cells in which tissue weight exceeded 1 g at the end of the reference run (i.e., one year of growth; Fig. S2), which is inclusive of locations of active oyster farms. All results shown are from the bottom level of the model, representing conditions similar to on-bottom or bottom cage aquaculture methods that are common in Virginia. Spatial variation in model outputs across grid cells in the model domain is reported using standard deviation.

3 Results

3.1 Model Skill Assessment

When compared to 2017 WQMP observations and seasonal decadal averages, model skill of ROMS-ECB is reasonably high (Table 2, Figs. 3, 4), and similar to other model implementations of the Chesapeake Bay (Irby et al., 2016). Temperature and salinity are reproduced relatively well year-round (Fig. 3a,b) with annual biases of only 0.2°C and -1.5, respectively (Table 2). Bottom O_2 and pH are slightly overestimated, exhibiting the greatest model-data misfit in the spring and summer months in the tributary channels (Fig. 3c,d). pH is overestimated by 0.2 units, which is within the accuracy of the electrode measurements. Observed POC concentrations in the York and upper Rappahannock are higher than simulated in the model and exhibit very high spatial variability (Fig. 4a). Despite the high spatial variability of the TSS observations (Fig. 4b), mean TSS (45 ± 54 mg L⁻¹) is captured within 1.1 mg L⁻¹ by the model.

394   Growth rates determined using the *EcoOyster* equations and environmental outputs from ROMS-ECB were compared
395 with oyster data collected in the York River (Paynter et al., 2008; Liddel, 2008; Kingsley-Smith et al., 2009; Degremont et al.,
396 2012; Callam et al., 2016). Specifically, the tunable parameter (*p*) that limits oyster filtration was adjusted to provide a best
397 match between the modeled oyster growth rates and the published rates. Multiple *p*-values were tested, and a value of *p*=0.15
398 resulted in modeled oyster growth that best matched published growth rates. The resulting shell growth predicted by the model
399 was found to be close to the *in situ* data (52.0 ± 1.1 mm y^{-1} and 51.3 ± 2.9 mm y^{-1} for the model and observation means and
400 standard deviations, respectively).

401 **3.2 Reference run results**

402   In the present-day reference run, the environmental variables used as inputs to the oyster parameterizations exhibit
403 substantial seasonal (Fig. 5a-f) and spatial (Figs. 6, S3) variability. As expected, bottom temperature is highest in summer,
404 reaching an average of 29.3 °C in July when averaged across grid cells that support oyster growth (Fig. 5a). Temperature is
405 higher in the shallower parts of the tributaries compared to the channels (Fig. S3a). Bottom salinity exhibits higher values in
406 the fall and winter, reaching a maximum average of 17.7 in October, and drops in the spring and summer to reach a minimum
407 average of 12.3 in June (Fig. 5b). Annual average bottom salinity ranges from 0 to 26 throughout the model domain (Fig. S3b),
408 with the highest values in the southern areas in closest proximity to the open-ocean. The seasonal cycle for bottom POC is
409 similar to that of temperature, peaking at 1.7 g C m^{-3} in June and dropping to 0.57 g C m^{-3} in January (Fig. 5c). Bottom POC
410 also varies widely throughout the model domain (Fig. 6a), with relatively higher values in the Rappahannock compared to the
411 York River, along the shoals of the tributaries, and along the western shoals of the mainstem Bay north of the Rappahannock.
412 Ω_{Ca} exhibits an annual cycle similar to that of temperature and POC, reaching a maximum average of 3.2 in August and a
413 minimum average of 1.1 in January (Fig. 5d). Annual mean bottom Ω_{Ca} also varies widely throughout the model domain (Fig.
414 6d). Generally, bottom Ω_{Ca} increases with salinity, with low to zero values in the tidal fresh portions of the upper tributaries
415 and higher values along the western shoals of the mainstem Chesapeake Bay. The opposite temporal pattern is seen in bottom
416 O_2, which peaks at 12.3 mg L^{-1} in February and drops to an average of 6.3 mg L^{-1} in August (Fig. 5e). O_2 concentrations are
417 highest along the shoals and lowest in the deep channels (Fig. S3c). Bottom TSS concentrations exhibit tidal variability
418 throughout the year and are highest in the York River with much lower concentrations observed in the other portions of the
419 model domain (Fig. S3d).

420   Tissue and shell weights increase modestly from July through April, and the highest rates of increase are seen in May
421 and June near the end of the one-year reference run (Fig. 5g,h). At the end of the reference run, the spatial patterns of shell and
422 tissue weight are nearly identical (Fig. 7), as tissue growth largely drives shell growth (Table S4). Both shell and tissue weights
423 are highest along the shoals of the York and Rappahannock Rivers (Fig. 7a,d) and low in the deeper waters where TSS
424 concentrations are high (Fig. S3d). A wider region of high shell and tissue weight appears in the Rappahannock, while the
425 highest weights in the York are confined to a very narrow and shallow strip along the coastline. Shell and tissue weights are

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431 higher along the southwestern than the northeastern coastlines of the tributaries, where the shoals are wider in both tributaries

432 (Fig. 1a).

433 **3.3 Results of Combined Future simulation**

434 All environmental variables examined exhibit change from the reference run in the *Combined Future* simulation.

435 Temperature and salinity are projected to increase across the entire model domain (Fig. S3a,b). When averaged over the model

436 domain, temperature is projected to increase by $1.5 \pm 0.26°C$, and salinity is projected to increase by 0.21 ± 0.11 (Table S5).

437 Bottom POC is projected to decrease by 0.07 ± 0.05 g C m^3 (Table S5), with POC reductions predicted to be most pronounced

438 in the mid- to upper tributaries (Fig. 6c). Mid-century bottom Ω_{Ca} is projected to be lower throughout most of the region (Fig.

439 6f), with an average reduction of 0.8 ± 0.19 over the whole model domain (Table S5). The spatial distribution of future Ω_{Ca} is

440 generally consistent with present-day Ω_{Ca} patterns, and the greatest decreases are projected to occur in regions with the highest

441 present-day Ω_{ca} (Fig. 6 d,e,f). An average reduction in O_2 of 0.3 ± 0.08 mg L^{-1} is predicted across the model domain (Table

442 S5), which will be mostly spatially uniform (Fig. S3c). TSS is projected to be reduced by 0.20 ± 0.25 mg L^{-1} with high spatial

443 variability in the projected change (Table S5; Fig. S3d).

444 Changes in environmental conditions do not occur uniformly throughout the year. Temporal changes in environmental

445 conditions averaged across grid cells that support oyster growth are provided in Figure 5. Annually averaged increases in

446 temperature and salinity are the same when averaged over only grid cells that support oyster growth as they are when averaged

447 across the entire model domain (Tables 3, S5). The greatest temperature increases are projected to occur in the warmer months,

448 with an average increase of 1.6 °C predicted for June through August and an average increase of 1.2 °C predicted for December

449 through February. Bottom temperatures are projected to surpass the optimal temperature for oyster filtration (27 °C) primarily

450 in July and August (Fig. 5a). Salinity increases are projected to be greatest at the beginning of the year, with an average increase

451 of 0.44 between January and March and an average increase of 0.20 for the remainder of the year (Fig. 5b). Bottom POC at

452 grid cells that support oyster growth is expected to decrease slightly less than the average for the entire region (Tables 3, S5),

453 by an annual average of 0.09 g C m^{-3}, with the greatest reductions in the spring and summer and little to no change in the

454 winter (Fig. 5c). For Ω_{Ca}, O_2, and TSS, projected reductions are slightly greater at oyster growth sites than for the entire domain.

455 Ω_{Ca} is projected to decrease by 0.9, with the greatest reductions expected to occur the warmer months (Fig. 5d). O_2 is projected

456 to decrease year-round, though with slightly greater reductions in the winter (Fig. 5e) and an annual average reduction of 0.4

457 mg O_2 L^{-1} (Table 3). TSS is projected to decrease annually by 0.3 mg L^{-1} (Table 3), mostly in the spring, due to lowered POC

458 (Fig. 5f).

459 Modeled shell and tissue weights after one year of growth are projected to decline in all regions that exhibit present-

460 day growth, with the most severe reductions (up to 100%) occurring along the York and Rappahannock River shoals (Figs.

461 7c,f, 8). One-year tissue weight will be reduced by 1.3 g, on average, representing a 60% reduction across grid cells that support

462 oyster growth (Table 4). Shell weight, which is largely driven by changes in tissue weight, is projected to be reduced by 11.4

463 g on average after one year of growth, representing a 68% reduction in average shell weight in regions that support oyster

11

growth (Table 4). The greatest reduction in shell and tissue growth rates will occur in the warmer months near the end of the one-year simulation (-0.1 g d^{-1} from May through June), whereas the smallest change will occur in the winter months (-0.02 g d^{-1} from December through February), as the least growth occurs during that time (Fig. 5g,h). Shell thickness, calculated as the ratio of shell weight to shell height, will be reduced by 61% on average (0.11 g mm^{-1}; Table 4).

Declines in year one shell weight will vary throughout the model domain (Fig. 8), following relative changes in bottom POC and Ω_{Ca} (Fig. 9). The mainstem has the most moderate reduction in shell weight relative to reference shell weight, with an average reduction of 31%, indicated by the slope of the scatterplot. Shell weights in the Rappahannock and York face the steepest reductions relative to reference, with average reductions of 86% and 96%, respectively, and a large portion of York oysters facing complete depletion of oyster tissue and shell in these locations (Fig. 9; indicated by proximity to 1:1 line). Proportional shell weight reductions in the mainstem are projected to correlate with POC reductions (Fig. 9a). For Ω_{Ca} in the mainstem, a group of sites face the greatest proportional reductions when Ω_{Ca} reductions are the greatest. However, for sites with lower proportional shell loss, the opposite trend is observed (Fig 9d). In the Rappahannock, higher POC reductions coincide with slightly lower proportional shell loss (Fig. 9b). Sites with the largest reductions in POC primarily occur in the York (Fig. 9c; see dark blue symbols on the 1:1 line), and the greatest proportional shell weight reductions coincide with the greatest POC and Ω_{Ca} reductions (Fig. 9c,f). Similar results are found for tissue weight (not shown).

3.4 Results of individual future sensitivity simulations

Four individual future sensitivity simulations were conducted to isolate the specific mechanisms (increased atmospheric CO_2, increased atmospheric temperature, and/or nutrient reductions) causing the projected changes described above in the *Combined Future* simulation. The *AtmCO2* sensitivity simulation produces substantial reductions in average bottom Ω_{Ca} (Fig. 10d) and, as expected, is not projected to substantially impact bottom temperature, salinity, POC, O_2, or TSS (Table 3; Figs. 10a, S4). The projected reduction in Ω_{Ca} is 0.9 when averaged over oyster growth sites (Table 3), 0.1 greater in magnitude than the average reduction for the entire model domain (Table S5), as greater reductions are expected along the shoals of the Rappahannock and mainstem shoal than the York and upper section of the Rappahannock (Fig. 10d). In this *AtmCO2* simulation, shell weight is predicted to be most steeply reduced in the Rappahannock, with less impact in the York and mainstem regions (Fig. 11a). At grid cells with oyster growth, *AtmCO2* produces a shell weight reduction of 6.3 g and a shell thickness reduction of 0.06 g mm^{-1} in comparison to the reference simulation, but no change in tissue weight (Table 4).

The *Temp* sensitivity simulation produces changes in all environmental variables impacting oyster growth, with the exception of Ω_{Ca} (Tables 3, S5). Average changes in temperature, salinity, and TSS will be identical to those from the *Combined Future* simulation (Tables 3, S5). Predicted reductions in POC and O_2 will be smaller in magnitude than in *Combined Future,* though more severe than any other single sensitivity experiment (Table 3, S5). Temperature and salinity will increase across the entire model domain, with a greater salinity increase occurring in the Rappahannock and along the mainstem shoal (Fig. S4a-c). TSS will decrease primarily in the channels of the lower York and Rappahannock and on the mainstem shoal (Fig. S4d). POC reductions are expected to cover the majority of the model domain, with larger reductions in

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504 the upper Rappahannock (Fig. 10b). Slight increases in Ω_{Ca} are observed in shallow tidal creeks (Fig. 10e); however, no

505 substantial change in average Ω_{Ca} is predicted (Table 3, S5). O_2 at oyster grid cells will exhibit a similar but slightly smaller

506 average reduction compared to *Combined Future* (Table 3). Patterns of change in shell weight in the *Temp* sensitivity

507 simulation resemble those in the *AtmCO2* simulation (Figs. 7c, 11b), with additional reductions along the mainstem shoal and

508 a greater predicted mean reduction of 7.1 g, a 42% decrease at grid cells with oyster growth (Table 4). Unlike *AtmCO2*, tissue

509 weight will decrease in *Temp*, by an average of 1 g, a 46% reduction (Table 4).

510   The *TMDL* sensitivity simulation produces a much smaller average change in environmental conditions than the

511 *AtmCO2* or *Temp* simulations (Tables 3, S5). *TMDL* does not substantially influence temperature, salinity, or O_2 (Tables 3, S4,

512 S5), but produces POC and TSS reductions close to the averages for *Temp* (Tables 3, S5). While POC change in the *Temp*

513 simulation is concentrated in the deeper portions of the tributaries (Fig. 10b), the POC reductions in the *TMDL* simulation are

514 concentrated along the shoals of the tributaries, with the greatest reductions in the upper Rappahannock (Fig. 10c). TSS changes

515 in *TMDL* are limited to the tributaries, occurring along the shoals of the Rappahannock and in patches throughout the York

516 (Fig. S4d). Future change in Ω_{Ca} in this simulation is less than for *AtmCO2* and is largely confined to the upper Rappahannock

517 shoals and in shallow tidal creeks throughout the study region (Fig. 10d, f). Patterns of change in shell weight will resemble

518 *AtmCO2* and *Temp* in the tributaries, but no change is predicted along the mainstem shoal (Fig. 11a-c). The *TMDL* simulation

519 produces reduced shell (3.7 g) and tissue (0.5 g) weights, with a smaller negative influence on shell and tissue weight than

520 *Temp* (Table 4; Fig. 11b-f).

521   Environmental conditions in the *AtmCO2 + Temp* simulation are nearly identical to those in the *Temp* simulation

522 (Tables 3, S5), with the exception of Ω_{Ca}, which is lower due to the influence of heightened atmospheric CO_2. As tissue growth

523 is unaffected by Ω_{Ca}, tissue weight in this simulation is identical to that of the *Temp* simulation (Table 4). Average shell weight

524 reduction in *AtmCO2 + Temp* is 10.2 g, greater than that from *AtmCO2* alone, due to the combined influences of lowered tissue

525 growth and lower Ω_{Ca} (Table 4).

526 **4 Discussion**

527   This study provides high-resolution projections for oyster growing conditions and corresponding oyster growth in the

528 Chesapeake Bay, with a specific focus on two Virginia tributaries. A high-resolution hydrodynamic-biogeochemical model

529 was coupled with an Eastern oyster bioenergetics model and forced with future projections for atmospheric CO_2, temperature,

530 and nutrient management. An overall reduction in Ω_{Ca} and oyster growth are predicted by mid-century throughout the study

531 region under the combined effects of all three future stressors. Specifically, the greatest reductions in oyster growth are

532 projected to occur in the York and Rappahannock Rivers, where unfavorable conditions for calcification will expand in the

533 future and where food availability will be strongly impacted by warming and nutrient reductions. Bottom conditions in the

534 York and Rappahannock rivers, particularly in the upper portions, will likely be unsuitable for aquaculture at mid-century on

535 average, indicating climate change preparedness is critical for the oyster aquaculture industry.

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4.1 Future projections of Ω_{Ca}

The magnitude of future change in Ω_{Ca} varies with present-day Ω_{Ca} conditions (Fig. 6d-f). Regions with high present-day Ω_{Ca}, primarily the mainstem shoals, are projected to experience the greatest reductions because of their low partial pressure of CO_2 (pCO_2) relative to fresher waters. Biologically driven low pCO_2 water on mainstem shoals has a greater capacity for CO_2 uptake from the atmosphere than high pCO_2 water, which is causing the fresher tributaries to experience smaller increases in DIC and smaller reductions in Ω_{Ca} (Fig. 6f). Acidic freshwater input often causes pCO_2 in the upper tributaries to exceed atmospheric pCO_2, causing outgassing (Cai et al., 2017; Shen et al., 2019b; St-Laurent et al., 2020; Cai et al., 2021). Despite the lower Rappahannock having a lower salinity than the lower York, it also has a lower DIC to TA ratio, so the rate at which the lower Rappahannock absorbs pCO_2 is higher (Da et al., 2021). Total alkalinity, or buffering capacity, is also lower in the lower Rappahannock than the lower York (Da et al. 2021), so the lower Rappahannock cannot resist changes in carbonate chemistry to the same degree as the York. As a result, we observe the Rappahannock changing faster than the York (Fig.6f, Fig. 9). Since higher Ω_{Ca} regions will experience greater reductions than lower Ω_{Ca} regions (Fig. 6d-f), the overall spatial variability of Ω_{Ca} will be reduced by mid-century, and more areas will experience conditions that are unfavorable for oyster shell-building (Fig. 5).

Although future atmospheric CO_2 and reduced nutrient loading will both contribute to Ω_{Ca} reductions, the modeling experiments conducted here highlight that increasing atmospheric CO_2 is the largest contributor to decreases in Ω_{Ca} throughout the study region. Increased atmospheric CO_2 will cause reductions in Ω_{Ca} across the model domain, while nutrient reductions are expected to mainly influence Ω_{Ca} in shallow and fresh coastal areas, with little influence in oyster growing regions. The effects of warming on Ω_{Ca} may slightly offset the influence of atmospheric CO_2 in certain areas, but this will likely only occur in fresh tidal creeks where oysters do not currently reside (Fig. 10e). Given the importance of atmospheric CO_2 in shaping future Ω_{Ca} conditions in the lower bay, reductions in anthropogenic carbon emissions will be necessary to lessen the projected impacts on carbonate chemistry in the Chesapeake Bay and globally.

Comparing our results to other studies examining the effects of acidification reveals that the Chesapeake Bay will likely acidify faster than the US West Coast. Siedlecki et al. (2021a) projected a decrease of 0.8-1.0 in Ω_{Ca} in the Northern California Current System between 2000 and 2100 under the same future emissions scenario used in our study, i.e., RCP8.5. Projections from the present work indicate a similar magnitude of reduction in the lower Chesapeake Bay over a shorter time period (50 years), suggesting a faster rate of acidification in the lower bay. Feely et al. (2009) also reported that projections for Ω_{Ca} reductions are slightly greater in the Atlantic than in the Pacific. The relative differences in rates of acidification should be considered, however, in the context of present-day Ω_{Ca}. The Pacific Ocean has a higher ratio of DIC:TA than the Atlantic, so present-day Pacific Ω_{Ca} is lower (Feely et al., 2004; Dunne et al., 2012). Therefore, while the Chesapeake Bay is acidifying faster, coastal Pacific waters may become undersaturated with calcite and aragonite sooner than in Chesapeake Bay. US West Coast shellfish mortality events associated with acidification or other climate change stressors may place increased pressure on US Atlantic fisheries to provide shellfish to the nation, highlighting the importance of climate change preparedness and resilience in the Chesapeake Bay region.

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592   While atmospheric CO_2 is primarily responsible for changes in Ω_{Ca}, nutrient reductions are also projected to worsen
593   carbonate chemistry conditions. Eutrophication can suppress acidification in surface waters by increasing primary production
594   (Borges and Gypens, 2010; Shen et al., 2019; Da et al., 2021), and when simulating a reduction in eutrophication via nutrient
595   management in our modeling study, the countering effect occurred. While the reduction in Ω_{Ca} from nutrient management is
596   minor compared to the projected impacts of CO_2-driven acidification, its small contribution may shift Ω_{Ca} conditions from
597   favoring net calcification to favoring net dissolution, demonstrating the importance of considering multiple drivers when
598   predicting exposure to ecologically relevant conditions of coastal acidification.

4.2 Future projections of oyster growth

600   Acidification, warming, and nutrient reductions are projected to affect shell and tissue growth of oysters in different
601   ways. Here, increased atmospheric CO_2 caused reductions in shell growth of Eastern oysters due to its negative effect on Ω_{Ca}
602   and thus calcification rates, which is consistent with experimental studies (Waldbusser et al., 2011; Gobler and Talmage, 2014;
603   Himes et al., 2024). Shell weight reductions from increased atmospheric CO_2 were driven by changes in calcification rate
604   alone, as tissue weight in *EcoOyster* is unaffected by Ω_{Ca} (Fig. 11d; Rivest et al., 2023). Experimental studies have identified
605   indirect physiological impacts of elevated CO_2 on juvenile/adult oyster metabolism, growth, and reproduction (Beniash et al.,
606   2010; Dickinson et al., 2012), suggesting that increased atmospheric CO_2 can sometimes influence tissue growth. Further
607   investigation is necessary in order to include the relationship between atmospheric CO_2 and oyster tissue growth in *EcoOyster*.
608   Biological and chemical reactions occur faster at higher temperatures, meaning calcification rates may be higher under future
609   warming conditions (Waldbusser et al., 2011), as long as Ω_{Ca} is still high enough to support calcification. Conversely, under
610   conditions of extreme low Ω_{Ca}, warming may exacerbate dissolution rates and shell weight reductions. Our results also show
611   that nutrient reductions will lead to reductions in shell weight, largely driven by a reduction in tissue weight resulting from
612   lower food availability (POC), rather than lower Ω_{Ca}.

613   While nutrient reductions are projected to have little influence on Ω_{Ca} in this study, their negative influence on food
614   availability may be detrimental to tissue growth in certain parts of the study region, particularly the York River. Our model
615   projections suggest that nutrient reductions may in some cases produce conditions that do not support any oyster growth along
616   the shoals of the York (Fig. 9c, f), a result of reductions in food availability that are predicted to be more substantial in the
617   tributaries than the mainstem region (Fig. 10c). Multiple studies have demonstrated that Eastern oysters and other calcifying
618   organisms perform better under acidification when they have sufficient food availability, as they are better able to keep up
619   with the energetic demands of environmental stress (Thomsen et al., 2015; Ramajo et al., 2016; Schwaner et al., 2023, Caillon
620   et al., 2025). Therefore, nutrient reductions will likely influence oyster growth under acidification stress by different
621   magnitudes in each tributary. When comparing the effects of local management actions to reduce nutrient runoff to the effects
622   of climate change (increased atmospheric CO_2 and warming), it is evident that, on average, climate change will have a much
623   greater negative influence on oyster growth (Table 4). However, the strong localized impacts of nutrient reductions in the York

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627 highlight the importance of examining the spatial variability of future changes in oyster growth. It is important for managers
628 to consider local conditions when assessing the effects of nutrient reductions on oyster production.

629   Increased water temperatures are projected to slow oyster growth in the future. Specifically, large reductions in tissue
630 weight are underpinned by three primary mechanisms: limitations on filtration at high temperatures (Loosanoff, 1958),
631 increased respiration rates (Dame, 1972), and reduced food availability. In *EcoOyster,* the optimal temperature for Eastern
632 oyster filtration is 27°C (Cerco et al., 2005; Jordan, 1987), and under warming, the frequency at which ambient summer
633 temperatures will surpass this optimal temperature will be higher (Fig. 5a), therefore causing more frequent declines in

634 filtration rate (Cerco et al., 2005; Fulford et al., 2007). There is no clear optimal temperature for oyster respiration, and
635 therefore it is assumed to increase exponentially with temperature (Hochachka and Somero, 2002). Thus, as oyster filtration
636 rates begin to decline at high temperatures, respiration rates will continue to rise and decrease the potential for tissue
637 accumulation (Fig. S1). Previous studies on juvenile Eastern oysters do not support a consensus on the relationship between
638 warming and tissue growth. Some report that growth is inhibited at higher temperatures (31°C, Stevens and Gobler, 2018;
639 30°C, Speights et al., 2017). In contrast, Talmage and Gobler (2011) found no significant influence of high temperature (28°C)
640 alone on tissue growth. The optimal temperature for oyster filtration may also vary among oysters, based on observations of
641 maximum filtration rates of adult Eastern oysters occurring between 28.1°C– 32°C (Loosanoff, 1958). Variation in
642 experimental design may have contributed to the contrast in results summarized here, in addition to the influence of local
643 adaptation (Burford et al., 2014). Other studies that incorporate higher temperature thresholds into their models predict
644 increases in oyster biomass under warming in Chesapeake Bay (Allen et al., 2023), underscoring the importance of properly
645 parameterizing growth models. Additional modeling studies should be conducted to test the sensitivity of oyster growth
646 changes to optimal temperature parameterizations. Due to a lack of consensus on temperature limits of Eastern oyster filtration,
647 further research is also needed to more robustly represent oyster filtration in bioenergetics models and improve predictions of
648 impacts of warming on oysters and their ecosystem services in the region.

649   Warming will likely have a negative effect on food availability for oysters. Compared to the effects of nutrient
650 reductions, warming will have a much more widespread influence on POC, causing reductions throughout the model domain
651 (Fig. 10b,c). Despite warming increasing rates of POC production via increased phytoplankton growth rates, factors such as
652 nutrient limitation and increased respiration rates will result in a net decrease in POC availability. In the tributaries, reductions
653 in food availability will be most widespread due to warming, but less extreme than those from nutrient reductions in the shallow
654 parts of the tributaries where oysters are affected. Remineralization of organic carbon in marine systems is temperature-
655 dependent (López-Urrutia et al., 2006), and as warming occurs, remineralization of detrital carbon to DIC in bottom waters
656 will occur at higher rates. As much of the lower bay is nutrient-limited (Zhang et al., 2021), phytoplankton growth rates will
657 not increase much from warming alone; therefore, increased remineralization will likely reduce the overall amount of food
658 available to oysters. Despite a similar average reduction in food availability being predicted for the future warming simulation
659 and managed nutrient reductions simulation, the influence of warmer temperatures will amplify the negative effects of reduced
660 food availability on growth. In this study, the critical temperature at which respiration rates exceed assimilation rates is

662 dependent on filtration. When food availability limits filtration, this critical temperature lowers, and the temperature threshold
663 for tissue loss is lowered. Experimental studies have demonstrated how organic carbon may be influenced by both warming
664 and acidification (Simone et al., 2021), but as these dynamics can differ based on nutrient availability, it is important to consider
665 how climate change will influence food webs and nutrient dynamics.

666 The projected mid-century reductions in oyster growth obtained from this analysis are consistent with the results of
667 other studies that examine oyster growth under similar projected climate change conditions. A study modeling oyster responses
668 to warming and low salinity co-stressors in Barataria Bay, LA, for example, predicts that under a warming and high flow
669 scenario (though without the effects of future nutrient reductions or atmospheric CO_2), oysters will experience widespread
670 mortality in fresher parts of the bay by the end of the century (Lavaud et al., 2021). Experimental studies have shown similar
671 negative effects of acidification, warming, lower food availability, and increased freshwater flow on oyster survival (La Peyre
672 et al., 2013; Rybovich et al., 2016; Lowe et al., 2019; Jones et al., 2019). Da (2023) found that the reductions in salinity and
673 Ω_{Ca} that result from high discharge events in the York River will increase in extent as climate change progresses and
674 increasingly threaten aquaculture production. In the Chesapeake Bay, extreme precipitation events are predicted to occur more
675 frequently with future climate change, however an overall decline in annual average precipitation is also predicted (St. Laurent
676 et al., 2021). As a result, the overall impact of freshwater from the land is not projected to change significantly in the future
677 (Hinson et al., 2023). Changes in precipitation were thus not simulated in this study, but future work could examine the
678 dynamics of climate change, salinity, Ω_{Ca}, and oyster growth in a year with more heavy rainfall events but lower annual rainfall.

679 **4.3 Influence of future changes in oyster growth on aquaculture**

680 Understanding the relative impacts of global climate change and local nutrient management actions on oyster growth
681 and survival will allow aquaculture producers to anticipate how their oyster stock may respond to these anthropogenic changes.
682 As the effects of climate change are subject to natural interannual variability, the magnitude of acidification and warming in a
683 given year will likely differ (Cai et al., 2021; Moore-Maley et al., 2016; Li et al., 2016), influencing oyster growth through
684 differing mechanisms. Smaller oysters resulting from slower growing times in a particularly warm year may present a different
685 challenge to growers than weak-shelled oysters in a year with lower Ω_{Ca} and average temperatures. Mortality may also become
686 a more urgent challenge as summer temperatures warm. A previous study examining commercial performance of Pacific
687 oysters in Brazil found that interannual variability in temperature, chl*a* abundance, and climate events influenced survival and
688 growth phase timing (Mizuta et al., 2012). High temperatures inhibited survival of oyster seed in that study, which frequently
689 occurs in Pacific oysters (*Crassostrea gigas)* during the summer months in Europe and California (Goulletquer et al., 1998;
690 Burge et al., 2007; Malhan et al., 2009). A similar phenomenon has been observed in Eastern oysters; however, mortality
691 events in this species have not been conclusively linked to warmer water temperatures (Guevelou et al., 2019; Biranik and
692 Allam, 2023), and the cause is yet to be resolved for either species, and remains an area of active research. Nonetheless, the
693 increasing occurrence of spring/summer mortality in Eastern oysters suggests that shifting the time of planting oysters on

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695 leases later in the year may help mitigate the risk of widespread mortality, although the economic tradeoffs involved in shifting
696 the growing season for oysters should be taken into account.

697 Future climate change and nutrient management are projected to worsen conditions for oyster growth, and the spatial
698 variation in these changes may unevenly influence aquaculture production. While reductions in shell and tissue growth are
699 predicted for nearly all regions where oysters grow, these changes will likely differ based on present-day environmental
700 conditions. Under present day conditions, the most oyster growth is projected to occur in regions with some of the highest
701 present-day Ω_{Ca} and the greatest projected Ω_{Ca} reductions, i.e., in the Rappahannock River and mainstem shoals. Some of the
702 most dramatic tissue and shell reductions are projected to occur in the York and upper Rappahannock, where reduced food
703 availability and low Ω_{Ca} will limit oyster filtration and shell growth. Oysters in parts of both the Rappahannock and York
704 Rivers will likely face mortality (represented by near complete depletion of oyster shell and tissue) by mid-century (Figs. 8,
705 9). However, these reductions will not be spatially uniform, underscoring the importance of oyster farm site selection within a
706 tributary. In contrast, oysters grown outside the tributaries are projected to exhibit a smaller decline in growth, indicating
707 greater future opportunity for oyster farming in these locations. Under the business-as-usual climate change trajectory analyzed
708 here, bottom conditions in the tributaries will be less suitable for oyster aquaculture by mid-century, and producers might
709 consider alternate farm locations or shifting production methods toward floating culture to avoid exposure to low Ω_{Ca}
710 conditions and access greater food availability.

711 Beyond reduced oyster growth, aquaculture operations may also be affected in the future by temporal changes in
712 optimal growing conditions. Due to the input of freshwater that lowers DIC and TA and increases pCO$_2$ (Cai et al., 2017; Cai
713 et al., 2021; Da et al., 2024), the greatest magnitude of Ω_{Ca} reductions occurs in spring. The majority of oyster growth is
714 projected to occur in the spring and summer (Fig. 5), so changes to growing conditions may be most consequential during

715 these warmer months. Deployment of oyster seed generally begins in the spring and continues into the summer, so it is
716 important for producers to be aware of ambient conditions being experienced by their newly deployed oysters. As spring
717 temperatures warm, phytoplankton blooms will likely occur earlier in the year, shifting the time when food availability is
718 highest (Da et al., 2021). Oysters deployed earlier in the year may benefit from greater food availability and perform better
719 than oysters deployed in July or August when waters are warmest. However, they may also face the challenge of spring/summer
720 mortality events, revealing the complexity of timing oyster deployment under worsening climate change conditions. For oyster
721 farms closer to freshwater sources, the combined effects of low Ω_{Ca}, low salinity, and high summer temperatures may severely
722 inhibit growth and extend time-to-market. This study focused on diploid oysters, since the laboratory experiments used to
723 develop the EcoOyster model parameterizations were conducted on diploids. Future laboratory experiments with triploids
724 would enhance our modeling efforts for triploid oysters, which are commonly used in aquaculture in the Chesapeake Bay
725 region.

4.4 Future work

Providing the aquaculture industry with the best existing estimates of climate change impacts to their operations will allow them to make more informed decisions about their future practices. This study used a 120-meter horizontal resolution model grid to examine near-lease-level effects of climate change and management actions on oyster growth in a section of the lower Chesapeake Bay. Similar studies with high resolution model grids in other systems will strengthen our understanding of how regional anthropogenic effects will influence the oyster aquaculture sector and could be used to identify areas of opportunity for new aquaculture practices (Swam et al., 2022; Palmer et al., 2021; Lavaud et al., 2024). The present study incorporated one Earth System Model and one emissions scenario; future work should quantify how these choices impact estimates of future Ω_{Ca} and oyster growth (e.g., Hinson et al., 2023). Future modeling studies should also incorporate other climate change impacts, such as sea level rise and increased storminess which are projected to influence conditions for oyster growth in the Chesapeake Bay region (Seneviratne et al., 2012; Lowe et al., 2019, Rybovich et al., 2016, Jones et al., 2019).

To improve estimates of shell and tissue growth of oysters under climate change, additional experimental studies should be conducted to reduce the data gaps that currently limit model formulations. Uncertainties in the functional relationships and rate parameters used in these models may lead to an inaccurate influence of some environmental variables on oyster growth. For example, results in this study may be particularly sensitive to the optimum temperature for filtration rate. Reductions in tissue weight are particularly dramatic when average temperature conditions at oyster lease sites remain above this optimal temperature from mid-June through late August, a vital time for oyster growth. Many physiological studies of temperature impacts on oyster filtration date back to the mid-to-late-20th century, and present-day seasonal extremes that coastal organisms experience may routinely exceed the maximum temperatures used in many of these earlier experimental designs. For example, Jordan (1987) used a maximum temperature of 27 °C, which was the ambient temperature when samples were collected in July from the Choptank River, MD. Between 1985 and 2014, bottom waters of the north mesohaline Bay main stem (closest to the Choptank) warmed 1.01 ± 0.13 °C during May to October (Hinson et al., 2022), and the present study predicts a 1.5 ± 0.26 °C increase across the entire model domain between 2017 and mid-century. To build stronger models of future climate impacts, and to expand scientific understanding of physiological limits of the Eastern oyster, future studies should re-examine temperature limitations on oyster filtration and respiration by using higher experimental temperatures.

As oyster growth is highly sensitive to food availability, improved measurements of particulate organic carbon in the region would fortify projections of oyster production under future climate change and nutrient management. Here, it is assumed that oysters feed on POC, a combination of plankton and detritus. However, average POC concentrations are highly spatiotemporally variable in the Chesapeake Bay due to eutrophication and algal blooms. In this study, POC was underestimated in the tributary channels; however, it is unclear how well POC was estimated in oyster growing areas, as *in situ* measurements are currently limited to stations in the channels during monthly or semi-monthly sampling cruises. More routine POC measurements, as well as measurements of POC in regions where oyster farming operations occur, are needed to

Deleted: As a result, growth in the model is sensitive to the simulation start date, and future studies should compare the influence of warming on growth in simulations that start at different times in the year.

19

763   verify the spatiotemporal dynamics of food availability. Improved measurements of oyster food availability would allow for

764   stronger model skill assessment and improved projections of oyster production.

765   **5 Conclusions**

766   This study predicts widespread reductions in Ω_{Ca} in the lower Virginia tributaries of the Chesapeake Bay by mid-

767   century, highlighting the use of high-resolution model projections to better understand present-day carbonate chemistry

768   conditions and to predict the effects of climate change on a region of high interest for aquaculture production. While similar

769   modeling studies have projected acidification conditions in coastal regions with 3D coupled models (Siedlecki et al., 2021a,b;

770   Fujii et al., 2023) or modeled oyster growth with remote-sensing data and dynamic energy budget models (Palmer et al., 2020;

771   Palmer et al., 2021; Bertollini et al., 2021), the present study projects both carbonate chemistry conditions and oyster

772   bioenergetics in the Chesapeake Bay with the highest resolution thus far. Specifically, widespread reductions in Ω_{Ca} will

773   negatively impact oyster growth, with implications for aquaculture operations and local and regional economies. As bottom

774   conditions worsen, altered site-selection for oyster farms or other adaptive measures will become imperative to sustain

775   production and reduce the impacts of low Ω_{Ca} on farmed oysters.

776   Increased atmospheric CO_2 and nutrient reductions are projected to inhibit oyster calcification, while warming and

777   nutrient reductions will reduce oyster tissue and shell growth due to limitations on filtration and lowered food availability.

778   While the effects of global climate change on oyster growth are projected to be much stronger overall than the effects of local

779   nutrient management, lowered food availability from nutrient reductions may have a strong influence on oyster growth in

780   certain parts of the study region. As a result, all areas will not be equally vulnerable to future changes in the atmosphere and

781   watershed. Understanding how individual drivers influence oyster growth is important for predicting effects on aquaculture

782   production in the context of interannual variability of climate change and nutrient management outcomes. While the negative

783   effects of temperature on growth were strong in this study, additional studies on Eastern oyster temperature limits are needed

784   to improve projections, particularly as summer mortality of oysters is already common. Increased *in situ* measurements of

785   biogeochemical variables and experimental studies on oyster physiology and bioenergetics will allow for improved projections

786   of mid-century conditions and their potential impacts on oyster growth and the aquaculture industry.

787

788   **Code Availability**

789   Model code will be available upon request.

790

791   **Data Availability**

792   Model output will be available with a DOI on William & Mary ScholarWorks.

793

794   **Author Contribution**

MAMF, EBR, MJB, and PS wrote the proposal and acquired the funding for the project; MAMF, PS, and FD developed the ROMS-ECB code; MJB developed the EcoOyster code; CC, MAMF, and EBR designed the experiment; CC ran model simulations, analyzed the output, created the figures, and wrote the manuscript draft; MAMF, EBR, MJB, PS, and FD reviewed and edited the manuscript.

Competing Interests

The authors declare that they have no conflict of interest.

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Table 1. Experimental design for future simulations conducted for comparison to reference run. Model forcings include a combination of 2017 (reference) and 2067 (future) inputs of atmospheric CO_2, atmospheric temperature, and terrestrial nutrient loadings.

Future Simulation Name	Atmospheric CO_2	Atmospheric Temperature	Terrestrial inputs
Combined Future	Future	Future	TMDL[a]
AtmCO2	Future	Reference	Reference
Temp	Reference	Future	Reference
TMDL	Reference	Reference	TMDL
AtmCO2 + Temp	Future	Future	Reference

[a]TMDL (Total Maximum Daily Load) forcing includes inputs of nitrate, ammonium, and dissolved and particulate organic matter under the assumption that the nutrient reduction goals (EPA, 2010) are met.

1335 Table 2. Model skill statistics (mean ± standard deviation) comparing bottom grid cells from the reference run to Chesapeake

1336 Bay Program observations from the same station location and time, within the bottom 10% of the water column.

Variable	Model	Observation	Model Bias	RMSD[a]
Temperature (°C) n = 130	17.0 ± 9	16.7 ± 9	+ 0.2	0.7
Salinity n = 127	13.9 ± 7	15.4 ± 7	-1.5	2.7
Oxygen (mg O_2 L^{-1}) n = 130	8.0 ± 2.3	7.2 ± 2.9	+0.9	1.3
pH n = 74	7.8 ± 0.4	7.6 ± 0.4	+ 0.2	0.4
TSS (mg L^{-1}) n = 74	44 ± 34	45 ± 54	-1.1	48.3
POC (g C m^{-3}) n = 74	0.7 ± 0.3	1.7 ± 2.1	-1.0	2.4

1337 [a]RMSD = root mean squared difference

Table 3. Bottom environmental variables for each model simulation (annual mean ± standard deviation) for grid cells that support oyster growth in the reference run (defined as those with greater than 1g dry tissue weight after one year of growth; Fig. S2). Analogous results averaged over all model grid cells are shown in Table S5.

Model Simulation	Temperature (°C)	Salinity	POC (g C m^{-3})	Ω_{Ca}	Dissolved Oxygen (mg O$_2$ L^{-1})	TSS (mg L^{-1})
Reference	17.0 ± 0.7	15.7 ± 2.1	1.12 ± 0.1	2.5 ± 0.49	9.1 ± 0.6	11.4 ± 5.8
Combined Future	18.5 ± 0.8	16.0 ± 2.1	1.03 ± 0.1	1.6 ± 0.35	8.7 ± 0.6	11.1 ± 5.9
AtmCO$_2$	17.0 ± 0.7	15.7 ± 2.1	1.12 ± 0.1	1.6 ± 0.35	9.1 ± 0.6	11.4 ± 5.8
Temp	18.5 ± 0.8	16.0 ± 2.1	1.07 ± 0.1	2.5 ± 0.41	8.8 ± 0.6	11.1 ± 5.9
TMDL	17.0 ± 0.7	15.7 ± 2.1	1.08 ± 0.1	2.4 ± 0.53	9.1 ± 0.6	11.2 ± 5.9
Temp + CO$_2$	18.5 ± 0.8	16.0 ± 2.1	1.07 ± 0.1	1.7 ± 0.33	8.8 ± 0.6	11.1 ± 5.9

1349 Table 4. Modeled oyster characteristics from the end of each simulation (mean ± standard deviation) over grid cells that support

1350 oyster growth in the reference run (defined as those with greater than 1g dry tissue weight after one year of growth; Fig. S2).

Model Simulation	Shell Weight (g)	Tissue Weight (g)	Shell Thickness (g mm^{-1})
Reference	16.8 ± 10.9	2.2 ± 1.5	0.18 ± 0.08
Combined Future	5.4 ± 5.7	0.9 ± 0.8	0.07 ± 0.05
AtmCO$_2$	10.5 ± 8.0	2.2 ± 1.5	0.12 ± 0.06
Temp	9.7 ± 9.1	1.2 ± 1.1	0.10 ± 0.07
TMDL	13.1 ± 8.2	1.7 ± 1.2	0.15 ± 0.06
Temp + CO$_2$	6.6 ± 7.1	1.2 ± 1.1	0.08 ± 0.06

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Figure 1. ROMS-ECBO model domain of Chesapeake Bay tributaries illustrating (a) bathymetry in meters and locations of Chesapeake Bay Program (CBP) water quality monitoring stations (red circles) and (b) bottom salinity zones. ¶
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Figure 2.¶

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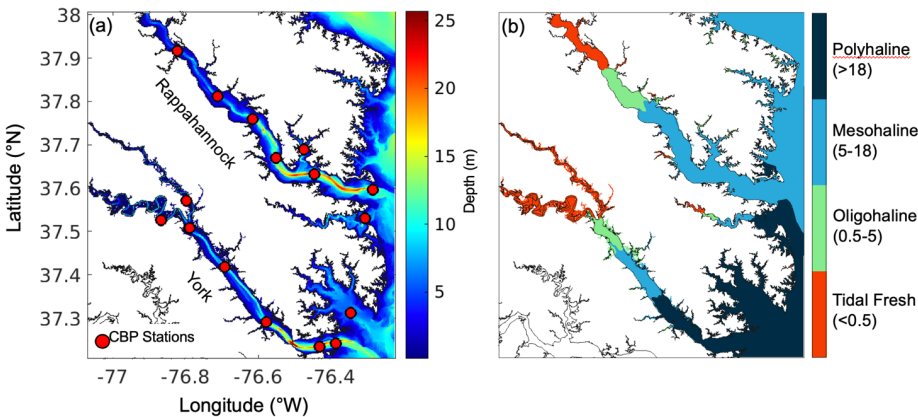
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1361 Figure 1. ROMS-ECBO model domain of Chesapeake Bay tributaries illustrating (a) bathymetry in meters and locations of
1362 Chesapeake Bay Program (CBP) water quality monitoring stations (red circles) and (b) bottom salinity zones.

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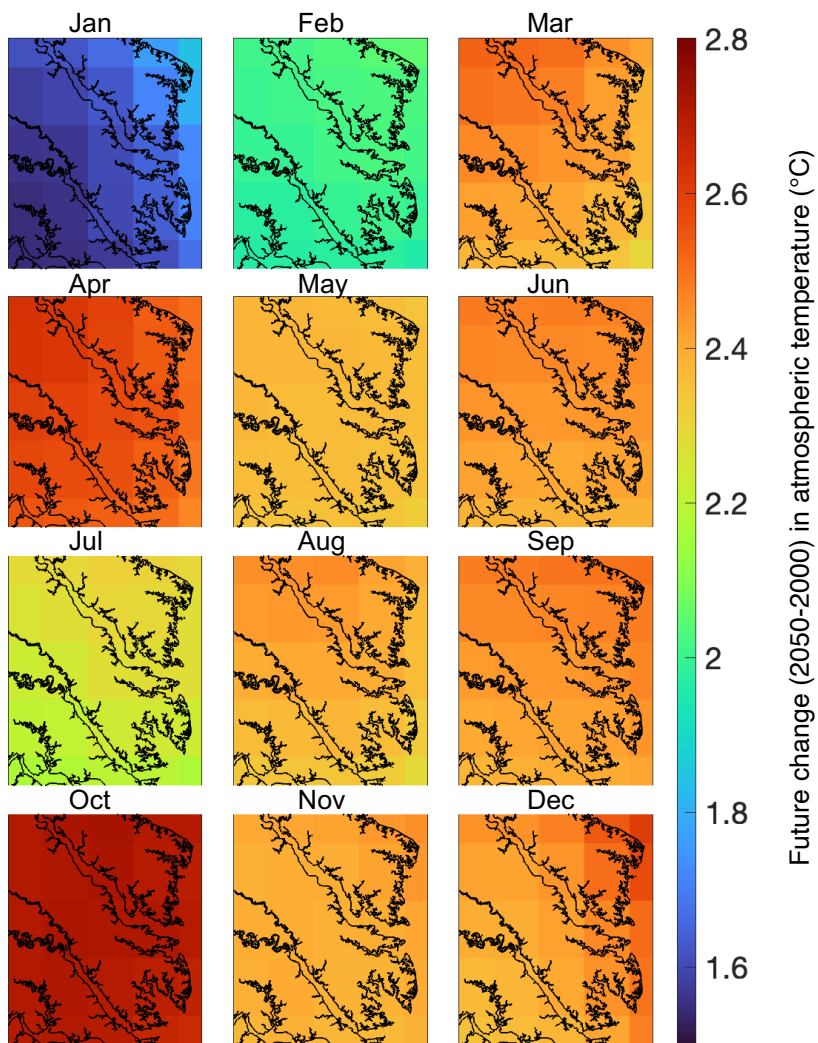
Figure 2. Monthly-averaged 50-year atmospheric temperature differences over the ROMS-ECBO model domain calculated as projections from 2050 minus those from 2000.
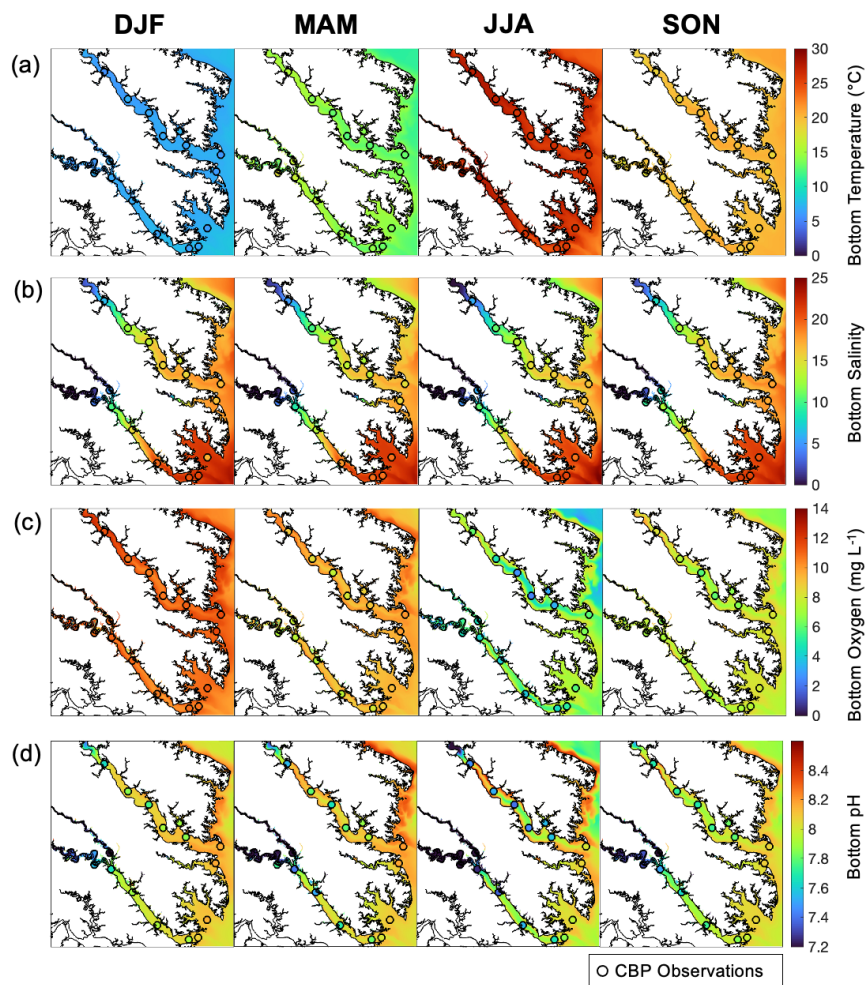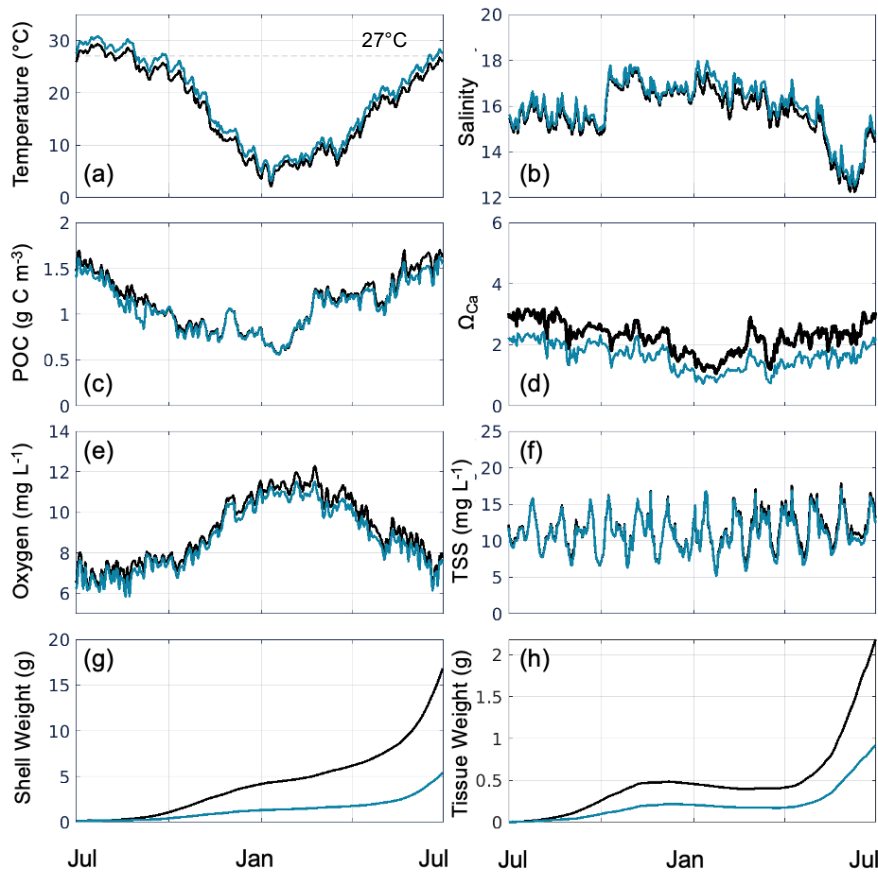
Figure 3. Seasonally-averaged bottom (a) temperature, (b) salinity, (c) dissolved oxygen, and (d) pH from the reference run. Circles represent seasonal decadal-averaged *in situ* observations at Chesapeake Bay Program stations (2010-2020). (DJF = winter, MAM = spring, JJA = summer, and SON = fall).

Figure 4. Seasonally-averaged bottom (a) POC and (b) TSS from output of the reference run of the ROMS-ECBO model. Circles represent seasonal decadal-averaged bottom measurements at Chesapeake Bay Program stations (2010-2020). (DJF = winter, MAM = spring, JJA = summer, and SON = fall).

Figure 5. Time series of daily bottom (a) temperature, (b) salinity, c) POC, (d) ΩCa, (e) oxygen, (f) TSS, (g) shell weight, and (h) tissue weight, averaged over grid cells that support oyster growth in the reference run, for the present-day reference run (black line) and Combined Future simulation (blue line). The horizontal dashed line in (a) represents the optimal filtration temperature in EcoOyster (27°C).

Figure 6. Annual mean bottom (a-c) POC and (d-f) Ω_{Ca} from (a,d) the present-day reference run, (b,e) the *Combined Future* simulation, and (c,f) the difference between (a) and (b), i.e., *Combined Future* minus reference.

Figure 7. (a-c) Shell weight and (d-f) tissue weight at the end of the one-year simulation from (a,d) the present-day reference run, (b,e) the *Combined Future* run, and (c,f) their difference, i.e., *Combined Future* minus reference.
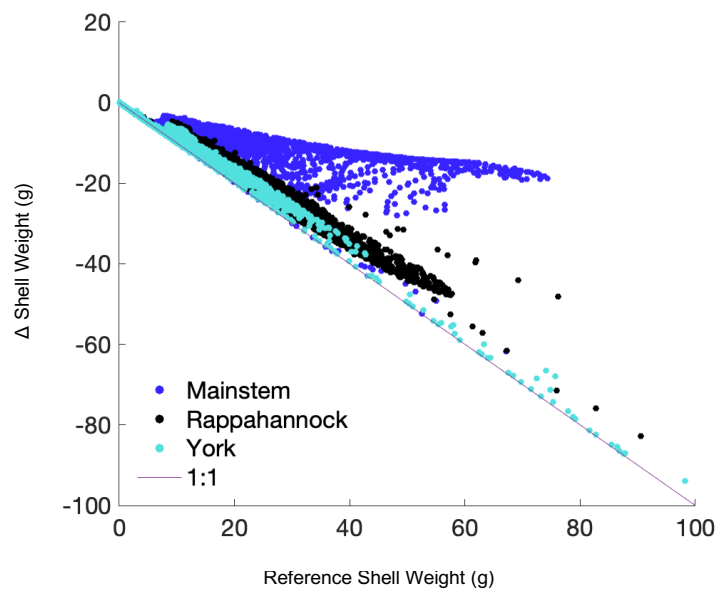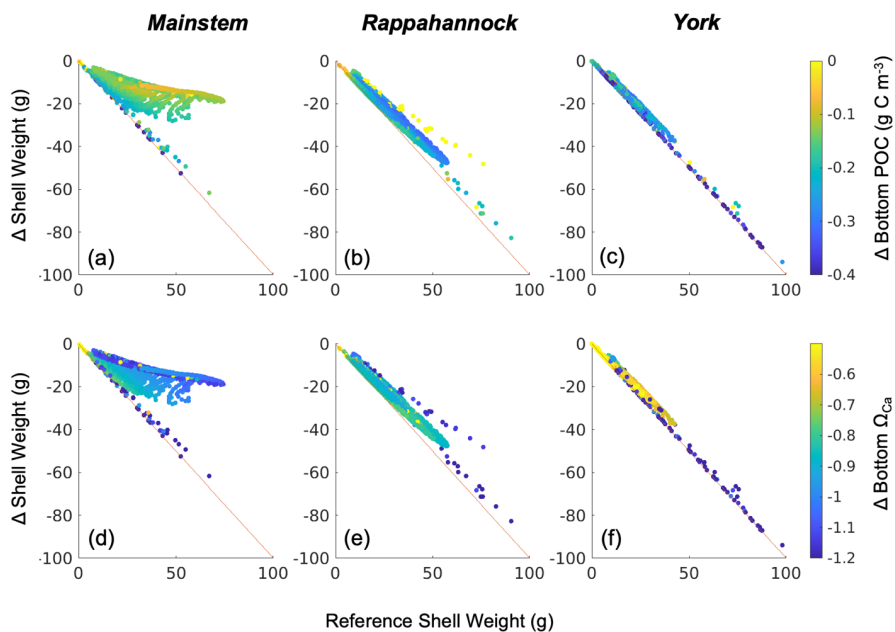
Figure 8. Difference in shell weight at the end of the one-year simulation between the *Combined Future* run and the reference run, i.e., *Combined Future* minus reference, colored by region. Each point represents a grid cell where oyster growth occurs in the reference run.

Figure 9. Difference in shell weight at the end of the one-year simulation between the *Combined Future* run and the reference run colored by (a-c) change in POC and (d-f) change in bottom Ω_{Ca} (i.e. *Combined Future* minus reference) for grid cells that support oyster growth. Results are presented for (a,d) the mainstem shoal only, (b,e) the Rappahannock River only, and (c,f) the York River only.
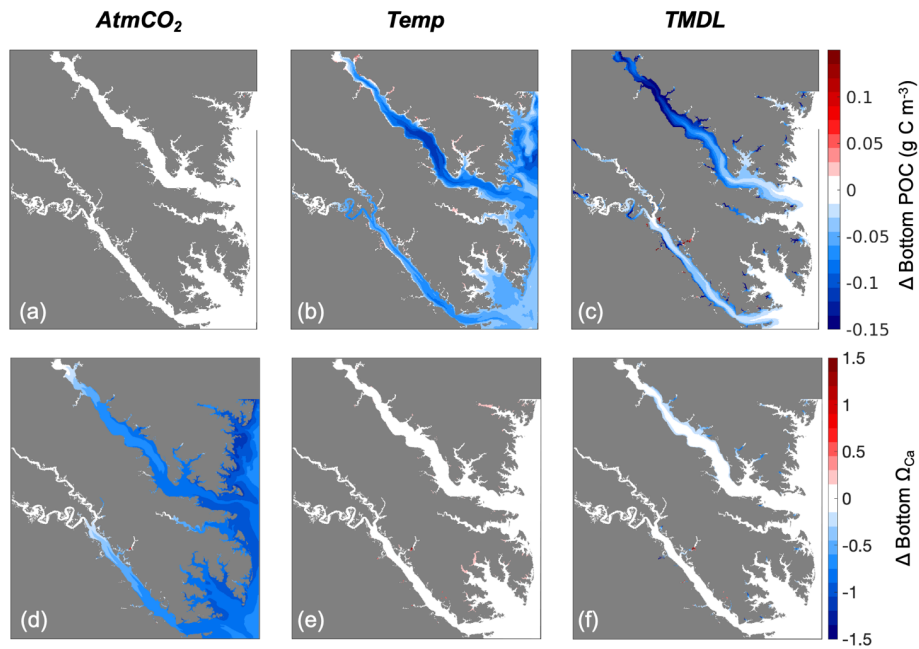
Figure 10. Differences in annual averaged (a-c) bottom POC and (d-f) bottom Ω_{Ca} (d-f) for three sensitivity experiments: (a) AtmCO$_2$, (b) Temp, and (c) TMDL. Differences represent future results minus those from the present-day reference run.
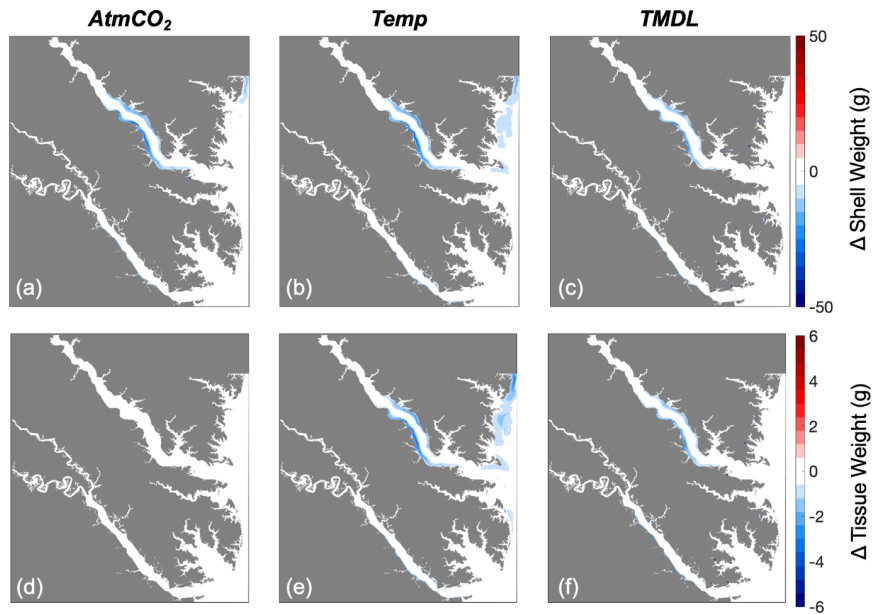
Figure 11. Differences in (a-c) shell weight and (d-f) tissue weight at the end of the one-year simulation for three sensitivity experiments: (a) AtmCO$_2$, (b) Temp, and (c) TMDL. Differences represent future results minus those from the present-day reference run.

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