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# Reviews and syntheses: Contribution of sulfate to methane oxidation in upland soils: a mini-review

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**Abstract.** Methane (CH<sub>4</sub>) is a potent greenhouse gas, its global warming potential is 25 times higher than carbon dioxide (CO<sub>2</sub>), and various environmental factors influence CH<sub>4</sub> oxidation in soil. Sulfate (SO<sub>4</sub><sup>2-</sup>) ion is the main component of atmospheric

- 20 deposition and has been increasing in recent years, it promotes CH<sub>4</sub> production and anaerobic CH<sub>4</sub> oxidation, however, the impact of SO<sub>4</sub><sup>2-</sup> on CH<sub>4</sub> oxidation remains inconclusive. Due to the limited research on the effects of SO<sub>4</sub><sup>2-</sup> on CH<sub>4</sub> oxidation, we synthesize current research on the effects of SO<sub>4</sub><sup>2-</sup> on CH<sub>4</sub> oxidation, examining both its direct impact and its influence on the dynamics of soil substances, and the potential
- 25 indirect effects of  $SO_4^{2-}$  on CH<sub>4</sub> oxidation. Through a literature review, we identified that  $SO_4^{2-}$  facilitates CH<sub>4</sub> oxidation within a range of 3-42%, moreover, it has been found that various physicochemical properties and processes in the soil are influenced by the addition of  $SO_4^{2-}$ , which in turn affects CH<sub>4</sub> oxidation. This review enhances our understanding of the role of  $SO_4^{2-}$  in promoting CH<sub>4</sub> oxidation and lays the foundation
- 30 for future studies aimed at validating these findings by quantifying CH<sub>4</sub> flux and oxidation rates, as well as elucidating the underlying microbial processes via experimental research. This review deepens the comprehension of atmospheric CH<sub>4</sub> flux and the global CH<sub>4</sub> cycle, particularly in the context of potential global environmental changes.





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## **1** Introduction

 $CH_4$  is an important greenhouse gas and has increased since pre-industrial times (Place, 2024; Praeg et al., 2016). Its global warming potential is 25 times higher than carbon dioxide (CO<sub>2</sub>) owing to its superior heat absorption efficiency (Yang et al., 2023).

- 40 Methanotrophs consume CH<sub>4</sub> under aerobic conditions (Le Mer and Roger, 2001), reducing CH<sub>4</sub> atmospheric concentration (Singh et al., 2010). Consequently, methanotrophs are crucial microbes that play an indispensable role in regulating and mitigating the greenhouse effect on Earth. Soil CH<sub>4</sub> oxidation is the sole known biological sink for atmospheric CH<sub>4</sub> (Ho et al., 2019; Murguia-Flores et al., 2018),
- 45 contributing to 5%–7% of the global annual atmospheric CH<sub>4</sub> uptake (Murguia-Flores et al., 2021). Upland soils are the primary biological CH<sub>4</sub> sink (Bodelier, 2011; Guo et al., 2023), owing to methanotroph-mediated CH<sub>4</sub> consumption reaching approximately 30 teragrams per year (Tg yr<sup>-1</sup>). This represents the second-largest atmospheric CH<sub>4</sub> consumption sink, surpassed only by hydroxyl radicals depletion (Deng et al., 2019).
- 50 CH<sub>4</sub> oxidation in soil is influenced by many factors, such as soil water content, soil texture, soil type, temperature, soil pH, soil inorganic nitrogen content, metal availability, etc. (Shukla et al., 2013). The ion SO<sub>4</sub><sup>2-</sup>, a significant ion component of acid deposition, may also significantly impact CH<sub>4</sub> oxidation.
- 55 Sulfur is transferred into the earth's surface through dry and wet deposition, with SO4<sup>2-</sup> being the dominant form. SO4<sup>2-</sup> deposition induces soil acidification (Huang et al., 2019), alters soil plant diversity (Li et al., 2022), affects microbial properties (Wang et al., 2018), limits grass yield potential (Klessa et al., 1989), and influences greenhouse gas emissions (Fan et al., 2017; Gauci et al., 2004; Schimel, 2004). The cycling of SO4<sup>2-</sup>
- 60 and CH<sub>4</sub> exhibits a significant inter-relationship.  $SO_4^{2-}$  suppresses CH<sub>4</sub> production (methanogenesis) primarily due to its thermodynamic and kinetic preference as an electron acceptor (Granberg. et al., 2001; Schimel, 2004), leading to decreased CH<sub>4</sub> emissions (Gauci et al., 2004).  $SO_4^{2-}$  has been shown to facilitate anaerobic CH<sub>4</sub>





oxidation in diverse ecosystems, such as oceans (Boetius et al., 2000), wetlands (La et
 al., 2022), and paddy fields (Fan et al., 2021), acting as a crucial electron acceptor.
 However, there is still a lack of a comprehensive review that summarizes the impact of SO4<sup>2-</sup> on soil CH<sub>4</sub> oxidation, a process that significantly contributes to reducing atmospheric CH<sub>4</sub> concentrations.

- 70 Previous studies indicated that SO<sub>4</sub><sup>2-</sup> addition may promote CH<sub>4</sub> oxidation by modulating the community structure or activity of methanotrophs in soils (Bradford et al., 2001b; Sitaula et al., 1995), while some studies indicated no significant impact on CH<sub>4</sub> oxidation (Bradford et al., 2001a). Upon reviewing the literature, we found that the enhancement of SO<sub>4</sub><sup>2-</sup> on CH<sub>4</sub> oxidation is prominent in numerous studies.
- 75 Furthermore, SO<sub>4</sub><sup>2-</sup> facilitates CH<sub>4</sub> oxidation within a range of 3-42% (Table 1). Thus, we hypothesize that SO<sub>4</sub><sup>2-</sup> may stimulate CH<sub>4</sub> oxidation. Nonetheless, the scarcity of data precludes a definitive conclusion regarding the direct effect of SO<sub>4</sub><sup>2-</sup> on CH<sub>4</sub> oxidation. Further, we reviewed references about the influence of SO<sub>4</sub><sup>2-</sup> on soil properties, substances, or biochemical processes, aiming to elucidate any indirect effect
- 80 on CH<sub>4</sub> oxidation through variations in soil substances or processes. Indeed, we find that SO<sub>4</sub><sup>2-</sup> may affect CH<sub>4</sub> oxidation both directly and indirectly. Based on the available literature, we infer that SO<sub>4</sub><sup>2-</sup> favors CH<sub>4</sub> oxidation. This review provides a comprehensive summary of the direct and potential indirect impacts of SO<sub>4</sub><sup>2-</sup> on CH<sub>4</sub> oxidation. The review underscores the viability of investigating the effect of SO<sub>4</sub><sup>2-</sup> on
- 85 CH<sub>4</sub> oxidation, providing a valuable reference for future experimental research. We will experimentally identify the enhancement of CH<sub>4</sub> oxidation following SO<sub>4</sub><sup>2-</sup> addition, along with elucidating the underlying microbial processes. This will contribute to a deeper understanding of the global CH<sub>4</sub> cycling in the context of increasing acid deposition in the future.

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- 2 The microbial CH<sub>4</sub> oxidation processes
- 2.1 CH<sub>4</sub> oxidation processes





Upland soils serve as sinks for atmospheric CH<sub>4</sub>, contributing about 6% to the total atmospheric CH<sub>4</sub> consumption (Liu and Greaver, 2009). Aerobic CH<sub>4</sub> oxidation,
mediated by aerobic methanotrophs (Chistoserdova et al., 2005), occurs in uplands, including grasslands (Kou et al., 2017; Zhang et al., 2020), forests (Jang et al., 2011; Mohanty et al., 2007), and agricultural soils (Ho et al., 2019), with oxygen serving as final electron acceptor. Soil CH<sub>4</sub> oxidation contributes the largest biological sink for

atmospheric CH4 (Pratscher et al., 2018). Moreover, the highest methanotrophic activity

- 100 is observed in the 6-10 cm layer of surface soil (King and Schnell, 1998; Nanba and King, 2000; Schnell and King, 1994). Methanotrophs oxidize CH<sub>4</sub> through dissimilatory and assimilatory pathways (Mancinelli, 1995). CH<sub>4</sub> is ultimately oxidized to CO<sub>2</sub> through the production of a series of intermediate substances (Fig. 1). In the dissimilatory pathways, CH<sub>4</sub> undergoes sequential oxidation: first, it is oxidized by
- 105 MMO to methanol (CH<sub>3</sub>OH, Fig. 1, path ①); methanol is oxidized by methanol dehydrogenase (MDH) to formaldehyde (HCOH, Fig. 1, path ②); formaldehyde is then oxidized by formaldehyde dehydrogenase (FADH) to formate (HCOOH, Fig. 1, path ③); and finally, formate is oxidized by formate dehydrogenase (FDH) to CO<sub>2</sub> (Fig. 1, path ④). Formaldehyde, a crucial intermediate in CH<sub>4</sub> oxidation, can be assimilated
- by methanotrophs to produce cellular biomass. The assimilating of formaldehyde occurs primarily through two pathways: the ribulose monophosphate (RuMP) cycle (Fig.1 path 5) and the serine cycle (Fig.1 path 6) (Kang and Lee 2016). As an aerobic CH<sub>4</sub> oxidation, CO<sub>2</sub>, water (H<sub>2</sub>O), and microbial biomass are ultimately produced, and the net reaction can be described as:

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#### $CH_4 + 2O_2 \rightarrow CO_2 + 2H_2O$ microbial biomass

## (1)

#### 2.2 Methanotrophs

Methanotrophs constitute a distinct subset of methylotrophs, uniquely dependent on the one-carbon compound CH<sub>4</sub> as their sole source of carbon and energy (Hanson and Hanson, 1996). Aerobic methanotrophs belong to various phyla, including *Proteobacteria, Verrucomicrobia* (Koo and Rosenzweig, 2021), and *Actinobacteria*. Proteobacterial methanotrophs are classified into type I, type II, and type X (Li et al.,





2020), based on their cell membrane arrangement, chemotaxonomic properties, physiological characteristics, and phylogenetic location. Type I and type X
125 methanotrophs belong to the *γ-Proteobacteria* class, while type II methanotrophs belong to the *α-Proteobacteria* class (Hanson and Hanson, 1996). Type I methanotrophs assimilate carbon using the RuMp pathway, type II methanotrophs using the serine pathway, and type X methanotrophs can use both pathways (Hanson and Hanson, 1996; Mancinelli, 1995). Verrucomicrobial methanotrophs, a newly discovered methanotrophs group (Op den Camp et al., 2009), are acidophilic and thermophilic extremophiles (Schmitz et al., 2021). They lack intracytoplasmic membranes, possess primarily saturated phospholipids, and use the serine and Calvin-Benson-Bassham for carbon fixation (Koo and Rosenzweig, 2021; Op den Camp et al., 2009). The

135 only recently been discovered and is preferred in low-pH and high-CH<sub>4</sub> environments (van Spanning et al., 2022). This group of methanotrophs expresses a comprehensive set of enzymes essential for aerobic CH<sub>4</sub> oxidation and uses the RuMp cycle to fix formaldehyde (van Spanning et al., 2022).

actinobacterial methanotrophs, Candidatus Mycobacterium methanotrophicum, has

- 140 Type I and type X methanotrophs belong to the *Methylococcaceae* family, wherear type II methanotrophs belong to two distinct families: *Methylocystaceae* and *Beijerinckiaceae* (Bowman, 2006; Yun et al., 2012). Verrucomicrobial methanotrophs are classified under the *Methylacidiphilaceae* family (Op den Camp et al., 2009). Up to now, a total of 21 genera and a species of methanotrophs have been identified, of which
- 145 20 genera belong to the Proteobacteria. Type I methanotrophs include Methylomonas, Methylobacter, Methylomicrobium, Methylosphaera, Methylosarcina, Methylohalobius, Methylothermus, Crenothrix, Clonothrix, Methylosoma, Methylovulum, Methylogaea, and Methylomarinum. Type II methanotrophs include Methylocystis, Methylosinus, Methylocella, Methylocapsa and Methyloferula. Type X methanotrophs include
- 150 Methylococcu and Methylocaldum. Verrucomicrobial methanotrophs consist of the genus Methylacidiphilum (Cai et al., 2022; Fenibo et al., 2023; Hanson and Hanson, 1996; Op den Camp et al., 2009). The recently discovered species Candidatus M.





methanotrophicum is classified with the *Mycobacterium* genus (van Spanning et al., 2022). There are two forms of MMOs, soluble cytoplasmic type monooxygenase

(sMMO) and particulate membrane-bound type monooxygenase (pMMO) (Koo and Rosenzweig, 2021). Figure 1 illustrates the differences between pMMO and sMMO, and their distribution across the identified 21 genera. The activity of methanotrophs is influenced by various environmental and climatic factors (Deng et al., 2019) and positively correlated with CH<sub>4</sub> oxidation (Tate, 2015).

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#### 2.3 The influence of sulfate

Acid rain pollution has been a significant environmental issue (Chen et al., 2020; Qi et al., 2022). SO4<sup>2-</sup> is the major ion in acid rain (Wright and Henriksen, 1978) and has profound impacts on substances and biochemical processes in soils. As a crucial

- 165 component of terrestrial ecosystems, soils serve as the ultimate receptor of acid deposition. The deposition of  $SO_4^{2-}$  can have numerous adverse effects on soil properties and functions. The application of  $SO_4^{2-}$  treatments led to a decrease in soil pH, as well as a reduction in the activities of soil enzymes such as cellulase, invertase, and polyphenol oxidase (Tie et al., 2020). The addition of  $SO_4^{2-}$  increased microbial
- 170 biomass carbon (MBC), microbial biomass nitrogen (MBN), and enzyme activities in the forest floor (Wang et al., 2018). SO4<sup>2-</sup> promoted methylation of mercury to form methylmercury (Jeremiason et al., 2006). However, the generation of methylmercury is influenced by various factors, including oxygen levels, temperature, pH, and the availability of labile organic carbon (Munthe et al., 2007). The significance of the
- 175 impact of  $SO_4^{2-}$  on the CH<sub>4</sub> cycle lies in its ability to influence both methanogenesis and anaerobic CH<sub>4</sub> oxidation. The addition of  $SO_4^{2-}$  to soil competitively displaced methanogens with sulfate-reducing bacteria (Granberg. et al., 2001), thus inhibiting CH<sub>4</sub> production (Eriksson et al., 2010; Granberg. et al., 2001). The anaerobic CH<sub>4</sub> oxidation process, fueled by  $SO_4^{2-}$ , is a crucial mechanism for mitigating methane
- 180 emissions from flooded ecosystems (Boetius et al., 2000; Fan et al., 2021; La et al., 2022), as sulfate (SO4<sup>2-</sup>) functions as the electron acceptor during anaerobic CH4 oxidation. However, a comprehensive review of the influence of SO4<sup>2-</sup> on aerobic CH4

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oxidation remains lacking. Therefore, synthesizing the literature regarding the influence of  $SO_4^{2-}$  on CH<sub>4</sub> oxidation is crucial for comprehending methane cycling in the context of  $SO_4^{2-}$ .

# 3 Soil CH<sub>4</sub> oxidation in response to SO<sub>4</sub><sup>2-</sup> addition

In a temperate mixed deciduous woodland, the cumulative uptake of CH<sub>4</sub> oxidation was observed to be 25% higher in the experimental group with H<sub>2</sub>SO<sub>4</sub> addition compared to 190 the control group during the final quarter of the study period (Bradford et al., 2001b). This increase in CH<sub>4</sub> oxidation could be attributed to alterations in the activity or community structure of methanotrophs caused by the addition of H<sub>2</sub>SO<sub>4</sub> (Bradford et al., 2001b; Sitaula et al., 1995). In Sitaula's study, the experimental group treated with H<sub>2</sub>SO<sub>4</sub> exhibited the lowest pH, specifically a pH of 3, at which the CH<sub>4</sub> oxidation rate

- 195 was observed to be the highest (Sitaula et al., 1995). However, given the addition of H<sub>2</sub>SO<sub>4</sub> in both experiments, it remains unclear whether the observed enhancement in CH<sub>4</sub> oxidation is primarily due to the decreased pH or the increased concentration of SO<sub>4</sub><sup>2-</sup> ions. Despite variable responses to pH in the temperate forest, a general trend indicates that CH<sub>4</sub> consumption trends increase with higher pH values (Brumme and
- 200 Borken, 1999; Silver et al., 1999). Furthermore, considering that the experimental soils were temperate forest and acidic (Bradford et al., 2001b; Sitaula et al., 1995), it is plausible that methanotrophs have adapted to these acidic conditions, potentially having a lower optimal pH range for CH<sub>4</sub> oxidation compared to non-acid soils. In a field experiment, a lower pH has been shown to increase CH<sub>4</sub> oxidation (Bradford et al.,
- 205 2001b; Sitaula et al., 1995). Additionally, if the hypothesis that  $SO_4^{2-}$  potentially facilitates CH<sub>4</sub> oxidation holds, high-pH soils are resilient to  $SO_4^{2-}$  adsorption post-sulfur deposition, thereby enabling acidic soils to absorb more  $SO_4^{2-}$  compared to alkaline soils. Consequently, more  $SO_4^{2-}$  is desorbed and transported in acidic soils, which could further augment CH<sub>4</sub> oxidation (Prietzel et al., 2004).

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The promotional effect of SO42- on CH4 oxidation can be further evidenced by





comparing its effects with those of other anions under the same cationic conditions. When  $NH_4Cl$  and  $(NH_4)_2SO_4$  were added to the soil at concentrations of 1µmol and 10

µmol NH4<sup>#</sup> per gram of fresh weight, respectively, both compounds caused the same

- 215 level of inhibition on net CH<sub>4</sub> oxidation. However, the inhibition observed at the lower concentration (1 µmol) was significantly lower than that at the higher concentration (10 µmol) (Adamsen and King, 1993; Bradford et al., 2001a; King and Schnell, 1998). NH<sub>4</sub><sup>+</sup> has been found to inhibit CH<sub>4</sub> oxidation (Bronson and Mosier, 1994; Dunfield and Knowles, 1995). However, the inhibitory effect of NH<sub>4</sub>Cl is greater than that of
- (NH<sub>4</sub>)<sub>2</sub>SO<sub>4</sub>, as SO<sub>4</sub><sup>2-</sup> may enhance the adsorption of NH<sub>4</sub><sup>+</sup> onto cation exchange sites in the soil (Bradford et al., 2001b; Gulledge and Schimel, 1998; King and Schnell, 1998). This reduced availability of NH<sub>4</sub><sup>+</sup> to compete with methanotrophs for MMO enzymes further intensifies the inhibitory effect of NH<sub>4</sub>Cl compared to (NH<sub>4</sub>)<sub>2</sub>SO<sub>4</sub>. In conclusion, SO<sub>4</sub><sup>2-</sup> served as a facilitator of CH<sub>4</sub> oxidation, mitigating the inhibitory effects of NH<sub>4</sub><sup>+</sup>
- 225 on this process. Benstead and King (2001) observed that, under equivalent soil acidic conditions, HNO<sub>3</sub> exerted a greater inhibitory effect on CH<sub>4</sub> oxidation than H<sub>2</sub>SO<sub>4</sub>. Similar results have been reported by Bradford et al. (2001a), who experimentally confirmed the inhibitory effect of nitrate (NO<sub>3</sub><sup>-</sup>) on CH<sub>4</sub> oxidation (Dunfield and Knowles, 1995; Wang and Ineson, 2003). When H<sub>2</sub>SO<sub>4</sub> and HNO<sub>3</sub> were added to the
- 230 soil to achieve H<sup>+</sup> concentrations of 10 and 1 µmol H<sup>+</sup> per gram of fresh weight (gfw), respectively, both acids inhibited CH<sub>4</sub> oxidation to a similar extent, albeit with H<sub>2</sub>SO<sub>4</sub> exhibiting a lesser inhibitory effect than HNO<sub>3</sub>. We hypothesize that SO<sub>4</sub><sup>2-</sup> may promote CH<sub>4</sub> oxidation, as evidenced by the findings of Benstead and King (2001) and Bradford et al. (2001a). Consequently, when H<sub>2</sub>SO<sub>4</sub> and HNO<sub>3</sub> are added to the soil, resulting in
- 235 equivalent acidic conditions, the inhibitory effect of H<sub>2</sub>SO<sub>4</sub> is less pronounced than that of HNO<sub>3</sub>. Nonetheless, future experiments are required to verify the direct effect of SO<sub>4</sub><sup>24</sup> on CH<sub>4</sub> oxidation.

Bradford et al. (2001a) observed no significant difference in CH<sub>4</sub> oxidation between
low (564 μM) and high (1408 μM) concentrations of H<sub>2</sub>SO<sub>4</sub> compared to the control.
Another possibility is that the H<sub>2</sub>SO<sub>4</sub> concentration, relative to the previous study, led





to  $SO_4^{2-}$  -stimulating CH<sub>4</sub> oxidation. Alternatively, Bradford et al. (2001a) suggested that the short duration of H<sub>2</sub>SO<sub>4</sub> addition to the soil might have prevented it from stimulating CH<sub>4</sub> oxidation (Bradford et al., 2001a). Similarly, Hu et al. (2018) reported

245 no significant effect of  $SO_4^{2-}$  on CH<sub>4</sub> oxidation. Due to the scarcity of studies investigating the direct effect of  $SO_4^{2-}$  on CH<sub>4</sub> oxidation, no definitive conclusion regarding its impact could be drawn. Therefore, it is necessary to conduct experimental studies to elucidate the influence of  $SO_4^{2-}$  on CH<sub>4</sub> oxidation and explore its underlying mechanisms.

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# 4 Indirect effects of SO4<sup>2-</sup> on CH<sub>4</sub> oxidation

The influence of  $SO_4^{2-}$  on CH<sub>4</sub> oxidation has been relatively understudied in research. Thus, we conducted a literature review focused on studies involving the addition of  $SO_4^{2-}$ . Our objective was to demonstrate that  $SO_4^{2-}$  can indirectly influence CH<sub>4</sub>

- 255 oxidation through its effects on soil physicochemical properties and changes in other soil constituents. Indeed, our review uncovered several significant findings that support this hypothesis. CH<sub>4</sub> oxidation is influenced by various environmental factors (Kou et al., 2017; Shukla et al., 2013), including SO<sub>4</sub><sup>2-</sup>. However, SO<sub>4</sub><sup>2-</sup> may affect CH<sub>4</sub> oxidation directly (Fig. 2 path c), stemming from variations in methanotrophs activity
- and community structure in response to SO4<sup>2-</sup> (Fig. 2 path e) (Bradford et al., 2001b;
  Sitaula et al., 1995). Alternatively, it may affect CH<sub>4</sub> oxidation indirectly by influencing soil physical properties (Fan et al., 2017), microbial substrates (Bjorneras et al., 2019;
  Palmer et al., 2013; Xu et al., 2017), soil nutrition content (Islam, 2012), and soil bacterial metabolism (Hu et al., 2013; Lv et al., 2014; Sogn and Abrahamsen, 1998;

265 Veraart et al., 2015) (Fig. 2).

First, the addition of  $SO_4^{2-}$  altered soil physical properties (Fig. 2 path a), i.e., specifically reducing soil pH (Fig. 2 path 1) (Fasth et al., 1991; Tie et al., 2020) and elevating soil redox potential (Eh) and soil oxidation capacity (OC) (Fig. 2 path 2)

270 (Fan et al., 2017). An increase in soil acidification, resulting from the augmented base





cation leaching associated with SO4<sup>2-</sup> addition (Hu et al., 2013), led to a decrease in the pH of forest soils (Fasth et al., 1991; Tie et al., 2020). Generally, CH4 consumption is greater at higher pH conditions in general forest soils (Brumme and Borken, 1999; Silver et al., 1999). Therefore, the reduction in soil pH due to SO4<sup>2-</sup> addition may result
in a decrease in CH4 oxidation. However, in acidic soils, a decrease in pH has been observed to increase CH4 oxidation (Sitaula et al., 1995). Consequently, when considering the impact of SO4<sup>2-</sup> addition on CH4 oxidation, it is crucial to account for the acid-base status of the study soils (Fig. 2 path i), as methanotrophs exhibit different pH preferences in acidic and alkaline environments (Shukla et al., 2013). SO4<sup>2-</sup> addition

- has been found to increase soil Eh and soil OC in subtropical forest soil (Fan et al., 2017). Furthermore, the amount of O<sub>2</sub> in the soil was closely correlated with soil Eh, and a decrease in O<sub>2</sub> content can lead to a reduction in soil Eh (Zausig et al., 1993). Therefore, an increase in O<sub>2</sub> content results in an elevation of soil Eh (Fig. 2 path g). Methanotrophs are obligate aerobes (Teh et al., 2005), requiring MMO to bind O<sub>2</sub> for
- the initial step of CH<sub>4</sub> oxidation (Mancinelli, 1995; Shukla et al., 2013). Additionally, CH<sub>4</sub> oxidation exhibited a significant positive correlation with O<sub>2</sub> levels (Mancinelli, 1995). Consequently, the increase in soil Eh and OC, resulting from the elevated O<sub>2</sub> content due to SO<sub>4</sub><sup>2<sup>2</sup></sup> addition, may enhance CH<sub>4</sub> oxidation (Fig. 2 path h).
- 290 Second, SO<sub>4</sub><sup>2</sup> could alter the soil microbial substrate (Fig. 2 path b), specifically by decreasing soil DOC concentration (Fig. 2 path ③) (Bjorneras et al., 2019; Palmer et al., 2013) and di-O-alkyl C content (Fig. 2 path ④) (Xu et al., 2017). Sullivan et al. (2013) reported that DOC stimulates CH<sub>4</sub> oxidation in semiarid soils; however, SO<sub>4</sub><sup>2-</sup> addition reduces DOC concentrations (Bjorneras et al., 2019; Palmer et al., 2013).
- 295 Consequently, it can be inferred that SO4<sup>2-</sup> addition may indirectly inhibit CH<sub>4</sub> oxidation by decreasing soil DOC concentration (Fig. 2 path 1). In a subtropical forest, SO4<sup>2-</sup> addition has been shown to increase the activity of gram-negative bacteria in soil by reducing the litter di-O-alkyl C (Fig. 2 path ④ and j) (Xu et al., 2017). Di-O-alkyl C is a component of SOC. SOC degradation occurs more readily with a higher percentage





of di-O-alkyl C (Huang et al., 2021). Conversely, when the content of di-O-alkyl C is low, SOC degradation proceeds relatively slowly in soil, resulting in a greater availability of substrate for utilization by microorganisms, including methanotrophs. Methanotrophs, which are gram-negative bacteria (Schimel and Gulledge, 1998), may experience an increase in activity when SO<sub>4</sub><sup>2-</sup> is added. This enhancement of methanotrophs activity (Fig. 2 path k), can ultimately promote CH<sub>4</sub> oxidation (Fig. 2 path q). However, further experiments are required to elucidate the mechanism of how SO<sub>4</sub><sup>2-</sup> affects the activity of methanotrophs.

Third, SO<sub>4</sub><sup>2-</sup> can alter soil nutrition content (Fig. 2 path d), specifically increasing soil

- Cu availability (Fig. 2 path <sup>(5)</sup>) (Islam, 2012), P content (Fig. 2 path <sup>(6)</sup>) by enhancing acid phosphatase activity (Lv et al., 2014; Veraart et al., 2015), and Al<sup>3+</sup> toxicity (Fig. 2 path <sup>(7)</sup>) (Hu et al., 2013; Sogn and Abrahamsen, 1998). Cu is a crucial component in CH<sub>4</sub> oxidation processes when it is utilized by methanotrophs in their molecular machinery, synthesized from metabolized CH<sub>4</sub> through the secretion of methanobactin
- 315 into the environment. This process facilitates the oxidization of CH4 to methanol (Dassama et al., 2016). It was anticipated that methanobactin secreted by methanotrophs during CH4 oxidation would facilitate Cu uptake (Knapp et al., 2007); however, the specific mechanism of how methanobactin affects Cu uptake remains unclear (Fig. 2 path n). An increase in Cu availability results in the overexpression of
- 320 protein-mediated steps in the conversion of CH<sub>4</sub> to CO<sub>2</sub>, including the synthesis of cell walls, lipids, and membranes (DiSpirito et al., 2016). Furthermore, Cu may act as a stimulant for CH<sub>4</sub> oxidation (Ho et al., 2013). Therefore, the addition of SO<sub>4</sub><sup>2-</sup> may indirectly enhance CH<sub>4</sub> oxidation by augmenting the availability of soil Cu (Fig. 2 path m). A positive correlation has been found between P and CH<sub>4</sub> oxidation in soils (Veraart
- 325 et al., 2015; Zhang et al., 2020). P can potentially enhance the activity of soil methanotrophs (Fig. 2 path p) (Zhang et al., 2011), with an increase in soil P content achieved through the hydrolysis of organic compounds, including nucleic acids, phospholipids, and phosphate esters, by acid and alkaline phosphatases (Veraart et al.,





2015). The addition of SO4<sup>2-</sup> accelerated acid phosphatase activity, thereby increasing
soil P content (Lv et al., 2014). Therefore, we hypothesize infer that SO4<sup>2-</sup> may indirectly enhance CH<sub>4</sub> oxidation through the augmentation of soil P content, subsequently promoting the activity of methanotrophs in the soil (Fig. 2 path p and q). It is well-established that Al<sup>3+</sup> inhibits CH<sub>4</sub> oxidation (Tamai et al., 2007; Tamai et al., 2003). Additionally, soil acidification resulting from SO4<sup>2-</sup> addition has been shown to
intensify the toxicity of Al<sup>3+</sup> in forest soils (Hu et al., 2013; Sogn and Abrahamsen,

1998). We hypothesize that the addition of  $SO_4^{2-}$  may indirectly inhibit CH<sub>4</sub> oxidation by enhancing the toxicity of soil  $Al^{3+}$  (Fig. 2 path o).

Finally, the addition of SO4<sup>2-</sup> alters soil microbial metabolism (Fig. 2 path f),

- specifically including an up-regulation of the tricarboxylic acid (TCA) cycle (Fig. 2 path <sup>®</sup>) (Wang et al., 2021). Methanotrophs utilize the TCA cycle during the metabolism of the serine cycle to assimilate CH<sub>4</sub> (Trotsenko and Murrell, 2008). A decrease in CH<sub>4</sub> oxidation rates leads to a decline in CH<sub>4</sub> assimilation (Roslev et al., 1997), whereas an increase in CH<sub>4</sub> assimilation in soils may trigger an up-regulation of the TCA cycle (Fig. 2 path s). Thus, the observed up-regulation of the TCA cycle
- following SO<sub>4</sub><sup>24</sup> addition can be attributed to the enhancement of CH<sub>4</sub> oxidation (Fig. 2 path c), which subsequently leads to an increase in CH<sub>4</sub> assimilation (Fig. 2 path r). However, the precise mechanisms of how SO<sub>4</sub><sup>2-</sup> ultimately impacts CH<sub>4</sub> oxidation and methanotrophs remain to be experimentally verified.

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#### **5** Implication

The microbiological processes of methanogenesis and CH<sub>4</sub> oxidation are influenced by numerous factors, significantly impacting the global CH<sub>4</sub> cycle.  $SO_4^{2-}$  plays a pivotal role in global acid deposition, with annual deposition rates ranging from 141.64  $\pm$  120.04 TgS a(year)<sup>-1</sup> (Gao et al., 2022). Various studies have demonstrated that  $SO_4^{2-}$ 

120.04 TgS a(year)<sup>-1</sup> (Gao et al., 2022). Various studies have demonstrated that SO<sub>4</sub><sup>2-</sup> inhibits soil methanogenesis (Eriksson et al., 2010; Granberg. et al., 2001; Schimel, 2004), reducing CH<sub>4</sub> emissions. SO<sub>4</sub><sup>2-</sup> also stimulates anaerobic CH<sub>4</sub> oxidation (Boetius)





et al., 2000; Fan et al., 2021; La et al., 2022), enhancing  $CH_4$  consumption and reducing emissions. Nevertheless, there is a scarcity of studies investigating the impact of  $SO_4^{2-}$ 

360 on aerobic CH<sub>4</sub> oxidation. By synthesizing the available literature, we observed that SO<sub>4</sub><sup>2-</sup> enhances CH<sub>4</sub> oxidation by up to 42%, suggesting that increased SO<sub>4</sub><sup>2-</sup> deposition may promote soil CH<sub>4</sub> oxidation, thereby reducing CH<sub>4</sub> emissions to the atmosphere and influencing the CH<sub>4</sub> cycle, ultimately affecting global warming.

## 365 6 Conclusions

Upon reviewing recent studies, we concluded that  $SO_4^{2-}$  promotes CH<sub>4</sub> oxidation and predicted that its definitive impacts on CH<sub>4</sub> oxidation may be observed in the coming decades. Mechanistically,  $SO_4^{2-}$  can either directly enhance CH<sub>4</sub> oxidation or influence it by modulating the activity of methanotrophs, pH, Eh, soil OC, DOC concentration,

di-O-alkyl C amount, Cu availability, P content, Al<sup>3+</sup> toxicity, and the soil TCA cycle. Therefore, SO<sub>4</sub><sup>2-</sup> reduces CH<sub>4</sub> emissions while enhancing its consumption, which is beneficial for CH<sub>4</sub> mitigation in the context of heightened acid deposition. Moreover, the relationship between SO<sub>4</sub><sup>2-</sup> and CH<sub>4</sub> oxidation may be mediated by intermediate substances, specifically, SO<sub>4</sub><sup>2-</sup> affects CH<sub>4</sub> oxidation by modulating methanotrophs or

375 including alterations in other soil components, requiring further experimental validation.

#### Data availability

All raw data can be provided by the corresponding authors upon request.

## 380 Author contribution

Rui Su finished writing; Kexin Li, Nannan Wang, Fenghui Yuan, Yunjiang Zuo, Ying Sun, and Liyuan He gave constructive comments and revised the structure and content of the article; Xiaofeng Xu and Lihua Zhang reviewed and edited the manuscript.

#### **Competing interests**

385 The authors declare that they have no conflict of interest.





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395 of Energy Office of Science.





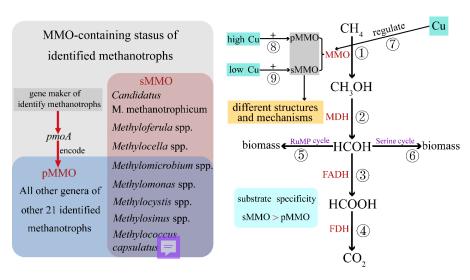


Figure 1: Diagram of the methane oxidation process and MMO-containing status of identified methanotrophs. ① CH4 is oxidized to methanol (CH3OH) by MMO;
2 CH3OH is oxidized to formaldehyde (HCOH) by methanol dehydrogenase (MDH);
3 HCOH is oxidized to formate (HCOOH) by formaldehyde dehydrogenase (FADH);
4 HCOOH is oxidized to CO2 by formate dehydrogenase (FDH); ⑤ HCOH is assimilated into cellular biomass via the RuMp cycle; ⑥ HCOH is assimilated into cellular biomass via the Serine cycle; ⑦ Cu controls two MMOs expression (Hakemian and Rosenzweig, 2007); ⑧ High Cu concentration regulates pMMO expression in soil (Hakemian and Rosenzweig, 2007); ⑨ Low Cu concentration

- regulates sMMO expression in soil (Hakemian and Rosenzweig, 2007). pMMO and sMMO have different structures and mechanisms, and sMMO has broder substrate specificity than pMMO. Most methanotrophs possess pMMO, so the *pmoA* gene is a
- 410 key gene for detecting methanotrophs in environment samples.





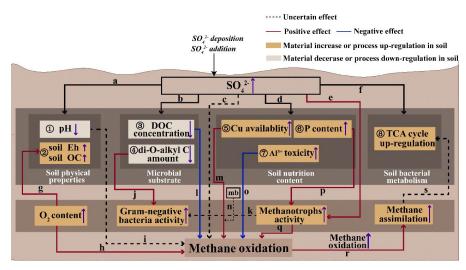


Figure 2. Conceptual diagram illustrating SO<sub>4</sub><sup>2-</sup> affected CH<sub>4</sub> oxidation through

- direct or indirect ways in soil. ①SO4<sup>2-</sup> decreases soil pH (Fasth et al., 1991; Tie et al., 2020);②SO4<sup>2-</sup> increases soil Eh (redox potential) and soil OC (oxidation capacity) (Fan et al., 2017); ③SO4<sup>2-</sup> decreases soil DOC concentration (Bjorneras et al., 2019; Palmer et al., 2013); ④SO4<sup>2-</sup> decreases soil di-O-alkyl C amount (Xu et al., 2017);⑤SO4<sup>2-</sup> increases soil Cu availability (Islam, 2012); ⑥SO4<sup>2-</sup> increases soil P content by
- 420 increasing soil acid phosphatase activity (Lv et al., 2014; Veraart et al., 2015); ⑦SO4<sup>2-</sup> increases soil Al<sup>3+</sup> toxicity (Hu et al., 2013; Sogn and Abrahamsen, 1998); ⑧SO4<sup>2-</sup> induced soil TCA cycle up-regulate (Wang et al., 2021); a. Changes in soil physical properties due to increased soil SO4<sup>2-</sup> content; b. Changes in soil microbial substrate due to increased soil SO4<sup>2-</sup> content; c. SO4<sup>2-</sup> may promote CH<sub>4</sub> oxidation; d. Changes
- in soil nutrition content due to increased soil SO4<sup>2-</sup> content; e. SO4<sup>2-</sup> affects the activity or community size of methanotrophs in soils (Bradford et al., 2001b; Sitaula et al., 1995); f. Change in soil bacterial metabolism due to increased soil SO4<sup>2-</sup> content; g. Elevated O<sub>2</sub> content increases soil Eh (Zausig et al., 1993); h. Elevated O<sub>2</sub> level stimulates soil CH<sub>4</sub> oxidation (Mancinelli, 1995); i. Decreased pH may inhibit or
- 430 stimulate soil CH<sub>4</sub> oxidation (Sitaula et al., 1995); j. Decreased di-O-alkyl C amount increases soil gram-negative bacteria activity (Xu et al., 2017); k. Increased activity of gram-negative bacteria may increase activity of methanotrophs; l. Decreased DOC concentration may inhibit soil CH<sub>4</sub> oxidation (Sullivan et al., 2013); m. Elevated Cu availability stimulates soil CH<sub>4</sub> oxidation (Ho et al., 2013); n. mb (methanobactin) is





- 435 expected to accelerate Cu uptake (Knapp et al., 2007); o. Elevated Al<sup>3+</sup> toxicity inhibits soil CH<sub>4</sub> oxidation (Tamai et al., 2007; Tamai et al., 2003); p. Elevated P content increases soil methanotrophs activity (Zhang et al., 2011); q. Elevated methanotrophs activity stimulates soil CH<sub>4</sub> oxidation (Bradford et al., 2001b; Sitaula et al., 1995); r. Increased CH<sub>4</sub> oxidation may increase CH<sub>4</sub> assimilation (Roslev et al., 1997); s.
- 440 Methanotrophs undergo the TCA cycle during CH<sub>4</sub> metabolism and may cause upregulation of the TCA cycle in soil.

Study site	Sulfate concentration	CH4	Effect	Reference
		concentration		
Perridge Forest	${ m H}_2{ m SO}_4~(50~{ m Kg~S~ha^{-1}})$	Ambient air	25 % increased	[28]
Perridge Forest	H <sub>2</sub> SO <sub>4</sub> (5mM)	Ambient air	no effect	[30]
	(NH4)2SO4 (5mM)	Ambient air	no effect	
Maine forest	Na2SO4 0.5µg S g <sup>-1</sup> soil	250ppm	3% increased	[37]
Norway Scots	H <sub>2</sub> SO <sub>4</sub> pH3	Ambient air	42% increased	[29]
Pine forest				
Birch taiga	Na2SO4 2.8 µmol S g <sup>-1</sup> soil	4ppm	no effect	[63]
	K <sub>2</sub> SO <sub>4</sub> 2.8 µmol S g <sup>-1</sup> soil	4ppm	no effect	

Table 1. Promotion effect of sulfate on methane oxidation in diverse biome soils.

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