



1 **Root zone in the Earth system**

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23 **Abstract:** The concept of “root zone” is widely used in hydrology, agronomy, and land surface
24 process studies. However, the root zone still lacks a precise definition. More essentially, the
25 importance of root zone in the Earth system science is largely under explored. In addition, the
26 methodology to estimate root zone is still controversial. In this study, we firstly attempted to
27 clarify the definition of the root zone by comparing with various “similar” terms, such as rooting
28 depth, soil depth, vadose zone, rhizosphere, and critical zone, to bridge the gaps within and
29 between traditional disciplinary boundaries. Secondly, we found that, from a hydrological and thus
30 water-centric perspective, the root zone holds profound implications across all the spheres of the
31 Earth system, including biosphere (living organisms), hydrosphere (water), pedosphere (soil),
32 lithosphere (rock), and atmosphere (air), through various exchange fluxes of mass and energy. The
33 role of the root zone in the Anthropocene is elaborated as well, including the intensifying impacts
34 of climate change and agriculture, along with implications for nature-based solutions and
35 planetary stewardship. Thirdly, for root zone estimation, we underscore that the theoretical
36 foundation of the traditional reductionist approach to understand and model the root zone is



37 problematic due to the complex and dynamic nature of root zone functions. We advocate for a
38 shift towards a holistic ecosystem-centered perspective, which offers a more realistic, simplified,
39 and dynamic representation of the root zone in Earth system science.

40

41 **1 Introduction**

42 During the Devonian Period, 416 to 360 million years ago, roots were an early development in the
43 plant's evolution history, even before leaves (Kenrick and Strullu-Derrien, 2014). The
44 development of root systems was a critical biological innovation that subsequently enabled land
45 plants to spread and colonize continental interiors, where previously, their limited simple rhizoid-
46 based rooting systems confined them to areas immediately adjacent to bodies of water (Smart et
47 al., 2023). The rapidly extending functionality and complexity of the root system performs
48 essential functions necessary for the survival and development of plants, such as anchoring them
49 to a substrate, absorbing water and nutrients, and storing photoassimilates (Shekhar et al., 2019).
50 The rapid expansion of terrestrial plants, enabled by the newly developed rooting system, caused a
51 dramatic increase in physical and chemical rock weathering and terrestrial photosynthesis
52 capability, which massively drew down the atmospheric CO₂. This likely triggered the ice age and
53 the mass extinction in the late Devonian (Smart et al., 2023). The rooting system, at its first
54 appearance, completely transformed the entire biosphere and their abiotic environment on the
55 planet.

56 Roots act as pioneers of biological activity, representing a definitive line between sterile sediments
57 and living soils. Fossil evidence showed that deeper soils and larger plants appeared almost
58 simultaneously (Kenrick and Strullu-Derrien, 2014). There is positive feedback between soils and
59 plants, mediated by the roots. Larger plants generally require more stable and comparatively
60 deeper soils, and plant roots contribute to soil formation. Thus, in some places, the deep roots for
61 groundwater access may have acted as a direct agent to accelerate weathering and soil formation
62 through physically widening rock fractures, inviting water and hastening other physical and
63 chemical weathering, and increasing the acidity of soil and groundwater that further enhanced
64 chemical weathering (Maeght et al., 2013). Thus, the evolution of plant root systems had not only
65 influenced the local-scale soil microenvironment and individual plant life cycles, but also the
66 weathering of the lithosphere and the formation of the pedosphere. Hence, root system
67 development has had far reaching impacts on the evolution of the atmosphere, the hydrosphere
68 and the entire Earth system (Hetherington, 2019).

69 The root zone, where plants developed a place to adapt to the environment through long-term
70 evolution, forms the substrate of the terrestrial ecosystem and hence is a crucial element of the
71 critical life zone on Earth (Banwart et al., 2017). It is well demonstrated that the soil forms an
72 ecosystem full of micro-biotic and macrobiotic life (Ponge, 2015). Fungi forming dense
73 underground networks live in symbiosis with vegetation, exchanging nutrients for carbon, which
74 makes them responsible for the larger part of subterranean carbon storage (Domeignoz-Horta et
75 al., 2021). Through evolution and natural selection, the ecosystem has found ways to make best
76 use of the climatic and geological resources.

77 Plants together with micro-organisms alter soil hydraulic properties and water-cycle processes



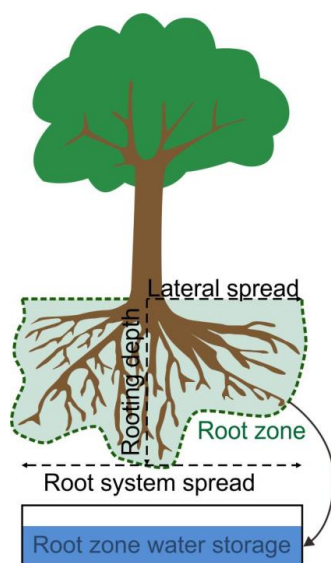
78 through creating soil macropores and adding soil organic matter. Macropores created through
79 living or decayed roots are large cracks or pores that facilitate infiltration and fast water drainage,
80 thus reducing the time to soil saturation that may cause root rotting. The cementation by organic
81 matter from dead roots affects the internal structure of soils. Thus, soils with high organic matter
82 commonly have high porosity and high field water holding capacity to constrain water in soil
83 pores. Thus, the root zone is the result of long-term co-evolution of the biosphere, hydrosphere,
84 and pedosphere, depending on the climatic and geological boundary conditions.

85 In the past decades, the root zone has attracted growing attention from different disciplines,
86 including hydrology, agronomy, plant biology, soil physics, atmospheric science, and landscape
87 engineering, for different reasons. However, there is still a lack of review on the basic concept of
88 the root zone, particularly in the context of Earth system science. Thus, there is an urgent need to
89 clarify the definition of the root zone and to bridge the knowledge gaps between traditional
90 disciplinary boundaries.

91 **2 What is the root zone?**

92 The first challenge lies in the difficulty of defining a common, unambiguous language to
93 accurately communicate between the variety of disciplines involved in root zone research and
94 subsequently between the broader disciplinary fields of hydrology, ecology, climatology,
95 pedology, agronomy, horticulture, forestry, etc.

96 In its most general form, the root zone is the upper part of the unsaturated zone that supports
97 vegetation rooting, where water and nutrients are potentially available to support plants (Sprenger
98 et al., 2019). The storage of moisture in the root zone is difficult to determine because of the
99 complex relation between water, air and the soil matrix with its ill-defined network of pores and
100 fissures. Generally, it is represented as a volume per unit area and thus as an average depth (Figure
101 1). Initially, the root zone has been considered as the substrate for agriculture. However, in a wider
102 context, for natural terrestrial ecosystems as a whole, the root zone storage is the buffer that the
103 ecosystem created to provide continuous access to water. Although these processes are confined to
104 a relatively thin layer, and only stores 0.13% of the total freshwater on Earth (McCartney et al.,
105 2021), it is a critical factor in controlling land surface hydrological response, land-atmospheric
106 moisture exchange, and biogeochemical processes (Zehe et al., 2019).



107

108 Figure 1. A sketch of the definitions of root zone and root zone water storage capacity (adapted
109 from Tumber-Dávila et al., 2022)

110 The hydrological and land surface relevant magnitude of the ecosystem's root zone can be
111 described by the root zone water storage capacity, i.e. S_{umax} , that represents the maximum
112 subsurface moisture volume that can be accessed by the vegetation's roots (Gao et al., 2014;
113 Stocker et al., 2023). Similar to artificial reservoirs designed by humans to bridge dry periods
114 requiring investment of materials and energy, ecosystems also need carbon, energy and nutrients
115 to develop their root zone water storage, as a buffer to overcome periods of drought. Conversely,
116 S_{umax} is the maximum water deficit in the root zone, which occurs when all available moisture has
117 been consumed after a critical drought period.

118 From a hydrological perspective, S_{umax} controls the partitioning of precipitation into infiltration to
119 meet soil evaporation and plant transpiration demand and runoff generation (i.e., subsurface storm
120 flow, percolation into groundwater, and discharge as baseflow). Thus, S_{umax} is the key parameter
121 determining ecosystems' resilience to drought and runoff generation.

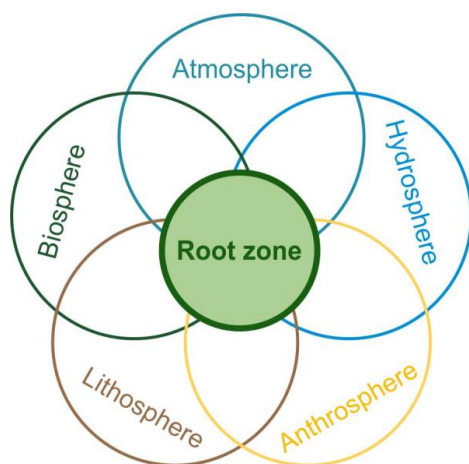
122 In Section 3, we will further clarify the definition of root zone and root zone storage capacity by
123 comparing it with other "similar" terms in different disciplines.

124 3 The role of the root zone in the Earth system

125 Earth System Science is a rapidly emerging interdisciplinary endeavor aimed at understanding the
126 structure and functioning of the Earth as an integrated, complex, and adaptive system (Steffen et
127 al., 2020). Inspired by early work on the biosphere–geosphere interactions and novel perspectives
128 such as the Gaia hypothesis (Lovelock, 1979), Earth System Science emerged in the 1980s
129 following demands for a new 'science of the Earth'. Interestingly, in the latest conceptual model of
130 the Earth System (Steffen et al., 2020), the root zone, as the interface among roots, soils,
131 vegetation, lithosphere, freshwater, biosphere, and the production and consumption in human



132 society, plays a central role, as the crucial element linking the Earth's geosphere and
133 anthroposphere (Figure 2).



134

135 Figure 2. Root zone is the cross-section of biosphere, lithosphere, atmosphere, hydrosphere, and
136 anthroposphere in the Earth system.

137 **3.1 Root zone and biosphere**

138 **3.1.1 Functions of roots for plants**

139 Roots are essential organs for plants to survive and reproduce in a great diversity of habitats.
140 Plants rely on roots to uptake water and nutrients—two fundamental elements for plant
141 performance; roots also anchor and support plants in the ground; roots store food made in the
142 leaves by photosynthesis. Roots' sensing of local resources, such as water, nutrients and
143 phosphate, enables better access to the heterogeneously distributed resources in the complex of
144 pores and fissures in the substrate. Roots combined with a fully integrated vascular system were
145 essential to the evolution of large plants, enabling them to meet the requirements of anchorage and
146 the acquisition of water and nutrients (Boyce, 2005).

147 The root zone determines the resilience of ecosystems to droughts and climate change. Natural
148 ecosystems optimally adjust their root systems to their environment, controlled mainly by climate
149 and topography (Gao et al., 2014; Fan et al., 2017). The vegetation, i.e., the ensemble individual
150 plants present at any moment, is a manifestation of their successful adaptation to past conditions -
151 as otherwise these specific individual plants would not have survived and would not be there. This
152 indicates that plants optimized (likely the results of long-term trial and error in evolution) their
153 root-accessible water storage according to a cost minimization strategy, i.e., to meet canopy water
154 demand as well as nutrient requirements with minimal carbon allocation to roots that in turn frees
155 resources for above-surface growth that is needed to survive in competition for sunlight with other
156 plants.

157 **3.1.2 Rooting depth**

158 There are many root traits, including root biomass, root turnover, root/shoot ratio, vertical root



159 distribution, and maximum rooting depth (Figure 3a). Rooting depth is one of the most basic plant
160 functional traits influencing ecosystem resilience, plant biogeography, pedogenesis, and carbon
161 cycle.

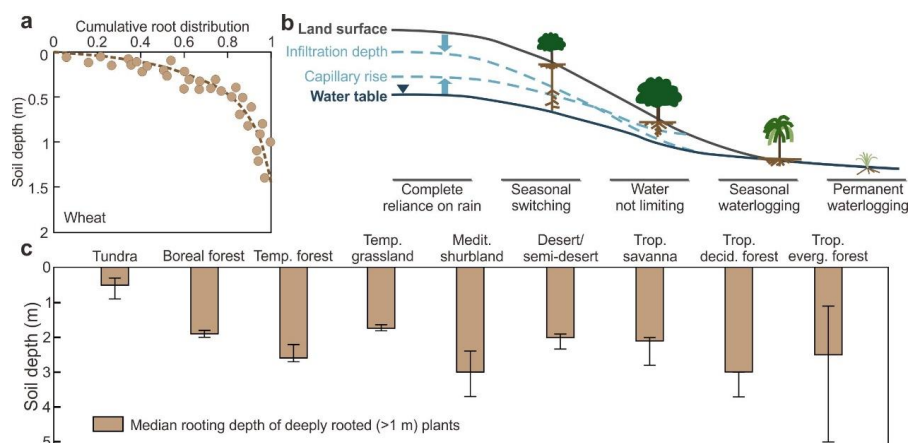
162 Measuring rooting depth and other traits accurately is one of the most time-consuming and
163 difficult activities in ecological and agricultural research, especially if it is a field study. The
164 methodology also varies depending on the plant community. For example, the rooting depth of
165 grass and agricultural crops is measured by collecting soil cores; while coarse roots of trees are
166 estimated by allometric equations, and fine roots of trees are estimated by soil cores. A global
167 rooting depth synthesis study covers 2200 world-wide root observations (Fan et al., 2017) reflects
168 still a small sample compared to the number of trees 3.04 trillion (Crowther et al., 2015), and the
169 even larger numbers of grass and agriculture crops. Thus, globally, we have very limited
170 knowledge of the root system.

171 From a limited number of studies, we know that rooting depth has non-unique associations with
172 climate and biome types (Jackson et al., 1996; Schenk, 2002). For instance, roots are shallow in
173 boreal biomes on thinly thawed soils; roots of annual crops starting from seeds each season reach
174 only shallow depths; and deep roots are found in arid, semiarid, and season-arid climates (Figure
175 3c). On average, woody plants such as trees and shrubs tend to be more deeply rooted than grasses
176 and forbs. Plant species differ in root growth and distribution in the soil profile, especially under
177 stress conditions. Interestingly, both the shallowest and the deepest roots are found in dry biomes
178 (Fan et al., 2017).

179 Topography also has a large impact on rooting depth. Along a topographic gradient, root–water
180 relation shifts systematically; on excessively drained uplands, the water table is deep or absent,
181 and rooting depth is limited to infiltration depth/frequency. With the increase of Height Above the
182 Nearest Drainage (HAND), rooting depth commonly increases with the increase of water table
183 depth due to water limitation (Figure 3b; Fan et al., 2017).

184 Generally, climate, topography, and root genetic code determine rooting depth. However,
185 accidental discoveries of >70-m-deep roots in wells and >20-m-deep roots in caves offer glimpses
186 of the enormous flexibility of root response to its environment.

187



188

189 Figure 3. (a) Cumulative root distribution as a function of soil depth for wheat as an example (Fan
190 et al., 2016). (b) Schematic of rooting depth and water table profile along a drainage gradient (Fan
191 et al., 2017). (c) Comparison between estimated rooting depths of global vegetation types and
192 sampling depths used in quantitative studies of vertical root distributions (Schenk and Jackson,
193 2002).

194 3.1.3 Root zone \neq rooting depth

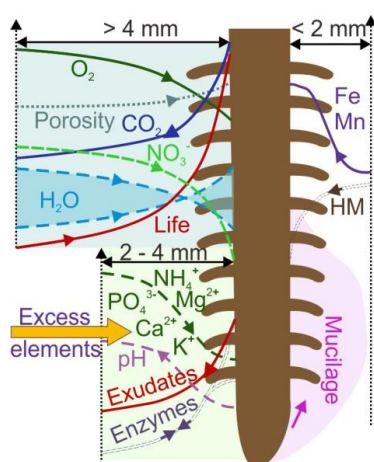
195 It is important to note that the root zone is not necessarily proportional to the rooting depth. While
196 rooting depth only describes the vertical extension, the root zone reflects the root distribution over
197 the entire profile. The root zone, moreover, accounts for lateral root extent and root density. Root
198 distribution patterns in the soil profile are important determinants of the ability of plants to acquire
199 nutrients and water necessary for growth. For example, an ecosystem covered by deep-rooting
200 vegetation with roots with low density may have a smaller root zone than one covered by
201 vegetation with shallow, high-density roots (Singh et al., 2020; van Oorschot et al., 2021). The
202 vertical and lateral distribution and the root density are essential elements of the root zone. Even
203 with detailed root distribution data, we cannot determine the root zone merely using root traits,
204 due to the complex pores and fissures in soils and bedrock, and the complex hydrodynamic
205 gradients, which will be discussed in Section 3.1.4 and 5.1.

206 3.1.4 Root zone \neq rhizosphere

207 The rhizosphere is the zone around the root where microorganisms and processes important for
208 plant growth and health are located (Hiltner, 1904; Hartmann et al., 2008). It includes roots, soil
209 and the space in between, which is usually 1-4 mm (Kuzyakov and Razavi, 2019) (Figure 4). The
210 plant root-soil interface is a dynamic region in which numerous biogeochemical processes take
211 place driven by physical activity, including water and nutrients (Bengough, 2011). The
212 rhizosphere supports microorganisms' growth (e.g., 10^7 - 10^{12} bacteria population in 1 g of
213 rhizosphere soil) and thus stimulates biochemical processes in the root zone (Hinsinger et al.,
214 2009). Root-associated microbial communities can strongly influence plant survival, phenology,
215 and expression of functional traits (Kuzyakov and Razavi, 2019). Also, as much as half of this
216 photosynthetic carbon can be lost from the soil by respiration within hours or days.



217 There are similarities between the rhizosphere and the root zone. Both zones are not stationary and
218 not fixed. However, the rhizosphere is the soil volume directly around the living roots and usually
219 covers just a few millimeters from the root surface. It is used to analyze direct interactions with
220 microorganisms and to assess nutrient depletion zones. The root zone, however, is used in
221 agronomy, hydrology, and land surface process issues, to assess the total amount of water and
222 nutrients potentially available for plants. So in volume, there is a 2-3 orders of magnitude
223 difference between the rhizosphere and the root zone. The rhizosphere is more like a biology and
224 pedology concept on a micro-scale (Kuzyakov and Razavi, 2019), while the root zone is used for
225 larger-scale hydrology and land surface studies in the Earth system.



226

227 Figure 4 Generalization of rhizosphere extents and gradient types for the most investigated
228 parameters: Gases, Root exudates, Nutrients and Excess elements, pH and Eh, Enzyme activities
229 and microorganisms (from Kuzyakov and Razavi, 2019).

230 3.2 Root zone, lithosphere, and pedosphere

231 3.2.1 Roots as active agents of soil formation

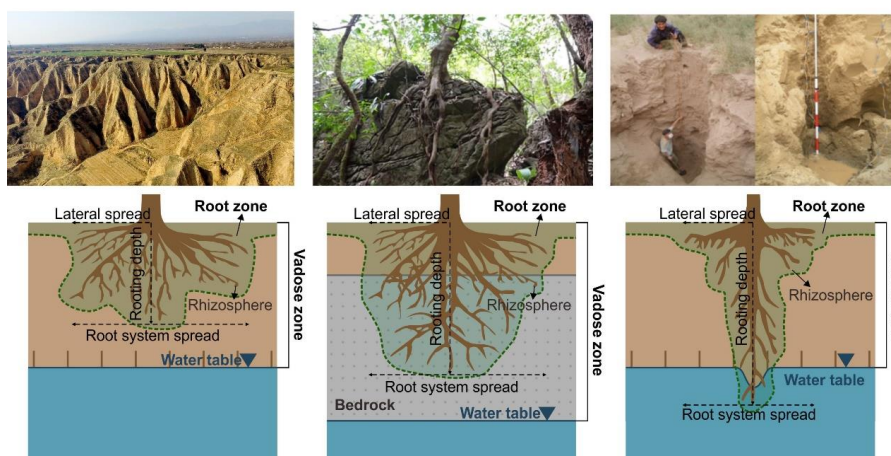
232 Roots can sense the surrounding heterogeneous structure and avoid obstacles, both at the
233 microstructural (volumes less than mm^3) and the macrostructural scales (Wang et al., 2020). Roots
234 can open up the substructure by acting like a “wedge” (Pawlik et al., 2016). Roots promote rock
235 weathering and play important roles in soil formation. Roots' associated fungi contribute to
236 weathering and enlargement of rock fractures as well (Schenk, 2008).

237 Roots together with climate, topography, parent material and time, determine soil properties
238 (Huggett, 2023). Roots prefer to grow in existing pore networks to penetrate deeper into soil and
239 bedrock fissures. Due to the physical, chemical and biological activity of roots, the lithosphere has
240 changed over time leading to increased root zone depth (Hetherington, 2019). Moreover, soils
241 receive organic matter through the interaction with roots and fungi, exchanging carbon for
242 minerals, fixing carbon underground and significantly altering soil properties (Cotrufo and
243 Lavalée, 2022).



244 **3.2.2 Root zone \neq soil depth**

245 The depth to bedrock forms the lower boundary of the soil, which determines soil depth. Although
246 the root zone is strongly connected to the soil, it does not necessarily correlate to soil depth
247 (Figure 5). For example, in most cases, especially in deep soil regions, such as the Loess Plateau
248 in China, the root zone is limited to the most active layer of the topsoil, where most hydrological
249 and biogeochemical processes occur, while the soil depth extends much deeper. In karst and other
250 mountainous regions, the root zone does penetrate the bedrock and roots access water in fissures
251 (McCormick et al., 2021). In arid climates, roots can even reach the deep groundwater; thus, in
252 this case, the root zone also includes part of the groundwater (see Singh et al., 2020).



253
254 Figure 5 Root zone is different from soil depth.

255 **3.2.3 Root zone \neq critical zone**

256 The definition of the critical zone is still controversial. The most widespread definition is from the
257 National Research Council in the US (NRC, 2001): the critical zone is the near-surface layer of the
258 Earth, including vegetation, soil, water, and rocks, all of which are essential elements for
259 supporting life. Based on this definition, the critical zone ranges "from the canopy top to the base
260 of the groundwater zone". There is high consensus on the top boundary of the critical zone, i.e.,
261 the top of the vegetation canopy. Its lower boundary, however, has many different versions, e.g.,
262 the base of the groundwater zone (NRC, 2001), the bottom of the weathering zone (Guo and Lin,
263 2016), the base of active groundwater (Fan, 2015), the storage of fresh groundwater (<50 years
264 old) (Gleeson et al., 2016). Whatever the definition of the critical zone, the root zone is part of it.
265 Moreover, the root zone is the most active layer in the critical zone, including all elements of
266 vegetation, soil, water, air, and rocks. It is the crucial element of the critical zone, connecting the
267 full range from canopy to groundwater.



268 3.3 Root zone and hydrosphere

269 3.3.1 Root zone and hydrological processes

270 In the terrestrial water cycle, there is a hierarchy of processes. Starting with precipitation of
271 moisture, it is first intercepted by vegetation and ground cover from which it can return to the
272 atmosphere by evaporation or infiltrate into the soil and percolate to the root zone, where it is
273 stored for plant utilization, and from where it can percolate deeper to the groundwater, or flow
274 laterally to the drain. The water that does not infiltrate can flow overland to the drain. All these
275 processes are to a larger or lesser extent dominated by the amount of water already present in the
276 root zone. One can say that the root zone soil moisture is the key agent determining the
277 partitioning of precipitation into evaporation, runoff, groundwater recharge and the build-up of
278 storage. Over longer timescales, under equilibrium conditions, storage change may be disregarded,
279 and precipitation goes to either runoff or evaporation, with the root zone exerting a strong control
280 on this partitioning (Figure 6). Most land surface water fluxes (including evaporation and runoff)
281 and water storages (including soil moisture, vegetation, and groundwater) are affected by the root
282 zone. Consequently, the root zone plays a major role in regulating the land surface water budget,
283 and much of this is due to the biophysical effects of plant roots on the surrounding soil
284 (Bengough, 2011). The moisture flow from the land surface to the atmosphere through vegetation
285 root water uptake (i.e. transpiration) is globally the largest water flux from terrestrial ecosystems
286 at about 60-70% of total evaporation (Schlesinger and Jasechko, 2014; Coenders-Gerrits et al.,
287 2014; Lian et al., 2018). This flux is largely responsible for the recycling of moisture in the
288 atmosphere where about 60% of all precipitation originates from terrestrial evaporation (Van der
289 Ent et al., 2010)

290 For runoff generation, there are two classical mechanisms, i.e., saturation excess flow and
291 infiltration excess flow (Beven, 2012). Root zone characteristics determine both. Saturation excess
292 flow is the dominant mechanism in humid catchments. Based on this theory, there is no runoff
293 generation when the root zone still has a moisture deficit, i.e. root zone moisture is below a critical
294 threshold (so called field capacity). Once the root zone moisture crosses this threshold, outflow
295 pathway becomes activated, and runoff is generated. This mechanism is also named fill-and-spill
296 (McDonnell et al., 2021) or store-and-pour (Phillips, 2022). Due to landscape heterogeneity
297 (mainly topography), the runoff threshold has a spatial distribution function so that saturation and
298 runoff generation does not happen all at the same time in an entire catchment. Usually, riparian
299 zones (close to the drain) are saturated first, where saturation excess flow first occurs.
300 Subsequently, the contributed areas expand with the increase of precipitation amount and root
301 zone moisture. Thus, the saturation excess flow mechanism is also called the variable contribution
302 area theory. The concept can be functioned as:

$$303 \quad Q(t) = f(P(t), S_u(t), S_{umax}, \beta) \quad (1)$$

304 where $Q(t)$ is the runoff generation in time t ; P represents precipitation, S_u represents root zone
305 moisture; S_{umax} represents root zone storage capacity; β represents the spatial distribution
306 function of the saturation threshold (Gao et al., 2017).

307 In contrast to saturation excess flow, infiltration excess flow occurs mostly in arid climates. Based
308 on this theory, the temporal and spatial dynamic of infiltration excess flow is determined by the



309 relationship between precipitation intensity and infiltration capacity. When rainfall intensity is
310 larger than infiltration capacity, runoff generation occurs; otherwise, all rainfall goes into
311 infiltration without runoff generation.

$$312 \quad Q(t) = f(P(t), f_c(t)) \quad (2)$$

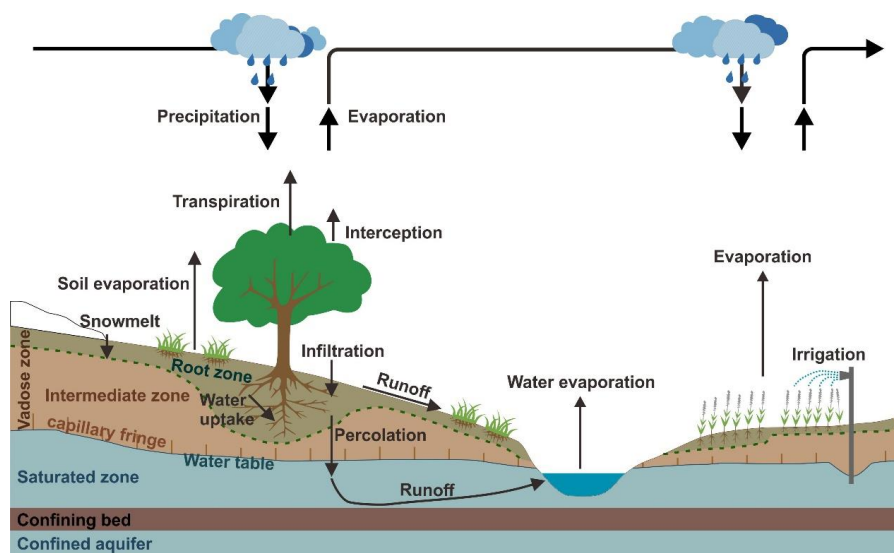
313 where P is precipitation intensity, f_c is infiltration capacity, including matrix infiltration capacity,
314 and preferential infiltration capacity.

315 Traditionally in hydrology, the infiltration process is considered to be dominated by matrix flow,
316 with the assumption that soil texture controls hydraulic properties and infiltration capacity (Beven,
317 2021). However, mounting evidence has revealed that preferential flow transmits free water quite
318 rapidly to the subsurface, dominating water movement even in arid regions (Liu et al., 2012; Gao
319 et al., 2023). Plant rooting system in water-limited environments correlates strongly with
320 infiltration capacity, velocity, and depth (Schenk & Jackson, 2002b; Schenk, 2008). Infiltration
321 recharges the root zone moisture through preferential patterns, and additional water can generate
322 runoff or percolate to the groundwater. Moreover, recent studies revealed that mucilage exuded
323 from roots and rhizosphere bacteria is critical in the formation of liquid bridges at the root–soil
324 interface, influencing mechanical stability and hydraulic properties in the root zone (Bengough,
325 2011). Root zone properties are essential in determining infiltration capacity and infiltration
326 excess flow generation for both matrix and preferential flow.

327 At landscape scale, topography emerges as a determining factor on the root zone size and runoff
328 generation mechanisms. Usually, saturation overland flow occurs on wetlands and riparian areas,
329 subsurface stormflow occurs on hillslopes, and infiltration overland flow usually happens on
330 plateaus or terraces (Savenije, 2010; Gao et al., 2014; Gharari et al., 2014). In summary, regardless
331 of saturation excess flow (also called variable contribution area) or infiltration excess flow, the
332 root zone plays a determining role in both hillslope and catchment hydrology.

333 The root zone also controls percolation to groundwater (Collenteur et al., 2021). In most
334 conditions, groundwater recharge and runoff generation are highly related, since runoff is
335 generated from groundwater in most circumstances (Ali et al., 2011). The root zone also
336 influences evaporation and transpiration from groundwater. Soil surface and root zone drying are
337 mitigated by upward capillary flow from the subsurface. The capillary rise sustains evaporation
338 during dry spells and decreases groundwater recharge on a longer timescale. A widely observed
339 phenomenon is the diurnal cycle of both streamflow and groundwater table fluctuation, mostly in
340 arid and semi-arid catchments (Lundquist and Cayan, 2002; Wang et al., 2014; Yue et al., 2016).
341 Vegetation, mostly in riparian areas, directly taps groundwater for transpiration, daytime root
342 water uptake causes water tables to decline during the day, while water tables recover at night
343 when transpiration essentially ceases (Yue et al., 2016). This shows the direct hydraulic link
344 between the root zone moisture, the groundwater and the runoff in streams,

345



346

347 Figure 6. Schematic diagram of the root zone in terrestrial hydrological processes and atmospheric
348 moisture cycling.

349 3.3.2 Root zone \neq vadose zone

350 The vadose zone is the unsaturated area, including the root zone, transition zone, and capillary
351 fringe above groundwater. From a hydrology perspective, the vertical profile of the critical zone
352 can be divided into different layers: canopy, litter layer, root zone, water transition zone,
353 unconfined groundwater, and confined groundwater. The most active layers in the critical zone
354 defining rainfall partitioning and related hydrological processes include the canopy, litter layer,
355 and root zone. Globally, the vegetation interception storage capacity of terrestrial ecosystems is
356 about 1-2 mm, estimated by remotes sensing LAI data (De Roo et al., 1996). The litter layer
357 storage capacity differs in different ecosystems, but it is likely to increase the total interception
358 storage capacity to around 2-5 mm (Shi et al., 2004). For determining the root zone storage
359 capacity of ecosystems, the mass curve technique appeared to work well locally and globally (Gao
360 et al., 2014; Wang-Erlandsson et al., 2016), whereby use was made of ERA-5 evaporation data,
361 resulting in a global average root zone storage capacity of approximately 180 mm. Hence, the root
362 zone storage capacity is significantly larger than the interception and litter layer storage capacities.
363 It is worth noting that soil evaporation in well vegetated regions is negligible due to the shielding
364 effect of the canopy and litter layer. In sparsely vegetated regions, e.g., deserts, soil evaporation
365 may play a more important role. Distinguishing the topsoil evaporation and the root zone
366 transpiration needs more complementary information, e.g., using isotope data. Both large-scale
367 isotope measurements and modeling studies confirmed that soil evaporation is around 6% of total
368 evaporation globally (Wang-Erlandsson et al., 2014; Good et al., 2015). Hence, the topsoil
369 evaporation is significantly less important than evaporation from the root zone.

370 Beneath the root zone, the transition zone is the layer between the lower limit of the root zone and
371 the upper limit of the capillary fringe, which is thick where the depth to the water table is deep and
372 may be absent where the water table is at or near the surface. This transition zone forms a



373 transport belt for percolation from the root zone to groundwater system, without water phase
374 changes, and does not have an impact on the land surface water balance. The percolation through
375 the transition zone recharges the groundwater and will be discharged eventually as runoff under
376 natural conditions. The deep confined groundwater has a limited connection to surface processes
377 and is not considered in surface runoff calculations in most cases. Hence, the root zone, as the
378 active layer within the vadose zone, plays the dominant role in catchment hydrology, determining
379 how catchments respond to precipitation.

380 **3.4 Root zone and atmosphere**

381 **3.4.1 Root zone and land surface processes**

382 The importance of the root zone on land surface and climate modeling is widely acknowledged
383 (Milly and Dunne, 1994). The moisture condition in the root zone is essential for water fluxes in
384 land-atmosphere interactions, mostly by increasing atmospheric water vapor through root water
385 uptake and transpiration. The contribution of transpiration to total land evaporation is regulated by
386 the interplay between the atmospheric water demand and the soil moisture within the reach of
387 vegetation's roots. Inversely, vegetation and land cover regulate the exchanges of root zone water,
388 energy and carbon with the atmosphere. Recent studies emphasized the potential of harnessing
389 climate-controlled root zone parameters to improve water flux simulations by the Hydrology Tiled
390 ECMWF Scheme for Surface Exchanges over Land (HTESSEL) land surface model (e.g., van
391 Oorschot et al., 2021).

392 The root zone is both affected by and affecting the Earth's climate (Jackson et al., 1996).
393 Compared with soil evaporation and canopy interception, vegetation can use deeper root zone
394 water for transpiration, allowing for long-term climate memory over longer distances (van der Ent
395 and Savenije, 2011). Land use impacts large scale moisture recycling and land use change in one
396 region can impact the precipitation in faraway regions downwind (van der Ent et al., 2010; Wang-
397 Erlandsson et al., 2018).

398 **3.5 Root zone and cryosphere**

399 The frontier of climate change studies is the terrestrial cryosphere, including glaciers, snow cover
400 and frozen ground. Glacier and bare rock/soil covered cold regions do not have vegetation cover
401 thus without root zone. Root zone and cryosphere have connections mostly in vegetation-covered
402 frozen ground. Snowfall and snowmelt, as an important cryospheric element, have significant
403 impacts on root zone hydrological processes in permafrost, with a time delay from snowfall on the
404 ground to snowmelt into the root zone. Snow cover as an isolation layer, reduces soil evaporation,
405 and likely increases soil temperature with a thick snow depth. Snow melt, in addition to rainfall,
406 increases infiltration, plant available water, plant growth, and runoff generation. In cold regions, it
407 is necessary to consider the effects of snowmelt on root zone water storage and the soil freeze-
408 thaw processes on hydrologic connectivity (Gao et al., 2020, 2022; Wang-Erlandsson et al., 2016;
409 Zhao et al., 2016; Dralle et al., 2021). The interaction between the root zone and the cryosphere,
410 especially snow cover and frozen soil, attracts increasing attention and research interest (Gao et
411 al., 2018; Leng et al., 2023) but still lacks systematic studies.



412 **3.5.1 Root zone ≠ active layer in permafrost**

413 In permafrost regions, the active layer, experiencing seasonal freeze/thaw processes, determines
414 the maximum thawing depth. Roots cannot grow in the ice layer and permanent frozen layer. The
415 rooting depth of deeper-rooting species, if present, is restricted to the active layer (the upper layer
416 of soil that thaws annually) because of the impermeability of the permafrost (Blume-Werry et al.,
417 2019). In climate change, when the permafrost underneath the active layer thaws, previously
418 unoccupied soil volumes become available for plant roots. Permafrost thaw leads to a larger active
419 layer and deeper groundwater table, which likely influence rooting depth and distribution.
420 Changes in rooting patterns and dynamics, as a direct response to permafrost thaw or indirectly
421 through changes in vegetation composition or soil moisture, have the potential to influence carbon
422 cycling in cold regions strongly (Wang et al., 2020).

423 **4 Root zone in the Anthropocene**

424 **4.1 Climate change**

425 Climate, as the upper boundary of the root zone, has the dominant impact on root zone dynamics.
426 Root growth is highly dynamic and responds quickly to changes in environmental conditions. The
427 time between photosynthesis and the transfer of carbon from leaves to roots and soil organisms is
428 fast, taking hours in grassland or days in forests. Changes in the root zone are generally
429 cumulative, which may be introduced by slow, gradual or abrupt changes.

430 Previous work demonstrated that root zone storage capacity (S_{umax}) varies spatially in response to
431 climatic conditions (Kleidon and Heimann, 1998; Gao et al., 2014; Wang-Erlandsson et al., 2016;
432 Stocker et al., 2023). It is, therefore, not implausible to assume that S_{umax} also varies temporally in
433 response to climate change. In the absence of longer time series of observed historical ecological
434 data, the 'space-for-time' assumption is a potential option to infer temporal ecological trajectories
435 from available spatial patterns (Singh et al., 2020). By comparing S_{umax} with aboveground tree
436 cover in several transitions across tropics and subtropics, Singh et al. (2020) found that in tropical
437 forests, with increased water stress, ecosystems tend to increase their S_{umax} and reduce tree cover.
438 Once the ecosystem passes across the tipping point, i.e. S_{umax} of 400 – 750 mm, and tree cover of
439 30-40%, rainforests in the Amazon appear to transition to savanna. This insight may be a
440 manifestation of an ecohydrological adaptation strategy, which offers potential for prediction
441 under future climate change.

442 At basin scale, Bouaziz et al. (2022) found that vegetation had adapted to climate change by
443 increasing its root zone storage capacity to offset the more pronounced hydro-climatic seasonality
444 in the Meuse basin. The increased vegetation water demand under global warming results
445 annually, and on average, in less streamflow, more evaporation and less groundwater recharge. At
446 global scale, Gao et al. (under review) analyzed observation-based ERA-5 reanalysis data to find
447 that the global average rootzone water storage capacity has increased by 8% over the past four
448 decades (1982-2020), in response to intensifying drought. The widespread increase of S_{umax} was
449 further validated by using the dynamic identifiability analysis (DYNIA) algorithm with a five-year
450 moving window to explore the dynamics of the S_{umax} parameter, forced by the observed catchment
451 hydrometeorological data of the USA (Liang et al., under review). At landscape scale, local
452 topography and groundwater also matter. In response to declines in the water table, rapid root



453 growth (up to 15 mm/d) toward the water table has been observed for desert phreatophytes
454 vegetation (Orellana et al., 2012; Kuzyakov and Razavi, 2019).

455 The net effect of the presence of plants with roots on land is to draw CO₂ out of the atmosphere,
456 which has a major effect on climate. Root zone dynamics are also important for carbon
457 sequestration, such as predicting how plants respond to elevated CO₂ under climate change (Nie et
458 al., 2013). Of carbon storage alternatives, root zone carbon has the largest uncertainty, with great
459 influence on carbon neutrality and sequestration, which still needs further experimental and
460 modeling studies (Friedlingstein et al., 2022).

461 **4.2 Agriculture**

462 Most agricultural activities are happening in the root zone, including but not limited to irrigation,
463 fertilization, tilling, non-point source pollution, and salinization. Root zone management is
464 essential for food production and food security. Originally, the term "root zone" has been used in
465 agronomy, e.g., to assess the total amount of water and nutrients available for plants. Ploughing as
466 an important agriculture management practice, which greatly changes and even determines the
467 root zone depth. The plough pan, distributed commonly from 15 to 30 cm beneath the soil surface,
468 caused by long-term cultivation, is formed under the root zone with a high clay content (Li et al.,
469 2019). This relatively impermeable layer limits water percolation and constrains root growth in the
470 relatively small soil layer above the plough pan. In such cases, it is indeed the moisture holding
471 capacity of the topsoil that determines the root zone storage capacity.

472 In cropland, where irrigation provides an extra water supply to the root zone during dry seasons,
473 the root zone water storage capacity is often smaller than under natural conditions with similar
474 climate conditions (Xi et al., 2021; Hauser et al., 2022; van Oorschot et al., 2023). Due to
475 irrigation (not considering climate change), rooting depth was reduced by 5%, or ~8 cm, or loss of
476 ~11,600 km³ of rooted volume (Hauser et al., 2022). In a global scale study, when considering
477 irrigation as an extra water supply, S_{umax} in cropland was found to consistently decrease (van
478 Oorschot et al., 2023).

479 Agricultural activities also dramatically changed root zone biogeochemistry cycles. All living
480 things need nitrogen (N) and phosphorus (P) to make amino acids, proteins, and DNA. However,
481 most of the nitrogen on Earth is in the atmosphere, in the form of N₂, which cannot be directly
482 used for most plants and all animals; and phosphorus is supplied by physical and biological rock
483 weathering, which is an extremely slow process. Thus, in natural terrestrial ecosystems, N and P
484 are often the most limiting nutrients for productivity. Chemical fertilizer revolutionized this
485 limitation, changing from nutrient scarcity to overloading, resulting in a series of environmental
486 issues, e.g., soil hardening in the root zone, groundwater contamination, surface water
487 eutrophication, and phosphate resource depletion (UNEP, 2019; Yao et al., 2020; Edixhoven et al.,
488 2014). Thus, agricultural root zone management is a pivotal issue for food security, human health,
489 water resources management, and environmental protection (Gu et al., 2023).

490 **4.3 Nature-based solutions**

491 **4.3.1 Large-scale ecological projects**

492 Large-scale land cover change will dramatically change hydrology, land surface processes, and



493 biogeochemistry cycle, between which the root zone is the crucial link. For example, intensive
494 deforestation greatly reduces root zone water storage capacity (Hrachowitz et al., 2021), requiring
495 over ten years to recover (Nijzink et al., 2016). In regions experiencing woody encroachment,
496 roots are deepening by ~38 cm compared to previous dominant vegetation (Hauser et al., 2022).

497 The root zone plays an essential role in evaluating the effects of large-scale ecological projects.
498 For instance, the “grain-to-grass” and “three norths afforestation projects” in China, have been
499 greatly beneficial for soil and water retention, biodiversity conservation, and carbon sequestration
500 (Chen et al., 2019; Li et al., 2021). But these projects still face great challenges in water
501 management of these arid regions, where severe competition exists between society's water
502 demand and ecosystem services (Feng et al., 2016). Moving from blue water management to
503 integrated water resources management, means bringing both blue- and green-water, surface and
504 groundwater, in the same framework (Falkenmark, 2000). Root zone moisture forms the essential
505 link between land surface fluxes, water resources, soil erosion, and carbon sequestration,
506 determining the success and sustainability of these grand projects (Sun et al., 2020).

507 **4.3.2 Water-soil conservation and landslide prevention**

508 Afforestation is regarded as an effective nature-based solution for water-soil conservation and
509 flood mitigation. Forests have higher evaporation due to deep roots and large biomass. Higher
510 evaporation results in higher soil water storage capacity to store more rain water for dry periods.
511 For example, afforestation in the Yangtze River basin was expected to reduce flood risk by
512 increasing its root zone water storage capacity as “green reservoirs”. However, there has been a
513 long dispute about the efficiency of forests in flood control, especially after the megaflood in
514 Changjiang in 1998. Interestingly, in Thailand, Sriwongsitanon et al. (2011) found that forests
515 have a large impact on small floods but no significant impact on extreme floods.

516 From a long-term evolutionary perspective, rooting increases weathering and soil formation. On
517 the short-term, roots' anchoring effect holds soil in place, reduces the erosion rate, increases slope
518 stability, and prevents landslides and debris flow catastrophes (Vannoppen et al., 2017). Forest
519 communities are considered to effectively minimize erosional processes and support hillslope and
520 riverbank stability (Pawlik et al., 2016). They also effectively contribute to soil strength, reducing
521 probability of shallow land sliding. The reduced sediment flow generation by roots also has basin
522 scale impacts on downstream river geomorphology. For example, rivers were changed from
523 braided river channels to predominately meandering river channels due to the stabilizing effects on
524 riverbanks that rooting systems provide (Ielpi et al., 2022). It changes both suspended and bed
525 load sediment fluxes, which has remarkable effects on infrastructural security and service life of
526 dams, dikes, and water uptake facilities (Cao et al., 2023).

527 **4.3.3 Sponge city**

528 Human alterations of the root zone have increasingly become pervasive and long-lasting, which is
529 particularly true in urbanized areas. Urbanization alters land surface by changing from permeable
530 to impermeable. This significantly changes the root zone hydrological process and land surface
531 energy budget, intensifying urban inundation and urban heat island effect (Fletcher et al., 2013).
532 There is also an increasing interest in using vegetation to manage catchment and regional
533 hydrology (Palmer et al., 2015), especially in the urban environment, as exemplified by the



534 “Sponge City initiative” in China (Xia et al., 2017). The root zone is the largest and most
535 widespread “sponge” in forests and grassland at large scale, whereas green roof and garden
536 provide a “sponge” at city and block scale. Increasing S_{umax} has potential to improve flood
537 prevention, water purification, relieving heat island effect, and aesthetic value in the urban
538 environment.

539 **4.4 Planetary stewardship**

540 As the key linkage between the natural geosphere and the anthroposphere (Figure 1, and Steffen et
541 al., 2020), understanding the root zone has broad relevance to planetary stewardship, such as green
542 water footprint, integrated water resources management, and carbon sequestration (Wang-
543 Erlandsson et al., 2022). The root zone plays a key role in partitioning precipitation into the
544 invisible green water (used for terrestrial ecosystem) and visible blue water (usable for human
545 society) (Falkenmark, 2000). The root zone is one of the most manageable variables in the
546 terrestrial water cycle and land system. Almost all human activities have impacts on the root zone,
547 like agriculture, urbanization, deforestation/afforestation, landuse change, and fertilization. Proper
548 root zone management is essential to keep both green and blue water within safe planetary
549 boundaries for Earth system resilience (Wang-Erlandsson et al., 2022; Stewart-Koster et al., 2023).
550 Root zone management is essential for numerous global sustainability development goals (SDGs),
551 including but not limited to SDG1 (no poverty), SDG2 (zero hunger), SDG3 (good health and
552 well-being), SDG6 (clean water and sanitation), SDG11 (sustainable cities and communities),
553 SDG13 (climate action), and SDG15 (life on land).

554 **5 Root zone estimation approaches**

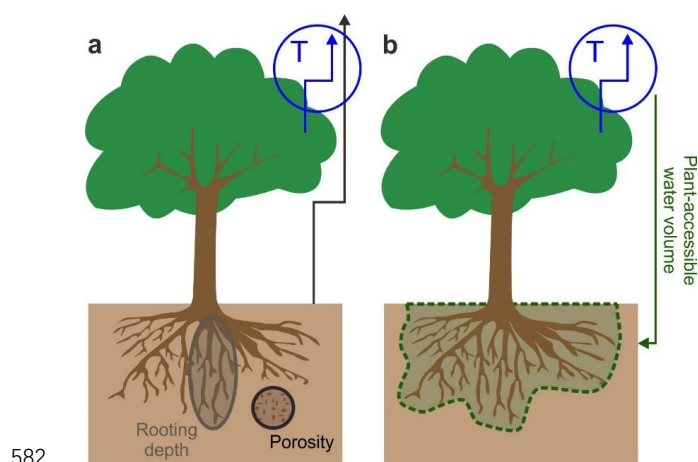
555 **5.1 Root zone water storage capacity \neq rooting depth \times soil water holding 556 capacity**

557 In most Earth system models, root zone storage capacity is obtained by combining rooting depth
558 with soil water holding capacity derived from soil texture data. The soil water holding capacity is
559 obtained in the laboratory and defined as the amount of water held by the soil between field
560 capacity and permanent wilting point, which is available for plant water use (van Oorschot et al.,
561 2021). In this type of models, the water exceeding field capacity is drained under gravity; and
562 below wilting point, vegetation starts to wilt, leading to plant mortality.

563 However, we argue that this traditional perspective, which believes that the sum of the parts is
564 equal to the whole, is problematic. Firstly, this type of method usually lacks detailed observations
565 of root density and vertical and lateral distribution, which has large uncertainty, and, in most cases,
566 lacks observations and heavily relies on parameter calibration. But even considering the root
567 density and vertical and lateral distribution, we cannot obtain reliable root zone water storage
568 capacity, from only one observation. Emerging ecohydrology studies revealed that hydraulic lift is
569 important for trees in the nighttime redistribution of water from wetter soil deeper in the soil
570 profile through roots to drier soil closer to the surface (Bleby et al., 2010; Nadezhdina et al.,
571 2010). This phenomenon is observed for individual trees and measured in exchanging water and
572 nutrients among neighboring trees (Hafner et al., 2021). Hence, it is extremely difficult if not
573 impossible to define the exact shape of the root zone of an ecosystem by a reductionist approach



574 (Figure 7). The root zone consists of a gradient in chemical, biological, and physical properties
575 that change radially and longitudinally along the root (McNear, 2013; Kuzyakov and Razavi,
576 2019). For example, the presence of arbuscular mycorrhizal fungi is known to enhance plant water
577 uptake, particularly important during water-stressed periods (Augé, 2001), by boosting root
578 growth and potentially also through direct water acquisition (Püschel et al., 2020). This complex
579 of pores and fissures in the substrate, extending both laterally and in depth, with complicated
580 extents and gradients of water and nutrients, cannot be simulated by the mechanical combination
581 of rooting depth and soil properties (Figure 7).



582

583 Figure 7. The reductionist (a) and holistic (b) perspective. (a) The reductionist perspective
584 estimates root zone storage capacity and transpiration based on rooting depth and soil properties,
585 which is a stationary view, mostly based on soil information. (b) The holistic perspective estimates
586 root zone storage capacity and transpiration based on plant water demand, which is a dynamic
587 view, mostly based on climate and plant adaptation.

588 5.2 Holistic perspective

589 Holistic perspectives view the root zone as an integrated and living system leading to self-
590 organization and fractal patterns. Reductionist modelers ask the question of "How deep is the root
591 zone?". Holistic modelers ask the questions "What is the size of the root zone?" and "How much
592 moisture can it buffer?".

593 The inverse modelling approach uses land surface water and energy fluxes, such as precipitation,
594 evaporation, and radiations, and biomass observation, such as net or gross primary production, to
595 derive root zone processes with different parameterizations (Kleidon, 2004). The mass curve
596 technique (MCT), as an inverse model, is an observation-based approach to estimate root zone
597 processes, based on the land surface water budget (Gao et al., 2014; Wang-Erlandsson et al.,
598 2016). Catchment and global scale studies have revealed that ecosystems tune to a drought period
599 with 10–40-year return periods depending on the type of ecosystem (evergreen, deciduous,
600 grassland, etc.). Successful applications have been made by Gao et al. (2014), Nijzink et al.
601 (2016), Wang Erlandsson et al. (2016), Sriwongsitanon et al. (2016), De Boer-Euser et al. (2019),
602 Mao and Liu (2019) and Bouaziz et al. (2022), which show large diversity in hydrological fluxes



603 globally, which cannot be obtained by traditional reductionist methods.

604 This powerful holistic method allows hydrologists and ecologists to derive underground processes
605 by large scale observations from space (Kühn et al., 2022). Another interesting implication is by
606 McCormick et al. (2021), who found that the water storage capacity of the topsoil is insufficient to
607 explain evaporation in dry seasons in many areas across the USA, and, therefore, rock moisture is
608 a critical component of terrestrial water. The MCT method, on the other hand, allows us to
609 quantify this amount of water use from rock fissures (Lapides et al., 2023).

610 At landscape scale, the spatial distribution of S_{umax} is mostly determined by topography (Gao et
611 al., 2019). Moreover, topography plays an essential role in determining runoff generation
612 mechanisms, such as by saturated overland flow in wetlands, subsurface storm flow in hillslope,
613 and infiltration excess overland flow in terrace (Savenije et al., 2010; Gharari et al., 2014).

614 The holistic perspective also needs holistic observations, including land surface fluxes like
615 lysimeter, eddy covariance, and satellite remote sensing (Gao et al., 2014; Wang-Erlandsson,
616 2016; Kühn, 2022). Weighing lysimeter provides a unique equipment to quantify the water
617 balance in the root zone (Seneviratne et al., 2012), especially, when comparing non-vegetated and
618 vegetated systems, demonstrating the role of vegetation dynamics in controlling the water cycle
619 (Scanlon et al., 2005). Satellite remote sensing provides more spatially and temporally extensive
620 data about land surface water fluxes, such as precipitation and evaporation. Remote sensing also
621 provides direct surface moisture measurement, such as SMAP and SMOS. However, satellites can
622 only measure water in a few centimeters of the topsoil. Root zone moisture was then derived from
623 surface moisture by the soil hydraulic model (Entekhabi et al., 2010). Interestingly,
624 Sriwongsitanon et al. (2016) found that the normalised difference infrared index (NDII) is a better
625 indicator for root zone moisture, which means that vegetation itself is a holistic indicator of root
626 zone moisture conditions rather than an obstacle for remote observation that needs to be removed.

627 Comparing the reductionist with the holistic approach, it is interesting that although the holistic
628 method is much simpler, its performance is even better than the most complex reductionist
629 approach in runoff modeling (Mao and Liu, 2019; Wang et al., 2021), especially in ungauged
630 basins (Hrachowitz et al., 2013), and evaporation flux simulations by land surface models (van
631 Oorschot et al., 2021). The holistic method has also been used in the widely used dynamic
632 vegetation model, i.e., LPJ-GUESS, by the inclusion of bedrock vadose zone, to improve its
633 representation of storage and hydrology (Lapides et al., 2023). Considering plant's water use from
634 bedrock improved evaporation estimation, especially in seasonally dry regions. Moreover, the
635 holistic model is much easier to be coupled with biogeochemistry processes, especially in the
636 sense that conceptual models are the dominant approach to simulate the carbon, nitrogen, and
637 phosphorus fluxes.

638 **6 Summary and future outlook**

639 The root zone is the crucial element linking multi-spheres of the Earth surface system, including
640 biosphere, lithosphere, hydrosphere, atmosphere, cryosphere, and anthroposphere. Although many
641 disciplines are studying the root zone from different angles, this is not yet done in a systematic
642 way. This study explored the differences and linkages between the root zone and other
643 terminologies. For example, the root zone does not equal rooting depth or rhizosphere in biology



644 and ecology. In pedology and lithology, the root zone is not equal to the soil depth or critical zone.
645 In hydrology, the root zone is not equal to the vadose zone. In atmospheric science, the root zone
646 storage capacity is not equal to the product of rooting depth with soil water holding capacity.
647 Clarifying the terminology is essential for future studies. Furthermore, we discussed the root zone
648 in the Anthropocene, including climate change, agriculture, nature-based solutions, and planetary
649 stewardship. In the end, we discussed the traditional reductionist perspective to understand and
650 model root zone, and we proposed to move forward to a holistic perspective to root zone in Earth
651 system science.

652 We argue that more experiments and measurements in diverse climates and landscapes, especially
653 holistic perspective methods, such as rhizobox and field control experiments (Nie et al., 2013),
654 and long-term lysimeter measurements (Scanlon et al., 2005), are especially needed to get a
655 deeper understanding of root zone patterns and dynamics. For example, rhizobox experiments
656 interestingly demonstrated that precise root distribution information is unnecessary to estimate
657 water budgets correctly (Maan et al., 2023). The booming development of root microbiome
658 studies in ecology, can greatly enhance our understanding of the rhizosphere biotic processes in
659 the root zone, its importance for the entire ecosystem, and even its impacts on the large scale
660 biogeochemistry cycle (McNear, 2013).

661 For modeling, the flexibility of plant roots and their ability to adapt to the environment, is still not
662 properly incorporated in land surface and dynamic vegetation models. Reductionist modeling is
663 still the mainstream methodology in Earth system models, but holistic modeling of the root zone
664 has large potential for simpler and more realistic large scale simulation. For example, Singh et al.
665 (2020) provided a spatial transient of S_{umax} changes as a function of climate. Based on space for
666 time exchange, this paradigm can be used to predict the dynamic of belowground and
667 aboveground biomass variations with climate change. Moreover, the holistic modeling approach
668 allows us to involve more physical laws in hydrological modeling, not only the Newtonian
669 approach but also evolution theory and energy efficiency, with the Carnot limit as its upper
670 constraint (Savenije, 2023). Additionally, we need a better representation of root zone
671 biogeochemistry processes, including carbon, nitrogen, phosphate, pollutants, and microbial
672 communities for broader implications in Earth system science. Finally, we need to consider the
673 impacts of human activities on the root zone in Earth System Models (ESMs) for better root zone
674 management and relevant decision-making.

675 Essential in the holistic approach is the ecosystem perspective. The ecosystem is the active agent
676 that manipulates the environment to sustain its survival. In doing so, it adjusts the substrate on
677 which it lives, creating a favorable hydrological, geochemical, and biological foundation for its
678 sustenance. Analysis of what the ecosystem has done over time to create conditions for survival
679 and rejuvenation holds the key to understanding hydrological processes and can help to predict
680 hydrological behavior under future climates.

681

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683 **Competing interests.** At least one of the (co-)authors is a member of the editorial board of
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