

Root zone in the Earth system

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Abstract: The root zone is a vital part of the Earth system and a key element in hydrology, ecology, agronomy, and land surface processes. However, its definition varies across disciplines, creating barriers to interdisciplinary understanding. Moreover, characterizing the root zone is challenging due to a lack of consensus on definition, estimation methods, and their merits and limitations. This Opinion paper provides a holistic definition of the root zone from a hydrology perspective, including its moisture storage, deficit, and storage capacity. We demonstrate that the root zone plays a critical role in the biosphere, pedosphere, rhizosphere, lithosphere, atmosphere, and cryosphere of the Earth system. We underscore the limitations of the traditional reductionist approach in modeling this complex and dynamic zone and advocate for a shift towards a holistic, ecosystem-centered approach. We argue that a holistic approach offers a more systematic, simple, dynamic, scalable, and observable way to describing and predicting the role of the root zone in Earth system science.

1 Introduction

Plant roots developed before leaves during the Devonian Period, 416 to 360 million years ago

36 (Kenrick and Strullu-Derrien, 2014). The development of root systems was a critical biological
37 innovation that enabled land plants to spread and colonize the continental interior. Before this
38 development, plants were limited to areas immediately adjacent to bodies of water due to their
39 simple rhizoid-based rooting systems (Smart et al., 2023). Roots acted as pioneers of biological
40 activity, representing a definitive line between sterile sediments and living soils. Fossil evidence
41 indicates that deeper soils and larger plants appeared almost simultaneously (Kenrick and Strullu-
42 Derrien, 2014).

43 There is a positive feedback between soils and plants, mediated by roots. Larger plants generally
44 require more stable and deeper soils, while plant roots contribute to soil formation. The rapid
45 expansion of terrestrial plants was largely enabled by newly developed rooting systems. Roots
46 promoted physical, chemical, and biological rock weathering and increased terrestrial
47 photosynthesis capability (Maeght et al., 2013), which significantly reduced atmospheric CO₂. This
48 reduction likely triggered the ice age and mass extinction in the late Devonian (Smart et al., 2023).

49 Thus, the evolution of plant root systems not only influenced the local soil microenvironment and
50 individual plant life cycles but also affected the weathering of the lithosphere, the formation of the
51 pedosphere, and atmospheric compositions. The initial appearance of rooting systems transformed
52 the entire biosphere and the abiotic environment on the planet (Hetherington, 2019).

53 In recent decades, the root zone has received increasing attention from various disciplines, including
54 hydrology, ecology, agronomy, biology, soil physics, atmospheric science, and landscape
55 engineering. Despite this growing interest, there remains a lack of synthesis on the basic concept of
56 the root zone, especially within the context of Earth system science. The root zone is arguably the
57 least understood portion of the ecosystem that controls land-surface processes (Wang et al., 2018).
58 Therefore, there is an urgent need to clarify its definition and bridge the knowledge gaps between
59 traditional disciplines.

60 This study has the following objectives:

- 61 1. Provide a definition of the root zone.
- 62 2. Propose a perspective of the root zone as a living, evolving, adapting, and essential part of
63 the Earth system.
- 64 3. Advocate for a shift from the traditional reductionist approach towards a holistic,
65 ecosystem-centered perspective of root zone hydrology in Earth system science.

66 **2 The definition of root zone**

67 **2.1 Root zone**

68 The *Root zone* is the upper part of the subsurface that supports vegetation rooting (Sprenger et al.,
69 2019), where water, air and nutrients are available to sustain plants (see Figure 1, and Rodríguez-
70 Iturbe & Porporato 2004). It typically extends through the unsaturated soil or parts thereof and may
71 also reach the groundwater (Fan et al., 2017) or penetrate weathered bedrock (McCormick et al.,
72 2021). The root zone is a hydrologically active layer, replenishing with water during wet periods
73 and supplying ecosystems with water to survive droughts in dry periods. It controls the partitioning
74 of precipitation into infiltration, soil evaporation, plant transpiration, percolation to groundwater
75 and runoff (including surface runoff and subsurface storm flow) (e.g.: Guswa, 2010; Lazarovitch et

76 al., 2018).

77 The root zone is a cross-scale concept, applicable at various scales including (Delcourt and Delcourt,
78 1988; Bloeschl and Sivapalan, 1995):

- 79 • Small scale ($1\sim 10^6$ m², plot, landscape, hillslope and sub-catchment scales),
- 80 • Meso-scale ($10^6\sim 10^{10}$ m², catchment scale), and
- 81 • Large scale (above 10^{10} m², river basin and continental scales).

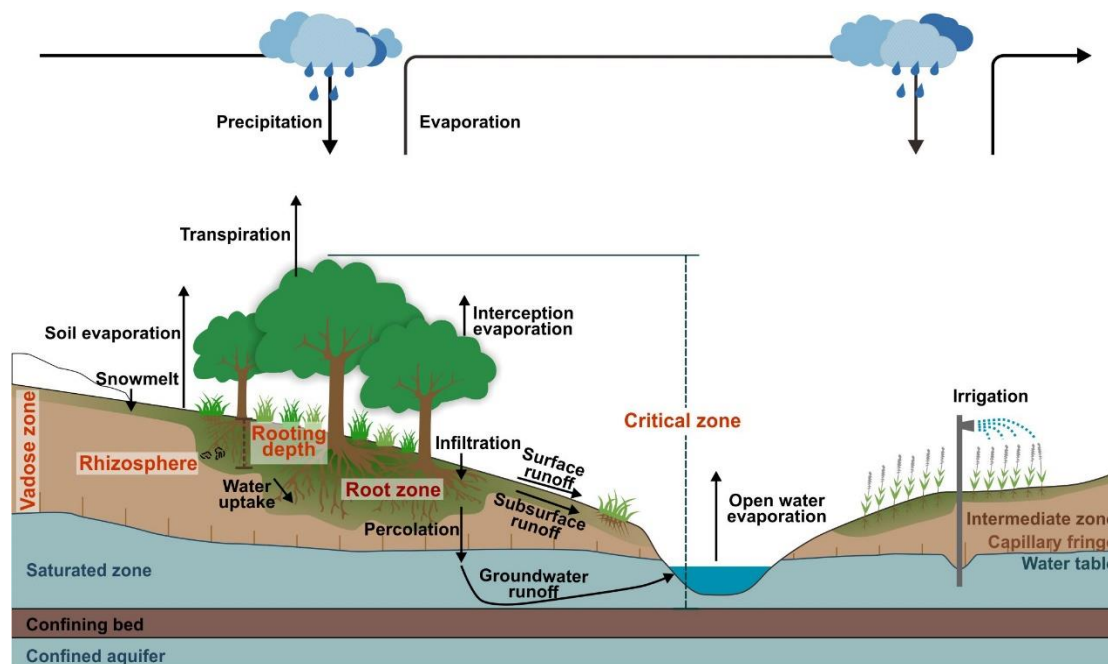
82 **2.2 Root zone water storage and root zone water deficit**

83 The root zone acts as a buffer, ensuring continuous access to water for an ecosystem to bridge dry
84 periods. Its shape, hydrological gradients, as well as internal water fluxes and processes, are highly
85 dynamic and complex. The interaction of plant roots, soil particles, bedrock fissures, water, nutrients,
86 and micro-organisms create an intricate system (McDonnell et al., 2007). From an ecosystem-
87 centered hydrology perspective, the complex root zone can, reflecting a “Darwinian” approach
88 (Harman and Troch, 2014), be viewed as a dynamic system within a living organism (Savenije and
89 Hrachowitz, 2017; Savenije, 2024).

90 Water storage in the root zone can be conceptualized as a volume of water per unit area, represented
91 as an average depth (e.g. in mm) (see Figures 1 and 4). The water deficit (S_d) is defined as the
92 difference between the maximum water volume and the actual root zone water storage at any time
93 $S(t)$. Water deficit and water storage are two sides of the same coin. During rainfall events, the root
94 zone wets up and refills as water is retained. Once the root zone moisture exceeds its water retention
95 capacity, excess water drains off by recharge to the groundwater (triggering groundwater flow), or
96 by lateral preferential flow, or as overland flow (by saturation excess or infiltration excess) directly
97 to the stream. These processes may occur at the same time (McDonnell, 2013). In dry periods,
98 ecosystems rely on the water stored in the root zone, leading to an increasing water deficit.

99 **2.3 Root zone water storage capacity (S_{Rmax})**

100 The root zone storage capacity (S_{Rmax}) is the maximum volume of water stored in the subsurface
101 that can be accessed by roots and that the root zone can contain, to allow plants to overcome critical
102 drought periods (Rodríguez-Iturbe & Porporato 2004; Gao et al., 2014a; Klos et al., 2018; Stocker
103 et al., 2023). S_{Rmax} is the core property of terrestrial hydrological systems and therefore also a key
104 parameter in process-based hydrological models, determining the partitioning between drainage and
105 evaporative fluxes. From a soil hydraulic perspective, S_{Rmax} represents the water volume stored in
106 the subsurface between permanent wilting point and field capacity and that is within reach of roots.
107 In other words, S_{Rmax} is equivalent to the maximum water deficit in the root zone, which defines the
108 volumetric extent of the root zone (Gao et al., 2014a; Lapidés et al., 2024). This concept can be
109 likened to the storage capacity of an artificial reservoir, which is dimensioned to balance societal
110 costs and benefits in the light of an acceptable risk (i.e. the design return period). Similar to human
111 reservoir design, ecosystems show evidence of optimizing the root zone, balancing guaranteed water
112 access (at a certain return period) against carbon expenditures for root growth and maintenance
113 (Kleidon and Heiman, 1998; Guswa, 2008; Schymanski et al., 2008; Gao et al., 2014a).



114

115 Figure 1. Schematic diagram of the root zone in terrestrial hydrological processes and atmospheric
 116 moisture cycling, showing the difference between root zone and other similar terms, such as vadose
 117 zone, rooting depth, rhizosphere and critical zone.

118 2.4 Root zone and “similar” concepts

119 2.4.1 Root zone vs vadose zone

120 The vadose zone (also known as the unsaturated zone) is the area underneath the land surface which
 121 forms the transition zone to the capillary fringe above the groundwater (Figure 1). The root zone is
 122 contained within the unsaturated zone, but is generally smaller. Below the root zone, the transition
 123 zone acts as a conduit facilitating percolation from the root zone to the groundwater with very
 124 limited phase changes of water (i.e. evaporation), and without directly affecting the water balance
 125 at the land surface. In other words, the water percolating into the transition zone from above will
 126 eventually recharge the groundwater (with some temporal delay depending on the depth of the
 127 groundwater and soil types) and finally contribute to river flow. Deep confined groundwater, which
 128 is minimally connected to surface processes, generally does not factor into surface runoff
 129 calculations (Fitts, 2002). Consequently, the root zone, functioning as the active layer that is located
 130 mostly within the vadose zone plays a pivotal role in catchment hydrology, determining in large part
 131 how catchments respond to precipitation events.

132 2.4.2 Root zone vs critical zone

133 The root zone, a key area where plants have evolved to adapt to their environment over time,
 134 constitutes the foundation of terrestrial ecosystems. As such, it is a vital *component* of the critical
 135 life zone on Earth (Banwart et al., 2017; Brantley et al., 2017). However, the definition of the critical
 136 zone remains a topic of debate. The most commonly accepted definition comes from the National
 137 Research Council in the U.S. (NRC, 2001): the critical zone is the near-surface layer of the Earth,
 138 encompassing vegetation, soil, water, and rocks, all of which are essential for sustaining life.
 139 According to this definition, the critical zone extends "from the canopy top to the base of the

140 groundwater zone." (Figure 1). While there is widespread agreement on the upper boundary of the
141 critical zone—the top of the vegetation canopy—the lower boundary is less defined and varies. For
142 instance, some definitions cite the base of the groundwater zone (NRC, 2001), the bottom of the
143 weathering zone (Guo and Lin, 2016), the base of active groundwater (Fan, 2016), or the storage of
144 fresh groundwater (less than 50 years old) (Gleeson et al., 2016). Regardless of the specific version,
145 the root zone is always included in critical zone definitions. Furthermore, the root zone is the most
146 dynamic layer within the critical zone, connecting the canopy to groundwater and encompassing all
147 components of the critical zone: water, vegetation, soil, air, and bedrock.

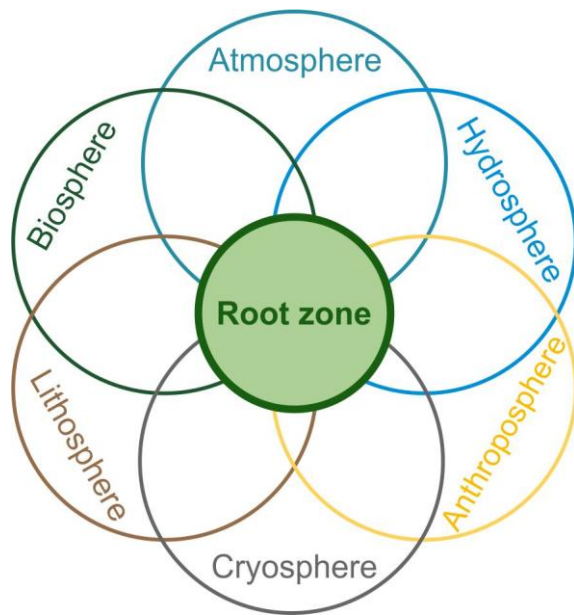
148 **2.4.3 Root zone vs rhizosphere**

149 The rhizosphere refers to the zone around roots where micro-organisms play crucial roles in
150 biological processes essential for plant growth and health (Hiltner, 1904; Hartmann et al., 2008). It
151 encompasses the roots themselves, the surrounding soil, and the narrow space between them,
152 typically ranging from 1 to 4 millimeters (Kuzyakov and Razavi, 2019) (see Figure 1 and 4). There
153 are similarities between the rhizosphere and the root zone. Both are dynamic and alive. However,
154 the rhizosphere specifically refers to the soil volume immediately surrounding living roots,
155 extending only a few millimeters from the root surface. It is primarily studied to understand direct
156 interactions with micro-organisms and to assess nutrient depletion zones. In contrast, the root zone
157 concept is applied in larger-scale hydrology, agronomy, and land surface studies to evaluate the total
158 water and nutrient resources available to plants. This results in a significant difference in volume,
159 spanning 2 to 3 orders of magnitude between the rhizosphere and the root zone. The rhizosphere is
160 more aligned with biological and soil science on a micro-scale (Kuzyakov and Razavi, 2019),
161 whereas the root zone is better suited for broader ecological and hydrological investigations within
162 Earth systems.

163 **3 Root zone in the Earth system**

164 Earth system science is an interdisciplinary field to describe the Earth's structure and operations as
165 a complex, integrated system (Steffen et al., 2020). Originating from early studies on interactions
166 between the biosphere and geosphere, and influenced by concepts such as the Gaia hypothesis
167 (Lovelock, 1979), Earth System science gained prominence in the 1980s in response to the need for
168 a comprehensive 'science of the Earth'. In the latest conceptual model (Figure 3 in Steffen et al.,
169 2020), the root zone assumes a central role as the interface between roots, soils, vegetation,
170 lithosphere, hydrosphere, biosphere, cryosphere, atmosphere and human society's production and
171 consumption (Figure 2). Despite its thin layer and containing only 0.13% of Earth's total freshwater
172 (McCartney et al., 2022), the root zone profoundly influences land surface hydrology, land-
173 atmospheric moisture exchange, and biogeochemical processes (Zehe et al., 2019).

174



175

176 Figure 2. Root zone is the interface between hydrosphere, biosphere, lithosphere, atmosphere,
 177 cryosphere, and anthroposphere in the Earth system.

178

179 **3.1 Roots and the Hydrosphere**

180 The terrestrial water cycle begins with precipitation, which is initially intercepted by vegetation and
 181 ground cover. From there, it either returns to the atmosphere through interception evaporation, runs
 182 off to drainage systems as overland flows, or infiltrates into the soil, eventually percolating into the
 183 root zone. Once infiltrated, water can be stored for plant use through transpiration, percolate deeper
 184 to recharge groundwater, or flow laterally, through preferential pathways, to the stream.
 185 Transpiration rates are regulated by plant characteristics, atmospheric water demand, and the
 186 availability of moisture in the root zone.

187 Overall, these processes—evaporation, transpiration, runoff, groundwater recharge—are largely
 188 controlled by the moisture content in the root zone (Rodriguez-Iturbe and Porporato, 2004).
 189 Therefore, root zone moisture plays a critical role in determining how precipitation is partitioned
 190 among these different pathways and affects the overall storage dynamics of the root zone.

191 In the global water balance over extended periods, precipitation is partitioned between runoff and
 192 evaporation, with the root zone playing a crucial role in this distribution. According to ERA-5
 193 reanalysis data, the average annual terrestrial precipitation is 745 mm/year, of which 295 mm/year
 194 (40% of precipitation) returns to the oceans as streamflow, while 440 mm/year (60% of precipitation)
 195 is evaporated back into the atmosphere by ecosystems (Gao et al., 2023a).

196 The largest terrestrial water flux to the atmosphere is through vegetation root water uptake, hence
 197 transpiration, which accounts for 60-90% of total terrestrial evaporation globally (Jasechko et al.,
 198 2013; Schlesinger and Jasechko, 2014; Coenders-Gerrits et al., 2014; Wang et al., 2014; Lian et al.,
 199 2018). Despite global average canopy interception having a storage capacity of only around 2 mm
 200 (compared to approximate 200 mm S_{Rmax} , see Gao et al., 2023a), interception plays a significant
 201 role in total forest evaporation due to its continuous activity in all rainfall events (Gerrits et al., 2010;

202 Sun et al., 2016).

203 In well-vegetated regions, soil evaporation is of minor relevance due to the protective effect of the
204 canopy and litter layer, but also due to the lack of turbulent exchange with depth (Brutsaert, 2014).
205 Conversely, in sparsely vegetated areas such as deserts, soil evaporation can be more substantial.
206 Globally, soil evaporation contributes approximately 6% of total evaporation, with its significance
207 being more prominent in arid than in humid regions (Wang-Erlandsson et al., 2014; Good et al.,
208 2015; Zhang et al., 2017).

209 Plant roots, along with microbiotic and macrobiotic life, profoundly influence soil hydraulic
210 properties and water-cycle dynamics by creating soil macropores and adding organic matter (Ponge,
211 2015). These macropores, formed through the presence of living or decaying roots but also by
212 animal activity enhanced by the presence of vegetation, act as large cracks or voids that reinforce
213 infiltration and facilitate rapid water drainage as preferential flow pathways. This capability reduces
214 the duration of soil saturation, thereby mitigating conditions that could lead to root rot.

215 Furthermore, organic matter from decomposed roots contributes to cementation within soils,
216 altering their internal structure. Soils rich in organic matter typically exhibit higher porosity and
217 water holding capacity, crucial for retaining water in soil pores—a vital adaptation for terrestrial
218 ecosystems in water-scarce environments.

219 The expansion of terrestrial ecosystems not only transforms hydrological processes within the root
220 zone but also exerts significant impacts on the larger-scale water cycle (see Figure 1 and Sect. 4.4).
221 Modeling studies indicate that terrestrial ecosystems enhance terrestrial precipitation by 40% (Xue
222 et al., 2010). Comparisons between a green planet (fully forest-covered) and a desert planet (without
223 vegetation) reveal that evaporation on continents triples and precipitation doubles in the former
224 scenario (Fraedrich et al., 1999; Kleidon et al., 2000). These findings underscore the profound
225 influence of root zone development in intensifying the global water cycle.

226 **3.2 Roots and the Biosphere**

227 Roots are indispensable organs enabling plants to thrive and reproduce across diverse habitats. They
228 serve crucial functions such as absorbing water and nutrients, which are essential for plant growth
229 and performance. Additionally, roots anchor and stabilize plants in the soil, store chemical energy
230 produced via photosynthesis in leaves in the form of carbohydrates and thus as biomass (Shekhar et
231 al., 2019), and play a pivotal role in sensing local resources like water, nutrients, and phosphate.
232 This sensing ability allows plants to efficiently access resources that are often heterogeneously
233 distributed in the substrate's complex network of pores and fissures.

234 The root zone represents the product of long-term co-evolution between the biosphere and its
235 inorganic environment, shaped by climatic and geological factors within the Earth system (Huggett,
236 2023). Similar to the extensive global aboveground vegetation, the root zone likely spans a vast land
237 area of approximately 103.9 million km² (70 % of total land area) (Ding et al., 2020).

238 Cutting-edge biological research has unveiled the root zone as a well-connected underground
239 network of mycorrhizal fungi, akin to a nervous system, through which resources and information
240 are exchanged among plants (Simard, 2018). This underscores that plants and their roots are integral
241 components of a cohesive ecosystem rather than isolated entities. Further exploration of these
242 interconnected relationships will be expanded upon in Section 4.1.5.

243 **3.3 Roots and the Lithosphere**

244 From a long-term evolutionary perspective, roots play pivotal roles in promoting rock weathering
245 and soil formation. Roots possess the ability to sense and navigate through heterogeneous structures,
246 circumventing obstacles at both microstructural (volumes less than mm³) and macrostructural scales
247 (Wang et al., 2020). They also act as "wedges," physically opening up substructures (Pawlik et al.,
248 2016) to access water stored in the bedrock (McCormick et al., 2021), while root-associated fungi
249 further contribute to rock weathering and the enlargement of fractures (Schenk, 2008).

250 In the short term, roots exert a stabilizing effect by anchoring soil, which reduces erosion rates,
251 enhances slope stability, and mitigates risks of landslides and debris flows (Vannoppen et al., 2017).
252 Forest communities are particularly effective in minimizing erosional processes and supporting
253 stability along hillslopes and riverbanks (Pawlik et al., 2016). Furthermore, roots significantly
254 enhance soil strength, lowering the likelihood of shallow landslides (Cohen and Schwarz, 2017).
255 The reduction in sediment flow facilitated by roots also yields basin-scale impacts on downstream
256 river geomorphology (Ielpi et al., 2022; Cao et al., 2023).

257 **3.4 Roots and the Atmosphere**

258 The root zone plays a pivotal role in land surface and climate modeling, a fact widely acknowledged
259 in scientific literature (Milly and Dunne, 1994). The moisture status within the root zone critically
260 influences water, energy and carbon exchange fluxes between the land surface and the atmosphere,
261 primarily through root water uptake and transpiration, as well as the associated latent heat flux,
262 which enhance atmospheric water vapor. Transpiration, as a component of total land evaporation,
263 depends on the balance between atmospheric water demand and the moisture supply in the
264 subsurface and accessible to plant roots. Vegetation and land cover thus actively regulate the
265 exchange of water, energy, and carbon between the root zone and the atmosphere.

266 The root zone is not only influenced by but also actively influences the Earth's climate and the large-
267 scale water cycle (Jackson et al., 1996). Firstly, evaporation entails a combined water and energy
268 flux. From that perspective, incorporating accurate root zone data has significantly improved
269 simulations of the land surface energy budget, particularly the latent heat flux (Zeng et al., 1998;
270 Zheng and Wang, 2007). Secondly, compared to soil evaporation and canopy interception, the root
271 zone of vegetation can access the deeper subsurface to extract water for transpiration. By supplying
272 vapor to the atmosphere, the root zone is therefore a major component to sustain atmospheric
273 moisture levels and the associated downwind precipitation, which shapes the long-term climate
274 pattern over considerable distances (van der Ent et al., 2010; van der Ent and Savenije, 2011; Wang-
275 Erlandsson et al., 2018). Approximately 56% of transpiration to return to land as precipitation
276 (Figure 1, and Van der Ent et al., 2014). Root-zone helps transpiration to be sustained during dry
277 periods, and plays a particularly important role for dry season length and buffers against variability
278 in precipitation (e.g., Keys et al., 2016, O'Connor et al., 2021).

279 **3.5 Roots and the Cryosphere**

280 The connection between the root zone and the cryosphere is particularly pronounced in regions
281 where vegetation covers frozen ground. In permafrost areas, the active layer undergoes seasonal
282 freeze-thaw cycles, defining the maximum depth of thawing. Roots are constrained to grow within
283 this active layer—the uppermost soil layer that thaws annually—because the permanent frozen layer

284 below it and the ice layer are impermeable to root penetration (Blume-Werry et al., 2019).
285 In cold climates, the effects of snowmelt on the inflow to the root zone and soil freeze-thaw
286 processes are crucial considerations (Gao et al., 2020, 2022; Wang-Erlandsson et al., 2016; Zhao et
287 al., 2016; Dralle et al., 2021). Climate change exacerbates these dynamics by causing permafrost
288 thawing, which expands the active layer and lowers the groundwater table. This thawing process
289 opens up previously inaccessible soil volumes for potential root growth. The expansion of the active
290 layer and changes in the groundwater table profoundly impact rooting depth and distribution in
291 permafrost regions.

292 The dynamics of the active layer and variations in root zone storage are intricately intertwined in
293 permafrost regions, influencing soil evaporation, plant transpiration, hydrologic connectivity, and
294 runoff generation (Sugimoto et al., 2002; Suzuki et al., 2021). Understanding these interactions is
295 essential for predicting how climate change will affect terrestrial ecosystems and hydrological
296 processes in cold regions.

297 **4 Heterogeneities within root zone**

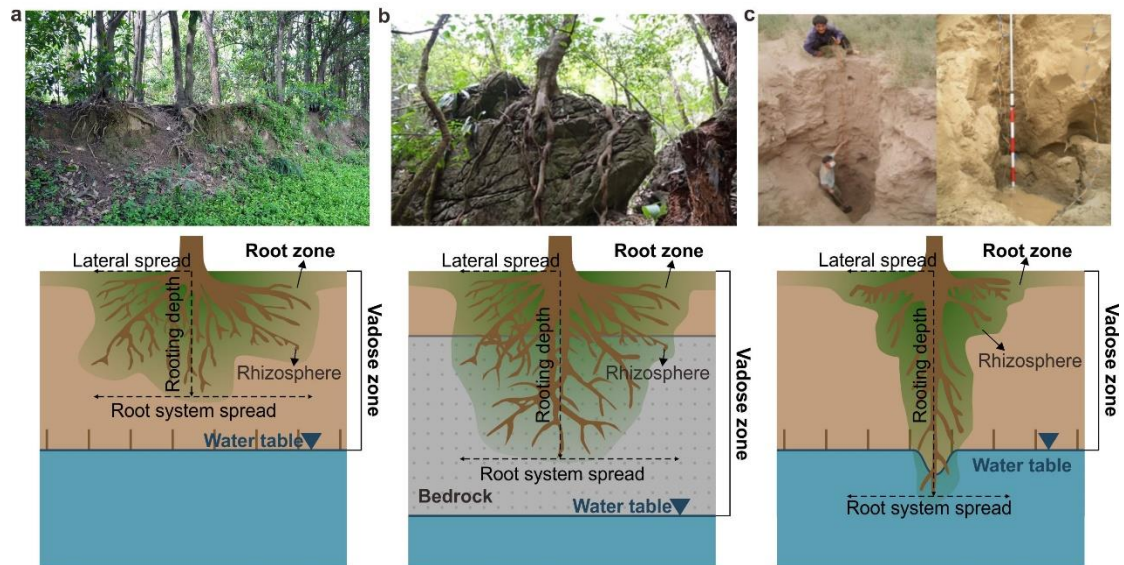
298 **4.1 Reductionist perspective: processes within the root zone**

299 From a reductionist perspective, enormous heterogeneity and complexity of root zone features are
300 characterized and catalogued by hydrologists, ecologists, pedologists and microbiologists
301 (McDonnell et al., 2007; Lin, 2010; Gao et al., 2018; Kuzyakov and Razavi, 2019). In this section,
302 we outline the complexities within the root zone across five key categories: soil heterogeneities,
303 rooting distribution heterogeneity, flux heterogeneities, plants' belowground zone of influence,
304 rhizosphere and mycorrhizal fungi.

305 **4.1.1 Soil heterogeneity**

306 Soil heterogeneity exhibits variability in both horizontal and vertical dimensions, characterized by
307 distinct horizons with diverse properties such as texture (varying proportions of sand, clay, and silt),
308 mineral composition, and organic content. Soil also possesses varied hydraulic characteristics
309 including porosity, moisture retention capacity, wilting point, and plant-available moisture. It
310 consists of a complex mixture of minerals, organic matter, living organisms, gases, and water,
311 earning it recognition as “the most complicated biomaterials on the planet” (NRC, 2009; Lin, 2010).

312 Despite its strong connection to soil, the root zone does not necessarily align with soil depth (Figure
313 3) (Hahm et al., 2024). In many cases, particularly in regions with deep soils, the root zone is
314 confined to the uppermost active layer of topsoil where the majority of hydrological and
315 biogeochemical processes occur, even though the soil may extend much deeper. In water-stressed
316 environments (Evaristo and McDonnell, 2017; Wang et al., 2022) or in regions with shallow
317 groundwater (Fan et al., 2017) roots can reach the groundwater, expanding the definition of the root
318 zone to include parts of the groundwater (see Singh et al., 2020). The depth to bedrock serves as the
319 lower boundary of the soil, defining its overall depth. However, in karst and other regions with
320 shallow soil, roots can penetrate bedrock fissures to access water despite the challenging substrate
321 conditions (McCormick et al., 2021; Lapidés et al., 2023).



322

323 Figure 3. Root zone is different from soil depth. (a) is the root zone in moderate climate. Root zone
 324 is the upper area of soil; (b) root zone contains not only soil but also bedrock; (c) root zone contains
 325 unsaturated zone and groundwater, for example the phreatophytic roots in arid regions (Wang et al.,
 326 2018).

327 4.1.2 Rooting distribution heterogeneity

328 Root distribution pattern in the soil profile plays a crucial role in determining how effectively plants
 329 can access water and nutrients. These patterns encompass the vertical and lateral spread of both
 330 coarse and fine roots, as well as the overall root density across the three-dimensional soil profile.
 331 Rooting depth, a key trait widely utilized in Earth system modeling, specifically refers to the depth
 332 to which roots extend vertically. However, it is important to note that root distribution does not
 333 always directly correlate with rooting depth (Zheng and Wang, 2007). For instance, an ecosystem
 334 dominated by deep-rooted plants with low root density may have a smaller overall root zone, and
 335 thus a smaller accessible water volume, compared to one with shallow-rooted plants that have higher
 336 root density (Singh et al., 2020; van Oorschot et al., 2021).

337 Despite the critical importance of root distribution and concerted ongoing research efforts that
 338 compile an impressive richness of observations (e.g. Schenk and Jackson, 2005; Tumber-Davila et
 339 al., 2022), upscaling root distributions remains challenging. This is due to the fact that although
 340 observations are available for several thousand of individual plants globally, these are essentially
 341 point/plot-scale observations for which spatial covariance fields remain largely unknown. It is
 342 worth emphasizing that even with detailed data on root distribution for individual plants,
 343 determining the precise extent of the root zone is challenging due to the complex network of pores,
 344 fissures in soils and bedrock, intricate hydrodynamic gradients, and mycorrhizal fungi network (e.g.
 345 Casper et al., 2003; see Figure 4).

346 4.1.3 Flux heterogeneity: preferential flow

347 Traditionally in hydrology, the infiltration process was believed to be primarily controlled by matrix
 348 flow, assuming soil texture dictates hydraulic properties and infiltration capacity (Beven, 2021).
 349 However, mounting evidence now indicates that preferential flow routes rapidly transmit free water
 350 through the subsurface, dominating water movement across various scales (Uhlenbrook, 2006; Liu

351 et al., 2012; Beven and Germann, 2013). In water-limited environments, the rooting systems of
352 plants strongly influence infiltration capacity, velocity, and depth (Schenk & Jackson, 2002; Schenk,
353 2008). Infiltration replenishes moisture in the root zone through preferential pathways, while excess
354 water either runs off or percolates into groundwater.

355 Recent research highlights the crucial role of mucilage exuded from roots and rhizosphere bacteria
356 in forming liquid bridges at the root–soil interface. These bridges significantly impact the
357 mechanical stability and hydraulic properties within the root zone (Bengough, 2012). The
358 complexity introduced by preferential flow channels and their influence on hydraulic properties
359 challenges the traditional soil-centric Darcy-Richards framework (Beven, 2012; Gao et al., 2023b).

360 **4.1.4 Plant’s belowground zone of influence**

361 The concept of a plant's belowground zone of influence refers to how roots absorb water and
362 nutrients not from a fixed area, but from a highly dynamic and irregular zone (Casper et al., 2003).
363 This zone exhibits complex and variable hydraulic and nutrient gradients between roots and their
364 surroundings. From a reductionist, root-centered micro-scale perspective, the belowground zone of
365 influence quantifies how the influence of roots decreases with distance from the stem. The
366 hydrological processes and shape of root zones are extremely variable, plastic, irregular, and
367 dynamic, making it effectively impossible to quantify or model this zone accurately.

368 However, viewing the root zone as a holistic system reveals emergent behaviors and spatial patterns
369 that enable us to move beyond the complexities of gradient-based models and the heterogeneity of
370 the plant's belowground zone of influence (see Figure 4).

371 **4.1.5 Rhizosphere and mycorrhizal fungi**

372 The rhizosphere, as the dynamic area at the plant root-soil interface, is characterized by a multitude
373 of biogeochemical processes driven by physical activities, such as water and nutrient dynamics
374 (Figure 4a, and Bengough, 2012; Daly et al., 2017). The rhizosphere supports microbial growth (e.g.,
375 populations of bacteria ranging from 10^7 to 10^{12} per gram of rhizosphere soil), thereby stimulating
376 biochemical processes crucial for plant nutrition and health (Hinsinger et al., 2009). Moreover, the
377 mucilages and exudates of plants’ roots, forming as soil aggregates, significantly alter rhizosphere’s
378 hydraulic properties, such as infiltration capacity, water holding capacity, and preferential flow
379 (Daly et al., 2017).

380 Mycorrhizal fungi form a complex network of roots and fungal hyphae, often referred to as the
381 'wood wide web' (Beiler et al., 2009). These fungi create an underground continuous network that
382 connects an estimated 90% of land plant species (Bonfante and Genre, 2010), facilitating the
383 exchange of resources such as carbon, water, and nutrients, as well as information about
384 environmental conditions and threats such as insect infestations (Bonfante and Genre, 2010).

385 The existence of mycorrhizal fungi dramatically increases the root zone water and nutrient
386 absorption capacity. The absorption fine roots are usually thicker than 0.2 mm (Strand et al., 2008;
387 Taylor et al., 2013), which means roots can only grow in macropores. Thus roots have limited
388 absorption capacity, can only use 4-7% of available soil volume (Guo et al., 2008). But one spoon
389 of soil can have mycorrhizal fungi as long as 1 km, and as thin as 2-10 microns (Allen, 2007), which
390 allows mycorrhizal fungi growing in micropores among fine minerals, which are unreachable for
391 roots. These fungi enhance plant water uptake, to be 7 times higher than without mycorrhizal fungi

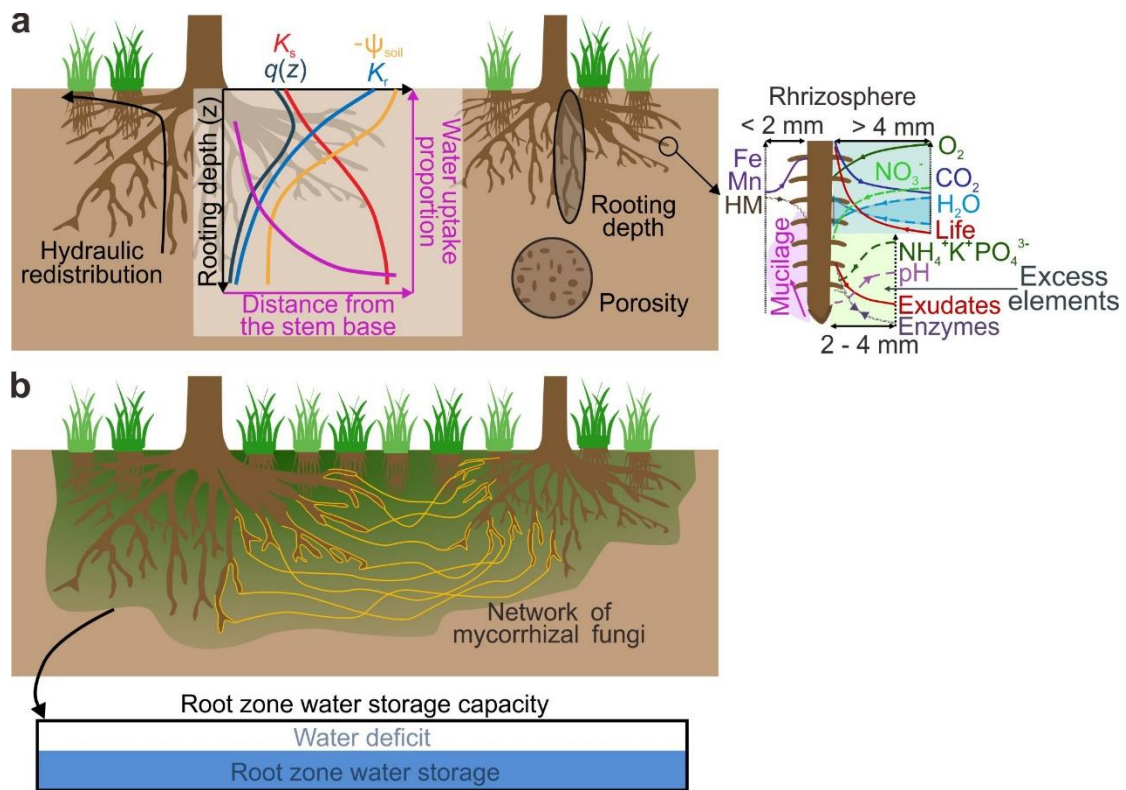
392 (Zhang et al., 2018), particularly beneficial during periods of water stress (Augé, 2001; Püschel et
393 al., 2020), by promoting water availability and transportation.

394 The intricate network of mycorrhizal fungi functions akin to neural tissue in animals, allowing the
395 root zone to respond to environmental changes—such as climate variations, water availability,
396 nutrient levels, and threats—in a holistic and predictable manner.

397 **4.2 Reductionists' root zone models**

398 Although the hydrology community has advocated moving beyond heterogeneities for some time
399 (McDonnell et al., 2007), reductionist modeling remains prevalent in describing the root zone across
400 hydrological models, dynamic vegetation models (DGVMs), and land surface schemes in Earth
401 system models. Rooting depth varies among plant functional types (PFTs) and is in these models
402 typically treated as a fixed parameter using lookup tables. Soil water holding capacity, determined
403 in laboratories, defines the amount of water retained under gravitational free drainage (van Oorschot
404 et al., 2021). Soil depth, ranging widely from 1 meter to over 10 meters or until bedrock, is a critical
405 parameter (Hidy et al., 2021; Wiltshire et al., 2021). Soil profiles are segmented into layers, varying
406 from 3 to more than 10 (Seneviratne et al., 2010). The maximum storage capacity of the root zone,
407 S_{Rmax} , combines rooting depth and soil water holding capacity derived from soil texture data. In
408 advanced models such LPJ-GUESS, root distribution follows an exponential function constrained
409 by maximum rooting depth, groundwater levels, and nutrient distribution (Smith et al., 2014).
410 Furthermore, DGVMs integrate dynamic root modules to simulate root growth, architecture, and
411 distribution in response to soil water and nitrogen availability (e.g., Lu et al., 2019).

412 However, we argue that this traditional perspective of the root zone, which assumes the whole is the
413 sum of its parts (see Table 1 and Figure 4a), is subject to several issues. First, the intricate network
414 of pores and fissures in the substrate, extending both laterally and in depth, with complex gradients
415 of water and nutrients, cannot be accurately simulated by mechanically combining rooting depth
416 and soil properties (Figure 4a). Second, this approach lacks detailed observations of root density,
417 vertical and lateral distribution, leading to significant uncertainties related to either the use of strong
418 assumptions that may not be supported by data or heavy reliance on parameter calibration. Third,
419 the method overlooks fractal patterns in soil structure, pore networks, root morphology, and
420 preferential flow paths. These fractal patterns embody organized complexity or simplicity that could
421 enhance the realism of models, but they are typically broken up in reductionist approaches. Fourth,
422 the root zone exhibits gradients in chemical, biological, and physical properties that vary radially
423 and longitudinally along roots (McNear, 2013; Kuzyakov and Razavi, 2019). For instance, hydraulic
424 redistribution is crucial for trees, redistributing water vertically from deeper, wetter soil layers to
425 shallower, drier layers during nighttime (Figure 4a) (Bleby et al., 2010; Nadezhdina et al., 2010),
426 but also laterally between individual plants (Hafner et al., 2021). Fifth, recent studies of the
427 rhizosphere highlight the extraordinary complexity of underground root systems (e.g. Daly et al.,
428 2017), which are difficult to adequately describe by reductionist methods. However, there is
429 evidence that from these small-scale heterogeneities relatively stable and simple pattern emerge on
430 larger scales. For example, small-scale studies of mycorrhizal fungi demonstrate that the root zone
431 system responds as a whole to changes, providing a microscale foundation for holistic approaches
432 to root zone modeling (Figure 4).



433

434 Figure 4. Comparison of reductionist (a) and holistic (b) views of root zone and root zone water
 435 storage capacity. (a) The reductionist perspective estimates root zone storage capacity and
 436 transpiration based on rooting depth and soil properties, assuming a static view primarily relying on
 437 soil information. Root water uptake as a function of depth ($q(z)$) is shown for a soil profile with a
 438 dry upper layer (leading to more negative soil water potential (ψ_{soil}) and low soil hydraulic
 439 conductivity (K_s)) and a deeper soil layer with higher water content (leading to less negative ψ_{soil} and
 440 higher K_s). At the whole-plant scale, root conductivity (K_r) is correlated with the root surface area,
 441 which decreases exponentially with depth. Proportion of plants that took up water as a function of
 442 distance from the stem based is shown on the right side of graph, which decreases with distance
 443 (adapted from Casper et al., 2003; Tumber-Dávila et al., 2022; Bachofen et al., 2024). Top right
 444 figure shows the generalization of rhizosphere extents and gradient types for the most investigated
 445 parameters: Gases, Root exudates, Nutrients and Excess elements, pH and Eh, Enzyme activities
 446 and microorganisms (from Kuzyakov and Razavi, 2019) (b) The holistic perspective views root
 447 zone storage as a volume per unit area, emphasizing an average depth. It estimates root zone capacity
 448 and transpiration based on plant water demand, reflecting a dynamic view influenced by climate
 449 and plant adaptation. The mycorrhizal fungi network enables the root zone to respond to
 450 environmental changes as a unified system.

451

452 5 Emergent behavior of root zone as a holistic system

453 5.1 Holistic perspective in hydrology

454 Holistic perspectives regard the root zone as an integrated living system shaped by fractal patterns,
 455 influenced by long-term self-organization and co-evolution. In contrast, reductionist models focus

456 on determining "How deep is the root zone?". Holistic approaches instead ask broader questions
 457 such as "What is the size of the root zone?" and "How much moisture can it buffer?". Viewing the
 458 root zone holistically allows for more predictable behavior, enabling simulation with simpler models,
 459 based on widely and readily available data.

460 5.1.1 Holistic root zone approach

461 The root zone, located beneath the surface and not directly observable, can be characterized using
 462 an inverse modeling approach to infer its systematic behavior. Various parameterizations driven by
 463 land surface water and energy fluxes, such as precipitation, evaporation, radiation, and biomass
 464 observations like net or gross primary production (Kleidon, 2004), play crucial roles in this process.

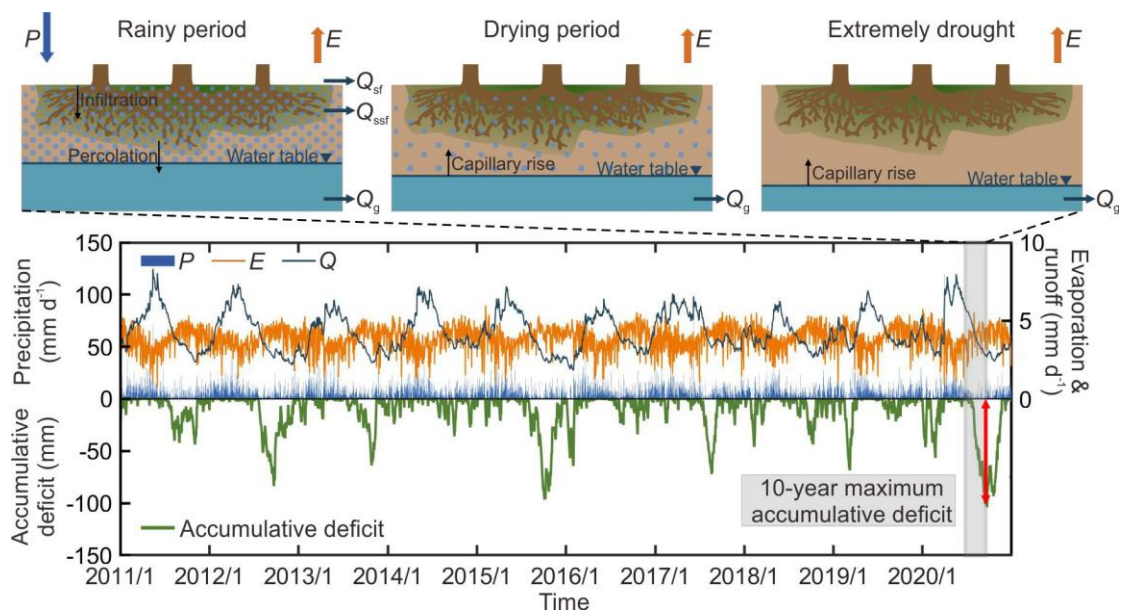
465 The mass curve technique (MCT), elsewhere also referred to as Memory Method (van Oorschot et
 466 al., 2021, 2024) serves as an inverse model that employs a water-deficit approach based on observed
 467 land surface water budgets to estimate root zone processes (see Figure 5) (Gao et al., 2014a; Wang-
 468 Erlandsson et al., 2016; Kühn et al., 2022). This holistic method allows hydrologists and ecologists
 469 to derive S_{Rmax} from ecosystem-scale observations of liquid water inflow (F_{in}), i.e. the sum of
 470 precipitation (de Boer-Euser et al., 2016), snowmelt (Dralle et al., 2021), and irrigation (van
 471 Oorschot et al., 2024) and water outflow (F_{out} , evaporation). Time series data of inflow and outflow
 472 are utilized to infer cumulative water deficits during dry periods (Equations 3 and 4). The largest
 473 cumulative deficit observed over a specific return period can be considered as representing the actual
 474 S_{Rmax} (Equation 5).

$$475 \quad A_{t_n \rightarrow t_{n+1}} = \int_{t_n}^{t_{n+1}} F_{out} - F_{in} dt \quad (3)$$

$$476 \quad D_{t_{n+1}} = \max(0, D_{t_n} + A_{t_n \rightarrow t_{n+1}}) \quad (4)$$

$$477 \quad S_{Rmax} = \max(D_{t_1}, D_{t_2}, \dots, D_{t_{end}}) \quad (5)$$

478 $A_{t_n \rightarrow t_{n+1}}$ is the water deficit in t_{n+1} day, $D_{t_{n+1}}$ is the accumulative water deficit in day t_{n+1} .



479
 480 Figure 5. The dynamic change of water deficit in root zone from the wet season to dry season. On

481 the upper panel, the conceptual illustration of the water deficit change in root zone is shown with
482 its effect on hydrology. P and E represent precipitation and evaporation respectively. Q_{sf} , Q_{ssf} and
483 Q_g represent surface, subsurface and groundwater runoff respectively. The lower panel shows the
484 precipitation (P), evaporation (E), runoff (Q) and accumulative water deficit in root zone from 2011
485 to 2020. Note that the real world root zone storage capacity cannot be smaller than the value S_{Rmax}
486 estimated with the above method, as otherwise there would not have been sufficient water available
487 and accessible for vegetation to sustain the observed evaporation (F_{out}). S_{Rmax} is therefore in any
488 case a minimum and thus a lower limit to the root zone storage capacity. In principle, it can be
489 argued that S_{Rmax} could be larger than that. However, increasing evidence from studies that estimate
490 S_{Rmax} with optimality approaches (e.g. Kleidon, 2004; Guswa, 2008; Schymanski et al., 2008) or as
491 model calibration parameter (e.g. Gao et al., 2014; Njzink et al., 2016; Hrachowitz et al., 2021;
492 Wang et al., 2024) suggests that it is not likely that the real world root zone storage capacity is larger
493 than the S_{Rmax} associated with an ecosystem-specific dry spell return period. Instead, the results of
494 these studies suggest that vegetation optimizes its below- and above-surface resource allocation to
495 allow sufficiently large root systems to access water (and nutrients) in dry periods while also
496 allowing sufficient above-surface growth to allow survival in the competition for light and for
497 increased strength against windfall. In other words, the root system and thus S_{Rmax} is as large as
498 necessary but not larger than that.

499

500 The holistic perspective requires integrated measurements (Reichstein et al., 2014), incorporating
501 methods such as lysimeters and eddy covariance for determining land surface fluxes, satellite remote
502 sensing, and catchment-scale water balance assessments (Gao et al., 2014a; Wang-Erlandsson et al.,
503 2016; Kühn et al., 2022; Stocker et al., 2023). Lysimeters are particularly valuable tools for
504 quantifying the water balance in the root zone, offering unique insights, especially in comparisons
505 between vegetated and non-vegetated systems, highlighting the influence of vegetation dynamics
506 on the water cycle (Seneviratne et al., 2012; Scanlon et al., 2005).

507 Satellite remote sensing provides extensive spatial and temporal data on land surface water fluxes,
508 including precipitation and evaporation, crucial for deriving root zone processes. Catchment-scale
509 water balance studies encompass variables such as precipitation and streamflow, offering valuable
510 insights into long-term average evaporation patterns. Additionally, S_{Rmax} serves as a key parameter
511 in conceptual hydrological models, often calibrated based on catchment precipitation and
512 streamflow data (Gao et al., 2014b; Liang et al., 2024; Wang et al., 2024).

513 **5.1.2 Climate and topography controls S_{Rmax}**

514 The holistic perspective in modeling frameworks can be traced back to the bucket model used in the
515 first global climate model by Manabe in 1969. In this early study, S_{Rmax} was globally set at 150mm,
516 providing a reasonable estimation of the global average (Wang-Erlandsson et al., 2016; Gao et al.,
517 2023a; Stocker et al., 2023). However, S_{Rmax} varies significantly across different climate zones and
518 ecosystems.

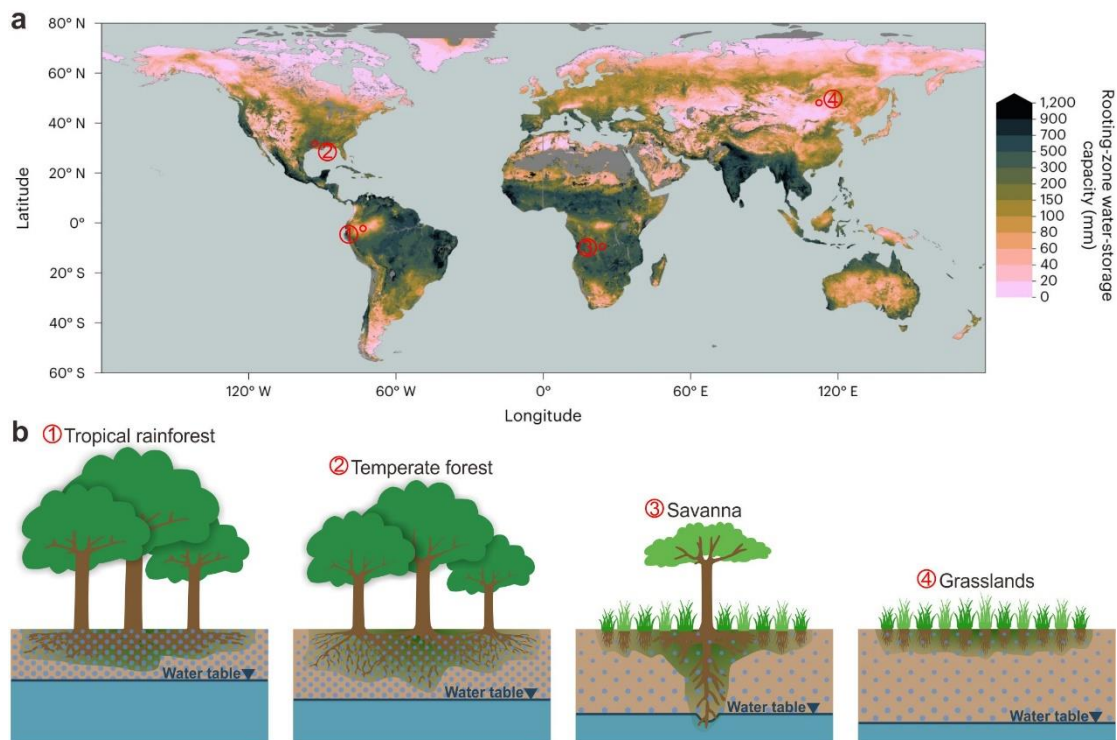
519 In tropical rainforests (see Figure 6), where rainfall is abundant throughout the year, S_{Rmax} tends to
520 be small because water availability is not a limiting factor, allowing forests to develop tall canopies
521 to compete for light. In temperate climates, where rainfall starts to constrain growth, ecosystems

522 typically allocate more resources underground, enhancing their root zone capacity. In savannas,
523 characterized by strong precipitation seasonality, ecosystems further expand their root zones to cope
524 with prolonged droughts, often at the expense of aboveground biomass. Grasslands, on the other
525 hand, may enter dormancy during droughts, exhibiting very limited S_{Rmax} (Figure 6).

526 Interestingly, research has shown that ecosystems optimally adjust their root systems to their
527 specific environments (Schymanski et al., 2008; Guswa, 2010), adapting to drought periods that
528 vary from 5 to 40-year return periods depending on the ecosystem type (evergreen, deciduous,
529 grassland, etc.) (Gao et al., 2014a; Wang-Erlandsson et al., 2016).

530 At the landscape scale, the spatial distribution of S_{Rmax} is largely influenced by topography (Fan et
531 al., 2017; Gao et al., 2019). Topography also plays a critical role in determining runoff generation
532 mechanisms, such as saturated overland flow in wetlands, subsurface storm flow in hillslopes, and
533 infiltration excess overland flow on terraces (Savenije et al., 2010; Gharari et al., 2014).

534



535

536 Figure 6. (a) The global root zone storage capacity (from Stocker et al., 2023). (b) The conceptual
537 illustration of aboveground biomass and root zone in different climate zones.

538 5.1.3 Holistic root zone impacts on groundwater

539 The influence of the root zone on catchment hydrology is prominently observed through its impact
540 on groundwater dynamics and baseflow recession. The root zone plays a crucial role in controlling
541 water percolation to groundwater (Collenteur et al., 2021). Groundwater recharge and streamflow
542 generation are closely linked, as runoff is typically derived from groundwater under most conditions
543 (Ali et al., 2011). Additionally, the root zone influences evaporation and plant transpiration from
544 groundwater sources. Surface soil and root zone drying are moderated by upward capillary flow

545 from the subsurface, sustaining evaporation during dry periods while reducing groundwater
546 recharge over longer timescales.

547 A well-documented phenomenon is the diurnal cycle of streamflow and groundwater table
548 fluctuations, particularly notable in arid and semi-arid catchments (Lundquist and Cayan, 2002;
549 Wang et al., 2014; Yue et al., 2016). Vegetation, particularly in riparian areas, directly accesses
550 groundwater for transpiration. Daytime root water uptake leads to declines in water tables, whereas
551 water tables recover at night when transpiration rates diminish (Yue et al., 2016). This illustrates the
552 direct hydraulic connection between root water uptake, groundwater dynamics, and river streamflow.

553 **5.2 Holistic root zone in hydrology modeling**

554 The concept of thresholds and hydrologic connectivity associated with runoff processes is crucial
555 for understanding hydrological responses at both hillslope and catchment scales (Saffarpour et al.,
556 2016), with the root zone playing a pivotal role as an integrated system. It is widely observed that
557 runoff generation does not occur until the root zone reaches a critical moisture threshold, known as
558 field capacity (Tromp-van Meerveld and McDonnell, 2006). Once this threshold is surpassed, the
559 outflow pathways are activated, leading to runoff generation. This mechanism is often referred to
560 as fill-and-spill (McDonnell et al., 2021) or store-and-pour (Phillips, 2022).

561 At the landscape scale, topography emerges as a primary factor influencing the mechanisms of
562 runoff generation. Due to landscape heterogeneity (including topography and vegetation cover), the
563 S_{Rmax} exhibits spatial variability. In many cases, S_{Rmax} increases with the elevation above the nearest
564 drainage (HAND) (Fan et al., 2017). Consequently, saturation and runoff generation do not occur
565 uniformly across an entire catchment (Hewlett and Hibbert, 1967; Ambroise et al., 2004; Beven,
566 2012; Gao et al., 2019). Typically, during the onset of wetting seasons, riparian zones near the
567 drainage experience the critical threshold first, initiating saturation excess flow. As precipitation and
568 root zone moisture increase, the contributing areas expand, a phenomenon known as the variable
569 contribution area theory (Ambroise et al., 2004). Accounting for this spatial heterogeneity of the
570 root zone significantly improves the performance of runoff simulations (Gao et al., 2019).

571 These conceptual models are widely applied in global hydrological prediction practices, albeit with
572 varied parameterizations and terminologies (e.g. Zhao, 1992; Perrin et al., 2003; Clark et al., 2008;
573 Beven, 2012; Fenicia et al., 2014; Gao et al., 2019). In land surface and global hydrology studies
574 comparing reductionist and holistic approaches, it is noteworthy that despite its simplicity, the
575 holistic method often outperforms more complex reductionist approaches in runoff simulation (Mao
576 and Liu, 2019; Wang et al., 2021), particularly in ungauged basins (Hrachowitz et al., 2013). Recent
577 studies highlight the potential of using climate-controlled root zone parameters to enhance water
578 flux simulations in land surface models such as HTESSSEL (Hydrology Tiled ECMWF Scheme for
579 Surface Exchanges over Land), resulting in improved discharge simulation correlations across
580 different regions (van Oorschot et al., 2021).

581 Moreover, the holistic approach to the root zone is gaining traction in studies of water quality, solute
582 transport, and transit times, presenting a promising new framework for integrating material transport
583 through the entire system (Harman and Fei, 2024).

584 **6 Root zone in the Anthropocene**

585 **6.1 The impacts of climate change**

586 Climate exerts significant control over root zone dynamics. The root zone plays a critical role in
587 determining ecosystem resilience to droughts and climate variability, alongside other factors such
588 as shifts in species composition and physiological adjustments like leaf shape and stomatal
589 regulation (Zhang et al., 2023). Root growth is highly dynamic and responsive to environmental
590 changes, with rapid carbon transfer from leaves to roots and soil organisms occurring within hours
591 to days across different ecosystems (Melkikh and Sutormina, 2022).

592 The spatial variability of $S_{R_{max}}$, reflecting climatic conditions, has been well documented (Kleidon
593 and Heimann, 1998; Gao et al., 2014a; Wang-Erlandsson et al., 2016; Stocker et al., 2023),
594 suggesting that temporal variations in $S_{R_{max}}$ are also plausible under changing climate regimes.
595 Local factors such as topography and groundwater dynamics further influence root zone dynamics,
596 with observations indicating rapid root growth towards declining water tables in desert phreatophyte
597 vegetation (Orellana et al., 2012; Wang et al., 2018; Kuzyakov and Razavi, 2019).

598 At the basin scale, temporal evolution of $S_{R_{max}}$ in response to climatic variability across multiple
599 decades, has been reported by several studies (Bouaziz et al., 2022; Tempel et al., 2024; Wang et al.,
600 2024). Globally, analyses based on ERA-5 reanalysis data indicate an 11% increase in global
601 average $S_{R_{max}}$ over the past four decades (1982-2020), driven by intensifying drought conditions (Xi
602 et al., 2024). This widespread increase in $S_{R_{max}}$ has been validated through dynamic identifiability
603 analysis (DYNIA) algorithms using observed hydrometeorological data from U.S. catchments
604 (Liang et al., 2024).

605 **6.2 The impacts of land-use and land management activities**

606 **6.2.1 Agriculture**

607 Agricultural activities exert significant influence on the root zone, impacting various aspects crucial
608 for food production and food security (Eisenhauer et al., 2024). Practices such as irrigation,
609 fertilization, tilling, non-point source pollution, and salinization predominantly occur within the root
610 zone, shaping its characteristics and functionality.

611 Ploughing, a fundamental agricultural practice, plays a pivotal role in altering and determining root
612 zone depth. Over time, intensive cultivation can lead to the formation of a plough pan, typically
613 situated 15 to 30 cm below the soil surface in clay-rich soils. This compacted layer acts as a barrier
614 beneath the root zone, limiting water percolation and restricting root growth to the relatively shallow
615 soil layer above it (Li et al., 2019). In such scenarios, the moisture holding capacity of the topsoil
616 becomes critical in defining the root zone storage capacity.

617 In irrigation agriculture, where additional water is supplied to the root zone during dry periods, the
618 $S_{R_{max}}$ often becomes smaller compared to natural conditions with similar climates (Xi et al., 2021;
619 Hauser et al., 2022; van Oorschot et al., 2024). Studies indicate that due to irrigation practices
620 (excluding considerations of climate change impacts), rooting depth can be reduced by
621 approximately 5%, equivalent to an approximate loss of 8 cm in depth or about 11,600 km³ of rooted
622 volume (Hauser et al., 2022). At a global scale, research has shown a consistent decrease in $S_{R_{max}}$ in

623 croplands when irrigation practices are factored in as an additional water supply (van Oorschot et
624 al., 2024).

625 **6.2.2 Deforestation / Afforestation**

626 Deforestation can significantly reduce S_{Rmax} (Hrachowitz et al., 2021). Recovery of S_{Rmax} following
627 deforestation can take more than a decade to restore to previous levels (Nijzink et al., 2016).
628 Conversely, regions experiencing woody encroachment often see root depths deepening by
629 approximately 38 cm compared to the previous dominant vegetation (Hauser et al., 2022).

630 The root zone's role extends to evaluating the outcomes of large-scale ecological initiatives. For
631 example, projects like China's "grain-to-grass" and "three norths afforestation projects" have proven
632 beneficial for soil conservation, flood control, biodiversity enhancement, and carbon sequestration
633 (Chen et al., 2019; Li et al., 2021). However, these initiatives have also increased water consumption
634 through evaporation, posing challenges for water management in arid regions where societal water
635 demands compete with ecosystem services (Feng et al., 2016). The underground root zone is the
636 key piece in the puzzle, to connecting the isolated dots, and explaining all these ecological and
637 hydrological phenomena in one framework.

638 Transitioning from traditional blue water management to integrated water resources management
639 involves incorporating both blue- and green-water considerations—surface and groundwater—into
640 a unified framework (Falkenmark, 2000). Root zone moisture plays a pivotal role in linking land
641 surface fluxes, water resources, soil erosion, and carbon sequestration, thereby influencing the
642 success and sustainability of such ambitious ecological projects (Sun et al., 2020).

643 **6.2.3 Urbanization**

644 Urbanization has profound effects on the root zone, transforming it from a permeable to an
645 impermeable structure, which leads to increased stormwater runoff and reduced vegetation
646 evaporation, as noted by Hao et al. (2018). Furthermore, urbanization significantly impacts the
647 underground mycorrhizal fungi network, altering the hydrological processes within the root zone
648 and catchment areas, as well as influencing the land surface energy budget. These changes
649 contribute to intensified urban inundation and exacerbate the urban heat island effect (Fletcher et
650 al., 2013).

651 There is growing interest in leveraging vegetation to manage hydrology in urban environments,
652 exemplified by initiatives like China's "Sponge City" initiative (Xia et al., 2017). The root zone acts
653 as a natural sponge in larger natural landscapes such as forests and grasslands, while green roofs
654 and urban gardens serve as smaller-scale sponges within cities and neighborhoods. Increasing the
655 S_{Rmax} has the potential to enhance flood prevention, improve water purification, mitigate the urban
656 heat island effect, and enhance the aesthetic value of urban areas (Palmer et al., 2015).

657 **7 Outlook**

658 **7.1 Observations**

659 Accurately measuring root traits and characteristics of the root zone is among the most challenging
660 tasks in ecological, agricultural, and hydrological research, particularly in field studies. For instance,
661 determining the rooting depth of grasses and crops typically involves soil core sampling, while

662 estimating coarse tree roots relies on allometric equations, and fine tree roots are assessed using soil
663 cores. Despite efforts such as the global rooting depth synthesis study covering 2200 observations
664 worldwide (Fan et al., 2017) and others (Tumber-Davila et al., 2022), this remains a relatively small
665 sample compared to the vast number of 3.04 trillion trees globally (Crowther et al., 2015), not to
666 mention the numerous grasses and agricultural crops. Consequently, descriptions of spatial
667 distribution of root systems remain uncertain at scales from hillslope to global.

668 Remote sensing products such as SMAP and SMOS provide relatively high-resolution estimates of
669 near-surface moisture, but their capabilities are confined to the top few centimeters of soil (Xu et
670 al., 2021). Estimating root zone moisture from surface measurements involves specific assumptions
671 and different types of models (Entekhabi et al., 2010; Reichle et al., 2019; Bouaziz et al., 2020; Kim
672 et al., 2023). In Earth System Models (ESMs), uncertainties in projections of dry-season water
673 availability are considerable, sometimes exceeding 200% of the ensemble mean, underscoring the
674 need for more accurate observational data (Dong et al., 2024, under review).

675 To deepen our quantitative descriptions of root zone hydrology, in particular from a holistic
676 perspective, additional experiments and measurements across diverse climates and landscapes are
677 crucial. Methods such as field control experiments, rhizobox studies (Nie et al., 2013; Zhou et al.,
678 2022; Maan et al., 2023), and long-term lysimeter measurements (Scanlon et al., 2005) are essential.
679 Interestingly, rhizobox experiments have shown that detailed root distribution information may not
680 be necessary to accurately estimate water budgets (Maan et al., 2023). This prompts the question:
681 which variables should we prioritize observing to enhance our understanding of root zone processes?

682 **7.2 Root zone biogeochemistry**

683 In this opinion paper, our focus has centered on root zone hydrology, recognizing water as the crucial
684 link among Earth system spheres. We contend that the root zone, viewed holistically, holds broad
685 implications for biogeochemistry studies, encompassing carbon, nitrogen, phosphate, pollutants,
686 and microbial communities, thereby influencing Earth system science on a global scale. The
687 burgeoning field of root microbiome research in ecology is poised to significantly deepen our
688 understanding of rhizosphere biotic processes within the root zone, their ecosystem-wide
689 importance, and their impacts on large-scale biogeochemical cycles (McNear, 2013).

690 The successful integration of the holistic root zone concept into hydrology and land surface models
691 sets the stage for incorporating more biogeochemical processes, such as carbon and nutrient
692 dynamics, as part of an integrated system. For instance, considering carbon dynamics, plants with
693 roots play a pivotal role in the global carbon budget by sequestering CO₂ through photosynthesis
694 and releasing CO₂ via respiration, influencing the climate system significantly (Bian et al., 2023).
695 Root zone dynamics are crucial for carbon sequestration strategies, particularly in predicting plant
696 responses to elevated CO₂ levels under climate change scenarios (Nie et al., 2013; Bian et al., 2023).
697 Among carbon storage alternatives, root zone carbon presents the highest uncertainty, yet it holds
698 substantial potential for carbon neutrality and sequestration strategies, necessitating further
699 experimental and modeling investigations (Friedlingstein et al., 2022).

700 The holistic modeling framework proves particularly adept at integrating biogeochemical processes,
701 leveraging conceptual models as the primary approach to simulate fluxes of carbon, nitrogen, and
702 phosphorus. This approach not only enhances our understanding of root zone processes but also
703 contributes to advancing our capabilities in predicting and managing Earth system dynamics in

704 response to global change (Violle et al., 2014; Reichstein et al., 2014).

705 **7.3 Making models alive for future prediction**

706 The call to integrate a more dynamic and "alive" root zone into Earth system models is urgent and
707 critical for improving predictions across various disciplines (Wang et al., 2018). Static root models
708 have shown significant discrepancies in simulating land surface water and energy dynamics (Jing et
709 al., 2014; Cai et al., 2018a; Drewniak, 2019; Liu et al., 2020b; Zheng and Wang, 2007), largely due
710 to their inability to capture the adaptability of root zones to changing environmental conditions.
711 Despite pioneering efforts to incorporate dynamic root behavior into land surface and dynamic
712 vegetation models (Wang Yuanyuan et al., 2018; Lu et al., 2019; Sakschewski et al., 2021),
713 reductionist modeling remains dominant in Earth system models.

714 Currently, holistic root zone modeling primarily serves as a diagnostic approach. However, its
715 potential lies in offering simpler and more realistic simulations. Moving forward, several
716 approaches can enhance the integration of holistic perspectives into Earth system models:

- 717 1. Space-for-Time Exchange: Analyzing spatial transitions in ecosystems can provide insights
718 into ecohydrological strategies and help predict hydrological responses under future
719 climate scenarios. For instance, using methods like the mass curve technique (MCT), Singh
720 et al. (2020) demonstrated how S_{Rmax} changes spatially with climate, offering a framework
721 to forecast belowground and aboveground biomass variations.
- 722 2. Optimization Approach: Leveraging optimality principles to integrate water, carbon,
723 nutrient dynamics, and vegetation responses based on ecological optimality can enhance
724 predictions of root dynamics and their impacts on environmental changes (Schymanski et
725 al., 2008; Wang Ping et al., 2018; Hunt et al., 2024).
- 726 3. Incorporating Physical Laws: Earth system modeling currently relies heavily on Newtonian
727 laws of mass, energy, and momentum conservation, typical of reductionist approaches.
728 However, broadening the scope to include additional physical laws, such as evolutionary
729 theory and the second law of thermodynamics, can reduce parameter calibration needs and
730 constrain model uncertainties (Savenije, 2024).

731 These approaches collectively aim to advance our understanding and prediction capabilities of Earth
732 system dynamics, emphasizing the need for more integrated and comprehensive models that capture
733 the dynamic interactions within the root zone and beyond.

734 **7.4 Planetary stewardship**

735 The root zone, as a critical interface between natural geosphere and anthroposphere (Fig. 3 in Steffen
736 et al., 2020), holds profound implications for planetary stewardship and sustainable development.
737 Human activities like agriculture, urbanization, deforestation/afforestation, land use change, the use
738 of pesticides and fertilization profoundly impact the root zone. Understanding the root zone is
739 pivotal for managing green water footprints, integrated water resources, and carbon sequestration
740 (Wang-Erlandsson et al., 2022). It serves a crucial role in dividing precipitation into green water
741 (used by terrestrial ecosystems) and blue water (available for human use), thereby influencing Earth
742 system resilience (Falkenmark, 2000).

743 The management of the root zone is essential for sustaining both green and blue water within safe
744 planetary boundaries, crucial for achieving global sustainability development goals (SDGs) such as

745 SDG2 (zero hunger), SDG6 (clean water and sanitation), SDG11 (sustainable cities and
746 communities), SDG13 (climate action), and SDG15 (life on land). Proper management practices are
747 therefore indispensable for maintaining Earth system resilience and achieving these SDGs (Stewart-
748 Koster et al., 2023; Wang-Erlandsson et al., 2022).

749 Integrating the impacts of human activities on the root zone into Earth System Models (ESMs) is
750 crucial for advancing scientific understanding, informed decision-making, and effective
751 management of the root zone. This integration will enhance our ability to predict and mitigate the
752 consequences of human actions on water resources, ecosystems, and global climate dynamics.

753 **8 Concluding remarks**

754 The root zone is the crucial element linking multi-spheres of the Earth surface system, including
755 hydrosphere, biosphere, lithosphere, atmosphere, cryosphere, and anthroposphere. Although many
756 disciplines are studying the root zone from different angles, this is not yet done in a systematic
757 way. This study explored the differences and linkages between the root zone and other “similar”
758 terminologies, such as vadose zone, critical zone, rhizosphere, and rooting depth. For the root
759 zone in Earth system studies, we underscored the heterogeneity within the root zone, including
760 complexities of soil, root distribution, preferential flow, plants’ belowground zone of influence,
761 rhizosphere and mycorrhizal fungi. However, viewing the root zone as an integrated living system,
762 influenced by long-term self-organization and co-evolution, allows for emergent and predictable
763 behavior, enabling simulation with simpler models, based on widely and readily available data.
764 We advocate for a paradigm shift towards ecosystem-centered root zone studies in Earth system
765 science, to develop ‘living’ models for more realistic future prediction, particularly in response to
766 climate change and intensifying human activities in the Anthropocene.

767

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