

Observation and modelling of atmospheric OH and HO₂* radicals at a subtropical rural site and implications for secondary pollutants

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Abstract

HO_x radicals (OH and HO₂) are crucial oxidants that determine atmospheric oxidation capacity and the production of secondary pollutants; however, their sources and sinks remain incompletely understood in certain rural, forest and maritime environments. This study measured HO₂* (HO₂ + parts of RO₂) and OH concentrations using a chemical ionisation mass spectrometer at a subtropical rural site in southern China from 12 November to 19 December 2022. The average peak concentrations were $3.50 \pm 2.47 \times 10^6 \text{ cm}^{-3}$ for OH and $1.34 \pm 0.93 \times 10^8 \text{ cm}^{-3}$ for HO₂*. Model estimated interference to HO₂ by RO₂ possibly contributed to 44%-69% of the HO₂*. Calculations based on an observation-constrained chemical model revealed an overestimation of HO₂ and OH concentrations during warm periods of the field study. Sensitivity tests suggest that adding HO_x sinks or a HO₂ recycle process to the model could improve the model performance. Over-simulation of HO_x in the model resulted in overestimations of midday production rates by more than 79% for ozone and a factor of 1.88 for nitric acid.

Our study highlights the need for further improving understanding of the sources/sinks of OH and HO₂ and representation of them in air quality models.

1. Introduction

The HO_x family, comprising hydroxyl (OH) and peroxy radicals (HO₂), plays a pivotal role in the Earth's atmosphere by driving photochemical processes that influence the air composition and chemistry. OH radicals are primarily produced by the photolysis of ozone (O₃), nitrous acid (HONO), and ozonolysis of alkenes. They initiate the oxidation of CO and most volatile organic compounds (VOCs), producing HO₂ and other peroxy radicals (RO₂, where R represents an alkyl group). HO₂ is also generated from the photolysis of oxygenated VOCs (OVOCs) and by reactions between OVOCs and OH. In the presence of NO, RO₂ radicals are converted to HO₂ and then to OH radicals, buffering OH concentrations and maintaining atmospheric oxidation capacity. (Stone et al., 2012). These interactions are crucial in the formation of photochemical smog and secondary organic aerosol (SOA), which generate NO₂, O₃, and highly oxygenated molecules. HO_x radicals are removed through reactions of OH with inorganic trace gases, self-reactions among radicals, peroxyacetyl nitrate (PAN) formation, and the heterogeneous uptake by aerosols. These processes subsequently contribute to atmospheric acidification and aerosol formation through the production of H₂SO₄ and HNO₃. See Figure S1 and Table S1 for detailed processes and chemical reactions.

The accuracy of model-predicted OH is a crucial indicator for assessing our understanding of atmosphere processes (Heard and Pilling, 2003). There is a longstanding debate regarding the discrepancies between simulated and observed radical concentrations under low NO_x conditions, which remains a significant issue (Hofzumahaus et al., 2009; Stone et al., 2012; Zou et al., 2023). Previous studies have shown that models generally predict OH levels well in polluted conditions (NO > 1 ppb), but notable overestimations were observed under low NO and aged conditions, such as coastal areas (Kanaya et al., 2007; Zou et al., 2023), marine boundary layers (Berresheim et al., 2002; Carslaw et al., 1999), and the rural area (Bottorff et al., 2023;

1 Kanaya et al., 2012). Missing OH sinks from both measurement and chemical
2 mechanism were proposed as the primary reason for the overestimation (Lou et al.,
3 2010; Yang et al., 2016; Hansen et al., 2014; Thames et al., 2020). Underestimations of
4 OH concentrations were also observed in high biogenic VOCs (BVOCs) and low NO
5 (<1 ppb) conditions which generally happen in the subtropical or tropical area
6 (Hofzumahaus et al., 2009; Lelieveld et al., 2008; Tan et al., 2001; Whalley et al., 2011).
7 After considering a new OH regeneration mechanism (Wennberg et al., 2018; Novelli
8 et al., 2020) and measurement interference (Feiner et al., 2016; Hens et al., 2014; Mao
9 et al., 2012; Novelli et al., 2014; Woodward-Massey et al., 2020), daytime OH
10 concentrations could be reasonably reproduced by the model in the high BVOC
11 conditions, with some unresolved underestimation in the evening (Jeong et al., 2022;
12 Lew et al., 2020; Tan et al., 2019). Those results call for more measurement and
13 modelling in subtropical and tropical rural areas.

14 HO₂ concentrations were consistently underpredicted in the polluted urban sites
15 (Ma et al., 2019; Yang et al., 2021; Ma et al., 2022) , with no clear trends in relatively
16 clean regions. Some studies reported good agreement between measurement and model
17 prediction (Feiner et al., 2016; Lew et al., 2020), whereas others indicated model
18 overprediction (Bottorff et al., 2023; Griffith et al., 2013) and underprediction (Whalley
19 et al., 2010; Kim et al., 2013; Mallik et al., 2018; Tan et al., 2017;). These discrepancies
20 may be attributed to several factors, including measurement interference caused by RO₂
21 recycling in environments rich in BVOCs or aromatics (Fuchs et al., 2011),
22 uncertainties associated with heterogeneous uptake in box models (Yang et al., 2022),
23 and the outflow (or entrainment) of reservoir species like PAN (Griffith et al., 2013)
24 (Whalley et al., 2010) . Despite these advances, it remains difficult to pin down the
25 exact causes of the model-measurement discrepancies in some of the previous studies.

26 In the present study, we measured concentrations of OH and HO₂* using a
27 quadrupole chemical ionization mass spectrometer (PolyU-CIMS) from November to
28 December 2022 at a subtropical rural site in southern China. We tested model's
29 capability in reproducing the radical concentrations and elucidate factors contributing

1 to discrepancy under varying temperatures, VOCs, and NO_x conditions. The
2 Methodology section describes the measurement site, the principles and the
3 configuration of PolyU-CIMS, and the setup of chemical box models. The Results and
4 Discussion section details our findings, providing an analysis of the radical
5 concentrations and exploring the discrepancies between observed data and model
6 predictions. By employing an observation-constrained box model, we analyzed the
7 radical budgets and investigated potential causes for these discrepancies. The study
8 concludes with a discussion of the implications of these findings.

9 **2. Methodology**

10 **2.1 Measurement Site**

11 The field campaign was conducted at the Conghua Liangkou Air Monitoring
12 Station (23°44'47"N, 113°47'06"E, 200m, above sea level) from November 12 to
13 December 19, 2022 (Figure 1). The site is located at the northern part of the Pearl River
14 Delta (PRD), approximately 80 kilometers from the densely populated areas. It is
15 nestled within the Liuxi River National Forest Park (an evergreen broad-leaf forest).
16 The site is situated just north of the G105 national highway and around 0.5 kilometers
17 east of Liangkou town. Even though it is close to the road, the traffic was generally
18 limited during the observation period due to the coronavirus disease pandemic
19 (COVID-19). The site is subject to BVOCs emission, predominantly isoprene, from the
20 surrounding forest when the daytime temperatures exceeds 20°C, and NO emissions
21 from the nearby national highway, particularly during periods of low wind speeds. The
22 measurements included trace gases such as O₃, NO, NO₂, CO, HONO, VOCs, OVOCs,
23 meteorological data such as relative humidity (RH), temperature, and photolysis
24 frequencies of HONO, NO₂, O₃, H₂O₂, and HCHO. Details about the instruments are
25 shown in Table S2.



Figure 1 Geographic location of the measurement site (Liangkou Air Monitoring station 23°44'47"N, 113°47'06"E, 200 m a.s.l. marked by the red inverted triangle) in Conghua, Guangdong Province, South China. The map is sourced from © Google Earth and © Amap.

2.2 Radical measurement principle

OH and HO₂ radicals were measured using the Hong Kong Polytechnic University's quadrupole Chemical Ionization Mass Spectrometer (PolyU-CIMS), which had been used in a previous study for OH measurement (Zou et al., 2023). The use of CIMS for OH measurement was pioneered by Eisele and Tanner (1991), with subsequent enhancements in measurement accuracy (Eisele and Tanner, 1993; Tanner et al., 1997; Tanner and Eisele, 1995) and the adoption of inlets for simultaneous measurements of HO₂ and RO₂ (Edwards et al., 2003; Sjostedt et al., 2007), H₂SO₄ (Mauldin III et al., 2004), and OH reactivity (Muller et al., 2018).

Figure 2 illustrates the measurement principle of the PolyU-CIMS used in this campaign. Briefly, the ambient OH radicals are converted to H₂SO₄ in the sample inlet system by reacting with SO₂ (R21 in the reaction Table S1). These are then transformed to HSO₄⁻ ion clusters in the ionization chamber by the reactions with a reagent gas in

sheath flow (HNO_3 , R24 to 27), and ultimately dissociated (R29) for detection by the mass spectrometer system at m/z 97 ($S_{97\text{SO}_2}$ in Figure 2). To mitigate interference and noise, scavenger gases (C_3F_6 in this study) were introduced to scavenge the ambient OH, creating a background signal (R23, $S_{97\text{ScaSO}_2}$ in Figure 2). See Text S3 in the Supplementary Information for details about the scavenging efficiency. The ambient OH radicals signal (S_{OH}) is then determined by subtracting $S_{97\text{ScaSO}_2}$ from $S_{97\text{SO}_2}$. The OH concentration is calculated using the following equation:

$$[\text{OH}] = \frac{1}{C_{\text{OH}}} \times \frac{S_{\text{OH}}}{S_{62}} \quad (\text{E1})$$

Where C_{OH} represents the calibration factors of OH, and S_{62} is the signal corresponding to the reagent ion (NO_3^-). The detailed calibration procedure for OH is outlined in previous studies (Kürten et al., 2012; Zou et al., 2023) and also in Text S1. Ambient and injected NO can interfere with OH measurements through reaction of $\text{NO} + \text{HO}_2$. The interference by ambient NO is estimated using a model (within a range of $10^4 - 10^5 \text{ cm}^{-3}$) and was subtracted from the ambient OH measurements (see Text S4.2). The interference from the injected NO after HO_2 measurement is discussed in Text S4.2 and is mitigated through the duty cycle configuration described in Text S5.

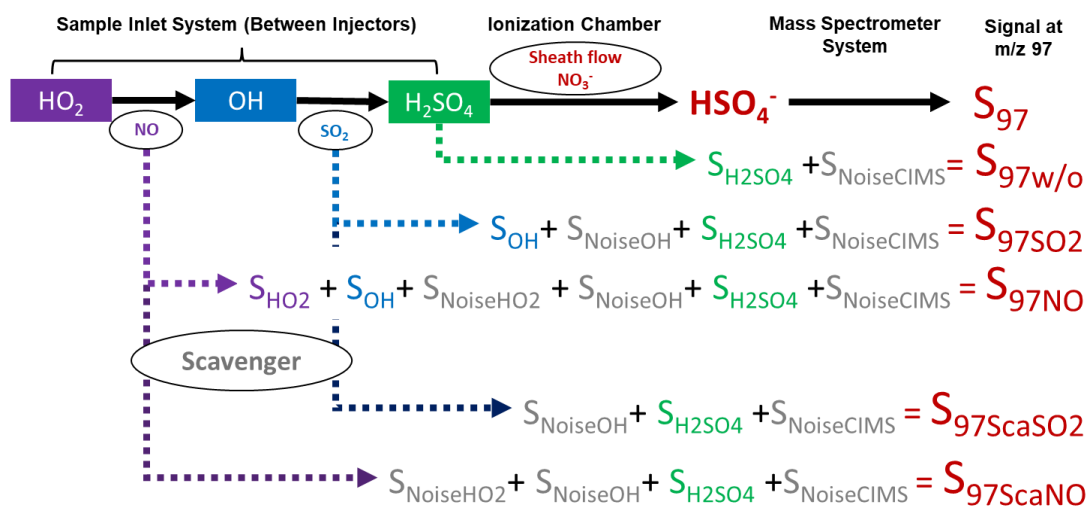


Figure 2 Flow chart depicts the relationship between measurement species and signal intensity at m/z 97 (S_{97}) with various gas injections. The color-filled grids label the ambient species, while oval shapes label the species injected into the sample flow. Signal intensities with different gas additions to the sample flow are represented by $S_{97\text{w/o}}$, $S_{97\text{SO}_2}$, $S_{97\text{NO}}$, $S_{97\text{ScaSO}_2}$, and $S_{97\text{ScaNO}}$. The signals correspond to ambient OH, HO_2 , H_2SO_4 and noise from OH measurement, HO_2 measurement and the CIMS denoted as S_{OH} , S_{HO_2} , $S_{\text{H}_2\text{SO}_4}$, S_{NoiseOH} , S_{NoiseHO_2} , and $S_{\text{NoiseCIMS}}$, respectively.

To measure ambient HO₂, NO is injected into the sample flow, converting HO₂ to OH (R11). This converted OH then follows the same reaction pathway (R21, R24 through R27, and R29) and is measured at m/z 97 (S_{97NO} in Figure 2). Similar to the OH measurement, the background signal for HO₂ (S_{97ScaNO} in Figure 2) is determined by introducing the scavenger gas. The corresponding signal for ambient HO₂ (S_{HO2}, as shown in Figure 2) is determined by subtracting S_{97ScaNO} and S_{OH} from S_{97NO}. The HO₂ concentration is calculated using an equation similar to E1, by replacing S_{OH}, and C_{OH} to S_{HO2} and C_{HO2}, respectively (E2).

$$[HO_2] = \frac{1}{C_{HO2}} \times \frac{S_{HO2}}{S_{62}} \quad (E2)$$

The procedure for determining the HO₂ calibration factor, C_{HO2}, is akin to that for C_{OH} (Text S1). The calibration tube generates equal amounts of radicals (R30 in SI, [OH]/[HO₂] = 1), allowing for simultaneous calibration of HO₂ and OH with and without NO addition to the sample flow.

In addition to HO₂, the added NO can also react with RO₂ and produce HO₂, which is a potential interference leading to an overestimation of ambient HO₂ levels (Edwards et al., 2003; Hanke et al., 2002; Fuchs et al., 2011; Whalley et al., 2013; Fuchs et al., 2014). For our CIMS configuration, the model estimated daytime interference increased HO₂ by 79% to 222% ([RO₂ interference]/[HO₂] × 100%) during the field study (Text S4.3 and Figure S9). Considering the uncertainties of RO₂ mechanisms in the MCM model, we opted not to use model results to correct RO₂ interference and denote our HO₂ measurement hereafter as HO₂* which is the maximum value of the ambient HO₂ concentrations.

Compared to the configuration of PolyU-CIMS in the previous campaign (Zou et al., 2023), the instrument has been upgraded for simultaneous HO₂* measurements. Refer to Figure S2 and Text S2 for modification to HO₂* measurement and to Text S5 for measurement duty cycle. Apart from these modifications, the settings and configurations of the PolyU-CIMS remained the same as those in the previous campaign (see Table S3). With the updated configuration, the PolyU-CIMS achieved the simultaneous measurement of OH, HO₂*, and H₂SO₄.

The calibration factors, detection limits and accuracies were $1.09 \times 10^{-8} \text{ cm}^{-3}$, $3 \times 10^5 \text{ cm}^{-3}$, and 44% for OH; $6.01 \times 10^{-9} \text{ cm}^{-3}$, $2 \times 10^6 \text{ cm}^{-3}$, and 46% for HO₂*; and $1.09 \times 10^{-8} \text{ cm}^{-3}$, $1 \times 10^5 \text{ cm}^{-3}$, and 40% for H₂SO₄, respectively (Table S3).

2.3 Box Model

HO_x concentrations in this study were simulated using the Framework for 0-D Atmospheric Modelling (F0AM, Wolfe et al., 2016) with the Master Chemical Mechanism (MCM) v3.3.1 (<http://mcm.leeds.ac.uk/MCM>), which encompasses over 6,700 species and 17,000 reactions. MCM v3.3.1 features a near-explicit chemical mechanism that includes isoprene degradation and OH regeneration mechanisms. This mechanism has previously been employed to investigate HO_x chemistry and conducting budget analyses (Slater et al., 2020; Tan et al., 2018; Zou et al., 2023). The gas-phase chlorine chemistry described by Xu et al. (2015) and Wang et al. (2019) was included in the model (Chen et al., 2022).

In the baseline scenario, the observation data were aggregated into one-hour intervals to provide input for the model, initially constraining it without incorporating observed OH and HO₂* data. For the assessment of ozone formation rates, the model was adjusted to include constraints based on the actual measured concentrations of OH and HO₂*. Observed VOCs were categorized into anthropogenic origin (AVOCs), including species from petroleum gas and industrial solvent evaporation (alkenes, alkenes, benzene, and TEXs - toluene, ethylbenzene, and xylenes), and OVOCs comprising aldehydes, ketones, and acids. The sole BVOC measured in this study was isoprene. Methacrolein (MACR), a derivative of isoprene, is distinctly classified among the biogenically sourced OVOCs for further discussion. Physical processes like deposition and entrainment in the model were represented by a first-order physical loss with a 24-hour lifetime for all species (Chen et al., 2022; Wolfe et al., 2016; Zou et al., 2023). The model also included the heterogeneous uptake of HO₂ by aerosols, represented as a pseudo-first order loss (Jacob, 2000):

$$\frac{d[HO_2]}{dt} = -k_{HO_2}[HO_2] \quad (E3)$$

$$k_{HO_2} = \frac{V_{HO_2} \times S_a \times \gamma_{HO_2}}{4} \quad (E4)$$

$$v_{HO_2} = \sqrt{\frac{8RT}{\pi \times M_{HO_2}}} \quad (E5)$$

Here, k_{HO_2} represents the first-order loss rate coefficient of HO_2 by aerosol uptake, determined by the effective HO_2 uptake coefficient γ_{HO_2} (0.1, Guo et al., 2019), the mean molecular velocity of HO_2 (v_{HO_2}), the aerosol surface area concentration (S_a) measured by the Scanning Mobility Particle Sizing (SMPS); and the molecular mass of HO_2 ($M_{HO_2} = 17$ g/mol). As aerosol and aqueous phase chemistry were not included in the model, it was assumed that the heterogeneous HO_2 loss would not lead to further reactions (Guo et al., 2019). For each day, a three-day spin-up was performed with constant inputs to establish stable model chemistry and reduce the uncertainty of unconstrained species. Refer to Text S4 for details on model setup for interferences assessment.

3. Results and Discussion

3.1 Results from Observations

3.1.1 Overview

Figure 3 illustrates a time series showing observed concentrations of radical and trace gases, along with meteorological parameters, from 12 November to 19 December 2022. In November, the conditions were characterised by warm temperatures ranging from 29°C to 19°C and high relative humidity averaging 86%. In contrast, December witnessed a significant decrease in temperature (ranging from 20°C to 9°C) and a reduction in relative humidity (averaging 72%). Wind speeds during the campaign were generally low, averaging 0.9 ± 0.6 m/s and typically remaining below 3.0 m/s, with higher speeds occurring towards the end of December. In November, daytime winds predominantly blew from the south, while nighttime winds came from the north. In December, northerly winds predominated both day and night. Detailed hourly wind speed and direction data are illustrated in Figure 3, and wind roses are shown in Figure S3. On days with low wind speeds (less than 0.5 m/s), NO_x emissions from the G105 national highway significantly influenced chemical measurements at the monitoring site, causing morning NO levels to peak at several parts per billion (ppb). Isoprene

1 concentrations peaked in the afternoons, ranging from 0.2 to 1.7 ppb in November and
2 dropping to less than 0.1 ppb in December. Other trace gases and particulate matter
3 levels were higher in November than in December.

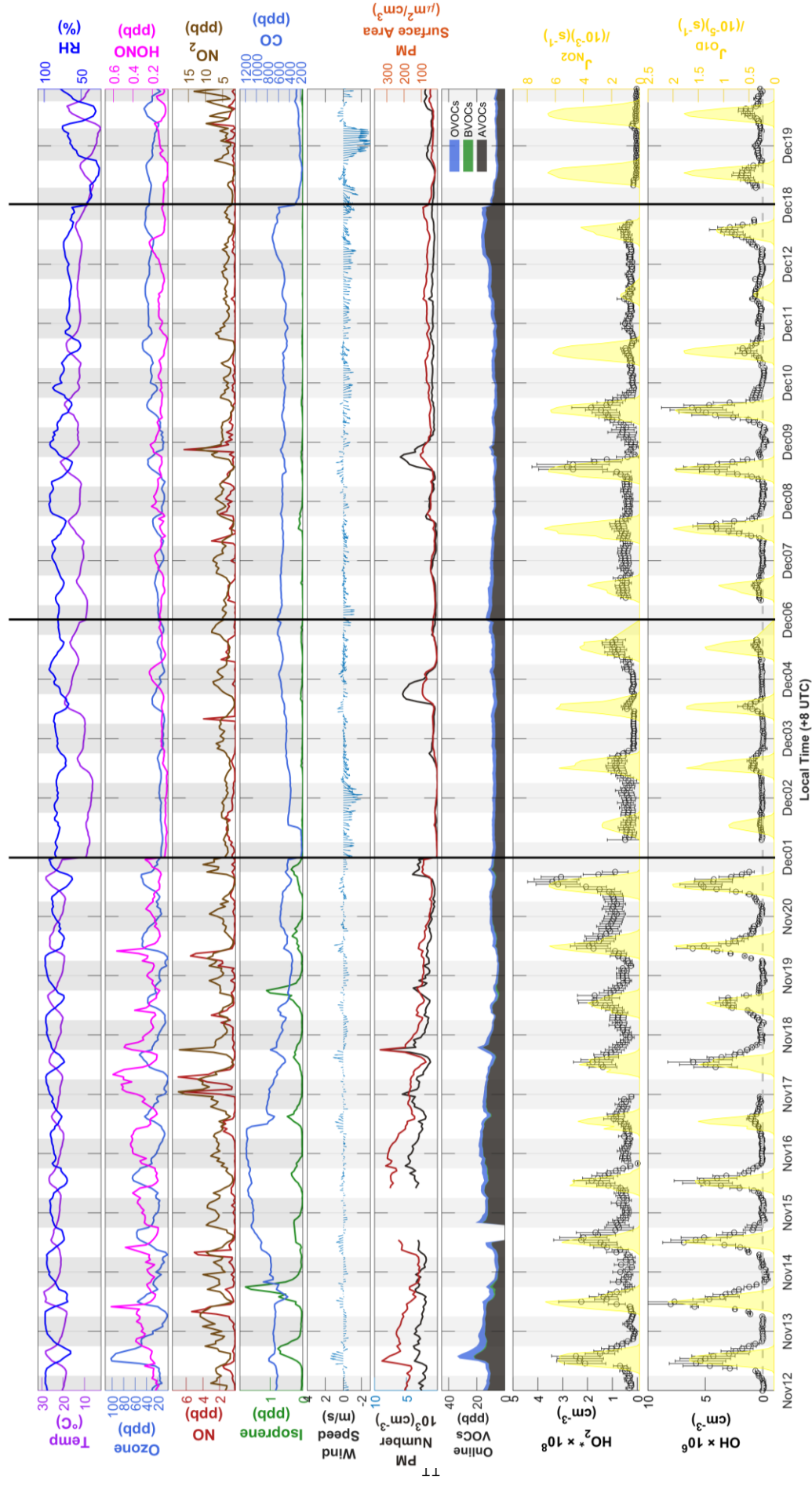


Figure 3 Time series of HO_2^* and OH radicals from 12 November to 19 December, showing measured weather conditions (temperature, RH, wind speed, and wind direction), primary sources of HO_x radicals (ozone, HONO), important sinks of the radicals (CO, isoprene, and VOCs), and photolysis frequencies of NO_2 (J_{NO_2}) and ozone (J_{O_3}). Non-continuous days during the campaign are delineated by a black line. The PRD, CEC, and CNC periods for further analysis were labelled in red, orange, and blue, respectively. The x-axis is in local time (+8 UTC).

Throughout the campaign, the daytime concentrations of OH and HO_2^* consistently exceeded detection limits and showed distinct diurnal patterns. The OH concentrations typically peaked around midday, while the HO_2^* levels reached their maximum approximately one to two hours later (Figure S4). The daily maximum concentration of OH varied significantly, ranging from $8.00 \times 10^6 \text{ cm}^{-3}$ to nearly the detection limit of $2.54 \times 10^5 \text{ cm}^{-3}$, with an average of $3.50 \pm 2.47 \times 10^6 \text{ cm}^{-3}$ (Table 1). Similarly, the daily maximum concentration of HO_2^* varied from $3.42 \times 10^8 \text{ cm}^{-3}$ to $2.17 \times 10^7 \text{ cm}^{-3}$, averaging $1.34 \pm 0.93 \times 10^8 \text{ cm}^{-3}$ (Table 1). At nighttime, while the HO_2^* levels generally remained above the detection threshold, the OH concentrations frequently approached the threshold. The average nighttime concentrations were $3.92 \times 10^7 \text{ cm}^{-3}$ for HO_2^* and $1.64 \times 10^5 \text{ cm}^{-3}$ for OH. We compared the observed OH and HO_2^* concentrations with those reported in previous studies conducted in urban, suburban, rural forest, and coastal sites. As illustrated in Figure S5, the OH concentrations were generally lower than those found in urban settings but similar to levels observed in suburban, rural, and forest environments. This suggests a moderate level of anthropogenic activity typical of mixed rural settings. In contrast, the HO_2^* concentrations during these periods were significantly lower than earlier observations in rural and forest environments, likely owing to reduced photochemical activity during our measurement period.

Table 1 Average concentrations and standard deviation of measured species throughout the entire campaign (Total) and the selected 3 days cases from each cluster (PRD, CEC and CNC).

Species (Unit)	Total	PRD	CEC	CNC
AveMax OH _{Obs} 10 ⁶ (cm ⁻³)	3.5±2.5	6.9±1.1	4.9±1.5	5.3±0.9
OH _{Obs} 10 ⁶ (cm ⁻³)	0.9±1.5	1.6±2.2	1.4±1.6	1.2±1.8
OH _{DL} 10 ⁶ (cm ⁻³)	0.5±0.3	0.4±0.3	0.4±0.2	0.9±0.6
AveMax HO ₂ _{Obs} 10 ⁸ (cm ⁻³)	1.34±0.93	2.32±1.25	2.36±0.92	1.82±1.02
HO ₂ * _{Obs} 10 ⁸ (cm ⁻³)	0.59±0.51	0.76±0.63	1.10±0.68	0.67±0.55
HO ₂ * _{DL} 10 ⁸ (cm ⁻³)	0.19±0.11	0.17±0.10	0.25±0.08	0.26±0.15
Pressure (hpa)	995±4	992±1	992±1	995±2
Temp (°C)	16±6.1	23±3.0	23±2.6	14±2.8
RH (%)	78±15	87±11	86±10	81±9.4
Wind Speed (m/s)	0.91±0.65	0.53±0.32	0.57±0.34	0.87±0.5
j _{OH} 10 ⁻⁶ (s ⁻¹)	3.2±5.4	3.5±6	3.6±5.9	4.0±6.6
j _{NO₂} 10 ⁻³ (s ⁻¹)	1.3±1.9	1.3±2.1	1.4±2	1.6±2.3
HONO	0.169±0.104	0.249±0.084	0.201±0.070	0.133±0.033
SO ₂	0.5±0.8	0.5±0.6	0.4±0.5	0.4±0.5
NO ₂	4.89±2.37	6.25±2.47	4.84±2.23	4.52±1.97
NO	0.57±0.86	0.73±1.09	0.69±1.00	0.73±0.85
CO	557.36±225.92	739.41±153.84	464.73±74.34	513.36±22.02
Ozone	25±14	32±23	24±13	19±9.4
Particle Surface Area (um ² /cm ³)	86±72	186±51	84±28	48±19
Isoprene	0.082±0.174	0.257±0.337	0.155±0.200	0.029±0.030
*OVOCs	2.218±1.056	3.163±2.324	1.755±0.379	1.730±0.330
*AVOCs	8.346±3.223	9.662±5.031	6.886±1.755	6.801±0.864
*TEXs	0.356±0.316	0.801±0.616	0.266±0.090	0.237±0.085

Notes: Concentrations are expressed in parts per billion (ppb) unless otherwise specified. Total VOCs concentrations are categorized by origin (AVOCs and OVOCs). For the average concentration of each measured VOC, refer to Table S5.

Figure 4 illustrates the results of the 24-hour backward trajectory analysis, revealing three distinct but sequentially occurring phases. In the initial phase (Figure 4a), air masses originated from the urban and industrial zones of the Pearl River Delta (PRD). This phase was characterised by intense photochemical activity, with ambient temperatures exceeding 20°C and relative humidity levels surpassing 70%. During this period, notably high concentrations of VOCs, ozone, and radicals were observed, with the average daily maximum concentrations of OH and HO₂* radicals reaching $6.50 \pm 1.19 \times 10^6 \text{ cm}^{-3}$ and $2.20 \pm 0.27 \times 10^8 \text{ cm}^{-3}$, respectively. The subsequent phase was characterised by air masses originating from Central East China (CEC, Figure 4b). This phase showed reduced photochemical reactivity and lower concentrations of the measured trace gases. The average daily maximum concentrations of OH and HO₂* during this phase were $4.35 \pm 2.19 \times 10^6 \text{ cm}^{-3}$ and $1.96 \pm 0.90 \times 10^8 \text{ cm}^{-3}$, respectively.

The final phase was influenced by air masses from Central North China (CNC, Figure 4c), which exhibited the lowest concentrations of trace gases and the least pronounced average daily maximum concentrations in OH and HO₂* concentrations, measured at $2.23 \pm 1.95 \times 10^6 \text{ cm}^{-3}$ and $7.63 \pm 7.66 \times 10^7 \text{ cm}^{-3}$, respectively. This phase coincided with an increase in cloudy days and a decrease in temperatures, indicating reduced photochemical conditions.

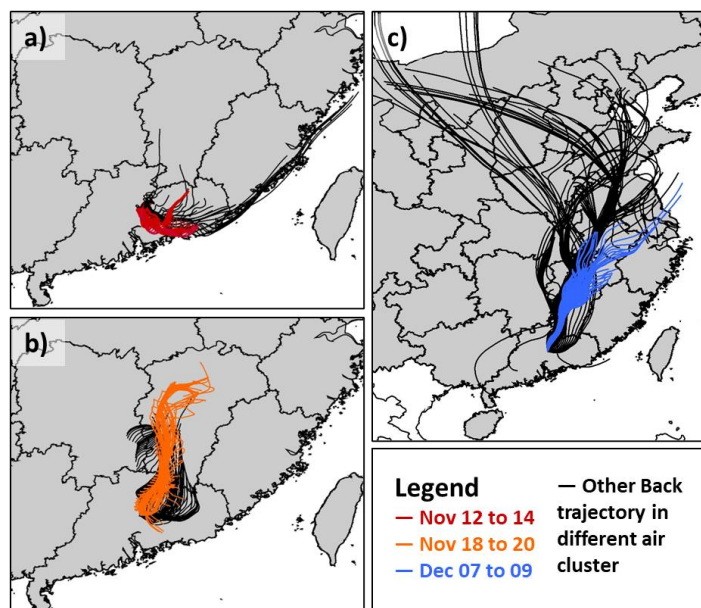


Figure 4 24-hour back trajectories for (a) Pearl River Delta (5 days), (b) Central East China (4 days), and (c) Central North China (14 days) cases. Three days selected from each cluster for model simulation are distinguished by different colours.

3.1.2 Selection of Cases

For each phase, a representative three-day period was selected for detailed analysis based on the availability of comprehensive data and sunny conditions (colored trajectories in Figure 4). In the subsequent analysis, 'PRD,' 'CEC,' and 'CNC' refer to the selected periods corresponding to the air masses originating from these regions. The average daily maximum concentrations of OH and HO₂* radicals for these periods are presented in Table 1. The average daily max OH concentrations were $6.89 \pm 1.10 \times 10^6 \text{ cm}^{-3}$ in PRD, $4.90 \pm 1.47 \times 10^6 \text{ cm}^{-3}$ in CEC, and $5.27 \pm 0.89 \times 10^6 \text{ cm}^{-3}$ in CNC, with a pronounced decrease from PRD to CEC (of $1.99 \times 10^6 \text{ cm}^{-3}$). The average daily max HO₂* concentrations were $2.32 \pm 1.25 \times 10^8 \text{ cm}^{-3}$ in PRD, $2.36 \pm 0.92 \times 10^8 \text{ cm}^{-3}$ in CEC, and $1.82 \pm 1.02 \times 10^8 \text{ cm}^{-3}$ in CNC, with a slight increase of $0.04 \times 10^8 \text{ cm}^{-3}$ from

PRD to CEC and a more substantial decrease of $0.54 \times 10^8 \text{ cm}^{-3}$ from CEC to CNC. These trends suggest a declining atmospheric oxidation capacity from PRD to CNC.

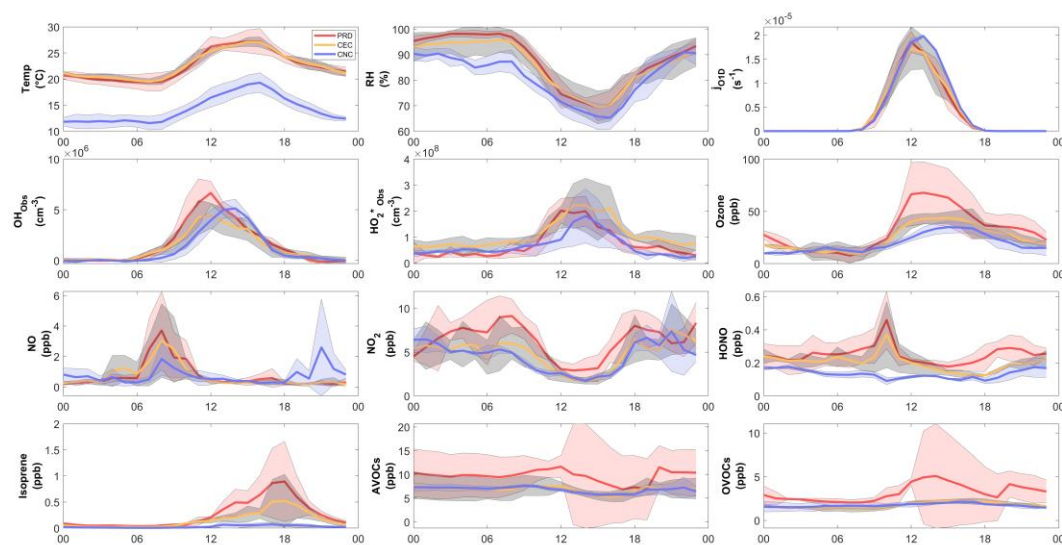


Figure 5 Average diurnal variations of (a) Temperature (b) Relative Humidity (c) J_{O1D} (d) OH (e) HO_2^{*} (f) Ozone (g) NO (h) NO₂ (i) HONO (j) Isoprene (k) AVOCs (l) OVOCs. The solid-colored lines represent selected cases: orange for PRD, green for CEC, and blue for CNC. The light band represents the standard deviations of the mean. The increase in the standard deviations of VOCs and OVOCs during the PRD case is a result of absence of data on the afternoon of November 14th and large variations in on November 12th and 13th.

The precursor concentrations and meteorological parameters also varied across cases in terms of statistics (Tables 1 and S4) and diurnal variations (Figure 5). In the PRD case, the average concentrations are characteristic of a rural environment, with AVOCs at 9.70 ± 5.00 ppb, OVOCs at 3.20 ± 2.30 ppb, isoprene at 0.26 ± 0.34 ppb, NO₂ at 6.3 ± 2.5 ppb, and NO at 0.73 ± 1.09 ppb. The NO concentration was affected by traffic sources, as there were no other fresh emission source nearby, and the diurnal variation of NO showed a morning peak in all three cases (Figure 5). In the CEC case, a general reduction in anthropogenic influence is evident. AVOCs, OVOCs isoprene and NO₂ drop significantly to 6.90 ± 1.80 ppb, 1.70 ± 0.38 ppb, 0.16 ± 0.20 ppb, and 4.84 ± 2.23 ppb respectively. Meanwhile, NO levels remained close to PRD levels at 0.69 ± 1.00 ppb. In the CNC case, the air mass was more aged with reduced biogenic emissions, reflected in further decreases in isoprene and NO₂ to 0.03 ± 0.04 ppb and 4.52 ± 1.97 ppb, respectively, due to colder weather conditions. The temperature

decreased significantly from PRD to CNC, whereas the average peak photolysis frequency was comparable between PRD and CNC, as shown in Table 1.

3.2 Chemical budgets of OH and HO₂

To investigate the OH and HO₂ chemical budgets during the three distinct periods, we employed a box model constrained by observed concentrations of NO_x, VOCs, and relevant meteorological parameters in the selected cases (base scenario in which OH and HO₂ concentrations were not constrained by the measurements). The resulting OH and HO₂ budgets, displaying typical bell-shaped patterns, are illustrated in Figure 6. During midday (10:00–15:00), the primary source of HO₂ was the recycling of RO species, with rates of 3.22 ppb h⁻¹ for PRD, 2.09 ppb h⁻¹ for CEC, and 1.08 ppb h⁻¹ for CNC. Additionally, HCHO photolysis contributed 0.75 ppb h⁻¹, 0.46 ppb h⁻¹, and 0.26 ppb h⁻¹ for PRD, CEC, and CNC, respectively. The sinks of HO₂ varied among the cases with minor contribution from the uptake process, driven by radical termination mechanisms. The rate of radical self-reactions decreased from PRD to CNC. In contrast, NO_x-radical reactions between CEC and CNC were comparable, with respective rates of 0.39 ppb h⁻¹ and 0.33 ppb h⁻¹, indicating a shift in radical termination mechanisms.

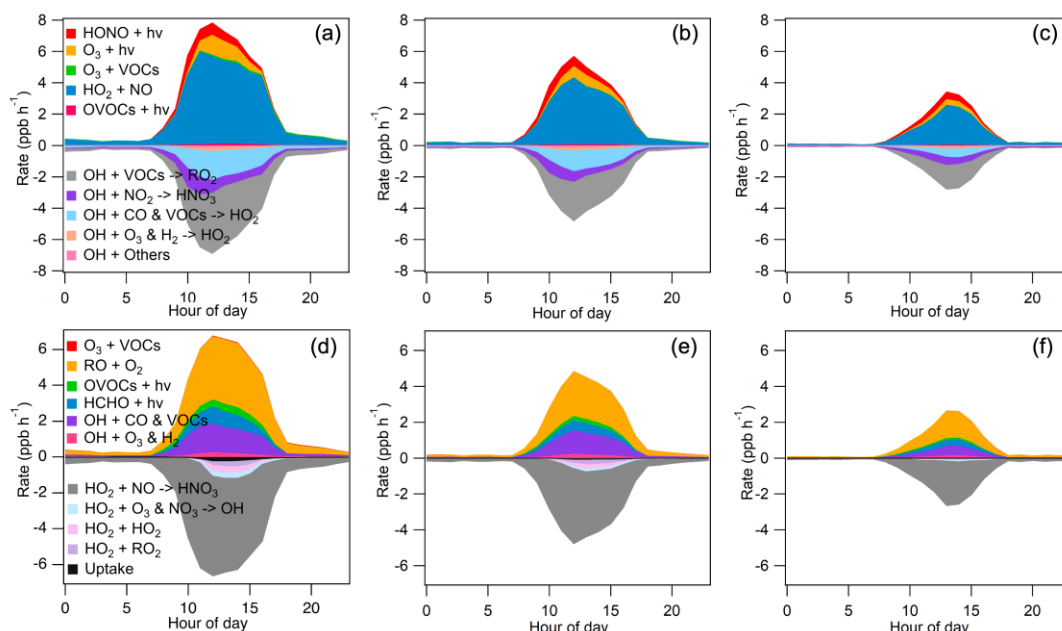


Figure 6 Chemical budgets of OH and HO₂ for PRD (a, d), CEC (b, e), and CNC (c, f) simulated using a chemical box model.

OH formation was predominantly driven by the $\text{HO}_2 + \text{NO}$ reaction, contributing 5.18 ppb h^{-1} , 3.51 ppb h^{-1} , and 1.81 ppb h^{-1} (average for 10:00 – 15:00, referred to hereafter in this section) for PRD, CEC, and CNC, respectively. Additionally, contributions from ozone photolysis and HONO increased from PRD to CEC and then to CNC, with rates of 21.4%, 22.7%, and 24.6%, respectively. The primary sinks for OH included reactions with VOCs to produce RO_2 , with rates of 3.31 ppb h^{-1} , 2.02 ppb h^{-1} , and 1.13 ppb h^{-1} , and reactions with CO and other VOCs to generate HO_2 , contributing 1.55 ppb h^{-1} , 1.06 ppb h^{-1} , and 0.38 ppb h^{-1} for PRD, CEC, and CNC, respectively. These findings highlight the critical role of OH + VOC reactions in the chemical budget of OH.

3.3 Comparison of model with observation results

Modeled OH and HO_2 concentrations were evaluated by comparing them against observations. The observed HO_2^* serves as an upper limit for ambient HO_2 due to the interference from RO_2 . The model calculated average daytime (08:00-16:00) RO_2 interference increased HO_2 by 127%, 117%, and 144% for PRD, CEC and CNC case, respectively. In the PRD case (Figure 7), which represents the most polluted and warmest environment among the three cases, the model slightly overestimated OH concentrations. However, the modeled HO_2 concentration was substantially higher than the observed HO_2^* , indicating an overprediction of HO_2 . Similar overestimations have been reported at a rural site (Kanaya et al., 2012). For the CEC case (Figure 8), the modeled OH and HO_2 concentrations were moderately higher than the observed values during the daytime, consistent with findings from a rural forest site (Bottorff et al., 2023). In the CNC case (Figure 9), the model generally reproduced the observed OH trend, with HO_2 comparable to the observed HO_2^* . However, on December 7, it exhibited daytime overestimation of HO_x , similar to the PRD case. At night, both CNC and CEC showed lower modeled HO_2 concentrations compared to observed HO_2^* , which can be explained by the expected contribution of RO_2 interference to HO_2^* in the observations. To further investigate the possible reasons for the model-observation

discrepancies, sensitivity tests were conducted. The modeled RO₂ interference ratio was used to correct HO₂* values, roughly estimating ambient HO₂ as input for the tests.

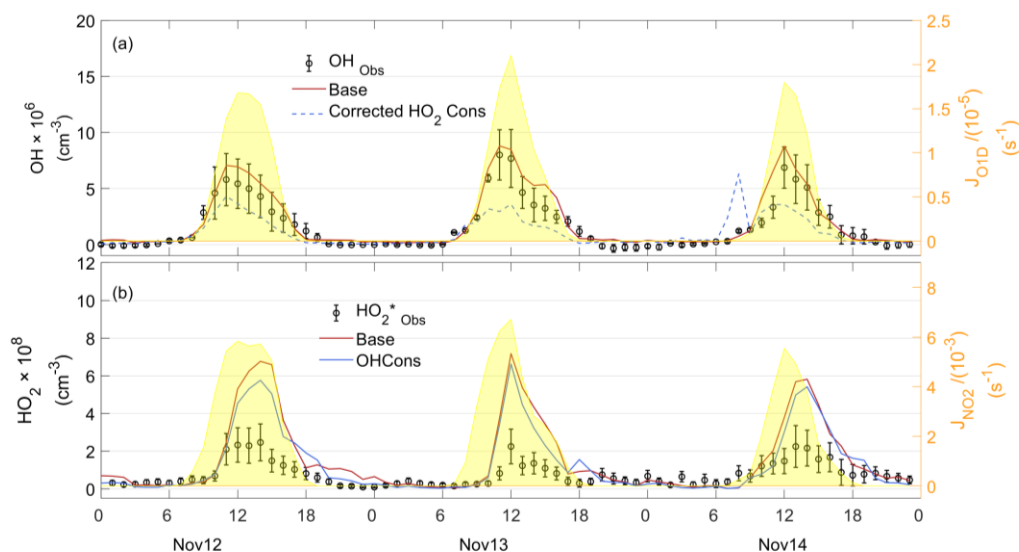


Figure 7 observed and simulated OH, observed HO₂*, and simulated HO₂ time series for the PRD case. The “Obs” subscript denotes the observation data. “Base” denotes the result of Baseline scenario as described in Box Model section. “Cons” denotes the results with additional constrained species compared to Base.

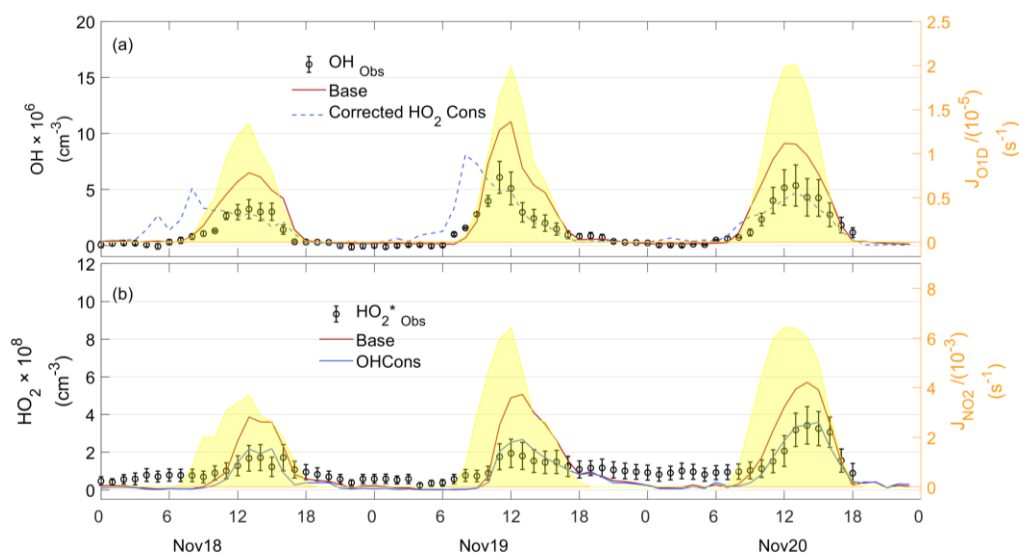


Figure 8 observed and simulated OH, observed HO₂*, and simulated HO₂ time series for the PRD case. The “Obs” subscript denotes the observation data. “Base” denotes the result of Baseline scenario as described in Box Model section. “Cons” denotes the results with additional constrained species compared to Base.

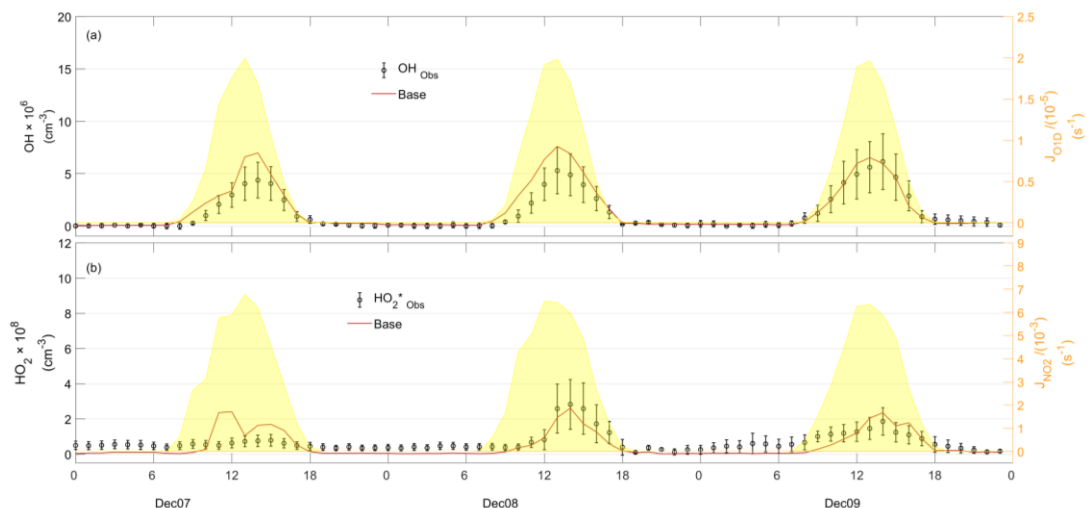


Figure 9 observed and simulated time series of OH and HO₂* for the CNC case. “Base” denotes the result of Baseline scenario as described in Box Model section.

3.3.1 Substantial overestimation of HO₂ in PRD case

To explain the HO₂ overprediction by the base model, we constrained OH or corrected HO₂* and compared these to the base case (without constraining OH and HO₂). Result shows that constraining corrected HO₂* causes the model to underestimate OH (blue line in Figure 7a), while constraining OH still leads substantially higher modeled HO₂ concentration than base model (blue line in Figure 7b). This result suggests that to align the modeled OH and HO₂ with observations, it may be necessary to introduce a strong, unknown process for HO₂ that efficiently recycles OH with a high yield (Kanaya et al., 2012).

3.3.2 Moderate overestimation of both OH and HO₂ radicals in CEC case

Unlike the PRD case, constraining either OH or corrected HO₂* in the CEC case generally reduces the daytime overestimation of both HO₂ and OH. These results indicate an additional sink for both OH and HO₂, as suggested by Bottorff et al. (2023). However, the OH concentration is overestimated in the morning when the corrected HO₂* was constrained, suggesting that there may be missing OH reactivity during this time. To further investigate the underlying causes, we examined the correlations between various pollutants. The significant negative correlation between CO and NO ($R^2=0.49$, $p=0.01$, Figure S6b) suggests that morning conditions in the CEC case may have been influenced by emission from fresh complete combustion during the CEC

case, while correlations in the PRD and CNC cases are not significant (Figures S6a and c). This indicates that the missing OH reactivity in the CEC case during the morning hours is possibly related to fresh vehicle emissions.

4. Implication for model overestimation of HO_x

OH and HO₂ are key species that determine the atmosphere's oxidative processes. Inaccurate modelling of their sinks can lead to significant overestimations of oxidation capacity, which may skew assessments of the impact of HO_x on air pollution and climate change. This problem is particularly pronounced in the case of ozone, a widespread photochemical pollutant. To demonstrate this issue, we compared results from two modelling scenarios: the first scenario was base case (as described in section 3.2), while the second scenario included constraints from all observational parameters, including measurements of OH and HO₂^{*}. Results for the scenario constrained by OH and corrected HO₂^{*} measurements are also showed in Figures 10 and 11.

As illustrated in Figure 10, not constraining free-radical measurement data in the chemical model (the orange lines) led to overestimates of ozone's photochemical production rates. Compared to scenarios constrained by OH and HO₂^{*}, in the PRD case, simulated midday O_x (O₃ + NO₂) formation rates in the base case were overestimated by 33% on average and 32% at peak O_x rates. In the CEC case, the overestimation was 79% on average and 89% at peak O_x rates, while in the CNC case, the smallest overestimation was 17% on average and 25% at peak. Since the ambient HO₂ concentration is lower than the constrained HO₂^{*} concentration, the impact of the overestimated HO_x on ozone production rates might be even larger in such environments (the green lines in Figure 10).

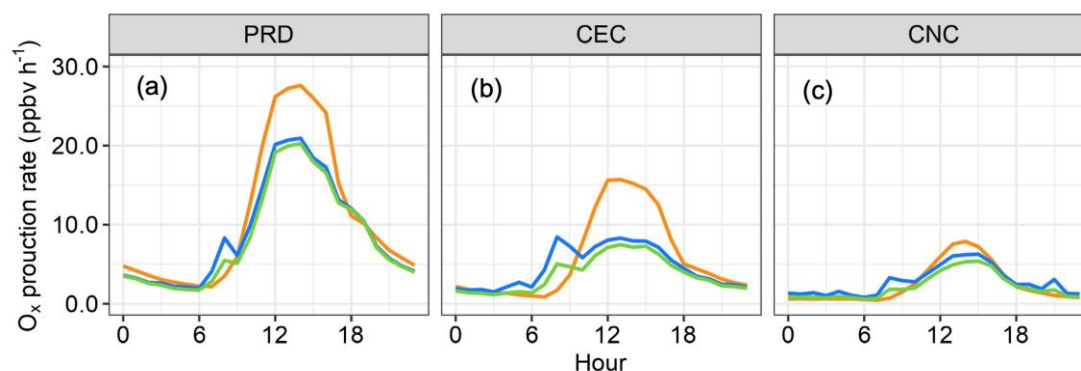


Figure 10 O_X (O_3+NO_2) photochemical production rates in three comparative cases: (a) PRD, (b) CEC, and (c) CNC. The orange, green, and blue lines represent rates modelled by excluding OH and HO_2^* , including OH with corrected HO_2^* , and including OH with uncorrected HO_2^* , respectively.

The overestimation of HO_X also significantly affected the simulated concentration of nitric acid (HNO_3), which is crucial for new particle formation and growth (Wang et al., 2020). Figure 11 illustrates that the chemical model moderately overestimated nitric acid production rates without the constraints of free-radical measurements (the orange lines). Compared to the scenarios constrained by OH and HO_2^* , the midday production rates of nitric acid were overestimated by 25%, 88%, and 31% in the PRD, CEC, and CNC cases, respectively. Such overestimations can considerably impact assessments of new particle formation and growth processes and their impact on air pollution and climate change. On the other hand, the impact on HNO_3 production is dominated by OH radical, therefore, measurement interference of HO_2 is negligible.

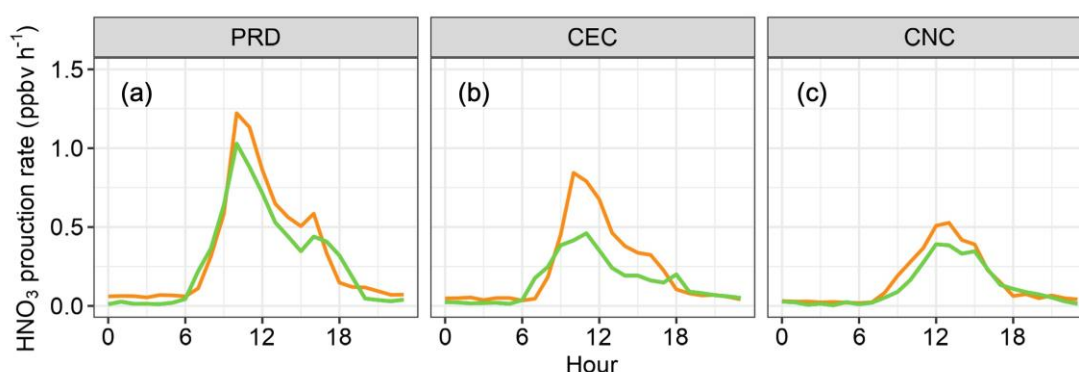


Figure 11 Modelled HNO_3 production rates in three comparative cases: (a) PRD, (b) CEC, and (c) CNC. The orange, green, and blue lines represent rates modelled by excluding OH and HO_2^* , including OH with corrected HO_2^* , and including OH with uncorrected HO_2^* , respectively. [The blue lines are covered by the green lines as the daytime HNO_3 production rate are determined by the constrained OH].

5. Conclusion

HO_2^* and OH concentrations were measured using a chemical ionization mass spectrometer at a subtropical rural site in southern China from 12 November to 19 December 2022. The measurements indicated generally lower concentrations of OH and HO_2^* than those observed in previous studies at various sites. The model estimated that interferences from RO_2 contributed to 44%-69% of the HO_2^* throughout the

1 **campaign.** Backward trajectory analysis revealed three distinct phases characterised by
2 sequentially decreasing pollution levels and temperatures. During the cold, clean period,
3 model simulations closely matched the observed OH and **HO₂*** concentrations.
4 However, during the warm, polluted period, the models overestimated HO₂ or both
5 radicals. Model sensitivity analysis indicates that adding an OH cycling reaction from
6 HO₂ or additional sinks of OH and HO₂ would largely reduce the model-observation
7 discrepancy in different cases of this study. However, the exact chemical reactions
8 remain to be identified. Our results are in line with previous studies indicating
9 substantial gaps in our understanding of the sources and sinks of OH and HO₂ in certain
10 environments. Our study provides additional evidence for the current incomplete
11 understanding of the HO_x sources and sinks and calls for more research to resolve the
12 model–observation mismatch found in this work and in previous studies.

13 The over-prediction of HO_x resulted in significant overestimation of the production
14 rates of other secondary pollutants such as ozone and nitric acid at the site. It is critical
15 to evaluate the capability of OH and HO₂ simulations in major chemical transport
16 models and earth system models, as inaccurate simulations of OH and HO₂ may
17 misguide the development of air pollution and global warming control strategies.

18 **Data availability.** All of the data used to produce this paper can be obtained by
19 contacting Tao Wang (tao.wang@polyu.edu.hk).

20 **Supplement.** The online supplement for this article is available at:

21 **Author contributions.** TW conceived the HO_x research. TW, XW and YZ planned and
22 organized the overall field campaign at Conghua. ZZ conducted the OH measurements
23 using CIMS, with contributions from TW and WS, QC, and SH. XF, ZR, XL, AG, QW,
24 CP, and XT performed the JNO₂ VOCs and OVOCs measurements. ZZ performed the
25 chemical box modelling with contributions from TC and QC. TC, ZZ, and TW analysed
26 the data and interpreted the result (ZZ analysed the time series and diurnal variations of
27 observation data; TC interpreted the results of box model, investigated the missing
28 sources; TW supervised and guided these processes). TC, ZZ, and TW wrote the paper.
29 All of the authors reviewed and commented on the paper.

Competing interests. One author (Tao Wang) is a member of the editorial board of Atmospheric Chemistry and Physics. The authors have no other competing interests to declare.

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Reference

Berresheim, H., Elste, T., Tremmel, H. G., Allen, A. G., Hansson, H.-C., Rosman, K., Dal Maso, M., Mäkelä, J. M., Kulmala, M., and O'Dowd, C. D.: Gas-aerosol relationships of H₂SO₄, MSA, and OH: Observations in the coastal marine boundary layer at Mace Head, Ireland, *Journal of Geophysical Research: Atmospheres*, 107, PAR 5-1-PAR 5-12, <https://doi.org/10.1029/2000JD000229>, 2002.

Bottorff, B., Lew, M. M., Woo, Y., Rickly, P., Rollings, M. D., Deming, B., Anderson, D. C., Wood, E., Alwe, H. D., Millet, D. B., Weinheimer, A., Tyndall, G., Ortega, J., Dusanter, S., Leonardis, T., Flynn, J., Erickson, M., Alvarez, S., Rivera-Rios, J. C., Shutter, J. D., Keutsch, F., Helmig, D., Wang, W., Allen, H. M., Slade, J. H., Shepson, P. B., Bertman, S., and Stevens, P. S.: OH, HO₂, and RO₂ radical chemistry in a rural forest environment: measurements, model comparisons, and evidence of a missing radical sink, *Atmospheric Chemistry and Physics*, 23, 10287–10311, <https://doi.org/10.5194/acp-23-10287-2023>, 2023.

Carslaw, N., Creasey, D. J., Heard, D. E., Lewis, A. C., McQuaid, J. B., Pilling, M. J., Monks, P. S., Bandy, B. J., and Penkett, S. A.: Modeling OH, HO₂, and RO₂ radicals in the marine boundary layer: 1. Model construction and comparison with field measurements, *J. Geophys. Res.*, 104, 30241–30255, <https://doi.org/10.1029/1999JD900783>, 1999.

Chen, Q., Xia, M., Peng, X., Yu, C., Sun, P., Li, Y., Liu, Y., Xu, Z., Xu, Z., Wu, R., Nie, W., Ding, A., Zhao, Y., and Wang, T.: Large Daytime Molecular Chlorine Missing Source at a Suburban Site

1 in East China, *JGR Atmospheres*, 127, <https://doi.org/10.1029/2021JD035796>, 2022.

2 Edwards, G. D., Cantrell, C. A., Stephens, S., Hill, B., Goyea, O., Shetter, R. E., Mauldin, R. L.,
3 Kosciuch, E., Tanner, D. J., and Eisele, F. L.: Chemical Ionization Mass Spectrometer Instrument
4 for the Measurement of Tropospheric HO₂ and RO₂, *Anal. Chem.*, 75, 5317–5327,
5 <https://doi.org/10.1021/ac034402b>, 2003.

6 Eisele, F. L. and Tanner, D. J.: Ion-assisted tropospheric OH measurements, *J. Geophys. Res.*,
7 96, 9295, <https://doi.org/10.1029/91JD00198>, 1991.

8 Eisele, F. L. and Tanner, D. J.: Measurement of the gas phase concentration of H₂SO₄ and
9 methane sulfonic acid and estimates of H₂SO₄ production and loss in the atmosphere, *Journal*
10 *of Geophysical Research: Atmospheres*, 98, 9001–9010, <https://doi.org/10.1029/93JD00031>,
11 1993.

12 Feiner, P. A., Brune, W. H., Miller, D. O., Zhang, L., Cohen, R. C., Romer, P. S., Goldstein, A. H.,
13 Keutsch, F. N., Skog, K. M., Wennberg, P. O., Nguyen, T. B., Teng, A. P., DeGouw, J., Koss, A.,
14 Wild, R. J., Brown, S. S., Guenther, A., Edgerton, E., Baumann, K., and Fry, J. L.: Testing
15 Atmospheric Oxidation in an Alabama Forest, *Journal of the Atmospheric Sciences*, 73, 4699–
16 4710, <https://doi.org/10.1175/JAS-D-16-0044.1>, 2016.

17 Fuchs, H., Bohn, B., Hofzumahaus, A., Holland, F., Lu, K. D., Nehr, S., Rohrer, F., and Wahner,
18 A.: Detection of HO₂ by laser-induced fluorescence: calibration and interferences from RO₂
19 radicals, *Atmospheric Measurement Techniques*, 4, 1209–1225, [https://doi.org/10.5194/amt-](https://doi.org/10.5194/amt-4-1209-2011)
20 [4-1209-2011](https://doi.org/10.5194/amt-4-1209-2011), 2011.

21 Fuchs, H., Acir, I.-H., Bohn, B., Brauers, T., Dorn, H.-P., Häseler, R., Hofzumahaus, A., Holland,
22 F., Kaminski, M., Li, X., Lu, K., Lutz, A., Nehr, S., Rohrer, F., Tillmann, R., Wegener, R., and Wahner,
23 A.: OH regeneration from methacrolein oxidation investigated in the atmosphere simulation
24 chamber SAPHIR, *Atmos. Chem. Phys.*, 14, 7895–7908, [https://doi.org/10.5194/acp-14-7895-](https://doi.org/10.5194/acp-14-7895-2014)
25 [2014](https://doi.org/10.5194/acp-14-7895-2014), 2014.

26 Griffith, S. M., Hansen, R. F., Dusanter, S., Stevens, P. S., Alaghmand, M., Bertman, S. B., Carroll,
27 M. A., Erickson, M., Galloway, M., Grossberg, N., Hottle, J., Hou, J., Jobson, B. T., Kammrath, A.,
28 Keutsch, F. N., Lefer, B. L., Mielke, L. H., O'Brien, A., Shepson, P. B., Thurlow, M., Wallace, W.,
29 Zhang, N., and Zhou, X. L.: OH and HO₂ radical chemistry during PROPHET 2008 and CABINEX
30 2009 - Part 1: Measurements and model comparison, *Atmospheric Chemistry and Physics*, 13,
31 5403–5423, <https://doi.org/10.5194/acp-13-5403-2013>, 2013.

32 Guo, J., Wang, Z., Tao Wang, and Zhang, X.: Theoretical evaluation of different factors
33 affecting the HO₂ uptake coefficient driven by aqueous-phase first-order loss reaction,
34 *Science of The Total Environment*, 683, 146–153,
35 <https://doi.org/10.1016/j.scitotenv.2019.05.237>, 2019.

36 Hanke, M., Uecker, J., Reiner, T., and Arnold, F.: Atmospheric peroxy radicals: ROXMAS, a new
37 mass-spectrometric methodology for speciated measurements of HO₂ and \sum RO₂ and first

1 results, International Journal of Mass Spectrometry, 213, 91–99,
2 [https://doi.org/10.1016/S1387-3806\(01\)00548-6](https://doi.org/10.1016/S1387-3806(01)00548-6), 2002.

3 Hansen, R. F., Griffith, S. M., Dusanter, S., Rickly, P. S., Stevens, P. S., Bertman, S. B., Carroll, M.
4 A., Erickson, M. H., Flynn, J. H., Grossberg, N., Jobson, B. T., Lefer, B. L., and Wallace, H. W.:
5 Measurements of total hydroxyl radical reactivity during CABINEX 2009 – Part 1: field
6 measurements, Atmos. Chem. Phys., 14, 2923–2937, [https://doi.org/10.5194/acp-14-2923-](https://doi.org/10.5194/acp-14-2923-2014)
7 2014, 2014.

8 Heard, D. E. and Pilling, M. J.: Measurement of OH and HO₂ in the Troposphere, Chem. Rev.,
9 103, 5163–5198, <https://doi.org/10.1021/cr020522s>, 2003.

10 Hens, K., Novelli, A., Martinez, M., Auld, J., Axinte, R., Bohn, B., Fischer, H., Keronen, P., Kubistin,
11 D., Nölscher, A. C., Oswald, R., Paasonen, P., Petäjä, T., Regelin, E., Sander, R., Sinha, V., Sipilä,
12 M., Taraborrelli, D., Tatum Ernest, C., Williams, J., Lelieveld, J., and Harder, H.: Observation and
13 modelling of HO_x radicals in a boreal forest, Atmospheric Chemistry and Physics, 14, 8723–
14 8747, <https://doi.org/10.5194/acp-14-8723-2014>, 2014.

15 Hofzumahaus, A., Rohrer, F., Lu, K., Bohn, B., Brauers, T., Chang, C.-C., Fuchs, H., Holland, F.,
16 Kita, K., Kondo, Y., Li, X., Lou, S., Shao, M., Zeng, L., Wahner, A., and Zhang, Y.: Amplified Trace
17 Gas Removal in the Troposphere, Science, 324, 1702–1704,
18 <https://doi.org/10.1126/science.1164566>, 2009.

19 Jacob, D.: Heterogeneous chemistry and tropospheric ozone, Atmospheric Environment, 34,
20 2131–2159, [https://doi.org/10.1016/S1352-2310\(99\)00462-8](https://doi.org/10.1016/S1352-2310(99)00462-8), 2000.

21 Jeong, D., Seco, R., Emmons, L., Schwantes, R., Liu, Y., McKinney, K. A., Martin, S. T., Keutsch,
22 F. N., Gu, D., Guenther, A. B., Vega, O., Tota, J., Souza, R. A. F., Springston, S. R., Watson, T. B.,
23 and Kim, S.: Reconciling Observed and Predicted Tropical Rainforest OH Concentrations, JGR
24 Atmospheres, 127, <https://doi.org/10.1029/2020JD032901>, 2022.

25 Kanaya, Y., Cao, R., Kato, S., Miyakawa, Y., Kajii, Y., Tanimoto, H., Yokouchi, Y., Mochida, M.,
26 Kawamura, K., and Akimoto, H.: Chemistry of OH and HO₂ radicals observed at Rishiri Island,
27 Japan, in September 2003: Missing daytime sink of HO₂ and positive nighttime correlations
28 with monoterpenes, J. Geophys. Res., 112, D11308, <https://doi.org/10.1029/2006JD007987>,
29 2007.

30 Kanaya, Y., Hofzumahaus, A., Dorn, H.-P., Brauers, T., Fuchs, H., Holland, F., Rohrer, F., Bohn,
31 B., Tillmann, R., Wegener, R., Wahner, A., Kajii, Y., Miyamoto, K., Nishida, S., Watanabe, K.,
32 Yoshino, A., Kubistin, D., Martinez, M., Rudolf, M., Harder, H., Berresheim, H., Elste, T., Plass-
33 Dülmer, C., Stange, G., Kleffmann, J., Elshorbany, Y., and Schurath, U.: Comparisons of
34 observed and modeled OH and HO₂ concentrations during the ambient measurement period
35 of the HO_xComp field campaign, Atmospheric Chemistry and Physics, 12, 2567–2585,
36 <https://doi.org/10.5194/acp-12-2567-2012>, 2012.

37 Kim, S., Wolfe, G. M., Mauldin, L., Cantrell, C., Guenther, A., Karl, T., Turnipseed, A., Greenberg,

1 J., Hall, S. R., Ullmann, K., Apel, E., Hornbrook, R., Kajii, Y., Nakashima, Y., Keutsch, F. N., DiGangi,
2 J. P., Henry, S. B., Kaser, L., Schnitzhofer, R., Graus, M., Hansel, A., Zheng, W., and Flocke, F. F.:
3 Evaluation of HO_x sources and cycling using measurement-constrained model calculations in
4 a 2-methyl-3-butene-2-ol (MBO) and monoterpene (MT) dominated ecosystem,
5 Atmospheric Chemistry and Physics, 13, 2031–2044, [https://doi.org/10.5194/acp-13-2031-](https://doi.org/10.5194/acp-13-2031-2013)
6 2013, 2013.

7 Kürten, A., Rondo, L., Ehrhart, S., and Curtius, J.: Calibration of a Chemical Ionization Mass
8 Spectrometer for the Measurement of Gaseous Sulfuric Acid, J. Phys. Chem. A, 116, 6375–
9 6386, <https://doi.org/10.1021/jp212123n>, 2012.

10 Lelieveld, J., Butler, T. M., Crowley, J. N., Dillon, T. J., Fischer, H., Ganzeveld, L., Harder, H.,
11 Lawrence, M. G., Martinez, M., Taraborrelli, D., and Williams, J.: Atmospheric oxidation capacity
12 sustained by a tropical forest, Nature, 452, 737–740, <https://doi.org/10.1038/nature06870>,
13 2008.

14 Lew, M. M., Rickly, P. S., Bottorff, B. P., Reidy, E., Sklaveniti, S., Léonardis, T., Locoge, N.,
15 Dusanter, S., Kundu, S., Wood, E., and Stevens, P. S.: OH and HO₂ radical chemistry in a
16 midlatitude forest: measurements and model comparisons, Atmospheric Chemistry and
17 Physics, 20, 9209–9230, <https://doi.org/10.5194/acp-20-9209-2020>, 2020.

18 Lou, S., Holland, F., Rohrer, F., Lu, K., Bohn, B., Brauers, T., Chang, C. C., Fuchs, H., Häsel, R.,
19 Kita, K., Kondo, Y., Li, X., Shao, M., Zeng, L., Wahner, A., Zhang, Y., Wang, W., and Hofzumahaus,
20 A.: Atmospheric OH reactivities in the Pearl River Delta – China in summer 2006: measurement
21 and model results, Atmos. Chem. Phys., 10, 11243–11260, [https://doi.org/10.5194/acp-10-](https://doi.org/10.5194/acp-10-11243-2010)
22 11243-2010, 2010.

23 Ma, X., Tan, Z., Lu, K., Yang, X., Liu, Y., Li, S., Li, X., Chen, S., Novelli, A., Cho, C., Zeng, L., Wahner,
24 A., and Zhang, Y.: Winter photochemistry in Beijing: Observation and model simulation of OH
25 and HO₂ radicals at an urban site, Science of The Total Environment, 685, 85–95,
26 <https://doi.org/10.1016/j.scitotenv.2019.05.329>, 2019.

27 Ma, X., Tan, Z., Lu, K., Yang, X., Chen, X., Wang, H., Chen, S., Fang, X., Li, S., Li, X., Liu, J., Liu, Y.,
28 Lou, S., Qiu, W., Wang, H., Zeng, L., and Zhang, Y.: OH and HO₂ radical chemistry at a suburban
29 site during the EXPLORE-YRD campaign in 2018, Atmospheric Chemistry and Physics, 22,
30 7005–7028, <https://doi.org/10.5194/acp-22-7005-2022>, 2022.

31 Mallik, C., Tomsche, L., Bourtsoukidis, E., Crowley, J. N., Derstroff, B., Fischer, H., Hafermann,
32 S., Hüser, I., Javed, U., Keßel, S., Lelieveld, J., Martinez, M., Meusel, H., Novelli, A., Phillips, G. J.,
33 Pozzer, A., Reiffs, A., Sander, R., Taraborrelli, D., Sauvage, C., Schuladen, J., Su, H., Williams, J.,
34 and Harder, H.: Oxidation processes in the eastern Mediterranean atmosphere: evidence from
35 the modelling of HO_x measurements over Cyprus, Atmospheric Chemistry and Physics, 18,
36 10825–10847, <https://doi.org/10.5194/acp-18-10825-2018>, 2018.

37 Mao, J., Ren, X., Zhang, L., Van Duin, D. M., Cohen, R. C., Park, J.-H., Goldstein, A. H., Paulot,
38 F., Beaver, M. R., Crounse, J. D., Wennberg, P. O., DiGangi, J. P., Henry, S. B., Keutsch, F. N.,

- 1 Park, C., Schade, G. W., Wolfe, G. M., Thornton, J. A., and Brune, W. H.: Insights into hydroxyl
2 measurements and atmospheric oxidation in a California forest, *Atmos. Chem. Phys.*, 12,
3 8009–8020, <https://doi.org/10.5194/acp-12-8009-2012>, 2012.
- 4 Mauldin III, R. L., Kosciuch, E., Henry, B., Eisele, F. L., Shetter, R., Lefer, B., Chen, G., Davis, D.,
5 Huey, G., and Tanner, D.: Measurements of OH, HO₂+RO₂, H₂SO₄, and MSA at the South Pole
6 during ISCAT 2000, *Atmospheric Environment*, 38, 5423–5437,
7 <https://doi.org/10.1016/j.atmosenv.2004.06.031>, 2004.
- 8 Muller, J. B. A., Elste, T., Plass-Dülmer, C., Stange, G., Holla, R., Claude, A., Englert, J., Gilge, S.,
9 and Kubistin, D.: A novel semi-direct method to measure OH reactivity by chemical ionization
10 mass spectrometry (CIMS), *Atmos. Meas. Tech.*, 11, 4413–4433, [https://doi.org/10.5194/amt-](https://doi.org/10.5194/amt-11-4413-2018)
11 [11-4413-2018](https://doi.org/10.5194/amt-11-4413-2018), 2018.
- 12 Novelli, A., Hens, K., Tatum Ernest, C., Kubistin, D., Regelin, E., Elste, T., Plass-Dülmer, C.,
13 Martinez, M., Lelieveld, J., and Harder, H.: Characterisation of an inlet pre-injector laser-
14 induced fluorescence instrument for the measurement of atmospheric hydroxyl radicals,
15 *Atmos. Meas. Tech.*, 7, 3413–3430, <https://doi.org/10.5194/amt-7-3413-2014>, 2014.
- 16 Novelli, A., Vereecken, L., Bohn, B., Dorn, H.-P., Gkatzelis, G. I., Hofzumahaus, A., Holland, F.,
17 Reimer, D., Rohrer, F., Rosanka, S., Taraborrelli, D., Tillmann, R., Wegener, R., Yu, Z., Kiendler-
18 Scharr, A., Wahner, A., and Fuchs, H.: Importance of isomerization reactions for OH radical
19 regeneration from the photo-oxidation of isoprene investigated in the atmospheric
20 simulation chamber SAPHIR, *Atmos. Chem. Phys.*, 20, 3333–3355,
21 <https://doi.org/10.5194/acp-20-3333-2020>, 2020.
- 22 Sjostedt, S. J., Huey, L. G., Tanner, D. J., Peischl, J., Chen, G., Dibb, J. E., Lefer, B., Hutterli, M. A.,
23 Beyersdorf, A. J., Blake, N. J., Blake, D. R., Sueper, D., Ryerson, T., Burkhardt, J., and Stohl, A.:
24 Observations of hydroxyl and the sum of peroxy radicals at Summit, Greenland during
25 summer 2003, *Atmospheric Environment*, 41, 5122–5137,
26 <https://doi.org/10.1016/j.atmosenv.2006.06.065>, 2007.
- 27 Slater, E. J., Whalley, L. K., Woodward-Massey, R., Ye, C., Lee, J. D., Squires, F., Hopkins, J. R.,
28 Dunmore, R. E., Shaw, M., Hamilton, J. F., Lewis, A. C., Crilley, L. R., Kramer, L., Bloss, W., Vu, T.,
29 Sun, Y., Xu, W., Yue, S., Ren, L., Acton, W. J. F., Hewitt, C. N., Wang, X., Fu, P., and Heard, D. E.:
30 Elevated levels of OH observed in haze events during wintertime in central Beijing,
31 *Atmospheric Chemistry and Physics*, 20, 14847–14871, [https://doi.org/10.5194/acp-20-](https://doi.org/10.5194/acp-20-14847-2020)
32 [14847-2020](https://doi.org/10.5194/acp-20-14847-2020), 2020.
- 33 Stone, D., Whalley, L. K., and Heard, D. E.: Tropospheric OH and HO₂ radicals: field
34 measurements and model comparisons, *Chem. Soc. Rev.*, 41, 6348,
35 <https://doi.org/10.1039/c2cs35140d>, 2012.
- 36 Tan, D., Faloona, I., Simpas, J. B., Brune, W., Shepson, P. B., Couch, T. L., Sumner, A. L., Carroll,
37 M. A., Thornberry, T., Apel, E., Riener, D., and Stockwell, W.: HO_x budgets in a deciduous forest:
38 Results from the PROPHET summer 1998 campaign, *J. Geophys. Res.*, 106, 24407–24427,

1 <https://doi.org/10.1029/2001JD900016>, 2001.

2 Tan, Z., Fuchs, H., Lu, K., Hofzumahaus, A., Bohn, B., Broch, S., Dong, H., Gomm, S., Häsel, R.,
3 He, L., Holland, F., Li, X., Liu, Y., Lu, S., Rohrer, F., Shao, M., Wang, B., Wang, M., Wu, Y., Zeng,
4 L., Zhang, Y., Wahner, A., and Zhang, Y.: Radical chemistry at a rural site (Wangdu) in the
5 North China Plain: observation and model calculations of OH, HO₂ and RO₂ radicals,
6 *Atmospheric Chemistry and Physics*, 17, 663–690, <https://doi.org/10.5194/acp-17-663-2017>,
7 2017.

8 Tan, Z., Rohrer, F., Lu, K., Ma, X., Bohn, B., Broch, S., Dong, H., Fuchs, H., Gkatzelis, G. I.,
9 Hofzumahaus, A., Holland, F., Li, X., Liu, Y., Liu, Y., Novelli, A., Shao, M., Wang, H., Wu, Y., Zeng,
10 L., Hu, M., Kiendler-Scharr, A., Wahner, A., and Zhang, Y.: Wintertime photochemistry in
11 Beijing: observations of RO_x radical concentrations in the North China Plain during the BEST-
12 ONE campaign, *Atmospheric Chemistry and Physics*, 18, 12391–12411,
13 <https://doi.org/10.5194/acp-18-12391-2018>, 2018.

14 Tan, Z., Lu, K., Hofzumahaus, A., Fuchs, H., Bohn, B., Holland, F., Liu, Y., Rohrer, F., Shao, M.,
15 Sun, K., Wu, Y., Zeng, L., Zhang, Y., Zou, Q., Kiendler-Scharr, A., Wahner, A., and Zhang, Y.:
16 Experimental budgets of OH, HO₂, and RO₂ radicals and implications for ozone formation in
17 the Pearl River Delta in China 2014, *Atmospheric Chemistry and Physics*, 19, 7129–7150,
18 <https://doi.org/10.5194/acp-19-7129-2019>, 2019.

19 Tanner, D. J. and Eisele, F. L.: Present OH measurement limits and associated uncertainties, *J.*
20 *Geophys. Res.*, 100, 2883, <https://doi.org/10.1029/94JD02609>, 1995.

21 Tanner, D. J., Jefferson, A., and Eisele, F. L.: Selected ion chemical ionization mass
22 spectrometric measurement of OH, *J. Geophys. Res.*, 102, 6415–6425,
23 <https://doi.org/10.1029/96JD03919>, 1997.

24 Thames, A. B., Brune, W. H., Miller, D. O., Allen, H. M., Apel, E. C., Blake, D. R., Bui, T. P.,
25 Commane, R., Crounse, J. D., Daube, B. C., Diskin, G. S., DiGangi, J. P., Elkins, J. W., Hall, S. R.,
26 Hanisco, T. F., Hannun, R. A., Hints, E., Hornbrook, R. S., Kim, M. J., McKain, K., Moore, F. L.,
27 Nicely, J. M., Peischl, J., Ryerson, T. B., St. Clair, J. M., Sweeney, C., Teng, A., Thompson, C. R.,
28 Ullmann, K., Wennberg, P. O., and Wolfe, G. M.: Missing OH reactivity in the global marine
29 boundary layer, *Atmospheric Chemistry and Physics*, 20, 4013–4029,
30 <https://doi.org/10.5194/acp-20-4013-2020>, 2020.

31 Wang, M., Kong, W., Marten, R., He, X.-C., Chen, D., Pfeifer, J., Heitto, A., Kontkanen, J., Dada,
32 L., Kürten, A., Yli-Juuti, T., Manninen, H. E., Amanatidis, S., Amorim, A., Baalbaki, R., Baccarini,
33 A., Bell, D. M., Bertozzi, B., Bräkling, S., Brilke, S., Murillo, L. C., Chiu, R., Chu, B., De Menezes,
34 L.-P., Duplissy, J., Finkenzeller, H., Carracedo, L. G., Granzin, M., Guida, R., Hansel, A., Hofbauer,
35 V., Krechmer, J., Lehtipalo, K., Lamkaddam, H., Lampimäki, M., Lee, C. P., Makhmutov, V., Marie,
36 G., Mathot, S., Mauldin, R. L., Mentler, B., Müller, T., Onnela, A., Partoll, E., Petäjä, T., Philippov,
37 M., Pospisilova, V., Ranjithkumar, A., Rissanen, M., Rörup, B., Scholz, W., Shen, J., Simon, M.,
38 Sipilä, M., Steiner, G., Stolzenburg, D., Tham, Y. J., Tomé, A., Wagner, A. C., Wang, D. S., Wang,
39 Y., Weber, S. K., Winkler, P. M., Wlasits, P. J., Wu, Y., Xiao, M., Ye, Q., Zauner-Wieczorek, M.,

1 Zhou, X., Volkamer, R., Riipinen, I., Dommen, J., Curtius, J., Baltensperger, U., Kulmala, M.,
2 Worsnop, D. R., Kirkby, J., Seinfeld, J. H., El-Haddad, I., Flagan, R. C., and Donahue, N. M.:
3 Rapid growth of new atmospheric particles by nitric acid and ammonia condensation, *Nature*,
4 581, 184–189, <https://doi.org/10.1038/s41586-020-2270-4>, 2020.

5 Wang, X., Jacob, D. J., Eastham, S. D., Sulprizio, M. P., Zhu, L., Chen, Q., Alexander, B., Sherwen,
6 T., Evans, M. J., Lee, B. H., Haskins, J. D., Lopez-Hilfiker, F. D., Thornton, J. A., Huey, G. L., and
7 Liao, H.: The role of chlorine in global tropospheric chemistry, *Atmospheric Chemistry and*
8 *Physics*, 19, 3981–4003, <https://doi.org/10.5194/acp-19-3981-2019>, 2019.

9 Wennberg, P. O., Bates, K. H., Crounse, J. D., Dodson, L. G., McVay, R. C., Mertens, L. A., Nguyen,
10 T. B., Praske, E., Schwantes, R. H., Smarte, M. D., St Clair, J. M., Teng, A. P., Zhang, X., and
11 Seinfeld, J. H.: Gas-Phase Reactions of Isoprene and Its Major Oxidation Products, *Chem. Rev.*,
12 118, 3337–3390, <https://doi.org/10.1021/acs.chemrev.7b00439>, 2018.

13 Whalley, L. K., Edwards, P. M., Furneaux, K. L., Goddard, A., Ingham, T., Evans, M. J., Stone, D.,
14 Hopkins, J. R., Jones, C. E., Karunaharan, A., Lee, J. D., Lewis, A. C., Monks, P. S., Moller, S. J.,
15 and Heard, D. E.: Quantifying the magnitude of a missing hydroxyl radical source in a tropical
16 rainforest, *Atmos. Chem. Phys.*, 11, 7223–7233, <https://doi.org/10.5194/acp-11-7223-2011>,
17 2011.

18 Whalley, L. K., Blitz, M. A., Desservettaz, M., Seakins, P. W., and Heard, D. E.: Reporting the
19 sensitivity of laser-induced fluorescence instruments used for HO₂ detection to an
20 interference from RO₂ radicals and introducing a novel approach that enables HO₂ and certain
21 RO₂ types to be selectively measured, *Atmospheric Measurement Techniques*, 6, 3425–3440,
22 <https://doi.org/10.5194/amt-6-3425-2013>, 2013.

23 Wolfe, G. M., Marvin, M. R., Roberts, S. J., Travis, K. R., and Liao, J.: The Framework for 0-D
24 Atmospheric Modeling (F0AM) v3.1, *Geosci. Model Dev.*, 9, 3309–3319,
25 <https://doi.org/10.5194/gmd-9-3309-2016>, 2016.

26 Woodward-Massey, R., Slater, E. J., Alen, J., Ingham, T., Cryer, D. R., Stimpson, L. M., Ye, C.,
27 Seakins, P. W., Whalley, L. K., and Heard, D. E.: Implementation of a chemical background
28 method for atmospheric OH measurements by laser-induced fluorescence: characterisation
29 and observations from the UK and China, *Atmospheric Measurement Techniques*, 13, 3119–
30 3146, <https://doi.org/10.5194/amt-13-3119-2020>, 2020.

31 Xu, Z., Xue, L., Wang, T., Xia, T., Gao, Y., Louie, P. K. K., and Luk, C. W. Y.: Measurements of
32 Peroxyacetyl Nitrate at a Background Site in the Pearl River Delta Region: Production
33 Efficiency and Regional Transport, *Aerosol Air Qual. Res.*, 15, 833–841,
34 <https://doi.org/10.4209/aaqr.2014.11.0275>, 2015.

35 Yang, X., Lu, K., Ma, X., Liu, Y., Wang, H., Hu, R., Li, X., Lou, S., Chen, S., Dong, H., Wang, F.,
36 Wang, Y., Zhang, G., Li, S., Yang, S., Yang, Y., Kuang, C., Tan, Z., Chen, X., Qiu, P., Zeng, L., Xie,
37 P., and Zhang, Y.: Observations and modeling of OH and HO₂ radicals in Chengdu, China in
38 summer 2019, *Science of The Total Environment*, 772, 144829,

- 1 <https://doi.org/10.1016/j.scitotenv.2020.144829>, 2021.
- 2 Yang, X., Lu, K., Ma, X., Gao, Y., Tan, Z., Wang, H., Chen, X., Li, X., Huang, X., He, L., Tang, M.,
3 Zhu, B., Chen, S., Dong, H., Zeng, L., and Zhang, Y.: Radical chemistry in the Pearl River Delta:
4 observations and modeling of OH and HO₂ radicals in Shenzhen in 2018, *Atmospheric*
5 *Chemistry and Physics*, 22, 12525–12542, <https://doi.org/10.5194/acp-22-12525-2022>, 2022.
- 6 Yang, Y., Shao, M., Wang, X., Nölscher, A. C., Kessel, S., Guenther, A., and Williams, J.: Towards
7 a quantitative understanding of total OH reactivity: A review, *Atmospheric Environment*, 134,
8 147–161, <https://doi.org/10.1016/j.atmosenv.2016.03.010>, 2016.
- 9 Zou, Z., Chen, Q., Xia, M., Yuan, Q., Chen, Y., Wang, Y., Xiong, E., Wang, Z., and Wang, T.: OH
10 measurements in the coastal atmosphere of South China: possible missing OH sinks in aged
11 air masses, *Atmospheric Chemistry and Physics*, 23, 7057–7074, [https://doi.org/10.5194/acp-](https://doi.org/10.5194/acp-23-7057-2023)
12 [23-7057-2023](https://doi.org/10.5194/acp-23-7057-2023), 2023.

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14

**Observation and modelling of OH and HO₂* radicals at a
subtropical rural site and implications for secondary
pollutants**

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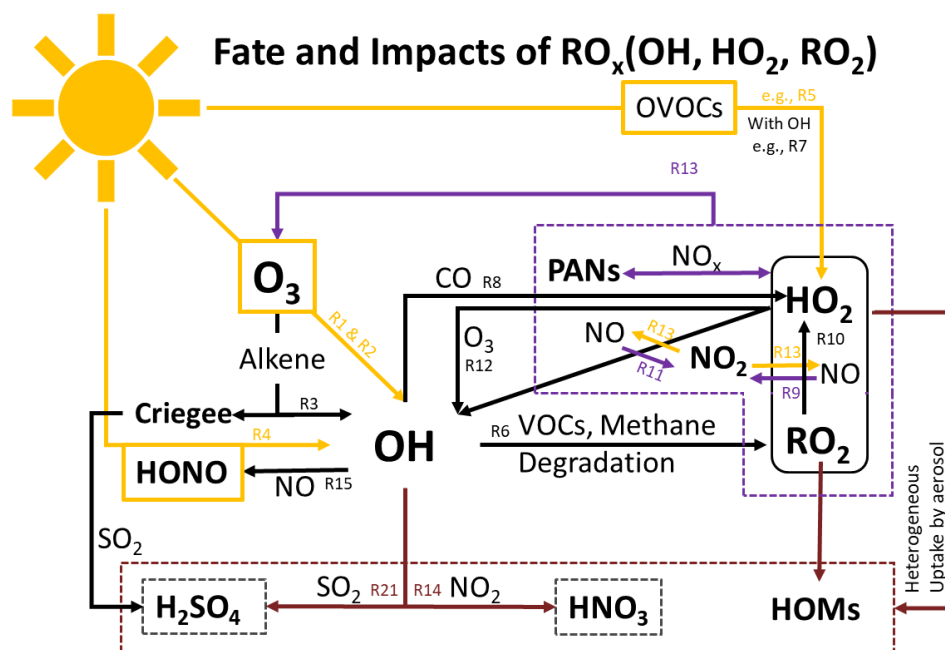


Figure S1 Schematic of the RO_x family's photochemical pathway. Photolysis reactions are highlighted in yellow, reactions contributing to secondary aerosol production are marked in brown; and reactions associated with photochemical pollution are indicated in purple. The chemical reactions (R1 to R21) referenced in Table S1.

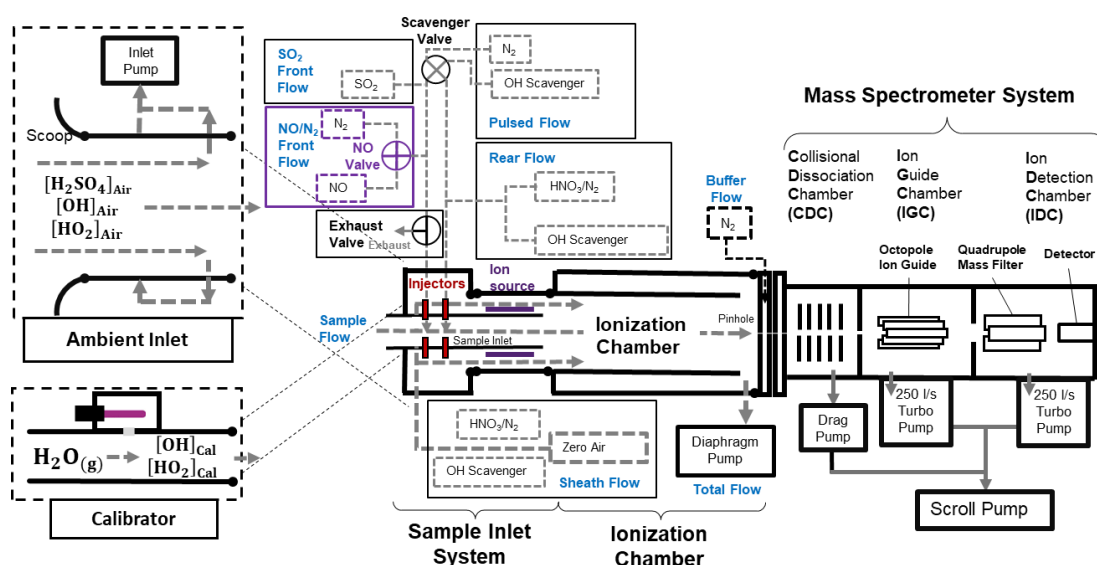
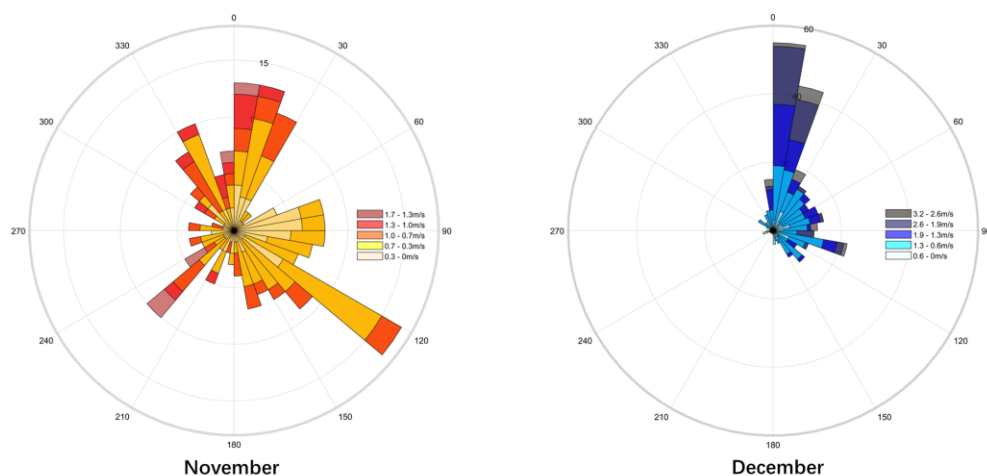
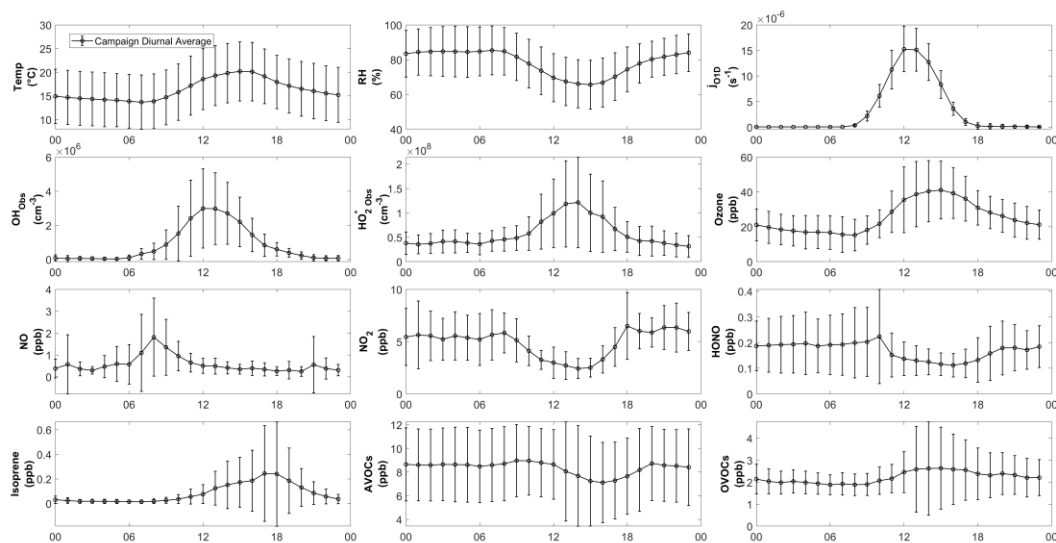


Figure S2 Schematic diagram of the PolyU-CIMS system. The CIMS is composed of two detachable components: the ambient inlet and the calibrator; and the main body, which includes the sample inlet system, ionization chamber, and the mass spectrometer system. The frames labeled in purple highlight the additional valve

1 incorporated for HO₂ measurement. Further details on setup, measurement principles
 2 of the CIMS are available in a previous study (Zou et al., 2023) and Text S1.



3
 4 **Figure S3** The wind rose for November and December.



5
 6 **Figure S4** Diurnal profiles of average concentration of HO₂*, OH, meteorological
 7 data and trace gases of the whole campaign. The shaded error bars represent standard
 8 deviations of the averaged data.

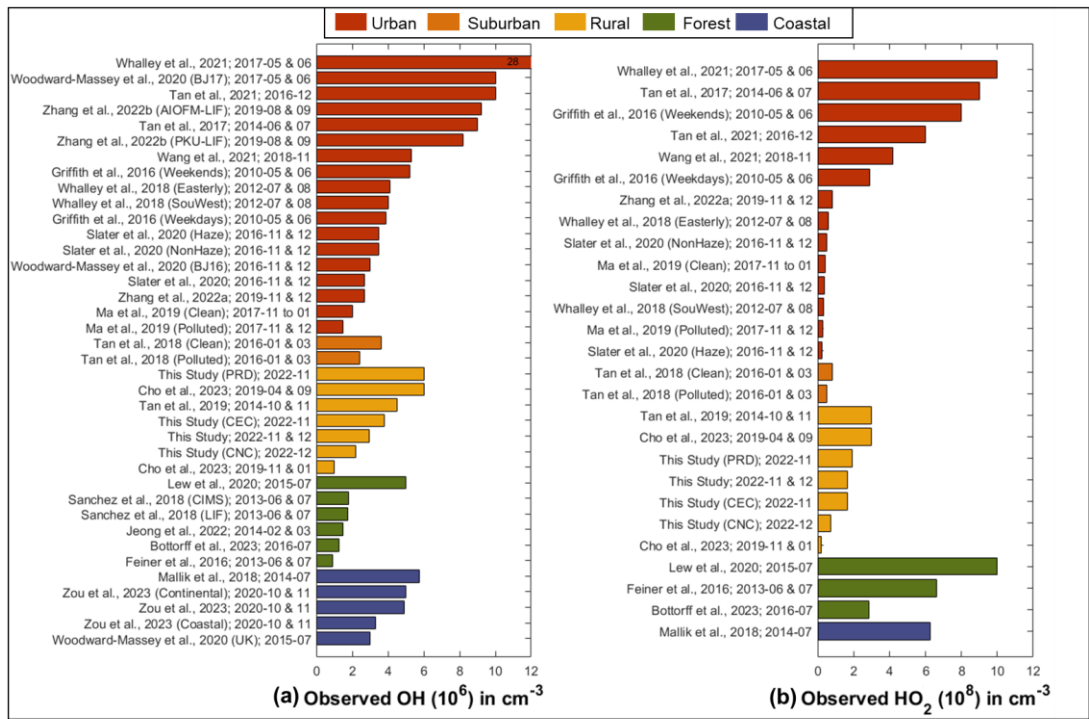


Figure S5 Typical daily averaged maximum concentration of (a) OH and (b) HO₂ observed in various geophysical regions, including coastal (blue), forest (green), rural (yellow) and urban (red).

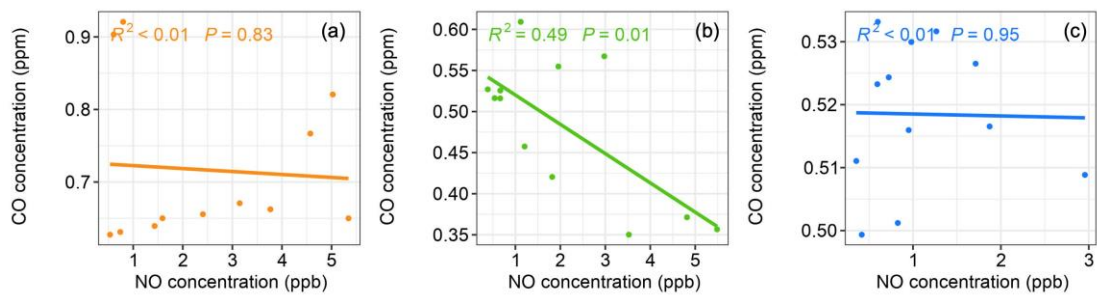


Figure S6 Relationship between NO and CO concentrations from 7:00 to 10:00 in (a) PRD, (b) CEC and (c) CNC. The solid lines depict the linear regression fits, with the corresponding equations R^2 and p-values annotated on the plot.

Table S1 The HO_x related reactions in the model.

Ambient: HO_x Productions	
(R1)	$O_3 + h\nu(<340\text{ nm}) \rightarrow O(^1D) + O_2$
(R2)	$O(^1D) + H_2O \rightarrow OH + OH$
(R3)	Alkenes + $O_3 \rightarrow RO_x$ + Products
(R4)	$HONO + h\nu(<400\text{ nm}) \rightarrow OH + NO$
(R5)	$HCHO + h\nu(<335\text{ nm}) + 2O_2 \rightarrow 2HO_2 + CO$
Ambient: HO_x Interconversions	
(R6)	$OH + RH + O_2 \rightarrow RO_2 + H_2O$
(R7)	$HCHO + OH + O_2 \rightarrow CO + H_2O + HO_2$
(R8)	$CO + OH + O_2 \rightarrow CO_2 + HO_2$
(R9)	$RO_2 + NO \rightarrow RO + NO_2$
(R10)	$RO + O_2 \rightarrow R'CHO + HO_2$
(R11)	$HO_2 + NO \rightarrow OH + NO_2$
(R12)	$HO_2 + O_3 \rightarrow OH + 2O_2$
(R13)	$NO_2 + h\nu(<420\text{ nm}) + O_2 \rightarrow NO + O_3$
Ambient: HO_x Loss	
(R14)	$OH + NO_2 \rightarrow HNO_3$
(R15)	$OH + NO \rightarrow HONO$
(R16)	$RO_2 + NO \rightarrow RONO_2$
(R17)	$RO_2 + RO_2 \rightarrow \text{products}$
(R18)	$RO_2 + HO_2 \rightarrow ROOH + O_2$
(R19)	$HO_2 + HO_2 \rightarrow H_2O_2 + O_2$
(R20)	$HO_2 + HO_2 + H_2O \rightarrow H_2O_2 + H_2O + O_2$
(R21)	$OH + SO_2 + O_2 + H_2O + M \rightarrow H_2SO_4 + HO_2 + M$
CIMS: Reactions in Sample Inlet System	
(R21)	$OH + SO_2 + O_2 + H_2O + M \rightarrow H_2SO_4 + HO_2 + M$
(R11)	$HO_2 + NO \rightarrow OH + NO_2$
(R22)	$RO_2 + NO + O_2 \rightarrow R'CHO + HO_2 + NO_2$
(R23)	Scavenger gas + OH \rightarrow Products
CIMS: Reactions in Ionization Chamber	
(R24)	$HNO_3 + e^- \rightarrow NO_2^- + OH$
(R25)	$HNO_3 + NO_2^- \rightarrow NO_3^- + HONO$
(R26)	$NO_3^- + (HNO_3)_m + (H_2O)_n + M \rightarrow NO_3^-(HNO_3)_m(H_2O)_n + M$
(R27)	$H_2SO_4 + NO_3^-(HNO_3)_m(H_2O)_n \rightarrow HSO_4^-(HNO_3)_m(H_2O)_n + HNO_3$
CIMS: Reactions in Collisional Dissociation Chamber	
(R28)	$NO_3^-(HNO_3)_m(H_2O)_n + M \rightarrow NO_3^- + (HNO_3)_m + (H_2O)_n + M$
(R29)	$HSO_4^-(HNO_3)_m(H_2O)_n + M \rightarrow HSO_4^- + (HNO_3)_m + (H_2O)_n + M$
CIMS: Calibration	
(R30)	$H_2O + h\nu(184.9\text{ nm}) + O_2 \rightarrow HO_2 + OH$

1
2

3 **Table S2** Summary of instruments used, and species measured during the field
4 campaign.

Instruments	Species	Resolution	Detection Limits	Accuracy
Q-CIMS (NO_3^-)	OH	1 hours	$3 \times 10^5 \text{ cm}^{-3}$	$\pm 46\%$
	HO_2^*	1 hours	$20 \times 10^5 \text{ cm}^{-3}$	$\pm 44\%$
Thermo 42i-TL	NO	1 min	60 ppt	$\pm 5.2\%$
Thermo 49i	O_3	1 min	1 ppb	$\pm 6.0\%$
NO_2 -11r-EP	NO_2	1 min	60 ppt	$\pm 6.0\%$
Online GC-MS	VOCs	1 hour	10 ppt	$\pm 20\%$
Thermo 43i	SO_2	1 min	0.5 ppb	$\pm 6.1\%$
Thermo 48i	CO	1 min	40 ppb	$\pm 7.4\%$
Thermo 17i	NH_3	2 mins	1 ppb	$\pm 8\%$
LOPAP-03	HONO	5 mins	1 ppt	$\pm 10\%$
SMPS	Aerosol Particles	5 mins	1 particle cm^{-3}	$\pm 10\%$

1

- 1 **Tabel S3** Configurations of the PolyU CIMS in Hok Tsui 2020 and Conghua 2022 campaigns. The changes from the last configuration were
- 2 labelled by red color.

a) Hok Tsui 2020								b) CongHua 2022								
Efficiency Related	Parameters	Gas	Values	Units	Specification for Measurement	Values	Units	Efficiency Related	Parameters	Gas	Values	Units	Specification for Measurement	Values	Units	
E _{Conv}	Front Injection	SO ₂ (0.9%)	5	sccm	Sample Flow [SO ₂]	12	ppm	E _{conv}	Front Injection	SO ₂ (0.9%)	5	sccm	Sample Flow [SO ₂]	12	ppm	
										NO (0.9%)	0.5		Sample Flow [NO]	1.2	ppm	
	Pulse Valve	N ₂	2	sccm	Cycle Duration (OH)	6	mins		Pulse Valve	N ₂	2	sccm	Cycle Duration (OH)	6	mins	
		C ₃ F ₆ (99.9%)	2	sccm	B/S Ratio for OH measurement	8%				C ₃ F ₆ (99.9%)	2	sccm	Cycle Duration (HO ₂)	60	mins	
	Rear Injection	C ₃ F ₆ (99.9%)	2	sccm	Sample Flow [C ₃ F ₆]	1072	ppm		Rear Injection	C ₃ F ₆ (99.9%)	2	sccm	Sample Flow [C ₃ F ₆]	1072	ppm	
		HNO ₃	10	sccm	Reaction Time	47	ms			HNO ₃	10	sccm	Reaction Time	47	ms	
Sample Flow			3.7	slpm	Sample Flow Speed	55	cm/s	Sample Flow			3.7	slpm	Sample Flow Speed	55	cm/s	
E _{Ion}	Sheath Flow	Zero Air	12.6	slpm	Reynolds Number in Ionization Chamber	>4000	Turbulent flows	E _{Ion}	Sheath Flow	Zero Air	12.6	slpm	Reynolds Number in Ionization Chamber	>4000	Turbulent flows	
		HNO ₃	10	sccm						HNO ₃	10	sccm				
		C ₃ F ₆ (99.9%)	2	sccm	Sheath Flow [C ₃ F ₆]	159	ppm			C ₃ F ₆ (99.9%)	2	sccm	Sheath Flow [C ₃ F ₆]	159	ppm	
	Total Flow		16.8	slpm	Sheath Flow Speed	25	cm/s		Total Flow		16.8	slpm	Sheath Flow Speed	25	cm/s	
	Sheath Voltages		-80	V	Voltages Difference for ionization	48	V		Sheath Voltages		-80	V	Voltages Difference for ionization	48	V	
	Sample Voltages		-32	V					Sample Voltages		-32	V				
E _{Trans}	Buffer Gas	N ₂	440	sccm	Voltages Difference for transmission	80	V	E _{Trans}	Buffer Gas	N ₂	440	sccm	Voltages Difference for transmission	80	V	
	Buffer Voltages		-70	V					Buffer Voltages		-70	V				
	Pinhole Voltages		-40	V					Pinhole Voltages		-40	V				
Cal	Calibration Flow		10	slpm	Calibration Factor			Cal	Calibration Flow		10	slpm	Calibration Factors		C _{OH}	1.09*10 ⁻⁸
	Flow Speed		65	cm/s	C _{OH}	1.21*10 ⁻⁸	cm ³		Flow Speed		65	cm/s	C _{HO2}	1.07*10 ⁻⁸	cm ³	
	Product It Value		8.8*10 ¹⁰	photon/cm ²	(Reagent ion: N ¹⁸ O ₃ ⁺)				Product It Value		8.8*10 ¹⁰	photon/cm ²	(N ¹⁸ O ₃ ⁺)		C _{H2SO4}	6.01*10 ⁻⁹
Uncertainties	Sigma		2		Detection Limit (×10 ⁵ cm ⁻³) (3σ)	In lab	1.7	Overall Uncertainties (2σ)	OH		44%	Detection Limit in Field Study (×10 ⁵ cm ⁻³) (3σ)	OH	3		
	Calibration		38%	Day		12	H ₂ SO ₄			40%	H ₂ SO ₄		1			
	Overall		44%	Night		8.5	HO ₂			46%	HO ₂		20			

- 3
- 4 Notes: B/S Ratio - background to signal ratio.

1 **Table S4** Average mixing ratios (ppbv) and standard deviations of measured VOCs
2 that are constrained in the box model in the entire campaign and in different cases.

Species	Total	PRD	CEC	CNC	MCM Abb.
Ethane	2.301±0.826	1.613±0.883	1.864±0.534	2.032±0.141	C2H6
Ethylene	0.691±0.331	0.622±0.257	0.582±0.206	0.581±0.162	C2H4
Propane	1.593±0.618	2.099±0.808	1.217±0.314	1.239±0.274	C3H8
Propene	0.072±0.077	0.127±0.173	0.059±0.023	0.067±0.035	C3H6
i-Butane	0.439±0.281	0.891±0.494	0.353±0.103	0.310±0.061	IC4H10
n-Butane	0.653±0.405	1.284±0.709	0.484±0.145	0.438±0.083	NC4H10
Acetylene	0.917±0.416	0.905±0.350	0.819±0.305	0.758±0.072	C2H2
trans-2-Butene	0.015±0.007	0.015±0.005	0.017±0.011	0.016±0.006	TBUT2ENE
cis-2-Butene	0.083±0.038	0.165±0.033	0.083±0.013	0.080±0.014	CBUT2ENE
Butene	0.044±0.021	0.047±0.049	0.037±0.014	0.048±0.010	BUT1ENE
Chloromethane	0.838±0.221	0.510±0.076	0.839±0.153	0.856±0.154	CH3CL
1,3-Butadiene	0.008±0.008	0.012±0.009	0.008±0.007	0.007±0.005	C4H6
Acetaldehyde	0.922±0.355	1.456±0.365	NaN	0.525±0.160	CH3CHO
Bromomethane	0.009±0.002	0.010±0.002	0.009±0.001	0.009±0.001	CH3BR
Chloroethane	0.023±0.012	0.015±0.011	0.021±0.009	0.020±0.004	CH3CH2CL
i-Pentane	0.338±0.167	0.610±0.224	0.311±0.059	0.237±0.042	IC5H12
1-Pentene	0.043±0.016	0.067±0.024	0.033±0.008	0.037±0.006	PENT1ENE
n-Pentane	0.187±0.102	0.339±0.181	0.121±0.034	0.136±0.022	NC5H12
trans-2-Pentene	0.003±0.005	0.011±0.005	0.002±0.002	0.001±0.001	TPENT2ENE
cis-2-Pentene	0.002±0.003	0.007±0.003	0.001±0.001	0.000±0.001	CPENT2ENE
Acrolein	0.060±0.031	0.092±0.043	0.053±0.019	0.043±0.015	ACR
Propanal	0.011±0.006	0.015±0.011	0.010±0.004	0.009±0.004	C2H5CHO
Vinylidene chloride	0.004±0.003	0.003±0.002	0.002±0.002	0.005±0.003	CCL2CH2
2,2-Dimethylbutane	0.017±0.015	0.046±0.025	0.012±0.003	0.010±0.002	M22C4
Dichloromethane	1.142±0.838	2.510±1.430	1.035±0.347	0.933±0.179	CH2CL2
2,3-Dimethylbutane	0.026±0.023	0.064±0.033	0.025±0.007	0.015±0.004	M23C4
2-Methylpentane	0.071±0.045	0.237±0.053	0.060±0.016	0.056±0.012	M2PE
3-Methylpentane	0.052±0.039	0.120±0.061	0.040±0.011	0.036±0.009	M3PE
Methyl tert-butyl ether	0.072±0.042	0.135±0.051	0.068±0.016	0.055±0.012	MTBE
1-Hexene	0.005±0.005	0.013±0.004	0.004±0.004	0.002±0.001	HEX1ENE
n-Hexane	0.066±0.043	0.131±0.063	0.049±0.016	0.042±0.012	NC6H14
Methacrolein	0.062±0.062	0.116±0.058	0.108±0.069	0.025±0.010	MACR
1,1-Dichloroethane	0.009±0.005	0.008±0.004	0.008±0.004	0.007±0.001	CHCL2CH3
Butyraldehyde	0.536±0.211	0.446±0.179	0.496±0.142	0.448±0.157	C3H7CHO
1,2-Dichloroethylene	0.049±0.076	0.138±0.171	0.032±0.014	0.026±0.014	DICLETH
2-Butanone	0.248±0.244	0.546±0.490	0.214±0.086	0.134±0.050	MEK
Ethyl acetate	0.267±0.386	0.707±0.868	0.174±0.076	0.148±0.076	ETHACET
Chloroform	0.082±0.032	0.125±0.042	0.087±0.015	0.069±0.010	CHCL3
Methylchloroform	0.002±0.001	0.004±0.000	0.001±0.000	0.002±0.000	CH3CCCL3
2-Methylhexane	0.015±0.017	0.046±0.028	0.010±0.004	0.007±0.002	M2HEX
Cyclohexane	0.019±0.015	0.041±0.021	0.011±0.005	0.011±0.004	CHEX
Tetrachloromethane	0.073±0.006	0.070±0.004	0.066±0.004	0.075±0.003	CCL4
3-Methylhexane	0.020±0.024	0.064±0.042	0.012±0.004	0.009±0.003	M3HEX
Benzene	0.346±0.139	0.288±0.106	0.315±0.102	0.311±0.029	BENZENE
Ethylene dichloride	0.359±0.170	0.259±0.112	0.342±0.185	0.414±0.097	CH2CLCH2CL
n-Heptane	0.035±0.023	0.072±0.038	0.024±0.005	0.022±0.003	NC7H16
Crotonaldehyde	0.446±0.139	0.464±0.008	0.480±0.007	0.501±0.008	C3MDBAL
Trichloroethene	0.021±0.023	0.061±0.037	0.018±0.013	0.013±0.003	TRICLETH
1,2-Dichloropropane	0.085±0.038	0.118±0.030	0.096±0.016	0.071±0.011	CL12PROP
Pantanal	0.018±0.011	0.033±0.017	0.019±0.008	0.013±0.006	C4H9CHO
1,3-Dichloro-1-propene	0.003±0.001	0.003±0.001	0.002±0.001	0.002±0.001	CLC3H4CL
4-Methyl-2-pentanone	0.005±0.007	0.019±0.004	0.001±0.001	0.000±0.001	MBK
Toluene	0.282±0.266	0.706±0.498	0.212±0.057	0.173±0.049	TOLUENE
n-Octane	0.009±0.007	0.022±0.007	0.005±0.001	0.005±0.001	NC8H18
1,1,2-Trichloroethane	0.014±0.010	0.011±0.008	0.013±0.009	0.015±0.005	CH2CLCHCL2
Tetrachloroethylene	0.015±0.013	0.040±0.019	0.014±0.004	0.009±0.001	TCE
2-Hexanone	0.050±0.025	0.087±0.029	NaN	0.038±0.011	HEX2ONE
Hexanal	0.041±0.022	0.076±0.025	NaN	0.030±0.009	C5H11CHO
1,2-Dibromoethane	0.002±0.002	0.004±0.001	0.001±0.001	0.002±0.001	DIBRET
Ethylbenzene	0.042±0.031	0.072±0.035	0.028±0.022	0.034±0.023	EBENZ
o-Xylene	0.039±0.030	0.077±0.039	0.027±0.018	0.031±0.017	OXYL
Styrene	0.020±0.013	0.034±0.008	0.012±0.005	0.013±0.007	STYRENE
Isopropylbenzene	0.006±0.006	0.016±0.003	0.003±0.001	0.003±0.001	IPBENZ
1,1,2,2-Tetrachloroethane	0.003±0.002	0.005±0.001	0.003±0.002	0.002±0.001	CHCL2CHCL2
n-Propylbenzene	0.005±0.004	0.013±0.003	0.003±0.001	0.002±0.001	PBENZ
m-Ethyltoluene	0.007±0.007	0.019±0.008	0.006±0.003	0.004±0.002	METHTOL
p-Ethyltoluene	0.005±0.005	0.013±0.005	0.003±0.002	0.002±0.002	PETHTOL
1,3,5-Trimethylbenzene	0.005±0.005	0.015±0.005	0.004±0.002	0.002±0.001	TM135B
n-Decane	0.003±0.003	0.009±0.003	0.002±0.001	0.001±0.001	NC10H22
Benzaldehyde	0.005±0.004	0.013±0.004	0.003±0.001	0.002±0.001	BENZAL
1,2,4-Trimethylbenzene	0.009±0.009	0.024±0.013	0.008±0.003	0.004±0.002	TM124B
1,2,3-Trimethylbenzene	0.004±0.003	0.010±0.004	0.003±0.002	0.002±0.001	TM123B
Undecane	0.002±0.002	0.006±0.001	0.002±0.001	0.000±0.001	NC11H24
Dodecane	0.009±0.003	0.015±0.003	0.007±0.001	0.010±0.002	NC12H26

3

4

Text S1 Calibration procedures

The calibration of the Chemical Ionization Mass Spectrometer (CIMS) involves the generation of OH and HO₂ radicals through photolysis of water vapor by 184.9 nm light, as outlined in Reaction R30. The concentration of radicals produced during calibration is determined from the known concentration of water vapor [H₂O], which is calculated from water vapor pressure, the relative humidity and temperature. Other essential parameters include the photolysis cross-section of water vapor ($\sigma_{H_2O} = 7.14 \times 10^{-20} \text{ cm}^2$; Cantrell et al., 1997), the photolysis quantum yield (Φ , assumed to be 1, Kürten et al., 2012) and the photon flux (It value, see details about It value determination on Kürten et al., 2012). The generated radical concentrations ([OH] and [HO₂]) are calculated using the following equation:

$$[OH] = [HO_2] = [H_2O] * \sigma_{H_2O} * \Phi * It$$

From these values, the calibration factors for OH and HO₂ (C_{OH} and C_{HO_2}) are calculated using the signals obtained during calibration (S_{OHcal} and S_{HO_2cal}), as expressed in the transformed form of Equations E1 and E2:

$$C_{OH} = \frac{1}{[OH]_{cal}} \times \frac{S_{OHcal}}{S_{62}} \text{ (E1, transformed)}$$

$$C_{HO_2} = \frac{1}{[HO_2]_{cal}} \times \frac{S_{HO_2cal}}{S_{62}} \text{ (E2, transformed)}$$

The calibrator produced OH and HO₂ concentrations in the range of 3×10^7 to $1 \times 10^9 \text{ cm}^{-3}$ depending on RH conditions in 10 LMP synthetic air. More detailed information on calculation procedures is given in our previous study (Zou et al., 2023).

Text S2 Modification for HO₂ measurement

To measure HO₂, a valve was added to switch the injection gas between NO and N₂ as indicated by the purple frame in Figure S2. When N₂ is added, the CIMS operates in OH measurement mode. When NO is added to the sample flow, the CIMS switches to HO₂ mode for total HO_x measurement.

During HO₂ measurement, ambient HO₂ is converted by NO to OH. It should be noted that in HO₂ mode, the increasing NO concentration can enhance HO₂ conversion to OH (R11). However, excessive NO levels trigger the HONO formation

when reacts with OH (R15), competing with the OH conversion process by SO₂ (R21) and lowering the detection efficiency for OH. Consequently, the NO-to-SO₂ concentration ratio is crucial for HO₂ measurements. Sensitivity tests revealed an optimal [NO]/[SO₂] ratio of 0.1 for the PolyU-CIMS, achieving 100% conversion of HO₂ in the laboratory ([H₂SO₄]/[HO₂] = 1), aligning with prior research recommendations (Edwards et al., 2003; Sjostedt et al., 2007). Because the concentrations of both SO₂ and NO injected into the sample flow are maintained at levels over 100-1000 times higher than ambient levels, and the injection flow rates are fixed, the efficiency of the HO₂-to-OH conversion remains stable and is believed to be complete.

Text S3 The Background mode and scavenge efficiency

In the background mode, the scavenger gas C₃F₆ is introduced into the sample flow along with SO₂. Since the concentration of C₃F₆ is 100 times higher than that of SO₂, ambient OH and any HO₂ converted to OH are scavenged by C₃F₆, rather than converted to H₂SO₄.

To determine the amount of C₃F₆ needed to achieve complete OH scavenging, we gradually increased the C₃F₆ concentration added to high levels of OH and HO₂ ([HO_x] ≈ 10⁹ cm⁻³) generated from the calibrator in synthetic air until no further reduction in the measured signal, which indicates complete scavenging of OH. This point defines the background noise, which is attributed to any Criegee intermediates and ambient sulfuric acid. In our setup, residual C₃F₆ is present after the CIMS switches from background to signal mode, but it does not affect the measurement results. As shown in Figure S7, after switching off C₃F₆, the measurement signals rapidly return to their initial levels within 20 seconds. Data affected by residual C₃F₆ is excluded to minimize its impact on the measurements.

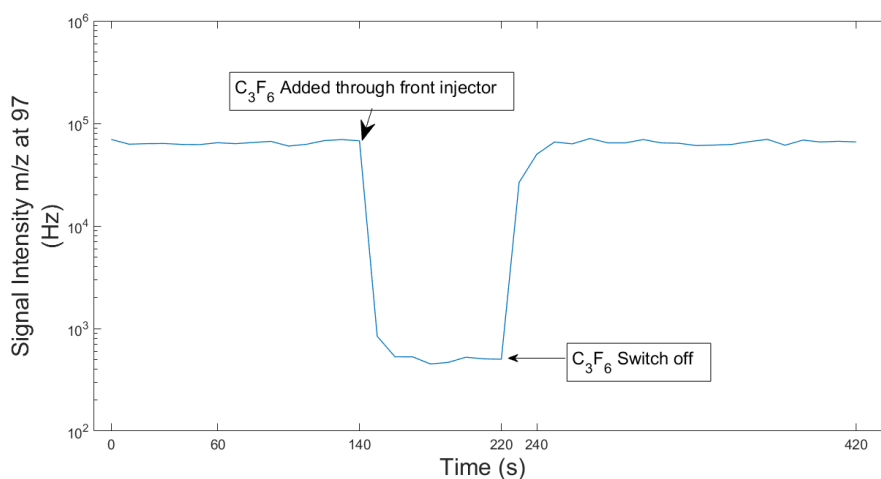


Figure S7 Variation of signal intensity m/z at 97 before C_3F_6 addition with time, after addition and switching off of C_3F_6 in synthetic air containing OH of $\sim 5 \times 10^8 \text{ cm}^{-3}$.

Text S4 Measurement interferences

4.1 Sampling Loss

Wall losses in the ambient inlet were evaluated by varying the distance between the calibration lamp and the inlet to assess potential signal attenuation. The instrument was calibrated in two distinct configurations: initially, the lamp was positioned close to the CIMS sample inlet (Figure S6a), and subsequently, moved away from the CIMS sample inlet (Figure S6b). By comparing the observed signals from these two configurations, we were able to calculate the wall losses associated with the ambient inlet. The results indicated no significant difference ($<1\%$) between the two measurements, suggesting negligible wall losses in the sampling system.

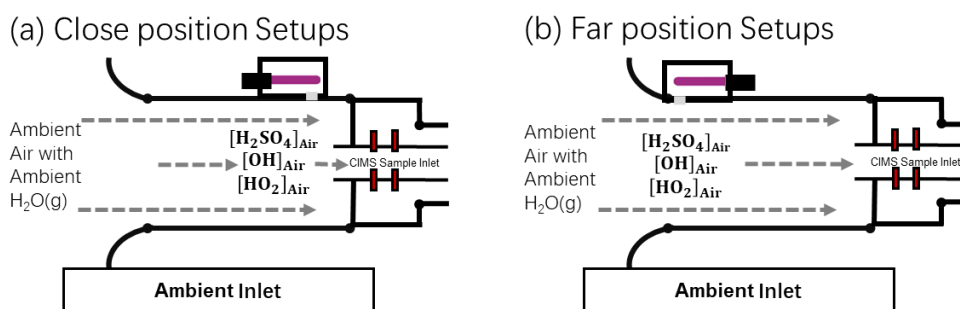


Figure S8 Calibration process during ambient sampling in (a) close and (b) far positions.

Furthermore, the potential for radical-radical loss after the radicals exit the calibrator and enter the sampling inlet was considered. Given the flow speed of the ambient inlet (12.2 m/s) the sample inlet (55 cm/s), along with the short distances involved—the calibration lamp is less than 1 cm from the sample inlet, and the sampling port to the front injectors is 1 cm—the calculated transport time for radicals to reach the front injectors is less than 20 ms. This short transport time is sufficiently short to prevent significant radical-radical losses. Additionally, since the sample inlet draws from the central part of the airflow within the ambient inlet, and the flow is laminar, wall losses at this stage are also considered negligible.

4.2 OH interference

The OH interference in the PolyU-CIMS, resulting from ambient NO + HO₂ recycling (R11) and ionization process (R24, artificial OH), was accounted for and included in the measurement uncertainty, as outlined by Zou et al. (2023). To assess and correct this interference, model simulations were conducted. Initially, we constrained all measured species (except OH) in the model, and a three-day spin-up was used to simulate the chemical conditions of the sampled air during measurement. Subsequently, the outputs (except OH) were utilized as inputs for another simulation that included the injection of 12 ppm SO₂, to simulate chemical reactions in the inlet with reaction time of 47 ms - matching the reaction time in the CIMS. Photolysis frequencies were maintained at zero to simulate the dark environment of the inlet. The modeled H₂SO₄ concentrations reflect interference from NO + HO₂ reactions, with an average value of 2.91×10^4 and range from 10^4 to 10^5 cm⁻³ throughout the whole campaign. These calculated interferences were subtracted from the ambient results.

In this study, the PolyU-CIMS encountered additional interference from residual NO in the injectors when switching from NO (used for HO₂ measurement) to N₂ (used for OH measurement). To prevent residual NO buildup, the inlet was cleaned daily, and a one-hour calibration was performed at both the start and end of day's measurements to monitor NO residuals. Monitoring results showed that the NO residual time for the PolyU - CIMS was approximately 26 mins, similar to values

1 reported in earlier studies (Edwards et al., 2003; Sjostedt et al., 2007). Consequently,
2 data collected during the residual period (30 mins after switching the measurement
3 target from HO₂ to OH) were discarded to eliminate NO residual interference from
4 the final dataset. Details about how the duty cycle was configured to monitor NO
5 residual interference are provided in Text S5.

7 **4.3 HO₂ interference**

8 The concentration of injected NO is the primary source of HO₂ measurement
9 interference in this study. High NO concentrations convert ambient RO₂, particularly
10 alkene and aromatic-related RO₂, into HO₂ and then OH, leading to a positive bias in
11 HO₂ measurements (Fuchs et al., 2014). To assess HO₂ interference, we conducted
12 model analysis using the same settings as those used for OH interference, including
13 observational data input (except HO₂* and OH), reaction time, and dark conditions,
14 while adjusting specific parameters to address HO₂-related reactions. The modeled
15 ambient RO₂ concentrations (blue line in Figure S9a) were used as inputs for
16 additional simulations assessing RO₂ interference with HO₂ via reaction with added
17 NO (1.2 ppm) and SO₂ (12 ppm) in the CIMS inlet. The H₂SO₄ concentrations
18 resulting from the model simulation represent the interference in HO₂ measurements
19 attributed to RO₂ (blue line in Figure S9b). By comparing these results with the
20 modeled H₂SO₄ concentrations (red line in Figure S9b) when modelled ambient HO₂
21 concentrations (red line in Figure S9a) are constrained, we determine that the
22 modelled average daytime (08:00-16:00) RO₂ interference increased HO₂ by 127%,
23 117%, and 144% PRD, CEC and CNC case, respectively, and in ranged of 79% to
24 222% throughout the campaign. These results indicate substantial positive
25 interference from RO₂ in HO₂ measurements under the CIMS configurations adopted
26 in the field study.

27 We note that the MCM model has inherent uncertainties on RO₂ chemistry, such
28 as incomplete representations of RO₂ species and the assumption that most of RO₂
29 species share the same reaction rate constant with NO. In addition, HO₂ was over-

simulated by the model (Figure S9a). For these reasons, we use the modeled interferences only for sensitivity tests and qualitative discussion rather than correcting the measured HO_2^* . Nonetheless, these values indicate the potential magnitude of positive interference in HO_2^* measurements caused by RO_2 .

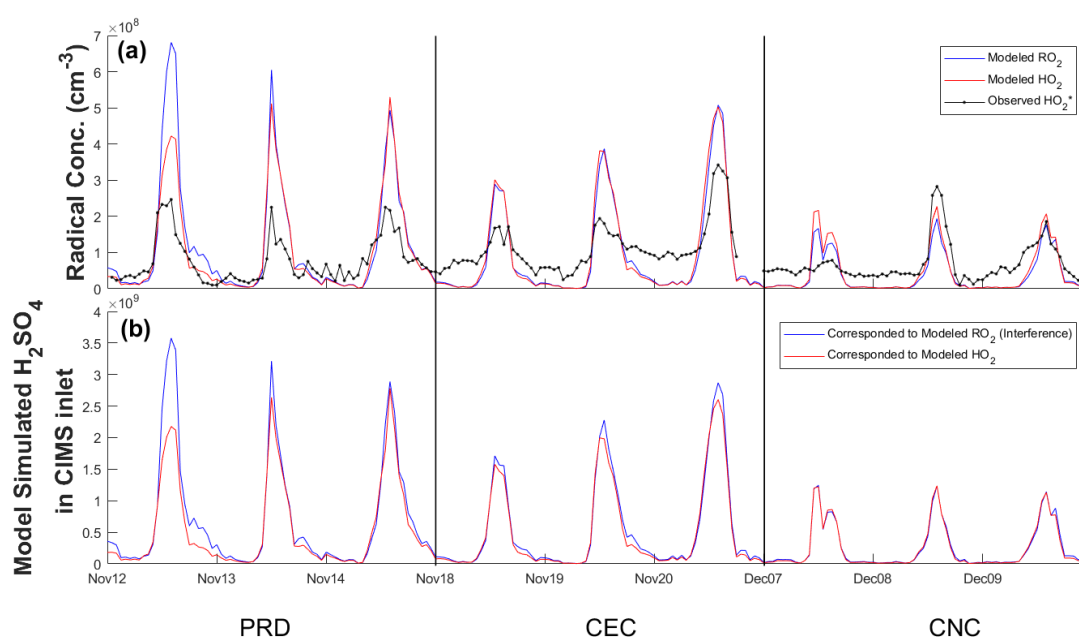


Figure S9: Modeled peroxy radical concentrations and resulting H_2SO_4 simulations. (a) Temporal profiles of modeled RO_2 and HO_2 concentrations alongside observed HO_2^* for the PRD, CEC, and CNC periods. (b) Simulated H_2SO_4 concentrations corresponding to RO_2 and HO_2 in the CIMS inlet for the same periods.

Text S5 The measurement duty cycle of CIMS

As detailed in Section 2.2, the PolyU-CIMS was configured to sequentially measure HO_2 , H_2SO_4 , and OH within each hour during the field study, corresponding to changes in the injection gases. Table S3 outlines the hourly measurement schedule and associated injection gases, while Figure S10 illustrates the 1-hour duty cycle. Table S5 Duty cycle and injection gases for targeted chemical analysis.

Purpose	Measure ment Mode	Signal 97 Label	Chemicals injected to the sample flow through different injectors		Duty time (s)	Repeat times	Total Duration (mins)
			Front Injectors	Rear Injectors			
HO ₂ *	SIG	S _{97NO}	NO, N _{2(p)} , SO ₂	Sca, Sca _(p)	60	3	6
	BKG	S _{97NOSca}	NO, Sca _(p) , SO ₂	Sca, N _{2(p)}	60		
H ₂ SO ₄	SIG	S _{97w/o}	-	Sca, Sca _(p)	60	2	4
	BKG	S _{97w/o}	-	Sca, N _{2(p)}	60		
OH	SIG	S _{97SO2}	N _{2(p)} , SO ₂	Sca, Sca _(p)	60	3	50
	BKG	S _{97SO2Sca}	Sca _(p) , SO ₂	Sca, N _{2(p)}	60		

1

Notes:

Front and Rear Injectors - The injector pairs as demonstrated in Figure

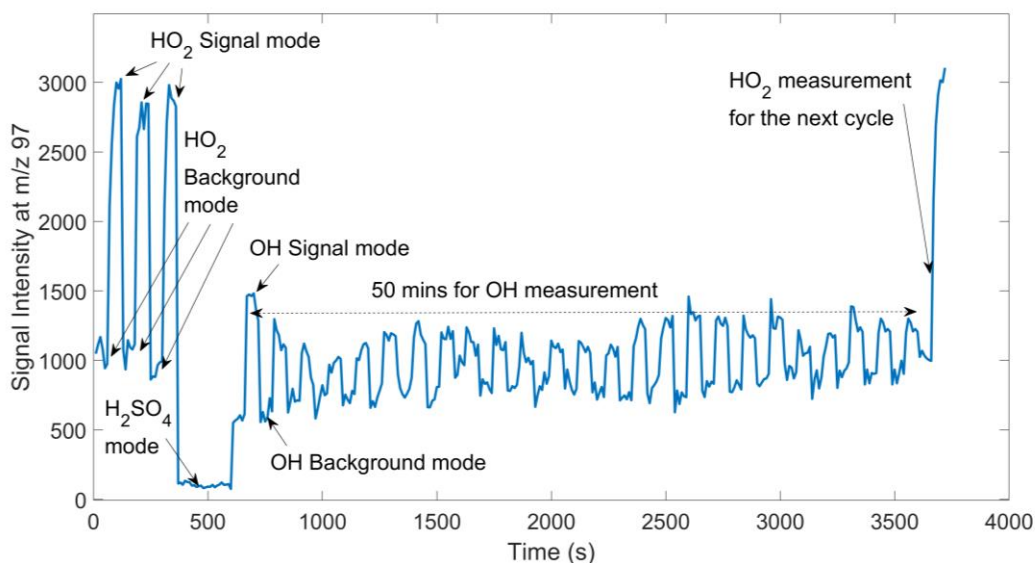
S2

SIG & BKG – the signal and background modes.

Sca - scavenger gas, C₃F₆ in this study.

Sca_(p) - scavenger gas added through the pulsed flow.

N_{2(p)}- nitrogen gas, added through the pulsed flow.



2

3 Figure S10 Variation of signal intensity at m/z 97 during a 1-hour duty cycle of CIMS
4 measurement.

5

6 **Reference**

7 Bottorff, B., Lew, M. M., Woo, Y., Rickly, P., Rollings, M. D., Deming, B., Anderson,

1 D. C., Wood, E., Alwe, H. D., Millet, D. B., Weinheimer, A., Tyndall, G., Ortega, J.,
 2 Dusanter, S., Leonardis, T., Flynn, J., Erickson, M., Alvarez, S., Rivera-Rios, J. C.,
 3 Shutter, J. D., Keutsch, F., Helmig, D., Wang, W., Allen, H. M., Slade, J. H., Shepson,
 4 P. B., Bertman, S., and Stevens, P. S.: OH, HO₂, and RO₂ radical chemistry in a rural
 5 forest environment: measurements, model comparisons, and evidence of a missing
 6 radical sink, *Atmospheric Chemistry and Physics*, 23, 10287–10311,
 7 <https://doi.org/10.5194/acp-23-10287-2023>, 2023.

8 Cho, C., Fuchs, H., Hofzumahaus, A., Holland, F., Bloss, W. J., Bohn, B., Dorn, H.-P.,
 9 Glowania, M., Hohaus, T., Liu, L., Monks, P. S., Niether, D., Rohrer, F., Sommariva,
 10 R., Tan, Z., Tillmann, R., Kiendler-Scharr, A., Wahner, A., and Novelli, A.:
 11 Experimental chemical budgets of OH, HO₂, and RO₂ radicals in rural air in western
 12 Germany during the JULIAC campaign 2019, *Atmospheric Chemistry and Physics*, 23,
 13 2003–2033, <https://doi.org/10.5194/acp-23-2003-2023>, 2023.

14 Edwards, G. D., Cantrell, C. A., Stephens, S., Hill, B., Goyea, O., Shetter, R. E.,
 15 Mauldin, R. L., Kosciuch, E., Tanner, D. J., and Eisele, F. L.: Chemical Ionization Mass
 16 Spectrometer Instrument for the Measurement of Tropospheric HO₂ and RO₂, *Anal.*
 17 *Chem.*, 75, 5317–5327, <https://doi.org/10.1021/ac034402b>, 2003.

18 Feiner, P. A., Brune, W. H., Miller, D. O., Zhang, L., Cohen, R. C., Romer, P. S.,
 19 Goldstein, A. H., Keutsch, F. N., Skog, K. M., Wennberg, P. O., Nguyen, T. B., Teng,
 20 A. P., DeGouw, J., Koss, A., Wild, R. J., Brown, S. S., Guenther, A., Edgerton, E.,
 21 Baumann, K., and Fry, J. L.: Testing Atmospheric Oxidation in an Alabama Forest,
 22 *Journal of the Atmospheric Sciences*, 73, 4699–4710, [https://doi.org/10.1175/JAS-D-](https://doi.org/10.1175/JAS-D-16-0044.1)
 23 16-0044.1, 2016.

24 Fuchs, H., Acir, I.-H., Bohn, B., Brauers, T., Dorn, H.-P., Häseler, R., Hofzumahaus, A.,
 25 Holland, F., Kaminski, M., Li, X., Lu, K., Lutz, A., Nehr, S., Rohrer, F., Tillmann, R.,
 26 Wegener, R., and Wahner, A.: OH regeneration from methacrolein oxidation
 27 investigated in the atmosphere simulation chamber SAPHIR, *Atmos. Chem. Phys.*, 14,
 28 7895–7908, <https://doi.org/10.5194/acp-14-7895-2014>, 2014.

1 Griffith, S. M., Hansen, R. F., Dusanter, S., Michoud, V., Gilman, J. B., Kuster, W. C.,
2 Veres, P. R., Graus, M., Gouw, J. A., Roberts, J., Young, C., Washenfelder, R., Brown,
3 S. S., Thalman, R., Waxman, E., Volkamer, R., Tsai, C., Stutz, J., Flynn, J. H., Grossberg,
4 N., Lefer, B., Alvarez, S. L., Rappenglueck, B., Mielke, L. H., Osthoff, H. D., and
5 Stevens, P. S.: Measurements of hydroxyl and hydroperoxy radicals during CalNex-LA:
6 Model comparisons and radical budgets, *J. Geophys. Res. Atmos.*, 121, 4211–4232,
7 <https://doi.org/10.1002/2015JD024358>, 2016.

8 Jeong, D., Seco, R., Emmons, L., Schwantes, R., Liu, Y., McKinney, K. A., Martin, S.
9 T., Keutsch, F. N., Gu, D., Guenther, A. B., Vega, O., Tota, J., Souza, R. A. F.,
10 Springston, S. R., Watson, T. B., and Kim, S.: Reconciling Observed and Predicted
11 Tropical Rainforest OH Concentrations, *JGR Atmospheres*, 127,
12 <https://doi.org/10.1029/2020JD032901>, 2022.

13 Lew, M. M., Rickly, P. S., Bottorff, B. P., Reidy, E., Sklaveniti, S., Léonardis, T., Locoge,
14 N., Dusanter, S., Kundu, S., Wood, E., and Stevens, P. S.: OH and HO₂ radical chemistry
15 in a midlatitude forest: measurements and model comparisons, *Atmospheric Chemistry
16 and Physics*, 20, 9209–9230, <https://doi.org/10.5194/acp-20-9209-2020>, 2020.

17 Ma, X., Tan, Z., Lu, K., Yang, X., Liu, Y., Li, S., Li, X., Chen, S., Novelli, A., Cho, C.,
18 Zeng, L., Wahner, A., and Zhang, Y.: Winter photochemistry in Beijing: Observation
19 and model simulation of OH and HO₂ radicals at an urban site, *Science of The Total
20 Environment*, 685, 85–95, <https://doi.org/10.1016/j.scitotenv.2019.05.329>, 2019.

21 Mallik, C., Tomsche, L., Bourtsoukidis, E., Crowley, J. N., Derstroff, B., Fischer, H.,
22 Hafermann, S., Hüser, I., Javed, U., Keßel, S., Lelieveld, J., Martinez, M., Meusel, H.,
23 Novelli, A., Phillips, G. J., Pozzer, A., Reiffs, A., Sander, R., Taraborrelli, D., Sauvage,
24 C., Schuladen, J., Su, H., Williams, J., and Harder, H.: Oxidation processes in the
25 eastern Mediterranean atmosphere: evidence from the modelling of HO_x measurements
26 over Cyprus, *Atmospheric Chemistry and Physics*, 18, 10825–10847,
27 <https://doi.org/10.5194/acp-18-10825-2018>, 2018.

28 Mo, Z., Shao, M., and Lu, S.: Compilation of a source profile database for hydrocarbon

1 and OVOC emissions in China, *Atmos. Environ.*, 143, 209–217,
 2 <https://doi.org/10.1016/j.atmosenv.2016.08.025>, 2016.

3 Sanchez, D., Jeong, D., Seco, R., Wrangham, I., Park, J.-H., Brune, W. H., Koss, A.,
 4 Gilman, J., de Gouw, J., Misztal, P., Goldstein, A., Baumann, K., Wennberg, P. O.,
 5 Keutsch, F. N., Guenther, A., and Kim, S.: Intercomparison of OH and OH reactivity
 6 measurements in a high isoprene and low NO environment during the Southern Oxidant
 7 and Aerosol Study (SOAS), *Atmospheric Environment*, 174, 227–236,
 8 <https://doi.org/10.1016/j.atmosenv.2017.10.056>, 2018.

9 Sjostedt, S. J., Huey, L. G., Tanner, D. J., Peischl, J., Chen, G., Dibb, J. E., Lefer, B.,
 10 Hutterli, M. A., Beyersdorf, A. J., Blake, N. J., Blake, D. R., Sueper, D., Ryerson, T.,
 11 Burkhardt, J., and Stohl, A.: Observations of hydroxyl and the sum of peroxy radicals at
 12 Summit, Greenland during summer 2003, *Atmospheric Environment*, 41, 5122–5137,
 13 <https://doi.org/10.1016/j.atmosenv.2006.06.065>, 2007.

14 Slater, E. J., Whalley, L. K., Woodward-Massey, R., Ye, C., Lee, J. D., Squires, F.,
 15 Hopkins, J. R., Dunmore, R. E., Shaw, M., Hamilton, J. F., Lewis, A. C., Crilley, L. R.,
 16 Kramer, L., Bloss, W., Vu, T., Sun, Y., Xu, W., Yue, S., Ren, L., Acton, W. J. F., Hewitt,
 17 C. N., Wang, X., Fu, P., and Heard, D. E.: Elevated levels of OH observed in haze events
 18 during wintertime in central Beijing, *Atmospheric Chemistry and Physics*, 20, 14847–
 19 14871, <https://doi.org/10.5194/acp-20-14847-2020>, 2020.

20 Tan, Z., Fuchs, H., Lu, K., Hofzumahaus, A., Bohn, B., Broch, S., Dong, H., Gomm, S.,
 21 Häsel, R., He, L., Holland, F., Li, X., Liu, Y., Lu, S., Rohrer, F., Shao, M., Wang, B.,
 22 Wang, M., Wu, Y., Zeng, L., Zhang, Y., Wahner, A., and Zhang, Y.: Radical chemistry
 23 at a rural site (Wangdu) in the North China Plain: observation and model calculations
 24 of OH, HO₂ and RO₂ radicals, *Atmospheric Chemistry and Physics*, 17, 663–690,
 25 <https://doi.org/10.5194/acp-17-663-2017>, 2017.

26 Tan, Z., Rohrer, F., Lu, K., Ma, X., Bohn, B., Broch, S., Dong, H., Fuchs, H., Gkatzelis,
 27 G. I., Hofzumahaus, A., Holland, F., Li, X., Liu, Y., Liu, Y., Novelli, A., Shao, M., Wang,

1 H., Wu, Y., Zeng, L., Hu, M., Kiendler-Scharr, A., Wahner, A., and Zhang, Y.:
 2 Wintertime photochemistry in Beijing: observations of RO_x radical concentrations in
 3 the North China Plain during the BEST-ONE campaign, *Atmospheric Chemistry and*
 4 *Physics*, 18, 12391–12411, <https://doi.org/10.5194/acp-18-12391-2018>, 2018.

5 Tan, Z., Lu, K., Jiang, M., Su, R., Wang, H., Lou, S., Fu, Q., Zhai, C., Tan, Q., Yue, D.,
 6 Chen, D., Wang, Z., Xie, S., Zeng, L., and Zhang, Y.: Daytime atmospheric oxidation
 7 capacity in four Chinese megacities during the photochemically polluted season: a case
 8 study based on box model simulation, *Atmos. Chem. Phys.*, 19, 3493–3513,
 9 <https://doi.org/10.5194/acp-19-3493-2019>, 2019.

10 Tan, Z., Ma, X., Lu, K., Jiang, M., Zou, Q., Wang, H., Zeng, L., and Zhang, Y.: Direct
 11 evidence of local photochemical production driven ozone episode in Beijing: A case
 12 study, *Science of The Total Environment*, 800, 148868,
 13 <https://doi.org/10.1016/j.scitotenv.2021.148868>, 2021.

14 Wang, G., Iradukunda, Y., Shi, G., Sanga, P., Niu, X., and Wu, Z.: Hydroxyl,
 15 hydroperoxyl free radicals determination methods in atmosphere and troposphere,
 16 *Journal of Environmental Sciences*, 99, 324–335,
 17 <https://doi.org/10.1016/j.jes.2020.06.038>, 2021.

18 Whalley, L. K., Stone, D., Dunmore, R., Hamilton, J., Hopkins, J. R., Lee, J. D., Lewis,
 19 A. C., Williams, P., Kleffmann, J., Laufs, S., Woodward-Massey, R., and Heard, D. E.:
 20 Understanding in situ ozone production in the summertime through radical observations
 21 and modelling studies during the Clean air for London project (ClearfLo), *Atmos. Chem.*
 22 *Phys.*, 18, 2547–2571, <https://doi.org/10.5194/acp-18-2547-2018>, 2018.

23 Whalley, L. K., Slater, E. J., Woodward-Massey, R., Ye, C., Lee, J. D., Squires, F.,
 24 Hopkins, J. R., Dunmore, R. E., Shaw, M., Hamilton, J. F., Lewis, A. C., Mehra, A.,
 25 Worrall, S. D., Bacak, A., Bannan, T. J., Coe, H., Percival, C. J., Ouyang, B., Jones, R.
 26 L., Crilley, L. R., Kramer, L. J., Bloss, W. J., Vu, T., Kotthaus, S., Grimmond, S., Sun,
 27 Y., Xu, W., Yue, S., Ren, L., Acton, W. J. F., Hewitt, C. N., Wang, X., Fu, P., and Heard,
 28 D. E.: Evaluating the sensitivity of radical chemistry and ozone formation to ambient

1 VOCs and NO_x in Beijing, *Atmospheric Chemistry and Physics*, 21, 2125–2147,
2 <https://doi.org/10.5194/acp-21-2125-2021>, 2021.

3 Woodward-Massey, R., Slater, E. J., Alen, J., Ingham, T., Cryer, D. R., Stimpson, L. M.,
4 Ye, C., Seakins, P. W., Whalley, L. K., and Heard, D. E.: Implementation of a chemical
5 background method for atmospheric OH measurements by laser-induced fluorescence:
6 characterisation and observations from the UK and China, *Atmospheric Measurement*
7 *Techniques*, 13, 3119–3146, <https://doi.org/10.5194/amt-13-3119-2020>, 2020.

8 Yang, X.: A Review of the Direct Measurement of Total OH Reactivity: Ambient Air
9 and Vehicular Emission, *Sustainability*, 15, 16246,
10 <https://doi.org/10.3390/su152316246>, 2023.

11 Yao, Z., Shen, X., Ye, Y., Cao, X., Jiang, X., Zhang, Y., and He, K.: On-road emission
12 characteristics of VOCs from diesel trucks in Beijing, China, *Atmos. Environ.*, 103, 87–
13 93, <https://doi.org/10.1016/j.atmosenv.2014.12.028>, 2015.

14 Zhang, G., Hu, R., Xie, P., Lu, K., Lou, S., Liu, X., Li, X., Wang, F., Wang, Y., Yang,
15 X., Cai, H., Wang, Y., and Liu, W.: Intercomparison of OH radical measurement in a
16 complex atmosphere in Chengdu, China, *Science of The Total Environment*, 838,
17 155924, <https://doi.org/10.1016/j.scitotenv.2022.155924>, 2022a.

18 Zhang, G., Hu, R., Xie, P., Lou, S., Wang, F., Wang, Y., Qin, M., Li, X., Liu, X., Wang,
19 Y., and Liu, W.: Observation and simulation of HO_x radicals in an urban area in
20 Shanghai, China, *Science of The Total Environment*, 810, 152275,
21 <https://doi.org/10.1016/j.scitotenv.2021.152275>, 2022b.

22 Zou, Z., Chen, Q., Xia, M., Yuan, Q., Chen, Y., Wang, Y., Xiong, E., Wang, Z., and
23 Wang, T.: OH measurements in the coastal atmosphere of South China: possible
24 missing OH sinks in aged air masses, *Atmospheric Chemistry and Physics*, 23, 7057–
25 7074, <https://doi.org/10.5194/acp-23-7057-2023>, 2023.

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