



1 Evaluating the accuracy of downwind methods for quantifying point source emissions

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5 **Abstract.** The accurate reporting of methane (CH<sub>4</sub>) emissions from point sources, such as fugitive leaks from oil and  
6 gas infrastructure, is important for evaluating climate change impacts, assessing CH<sub>4</sub> fees for regulatory programs,  
7 and validating methane intensity in differentiated gas programs. Currently, there are disagreements between emissions  
8 reported by different quantification techniques for the same sources. It has been suggested that downwind CH<sub>4</sub>  
9 quantification methods using CH<sub>4</sub> measurements on the fence-line of production facilities could be used to generate  
10 emission estimates from oil and gas operations at the site level, but it is currently unclear how accurate the quantified  
11 emissions are. To investigate model accuracy, this study uses fence-line simulated data collected during controlled  
12 release experiments as input for eddy covariance, aerodynamic flux gradient and the Gaussian plume inverse methods  
13 in a range of atmospheric conditions. The results show that both the eddy covariance and aerodynamic flux gradient  
14 methods underestimated emissions in all experiments. Although calculated emissions had significant uncertainty, the  
15 Gaussian plume inversion method performed better. The uncertainty was found to have no significant correlation  
16 with most measurement variables (i.e. downwind measurement distance, wind speed, atmospheric stability, or  
17 emission height), which indicates that the Gaussian method can randomly either underestimate or overestimate  
18 emissions. For eddy covariance, downwind measurement distance and percent error had negative correlation  
19 indicating that far away emissions sources were likely underestimated or be undetected. The study concludes that  
20 using fence-line measurement data as input to eddy covariance, aerodynamic flux gradient or Gaussian plume inverse  
21 method to quantify CH<sub>4</sub> emissions from an oil and gas production site is unlikely to generate representative emission  
22 estimates.

## 23 **1 Introduction**

24 Methane (CH<sub>4</sub>), the primary component of natural gas (NG), is a potent greenhouse gas with a global warming  
25 potential of 27 carbon dioxide (CO<sub>2</sub>) equivalent over 100 years (US EPA, 2016). Methane emissions reduction is a  
26 key part of global initiatives to reduce climate change (Chung, 2021). The 2021 Global Methane Assessment by the  
27 Climate and Clean Air Coalitions (CCAC, 2024) and the United Nations Environment Programme (UN Environment  
28 Programme, 2024) state that reducing CH<sub>4</sub> emissions from anthropogenic sources by 45% in 2030 would result in  
29 avoiding a global atmospheric temperature increase of 0.3°C in 2045 (Chung, 2021). Such measures would align with  
30 the Paris Agreement goal of limiting global temperature rise to 1.5°C by 2030 (United Nations Climate Change, 2015).  
31 The US is one of the countries that reports its total greenhouse gas emissions to the Intergovernmental Panel on  
32 Climate Change as part of the Paris Agreement (United Nations Climate Change, 2015).  
33 Currently, the amount of CH<sub>4</sub> emitted from US oil and gas production is calculated by the US Environmental  
34 Protection Agency (EPA) using a bottom-up inventory approach. The inventory approach multiplies emission factors



35 (CH<sub>4</sub> emissions per equipment e.g., separator or emissions per event e.g., liquid unloading) by activity factors (total  
36 number of pieces of equipment or events (OAR US EPA, 2023)). This quantification approach has several  
37 shortcomings, including: 1. It separately calculates CH<sub>4</sub> emissions from natural gas and petroleum systems, which  
38 practically are not independent systems, and can result in bias based on changes in gas to oil ratios throughout a basin  
39 (Riddick et al., 2024a); 2. Some emission factors used are outdated (Riddick et al., 2024b) and others do not account  
40 for the temporal and spatial variation in emissions (Riddick and Mauzerall, 2023); and 3. Emission factors do not  
41 account for the long-tail distributions in emissions distributions (Riddick et al., 2024b). Recently, mechanistic models,  
42 such as the Colorado State University's Mechanistic Air Emissions Simulator (MAES), have been developed to  
43 address shortcomings in bottom-up CH<sub>4</sub> reporting (Colorado State University, 2021) but these still depend on direct  
44 measurements to inform emission factors.

45 Top-down methods, including using aircraft and satellites, can also be used to infer emissions. For example, Carbon  
46 Mapper satellites can locate and quantify CH<sub>4</sub> emissions using absorption spectra taken from space (Carbon Mapper,  
47 2024). However, these survey methods only quantify emissions over a very short period of time (< 10 s) and  
48 observations are typically made during the day which can often coincide with maintenance activities that can bias  
49 emissions and result in overestimation (Riddick et al., 2024a; Zimmerle et al., 2024). Additionally, different top-down  
50 technologies measuring the same source have disagreed in their reported emissions which has called into question the  
51 credibility of these methods (Brown et al., 2023; Conrad et al., 2023). As a result, ensuring accuracy in models and  
52 technologies used in CH<sub>4</sub> emissions quantification has been a complex issue.

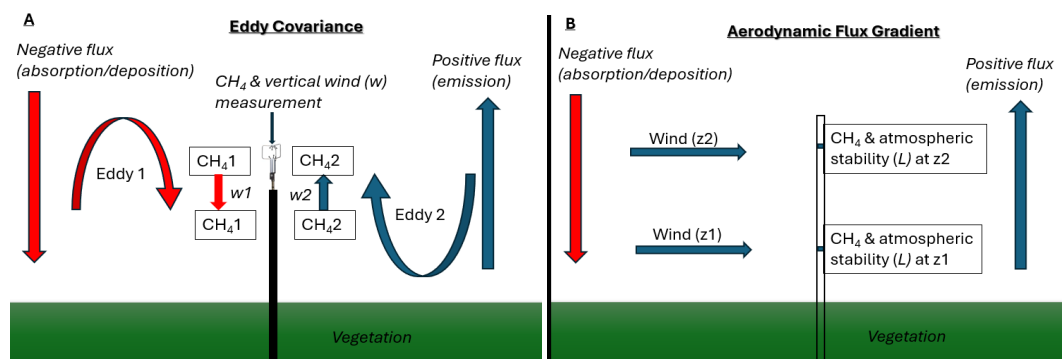
53 The accurate reporting of CH<sub>4</sub> from fugitive emissions at oil and gas production sites is important for evaluating  
54 potential effects on climate change, correctly assessing CH<sub>4</sub> fees on companies as part of the Methane Emissions  
55 Reduction Program created under the 2022 Inflation Reduction Act (OA US EPA, 2023), and validating CH<sub>4</sub> content  
56 of reported differentiated gas composition where NG companies differentiate their market products based on the  
57 environmental impact (CO2EFFICIENT, 2022). Direct measurements have been recommended to augment/update  
58 emissions factors used in bottom-up inventories and for better understanding temporal/spatial variability of emissions  
59 (Riddick et al., 2024). Downwind methods are widely used to directly measure CH<sub>4</sub> emissions from area and point  
60 sources at site/basin levels due to their low cost and wide coverage within a short time (Caulton et al., 2018;  
61 Heimburger et al., 2017; Riddick et al., 2020, 2022a; Sonderfeld et al., 2017). Commonly used downwind  
62 quantification methods include the Gaussian plume inversion method, eddy covariance, backward Lagrangian  
63 stochastic models, aerodynamic flux gradient, mass balance method, the EPA Other Test Method (OTM 33) and the  
64 Gaussian puff modelling approach (Denmead, 2008; Edie et al., 2020; Foster-Wittig et al., 2015; Jia et al., 2023; Kamp  
65 et al., 2020; Nemitz et al., 2018; Shaw et al., 2021).

66 Currently, fence-line methods are used to detect, localize and quantify emissions. This approach uses point sensors  
67 fixed to the fence-line of the production site and emissions detected when the measured concentration exceeds a  
68 threshold, localized by triangulating multiple detections and quantified using a simple dispersion modelling  
69 framework, usually based on a Gaussian plume approach (Bell et al., 2023; Day et al., 2024; Jia et al., 2023; Riddick  
70 et al., 2022a). The detection and localization of simulated fugitive emission have been successful, with controlled  
71 release testing against point sensors and scanning/imaging solutions reporting a 90% probability of detection for



72 emission of between 3.9 and 18.2 kg CH<sub>4</sub> h<sup>-1</sup> (Ilonze et al., 2024). Major shortcomings have been identified using a  
73 fence-line approach with quantified emissions reported at between a factor of 0.2 to 42 times for emissions between  
74 0.1 and 1 kg CH<sub>4</sub> h<sup>-1</sup>, and between 0.08 and 18 times for emissions greater than 1 kg CH<sub>4</sub> h<sup>-1</sup> (Ilonze et al., 2024). As  
75 a result, questions have arisen if other approaches, such as the eddy covariance or aerodynamic flux gradient would  
76 generate more accurate results. These methods have been suggested as they have been used to quantify emissions  
77 from other sectors, i.e. agriculture (Denmead, 2008; Morin, 2019) and landfills (Xu et al., 2014), have been used to  
78 quantify emissions in large downwind areas (Vogel et al., 2024), and quantification does not require assumptions  
79 made on downwind dispersion coefficients or micrometeorology that are often required for dispersion modelling  
80 (Denmead, 2008).

81 Due to interest in using a subset of these methods to quantify emissions from oil and production sites, this study will  
82 evaluate the quantification accuracy of the eddy covariance, aerodynamic flux gradient, and Gaussian plume inverse  
83 methods. Eddy covariance is a vertical flux gradient measurement that measures CH<sub>4</sub> emissions based on the  
84 covariance between CH<sub>4</sub> concentrations measured using a fast-response analyzer (> 10 Hz) and vertical wind vector  
85 measured by a fast-response sonic anemometer (>10 Hz) (Figure 1A; Morin, 2019). It is typically implemented over  
86 long homogenous fetches where eddy mixing scale is a small fraction of the distance from the site providing more  
87 predictable vertical transport. The aerodynamic flux gradient method quantifies CH<sub>4</sub> emissions from a source by  
88 comparing CH<sub>4</sub> concentrations at two heights (Figure 1B; Querino et al., 2011). The Gaussian Plume Inverse method  
89 calculates CH<sub>4</sub> mole fraction at a point in space (x, y, z) as a function of the downwind distance, perpendicular distance  
90 (crosswind), mean wind speed and atmospheric stability (Jia et al., 2023; Riddick et al., 2022b). These approaches  
91 were developed to quantify emissions from single-point or area emission sources and have not been tested against a  
92 controlled release to evaluate their quantification performance. The aerodynamic flux gradient and eddy covariance,  
93 for example, have been used to measure trace gas, e.g., nitrogen oxide and carbon dioxide, fluxes from large croplands  
94 (Kamp et al., 2020).



95  
96 Figure 1: Illustrations of eddy covariance (A) and flux gradient measurements (B) where CH<sub>4</sub> is methane  
97 concentrations,  $w$  is the vertical wind speed,  $L$  is the Monin-Obukhov length (measure of atmospheric stability), and  
98  $z$  is the measurement height.

99 The Gaussian plume inversion method has been used to quantify emissions from oil and gas production sites (Caulton  
100 et al., 2014; Riddick et al., 2022b) but it assumes a homogenous, steady state flow, uniform dispersion of gas in an

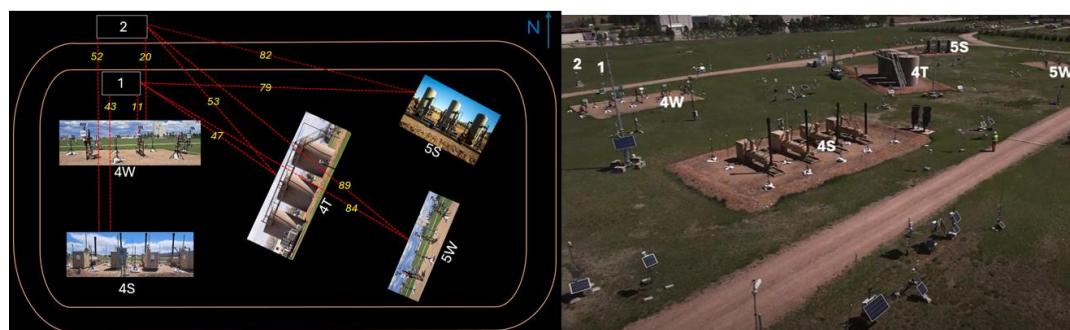


101 open area free of obstructions (Hutchinson et al., 2017). Oil and gas emissions are characterized by intermittent, non-  
102 uniform, single or multiple point source emissions, varying in leak size, location, height and distance between the  
103 source and sensor, and are typically in complex aerodynamic environments (i.e. not flat). The need for accurate CH<sub>4</sub>  
104 quantification and reporting necessitates evaluating the performance of these downwind quantification approaches in  
105 different controlled release and characterized meteorological conditions, to ensure credibility.  
106 This study aims to investigate the performance of these methods in quantifying emissions for known gas release rates  
107 and evaluating uncertainties that could result in incorrect CH<sub>4</sub> reporting. Specifically, the study will (1) evaluate the  
108 overall quantification accuracy of eddy covariance, aerodynamic flux gradient, and the Gaussian plume inverse  
109 method in quantifying single-point and multi-point emissions that simulate oil and gas emissions, (2) evaluate the  
110 probability of these models quantifying within a defined range (i.e.  $\pm 30\%$ ), and (3) investigate which variables have  
111 the largest effect on quantification uncertainty.

## 112 **2 Methods**

### 113 **2.1 Experimental Setup**

114 Controlled release experiments were conducted at the Colorado State University's Methane Emissions Technology  
115 Evaluation Center (METEC) in Fort Collins, CO, USA, between February 8, and March 27, 2024. The weather  
116 conditions during the test period were mostly sunny but precipitation was also observed (32 sunny, 7 snowy, 12 rainy,  
117 7 cloudy and 1 foggy day; Supplementary Information Section 1). Wind speeds were between 0 and 25 m s<sup>-1</sup> and  
118 temperatures ranged between -15 and +19 °C (Supplementary Information Section 1). Two stationary masts holding  
119 the instrumentation were setup on the North-West corner of METEC to take advantage of the predominant wind  
120 direction, avoid the largest aerodynamic obstructions and to simulate the likely placement of a fence line instrument  
121 (Figure 2; Day et al., 2024; Riddick et al., 2022a). Fenceline sensors are typically placed within the oil and gas  
122 perimeter (~30 m) (Riddick et al., 2022a). This study collected data for both close and far away releases, distances  
123 between 9 and 94 m.

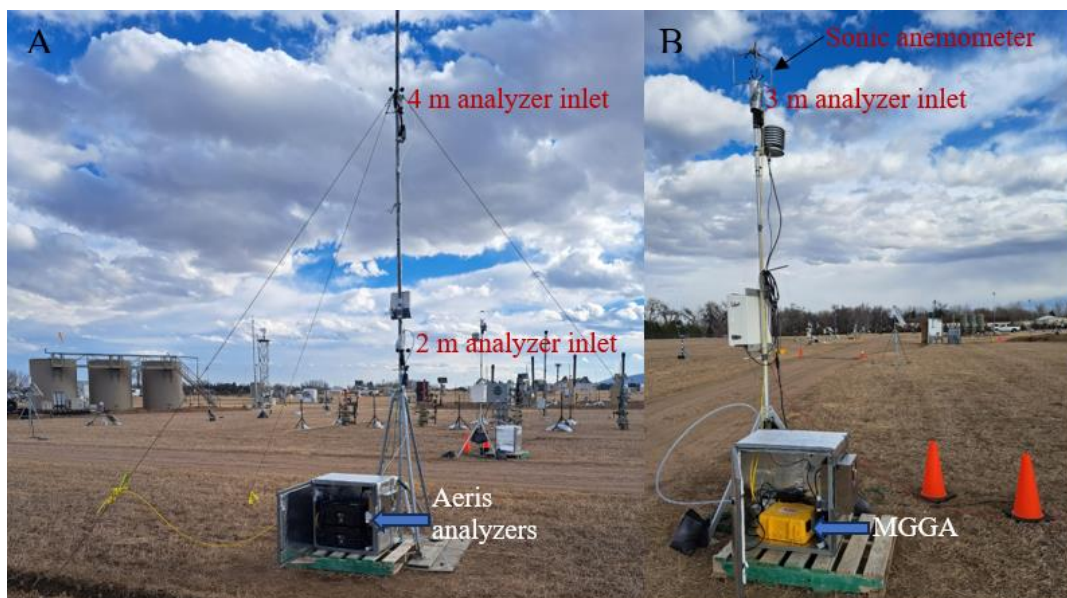


124  
125 Figure 2: Left pane: Map illustration of major pieces of equipment and the measurements points at Colorado State  
126 University's Methane Emissions Technology Evaluation Center (METEC) in Fort Collins, CO, USA. 4S denotes the  
127 location of horizontal separators, 4W are well heads, 4T are tanks, 5S are vertical separators and 5W are well heads.  
128 1 is the measurement point for the Microportable Greenhouse Gas Analyzer and 2 is the measurement point for the  
129 AERIS analyzers. The red dotted lines with yellow numbers show the average distances (meters) between emission  
130 equipment and measurement points. Right pane: Image of METEC showing relative heights of equipment ("METEC  
131 | Colorado State University," 2024).  
132

133 To calculate emissions using the aerodynamic flux gradient approach, two sampling inlets were mounted at 2 and 4  
134 m heights on mast 2 and connected to the inlets of two AERIS (Hayward, CA, USA) MIRA Ultra Series analyzers  
135 (Figure 3A). The analyzers were housed in a temperature-controlled unit and sampled at 5 Hz. Data from the 2 m  
136 analyzer were also used as input for the Gaussian Plume Inverse method analysis. To collect CH<sub>4</sub> concentration data  
137 for the eddy covariance method, the inlet tubing of the ABB (Zurich, Switzerland) GLA131 Series Microportable  
138 Greenhouse Gas Analyzer (MGGA) sampling at 10 Hz was collocated with an R. M. Young (Traverse City, MI, USA)  
139 81000 sonic anemometer (R.M. Young Company, 2023) which measured micrometeorology at 10 Hz, 3 m height  
140 above ground level on mast 1 (Figure 3B).



141



142

143 Figure 3: A is the aerodynamic flux gradient and Gaussian plume inverse sampling points and B is the eddy covariance  
144 sampling point. The two sampling points are 9.4 m apart.

## 145 2.2 Controlled Methane Releases

146 At METEC, natural gas of known CH<sub>4</sub> content was released from above-ground emission points attached to equipment  
147 typically present in an oil and gas facility (tanks, separators and well pads). The gas release rates ranged between  
148 0.005 kg h<sup>-1</sup> and 8.5 kg h<sup>-1</sup>, and the release durations ranged from 10 seconds to 8 hours, simulating both fugitive and  
149 large emission events. The releases were run both during the day and night. The distance from the release points to  
150 the measurement points ranged between 9 and 94 m, and emission heights were between 0.4 and 6.9 m (Figure 2).  
151 Emission points simulate the realistic size and locations of typical emission from components such as the thief hatches,  
152 pressure relief valves, flanges, bradenheads, pressure transducers, Kimray valves and vents. The releases included  
153 both single point emissions (single releases) and multi-point emission events (multiple simultaneous releases).

## 154 2.3 Data Processing

155 Methane concentrations data from the analyzers were aggregated with the meteorological data from the sonic  
156 anemometer. For aerodynamic flux gradient and Gaussian plume inverse method data were averaged to 1 Hz, for the  
157 eddy covariance the raw CH<sub>4</sub> 10 Hz data was used. The aggregated meteorological-concentration data were then  
158 merged with METEC's release data and metadata, and event tables created. The meteorological-concentration-release  
159 event data were then separated into single-point and multi-point events. The event tables were split into 20-minute  
160 emission events for aerodynamic flux gradient and Gaussian plume inverse method as they are dependent on  
161 atmospheric stability that is typically determined in time durations of 15 to 30 minutes. Shorter duration measurements  
162 (i.e. <15 minutes) may not represent the mean atmospheric state, while longer periods (> 30 minutes) may cause errors  
163 especially during rapid transitions in weather conditions (Crenna, 2006). 30-minute events were used for eddy



164 covariance processing following published typical averaging times of eddy covariance measurements (Nemitz et al.,  
165 2018), and its quantification is assumed to be independent of atmospheric stability (Denmead, 2008).  
166 For eddy covariance and aerodynamic flux gradient, Monin-Obukhov length ( $L$ ) was calculated as the measure of  
167 atmospheric stability for every 20 or 30-minute time period, depending on the method, using output from the sonic  
168 anemometer.  $L$  was calculated from the surface friction velocity ( $u_*$ ,  $\text{m s}^{-1}$ ), mean potential temperature ( $\theta$ , K), von  
169 Kármán's constant ( $k$ , 0.41), gravitational acceleration ( $g$ ,  $9.8 \text{ m s}^{-1}$ ) and the surface (kinematic) turbulent flux of  
170 sensible heat  $w'\theta'$  (Eq. 1 and 2) (Kljun et al., 2015; Stull, 1988).

$$L = - \frac{u_*^3 \theta}{k_v g w' \theta'} \quad (1)$$

171

$$u_* = \left[ (\overline{u'w'})^2 + (\overline{v'w'})^2 \right]^{1/4} \quad (2)$$

172 For the Gaussian method, atmospheric stability was calculated based on the EPA standard operating procedure for  
173 point source Gaussian method (US EPA, 2013). The average local wind stability class ( $pgi$ ) was calculated as the  
174 average of atmospheric stability determined using the standard deviation of the wind direction, and the stability  
175 calculated from turbulent intensity (ratio of the standard deviation of the wind speed to the average wind speed). The  
176 dispersion coefficients used for Gaussian quantification were extracted from the EPA operating procedure that  
177 provided coefficients for distances ranging from 1 to 200 m from source (US EPA, 2013).

178 The wind direction ( $WD$ ) and speed ( $WS$ ) were calculated from the wind vectors  $u$  and  $v$ , based on the manufacturer's  
179 configuration:  $+u$  values = wind from the east,  $+v$  values = wind from the north, and  $+w$  = updraft (Eq. 3 and 4).

180

$$WD = \text{mod}(90 - \text{atan2d}(v, u), 360) \quad (3)$$

$$WS = \sqrt{u^2 + v^2} \quad (4)$$

181 The bearing of each release point to the masts' location was calculated using the latitudes and longitudes of the release  
182 points provided in the METEC metadata. This bearing was used to determine when the masts were downwind of the  
183 release points during the 20/30-minute period. The models' quantification accuracies were tested in three downwind  
184 ranges:  $\pm 10^\circ$ ,  $\pm 20^\circ$ , and  $\pm 30^\circ$ . A mast was considered downwind when the wind direction was within the specified  
185 range for 30% of the 20/30-minute duration. Results for the 20-degree range are presented in the Results section, while  
186 the 10- and 30-degree results are included in the Supplementary Material. The 30% threshold was chosen to ensure  
187 sufficient data points for evaluating the models. The data were categorized into single release single emission (single  
188 emission at the site and the mast was downwind of the release point), multi release single emission (multiple emissions  
189 at the site level, but the mast was downwind of a single release point), and multi release multi emission (multiple  
190 emissions at the site level, but the mast was downwind of more than one release point).



## 191 2.4 Methane Emissions Quantification

### 192 2.4.1 Background Concentration

193 Background concentration was determined for each of the sensors to calculate CH<sub>4</sub> enhancement. Due to inherent  
194 variation in sensors that were used in this study, CH<sub>4</sub> background was calculated for each sensor separately. CH<sub>4</sub>  
195 background was calculated as the average of the lowest 5<sup>th</sup> percentile of all continuous concentration readings (US  
196 EPA, 2013). Methane enhancement was determined as CH<sub>4</sub> concentration measurement minus the background  
197 concentration measurement.

### 198 2.4.2 Eddy Covariance

199 Emissions were quantified using the eddy covariance method for all three emissions scenarios (single release single  
200 emission, multi release single emission and multi release multi emission). Methane flux ( $F$ , kg m<sup>-2</sup> s<sup>-1</sup>) was calculated  
201 as the covariance between the vertical wind speed ( $w$ , m s<sup>-1</sup>) and CH<sub>4</sub> enhancement ( $c$ , g m<sup>-3</sup>) over 30 minutes (Eq. 5;  
202 Denmead, 2008).

$$F = w'c' \quad (5)$$

### 203 2.4.3 Aerodynamic Flux Gradient

204 Aerodynamic flux gradient quantification was also tested in all three cases. Methane flux ( $F$ , kg m<sup>-2</sup> s<sup>-1</sup>) was calculated  
205 based on surface friction velocity ( $u_*$ , m s<sup>-1</sup>), von Kármán's constant ( $k_v$ , 0.41), the difference in the average CH<sub>4</sub>  
206 enhancement between the higher and lower height ( $g$ , m<sup>-3</sup>), natural log of the higher and lower height, and stability  
207 correction factors  $\Psi$  (Eq. 6; Denmead, 2008; Kamp et al., 2020).

$$F = \frac{u_* k_v (c_2 - c_1)}{\ln\left(\frac{z_2}{z_1}\right) - \Psi_{c,2} + \Psi_{c,1}} \quad (6)$$

### 208 2.4.4 Determining the Area of Vertical Flux Contribution

209 Eddy covariance and aerodynamic flux gradient measurements at a point (0, 0,  $z$ ) generate vertical fluxes in kg m<sup>-2</sup> s<sup>-1</sup>.  
210 In this study, these fluxes represent emissions from single-point or multi-point sources distributed over an area (m<sup>2</sup>).  
211 The Kljun et al. (2015) footprint model, was used to calculate footprint, and determine the area that contributed 80%  
212 ( $r = 80$ ,  $10 \leq r \leq 90$ ) of the vertical flux measured by the eddy covariance and aerodynamic flux gradient systems. In  
213 previous studies, 80% footprints have been used due to the difficulty of reproducing 90% of the sources under neutral  
214 and stable conditions, where footprints tend to be long. The difference between the 80% and 90% contours is typically  
215 excessively large, despite minimal flux contributions in that area (Rey-Sanchez et al., 2022). The Kljun et al. (2015)  
216 model calculates footprint as a function of effective height ( $z_m =$  sensor height ( $z$ ) – displacement height ( $m$ )),  
217 roughness length ( $z_o$ , m) / mean wind speed ( $u_{means}$ , m s<sup>-1</sup> - used in this study), height of the boundary layer ( $h$ , m),  
218 Obukhov length ( $L$ , m), standard deviation of the lateral velocity ( $\sigma_v$ , m s<sup>-1</sup>), and friction velocity ( $u^*$ , m s<sup>-1</sup>) (Kljun et  
219 al., 2015). The roughness sublayer in the model was set to 1 (footprint is calculated even if  $z_m$  is within the roughness  
220 layer). The area of vertical flux contribution was calculated as the polygon area covered by the contour. Due to the  
221 limitations of the flux footprint model for the measurement height and stability (Kljun et al., 2015), 20/30-minute files  
222 flagged by the footprint model when  $z_m/L < -15.5$ , were excluded from further analysis.





223 **2.4.5 Gaussian Plume Inverse Method**

224 The Gaussian plume inverse method was used to quantify single release single emission and multi release single  
 225 emission. The quantified emission ( $Q$ ,  $\text{kg h}^{-1}$ ) was calculated from the  $\text{CH}_4$  enhancement ( $X$ ,  $\text{g m}^{-3}$ ), wind speed ( $u$ ,  $\text{m}$   
 226  $\text{s}^{-1}$ ), horizontal dispersion coefficient ( $\sigma_y$ ,  $\text{m}$ ), vertical dispersion coefficient ( $\sigma_z$ ,  $\text{m}$ ), crosswind distance ( $y$ ,  $\text{m}$ ),  
 227 sampling height ( $z$ ,  $\text{m}$ ), emission height ( $h_s$ ,  $\text{m}$ ), and the height of the boundary layer (Equation 7; Riddick et al.,  
 228 2022b).

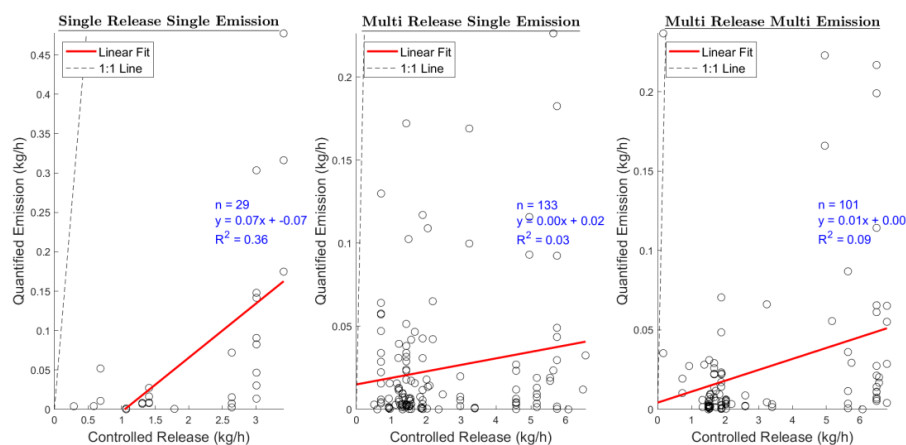
$$X(x, y, z) = \frac{Q}{2\pi u \sigma_y \sigma_z} e^{-\frac{y^2}{2\sigma_y^2}} \left( e^{-\frac{(z-h_s)^2}{2\sigma_z^2}} + e^{-\frac{(z+h_s)^2}{2\sigma_z^2}} + e^{-\frac{(z-2h+h_s)^2}{2\sigma_z^2}} + e^{-\frac{(z+2h-h_s)^2}{2\sigma_z^2}} + e^{-\frac{(z-2h-h_s)^2}{2\sigma_z^2}} \right) \quad (7)$$

229 **3 Results**

230 **3.1 Methane Emission Quantification**

231 **3.1.1 Eddy Covariance**

232 For stable, continuous 30-minute release events, emissions calculated using the eddy covariance method were an  
 233 underestimate for single release single emission, multi release single emission and multi release multi emission events  
 234 (Figure 4). All data points were below the 1:1 line. A plot of the quantified emission versus controlled release ( $\text{kg h}^{-1}$ )  
 235 did not show a linear correlation ( $R^2$  between 0.03 and 0.36), as all emissions were largely underestimated. The  
 236 eddy covariance method reported emissions of between 0 and 0.5  $\text{kg h}^{-1}$  overall, despite actual emissions being  
 237 between 0 and about 7  $\text{kg h}^{-1}$  (Figure 4). The underestimation was consistent across all downwind ranges, 10, 20 and  
 238 30 degrees (Supplementary Material Section 2.1).



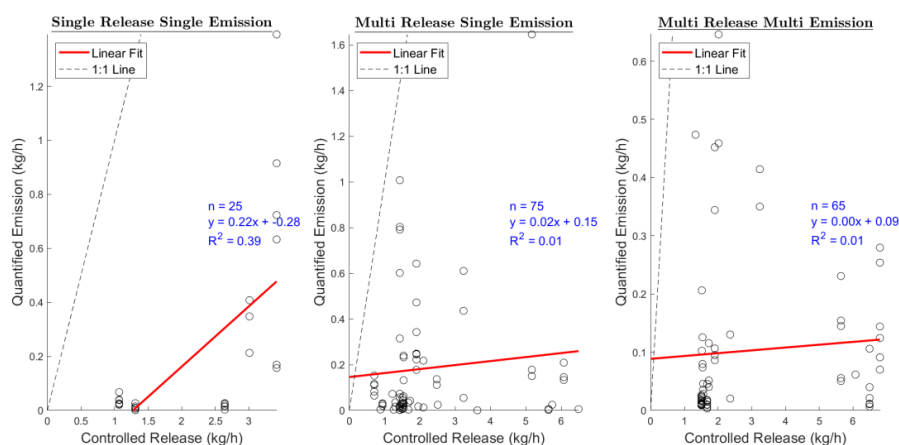
239

240 Figure 4: Quantified emission calculated using the eddy covariance method. Left pane shows a scatter plot of  
 241 quantified emission versus total controlled release for a single release at the site level and the mast was downwind of  
 242 the release point. Center pane shows a scatter plot of quantified emission versus total controlled release for multiple  
 243 releases at the site level, but the mast was downwind of a single release point. Right pane shows a scatter plot of  
 244 quantified emission versus total controlled release for multiple releases at the site level and the mast was downwind  
 245 of more than one release point. The dashed line represents the 1:1 line (points below the line were underestimated),  
 246 the red line is the linear regression fit of the data, and  $n$  is the number of data points.



### 247 3.1.2 Aerodynamic Flux Gradient

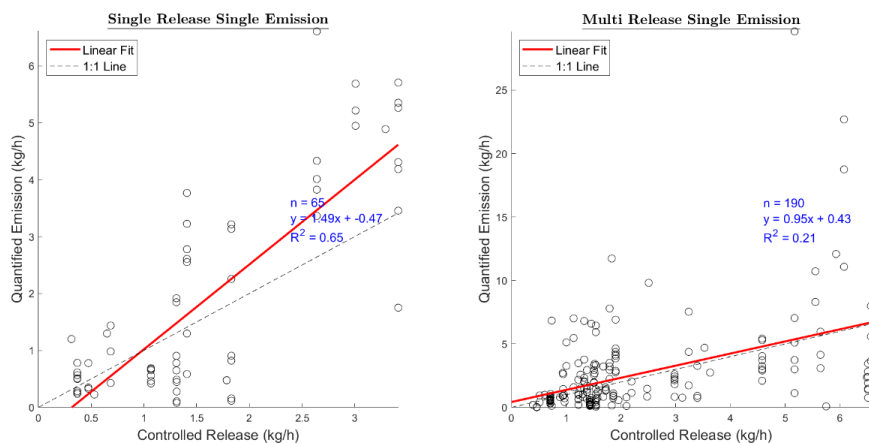
248 The aerodynamic flux gradient method also largely underestimated emissions for single release single emission, multi  
249 release single emission and multi release multi emission (Figure 5). A plot of quantified emission versus actual release  
250 did not show a linear relationship ( $R^2$  between 0.01 and 0.39), and most data points were below the 1:1 line (Figure  
251 5). The aerodynamic flux gradient quantified emissions were between 0 and about  $1.6 \text{ kg h}^{-1}$  despite actual emissions  
252 being between 0 and about  $7 \text{ kg h}^{-1}$  (Figure 5). The underestimation was also consistent across all downwind ranges,  
253 10, 20 and 30 degrees (Supplementary Material Section 2.2).



254  
255 Figure 5: Quantified emission calculated using the aerodynamic flux gradient method. Left pane shows a scatter plot  
256 of quantified emission versus total controlled release for a single release at the site level and the mast was downwind  
257 of the release point. Center pane shows a scatter plot of quantified emission versus total controlled release for  
258 multiple releases at the site level, but the mast was downwind of a single release point. Right pane shows a scatter  
259 plot of quantified emission versus total controlled release for multiple releases at the site level and the mast was  
260 downwind of more than one release point. The dashed line represents the 1:1 line (points below the line were  
261 underestimated), the red line is the linear regression fit of the data, and  $n$  is the number of data points.

### 262 3.1.3 Gaussian Plume Inverse Method

263 The Gaussian plume inverse method was tested for single release single emission and multi release single emission as  
264 the method is only used for single-point sources and preliminary results showed the method provided reasonable  
265 results within 20 degrees downwind range (Figure 6; Supplementary Material Section 1.3). For single release single  
266 emission, the method quantified emissions within a factor of 1.5 (Figure 6) and showed reasonably linear relationship  
267 ( $R^2$  of 0.65) (Figure 6). For multi release single emission, the gradient (m) of the linear regression was 0.95 and  $R^2$  of  
268 0.21. This suggests that the linear relationship cannot be well explained due to a random scatter of calculated  
269 emissions.



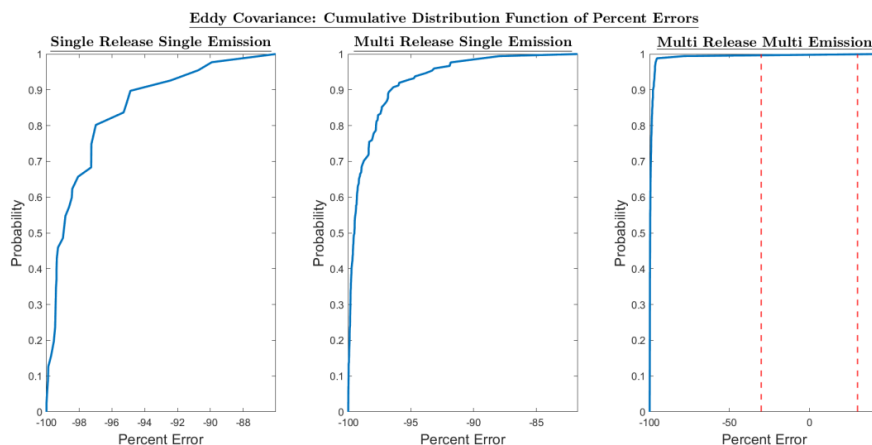
270

271 Figure 6: Quantified emission calculated using the Gaussian plume inverse method. Left pane shows a scatter plot of  
272 quantified emission versus total controlled release for a single release at the site level and the mast was downwind of  
273 the release point. Right pane shows a scatter plot of quantified emission versus total controlled release for multiple  
274 releases at the site level, but the mast was downwind of a single release point. The dashed line represents the 1:1 line  
275 (points below the line were underestimated), the red line is the linear regression fit of the data, and  $n$  is the number  
276 of data points.

### 277 3.2 Quantification within 30% Uncertainty

#### 278 3.2.1 Eddy Covariance

279 The eddy covariance method showed a very low probability of quantifying emissions within 30% uncertainty ( $\pm 30\%$ )  
280 (Figure 7). Only a single measurement in the multi release multi emission category showed an approximately 0.01  
281 probability of quantifying within 30% (Figure 7). The errors for eddy covariance were between -100 and -86% for  
282 single release single emission, between -100 and -82% for multi release single emission, and between -100 and about  
283 +30% for multi release multi emission (Figure 7). This shows that using eddy covariance to quantify single-point and  
284 multi-point emissions will largely underestimate emissions.

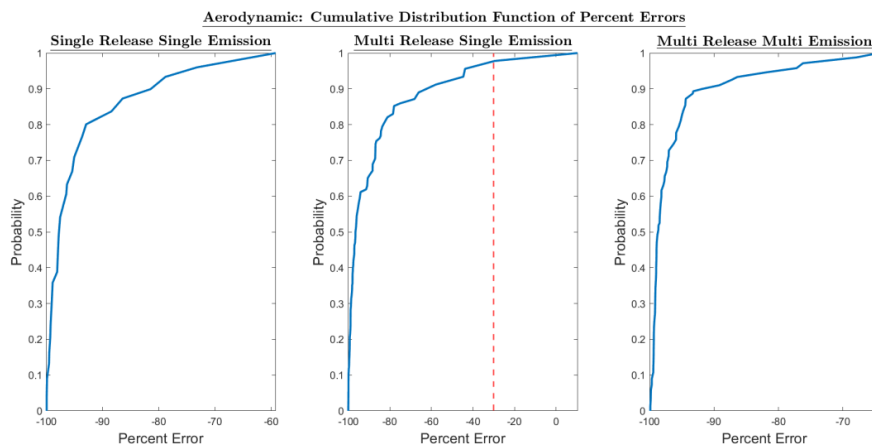


285

286 Figure 7: Cumulative distribution function (cdf) of percent errors for eddy covariance. Left pane shows a cdf plot for  
287 a single release at the site level and the mast was downwind of the release point. Center pane shows a cdf  
288 for multiple releases at the site level, but the mast was downwind of a single release point. Right pane shows a cdf for  
289 multiple releases at the site level and the mast was downwind of more than one release point. The area bounded by  
290 the red dotted line shows the region within  $\pm 30$  uncertainty.

### 291 3.2.2 Aerodynamic Flux Gradient

292 The aerodynamic flux gradient also showed a very low probability of quantifying within 30% uncertainty (Figure 8).  
293 In the multi release single emission category results indicate a 0.02 probability of quantifying within 30% (Figure 8)  
294 of the true value. The errors for aerodynamic flux gradient were between -100 and -60% for single release single  
295 emission, between -100 and 0% for multi release single emission, and between -100 and -70% for multi release multi  
296 emission (Figure 8). These data show that the aerodynamic flux gradient will underestimate a point emission. Similar  
297 to eddy covariance, quantifying an emission within 30% uncertainty using aerodynamic flux gradient for point sources  
298 is highly unlikely.



299

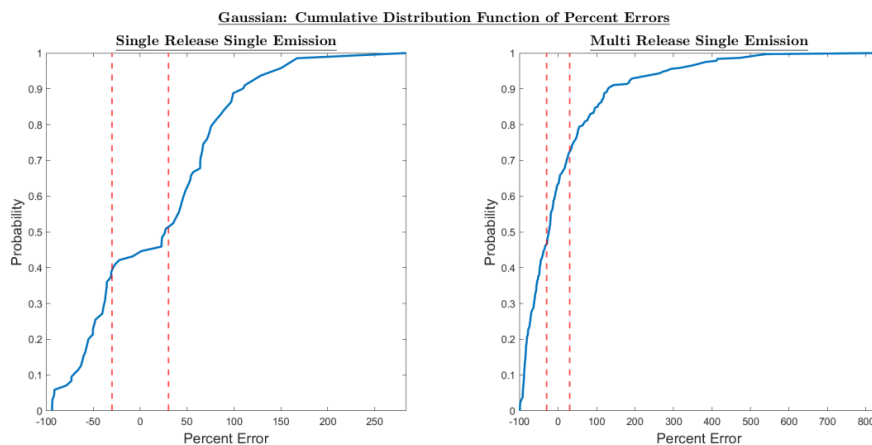
300 Figure 8: Cumulative distribution function (cdf) of percent errors for aerodynamic flux gradient method. Left pane  
301 shows a cdf plot for a single release at the site level and the mast was downwind of the release point. Center pane



302 shows a cdf for multiple releases at the site level, but the mast was downwind of a single release point. Right pane  
303 shows a cdf for multiple releases at the site level and the mast was downwind of more than one release point. The  
304 area bounded by the red dotted line shows the region within  $\pm 30$  uncertainty.

### 305 3.2.3 Gaussian Plume Inverse Method

306 The Gaussian plume inverse method showed a higher probability of quantifying an emission correctly within 30%  
307 uncertainty than eddy covariance and aerodynamic flux gradient methods (Figure 9);  $\approx 0.12$  for the single release single  
308 emission and  $\approx 0.25$  for the multi release single emission categories (Figure 9). Percent errors of the Gaussian method  
309 calculated emissions are between -100 and +250% for single release single emission and between -100 and +800%  
310 for multi release single emission (Figure 9). This shows that even though the Gaussian method is designed for point  
311 sources, it is highly likely to miss, underestimate or overestimate an emission. Similar to eddy covariance and  
312 aerodynamic flux gradient, it is a challenge to correctly quantify a single emission event (single release or multiple  
313 release) using the Gaussian plume inverse method.



314  
315 Figure 9: Cumulative distribution function (cdf) of percent errors for the Gaussian plume inverse method. Left pane  
316 shows a cdf for a single release at the site level and the mast was downwind of the release point. Right pane shows a  
317 cdf for multiple releases at the site level, but the mast was downwind of a single release point. The area bounded by  
318 the red dotted line shows the region within  $\pm 30$  uncertainty.

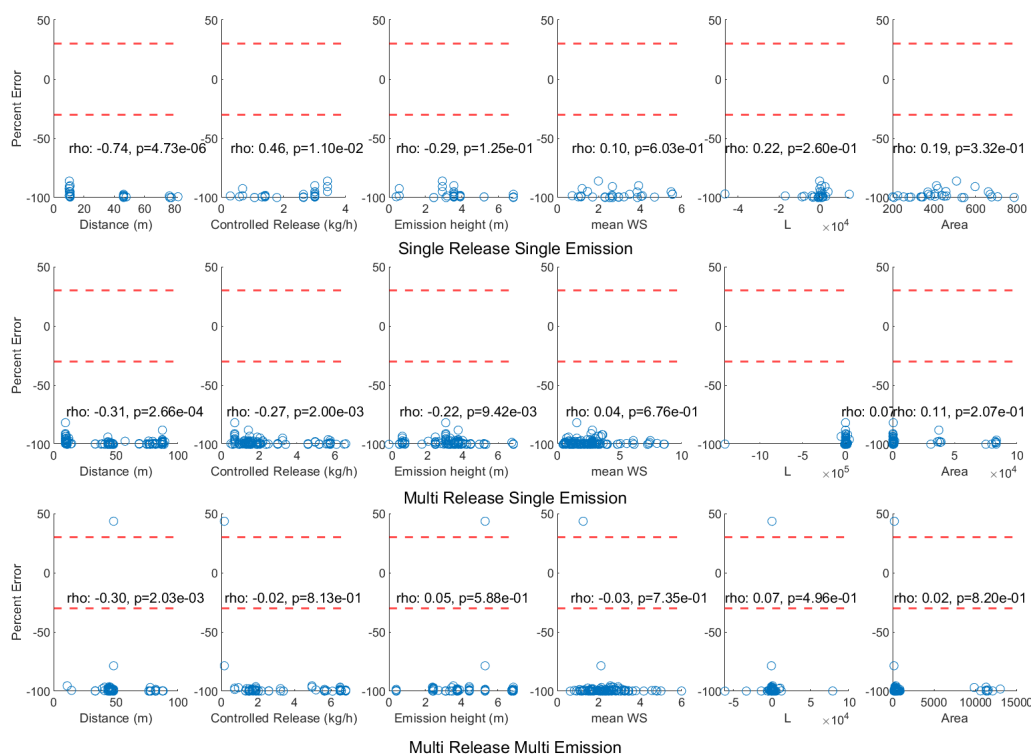
## 319 3.3 Variables Affecting Quantification

### 320 3.3.1 Eddy Covariance

321 A Spearman's rank correlation analysis of measurement and environmental variables (distance, controlled release,  
322 emission height, mean wind speed ( $WS$ ), Monin-Obukhov length ( $L$ ) and contribution area) to percent error in  
323 quantification as calculated by the eddy covariance method, showed that downwind distance had significant impact  
324 on quantification for the single release single emission ( $p = 4.73e-6$ ), multi release single emission ( $p = 2.66e-4$ ), and  
325 multi release multi emission ( $p = 2.00e-3$ ) categories for  $p < 0.01$  significance level (Figure 10). The correlation  
326 coefficients were -0.74 for single release single emission, -0.31 for multi release single emission, and -0.30 for multi  
327 release multi emission. The negative correlation in all three categories suggests that the percent error became more  
328 negative as distance increased i.e., far away emission sources were likely underestimated or undetected. Also,



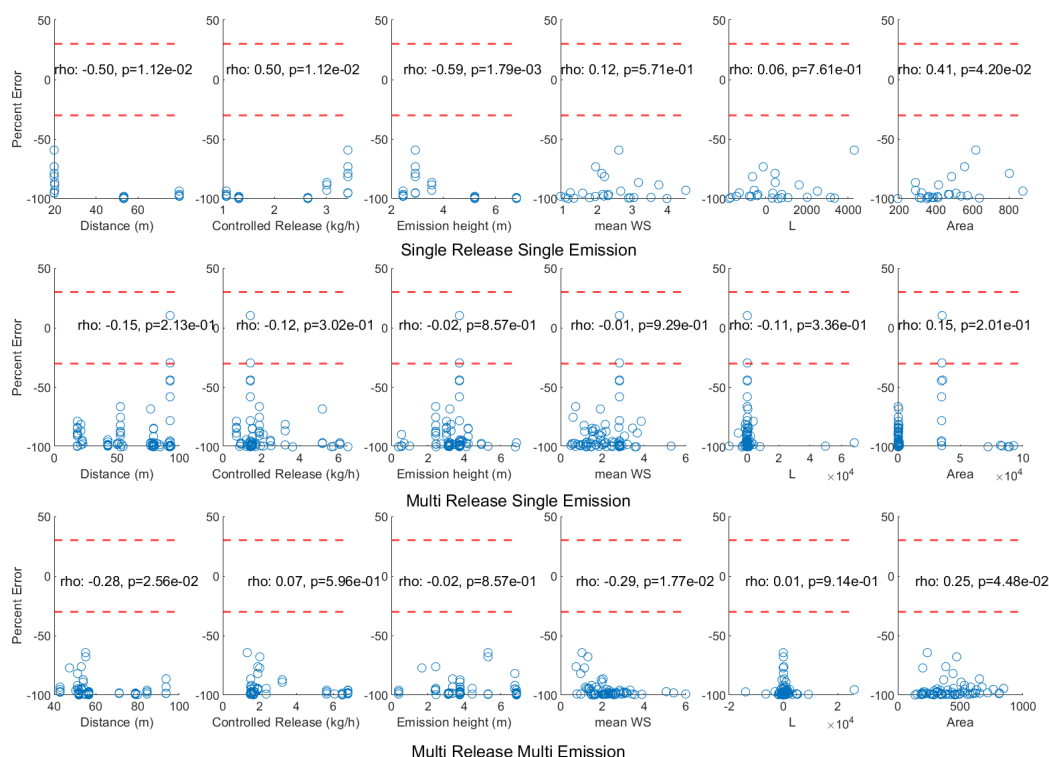
329 controlled release and emission height had significant impact on quantification only in the multi release single  
 330 emission category,  $p = 2.00e-3$  and  $9.42e-3$  respectively, (Figure 10) but this correlation was inconsistent across the  
 331 three categories. Due to inconsistent correlation, and the errors being close to -100%, the results show that generally,  
 332 quantifying emissions using an eddy covariance approach will not work for emissions typically observed at oil and  
 333 gas production sites.



334  
 335 Figure 10: Correlation analysis for eddy covariance in the three release categories. The area bounded by the red  
 336 dotted line shows the region within  $\pm 30$  uncertainty.

### 337 3.3.2 Aerodynamic Flux Gradient

338 A Spearman's rank correlation analysis between the environmental and measurement variables and emissions  
 339 calculated using the aerodynamic flux gradient method showed that only emission height in the single release single  
 340 emission category had significant impact on model quantification ( $p = 1.79e-3$ ) (Figure 11). The correlation between  
 341 emission height and percent error in this category was -0.59 suggesting percent error became more negative as  
 342 emission height increased. However, the correlation between emission height and percent error in the multi release  
 343 single emission and multi release multi emission categories is approximately zero, meaning no correlation. Similar to  
 344 eddy covariance, there is inconsistent correlation, and most errors are close to -100% (Figure 11). The results show  
 345 that generally, quantifying emissions using an aerodynamic flux gradient approach will not work for emissions  
 346 typically observed at oil and gas production sites.

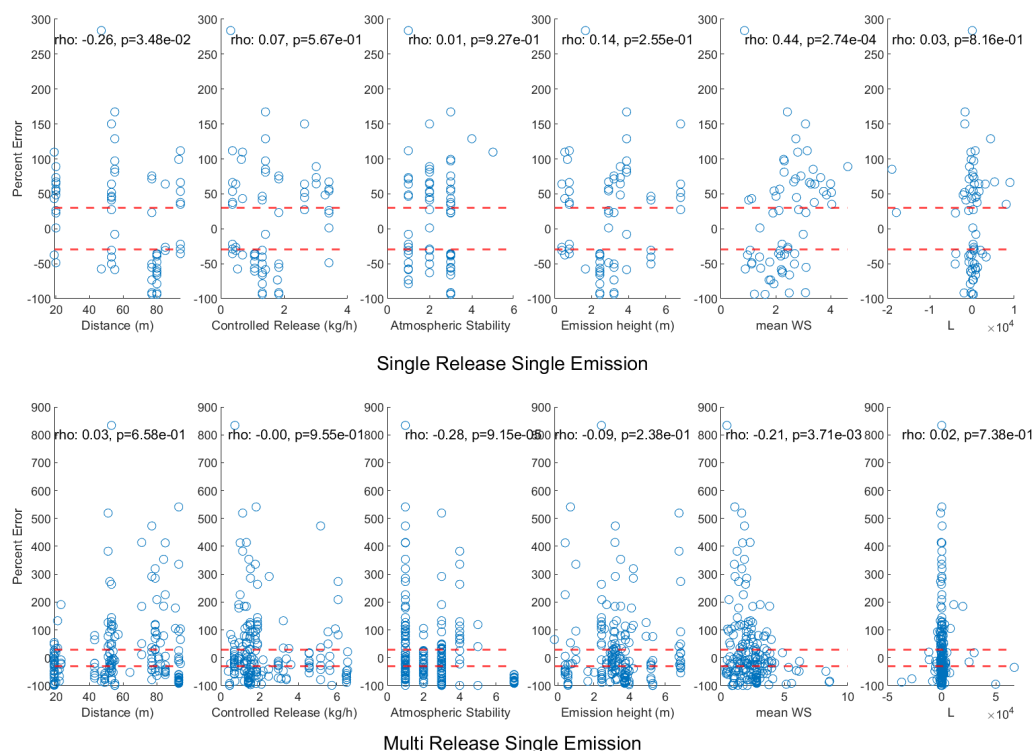


347

348 Figure 11: Correlation analysis for aerodynamic flux gradient in the three release categories. The area bounded by  
 349 the red dotted line shows the region within  $\pm 30$  uncertainty.

### 350 3.3.3 Gaussian Plume Inverse Method

351 The Spearman's rank correlation analysis between the emissions calculated using the Gaussian plume inverse method  
 352 and measurement/environmental variables showed that only the mean wind speed and atmospheric stability had  
 353 significant impact on the model quantification (Figure 12). In the single release single emission category, mean wind  
 354 speed and percent error had a positive correlation ( $0.44, p = 2.74e-4$ ) indicating that an increase in WS increased the  
 355 model's positive error. However, in the multi release single emission category, the correlation is opposite (a negative  
 356 correlation of  $-0.21, p = 3.71e-3$ ) (Figure 12). Atmospheric stability had significant impact on model quantification  
 357 in the multi release single emission category ( $p = 9.15e-5$ ) but not in the single release single emission category (Figure  
 358 12). The correlation analysis for the Gaussian plume inverse model was inconsistent suggesting random errors in  
 359 quantification. This shows that the model could either underestimate or overestimate an oil and gas emission at  
 360 random.



361

362 Figure 12: Correlation analysis for the Gaussian plume inverse method in the three release categories. The area  
 363 bounded by the red dotted line shows the region within  $\pm 30\%$  uncertainty.

#### 364 4 Discussion

365 Methane emissions quantification from oil and gas is a complex system comprising of gas emissions from different  
 366 heights, different locations, encountering aerodynamic obstacles of different sizes, and of varying duration, amongst  
 367 others. The ability to precisely quantify an emission using data collected by a point sensor, downwind of a source is  
 368 directly influenced by plume dynamics. The CH<sub>4</sub> plume downwind of a source will change in size and shape in  
 369 different atmospheric conditions, in open areas versus areas with obstacles, diurnally, and in different seasons (Casal,  
 370 2008). In this study, the precision to which downwind models (eddy covariance, aerodynamic flux gradient and  
 371 Gaussian plume-based) could quantify the emission rate of point source(s) were tested in different atmospheric  
 372 conditions (rain, sunny, snow, windy, calm etc.), and aerodynamic scenarios (emissions sources in open areas, behind  
 373 obstacles, changing atmospheric stability, and day/night). As a result, testing the models' predicted emission rates to  
 374 controlled release rates in different conditions introduced real-world scenarios that have not previously been tested,  
 375 hence better understanding model uncertainty in the application of quantifying emissions from oil and gas production  
 376 infrastructure.

#### 377 4.1 Eddy Covariance

378 Eddy covariance underestimated or failed to observe almost all emissions released during this study (linear regression  
 379  $m$  between 0 and 0.07, and  $R^2$  between 0.03 and 0.36) (Figure 4). The method measures CH<sub>4</sub> atmospheric fluxes for





380 area sources transferred as eddies of different sizes as caused by turbulence within the atmospheric boundary layer  
381 (Babaeian and Tuller, 2023). Assumptions governing eddy covariance include: (1) the terrain is homogenous and  
382 horizontal, (2) CH<sub>4</sub> fluxes are turbulent, (3) measurements at a point are from an upwind area, (4) measurements are  
383 within the boundary layer and in the constant flux layer, (5) instruments can capture small fluctuations at high  
384 frequency, (6) fluctuations in air density are negligible (Babaeian and Tuller, 2023), and (7) upward fluxes represent  
385 emissions and downward fluxes represent depositions (Zinke et al., 2024). Nemitz et al., (2018) adds that eddy  
386 covariance is frequently deployed to target large fluxes in high-emission ecosystems, which is not typical in oil and  
387 gas, and that data where wind direction includes obstructed wind sectors should be flagged (Nemitz et al., 2018).  
388 For oil and gas point sources, the measured gas concentration is dependent on plume dynamics as opposed to mass  
389 transfer and eddy covariance methods using fence-line measurements are unlikely to work because:

- 390 • Oil and gas point sources violate assumptions (1), (2), and (4) as these sources are heterogenous and  
391 emissions are collimated plumes instead of turbulent fluxes.
- 392 • As the measurement by a point sensor is dependent on being inside the plume, which changes in different  
393 atmospheric conditions, placing the sensor high enough, and/or far enough downwind, to where the flux  
394 layer is constant, is impractical.
- 395 • Even though current eddy covariance application assumes the vertical flux at a point is independent of  
396 atmospheric stability (Denmead, 2008), atmospheric stability has impact on point source gas dispersion  
397 at fence line distances and hence needs to be accounted for even for eddy measurements.
- 398 • Footprint models are designed for area sources that require horizontal homogeneity of the flow (Kljun  
399 et al., 2015). As a result, the area of contribution generated by the models do not accurately represent  
400 the area between the point sources and the measurement location at fence line distances.

401 In summary, this study shows that eddy covariance is not applicable for oil and gas point source quantification.

#### 402 **4.2 Aerodynamic Flux Gradient**

403 Overall, aerodynamic flux gradient method underestimated the emission rate of all controlled releases during this  
404 experiment with high variability. The slope of the linear regression and R<sup>2</sup> were both very small (linear regression m  
405 between 0 and 0.22, and R<sup>2</sup> between 0.01 and 0.39) (Figure 5). The aerodynamic flux gradient model quantification  
406 is used to quantify emissions from area sources and relies on differences in CH<sub>4</sub> concentrations between the higher  
407 and lower height, and stability correction factors. Assumptions of flux-gradient approach using Monin-Obukhov  
408 similarity theory include: (1) measurements require steady state conditions of wind direction and speed, (2)  
409 measurements should be done above the roughness sub-layer, (3) sufficiently large homogenous area for development  
410 of an adequately equilibrated layer of air, and for constant equilibrium during measurement (Prueger and Kustas,  
411 2015), and (4) positive fluxes represent emissions and downward fluxes represent absorptions (Kamp et al., 2020).  
412 Similar to eddy covariance, aerodynamic flux gradient methods at fence-line distances are unlikely to work because  
413 point sources typical of oil and gas emissions violate the following assumptions:

- 414 • Obstacles at an oil and gas facility affects wind direction and speed, and these impacts may also vary  
415 substantially with small changes in wind direction. Therefore, wind conditions are unlikely to attain  
416 steady state during the measurement period, as directed by assumption (1) above.



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- The emission height of oil and gas sources in typical upstream field conditions can be as low as 0.4 m and as high as 6.9 m and measurements are unlikely to be made by fence-line sensor above the roughness sublayer (2 above), i.e. twice the height of the mean obstacle height for ~30 m downwind.
  - Oil and gas sources are heterogeneous (i.e. varying source distance and height) and can last a short time (e.g. a short maintenance event) or a long time ('normal' fugitive emissions) hence, achieving constant equilibrium, as stated in (3) above, is unlikely.
  - Footprint models used to generate the area of contribution between the source and the measurement location are designed for area sources with horizontal flow homogeneity (Kljun et al., 2015). Thus, the area of contribution generated for oil and gas point sources is likely inaccurate.

#### 426 **4.3 Gaussian Plume Inverse Method**

427 In contrast to the other methods in this study, the Gaussian plume inverse model both underestimated and  
428 overestimated emissions in this study. Linear regression gradient and coefficient of correlation (m between 0.95 and  
429 1.49, and  $R^2$  between 0.21 and 0.65; Figure 6) was better than either eddy covariance or aerodynamic flux gradient.  
430 The main assumption of the Gaussian plume model is that  $\text{CH}_4$  emitted from a point source enters the air flow,  
431 disperses vertically and laterally, forming a conical plume (Riddick et al., 2022b; US EPA, 2013). However, the  
432 formation of a conical plume is hindered at oil and gas facilities by obstacles (equipment) and is affected by  
433 atmospheric stability. Atmospheric stability in the Gaussian plume inverse model is based on Pasquill-Gifford  
434 classification system which accounts for daytime solar insolation (slight, moderate and strong), nighttime cloud cover  
435 and surface wind speed at 10 m (Kahl and Chapman, 2018). Solar insolation and cloud cover are not typically  
436 measured, and if measured, dispersion parameter models currently available do not use this data, therefore, it is  
437 difficult to calculate for continuous fence-line measurements. The modified dispersion parameters developed by EPA  
438 (US EPA, 2013) only account for wind conditions i.e., speed and deviation in direction. As a result, plume dynamics  
439 during diverse atmospheric conditions such as during snow versus rain or sunny conditions are unaccounted for.  
440 In this study, despite the Gaussian model having been developed for point sources, the model did not show consistent  
441 correlation with the measurement and atmospheric variables. This showed that there are complexities in continuous  
442 monitoring quantification compared to survey solutions where the model is widely applied, that introduce significant  
443 uncertainties in quantification. It is suggested that one problem with the Gaussian plume model is that the dispersion  
444 coefficients are simply not representative as they were developed for longer distances, in different climatological  
445 conditions, and do not transfer well to current applications (Riddick et al., 2022a). We conclude that, while it is better  
446 suited than eddy covariance or aerodynamic flux gradient, a Gaussian plume inverse approach will likely have  
447 significant uncertainties when used to quantify emissions from oil and gas production sites using data collected at a  
448 fence line (~ 30 m away).

#### 449 **4.4 Implications**

450 In the recent years, there has been growing interest and need for accurate  $\text{CH}_4$  quantification from oil and gas sites.  
451 This is generally done through survey methods and continuous monitoring using fence-line sensors. Continuous  
452 monitoring involves having stationary sensors measuring meteorology and  $\text{CH}_4$  mixing ratios, which are then used to  
453 infer emission rate. For point sources, downwind methods such as the Gaussian plume inverse method have been



454 widely used, especially for survey quantification. Continuous monitoring is relatively new but fast growing. This  
455 study's design replicated a continuous monitoring setup.

456 Oil and gas point sources could either be single emissions or multiple emissions occurring concurrently. In cases of  
457 multiple emissions with more than one release point being downwind, the Gaussian model is limited, as it can only  
458 quantify one source at a time (dispersion coefficients are generated as a function of emission height and source  
459 distance). As a result, models used in other applications such as eddy covariance and aerodynamic flux gradient have  
460 been proposed as the solution. However, as this study has shown, eddy covariance and flux gradient approaches are  
461 unlikely to quantify realistic emission estimates using fence-line measurements. Here, we strongly advise that  
462 controlled tests under controlled environments are crucial to evaluate modelling approaches' precision and accuracy,  
463 and associated uncertainties before applying them in the real world. Even though these modelling approaches have  
464 been reported to work elsewhere (e.g., agricultural and landfill emissions), it does not necessary mean it could work  
465 in the intended area of application.

#### 466 **Author contributions**

467 Mercy Mbua: Conceptualization, Data curation, Methodology, Formal analysis, Investigation, Writing original draft,  
468 Review, and Editing. Stuart N. Riddick: Conceptualization, Methodology, Review, Editing, Project Administration,  
469 Supervision, and Funding Acquisition. Elijah Kiplimo: Investigation, Review, and Editing. Daniel J. Zimmerle:  
470 Review, Editing, Funding acquisition, and Project administration.

#### 471 **Declaration of competing interest**

472 The authors declare that they have no known competing financial interests or personal relationships that could have  
473 appeared to influence the work reported in this paper.

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#### 481 **Data availability**

482 Data sets for this research are available in the in-text data citation reference: Mercy, Mbua; Riddick, Stuart N.;  
483 Kiplimo, Elijah, and Zimmerle, Daniel J. Dataset for evaluating the accuracy of downwind methods for quantifying  
484 point source emissions. [Dataset]. Dyrad.

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