Extraordinary bloom of toxin-producing phytoplankton enhanced by strong retention on the~~in~~ offshore ~~continental~~ Patagonian ~~s~~Shelf ~~waters~~

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Abstract:

The extensive Patagonian continental shelf in the Atlantic Ocean is renowned for its high productivity associated with nutrient-rich waters that fertilise massive phytoplankton blooms, especially along the shelf-break frontal system. Growing evidence reflects this ecosystem as a hotspot for harmful algal blooms (HABs). Whether these HABs reach coastal areas or are exported to the adjacent ocean basin by energetic edge currents remains unexplored. During two oceanographic cruises in spring 2021, a bloom of dinoflagellates of the Amphidomataceae family was sampled over the outer shelf with a ten-day interval, at stations 40 km apart. The bloom was first sampled on November 16, with 32 million cells L^{-1}, and was still persistent on November 25, with 14 million cells L^{-1}. The magnitude of this bloom is a global record for this group so far reported in the literature. The toxin azaspiracid-2 was detected in both stages of the bloom, with values up to 2122 pg L^{-1}. The most likely source of AZA-2 was *Azadinium spinosum* ribotype B. The bloom developed in vertically stable waters (60 m mixed layer depth) with elevated chlorophyll concentration. Water retention and the presence of fronts induced by horizontal stirring controlled the persistence and trajectory of the bloom in a localised area over the continental shelf, as evidenced by analysis of geostrophic surface currents, Lyapunov coefficients, and particle advection modelling. These findings underscore the importance of monitoring HABs in offshore environments, and the need to understand bio-physical interactions that govern bloom taxa assemblages and transport pathways.

1. Introduction

In marine environments, dinoflagellates are the primary toxin-producing group of protistan plankton and key causative agents of harmful algal blooms (HABs). As the most diverse group of toxic microorganisms, -e.g. *Alexandrium* spp., *Karenia* spp., *Dinophysis* spp., *Azadinium* spp., *Amphidoma* spp.-, dinoflagellates produce a wide range of toxins. Phycotoxins are natural intracellular metabolites synthesised by certain microalgae that can be transferred through the food web, having severe impacts on marine biota, ecosystems and human health (Anderson et al., 2015; Sunesen et al., 2021). In the Argentine Sea, records of HABs caused by different plankton species have risen since the first documentation of human poisoning in spring 1980 caused by paralytic shellfish toxins (PST), produced by *Alexandrium catenella* in coastal areas (reviewed in Ramírez et al., 2022). Broadly, HABs were long thought to occur exclusively in coastal regions, due to their visible impacts on water quality and human-related activities, as documented for instance in the Argentine Patagonian Gulfs (Wilson et al., 2015; D'Agostino et al., 2019) and the Beagle Channel (Almandoz et al., 2019; Cadaillon et al., 2024). However, the perception of HABs as solely coastal events was biased, primarily due to greater monitoring efforts in coastal areas compared to the fewer studies conducted offshore (Hallegraeff et al., 2021; Sunesen et al., 2021; Anderson et al., 2021). In line with this trend, the expansion of the monitored area in the Argentine Sea from the coast to the outer continental shelf over recent decades has~~ve~~ confirmed that toxic species are indeed common in offshore waters ~~in the Argentine Sea~~ (Ramírez et al., 2022), especially ~~in the outer continental shelf~~ associated with~~to~~ the shelf-break front (reviewed in Guinder et al., 2024). Furthermore, the increase in oceanographic studies focused on detecting HABs along the outer Patagonian shelf has led to several new records of toxin-producing species and phycotoxins in the South Atlantic (Akselman et al., 2015; Guinder et al., 2018; Tillmann et al., 2019). In particular, large HABs formed by the nano-dinoflagellates of the Amphidomataceae family — producers of the toxin azaspiracids — have emerged as important hazards in the productive Patagonian shelf-break frontal ecosystem (Guinder et al., 2024). It is well known that dinoflagellates possess advantageous strategies for thriving in frontal systems, such as effective swimming, mucus and cyst formation, mixotrophy, and toxin production (Smayda, 2002; Glibert, 2016). However, the bio-physical mechanisms explaining the development of large harmful blooms on hydrographically complex shelves are still not fully understood mainly due to the lack of simultaneous taxonomic data and velocity fields at synoptic scales.

In oceanic waters, the permanence and spatial extent of discrete phytoplankton blooms are influenced by dispersal mechanisms that rely on diffusion and horizontal advection (Abraham et al., 2000; Mahadevan, 2016; Lehahn et al., 2017). Typically, the dispersion and stirring of phytoplankton blooms in the ocean have been studied using remote sensing of chlorophyll-*a* and models (Lehahn et al., 2007; Lévy et al., 2018; Ser-Giacomi et al., 2023), with few studies considering *in situ* sampling (Abraham et al., 2000; Giddings et al., 2014; Hernández-Carrasco et al., 2020) to assess the bio-physical couplings of bloom development and persistence. In addition to the considerable complexity of studying the dynamics of phytoplankton blooms in the ocean, which involve multiple physical (temperature, salinity, and mixing of water masses), biological (species competition, grazers), and chemical (nutrients) factors, monitoring HABs presents an even greater challenge (Pitcher et al., 2010; Wells et al., 2020; Iriarte et al., 2023). This includes the identification of the species involved in these blooms. It is precisely in this area where detailed studies, employing multiple approaches such as microscopy, molecular techniques, and toxin profiling (Smayda et al, 1997; Wells et al., 2020), make a significant contribution to understanding HABs' development and potential impacts.

The extensive Patagonian shelf-break front (35-55°S) in the SW Atlantic Ocean is a high productivity ecosystem, located ~200 to ~900 km offshore (Martinetto et al., 2019; Guinder et al., 2024). This permanent termohaline front is associated with the upwelling of nutrient-rich waters of the westerly edge of the Malvinas Current, which fertilises the surface waters ~~over the shelf~~ (Palma et al., 2008; Matano et al., 2010). ~~Additionally, over the mid shelf, the Patagonian Current transports diluted~~

subantarctic waters northwards, also loading with nutrients the region. Hence, mMassive phytoplankton proliferations occur over the outer shelf in spring and summer (García et al., 2008; Carreto et al., 2016; Ferronato et al., 2023), including a variety of HAB-forming taxa and associated phycotoxins (Ramírez et al., 2022; Guinder et al., 2024). In fact, the increase in oceanographic studies focused on the detection of HABs along the outer Patagonian shelf has led to several new records of toxin-producing species and phycotoxins in the South Atlantic (Akselman et al., 2015; Guinder et al., 2018; Tillmann et al., 2019). The most conspicuous HABs are those formed by the nano-dinoflagellates of the clade Amphidomataceae family (Akselman and Negri, 2012; Fabro et al., 2019; Guinder et al., 2020; Tillmann et al., 2018; 2019), producers of the toxin azaspiracids (AZAs). These large HABs have emerged as important hazards in the productive Patagonian shelf-break frontal ecosystem, and which include the four azaspiracids (AZAs)-producing species, i.e. *Azadinium dexteroporum*, *Az. poporum*, *Az. spinosum*, and *Amphidoma languida* (Krock et al. 2019). Amphidomataceans have been reported from different marine regions globally (Tillmann, 2018; Salas et al., 2021; Liu et al., 2023) but so far, the maximum bloom abundances reported in the literature are from the Argentine Sea (Akselman and Negri, 2012). During the springs of 1991 and 1992, these dinoflagellates reached between 3 and 9 million cells per litre and caused water discoloration in the northern area (38-42°S; 58°-56°W) of the Patagonian shelf (Akselman and Negri, 2012). No toxin screening was performed at that time, but in spring 2015, an AZA-producing *Azadinium* strain was isolated from the production of AZAs was confirmed in another large bloom in the area (Tillmann et al., 2019). Moreover, AZAs have been detected in the tissue of the scallop *Zygochlamys patagonica* since the early 90ies (Turner and Goya, 2015). These scallops form large seabed banks along the 100-metre isobath between 38°S-48°S (Alemany et al., 2024), associated with the high phytoplankton productivity over the outer Patagonian continental shelf.

Despite the limited synoptic sampling in offshore waters, the prevalence of HABs in the Patagonian front highlights this ecosystem as a hotspot that requires further monitoring. The notably high abundance of Amphidomataceans over the outer shelf holds greater significance when assessing the potential risks that bloom dissemination may posed to both regional and global ecosystems. The evolution and transport of HABs remain poorly understood, as does the question of whether they may reach coastal areas or be exported offshore into the stirring Atlantic Ocean. The accumulation and dispersion of phytoplankton patches in complex shelf waters need to be addressed from an interdisciplinary perspective, combining remote sensing, modelling, and in situ sampling (Ferronato et al., 2025). In a recent study conducted on the outer Patagonian Shelf, the core of the Malvinas Current, and adjacent open ocean waters, we emphasized Aside from the key role of mesoscale energetic variability in modulating phytoplankton spring blooms (Ferronato et al., 2025), community,However, these observational studies combining multiple approaches at synoptic scale are stillremain scarce in the region.

In this study, we characterised the bio-physical aspects of a large multispecific spring bloom of Amphidomataceans, detected through an unusual sampling effort that involved two research expeditions in November 2021. This HAB was observed at two sampling sites 40 km apart within a span of 10 days. In oceanic waters, the permanence and spatial extent of discrete phytoplankton blooms are influenced by dispersal mechanisms that rely on diffusion and horizontal advection (Abraham et al., 2000; Mahadevan, 2016; Lehahn et al., 2017). Typically, the dispersion and stirring of phytoplankton blooms in the ocean have been studied using remote sensing of chlorophyll-*a* and models (Lehahn et al., 2007; Lévy et al., 2018; Ser-Giacomi et al., 2023), with few studies considering *in situ* sampling (Abraham et al., 2000; Giddings et al., 2014; Hernández-Carrasco et al., 2020) to assess the bio-physical couplings of bloom development. Aside from the key role of mesoscale energetic variability in modulating phytoplankton community, observational studies combining multiple approaches at synoptic scale are still scarce. WHere, we combined field observations of protistan plankton species composition and associated toxins, with remotely sensed ocean colour images of chlorophyll-*a* and geostrophic surface currents, particle tracking experiments, and Lyapunov coefficient analysis to assess the horizontal displacement and retention of the Amphidomataceae bloom within a mesoscale eddy. Furthermore, we aim to explore whether this HAB that developed in offshore shelf waters might reach coastal areas or be advected by the Malvinas Current, facilitating the dispersal of toxic species to other shelves and ocean basins.

2. Materials and Methods

2.1 Hydrography and productivity in the Patagonian continental shelf

Along the external margin of the Patagonian continental shelf, between 35°S and 55°S, a thermohaline front develops throughout the year, characterised by high biological productivity. The development of this front and the associated upwelling is due to the interaction of the energetic western boundary current system with the steep slope and the waters of the shelf, as well as the effect of winds and tides (reviewed in Piola et al., 2024). The Malvinas Current originates at ~55°S as a branch of the Antarctic Circumpolar Current (**Fig. 1**) and runs northwards at high velocity (mean surface velocities 45 cm/s, Piola et al., 2024) along the shelf break in two jets that meet at ~44°S (Frey et al., 2023). Then, at ~38°S, the Malvinas Current meets the warm and oligotrophic Brazil Current, which runs southwards, in the so-called Brazil-Malvinas Confluence, and waters are exported eastwards into the South Atlantic Ocean basin (**Fig. 1**). In addition, another branch of the Antarctic Circumpolar Current gives origin to the Patagonian Current which runs northwards over the continental shelf, carrying diluted subantarctic waters (**Fig. 1**).

The interaction of the Malvinas Current with the irregular bottom topography generates upwelling of cold, nutrient-rich waters that fertilise phytoplankton over the shelf, together with the Patagonian Current (reviewed in Guinder et al., 2024). Phytoplankton blooms expand over the mid and outer shelf as reflected by a persistent satellite chlorophyll-*a* band, wider in spring and narrower in summer along the shelf-beak front (Guinder et al., 2024). The magnitude of the upwelling has low seasonal variability and is heterogeneous along the extensive latitudinal range of the slope (Combes and Matano, 2018). Hence, productivity over the shelf varies spatially and temporally, and in consequence, multiple bioregions emerge, each characterised by unique phytoplankton phenological patterns, as revealed by climatological analysis of satellite-derived chlorophyll *a* (Delgado et al., 2023).

2.2 Research cruises

Two oceanographic expeditions were carried out during late-spring (November 2021) to study the microbial plankton communities in the Argentine continental shelf and adjacent ocean basin (**Fig. 1**). The first research cruise was the Ana María Gayoso (hereafter Gayoso, GA) onboard the R/V Bernardo Houssay (PNA and CONICET, Arg.), which covered a sampling period between 16 and 22November 16 and 22, 2021, at 10 stations along the outer Patagonian shelf, the core of the Malvinas Current, and adjacent open ocean waters (**Fig. 1**). This cruise was planned to analyse the pre-bloom conditions of the coccolithophore *Gephyrocapsa huxleyi* which blooms in early summer (December) on the Patagonian shelf. The location of the sampling stations was delineated tracking daily satellite-derived signals of chlorophyll-a and particulate inorganic carbon (see Ferronato et al., 2025). The second cruise was the Agujero Azul (AA) onboard the R/V Victor Angelescu (INIDEP and CONICET, Arg.) which covered a sampling period between 25November 25 and 3December 3, 2021, at 23 stations aligned in two cross-shelf transects in the so-called Agujero Azul area (44-47°S; 62-57.5°W) (**Fig. 1**). This research expedition was planned as part of a multidisciplinary national project to analyse the pelagic and benthic biodiversity of this productive area, with a specific focus on species of commercial interest (Alemany et al., 2024). We first provide a general overview of the same in situ measurements (CTD) and water sampling conducted in both cruises for plankton and nutrients. . Tbut then we focus on the Amphidomataceae bloom observed over the shelf at station GA01 on 16 November 16 and at station AA09 on 25November 25, separated by *ca.* 40 km (**Fig. 1**).

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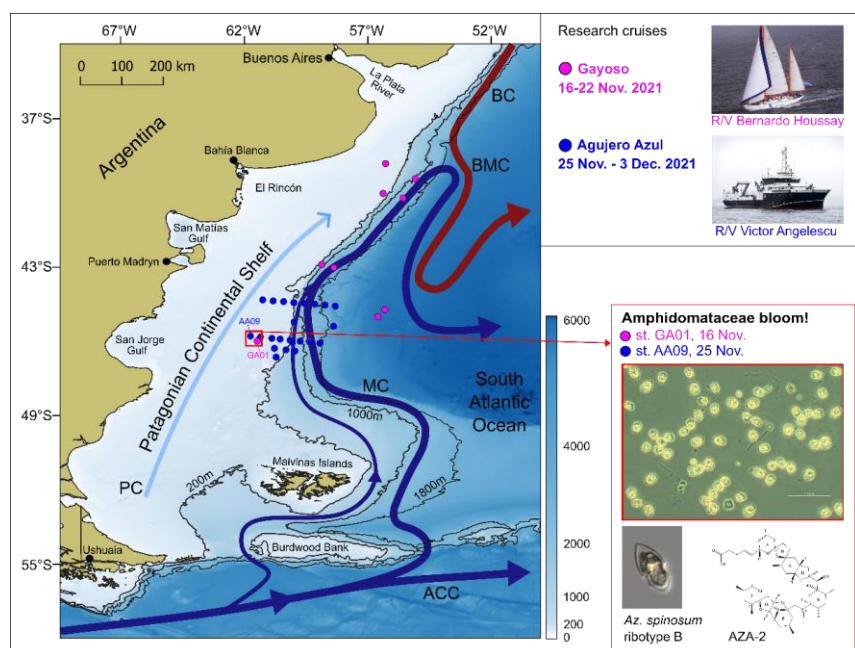
170 **Figure 1: Sampling stations of the research cruises Gayoso (pink dots, onboard R/V Houssay) and Agujero Azul (blue**

171 **dots, onboard R/V Angelescu), and the main circulation pattern in the Argentine continental shelf and shelf break. The**

172 **stations where the bloom of Amphidomataceae was observed are indicated with a red square: GA01 and AA09, sampled**

173 **on November 16 and 25, respectively. Micrographs show the Amphidomatacean bloom and a cell of *Azadinium***

174 ***spinosum* ribotype B, producer of azaspiracid-2 (AZA-2). ACC: Antarctic Circumpolar Current, MC: Malvinas**

175 **Current, BC: Brazil Current, BMC: Brazil-Malvinas Confluence and PC: Patagonian Current. Bathymetry from**

176 **GEBCO, 2021. Isobaths of 200 m, 1000 m and 1800 m are displayed.**

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178 **2.3 Remote sensing of surface Chl-*a* and, SST ~~and ADT~~**

179

180 In order to contextualise the discrete observations of plankton within broader spatio-temporal dynamics, we explored satellite-

181 derived surface chlorophyll-a concentration (Chl-*a* sat) data from the days of the cruises. In this sense, daily merged product

182 with a 4 km resolution, provided by the GlobColour project (distributed by ACRI ST, France: https://hermes.acri.fr/), were

183 downloaded. This ocean color product is generated from the fusion of the SeaWiFS, MERIS, MODIS-Aqua, and OLCI sensors

184 and estimates the average Chl-*a* concentration in the surface layer (Maritorena et al., 2010). The fusion of data from different

185 satellite sensors, combined with the quality control criteria used by GlobColour, enables enhanced spatial and temporal

186 coverage. Eight-day temporal averages were calculated for the periods of November 12–19 and November 20–27, 2021.

187 Additionally, daily Chl-*a* images were assessed from November 14 to 27 focused on the bloom area to track its short-term

188 evolution. Due to high cloud covering, data is missing from ~~22 to 24~~ November 22 to 24. To analyse Sea Surface Temperature

189 (SST) during the sampling periods, daily NSST MODIS-Aqua level L3 images with a 4 km resolution were downloaded from

190 https://oceancolor.gsfc.nasa.gov/. Eight-day temporal averages were also constructed for the periods of November 12–19 and

191 November 20–27, 2021. All images were processed using SeaDAS v8.3 and QGIS 3.38, mapped to the WGS84 reference

192 system (datum WGS84, ellipsoid WGS84), and restricted to the study area. The images were smoothed using a 'Non-Linear
193 Mean 3x3' filter.
194 To compute trajectories of virtual particles (see Sect. 2.6 below) and Finite Size Lyapunov Exponents (FSLE, Sect. 2.7) we
195 used Absolute Dynamic Topography (ADT) maps and geostrophic velocities derived from the ADT maps. Gridded ADT maps
196 of daily temporal resolution and ¼ of degree spatial resolution maps were obtained from CMEMS
197 (https://marine.copernicus.eu/). FSLE images with a spatial resolution of 1/25° grid were downloaded from AVISO
198 (https://www.aviso.altimetry.fr).

2.4 *In situ* measurements and sample collection

202 At each sampling station, continuous vertical profiles of temperature, salinity and fluorescence were measured. In Gayoso
203 cruise, a Sea-Bird 9 plus CTD and a fluorometer sensor Wet Labs FLRTD-5105 were used. In Agujero Azul cruise, a Seabird
204 SBE 9 Digiquartz CTD coupled with an ancillary Seapoint SCF chlorophyll fluorometer were used. To assess the vertical
205 stability of the water column, the Brunt-Väisäla buoyancy frequency (cyc h^{-1}) was computed using the function swN2 of the
206 package oce (Kelley et al. 2022) in R statistical software. Thereafter, the mixed layer depth (MLD, in metres) was defined at
207 the depth where the maximum value of the Brunt-Väisäla frequency was detected.
208 Niskin bottles attached to the CTD-rosette were used to collect water samples at the surface (5 m depth) for the analysis of
209 chlorophyll *a,* dissolved inorganic macronutrients, protistan plankton by microscopy, genetic analysis of the species diversity,
210 and phycotoxins in field samples.
211 For chlorophyll *a*, a volume of 400 mL was filtered through filter GF/F fiber-glass filters pre-combusted at 450 °C for 4 h. A
212 volume of 10 mL of 90 % acetone was used for pigment extraction during 24 hs (4 °C), and thereafter quantified using an
213 Agilent Cary 60 UV-Vis spectrophotometer. Concentration was estimated using the equations developed by Jeffrey and
214 Humphrey (1975).
215 For inorganic nutrients, the water samples filtered through Whatman GF/F fiber-glass filters pre-combusted at 450 °C for 4 h,
216 were stored at -20 °C in alkali-rinsed (NaOH, 0.1 M) polyethylene bottles. Nitrite, nitrate, ammonium, silicate and phosphate
217 were measured using a spectrophotometer Agilent Cary 60 UV-Vis following standard seawaterthe methods outlined by
218 Hansen and Koroleff (1999). These protocols for chlorophyll and nutrient determinations have been applied and intercalibrated
219 for other areas of the Patagonian Shelf (e.g. Ferronato et al., 2025; Gilabert et al., 2025).
220 Duplicate samples for plankton counts collected with Niskin bottles were preserved with Lugol (1% f/c) and formaldehyde
221 (1% f/c) in glass bottles (250 mL) and kept in dark and at 4 °C for their analyses under microscopy. Similarly, duplicate water
222 samples were collected by three vertical net tows (20 µm size pore) integrating the first 30 m depth for the identification of
223 protists' taxa.
224 For the quantification of azaspiracids (AZAs) as well as for genetic analysis of field species diversity, the same sampling
225 protocol was applied. A volume of 4-5 L of seawater from the Niskin bottles was pre-screened through a 20 µm mesh-size,
226 and subsequently filtered through 5 µm pore-size polycarbonate filters (Millipore, Eschborn, Germany) under gentle vacuum
227 (< 200 mbar). Filters were placed in 50 mL centrifuge tubes and preserved at -80 °C for further analyses in the laboratory.

2.5 Microscopy analysis of protistan Pplankton diversity analyses

231 Morphological aspects of plankton cells were carefully observed under different light microscopes all equipped with
232 epifluorescence and differential interference contrast optics: a Nikon Eclipse E-400 microscope, a Zeiss Axioskop 2
233 microscope, and an inverted Axiovert 200 M. In order to measure length and width of cells, micrographs were taken at 1000
234 magnification under a Zeiss Axio Vert.A1 equipped with a digital camera AxioCam 208 Color, and under an Axioskop 2

equipped with a Axiovision digital camera. Thereafter, ~~they~~ micrographs were processed with the software ZEN (v.2.7, Zeiss) and Axiovision (v.4.8, Zeiss). Further, scanning electron microscopy (SEM, FEI Quanta FEG 200) was used to assess detailed taxonomic features of the dinoflagellate species (e.g. arrangement of thecal plates, presence of pores and spines, etc). SEM samples were treated following the protocol described in Tillmann et al. (2017). For the estimation of total protists' abundance (in cells L^{-1}), seawater samples collected with Niskin bottles and fixed with Lugol were settled in sedimentation chambers and single cells were counted under inverted microscope using a magnification of 400 following traditional techniques (Hasle, 1987). All protists larger than 5 µm in cell size were counted and classified into species or genera taxonomic levels, or merged into taxonomic/functional groups organized in size ranges (e.g. ciliates between 10-20 µm, cryptophytes <10 µm, *Gymnodinium*-type cell, Kareniaceae-type cell, etc.). In addition, to assess the relative abundance of the Amphidomataceae species responsible for the multispecific bloom of this clade, subsamples (10-mL) were carefully counted with high taxa resolution.

2.6. Genetic analysis

~~For a broad detection of amphidomatacean~~ As an additional and complementary mean for analysing species diversity in the field samples, metarbarcoding was performed specifically targeting the internal transcribed spacer (ITS1) region, following Liu et al. (2023). This information was used as a complement to the exhaustive morphological taxonomy performed under light microscopy and SEM~~.~~ In particular, metabarcoding also allows the specific identification of ribotypes of Amphidomatacean species. For example, this is important because for *Az. spinosum* different ribotypes have been shown to differ in toxin production and toxin profile (Tillmann et al. 2019). A detailed rationale for the Amphidomatacean species and ribotype designation is given in the Appendix B~~especially for the accurate identification of ribotypes~~.

2.6~~7~~ Toxin identification and quantification

Filters were repeatedly rinsed with 500 µL methanol until complete discoloration of the filters. The methanolic extracts were transferred to a spin-filter (0.45 µm pore-size, Millipore) and centrifuged at 800 × g for 30 s, followed by transfer to autosampler vials and stored at −20 °C until analysis. Toxin analyses were performed using high performance liquid chromatography coupled to tandem mass spectrometry HPLC-MS/MS in the selected reaction monitoring (SRM) mode for the detection of known AZA variants. In addition, precursor experiments of the ions *m/z* 348, 350, 360, 362, and 378 were carried out to find potentially new AZA variants. Screened mass transitions and instrument parameters are detailed in Tillmann et al. (2021).

~~2.6~~ **2.7 Lagrangian ~~analysis~~simulations and Finite Size Lyapunov Exponent analysis**

To explore the physical mechanisms that might explain the concentration of amphidomatacean measured in the two locations sampled we used two complementary analysis: Lagrangian advection of virtual particles and Finite Size Lyapunov Exponents (FSLEs).
The first technique consists on the analysis of trajectories of virtual neutrally buoyant particles that were obtained with an algorithm that represents the advection process caused by surface currents. The advection equation:

$$\boldsymbol{X}(t + \Delta t) = \boldsymbol{X}(t) + \int_{t}^{t+\Delta t} \boldsymbol{v}(\boldsymbol{x}, \tau)\, \mathrm{d}\tau \tag{1}$$

where X is the three-dimensional position of a particle, $v(x, \tau)$ is the three-dimensional velocity field, is integrated using a fourth-order Runge–Kutta scheme. Particles were released at the surface along 46°S and every 0.05 grades in the four regions indicated in the **Appendix D**. The algorithm computes the particle positions based on initial location and knowledge of the velocity field. A time step of one hour was considered. The accuracy of the trajectories obtained relies on the accuracy of the velocity field used. For this experiment we considered geostrophic velocities obtained from satellite altimetry. In the northern portion of the Argentine continental shelf, such surface velocities showed to be well correlated with in situ current measurements (Lago et al., 2021). We therefore assume that the surface dynamics can be represented by satellite altimetry derived data and use it as the input velocity field for the algorithm to advect the virtual particles. ~~To compute trajectories of virtual particles (see Sect. 2.6 below) and Finite Size Lyapunov Exponents (FSLE, Sect. 2.7) we used Absolute Dynamic Topography (ADT) maps and~~ Ggeostrophic velocities derived from gridded Absolute Dynamic Topography (ADT) ~~the ADT maps. Gridded ADT maps~~ of daily temporal resolution and ¼ of degree spatial resolution maps were ~~obtained~~downloaded from CMEMS (https://marine.copernicus.eu/).

The second technique consists on the analysis of FSLEs images with a spatial resolution of 1/25° grid that were obtained from AVISO (https://www.aviso.altimetry.fr). FSLEs ridges approximate the so-called Lagrangian Coherent structures which are the generalization of stable hyperbolic trajectories of time independent flow. They are defined as the larger eigenvalues of the Cauchy-Green strain tensor of the flow map. FSLEs are strongly linked with the exponential rate λ of separation of two neighbouring particles during a time advection t:

$$\lambda = t^{-1} \log(\delta f/\delta 0) \hspace{3cm} (2)$$

where δ0 and δf are the initial and final separation distance which are fixed before computation. FSLEs are commonly used as an indicator of frontal activity and stirring intensity (d'Ovidio et al., 2004). Relatively large FSLEs values are associated with formerly distant water masses, whose confluence creates a transport front (d'Ovidio et al., 2004; d'Ovidio et al., 2009). Fronts identified as maxima (ridges) of FSLEs have a convergent dynamics transverse to them, so that passive particles in their neighbourhood are attracted to the front and then advected along it (Della Penna et al., 2015).

~~We used the geostrophic currents computed from satellite altimeter data (see Sect. 2.3) and an algorithm that represents the advection process caused by those currents to represent the trajectories of virtual neutrally buoyant particles. Particles were released at the surface along 46°S every 0.05 grades and in the four regions indicated in the **Appendix D**. The algorithm used for the advection process is fully described in Haller and Beron-Vera (2012). The algorithm computes the particle positions based on initial location and knowledge of the velocity field. Therefore, the accuracy of the trajectories obtained relies on the accuracy of the geostrophic velocity field obtained from satellite altimetry. In the northern portion of the Argentine continental shelf, such surface velocities showed to be well correlated with in situ current measurements (Lago et al., 2021). We therefore assume that the surface dynamics can be represented by satellite altimetry derived data and use it as the input velocity field for the algorithm to advect the virtual particles.~~

~~**2.7** **Lyapunov coefficient analysis**~~

In order to examine meso- and submesoscale frontal structures during phytoplankton blooms, daily ~~finite size Lyapunov exponent (~~FSLEs~~)~~ images from November 10 to 25 were analysed. ~~Additionally, T~~the daily images were used to create a video ~~with Filmora v.11~~ (available in the **Appendix D**) to illustrate the daily evolution of the FSLEs in the area where the phytoplankton bloom developed. ~~The FSLE is obtained by measuring the backward in time divergence of initially nearby particles and it is commonly used as an indicator of frontal activity and stirring intensity. Relatively large FSLE values are associated with formerly distant water masses, whose confluence creates a transport front (d'Ovidio et al., 2004; d'Ovidio et~~

al., 2009). Fronts identified as maxima (ridges) of FSLEs have a convergent dynamics transverse to them, so that passive particles in their neighbourhood are attracted to the front and then advected along it (Della Penna et al., 2015).

3. Results

3.1 Satellite-derived chlorophyll *a* during the sampling period

During the Gayoso cruise (12-19 Nov. 12-19, **Fig. 2a**), a uniform, large band of high surface Chl-*a* concentration (between 4.0 and 32 mg m^{-3}) expanded over the mid and outer shelf. During the Agujero Azul cruise (20-27 Nov. 20-27, **Fig. 2b**), the band of Chl-*a* disaggregated and showed lower intensity (1.5 and 34 mg m^{-3}), but the Chl-*a* concentration was still high in the area of the sampling stations **(Fig. 2b)**. The distribution of Chl-*a* concentration during the time of the cruises is indicative of mid-late spring phytoplankton bloom (November) over the Patagonian shelf which typically exhibits higher concentration south of 43°S following the thermal stratification (**Fig. 2a, b**). The SST showed warming of the inner-mid shelf waters north of 44°S over the 8-day average periods (around 14.5 °C in mid shelf) **(Fig. 2c, d),** but the SST remained yet constant at the sampling stations' area around 11 °C) (see **Fig. 4** and **Table 1**). The daily images of Chl-*a* from 14 to 25 November 14 to 25 (**Fig. 3**) showed an abrupt proliferation of phytoplankton on November 15 which notably intensified on November 16 when it reached the maximum concentrations during the studied period (up to 35 mg m^{-3}). On this date, the extraordinary bloom of Amphidomataceae was sampled at GA01. During the following days, the chlorophyll levels remained high in the area but became more disaggregated into variable patches. On November 25, the sampling day at station AA09, Chl-*a* had decreased at station GA01 but remained intense in the area at AA09, **(Fig. 3)** with still extraordinary densities of Amphidomataceans at station AA09.

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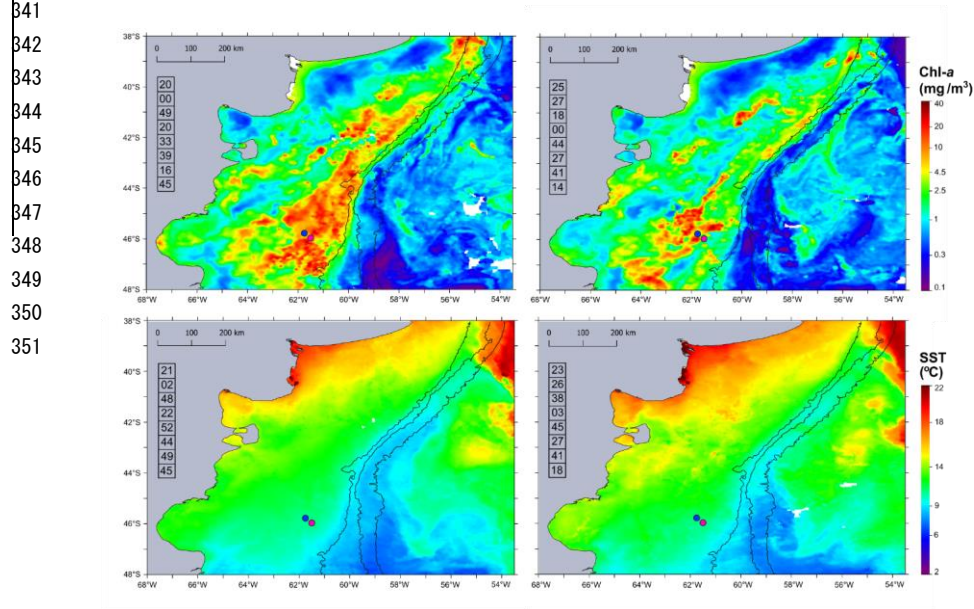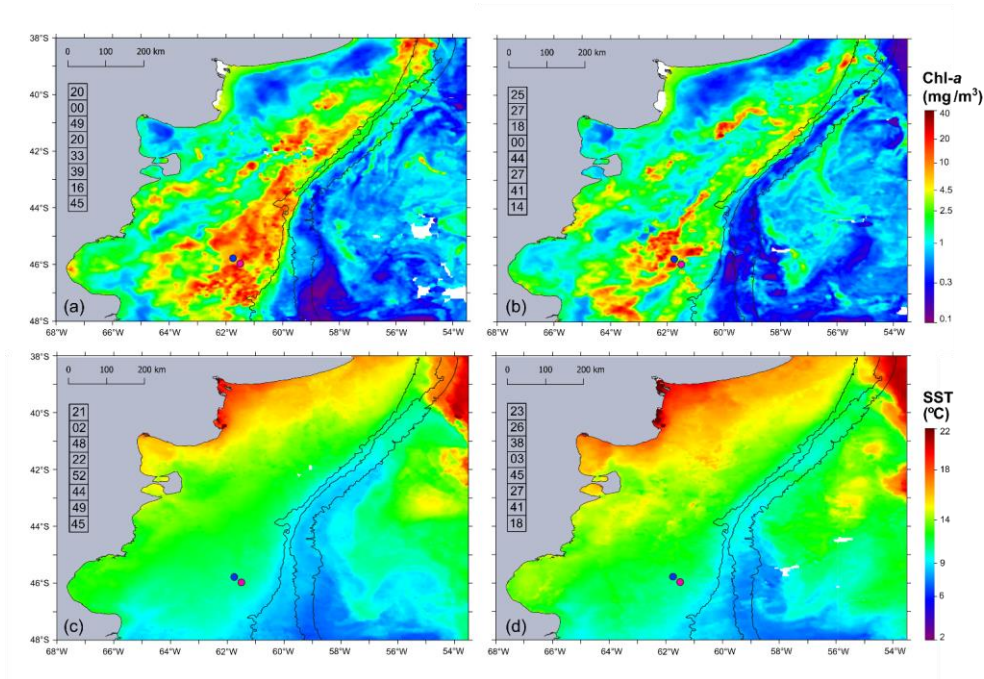
Figure 2: **Eight-day time mean of (a, b) satellite chlorophyll *a* (Chl-*a* in ~~mg m⁻³~~mg/m3) and (c, d) sea surface temperature (SST in °C) during the sampling period: (a) and (c) ~~12-19~~ November 12-19, 2021, and (b) and (d) ~~20-27~~ November 20-27, 2021. From left to right, isobaths of 200 m, 1000 m and 1800 m are shown. The numbers in the column on the left side of each panel correspond to the percentage of cloud-free pixels in each daily satellite image. The sampling stations GA01 (pink dot) and AA09 (blue dot) are shown.**

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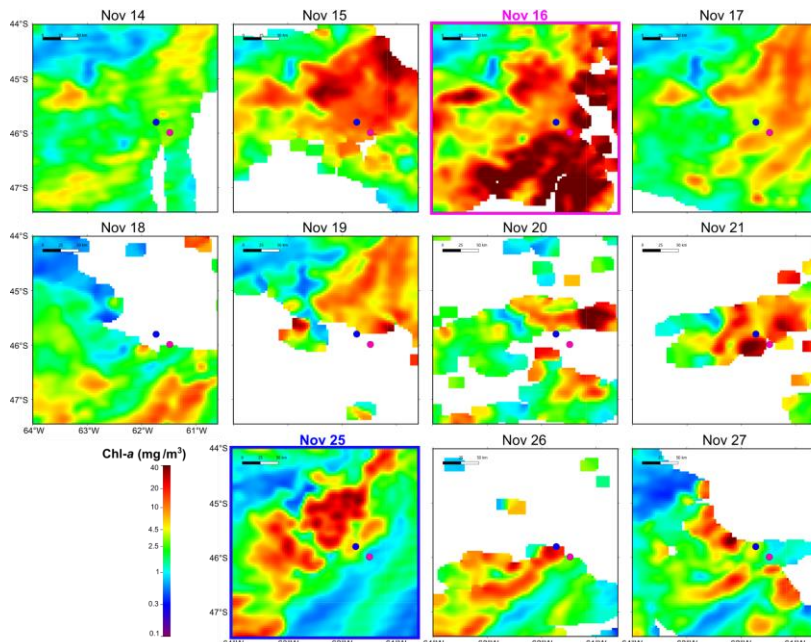
Figure 3: Daily satellite-derived surface chlorophyll-*a* in the area of the sampling stations: GA01 (pink dot, sampled on November 16, 2021), and AA09 (blue dot, sampled on November 25, 2021).

3.2 *In situ* biogeochemical properties and water column structure

A reddish water discoloration was observed in the bloom area (45.5-46°S, 62-61°W) during the cruises. Surface water temperature and salinity remained similar at both stations GA01 and AA09 sampled with a ten-day interval **(Fig. 4, Table 1)**. Moreover, both stations displayed the same vertical structure in terms of temperature and salinity **(Fig. 4)**, the mixed layer depth (MLD), and the subsurface chlorophyll maximum (SCM) **(Table 1)**. *In situ* chlorophyll-*a* concentration at the surface was 20 µg L^{-1} at GA01 and 4.5 µg L^{-1} at AA09 **(Table 1)**. Dissolved inorganic macronutrients in surface waters were similar in both stages of the bloom, except for nitrate and silicate, which were higher at AA09. In particular, the silicate recovered by ~5 times towards the advanced stage of the bloom.

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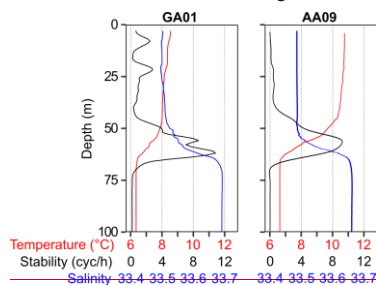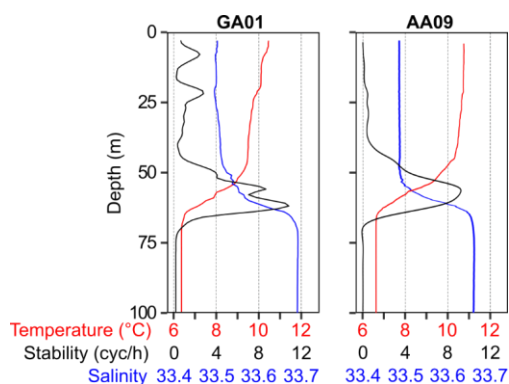
Figure 4: Vertical profiles of temperature, salinity, and stability measured at the two blooming stations, GA01 (sampled on ~~16~~ Nov. 16) and AA09 (sampled on ~~25~~ Nov. 25). Strong stratification is denoted by the water column stability (Brunt-Väisäla buoyancy frequency) which indicates a mix layer depth (MLD) around 60 m.

Table 1: Surface values of physical and chemical variables measured at stations GA01 and AA09. Sea surface temperature (SST), sea surface salinity (SSS), and concentration of *in situ* chlorophyll-*a* (Chl-*a*), and macronutrients. The depth of the subsurface chlorophyll maxima (SCM) and the mix layer depth (MLD) is also displayed.

	GA01	**AA09**
Date	16-Nov	25-Nov
SST (°C)	10.5	10.9
SSS	33.5	33.5
SCM (m)	15	10-25
MLD (m)	60	55
Chl-a (μg L^{-1})	20.0	4.5
Nitrite (μM)	0.48	0.33
Nitrate (μM)	1.77	3.57
Ammonium (μM)	1.32	1.32
Silicate (μM)	4.54	22.35
Phosphate (μM)	0.33	0.44

3.3 Multispecificity of the Amphidomatacean bloom and azaspiracids

Total protistan plankton of cell size larger than 5µm reached up 31.68 x10^6 cells L^{-1} at station GA01, and 10-days later at station AA09, the abundance was 13.69 x10^6 cells L^{-1}. Of all this total protist abundance, the Amphidomataceae clade represented up to 99 % and 98 %, respectively (**Fig. 5**), mostly dominated by the non-toxigenic species *Azadinium spinosum* ribotype C and *Azadinium dalianense* (**Fig. 5 and 6**)*, representing together >95 % of total Amphidomataceae (**Fig. 5**). Taxonomic identification up to species level was possible after exhaustive morphological examination of cells under light microscopy (**Fig. 5**) and scanning electron microscopy (**Fig. 6**). In the **Appendix A**, more micrographs of Amphidomataceans are shown taken under light microscopy and scanning electron microscopy (**Figs. A01 to A10**), along with the rationale for the Amphidomatacean species designations. The ITS-based metabarcoding of species diversity detected in the field samples at GA01 and AA09 are also shown in the **Appendix B**, which also support the dominance of *Azadinium spinosum* ribotype C and *~~Azadinium~~ Az. dalianense*. In the detailed counting of protistan species in 10-mL subsamples, the well-recognized Amphidomataceae species under inverted light microscopy for their individual quantification were: *~~Azadinium~~ Az. spinosum*

ribotype B and ribotype C, *Az. dalianense*, *Az. obesum,* and the smaller taxa *Az. dexteroporum, Amphidoma parvula* and *Am. languida* (**Fig. 5**). This distinction was based on morphological aspects combining the cell size and shape, such as the length/wide relation, and other taxonomic aspects. For instance, a slender shape: *Az. spinosum* ribotype B; round: *Az. obesum*, short, tiny: *Az. dexteroporum*, ~~*Amphidoma*~~ *Am. parvula* and *Am. languida*; with a bump in the hypotheca and a pyrenoid in the hyposome: *Az. dalia~~na~~ense*; with a spine in the hypotheca: ~~*Azadinium*~~ *Az. spinosum* ribotype C, *Az. dalianense*, and *Az. dexteroporum*. These and other taxonomic features were further examined by SEM (**Fig. 6**), for example the number and arrangement of the thecal plates, the presence of thecal pores, etc. (see the **Appendix A**). Finally, other protists than Amphidomataceans (**Fig. 7**) contributed up to 1.0 and 2.2 % of the total abundance at stations GA01 and AA09, respectively. Most of the other protists were heterotrophic and mixotrophic dinoflagellates and ciliates. No diatoms were observed in the field samples. For an overview of the pure Amphidomataceae bloom in the field samples (e.g. no mucus formation, no aggregates, cells undergoing cell division), low-magnification micrographs obtained through light microscopy are presented in the **Appendix C**. The screen of all known and potentially novel variants of azaspiracids, which are produced by Amphidomataceae, revealed the presence of solely azaspiracid-2 (AZA-2) in both stages of the bloom, with field values of 2122 pg L^{-1} at GA01 and 620 pg L^{-1} at AA09.
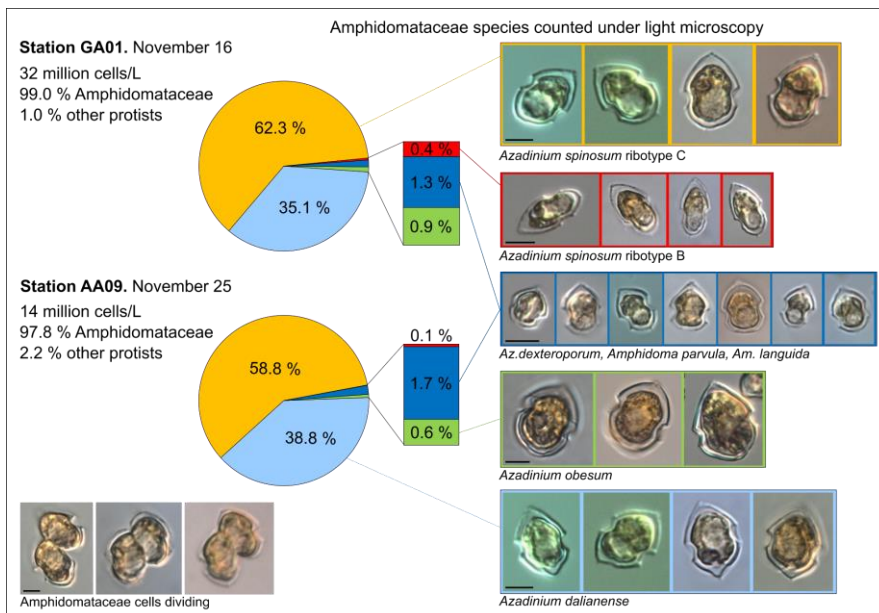
Figure 5: Relative abundance (in %) of the Amphidomataceae species identified under light microscopy at stations GA01 and AA09. The total abundance of protists at each station was 32 million cells L^{-1} and 14 million cells L^{-1}, respectively. From the total cells counted, Amphidomataceae represented up to the 99.0 % and 97.8 %, respectively. The colours in the pie charts correspond to the same species at both stations. Scale bar: 5 μm.

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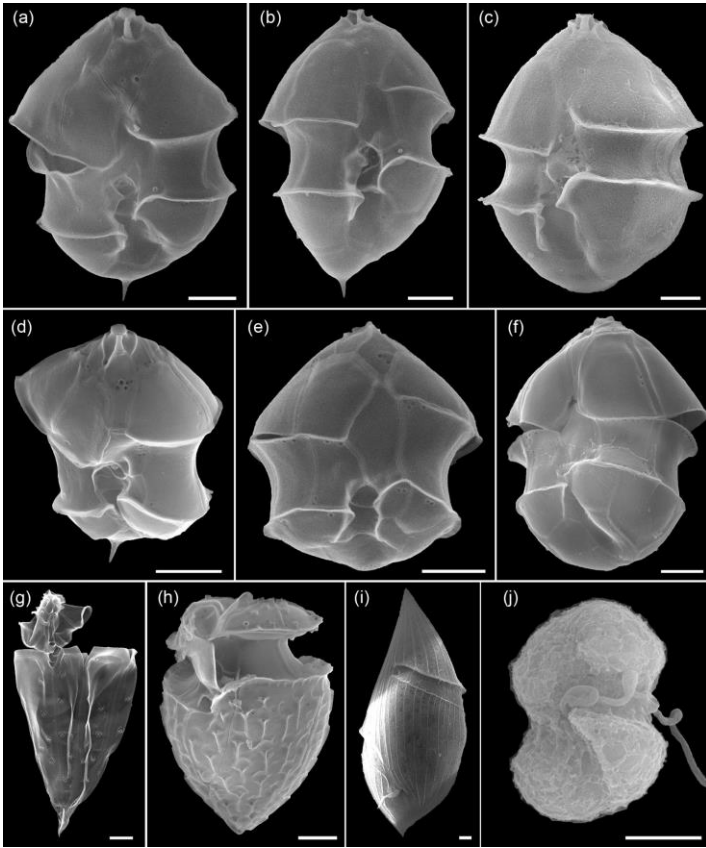
Figure 6: Scanning electron microscopy of Amphidomataeae species (a-f) and other dinoflagellates (g-j) at stations GA01 and AA09. (a) *Azadinium spinosum,* **(b)** *Az. dalianense*, **(c)** *Az. obesum*, **(d)** *Az. dexteroporum*, **(e)** *Amphidoma parvula*, **(f)** *Am. languida*, **(g)** *Oxytoxum gracile*, **(h)** *Oxytoxum laticeps*, **(i)** *Gyrodinium* **sp., (j) unidentified gymnodinoid species. Scale bars = 2 μm. See the Appendix A for more micrographs of Amphidomataceans and for evidence and rationale for the Amphidomatacean species designations.**

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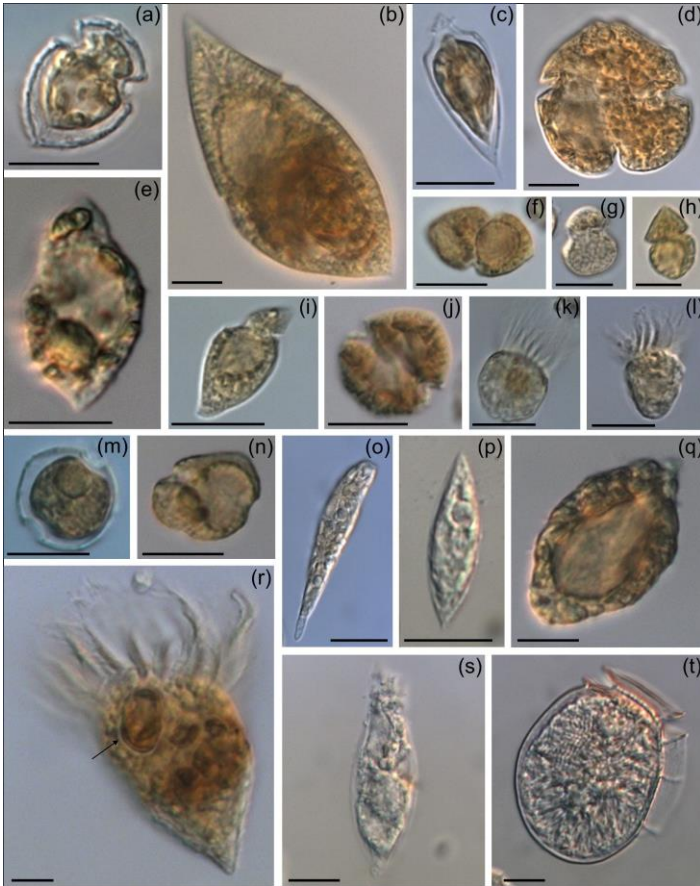
Figure 7: Micrographs of other protists (1.0 - 2.2 % of total abundance) present in the bloom. (a) *Oxytoxum laticeps*, **(b)** *Gyrodinium spiralis*, **(c)** *O. gracile*, **(d)** *Karenia* **sp., (e) unidentified dinoflagellate, (f)** *Gyrodinium sp.*, **(g-i) unidentfied gymnodinoid cells, (k-l) naked ciliates, (m)** *Peridinella* **sp., (n) Kareniacea-type cell, (o) Euglenophyte, (p)** *Lessardia elongata*, **(q)** *Torodinium robustum*, **(r) ciliate with an Amphidomataeae cell (arrow), (s) ciliate, (t)** *Dinophysis* **sp. Scale bars~~:~~ = 10 μm, except in photos (g) and (h), scale bar~~s~~ = 5 μm.**

3.4 Surface currents, particle advection model and ~~Lyapunov~~ description of the frontal systems

The mean and standard deviation of the ADT during the sampling period (~~16 to 28~~ November 16 to 28, 2021) evidenced a mesoscale anticyclonic eddy of about 100 km in diameter in the area where the Amphidomataceae bloom was observed at the two sampling locations (**Fig. 8**). In addition, the modelled trajectory of the particles released along the zonal transect at 46°S on November 10 showed high retention at the blooming area over the continental shelf after running for 20 days. On the contrary, high advection within the flow of the Malvinas Current was evidenced (**Fig. 9**). Notably, in the eddy area, particles slightly displaced southwards, remaining trapped in the area after 20 days since their release. The particles advected by geostrophic velocities suggest that the anticyclonic eddy acted as a potential mechanism to retain the Amphidomataceae bloom within the same location during the two synoptic samplings (**Fig. 9**). All the other particles released East and West of the eddy

displayed a different behaviour (**Fig. 9**), also shown in the **Appendix D** where four parcels of particles were advected from ~~16 to 25~~ November 16 to 25. West of the eddy they were advected northward at rather slow speed (average 5 cm/s) while East of the eddy they increased their speed towards the north as they approached the continental slope. Within the core of the Malvinas Current, speeds were as large as 0.8 m s^{-1}~~80 cm/s~~ (**Fig. 9**).
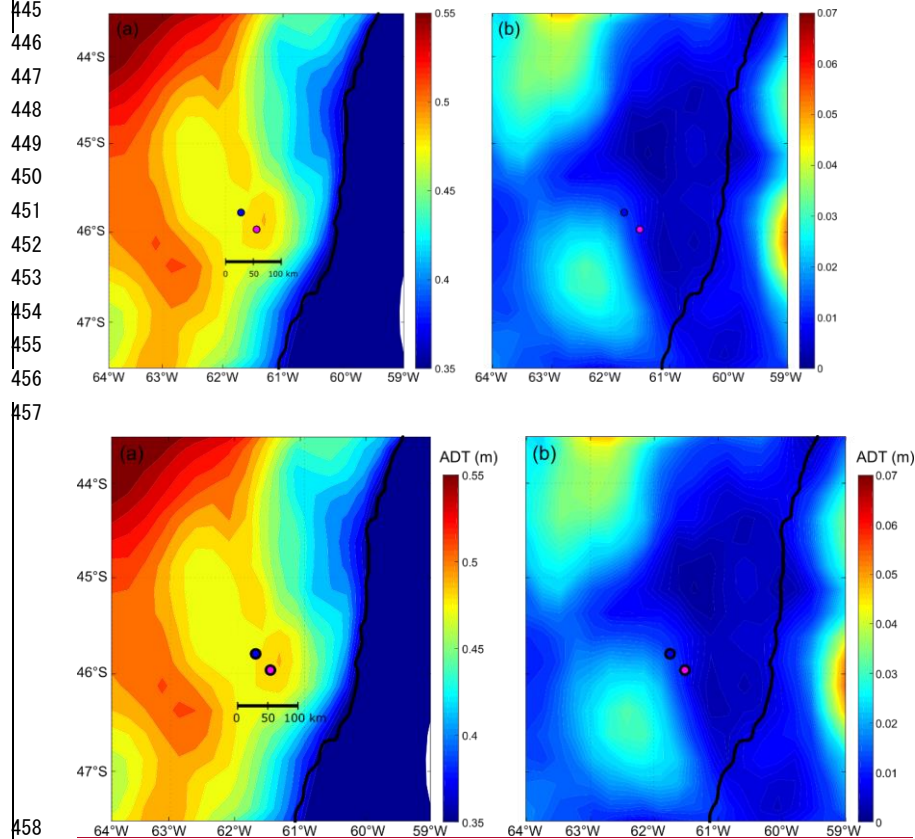
Figure 8: (a) Mean and (b) standard deviation (std) of the Absolute Dynamic Topography (ADT, in meters) displayed in colour scale, during the period November 16 to November 28, embracing the sampling period on both bloom stations: GA01 (pink dot, Gayoso cruise) and AA09 (blue dot, Agujero Azul cruise).

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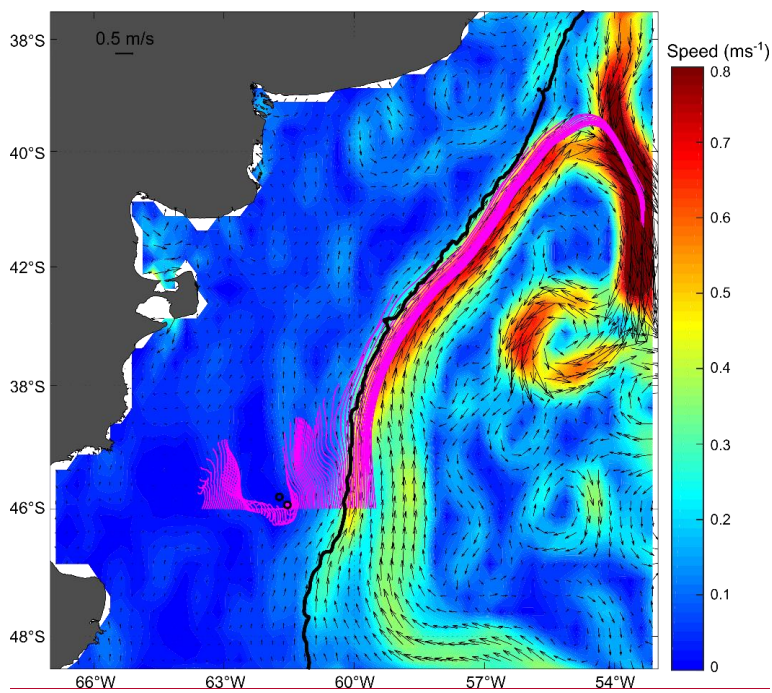
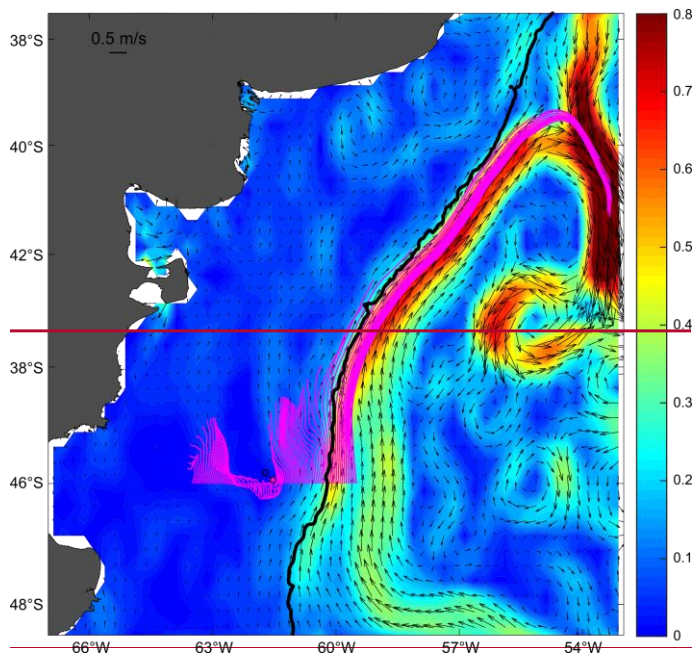
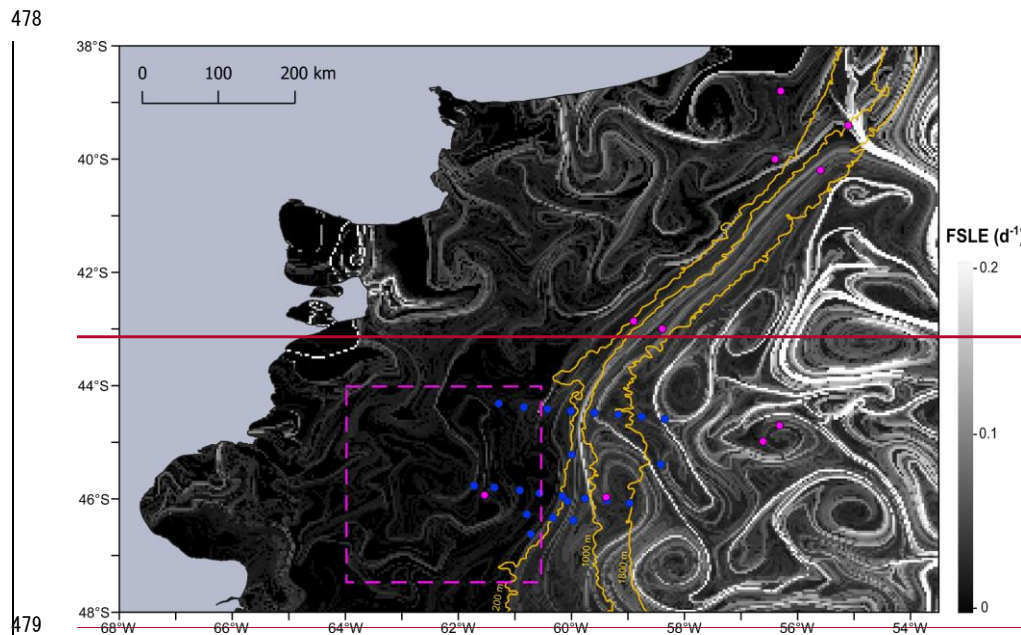


Figure 9: **Particle advection after 20 days since particle release on November 10, 2021. Initial points: one particle every 0.05 grades along 46°S. Note the high retention and the distinct behaviour of particles in the area of the sampling stations: GA01 (pink dot) and AA09 (blue dot). Satellite altimetry geostrophic velocities averaged for the same 20 days are also displayed with black vectors; the magnitude of the averaged speeds is represented with colours in the background (m/s).**

Moreover, the finite-size Lyapunov exponent (FSLE) ridges highlighted the stirring and hydrologically complex nature of the Southwestern Atlantic Ocean, associated with the high hydrographical heterogeneity of the oceanic waters (**Fig. 10**). Although FSLE were less intense over the shelf than in the adjacent oceanic waters, in the area where the Amphidomataceae bloom was sampled (pink dashed square in **Fig. 10**), two relatively strong FSLE ridges consistently kept both bloom stations within the same water mass during the period between the two synoptic samplings (**Fig. 11**).

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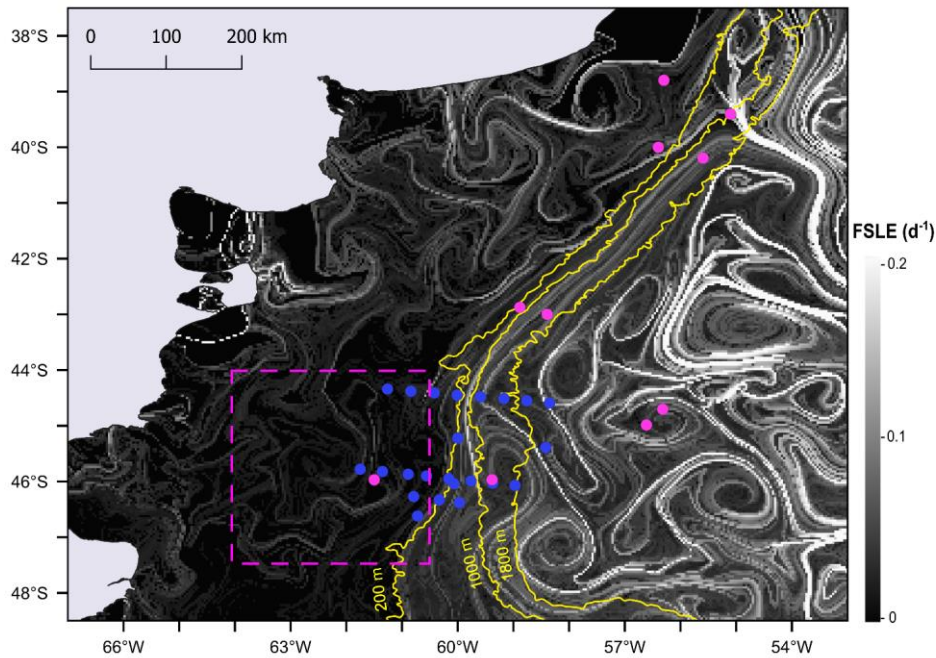
Figure 10: Fronts identified as Finite-size Lyapunov Exponent (FSLE) ridges computed for November 16, shown against a grayscale background. All sampling stations from the Gayoso cruise (pink dots) and Agujero Azul cruise (blue dots) are indicated. November 16 corresponds to the sampling of the Amphidomataceae bloom at station GA01 (pink dot within the square marked by a dashed pink line). The square highlights the area shown in Fig. 11.

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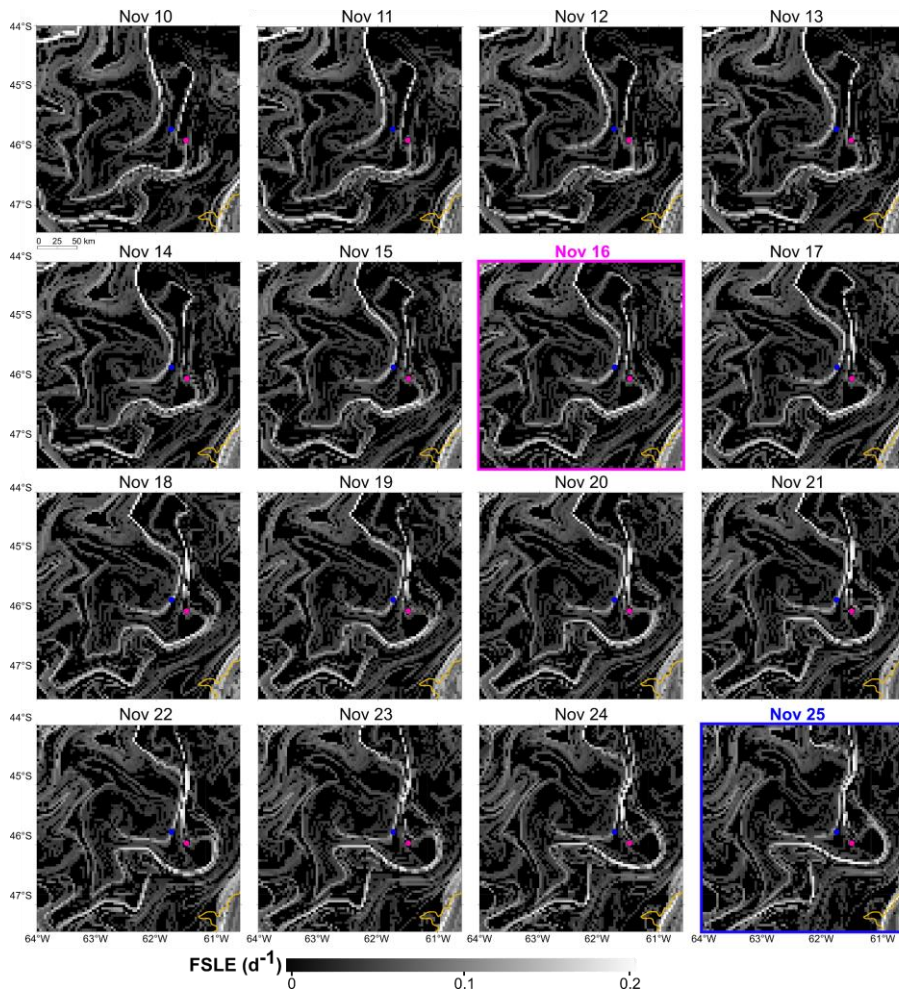

488 **Figure 11: Finite-size Lyapunov Exponent (FSLE) ridges in the area of the two locations with the Amphidomataceae**
489 **bloom: GA01 (pink dot) sampled on November 16, and AA09 (blue dot) sampled on November 25. See the Video of**
490 **these daily images in the Appendix D. The 200 m isobath is indicated in yellow.**

491

492 **4. Discussion**

493

494 This study is unique from both biological and physical perspectives due to the following factors: (i) the Amphidomataceae
495 bloom observed in spring 2021 in the Argentine Sea, with up to 32 million cells per litre, represents the largest bloom of this
496 clade ever recorded globally; (ii) unusual sampling in offshore shelf waters with two vessels over a ten-day interval allowed
497 for synoptic observations of the bloom at two active developmental stages; (iii) simultaneous ecological characterization of
498 the bloom and surface currents and fronts provided insights into patch stirring and the short-term evolution of the bloom; (iv)
499 field quantitative abundance data for Amphidomataceae are rare, and to our knowledge, this is the first detailed description of
500 species abundance in field samples, combining light microscopy, electron microscopy, and metabarcoding; (v) the relatively

low abundance of *Azadinium spinosum* Ribotype B indicated high AZA-2 cell quotas; (vi) the fine-taxon assessment of the Amphidomataceae bloom revealed biogeographical patterns and strain-specific toxic potential; and (vii) the use of interdisciplinary approaches sheds light on the bio-physical coupling underlying the persistence and horizontal transport of this extraordinary bloom in offshore shelf waters.

4.1. Amphidomataceae blooms ~~in~~over the Patagonian shelf

The phytoplankton spring and summer blooms in the Patagonian shelf display a southward progression related to the seasonal thermal cycle. In early spring (September-October), the water column stratifies north of ~45°S, favouring the proliferation of opportunistic micro-diatoms in the nutrient-rich, well-lit surface layers (Ferronato et al., 2023). South of ~45°S, the bloom initiates later in spring-early summer (December-January) and continues until autumn (March) (revised in Guinder et al., 2024). The distribution of surface Chl-*a* ~~concentration during the time of the cruises~~ was indeed indicative ~~is indicative~~ of mid-late spring phytoplankton bloom (November) over the Patagonian shelf, ~~which typically exhibits higher concentration south of 43°S~~ following the thermal stratification (**Fig. 2a, b**). Here, blooms of nanoflagellates and dinoflagellates are triggered by combined vertical stability and nutrient-depleted surface waters (especially silicates) after the early-spring blooms of large diatoms (Balch et al., 2014; Carreto et al., 2016~~; Ferronato et al., 2024~~). During the Gayoso cruise, detailed analyses of plankton communities and microbial carbon fluxes at all sampling stations revealed a latitudinal shift from more heterotrophic to more phototrophic activity toward the south (Ferronato et al., 2025; Gilabert et al., 2025). The massive proliferation of Amphidomataceans in mid-November 2021 was in line with this successional pattern. These nano-dinoflagellates bloomed in a stratified water mass with a deep mixed layer depth (~60 m). While dinoflagellates are less effective at nutrient resorption compared to diatoms, they can move throughout the stable water column to find light and nutrients (Glibert, 2016), especially at low phosphate levels (Lin et al., 2016) as observed at the stations GA01 and AA09 with high nitrate-to-phosphate ratios.

The success of the multispecific bloom of Amphidomataceae may be attributed to a combination of multiple intrinsic and extrinsic factors. For instance, their small cell-size, unique swimming modes and the production of azaspiracids may have alleviated grazing pressure (Tillmann et al., 2019). In fact, in the fixed samples, it was observed that many cells were obviously active and undergoing cell division and that these nanoflagellates were overwhelmingly the predominant photosynthetic protists responsible for the high chl-*a* levels, with negligible abundance of micro-grazers accompanying the bloom development. The presence of less than 2% of other protists (mixotrophs and heterotrophs) could be also related to a delay in the recovery of predators following the early blooms of microdiatoms, as well as to an abrupt development of the Amphidomataceae bloom, which may have prevented micrograzers from taking advantage of the available food. Another observation supporting the active persistence of the bloom was the pristine condition of the microenvironment surrounding the dense populations of Amphidomataceae (see **Appendix C**), with no aggregates or mucus formation typically observed in the late stages of blooms (Genitsaris et al., 2021). Furthermore, no competitors for light and nutrients were detected in the samples; specifically, no diatoms were found during microscopic examination. This may explain the rapid recovery of silicates (from 5 to 22 µM) over the 10-day persistence of the Amphidomataceae bloom, as these silicates were not being utilised by silicate-requiring species. A similar observation was noticed during a bloom induced by iron-fertilization in the Southern Ocean, where diatoms predominated and silicate levels decreased from 10 to 6 µM within the bloom patch over the course of 12-days (Abraham et al., 2000).

4.2. Highest abundance ever recorded for a bloom of Amphidomataceans

With 17 described species of *Azadinium* and 14 of *Amphidoma*, the Amphidomataceae represent a small but diverse group of dinoflagellates. Most of these species are very similar in size and shape, which makes the qualitative identification and, in

particular, the species specific quantification in field samples difficult. Hence, characterising the cryptic species of the multispecific bloom of Amphidomataceae in this study represented a major challenge, where a reliable species identification was achieved by the combination of several diagnostic details using electron microscopy. Light microscopic counting of fixed samples provided high quantitative accuracy, but reliable species identification was only possible in a few cases. Complementing the identification by microscopy with metabarcoding specifically targeting the internal transcribed spacer (ITS1) region (Liu et al., 2023), allowed for the detailed characterization of Amphidomatacean species diversity in the field samples. By combining the three approaches we were able to identify ribotypes and the toxic species, including previously described Amphidomataceae species for the Argentine Sea (Fabro et al. 2019; Tillmann et al., 2019; 2021) or species still undescribed in the global seas (see the **Appendix A**).

In the North Atlantic, AZA-1 (and its producing species, *Az. spinosum* ribotype A), has been identified as one of the most prevalent toxins among a wide range of AZA variants (Tillmann et al. 2021). Bloom density of Amphidomataceae around Ireland have been reported as 8.3 x 10^4 cells L-1 for *Az. spinosum* (Wietkamp et al. 2020) and 47 x 10^5 cells L^{-1} for *Amphidoma languida* (McGirr et al. 2022) and a small bloom of *Am. languida* in the North Sea with 1.2 10^5 cells L^{-1} has also been described (Wietkamp et al. 2020). In this region, cases of human intoxication with azaspiracids have been linked to the consumption of contaminated mussels from the Irish West coast, where blooms in the shelf-break area can reach coastal shellfish beds through wind-driven advection (Raine, 2014). Notably, in the Argentine Sea, only AZA-2 has been detected in field and culture samples (Turner and Goya, 2015; Fabro et al., 2019; Tillmann et al., 2019; Guinder et al., 2020), and so far, no poisoning events have been attributed to AZAs.

In this study, relatively high levels of solely AZA-2 were detected in bloom samples. A toxin profile of solely AZA-2 is up to now only known for the Argentine strain H-1-D11 of *Azadinium spinosum* ribotype B (Tillmann et al., 2019), and this ribotype was also identified in the present bloom. Relating AZA quantities to the relatively low abundance (0.1 to 0.4 % of total Amphidomataceae) of *Az. spinsoum* ribotype B revealed AZA-2 cell quotas of 17-42 fg per cell, which is an order of magnitude higher that the cells quota of 2 to 9 fg per cell for strain H-1-D11 grown under laboratory conditions (Tillmann et al. 2019). In fact, *Az. spinosum* of ribotype A producing AZA-1 and -2 is the primary causative agent of AZA poisoning in Europe (Tillmann, 2018). However, the large majority of cells of *Az. spinosum* in the present bloom sample is from ribotype C, which is, based on analyses of several strains from Argentina, non-toxigenic (Tillmann et al 2019), and all globally available strains (including strains from Argentina) of the other co-dominant species in the bloom, *Az. dalianense*, (Tillmann et al., 2019), also do not produce azaspiraids. In the Chilean continental shelf in the SE Pacific, AZAs have been detected in scallops and mussels (López-Rivera et al., 2010), but no intoxication events or large blooms of this clade have been documented (Iriarte et al., 2023). Other *Azadinium* cells have been observed widely distributed in inner seas along the coast of Chile, but with no further characterization of their potential toxicity (Iriarte et al., 2023; Corredor-Acosta et al., 2025). Only *Az. poporum* has been described as an AZA producer in Chilean waters (Tillmann et al., 2017a). Likewise in Peru, a relatively high bloom (up to one million cells per liter) of *Az. polongum* was detected in the summer of 2014, with no AZA production (Tillmann et al., 2017b). Although the continental shelves of Chile-Perú and Argentina have different hydrology, both span similar latitudinal gradients along the South American coasts and are influenced by the Humboldt and Malvinas Currents, which share a common origin in the Circumpolar Antarctic Current. Strikingly, both shelves exhibit different populations of Amphidomataceae, despite the expectation that ocean currents could serve as transport pathways for HAB species, promoting their dispersion (Giddings et al., 2014). This suggests that additional factors, such as coastline configuration and bottom topography, play a role in water mixing and retention processes, and ultimately the development of HABs (Pitcher et al., 2010).

4.3. Spatio-temporal evolution of the bloom: retention and stirring

Phytoplankton bloom initiation, magnitude, and persistence rely on a host of biogeochemical and physical processes. As discussed in previous sections, the explosive onset of the bloom of multiple species of Amphidomataceae was associated with a combination of water column stability, the negligible presence of micrograzers, and the ecological traits of this group that facilitate massive proliferation. However, these conditions alone do not fully explain the persistence of this bloom, which was sampled 10 days later at a location 40 km away in the offshore waters of the Patagonian shelf, where strong surface currents were expected to disperse plankton blooms. The persistence of this extraordinary bloom, characterized by its remarkable magnitude and consistent species composition, indicates the retention and accumulation of the bloom patch within the same water mass. In addition to the biological evidence confirming the presence of the same Amphidomataceae bloom at both sampling stages, analyses of circulation through satellite altimetry showed that an anticyclone of about 100 km of diameter was the responsible feature to retain the bloom within the two sampling stations. The retention that the eddy caused was evidenced by two Lagrangian experiments. In the first one, the trajectories of virtual particles released within the eddy show that almost none of the particles escaped from the eddy during the 10 days that separated the two sampling stations. In the second one, FSLEs maps showed that no fronts separated the two sampling stations during those days. In addition to the biological evidence confirming the presence of the same Amphidomataceae bloom at both sampling stages, analyses of circulation through altimetry, particle experiments, and FSLE—an indicator of frontal activity and stirring intensity—support the conclusion that the same bloom patch was captured at both locations. Two potential scenarios could explain this: (1) the same patch remained in the area over the ten days of sampling, occupying a space area of 40 km or larger (**Fig. 12a and b**), or (2) a smaller bloom patch was initially detected at station GA01 and then transported by stirring towards AA09 (**Fig. 12c**). Sampling with higher spatial and temporal resolution of bloom initiation, development, species succession, and collapse would have been necessary to distinguish between scenarios (1) and (2). A less likely scenario is that (3) two Amphidomataceae blooms developed independently at both locations (**Fig. 12d**). This situation is improbable due to the complex physical-biological interactions that drive different bloom developments, such as variable stirring (Abraham et al., 2000; Lehan et al., 2007; Della Penna et al., 2015) and changing environmental conditions that select for different species and functional groups across diverse spatio-temporal scales (Levy et al., 2018; Hernández-Carrasco et al., 2020; Mangolte et al., 2023). Moreover, no dormant cysts of Amphidomataceae are known, which could explain population outbreaks in specific locations, as observed in frontal areas for other dinoflagellates forming HABs (Smayda, 2002; Akselman et al., 2015).

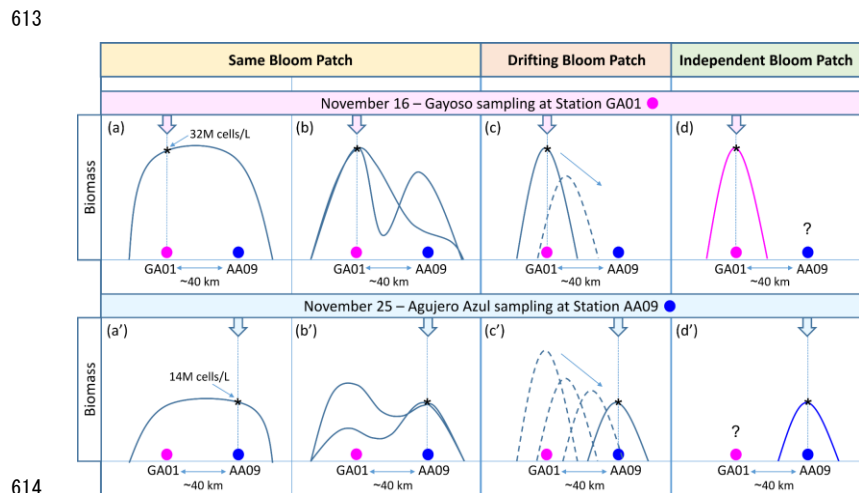



Figure 12: Hypothetical scenarios of the spatio-temporal evolution of the Amphidomataceae bloom during the 10-day

period between the synoptic sampling at the two stations 40 km apart: GA01 (pink circle) and AA09 (blue circle). The bloom's biomass was 32 million cells L^{-1} at GA01 and 14 million cells L^{-1} at AA09. In the Same Bloom Patch (a-a' and b-b'), the bloom covered an area encompassing both sampling stations. The bloom developed such that the biomass was either distributed homogeneously across the patch (a-a') or heterogeneously (b-b'), resulting in variable biomass patterns over time and space. In the Drifting Bloom Patch (c-c'), the water mass with the Amphidomataceae bloom detected at station GA01 was transported by currents towards station AA09, where the bloom was detected with lower biomass but still with high intensity. The Independent Bloom Patch (d-d') suggests that two discrete, autonomous Amphidomatacean blooms developed locally at each sampling station.

Additionally, the hydrographically complex Southwestern Atlantic creates a variety of microhabitats at the meso- and submesoscale. These include areas of upwelling, downwelling, eddies, retention, and dispersion (Becker et al., 2023; Beron-Vera et al., 2020; Salyuk et al., 2022; Saraceno et al., 2024). This spatial heterogeneity enhances the development of variable nutrient patches and phytoplankton productivity (Lehahan et al., 2017; Levy et al., 2018; Hernández-Carrazsco et al., 2020; Ser-Giacomi et al., 2023). During the Gayoso cruise, contrasting phytoplankton assemblages and bloom types were observed at all sampling locations, including distinct blooms of dinoflagellates, coccolithophores, diatoms, and nanoflagellates. These variations were related to substantial heterogeneity in surface velocities and environmental conditions across the region (Ferronato et al., 20254). It is important to highlight that the bloom of Amphidomataceae at GA01 and AA09 was not observed at any other station during the Agujero Azul cruise. In the Gayoso cruise, this clade was only observed in relatively high abundance at GA10, located in the northern area over the shelf, where species assemblages were significantly different (Ferronato et al., 2025). In this study, while the estimation of the retention of the Amphidomataceae bloom is certainly limited by the accuracy of satellite altimetry maps, the documented 100 km diameter of the eddy is a reasonable size that can be distinguished using the gridded satellite altimetry maps produced by CMEMS. Although this retention is transient, this particular circulation facilitated the massive development of the Amphidomataceae bloom, with no evidence that this patch was advected through to the Malvinas Current, as observed with drifters released east of the bloom area in spring 2021 (Saraceno et al., 2024). Our results highlight the importance of studying the evolution of toxic phytoplankton blooms on continental shelves, focusing on the bio-physical coupling that drives their patchy nature, persistence, and transport, in order to capture short-lived blooms and their potential to cause toxic outbreaks.

5. Conclusions:

We present an example where mesoscale (100 km) circulation on the Patagonian Shelf acted as the physical driver that prevented the dispersion of an Amphidomatacean bloom. As a result, the bloom persisted for 10 days in the same area. While our study provides valuable insights, it has limitations, as it only explores physical processes at a mesoscale resolution, whereas sub-mesoscale processes are crucial for understanding divergence, convergence, and the mixing of plankton and nutrients within and between water masses. We also note that the collection of samples from only two synoptic stations, along with the incidental detection of the Amphidomatacean bloom during two cruises with different ecological objectives, limited our ability to track the detailed spatial and temporal evolution of the bloom. One strategy to overcome these limitations would be continuous sampling of the bloom, which could be achieved through a Lagrangian experiment. In this approach, water samples would be collected following drifters that track surface currents. At the same time, detailed biodiversity analyses of the bloom should be conducted, including the succession of species and functional groups in relation to their differential uptake of nutrients, grazing pressure, and the potential exchange of

658 plankton and nutrients between neighboring water masses.

659

660 Our results highlight the importance of studying the evolution of phytoplankton blooms on continental shelves, focusing on
661 the bio-physical coupling that drives their patchy nature, persistence, and transport, in order to capture short-lived blooms and
662 their potential to cause toxic outbreaks.
663 Overall, this study is unique from both biological and physical perspectives due to the following factors: (i) the
664 Amphidomataceae bloom observed in spring 2021 in the Argentine Sea, with up to 32 million cells per litre, represents the
665 largest bloom of this clade ever recorded globally; (ii) unusual sampling in offshore shelf waters with two vessels over a ten-
666 day interval allowed for synoptic observations of the bloom at two active developmental stages; (iii) simultaneous ecological
667 characterization of the bloom and surface currents and fronts provided insights into patch stirring and the short-term evolution
668 of the bloom; (iv) field quantitative abundance data for Amphidomataceae are rare, and to our knowledge, this is the first
669 detailed description of species abundance in field samples, combining light microscopy, electron microscopy, and
670 metabarcoding; (v) the relatively low abundance of *Azadinium spinosum* Ribotype B indicated high AZA-2 cell quotas; (vi)
671 the fine taxon assessment of the Amphidomataceae bloom revealed biogeographical patterns and strain-specific toxic
672 potential; and (vii) the use of interdisciplinary approaches sheds light on the bio-physical coupling underlying the persistence
673 and horizontal transport of this extraordinary bloom in offshore shelf waters.

674

675 **Appendix A**

676

677 **A1. Estimation and categorization of Amphidomatacean species diversity**
678 For the present study, we combined light microscopy (LM) quantification, scanning electron microscopy (SEM) examination,
679 and metabarcoding to characterize the field samples as accurately as possible, both qualitatively and quantitatively. Generally,
680 the following species of Amphidomataceae were identified:

681

682 **A1.1. *Azadinium spinosum***
683 Specimens were identified with SEM as *Az. spinosum* based on the combination of (1) presence of an antapical spine, and (2)
684 presence of a ventral pore located on the right side of the suture of Plate 1' and 1" (Fig. A01). The vast majority of cells thus
685 identified as *Az. spinosum* had a somewhat broader cell shape. Generally, identification of *Az. spinosum* is complicated as there
686 are several different ribotypes (Tillmann et al., 2021), which notably differ in azaspiracid toxin presence and profile. In a
687 previous study from the Argentine shelf region, it was shown that 23 out of 24 isolated strains of *Az. spinosum* were assigned
688 to the non-toxigenic ribotype group C, and cells from these strains also had a somewhat broader cell shape (Tillmann et al.,
689 2019). Metabarcoding of the present bloom sample revealed that the most common sequences showed a high match with
690 sequences of these Argentine ribotype C strains. A dominance of non-AZA-producing *Az spinosum* (ribotype C) also aligns
691 with the finding that no AZA-1, the marker toxin of the toxigenic *Azadinium spinosum* ribotype A strains (Tillmann et al.,
692 2021), was detected in the field samples. In the quantitative light microscopy (LM) counts, all medium-sized cells (length >
693 12 µm) with a rounded hypotheca and an antapical spine were thus categorized as *Azadinium spinosum* ribotype C.

694

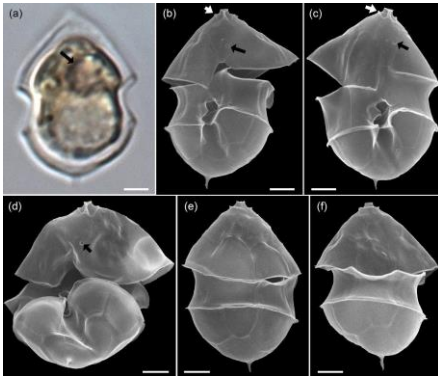



Figure A01: LM (a) and SEM (b-f) of cells of the Amphidomatacean bloom stations identified as *Azadinium spinosum* ribotype C. (b–d) Cells in ventral view. (e, f) Cells in dorsal view. Note the pryrenoid (black arrow in a), the position of the ventral pore (black arrows in b, c, d), the rim around the pore plate (white arrows in b, c) and the distinct antapical spine. Scale bars = 2 μm.

Recent studies have revealed the presence of a new, molecularly distinct species of *Azadinium* in the North Atlantic, which is morphologically indistinguishable from *Az spinosum* and is currently provisionally referred to as *Azadinium* cf. *spinosum* (Tillmann et al., 2021). Therefore, it cannot be ruled out that this species was also present in the samples, but metabarcoding showed no evidence of the presence of *Az.* cf. *spinosum* in the bloom samples.

In addition to these broader cells of *Az. spinosum* ribotye C cells, LM analysis revealed a (much rarer) presence of distinctly slender cells with an antapical spine (Fig. A02a). Such cells perfectly correspond to cell shape of a single strain H-1-D11 from Argentina identified as a ribotype B strain of *Az spinosum* (Tillmann et al., 2019), and this strain is depicted here for comparison (Fig. A02c, d). In SEM, specimen of slender shape lacked the rim around the pore plate (Fig. A02b), which is the morphological diagnostic feature differing in ribotype B strains from ribotype A and C strains, which all have a thick rim. Additionally, metabarcoding showed conformity of some sequences with other ribotype B *Az. spinsum* strains (e.g. 99% similarity with the Argentinean strain H-1-D11). Consequently, all slender cells with an antapical spine quantified in LM were categorized here as *Az. spinosum* ribotype B. Strain H-1-D11 from Argentina was shown to produce solely AZA-2 (Tillmann et al., 2019). As this was the only AZA detected in our field sample, this is additional support for this *Az. spinosum* ribotype B designation.

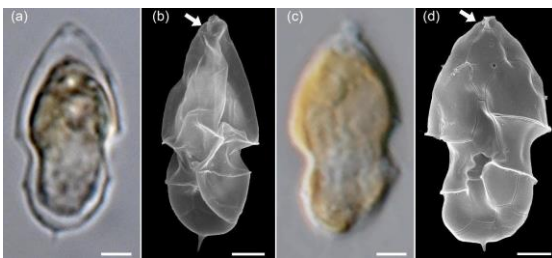
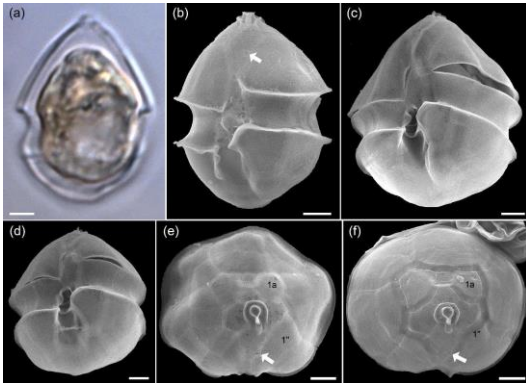



Figure A02: LM (a), and SEM (b) of cells of the amphidomatacean bloom stations identified as *Azadinium spinosum*. For comparison, LM (c) and SEM (d) of cells of strain H1-D11 of *Azadinum spinosum* ribotype B isolated from the Argentine shelf in 2015. Note the elongated cells shape, the distinct antapical spine, and the lack of a rim around the pore plate (white arrows in b and d). Scale bars = 2 μm.

721

722 **Figure A03: LM (a) and SEM (b-f) of cells of the Amphidomatacean bloom stations identified as *Azadinium obesum*.**

723 **(b–d) Cells in ventral view. (e, f) Epitheca in apical view. Note the lack on an antapical spine, the position of the ventral**

724 **pore (white arrows in b, e, f) and the lack of contact between Plates 1a and 1'' (kofoidian plate label notation) visible in**

725 **e) and f). Scale bars = 2 μm.**

726

727 **A 1.2. *Azadinium obesum***

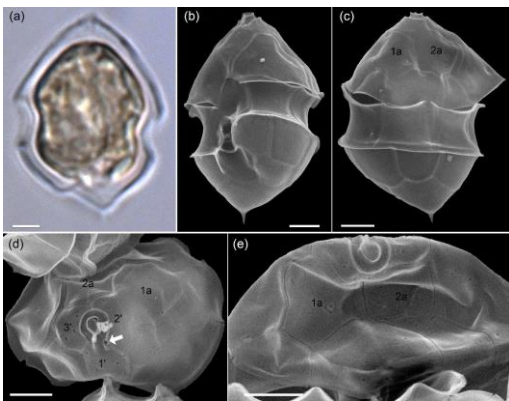
728 Cells of the non-toxigeneic species *Az. obesum* were identified in the SEM samples based on the combination of the following

729 features: (1) no antapical spine, (2) ventral pore on the right side of Plate 1', and (3) no contact between Plates 1a and 1'' (Fig.

730 A03). All such cells had a distinctly broad oval shape and were relatively large. In the light microscope, all relatively large

731 oval Amphidomataceae cells with a rounded hypotheca and no visible spine were therefore categorized as *Azadinium obesum*.

732

733

734 **Figure A04: LM (a) and SEM (b-f) of cells of the amphidomatacean bloom stations identified as *Azadinium dalianense*.**

735 **(b) Cell in ventral view. (c) Cell in dorsal view. (d) Epitheca in apical view. (f) Epitheca in dorsal view. Note the distinct**

736 **apical spine on a triangular bumpy hypotheca (a–c), the position of the ventral pore on the left side of the pore plate**

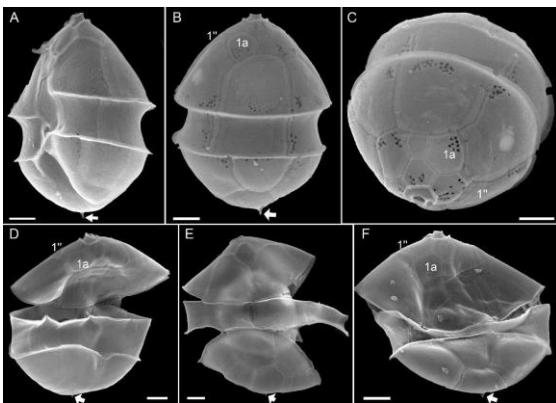
737 **(white arrow in d), the presence of only two large anterior intercalary plates 1a and 2a (c–e) and presence of only 3**

738 **apical plates in (d) (kofoidian plate label notation). Scale bars = 2 μm.**

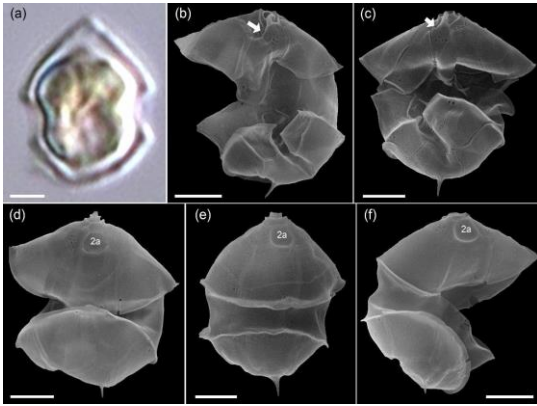
739

740 **A 1.3. *Azadinium dalianense***

741 With the rounded hypotheca, cells of *Az. obe*sum and *Az. spinosum* ribotype C were clearly distinguishable in the light

742 microscope from cells with a distinctly protruding bump, at the tip of which a small spine was present (Fig. A04a).
743 Corresponding cells detected in SEM preparations (Fig A04b–e) were identified as *Az. dalianense*, based on the combination
744 of the following features: (1) ventral pore on the left side of the pore plate, (2) asymmetrical hypotheca with a bump and a
745 distinct antapical spine, and (3) presence of only 3 apical plates and 2 anterior intercalary plates. Since all three features were
746 only rarely visible simultaneously due to the cell's orientation, the presence of the somewhat similar species *Az. perfusorium*
747 cannot be ruled out. *Az. perfusorium* also has a posterior small bump with a spine and a ventral pore located on the left side of
748 the pore plate, but it possesses 4 apical plates and 3 intercalary plates (Salas et al., 2021). However, neither the SEM nor
749 metabarcoding provided any indication of its presence in the samples. The occurrence of *Az. dalianense* in the region is well
750 documented by a series of strains isolated from the Argentine shelf in 2015 (Tillmann et al., 2019), and *Az. dalianense* was
751 also identified as part of *Azadinium* blooms in 1991(Tillmann and Akselman, 2016). Metabarcoding additionally indicated the
752 presence of two differen ribotypes of *Az. dalianense*, namely E and B as defined in Tillmann et al. (2019), where all previous
753 strains from Argentina belong to the ribotype E clade. Accordingly, the majority of reads from the bloom station *Az. dalianense*
754 were from ribotype E (represented by strains H-4-E8 and N-12-04 in the reference dataset), whereas reads of ribotybe B
755 (represented by strain IFR-ADA-01C) made only ca. 0.02 % of all *Az. dalianense* reads. All strains of *Az. dalianese*
756 representing different ribotypes collected from various regions analysed so far were non-toxigenic.
757

759 **Figure A05: SEM of yet unidentified cells of *Azadinium* sp. 1 of the Amphidomatacean bloom stations. (a) Cell in left-**
760 **lateral ventral view. (b) Cell in dorsal view. (c) Cell in apical view. (d–f) Cells in dorsal view. Note the small indistinct**
761 **apical spine (white arrow in a, b, d–f) and the lack of contact between plates 1a and 1'' (kofoidian plate label notation)**
762 **visible in b, c, d, f. Scale bars = 2 μm.**
763

764 The classification and quantification of *Az. obesum* and *Az. dalianense*, however, is complicated by the fact that a number of
765 cells of an unclear assignment were found in the samples (Fig. A05). These cells, like *Az. obesum*, (1) had a ventral pore on
766 the right side of Plate 1' and (2) no contact between Plates 1a and 1'', but unlike *Az. obesum*, they had a distinct, albeit small,
767 antapical spine. This combination of features is not known from any described *Azadinium* species, suggesting that this may be
768 a new species. However, for a complete description as a new *Azadinium* species, further investigations are necessary, ideally
769 adding sequence data and analyses of toxin production. In any case, it is clear that cells of this type will have been included in
770 the categories *Az. obesum* or *Az. spinosum* during the light microscope analyses and quantifications.

771

772 **Figure A06: LM (a) and SEM (b-f) of cells of the Amphidomatacean bloom stations identified as *Azadinium***

773 ***dexteroporum*. (b, c) Cells in ventral view. (d–f) Cells in dorsal view. Note the small size, the distinct antapical spine, the**

774 **position of the ventral pore on the right side of the pore plate (white arrows in b, c,), and the concave central intercalary**

775 **plate 2a visible in d–f. Scale bars = 2 μm.**

776

777 **A 1.4. Smaller Amphidomatacean species: *Azadinium dexteroporum*, *Amphidoma parvula*, and *Amphidoma languida***

778 While *Az. spinosum*, *Az. obesum*, and *Az. dalianense* fall into a slightly larger size class, a number of smaller *Azadinium* species

779 were identified and categorized in the samples. One of them was *Az. dexteroporum* (Fig. A06), which was identified in the

780 SEM by the following combination of features: (1) relatively small size, (2) presence of a distinct antapical spine, (3) a slightly

781 posteriorly positioned ventral pore on the right side of the ventral plate, and (4) a distinctly concave central intercalary Plate

782 2a. Metabarcoding revealed a number reads for an *Azadinium* sp. 1 with *Az. dexteroporum* as closest species suggestion (line

783 7 in Table S01), however only with rather low similarity (90-95%) compared to the reference database. Global wise, there are

784 only three available strains of *Az. dexteroporum*, and of those only one strain from the Mediterranean was identified as a

785 producer of AZA (Rossi et al. 2017). In contrast, two additional strains from the North Atlantic, which also had marked

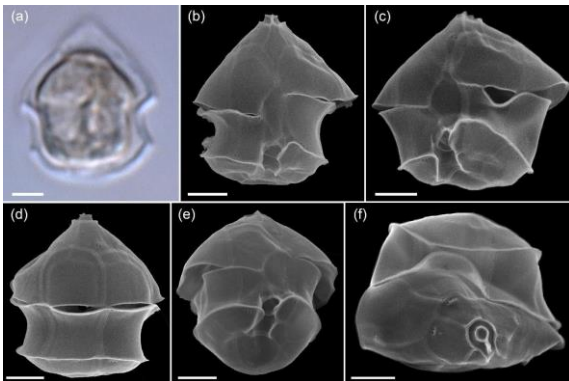
786 sequence differences compared to the Mediterranean strain, did not produce AZAs (Tillmann et al. 2020). The low similarity

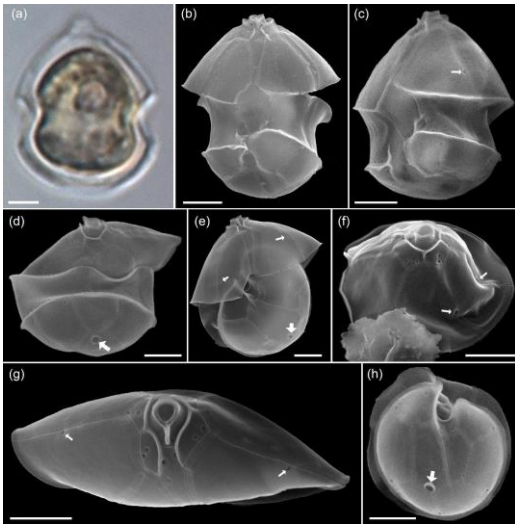
787 of *Az. dexteroprum* reads from the present bloom samples thus indicate that the local population may represent a new ribotype

788 quite distinct from the AZA-producing ribotype, and strain isolation of local *Az. dexteroporum* is needed to clarify its identity

789 and toxin production potential.

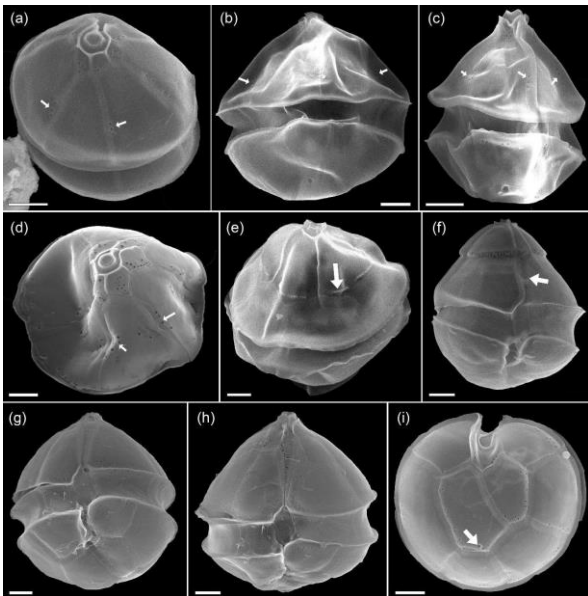
790

791

792 **Figure A07: LM (a) and SEM (b-f) of cells of the Amphidomatacean bloom stations identified as *Amphidoma parvula*.**
793 **(b, c) Cells in ventral view. (d) Cell in dorsal view. (e) Cell in ventral antapical view. (f) cell in apical view. Note the**
794 **small size, the flat hypotheca, the shape of the 1' plate visible in a, b, the group of pores in the second antapical plate**
795 **(white arrow in e), and the relatively long apical plates visible in d and f. Scale bars = 2 μm.**

797 In the same size class, cells were also observed that were identified in SEM as *Amphidoma parvula* (Fig. A07) by (1) their flat
798 hypotheca and (2) a characteristically shaped 1' plate. This non-toxigenic species was described in 2018 based on a culture
799 isolated from the Argentine shelf (Tillmann et al., 2018). In accordance, a low number of reads with high similarity to *Am.*
800 *parvula* strain H-1E9 (>98 %) recorded by metabarcoding. With its relatively long apical plates, *Am. parvula* could also be
801 easily distinguished in the SEM from the similarly small *Am. languida*, which was also identified in SEM (Fig. A08). In *Am.*
802 *languida*, (1) the small apical plates, and (2) the presence of a large characteristic antapical pore is a distinguishing feature.

804 **Figure A08: LM (a) and SEM (b-f) of cells of the Amphidomatacean bloom stations identified as *Amphidoma languida*.**
805 **(b, c) Cells in ventral view. (d) Cell in dorsal view. (e) Cell with epitheca in ventral view and hypotheca in antapical**
806 **view. (f) Epitheca in lateral apical view. (g) epitheca in apical view. (h) Hypotheca in antapical view. Note the shape of**
807 **the 1' Plate visible in a), the distinct antapical pore in the second antapical plate (white arrows in d, e, h), and the**
808 **relatively short antapical plates visible in d, f, g. Also note that there are only single pores on precingular plates (small**
809 **white arrows in c, e, f, g). Scale bars = 2 μm.**

811 However, in 2024, a new species of *Amphidoma*, *Am. fulgens*, was described, which is morphologically almost identical to
812 *Am. languida* but shows significantly different sequence data and, different to *Am. languida*, does not produce azaspiracids
813 (Kuwata et al., 2024). *Amphidoma fulgens* was found to be widely distributed in the Pacific, but there are no records yet from
814 the Atlantic Ocean. Despite its presence in the bloom samples, there were no reads related with higher similarities to *Am.*
815 *languida* in the ITS metabarcoding data set, but only very few ready with rather low similarity to *Amphidoma entries in the*
816 *database*- This seem to be in line with the general notion that ITS sequencing of cultured strains of *Am. languida* in many
817 cases failed (Wietkamp et al. 2019). In fact, data of a previous metabarcoding study (Liu et al. 2023) showed that *Am. languida*
818 hits were abundant in Chinese waters in the LSU dataset but absent in the ITS-1 data set. While *Am. languida* and also *Am.*
819 *fulgens*, according to their original descriptions, only have one small pore on each precingulate plate (Tillmann et al., 2012;

820 Kuwata et al., 2024), there were also several cells of *Amphidoma* with three or more small pores on individual precingular
821 plates (Fig. A09a–d). To what extent these cells represent *Am. languida* or other yet undescribed and closely related species,
822 requires further clarification.

823

824

825 **Figure A09: SEM of unidentified cells of *Amphidoma* spp. of the Amphidomatacean bloom stations. (a–d) Cells in apical**
826 **(a, c) or dorsal (b, c) view resembling *Amphidoma languida*, but with multiple pores in precingular plates (small white**
827 **arrows in a–d). (e) A cell of *Amphidoma* sp. in dorsal view. Note the very long apical plates. The row of pores with a**
828 **distinct rim on the apical plates (white arrow in e) resemble *Amphidoma alata*. (f) Cell of an unidentified *Amphidoma***
829 **sp. 1 in ventral view. Note the long apical plates, the ventral depression (white arrow in f), and the row of pore on the**
830 **posterior suture of apical plates. (g, h). Two cells of *Amphidoma* sp. resembling *Amphidoma trioculata*. (i) Hypotheca in**
831 **antapical view of an unidentified *Amphidoma*. Note the multiple pores on the plates and the presence of a very small**
832 **antapical pores on the second antapical plate (white arrow in i). Scale bars = 2 µm.**

833

834 Due to the very similar size of *Az. dexteroporum*, *Am. parvula*, and *Am. languida*, all three species as identified with SEM are
835 present in the one category of small amphidomatcean cells used for light microscopic analysis and quantification.

836 **A 1.5. Other Amphidomataceae**

837 In the SEM, a few other cells were observed, which can also be assigned to the genus *Amphidoma*. The epitheca found in
838 dorsal view in Fig. A09e with its characteristic pore ridges on the large apical plates corresponds to *Am. alata*, a species which
839 was described from the Argentina shelf (Tillmann, 2018). The cell depicted in Fig. A09f in ventral view likely represents a
840 new species of *Amphidoma*. The cells in Figs. A09g-i likely correspond to *Am. trioculata*, another species described from
841 Argentina (Tillmann, 2018), though assigning the isolated hypotheca (Fig. A09i) is difficult.

842

843

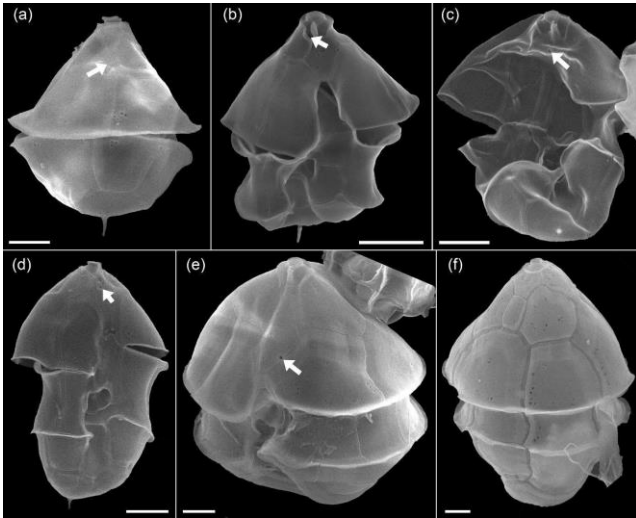



Figure A10: SEM of unidentified cells of *Azadinium* spp. of the Amphidomatacean bloom stations. (a) Cell in dorsal view. Note the very small and 5-sided intercalary plate (white arrow in a). (b) Cell of *Azadinium* sp. 1 in ventral view. Note the very short 1' Plate and the position of the ventral pore inside the pore plate (white arrow in b). It may be assumed but is not clear if (a) is the dorsal view of such an *Azadinium* sp. 1. (c. d), Two cells in ventral view of *Azadinium* sp. 2. Note the position of the ventral pore in apical position inside of Plate 1' (white arrows in c, d). (e) Cell of *Azadinium* sp. 3 in lateral ventral view. Note the position of the ventral pore (white arrow in e) and the rather long apical plates. (f). Dorsal view of an *Azadinium* sp. 4 resembling in size and shape *Azadinium asperum*. Scale bars = 2 μm.

Moreover, for a more complete description of the diversity of Amphidomataceae in the bloom sample, the following individual findings (Fig. A10) should also be mentioned: A cell in dorsal view (Fig. A10a) had a very distinct antapical spine, relatively large apical plates, and a small six-sided central intercalary plate. This combination of features has not been described in any known *Azadinium* species, suggesting that this may represent a new species. The cell depicted in Fig. A10b had a distinct antapical spine, a ventral pore on the right side of the pore plate, and a very short first apical plate, with the anterior sulcal plate extending far into the epitheca. Both cells in Figs. A10b and c resemble *Az. spinosum* but differ in that the ventral pore is centrally located within Plate 1' in the apical area. They likely correspond to the cells designated as *Azadinium* sp. 3 in Fig. 14 b, c, in Tillmann (2018). The cell in Fig. A10e had a ventral pore on the right side of Plate 1' (like *Az. spinosum* and *Az. obesum*), but here the lateral apical plates were significantly larger than in these species. The cell in Fig. A10f in dorsal view in terms of size and shape might correspond to *Az. asperum* described from the Argentine shelf (Tillmann, 2018), undoubtedly an *Azadinium* species due to the apical pore and intercalary plates, does not match any previously described species based on size (ca 20 μm cell length) and shape. This, along with the other cells in Fig. A09 which likely represents previously undescribed species due to the unique combination of features, highlights the great diversity of Amphidomataceae in this 2021 bloom sample.

Appendix B

Species diversity based on ITS1-based metabarcoding

871 **Table B01: Species detected by Amplicon Sequence Variant (ASV) reads at stations GA01 and AA09.**

872

Nr	species	Genbank	Identity %	GA01	AA09
1	*Amphidoma parvula*	KY996792	**>98**	19	0
2	*Amphidoma* sp. 1	OQ360107	90-95	158	200
3	*Amphidoma* sp. 2	LC788745	90-95	16	0
5	*Azadinium dalianese* ribotype E	LS974150	90-95	41	63
			95-98	33	58
			>98	93.775	98.503
6	*Azadinium dalianense* ribotype B	MF033117	95-98	9	0
			>98	0	19
7	*Azadinium* sp. 1	OQ360091	90-95	36	41
8	*Azadinium* sp. 2	OQ360094	**>98**	6	5
9	*Azadinium spinosum* ribotype C	MK405512	**>98**	17.006	7.589
10	*Azadinium spinosum* ribotype B	LS974169	**>98**	0	37
11	*Ansanella* sp.	MN604385	90-95	4	13
12	*Bicheleria* sp.	KC895487	90-95	12	52
13	*Bicheleria cincta*	KC895487	**>98**	6	17
14	*Blastodinium oviforme*	JX473680	95-98	2	0
15	*Karlodinium decipiens*	LC521288	95-98	0	7
			>98	885	106
16	*Karlodinium digitatum*	MN133932	**>98**	3	0
17	*Kirithra asteri*	MW267275	**>98**	41	186
18	*Pelagodinium beii*	KP843723	**>98**	48	77

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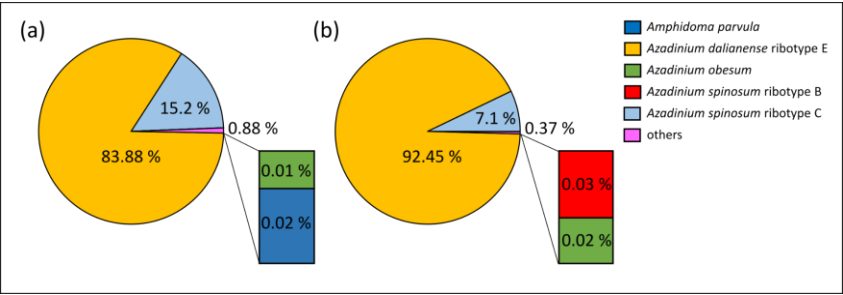
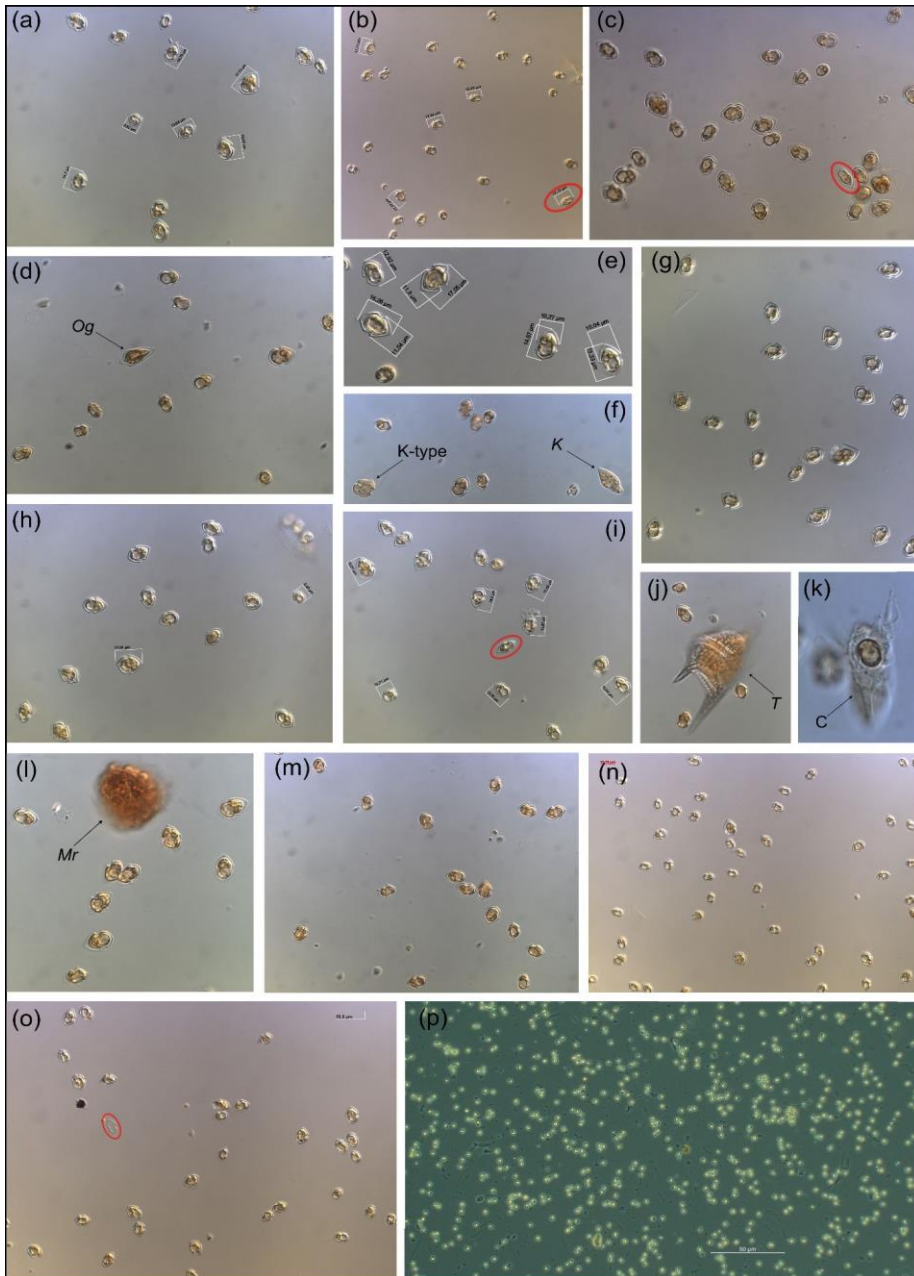

875

876 **Figure B01: Relative abundance (in %) of the Amphidomataceae species detected with metabarcoding targeting ITS1**
877 **regions at stations GA01 (a) and AA09 (b).**

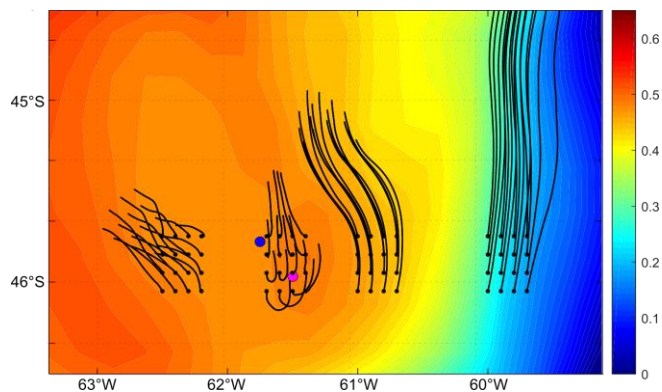
878

879 **Appendix C**

880

881 **Overall appearance of the pure Amphidomataceae bloom in fixed field samples**

882
883 **Fig. C01: Micrographs of the Amphidomataceae bloom taken under light microscopy at low magnification: 200x. Only**
884 **(e) was taken under 400x and (p) was taken using fluorescence in 100x. The toxic species *Azadinium spinosum* ribotype**
885 **B is indicated in a red circle. The arrows indicate other protists: *Og*: *Oxytoxum gracile*, K-type: Kareniacea-type cell,**
886 **K: *Katodinium* sp., *T*: *Tripos* sp., C: ciliate with a cell of Amphidomataceae inside, *Mr*: *Mesodinium rubrum*.**

887

888 **Appendix D**
889
890 **Retention of particles in the blooming area**

891
892 **Fig. D01: Background colours: Absolute Dynamic Topography (m) averaged from ~~16 to 25~~ November 16 to 25. Black**
893 **lines correspond to the advection of particles for the same period after release at the points indicated with a black dot**
894 **on ~~16 of~~ November 16.**
895
896 **Supplementary material:** Video showing the daily evolution from ~~10 to 25~~ November 10 to 25, 2021 of the Finite-size
897 Lyapunov Exponent (FSLE) ridges in the area of the two locations with the Amphidomataceae bloom: GA01, pink dot,
898 sampled on November 16, and AA09, blue dot, sampled on November 25. The two stations remained within the same water
899 mass separated by two maxima FSLE.
900
901 **Data availability:** The CTD data and the abundance of protistan species counted under light microscopy at the sampling
902 station GA01 and AA09, will be publicly available in the repository Pangaea Data Publisher (www.pangaea.de)**.** Data from
903 the Gayoso cruise is available at: https://doi.pangaea.de/10.1594/PANGAEA.971564. Sequences obtained in this study are
904 available in the National Center for Biotechnology Information, Sequence Read Archive (http://www.ebi.ac.uk/ena).
905
906 **Author Contribution:** VAG conceptualised and designed the study, coordinated the planning and execution of field and
907 laboratory work, and secured funding. VAG and UT analysed the plankton samples using LM and SEM. MR processed the
908 satellite-derived chlorophyll data and the Lyapunov coefficients. CF and FR conducted field research and processed the CTD
909 data. VAG and BK processed the toxin samples. HG processed the DNA samples. MS analysed the geostrophic currents and
910 performed the particle tracking modelling. All co-authors contributed to the interpretation of the results. VAG prepared the
911 manuscript with contributions from all co-authors.
912
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914

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923

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