

1 **Impacts of Sea Ice Leads on Sea Salt Aerosols and Atmospheric Chemistry in the Arctic**

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28 **Abstract.** The processes contributing to Arctic cold season (November-April) sea salt aerosols
29 (SSA) remain uncertain. Observations from coastal Alaska suggest emissions from open leads in
30 sea ice, which are not included in climate models, may play a dominant role. Their Arctic-wide
31 significance has not yet been quantified. Here, we create an emissions parameterization of SSA
32 from leads by combining satellite data of lead area (the AMSR-E product) and a chemical
33 transport model (GEOS-Chem) to quantify pan-Arctic SSA emissions from leads during the cold
34 season from 2002-2008 and predict their impacts on atmospheric chemistry, evaluating the
35 results of our simulated SSA against in-situ observations. The AMSR-E product detects large
36 leads with certainty (>3 km in size) and hence our study is limited to quantifying emissions from
37 large leads. Lead emissions vary seasonally and interannually. Simulated total monthly SSA
38 emissions increase by 1.1-1.8% ($\geq 60^\circ\text{N}$ latitude) and 5.6-7.5% ($\geq 75^\circ\text{N}$) for the 2002-2008 cold
39 season. SSA concentrations increase primarily at the location of leads, where standard model
40 concentrations are low. GEOS-Chem overestimates SSA concentrations at Arctic sites compared
41 to ground observations even when lead emissions are not included, suggesting underestimation
42 of SSA sinks and/or uncertainties in SSA emissions from blowing snow and open ocean. Multi-
43 year monthly mean surface bromine atom (Br) concentrations increase 2.8-8.8% due to SSA from
44 leads for the 2002-2008 cold season. Changes in ozone concentrations are negligible. While
45 leads contribute <10% to Arctic-wide SSA emissions in the years 2002-2008, these emissions
46 occur in regions of low background aerosol concentrations. Leads may increase in frequency
47 under future climate change, which could increase SSA emissions from leads.

48

49 **Short Summary**

50 There is uncertainty in the sources of Arctic cold season (November-April) sea salt aerosols.
51 Using a chemical transport model and satellite observations, we quantify Arctic-wide sea salt
52 aerosol emissions from fractures in sea ice, called open sea ice leads, and their atmospheric

53 chemistry impacts for the cold season. We show sea ice leads contribute to Arctic sea salt
54 aerosols and bromine, especially in under-observed regions.

55

56 **1. Introduction**

57

58 Sea salt aerosols (SSA) affect Arctic climate by scattering incoming solar radiation and acting as
59 cloud condensation nuclei and ice nuclei (DeMott et al., 2016; Pierce and Adams, 2006; Quinn et
60 al., 1998). While in the Arctic there is no sunlight during polar night to scatter radiation, cloud
61 condensation nuclei and ice nuclei can still have impacts on clouds and longwave radiation. Long-
62 term measurements have shown that peak SSA concentrations in the Arctic occur during the cold
63 season (Leaitch et al., 2018; Quinn et al., 2002; Schmale et al., 2021). However, the sources and
64 mechanisms of cold season SSA emissions are uncertain, which hinders atmospheric chemistry
65 and climate models from accurately representing polar regions. Recent observations from
66 Utqiagvik, Alaska have suggested that open leads, or open sea ice fractures, are an important
67 source of cold season SSA emissions (Kirpes et al., 2019; May et al., 2016). Climate change has
68 impacted the Arctic by rapidly decreasing sea ice age and thickness (Intergovernmental Panel
69 On Climate Change, 2023; Sumata et al., 2023; Vaughan et al., 2013), and future projections
70 indicate this will continue (Intergovernmental Panel On Climate Change, 2023), suggesting the
71 amount of open leads will increase in the future due to thinner ice that is prone to fracture. More
72 work is needed to discern the Arctic-wide importance and impacts of SSA emissions from sea ice
73 leads (“lead emissions”) on atmospheric chemistry and climate. By combining satellite
74 observations and chemical transport modeling, we quantify the significance and impacts of lead
75 emissions on atmospheric concentrations of SSA and bromine and evaluate simulated SSA
76 against in-situ observations.

77

78 While global models have not yet included SSA emissions from leads, several observational
79 studies largely based in Utqiagvik, Alaska suggest emissions of SSA from leads may be
80 important. Key early observations in the 1970s in Utqiagvik by Scott & Levin (1972) and Radke et
81 al. (1976) demonstrated an increase in sodium-containing particles in the presence of open water
82 leads. Since then, more recent measurement studies have quantified SSA emissions from leads.
83 Nilsson et al. (2001) estimate that leads contribute an order of magnitude less than the open
84 ocean to the Arctic SSA flux during the summer months. A multi-year study of observed SSA at
85 Utqiagvik (May et al., 2016), conducted over all seasons, found that leads are a significant
86 contributor to SSA through wind-driven production, increasing the supermicron range in particular,

87 but to a lesser extent than wind-driven production from the open ocean. Willis et al. (2018) suggest
88 that lead emissions are more important in winter and early spring as winds over the Northern
89 oceans are at their highest. Kirpes et al. (2019) also convey the importance of seasonality,
90 identifying SSA produced by local leads as the dominant aerosol source in the coastal Alaskan
91 Arctic during winter months. Chen et al. (2022), focusing on the spring at Utqiaġvik, shows leads
92 were present locally throughout the study and contributed to sea spray aerosol production. As
93 ground-based observations in the Arctic are mainly limited to coastal stations, such as Utqiaġvik,
94 it is difficult to estimate the significance and impacts of lead emissions over the entire Arctic.
95 Representing Arctic-wide emissions from leads in a global chemical transport model, especially
96 during the cold season, will help discern whether lead emissions and their impacts on atmospheric
97 chemistry are significant enough to warrant inclusion in chemistry as well as climate models.

98
99 Other modeling studies in the Arctic and observations primarily from Antarctica suggest blowing
100 snow is a potential major contributor of cold season SSA in polar regions. Blowing snow SSA
101 comes from saline snow over sea ice that is swept up by wind; the snow becomes salty through
102 the upward movement of brine from sea ice to the snow surface, incorporation of frost flowers,
103 and deposition of SSA from the nearby open ocean (Domine et al., 2004). In two chemical
104 transport models, the inclusion of additional SSA emissions from blowing snow brought simulated
105 SSA mass concentrations closer to what was observed (Confer et al., 2023; Huang et al., 2018;
106 Huang and Jaeglé, 2017; Rhodes et al., 2017). Other potential sources of cold season SSA, such
107 as frost flowers, have been found to be insignificant (Alvarez-Aviles et al., 2008; Roscoe et al.,
108 2011; Yang et al., 2017). Incorporating blowing snow SSA emissions into models has shown how
109 missing sources of SSA in the Arctic can have a significant impact on atmospheric chemistry; for
110 example, Huang et al. (2020) show bromine released by blowing snow impacts modeled
111 springtime bromine activation and ozone depletion events. The strong observational evidence
112 that leads contribute to cold season SSA and the impact of blowing snow SSA on modeled Arctic
113 atmospheric chemistry suggests there is a need to assess the potential impacts of lead emissions,
114 which are currently missing from global chemistry and climate models. One study incorporated
115 SSA emissions from leads in a chemical transport model (WRF-Chem), but the study was limited
116 to the 400 km² area surrounding Utqiaġvik, Alaska and used ERA-5 reanalysis sea ice fraction to
117 define the presence of leads (Ioannidis et al., 2023). They find open leads are the primary source
118 of fresh and aged SSA in Utqiaġvik, Alaska during the cold season, consistent with the
119 observational analyses by May et al. (2016) and Kirpes et al. (2019).

120

121 SSA play a critical role in Arctic tropospheric chemistry. SSA debromination is the main global
122 source of reactive bromine in the troposphere (Wang et al., 2021). Reactive bromine chemistry
123 has been attributed to the rapid depletion of ozone in the Arctic springtime, which reaches a
124 maximum in March-April (Simpson et al., 2007). In particular, bromine atom (Br) is key to these
125 ozone depletion events; it is produced through the photolysis of Br₂, which is sourced from SSA
126 debromination and snowpack chemistry (Abbatt et al., 2012; Dibb et al., 2010; Pratt et al., 2013;
127 Stutz et al., 2011). Swanson et al. (2022) show improved springtime model-observation
128 agreement of BrO by including a snowpack photochemistry mechanism based on multiple field
129 observations in a global chemical transport model. While on a global scale, reaction of OH with
130 other SSA-sourced bromine species can also produce Br (Wang et al., 2021), this is minor in polar
131 regions due to low OH concentrations. Br rapidly depletes ozone through heterogeneous
132 reactions, which produces BrO that can photolyze to reform Br, creating a catalytic ozone-
133 depletion cycle (Simpson et al., 2007).

134

135 Here, we estimate the pan-Arctic contribution of leads to total SSA emissions during the cold
136 season for the years 2002-2008, by using satellite observations of lead area to parameterize lead-
137 based SSA production in the global chemical transport model GEOS-Chem. We evaluate
138 simulated SSA concentrations against observations and predict the impacts of lead SSA
139 emissions on atmospheric chemistry, including concentrations of Br and ozone.

140

141 **2. Methods**

142

143 *2.1 Satellite Data of Lead Area Fractions*

144

145 In this study, we use satellite data of lead area fractions to inform the GEOS-Chem chemical
146 transport model (next section) of where leads are present. The Advanced Microwave Scanning
147 Radiometer-Earth Observation System (AMSR-E) sensor aboard NASA's Aqua satellite recorded
148 brightness temperatures from Earth from 2002-2011 at six different frequencies
149 (<https://www.cen.uni-hamburg.de/en/icdc/data/cryosphere/lead-area-fraction-amsre.html>)
150 (Integrated Climate Data Center (ICDC) et al., n.d.), which are converted to lead area fractions
151 following the algorithm of Röhrs and Kaleschke (2012). This method of detection can only be
152 applied to the Arctic freezing season (November-April) due to surface melt of the sea ice modifying
153 the sea ice emissivity from May-October, which affects the lead detection algorithm. Daily data is
154 available at 6.25 km horizontal resolution, as the algorithm is not limited by cloud cover. The

155 AMSR-E satellite data is regridded to $0.5^{\circ} \times 0.625^{\circ}$ from 6.25×6.25 km using a distance-weighted
156 average remapping for consistency with the emission model's resolution (see Sect. 2.2 below for
157 model details). For the rare individual days with missing data in the dataset (0.8%), we use the
158 average lead area fraction for that month. The lead area fraction includes open water leads and
159 thin ice-covered leads 3 km and wider. The data spans latitudes 41° to 90° N, though a majority
160 of Arctic sea ice lies above 60° N and leads are therefore unlikely to be present at lower latitudes.

161
162 We use the AMSR-E lead area product for this study as it avoids cloud interference when
163 detecting leads and provides nearly consistent daily resolution. A limited quantitative validation
164 by Röhrs and Kaleschke (2012) of one day (March 21, 2006) of the AMSR-E product against
165 Moderate Resolution Image Spectroradiometer (MODIS) showed 50% of the total lead area
166 visible in 500 m MODIS images was detected in the AMSR-E product. Leads greater than 3 km
167 in size ("large leads") were detected with certainty for the AMSR-E product (Röhrs and Kaleschke,
168 2012), so our results effectively estimate emissions from large leads only.

169
170 *2.2 GEOS-Chem: global chemical transport model*

171
172 Here we use the 3-D atmospheric transport model GEOS-Chem (geos-chem.org) version 13.2.1
173 (<https://doi.org/10.5281/zenodo.5500717>). Within GEOS-Chem, the Harmonized Emissions
174 Component (HEMCO) computes emissions for different sources, regions, and species (Keller et
175 al., 2014). GEOS-Chem and HEMCO are driven by Modern-Era Retrospective Analysis for
176 Research and Applications (MERRA-2) (Gelaro et al., 2017) meteorological fields from the NASA
177 Global Modeling and Assimilation Office (GMAO), which is reanalysis meteorological data
178 assimilated from various observational sources (i.e., satellite, aircraft campaigns, and ground
179 stations) providing variables such as temperature, wind, precipitation, and humidity. GEOS-Chem
180 represents one-way interactions between the MERRA-2 meteorology and chemical constituents,
181 meaning the meteorological conditions can affect the concentration of chemical species but not
182 vice versa.

183
184 SSA emissions calculations for the open ocean use a wind- (Gong, 2003; Monahan et al., 1986)-
185 and sea-surface-temperature-dependent (Jaeglé et al., 2011) source function. In polar regions,
186 SSA emissions from blowing snow are also included (Huang and Jaeglé, 2017). SSA have two
187 size bins: coarse mode (SALC; $r = 0.5$ to $10 \mu\text{m}$) and accumulation mode (SALA; $r = 0.1$ to 0.5
188 μm). For gas and aerosol species, wet deposition (both rain and snow) includes washout and

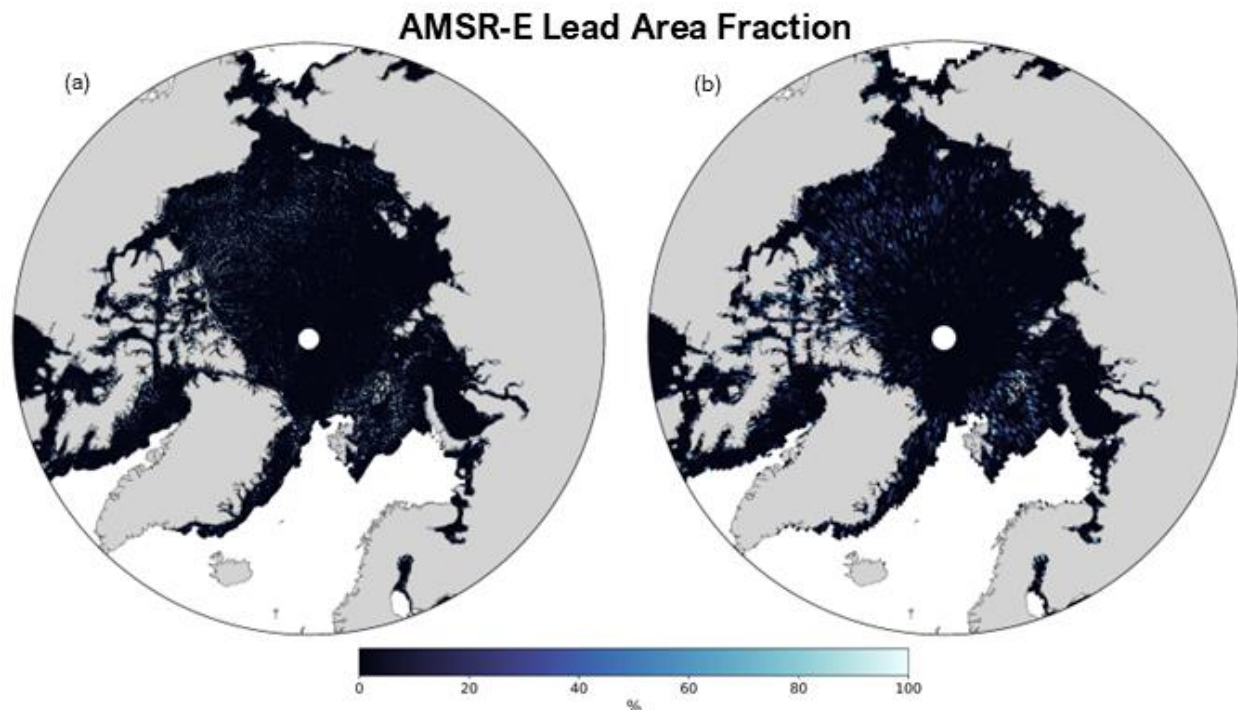
189 rainout in convective and large-scale stratiform precipitation (Amos et al., 2012; Liu et al., 2001;
190 Wang et al., 2014). From November to April in the Arctic, wet deposition is mainly in the form of
191 snow (Screen and Simmonds, 2012). Dry deposition of gas and aerosol species follows a
192 resistance-in-series approach, and includes gravitational settling of sea salt (Jaeglé et al., 2011;
193 Pound et al., 2020; Wang et al., 1998; Zhang et al., 2001). Coupled gas- and multiphase-reactive
194 halogen chemistry, including sea salt debromination, acid displacement, and photolysis and
195 oxidation of gas-phase inorganic bromine and chlorine species, is described in Wang et al. (2021).
196 This version of GEOS-Chem does not include snowpack chemistry as a source of reactive
197 bromine in the standard model.

198
199 We parameterize SSA emissions from leads with the same function as the open ocean emissions
200 from Jaeglé et al. (2011) (Eq. (S1) in the Supplemental Information (SI)), scaled by the fractional
201 area of leads in each grid cell from the AMSR-E satellite data. The Jaeglé et al. (2011) function is
202 empirically derived to best match global observations in GEOS-Chem. We assume that leads emit
203 SSA at an equal rate as a function of lead area. This lead emissions parameterization is a unique
204 wind- and SST-dependent source function for calculating lead emissions, driven by satellite
205 observations defining the presence of leads. Figure 1 shows an example of the daily temporal
206 frequency and spatial resolution of the AMSR-E satellite data (both the raw (a) and regridded (b))
207 used to drive the model.

208

209

210



211
 212 **Figure 1-** Map of AMSR-E daily lead area fraction in percent (%) for November 1, 2002, both raw
 213 (6.25-km resolution) (a) and re-gridded (0.5°x0.625° resolution) (b).

214
 215 We first calculate SSA emissions at the highest resolution of HEMCO (0.5°x0.625°), which is the
 216 native resolution of MERRA-2. Two sets of emissions are calculated: (1) the standard emissions
 217 only (i.e., open ocean and blowing snow SSA emissions, the “standard” case); (2) SSA emissions
 218 with lead emissions added (“standard + leads” case). Each set of emissions are then implemented
 219 separately into GEOS-Chem “offline” to ensure total SSA emissions are properly scaled and
 220 distributed and not influenced by the resolution-dependence of the wind speed (Lin et al., 2021).
 221 GEOS-Chem is run at the highest global horizontal (2° latitude x 2.5° longitude) and vertical (72
 222 vertical levels) resolution. The absolute difference between the standard + leads and standard
 223 simulations is the change in SSA emissions or concentration from leads, and we present the
 224 percent change due to leads (%) as calculated with Eq. (1).

225
 226
$$\text{Percent change due to leads (\%)} = 100 \times \frac{(\text{Standard+leads})_{\text{simulation}} - (\text{Standard})_{\text{simulation}}}{(\text{Standard})_{\text{simulation}}} \quad (1)$$

227
 228 Simulations are performed for the years 2002-2008, when there is overlap between the AMSR-E
 229 satellite data and available observed Arctic SSA concentrations at multiple sites, following one
 230 year of initialization. Because satellite observations of lead area fractions begin November 1,

231 2002, we initialize the standard + leads case for GEOS-Chem with standard + leads SSA
232 emissions for one year (November 1, 2002 to November 1, 2003) and then start the simulation
233 for analysis on November 1, 2002, with the spun-up November 1, 2003, initial conditions. For the
234 standard case, the initialization year begins November 1, 2001. For both cases, we simulate SSA
235 concentrations, evaluate against observed concentrations, and assess the impacts of additional
236 lead emissions on atmospheric chemistry. This includes analysis of the change in atmospheric
237 concentrations of bromine atom (Br) and ozone (O₃). For model evaluation, GEOS-Chem does
238 not track sodium (Na⁺) content for SSA, so we convert simulated SSA to Na⁺ mass concentrations
239 using a factor of $\frac{1}{3.256}$, which is based on the mass ratio of Na⁺ in seawater (Confer et al., 2023;
240 Huang and Jaeglé, 2017; Riley and Chester, 1971).

241

242 *2.3 In-Situ Observations of Arctic Sea Salt Aerosol Concentrations*

243

244 We evaluate simulated concentrations of SSA from GEOS-Chem, converted to Na⁺
245 concentrations, against in situ observations of Na⁺ concentrations at 4 Arctic sampling sites:
246 Utqiagvik, Alaska (71.3°N, 156.6°W; 11m a.s.l.) (Quinn et al., 2002); Zeppelin Mountain, Svalbard,
247 Norway (78.9°N, 11.9°E; 475m a.s.l.) (World Meteorological Organization (WMO), 2003); Alert,
248 Nunavut, Canada (82.5°N, 62.5°W; 210m a.s.l.) (World Meteorological Organization (WMO),
249 2003); Pallas (Matorova), Helsinki, Finland (68 °N, 24.24 °E; 340m a.s.l.) (Salmi, 2018). These
250 observations are available for the time period of this study (November-April from 2002-2008,
251 except for Pallas station, 2003-2008). In winter months, the Utqiagvik, Zeppelin, and Alert coastal
252 sites border mostly ice-covered ocean (Huang and Jaeglé, 2017). At Utqiagvik, mass
253 concentrations of Na⁺ for submicron and supermicron aerosols are separated, while the other two
254 sites measure the total mass concentration without size distinction. The Na⁺ mass concentrations
255 are determined from ion chromatography with uncertainties of 5-11%, or an absolute uncertainty
256 of 0.01 µg/m³ (Quinn et al., 2000; World Meteorological Organization (WMO), 2003). The aerosol
257 sampling frequency is daily at Zeppelin, Utqiagvik (submicron), and Pallas and weekly at Alert
258 and Utqiagvik (supermicron).

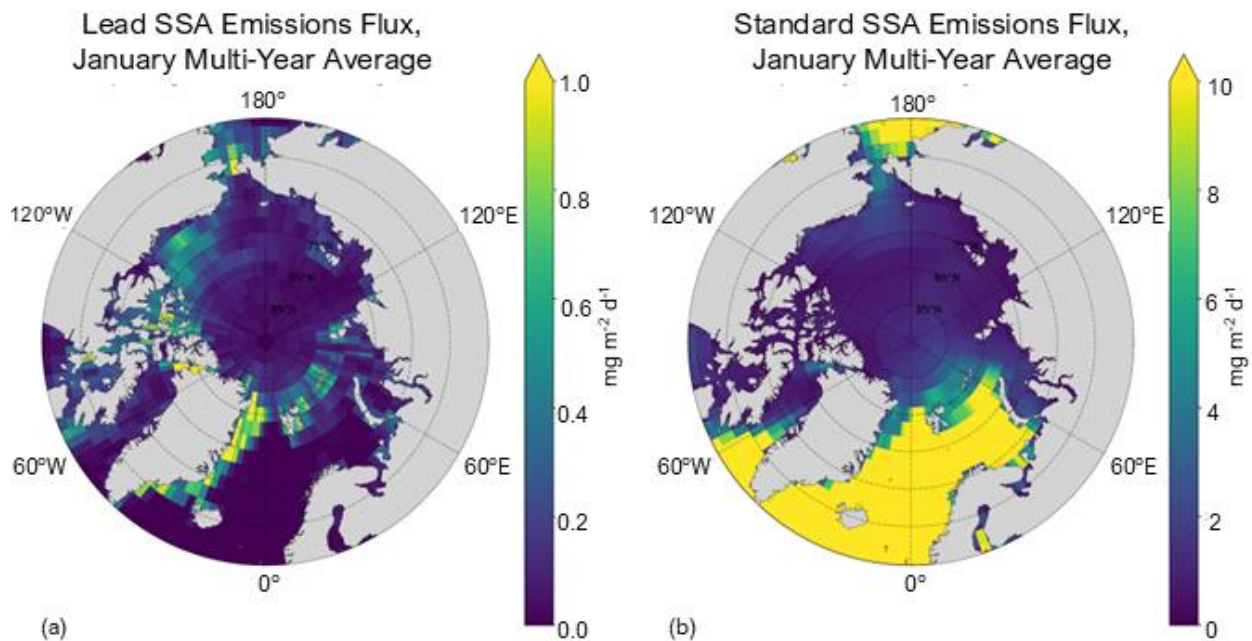
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260 **3. Results**

261

262 *3.1 Emissions of Sea Salt Aerosols from Leads*

263



264
 265 **Figure 2-** Total (coarse + accumulation mode) lead SSA emissions (a) and standard SSA
 266 emissions (b), averaged over 2002-2008 for January. Note the difference in magnitude of the
 267 colorbar of (a) and (b).

268
 269 Figure 2a shows the spatial distribution of multi-year (2002-2008) average lead emissions for the
 270 month of January, which is a climatology based on model simulations that use daily resolution
 271 lead data (e.g., Fig. 1). We focus Figs. 2 and 4 on the month of January as an example. January
 272 is tied for highest lead emissions for latitudes 60°N and greater and second highest for latitudes
 273 75°N and greater (Table 1), and also has the second largest multi-year average lead area (see
 274 Fig. S3b in SI). Alongside Fig. 2a is the standard model, which includes open ocean and blowing
 275 snow emissions (Fig. 2b; see Sect. 2.2). Total emissions are resolution independent and are
 276 shown in Fig. 2 for the 2.0°x2.5° resolution of the online atmospheric chemistry simulation. We
 277 find the lead emissions and lead area are spatially consistent (Figs. 1 and 2a) and occur in regions
 278 where the standard SSA emissions are low (e.g., in the Greenland Sea and parts of the Barents
 279 Sea). The percent change in SSA emissions due to leads (calculated with Eq. (1)) is detailed in
 280 Table 1; Figs. 4, 5, and S4 show the percent change in SSA concentration due to leads. Generally,
 281 emissions tend to be higher from 70° to 80° N and more concentrated within the Bering Strait,
 282 Nares Strait, Wynniatt Bay in the Canadian archipelago, and the eastern Greenland Sea, as
 283 opposed to off the coast of Northern Russia and Europe. Month to month, regions where
 284 emissions are higher remains similar while the magnitude varies (see Fig. S1 in SI).

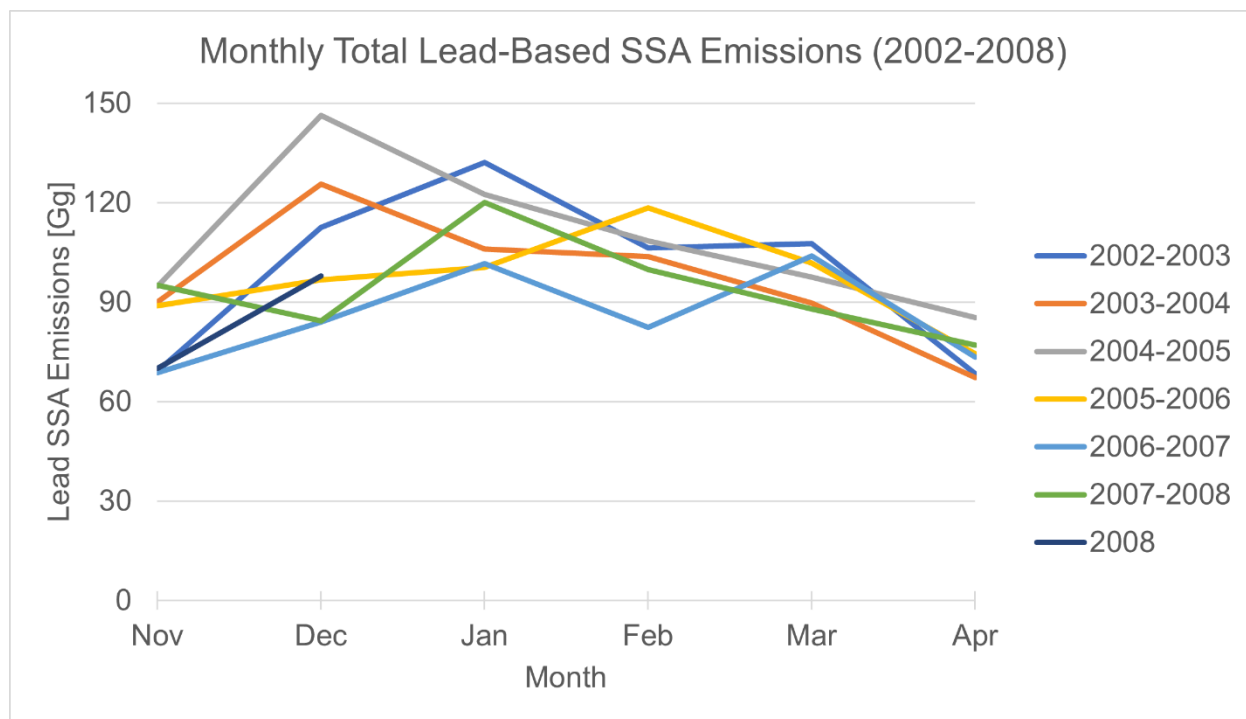
Month	Multi-Year (2002-2008) Average Standard Emissions [Gg]		Multi-Year (2002-2008) Average Lead Emissions [Gg] (and corresponding Monthly Percent Change in SSA Emissions due to Leads)	
	$\geq 60^{\circ}\text{N}$	$\geq 75^{\circ}\text{N}$	$\geq 60^{\circ}\text{N}$	$\geq 75^{\circ}\text{N}$
November	7800 \pm 1000	610 \pm 210	82 \pm 0.14 (1.1% \pm 0.14%)	42 \pm 0.13 (6.9% \pm 0.13%)
December	8700 \pm 1400	640 \pm 140	110 \pm 0.20 (1.2% \pm 0.20%)	48 \pm 0.32 (7.5% \pm 0.32%)
January	8400 \pm 1100	670 \pm 290	110 \pm 0.10 (1.3% \pm 0.10%)	46 \pm 0.15 (6.9% \pm 0.15%)
February	6700 \pm 850	510 \pm 90	100 \pm 0.11 (1.5% \pm 0.11%)	37 \pm 0.17 (7.2% \pm 0.17%)
March	6000 \pm 1000	470 \pm 66	98 \pm 0.074 (1.6% \pm 0.07%)	34 \pm 0.26 (7.2% \pm 0.26%)
April	4200 \pm 330	400 \pm 61	74 \pm 0.081 (1.8% \pm 0.08%)	23 \pm 0.17 (5.6% \pm 0.17%)

286 **Table 1-** Multi-year (2002-2008) monthly average standard emissions and lead emissions ± 1
287 standard deviation [Gg] and percent change in SSA emissions due to leads ± 1 standard deviation
288 in parentheses (calculated using Eq. (1)), averaged for $\geq 60^{\circ}\text{N}$ and $\geq 75^{\circ}\text{N}$.

289
290 Table 1 shows the standard and lead emissions in Gg as well as the percent change in multi-year
291 monthly average SSA emissions due to leads for 60° to 90°N latitude ($\geq 60^{\circ}\text{N}$) and 75° to 90°N
292 ($\geq 75^{\circ}\text{N}$). The standard deviations in Table 1 represent the year-to-year variability in emissions, as
293 the calculation is performed across the 7-year simulation time period for each month. Leads are
294 relatively more important to total SSA emissions at higher latitudes due to large open ocean
295 emissions in the North Atlantic at lower latitudes (Table 1; Fig. 2b) and the spatial variability of the
296 lead emissions (Fig. 2a). The month with the highest contribution to SSA emissions from leads
297 varies with the region being analyzed. The smaller magnitude of standard emissions later in the
298 cold season poleward of 60°N make lead emissions relatively more important, with the largest

299 percent increase $\geq 60^\circ$ N in SSA emissions due to leads occurring in April. Poleward of 75° N, the
 300 lead emissions represent a larger fraction of the standard emissions, resulting in higher percent
 301 increases due to leads ($\sim 4\text{-}6\%$ higher than for $\geq 60^\circ$ N). Absolute lead emissions peak in December
 302 for $\geq 75^\circ$ N latitude, which is also the month with the highest percent increase due to leads $\geq 75^\circ$
 303 N, and decrease more than twofold by April. Controlling factors of the lead emissions are
 304 discussed in the next paragraph.

305



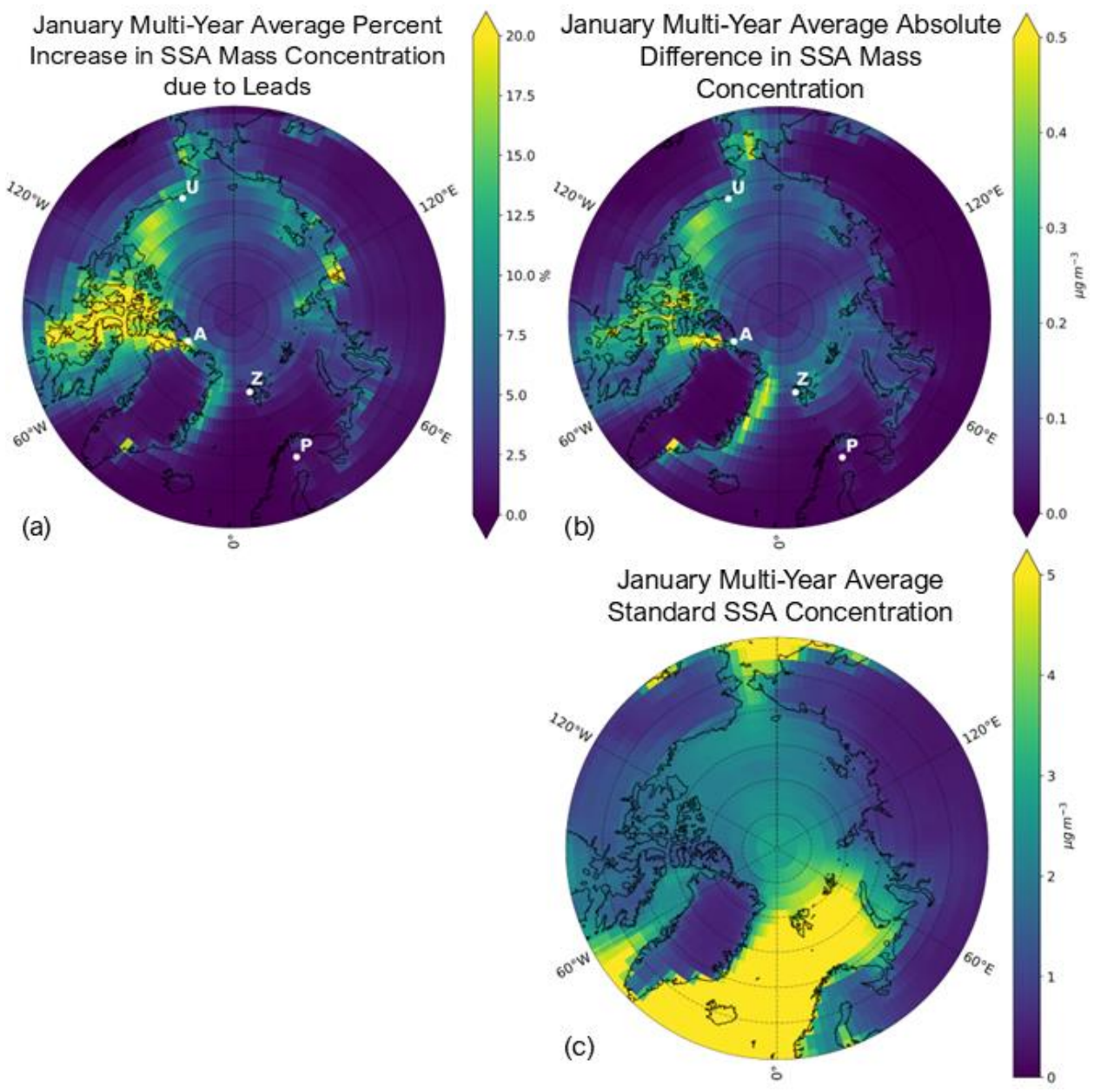
306
 307 **Figure 3-** Monthly variations of total (coarse + accumulation mode) lead emissions of SSA during
 308 the cold season for 2002-2008. Each line includes November and December of the first year and
 309 January through April of the following year, except for the year 2008, which only includes
 310 November and December of 2008.

311
 312 We find that the magnitude of lead emissions varies by month and year, as well as seasonally
 313 (see Fig. 3 and Figs. S1 and S2). Monthly total lead emissions and lead area have low correlation
 314 ($R^2= 0.13$, see Fig. S3), indicating the variance in monthly total lead emissions is dominated by
 315 the nonlinear dependencies on wind speed and sea surface temperature (Eq. S1 in SI), as the
 316 lead emissions are calculated with the Jaegle et al. (2011) wind speed and sea surface
 317 temperature source function. In most years, lead emissions decrease from January-April, but
 318 there is no single month when lead emissions peak each year (Fig. 3). There is also no clear
 319 interannual trend in cold season total lead emissions (see Fig. S2). Lead emissions are lowest in

320 the 2006-2007 cold season and highest in the 2004-2005 cold season (Fig. S2). In the future,
321 climate models predict that Arctic sea ice will continue to thin (high confidence) and the presence
322 of first-year vs. multi-year sea ice will increase (very high confidence) (Intergovernmental Panel
323 On Climate Change, 2023), suggesting a possible future increasing trend in lead area and
324 therefore lead emissions.

325
326
327

3.2 Atmospheric Chemistry Impacts of Sea Ice Leads



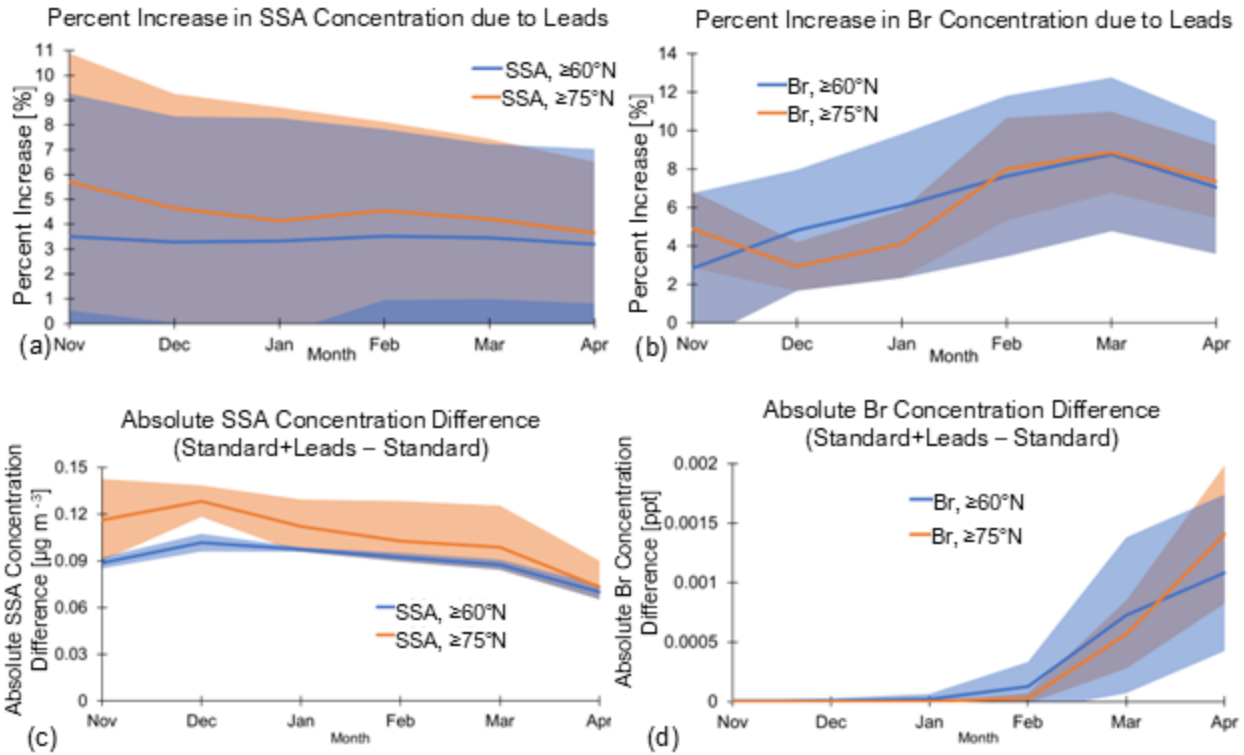
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329 **Figure 4-** Percent change due to leads (calculated with Eq. (1)) in SSA mass concentration (a),
330 absolute difference between the standard+leads and standard SSA mass concentrations in $\mu\text{g m}^{-3}$
331 ³ (b), and the standard surface SSA mass concentration in $\mu\text{g m}^{-3}$ (c) for the January multi-year
332 (2002-2008) average. White points in (a) and (b) represent the respective locations of each
333 observational site: Alert, Nunavut, Canada (A); Utqiaġvik, Alaska (U); Zeppelin Mountain,
334 Svalbard, Norway (Z); Pallas (Matorova), Helsinki, Finland (P). Note the difference in magnitude
335 of colorbars (b) and (c).

336
337 Figure 4 shows the spatial distribution of the multi-year (2002-2008) average percent change due
338 to leads in surface SSA mass concentration (4a) and the absolute difference in SSA mass
339 concentration between the standard + leads and standard simulations (4b), as well as the
340 standard simulated SSA mass concentration (4c) for the month of January. With the addition of
341 leads, the average Arctic-wide ($\geq 60^\circ\text{N}$) percent increase in multi-year mean January SSA mass
342 concentrations is 3.3%, and the maximum percent increase in an individual model gridbox is
343 60.5%. We find that the greatest percent increases due to leads in SSA mass concentrations
344 occur at the location of lead emissions (see Fig. 2a), where the standard concentrations are also
345 very low, except off the eastern coast of Greenland, where the percent increase is reduced due
346 to the high background SSA concentrations in the Greenland Sea (Fig. 4c) from open ocean
347 emissions (Fig. 2b).

348
349 Figure 5a shows the average Arctic-wide percent increase and 5c shows the absolute difference
350 due to leads in multi-year monthly mean SSA mass concentration for each cold season month.
351 Averaged poleward of 60°N , the percent increase and absolute difference due to leads in SSA
352 mass concentration remains relatively constant throughout the cold season (Fig. 5a and c).
353 Changes in monthly mean SSA mass concentrations are also higher poleward of 75°N . However,
354 the percent increase in SSA mass concentration for both latitudinal ranges have large spatial
355 variability, as seen in the standard deviation in Fig. 5a. The spatial distribution of the percent
356 increase and absolute difference in SSA mass concentration due to leads remains similar month
357 to month (see Fig. S4 and S5 in SI).

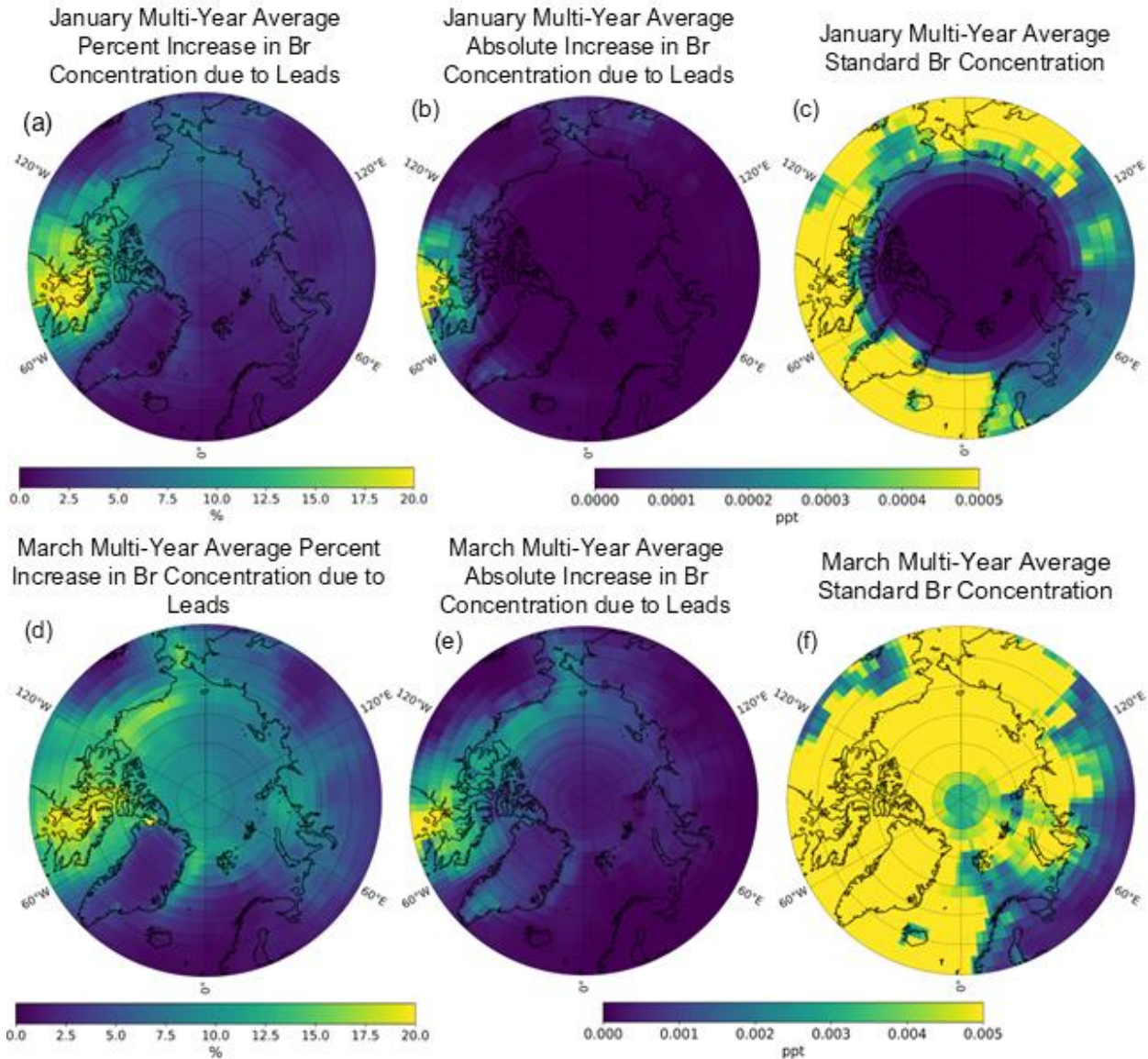
358
359



360
 361 **Figure 5.** Multi-year (2002-2008) monthly mean percent increase due leads (calculated with Eq.
 362 (1)) in surface (a) SSA and (b) Br concentrations averaged across different Arctic regions (blue
 363 line: $\geq 60^\circ\text{N}$; orange line: $\geq 75^\circ\text{N}$). Shaded area represents ± 1 standard deviation.

364
 365 As described in Sect. 1, SSA contribute to the production of tropospheric reactive bromine and
 366 thereby bromine atom (Br). Here we examine changes in Br due to its role in ozone depletion
 367 events.

368



369
 370 **Figure 6-** Multi-year (2002-2008) mean January (a, b, and c) and March (d, e, and f) percent
 371 increase due to leads in surface Br concentration (a and d), absolute increase in surface Br
 372 concentration due to leads (b and e), and the standard model surface Br concentration in ppt (c
 373 and f). Note the scale of the absolute difference and standard Br concentrations for the January
 374 and March multi-year averages are an order of magnitude difference.

375
 376 Figure 6 shows the multi-year (2002-2008) mean percent increase and absolute difference due
 377 to leads in surface Br concentrations, and the standard Br concentration (in parts per trillion, or
 378 ppt) for the months of January (a-c) and March (d-f), respectively. Increased SSA from leads
 379 increases surface levels of Br across all months during the cold season (Fig. 6a, b, d, and e; Figs.
 380 S6 and S7 in the SI for other months). These increased concentrations spatially follow the

381 increased SSA mass concentrations from leads (Fig. 4a; Figs. S4 and S5 in SI for other months)
382 with differences due to where Br can be produced photochemically from the precursors released
383 from SSA. The spatial distribution of the percent increases in Br due to leads remains relatively
384 similar month to month during the cold season (see Fig. S6 in SI), but with varying magnitude
385 (Fig. 6). The changes in Br concentration in February to April occur over a larger area (Fig. 6d
386 and e and Figs. S6 and S7), likely due to the seasonality of Arctic bromine chemistry, which is
387 influenced by increasing area where sunlight is available to photolyze Br-sourced SSA species.
388 The average Arctic-wide ($\geq 60^\circ\text{N}$) percent increase due to leads in multi-year January mean
389 surface Br concentration is 6.1% and the maximum increase in an individual gridbox is 35%; for
390 March, it is 8.8% and 20.4%, respectively. Overall, the average monthly percent increase in Br
391 concentration is higher than the corresponding increases in SSA concentration, particularly after
392 January, and reaches a maximum in March (see Fig. 5). The percent change due to leads in Br
393 concentrations increases from November-March poleward of 60°N and from December-March
394 poleward of 75°N (Fig. 5b). This does not strictly follow the seasonality of lead emissions (Fig. 3)
395 or the percent increase in SSA concentrations due to leads (Fig. 5a), likely due to more available
396 sunlight for photochemical reactions that produce Br later in the cold season. Increases in surface
397 Br concentration could lead to decreased surface ozone concentrations. We find that the percent
398 decrease due to leads in average surface ozone concentrations during the Arctic cold season,
399 however, are negligible ($< -0.25\%$).

400

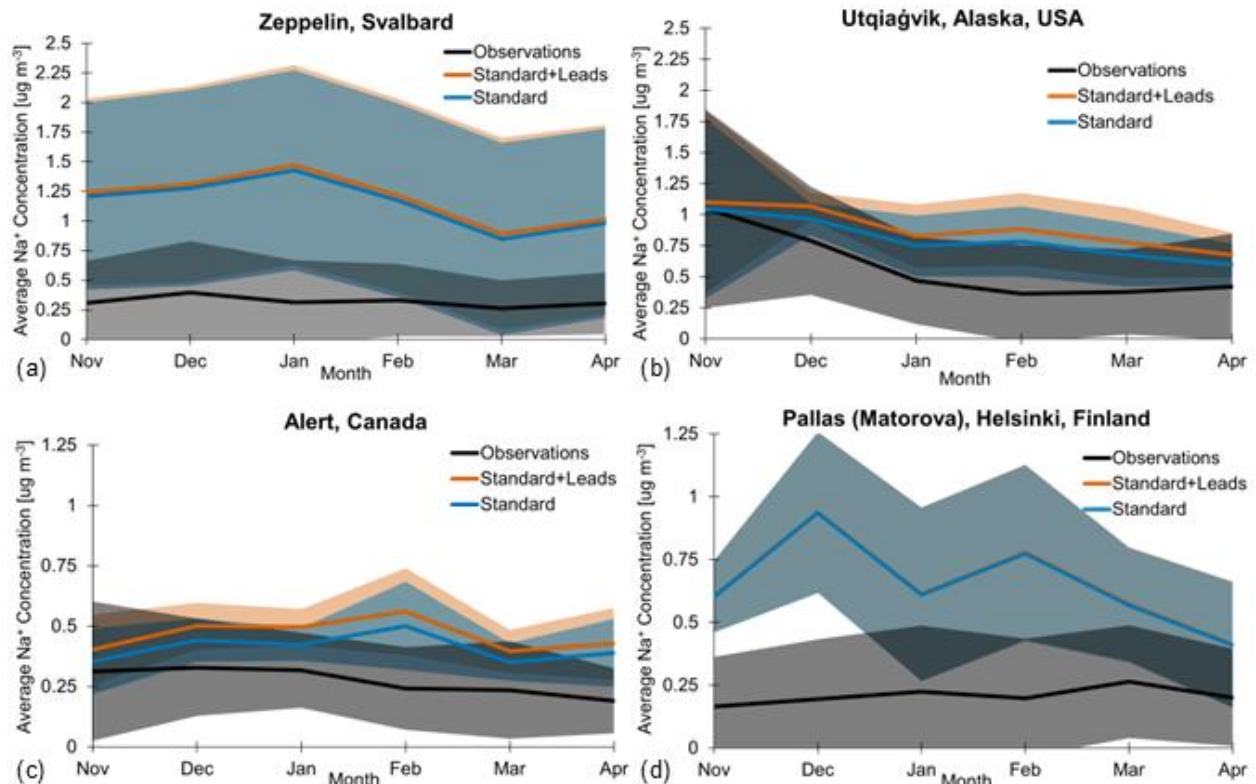
401 *3.3 Evaluation Against Sea Salt Aerosol Observations*

402

403 We compare modeled and observed sodium (Na^+) mass concentrations at four long-term
404 monitoring stations to evaluate the performance of the simulation with and without additional lead
405 emissions. The locations of each observational site are shown in Fig. 4a.

406

407



408
 409 **Figure 7-** Observed (blackline) and simulated (blue and orange lines) multi-year monthly mean
 410 sodium mass concentrations at (a) Zeppelin, Norway, (b) Utqiagvik, Alaska, (c) Alert, Canada,
 411 and (d) Pallas (Matorova), Helsinki, Finland for the cold seasons of 2002-2008 for (a)-(c) and
 412 2003-2008 for (d). Shaded regions are ± 1 standard deviation. Note the y-axis for Alert (c) and
 413 Pallas (d) are half as large as Zeppelin (a) and Utqiagvik (b).

414
 415 Figure 7 shows multi-year monthly mean Na^+ concentrations in the observations (black), standard
 416 + leads simulation (orange), and standard simulation (blue) for Zeppelin (a), Utqiagvik (b), Alert
 417 (c), and Pallas (d) during the cold season for 2002-2008 (a-c) and 2003-2008 (d). We sample the
 418 model simulations in the gridbox that encompasses the latitude, longitude, and altitude of each
 419 monitoring station (see Sect. 2.3) and convert the simulated SSA to Na^+ concentrations. For all
 420 sites and months during the cold season, the simulated and observed Na^+ mass concentrations
 421 overlap within ± 1 standard deviation (shaded regions in Fig. 7), except in November and
 422 December at Pallas. We find mean concentrations are overpredicted in both the standard and
 423 standard + leads simulations at all sites and months during the cold season, apart from the
 424 standard model at Utqiagvik and Alert in November which agree closest with observations.
 425

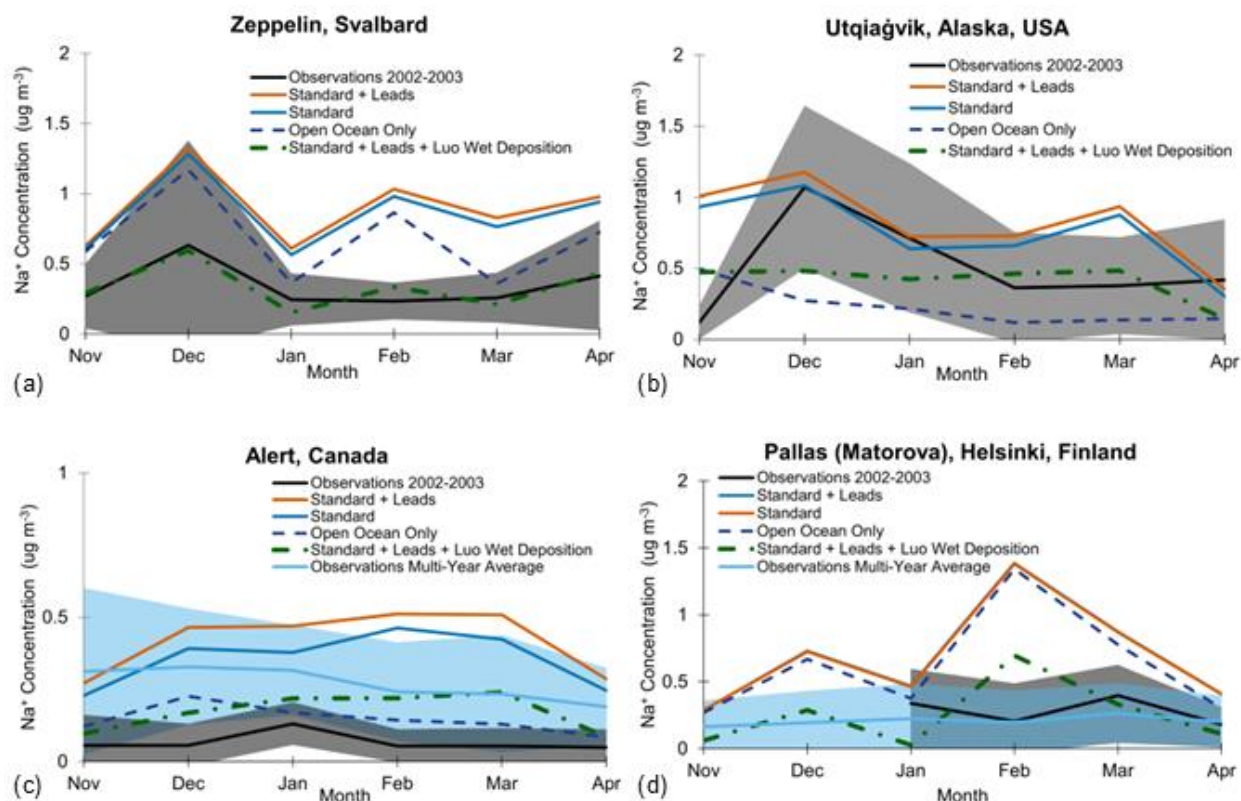
426 The model overpredicts Na⁺ concentrations the most at Zeppelin and Pallas, with the standard +
427 leads and standard mean concentrations a factor of 3.2 to 4.71 and 2.0 to 4.8 higher, respectively,
428 than observations across all months during the cold season. Confer et al. (2023) similarly find an
429 overprediction of SSA at Zeppelin, which they find is exacerbated by including blowing snow
430 emissions. Additionally, Zeppelin is at high elevation (located on a mountain at 475m) and has
431 been found to be more impacted by the free troposphere and aerosol-cloud interactions than other
432 Arctic sites (Freud et al., 2017); the chemical transport model cannot represent two-way aerosol-
433 cloud interactions. The model overestimate is less at Utqiaġvik, where the standard + leads
434 simulation still overpredicts observed concentrations by a factor of 1.0 to 2.4, and least at Alert,
435 with observed concentrations overestimated by a factor of 1.3 to 2.3 for the standard + leads
436 model. Lead emissions do not change the simulated seasonality of cold season surface SSA
437 concentrations. The timing of cold season maximum and minimum concentrations at Zeppelin,
438 Alert, and Pallas differs between the observed and simulated, for both the standard + leads and
439 standard models. At Utqiaġvik, the maximum mass concentration in the observations and both
440 model simulations occurs in November. However, the minimum observed cold season mass
441 concentration occurs in February at Utqiaġvik, whereas the standard + leads and standard mean
442 concentrations reach a minimum in April.

443
444 Figure 4a and b places the differences seen at each of the three sites in Fig. 7 into a broader
445 context, with maps of the relative and absolute increases in SSA mass concentrations for the
446 month of January. There is minimal change in SSA concentrations where Pallas is located,
447 explaining the near equal Na⁺ concentrations for the standard + leads and standard simulations
448 which results in the overlapping lines in Fig. 7d, suggesting minimal influence from leads at this
449 site likely due to its inland location. The most significant relative increase in SSA concentration
450 from leads out of the four sites occurs at Alert (Fig. 7a). However, regions with the largest changes
451 in SSA mass concentration due to leads in Fig. 4a and b for the month of January (i.e., parts of
452 Northern Canada southwest of Alert), which are consistent throughout the cold season (Fig. S4
453 and S5 in SI), are not sampled by long-term ground monitoring sites, which would help constrain
454 lead impacts on SSA. In our simulation, lead emissions have the same size distribution as the
455 open ocean, with most of the mass in the coarse mode (82-90%). Despite this, there are increases
456 in SSA concentration over land (Fig. 4a and b) indicating transport (see also Text S2 and Fig.
457 S8). This is consistent with observed inland transport of SSA across the North Slope of Alaska
458 (Simpson et al., 2005). It is likely that leads emit smaller SSA particles relative to open ocean
459 emissions (Nilsson et al., 2001), which would increase their lifetime, so non-local impacts from

460 leads may be greater than simulated here. This further highlights the need for observations in
461 other regions to better understand the impacts of lead emissions.

462
463 There is strong observational evidence that lead emissions contribute to cold season SSA (see
464 Sect. 1), but the standard model consistently overpredicts observed SSA concentrations prior to
465 inclusion of additional lead emissions. This suggests other sources of SSA may be overpredicted
466 or sinks of SSA may be underpredicted. Ongoing work to improve the treatment of aerosol wet
467 removal processes in GEOS-Chem has not specifically investigated the impacts on sea salt
468 aerosol (Luo et al., 2020; Luo and Yu, 2023). Additionally, a recent observational study (Chen et
469 al., 2022) suggests that the GEOS-Chem blowing snow emissions parameterization may
470 overpredict the frequency of blowing snow events, therefore possibly contributing to the
471 overprediction of Arctic SSA mass concentrations.

472



473
474 **Figure 8-** Model evaluation for the cold season 2002-2003 at Zeppelin (a), Utqiagvik (b), Alert (c),
475 and Pallas (d). Observed Na⁺ concentrations are included as monthly averages for 2002-2003
476 (black + standard deviation margin), and the multi-year monthly averages (light blue + standard
477 deviation margin) (note: no 2002 data is available at Pallas). We show monthly average modeled
478 Na⁺ concentrations for 2002-2003 for the standard + leads (orange) and standard (blue)

479 simulations with two additional sensitivity studies: open ocean only emissions contributing to Na⁺
480 concentrations (dark blue with dashes) and the standard + leads emissions with Luo et al. (2020)
481 wet deposition applied (green line with dashes + dots). Note the different y-axis for Alert (c), as
482 concentrations are much lower at this site.

483
484 To test these possible sources of uncertainty, we run two additional sensitivity simulations for one
485 cold season (November 2002-April 2003): (1) using the Luo et al. (2020) wet deposition scheme
486 with the standard + leads SSA emissions (“standard + leads + Luo Wet Deposition”) and (2)
487 turning off blowing snow emissions in the standard model for an “open ocean only” case (see Text
488 S3 for further description). We find that the Luo wet deposition scheme improves model
489 agreement most at Zeppelin (see Fig. 8a), especially in the months of November, December,
490 March, and April. At Utqiaġvik, the Luo wet deposition scheme results in underestimates in Na⁺
491 concentrations compared to observations (Fig. 8b) in December, January, and April and
492 overestimates in November, February, and March; however, the overestimated months are closer
493 to the observed concentrations than the standard + leads and standard simulations. Additionally,
494 the standard model at Utqiaġvik agrees with observations in December and the standard + leads
495 model agrees with observations in January and April.

496
497 At Alert, the Luo wet deposition scheme decreases the model overestimate of the standard +
498 leads simulation when compared to the observations for the 2002-2003 cold season (Fig. 8c) but
499 still overestimates Na⁺ concentrations in each month. As the 2002-2003 observations at Alert are
500 particularly low, we also include the observed multi-year (2002-2008) monthly average Na⁺
501 concentrations for comparison. The Luo wet deposition scheme improves model evaluation from
502 February-March compared to the multi-year average observed concentrations at Alert, but
503 otherwise underpredicts concentrations. The Luo wet deposition scheme decreases
504 overprediction at Pallas in February compared to observations from the 2003 cold season and
505 improves model agreement in March and April, but underpredicts Na⁺ concentrations in January
506 (Fig. 8d). As there are no available observations in 2002 at Pallas, we also include the observed
507 multi-year (2003-2008) monthly average Na⁺ concentrations for comparison. The Luo wet
508 deposition scheme underpredicts Na⁺ concentrations in November, January, and April and
509 overpredicts concentrations in December and March compared to the multi-year average
510 concentrations at Pallas.

511

512 At Utqiagvik, under predicted Na⁺ concentrations with only open ocean emissions (except in
513 November) suggest that this site is influenced by blowing snow emissions and/or lead emissions.
514 Of the four sites, blowing snow is most important and well-represented here, as it also improves
515 the modeled seasonality by correctly representing the December peak in Na⁺ concentrations in
516 the standard + leads and standard model; there may be larger uncertainty in the emissions
517 parameterization in other regions. At Zeppelin, Alert, and Pallas, even with open ocean emissions
518 only and the standard wet deposition, the model overestimates Na⁺ concentrations for all months
519 during the cold season for 2002-2003, except at Pallas in January, where only open ocean
520 emissions more closely match observations. Moreover, the open ocean only Na⁺ concentrations
521 are close in value to the standard + leads and standard concentrations, indicating Pallas is largely
522 influenced by open ocean emissions, rather than blowing snow and lead emissions.

523
524 The results of these sensitivity tests suggest that changes to wet scavenging may be more
525 important at higher altitudes, given the improvement in model evaluation at Zeppelin. Yet, the
526 inclusion of the Luo wet deposition scheme to the standard + leads simulation still overestimates
527 concentrations at Alert and generally leads to disagreement with observations at Utqiagvik and
528 Pallas (except in March and April at Pallas).

529 530 **4. Uncertainties Discussion**

531
532 Our model evaluation reveals SSA is overestimated in the standard and standard + leads model
533 at each of the 4 Arctic sampling sites, pointing to possible sources of uncertainty. First, we use
534 the Jaeglé et al. (2011) open ocean function for our lead emissions parameterization as it is the
535 standard SSA emission function in GEOS-Chem that has been previously evaluated across global
536 oceans. However, there are possible differences in the mechanisms and meteorological
537 dependencies of SSA emission from leads vs. the open ocean which could impact the magnitude
538 and spatial patterns of lead emissions. Some potential differences were investigated in a
539 summertime measurement study (Nilsson et al., 2001), where they derive an empirical lead
540 emissions flux equation with an exponential dependence on windspeed and no consideration of
541 SST (Eq. (S2) in SI). They found the emissions rate per area from leads is smaller than that of
542 the open ocean due to lower fetch in leads, which suggests the lead emissions estimated in our
543 study may be an upper limit when considering large leads only (>3km in size); however, this lead
544 fraction detected by AMSR-E may only include 50% of total lead area (Röhrs and Kaleschke,
545 2012). Additionally, Nilsson et al. (2001) suggest leads emit smaller SSA particles relative to the

546 open ocean, which would increase their lifetime and transport distance. To create a more robust
547 understanding of the different SSA emission mechanisms from leads vs. the open ocean, more
548 studies using size-resolved observations could be conducted within the areas we predict the
549 highest lead emissions, such as within the Bering Strait, Nares Strait, Wynniatt Bay in the
550 Canadian archipelago, and the eastern Greenland Sea.

551
552 Our sensitivity study results do not ultimately confirm the source(s) of overprediction within the
553 GEOS-Chem model. Blowing snow emissions are included as a standard source of SSA
554 emissions in the Arctic, but remaining uncertainties about the GEOS-Chem blowing snow
555 emissions parameterization (Chen et al., 2022) suggest a need for refinement. The results of the
556 standard + leads + Luo wet deposition simulation highlight there are also remaining uncertainties
557 associated with wet deposition schemes as the Luo et al. (2020) mechanism does not lead to
558 consistent improvement of simulated SSA concentrations. Luo & Yu (2023) find that the scheme
559 overestimates wet scavenging on a global scale, so continued improvement in the model
560 deposition processes may resolve SSA overestimates. In addition, there are sparse ground
561 observations of precipitation in the Arctic, and while the MERRA-2 reanalysis uses both model
562 and satellite data to fill these gaps, Arctic cloud properties and precipitation can still be difficult to
563 predict (Barrett et al., 2020; Taylor et al., 2019), which could affect the accurate simulation of
564 aerosol deposition and, in turn, our simulated SSA concentrations.

565

566 5. Conclusions

567

568 Observational evidence (Chen et al., 2022; Kirpes et al., 2019; May et al., 2016; Radke et al.,
569 1976; Scott and Levin, 1972; Willis et al., 2018) and one modeling study of the 400 km² region
570 around Utqiagvik, Alaska (Ioannidis et al., 2023) have shown that leads may be an important
571 source of cold season SSA for the coastal Arctic. Here, we evaluate their importance as an Arctic-
572 wide source of cold season SSA emissions and their potential atmospheric chemistry impacts in
573 the global chemical transport model GEOS-Chem.

574

575 We find that lead SSA emissions occur primarily in regions where other SSA emissions sources
576 are very low, mainly within the Bering Strait, Nares Strait, Wynniatt Bay in the Canadian
577 archipelago, and the eastern Greenland Sea. Poleward of 75° N, leads increase total monthly
578 cold-season SSA emissions by 5.6 to 7.5%, with the highest contribution of SSA emissions from
579 leads in January and the lowest in April. Lead emissions vary in magnitude by month and year,

580 mainly due to variations in lead area. Future trends in Arctic sea ice predicted by climate models
581 suggest a possible future increasing trend in lead area (Intergovernmental Panel On Climate
582 Change, 2023), which would increase lead emissions. The additional SSA from leads in regions
583 where the background aerosol concentrations are low could also affect local aerosol-cloud
584 interactions, but the overall warming or cooling effect of these additional aerosols remains
585 uncertain (Cox et al., 2015; Schmale et al., 2021; Stramler et al., 2011; Tan et al., 2023; Villanueva
586 et al., 2022).

587
588 SSA mass concentrations increase primarily at the location of lead emissions, in regions where
589 the standard SSA mass concentration is very low ($\leq 1.2 \mu\text{g m}^{-3}$). Throughout the cold season, the
590 increased SSA mass concentrations from leads remain relatively constant in magnitude and
591 spatial distribution. The highest increase in multi-year average SSA mass concentrations due to
592 leads, spatially averaged for $\geq 75^\circ\text{N}$, occurs in November ($5.7\% \pm 5.2\%$) and the lowest occurs in
593 April ($3.7\% \pm 2.9\%$). Increased SSA from leads increases surface Br concentrations during the
594 cold season in corresponding locations. We find total Arctic-wide ($\geq 60^\circ\text{N}$) increases in multi-year
595 mean surface Br concentration range from 2.8 to 8.8%. The increases in Br are not sufficient to
596 have an impact on ozone; subsequent decreases in average surface ozone concentrations in the
597 Arctic are negligible ($< -0.25\%$).

598
599 Overall, we predict sea ice leads may impact Arctic-wide cold-season SSA concentrations and Br
600 concentrations by up to 5-10% on average during the 2002-2008 period. As leads are likely to
601 increase in prevalence under climate change, including this source of SSA in chemistry and
602 climate models may become more important for future predictions.

603
604 **Code and Data Availability**
605 Standard model code: <https://doi.org/10.5281/zenodo.5500717>; AMSR-E data:
606 <https://www.cen.uni-hamburg.de/en/icdc/data/cryosphere/lead-area-fraction-amsre.html>);
607 observational site data (Alert, Pallas, and Zeppelin): <https://ebas-data.nilu.no/Default.aspx>;
608 observational data (Utqiagvik): <https://saga.pmel.noaa.gov/data/stations/>; model data shown in
609 paper: <https://doi.org/10.5281/zenodo.14611355>.

610
611 **Author Contribution**

612 EJE was responsible for data curation, model simulations, validation, visualization, and analysis
613 with expert advice from HMH. HMH is responsible for conceptualization. EJE drafted the
614 manuscript which was revised by HMH.

615

616 **Competing Interests**

617 The authors declare that they have no conflict of interest.

618

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623 Urbana-Champaign.

624

625 **Supplemental Information**

626 Equations of SSA flux from Jaegle et al. (2011) and Nilsson et al. (2001); Additional figures of lead
627 SSA emissions for months other than January during the cold season; Cold season total lead
628 SSA emissions; Description and figure of the correlation between lead area and lead SSA
629 emissions; long-term trends in lead area (2002-2011) and relevant statistical testing; additional
630 figures of multi-year (2002-2008) mean percent increase due to leads in SSA and bromine
631 concentration for months other than January during the cold season; Description and figures of
632 correlation between lead emissions and coarse and accumulation mode SSA concentration;
633 Sensitivity simulations.

634

635 **6. References**

636

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