Dear Prof. Delphine Bosch,

Firstly, we would like to thank you for your valuable comments. Detailed below are our comments and revisions to the manuscript based on your recommendations and comments.

The 1st bullet point: i- provide, more consistently, evidence that the trace element balance of the different samples was not significantly affected by crystallisation fractionation processes. In other words, that the different spectra observed do not result from a co-precipitation effect of the minerals. This can be achieved by drawing new figures combining certain LILE elements according to a differentiation index such as SiO_2 wt% or MgO wt%. This has major consequences for the author's interpretations, because if the distribution of trace elements can be proven to be linked to fractionation processes, then the different batches of samples proposed by the authors are not strongly demonstrated to be effective, nor is their origin ;

Response: This reviewer has suggested adding some plots that include LILE, in order to demonstrate the independence of the major rock series (i.e., namely the LILE-rich and LILE-poor series). A proof that one of these series is not developed by the differentiation of the other series. To carry out this task, the SiO_2 versus Ba/Nb and SiO_2 versus Ba/La plots are now added to the manuscript (Fig. 6.a, b). On these plots the more mafic samples show distinct compositional spectra that support different origins for the LILE-rich and LILE-poor series.

The following text is now added to the 3rd paragraph of section 6.3 (Does fractional crystallization provide an interseries link?):

"Further assurance as to the primitive nature of the ternary classification of the volcanic rocks from the study area (i.e., that the geochemical distinction between these three rocks series are not an artifact of fractional crystallization) are provided by a series of plots where the mafic samples define distinct groupings/series. SiO_2 versus Ba/Nb or Ba/La help distinguish the LILE-rich from the LILE-poor series (Fig. 6.a, b) whereas SiO_2 versus Nb or Zr help distinguish the ITE-rich series from the other two series (Fig. 6.c, d)."

The 2nd bullet-point comment: better demonstrate the validity of the proposed subdivision into three categories of samples, namely ITE-rich samples, LILE-poor samples and LILE-rich samples. Looking at Figure 4, the distinction between LILE-poor and LILE-rich samples is not obvious, even for Kahak or Ardestan samples. Please add additional figures using individual LILE and inter-element variations to unequivocally demonstrate the existence of the three batches of samples proposed on the basis of strong discriminating features;

Response: The reviewer requested a better demonstration of the validity of the subdivision of samples from this study into the three categories; the LILE-rich, LILE-poor, and ITE-rich series. By adding the above mentioned plots; SiO_2 versus Ba/Nb, Ba/La, Nb, and Zr, this has now been carried out (see Fig. 6).

The 3rd bullet-point comment: the discussion of fluids and melts (wet or dry) from the subducted slab merit further development as, as it stands, the authors first suggest the contribution of wet slab

melts and dry slab melts to explain the LILE-rich and LILE-poor series, respectively. However, later in the text, they also suggest the involvement of fluids. To help the reader follow the proposed scenario, it would be easier to present the different processes in the order in which they occur, i.e. dehydration first, then wet slab melting and eventually dry slab melting. Fluid or melt input can also be controlled using ratios of elements with similar partition coefficients during mantle melting but contrasting behaviour in aqueous fluids [e.g. Ba/La, U/Th and Pb/Ce]. In the discussion, the distinction between the contribution of fluids and the contribution of melts is not clear. Th is an element considered to be immobile, but if there is a melting process, even a wet one, it will be enriched in the corresponding rock. The conditions required to produce melting in a subduction zone impose strong changes in the subduction parameters and in particular a higher plate temperature of the order of 700-900°C at depths less than the arc, conditions required to transfer significant quantities of trace elements from the plate to the mantle (see Hermann and Spandler, 2008; Hermann et al., 2006 and others);

Response: The term wet-slab-melt and dry-slab-melt have now been modified into fluid-rich slab melt and fluid-poor slab melt to better match the realities. The text has also been rechecked to make sure that the term "fluid-involvement" communicates the degree of contribution by one of these two end-members (i.e., fluid-rich slab melt and fluid-poor slab melt).

As per slab melting conditions, the following revision has been made. This includes an emphasis on the spatial distinctions of the LILE-rich and the LILE-poor series that furnishes further explanations as to their distinct petrogenesis.

In section 6.4.2 (i.e., Simultaneous fluid-rich slab melts and fluid-poor slab melts mantle metasomatism scenario) the following text:

"At first, hydrous slab melting at a given depth lead to the mantle metasomatism by hydrous slab partial melts. It is shortly followed by a modest change in the physiochemical conditions that governed the slab-mantle wedge system. The new conditions prompted partial melting of the slab that has already undergone dehydration, leading to the development of second type of mantle metasomatism by anhydrous slab partial melts"

has been replaced by the following text:

"Young hot slabs are capable of yielding partial melts at subarc depths that correspond to the volcanic front (Hermann and Spandler, 2008; Manea et al., 2014; Zheng, 2019). Hot slab subduction of then narrow/almost vanished Neotethyan oceanic plate has likely been responsible for petrogenesis of the slab melts that metasomatized the source mantles of the three series; the LILE-rich, LILE-poor and ITE-rich series. Disappearance and termination of the Neotethyan oceanic slab is estimated to have predated the early Cenozoic (Agard et al., 2005; Horton et al., 2008; Dargahi et al., 2010). The LILE-rich and LILE-poor series rocks are mainly spatially constrained towards the northern and southern parts of the study area. This implies a more hydrated, and more altered nature of the northern segment of the slab whereas the southern segment of the slab is deemed to have been of higher T (i.e. they might have been thicker and thinner slabs respectively)."

The 4ᵗʰ bullet-point comment: the ternary mixing model used to calculate the different proportions of the various end-members involved in the mixing lacks details on the procedure used to obtain such proportions such as the formulae, what type of mixing? mass to mass?, KD, the errors in the percentages obtained?;

Response: In the revised Table S5, we have now included a more comprehensive explanation of the procedure used to calculate the proportions of the different end-members. Details have also been added to the modified Table:

"The mixing proportions were determined using a simple mass-balance approach, adapted from the two-end-member mixing equations proposed by Powell (1984). For each element, the concentration in the mixed magma (C_m) can be expressed as:

$$C_m = X\,(C_a - C_b) + C_b$$

Where C_a and C_b are the concentrations of the element in the two end-member magmas, and X represents the degree of mixing. For isotopic compositions, the equation is:

$$I_{C_m} = I_{C_a}\left(\frac{C_a X}{C_m}\right) + I_{C_b}\left(\frac{C_b\,(1 - X)}{C_m}\right)$$

where I_{C_a} and I_{C_b} are the isotopic ratios in the two end-member magmas, and I_{C_m} is the isotopic ratio in the mixed magma.

The 5ᵗʰ bullet-point comment: section 7.6 is unclear as it stands. I would therefore recommend to remove it.

Response: As suggested by the Reviewer #2, Section 7.6 that include a figure (i.e., Age vs. ƐNd) is removed. As a result, the following modifications have been made:

In Fig. 5: The following text is added in the end of the figure caption:

"Gray areas/bands marks restricted Sr-Nd isotopic variation for most of the mafic rocks. One sample from the ITE-rich series plots towards less radiogenic Sr-Nd isotopic ratios (see Section 6.5 and Fig. 4.a for details). Two samples from the LILE-poor series plot towards more radiogenic Sr-Nd isotopic ratios; this is due to the higher slab components (i.e., slab + sediment) involvement (see Section 6.6 and Fig. 4.a)."

The related text is removed from Abstract.

The related text is removed from Conclusion.

In Section 6.3 (Does fractional crystallization provide an interseries link?), second paragraph, the phrase:

"….. see Sections 7.5 and 8)"
is replaced by:
"….. see Sections 6.5 and 6.6 and Fig. 4)"

In Section 6.5 (The incompatible trace elements-rich (ITE-rich) series), the following is added at the very end of second paragraph:

"(see caption in Fig. 5)"

Miscellaneous

Comment: l.145: "Accessory minerals are trace element repositories that resist the mixed acid digest method." This is true for zircon or spinel but not for other accessory minerals.

Revision made: The term "Accessory minerals" is now replaced by "Some Accessory minerals".

Comment: Section 3 "E_1-E_6 subdivisions versus age dated volcanic succession »: in its present state, I think this paragraph fits better in the Geological background section.

Revision made: As suggested by the referee, the content of this section is now distributed in three parts of the section "Geology of the study area".

Comment: Section 5.2: please indicate the error level for Nd isotopes. This data is also missing in Table S2. Please fill in.

Revision made: The error level for Nd isotopes is now added to Table S3.

Comment: l.155: precision on U content is mentioned but this element is not indicated in Table S1 nor in the trace elements patterns. Why?

Revision made: U is now added in the trace element patterns (Fig. 3) and in Table S2. The following text is also added as the 3rd paragraph in Section 6.2 (Primitive signatures):

"Most of the mafic volcanic rocks from the study area show deep negative anomalies for Th but mild or no anomaly for U similar to the island arc volcanic (Hawkesworth et al., 1997). It is because Th behaves like HFSE whereas U is mobilised by fluids. However, a few samples from the LILE-poor series show positive Th-U anomalies (Fig. 3). One of these samples show the highest LREE and P abundances amongst the mafic rocks (sample SA.22; Fig. 3). These might indicate apatite involvement in petrogenesis of the mafic rocks with positive Th-U anomalies. Apatite show extremely high partition coefficients for Th, U, and LREE (O'Reilly et al., 1991). Morishita et al. (2003) found that apatite development in the Finero peridotite is due to the metasomatizing agent derived from subducting 'slab possibly containing small quantity of sediment', the condition supported by the isotopic modelling of source end-members for the mafic volcanic rocks from the study area (see Section 6.6)."

Comment: Section 7.3.2 "Simultaneous slab fluids and slab melts mantle metasomatism scenario": I do not follow how it will be possible to produce simultaneously dehydration and melting processes at a same place?

Revision made: Due corrections are made in this section of the manuscript. These include the following:

"The LILE-rich and LILE-poor series rocks are mainly spatially constrained towards the northern and southern parts of the study areas, namely the Kahak and Ardestan areas respectively."

Thereby it is made clear that different styles (compositions/affinities) of magmatism occurred concurrently in different parts of the study area.