

Intertidal Regions Regulate Seasonal Coastal Carbonate System Dynamics in the East Frisian Wadden Sea

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Abstract. Seasonal and regional changes in carbon dynamics in the Wadden Sea, the world's largest intertidal sand and mud flats system, were analysed to quantify the influence of biogeochemical processes on the carbonate system at the land-sea interface. With a focus on the East Frisian Wadden Sea (EFWS), we successfully used the difference between total alkalinity (TA) and dissolved inorganic carbon (DIC) ([TA-DIC]), $\Delta\text{TA}_{\text{excess}}$, $\Delta\text{DIC}_{\text{excess}}$ (deviations beyond conservative mixing) and
15 $\Delta\text{TA}_{\text{P}}$ (alkalinity production linked to primary production) to quantify local biogeochemical influences on carbonate system dynamics.

In spring, a phytoplankton bloom with high biological activity, indicated by (a) supersaturated oxygen (up to 180 in % saturation), (b) elevated chlorophyll *a* (up to 151.7 $\mu\text{g L}^{-1}$) and (c) low $p\text{CO}_2$ (as low as 141.3 μatm), resulted in decrease in nitrate (NO_3^- , $19.29 \pm 18.11 \mu\text{mol kg}^{-1}$) and DIC ($159.4 \pm 125.4 \mu\text{mol kg}^{-1}$), and a slight increase in TA ($9.1 \pm 29.2 \mu\text{mol kg}^{-1}$).
20 The regression analysis of the differences between March and May 2022 ($\Delta\text{March-May}$) in NO_3^- concentrations (ΔNO_3^-) against the differences in DIC (ΔDIC) yielded a slope of 6.90, matching the Redfield C:N ratio, and suggesting that uptake of nitrate by primary producers increased total alkalinity during the spring bloom.

In summer, we assume that organic matter remineralization, along with CaCO_3 dissolution in sediments, enhances TA production in the coastal and nearshore regions of the Western EFWS (up to 2400 $\mu\text{mol kg}^{-1}$). In the Eastern EFWS, enhanced
25 CaCO_3 formation may consume TA ([TA-DIC] < 200 $\mu\text{mol kg}^{-1}$), but the region still acts as a net source of TA, likely due to sedimentary processes such as organic matter decomposition, which follow the time of increased biological activity during the spring bloom. The increase of TA enhances the coastal ocean's ability to absorb and store CO_2 through buffering and suggests that the EFWS can be a source of TA to the coastal regions during the warm productive seasons. This study highlights the complex relationships between these factors, emphasizing the need for a comprehensive understanding of regional and seasonal
30 variations to better assess the role of coastal systems in carbon cycling and storage, and climate regulation.

1 Introduction

Coastal oceans are biogeochemically active regions, which play a significant role in biogeochemical cycles, despite covering less than 10 % of the oceanic realm (Gattuso et al., 1998). Coastal regions are directly affected by terrestrial organic matter and nutrients through river run-off, groundwater discharge, atmospheric deposition, and exchange of large amounts of matter and energy with the open ocean (Borges et al., 2006; Gattuso et al., 1998). Overall, coastal oceans also support approximately 14 - 23 % of the ocean carbon dioxide uptake, 10 - 30 % of the primary production, 80 % of organic matter burial, 90 % of sedimentary mineralization and 75 - 90 % of the oceanic sink of suspended river loads (Bauer et al., 2013; Gattuso et al., 1998). Since the start of the industrial era, the levels of CO₂ in the atmosphere have increased from ~280 ppm to over ~419 ppm due to human activities (Friedlingstein et al., 2023). About 30 % of anthropogenic CO₂ emissions since the industrial period have been absorbed by the ocean (Friedlingstein et al., 2023). The uptake increases the concentration of protons [H⁺] and decreases the carbonate ion concentration [CO₃²⁻], leading to lower pH and a reduced saturation state of calcium carbonate, a process known as ocean acidification (Orr et al., 2005). Depending on the different model scenarios, it is predicted that surface pH in the ocean might decline by about 0.3 – 0.4 pH units by 2100, corresponding to a decrease of about 40 – 50 % of carbonate ions in the seawater (Feely et al., 2009; Orr et al., 2005). The future capacity of the ocean to take up CO₂, can affect the precipitation and dissolution of the carbonate minerals, as well as the survival of marine organisms (Duan et al., 2023; Kroeker et al., 2013; Liang et al., 2023; Ricour et al., 2023). Whereas oceans are a significant sink for anthropogenic CO₂ it is not well known how this uptake will further change under the continual increase of anthropogenic CO₂ in the atmosphere (Lorkowski et al., 2012; Sabine et al., 2004; Thomas et al., 2007). However, rising atmospheric CO₂ will influence carbon stocks and fluxes in the pelagic, benthic and coastal zone, particularly in shelf seas, which are annually mixed and ventilated (Legge et al., 2020).

Total alkalinity (TA) represents the buffering capacity of the ocean and is controlled by many factors including erosion/weathering pathways on land (Lehmann et al., 2023) and respiration of organic matter (OM) along anaerobic metabolic pathways, mostly generated in shallow marine and shelf sediments (Dickson, 1981). These processes are directly influenced by terrestrial and anthropogenic nutrient inputs (Burt et al., 2016; Thomas et al., 2009; Van Beusekom & De Jonge, 2002), as well as increased sedimentation of reactive organic matter (OM) (Al-Raei et al., 2009; Böttcher et al., 1998). Coastal seas and shelf seas, like the North Sea often have relatively high rates of primary production in spring, leading to drawdown in DIC, *p*CO₂ and a consequent pH increase (Macovei et al., 2021; Thomas et al., 2005). In addition, nutrient loads from land contribute to enhanced primary production, and increased carbon remineralization (Prowe et al., 2009; Thomas et al., 2009), and subsequent changes in the carbonate system. Previous studies suggested that the Wadden Sea, a large network of intertidal sand and mudflats bordering the North Sea along the Dutch, German and Danish coasts (Staneva et al., 2009), plays a significant role in modulating local carbonate system dynamics (Thomas et al., 2009; Voynova et al., 2019). The seasonal TA production in the southern North Sea, exhibited a regional TA variability, which was attributed to the influence of the Wadden Sea (Voynova et al., 2019), and to benthic TA production (Brenner et al., 2016). The observed seasonal changes of TA can

affect the coastal ocean capacity to absorb carbon from the atmosphere (Burt et al., 2016; Gruber et al., 2019; Li et al., 2024; Schwichtenberg et al., 2020).

65 This study offers a detailed analysis of seasonal and inter-annual carbonate system dynamics in the EFWS, a key region within the world's largest intertidal sand and mudflat system (UNESCO World Heritage Centre, n.d.). The research focuses on seasonal and regional variations, and particularly on the drivers of TA production in spring-summer. By examining changes in TA, DIC, and other biogeochemical parameters, this study identifies key mechanisms affecting carbonate chemistry, including calcium carbonate (CaCO_3) dissolution and formation, photosynthesis and respiration. Our findings highlight the
70 complex interactions among these factors, emphasizing the need for a comprehensive understanding of regional and seasonal variations to better assess the role of coastal systems in carbon cycling, storage, and climate regulation. Moreover, observed TA and DIC distributions provide valuable insights into potential carbon and TA sources within this dynamic environment.

2 Material and Methods

2.1 Study Side

75 The German Bight region is bordered by Germany, Denmark and the Netherlands, situated in the southeastern corner of the North Sea (Fig. 1a). The East Frisian Wadden Sea (EFWS) is one of the shallowest regions of the German Bight, characterized by a series of barrier islands (Fig. 1b), each 5 – 17 km long and 2 – 3 km wide (Staneva et al. 2009). The system is an intertidal region, influenced by semidiurnal tides with a tidal range from approximately 2.2 - 2.8 m (Grunwald et al. 2009; Staneva et al. 2009) and up to ~ 3.5 m in the Elbe River mouth (Staneva et al. 2009). Main freshwater sources to the German Bight are the
80 rivers Elbe, Weser, and Ems, with additional river discharge values of $860 \text{ m}^3 \text{ s}^{-1}$ (Elbe), $323 \text{ m}^3 \text{ s}^{-1}$ (Weser), and $80 \text{ m}^3 \text{ s}^{-1}$ (Ems) (Luebben et al., 2009; Schmidt et al., 2011 and references herein). However, model estimates indicate that only about 5% of the TA variability in the German Bight is attributable to effective river loads (Schwichtenberg et al., 2020). Seasonally, the river inflow and hydrodynamics can affect the carbonate system in the German Bight. During winter, high freshwater input (e.g., Elbe ~ $1420 \text{ m}^3 \text{ s}^{-1}$ in February) leads to reduced salinities, whereas in summer discharge is considerably lower (e.g., Elbe
85 ~ $300 \text{ m}^3 \text{ s}^{-1}$ in August), resulting in comparatively more marine conditions (Rewrie et al., 2023; Schwichtenberg, 2013; Schwichtenberg et al., 2020; Voynova et al., 2017).

The study site is also characterized by complex local freshwater dynamics. Local sluices (e.g., in Greetsiel, Bengersiel, Harlesiel, Carolinensiel, or Neuharlingersiel) regulate freshwater inflow in intertidal regions and can thereby affect salinity levels and nutrient dynamics in the Wadden Sea (Luebben et al., 2009). The opening of such sluices is primarily controlled by
90 tidal cycles but can also vary seasonally depending on precipitation and drainage needs. As shown by Beck et al. (2012) freshwater discharged through a sluice along the East Frisian coast, can be enriched in terrestrial organic matter and trace metals and undergoes strong geochemical modifications upon mixing with seawater. Such processes, comparable to those in estuaries, are likely to influence not only trace metal dynamics but also carbonate chemistry and nutrient availability in the

Wadden Sea. Terrestrial run-off enters the back-barrier tidal flats of Spiekeroog via the sluice located in Neuharlingsiel (Fig. 1b), generally about 2 h before low tide when the water level of the hinterland is above sea level, with drainage volume depending on precipitation in the catchment area (ca. 125 km²) and averages 15×10^6 m³ in winter and 9×10^6 m³ in summer (Luebben et al., 2009 and references therein).

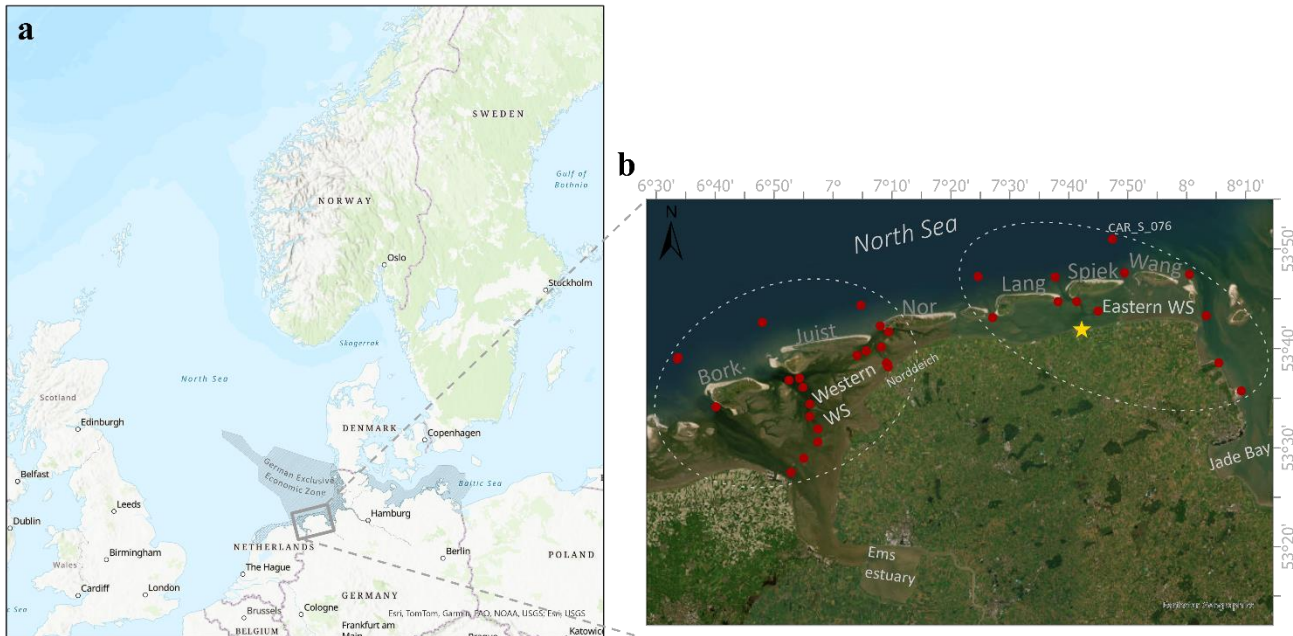


Figure 1: Study area, (a) The German Wadden Sea (=WS), (shown in grey) with the German Exclusive Economic Zone (hatched region) and the East Frisian Wadden Sea (EFWS) region (grey box). (b) Zoom into the EFWS, showing the sampling stations (red dots) in the different regions (Western WS and Eastern WS) considered in this study. The yellow star shows the Neuharlingsiel sluice location. Station CAR_S_076, used as the North Sea endmember, is explicitly labelled (see section 2.5). The different Islands are labelled with the first letters of their names (Bork= Borkum, Juist, Nor= Norderney, Lang= Langeoog, Spiek= Spiekeroog, Wang= Wangoog). The map in this Figure was generated using ArcGIS. Data sources: (a) Esri, TomTom, Garmin, FAO, NOAA, USGS; (b) Earthstar Geographics. © 2024 Julia Meyer.

We separated the study area into two regions (Western WS & Eastern WS; Fig. 1b) due to the differences in the tidal dynamics and hydrodynamic properties (Herrling & Winter, 2015), which drive the area's carbon dynamics, sediment transport, and overall ecological functioning. The tidal range increases from 2.4 m at Borkum to 3.0 m at Wangeroog, with the Eastern region experiencing stronger tidal influences (Herrling & Winter, 2015). In the Langeoog basin, wind effects cause the largest relative increase in residual discharge, while Norderney experiences the largest absolute increase in water flux. Westerly winds influence residual circulation and sediment transport differently in each region. The flow dynamics between the basins of Baltrum and Langeoog are interconnected, whereas the flow regime between Borkum and Norderney is more independent from the dominant circulation patterns and inter-basin exchange (Herrling & Winter, 2015).

Seasonal cruises were completed in the Wadden Sea (WS) and North Sea around the East Frisian Islands (and the Ems River estuary) with the research vessel (RV) *Burchana* (NLWKN; Fig. 1b). The July 2021 and October 2021 cruises focused exclusively on the intertidal mudflats of the EFWS. The RV *Burchana*, with a draft of 1.3 meters, allowed sea water sampling

even during low tide, providing the opportunity to collect samples at various tidal stages. Sampling took place during daylight hours, typically starting in the morning at low tide and continuing throughout the day, with no nighttime samples collected. Both shallow intertidal areas (accessible due to the vessel's low draft) and deeper subtidal channels were sampled to ensure comprehensive spatial coverage. Later cruises extended into the Ems River Estuary, from the island of Borkum (Bork, Fig. 1b) to Weener (53°09'55.4"N 7°20'39.9"E), a town located upstream the Ems River Estuary. These additional data from the Ems River will be presented in a subsequent paper.

2.2 FerryBox measurements

A FerryBox system (4H-JENA engineering GmbH, Jena, Germany) was operated during all cruises on board the *RV Burchana* (NLWKN), measuring the following parameters every minute: temperature (SBE38, Sea-Bird Scientific, USA), salinity (SBE45, Sea-Bird Scientific, USA), dissolved oxygen (DO) (Optode 4835, Aanderaa, Bergen, Norway), chlorophyll fluorescence to estimate chlorophyll a concentrations (AlgaeOnlineAnalyser, bbe moldaenke, Germany), pH (electrode, Xylem; measured on total scale, USA) and turbidity (Solitax inline SC, Hach Lange, Germany). The pH electrode was calibrated using standard DuraCal buffer solutions at pH 7 and pH 10 (Hamilton Company, USA). The partial pressure of CO₂ (*p*CO₂) was measured using a sensor (HydroC CO₂-FT, 4H-JENA engineering, Germany) attached to the flow-through system of the FerryBox. These data were corrected using the data processing manual of 4H - Jena Engineering GmbH (4H Jena, 2021). Regions near the ports were excluded in all datasets, to remove any influence from the port, or from the cleaning cycles.

2.3 Discrete samples

Samples for Winkler titration (dissolved oxygen), nutrients, salinity and turbidity were collected from each cruise to crosscheck the measurements of the FerryBox. Therefore, dissolved oxygen (DO) samples were collected on the last day of each cruise by filling Biological Oxygen Demand (BOD) bottles from the FerryBox outflow. Three replicates collected underway were treated according to the standard Winkler method and measured within 24 hours in the lab using a Metrohm 870 KF Titrino Plus (Germany). The Winkler titrations were used to correct the FerryBox dissolved oxygen measurements from the seasonal cruises. A single regression equation of the Winkler titrations and the FerryBox data was used to apply a dissolved oxygen correction of the entire data set ($y = 1.17x - 31.89$, $R^2 = 0.98$, $n = 46$). Here, x is the measured raw data of the FerryBox optode in $\mu\text{mol L}^{-1}$, and y is the corrected dissolved oxygen value based on Winkler titrations, also in $\mu\text{mol L}^{-1}$.

In addition, the apparent oxygen utilization (AOU) was calculated using the corrected oxygen measurements, as according to:

$$\text{AOU} = \text{O}_2' - \text{O}_2 \quad (1)$$

Where O_2' is the expected oxygen in $\mu\text{mol L}^{-1}$ at equilibrium with the atmosphere, at the measured temperature and salinity according to Grasshoff et al., (1999), and O_2 is the oxygen concentration measured by the FerryBox optode.

Duran bottles (~ 300 mL) were filled with sample water during all seasonal cruises using the FerryBox outflow to measure turbidity and salinity in the laboratory. Salinity was measured using an OPTIMARE High Precision Salinometer (Optimare Systems GmbH, Bremerhaven, Germany) and turbidity was measured using a Hach 2100 turbidimeter.

Nutrient samples were collected at each station using the onboard water sampler. A sample volume of 250 mL was filtered through pre-combusted GF/F filters, and the samples were collected in clean centrifuged tubes, frozen and stored at -20 °C. The concentrations of nitrite (NO_2^-), nitrate (NO_3^-) ammonium (NH_4^+) and silicate (SiO_2) were also measured using a MicroMaC analyser from SYSTEA (Anagni (FR), Italy). The system, which induces a colour reaction, is coupled with a photometer (NO_2^- , NO_3^- , SiO_2) and a fluorometer (NH_4^+), using a one-point calibration. NO_3^- was determined with sulfanilamide and N- (1- Naphthyl)ethylenediamine, NH_4^+ with orthophthalaldehyde and NO_2^- with diethylenetriaminepentaacetic acid and automatic UV reduction (Luitjens, 2019).

Water samples for TA and DIC were collected in 300 mL BOD bottles at all stations using the FerryBox outflow according to the Standard Operating Procedure (SOP) (Dickson et al. 2007). The samples were poisoned with saturated mercury chloride and measured in the laboratory with a VINDTA 3C (MARIANDA, Kiel, Germany), using certified reference material (CRM) (Dickson et al. 2003). The method is precise, with a reproducibility better than $1 \mu\text{mol kg}^{-1}$ and an accuracy within $2 \mu\text{mol kg}^{-1}$ (Dickson et al. 2003). The results within this study were plotted with R Project (ggplot package) and maps were created with ArcGIS Pro.

In addition, the saturation state of calcite (Ω_{cal}) was calculated using the CO2SYS program developed by (Lewis & Wallace, 1998). The dissociation constants for carbonic acid (K_1 and K_2) were taken from Lueker et al. (2000), which are widely used for seawater CO_2 chemistry and generally provide reliable results, with uncertainties in K_1 and K_2 translating to about $\pm 1 \%$ in calculated $p\text{CO}_2$. For bisulfate (KHSO_4), the constant of (Dickson, 1990) was applied, and for the boron-to-chlorinity ratio of Lee et al. (2010).

2.4 [TA-DIC] as a proxy for Biogeochemical Processing

[TA-DIC] is a good proxy for biogeochemical processes such as CaCO_3 precipitation / formation, photosynthesis, respiration and therefore CO_2 uptake and release, even in coastal oceans (Xue and Cai, 2020). The parameter is independent of ocean mixing and not sensitive to temperature and pressure changes, and this makes it a good tracer for larger-scale oceanographic studies, and suitable for seasonal observations of biogeochemical processes and carbonate dynamics of an ecosystem. In addition, it is assumed that [TA-DIC], can better reflect variations of $[\text{CO}_3^{2-}]$ compared to the ratio of TA and DIC (Xue et al., 2017).

The difference between TA and DIC is often expressed as:

$$[\text{TA} - \text{DIC}] = \text{TA} - \text{DIC} \quad (2)$$

However, [TA-DIC] should not be used at low salinity (e.g. < 20) and when [TA-DIC] is $< \sim 50 \mu\text{mol kg}^{-1}$, where the relationships of [TA-DIC] with pH and/or ocean acidification are nonlinear, these low values also occur in oxygen minimum zones (Xue and Cai 2020).

2.5 Calculations of estuarine DIC and TA contributions

180 To estimate the contributions of estuarine DIC and TA in the Western and Eastern EFWS, we used DIC and TA measurements from the lowest salinity station located in the Wadden Sea, just before the Ems Estuary (where the river enters the Wadden Sea) as endmembers for our mixing model. The stations chosen for these measurements were in areas of low salinity within the estuary, with considerable freshwater influence (DIC_{estuary} ; TA_{estuary} ; Table 1). These values (DIC_{estuary} ; TA_{estuary}) were used to calculate the $DIC_{\text{mixing w/est}}$ and $TA_{\text{mixing w/est}}$ (equation 3 and 4) for the different regions. We hereafter denote these
185 conservative mixing values as $DIC_{\text{mixing w/est}}$ and $TA_{\text{mixing w/est}}$, i.e. the expected concentrations under two-endmember mixing between the North Sea (Wadden Sea) and estuarine sources.

Table 1: DIC_{estuary} , DIC_{NorthSea} , S_{NorthSea} , TA_{estuary} and TA_{NorthSea} used for the calculation of the seasonal $DIC_{\text{mixing w/R}}$ concentrations of each season.

Season and Region	DIC_{NorthSea}	DIC_{estuary}	S_{NorthSea}	TA_{NorthSea}	TA_{estuary}
July 2021	2144.1	2261.2	31.8	2273.7	2465.1
October 2021	2188.5	2224.9	31.0	2369.3	2411.7
March 2022	2205.7	2238.1	31.5	2338.4	2318.9
May 2022	2199.9	2205.5	32.1	2332.5	2447.8
July 2022	2144.4	2310.2	31.8	2356.6	2489.6

In this study, DIC_{NorthSea} (and TA_{NorthSea}) refers to the DIC (and TA) values at the station (CAR_S_076) located farthest from
190 land (Fig. 1b), which we used as our North Sea endmember, as it is situated farthest offshore of the Ems and experiences the highest salinity (S_{NorthSea}) levels during almost each season. The DIC_{NorthSea} and TA_{NorthSea} endmembers applied for each season are shown in Table 1.

Starting from these endmembers, the expected DIC and TA at each station due to conservative mixing were calculated following the equations of Jiang et al. (2008) and Joesoef et al. (2015):

$$195 \quad DIC_{\text{mixing w/est}} = \frac{S_i}{S_{\text{NorthSea}}} * DIC_{\text{NorthSea}} + \left(1 - \frac{S_i}{S_{\text{NorthSea}}}\right) * DIC_{\text{estuary}} \quad (3)$$

$$TA_{\text{mixing w/est}} = \frac{S_i}{S_{\text{NorthSea}}} * TA_{\text{NorthSea}} + \left(1 - \frac{S_i}{S_{\text{NorthSea}}}\right) * TA_{\text{estuary}} \quad (4)$$

where S_i represents the salinity at each related station i . The ratio $\frac{S_i}{S_{\text{NorthSea}}}$ normalizes the influence of the salinity at the specific station by the salinity of the North Sea (S_{NorthSea}).

To quantify the net contributions of estuarine processes beyond conservative mixing, $\Delta DIC_{\text{excess}}$ and $\Delta TA_{\text{excess}}$ were calculated,
200 using the equation of (Jiang et al., 2008; Van Dam et al., 2018):

$$\Delta DIC_{\text{excess}} = DIC_i - DIC_{\text{mixing w/est}} \quad (5)$$

$$\Delta TA_{\text{excess}} = TA_i - TA_{\text{mixing w/est}} \quad (6)$$

These calculations use the measured DIC and TA concentrations at station i , together with the $DIC_{\text{mixing w/est}}$ and $TA_{\text{mixing w/est}}$ values from equations 3 and 4. Here, $\Delta DIC_{\text{excess}}$ represents the deviation of the measured DIC from the expected conservative

205 mixing value, while $\Delta TA_{\text{excess}}$ quantifies the corresponding deviation for TA. Values close to zero indicate that concentrations are consistent with mixing alone. Positive or negative deviations reflect local biogeochemical processes (e.g., photosynthesis, respiration, carbonate precipitation or dissolution) or additional inputs and losses not represented by the two endmembers. By accounting for conservative mixing, this approach isolates the influence of biogeochemical processes and external contributions from the physical mixing of North Sea and estuarine waters.

210 During productive seasons, primary production will decrease DIC, while consuming CO_2 and nutrients (Xue & Cai, 2020). The uptake of NO_3^- can increase TA, while the uptake of NH_4^+ can decrease TA in an ecosystem (Brewer & Goldman, 1976; Wolf-Gladrow et al., 2007). The production of 1 mol organic matter ($(\text{CH}_2\text{O})_{106}(\text{NH}_3)_{16}\text{H}_3\text{PO}_4$) will generally increase TA by 17 mol TA (ΔTA_P) and decrease DIC by 106 mol and will change the TA and DIC concentrations (Chen, 1978). Consequently, ΔTA_P is used to calculate the expected amount of TA produced, according to the equation of ΔTA_{bio} from (Xue and Cai 2020),
 215 which was modified using the calculated $\Delta \text{DIC}_{\text{excess}}$:

$$\Delta TA_P = (-17/106) * \Delta \text{DIC}_{\text{excess}} \quad (7)$$

It is important to note, that the calculation of ΔTA_P is an overestimation assuming the measured TA is modulated by photosynthesis (NO_3^- - fuelled) or respiration, such that the presence of other non-photosynthetic DIC sinks will cause ΔTA_P to be overestimated.

220 3 Results

3.1 Regional and seasonal variation in the EFWS

In July 2021, salinities were lowest on the route from Norddeich harbour to Norderney (25.2) in the Western WS (Table 2; Fig. S1). In summer 2022, the salinity on July 12 (31.16 - 32.03) increased by 2 - 3 salinity units compared to July 11 (28.28 - 29.43) in the Eastern WS (Table 2; Fig. S1). This indicates that during July 2022 cruise, there was a change in salinity
 225 range between the first and second leg of the cruise (in roughly the same region). A possible reason for these salinity differences could be the Neuharlingersiel sluice opening in the early morning of July 11, 2022, which could have influenced the salinity on this day. The turbidity in October 2021 showed a completely different pattern compared to all other seasons, with values from 96.08 – 306.46 NTU (Fig. S1), caused by rougher weather conditions during the campaign.

A land-to-sea gradient in pH was observed in October 2021 (7.6–8.1; Table 2; Fig. S2), with higher values in offshore regions
 230 and lower values near the coast. Regional differences with maximum pH values > 8 in higher-salinity waters were observed in all seasons, with particularly high values in the Western Wadden Sea in May 2022 (Fig. S2; Table 2).

The maximum concentrations of chlorophyll a were observed in May 2022, displaying the most substantial fluctuations (Table 2), especially in the Western part (up to $151.7 \mu\text{g L}^{-1}$) (Fig. S2). In July 2021, the highest chlorophyll measurements (up to $74.7 \mu\text{g L}^{-1}$) were detected underway between Norderney to Spiekeroog (Table 2; Fig. S2). A similar pattern can be seen
 235 for July 2022 (Fig. S2), although measurements were not available for all transects.

Table 2: Overview of the different measured parameters: Temperature (°C), Salinity, pH, oxygen (% saturation), AOU ($\mu\text{mol L}^{-1}$), Chlorophyll a (Chl a in $\mu\text{g L}^{-1}$), $p\text{CO}_2_{\text{obs}}$ (μatm) and Calcite saturation Ω_{cal} of all seasonal samplings in the EFWS.

Parameter	July 2021	October 2021	March 2022	May 2022	July 2022
Temperature (°C)	18.59 – 23.15	12.66 – 16.57	3.89 – 6.47	9.65 – 18.19	17.36 – 21.04
Salinity	25.21 – 32.33	25.21 – 31.32	18.32 – 31.97	23.59 – 32.28	23.35 – 32.04
pH	7.71 – 8.16	7.64 – 8.13	7.66 – 8.13	7.95 – 8.55	7.71 – 8.08
Oxygen (% sat)	109.7 $\pm$ 9.5	100.9 $\pm$ 5.3	105.5 $\pm$ 3.1	132.9 $\pm$ 13.0	102.4 $\pm$ 5.2
AOU ($\mu\text{mol L}^{-1}$)	-22.7 $\pm$ 22.2	-2.3 $\pm$ 14.4	-17.5 $\pm$ 10.0	-88.9 $\pm$ 33.9	-5.3 $\pm$ 12.4
Chl a ($\mu\text{g L}^{-1}$)	25.6 $\pm$ 12.9	16.9 $\pm$ 8.9	13.4 $\pm$ 6.5	52.1 $\pm$ 29.4	23.2 $\pm$ 19.3
$p\text{CO}_2_{\text{obs}}$ (μatm)	–	521.6 $\pm$ 72.2	468.2 $\pm$ 54.5	269.0 $\pm$ 63.7	536.0 $\pm$ 116.5
Ω_{cal}	3.8 $\pm$ 0.9	3.21 $\pm$ 0.6	2.5 $\pm$ 0.4	5.2 $\pm$ 1.1	3.7 $\pm$ 0.4

In July 2021 the measured oxygen ranged from 72.4 – 112.01 in % saturation (Fig. S2), with the highest values in the Western WS. Oxygen decreased from July 2021 to October 2021, on average by 8.8 ± 10.9 % saturation (Table 2; Fig. S2). The lowest oxygen saturation was measured from Norddeich to Norderney in October (down to ~ 57 % saturation). Oxygen saturation then increased continuously until May, reaching maximum values up to ~ 180 % saturation in the Western EFWS and up to ~ 152 % saturation in the Eastern EFWS (Table 2; Fig. S2). Overall, oxygen decreased by a mean value of 28.4 ± 16.4 % saturation from May to July 2022 (Fig. S2), resulting in slightly lower oxygen saturation in July 2022 compared to the previous year (Table 2). For the AOU (Table 2), the same inverse picture was obtained. The observed $p\text{CO}_2$ ($p\text{CO}_2_{\text{obs}}$) was highest in July 2022 and lowest in May 2022 (down to 141.3 μatm) in the Western WS. The average decrease in $p\text{CO}_2_{\text{obs}}$ was 166.2 ± 276.1 μatm from March to May 2022 (Table 2; Fig. S3).

During all seasons, Ω_{cal} was supersaturated, ranging from 1.5 – 7.5 (Table 2), but with a pronounced seasonal pattern. Higher, supersaturated Ω_{cal} values (> 2.5) were observed during more productive seasons (July 2021, July 2022, May 2022). However, the highest variability of Ω_{cal} was found in May 2022 with values > 1 in 80 % of the stations and reaching up to 7.52 in the Western WS (Table 2; Fig. S3). In summer (July 2021, July 2022), a decrease of Ω_{cal} in the EFWS from West to East was observed, regionally (Fig. S3).

3.2 TA and DIC Variability in the EFWS

Large seasonal and regional variability in TA and DIC was observed along the estuarine land–sea gradient of the EFWS (Table 3; Fig. 2), with particularly high TA values in summer (Table 3; Fig. 3b). In July 2021, TA ranged from ~ 2273 $\mu\text{mol kg}^{-1}$ in the Jade Bay (Eastern WS) to ~ 2465 $\mu\text{mol kg}^{-1}$, following a negative mixing line with a slope of -24.9 $\mu\text{mol kg}^{-1}$ per salinity unit (Fig. 3b). Most data points were close to or slightly above this line, consistent with net TA production in this tidally driven system. In July 2022, TA was lower at reduced salinities (28–29) in the Eastern WS, although the slope versus salinity remained similar (-23.3 $\mu\text{mol kg}^{-1}$ per salinity unit). At higher salinities (31–32) in the Western WS, values were comparable to July 2021, again closely following the conservative mixing line (-25.1 $\mu\text{mol kg}^{-1}$ per salinity unit). The DIC concentrations also

showed a similar pattern during both summer cruises, with negative slopes of - 49.0 $\mu\text{mol kg}^{-1}$ and - 44.3 $\mu\text{mol kg}^{-1}$ per salinity unit in July 2021 and 2022, respectively, in the Eastern WS (Table 3; Fig. 3a).

Table 3: Overview of different the parameters TA ($\mu\text{mol kg}^{-1}$), DIC ($\mu\text{mol kg}^{-1}$) and [TA-DIC] ($\mu\text{mol kg}^{-1}$) of all seasonal samplings in average with the standard deviation.

Parameter	July 2021	October 2021	March 2022	May 2022	July 2022
TA ($\mu\text{mol kg}^{-1}$)					
Western WS	2403 $\pm$ 27	2389 $\pm$ 16	2359 $\pm$ 28	2379 $\pm$ 39	2380 $\pm$ 43
Eastern WS	2284 $\pm$ 10	2237 $\pm$ 5	2347 $\pm$ 12	2350 $\pm$ 10	2337 $\pm$ 35
DIC ($\mu\text{mol kg}^{-1}$)					
Western WS	2186 $\pm$ 48	2257 $\pm$ 91	2238 $\pm$ 38	2066 $\pm$ 107	2184 $\pm$ 43
Eastern WS	2124 $\pm$ 10	2183 $\pm$ 18	2212 $\pm$ 21	2089 $\pm$ 38	2142 $\pm$ 28
[TA-DIC] ($\mu\text{mol kg}^{-1}$)					
Western WS	218 $\pm$ 52	150.5 $\pm$ 47.1	121 $\pm$ 30	313 $\pm$ 76	197 $\pm$ 27
Eastern WS	159 $\pm$ 16	185 $\pm$ 20	138 $\pm$ 11	261 $\pm$ 35	195 $\pm$ 21

265 In October 2021, TA peaked at Norddeich in the Western WS (up to 2571 $\mu\text{mol kg}^{-1}$; Fig. 2). DIC values reached 2516 $\mu\text{mol kg}^{-1}$ and decreased eastward to 2158 $\mu\text{mol kg}^{-1}$ (Fig. 2). The TA mixing line in the Western WS had a negative slope ($-25.4 \mu\text{mol kg}^{-1}$ per salinity unit), while the Eastern WS showed a near-linear positive slope ($+13.1 \mu\text{mol kg}^{-1}$ per salinity unit) (Fig. 3d). DIC conservative mixing lines revealed a similar pattern, with a steep negative slope ($-46.9 \mu\text{mol kg}^{-1}$ per salinity unit) in the Western WS (Fig. 3c). This consistent difference across both parameters highlights strong spatial

270 heterogeneity in carbonate system dynamics

Relatively low DIC values ($<2200 \mu\text{mol kg}^{-1}$ in the Western WS) were observed outside the barrier islands in March, with slight increase towards the coast (Fig. 2). During this season, TA concentrations showed little deviation from conservative mixing during this season. The mixing lines exhibited weak negative slopes in both subregions ($-7.5 \mu\text{mol kg}^{-1}$ per salinity unit in the Western WS and $-6.8 \mu\text{mol kg}^{-1}$ per salinity unit in the Eastern WS; Fig. 3f), indicating that TA dynamics were

275 largely controlled by physical mixing with only a minor signal of net production. Similarly, DIC mixing lines revealed shallow negative slopes (-18.2 and $-15.9 \mu\text{mol kg}^{-1}$ per salinity unit in the Western and Eastern WS, respectively; Fig. 3e).

Mean TA values in the Western WS reached 2379 $\mu\text{mol kg}^{-1}$ in May (Table 3; Fig. 2), reflecting a moderate seasonal increase, with concentrations generally close to or slightly above the mixing line and showing similar negative slopes in the Western ($-15.6 \mu\text{mol kg}^{-1}$ per salinity unit) and Eastern WS ($-15.3 \mu\text{mol kg}^{-1}$ per salinity unit) (Fig. 3h). During this period, the signal

280 of TA production was most clearly expressed in the Western WS. The lowest DIC concentrations of the study (down to 1872 $\mu\text{mol kg}^{-1}$) were observed in spring (Fig. 3g).

We note that conservative mixing lines in intertidal areas are often difficult to resolve visually, especially when data cluster closely. To better capture non-conservative variability, we calculated $\Delta\text{DIC}_{\text{excess}}$ and $\Delta\text{TA}_{\text{excess}}$ for each season and region, reflecting deviations from mixing. In addition, $\Delta\text{TA}_{\text{P}}$ was derived as an upper estimate of photosynthetically driven TA

285 changes. These metrics reduce uncertainty and provide a clearer assessment of net production and consumption than graphical mixing lines alone.

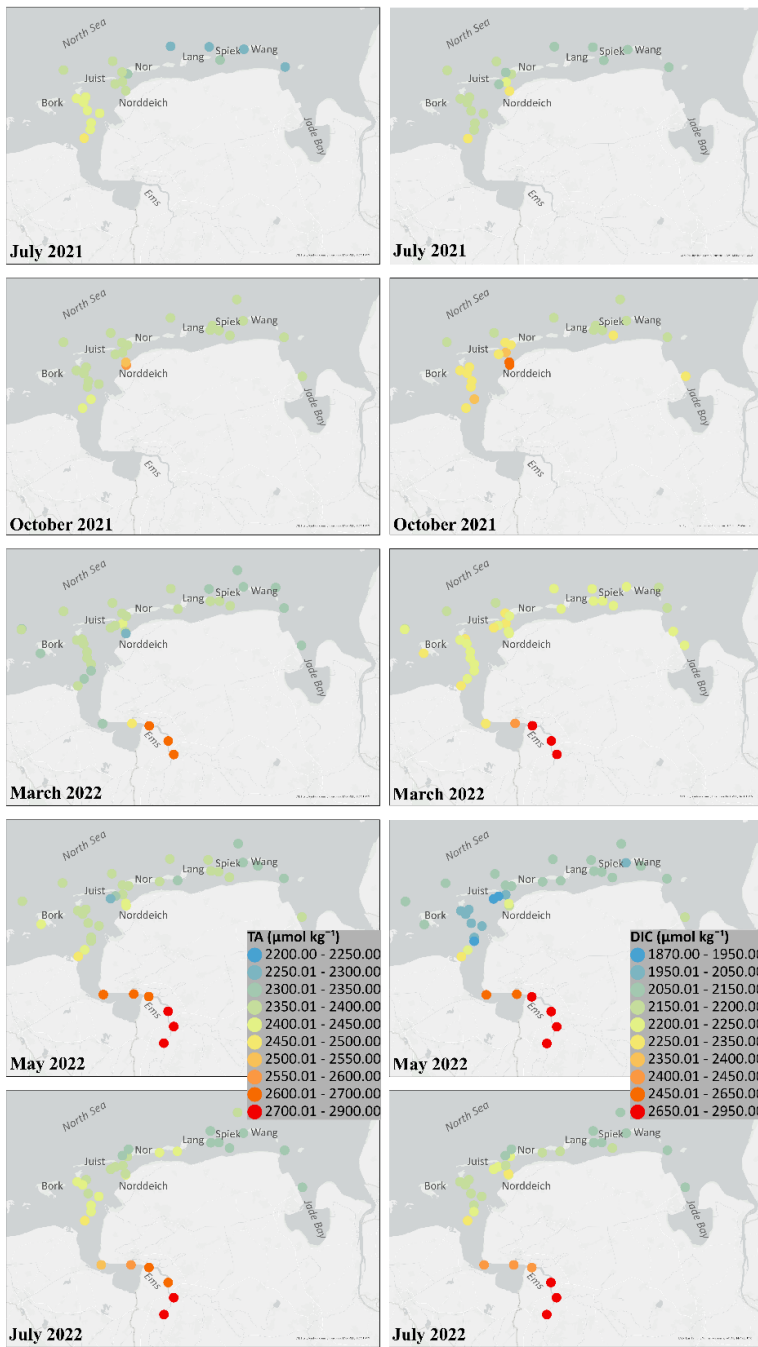


Figure 2: Measured results of TA and DIC of all seasonal cruises of July 2021, October 2021, March 2022, May 2022 and July 2022. All TA and DIC values are in $\mu\text{mol kg}^{-1}$. The map in this Figure was generated using ArcGIS. Data sources: LGLN, Esri, TomTom, Garmin, Foursquare, FAO, METI/NASA, USGS. © 2024 Julia Meyer.

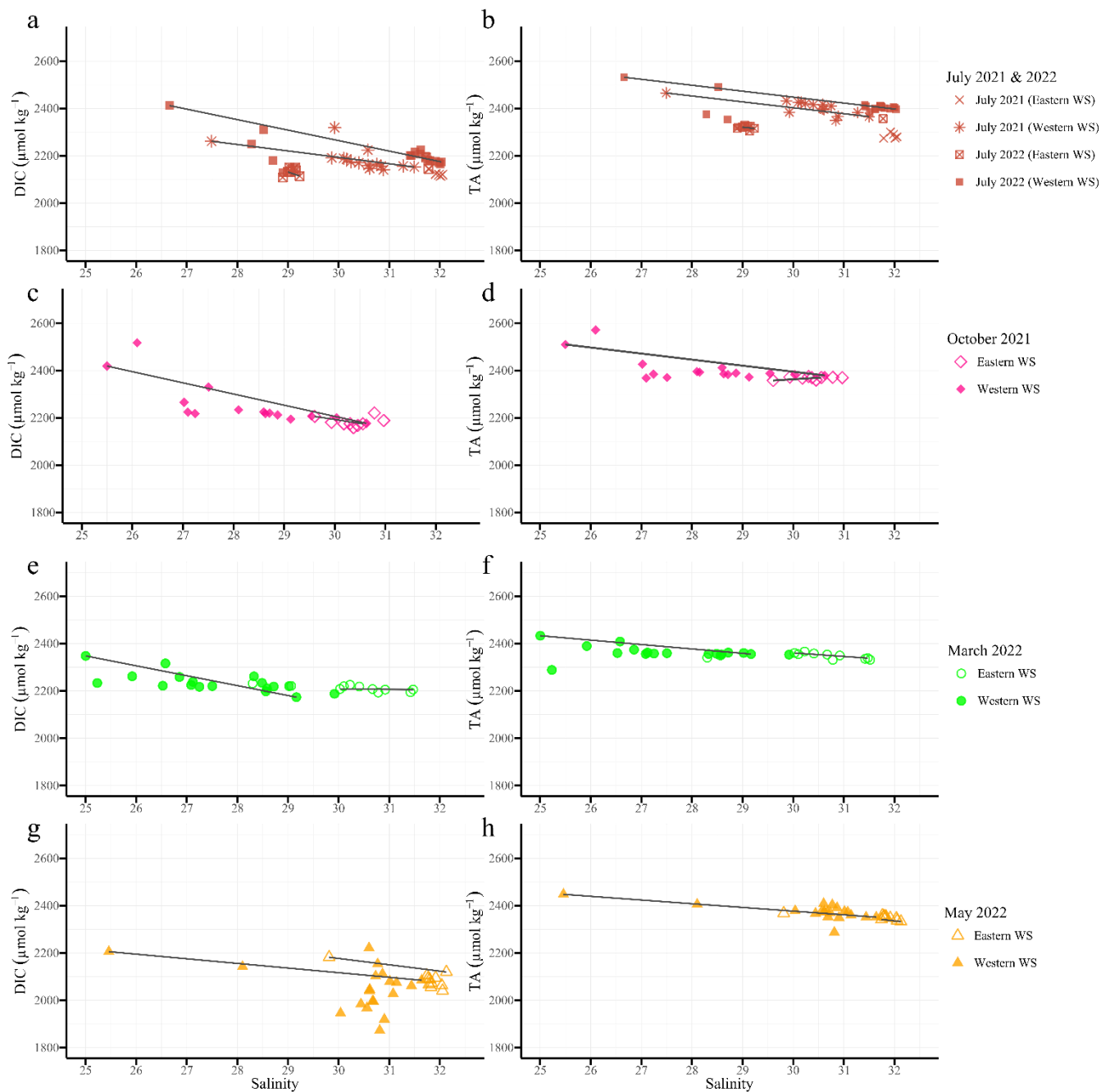


Figure 3: Conservative mixing plots of DIC (left column) and TA (right column) in $\mu\text{mol kg}^{-1}$ versus salinity in the EFWS, separated by cruise and region. Each plot includes data from both Eastern and Western regions (different symbols). Rows correspond to cruises: (a, b) July 2021 and July 2022, (c, d) October 2021, (e, f) March 2022, and (g, h) May 2022. Solid lines represent conservative mixing lines. Only data with salinity >25 are shown, illustrating seasonal and regional variability along the land-to-sea gradient in the EFWS.

3.3 Seasonal Dynamics of DIC and TA Excess and Primary Production Effects

Stations near Norddeich harbour, closer to the mainland, consistently exhibited the highest TA (up to 2571 $\mu\text{mol kg}^{-1}$ in October) and DIC (up to 2516 $\mu\text{mol kg}^{-1}$ in October) concentrations across all seasons (see section 3.2; Fig. 2). These locally high values were therefore excluded from the calculation of mean $\pm$ standard deviation to provide a more representative assessment of $\Delta\text{DIC}_{\text{excess}}$, $\Delta\text{TA}_{\text{excess}}$, and ΔTA_p . However, these values are still displayed in the scatter plots (Fig. 4) to illustrate the range of variability, including highly positive and negative deviations.

Table 4: Mean $\pm$ Standard Deviation of $\Delta\text{DIC}_{\text{excess}}$, $\Delta\text{TA}_{\text{excess}}$, and ΔTA_p ($\mu\text{mol kg}^{-1}$) in the Eastern and Western EFWS of all seasons and regions. One-sample t-tests against zero (representing conservative mixing between North Sea and estuarine waters) were performed. Significant deviations from zero are indicated with symbols (* $p < 0.05$, ** $p < 0.01$, * $p < 0.001$).**

Season	July 2021	October 2021	March 2022	May 2022	July 2022
$\Delta\text{DIC}_{\text{excess}}$					
Western WS:	23.4 $\pm$ 27.6 *	47.7 $\pm$ 84.9	20.5 $\pm$ 23.8 **	-69.3 $\pm$ 107.0 *	54.6 $\pm$ 33.3 ***
Eastern WS:	-19.3 $\pm$ 11.2 *	-6.4 $\pm$ 19.4	5.6 $\pm$ 11.2	-30.1 $\pm$ 35.8 *	-16.0 $\pm$ 26.2
$\Delta\text{TA}_{\text{excess}}$					
Western WS:	123.0 $\pm$ 27.8 ***	29.9 $\pm$ 51.9	16.6 $\pm$ 21.9 **	36.6 $\pm$ 36.1 ***	50.9 $\pm$ 21.9 ***
Eastern WS:	10.9 $\pm$ 11.5	-2.4 $\pm$ 5.9	10.7 $\pm$ 12.5 *	18.8 $\pm$ 13.8 ***	-34.7 $\pm$ 32.3 **
ΔTA_p					
Western WS:	-3.6 $\pm$ 4.4 *	-7.7 $\pm$ 13.6	-3.3 $\pm$ 3.8 **	11.1 $\pm$ 17.2 *	-8.8 $\pm$ 5.4 ***
Eastern WS:	3.1 $\pm$ 1.8 *	1.0 $\pm$ 3.1	-0.85 $\pm$ 1.7	4.8 $\pm$ 5.7 *	2.6 $\pm$ 4.2

Seasonal deviations in $\Delta\text{DIC}_{\text{excess}}$, $\Delta\text{TA}_{\text{excess}}$, and ΔTA_p were evaluated using one-sample t-tests against a null hypothesis of zero for each Season and Region, where zero represents the scenario in which all DIC and TA at a station are solely the result of conservative mixing between North Sea and estuarine waters, with no additional contributions from estuarine processes or biological activity.

The Western WS generally exhibits positive $\Delta\text{DIC}_{\text{excess}}$ values, with a peak in October 2021 (47.7 $\pm$ 84.9 $\mu\text{mol kg}^{-1}$, $p = 0.078$) and summer (July 2021: 23.4 $\pm$ 27.6 $\mu\text{mol kg}^{-1}$; July 2022: 54.6 $\pm$ 33.3 $\mu\text{mol kg}^{-1}$, $p < 0.001$), suggesting an excess of DIC beyond the expected mixing (Fig. 4a). In contrast, during May, $\Delta\text{DIC}_{\text{excess}}$ in the Western WS was negative (-69.3 $\pm$ 107.0 $\mu\text{mol kg}^{-1}$, $p = 0.014$), suggesting significant DIC consumption, due to enhanced primary production during spring bloom conditions. Similarly, negative values were observed in the Eastern WS in May 2022 (-30.1 $\pm$ 35.8 $\mu\text{mol kg}^{-1}$, $p = 0.014$) and July 2021 (-19.3 $\pm$ 11.2 $\mu\text{mol kg}^{-1}$, $p = 0.018$) (Table 4; Fig. 4a).

$\Delta\text{TA}_{\text{excess}}$ also exhibited seasonal and regional patterns: in the Western WS, high positive and significant deviations were observed in July 2021 (123.0 $\pm$ 27.8 $\mu\text{mol kg}^{-1}$, $p < 0.001$) and July 2022 (50.9 $\mu\text{mol kg}^{-1}$, $p = 2.36 \times 10^{-6}$), reflecting additional alkalinity sources. In the Eastern WS, $\Delta\text{TA}_{\text{excess}}$ was negative in July 2022 (-34.7 $\mu\text{mol kg}^{-1}$, $p = 0.005$), whereas July 2021 showed a slight, non-significant positive deviation (10.9 $\mu\text{mol kg}^{-1}$, $p = 0.102$) (Fig. 4b).

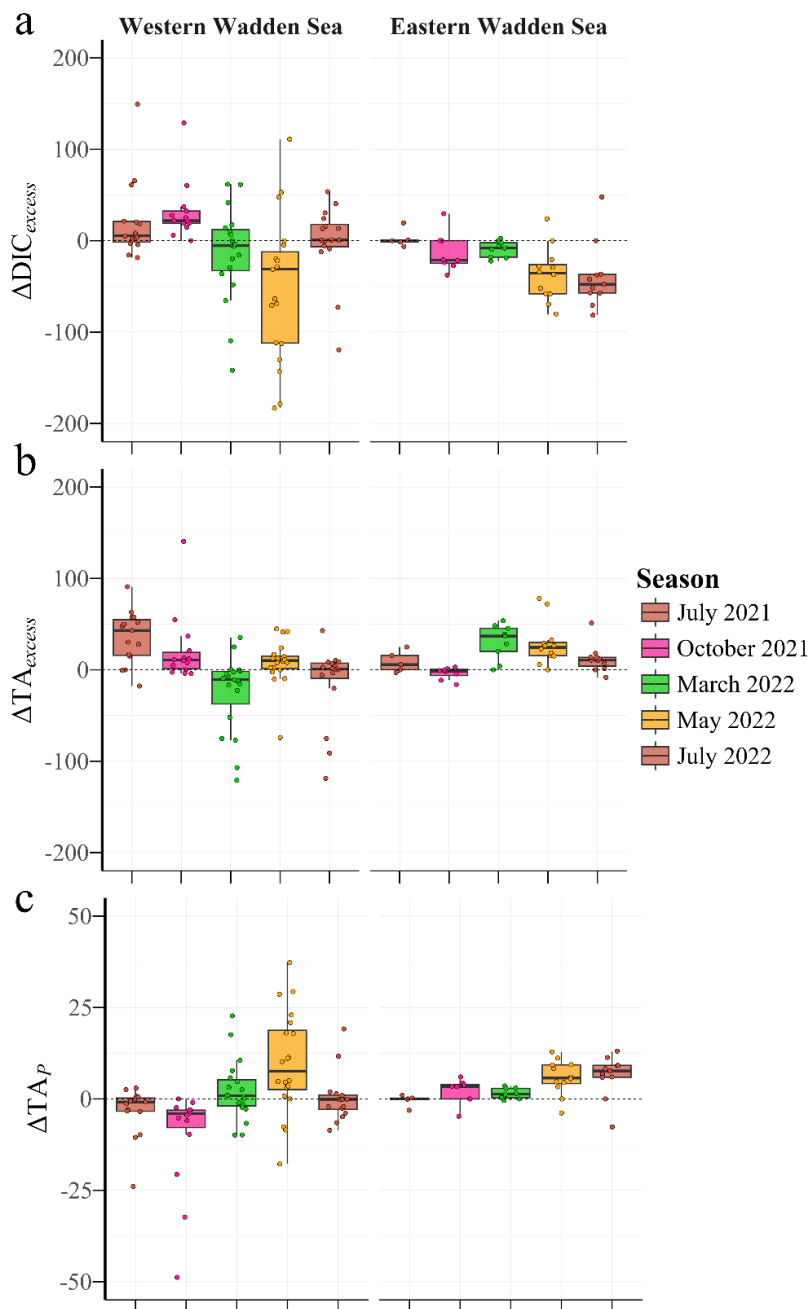


Figure 4: Boxplots of ΔDIC_{excess} , (a) ΔTA_{excess} , (b) and ΔTA_P (c) in $\mu mol\ kg^{-1}$ in the Western WS (left panels) and Eastern WS (right panels) from each season. Each colour represents different seasons. Each box represents the interquartile range (IQR), with the median shown by the thick line.

325 ΔTA_P , representing primary production effects, was positive in May in both the Western ($11.1\ \mu mol\ kg^{-1}$, $p = 0.014$) and Eastern WS ($4.83\ \mu mol\ kg^{-1}$, $p = 0.014$), indicating a significant impact on TA, during times with net uptake of CO_2 through primary production with nitrate uptake, and a corresponding significant decrease in DIC (Fig. 4a, c). In other seasons, ΔTA_P

exhibited contrasting patterns between regions: ΔTA_P in July in the Western WS was negative (July 2021: $-3.8 \mu\text{mol kg}^{-1}$, $p = 0.010$; July 2022: $-8.75 \mu\text{mol kg}^{-1}$, $p < 0.001$), whereas in the Eastern WS ΔTA_P was positive in July 2021 ($3.1 \mu\text{mol kg}^{-1}$, $p = 0.018$), highlighting regional differences (Fig. 4c). To sum up, the Western WS exhibited higher and more variable deviations in $\Delta DIC_{\text{excess}}$ and $\Delta TA_{\text{excess}}$, with significant summer TA production signal. In the Eastern WS, $\Delta DIC_{\text{excess}}$ followed a similar seasonal pattern with smaller magnitudes, while $\Delta TA_{\text{excess}}$ was positive in 2021 but turned significantly negative in July 2022. In the Eastern WS, $\Delta DIC_{\text{excess}}$ followed a similar seasonal pattern with smaller magnitudes, while $\Delta TA_{\text{excess}}$ was positive in 2021 but turned significantly negative in July 2022, which may indicate a stronger influence of sluice openings on TA.

3.4 Influence of Biogeochemical Processes on Seasonal [TA–DIC] and AOU in the EFWS

The relationship between [TA–DIC] and AOU provides valuable insights into the biogeochemical processes that drive carbonate dynamics in coastal systems (Xue & Cai, 2020). A critical threshold for interpreting this relationship is that [TA–DIC] cannot be used when values are $< 50 \mu\text{mol kg}^{-1}$, as proposed by Xue and Cai (2020). This threshold applied only to one station near Norddeich harbor in October (Fig. 5), so these data were excluded from the seasonal calculations (Table 3).

Figure 5a illustrates the AOU relationship to [TA–DIC] from all stations sampled during various seasons in the Wadden Sea. The regression analysis highlights a negative correlation, with a slope of $-1.42 \pm 0.08 \mu\text{mol kg}^{-1}$ per $\mu\text{mol L}^{-1}$ of AOU. When accounting for density-normalized O_2 (i.e., converting AOU from $\mu\text{mol L}^{-1}$ to $\mu\text{mol kg}^{-1}$ using the potential density of seawater), the slope becomes $-1.44 \pm 0.09 \mu\text{mol kg}^{-1}$ per $\mu\text{mol kg}^{-1}$. The grey dashed line in Fig. 5a represents the theoretical Redfield ratio slope ($-123/138 \approx -0.89$), which Xue and Cai (2020) summarize as a reference for expected [TA–DIC] changes per unit AOU in subsurface waters. In their framework, [TA–DIC] serves as a proxy for ocean acidification, directly linking biological production and respiration to changes in DIC and TA. Biological production consumes CO_2 and nutrients, decreasing DIC and increasing TA, while respiration has the opposite effect, increasing DIC and decreasing TA; thus, the decomposition or formation of 1 mol of organic matter ($(CH_2O)_{106}(NH_3)_{16}H_3PO_4$) alters [TA–DIC] by ± 123 mol, which can be approximated from the slope of ΔAOU . Deviations from the Redfield slope, for example in deep waters, may result from processes such as $CaCO_3$ dissolution, which increases both DIC and TA. Our measured slope in the Wadden Sea is steeper than the Redfield reference, indicating a stronger local biological influence on [TA–DIC].

Based on the distribution of our dataset (Fig. 5), we use $\sim 200 \mu\text{mol kg}^{-1}$ [TA–DIC] as an empirical threshold that separates respiration-dominated conditions (lower [TA–DIC], positive AOU) from photosynthesis-dominated conditions with net TA production (higher [TA–DIC], negative AOU). We emphasize that this threshold is not a universal constant, but a pragmatic value derived from our dataset to interpret biogeochemical dynamics in the EFWS.

In spring, negative AOU values (down to $-169 \mu\text{mol L}^{-1}$; Fig. 5a) coincided with [TA–DIC] values exceeding $200 \mu\text{mol kg}^{-1}$. In the summer (2021; 2022), [TA–DIC] values from the Eastern WS are located under the respiration/photosynthesis line ($< 200 \mu\text{mol kg}^{-1}$; Fig. 5b). In contrast, the measured [TA–DIC] concentrations in the Western WS in summer are

> 200 $\mu\text{mol kg}^{-1}$, above the respiration/photosynthesis fit (Fig. 5b). In July 2022, the salinity in the same region changed by 2 - 3 units from one day (July 11) to another (July 12), where the salinity was higher in the East. This short-term change in salinity is relevant to the observed TA and DIC because both parameters are partially controlled by conservative mixing, and slight increases in salinity can lead to proportionally higher TA values. Indeed, TA was slightly higher in summer 2022 ($+53.7 \pm 36.7 \mu\text{mol kg}^{-1}$) compared to the previous year (Table 3).

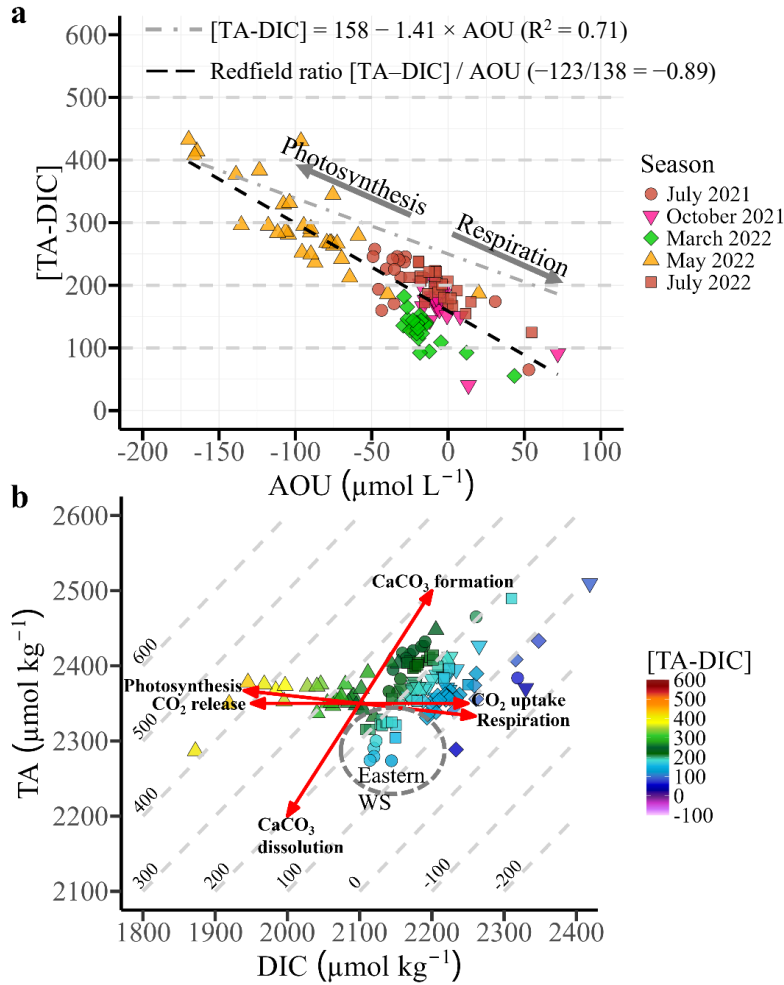


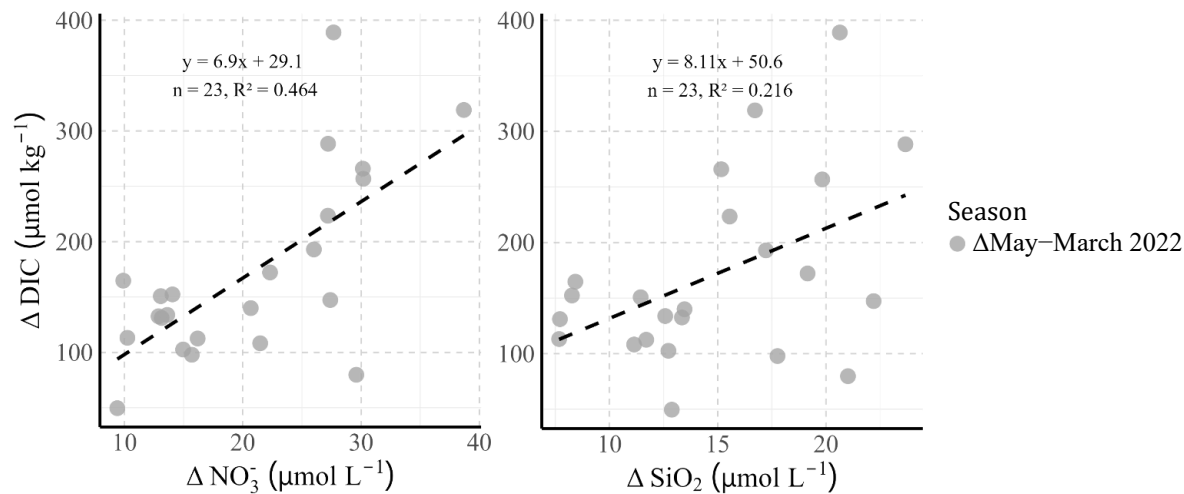
Figure 5: (a) AOU vs. [TA-DIC] (both in $\mu\text{mol kg}^{-1}$) across all seasonal cruises, with regression analysis (black dashed line). Arrows indicate potential processes affecting [TA-DIC] and AOU. The grey dashed line represents the regression slope ($-123 / 138 = -0.89$) proposed by Xue and Cai (2020). (b) DIC versus TA plot from all seasons, with the coloured values of [TA-DIC] for the EFWS at salinity 25 – 32.5. Isoclines represent [TA-DIC] values. Red lines indicate key biogeochemical processes (photosynthesis / respiration, CO₂ exchange, and CaCO₃ formation / dissolution). The Eastern EFWS is highlighted with a grey circle.

3.5 Regional and seasonal variation of Nutrients

A clear seasonal pattern can also be seen in the measured nutrients (Fig. S4). For NO₂⁻ and NH₄⁺ a decrease in concentrations was observed from October 2021 to May 2022, except in the East part of the EFWS (Fig. S4), where the concentration did not

375 change much. NO_3^- concentration ranged seasonally, with an increasing trend from summer to March 2022 (up to $66.28 \mu\text{mol L}^{-1}$) in the EFWS. In May 2022, the NO_3^- concentrations decreased again below the detection limit of the instrument ($> 0.01 \mu\text{mol L}^{-1}$) at some stations, mainly in the Western part (Fig. S4). This analogous seasonal tendency possibly will point to common sources and sinks.

Overall, from March 2022 to May 2022 there was an average decrease in NO_3^- of $19.29 \pm 18.11 \mu\text{mol kg}^{-1}$ (Fig. 6). TA slightly
 380 increased by $9.1 \pm 29.2 \mu\text{mol kg}^{-1}$ during this time (Fig. 4 - 6), while DIC decreased on average by $159.4 \pm 125.4 \mu\text{mol kg}^{-1}$. In figure 6, the changes from March to May ($\Delta \text{May} - \text{March 2022}$) of NO_3^- and SiO_2 at all stations are plotted against the corresponding change in DIC (ΔDIC). For ΔNO_3^- , a regression line was fitted with equation of $y = 6.90x + 29.1$ ($R^2 = 0.464$, $p < 0.05$, Fig. 6), while the ΔSiO_2 regression has an equation of $y = 8.11x + 50.6$ ($R^2 = 0.216$, $p < 0.05$; Fig. 6). The slopes of these regressions represent the calculated C:N (6.90) and C:Si (8.11) ratios. Both values are close to the Redfield
 385 ratios of $106:16 = 6.625$ (C:N) and $106:15 = 7.067$ (C:Si) (Redfield et al., 1963), indicating that enhanced primary production in spring, along with nitrate and silicate uptake, most likely led to the concomitant decrease in DIC (Fig. 6) and the observed increase in TA.



390 **Figure 6: Differences between May and March 2022 for nitrate (ΔNO_3^-) and silicic acid (ΔSiO_2) plotted against the corresponding differences in DIC (ΔDIC) for all measured stations in the Western and Eastern Wadden Sea. Each point represents the change between the two months at a given station ($\Delta \text{May} - \text{March 2022}$). Black dashed lines show linear regressions through the data, and the slopes correspond to the calculated C:N and C:Si ratios at the measured stations, derived from these differences.**

Discussion

4.1 Regional and seasonal Differences of the Carbonate Dynamics in the EFWS

395 This study highlights regional and seasonal variability in the carbonate system, with a notable West-to-East gradient in DIC and TA, showing considerable fluctuations across seasons (Table 3; Fig. 2). This is the first study to investigate $\Delta \text{DIC}_{\text{excess}}$, $\Delta \text{TA}_{\text{excess}}$, ΔTA_p and $[\text{TA-DIC}]$ in the EFWS to provide insights into the respective biogeochemical processes and to determine

the source and sink dynamics of carbon on such a large spatial and temporal scale (Table 4; Fig. 4). We successfully demonstrated that TA increased in the spring because of intense primary production, most likely driven by nitrate uptake (Fig. 4; Fig. 6), highlighting the impact of biological activity on carbonate system dynamics. During the other seasons, the system acted as a source of DIC, indicating seasonal shifts in carbon cycling. Similar patterns of intense production periods and TA increases have also been reported in the North Sea adjacent to the Wadden Sea (Voynova et al., 2019). Additionally, we observed TA production during summer, which was more prominent in the Western WS, further suggesting that these intertidal regions act as a stronger source of TA and DIC during all seasons except in spring, compared to the Eastern EFWS.

4.2 Seasonal TA Production and DIC Dynamics in the EFWS

The data reveal important differences in the patterns between years and regions in summer (Table 4; Fig. 2 – 5). The observed summer dynamics in the Eastern and Western EFWS show distinct patterns in the biogeochemical processes related to TA and DIC production. Positive $\Delta\text{TA}_{\text{excess}}$ (e.g., $123.0 \pm 27.8 \mu\text{mol kg}^{-1}$ in July 2021) and $\Delta\text{DIC}_{\text{excess}}$ (23.4 ± 27.6 ; Table 4) suggest the presence of additional sources of TA and DIC beyond what would be expected from conservative mixing between North Sea and estuarine waters. One key biogeochemical process, which can contribute to increasing TA, without a proportional increase in DIC, is the dissolution of CaCO_3 , which results in elevated [TA-DIC] values ($>200 \mu\text{mol kg}^{-1}$) (Chen & Wang, 1999; Hoppema, 1990). However, CaCO_3 dissolution can be ruled out as a major TA source in the water column, since calcite is supersaturated ($\Omega_{\text{cal}} > 1$) during the study period (Brasse et al., 1999; Norbistrath et al., 2024). Nevertheless, in the sediments large amount of DIC can be produced by carbonate shells and transported by erosion processes into the water column (Brasse et al., 1999). Previous studies have estimated that TA production in the Wadden Sea is driven by anaerobic processes such as denitrification and/or CaCO_3 dissolution in sediments (Norbistrath et al., 2024; Thomas et al., 2009). Anaerobic degradation of organic matter—via denitrification and sulfate reduction—releases both TA and DIC (Brasse et al., 1999; Hu & Cai, 2011; Norbistrath et al., 2024; Thomas et al., 2009), with TA concentrations reaching up to $2310 \mu\text{mol kg}^{-1}$ in summer (Thomas et al., 2009). Other studies suggest that permeable Janssand sediments (a tidal sand flat in the back-barrier area of Spiekeroog Island, Eastern EFWS) exhibit the highest potential denitrification rates (Gao et al., 2010).

In the summer months, the production of TA in both the Eastern and Western EFWS can be attributed to the decomposition of organic matter, especially following a productive spring season, such as the one observed in May 2022 (see section 4.3), which likely led to an elevated input of organic matter into the system (Borges et al., 2017). This organic matter subsequently undergoes both aerobic and anaerobic decomposition in the sediments during the summer months, where anaerobic processes like denitrification and sulphate reduction are drivers in generating both TA and DIC (Brasse et al., 1999; Norbistrath et al., 2024). These processes can significantly contribute to the enhancement of TA production (Fig. Fig. 4b; 5a, b), especially in the Western EFWS, where sedimentary anaerobic processes dominate in the summer (Al-Raei et al., 2009; Böttcher et al., 1998; Hu & Cai, 2011; Kamyshny & Ferdelman, 2010; Norbistrath et al., 2024; Thomas et al., 2009; Wu et al., 2015). Furthermore, with oxygen saturation remaining high ($> 100 \%$ saturation) and apparent oxygen utilization (AOU) being negative (Fig. 5a), it is evident that the water column is net autotrophic. This suggests that TA is primarily produced in the

sediments and subsequently transported into the water column, rather than being generated in situ via remineralization (Beck & Brumsack, 2012; Postma, 1981).

Organic matter remineralization varies regionally across the North Frisian Wadden Sea (NFWS), EFWS, and Jade Bay (Kowalski et al., 2013; Van Beusekom et al., 2012). The NFWS favours aerobic degradation due to its wide tidal basins and
435 high-water exchange ($\sim 8.1 \text{ km}^3 \text{ d}^{-1}$), resulting in lower organic matter accumulation, reduced eutrophication, and lower TA production relative to DIC (Kowalski et al., 2013; Schwichtenberg et al., 2020; Van Beusekom et al., 2012). In contrast, the EFWS experiences higher eutrophication and organic matter accumulation due to its narrower basins and limited exchange, promoting anaerobic degradation processes (e.g., sulfate reduction), which enhance TA production (Kowalski et al., 2013; Schwichtenberg et al., 2020; Thomas et al., 2009; Van Beusekom et al., 2012). Jade Bay, with the lowest water exchange (~ 0.8
440 $\text{km}^3 \text{ d}^{-1}$), exhibits sporadic high TA/DIC ratios, likely due to short-term iron reduction processes (Brasse et al., 1999).

A similar pattern was also seen in the Eastern EFWS compared to the Western in [TA-DIC] dynamics (Fig. 2; Fig. 5b). In comparison to the Western, the Eastern EFWS consistently exhibits lower [TA-DIC] values ($< 200 \text{ } \mu\text{mol kg}^{-1}$) during the summer months of 2021 and 2022, while the Western EFWS shows slightly higher [TA-DIC] values ($> 200 \text{ } \mu\text{mol kg}^{-1}$; Fig. 5b). This difference can be attributed to enhanced CaCO_3 formation, as indicated by its Ω_{cal} (1.9 ± 0.8) in the Eastern CaCO_3
445 formation likely consumes TA (Chen, 1978), leading to the observed shift toward negative $\Delta\text{TA}_{\text{excess}}$ values (Fig. 4b). The formation of CaCO_3 , particularly in sediments or shells, contributes to the lower observed source of TA in the Eastern compared to the Western WS. Despite this, the region still acts as a net source of TA, which aligns with the assumption that the Eastern EFWS can contribute TA to the coastal system during the summer months, especially from anaerobic degradation processes (Schwichtenberg et al., 2020; Thomas et al., 2009).

450 Additionally, the salinity change observed in July 2022 was likely due to local anthropogenic influences, such as the sluice opening in summer 2022. On July 11, 2022, the sluice in Neuharlingersiel was opened before we started our sampling. This anthropogenic intervention likely caused sudden changes in salinity and local hydrological conditions (Fig. S1; Fig. 3a, b). Luebben et al. (2009) investigated the influence of the sluice in the back-barrier area of Spiekeroog and found that the patterns of salinity, CDOM fluorescence, and DOC are strongly affected by the discharge of organic-rich freshwater via the sluice in
455 Neuharlingersiel. Local freshwater inputs from sluices in the EFWS, such as Neuharlingersiel—which can induce a salinity decrease of at least 1 unit (Luebben et al., 2009)—and potentially sluices in the Jade Bay (e.g. Dangaster Siel), as well as coastal circulation, may therefore play a more important role, although the exact contribution of each factor remains uncertain. While, the observed salinity drop extends beyond Norderney into Jade Bay, the low river discharge in summer (Pätsch & Lenhart, 2004; Schwichtenberg et al., 2020) suggests that large riverine inputs (e.g., from the Weser) are unlikely to be the
460 primary cause. Importantly, lower [TA-DIC] values in the Eastern EFWS were observed both on the day of sluice opening in July 2022 and when the sluice was closed in July 2021, suggesting that sluice activity did not substantially control the [TA-DIC] dynamics in July 2022.

The seasonal shifts in TA can influence the coastal ocean's ability to absorb carbon from the atmosphere (Burt et al., 2016; Gruber et al., 2019; Li et al., 2024; Schwichtenberg et al., 2020). In summer, the generation of TA and DIC may alter the

465 region's buffering capacity, with the Western WS possibly storing or taking up carbon (Fig. 2; Fig. 3-5) (Gruber et al., 2019; Li et al., 2024). These findings align with previous studies (Thomas et al., 2009; Voynova et al., 2019) suggesting that the intertidal regions of the EFWS act as a source of both TA and DIC to the coastal system during the summer months.

Particularly, the tidal WS plays an important role in the biogeochemical cycling of the North Sea (Santos et al., 2015; Thomas et al., 2009), because many European rivers are empty in the WS (Thomas et al., 2009). A few studies discussed the generation
470 in of TA in summer before (Schwichtenberg, 2013; Voynova et al., 2019) and the contribution of rivers (Pätsch and Lenhard, 2004), however in summer the riverine inflow was lowest, which could not explain an increase of TA in the WS. The highest riverine contribution of TA is expected from January to April (Pätsch & Lenhart, 2004; Schwichtenberg, 2013; Schwichtenberg et al., 2020).

The moderate rainfall and cooler-than-average weather in October 2021 may have influenced the hydrology of the region,
475 potentially causing increased terrestrial runoff and enhanced delivery of alkalinity-rich water to the coastal system. The slight source of DIC in October (Table 3; Fig. 3c) suggests that organic matter remineralization and sediment-water exchange continue to play a role during this period (Borges et al., 2017). In addition, pore waters enriched with remineralized nutrients are actively released into the overlying water column (Beck & Brumsack, 2012) and organic matter-enriched water masses transported from the North Sea contribute to the availability of degradable material in the Wadden Sea, sustaining
480 biogeochemical activity into autumn (Van Beusekom et al. 1999). This mechanism aligns with studies highlighting the importance of tidally-driven nutrient and carbon fluxes in permeable sediments (Postma, 1981), where advective transport processes facilitate the continuous exchange of dissolved carbon species between sediments and the water column (Santos et al., 2015). Model simulations suggest that effective riverine inputs account for only about 5% of the total TA variability in the German Bight, implying that other processes—particularly tidal pumping—likely play a key role in sustaining elevated TA
485 and DIC concentrations (Moore et al., 2011), as observed in October. In the Dutch Wadden Sea, De Groot et al. (2023) showed that tides control methanotrophic activity, which increases by ~40 % at low tide compared to high tide. By analogy, tidal cycles in the EFWS likely enhance advective transport of dissolved carbon from sediments to the water column, with Ra, TA, and DOC peaking at low tide and potentially driving substantial export to the adjacent North Sea (Moore et al., 2011), reinforcing the role of sediments as an active source of carbon and nutrients, consistent with the observations in October 2021.

490 **4.3 Nitrate Assimilation and Carbonate Dynamics in Spring**

The highest rates of photosynthesis were measured in our study in the spring of 2022 (May 2022). This is evidenced by high O₂ levels (up to 180 % saturation), low $p\text{CO}_2$ obs (Table 2) and high chlorophyll a levels, along with negative AOU values (down to -169 $\mu\text{mol L}^{-1}$; Fig. 5a) (Artioli et al., 2012; Thomas et al., 2005). These findings highlight the dominant role of
495 photosynthesis in modulating the carbonate system, particularly in the Western EFWS, where carbon fixation via photosynthesis is a key factor in this study. The resulting changes in [TA-DIC] during this period are reflected in the AOU to [TA-DIC] relationship shown in Figure 4a, with a steeper negative slope of -1.416 $\mu\text{mol kg}^{-1}$ per $\mu\text{mol L}^{-1}$ AOU, deviating from the Redfield ratio of -0.89 (Xue & Cai, 2020). This deviation suggests that the EFWS may differ from the typical

conditions assumed in the Redfield model (Redfield et al., 1963), which is primarily based on aerobic respiration and production (Xue & Cai, 2020).

500 Furthermore, [TA-DIC] increased due to the significant decrease in DIC in May 2022 (Fig. 3c). Together with the positive ΔTA_P and strongly negative $\Delta\text{DIC}_{\text{excess}}$ values after removal of mixing, this supports the findings of intense spring primary production generating TA, while drawing down DIC (Fig. 4). This was particularly evident in the Western EFWS, but also in the Eastern, indicating the Western EFWS is a stronger sink for CO_2 due to carbon fixation and nitrate assimilation (Borges et al., 2005). The May measurements follow roughly the photosynthesis/respiration line, indicated by the progressively low DIC
505 values in both regions, and slight increase in TA ($9.1 \pm 29.2 \mu\text{mol kg}^{-1}$ from March 2022 to May 2022; Fig. 5b), above the mixing line (Fig. 3h). The relationship in Figure 6 also indicates that primary production is influenced by nitrate and silicate availability. Nitrate concentrations decreased substantially from $65 \mu\text{mol kg}^{-1}$ in March to $22 \mu\text{mol kg}^{-1}$ in May 2022 (Fig. 6; Fig. S4), coinciding with the period of intense primary production. This significant drawdown in nitrate concentrations suggests that nitrate assimilation was the primary driver of the observed decrease in DIC and the slight increase in TA during
510 this period. Nitrate assimilation involves the consumption of hydrogen ions (H^+) and the release of hydroxide ions (OH^-), which leads to an increase in pH and, consequently, in TA (Brenner et al., 2016; Wolf-Gladrow et al., 2007). Therefore, nitrate assimilation played a crucial role in shaping the carbonate system dynamics by decreasing CO_2 and DIC, while increasing TA. The regression analysis of ΔNO_3^- and ΔDIC between March and May 2022 revealed a slope of 6.90 for the changes in nitrate concentrations (Fig. 6), which is close to the Redfield ratio of 6.625 for the C:N ratio (Redfield et al., 1963). This close match
515 suggests nitrate assimilation during the spring bloom, which is closely linked to a reduction in DIC. Similarly, the ΔSiO_2 regression produced a slope of 8.11 (Fig. 6), which is also close to the Redfield ratio for SiO_2 (7.067), further supporting the conclusion that the enhanced primary production in this region during the spring bloom contributed to nutrient uptake, particularly nitrate and silicate.

We therefore propose that assimilation of NO_3^- during the time of intense primary production in May 2022 could explain the
520 local increase of total alkalinity during the spring bloom. First, the maximum concentration of NO_3^- was captured in March before the high biologically productive season started in May 2022 (Fig. 3 – 6

), with maximum concentrations measured in the Western WS. A significant drawdown of NO_3^- from the maximum value of $65 \mu\text{mol kg}^{-1}$ in March 2022 to $22 \mu\text{mol kg}^{-1}$ in May 2022, resulted in an average decrease of $19.2 \pm 9.6 \mu\text{mol kg}^{-1}$ of NO_3^- (Fig. 6; Fig. S4). Brewer and Goldman (1976) also documented that nitrate assimilation increases TA ($9.1 \pm 29.2 \mu\text{mol kg}^{-1}$).
525 An uptake of NH_4^+ was not obvious, because NH_4^+ concentrations were much lower during this period (Fig. S4), suggesting only a small impact on TA patterns. This further emphasizes the central role of nitrate assimilation via primary production in shaping the observed patterns in the carbonate system in the spring-early summer, and the role of land-based nitrate inputs to the coast. Nitrification, the process by which NH_4^+ is converted into NO_2^- , causes a decrease in TA by 2 moles per mole of nitrogen (Wolf-Gladrow et al., 2007; Xue & Cai, 2020), thus has the opposite effect compared to nitrate assimilation, which
530 increases TA. The decrease in NO_3^- from May to July 2022 was less pronounced ($1.51 \pm 5.16 \mu\text{mol kg}^{-1}$, Fig. 6), which is

consistent with the general trend of lower NO_3^- concentrations in summer, likely driven by higher turnover rates (Kieskamp et al., 1991).

5 Conclusion

The findings highlight significant regional and seasonal variations in the EFWS carbonate system, reflecting broader carbon dynamics in coastal and shelf seas. Both TA and DIC exhibit substantial variations across regions and seasons, with a notable decrease in DIC from East to West and an increase in TA during biologically productive periods, such as spring and summer. In spring 2022, a significant drawdown of NO_3^- was observed, correlating with a slight increase in TA, likely due to nitrate assimilation during primary production. Primary production could explain up to 88 % of the $\Delta\text{DIC}_{\text{excess}}$ in the Western WS and up to 92 % in the Eastern WS, contributing to the significant drop in DIC, slight TA production, and NO_3^- drawdown during this period.

In the summer, it is likely that the remineralization of organic matter, combined with the dissolution of CaCO_3 in sediments, contributes to higher TA production, especially in the coastal and nearshore areas of the Western EFWS. On the other hand, our work suggested possible CaCO_3 formation in the Eastern EFWS, which may reduce TA levels ($[\text{TA-DIC}] < 200 \mu\text{mol kg}^{-1}$). However, the region still acts as a net source of TA, in part due to the known high rates of benthic anaerobic respiration, such as organic matter decomposition and associated TA production, particularly after the high productivity of the spring season. This TA generation may enhance the region's capacity to absorb CO_2 , despite the broader southern North Sea generally being considered a carbon source to the atmosphere. These findings emphasize the complexity of the biogeochemical processes driving regional and seasonal variability in the EFWS carbonate system, particularly those influenced by tidal cycles, river inputs, and sediment interactions.

To gain a comprehensive understanding of these intricate interactions and their impact on carbon storage and marine biogeochemistry in this ecologically important region, further research, including sediment studies and continuous tidal monitoring, is essential. This study is the first to combine the analysis of $[\text{TA-DIC}]$ with other parameters such as $\Delta\text{TA}_{\text{excess}}$, $\Delta\text{DIC}_{\text{excess}}$, and ΔTA_p to infer underlying biogeochemical processes - such as biological productivity and nutrient availability. This innovative approach offers a new way to examine how various environmental factors interact and influence the carbonate system. The results have the potential to refine existing models of the biogeochemical cycle, providing valuable insights for more accurate climate predictions and improved strategies for managing coastal systems in response to environmental change.

Code and data availability

The data supporting our findings in this study have been submitted by Lara Luitjens (NLWKN) to the PANGAEA data repository and are currently undergoing final editorial processing. A DOI for the dataset will be available upon completion of

560 the review process (Luitjens et al., in review). The datasets can be accessed through the following links:
https://doi.pangaea.de/10.1594/PANGAEA.974424
https://doi.pangaea.de/10.1594/PANGAEA.974426
https://doi.pangaea.de/10.1594/PANGAEA.974427
https://doi.pangaea.de/10.1594/PANGAEA.974428

565 **Competing interests**

The contact author has declared that none of the authors has any competing interests.

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