

Enhancing Single-Precision with Quasi Double-Precision: Achieving Double-Precision Accuracy in the Model for Prediction Across Scales-Atmosphere (MPAS-A) version 8.2.1

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Abstract. The development of numerical models is significantly constrained by the limitations of high performance computing (HPC). While low precision computations can substantially reduce computational costs, they may introduce round-off errors that can affect computational accuracy under certain conditions. Quasi double-precision algorithm (QDP) can compensate for round-off errors by keeping corrections, thereby achieving the low numerical precision while maintaining result accuracy. This paper applies the algorithm to the Model for Prediction Across Scales-Atmosphere (MPAS-A) and evaluate its performance across two idealized and two real data cases. The results demonstrate that, after reducing numerical precision to single precision (from 64 bits to 32 bits), the application of QDP can achieve accuracy comparable to double-precision computations. Comparing to single-precision computations, bias of surface pressure are reduced respectively by 68%, 75%, 97% and 96% in cases, the memory has been reduced by almost half, while the runtime increases only 6.0%, 0.3%, 2.2%, and 17.8% respectively, significantly reducing computational cost. This research demonstrates that the QDP provides both effectiveness and cost-efficient computational capabilities for numerical models.

1 Introduction

Since the advent of modern computers in the 1950s, numerical simulation-based weather and climate modeling has emerged as one of the most effective methods for exploring weather and climate systems, providing a new platform for numerical model research (Bauer et al., 2015). However, in order to achieve more accurate and precise simulation results, numerical weather and climate models are evolving towards higher resolutions and more complex physical parameterization schemes (Bauer et al., 2015). With the integration of increasingly complex modules to meet diverse requirements, numerical weather and climate models have developed rapidly, and the next generation of these models will feature unprecedented resolution and complexity (Hatfield et al., 2019). In light of these circumstances, the demand for more powerful HPC systems and more efficient computational methods has become particularly urgent. As noted by Bauer et al.(2015), the computational tasks of

future numerical model prediction (NMP) systems are expected to be 100 to 1000 times greater than those of 2015's systems. To bridge the gap between hardware advancements and application performance, the design of code and the selection of algorithms must focus on the optimization of floating-point operations and memory usage (Hatfield et al., 2019).

Mixed precision is a promising research direction in optimizing computational resources within numerical models. By reducing the bit-width required for number representation and thereby lowering the precision of floating-point numbers, mixed precision methods enable storage and computations to be performed with fewer bits. This approach reduces the computational and communication costs in numerical simulations such as climate modeling. Lower precision numerical representations is a feasible approach to reducing computational costs in complex numerical models (Dawson et al., 2017). Low precision computations, defined as operations utilizing fewer than 64 bits of significance, remarkably decrease resource requirements but may introduce round-off errors. To address this challenges, the study of mixed precision techniques has emerged.

In recent years, notable advancements have been made in the application of mixed-precision computing in numerical weather and climate models. Váňa et al. (2016) investigated the implementation of mixed-precision computing in the Integrated Forecast System (IFS) prediction model. They employed double precision in certain regions while utilizing lower precision in others. This approach significantly enhanced computational efficiency by an average of 40% while maintaining acceptable error margins. Dawson et al. (2018) expanded the scope of mixed-precision methods, demonstrating their applicability to simple thermal diffusion models, while key state variables are stored and updated with higher precision. For more complex real-world land surface schemes, they showed that using lower precision for the majority of computations while ensuring high-precision processing of state variables could still meet the requisite accuracy standards. Concurrently, Nakano et al. (2018) conducted an in-depth study on the dynamical core of the global compressible non-hydrostatic model, particularly in the baroclinic wave tests by Jablonowski and Williamson. Nakano et al. (2018) opted to use double precision for grid geometry calculations and single precision for other components. The results indicated that this strategy not only successfully simulated the growth of baroclinic waves with minimal error also reduced runtime by 46%. This study further corroborated the efficacy of mixed-precision computing in dynamical core calculations. Hatfield et al. (2019) applied mixed-precision computing to the Legendre transform in the IFS, successfully implementing half-precision computations. Remarkably, this modification reduced the computational cost to 25% of that in the double-precision reference test, significantly lowering computational overhead. This achievement underscored the substantial potential of mixed-precision computing in large-scale numerical prediction models. Oriol Tintó et al. (2019) applied mixed-precision methods to the Nucleus for European Modelling of the Ocean (NEMO). They discovered that 95.8% of the 962 variables could be computed using single precision. Additionally, in the Regional Ocean Modeling System (ROMS), all 1146 variables could be computed using single precision, with 80.7% of them even using half precision. This finding suggests that mixed-precision methods have extensive applicability in ocean modeling. Cotronei et al. (2020) converted the radiation component of the atmospheric model ECHAM to a single-precision algorithm, resulting in an approximately 40% acceleration in radiation calculations. This result indicates that applying single-precision computing in atmospheric models can significantly enhance computational efficiency while preserving computational accuracy to a reasonable extent. Paxton et al. (2022) further investigated the feasibility of reduced-precision computing. He conducted tests in the Lorenz system, shallow water approximation over a ridge, and the simplified parameterized coarse-resolution spectral global atmospheric

model (SPEEDY). The findings revealed that single precision (23 bits) sufficed for most computational needs, and in numerous cases, half precision (10 bits) could also achieve the desired results. This provides an important reference for adopting lower-precision computing in various models in the future. This year, Hugo et al. (2024) further substantiated the effectiveness of mixed-precision methods in the regional weather and climate model COSMO. The study found that the differences between double-precision and single-precision simulations were minimal, typically detectable only in the initial few hours or days of the simulation. However, single-precision simulations reduced computational costs by approximately 30%. In the same year, Chen et al. (2024) applied the principle of limited iterative development to identify equations that were insensitive to precision in weather and climate modeling tests, modifying them from double precision to single precision. This optimization resulted in a reduction of the runtime of the model's hydrostatic solver, non-hydrostatic solver, and tracer transport solver by 24%, 27%, and 44%, respectively, thereby substantially enhancing computational efficiency. In summary, mixed-precision computing exhibits broad application prospects and potential advantages in numerical weather and climate modeling. By flexibly applying varying precision computing methods while ensuring predictive accuracy, it is feasible to significantly enhance computational efficiency and reduce computational costs.

When utilizing mixed-precision computation, low-precision calculations inevitably introduce rounding errors, particularly when adding numbers with significantly different magnitudes. In such scenarios, the limited precision can cause the larger number to effectively "swallow" the smaller number, thereby compromising the accuracy of the result. For instance, consider the variables $A = 0.7315 \times 10^3$ (a large number) and $B = 0.4506 \times 10^{-5}$ (a small number). If the precision of the result is reduced to 4 significant digits, the outcome will be 0.7315×10^3 , with the large number effectively overshadowing the small one. This phenomenon is especially pertinent in numerical modeling, where the introduction of biases into fundamental fields often necessitates the addition of large and small numbers, inherently causing round-off errors. These errors can accumulate over successive computations, leading to a degradation in model accuracy or even complete failure. Therefore, this issue can not be overlooked.

There have some methods to address the round-off errors. In an early study, Gill (1951) proposed a fourth-order, four-step explicit Runge-Kutta method aimed at correcting round-off errors during computation. This method constructs auxiliary variables at each step to compensate for the rounding errors generated, thereby further refining the results to achieve higher precision. However, this method is not applicable to other forms of numerical solutions. In addition to this, compensated summation methods can enhance the accuracy of summation by utilizing the floating-point precision supported by lower-level hardware (Higham, 1996). These methods rely on recursive summation and incorporate correction terms to reduce rounding errors. Subsequently, Møller (1965) and Kahan (1965) respectively proposed the QDP method and the Kahan method. The primary idea behind both methods is to make slight adjustments to the total sum to avoid the precision loss caused by adding a small, precise value to a much larger one in floating-point addition. The QDP method has been validated in solving ordinary differential equations using the fourth-order Runge-Kutta method (Møller, 1965), where the error after precision reduction is essentially minimized to zero.

Currently, methods for compensating round-off errors are primarily employed in the step-by-step integration of ordinary differential equations (Thompson et al., 1970; Tomonori et al., 1995; Dmitruk et al., 2023). However, their validation in

numerical models remains uncertain. Considering the broader applicability of the QDP method, which can be utilized for recursive summation in any format, and its superior performance in HPC environments compared to the Kahan method (Kahan, 1965), this study aims to implement the QDP method in MPAS-A model. The application of the QDP method to a realistic numerical model, as presented in this study, represents a novel contribution to the field, with no prior research exploring this specific implementation.

Most works involving numerical models that reduce numerical precision adopt a mixed-precision scheme, where some variables use single precision while others remain in double precision to ensure integration stability, as demonstrated in the work of Chen et al. (2024). Currently, there are very few studies that almost entirely employ low precision (32-bit) in numerical models, only applied in IFS by Váňa et al. (2016). However, they only utilize single precision without considering error compensation for it. In this study, all variables in the numerical model were implemented using single precision, and error compensation was applied to key variables. By using error compensation methods (QDP), we can maintain integration stability comparable to that applying double precision scheme while significantly reducing memory requirements by lowering the numerical precision of all variables and improved the accuracy comparable to that applying the single precision. This approach not only reduces communication pressure but also allows for substantial increases in computational speed through vectorization optimization. The structure of this paper is as follows: Section 2 introduces the QDP algorithm, the MPAS model, application of QDP algorithm in MPAS-A, and the experimental design and configuration. Section 3 provides case study in MPAS. Section 4 presents conclusions and discussion of the experiments.

2 Methodology, model and experiments

2.1 Quasi double-precision algorithm

The QDP algorithm, proposed by Møller et al. (1965), aims to address the precision loss that occurs when adding small values to large values in floating point arithmetic. This precision loss typically arises from coarse truncation operations. The QDP algorithm reduces round-off errors by keeping corrections. Primarily applied in the step-by-step integration of ordinary differential equations, the algorithm significantly corrects rounding errors in sum, particularly in computers where truncation operations are not followed by proper rounding.

A brief introduction to the algorithm is as follows, with a detailed derivation available follows Møller et al. (1965). Define the floating-point numbers u , v , s , and c , where in each step of the time integration, $s = u + v$. By introducing a correction variable c before computing sum (s) of u and v in each step, the final s is adjusted to reduce round-off errors. This algorithm is illustrated in Fig. 1

The process can be viewed as v being continuously accumulated onto u ; however, in numerical models' computations, it is impossible to ensure that u is always greater than v . To enhance the precision of the correction process, a precondition of magnitude comparison is added to the algorithm, as shown in Fig. 2.

It is important to note that the applicability of the QDP algorithm has been thoroughly analyzed (Møller et al., 1965), cases of inapplicability are exceedingly rare. Considering the numerous summation operations involved in numerical models, even

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1 :    $u := \text{initial } u;$ 
2 :    $c := 0;$ 
3 : L:  $v := (\langle \text{evaluation of } v \rangle) + c;$ 
4 :    $s := u + v;$ 
5 :    $c := (v - (s - u)) + (u - (s - (s - u)))$ 
6 :    $u := s$ 
7 :   go to L:

```

Figure 1. The QDP algorithm in case of a step-by-step integration. The line 3 to 7 represent the iterative process of applying the QDP method.

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1 :    $c := 0;$ 
2 : L: if  $u > v$  then;
3 :    $u := \text{initial } u;$ 
4 :    $v := (\langle \text{evaluation of } v \rangle) + c;$ 
5 :    $s := u + v;$ 
6 :    $c := (v - (s - u)) + (u - (s - (s - u)))$ 
7 :   else
8 :    $v := \text{initial } v;$ 
9 :    $u := (\langle \text{evaluation of } u \rangle) + c;$ 
10 :   $s := u + v;$ 
11 :   $c := (u - (s - v)) + (v - (s - (s - v)))$ 
12 :  go to L:

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Figure 2. The QDP algorithm adding a precondition of magnitude. This step is designed to ensure that large and small numbers can always be identified during the addition process.

if a few inapplicable instances occur, their impact on the overall result is negligible. Therefore, in practical applications, these
125 infrequent cases are typically not considered.

2.2 MPAS-A

MPAS-A is a compressible, non-hydrostatic atmospheric numerical model developed by NCAR. It employs an unstructured
centroidal Voronoi grid (mesh or tessellation) and a staggered C-grid for state variables as the basis for horizontal discretization
in the fluid flow solver. MPAS-A consists of two main components: the model, which includes atmospheric dynamics and
130 physics, and the initialization component, which generates initial conditions for the atmosphere and land surface, updates for
sea surface temperature and sea ice, and lateral boundary conditions. Both components (model and initialization) are integral
constructs within the MPAS software framework and utilize the same drivers and software infrastructure.

The MPAS-A solves the fully compressible, nonhydrostatic equations of motion (Skamarock et al., 2011). The spatial dis-
cretization uses a horizontal (spherical) centroidal Voronoi mesh with a terrain-following geometric-height vertical coordinate
135 and C-grid staggering for momentum. The temporal discretization uses the explicit time-split Runge–Kutta technique from
Wicker and Skamarock (2002) and Klemp et al. (2007).

The algorithm applied here primarily addresses the rounding error compensation between large and small numbers in addition. Currently, it is only applicable to the time integration process and has not been implemented in the spatial discretization process. Therefore, this section will provide a detailed introduction to the time integration scheme. For the spatial discretization scheme, please refer to Skamarock et al. (2011), and it will not be introduced upon here.

The formulation of the scheme can be considered in one dimension as Eq. (1)

$$\frac{\partial \phi}{\partial t} = \text{RHS}_{\phi} \quad (1)$$

The variable ϕ represents any prognostic variable in the prognostic equations, while RHS represents the right-hand side of the prognostic equations (i.e., the spatial discretization equation). In MPAS-A, a forward-in-time finite difference is used, and it can be written as Eq. (2):

$$\frac{\phi_i^{n+1} - \phi_i^n}{\Delta t} = \text{RHS}_{\phi}. \quad (2)$$

Where superscript represent the time step, and subscript represent the position of grid zone.

The two-order Runge-Kutta time scheme is used in MPAS-A as described in Gear et al. (1971) as Eq. (3), (4) and (5):

$$\phi^* = \phi^t + \frac{\Delta t}{2} \cdot \text{RHS}(\phi^t), \quad (3)$$

$$\phi^{**} = \phi^t + \frac{\Delta t}{2} \cdot \text{RHS}(\phi^*), \quad (4)$$

$$\phi^{t+\Delta t} = \phi^t + \Delta t \cdot \text{RHS}(\phi^{**}). \quad (5)$$

In this study, the version 8.2.1 of MPAS-A was used for the following reasons:

1. This research primarily focuses on the accumulation of variables in time integration, specifically the accumulation of time integration variables within the dynamical core. Version 8.2.1 supports the option to close physical processes during model construction, preventing the influence of physical processes on the results of the dynamical core. Therefore, this version was chosen. It should be noted that all cases in this study have closed physical processes.
2. This version supports single-precision operations, reducing the repetitive work of code modification. It is not the only version that supports single precision, but the latest version currently released.

2.3 Application of QDP algorithm in MPAS-A

According to Eq. (3), (4) and (5), it can be observed that in the time integration scheme, each step involves the process of adding tends on the basic field ϕ^t . In numerical models, the basic field is generally much larger than the tends, which aligns with the principles of numerical computation regarding the addition of large and small numbers, as well as the time integration process. It is important to note that the QDP algorithm currently only addresses time integration and has not been validated during the spatial discretization process. The spatial discretization primarily involves subtraction, specifically the subtraction of a small number from a large number or the subtraction of two close values. Whether this algorithm is applicable in spatial discretization remains uncertain, therefore, we will not apply it in this context.

Based on the application principles of the algorithm, which involve the processes of adding large and small numbers as well as the time integration process, we have established a strategy for applying the QDP algorithm within the MPAS-A. Specific improvements are provided based on Eq. (6), (7), (8) and (9):

$$170 \quad \frac{\partial V_H}{\partial t} = -\frac{\rho_d}{\rho_m} \left[\nabla_\zeta \left(\frac{p}{\zeta z} \right) - \frac{\partial z_{HP}}{\partial \zeta} \right] - \eta \mathbf{k} \times \mathbf{V}_H - \nu_H \nabla_\zeta \cdot \mathbf{V} - \frac{\partial \Omega \nu_H}{\partial \zeta} - \rho_d \nabla_\zeta K - eW \cos \alpha_r - \frac{\nu_H W}{r_e} + \mathbf{F}_{V_H}, \quad (6)$$

$$\frac{\partial W}{\partial t} = -\frac{\rho_d}{\rho_m} \left[\frac{\partial p}{\partial \zeta} + g \tilde{\rho}_m \right] - (\nabla \cdot \mathbf{v} W)_\zeta + \frac{uU + vV}{r_e} + e(U \cos \alpha_r - V \sin \alpha_r) + F_W, \quad (7)$$

$$\frac{\partial \Theta_m}{\partial t} = -(\nabla \cdot \mathbf{V} \theta_m)_\zeta + F_{\Theta_m}, \quad (8)$$

$$\frac{\partial \tilde{\rho}_d}{\partial t} = -(\nabla \cdot \mathbf{V})_\zeta. \quad (9)$$

The meaning of each variable in the equations exactly follows Skamarock et al. (2012), so that we don't repeating explanation. For a numerical model, the most crucial variables are the prognostic variables. Therefore, In the MPAS-A model we applied the QDP algorithm to the time integration process of these prognostic variables, including horizontal momentum (V_H), dry air density ($\tilde{\rho}_d$), potential temperature (Θ_m) and vertical velocity (W), that is, the process in red of Eq. (6), (7), (8) and (9) (Only the predictive equations for the dynamic core are presented here, without the scalar transport). This study focuses on dynamic core, involving the gravity wave and acoustic wave, so we turned off the scalar transport in all cases. In order to be understood well, we provide the pseudo-code in the supplement.

2.4 Experimental design and configuration

This study aims to investigate whether the QDP algorithm can effectively compensate for the round-off errors that caused by reduced numerical precision. Setting the double-precision version (DBL) as the benchmark experiment. Two control experiments are also established: the first control experiment uses the single-precision (SGL), and the second control experiment applies the QDP algorithm to the single-precision (QDP) (Since all the QDP experiments in this paper are set up as single-precision experiments, the experiment names for all QDP applications in single precision will also be referred to as "QDP" in the following sections). By comparing the spatial root-mean-square error (Spatial RMSE) and Mean Absolute Error (MAE) between these two control experiments and the benchmark experiment, this study evaluates the effectiveness of the QDP algorithm in reducing round-off errors.

To assess the application effect of the quasi double-precision algorithm, We selected these two ideal cases and the real-data case because they are the only complete datasets available for download on the MPAS website:

1. **Jablonowski and Williamson baroclinic wave:** A deterministic initial-value test case (Jablonowski and Williamson, 2006) for dry dynamical cores of atmospheric general-circulation models, is presented that assesses the evolution of an idealized baroclinic wave in the northern hemisphere. The primary objective is to assess the model's efficacy in replicating the typical dynamics of moist atmospheric conditions across various precision settings.

2. **Super-cell:** A reduced-radius sphere (Klemp et al. 2015) can be used to assess the behavior of nonhydrostatic processes in global atmospheric dynamical cores, as long as the simulated cases demonstrate good agreement with the corresponding flows in Cartesian geometry, for which analytical solutions are available.

3. **Real data:** with initial conditions generated using GFS data at 2014-09-10_00) using two different resolutions (total domain size of 120 km × 120 km and 240 km × 240 km).

To prevent the influence of other factors, the basic parameters of all cases are kept consistent, including the Number of acoustic steps per full RK step, config dynamics split steps, and config number of sub steps (integer), among others.

3 Results and analysis

In this section, we introduce Spatial RMSE and MAE, and show results across four cases, include two ideal scenarios: Jablonowski and Williamson baroclinic wave and super-cell, as well as two real case (with initial conditions generated using GFS data) using two different resolutions. By using the Spatial RMSE and MAE for quantitative comparison, the differences between the benchmark and control experiments are used to evaluate the effectiveness of the QDP algorithm in reducing round-off error.

3.1 Spatial RMSE and MAE

To quantify the difference between the simulations using SGL, QDP, and DBL, (used as the benchmark), we calculate the Spatial RMSE. First, for each grid point, the temporal averages of the variables (e.g., surface pressure, 500hPa height) are computed across the entire simulation period for each experiment (SGL, QDP, and DBL). Then, the spatial RMSE is calculated as the root-mean-square difference between the temporally averaged fields of the control experiment (SGL or QDP) and the benchmark double-precision experiment (DBL), following Eq. (10):

$$\text{Spatial RMSE} = \sqrt{\frac{1}{N} \sum_{i=1}^N (M_i - C_i)^2}. \quad (10)$$

Where, N is the total number of grid points, M_i is the temporally averaged value at grid point i for the benchmark double-precision experiment, C_i is the temporally averaged value at grid point i for the control experiment (SGL or QDP).

In addition to the spatial RMSE, we also calculate the MAE to assess the magnitude of the difference between the control experiments (SGL and QDP) and the benchmark double-precision experiment (DBL), irrespective of the direction of the difference. Like the spatial RMSE calculation, we first compute the temporal average for each grid point across the entire simulation period for each experiment. The MAE is then calculated as the average absolute difference between the temporally averaged fields of the control experiment and the benchmark experiment, following Eq. (11):

$$\text{MAE} = \frac{1}{N} \sum_{i=1}^N |M_i - C_i|. \quad (11)$$

where N is the total number of grid points, M_i represents the temporally averaged value at grid point i for the benchmark double-precision experiment, and C_i represents the temporally averaged value at grid point i for the control experiment (either SGL or QDP).

RMSE is primarily used to measure the difference between predicted and actual values and is more sensitive to large errors. MAE calculates the average absolute prediction error, and is less sensitive to outliers than RMSE, making it more suitable for conventional error measurements. Therefore, the combination of RMSE and MAE provides a more comprehensive evaluation. When comparing the performance of different experiments, RMSE may be used to quantify differences in extreme values (such as temperature fluctuations, ocean current speeds, etc.), while MAE is used to assess the accuracy of the model's overall trend. Combining both provides a better reflection of the algorithm's performance advantages.

As shown in Table 1 (spatial RMSE) and Table 2 (MAE), the addition of the QDP algorithm consistently improves accuracy (compared to single precision) across all cases. For specific analysis, please refer to the following contents(The results of spatial RMSE and MAE are consistent, so to avoid duplication, only the results of spatial RMSE are analyzed in the following text).

Table 1. The spatial RMSE values of surface pressure compared to DBL for cases, unit: Pa. Note: JW wave = Jablonowski & Williamson baroclinic wave;SC = Super-cell; RD-120/240 = Real data with total domain size of 120/240 km.

Case name	SGL	QDP (Proposed)
JW wave	3.42×10^{-2}	1.09×10^{-2}
SC	8.80×10^{-4}	2.27×10^{-4}
RD-120	6.33×10^{-2}	2.25×10^{-3}
RD-240	6.68×10^{-2}	2.25×10^{-3}

Table 2. The MAE values of surface pressure compared to DBL for cases, unit: Pa. Note: JW wave = Jablonowski & Williamson baroclinic wave;SC = Super-cell; RD-120/240 = Real data with total domain size of 120/240 km.

Case name	SGL	QDP (Proposed)
JW wave	1.29×10^{-2}	3.81×10^{-2}
SC	8.79×10^{-4}	2.26×10^{-4}
RD-120	5.38×10^{-2}	1.95×10^{-3}
RD-240	5.52×10^{-2}	1.94×10^{-3}

3.2 Jablonowski and Williamson baroclinic wave

This case is a deterministic initial-value test case for dry dynamical cores of atmospheric general-circulation models(Jablonowski and Williamson, 2006), assesses the evolution of an idealized baroclinic wave in the northern hemisphere. The initial zonal state is quasi-realistic and entirely defined by analytical expressions, which are steady-state solutions of the adiabatic, invis-

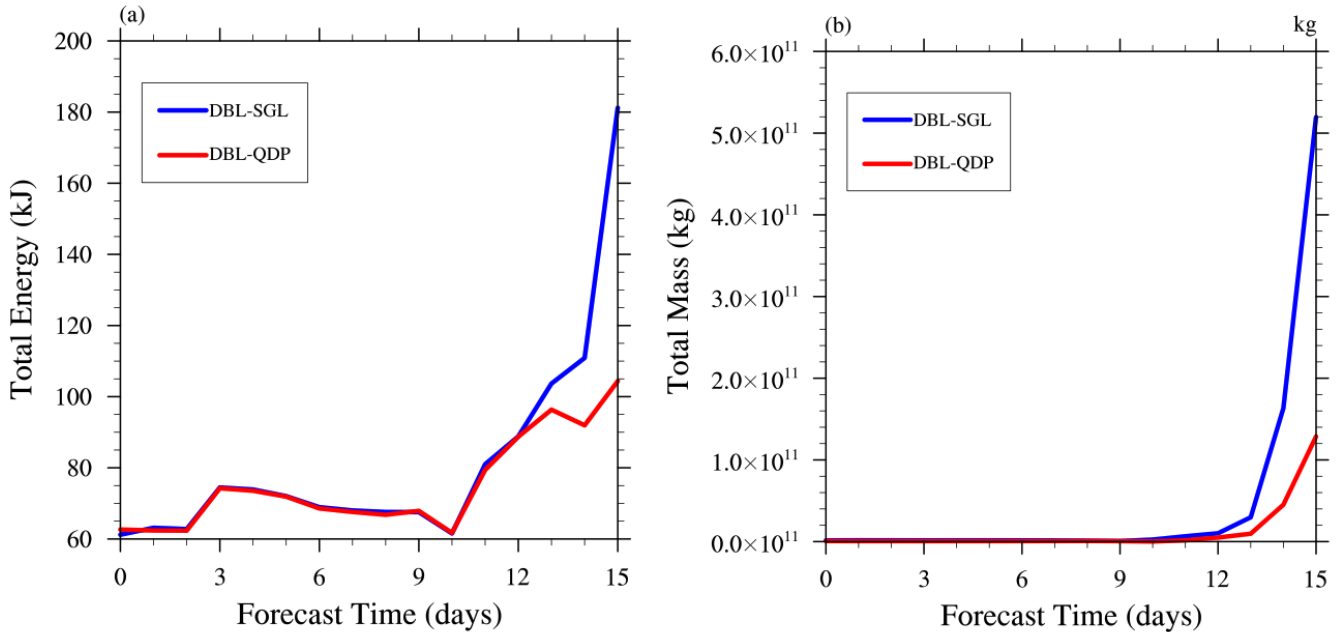


Figure 3. The time evolution of difference between DBL and SGL, as well as difference between DBL and QDP (proposed) of (a) Total Energy, (b) Total Mass in case of Jablonowski and Williamson baroclinic wave. Both figures (a) and (b) demonstrate the advantages of QDP in compensating for errors, with this advantage becoming more apparent after integrating for a period of time.

cid primitive equations in a pressure-based vertical coordinate system (Jablonowski and Williamson, 2006). The experimental configuration is consistent with the test case presented by Jablonowski and Williamson (2006), with a time step of 450 seconds, 26 vertical levels, total domain size of 120 km \times 120 km, and an integration period of 15 days.

The bias begins to appear at the tenth day. Starting from the tenth day, the bias of total energy and total mass caused by SGL can be reduced by using QDP (Figs. 3a and 3b). Unlike SGL, where the bias increases rapidly after more than 10 days, QDP has a very small bias compared to double precision. Therefore, it can be considered that QDP can be used to replace double precision in medium range weather forecast.

It can be found that SGL can increase the round-off error in all regions (Fig. 4a), especially in high-latitude regions, such as Southern Ocean westerly belt, its high wind speed increase error caused by SGL, but instability caused by high wind speeds is more important. Surprisingly, the bias can be reduced significantly in QDP (4b), it means that QDP can improve stability compared to SGL. It should be emphasized that, this does not mean that the higher the wind speed, the better the improvement effect. Instead, the improvement effect is more pronounced in areas with larger errors. The spatial RMSE of surface pressure between DBL and SGL is 3.42×10^{-2} Pa, as well as 1.09×10^{-2} Pa between DBL and QDP, the error reduced by 68%.

The sources of unpredictability, as noted by Bauer et al. (2015), include instabilities that inject chaotic ‘noise’ at small scales and the upscale propagation of their energy. For the cases examined, both SGL and QDP begin to exhibit errors after 10 days of

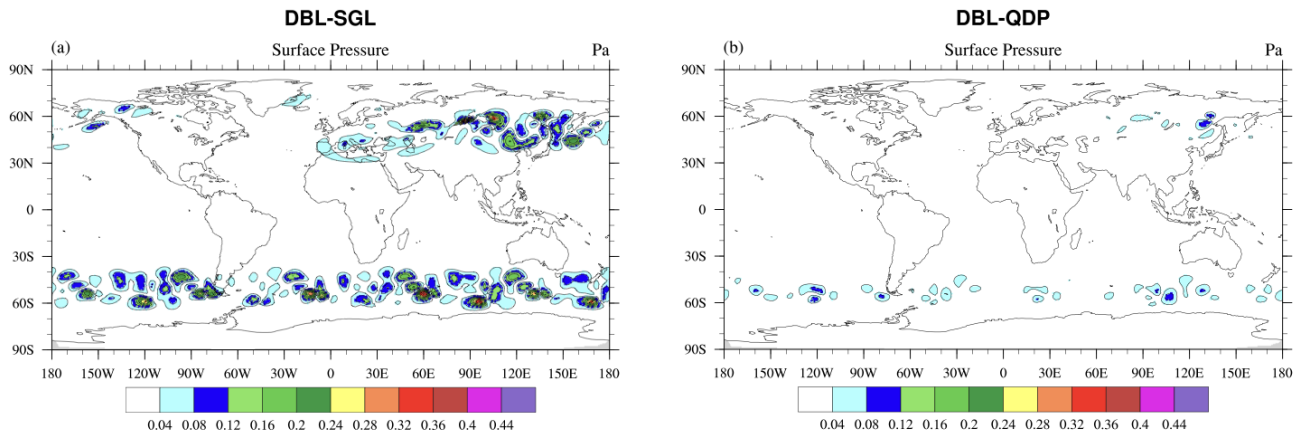


Figure 4. Spatial distributions of averaged (1-15days) difference of surface pressure (units: Pa) between DBL and (a) SGL simulations, (b) QDP simulations(proposed) in case of Jablonowski and Williamson baroclinic wave. The figure shows that the QDP method improves the overall error, with more pronounced effects in the mid- to high-latitude regions.

integration. These errors arise from factors such as rounding errors due to reduced numerical precision and energy loss during the propagation process. The QDP algorithm can reduce the impacts of these errors.

While we acknowledge other potential sources of uncertainty, such as initial condition errors, we have not conducted an in-depth study on them in this research. Our primary focus remains on evaluating the improvements provided by the compensation
 260 algorithm in addressing rounding errors.

3.3 Super-cell case

The test case (Klemp et al., 2015) is on a reduced-radius sphere, can evaluate the behavior of nonhydrostatic processes in nonhydrostatic global atmospheric dynamical cores provided the simulated cases exhibit good agreement with corresponding flows in a Cartesian geometry, and for which there are known solutions. The settings include a time step of 3 seconds, 40
 265 vertical levels, the total domain size is $84 \text{ km} \times 84 \text{ km}$, and an integration period of 2 hours.

In this case, the reduction of Total energy in error is not significant in QDP (Fig. 5a), except for the initial time, all others showed larger errors than SGL. But the errors of both are negligible. For total mass (Fig. 5b), the error caused by SGL can obtain effective improvement in QDP. This improvement exists throughout the entire integration period.

Figure 6 shows the spatial distribution of perturbation theta, an important variable in numerical models, when reducing
 270 the numerical precision from double (Fig. 6a) to single (Fig. 6b), it displays differences, it indicates a significant increase in round-off error. In QDP, this difference can be compensated(Fig. 6c). The spatial RMSE of surface pressure between DBL and SGL is $8.95 \times 10^{-4} \text{ Pa}$, as well as $2.19 \times 10^{-4} \text{ Pa}$ between DBL and QDP, the error reduced by 75%.

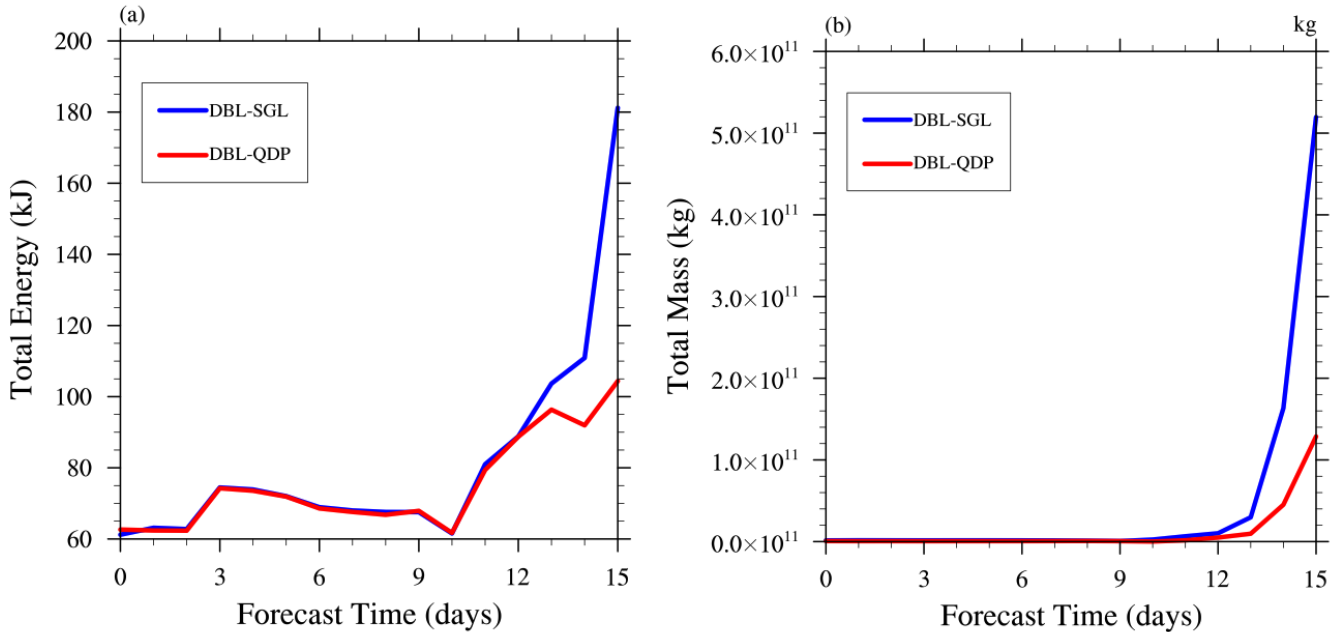


Figure 5. The time evolution of difference between DBL and SGL, as well as difference between DBL and QDP(proposed) of (a) Total energy, (b) Total mass in case of super-cell. Both figures (a) and (b) demonstrate the advantages of QDP in compensating for errors, with this advantage becoming more apparent after integrating for a period of time.

3.4 Real data cases

In this section, we will show the results from two cases using different resolutions. The settings include a time step of 720
 275 seconds, 55 vertical levels, the total domain size are $240 \text{ km} \times 240 \text{ km}$ and $120 \text{ km} \times 120 \text{ km}$, and an integration period of 15 days. (Except for the resolution, all other configurations are exactly the same)

Consistent with the analysis presented in Section 3.2, errors are relatively small in the early stages and begin to emerge after 140 hours. This increase is attributed to the accumulation of round-off errors and energy loss over time. The effects become more pronounced beyond 140 hours. Overall, the QDP algorithm demonstrates a certain level of improvement in addressing
 280 these errors. The case with the total domain of $240 \text{ km} \times 240 \text{ km}$ (Fig. 7a) show the larger error than $120 \text{ km} \times 120 \text{ km}$ (Fig. 7b), and the error can be reduced in QDP caused by SGL.

Figure 8 and 10 show spatial distributions of surface pressure with different resolution. The error has reduced throughout the all region, and the improvement effect is very obvious. From a spatial perspective, the case of SGL with the total domain size of $240 \text{ km} \times 240 \text{ km}$ (Fig. 8a) show the larger error than $120 \text{ km} \times 120 \text{ km}$ (Fig. 10a), and the errors both can be reduced by
 285 QDP (Figs. 8a and 10a). The spatial RMSE of surface pressure with $240 \text{ km} \times 240 \text{ km}$ between DBL and SGL is 6.68×10^{-2} Pa, as well as 2.25×10^{-3} Pa between DBL and QDP, the error reduced by 97%. The spatial RMSE of surface pressuse with

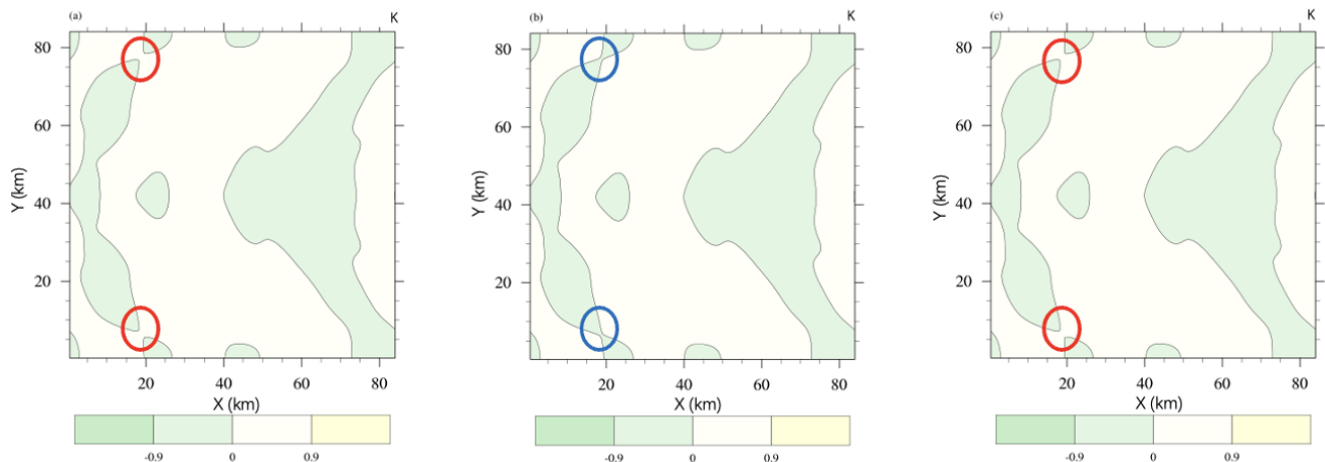


Figure 6. Perturbation theta in super-cell development at 5400s in the (a) DBL simulation, (b) SGL simulation and (c) QDP simulation (bias has reduced)(proposed), unit: K, the circle represents the pattern bias (the same color means the consistent value). In figure (a), after reducing the double-precision experiment to single precision (b), error regions (in blue) appear. The application of the QDP method is able to improve the errors in these regions (c).

120 km \times 120 km between DBL and SGL is 6.33×10^{-2} Pa, as well as 2.25×10^{-3} Pa between DBL and QDP, the error reduced by 96%.

Figure 9 and 11 show spatial distributions of 500 hPa height with different total domain size, 240 km \times 240 km (Fig. 9) and 290 120 km \times 120 km (Fig. 11), The error improvement effect is consistent with surface pressure. The spatial RMSE of 500 hPa height with 240 km \times 240 km between DBL and SGL is 2.80×10^{-1} m, as well as 1.40×10^{-1} m between DBL and QDP, the error reduced by 50%. The spatial RMSE of with 120 km \times 120 km between DBL and SGL is 4.35×10^{-3} Pa, as well as 1.90×10^{-3} Pa between DBL and QDP, the error reduced by 56%.

In this research, we focus on the processes of summing the basic field and trends. When the resolution is increased, the basic 295 field remains relatively unchanged; however, the trends become smaller. This characteristic aligns with the nature of adding large and small numbers, making the advantages of the QDP algorithm more pronounced. Thus, it is evident from Figure 8 that as the resolution increases, the improvement achieved by QDP algorithm also enhances.

On the other hand, it is important to note that the propagation of round-off errors is not immediately apparent over short time scales. However, as the number of iterations increases, these errors can become more significant. The QDP algorithm employs 300 compensation mechanisms that help mitigate the propagation of these errors.

Due to the current limitations on the MPAS-A website, which only provides a single set of terrain and initial condition fields for different experiments, our future plan is to request assistance from the MPAS-A website to construct different terrain and initial condition fields for a specific experiment. We aim to conduct sensitivity analysis, particularly for real data experiments.

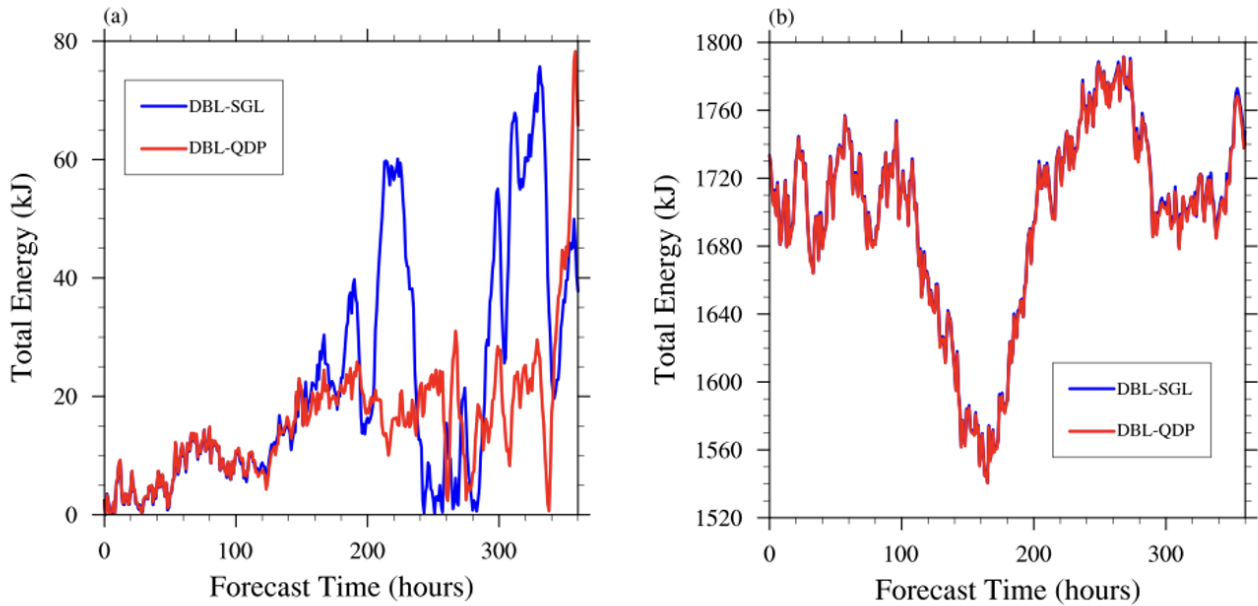


Figure 7. The temporal evolution of spatially averaged difference of total energy between DBL and SGL, as well as difference between DBL and QDP(proposed) in case of real data, with resolution of (a) $240 \text{ km} \times 240 \text{ km}$, (b) $120 \text{ km} \times 120 \text{ km}$. Different resolutions exhibit varying degrees of error improvement. However, overall, the application of QDP (red line) is able to reduce the errors generated by the loss of numerical precision (blue line).

Provided that computational resources allow, we plan to carry out simulations with different resolutions and initial conditions.
 305 This will help lay the data foundation for future uncertainty analysis.

3.5 Computational performance

In comparison with the SGL, although there is a slight increase in runtime, it is minimal, at only 6.0% (Jablonowski and Williamson baroclinic wave), 0.3% (Super-cell), 2.2% (Real data with resolution of 120km) and 17.8% (Real data with resolution of 240km) (Table 3). This slight increase is attributed to the addition of a small number of global variable arrays when
 310 using quasi double-precision. And compared to DBL, QDP demonstrated relatively better performance across different cases, reducing the runtime by 28.6% (Jablonowski and Williamson baroclinic wave), 28.5% (Super-cell), 21.1% (Real data with total domain size of 120km) and 5.7% (Real data with total domain size of 240km) (Table 3).

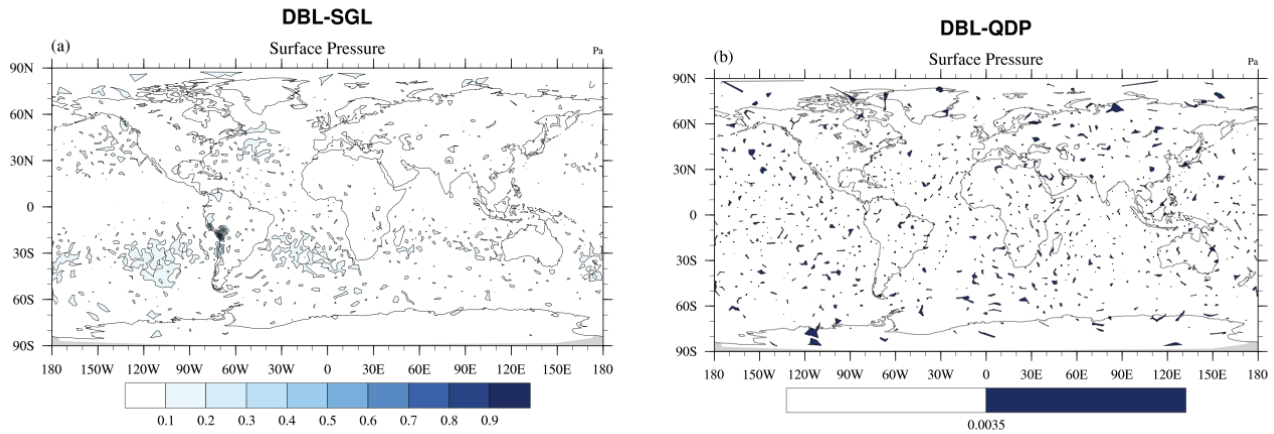


Figure 8. Spatial distributions of averaged (1-15days) difference of surface pressure (units: Pa) between DBL and (a) SGL simulation, (b) QDP simulation (the total domain size: $240 \text{ km} \times 240 \text{ km}$) (proposed). The RMSE of surface pressure between DBL and (a) SGL simulation is $6.68 \times 10^{-2} \text{ Pa}$, (b) QDP simulation is $2.25 \times 10^{-3} \text{ Pa}$. The reduction in error is quite significant, with each region showing a decrease of several orders of magnitude spatially, demonstrating a substantial improvement in error correction.

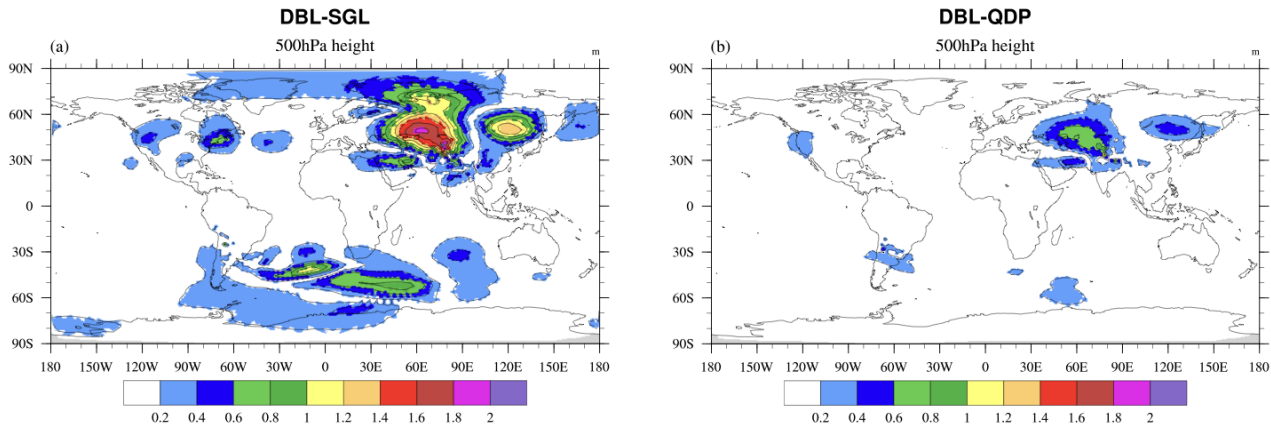


Figure 9. Spatial distributions of averaged (1-15days) difference of 500 hPa height (units: m) between DBL and (a) SGL simulation, (b) QDP simulation (the total domain size: $240 \text{ km} \times 240 \text{ km}$). The RMSE of 500 hPa height between DBL and (a) SGL simulation is $2.80 \times 10^{-1} \text{ m}$, (b) QDP simulation (proposed) is $1.40 \times 10^{-1} \text{ m}$ (round-off error has reduced). From a spatial perspective, the errors in each region have shown notable improvement.

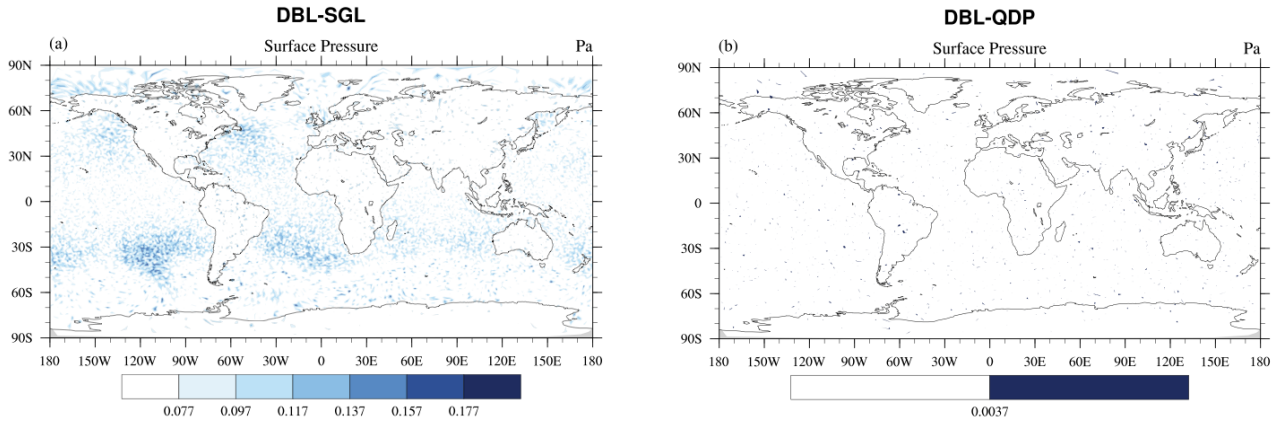


Figure 10. Distributions of averaged (1-15days) difference of surface pressure (units: Pa) between DBL and (a) SGL simulation, (b) QDP simulation (the total domain size: $120 \text{ km} \times 120 \text{ km}$)(proposed). The RMSE of surface pressure between DBL and (a) SGL simulation is $6.33 \times 10^{-2} \text{ Pa}$, (b) QDP simulation(proposed) is $2.25 \times 10^{-3} \text{ Pa}$ (The color bars in (a) and (b) are different). With the addition of QDP, the errors have shown noticeable improvement across all regions from a spatial perspective.

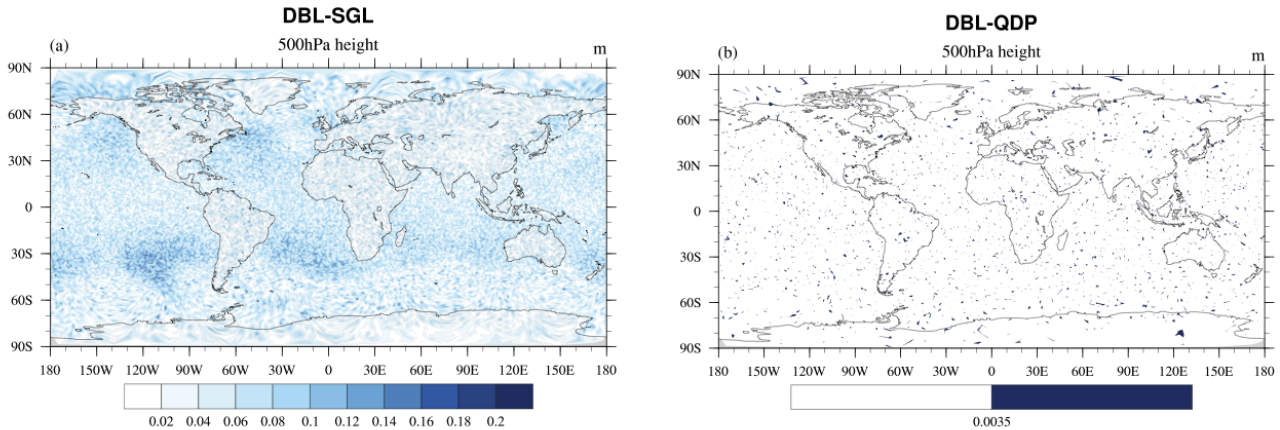


Figure 11. Spatial distributions of averaged (1-15days) difference of 500 hPa height (units: m) between DBL and (a) SGL simulation, (b) QDP simulation (the total domain size: $120 \text{ km} \times 120 \text{ km}$)(proposed). The RMSE of 500 hPa height between DBL and (a) SGL simulation is $4.35 \times 10^{-3} \text{ m}$, (b) QDP simulation(proposed) is $1.90 \times 10^{-3} \text{ m}$, This is consistent with the overall improvement shown in Fig. 10.

4 Conclusions and discussion

Although the QDP method has been extensively employed for the temporal integration of ordinary differential equations, its application in the context of realistic numerical models remains unexplored. This study bridges this gap by introducing a novel

315

Table 3. Comparative Analysis of Computational Efficiency: DBL vs SGL vs **QDP**

Case name	DBL	SGL	QDP (Proposed)		
	Runtime	Runtime	Runtime	vs DBL	vs SGL
JW wave	1768 s	1191 s	1263 s	-28.6%	+6.0%
SC	1507 s	1073 s	1077 s	-28.5%	+0.3%
RD-120	19126 s	14765 s	15092 s	-21.1%	+2.2%
RD-240	1397 s	1118 s	1317 s	-5.7%	+17.8%

Note: JW wave = Jablonowski & Williamson baroclinic wave;
SC = Super-cell; RD-120/240 = Real data with total domain size of 120/240 km

implementation of the QDP method, thereby expanding its scope and potential impact within the field. The algorithm can compensate for round-off errors by keeping corrections in addition of large and small numbers. And in numerical models, the basic field is generally much larger than the tends, which aligns with the principles of QDP, as well as the time integration process. Based on the it, we have established a strategy for applying the QDP algorithm within the MPAS-A. Through the
320 implementation of QDP methods, we maintain accuracy similarly to the DBL and achieve comparable integration stability to the tests comparing to SGL. The error of surface pressure of 4 cases are reduced by 68%, 75%, 97%, 96% (see Section 3). Overall, QDP demonstrates higher accuracy than the SGL, suggesting the potential for applying the QDP algorithm in numerical models.

As spatial discretization process, there involves the subtraction of two close values. Whether this algorithm is applicable in
325 subtraction remains uncertain, so in this study, we don't consider it.

While mixed-precision approaches, where certain variables retain double precision for stability (e.g., Chen et al., 2024), are common for reducing numerical precision in models, and they don't consider the error compensation. This study distinguishes itself by implementing single precision for all model variables and applying error compensation for critical variables.

The QDP algorithm significantly reduces memory usage in MPAS-A while maintaining computational efficiency, making
330 it highly suitable for large-scale numerical simulations. Although increased few local variables and arrays in every time-integration variable, we achieved to reduce the all variables'precisions from double to single. We experimentally show that memory has been reduced by almost half(compared to DBL), with a runtime increase of just 6.0%, 0.3%, 2.2%, and 17.8% (compared to SGL) in the respective cases (see section 3.4). This substantial memory reduction and relatively modest increase in computational cost can be attributed to the inherent trade-off between precision and performance. In large-scale numerical
335 simulations, the impact of rounding errors cannot be ignored, which is why double precision is commonly employed to maintain the accuracy of results. However, double-precision computations significantly increase computation time and resource consumption, which is often impractical for large-scale simulations. By using QDP, we not only reduced memory and communication overhead but also enhanced scalability, particularly in ultra-large parallel simulations where inter-node communication can become a major performance bottleneck. Moreover, while our experiments did not include vectorization techniques, there

340 is considerable potential for further performance improvements. In similar computational environments, vectorization optimization combined with QDP has been shown to offer significant computational efficiency advantages over double precision (Dmitruk et al., 2023). Although this study is limited by hardware and scale, vectorization will be a key area for future research, especially in fully exploiting the potential of HPC architectures.

Nevertheless, there are some limitations to the application of QDP algorithm. Firstly, the algorithm relies on iterative process
345 of time integration, its effectiveness is dependent on the number of time iterations, the more iterations, the better the error compensation. Secondly, although the QDP algorithm partially reduces the round-off errors of low-precision calculations, it still shows error compared to the DBL, making it less suitable for experiments requiring high precision. Additionally, applying QDP algorithm must bring other variables, increasing the complexity to a certain degree.

Looking ahead, we plan to extend the application of the QDP algorithm to additional components of the MPAS-A model.
350 Currently, the QDP algorithm is implemented only within the time integration scheme of the dynamic core, with no consideration for its application to tracer transport. Tracer transport processes involve numerous operations where both large and small values are added together, making them particularly sensitive to precision requirements. In addition, the QDP algorithm has been applied to ideal and real data tests at low and medium resolutions, while its performance at high resolutions has not yet been studied. In future work, we will focus to these fields.

355 *Code and data availability.* Model code and plotting data related to this manuscript is available at: <https://doi.org/10.5281/zenodo.14537107>. Details regarding the code structure and instructions for running the code are provided in the supplementary material, which can be downloaded and viewed in Fig. S1. This figure provides a visual overview of the code organization. The information of steps how to execute the simulations can be found in README file in each test case folder.

Author contributions. JYL and LNW developed the code of applying quasi double-precision algorithm to MPAS-A and design structure of
360 the manuscript. JYL carried out the simulations and analyzed the results with help from LNW, YZY, FW, QZW and HQC. All authors gave comments and contributed to the development of the paper.

Competing interests. The authors declare that they have no conflict of interest.

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