Factors influencing lake surface water temperature variability in West Greenland and the role of the ice-sheet

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Abstract. Subarctic West Greenland is populated by thousands of seasonally ice-free lakes. Using remotely sensed observations, we analyse the surface water temperatures of six lakes during 1995-2022 to identify factors influencing their variability. The connectivity to the Greenland Ice Sheet (GrIS) has a clear influence on lake surface temperature, with ice-sheet marginal lakes experiencing smaller average summer maximum temperature ($< 6°C$) and minimal inter-annual variability. A lake fed by a GrIS-originating river has the fastest seasonal response and largest seasonal amplitude with average maximum temperatures above $13°C$. The seasonal cycle of surface water temperature for all studied lakes is asymmetrical, with faster warming observed after ice off, and a slower cooling of water towards winter freezing. We find that during the study period, the onset of positive stratification has occurred earlier, at rates of up to 0.5 days year^{-1}, and that July-August temperatures have increased at rates up to 0.1°C year^{-1}, although the GrIS-connected lakes show smaller increases. Our analysis suggests that the main meteorological factor determining interannual variability of surface water temperature in the studied lakes is air temperature. This study highlights the important role of remote sensing for long-term monitoring of Greenlandic lakes under climate change.

1 Introduction

The subarctic expanse of West Greenland is home to a multitude of lakes (How et al., 2021), their seasonal ice-free periods providing a window into the region's dynamic environmental processes (Anderson et al., 2017; Fowler et al., 2020). Influenced by climatic and glacial factors, these water bodies offer invaluable insights into the intricate interplay of climatic forces shaping high-latitude ecosystems (Saros et al., 2019; Prater et al., 2022; Saros et al., 2023). Among the myriad consequences of climate change on these lakes, one of the most pressing concerns is the direct and indirect impacts of rising lake surface water temperature (LSWT). This phenomenon not only affects various ecological states, structures, and processes (Williamson et al., 2009; Woolway et al., 2022) but also holds implications for ice-sheet mass balance (Zhang and Bolch, 2023). Specifically, the interaction between lakes and ice margin involving lake-ice processes such as calving (Mallalieu et al., 2021; Minowa et al., 2023), submarine melting (Sugiyama et al., 2021), lake temperature (Dye et al., 2021), and glacial lake outburst floods (GLOFs) (Kjeldsen et al., 2017; Grinsted et al., 2017) can accelerate glacier mass loss, particularly in areas where lakes abut ice fronts. Studies suggest that such lakes contribute to enhanced thermo-mechanical erosion and reduce basal resistance, resulting

in heightened ice flux towards the margins (King et al., 2019; Maurer et al., 2019; King et al., 2020; Benn et al., 2007; Tsutaki et al., 2011; Pronk et al., 2021). Given the ongoing trajectory of global climate change, understanding the factors driving LSWT variability in Greenland is paramount for understanding broader environmental shifts. Traditionally, global studies investigating variations in LSWT have relied on in situ measurements, yet such data are notably sparse in West Greenland (Kettle et al., 2004). Few studies have been found in literature reporting temperature measurements on lakes in Greenland which have been collected as part of three main efforts:

1. Measurements taken between 2011 and 2022 on small similar lakes (from a minimum of 13 lakes and a maximum of 22 of surface area between 0.022 and 0.824 km^2) in southwest Greenland around Kangerlussuaq have been reported in studies on thermal stratification (Saros et al., 2016), under ice phytoplankton dynamics (Hazuková et al., 2021) and on the relation between ice melt timing and dissolved oxygen concentration (Hazuková et al., 2024). None of these lakes are connected to the GrIS.

2. Water temperature measurements were collected from approximately 30 lakes along a east-west transect from the ice-sheet near Kangerlussuaq to the outer coast south of Sisimiut during a campaign conducted between 1998 and 2000. The studies reporting the dataset focussed on empirical modelling of summer lake water temperature in Greenland lakes (Kettle et al., 2004), and on limnology and paleolimnology of lakes in Greenland (Anderson et al., 1999; Brodersen and Anderson, 2000, 2002).

3. The Greenland Ecosystem Monitoring BioBasis programme (GEM, 2025) recorded water temperature from 2009 to 2019 on two small lakes in Kobbefjord, Badesø / Kangerluarsunnguup Tasia and Qassi Sø, at 2m and 10m respectively.

However, monitoring especially over decadal time scales is limited and happens in areas close to settlements. Access to remote regions (i.e. far from a well-connected settlement) remains challenging. Consequently, most lakes in West Greenland remain largely unexplored.

The absence of widespread, long-term monitoring has left the thermal dynamics of many lakes poorly quantified, impeding accurate estimations of their role in regional energy and mass exchange with the atmosphere and interactions with the GrIS margins, particularly in the context of climate change impacts (Anderson et al., 2001). Moreover, in situ measurements are point measurements and generally spatially resolved LSWTs can better give a more representative picture of the temperature change (Woolway and Merchant, 2018).

Recent advancements in remote sensing have transformed our ability to observe environmental variables across vast spatial scales and extended time frames. Leveraging these capabilities, our study investigates the thermal behaviours of six representative lakes (Figure 1) across West Greenland using satellite-derived time series data of LSWT and Lake Ice Cover (LIC) from the European Space Agency (ESA) Climate Change Initiative (CCI) Lakes dataset specifically created for climate studies according to the requirements of the Global Climate Observing System (GCOS) of the World Meteorological Organisation (WMO) (Buontempo et al., 2022).

Spanning from 1995 to 2022, our study explores the intricate dynamics of these aquatic ecosystems in response to climatic and glacial influences using data that are consistent in time and space. These six lakes were carefully chosen to encapsulate

the diverse environmental conditions prevalent in West Greenland. Notably, two lakes are directly linked to the GrIS, serving as sentinel sites for evaluating the impacts of glacial processes on lake thermal regimes (see Methods). Additionally, one lake, indirectly connected via a river fed by the GrIS, provides valuable insights into the downstream ramifications of glacial meltwater discharge on aquatic systems.

Our investigation extends beyond simple characterization of seasonal temperature fluctuations by exploring the complex interrelationships between climatic variables and lake dynamics. By analyzing LSWT and LIC data alongside concurrent meteorological parameters such as air temperature and shortwave radiation, we investigate the underlying mechanisms governing thermal variability in these lakes in a consistent way. We have included in our study lakes that are connected to the GrIS, as well as those that are not, in order to gain insights into the influence of the ice-sheet on lake thermal properties. Central to our investigation is the detection of climatic signals imprinted within the temporal evolution of LSWT, particularly against the backdrop of observed global warming trends. By examining the inter-annual variability and trends in surface temperature, we aim to investigate the extent to which climatic warming influences the thermal regimes of these high latitude lakes. Furthermore, by scrutinizing the phenology of warm-stratified conditions, we aim to discern alterations in the seasonal dynamics of these aquatic environments.

In summary, this work aims to provide a substantial initial contribution to understanding lake thermal dynamics in this critical region, with the potential to enhance predictive modelling and support the strategic selection of monitoring locations. Lakes are important sentinels of climate change, as their thermal regimes respond rapidly to atmospheric forcing and reflect broader environmental shifts—particularly in this region. However, consistent in situ observations in this area remain limited. Remote sensing offers a powerful tool for capturing spatio-temporal patterns in surface water temperature at otherwise unattainable scales. By addressing this gap and leveraging satellite-based observations, this study also intends to demonstrate the value of remote sensing in supporting climate-related assessments of freshwater systems. Our findings contribute to ongoing efforts to develop robust, observation-driven frameworks for monitoring lake responses to climatic shifts and offer valuable insights for enhancing lake modelling in the region.

2 Data and Methods

2.1 Study region

The west of Greenland has its most substantial ice-sheet free margin between 63.5°N and 68.5°N. The seasonally ice-and-snow-free parts of this region generally have an elevation between 0 and 1100 m above sea level (m a.s.l.). The margin of the GrIS is mostly found at elevations ranging from ∼600 m to ∼1200 m, and at distances up to 180 km from the west coast (excluding fjords). The intervening area is characterised by thousands of inland water bodies exhibiting a range of limnological conditions (Ryves et al., 2002). Specifically, lakes in this region vary in area, depth, elevation, components of water balance (groundwater, precipitation, surface runoff and stream feeding), surface connectivity to other lakes via streams (which may be seasonal), chemistry, connectivity to the ice-sheet margin and water clarity. The latter two aspects are linked, as glacial meltwater contains fine particulates in suspension, particularly minerogenic glacial flour (Burpee et al., 2018). The particulates cause high turbidity,

with associated ecosystem impacts through reduction in photosynthetically active radiation (Sommaruga, 2015). Turbid lakes may respond not only directly to climatic changes in weather-driven forcing, but also indirectly via modification of albedo associated with changing glacial fluxes, since water clarity directly influences lake heat budgets and stratification (Saros et al., 2019).

95   The climate of the studied region also varies considerably, principally from the coast to the area next to the GrIS (Cappelen and Jensen, 2021), with consequent changes in vegetation and lake properties. Areas close to the GrIS have a Low Arctic continental climate, with mean annual surface air temperature of -5°C, an annual temperature range of $\sim$30°C, continuous permafrost, and low precipitation $<$ 170 mm year^{-1}. The zone from the GrIS to mid-way towards the coast is characterized by very low (negative) effective precipitation. The coast is characterised by a Low Arctic maritime climate, with a lower annual

100 temperature range ($\sim$22°C) and higher precipitation ($\sim$500 mm year^{-1}). The moderated summer maximum temperatures and the presence of coastal fog banks associated with the more maritime climate imply a longer presence of snow and snow packs in this region (Cappelen and Jensen, 2021).

    In this study, the thermal behaviour of six relatively large lakes from 1995 to 2022 is investigated using satellite-derived time series of LSWT and LIC from ESA CCI (Hollman et al., 2013) project dedicated to lakes (Carrea et al., 2023) - see below.

105 The six lakes have been selected among the sixteen lakes included in the ESA CCI Lakes portfolio based on their location and characteristics. The six lakes have diverse properties that are representative of those distributed throughout West Greenland. Two lakes are directly connected to the GrIS (they are ice marginal lakes), and one lake is indirectly connected by a river of length $\sim$92 km fed by the GrIS. These three lakes are at elevations of 300 m or higher. The other three lakes are unconnected to the GrIS and at elevations less than 50 m, with one very close to the coast (Figure 1).

110 **2.2   Digital elevation model**

To explore the topography of the study region, and that near the vicinity of the studied lakes, we used the Terra Advanced Spaceborne Thermal Emission and Reflection Radiometer (ASTER) Global Digital Elevation Model (GDEM) Version 3 (ASTGTM) (NASA/METI/AIST/Japan Spacesystems and U.S./Japan ASTER Science Team, 2019; Abrams et al., 2020), jointly created and developed by National Aeronautics and Space Administration (NASA) and Japan's Ministry of Economy, Trade, and In-

115 dustry (METI). It provides a global digital elevation model (DEM) of land surface on Earth that extends from 83°N to 83°S. The data are projected onto the 1984 World Geodetic System (WGS84)/1996 Earth Gravitational Model (EGM96) geoid and the spatial resolution is 1 arc second (about 30 metres at the equator) and they have been resampled at 50 metres resolution to be used with Landsat 8 data for this work. The ASTER GDEM version 3 dataset can be downloaded at the Land Processes Distributed Active Archive Center (LP DAAC) (https://doi.org/10.5067/ASTER/ASTGTM.003).

120 **2.3   Lake bathymetry**

Bathymetry data (maximum depth, mean depth, area and volume) for each of the studied lakes were extracted from the GLO-Bathy dataset (Khazaei et al., 2022) to characterise the lakes in the region. GLOBathy provides estimates of bathymetry for lakes worldwide with reasonable accuracy, given the complexity of estimating underwater topography, as reported in the ar-
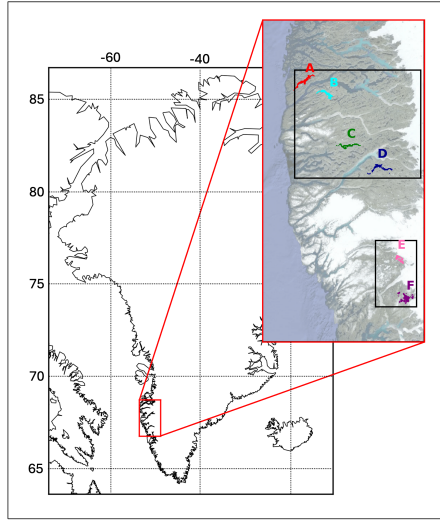
Figure 1. (a) Map of Greenland showing the general location of the main western ice-free margin. (b) Map of the western ice-free margin, showing the north and south study domains (boxes) and (colour-coded) locations of six major lakes for which temperature time-series are available. The lake names and characteristsics are listed in Table 2. The lake mask shown in the plot is from the ESA CCI Lakes dataset (Carrea et al., 2022b) and it is used for this work.

ticle. The bathymetry is derived with a model which has been selected among few candidates. The selection is based on the

125   comparison of the predicted maximum depth with the observed value for about 1500 lakes giving the root mean squared error normalized with standard deviation, NRMSE = 0.17, and the Spearman's Rho correlation coefficient, $\rho = 0.94$ for the selected model. Also, a cross validation has been carried out. The actual validation of the bathymetry reported on the paper consists of a visual comparison of the predicted bathymetry with the observed bathymetry for eight lakes. Regarding the six lakes of this paper, we have found that the maximum depth of Eqalussuit Tasiat (Lake A) is unrealistic (see Section 3.2). The dataset is

130   available at *figshare* (https://doi.org/10.6084/m9.figshare.c.5243309.v1).

2.4   Landsat 8

To characterise the studied region and the spatial variability of the lakes in the region, we analysed data from the Thermal Infrared Sensor (TIRS) and the Operational Land Imager (OLI) instruments on Landsat 8 (www.usgs.gov). The two sensors provide a spatial resolution of 30 meters for the visible, near-infrared (NIR), short-wavelength infrared (SWIR) and of 100

135   meters for the thermal channels. The repeat coverage is about 16 days and the scenes are provided by the Worldwide Reference System-2 (WRS-2) path/row system, the size of which is 185 km x 180 km, and with a swath overlap percentage increasing from the equator to extreme latitudes (Vanhellemont and Ruddick, 2015). The Landsat 8 data are available at the Landsat Data Access web page (https://www.usgs.gov/landsat-missions/landsat-data-access) within the United States Geolog-

ical Survey (USGS) data portal. The Landsat no-cost open access data policy of the USGS remains intact since its inception in 2008.

2.5 ESA CCI Lakes

For most of this work, we utilized LSWT and LIC from the ESA CCI Lakes dataset v2.1 (Carrea et al., 2024) and v2.0.2 (Carrea et al., 2022a) respectively (Carrea et al., 2023). LSWT observations are derived from the following visible and infrared radiometers on polar orbiting satellites: ATSR-2 (1995–2002), AATSR (2002–2012), MODIS on Terra (2000-2022), AVHRR on MetOpA (2007-2019), AVHRR on MetOpB (2012-2019), SLSTR on Sentinel3A (2016-2022), and SLSTR on Sentinel3B (2018-2022). The LIC observations are derived from MODIS on Terra and Aqua (2000-2020). The LSWT and LIC have been retrieved at a spatial resolution of ∼1 km at nadir and then gridded to a common grid 1/120° cells on a global lake mask at 1/120° resolution (Carrea et al., 2023). The lake mask is available at *zenodo* (Carrea et al., 2022b) and it was derived from the GloboLakes mask (Carrea et al., 2015). For LSWT, the optimal estimation retrieval method of MacCallum and Merchant (2012) was applied on image pixels identified as water according to both the lake mask and a reflectance-based water detection scheme (Carrea et al., 2023), which was specifically designed to distinguish water from non-water pixels with clouds, ice or land-contaminated. LSWT are accompanied by per-pixel quality levels which summarise the confidence in the retrieval and by the uncertainty. Only LSWT of quality levels 4 and 5 (the highest) have been used for this study. Given that the best resolution of the instruments used for the retrieval of LSWT is 1 km, obtaining reliable information for each of the six lakes is challenging as their maximum distance to land is between 1.2 km and 1.8 km. In these cases, the LSWT retrieval is very likely to be available only for that part of the lake and only for a limited proportion of almost daily overpasses (clear sky and observations relatively central within the swath). The LIC is retrieved using reflectance channels, therefore the illumination at a local time of ∼11 am becomes marginal or absent from November until the middle of January. The LSWT and LIC of the ESA CCI Lakes dataset have been validated through comparison with in situ data and through manual inspection respectively (Carrea et al., 2023). Because the LSWT retrieval algorithm is based on physics, a stable performance is expected across domains in times and spaces where LSWT cannot be directly easily validated through comparison with in situ data, such as for lakes in Greenland. The ESA CCI Lakes dataset has been created as a climate data record following the requirements of GCOS of the WMO (Buontempo et al., 2022). This is the first consistent long term LSWT and LIC time-series from satellite for lakes in Greenland. The ESA CCI Lakes v2.0.2 and v2.1 dataset are available for download at the Centre for Environmental Data Analysis (CEDA - United Kingdom) archive (https://dx.doi.org/10.5285/7fc9df8070d34cacab8092e45ef276f1).

2.6 ESA CCI Clouds

The bottom of atmosphere downwelling shortwave radiation from the ESA CCI Clouds dataset (Stengel et al., 2020) has been used in this study for establishing if any correlation between LSWT and shortwave solar radiation is present. The dataset is available at https://doi.org/10.5676/DWD/ESA_Cloud_cci/AVHRR-AM/V003.

2.7 ESA CCI Sea Surface Temperature

The sea surface temperature dataset used in this paper to determine the marine influence on the lake close to the coast. We use the ESA CCI Sea Surface Temperature (SST) dataset (Embury et al., 2024) and it can be downloaded at the CEDA archive (http://dx.doi.org/10.5285/62c0f97b1eac4e0197a674870afe1ee6). The dataset utilised for this work is monthly L4 SST (analysis data) from 1995 to 2021. The monthly aggregated dataset can be downloaded at the SST regridding service of the University of Reading (https://surftemp.net/regridding/index.html).

2.8 ECWMF ERA5-Land data

Data for two atmospheric variables, 2m air temperature (Tair) and solar radiation (surface solar radiation downwards), used in this study are from the ERA5-Land reanalysis dataset of the European Centre for Medium-Range Weather Forecasts (ECMWF) (Hersbach et al., 2020; Muñoz-Sabater, 2019). Data at three hourly intervals have been used to create a climatology and monthly values. The values of Tair have been corrected because of the difference in the nominal altitude of the ERA5-Land dataset and the high-resolution DEM from ASTER. ERA5-Land is a high-resolution version of the land component of the ERA5 reanalysis (Hersbach et al., 2020). It is run at 9 km resolution, compared to ERA5 at 31 km, before being interpolated onto a regular $0.1°$ lat/lon grid. This allows for representation of smaller-scale variability than in ERA5 and other reanalysis products. However, there are important limitations to ERA5-Land. Only the land component of the ECMWF Integrated Forecasting System (IFS) is run, with interpolated ERA5 fields providing forcing from other components (i.e. atmosphere, ocean) without two-way coupling. Variables such as surface pressure, wind, precipitation and radiative fluxes therefore do not capture variability at sub-31 km scales, despite being available on the ERA5-Land grid. There is no data assimilation, although observations do indirectly affect the simulation via the forcing from ERA5. The dataset can be downloaded at the Copernicus Climate Data Store (https://doi.org/10.24381/cds.e2161bac).

2.9 Analysis

2.9.1 Characterisation of lakes in the study region

To verify that the six studied lakes are representative of most lakes in the study region, we quantified the physical characteristics of observable lakes in the region using satellite instruments on Landsat 8. We found that most lakes are smaller than those detectable with meteorological satellite sensors (used for the ESA CCI Lakes dataset), which provide quality data and offer frequent revisit time ($\sim$2 days) but have a spatial resolution of (nominally) 1000 m. Consequently, to explore the spatial variability of the lakes in the region, we utilised higher-resolution remote sensing observations from Landsat 8. Even with Landsat 8's finer thermal resolution (nominally $\sim$100 m), careful measures were taken to minimize the influence of mixed land-lake pixels around lake peripheries on the derived lake-averaged properties.

For this analysis, only lakes with an area exceeding 2.5 km^2 were characterised using Landsat 8 thermal observations. The characterisation involved maximum brightness temperature, elevation, water albedo and lake area. The elevation and a limited

(in time) ultra-clear-sky sample of visible reflectance and infrared brightness temperature data from 2015 and 2021 (see Data) were resampled onto a common 50 m-resolution latitude-longitude rectilinear grid to resolve smaller lakes. A composite image was created from several clear-sky Landsat 8 images taken between 2015 and 2021. This composite image consisted of the time-minimum reflectance in the red, blue and green channels in these images.

205   To identify pixels with liquid water, the modified normalised difference water index (NDWI) (Yang and Du, 2017; McFeeters, 2021; Xu, 2006) was applied to the minimum reflectance composite. Shadows north of cliffs were additionally filtered out from the resulting inland water mask using the slope of the ASTER DEM. The DEM is also used to compare the thermal/ice characteristics of the study region to elevation. The area of the lakes was estimated using the water-filled pixels derived from the modified NDWI. Water albedo was estimated by adapting method 1 of Liang (Liang, 2001) to the corresponding bands of Land-
210   sat 8. Finally, for each selected lake, we calculated the following properties: lake-median elevation, lake-mean time-maximum temperature, lake-mean time-minimum albedo and lake area. The statistical estimator used for each property was chosen based on its appropriateness to the property and the nature of possible outliers. For example, taking the minimum across the summer Landsat samples ("time-minimum") of albedo ensures it represents the albedo of the ice-free water body, avoiding skewing of the results by any samples where the lake is only partially ice free.

215   **2.9.2   Determination of the climatological curves**

The climatology of all variables in this study, except for LIC, has been calculated at the lake centre for each day of the year using a 15-day smoothing window (7 days on either side of the central day). This analysis is based on data collected from 1995 to 2020. Additionally, the standard deviation and the number of observations has been evaluated to assess the variability of temperature on specific days across the years, with particular focus on the number of observations for LSWT and LIC.
220   The LIC climatology has been computed as the percentage of ice relative to the total number of ice or water observations, based on data from 2000 to 2020. This climatology is accompanied by the number of days when ice was observed. To reduce false positives, only days with at least 50 observations and confirmed ice presence were included in the LIC climatology. The climatological curves for LSWT and LIC have been derived from the ESA CCI Lakes dataset and the curves for Tair and the solar downwelling flux have been derived from the ERA5-Land dataset. Tair and the solar radiation have been extracted at the
225   closest ERA5-Land resolution cell to the lake centre.

2.9.3   A proxy for stratification phenology

Using observations for LSWT and LIC (ESA CCI Lakes dataset), inferences about the mixing state of lakes can be made (Woolway et al., 2021). In this study, we use LSWT for estimating lake stratification phenology, assuming that the six lakes are dimictic (or monomictic) such as most of the lakes in southwest Greenland (Saros et al., 2016). The day of year in which
230   LSWT crosses the threshold of 3.98°C, the temperature of maximum density of freshwater, is taken as a phenological indicator of stratification (Woolway et al., 2021; Fichot et al., 2019). Where consecutive days with observations are not available, the available days are linearly interpolated in time before the threshold-crossing day is calculated. We have used the temperature of maximum density as a proxy for stratification onset and breakup. However, stratification depends on the lake's depth, size

Table 1. Data usage for each of the aspects analysed in this paper.

	Data used	Results used from
Region characterisation	DEM Landsat 8	
Lake characterisation	Bathymetry ERA5-Land	Region characterisation
Lake seasonal cycle	ESA CCI Lakes LSWT ESA CCI Lakes LIC ERA5-Land	Lake characterisation
Temporal variability in lake mean	ESA CCI Lakes LSWT ESA CCI SST	Lake characterisation
Lake stratification phenology	ESA CCI Lakes LSWT	Lake characterisation
Factor influencing temporal variability and stratification phenology	ERA5-Land ESA CCI Clouds	Temporal variability in lake mean Lake stratification phenology

and shape. ~~Some~~ , but also wind. For example, some small and shallow lakes ~~wind~~ may not experience thermal stratification
since the wind could be strong enough to contrast it and mix the entire lake (Boehrer and Schultze, 2008).

2.9.4 Trend and correlation anaylsis

Trends have been calculated using a linear regression model and the standard uncertainty are reported. The correlations are
evaluated as Spearman correlations, assuming a monotonic relationship and statistically significant correlations (with a p-value<0.05) are highlighted. The relation between LSWT and Tair is evaluated as a function of time to investigate a potential
lag between Tair and LSWT. The LSWT monthly lake mean for each of the ice-free months is correlated with the monthly Tair
values for all the months of the year. The values of Tair and the solar radiation to be compared with the lake monthly mean
have been extracted over a box at the ERA5-Land resolution cell covering the full lake and have been averaged in space and
per month.

2.10 Summary of data usage

Table 1 outlines the data usage (described in Section 2.2 - 2.8) for each of the aspects of the present study (described in details
in Section 3) and relations between them.

3 Results

3.1 Physical characterisation of the study region

In this study, we first characterise the physical environment of the studied lakes in terms of the landscape in which they are
located, as well as their elevation, surface reflectance and surface temperature variability. Two domains are addressed: (i) the

northern domain, which includes four of the six target lakes, and (ii) the southern domain, which includes two lakes (Figure 1). The colour coding of these water bodies in Figure 1(b) are used throughout the results section.
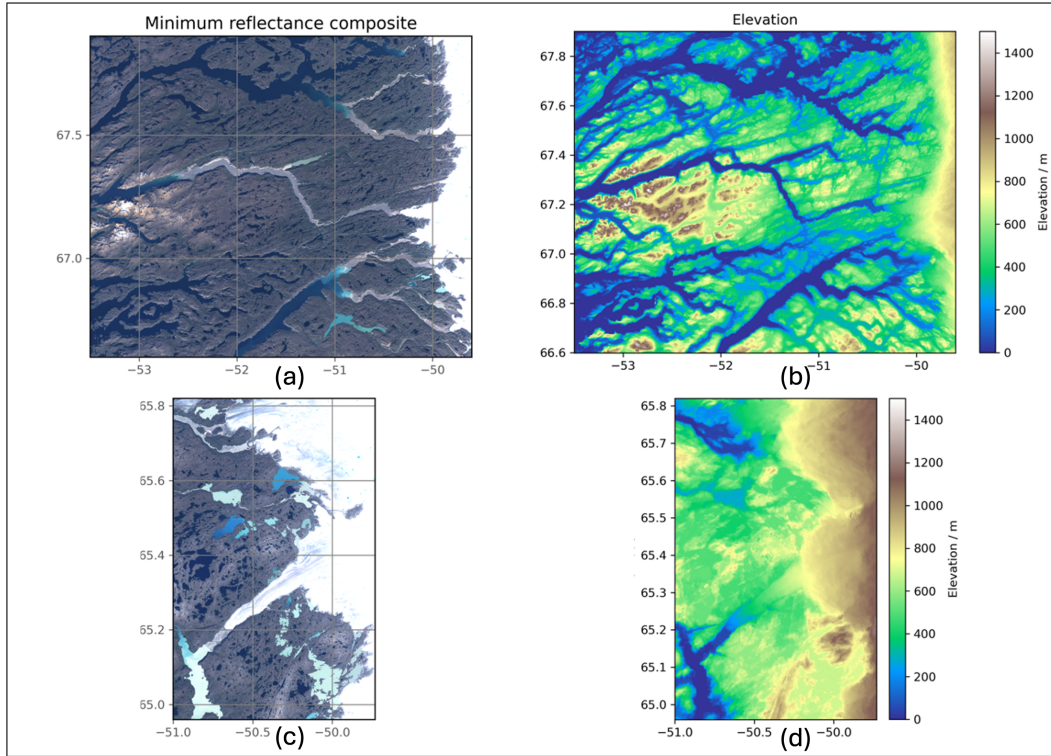
Figure 2. On the left hand side, natural-colour composite of the minimum reflectance in red, green and blue bands of Landsat 8, for the northern (a) and southern (c) domains showing the clear-sky appearance of the study areas in summer at a time of minimum snow and ice cover. The images are composited from several clear-sky Landsat 8 images occurring between 2015 and 2021 inclusively (see Data and Methods). On the right hand side, elevation from the ASTER DEM resampled at 50m resolution for the northern (b) and southern (d) domains.

Comparing thermal/ice characteristics (Figure 2(a) and 2(c)) to elevation (Figure 2(b) and 2(d)), we find that ice and snow persist throughout summer at elevations exceeding ∼1200 m (mountains to west in the northern domain) while the ice-sheet margin is seen at elevations from ∼600 m. The modal elevation of the domain is ∼400 m. Many dark lakes are visible across the landscape of seasonally snow-free permafrost, as well as several prominent bright-blue ice-free water bodies (most obviously in the southern domain) (Figure 2). The latter are directly adjacent to glacier termini or are connected to glacial water inputs by streams or river. Lakes situated in the northern domain are generally more elongated, as is typical of lakes formed in glacially eroded valleys. The lakes of the southern domain more often represent flooded basins in the landscape, sometimes joined by rivers. Using Landsat 8 thermal observations, we identify 90 lakes in the northern domain and 48 in the southern. These include the six lakes for which timeseries data (from the ESA CCI Lakes) are available.

Figure 3 shows the minimum albedo and, in both domains, the higher albedo lakes are either directly connected to the ice-sheet or connected to other bright lakes by rivers (compare Figures 2 and 3). The physical properties of these 138 water bodies of both domains are shown in Figure 4. Most (79%) lakes are dark with albedo of <5%. Of these, 86% were observed with a lake-mean time-maximum temperature exceeding 287.15 K (14°C); however, these warmer lakes tend to be smaller, representing only 66% of the low-albedo water surface area. There is no apparent role for elevation in determining the maximum observed LSWT. This is because higher air temperatures can occur across the central band of the margin at middle elevations, which are maximally distant from the temperature-moderating effects of either the sea or the ice-sheet.

High albedo lakes are not observed at the lowest elevations (<100 m) but occur across the full range of lake surface areas. High-albedo lakes tend to attain less extreme maximum temperatures. This can be partly explained by these lakes reflecting more incoming solar irradiance. Additionally, the inflow of meltwater from the ice-sheet (including indirectly via rivers), which correlates with higher albedo, is likely a contributing factor. Meltwater initially has a temperature just above freezing and likely remains cooler than the lake temperature when it reaches the lakes. This influx of colder water corresponds to a negative heat flux to the lake energy balance.

The six target lakes with time-series data mostly lie within the distribution of other lakes in the landscape, except for Eqalussuit Tasiat (Lake A). Eqalussuit Tasiat (Lake A) has low albedo (indicating good solar absorption), yet its maximum temperature in the Landsat 8 observations is only 281 K ($\sim$ 8°C). This particularly low temperature is partly an artefact of temporal sampling, as the corner of the domain where Eqalussuit Tasiat (Lake A) is located was not observed on the same day that the lakes in the rest of the domain were observed at their maximum temperature. Nonetheless, the general maritime climate of Eqalussuit Tasiat (Lake A)'s location may also play a role.
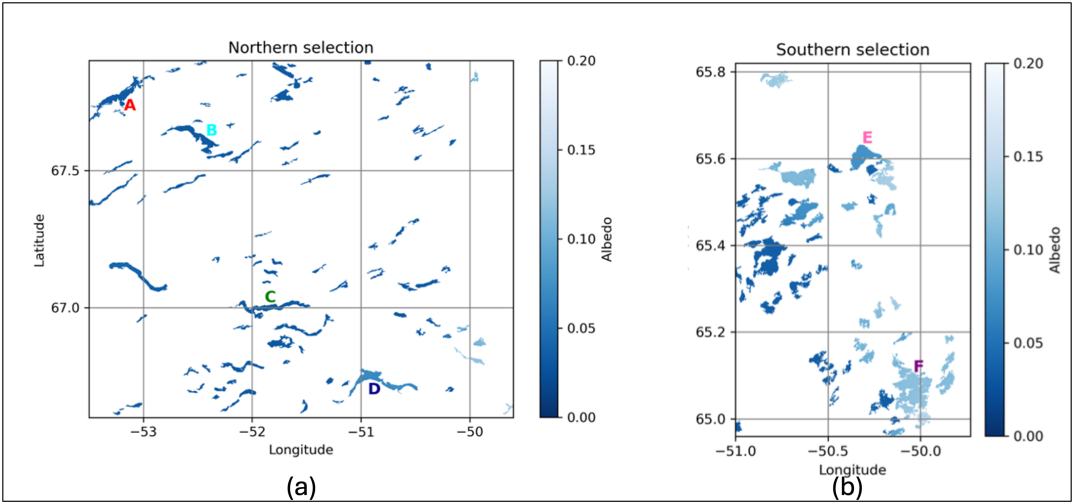
Figure 3. Lake pixels selected for remote sensing of albedo and temperature characteristics in the (a) northern and (b) southern domains. The lake pixels are coloured according to their lake-mean minimum albedo in the Landsat 8 sample (see Data and Methods). The lake names and characteristics of the six lakes A-F are listed in Table 2.
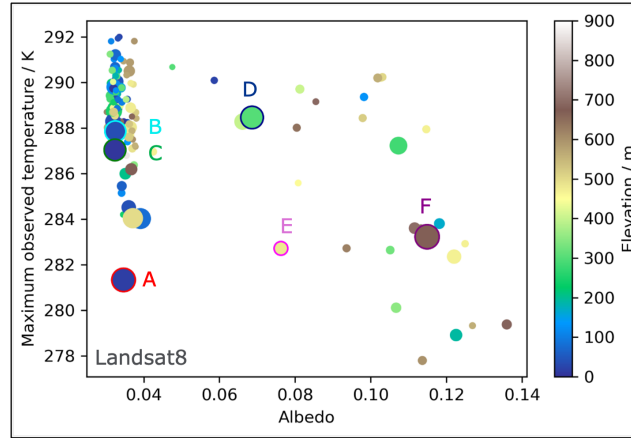
Figure 4. Distribution of physical characteristics of remotely sensible lakes across both study domains. The areas of the scatter points are in proportion to the lake areas. The points corresponding to the six lakes with time-series data are distinguished by borders in accordance with the colour scheme of Figure 1. Their names and characteristics are listed in Table 2. (b). Area, albedo and minimum observed brightness temperature have been derived from Landsat 8 observations as described in Sec 2.9.1. The elevation is from the ASTER DEM.

3.2 Characterisation of the six studied lakes

The six studied lakes (Table 2) in west Greenland represent the variety of larger lakes in the region. Each of the six studied lakes have outflows that at least seasonally connect to fjords or the coastal sea. All six lakes have a maximum distance from the closest land (Carrea et al., 2015) of less than 2 km. Consequently, the six lakes can be resolved in the satellite imagery, given that the best nadir resolution of the meteorological instruments used for LSWT retrieval is 1 km. However, observing these six lakes remains challenging.

Ammalortoq (Lake E) and Tarsartuup Tasersua (F) are situated very close to the GrIS, characterised by a Low Artic continental climate and directly connected to the ice-sheet (Figure 2(c)). They have high albedo due to the presence of fine particulates in the meltwater deriving from the GrIS. While Tarsartuup Tasersua (Lake F) shows a relatively uniform spatial distribution of the albedo, Ammalortoq (Lake E) is composed of three sub-basins with very different albedos (0.12, 0.08 and 0.04), suggesting the inter-basin mixing and flows are weak and/or episodic (depending on lake water levels). Nonetheless, for the purpose of this study, we consider it as a single lake. Most of the LSWT observations for Ammalortoq (Lake E) are from the northern sub-basin where the albedo is 0.08, similar to that of Tasersuaq Aallaartagaq (Lake D).

Figure 5 shows the number of observations across the time-series for each of the cells of Ammalortoq (Lake E) present in a raster of inland water locations used to assist the identification of water-only pixels associated with the lake in meteorological satellite processing (Carrea et al., 2022b). It illustrates that the lake mean temperature obtained is dominated by the centre of the largest sub-basin away from the lake margins where pixels are often detected as partially overlapping land, and therefore not used for LSWT. Some narrow portions of Ammalortoq (Lake E) are never observed with the meteorological sensors for LSWT, which is also true of the other lakes.

Table 2. List of the six lakes present in the ESA CCI Lakes dataset and their characteristics.

CCI Lake identifier	300002493	300002545	300002685	2199	3219	1099
Name[a]	Eqalussuit Tasiat	Nassuttuutaata Tasia	Itinnerup Tasersua	Tasersuaq Aallaartagaq	Ammalortoq	Tarsartuup Tasersua
Identifier in this paper	A (red)	B (cyan)	C (green)	D (blue)	E (pink)	F (purple)
Latitude lake centre	67.779	67.607	66.999	66.754	65.615	65.087
Longitude lake centre	-53.179	-52.429	-51.887	-50.937	-50.318	-50.040
Maximum distance to land (km)	1.6	1.4	1.2	1.8	1.8	1.8
Max depth[b] **(m)**	95	38	44	28	23	302
Mean depth[b] **(m)**	27	11	13	8	6	59
Volume[b] **(km^3)**	2.0	0.7	0.9	0.6	0.3	10.0
Area[b] **(km^2)**	76	64	65	82	55	170
Area L8[c] **(km^2)**	68	54	58	63	48	72
Albedo L8[d]	0.03	0.03	0.03	0.07	0.08	0.11
Elevation DEM[e] **(m)**	15	36	5	299	479	671
NWP nominal elevation[f] **(m)**	150	335	318	534	590	904

[a]Oqaasilerffik placename dataset, distributed with the QGreenland spatial dataset suite (https://qgreenland.org/) (Moon et al., 2023)
[b]From GLOBathy (Khazaei et al., 2022)
[c]Sensible area, Landsat 8
[d]Derived from Landsat 8
[e]From ASTER (Abrams et al., 2020; NASA/METI/AIST/Japan Spacesystems and U.S./Japan ASTER Science Team, 2019)
[f]From ERA5-Land (Hersbach et al., 2020; Muñoz-Sabater, 2019)

Ammalortoq (Lake E) and Tarsartuup Tasersua (F) are, among the six studied lakes, situated at the highest elevation, and at the lowest latitudes. Tarsartuup Tasersua (Lake F) has the largest area and is most likely the deepest lake according to GLOBathy.

Tasersuaq Aallaartagaq (Lake D) is located about 50 km (directly) from the GrIS with an albedo close to the one of Ammalortoq (Lake E). Tasersuaq Aallaartagaq (Lake D) is connected to the ice-sheet by a ∼92 km long stream. This stream transports particulates from small glacier-terminal lakes at an elevation of about 900 m to Tasersuaq Aallaartagaq (Lake D), at an elevation of about 300 m. We expect that this distance is sufficient for the water from the melting ice-sheet to be modified, in terms of its temperature, before reaching Tasersuaq Aallaartagaq (Lake D), despite the albedo of the lake demonstrating that significant glacial flour remains suspended and is transported to the lake.

Eqalussuit Tasiat (Lake A), Nassuttuutaata Tasia (B) and Itinnerup Tasersua (C) absorb a greater proportion of insolation than Tasersuaq Aallaartagaq (Lake D), Ammalortoq (E), Tarsartuup Tasersua (F), having low albedo. They are also at lower elevation. Eqalussuit Tasiat (Lake A) is the lake located at the highest latitude, the lowest elevation, and nearest to open sea. The maximum depth for Eqalussuit Tasiat (Lake A), as reported by the GLOBathy database, is 95 m, but this may be an over-

estimate for a lake so close to the sea and at an elevation of 15 m. Nassuttuutaata Tasia (Lake B) is further inland, separated from Eqalussuit Tasiat (Lake A) by a fjord and mountains reaching a height of about 600 m. Itinnerup Tasersua (Lake C) is
315 roughly midway between open sea and the GrIS.
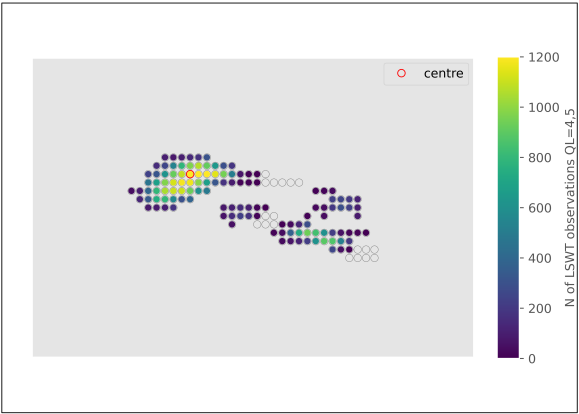
Figure 5. Number of LSWT observations from 1995 to 2022 for Ammalortoq (Lake E) of the ESA CCI Lakes LSWT dataset. Each dot in the plot represents a 1/120°x 1/120° resolution cell of the common regular grid for all the variables of the ESA CCI Lakes dataset.

3.3   Seasonality of LSWT and LIC

3.3.1   Relation between LIC and LSWT seasonality

All the six studied lakes have at least partial ice-free water surfaces in May (the timing of which differs between lakes) to at least mid-November (Figure 6). Figure 6(a) shows the climatology for the centre location of the lakes reported in Table 2.
320 LSWTs are retrieved through to the middle of November, by which time all lakes have cooled below 4°C, such that water column mixing has occurred. All six lakes have complete ice cover at the lake centre for a period starting after mid-November (timing not observed due to no daylight) and ending in May, Figure 6(b). However, the timing of full freezing (on average) of the water surface is uncertain in this dataset because of the observational limits LIC from November to mid-January due to marginal or absent during this period (see Data section). An increase of ice presence can be observed already from mid-October
325 for the lakes directly or indirectly connected to the GrIS (Tasersuaq Aallaartagaq (Lake D), Ammalortoq (E) and Tarsartuup Tasersua (F)), and it can be assumed that these lakes fully freeze shortly after the last LSWT measurements are available. During summer, the reported LIC is less than 0.2% but greater than zero for all the lakes, even at the time of maximum LSWT. This is thought to reflect a base level of misclassification (false ice detection) in these observations rather than ice being present, although, for Ammalortoq (Lake E) and Tarsartuup Tasersua (F), there might be some detection of ice detached from the GrIS
330 (Mallalieu et al., 2020). An unambiguous feature is that Ammalortoq (Lake E) and Tarsartuup Tasersua (F) have later and slower loss of ice cover during May and June. The earliest ice breakup (on average) happens for Itinnerup Tasersua (Lake C), which is the lake with the smallest maximum distance to land (Table 2). For Eqalussuit Tasiat (Lake A), Nassuttuutaata Tasia

(B) and Itinnerup Tasersua (C), the seasonal cycle of the ice cover (Figure 6(b) shows that the percentage of ice pixels is still low by mid-November. LSWTs may therefore also still be obtained in November for ice-free parts of the lakes.
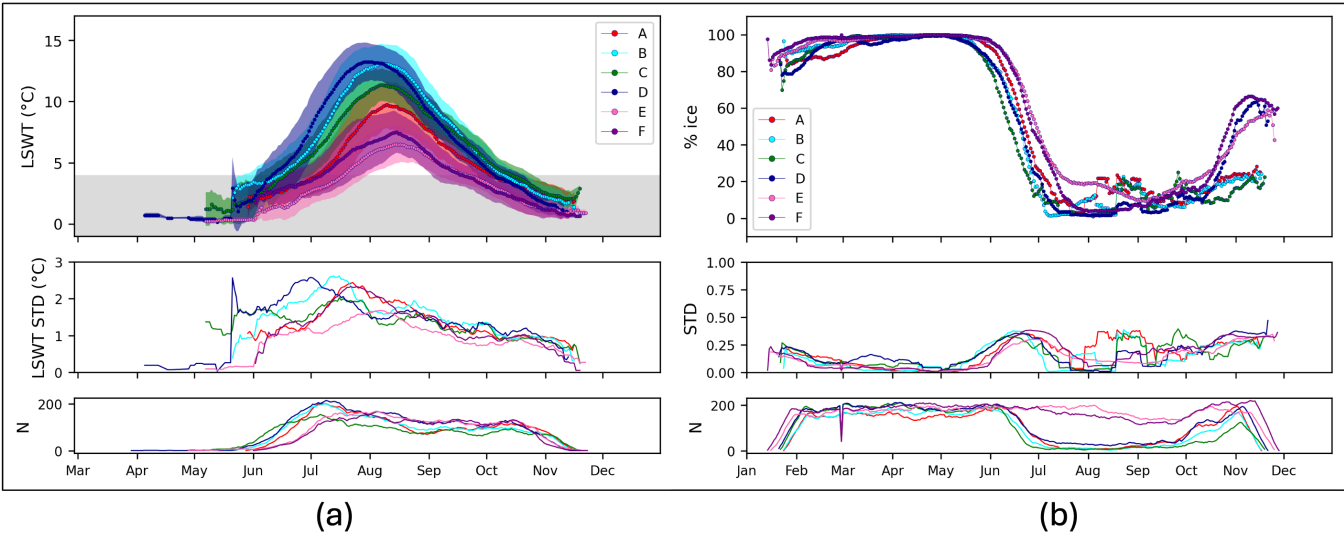
Figure 6. (a) Seasonal cycle of LSWT, calculated as day-of-year mean over 1995 to 2020 of a centred 7-day aggregation window temperatures observations at the lake centre location in °C; the shaded area represents one standard deviation away from the mean (upper panel) and the grey shaded area features temperatures below 4°C. (b) Seasonal cycle of LIC as the percent of cells covered by ice with respect to the total number of valid (ice and water) observations on the lake. Only cases where the number of observations were greater than 50 (ice or water) are included here. The seasonal cycle is plotted for the six lakes (upper panel) and the values of the standard deviation are reported in the middle panel. The number of days with observations is shown in the lower panel.

3.3.2 Maximum average temperatures

The climatological peak temperatures are in broad correlation with the maximum temperature from Landsat 8 reported in Figure 4. Each of the six lakes warm, on average, to a temperature greater than 4°C. The maximum depths estimated for the studied lakes make it plausible they are stratified during periods when LSWT exceeds 4°C although potential stratification is only brief in the case of Ammalortoq (Lake E) and Tarsartuup Tasersua (F). As the studied lakes are covered with ice for approximately half of the year, they can most likely be characterised as dimictic. The coolest studied lakes are Ammalortoq (Lake E) and Tarsartuup Tasersua (F), which are not only situated at the highest elevation but are also directly influenced by the GrIS meltwater. Ammalortoq (Lake E) and Tarsartuup Tasersua (F) reach their climatological maximum value of LSWT of 6.5±1.6°C and 7.5±1.7°C on the 14th and 15th of August respectively. Between May and the end of August, LSWT in Tarsartuup Tasersua (Lake F) is between 1°C and 2°C warmer on average than Ammalortoq (Lake E). Ammalortoq (Lake E) is likely colder than Tarsartuup Tasersua (Lake F) due to its smaller size and shallower depth, and perhaps due to it being more easily cooled by the meltwater from the GrIS. However, the two lakes have very similar temperatures on average after the end

Table 3. Average maximum LSWT and 2m air temperature (Tair); lakes are listed in order of decreasing maximum LSWT.

Lake name	Tasersuaq Aallaartagaq	Nassuttuutaata Tasia	Itinnerup Tasersua	Eqalussuit Tasiat	Tarsartuup Tasersua	Ammalortoq
Lake identifier	D (blue)	B (cyan)	C (green)	A (red)	F (purple)	E (pink)
LSWT max (°C)	13.23	13.11	11.37	9.75	7.49	6.50
DoY LSWT max	30/07	07/08	07/08	10/08	15/08	14/08
Tair max (°C)	11.69	9.76	11.33	8.93	5.80	5.68
DoY Tair max	15/07	14/07	14/07	24/07	16/07	15/07

of August, during the cooling period. Tasersuaq Aallaartagaq (Lake D) and Nassuttuutaata Tasia (B) are the warmest lakes with a very similar maximum average LSWT 13.2±1.6°C and 13.1±1.6°C on the 30th of July and the 7th of August respectively, despite being lakes with different characteristics. As reported in Table 3 the four lakes that are not directly connected to the GrIS peak on average above 10°C. Similar summer maxima have been reported for smaller lakes, not connected to the GrIS, in previous studies (Anderson et al., 2001; Saros et al., 2016).

3.3.3 LSWT in relation with Tair and insolation seasonal cycle

Tair and insolation are two very important drivers of LSWT (Woolway et al., 2017). Despite being the coolest lakes among the six, Ammalortoq (Lake E) and Tarsartuup Tasersua (F) receive the highest insolation (Figure 7(b)) being the lakes at the lowest latitude but Tair is the lowest, being the lakes at the highest elevation showing that very likely the insolation will result in a lower temperature (on average) contributing to the melt of the GrIS which in turn generates a cold water influx into the lake which in turn masks the contribution of the insolation. Tasersuaq Aallaartagaq (Lake D) and Nassuttuutaata Tasia (B) experience very similar maximum temperatures. Tasersuaq Aallaartagaq (Lake D) is likely shallower than Nassuttuutaata Tasia (Lake B). Tasersuaq Aallaartagaq (Lake D) is situated at the lowest latitude among Eqalussuit Tasiat (Lake A), Nassuttuutaata Tasia (B), Itinnerup Tasersua (C) and Tasersuaq Aallaartagaq (D), and (Figure 7(a)) has the warmest seasonal cycle of Tair. Moreover, ERA5-Land indicates higher insolation for Tasersuaq Aallaartagaq (Lake D) than for Nassuttuutaata Tasia (Lake B) (Figure 7(b)). Based on depth, Tair and insolation, we would expect a higher maximum temperature for Tasersuaq Aallaartagaq (Lake D) compared to Nassuttuutaata Tasia (Lake B). The fact that the observed difference is insignificant raises the speculation that the river inflow of Tasersuaq Aallaartagaq (Lake D), fed by a colder proglacial lake, is a significant negative heat flux and that this damps the maximum LSWT.

Figure 6(a) suggests that warmer lakes typically experience an earlier timing of maximum temperature while the correspondent Tair and insolation maximums happen at the similar timing. Table 3 reports the maximum annual LSWT and Tair and the days of the year of these for all lake centres. The maximum Tair occurs, on average, in mid-July, except for Tair near Eqalussuit Tasiat (Lake A) which occurs about ten days later. Eqalussuit Tasiat (Lake A) is located at the highest latitude and is closest to the open sea. The maximum LSWT is higher on average and occurs two to four weeks later than the maximum value of Tair.

The longest lag between Tair and LSWT are calculated for the coldest lakes, Ammalortoq (Lake E) and Tarsartuup Tasersua (F), which are connected to the GrIS.

(a)   (b)
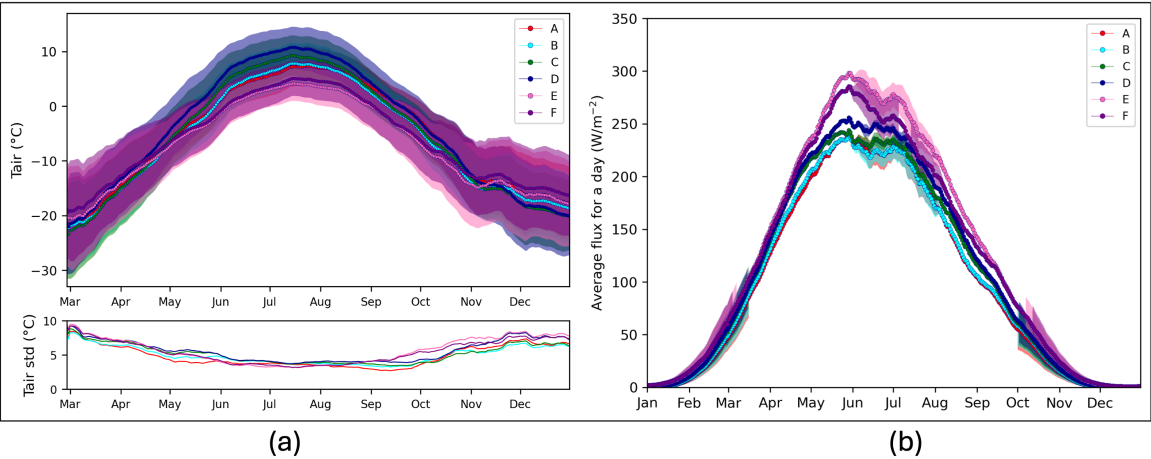
Figure 7. Seasonal cycle (daily means from 1995 to 2020 with a 7-day smoothing window) at the lake centre of Tair ($^\circ$C) (a) and solar downwelling flux (Wm^{-2}) from ERA5-Land reanalysis (b). The seasonal cycle is plotted for the six lakes (upper panel) and for (a) in the lower panel the values of the standard deviation are reported.

3.3.4 Asymmetry of the LSWT seasonal cycle

Figure 6(a) (upper plot) shows that all six lakes warm faster and cool more slowly while the cycles of Tair and insolation are
375   symmetrical. The asymmetry in LSWT over the summer season is due to the effect of ice cover. Upon ice break up in May, cooler LSWTs are in relatively large disequilibrium with the prevailing Tair and insolation, and the LSWT rapidly increases in response. During the cooling phase, in contrast, the lake has a steadier lagged response to the progressive decrease in Tair and insolation through late summer and autumn. On average, the earliest retrieved LSWTs at the lake centre are a few degrees above 0°C. This is most likely a sampling effect arising because the first cloud-free observation of the lake centre after ice melt
380   is not immediately after the open water first forms. However, we note that warmer-than-freezing temperatures at the time of ice melt has been reported for lakes in the Tibetan plateau (Kirillin et al., 2021) and Ontario in Canada (Williams, 1969). In the circumstances of dry, cold atmospheric conditions, solar radiation may be relatively strong and the ice snow free, enabling a sufficient flux of insolation through the ice to increase the temperature of the water below, even above 4°C such that there is convective mixing before ice melt. It is uncertain from the satellite observations alone whether such effects are relevant to
385   these lakes.

3.4 Temporal variability in lake-mean LSWT

The long term LSWT variations for the six lakes are shown in Figure 8 for the months of June, July, August and September when the lakes are (mostly) ice-free. The warmest months are July and August, when also the temperatures among the lakes differ most. The ranking with respect to temperature tends to be maintained throughout the seasonal cycle.

(a)   (b)
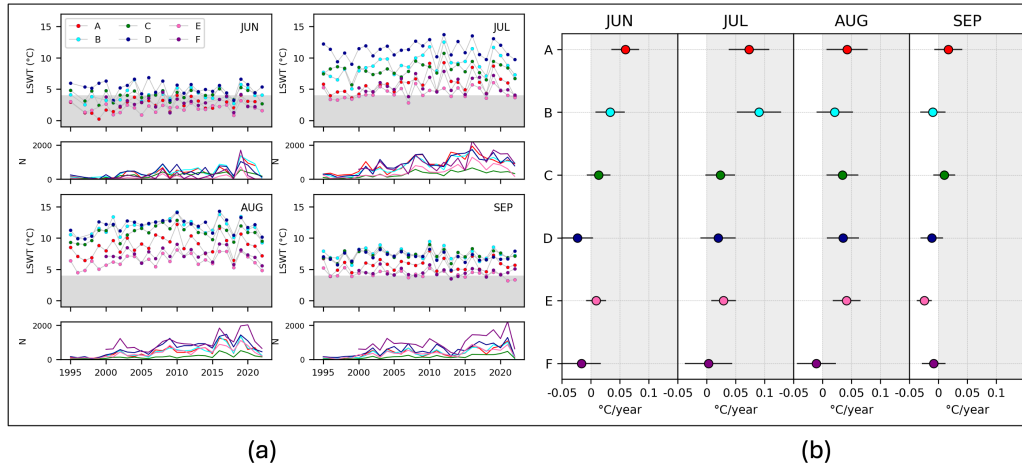
Figure 8. (a) Monthly spatial lake mean time series for the months of June, July, August and September for the six lakes (upper plot for each of the months) and the number of observations to create the means (lower plot for each of the months). The grey shaded area features temperatures below 4°C. (b) Trends (linear regression slopes) with standard uncertainty of trend (black bar) of the monthly spatial lake mean time series across years for each month and for each lake. The grey shaded area highlights positive values of the trends.

390   In June and July, monthly lake-mean LSWT in Tasersuaq Aallaartagaq (Lake D) is noticeably higher than the other lakes across the years, while in August and September the LSWT of Tasersuaq Aallaartagaq (Lake D), Nassuttuutaata Tasia (B) and Itinnerup Tasersua (C) are comparable. Ammalortoq (Lake E) and Tarsartuup Tasersua (F), being connected to the GrIS, do not exhibit a considerable difference in lake mean temperature throughout the years for any of the months. The six time series of monthly lake-mean LSWTs in Figure 8(a) reveal temporal trends whose values are displayed in Figure 8(b). Twelve of the 395   15 statistically significant trends in Figure 8(b) are positive (warming over time). The trends for Eqalussuit Tasiat (Lake A) and Nassuttuutaata Tasia (B) in July are respectively 0.090.07±0.03°C year^{-1} and 0.09±0.04°C year^{-1}. Comparing the six lakes, Eqalussuit Tasiat (Lake A) has consistently the highest trends (lake warming at highest rate) across the months. The highest values for the trend of Eqalussuit Tasiat (Lake A) are in July (0.060.07±0.020.03°C year^{-1}) and August (0.070.04±0.03°C year^{-1}). This can be related to the fact that Eqalussuit Tasiat (Lake A) is not connected to the GrIS, it is close to the coast and 400   therefore possibly more exposed to meteorological drivers than the other lakes.

    To consider the marine influence on Eqalussuit Tasiat (Lake A), which is close to the coast, SSTs for June to September has been computed from the ESA CCI SST dataset (Merchant et al., 2019), covering a 2 degree box centred on Eqalussuit Tasiat's latitude immediately to the west. LSWT is 2°C warmer than the nearby SST during July and August, and warmer by less a 1°C

Table 4. Trend comparison between Eqalussuit Tasiat (Lake A) and the closest sea for June, July, August and September.

	Jun	Jul	Aug	Sep
correlation	0.44	0.62	0.69	0.72
p-value	0.02	0.0004	10^{-5}	10^{-5}
trend lake A (°C year^{-1})	0.06±0.02	0.07±0.03	0.04±0.03	0.02±0.02
trend sea (°C year^{-1})	0.00±0.01	0.00±0.02	0.03±0.02	0.02±0.01

in June and September. The inter-annual correlation between LSWT and SST is positive and statistically significant for all four
405 months. The mean trend across the four months is 0.01±0.01°C year^{-1} in the offshore SST, and is 0.05±0.01°C year^{-1} in
the LSWT. Together, the statistics suggest a coupling to the marine climate and an amplified response. The monthly statistics
are shown in Table 4. The lakes connected to the GrIS (Tasersuaq Aallaartagaq (Lake D), Ammalortoq (E) and Tarsartuup
Tasersua (F)) tend to show smaller interannual trends than the unconnected lakes, although the distinction is not so clear in
August. Ammalortoq (Lake E) is overall warming faster than Tarsartuup Tasersua (Lake F). Tasersuaq Aallaartagaq (Lake D)
410 experiences contrasting trends, negative in June and September and positive in July and August. Speculatively, the overall
distinction between Eqalussuit Tasiat (Lake A), Nassuttuutaata Tasia (B), and Itinnerup Tasersua (C) compared to Tasersuaq
Aallaartagaq (Lake D), Ammalortoq (E) and Tarsartuup Tasersua (F) may reflect a negative feedback for the connected lakes
via the flux of meltwater. However, Tasersuaq Aallaartagaq (Lake D) is connected only via a river, and, speculatively, warming
of the meltwater during their passage to Tasersuaq Aallaartagaq (Lake D) could act in opposition to this negative feedback,
415 particularly in July and August when the landscape is warmest.

3.5 Lake stratification phenology

We estimate lake stratification phenology as the day of the year when the LSWT crosses 3.98°C. On average, warmer lakes
in the study region experience an earlier onset and later breakup of thermal stratification. In turn, the cooler lakes, which
are directly connected to the GrIS, experience the latest onset and earliest breakup of stratification (Figure 9(c)). Tarsartuup
420 Tasersua (Lake F) on average reaches ∼4°C on the 1st of July and Ammalortoq (Lake E) on 7th of July, while for the other
lakes this happens between the 11th and 28th of June. Following the warm season, LSWTs in Tarsartuup Tasersua (Lake F)
cools, on average, below ∼4°C on 22nd of September and Ammalortoq (Lake E) on 18th of September, while for the other
lakes this happens between the 1st and the 12th of October. The time series of the days of the year of onset and breakup of
stratification follow an asymmetric cycle: the break up date is relatively stable across the six lakes, while the onset tends to
425 occur substantially earlier in the year during the study period, other than for Tarsartuup Tasersua (F) (Figure 9(a)). The largest
rates of change in stratification onset are -0.5±0.2 days year^{-1} for Eqalussuit Tasiat (Lake A) and Nassuttuutaata Tasia (B)
(Figure 9(b)). The breakup of stratification only occurs later in the year for Eqalussuit Tasiat (Lake A) (0.2±0.2 days year^{-1}
). The warm stratified season is lengthening for the three lakes that are not connected to the GrIS (Figure 9), while this is less
clear for the connected lakes given the trend uncertainties.
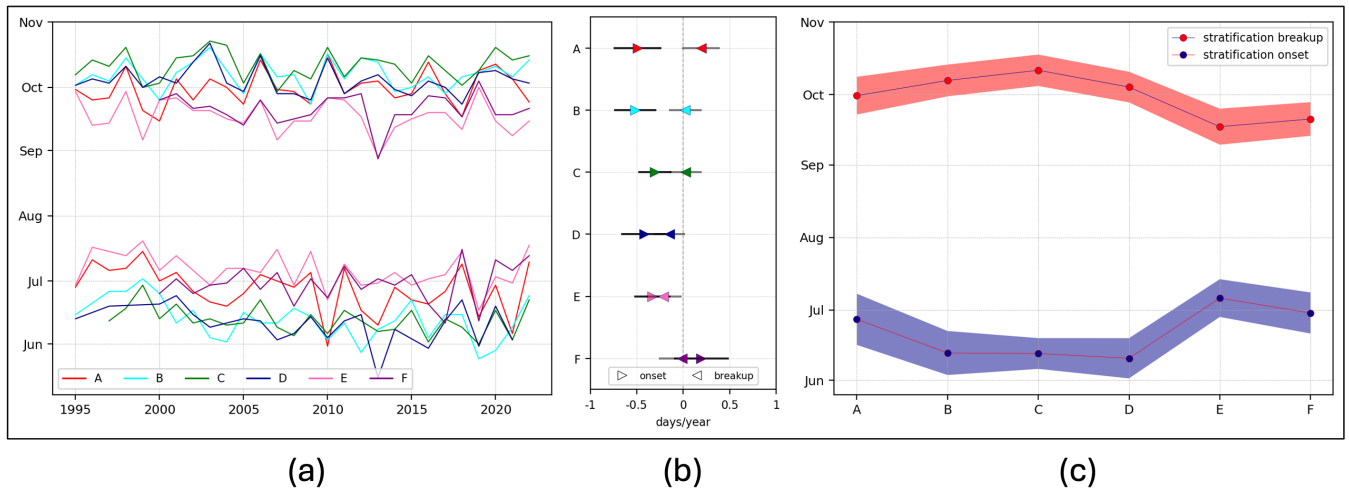
Figure 9. (a) Time series of the stratification onset and breakup timing (day of the year) for the six lakes. (b) Trends (linear regression slopes) with standard uncertainty (black bar) of the stratification onset and breakup timing (days per year) for the six lakes. (c) Average day of the year for the stratification onset and breakup, where the shaded area represents one standard deviation from the mean, expressing the variability of the thermal stratification phenology across the years.

3.6 Factors influencing the variability in LSWT and stratification phenology

Tair is typically considered the most important atmospheric driver of LSWT (Edinger et al., 1968), but their relationships differ across seasons. Here, we calculated the correlation (Spearman) between Tair and LSWT anomalies across the studied lakes (Figure 10) in time. Figure 10(a) and 10(b) show the interannual correlation between the June and August (respectively) LSWT anomaly and the Tair anomalies of each of the months of the year. For all six lakes, LSWT anomalies in June are best correlated with Tair anomalies in May (Figure 10(a)). This seems to arise because the lake ice-off date strongly influences the June LSWT anomaly. Earlier ice-off results in a rapid initial rise of LSWT during summer, thus influencing the June LSWT. Warmer Tair in May can lead to earlier ice-off. This interpretation is supported by the observation that the stratification onset date is better correlated with the May Tair anomaly (mean of correlation coefficient across the six lakes: 0.7±0.1) than with that of June (0.5±0.1), despite stratification onset usually occurring during June. In contrast, LSWT anomalies in August (Figure 10(b)) are most significantly correlated with the August Tair anomalies. During this time the lakes are likely thermally stratified and the volume of water interacting with the atmosphere is relatively small. The exceptions to the high correlation for August are Ammalortoq (Lake E) and Tarsartuup Tasersua (F) that directly receive GrIS meltwater. Since warm Tair is expected to increase meltwater discharge (but not, for an ice-marginal lake, meltwater temperature), it appears that LSWT-Tair correlation for August is in these cases suppressed by a negative feedback via enhanced cooling by water influx from the GrIS.

The most important driver of the stratification phenology is air temperature (Woolway et al., 2021). The correlation between stratification phenology and Tair is found to be relevant and significant only for the onset, while for the break up only mild

(a)   (b)
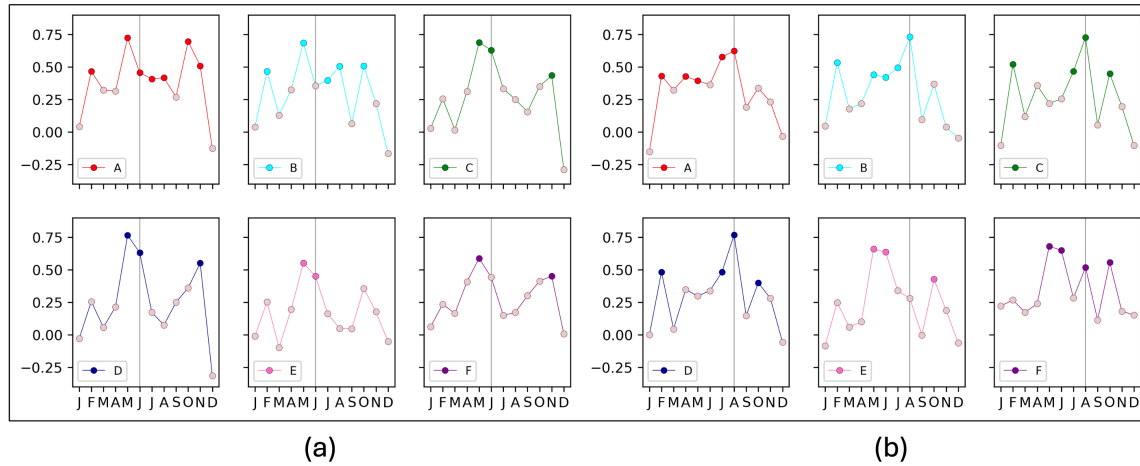
Figure 10. Correlations (Spearman) between LSWT anomalies for the month of June (a) and August (b) and the Tair anomalies from ERA5-Land for all the months in the year. The grey circles indicate a p-value greater than 0.05.

correlations appear for the lakes connected to the GrIS and Eqalussuit Tasiat (Lake A). For Tasersuaq Aallaartagaq (Lake D), which is connected to the GrIS through a long stream, the correlations with Tair are not significant, indicating that the changes in the stratification timing are driven by other factors. The highest correlations are found for stratification onset for the lakes directly connected to the GrIS (Ammalortoq (Lake E), Tarsartuup Tasersua (F)) and for Eqalussuit Tasiat (Lake A)). Correlations with the shortwave solar radiation anomalies have been investigated using the surface solar irradiance values from Era5-Land and using the ESA CCI Clouds observations (Stengel et al., 2020). No significant correlation has been found for any of the lakes with either dataset.

4 Discussion

4.1 Spatial thermal characteristics of lakes in the region with high resolution sensor

High-resolution thermal remote sensing allowed to characterise the lakes in the region spatially only since currently high-resolution sensors have low revisiting time. Analyzing data from 138 lakes observable in Landsat 8 imagery, we observed a wide spectrum of maximum LSWTs varying from 5°C to 20°C, alongside varying surface albedo values spanning 0.03 and 0.14. Notably, high-albedo lakes, often connected directly to the GrIS or indirectly through river systems, tend to exhibit cooler temperatures. This cooling effect is attributed to their enhanced reflectivity, which reduces solar absorption, coupled with the influence of cold glacial meltwater inflows. Conversely, smaller and typically shallower lakes with lower albedo tend to absorb more solar radiation and consequently reach warmer temperatures (Read and Rose, 2013; Heiskanen et al., 2015). These findings underscore the complex relationship between lake morphology, solar radiation dynamics, and climatic influences, essential for understanding the response of these ecosystems to climate change. The six lakes that have been selected

465 specifically for this study are representative of the lakes in the region exhibiting different albedos, temperatures and elevations. The only limitation to the variety of lakes was the size since meteorological satellites used to create the ESA CCI Lakes dataset have a spatial resolution of about 1km.

4.2 Seasonal trends in LSWT and LIC

Meteorological sensors provided critical time series of LSWT and LIC observations that enabled us to uncover nuanced insights
470 into the thermal dynamics of these water bodies. These six studied lakes had contrasting elevation and proximity to the GrIS, as well as to the sea. Seasonally, these lakes exhibit significant variability in ice cover, typically remaining ice-free from June through to November, although the exact timing of complete ice reformation is affected by some observational uncertainty. This seasonal pattern, notably asynchronous with the solar insolation cycle, underscores the complex interplay of climatic factors influencing lake dynamics. The detailed examination of these lakes reveals distinct thermal dynamics and responses to
475 environmental stimuli. Lakes not directly connected to the GrIS typically reach peak temperatures ranging from 9°C to 14°C during the summer months, exceeding average air temperatures due to direct solar heating. In contrast, ice-marginal lakes exhibit cooler peak temperatures around 7°C, a phenomenon not fully explained by their elevated albedo alone. The influx of cold meltwater from nearby glaciers likely plays an important role impacting the thermal balance and seasonal variations of these water bodies. Analysing the seasonal cycle, we found that the warm stratified season of the studied lakes typically spans
480 from mid-June to early October, though slightly shorter for cooler ice-marginal lakes and that all six lakes warm faster and cool more slowly due to the effect of the ice cover. In contrast, the forcing cycles of Tair and insolation are relatively symmetrical. When the ice breaks up in May, the cooler LSWTs are significantly out of balance with the prevailing Tair and solar radiation, leading to a swift rise in LSWT. In contrast, during the cooling phase, the lake reacts more slowly, with a lagged response to the gradual decline in Tair and solar radiation throughout late summer and autumn.

485 4.3 Temporal variability in LSWT

The observed interannual variability in these processes highlights the sensitivity of West Greenland lakes to fluctuations in atmospheric conditions, emphasizing their role as indicator of broader climatic trends in the region. Over the months of June to September, LSWT trends varied from insignificantly negative to warming trends of 0.5°C to 1°C decade^{-1} across different lakes with clear differences for ice-marginal lakes, coastal lakes and inland lakes. Stratification onset appears to be advancing
490 by approximately 5 to 8 days per decade, while the end of stratification shows more stability over time, consistent with studies from other regions worldwide (Woolway et al., 2021). Our findings suggest that ice-marginal lakes may increasingly interact with the dynamics of the GrIS margin, potentially becoming more important in the future (Mallalieu et al., 2021; Carrivick et al., 2022). Changes in water temperature and stratification phenology substantially impact light availability, nutrient cycling, and oxygen levels crucial for lake ecosystems (Woolway et al., 2021). Surface water temperature serves as a 'sentinel' of
495 climate change, integrating various climatic and inlake drivers that influence the lake's surface energy budget (Woolway et al., 2021; Schneider and Hook, 2010). Understanding these dynamics is crucial for assessing the resilience of West Greenland's lakes to ongoing environmental changes and their implications for ecosystem health and function.

4.4 Influence of air temperature and insolation on LSWT and lake stratification

Air temperature conditions leading up to the ice-free season, particularly in May, emerge as a pivotal factor influencing both stratification onset and the timing of ice-off events. The onset of stratification correlates closely with air temperature in the preceding month, influencing the rate of lake ice melt. Our study found no significant influence of insolation on this process. Warmer temperatures during the period leading to the ice-free season accelerate the melting of lake ice, influencing stratification dynamics and subsequent thermal regimes. The influx of cold meltwater from nearby glaciers likely acts as a significant cooling mechanism, influencing the thermal balance and seasonal variability of these water bodies. The high albedo of ice-marginal lakes suggests substantial influx of mass and meltwater from the GrIS. This meltwater introduces a negative heat flux that correlates with August air temperatures, affecting glacial melting dynamics. The lack of correlation with August air temperature for ice-marginal lakes arises from competing mechanisms, notably warming from solar radiation and cooling from meltwater influx associated with warmer summer air temperatures, complicating the LSWT response to climate change.

4.5 Implications for LSWT and LIC studies in Greenland and the Arctic

Observations of the temperature of pro-glacial lakes in this landscape are sparse. It has been assumed in some mass balance models of glaciers terminating in lakes that the lakes do not warm to more than 1°C (Truffer and Motyka, 2016; Chernos et al., 2016), a notion contradicted by the remotely sensed data analysed in this study. This discrepancy underscores the necessity of leveraging remote sensing to complement sparse in situ data in regions like West Greenland. However, only large lakes (of area at least 3km^2) can be observed with good temporal resolution by meteorological satellites (used for ESA CCI Lakes dataset) which offer a sufficiently long time series to detect climatic signals and well-characterised uncertainty and stability. For smaller lakes, higher spatial resolution satellites, such as Landsat 8, can offer observations at a finer spatial resolution (100m) but a much longer revisiting time is often insufficient for a good description of the temporal variations. An example is the use of the high spatial resolution sensor ASTER to characterise the thermal behaviour of lakes in Arctic Sweden where the measurements considered were very sparse in time (Dye et al., 2021). These are current limitations of remote sensing data, which will be partly addressed with improvements in spatial and temporal resolution such as those provided by the satellites for thermal remote sensing such as TRISHNA (Lagouarde et al., 2018) to be launched after 2025. Furthermore, lake-glacier interaction mechanism could be better investigated through new field work. Collecting ground truth data would be very useful to more precisely quantify the interactions between lakes and the ice-sheet and to unravel the complex thermal structure at various depths of these lakes and how it varies in space and time. Such approaches are essential for elucidating the complex relationships between climatic variables, ice dynamics, and lake behaviours, providing critical insights into ecosystem responses to climate change in such a key region for local and global climate.

5 Conclusions

Lakes are important features of West Greenland's seasonally ice-free landscape, their biogeophysical interactions in summer being in part determined by their energy budget and the LSWT they attain. In this study, we presented new, long-term and consistent LSWT information for lakes across this landscape using satellite data.

A broader analysis across the region to characterise thermally the lakes in space was achieved with high spatial resolution data from Landsat 8. Investigations on specific dynamics were conducted on six lakes in West Greenland using high temporal resolution meteorological sensors data, the ESA CCI Lakes data. The six lakes are located at contrasting elevations and proximity to the GrIS and to the coast, allowing us to unveil their long term thermal behaviour (seasonal climatology and interannual variability) for the first time and to attempt to disentangle the complex interaction between lake morphology, solar radiation dynamics, and climatic influences the influence of atmospheric variables on the lake dynamics in details and in a consistent manner. Contrary to some mass balance models of glaciers terminating in lakes, we found that these lakes warm to much more than $1°C$ reaching on average a water temperature between 6.6 and $7.5°C$. The lakes not connected to the ice-sheet reach an average surface water temperature of more than $13°C$.

The most important atmospheric driver is air temperature. However, the ice-sheet significantly impacts the thermal behaviour of the lakes, frequently overshadowing the atmospheric signatures. For example, no correlation is found in summer between LSWT anomalies and Tair only for the ice marginal lakes, indicating a mitigation of Tair impact on the lakes by an enhanced cooling of water influx from the GrIS. The vicinity of the coast also influences lake thermal state. The lake temperature is influenced by the marine climate and a comparison with nearby SST shows that the lake amplifies the response, exhibiting consistently the largest trend in temperatures (increase at a rate of more than half a degree per decade) and in stratification onset and breakup (half a day per year earlier) and breakup (a quarter of a day later per year) timing among the six lakes.

Advancements in satellite remote sensing have enhanced the radiometric, spatial, and temporal quality of observations, enabling long-term thermal characterisation of lakes covering large areas in a consistent way, as demonstrated in this study. Satellite data are especially crucial in remote and data-sparse regions such as Greenland. However, current Earth observation missions often face trade-offs between spatial and temporal resolution. These limitations are expected to be addressed by upcoming missions such as TRISHNA, which aim to provide high-resolution thermal and optical data with improved revisit times. Complementary in situ data collection remains critical for quantifying the hydrological and cryospheric interactions, particularly those between proglacial lakes and the Greenland Ice Sheet.

Data availability. All the datasets utilized for the current study are all available online as open data and in particular: the digital elevation model ASTER GDEM Version 3 (ASTGTM) is available at the Land Processes Distributed Active Archive Center (LP DAAC) https://doi.org/10.5067/ASTER/ASTGTM.003, the bathymetry GLOBathy dataset is available at figshare https://doi.org/10.6084/m9.figshare.c.5243309.v1, the ESA CCI Lakes - LWST and LIC dataset available at the Centre for Environmental Data (CEDA) archive http://dx.doi.org/10.5285/a07deacaffb8453e93d57ee214676304, the ESA CCI Clouds dataset is available at the Deutscher Wetterdienst archive https://dx.doi.org/doi:10.5676/DWD/ESA_Cloud_cci/AVHRR-AM/V003, the ESA CCI SST dataset is available at the CEDA archive http://dx.

560 doi.org/10.5285/62c0f97b1eac4e0197a674870afe1ee6, and the ECMWF ERA 5 Land data are available at the Copernicus Climate Data Store https://doi.org/10.24381/cds.e2161bac. The Landsat 8 data are available at the Landsat Data Access web page https://www.usgs.gov/landsat-missions/landsat-data-access where to discover how to search and download all Landsat products from United States Geological Survey (USGS) data portals. The USGS Landsat no-cost open access data policy remains intact since its inception in 2008. Also, all the datasets used and/or analysed during the current study are available from the corresponding author on reasonable request.

565 *Author contributions.* L.C. and C.M. designed the study, performed the analysis, interpreted the results, and wrote the manuscript. R.I.W. contributed in writing and revising the manuscript. N.M. dealt with the extraction of the Landsat 8 data.

Competing interests. The authors declare no competing interests.

Acknowledgements. The authors acknowledge the European Space Agency for this work with gratitude. The European Space Agency supported the Climate Change Initiative - New ECVs for Lakes, which in turn has has provided the majority of support leading to the outcomes
570 herein described, via grant reference 4000125030/18/I-NB. The authors acknowledge the United Kingdom National Center for Earth Observation (NCEO) for funding the development of some of the software used to process Landsat 8 data, courtesy of the U.S. Geological Survey and also acknowledge the United Kingdom National Environment Research Council (NERC) for funding part of this work within the project on Greenland ice marginal lake evolution as driver of ice-sheet change (NERC reference: NE/X013537/1).

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