https://doi.org/10.5194/egusphere-2024-2913
Preprint. Discussion started: 23 September 2024

Technical Note:

A note on stabilization mechanisms of, e.g., Atlantic Ocean meridional overturning circulation

by Hans van Haren

Royal Netherlands Institute for Sea Research (NIOZ), P.O. Box 59, 1790 AB Den Burg, the Netherlands.
e-mail: hans.van.haren@nioz.nl

https://doi.org/10.5194/egusphere-2024-2913
Preprint. Discussion started: 23 September 2024

Short summary. The extent of mankind's influence on Earth's climate warrants ocean-studies. A supposed major heat-transporter is the Atlantic Meridional Overturning Circulation (AMOC). As AMOC is a complex nonlinear dynamical system, mathematical models may predict its potential collapse using single parameters like surface temperature. However, physical processes such as (sub-)mesoscale eddy transport and turbulent mixing by internal wave breaking will alter the estimators, so that the AMOC may not collapse.

Abstract. The extent of anthropogenic influence on the Earth's climate warrants studies of the ocean as a major player. The ocean circulation is important for transporting properties like heat, carbon and nutrients. A supposed major conduit is the Atlantic Meridional Overturning Circulation (AMOC). As the AMOC is a complex nonlinear dynamical system, it is challenging to predict its potential to collapse and/or reversal of direction from a statistical viewpoint using a single parameter like sea-surface temperature or freshwater influx in numerical models. However, as is argued in this note supported by spectra from ocean observations, physical processes such as transport by sub-mesoscale eddies and turbulence-generating breaking of internal waves that are not incorporated in these models will alter such parameters, and thereby statistical analyses. This may lead to feed-back mechanisms on property gradients such as density stratification so that the AMOC may not collapse.

1 Introduction

Schematically, the Atlantic(-Ocean) Meridional Overturning Circulation (AMOC) transports heat from the equator to the poles near the surface and carbon in the abyssal return (e.g., Aldama-Campino et al., 2023). It includes physical processes like 'deep dense-water formation' in the polar region. Recent mathematical and numerical modelling such as based on varying single parameters like sea-surface temperature (e.g., Ditlevsen and Ditlevsen, 2023) and freshwater influx (e.g., van Westen et al., 2024) suggest a potential future collapse of the AMOC. It is argued that this may have consequences for Northwest-European climate.

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Whilst the modelling might be robust mathematically, it lacks physical processes of the drivers of the AMOC and observational evidence thereof. This will have consequences for the feed-back mechanisms at work in the nonlinear dynamical system of ocean circulation. As has been reviewed for AMOC numerical models (Gent, 2018), important feed-back mechanisms are vertical (turbulent) mixing, (sub-)mesoscale gyre (eddy) transport, and the coupling with the atmosphere. Here we elaborate on the importance of turbulence induced by internal wave breaking, possibly coupling with sub-mesoscale eddies, and stability variations in vertical density stratification for such feed-back, by reviewing insights from recent modeling and deep-sea observations. In particular, the core of ocean motions is spectrally investigated focusing on most energetic mesoscale, internal wave, and turbulence scales.

In contrast with the atmosphere, the ocean is not a heat engine (Wunsch and Ferrari, 2004). As a result, the AMOC is not buoyancy-driven via push by deep dense-water formation near the poles, which notably occurs in sporadic pulses rather than continuously. Instead, the AMOC is wind- and tide-driven, with turbulent mixing by internal wave breaking being considered an important physics process of pull. Winds, near the ocean surface, and tides, via interaction with seafloor topography deeper down, contribute about equally to generate internal waves that are found everywhere in the ocean interior. Such waves break predominantly at ubiquitous underwater seamounts and continental slopes.

Without turbulent mixing, the AMOC would be confined to a 100-m thick near-surface layer and the deep-ocean would be a stagnant pool of cold water (Wunsch and Ferrari, 2004). This is not the case however, and the solar heat is mixed from the surface downward so that the ocean is stably stratified in density all the way into its deepest trenches, as has been shown in hydrographic deep-ocean observations (Taira et al., 2005; van Haren et al., 2021a). Although turbulent mixing by internal wave breaking in the ocean-interior is insufficient by at least a factor of two to maintain the vertical density stratification (e.g., Gregg, 1989, Polzin et al., 1997), such breaking along ocean boundaries has been suggested to be more than sufficient (Munk, 1966; Polzin et al., 1997). Especially large internal wave breaking is expected to occur above steeply sloping topography (Eriksen, 1982; Thorpe, 1987; Sarkar and

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Scotti, 2017). Because there are more and larger seamounts than mountains on land, equally abundant sloping seafloors lead to abundant turbulent mixing, as has been charted from recent observations and modelling results summarized below.

2 Recent internal wave breaking results

Detailed observations and numerical modeling have revealed the extent of internal tide breaking processes above ocean topography (van Haren and Gostiaux, 2012; Winters, 2015; Wynne-Cattanach et al., 2024). Quantification of the turbulent mixing shows that it occurs with typical tidal-period-average values that are more than 100 times larger over (just) super-critical slopes than open-ocean values. A super-critical seafloor slope is steeper than the slope of internal wave characteristics. While ocean-wide tides energetically dominate internal waves, not all seafloor slopes are super-critical for these waves. In contrast, nearly all seafloor slopes are super-critical for (at least one component of) secondly energetic near-inertial waves, which are generated via geostrophic adjustment following the passage or collapse of a disturbance such as fronts or atmospheric storms on the rotating Earth. Under common stratification, near-inertial waves are at the lowest frequency of freely propagating internal waves. The highest frequency propagating internal waves, near the buoyancy frequency, experience only vertical walls as super-critical seafloor slopes.

Within a tidal period, turbulence peaks in bursts of shorter duration than half an hour when highly nonlinear internal waves propagate as internal bores up a slope, once or twice a tidal cycle. The breaking of bores leads primarily to convective, buoyancy-driven turbulence, rather than frictional shear-turbulence over the seafloor (van Haren et al., 2013; van Haren et al., 2024). Between bores, the turbulent mixing varies by an order of magnitude in intensity, with effects extending about 100 m vertically and several kilometers horizontally from the seafloor. Although intermittently occurring at a given position of the sloping seafloor and about 10% varying in arrival time, the turbulence is generated internally by the tide, for about 60% (Wunsch and Ferrari, 2004), and by winds, for about 40%, in a stratified ocean-

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environment. The turbulent bores also resuspend sediment and thereby replenish nutrients away from the seafloor (Hosegood et al., 2004), important for deep-sea life.

Question is whether the intensity of internal-wave induced deep-ocean turbulence is affected by variations in sea-surface temperature or salinity, with what consequences for the AMOC. In considering these it should be noted that various properties determine different equilibria. For example, deep dense-water formation does not only occur in polar seas, but occasionally also in the at least 10°C warmer Mediterranean (Gascard, 1978), with an important contribution of atmospheric exchange due to orographic generated winds affecting the preconditioning by cooling and drying of near-surface waters. Similarly, internal waves occur in oceans and in the Mediterranean, but tides are relatively weak in the latter, and yet 'sufficient' turbulent mixing is generated via near-inertial motions mainly (van Haren et al., 2013).

3 Revisiting Mediterranean observations as a proxy for ocean conditions

In many physical oceanographic aspects of heat and salt budgets, large-scale water-flow circulation, eddies at sub-mesoscales, near-inertial motions including gyroscopic waves and internal wave turbulence, the Mediterranean Sea can be considered a sample for the state of the much larger oceans (e.g., Gascard, 1973; Garrett, 1994; Milllot, 1999; van Haren and Millot, 2004; Testor and Gascard, 2006).

In the Northwest Mediterranean, vertical density stratification varies markedly with seasons and years, having relatively large near-surface values in summer and relatively low values in winter. The proximity of extensive mountain ranges on land generates highly variable winds that can cool and dry surface waters. In winter in weaker stratified waters, this may lead to unstable conditions of buoyancy driven convection in an exchange of dense-water sinking down, and less dense-waters up. Like in the polar regions, such exchange can be observed daily in the upper 10 m from the sea-surface, regularly down to a few 100 m from the surface, and seldom, once every 5-8 years (e.g., Rhein, 1995; Mertens and Schott, 1998), down to the abyssal seafloor at about 2500 m. In contrast, horizontal density gradients

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associate with forcing of a dynamically unstable boundary current and eddies at multiple 1-100 km (sub-)mesoscales (e.g., Testor and Gascard, 2006). These eddy motions may push relatively warm waters down, thereby increasing the weak stratification in the deep-sea.

In summer, atmospheric disturbances are less intense, near-surface stratification is large due to solar heating, and eddy activity associated with some continental boundary flows is weaker (Albérola et al., 1995). This opens the possibility for detection of near-inertial wave dominance in kinetic energy. In relatively strong stratification, mainly gravity-driven parts of near-inertial waves generate largest vertical current differences 'shear' that destabilizes stratification due to their relatively short vertical length-scale, not only in the Mediterranean but also as observed in the Atlantic Ocean (van Haren, 2007). In near-homogeneous water-layers with weak stratification, their gyroscopic, Earth-rotation-driven, parts dominate and result in 0.1-1 km diameter sub-sub-mesocale tubes of slantwise rather than vertical convection (Straneo et al., 2002; van Haren and Millot, 2004). Hence, one may expect frequency spectra of non-tidal dominated data from instruments moored in the Mediterranean reveal convection and thus deep transport under winter and summer conditions.

It is noted that ocean-spectra may show peaks such as at narrowband tidal and at, broader band, inertial frequencies, but they lack gaps. This lack of spectral gaps potentially couples motions at sub-inertial with inertial-buoyancy internal wave with super-buoyancy turbulence frequency ranges. However, it is unclear how such a coupling may work as some motions represent two-dimensional (2D) eddies, some linear waves, some non-linear waves, some anisotropic stratified turbulence, and some isotropic 3D turbulence.

Kinetic energy (KE) spectra from historic current meter observations down to mid-depth z = -1100 m in the Ligurian Sea under upper-sea strongly stratified 'summer' and weakly stratified 'winter' conditions surely lack gaps (Fig. 1). Although these hourly sampled data barely resolve the turbulence ranges at frequencies higher than the buoyancy frequency, the internal wave continuum was suggested to scale with frequency ω like ω^p, with, on a log-

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log plot, 'spectral slope' p = -2.2±0.4 (van Haren and Millot, 2003), independent of location and season albeit with different KE (power) levels.

Within the uncertainty range, several possible explanations can be given for the observed spectral slope. Internal gravity waves have been fitted to p = -2±0.5 but only for f << ω << N (Garrett and Munk, 1972), where f denotes the inertial frequency involving Earth rotation and N denotes the buoyancy frequency reflecting the square-root of vertical density stratification. Considering that the data in Fig. 1 are from a site where locally N = 3±2f, irrespective of season (van Haren and Millot, 2003), alternative explanations were sought for observed spectral slopes at sub-inertial frequencies 0.2 cpd < ω < f. Cpd is short for 'cycles per day'. An obvious candidate is 'fine-structure contamination' of step functions passing sensors which gives a theoretical vale of p = -2 (Phillips, 1971; Reid, 1971). For their winter data, van Haren and Millot (2003) attributed such a slope to evidence intense mesoscale activity, because of the continuation of slope up to ω = 5 cpd before rolling off to white noise (slope 0). However, they did not elaborate. Below, the data in Fig. 1 are re-analyzed from a perspective of convection-turbulence.

Theoretical considerations of non-zero-mean flow convection-turbulence suggest a spectral scaling in the buoyancy range having p = -11/5 = -2.2 for KE, and p = -7/5 for a(n active) scalar quantity. This 'BO'-scaling follows atmospheric and theoretical works by Bolgiano (1959) and Obukhov (1959). The scaling was set-up for a stably stratified (atmospheric) environment for the anisotropic part in which turbulent kinetic energy is partially transferred to potential energy leading to turbulent convection. Later works extended BO-scaling to purely buoyancy-driven turbulence, e.g. for Rayleigh-Bénard convection (Lohse and Xia, 2010).

Laboratory experiments on such gravitationally driven convection are inconclusive on BO-scaling. This scaling is confirmed for both KE and temperature in experiments by Ashkenazi and Steinberg (1999), while only for scalars by Pawar and Arakeri (2016) who found a slope of p = -5/3 for KE. The p = -5/3-slope suggests dominance of shear-induced

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turbulence of the inertial subrange for equilibrium (isotropic) turbulence cascade in the 'KO'-scaling (Kolmogorov, 1941; Obukhov, 1949), but should also be found in spectra of scalars that are passive in this range. Obviously, scalars cannot be passive and active at the same time and space. This discrepancy between (types of) scaling between scalars and KE may be because the laboratory experiments of Pawar and Arakeri (2016) were in zero mean flow. Also, under sufficiently stable conditions without shear, no inertial subrange is expected (Bolgiano, 1959). However, the spectral extent of BO-scaling is largely unknown. While KO-scaling is based on a forward cascade of energy, the direction of energy cascade is inconclusive for BO-scaling and may be partially forward and partially backward, at least as reasoned for pure buoyancy-driven convection-turbulence (Lohse and Xia, 2010). Probably, directions of cascade change with locality in the flow, and perhaps depend on scale, which would also imply that KO- and BO-scaling cannot be found at the same site.

Revisiting data from non-zero mean flow and (weakly) stratified deep-sea in Fig. 1 demonstrates the possibility of fit of $p = -11/5$ outside near-inertial harmonic peaks. In winter, such a fit is observed consistently through the entire range of $0.2 < \omega < 5$ cpd. In traditional terms, this frequency range covers the transition from mesoscale $\omega < f$, via internal wave $f < \omega < N$, to turbulence $\omega > N$ motions. In summer, the $p = -11/5$-slope is found at two different KE levels for bands $0.2 < \omega < \omega_{min}$ and $2\Omega < \omega < 5$ cpd at sub- and super-IGW frequencies, respectively. Here, $\omega_{min} < f$ denotes the minimum frequency bound for inertio-gravity waves IGW (LeBlond and Mysak, 1978) and Ω the Earth rotational frequency. Maximum IGW is denoted by $\omega_{max} > 2\Omega$. The plotted IGW-bounds $[\omega_{min}\ \omega_{max}]$ are for weakly stratified, near-homogeneous layers in which $N = f$.

The bridge between the KE-levels at sub- and super-IGW is formed by the finitely broad near-inertial peak. The base of this peak is proposed to slope like $p = -1$ reaching super-IGW BO-scaling at about $\omega \approx 4$ cpd $\approx N$. Such $p = -1$-slope has been observed for the KE-spectral continuum between [f N] from the deep Bay of Biscay, Northeast Atlantic Ocean (van Haren et al., 2002). Theoretically, this slope represents spectral scaling of intermittency

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of a weakly chaotic nonlinear system (Schuster, 1984), i.e., 3D dynamical systems that evolve into self-organized critical structures of states which are minimally stable (Bak et al., 1987).

These observations suggest a dominance of convection cascade from sub-meso- via IGW- to, probably because unresolved, turbulence-scales under high-energetic winter-conditions. It is also observed under quieter summer conditions when, however, such a cascade is masked by IGW that lead a cascade at $\omega > \omega_{min}$. Especially the sub-inertial range of apparent BO-scaling seems out of the turbulence range, unless waters are near-homogeneous $N \rightarrow 0$ so that $\omega_{min} \rightarrow 0$. This would extend not only IGW, notably gyroscopic waves, but also turbulence, probably in the form of slantwise convection, to the (sub-)mesoscale range.

For the mesoscale range, the observations in Fig. 1 are supported by numerical modeling results that have suggested eddy-KE has a broad range of spectral slopes between $-3 < p < -5/3$ (Storer et al., 2022), and by satellite altimetry observations that indicated, after noise-correction and transfer to KE, a best-fit of $p = -2.28$ (Xu and Fu, 2012). No mention was made of BO-scaling, but the correspondence seems evident.

As the KE in Fig. 1 is at least one order of magnitude larger in winter than in summer, a near-inertial peak, if existent, will be part of the spectral continuum during the former. Saint-Guily (1972) proposed that winter-time inertial KE is spread over a broad featureless band, like quasi-gyroscopic waves may be present between IGW-bounds $[\omega_{min}, \omega_{max}]$ for $N \sim f$ (LeBlond and Mysak, 1978; Gerkema et al., 2008). However, observations from the year-round upper-layer-stratified central Western Mediterranean demonstrate that, also in deep homogeneous $N = 0$ waters, a near-inertial peak is observed in KE-spectra (van Haren and Millot, 2004). This may be attributed to a year-round source of atmospheric-generated inertial waves that are the only internal waves that can propagate without reflection from well-stratified to near-homogeneous layers and back (van Haren, 2023b).

Based on limited spectral observations, Gascard (1973) suggested the generation of 12-h stability waves (close to the buoyancy frequency of very weak stratification) that may briefly force dense-water formation, thereby implicitly suggesting a link between internal

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waves and (sub-)mesoscale eddies. As such eddies have estimated relative vorticity of $|\zeta| = f/2$ in the Western Mediterranean (Testor and Gascard, 2006), this addition to the planetary vorticity (f) automatically widens the 'effective' near-inertial band $0.5f < f_{eff} < 1.5f$, which bounds are close to IGW-bounds for N = 0.8f. One of the properties can be a modification of near-inertial frequency (Perkins, 1976), and trapping of near-inertial waves in anticyclonic eddies (Kunze, 1985). Although found to be limited to the rather flat KE-spectral dip in the immediate half-order-of-magnitude sub-inertial frequency band, standing vortical modes (low-frequency non-propagating motions) of vertical length-scale <10 m are suggested to be as energetic as internal waves (Polzin et al., 2003). Alternatively, it has been suggested for North-Atlantic observations that vortical modes may interact with internal waves, affecting internal-wave shear that was peaking over O(10) m vertical scales at IGW-frequencies in a band with limits determined by weak stratification as in N = f (van Haren, 2007).

For hypothetical ω_{min} = 0.2 cpd, at which the observed spectral slope changes away from p = -11/5 (Fig. 1), one would require N = 0.21f, which is almost unmeasurable and not existent for any prolonged period even in the deep Northwestern Mediterranean, to the knowledge of the author. However, it may reflect ω_{min} computed using f_{eff} = 0.5f and N = f_{eff}, noting that such conditions can only apply for part of the record. If so, it would reflect a direct coupling between sub-mesoscale and IGW-motions with slantwise convection (Straneo et al., 2002; van Haren and Millot, 2004; Gerkema et al., 2008). The p = -11/5 is significantly distinguishable from -2 over a frequency range of nearly two orders of magnitude, and from -5/3 over a range of just over half an order of magnitude (Fig. 1). The roll-off to noise (slope 0), for ω > 5 cpd, may partially be seen as following a slope of p = -5/3 before 0. The roll-off around 0.1 cpd suggests an unresolved broad mesoscale peak-value between 0.01 and 0.1 cpd. While these 1980's moored current meter data barely resolved the turbulence part of the KE-spectrum, their temperature sensors were too poor to simultaneously verify any spectral scaling for scalars.

About 40 years later, high-resolution moored temperature sensor ‘’T-’ data provided opportunity to verify scalar spectral scaling in the area. These T-data evidenced occasional warming of the deep Northwest Mediterranean seafloor (Fig. 2a), which, after comparison with data from higher-up appeared to be coming from above, or slanted sideways, under relatively stratified conditions (van Haren, 2023a). The data were collected during mid-fall, when waters above were sufficiently stratified and no (cooler) dense-water convection was formed. The broad two-day warming around day 308 is most stratified, whilst during other periods waters are only weakly stratified, including the quasi-inertial variations between days 316 and 322. These weakly stratified near-inertial, or near-buoyancy as $N \approx f$, temperature variations may evidence slantwise quasi-gyroscopic near-inertial waves, which can have a large vertical component (LeBlond and Mysak, 1978), as opposed to more common near-horizontal near-inertial waves in strongly stratified waters that are barely noticeable in temperature records.

The 18-day average spectra of the 2-s sampled data poorly resolve sub-mesoscales, but show a well-resolved slope of $p = -1.4\pm0.025$ between $0.5 < \omega < 6000$ cpd, across the IGW band and well into the turbulence band (Fig. 2b). No transition to a -5/3-slope is observed before roll-off to noise. The observed $p = -7/5$-slope is found significantly different from $p = -2$ and $-5/3$ over the indicated frequency range of four orders of magnitude and over the turbulence range between $100 < \omega < 10^4$ cpd. Over a frequency range of half an order of magnitude the slope-error is about ±0.1. Albeit not greatly resolved, the range between $\omega_{max} < \omega < 10$ cpd falls-off steeper roughly at $p = -2$ and the range between $10 < \omega < 100$ cpd shows a reduced variance that may partially be characterized by intermittency ($p = -1$; Schuster, 1984), but which is not yet explained. Here, it is observed to bridge between $p = -2$ and super-IGW BO-scaling $p = -7/5$. This would be further observation of a marginally ocean-state to the -1-scaling in KE-spectra (present Fig. 1 and van Haren et al., 2002) and in the continuum of the band [f N] in open-ocean T-spectra (van Haren and Gostiaux, 2009).

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Whilst more extended work with longer data sets is to be done, these high-resolution temperature observations suggest a direct coupling between sub-mesoscale motions, IGW motions, comprising internal gravity and gyroscopic waves, and convection turbulence. The temperature spectra are also consistent with the limited KE-spectra of Fig. 1 from roughly the same area, and both indicate a dominance of non-isotropic, stratified-turbulence convection between sub-mesoscales and largest turbulent overturning scales in extended BO-scaling suggesting cross-spectral coupling. The discrepancy with KE-spectra in laboratory experiments of Pawar and Arakeri (2016) may be due to difference of settings. In a non-zero mean flow turbulence convection experiment near the gas-liquid critical point, BO-scaling was observed for both KE and temperature (Ashkenazi and Steinberg, 1999). We recall that our deep-sea conditions are non-zero mean flow, weak tides, very high bulk Reynolds numbers O(10^5) given the large scales, and varying non-zero vertical density stratification.

The mesoscale-IGW-turbulence motions transport and mix warm waters downward. This contrasts with the process of buoyancy-driven dense-water formation that is thought to bring cooler waters downward during short periods of time, but for which no evidence exists in the 18-day T-sensor data set.

4 How robust is the system of ocean circulation and stratification?

Any variation to the nonlinear system of ocean circulation may encounter several complex feedback mechanisms, of which the effects are not yet fully understood for the present-day ocean. Although stable density stratification hampers vertical exchange by turbulent mixing, it does not block it. While stratification supports internal waves and their destabilizing shear, turbulent mixing during particular phase of a wave may decrease or destroy it locally in time and space. However, a subsequent internal wave-phase will restratify the mixed patch, thereby maintaining its own support of stable stratification. Such a feed-back system may be at work, for example when the ocean absorbs more heat.

Increased sea-surface temperature may lead to increased vertical density stratification, which may lead to less turbulent exchange as vertical overturning is suppressed.

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However, it will also lead to more internal waves through the extension of their spectral band to higher frequencies, with the potential to increased interaction, non-linearity, and turbulence-generating wave breaking. As particular internal waves can propagate deep into the ocean interior, they can cause enhanced turbulent mixing elsewhere.

Limited observations have thus far not provided evidence for an inverse correspondence between changes in turbulent mixing and changes in temperature across the near-surface photic zone along a longitudinal section of the Northeast Atlantic Ocean (van Haren at al., 2021b). This lack of correspondence suggests a feedback mechanism at work mediating potential physical environment changes so that global warming may not affect vertical turbulent fluxes of heat, and thereby also of, e.g., carbon. One such feedback mechanism may be convection-turbulence induced by internal waves and sub-mesoscale eddies. Re-analysis of moored current meter data from the Irminger Sea (North-Atlantic Ocean) demonstrate a significant p = -11/5 spectral slope at sub- and at super-inertial frequencies (Fig. 3). As was outlined in van Haren (2007), the area showed an IGW-band (for N = f) with dominant sub-inertial shear at small 8-m vertical scales despite the dominant internal tidal KE. The correspondence with the Mediterranean data of Fig. 1 is striking, including the one order of magnitude change in KE between sub- and super-IGW p = -11/5-slopes with similar p = -1 bridge albeit uncertain crossing level, and similar heights of near-inertial peak despite the tidal peak in Fig. 3.

While few ocean observations have been presented of BO-scaling thus far, coupling has not been established between convection and stratified small-scale turbulence with mesoscale motions. Likewise, complexing factors are spectral interruption by internal waves. However, internal wave trapping by mesoscale eddies has been well described (e.g., Kunze, 1985), and which thus provides an obvious coupling between these motions. It is expected that such coupling may lead to strong nonlinearity (of the internal waves) that leads to turbulent mixing produced by wave breaking.

As demonstrated using Mediterranean observations, not only convectively unstable cooler and/or saltier waters potentially lead to downward motions from the surface. Also sub-

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mesoscale eddies and near-inertial waves can convectively push stratified waters to the deep sea. Such downward push can be fast to transport materials from surface to 2500-m deep seafloor in a day (van Haren et al., 2006), and which speed is of the same order of magnitude as attributed to dense-water convection (Schott et al., 1996). It can also be more turbulent compared to shear-induced motions in the stratified ocean-interior, whereby turbulence reaches the seafloor according to few observations from the abyssal Pacific (van Haren, 2020) and alpine freshwater Lake Garda (van Haren and Dijkstra, 2021). Further extended observational evidence is urgently needed.

Although the anthropogenic influence on the Earth's climate is without doubt, the impact on the ocean circulation is not fully known because we lack sufficient, notably observational, information of the relevant processes that can thus not be properly modeled yet. Therefore, we should be cautious in making predictions such as in (e.g., Ditlevsen and Ditlevsen, 2023; van Westen et al., 2024) on future ocean circulation based on single parameters like ocean-surface temperature or fresh-water flux that are uncertain proxies. Because no observational (van Haren et al., 2021b), modeling (Little et al., 2020) or paleo-proxy validation (Cisneros et al., 2019) physics evidence exists that sea-surface temperature is a solid estimator of AMOC-strength variations, other properties like vertical density gradients (stratification), and turbulence intensity may be considered.

Variability of the ocean in space and time is a key to its dynamics, but it is unclear how robust such variations can be, e.g., whether shifting sites for deep dense-water formation (Gou et al., 2024) may be part of the same system. Observational evidence verifying numerical simulations' outcome, not only predictions but also present-day, of ocean-state is needed. Observations are also required to evidence variability in relevant physics processes for model-implementation. Besides eddies and coupling with atmosphere (e.g., Gent, 2018), numerical models of complex nonlinear ocean circulation should contain internal-wave turbulence with appropriate space and time dependency.

As for the ocean circulation in the horizontal plane near its surface with most impact on mankind, wind will remain the main driver rather than the AMOC. As long as the Earth

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rotation does not alter direction, wind will maintain its general course (Wunsch, 2004). The atmosphere remains the key player in the global heat transport across mid-latitudes rather than the ocean. Simultaneously, the importance of processes like stratification and turbulent mixing induced by, e.g., internal wave breaking with or without sub-mesoscale coupling cannot be underestimated for life near the ocean-surface as well as in the -deep, because it will come to a halt without such processes.

Data availability. No new data were created or analyzed in this study: replot and re-analysis of data presented in van Haren and Millot (2003), in van Haren (2007) and in van Haren (2023a).

Competing interests. The author declares that he has no conflict of interest.

Acknowledgments. I thank L. Gerringa for commenting a previous draft of the manuscript.

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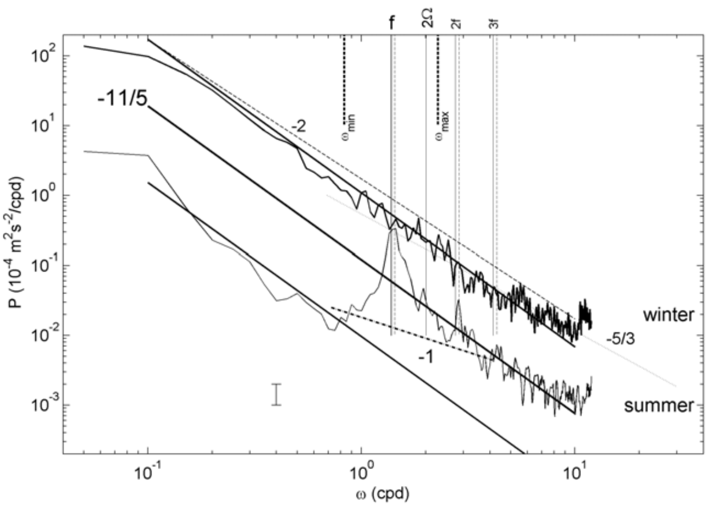

Fig. 1. Moderately smoothed (20 degrees of freedom, dof) kinetic energy (KE) spectra over 100 days of data from 600-s sampled Aanderaa mechanical current meter moored in 1981/1982 at z = -1100 m over the continental slope in the Ligurian Sea at 43° 28.32′ N, 7° 46.10′ E, 2250 m water depth. For details on these data, see van Haren and Millot (2003). The 'summer' spectrum is an average from data between days 190 and 290 (in 1981), the 'winter' between days 375 and 475 (adding +365 for days in 1982). Several frequencies are indicated including inertial frequency f, Earth rotational Ω and inertio-gravity wave bounds [ω_{min}<f, ω_{max}>N] for buoyancy frequency N = f. The dashed lines indicate (harmonics of) 1.04f. Four spectral slopes ω^p are indicated by their exponent: p = -11/5 (solid slope in the log-log plot) for Bolgiano-Obukhov 'BO' scaling reflecting the buoyancy subrange of convective turbulence (e.g., Pawar and Arakeri, 2016), p = -5/3 (dotted slope) for Kolmogorov-Obukhov 'KO' scaling reflecting the equilibrium inertial subrange for dominant shear-induced turbulence (Kolmogorov 1941; Obukhov, 1949), p = -1 (dash-dotted slope) for intermittency of self-organized criticality (Schuster, 1984; Bak et al., 1987) and p = -2 (dashed slope) for internal wave scaling (Garrett and Munk, 1972) or finestructure contamination (Phillips, 1971; Reid, 1971).

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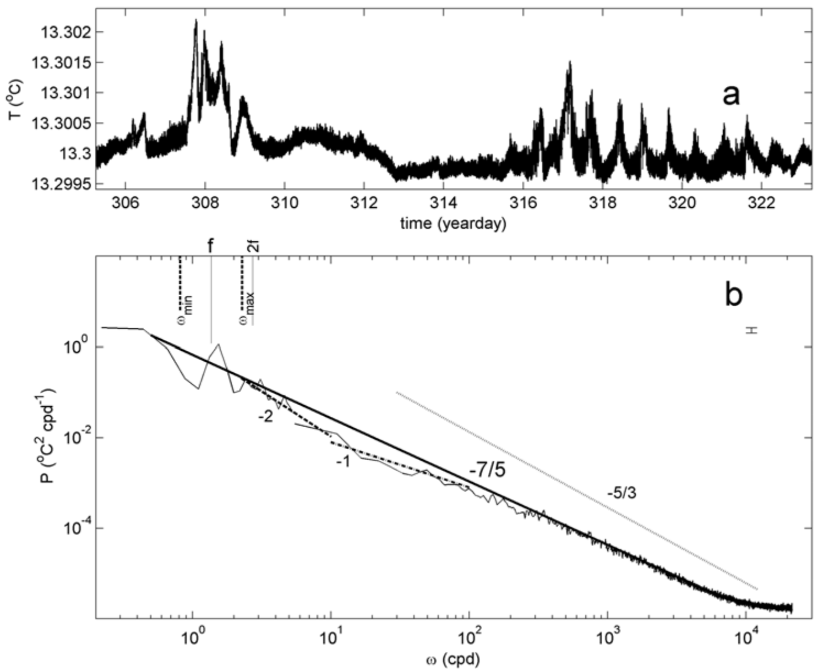

Fig. 2. Eighteen days of high-resolution 2-s sampled temperature (T) data from a NIOZ T-sensor fallen off a mooring-line in 2020, and lying 0.01 m above a flat seafloor about 10 km south of the foot of the continental slope at 42° 49.50′ N, 6° 11.78′ E, 2458 m water depth, about 100 km WSW from the site in Fig. 1. For details on these data see van Haren (2023a). (a) Time series of 18 days of raw temperature data. (b) Weakly smoothed (10 dof; ω<5 cpd) and heavily smoothed (250 dof; ω>5 cpd) temperature variance spectra of data in a. Frequency and spectral slope indications as in Fig. 1, while -7/5 (solid slope) indicates BO-scaling of an active scalar (e.g. Pawar and Arakeri, 2016), and -1 (dash-dotted slope) for scaling of intermittency of a weakly chaotic nonlinear system (Schuster, 1984). Note the different axes-ranges compared with Fig. 1.

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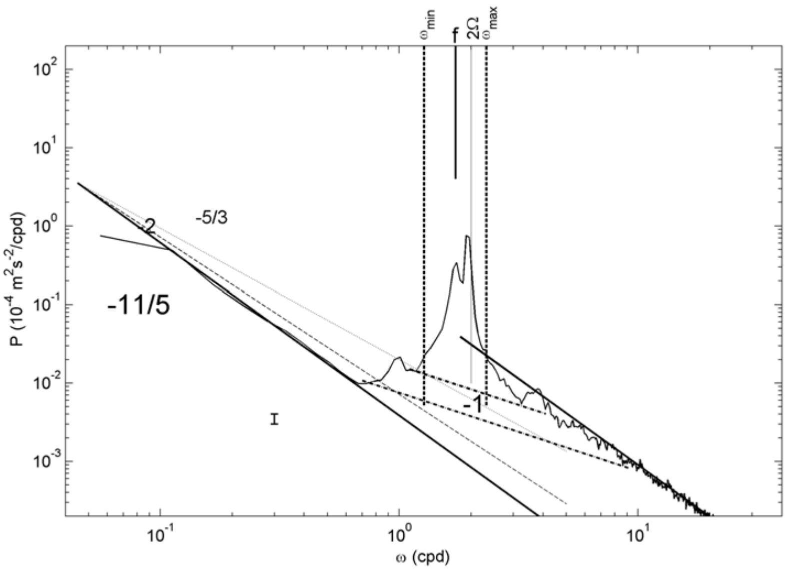

Fig. 3. Like Fig. 1, but for strongly smoothed (50 dof) KE spectra averaged over 400 days of data from 600-s sampled Valeport mechanical current meter moored at z = -1000 m over the Mid-Atlantic Ridge at 58° 59.67′ N, 33° 56.12′ W, 2540 m water depth in 2003/2004, within the project discussed in van Haren (2007).