On the role of moist and dry processes for atmospheric blocking biases in the Euro-Atlantic region in CMIP6

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Abstract. Synoptic- and large-scale features such as extratropical cyclones, Rossby wave packets, and atmospheric blocking modulate the mid-latitude weather and climate. However, several studies have shown strong biases in the frequency of these features in state-of-the-art global climate models. One notable and persistent bias is an underestimation of the atmospheric blocking frequency in the Euro-Atlantic region. In this study, we validate the representation of synoptic- and large-scale features
5 of the North Atlantic flow in eight climate models of the Coupled Model Intercomparison Project 6 (CMIP6), taking the ERA5 reanalysis as a reference. Validation includes atmospheric blocking, storm tracks, eddy heat and moisture fluxes, and warm conveyor belts (WCBs).

The selected CMIP6 models underestimate the atmospheric blocking frequency over the eastern North Atlantic and Europe in winter (December to February) by up to 80%. The frequency biases result from combined biases at different spatial and
10 temporal scales described in the following. First, we define the background flow as the most frequent value of the latitudinal gradient of the geopotential at 500hPa. In the CMIP6 models, the strongest latitudinal geopotential gradients are equatorward shifted in the North Atlantic basin. This shift favours more zonal and stronger winds to the south of the climatological jet. The differences in the background flow affect Rossby wave breaking and blocking onset and persistence, as illustrated by analysing the eddies in the Euro-Atlantic region. We find an equatorward shift in the eddies in CMIP6 that accelerates the
15 mean flow in the exit region of the Atlantic jet, as indicated by a reduction of the divergence of E-Vectors. The shift in the eddies leads to a less diffluent flow in the east Atlantic and, thus, a less favourable flow for blocking formation. Second, we find a negative bias in WCB outflow frequency in the CMIP6 models in the North Atlantic. Reduced WCB outflow indicates weaker transport of low potential vorticity (PV) from the lower to the upper troposphere by moist diabatic processes and consequently weaker downstream ridge amplification and, therefore, a less diffluent flow over the eastern Atlantic and weaker
20 diabatic contributions to blocking. The negative WCB outflow bias can be linked to an underestimation of the meridional moisture transport at low levels in the climatological WCB inflow area in the western Atlantic. Thus, the misrepresentation of moist processes contributes to the negative blocking biases. Accordingly, an improved representation of the moist processes in the next generation of climate models could improve the blocking representation.

1 Introduction

25 Atmospheric blocks are quasi-stationary and persistent anticyclones often associated with up- and/or downstream Rossby wave breaking (Altenhoff et al., 2008; Trevisiol et al., 2022). Blocks may lead to high-impact surface weather (see Kautz et al., 2022, for a review). In summer, blocks can be associated with heatwaves (Jeong et al., 2022), wildfires (Antokhina et al., 2023), and thunderstorms (Mohr et al., 2019) and in winter, blocks can be associated with cold spells (Brunner et al., 2018).

Upstream eddies and wave breaking are essential for amplifying and maintaining blocks (Shutts, 1983; Woollings et al.,
30 2018). In recent years, research has shown that moist processes also play a major role in the initiation, amplification, and maintenance of blocking (Pfahl et al., 2015; Steinfeld and Pfahl, 2019; Steinfeld et al., 2020). For example, latent heat release upstream greatly contributed to the maintenance of the block and the associated heat wave in June 2021 in Canada (White et al., 2023). Pfahl et al. (2015) followed a Lagrangian perspective to analyse the development of blocking and showed that about 50% of the air parcels ending at the upper-level anticyclone undergo diabatic processes. Moist diabatic processes contribute
35 to block development prominently in the western North Atlantic region in winter. At the same time, moist diabatic processes are an important contributor to blocking formation and intensification throughout the North Pacific and parts of eastern North America (Steinfeld and Pfahl, 2019). In contrast, contributions from latent heat release are small ($< 30\%$) to blocks over the Eurasian continent (Steinfeld and Pfahl, 2019). The impact of moist diabatic processes on large-scale circulation might increase as the atmosphere warms. For instance, extratropical cyclones may respond to the increased latent heat release with
40 faster ascending warm conveyor belts (WCBs) (Binder et al., 2023; Dolores-Tesillos and Pfahl, 2024), which may impact the block properties. To capture the effects of climate change on atmospheric blocking both dry and moist dynamical processes thus need to be accurately captured by climate models.

In general, Earth System Models (ESM) still struggle to represent blocks accurately (Schiemann et al., 2017, 2020). There
45 are significant biases in the representation of blocking properties and frequency (Davini and D'Andrea, 2016; Schiemann et al., 2017, 2020; Eyring et al., 2020). A recent overview of blocking biases in various climate models (Woollings et al., 2018) summarises some known sources of the biases. They include:

- Errors in the mean state or background flow that affect the waviness and strength of the jet and thereby Rossby wave propagation (Berckmans et al., 2013; Kleiner et al., 2021). Indeed, the representation of the mean state can be one of the
50 leading causes of blocking bias (Branstator, 2002; Branstator and Teng, 2017; Wirth et al., 2018; Kleiner et al., 2021). Luo et al. (2023) show how the meridional gradient of the background of potential vorticity (PV) modifies the lifetime of blocks.

- Sea surface temperature biases have been linked to block frequency biases. Specifically in the North Atlantic, a negative sea surface temperature (SST) bias is found in several climate models (Scaife et al., 2011; Athanasiadis et al., 2022). The
55 colder ocean increases the low-level temperature gradient, which results in a stronger low-level baroclinicity and reduced surface evaporation (Scaife et al., 2011; Athanasiadis et al., 2022; Cheung et al., 2023; Wenta et al., 2024). Other studies show that blocking frequency is also sensitive to the tropical SST biases (e.g. Hinton et al., 2009).

– Other studies suggest a misrepresentation of the tropopause height as a cause for blocking biases; however, this link is found when applying anomaly block indices (Attinger et al., 2021).

According to Woollings et al. (2018), finer horizontal and/or vertical resolution can help reduce the blocking biases mentioned above. The finer horizontal resolution could improve blocking representation by better-resolving eddies and also a better representation of the orography, resulting in more realistic planetary waves and mean state (Berckmans et al., 2013; Davini et al., 2017). Woollings et al. (2018) also mentioned the relevance of increasing the vertical resolution, affecting the dynamics of the tropopause and the ascent of airstreams. Improving convection and drag parameterisations could also reduce blocking biases (Woollings et al., 2018; Martínez-Alvarado et al., 2018).

One aspect that has received limited attention is the moist processes. Misrepresentation of moist processes generates large-scale circulation biases in current climate models. For example, the underestimation of latent heating has been mentioned as the primary source of bias in stationary waves in CMIP5 models (Park and Lee, 2021). Idealised experiments have shown that increasing latent heat release in cyclones helps reduce biases in large-scale flow representation by improving the often too-zonal storm tracks and too-weak cyclones. A better representation of latent heating can be achieved by increasing horizontal resolution on a scale of storm resolution (Schemm, 2023).

Latent heat release through moist diabatic processes is very intense in WCBs. The LH release in WCBs is associated with cloud formation processes, and its maximum is usually located downstream to the extratropical cyclone centre, where heavy precipitation is occurring (Madonna et al., 2014). The moist diabatic processes in a WCB generate anticyclonic PV anomalies in the WCB outflow at the tropopause level. The anticyclonic PV anomaly influences the extratropical waveguide and downstream atmospheric flow (e.g. Grams et al., 2011; Binder et al., 2016; Wirth et al., 2018). Thus, the WCBs may contribute to blocking formation and maintenance (Steinfeld and Pfahl, 2019; Steinfeld et al., 2020). However, the evaluation of WCBs in climate model data is not straightforward. Identifying WCBs traditionally relies on calculating air parcel trajectories (Wernli and Davies, 1997; Madonna et al., 2014). For the trajectory calculation, 3D wind fields are necessary that are not always available from ESMs, and the calculations require high computational resources (Joos and Wernli, 2012; Madonna et al., 2014; Sprenger and Wernli, 2015; Binder et al., 2016; Schielicke and Pfahl, 2022). A novel and alternative approach is to apply deep learning to identify WCBs such as the EuLerian Identification of ascending AirStreams (ELIAS2.0) diagnostic (Quinting and Grams, 2022; Quinting et al., 2022). ELIAS2.0 is a deep learning tool that predicts the probability of WCB inflow, ascent and outflow on the basis of atmospheric state variables and derivations thereof. ELIAS2.0 has been successfully implemented in ERA5 and for numerical weather prediction data (Oertel et al., 2023b). Here, we apply ELISA2.0 for the first time to climate model data. Validating the representation of WCBs in the Coupled Model Intercomparison Project 6 (CMIP6) models contributes to the assessment of the role of moist diabatic processes in the blocking biases.

The aim of this study is thus, for the first time, to estimate to which extent the representation of the background flow, storm tracks, eddy fluxes and WCBs contribute to the known Euro-Atlantic blocking biases in the CMIP6 models. In Sect. 2, we describe the dataset used for the analysis. The method is presented in Sect. 3. We will show the results and discuss them in Sects. 4 and 5. Finally, the conclusions are provided in Sect. 6.

2 Data

We select a sub-set of CMIP6 models based i) on the models' ability to represent the climatological atmospheric blocking occurrence frequency in the Euro-Atlantic region as quantified by Palmer et al. (2023) and ii) on data availability needed
95 to identify the various synoptic-scale features. We ensure that for each model, more than 30 years of data are available (on average, the historical simulations cover a period of 164 years) and that the data include the below-described variables. We aim for some model independence by considering models from different research institutes and model cores (Brunner et al., 2020). Of the possible qualitative levels for model performance indicated by Palmer et al. (2023), we analyse only models with adequate performance, excluding the inadequate models.

100   A model overview is shown in Table 1. We use daily data remapped to $1°x1°$ of geopotential height (Z), the horizontal wind components (u and v), temperature (T), and specific humidity (q) at the following pressure levels: 1000, 850, 700, 500, 300, 250, 200 hPa. We focus on the boreal winter (December to February). The historical simulations cover the period from 1979 to 2014 and were retrieved from the DKRZ data pool (Eyring et al., 2016).

Model name	Institution	Member	Blocking DJFM	Nominal resolution
MRI-ESM2-0	MRI (Japan)	r1i1p1f1	adequate (s)	100 km
ACCESS-CM2	CSIRO-ARCSS (Australia)	r1i1p1f1	adequate (s)	250 km
EC-Earth3	EC-Earth-Consortium	r1i1p1f1	adequate (s)	100 km
MPI-ESM1-2-HR	MPI (Germany)	r1i1p1f1	adequate (s)	100 km
CESM2-WACCM	NCAR (USA)	r1i1p1f1	adequate (s)	100 km
MIROC6	MIROC (Japan)	r1i1p1f1	adequate (u)	250 km
MPI-ESM1-2-LR	MPI (Germany)	r1i1p1f1	adequate (u)	250 km
CESM2	NCAR (USA)	r11i1p1f1	adequate (u)	100 km

Table 1. List of CMIP6 models and their performance. The ability of models to reproduce Atmospheric blocking in the Euro-Atlantic region is shown from Palmer et al. (2023). Qualitative criteria are adequate (satisfactory (s), unsatisfactory (u)), inadequate, and not available. Models with the same colour indicate models with dependencies, for example, shared components (Brunner et al., 2020).

We evaluate the CMIP6 models against ERA5 reanalysis (Hersbach et al., 2020). The same variables as for CMIP6 were
105 remapped to the exact spatial resolution $(1°x1°)$ from 1979 to 2014. Note that the grid spacing $1°x1°$ is chosen to identify CMIP6 WCBs in ELIAS2.0 (similar to ERA5 WCBs).

3 Methods

3.1 Blocking index

Instantaneous blocks (IB) are identified based on the reversal of the flow (Dole and Gordon, 1983; Tibaldi and Molteni, 1990;
110 Scherrer et al., 2006; Davini et al., 2012; Davini and D'Andrea, 2016; Brunner and Steiner, 2017; Marco Rohrer and Wild,

2019; Rohrer et al., 2018). We follow the methodology of Brunner and Steiner (2017) that consists of the calculation of geopotential height gradients to the north of a central latitude ϕ (ΔZ_N), to the south of ϕ (ΔZ_S) and a third zonal gradient towards the Equator to remove low-latitude blocking (ΔZ_E):

$$\Delta Z_N\left(\lambda, \phi\right) = \frac{Z\left(\lambda, \phi + \Delta\phi\right) - Z\left(\lambda, \phi\right)}{\Delta\phi} \tag{1}$$

$$\Delta Z_S\left(\lambda, \phi\right) = \frac{Z\left(\lambda, \phi - \Delta\phi\right) - Z\left(\lambda, \phi\right)}{\Delta\phi} \tag{2}$$

$$\Delta Z_E\left(\lambda, \phi\right) = \frac{Z\left(\lambda, \phi - 2\Delta\phi\right) - Z\left(\lambda, \phi - \Delta\phi\right)}{\Delta\phi} \tag{3}$$

where for the Northern Hemisphere, the IB follows the next three conditions:

$\Delta Z_N\left(\lambda, \phi\right) <$ -10 m ($°$lat.$^{-1}$)

$\Delta Z_S\left(\lambda, \phi\right) < 0$ m ($°$lat.$^{-1}$)

$\Delta Z_E\left(\lambda, \phi\right) > 5$ m ($°$lat.$^{-1}$)

λ is longitude and goes from 180° W to 179° E and ϕ is latitude and goes from 75° S to 75° N. Gradients are computed at each latitude point with $\Delta\phi$ = 15°.

We select only large-scale blocks by filtering out IBs smaller than 15° longitude. To ensure quasi-stationarity and persistence, we consider only blocks lasting five days or more in a box of 6° x 4° longitude and latitude. This method captures blocks in their mature stage, featuring a Rossby Wave Breaking (RWB).

3.2 Background flow

Defining the background flow is not straightforward (e.g. Wirth et al., 2018). Often, a time mean is used to represent the background flow. However, Swanson (2001) shows that the frequency distributions of upper-tropospheric Z values are positively skewed poleward of the midlatitude jets and negatively skewed equatorward. They, therefore, recommend using the mode (most frequent value) rather than the time mean to describe the background flow. We calculate the mode of Z and that of the absolute meridional gradient of Z ($|\frac{dZ}{dy}|$) for each grid point to capture the background flow. Bins of 5 $m\,km^{-1}$ for the $|\frac{dZ}{dy}|$ time series (bins of 50 m for the Z time series) of each grid point are computed. After that, we calculate the mode for each model. We focus on the gradient of Z rather than the absolute values of Z because the gradient values are more closely associated with the position of the zonal jet, and blocks are identified using a gradient criterion.

3.3 Storm tracks and eddy fluxes

We apply a bandpass filter (2 and 6 days) to different fields (Z, u, v, T, q) to capture synoptic weather variability (Hoskins and Hodges, 2002; Greeves et al., 2007; Davini et al., 2017). The filtered variables are hereafter indicated by prime ′. Storm tracks

are identified using the standard deviation of the bandpass filtered daily Z at 500 hPa (Z'). We quantify the influence of the eddies on the mean flow by computing the divergence of the E-Vector (Hoskins et al., 1983; Trenberth, 1986; Raible et al., 2010). We calculate the E-vector from the bandpass filtered u and v at 250 hPa following Hoskins et al. (1983):

$$\mathbf{E} = (\overline{v'^2 - u'^2}, \overline{-u'v'}) \tag{4}$$

E-vectors tend to have an eastward direction of the wave energy propagation, and $\nabla{\cdot}\mathbf{E}$ provides a measure of the eddies forcing of the mean flow, divergent E-vectors indicate an acceleration of the mean flow by the eddies. In contrast, the convergence of the E-vectors indicates a deceleration of the mean flow by the eddies (Hoskins et al., 1983; Trenberth, 1986; Strong and Magnusdottir, 2008; Cheung and Zhou, 2015). E-vectors are meridionally elongated upstream of blocks and converge around the blocked area (Berckmans et al., 2013; Drouard et al., 2021). During a European blocking, a poleward-shifted jet stream is characterized by an anomalous divergence of E-vectors to the north of the climatological jet. The displaced jet is forced into its position by the eddy forcing, which also helps to maintain the blocks by keeping the jet in its position (Berckmans et al., 2013).

Finally, we also calculate eddy moisture fluxes ($\overline{v'q'}$) from the bandpass filtered daily specific humidity and the bandpass filtered meridional wind at 850 hPa and eddy heat fluxes ($\overline{v'T'}$) from bandpass filtered daily temperature and bandpass filtered meridional wind at 850 hPa.

3.4 Warm conveyor belt identification

We use ELIAS2.0 to detect WCBs. ELIAS2.0 is a deep learning tool, which has been shown to work in reanalysis data and operational models (Quinting and Grams, 2022; Quinting et al., 2022). ELIAS2.0 has been employed to analyse extreme events (Oertel et al., 2023b), to investigate microphysical processes in WCBs (Oertel et al., 2023a), and to study the role of Madden–Julian Oscillation in the occurrence of North Pacific WCB (Quinting et al., 2024).

ELIAS2.0 is trained on the basis of ERA-interim (Dee et al., 2011). The predictors are derived from T, q, Z, u and v on a limited number of pressure levels. Using a U-Net convolutional neural network, ELIAS2.0 outputs conditional probabilities of WCB inflow, ascent, and outflow. WCB masks in ERA5 are obtained at a frequency of 6 hours [0, 6, 12, 18 hours]. We compute WCB masks in the CMIP6 models from daily data. Therefore, we select only the 12 UTC timestamp of ERA5 to compute the WCB reference climatology.

4 Results

We first compare the blocking frequencies in the CMIP6 models against ERA5 to identify the largest and most significant biases; we then link these biases to biases in the background flow, dry processes, and moist processes.

4.1 Atmospheric blocking and the background flow

The biases in atmospheric blocking frequency are shown in Figure 1. The contours show the climatological blocking frequency for ERA5 and the shaded colours the difference between each CMIP6 model and the ERA5 (Fig. 1a-h), and between the CMIP6 multi-model mean and ERA5 (Fig. 1i). By analysing the biases in each CMIP6 model, one can identify that the underestimation

170 of the Euro-Atlantic blocking is present in all models (Figs. 1a-i). The strongest negative biases are found in MPI-ESM1-2-HR, MPI-ESM1-2-LR, and CESM2 (Figs. 1d,g,h). On the contrary, MRI-ESM2-0 shows the lowest negative biases (Figs. 1a). Note that in the analysis of the CMIP6 performance by Palmer et al. (2023), the MRI-ESM2-0 (compared to the other models selected here) exhibits the lowest blocking bias in the Euro-Atlantic region. Ural blocks are also underestimated in most models except for MIROC6 and CESM2. However, there is a disagreement over the Greenland region. MPI-ESM1-2-LR, MPI-ESM1-2-

175 HR, MIROC6, MRI-ESM2-0, and ACCESS-CM2 simulate more Greenland blocking than the ERA5 climatology. In contrast, CESM2, CESM2-WACCM, and EC-Earth3 underestimate Greenland block frequency. In the North Pacific, most of the models simulate fewer blocks, but the location varies; for instance, CESM-WACCM underestimates blocks to the south of the main North Pacific blocking region, while MRI-ESM2-0 underestimates blocks to the north of the main North Pacific blocking region. We find a robust underestimation of the Euro-Atlantic and Ural blocks in the CMIP6 multi-model mean bias (Fig. 1 i).

180 The selected CMIP6 models underestimate the Euro-Atlantic blocking frequency by more than 80% over the United Kingdom and parts of Scandinavia (Fig. 1). Another region of underestimation of blocking frequency is over the East Siberian Sea (Figs. 1 i and 2 e). Lower blocking biases of positive signs are found over Baffin Bay, Alaska, Northern Kazakhstan, Barents Sea, and Kara Sea (Fig. 1 i). Our results are in line with previous studies looking at both CMIP5 and CMIP6 models (Masato et al., 2013; Davini et al., 2014; Davini and D'Andrea, 2016; Davini et al., 2017; Woollings et al., 2018; Davini and D'Andrea, 2020;

185 Trevisiol et al., 2022), that find a negative blocking frequency bias over the Euro-Atlantic region. Contrary to previous studies using CMIP6 models, the negative frequency biases do not extend into central Europe and are smaller in the North Pacific region (e.g. Davini and D'Andrea, 2020). This improvement in the blocking frequency may be due to selecting models with the lowest bias (see discussion in the data section).

The Euro-Atlantic region shows the most considerable blocking bias. Thus, we focus only on this region when analysing the

190 biases in background flow, eddy transient activity, and WCB outflow. Biases are obtained for each climate model. However, we present the multi-model CMIP6 mean biases for concision. In the following, the biases of CMIP6 models in the background flow are analysed. First, the mean and mode of Z for ERA5 are shown in Figure 2a,b. We identify a closer spacing of Z contours when considering the mode rather than the mean value (Fig. 2d).
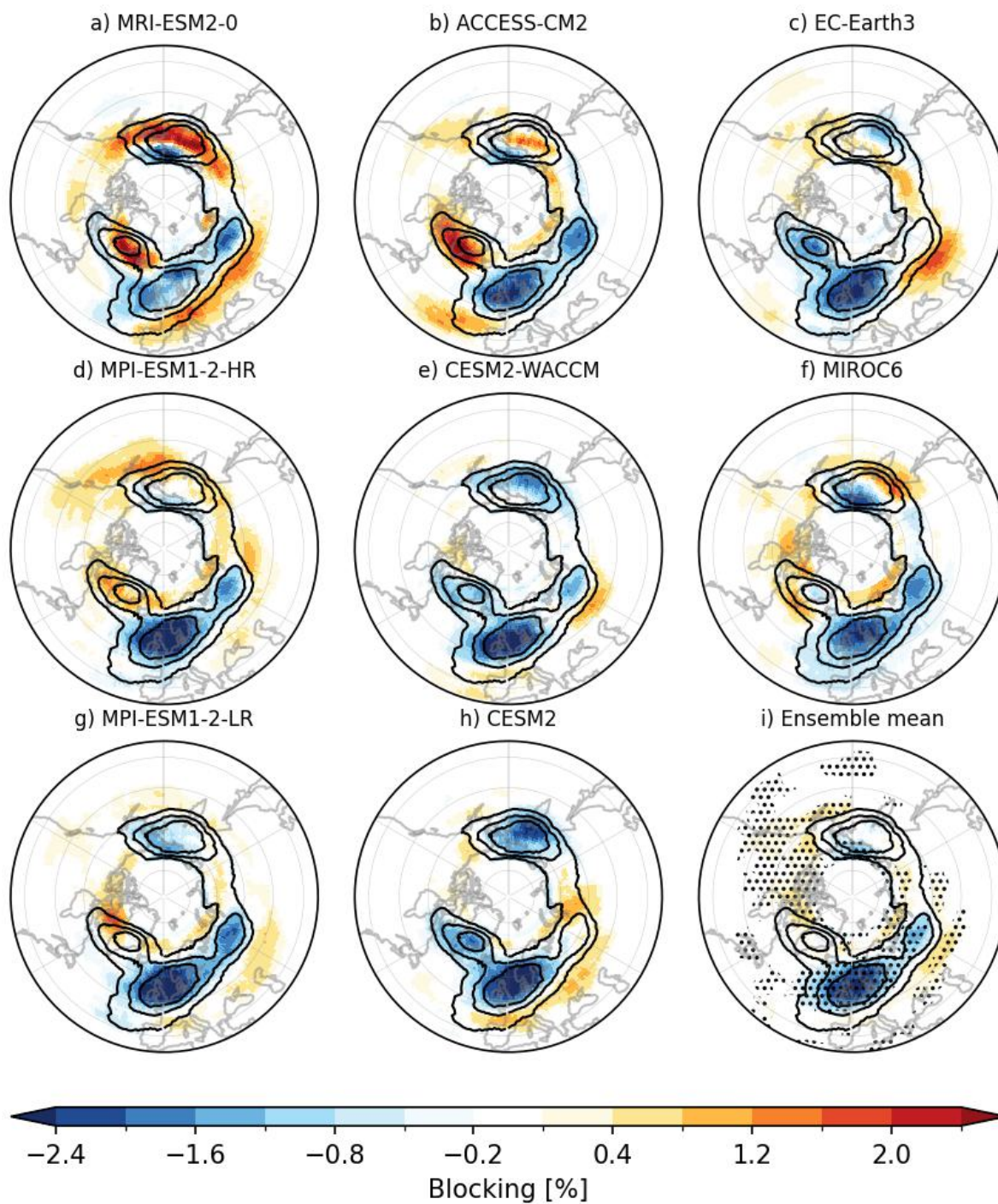
Figure 1. a-h) Blocking frequency biases of individual CMIP6 models and i) multi-model mean against ERA5 (1979-2014). The ERA5 blocking frequency is indicated by contours (1% intervals starting at 1%). Stippling denotes regions where more than 80% (6 models) of the ensemble members indicate a bias of the same sign.
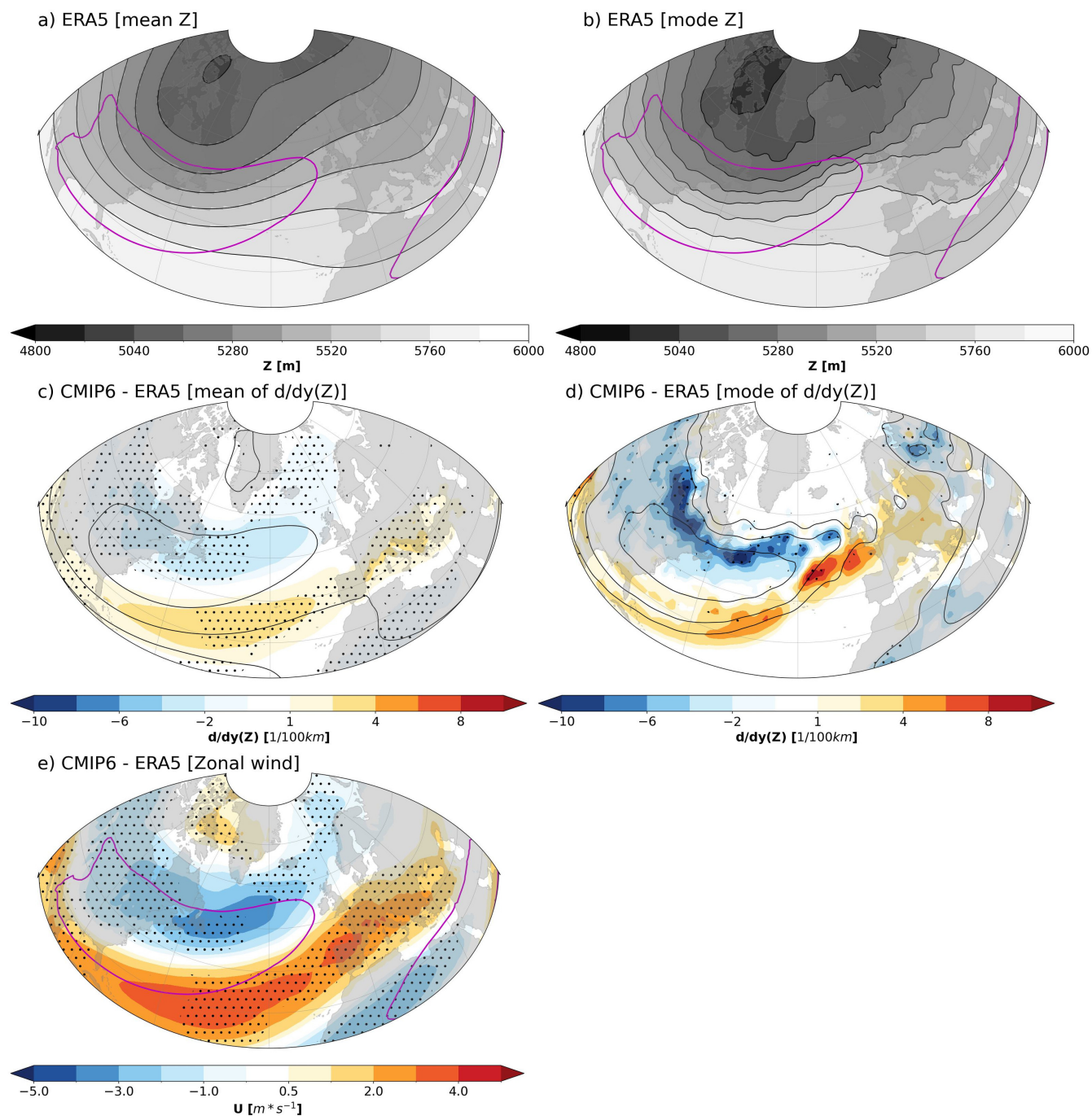
Figure 2. a,b) Climatological a) mean and b) mode of Z at 500 hPa for ERA5. The zonal wind at 500 hPa is shown as a purple contour ($15\ m\,s^{-1}$). c,d,e) CMIP6 multi-model mean biases of c) mean $\left|\frac{dZ}{dy}\right|$, d) mode of $\left|\frac{dZ}{dy}\right|$, and e) zonal wind at 500 hPa against ERA5. Black contours depict the ERA5 $\left|\frac{dZ}{dy}\right|$ climatology in c) and d) (5, 10, 20, 30, 40 $m\,km^{-1}$ intervals). Stippling denotes regions where more than 80% (6 models) of the ensemble members indicate a bias of the same sign.

195   Using the mode results in sharper gradients in the mid-latitudes as Swanson (2001) described. Note that the strong gradient of Z follows the location of the midlevel jet (zonal wind at 500 hPa, purple contour in Fig. 2a,b). As discussed in the methods, we chose to look at the CMIP6 $|\frac{dZ}{dy}|$ biases for understanding biases in the background flow due to the gradient index used to identify blocks. Thus, we show the CMIP6 biases of mode and mean $|\frac{dZ}{dy}|$ in Figure 2c,d. Again, note that there are sharper gradients in the mode of $|\frac{dZ}{dy}|$ (contours in Fig. 2 b,c). Both the mean and mode of $|\frac{dZ}{dy}|$ indicate an equatorward shift of strongest $|\frac{dZ}{dy}|$ in the North Atlantic (Fig. 2 c,d). The equatorward shift of $|\frac{dZ}{dy}|$ is more pronounced downstream of the ERA5

200   climatological maximum $|\frac{dZ}{dy}|$ in the north Atlantic, indicating a too-zonal background flow. The biases in the mode of $|\frac{dZ}{dy}|$ also indicate an eastward shift of the strongest gradient of $|\frac{dZ}{dy}|$ in the North Atlantic reaching the western European countries.

   In general, the background flow in the CMIP6 models may favour a shifted equatorward and too-zonal jet and is confirmed by the biases in the zonal wind at 500 hPa (Fig. 2 e). Over the eastern Atlantic, the biases of the mode of $|\frac{dZ}{dy}|$ hint at an eastward extension of the jet compared to the climatological ERA5. The mode hints at a strong jet over Western Europe (e.g. the UK),

205   while the mean shows a jet extension to South Europe (Fig. 2 d). A stronger $|\frac{dZ}{dy}|$ (stronger jet) in Western Europe coincides with the low frequency of blocking simulated by the CMIP6 models (Fig. 1 i). We can find the strongest jet (positive bias in mode/mean $|\frac{dZ}{dy}|$) over the eastern Atlantic for individual CMIP6 models that show the highest blocking underestimation in the Europe-Atlantic region (e.g. CESM2, not shown). Negative biases of mode and mean $|\frac{dZ}{dy}|$, thus, weaker gradients of background flow are found to the northeast of the USA. The less intense jet may be associated with slightly overestimated

210   blocking frequency in North America (Fig. 1 i).

4.2 Transient wave activity and the role of dry processes

High-frequency eddies are crucial for the formation and maintenance of blocking; the associated PV fluxes contribute to the amplification and maintenance of the blocking anticyclones (e.g. Hoskins et al., 1983; Shutts, 1983; Zappa et al., 2014; Cheung and Zhou, 2015; Woollings et al., 2018). Therefore, we next analyse the CMIP6 biases in storm track positions, transient wave

215   activity, and blocking occurrence.

   The Eulerian storm track diagnostic captures the main storm track regions, with the maxima in the western Atlantic (Fig. 3a). The main storm tracks are found on the poleward side of the climatological jet (purple contours in Fig. 3a). The storm track biases in the CMIP6 models are shown as shaded colours in Fig. 3c. CMIP6 models show an equatorward shift of the storm tracks in the North Atlantic region. We also observe an overestimation of the magnitude of storm tracks over the Mediterranean

220   and Eastern Europe. Similar storm track biases have been described for other CMIP6 and CMIP5 models based on Lagrangian metrics (Priestley et al., 2023; Zappa et al., 2013) and Eulerian metrics (Harvey et al., 2020). Priestley et al. (2023) identified an underestimation of cyclone frequency over the Mediterranean region that is not captured in our analysis. This discrepancy may be due to our model selection or the Lagrangian approach. Even when we can not make quantitative statements due to the feedback between background flow and storm tracks, the biases in storm tracks are consistent with background flow biases

225   (Fig. 1 b,c). For example, the more equatorward and zonal passage of extratropical storms in the North Atlantic aligns with the equatorward displacement of the most substantial gradients in the background flow.
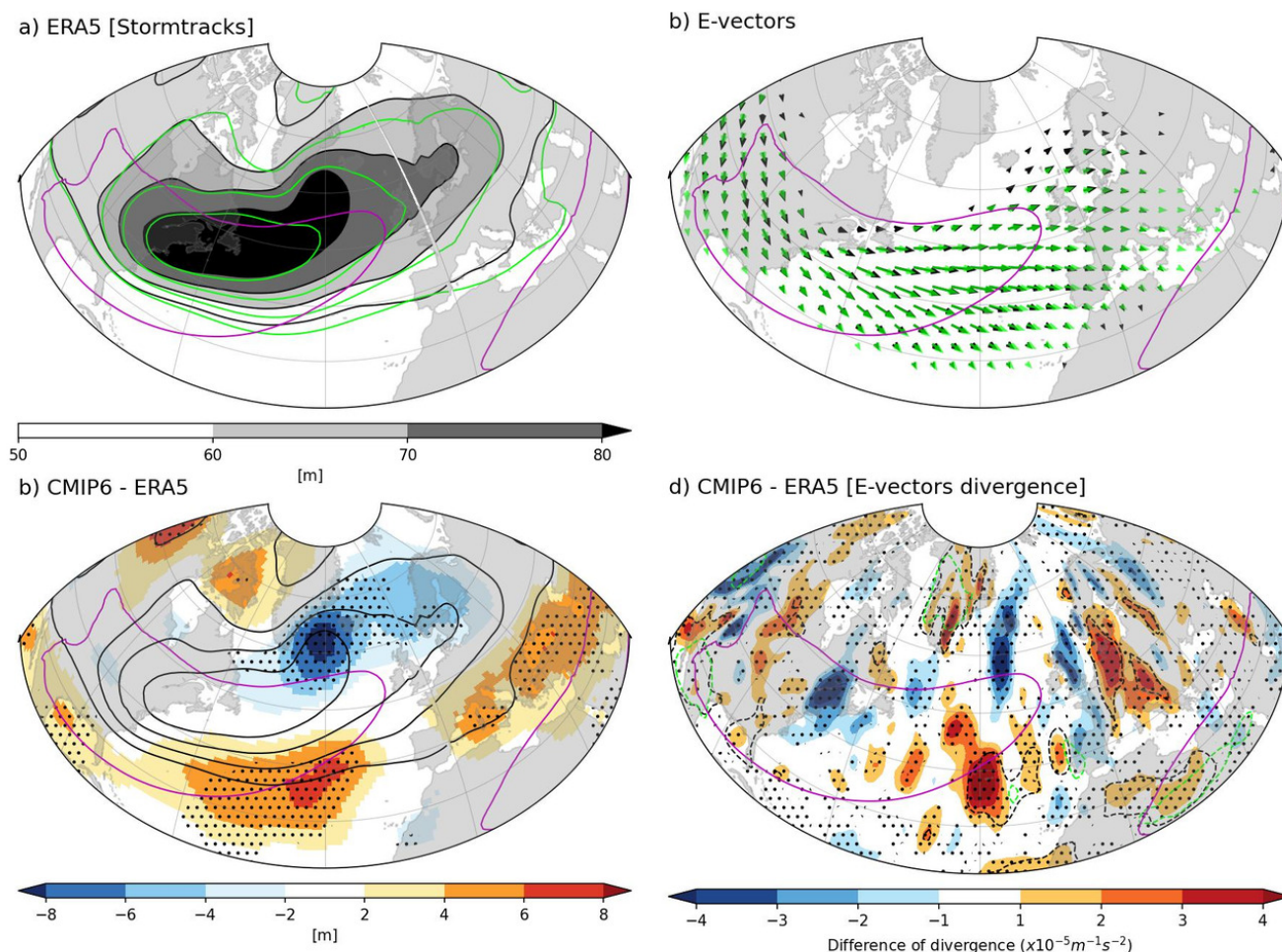
Figure 3. a) Climatological storm tracks (standard deviation of the band-pass filtered Z at 500 hPa in units of m) for ERA5 (black shading) and the CMIP6 multi-model mean (green contours). The ERA5 zonal wind at 500 hPa is the purple contour ($15\ m\,s^{-1}$). b) E-vectors at 250 hPa for ERA5 (black arrows) and CMIP6 (green arrows). Only E-vectors are drawn where the wind speed is larger than $30\ m\,s^{-1}$ (at 250 hPa). c) CMIP6 multi-model mean biases in storm tracks (band-pass filtered Z) against ERA5. Black contours depict the ERA5 storm tracks (every 10 m, starting from 40 m). d) CMIP6 multi-model mean biases in E-vector divergence (colour shading). E-vector divergence of -2.0 x $10^{-5}m^{-1}s^{-2}$ is shown for ERA5 (black contour) and CMIP6 (green contour). Stippling denotes regions where more than 80% (6 models) of the ensemble members indicate a bias of the same sign.

To further understand the interaction between the eddies and the mean flow, we show the E-vectors in Figure 3b and their divergence in Figure 3d. The E-vectors indicate the eastward propagation of Rossby waves in the North Atlantic as described in

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previous studies (e.g. Hoskins et al., 1983; Strong and Magnusdottir, 2008; Drouard et al., 2021). E-vectors exhibit the greatest

230 magnitude over the central North Atlantic, and they reduce their magnitude in the Northeastern Atlantic. This reduction in magnitude coincides with the diffluent flow, usually found upstream of blocks (Hoskins et al., 1983; Drouard et al., 2021). Eddy activity associated with the storm tracks decelerates the background flow and can be identified as the convergence of E-vectors, and it is shown as dashed black (green) contours for ERA5 (CMIP6) in Fig.3d. In ERA5, convergence regions are found in the storm tracks, for instance, in the northeast Atlantic, central Europe, and Greenland. Convergence of E-vectors is

235 prominent in the Euro-Atlantic region, in the exit region of the jet, where we have found the most considerable underestimation in blocking frequency.

We investigate the interaction of the shifted eddies and the background flow in CMIP6 models by looking at the biases of the E-vector divergence (colours in Fig. 3d). Note that CMIP6 E-vectors have a larger zonal component vector on the equator side of the jet in the western Atlantic (green arrows in Fig. 3b), which indicates larger positive momentum transport northward

240 that accelerate the jet to the south of the climatological jet (Drouard et al., 2021). CMIP6 models overestimate the divergence at the jet's exit over regions where the ERA5 climatological mean shows convergence (i.e. upstream of the Iberian Peninsula). This bias may result in a positive contribution from the eddies to the mean flow, speeding up the mean flow and reducing the wave breaking and the blocking onset.

Underestimation of divergence is observed on the polar side of the jet, which could result in a negative contribution to the

245 mean flow and suggest an equatorward shift of the eddy wave activity. Note that the equatorward shift of the eddy wave activity coincides with the shift of the storm tracks towards the equator, showing the vital link between eddies and storm tracks (Fig. 3b). Thus, we see strong feedback between the transient eddy activity and the background flow in the Euro-Atlantic region. In tandem with the biases in the background flow, the transient eddy activity is shifted equatorward in the main storm track region and the associated differences in E-vector divergence speed up the westerly flow upstream of Europe. The advection by the

250 background flow is too strong to balance the westward propagation of Rossby waves. As a consequence, the jet extends further eastward.

4.3 WCBs and the role of moist processes

Recent studies have demonstrated the impact of moist processes on the development and maintenance of blocking. A midlatitude weather system unambiguously linked to moist processes is the WCB (Madonna et al., 2014). Latent heat release due to

255 cloud formation in these ascending air systems leads to the net transport of low PV air into the upper troposphere, amplifying the downstream ridge (Grams et al., 2011). Further, the ascent and latent heat release are associated with divergent outflow in the upper troposphere, further amplifying the downstream ridge due to a poleward advection of low-PV air (Steinfeld and Pfahl, 2019). The airmass of the WCB that reaches the upper troposphere is called WCB outflow. The described process chain associated with WCBs indicates that an underestimation of WCB outflow may indicate a misrepresentation of moist diabatic

260 processes and that it may cause a reduction of blocking frequency. In the following, we analyse the WCB inflow and outflow biases in the CMIP6 models. Subsequently, we associate WCB outflow biases with blocking biases.

Figure 4 shows the ERA5 mean WCB inflow and WCB outflow frequency obtained from ELIAS2.0 and the biases in the CMIP6 models. We show biases as an absolute difference (Fig. 4c,d) and relative to ERA5 climatological mean (Fig. 4e,f). The WCB inflow area indicates the location of air parcels in a WCB prior to the main ascent (Madonna et al., 2014; Joos et al.,
265   2023). The WCB inflow frequency in ERA5 reaches its maxima over the Atlantic just off the east coast of North America (Fig. 4 a). In the North Atlantic, the area with a WCB inflow frequency equal to or larger than 4% extends from the Gulf of Mexico to the Central North Atlantic. CMIP6 models systematically underestimate WCB inflow frequency in the western North Atlantic (Fig. 4c,e). The relative biases show an underestimation of around 30% over the main WCB inflow region in the North Atlantic (Fig. 4e). Also, in the North Atlantic, there is a slight equatorward shift of the main WCB inflow region (see green contour
270   farther to the black on the polar side than on the equator side).

The WCB outflow frequency distribution from ELIAS2.0 is comparable with the locations of WCB trajectories 48h after they started their ascent in a Lagrangian analysis (Madonna et al., 2014; Joos et al., 2023). The regions with a high frequency of WCB outflow are from the North American east coast to the south of Greenland (Fig. 4b). The WCB outflow biases in the CMIP6 models are shown in Figures 4d,f. In general, in the North Atlantic, the models tend to produce less WCB outflow, which
275   is consistent with an underestimation of WCB inflow. In the North Atlantic, the WCB outflow is more concentrated at mid-latitudes and extends less to high latitudes. The WCB outflow rarely reaches southern Greenland, where the underestimation is around 80%. We may interpret this as a weaker WCB ascent that does not reach the upper atmosphere or a weaker poleward displacement of the WCB air parcels.

These results are similar to those results obtained using a Lagrangian analysis (i.e. the LAGRANTO tool) in the Comunity
280   Earth System Model Large Ensemble (CESM-LE) (Joos et al., 2023). Some differences are the location of maximum bias; CESM-LE has the most significant bias over eastern Europe and southern Greenland (see Fig. S1 in Joos et al., 2023), while the CMIP6 bias is most prominent over the central Atlantic and south of Greenland.

The too-zonal storm tracks in the CMIP6 models may explain the lack of WCB outflow at higher latitudes. Similarly, the underestimation of WCB inflow is ultimately linked to the underestimation of WCB outflow. However, the fact that the
285   WCB inflow underestimation is about 40% hints that other processes are involved in the WCB outflow underestimation. Such underestimation of WCB outflow can be associated with misrepresenting moist diabatic processes in the cloud formation and its associated circulation. Less heating leads to weaker cross-isentropic ascent and a weaker upper-tropospheric divergent outflow. As discussed, ascending trajectories by the WCB activity are crucial for blocking amplification and maintenance. CMIP6 models fail to reproduce the WCB inflow and outflow frequency. The lack of WCBs reaching upper levels south of
290   Greenland may thus explain a large portion of blocking biases in the Europe-Atlantic region.
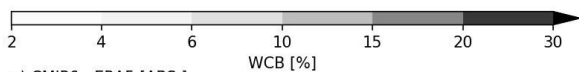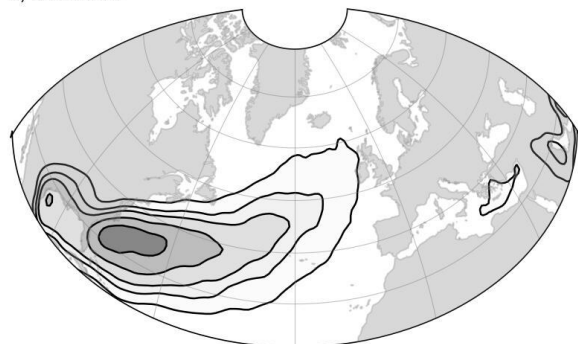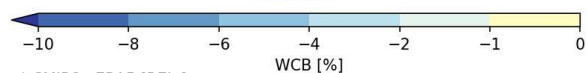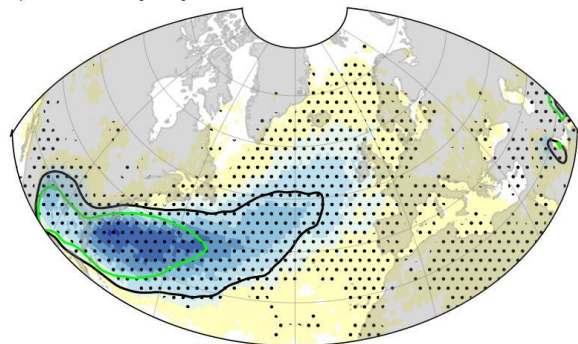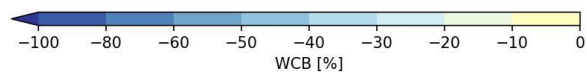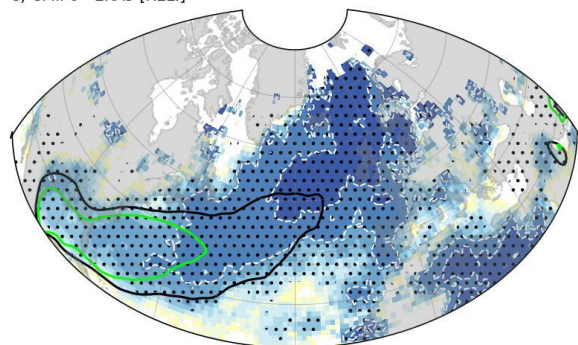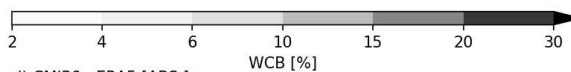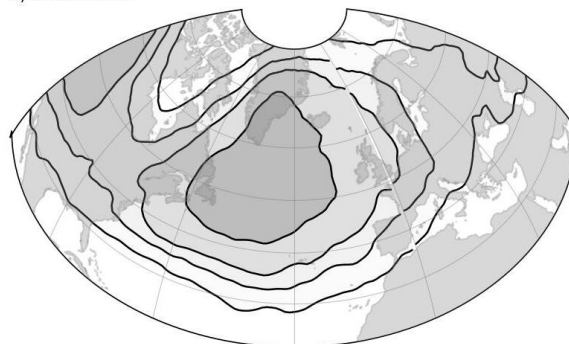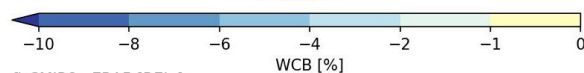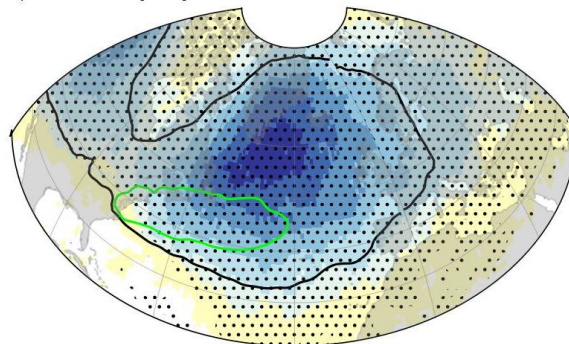
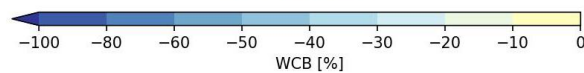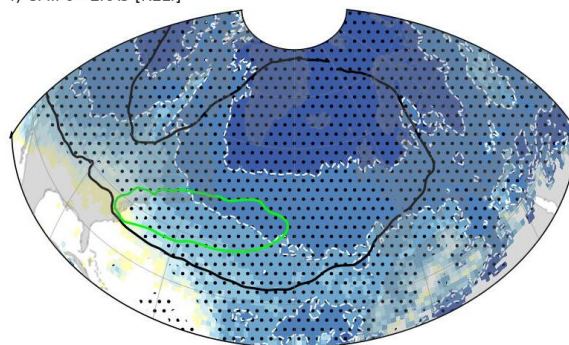**Figure 4.** a,b) Climatological frequency of a) WCB inflow and b) WCB outflow for ERA5. c, e) CMIP6 multi-model mean biases of WCB inflow and d,f) outflow against ERA5. c,d) show the absolute biases in % frequency, and e,f) show the relative biases. In panels c,d,e, and f), a contour of 4% frequency is depicted in black for ERA5 and green for the CMIP6. The white contour in e,d) depicts the biases of 60% and 80%. Stippling denotes regions where more than 80% (6 models) of the ensemble members indicate a bias of the same sign.

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5   Discussion

This study links Euro-Atlantic blocking biases in CMIP6 models to biases of the background flow, dry, and moist diabatic processes. The CMIP6 models analysed here exhibit substantial biases in the blocking frequency over the North Pacific, Greenland and the Euro-Atlantic region. Over the North Pacific and Greenland (i.e. in blocking regions north of the storm track), blocking

295   biases vary across the models. In the Euro-Atlantic region, where blocking occurs downstream of the storm track, the underestimation is particularly pronounced, reaching 80% relative to ERA5. Accordingly, we focus on that region in the following. The underestimation is limited to a smaller area than reported by previous studies looking at the CMIP6 models (e.g. Davini and D'Andrea, 2020; Schiemann et al., 2020). This improvement may come from selecting only models with a reasonable performance according to Palmer et al. (2023). In the following, we further discuss the role of background flow and dry and

300   moist processes.

Regarding the background flow, CMIP6 models show an equatorward shift of the area of most substantial meridional geopotential height gradients in the North Atlantic (see Figure 2). The stronger winds to the south of the climatological jet lead to a faster eastward translation of Rossby waves. Accordingly, Rossby waves in these areas need larger zonal wavelengths to become stationary. The storm tracks and the eddy wave activity (defined by the E-vectors) are too zonal in the CMIP6 models

305   (Figure 3). Usually, diffluent flow, from a strong zonal flow to a more meridional flow, favours blocking development (Gabriel and Peters, 2008; Nakamura and Huang, 2018). However, the diffluent flow is unclear in the Euro-Atlantic region, where a stronger $|\frac{dZ}{dy}|$ tends to dominate. The momentum forcing of eddies, quantified by the E-vectors, strengthens the southward shift and zonal extension of the jet (Figure 3). In the Euro-Atlantic region, the divergence of E-Vectors is stronger in the CMIP6 models and associated with accelerating the background flow. The resulting flow is less diffluent over the east Atlantic and thus

310   less favourable for blocking formation in CMIP6.

Moist processes (e.g. cloud formation) interact with atmospheric blocking and can impact their onset and maintenance (e.g. Pfahl et al., 2015). Here, we consider moist processes by analysing WCB outflow. We find a substantial underestimation of the WCB outflow in the North Atlantic in the CMIP6 models. We argue that two factors can explain these biases: i) misrepresenting diabatic processes and ii) biases in the WCB inflow. We find a relatively higher underestimation of the WCB outflow frequency

315   than the WCB inflow frequency (see Figure 4). The weaker ascent of air parcels in WCBs in CMIP6 might explain part of the WCB outflow biases in the CMIP6 models (see also Woollings et al., 2018; Steinfeld et al., 2020; Joos et al., 2023). Moreover, the WCB outflow is constrained to lower latitudes (no poleward displacement), which can be linked to the too-zonal and equatorward-shifted storm tracks. To better understand WCB inflow biases, we analyse lower tropospheric heat and moisture fluxes. Lower tropospheric moisture fluxes are essential during the initial stage of WCB and may indicate the presence of

320   the WCB inflow (Dacre et al., 2019; Quinting et al., 2024). We hypothesise that underestimating moisture and heat eddy flux reduces WCB frequency and ascent. Indeed, the meridional moisture flux is one predictor in ELIAS2.0 (Quinting and Grams, 2022). Figure 5 shows the meridional eddy moisture and heat fluxes at 850 hPa in ERA5 and the CMIP6 biases. In ERA5, the climatological eddy moisture flux maxima are in similar locations as the WCB inflow climatology with a slight poleward

displacement in the North Atlantic (Fig. 5a). The maxima in eddy heat fluxes are at higher latitudes than the WCB inflow

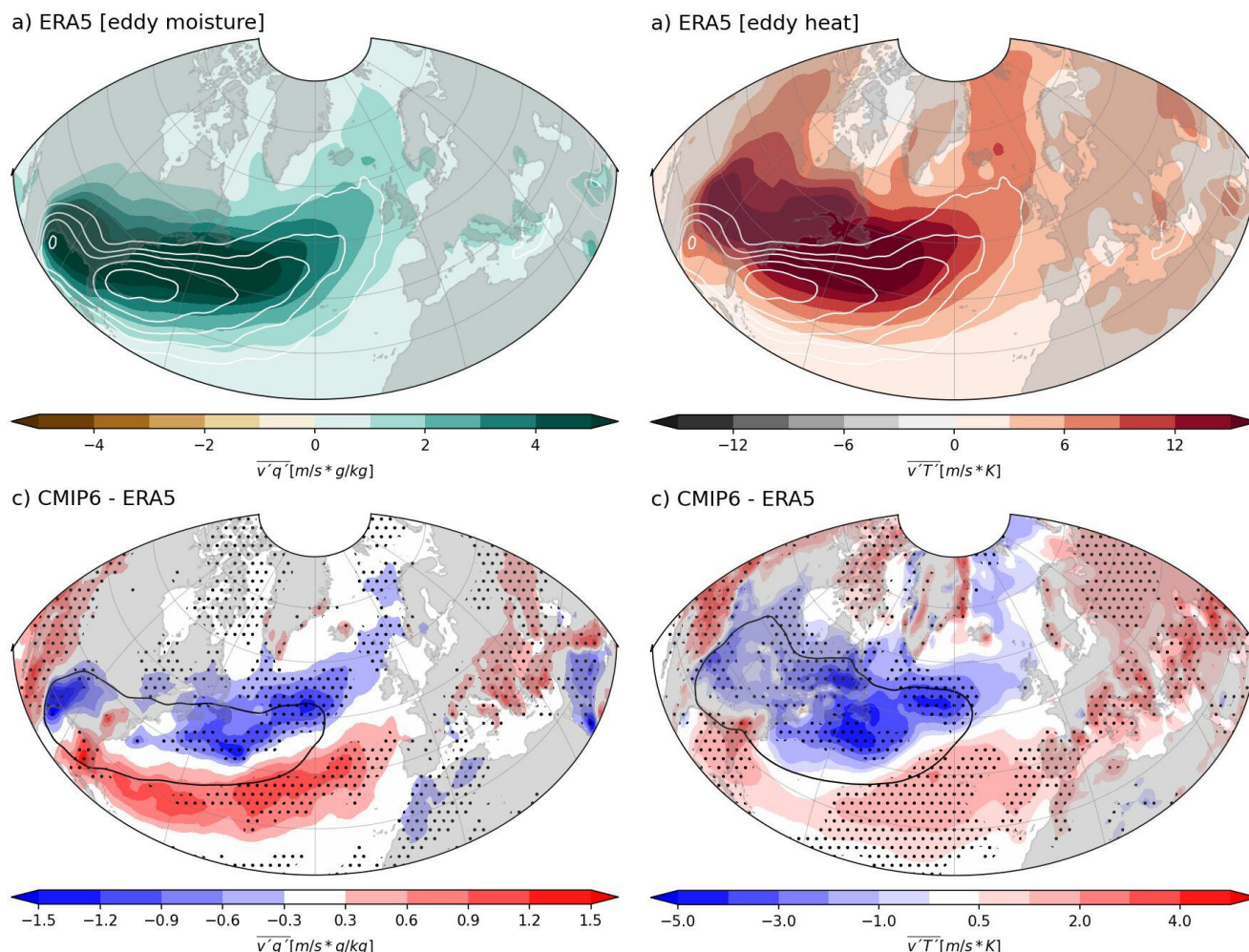325 regions closer to the storm tracks (Fig. 5b).

Figure 5. a) Climatological eddy moisture flux in ERA5 (colour shading). The white contours are the climatological WCB inflow frequency for ERA5 [2,4,6,10,15,20,30%]. b) Climatological eddy heat flux in ERA5. c,d) CMIP6 multi-model mean biases of eddy c) moisture flux and d) heat flux against ERA5. We show the contour of 3.5 $m\,s^{-1}*g\,kg^{-1}$ in panel c and 10 $m\,s^{-1}*K$ in panel d of the ERA5 fluxes. Stippling denotes regions where more than 80% (6 models) of the models indicate a bias of the same sign.

In the North Atlantic, the maximum eddy moisture flux from the CMIP6 models is equatorward-shifted (Fig. 5c). This is consistent with the equatorward-shifted storm tracks but does not explain the underestimation of WCB inflow in the North

Atlantic. Although, we observe an eddy moisture flux underestimation in the Gulf of Mexico, which can lead to WCB inflow biases. Nevertheless, we observe a clear underestimation (50% relative to ERA5) of eddy heat fluxes in the USA west coast (Fig. 5d). Thus, the biases in the WCB inflow over the North Atlantic are associated with biases in eddy moisture and heat fluxes.

In summary, the CMIP6 models exhibit two systematic biases: 1) a southward shift of the background flow and 2) an underestimation of eddy activity in the North Atlantic region, associated with WCB inflow and outflow biases and blocking underestimation. Moreover, biases can also arise from the skill of ELIAS2.0 to identify WCB inflow in CMIP6 models. One limitation is that the ELIAS2.0 model was trained with ERA-Interim and here is applied to ERA5 and climate models. This is the first time ELIAS2.0 has been applied to climate models, which generally have a lower resolution than ERA-Interim. Validation with explicit WCB trajectories in climate models (e.g. CESM-LE) could help better understand potential biases in ELIAS2.0. Finally, we have used earth system models interpolated from different resolutions to 1^ox1^o degrees; smoothed values can arise from the interpolation processing challenging ELIAS2.0 to identify the WCBs.

6 Conclusions

This study evaluates the representation of atmospheric blocking in eight CMIP6 models (see table 1 for details). We base the model selection on information from the literature: i) on the models' skill to simulate atmospheric blocking and ii) on the independence of the models in terms of shared model components. The CMIP6 multi-model mean biases reveal that the frequency of Euro-Atlantic blocking is the most severely underestimated of the major blocking regions in the Northern Hemisphere. During winter, the underestimation reaches 80%, confirming the results of previous studies.

We then focus on identifying different potential sources of the Euro-Atlantic blocking biases: the background flow, dry processes, and moist processes. We use the mean and mode of the latitudinal gradient of the geopotential height at 500 hPa to describe the background flow. In the multi-model CMIP6 mean, the jet (area with strong gradients and depicted as purple contour in Figs. 2a,b,e) is shifted equatorward, more zonal and extends further downstream. The storm tracks, and their associated eddy activity is also shifted equatorward in the North Atlantic in the CMIP6 models. CMIP6 E-vectors show that diffluent flow is reduced in the Euro-Atlantic region. Moreover, the contribution of the eddies to mean flow is examined by the $\nabla \cdot \mathbf{E}$, where the convergence indicates a deceleration of mean flow by the eddies. In CMIP6, we find a reduced convergence or divergence in the Euro-Atlantic region. Thus, the eddies accelerate the background flow and negatively contribute to the block development in the eastern North Atlantic.

We have applied deep learning WCB detection (ELIAS2.0) to the CMIP6 models. Both WCB inflow and WCB outflow frequency are underestimated in the CMIP6 models. A similar underestimation of WCB frequency is found using a Lagrangian approach in climate models (Joos et al., 2023). The WCB outflow frequency is 80% lower in the Atlantic in CMIP6 compared to ERA5. The WCB outflow biases are partly linked to the underestimation of WCB inflow in the western Atlantic. The underestimation of the WCB inflow is likely linked to the misrepresentation of eddy moisture and heat fluxes in the lower troposphere. In the North Atlantic, WCB outflow frequency is underestimated by 80% in the region of maximum WCB outflow

(south of Greenland), while inflow is underestimated by 40%. Thus, WCB inflow may explain part of the WCB outflow underestimation but not the whole. CMIP6 models lack WCB outflow at higher latitudes and may be associated with air masses that do not reach as far polewards as in ERA5 or do not reach the upper troposphere. One can interpret this misrepresentation as a consequence of weaker latent heat release, reflecting an underestimation of moist processes in CMIP6 models. Thus,
365   the misrepresentation of moist processes in CMIP6 is likely one factor contributing to reducing Euro-Atlantic blocking by depositing less anticyclonic PV air in the upper troposphere.

We summarise the background flow, dry and moist processes biases, and their link to the blocking biases in Figure 6. In the North Atlantic, a too-zonal background flow is associated with a southward migration of the eddy activity, and momentum transport from eddies accelerates the mean flow in the Euro-Atlantic region (yellow region in Fig. 6) and reduce diffluent
370   flow and blocking formation. CMIP6 models also underestimate moisture fluxes in the lower troposphere, linked to a lower frequency of WCB inflow and WCB outflow and blocking.
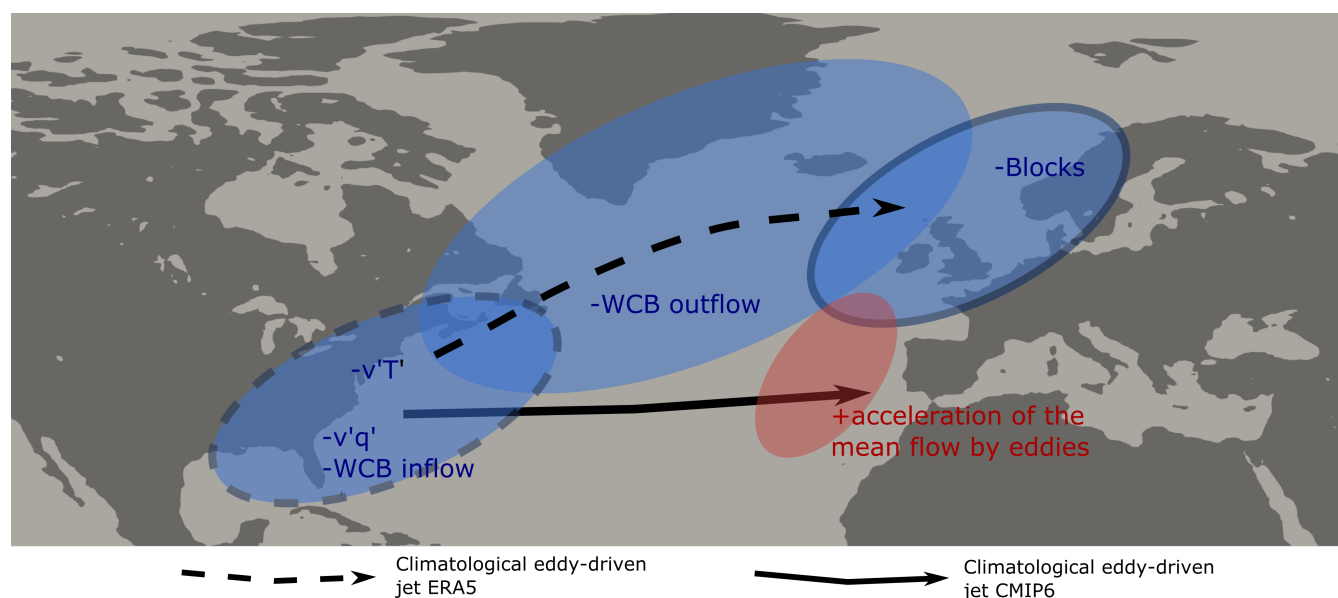

Figure 6. Schematic summary of the identified sources of the Euro-Atlantic blocking biases in the CMIP6 models. Red shaded colour indicates positive biases and blue indicates negative biases. There is a negative blocking bias over the eastern North Atlantic and Europe in the CMIP6 simulations (blue area and solid grey contour). The eddy-driven jet over the North Atlantic is shifted equatorward and extends farther into the eastern North Atlantic in the CMIP6 simulations. The associated shift in the storm track results in a positive eddy-mean flow feedback, with the transient wave activity accelerating the jet over the eastern North Atlantic. Weaker moisture and heat fluxes in the western Atlantic are associated with a negative WCB inflow frequency bias (blue area and dashed grey contour). A negative frequency bias in the WCB outflow over the central and eastern North Atlantic contributes to a weaker downstream ridge.

Code availability. The code of ELIAS2.0 is available from https://gitlab.kit.edu/julian.quinting/elias-2.0.git.

Data availability. CMIP6 data are available via easy.gems.dkrz.de.

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Author contributions. OM and EDT designed the study. EDT performed the analysis, with the support of JQ for the ELIAS2.0, produced the figures, and drafted the manuscript. All authors discussed the results and edited the manuscript.

Competing interests. The authors declare no competing interests

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