

Emulating the ~~future distribution~~ expansion of Antarctic perennial firn aquifers in ~~Antarctica~~ the 21st century

Sanne B. M. Veldhuijsen¹, Willem Jan van de Berg¹, Peter Kuipers Munneke¹, Nicolaj Hansen², Fredrik Boberg², Christoph Kittel^{3,4,5}, Charles Amory³, and Michiel R. van den Broeke¹

¹Institute for Marine and Atmospheric Research Utrecht, Utrecht University, Utrecht, The Netherlands

²Department of National Centre for Climate Research (NCKF), Danish Meteorological Institute, Copenhagen, Denmark

³Institut des Géosciences de l'Environnement (IGE), Univ. Grenoble Alpes/CNRS/IRD/G-INP, Grenoble, France

⁴Département de Géographie, SPHERES, ULiège, Liège, Belgique

⁵Physical geography research group, Department of geography, Vrije Universiteit Brussel

Correspondence: Sanne B. M. Veldhuijsen (s.b.m.veldhuijsen@uu.nl)

Abstract. Perennial firn aquifers (PFAs) are year-round bodies of liquid water within firn, which modulate meltwater runoff to crevasses, potentially impacting ice-shelf and ice-sheet stability. Recently identified in the Antarctic Peninsula, PFAs form in regions with both high surface melt and snow accumulation rates, and are expected to expand due to the anticipated increase in surface melt and snowfall. Using a firn model to predict future Antarctic PFAs for multiple climatic forcings is relatively computationally expensive. To address this, we develop an XGBoost perennial firn aquifer emulator, a fast machine learning model ~~trained~~. It was trained, using a scenario and spatial blocking evaluation approach, on PFA output of simulations from the firn densification model IMAU-FDM. ~~The firn simulations were,~~ which was forced by the combined regional climate model RACMO2.3p2 and global climate model CESM2 for three emission scenarios (SSP1-2.6, SSP2-4.5 and SSP5-8.5). ~~Using a scenario and spatial blocking evaluation approach, we found that the emulator successfully explains at least 89 % of PFA presence and meltwater storage variance. Using the PFA emulator, we predict future PFAs (2015-2100) for~~ The trained emulator was applied on nine additional forcings (2015-2100) from the regional climate models MAR and HIRHAM in combination with five global climate models. ~~Under~~ We show that the emulator is robust, explaining at least 89% of the variance in PFA presence and meltwater storage. Our results indicate that for the SSP1-2.6 and SSP2-4.5 scenarios, PFAs remain mostly restricted to the Antarctic Peninsula. For SSP5-8.5, PFAs expand to Ellsworth Land in six out of the seven simulations and to Enderby Land in East Antarctica in five out of the seven simulations. ~~The emulator results for RACMO and MAR agree on the~~ Furthermore, the emulator predicts PFAs for similar surface melt and accumulation conditions ~~required for PFA formation. While these results for HIRHAM when forced with MAR or RACMO data. For HIRHAM these conditions are slightly different, caused by different modelled relations due to the different relationship between temperature, accumulation and melt in HIRHAM compared to RACMO. Overall, our findings show that PFAs will likely are likely to expand in a warmer Antarctica, irrespective of the emission scenario,~~ increasing the risk that an ice shelf collapses due to hydrofracturing.

1 Introduction

The Antarctic Ice Sheet (AIS) has been losing mass since at least 2002 (Shepherd et al., 2019), contributing $\sim 10\%$ to global average sea level rise since 1993 (Oppenheimer et al., 2019). The primary factor currently contributing to AIS mass loss is enhanced basal melt beneath ice shelves (Smith et al., 2020), leading to their thinning, enhanced iceberg calving or collapse.

25 This, in turn, reduces the ice-shelf buttressing effect, allowing inland ice to flow faster into the ocean. Another process that can potentially indirectly reduce the buttressing effect of ice shelves is melt pond-driven hydrofracturing. This mechanism contributes to ice-shelf disintegration thereby reducing the ice-shelf volume, which in turn reduces the ice-shelf buttressing. This process is expected to increase under future warming (Lai et al., 2020; van Wessem et al., 2023). In the case of the disintegration of Larsen A and B ice shelves in 1995 and 2002, respectively, depletion of the firn air content and subsequent
30 meltwater ponding and infilling of crevasses is generally thought to have caused hydrofracturing (Scambos et al., 2000; Banwell et al., 2013). Surface meltwater accumulation on ice shelves can trigger ice-shelf collapse when tensile stresses are sufficiently large (Banwell et al., 2013). Currently, 60% of the ice shelf area buttresses upstream ice and is vulnerable to hydrofracturing if inundated by meltwater (Lai et al., 2020).

While much attention has been directed to melt accumulation on the ice shelf surface following firn air depletion, meltwater
35 storage within the firn layer can also modulate meltwater drainage to crevasses (e.g., Scambos et al., 2009). Under climatic conditions characterized by high surface melt rates and snow accumulation, liquid water can persist year-round within the firn's pore space, forming perennial firn aquifers (PFAs). As snow has a low thermal conductivity (Calonne et al., 2019) high snowfall rates in fall and winter rapidly cover and insulate the summer meltwater and thereby prevent it from refreezing in winter.

40 PFAs are common across southeast Greenland (e.g., Forster et al., 2014; Miège et al., 2016). On the AIS, firn aquifers have been observed on the Wilkins and Müller ice shelves in the western Antarctic Peninsula (AP) (Montgomery et al., 2020; MacDonell et al., 2021), the warmest and wettest region of Antarctica. An exploratory firn modelling study predicted PFAs to occur also on the grounded ice along the north-western coast of the AP (van Wessem et al., 2021). Combining regional climate model (RCM) output with satellite data ~~confirm~~ confirms a high probability of PFA occurrence in all these regions (Di Biase
45 et al., 2024). Increasing surface melt and snowfall may result in future PFA expansion, also to other regions of Antarctica. Similarly, an inland expansion of aquifers occurred over the last decades in Greenland (Horlings et al., 2022). However, the future distribution of PFAs in Antarctica has not yet been explored.

~~Firn aquifers provide sufficient water storage capacity to contribute to ice shelf disintegration events, as evidenced by the partial break-ups of the northern and northwestern Wilkins ice shelf in 1993, 1998, and 2008–2009, where the detection of
50 bright reflectors in airborne radar surveys indicated the presence of a firn aquifer (e.g., Braun et al., 2009; Scambos et al., 2000; Montgomery et al., 2009).~~
~~The year-round availability of water at depth in firn aquifers can lead to hydrofracturing when stress conditions shift to favor tensile extensions-extension or reduced compression, even during the winter (Scambos et al., 2009). This mechanism can initiate cascading drainage events that cause rapid and large-scale disintegration, such as observed for the ice shelf disintegration. The partial break-ups of the northern and northwestern Wilkins ice shelf in 1993, 1998, and 2008 (Scambos et al., 2009).~~

55 ~~Another ice shelves with an observed firn aquifer~~ 2009 could well exemplify this process, as the detection of bright reflectors in
airborne radar surveys indicated the presence of a firn aquifer (e.g., Braun et al., 2009; Scambos et al., 2000; Montgomery et al., 2020)
. These events highlight how water stored in firn aquifers can potentially accelerate disintegration. Similarly, the Müller ice
~~shelves, shelf, which also contains firn aquifers,~~ has lost a considerable portion of ~~their surface areas~~ its surface area (49%)
since the 1950s (Cook and Vaughan, 2010) , while neighbouring Jones and Wordie ice shelves, with similar climatic conditions,
60 have completely disintegrated (Cook and Vaughan, 2010).

PFAs can also play an important role over the grounded ice. In Greenland, liquid water from PFAs drains through crevasses
to the bed (Poinar et al., 2017), leading to basal lubrication, at least temporarily increasing ice velocity and ice discharge into
the ocean (Zwally et al., 2002). Warming through the release of latent heat during refreezing changes the ice rheology (Hubbard
et al., 2016). Hence, understanding the future evolution of PFAs is relevant when assessing the future stability of ice shelves
65 and grounded ice.

Offline firn models forced by output of RCMs have proven to be useful tools to simulate the ice-sheet wide transient evolution
of firn. They have already been used to simulate historical firn aquifers over Greenland (Brils et al., 2024), and the AP (van
Wessem et al., 2021). RCMs are better suited for this than global climate models (GCMs), whose resolution is too coarse
to correctly represent strong coastal gradients in precipitation and surface melt (Bozkurt et al., 2021), which are particularly
70 important for PFA formation (van Wessem et al., 2021). In addition, most GCMs do not properly represent important physical
processes, such as the snowmelt-albedo feedback (Jakobs et al., 2019).

For future projections of PFA extent, multiple climatic scenarios, and combinations of RCMs and GCMs are ideally used, as
projections of near-surface climatic conditions and surface melt extent over Antarctica vary widely among models (Carter et al.,
2022; Kittel et al., 2021). However, running a firn model for multiple forcings is relatively computationally demanding and
75 requires 3-hourly model output, which is regularly unavailable. ~~Therefore~~ To enable projections under a wide range of climate
forcings, we developed an ~~XGBoost PFA emulator~~ XGBoost-based PFA emulator to estimate future PFA extent. Emulators
are statistical or machine learning models that are trained on specific output of a more complex model. Emulators can be based
on simple statistical models, such as linear regression, or use more advanced machine learning approaches, such as Random
Forest or XGBoost regressors. Several studies have used emulators to simulate firn properties (Dunmire et al., 2024; Verjans
80 et al., 2021; Jourdain et al., 2024), as well as other ice sheet processes (Van Katwyk et al., 2023). XGboost is a decision
tree-based methods (Chen and Guestrin, 2016) that has been shown to outperform other machine learning algorithms, such
as neural networks, random forest regression and linear regression methods, for various glaciological applications (Anilkumar
et al., 2023; Iban and Bilgilioglu, 2023; Sun et al., 2024).

We train and evaluate the PFA emulator using output from the IMAU Firn Densification model v1.2AD (IMAU-FDM
85 v1.2AD), which was recently updated by Veldhuijsen et al. (2024d) to better simulate transient firn densification in a changing
climate. IMAU-FDM was forced by future climate realisations of the GCM CESM2, dynamically downscaled over the entire
AIS by the RCM RACMO2.3p2 to a 27 km resolution, for emission scenarios SSP1-2.6, SSP2-4.5 and SSP5-8.5 (Veldhuijsen
et al., 2024d). We subsequently apply our emulator to 12 climate projections from three polar RCMs (RACMO2.3p2, MAR3.11
and HIRHAM5) forced by several GCMs. We then report the resulting PFA distribution and expansion on the AIS.

2.1 IMAU-FDM firn model and forcing

IMAU-FDM version v1.2AD is a semi-empirical 1D firn densification model that simulates the time evolution of firn depth, density, temperature, grain size and liquid water content. Firn compaction is calculated based on the semi-empirical dry-snow densification equations of Arthern et al. (2010) with an updated dynamical densification expression to cope with changing climate forcing (Veldhuijsen et al., 2024d). The updated ~~densification rate depends~~ expression lets densification rate depend on firn temperature, grain size and overburden pressure ~~instead of firn temperature and~~, whereas previously densification rate depended on firn temperature along with averages over the past 40 years of accumulation and surface temperature. The vertical percolation of liquid water from melt or rain is simulated using the bucket method, whereby liquid water is only retained through capillary forces (i.e. irreducible water). The maximum irreducible water decreases with increasing density (Coléou and Lesaffre, 1998). The meltwater can percolate through all layers in a single time step and (partly) refreeze when it reaches a layer with nonzero pore space and a temperature below the freezing point. Once the liquid water content of the lowermost firn layer exceeds the maximum irreducible water content, the surplus is assumed to instantaneously leave the firn column as runoff. The bucket method is computationally efficient but does not allow for saturated pore spaces, preferential flow, standing water over ice layers or horizontal flow. Since perennial firn aquifers (PFAs) are defined as saturated firn, ~~this is something different from what~~ IMAU-FDM ~~can simulate~~ cannot simulate them. However, the model does simulate the insulation of downward percolating liquid water by sufficiently high accumulation rates, that prevent the meltwater from refreezing in the winter. Therefore, the presence of year-round or spring liquid water in the firn has been successfully used to identify the presence of aquifers (Forster et al., 2014; van Wessem et al., 2021; Brils et al., 2024).

In the simulations presented here, IMAU-FDM is forced at the upper boundary, the snow surface, with 3-hourly values of snowfall, sublimation, snowdrift erosion, 10-m wind speed, surface temperature, surface melt and rainfall from simulations of the regional climate model (RCM) RACMO2.3p2 at a 27 km resolution (van Wessem et al., 2023). RACMO2.3p2 was in turn forced by the Community Earth System Model version 2 model output (CESM2) (Danabasoglu et al., 2020). These simulations consist of a historical climate realisation (1950–2014) and climate realisations for low-, middle- and high-emission future (2015–2100) scenarios (SSP1-2.6, SSP2-4.5 and SSP5-8.5, respectively). The projected Antarctic end-of-century (2090–2100) warming in SSP5-8.5 in CESM2 (+6.7 K) compared to 2005–2015 is stronger than the mean Antarctic warming in CMIP6 models (+4.8 K), which enables us to train the firn model on a wide range of temperatures (Dunmire et al., 2022; Kittel et al., 2021). The coupling between RACMO2.3p2 and IMAU-FDM is unidirectional. The stand-alone approach of IMAU-FDM allows for a more realistic initialisation and higher vertical resolution than RACMO’s built-in firn model, which uses similar physics. The disadvantage is that interaction with the atmosphere is not possible. An initial firn layer is obtained by looping over the forcing of the 1950–2000 reference period until the firn layer is in equilibrium with the surface climate. The model employs up to 3000 layers of 3 to 15 cm thickness, which represent the firn properties in a Lagrangian way. For further details on the model set up and forcing, we refer to Veldhuijsen et al. (2024d).

2.2 PFA emulator

We develop a PFA emulator that mimics how IMAU-FDM transiently simulates PFAs. We use the annual perennial amount of liquid water content (LWC) as the target variable (see Section 2.2.1). To emulate PFAs we use an extreme gradient boosting (XGBoost) machine learning model (Chen and Guestrin, 2016). The model is developed in Python using the open-source scikit-learn and XGBoost packages. XGBoost sequentially builds an ensemble of weak decision trees. In each iteration, it fits a new tree to the residuals of the previous iteration, optimizing a specific objective function to minimize the overall prediction error.

XGBoost is chosen for its high predictive accuracy and [because it](#) has been shown to outperform other machine learning algorithms, including neural networks, random forest regression and linear regression methods, in predicting e.g., snow water equivalent, snow avalanche susceptibility and glacier mass balance (Anilkumar et al., 2023; Iban and Bilgilioglu, 2023; Sun et al., 2024) and in general for medium and large sized (>10,000) tabular datasets (Grinsztajn et al., 2022). Additionally, XGBoost is highly scalable, meaning it can handle large datasets and complex models, ~~it~~. [Furthermore, XGBoost](#) incorporates regularizations techniques to prevent overfitting and it has built-in mechanisms to estimate feature importance. Recently, a Random Forest model, a bagging algorithm that trains decision trees in parallel, has been used to emulate firm air content over Antarctica (Dunmire et al., 2024). While Random Forest is known for its simplicity, ease-of-use, and resistance to overfitting, XGBoost often better captures complex nonlinear relationships, especially in scenarios where there are interactions between features or when the dataset is high-dimensional (Fatima et al., 2023).

2.2.1 Target variable and input features

Our target variable is the annual perennial LWC, which is defined as the minimum vertically integrated LWC over a year, based on model output at 10-day intervals. An annual perennial LWC of zero indicates the absence of LWC at some point during a year, meaning there is no PFA. A drawback of this approach is that it does not differentiate between years with brief periods without LWC (e.g. 10 days) and years with long periods without LWC (e.g. 10 months). Since the target variable, the annual perennial LWC, cannot take negative values, this prevents negative values from counterbalancing positive biases in the predictions (e.g. an asymmetric distribution of the errors). To address this, we introduce negative values to represent the number of days without LWC during periods when PFAs are absent. This enables us to capture both the presence and absence of LWC in a unified way. For example, if LWC is absent for 10 days, the target value is -10; if LWC is absent for 10 months, the target value is approximately -300. Although the negative and positive quantities are conceptually distinct, their absolute magnitudes are comparable in scale, providing a relatively smooth transition around zero. To further clarify, this method does not involve averaging LWC values with negative values. Instead, the target variable is a single, unified measure that can take either positive values (indicating LWC amount) or negative values (indicating days without LWC). If the target variable exceeds zero, it is considered to indicate the presence of a PFA.

The climate variables used as input features for our emulator are: (1) total annual snow accumulation (snowfall minus evaporation/sublimation), (2) total autumn (MAM) snow accumulation, (3) total summer (DJF) snowfall, (4) total annual liquid

water input (surface melt plus rainfall), (5) annual Melt-over-Accumulation (MoA) ratio, (6) mean annual surface temperature, (7) mean summer (DJF) surface temperature and (8) annual seasonal temperature amplitude. These input features describe the most important mass fluxes and boundary conditions governing ~~the~~ firn density, temperature and LWC, and ~~therewith~~ ultimately the presence of PFAs (Kuipers Munneke et al., 2014; van Wessem et al., 2021). The annual seasonal temperature amplitude is defined as the difference between the average temperature of the warmest month and the coldest month. For the summer, we consider snowfall instead of snow accumulation, due to the likely presence of evaporation alongside sublimation, which decreases liquid water input compared to sublimation which decreases snow accumulation. To account for the time it takes for firn to adjust to climatic conditions, we consider averages over the past 5, 10 and 30 years of the input features. Additionally, we include surface elevation and surface slope as input features, as the topography influences short-term climate variability, such as extreme precipitation (González Herrero et al., 2023). This amounts to 23 input features in total.

2.2.2 Dataset selection

The emulator is trained to predict annual perennial LWC across Antarctica. The dataset consists of 18,136 IMAU-FDM grid cells with 85 time steps each (years between 2015-2100), amounting to 4.6 million data points for three climate scenarios. However, for most of the AIS, the climate conditions are not favorable for aquifer formation, i.e. too dry and/or too cold, even under strong future warming. For instance, RACMO forced by CESM2 predicts that 76% of the AIS does not experience any melt by the end of the century even for the SSP5-8.5 scenario (Veldhuijsen et al., 2024d). To minimize the influence of non-aquifer locations, we limit the training dataset to aquifer-favorable locations. For each scenario, we select data points where PFAs are present or where the duration without LWC is shorter than 150 days. In addition, we select data points with aquifer-favorable conditions, defined as having more than 400 mm w.e. yr⁻¹ of accumulation and more than 200 mm w.e. yr⁻¹ of melt (van Wessem et al., 2021; Brils et al., 2024). From the remaining dataset, we randomly select 5,000 points from each scenario. In total, our dataset used for training consists of 123,000 data points (2.6% of the total AIS data points).

2.2.3 Emulator training

Our dataset has a spatial and temporal structure, and to ensure independence between training and testing data, we split the training and testing data strategically rather than randomly (Roberts et al., 2017). Firstly, we evaluate the performance using spatial blocking cross-validation, with R^2 as the scoring metric. For this, we divide the area in which firn aquifers occur into eight regions. During the cross-validation, we leave out one region at a time. Since firn aquifers are rare on the AIS, cross-validation ensures that the model is evaluated on the full dataset. Additionally, we also evaluate the performance by leaving out each of the three emission scenarios at a time, referred to as scenario blocking cross-validation. The initial 30 years of the future ~~part of the~~ simulations (2015-2045) were excluded from the validation scores to avoid dependency on the other scenarios. Finally, we use all scenarios and regions for training the emulator, to ensure that the widest possible range of warming trends and climatic conditions are included in the emulator.

The hyperparameters are tuned using the BayesSearchCV function (built-in scikit-learn) during the spatial blocking cross-validation. Eight key hyperparameters involved in the XGBoost algorithm were optimized. Firstly, the number of estimators

(`n_estimators`) refers to the number of trees. The learning rate (`learning_rate`) controls the step size at which the optimizer makes updates to the weights. The regularisation techniques included in XGBoost are the maximum depth of the trees (`max_depth`), the minimum weight needed in a leaf node (`min_child_weight`), the subsample ratio of the training instances (`subsample`), subsample ratio of columns when constructing each tree (`colsample_bytree`), the minimum loss reduction (`gamma`), and the regularisation degree (`reg_lambda`).

2.3 Emulator training and forcing data

IMAU-FDM simulations forced by RACMO-CESM for three scenarios (SSP1-2.6, SSP2-4.5 and SSP5-8.5) were used to train the PFA emulator. By applying the emulator to output from three polar RCMs, an ensemble of 12 scenarios of perennial LWC predictions for 2015-2100 is produced (Table 1). CESM2-forced runs of MAR (Jourdain et al., 2024) and HIRHAM (Hansen et al., 2021) RCMs are available for SSP1-2.6, SSP2-4.5 and SSP5-8.5, and SSP1-2.6 and SSP5-8.5, respectively. In addition, MAR simulations driven by global climate models (GCMs) ACCESS-1.3, NorESM1-M and CNRM-CM6-1 for SSP5-8.5 (Kittel et al., 2021) and HIRHAM simulations driven by GCMs EC-Earth3 for SSP5-8.5 (Boberg et al., 2022) were also used. The horizontal resolution of each climate forcing is preserved in the emulator simulations. Although the probability of SSP5-8.5 scenarios is conceived as low, they are useful to assess low-probability/high-impact futures, and they characterize the climate of the 22nd century (the probability for these conditions increases to about 7% by 2150) (Sarofim et al., 2024). In addition, we are limited by the unavailability of downscaled forcing datasets over Antarctica for the alternative emission scenario SSP3-7.0.

Even for the contemporary AIS climate, significant differences in near-surface air temperature, snowfall and melt exist between these three RCMs (Mottram et al., 2021; Carter et al., 2022). An important difference between the RCMs lies in the handling of precipitation advection: MAR allows precipitation to be advected through the atmospheric layers until reaching the surface, whereas RACMO and HIRHAM deposit precipitation instantaneously. Furthermore, RACMO and MAR use sub-surface schemes optimized over snow and ice for Antarctica, while HIRHAM applies simpler surface snow physics for ice surfaces. In addition, notable climatic differences between the GCM simulations exists; for instance, the 21st century AIS warming for the four selected GCMs for MAR for SSP5-8.5 ranges from +3.2 K (NorESM1-M) to +8.5 K (CNRM-CM6-1) (Kittel et al., 2021).

3 Emulator tuning and evaluation

3.1 Emulator tuning

The eight regions used for the spatial blocking cross-validation are shown in Fig. 1. The optimal hyperparameters obtained during the spatial blocking cross-validation are listed in Table 2. We found that removing the 30-year average total annual liquid water input (surface melt plus rainfall) and surface temperature input features improved the performance. Therefore, we leave out the 30-year total annual liquid water input feature. The 30-year surface temperatures were replaced by the 30-year

Table 1. List of available forcing datasets of regional climate models (RCMs) driven by global climate models (GCMs) along with the GCM era, GCM ensemble member number, RCM horizontal resolution and the available scenarios. We use the period 1980-2100 of all forcing datasets.

RCM	GCM	GCM era	GCM member	RCM horizontal resolution	Scenarios
RACMO	CESM2	CMIP6	r11i1p1f1	27 km	SSP1-2.6, SSP2-4.5, SSP5-8.5
MAR	ACCESS-1.3	CMIP5	r1i1p1	35 km	SSP5-8.5
MAR	CESM2	CMIP6	r11i1p1f1	35 km	SSP1-2.6, SSP2-4.5, SSP5-8.5
MAR	NorESM1-M	CMIP5	r1i1p1	35 km	SSP5-8.5
MAR	CNRM-CM6-1	CMIP6	r1i1p1f2	35 km	SSP5-8.5
HIRHAM	CESM2	CMIP6	r11i1p1f1	12 km	SSP1-2.6, SSP5-8.5
HIRHAM	EC-Earth3	CMIP6	r5i1p1f1	12 km	SSP5-8.5

220 averaged temperature trend. This led to a small improvement in the performance (3% less false negatives and 1% less false positives).

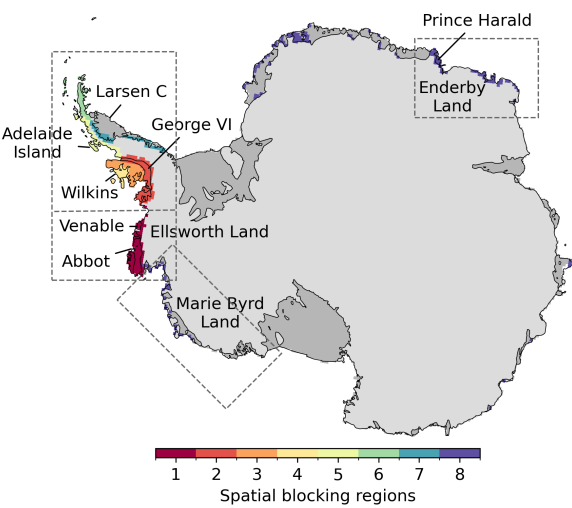


Figure 1. Regions used for spatial blocking cross-validation. The names indicate locations referred to in the text and the dashed boxes indicate the regions in Figs. 3, 5, 7, 8, 10, S1 and S4.

3.2 Emulator evaluation

During the spatial blocking cross-validation, the emulator explains 89% of the variance in the IMAU-FDM perennial liquid water content (LWC) ($R^2 = 0.89$; Fig. 2a). The RMSE is 86 (mm or days for positive values, days for negative values), and

Table 2. Hyperparameter ranges used for tuning and their optimal values.

Parameter	Tuning range	Optimal value
n_estimators	0-300	179
learning_rate	0.001-1	0.034
min_child_weight	0-50	44
max_depth	2-15	6
subsample	0.5-1	0.58
colsample_bytree	0.5-1	0.77
gamma	0-10	4.2
reg_lambda	0-10	2.8

the bias is 0.4 (mm ~~or days~~ for positive values, days for negative). However, in this work we aim to predict the presence or absence of aquifers rather than perennial LWC. When employing a positive perennial LWC threshold to identify aquifers, the emulator successfully predicts 88% of the IMAU-FDM aquifers (12% false negatives) with 9% false positives. The emulator yields an R^2 value of 0.90 for the years with perennial firm aquifers (PFAs) of each grid point (Fig. 2c). Maps of PFA years of the Antarctic Peninsula and Ellsworth Land for the spatial blocking for SSP5-8.5 are shown in Fig. 3. Although Figure 2c indicates that there is some spread and outliers, Figure 3 demonstrates that the firm model and emulator largely agree on the firm aquifer distribution, and the differences are rather subtle. The performance, while using spatial blocking, is poorest on and around the Larsen C and D ice shelves (region 7, RMSE = 8 years), with under- and overestimations. When randomly, instead of strategically, splitting our training and testing data, the emulator yields a misleadingly high R^2 value of 0.98, which highlights the importance of ~~using~~ strategically splitting the data to prevent overfitting.

During the validation in which SSP2-4.5 is left out, the emulator yields an R^2 value of 0.96 for both the perennial LWC and the years with PFAs (Figs. 2b,d). For this, we include all data of all locations where either the emulator or the model forecasts the presence of an aquifer or predicts fewer than 150 days without LWC in at least one year. The emulator successfully predicts 88% of the IMAU-FDM aquifer years of the SSP2-4.5 scenario (12% false negatives), with 7% false positives. The results for leaving out SSP1-2.6 are similar, with the emulator predicting 92% of the aquifer years with 11% false positives ($R^2 = 0.96$ and 0.97 for perennial LWC and years with PFAs, respectively). When leaving out the SSP5-8.5 scenario, the emulator predicts 89% of the aquifer years, with 15% false positives ($R^2 = 0.9$ and 0.91 for perennial LWC and years with PFAs, respectively). The larger error can be attributed to the large differences in warming trend.

Four example time series from the spatial blocking and scenario blocking evaluations are shown in Fig. 4. These figures show that the emulator can mimic the growth as well as the shrinking of aquifers. They also illustrate that deviations from the firm model mainly occur during the onset of ~~aquifers~~ aquifer formation, when aquifers are relatively small.

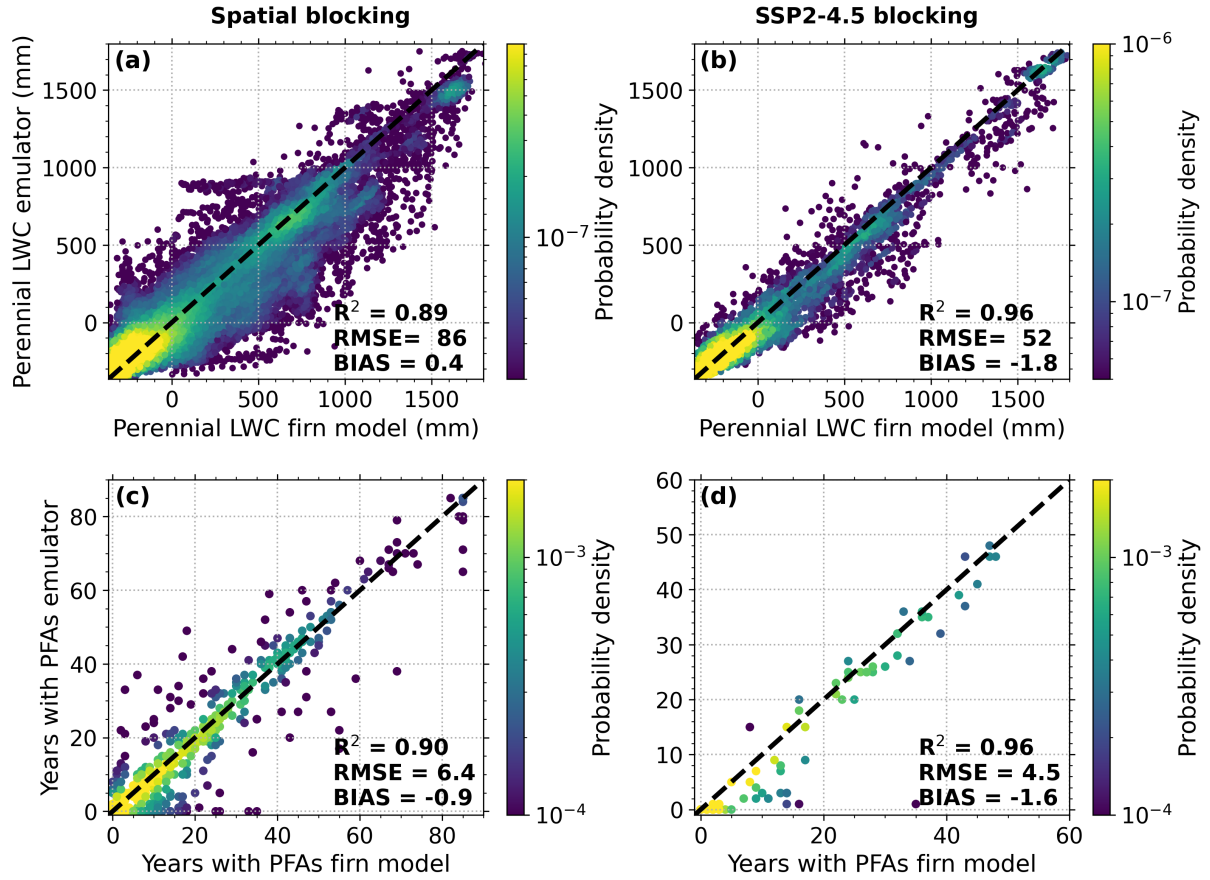


Figure 2. (a,b) Perennial liquid water content (LWC) and (c,d) number of years with perennial firm aquifers (PFAs) from the (a,b) spatial blocking cross-validation and (c,d) SSP2-4.5 block validation. Panels (a) and (b) include the data points used for training and evaluation described in Sect. 2.2.2. Panels (c) and (d) include all locations where either the emulator or the model forecasts the presence of an aquifer or predicts fewer than 150 days without LWC in at least one year in the associated scenario.

4 Results

4.1 ~~PFAs~~PFA predictions under scenarios SSP1-2.6 and SSP2-4.5

The trained emulator is used to predict perennial firm aquifers (PFAs) over the period 2015-2100 for the 12 forcings listed in Table 1. It should be noted that all available SSP1-2.6 and SSP2-4.5 regional climate model (RCM) forcings are from CESM2. For SSP1-2.6 and SSP2-4.5, PFAs are projected to remain mostly restricted to the Antarctic Peninsula (AP), except for expansion to Enderby Land (East Antarctica) under SSP2-4.5 for MAR-CESM (Fig. 5). The emulator predicts PFAs along the north-western coast of the AP for all forcings. Additionally, the emulator predicts PFAs on Wilkins ice shelf for SSP1-2.6 and SSP2-4.5, and on George VI ice shelves for SSP2-4.5 for RACMO-CESM forcings. Emulator predictions for MAR-CESM

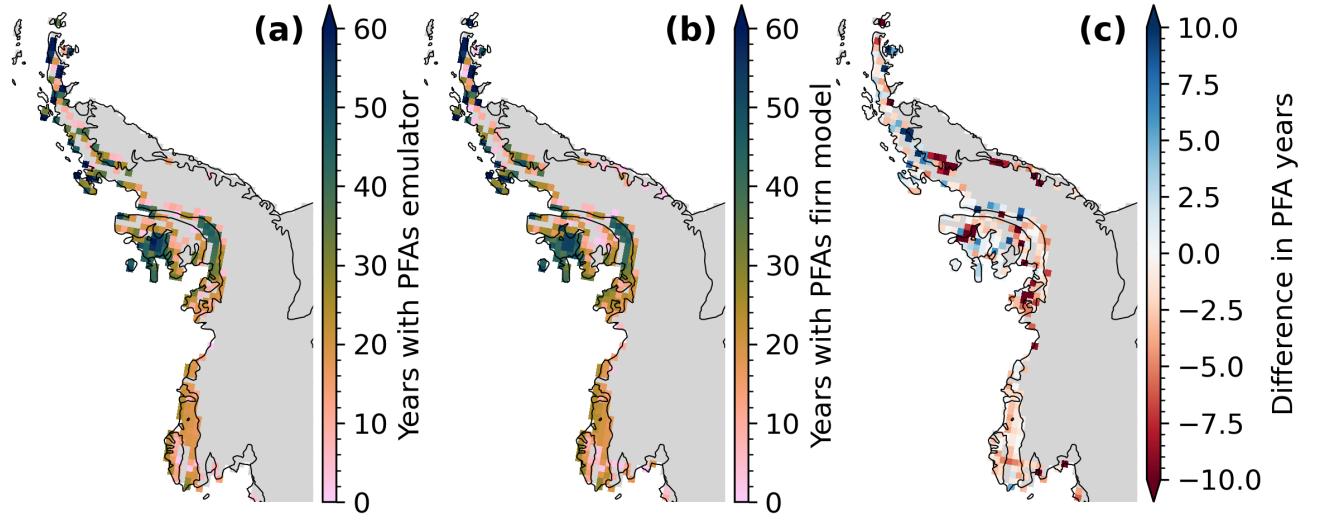


Figure 3. (a) Number of years with perennial firn aquifers (PFAs) under SSP5-8.5 from the emulator during the spatial blocking validation and (b) from the IMAU-FDM firn model and (c) the difference. The location of the depicted region is indicated in Fig. 1.

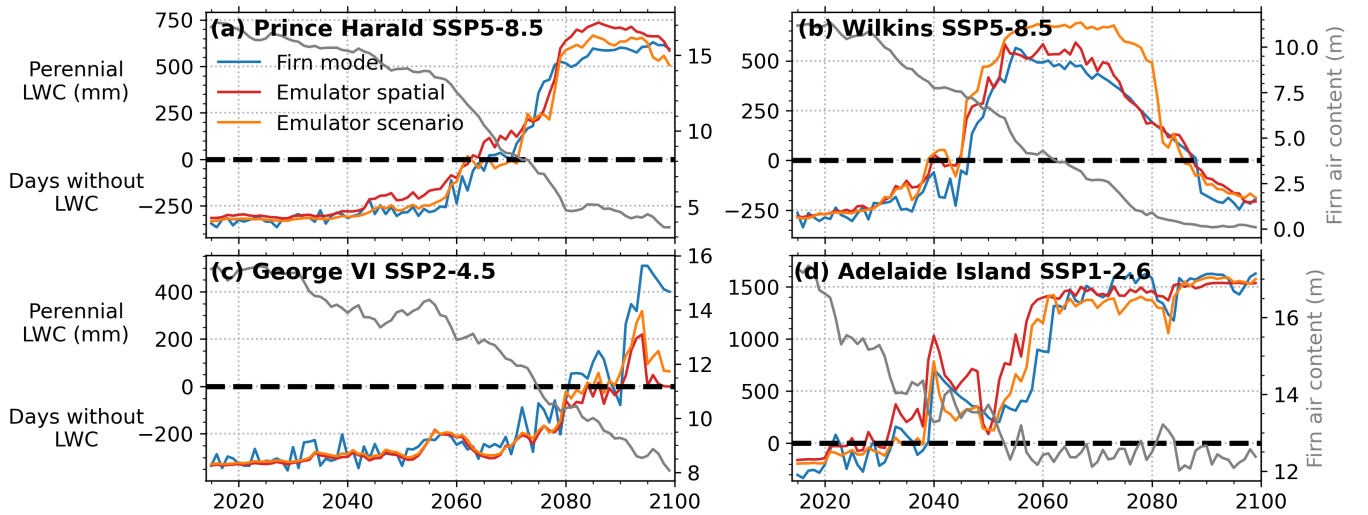


Figure 4. Example time series of perennial liquid water content (LWC)/days without LWC for the firm model (blue) and the emulator for spatial blocking (red) and for scenario blocking (orange), as well as firn air content for the firm model (gray), for individual grid points and different scenarios. (a) and (b) are SSP5-8.5, (c) is SSP2-4.5, and (d) is SSP1.2-6. The locations of the grid points are indicated in Fig. 1.

show aquifers on Wilkins and George VI ice shelves for SSP2-4.5, and both MAR forcings predict aquifers along the grounding
 255 lines of Larsen C ice shelf. Figure 6a shows time series of PFA extent for all SSP1-2.6 and SSP2-4.5 emulator simulations. As
 all RCMs are forced with the same two CESM2 realisations, peaks and minima in modelled PFA extent align. The emulator

forced by MAR predicts the highest PFA extent by 2100 for SSP1-2.6 (50,000 km²) and SSP2-4.5 (87,000 km²), followed by RACMO (31,000 and 43,000 km²), while HIRHAM-CESM only predicts a PFA extent of 19,000 km² for SSP1-2.6. Part of the difference between the RCMs arises by from the distinct horizontal resolutions. Overall, PFA expansion starts to accelerate around 2050 in these runs.

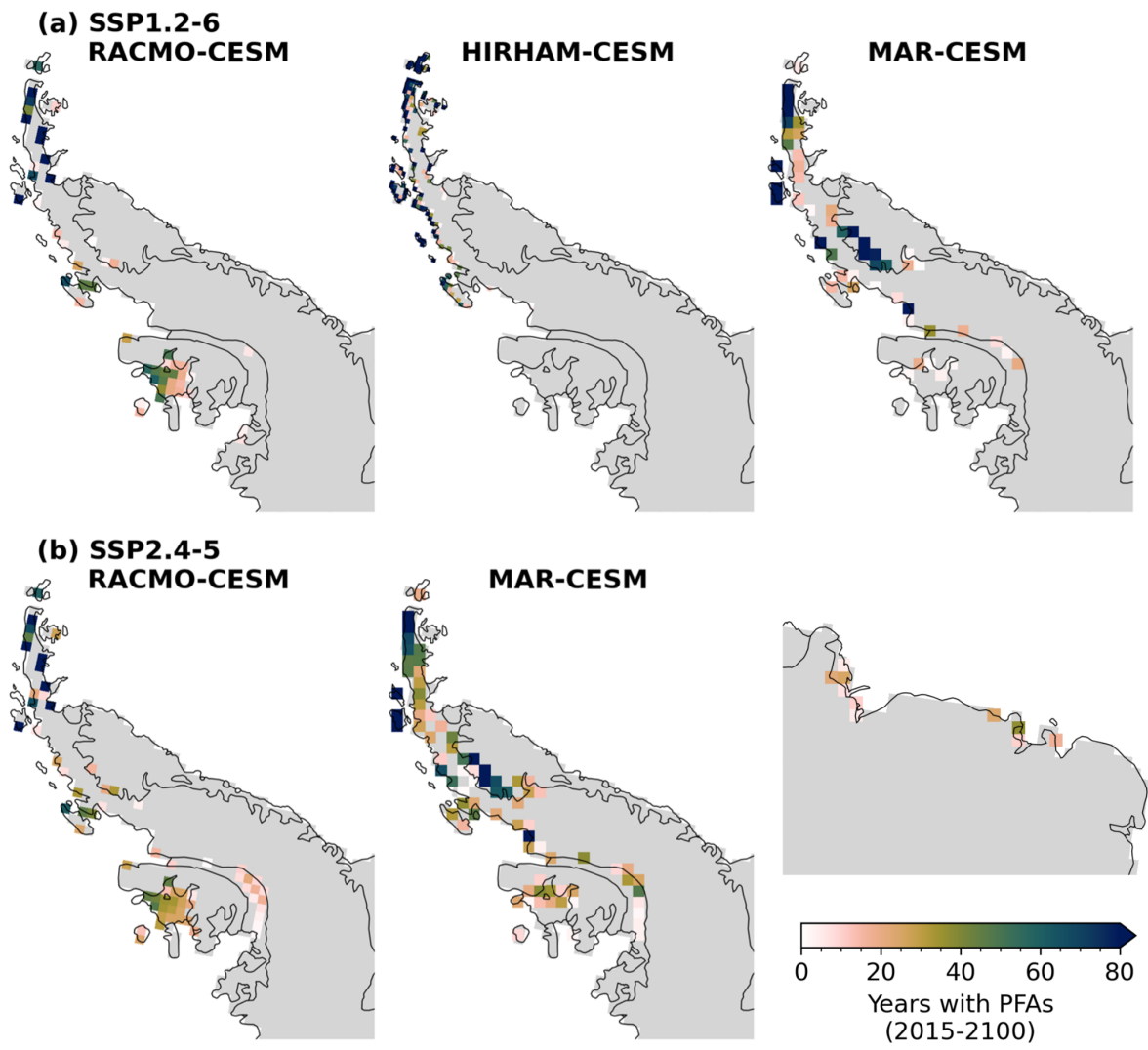


Figure 5. Number of years with perennial firn aquifers (PFAs) simulated by the emulator for (a) SSP1-2.6 and (b) SSP2-4.5 forcing datasets over the period 2015-2100 for the Antarctic Peninsula. The MAR-CESM SSP2-4.5 forcing also includes Enderby Land. The locations of the depicted regions are indicates in Fig. 1.

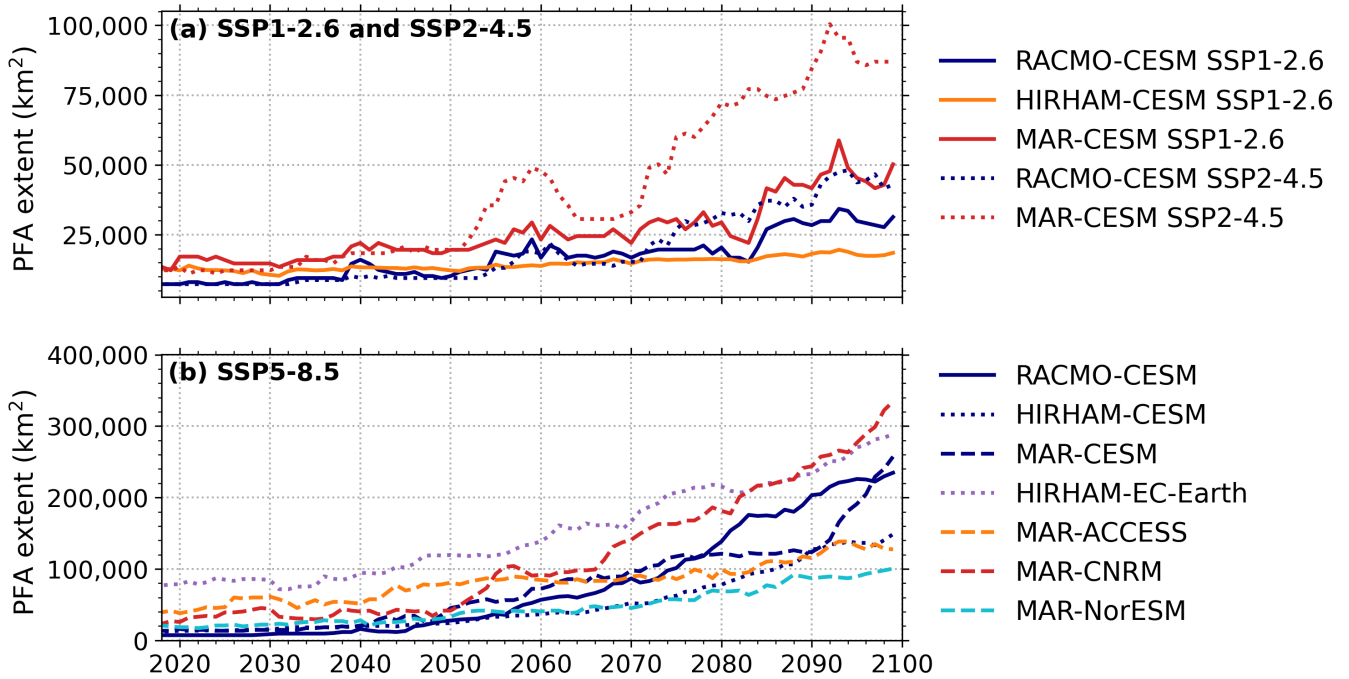


Figure 6. Time series of perennial firn aquifer (PFA) extent simulated by the emulator for all climatic forcing datasets for (a) SSP1-2.6 and SSP2-4.5 and (b) SSP5-8.5 scenarios.

4.2 PFAs predictions under scenario SSP5-8.5

For SSP5-8.5, the emulator predicts that PFAs expand within the AP, and to Ellsworth Land, on and around Abbot and Venable ice shelves, except for MAR-NorESM, which only expands within the AP (Fig. 7). The emulator predicts aquifers along the north-western coast of the AP, and along the grounding lines of Larsen C, George VI and Wilkins ice shelves for all forcings. In addition, the simulation forced by RACMO-CESM stands out, as it predicts extensive PFA coverage on Wilkins, George VI and Abbot ice shelves (between 70-100% of the ice shelf area). In contrast, all MAR and HIRHAM-EC-Earth driven simulations show extensive PFA coverage on Larsen C ice shelf ($>12,000 \text{ km}^2$). PFAs are also predicted for 5 out of the 7 forcings in Enderby Land in East Antarctica (Fig. 8a). MAR-ACCESS and MAR-NorESM are the two forcings for which PFAs are not predicted in this region, which also yields the lowest overall PFAs extent ($100,000 \text{ km}^2$ and $127,000 \text{ km}^2$ by 2100, respectively, Fig. 6b). For HIRHAM-EC-Earth, and MAR-CNRM, the emulator predicts PFAs in Marie Byrd Land. For MAR-CNRM the largest PFA extent by 2100 is predicted to be $334,000 \text{ km}^2$, followed by HIRHAM-Ec-Earth, $289,000 \text{ km}^2$. Notably, for HIRHAM-EC-Earth the emulator predicts at least twice the initial (2015) PFA extent ($72,000 \text{ km}^2$) compared to the other simulations, related to the higher initial temperatures (see Section 4.3). Using the same global climate model (GCM) forcing (CESM2), RACMO and MAR yield comparable PFA extent ($257,000$ and $234,000 \text{ km}^2$), while HIRHAM-CESM predicts a considerably lower extent ($148,000 \text{ km}^2$). On the other hand, for SSP1-2.6 and SSP2-4.5, for MAR the emulator yields higher

PFA a higher PFA extent compared to RACMO for the same GCM forcing. PFA expansion starts to accelerate around 2050 for the CESM2 and CNRM-CM6 forcings, similarly as for SSP1-2.6 and SSP2-4.5 (Fig. 6).

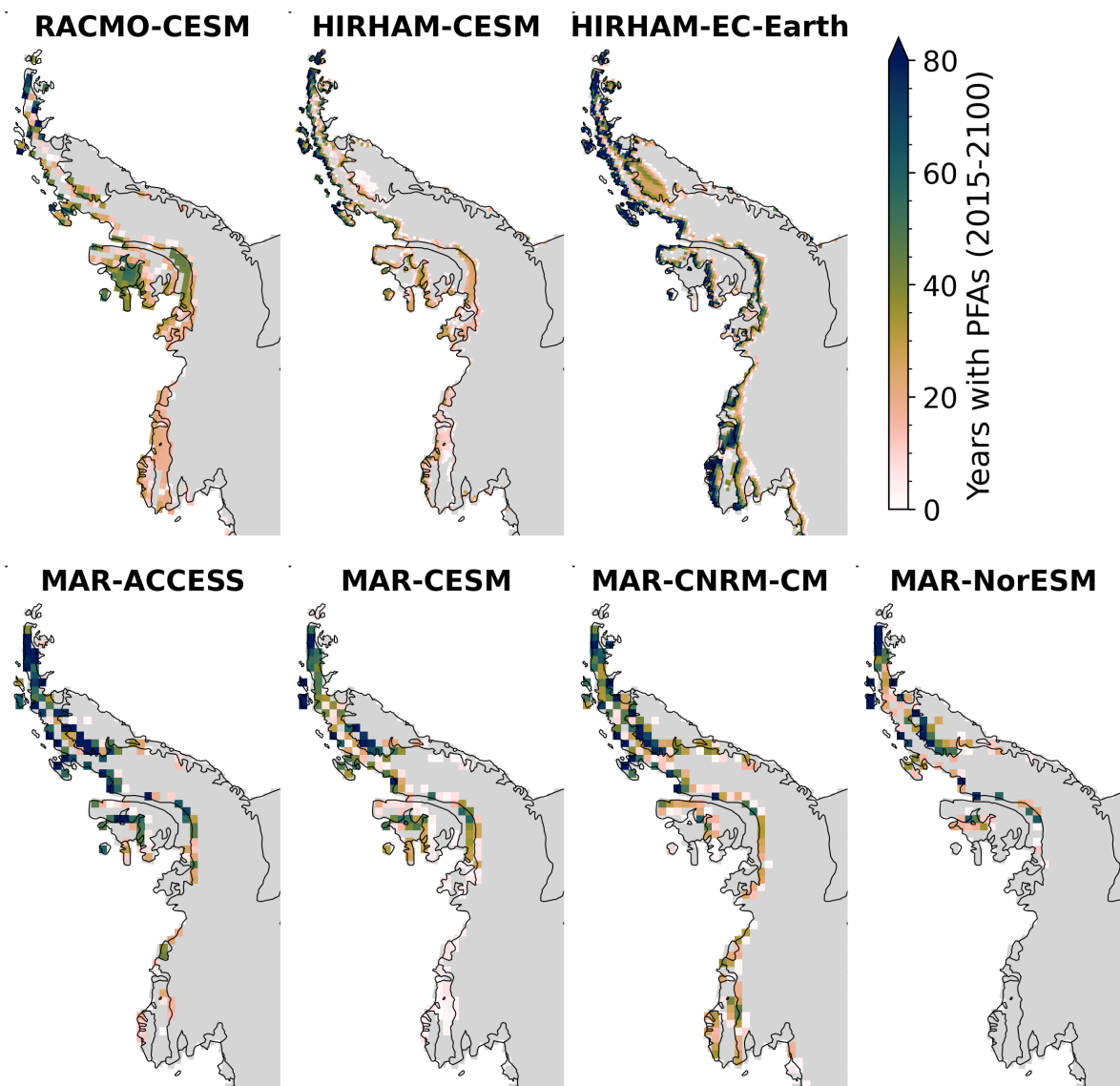


Figure 7. Number of years with perennial firn aquifers (PFAs) simulated by the emulator for SSP5-8.5 forcing datasets for the Antarctic Peninsula and Ellsworth Land. The location of the depicted region is indicated in Fig. 1.

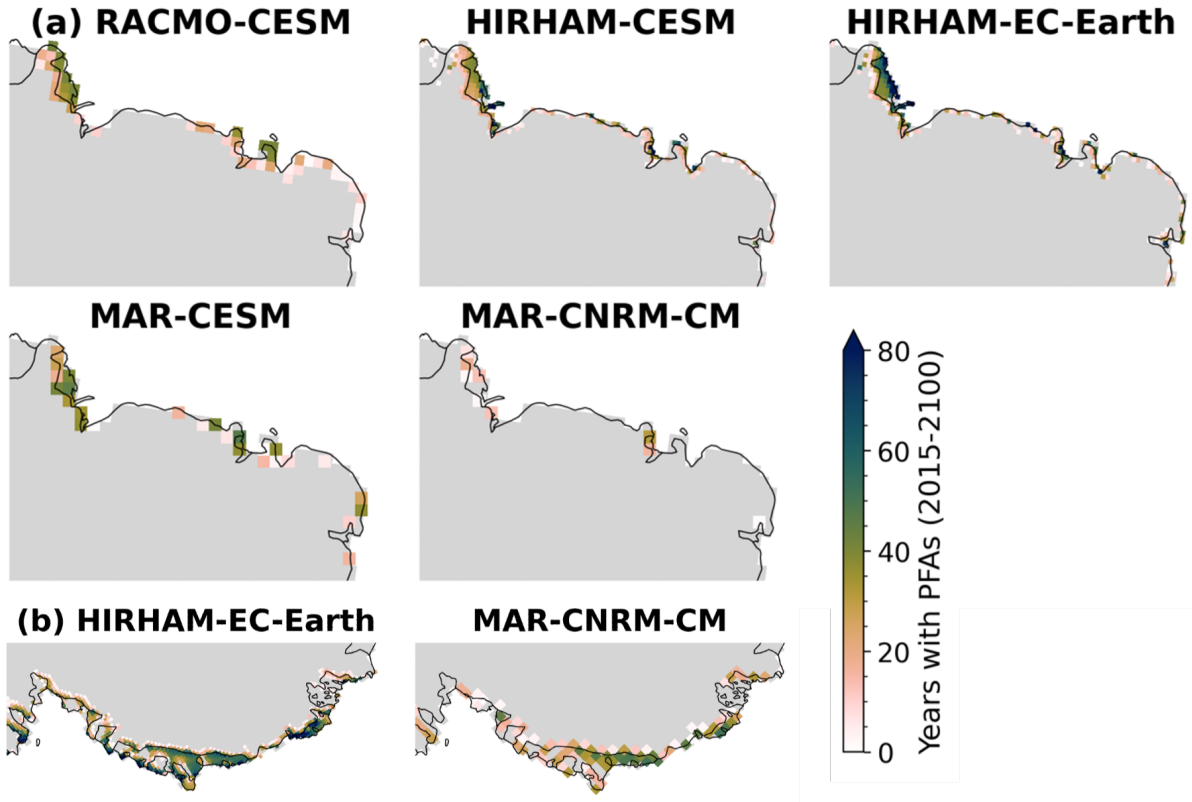


Figure 8. Number of years with perennial firn aquifers (PFAs) simulated by the emulator for SSP5-8.5 forcing datasets for (a) Enderby Land and (b) Marie Byrd Land. The locations of the depicted regions are indicated in Fig. 1.

4.3 Climatic drivers

The previous sections show a large range of PFA predictions by the emulator, caused by differences in the climatic forcing. Firstly, differences in initial surface temperature and warming between the forcing datasets explain part of the spread (Fig. 9b). For example, for MAR-NorESM the smallest PFA extent for SSP5-8.5 by 2100 is predicted, as it has the smallest AIS end-of-century (2090-2100) warming (+3.3 K) compared to 2005-2015, and lowest end-of-century surface temperature (241.6 K). On the other hand, HIRHAM-EC-Earth and MAR-CNRM-CM have the highest temperatures by 2090-2100 for SSP5-8.5 (245.9 and 244.9 K, respectively) with a warming of +4.6 and +6.6 K, respectively, and the emulator predicts the largest PFA extent. For HIRHAM-EC-Earth, the initial temperature is high, related to the warm bias over Antarctica in EC-Earth3 (Boberg et al., 2022), which explains the high PFA extent at the start of the emulator simulation. For SSP1-2.6 and SSP2-4.5, the largest PFA extent is predicted for MAR-CESM, which is also the warmest model combination (Fig. 9a).

Significant regional differences also exist between RCMs for the same GCM forcing, in this case CESM2. For example, on Larsen C ice shelf the emulator predicts more aquifers for MAR compared to RACMO and HIRHAM, which is related

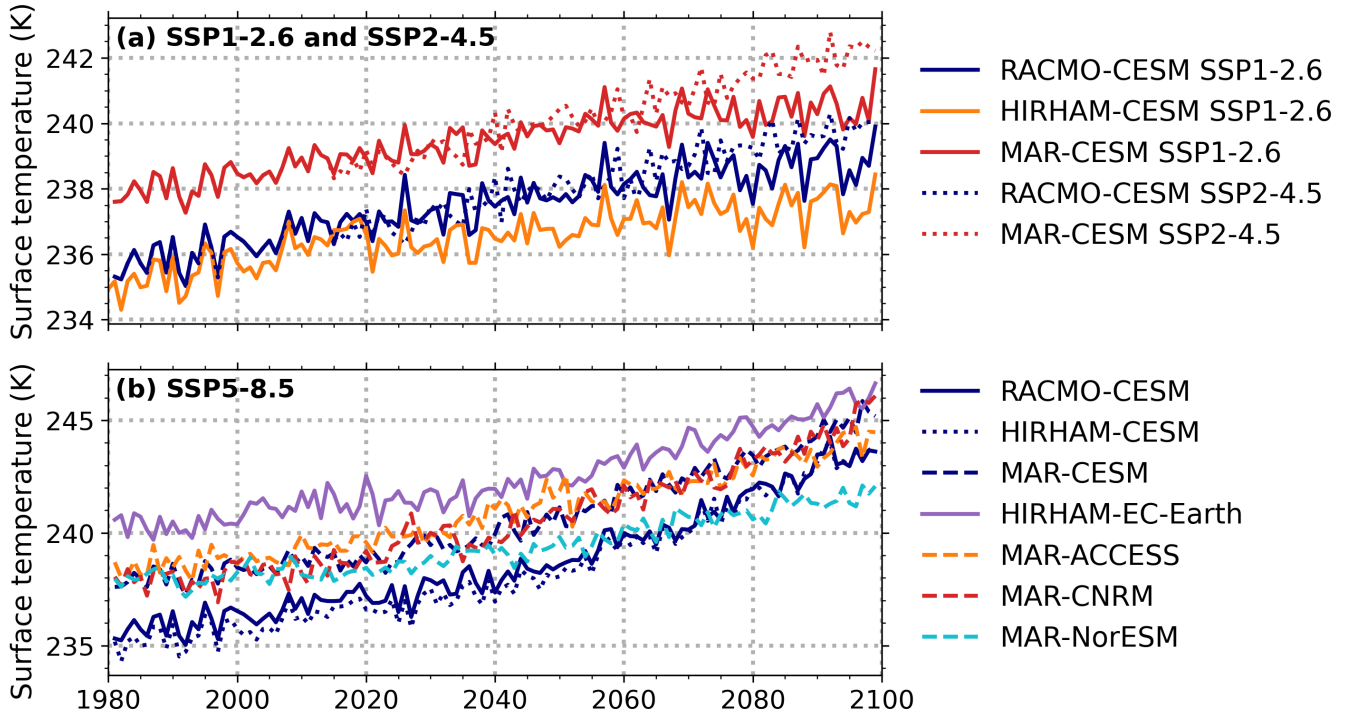


Figure 9. Time series of annual average AIS surface temperature for all climatic forcing datasets for (a) SSP1-2.6 and SSP2-4.5 and (b) SSP5-8.5 scenarios.

to high surface melt and high accumulation at the foot of the mountains, which are absent in the other two regional climate models (Fig. 10; see Fig. S2 for surface temperature). On the other hand, the absence of extensive PFAs on Wilkins and George VI ice shelves for MAR and HIRHAM is due to lower accumulation compared to RACMO. Furthermore, HIRHAM-CESM calculates lower surface melt rates than the other two models, despite mean surface temperatures in HIRHAM-CESM being comparable to those in RACMO-CESM.

Figure 11a shows that the presence of PFAs for RACMO and MAR driven emulator simulations is strongly governed by the rates of accumulation and melt (plus rain), here shown as averages for the preceding 10 years. It shows that at least 300 mm w.e. yr^{-1} of surface melt and 600 mm w.e. yr^{-1} of accumulation are required for aquifers to form. When melt and accumulation exceed 600 mm w.e. yr^{-1} and 1000 mm w.e. yr^{-1} , respectively, an aquifer is predicted in nearly all cases, in line with Brils et al. (2024). The same figure for HIRHAM (Fig. S2a) suggests that PFAs also form at slightly lower surface melt and accumulation rates. This is because HIRHAM models considerably less surface melt and accumulation for a given temperature than MAR and RACMO (Fig. S3). The results of the emulator for HIRHAM appear thus less reliable, as it has a different modelled relations between temperature, accumulation and melt compared to RACMO. Therefore, in the remainder we focus on the emulator results from RACMO and MAR forcings.

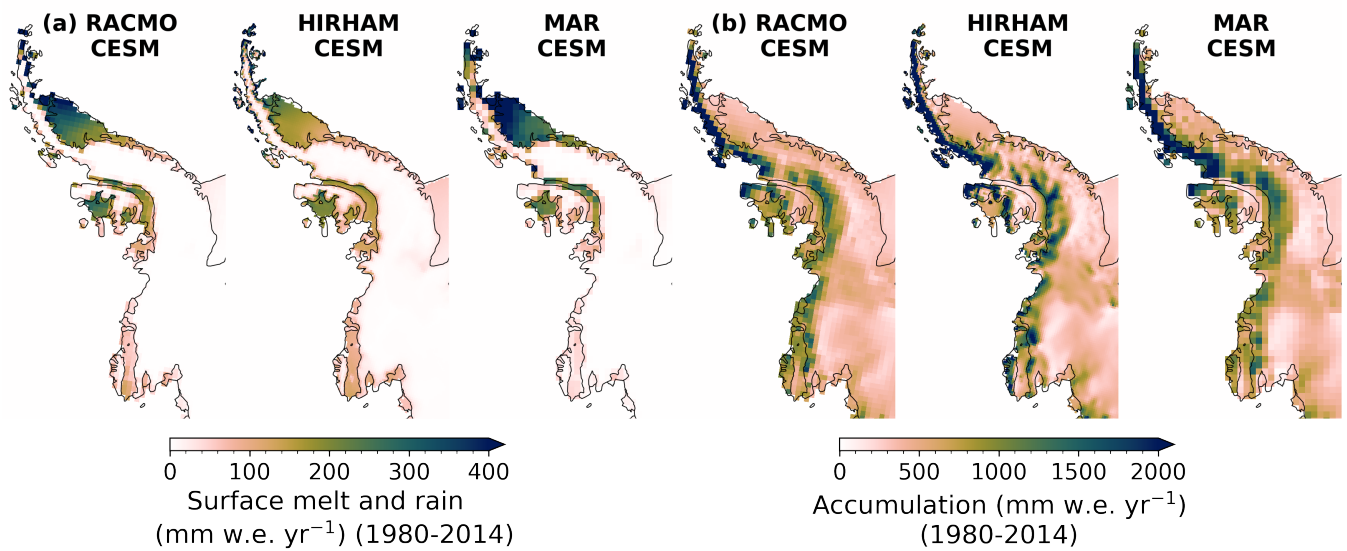


Figure 10. Annual average (a) surface melt and rainfall, and (b) snow accumulation over the historical period (1980-2014) for RACMO-CESM, HIRHAM-CESM and MAR-CESM. The location of the depicted region is indicated in Fig. 1.

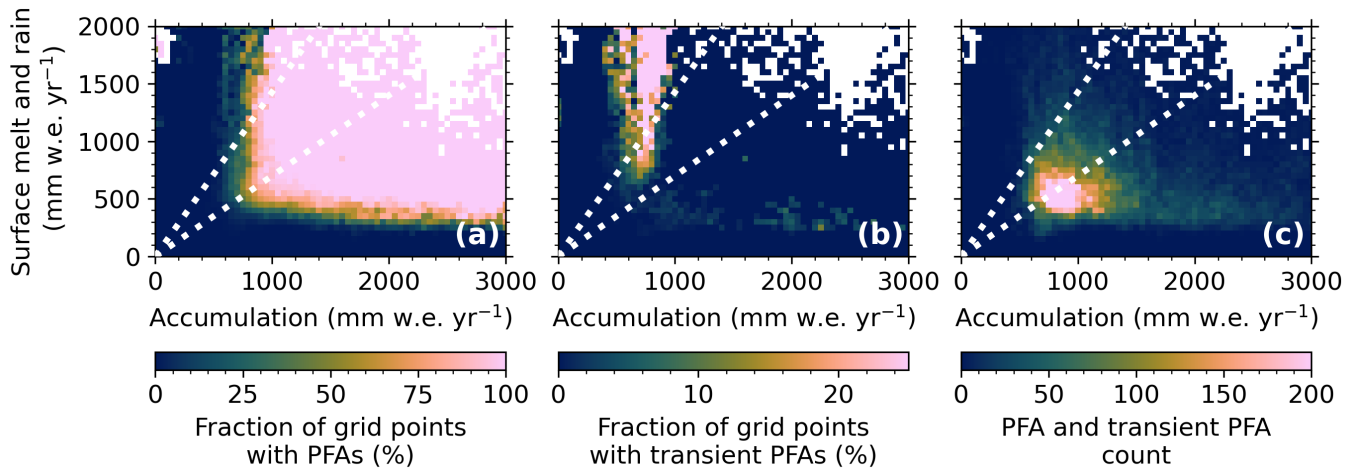


Figure 11. Percentage of (a) perennial firm aquifer (PFA) occurrence and (b) transient PFA occurrence, and (c) total amount of PFA and transient PFA occurrence, simulated by the emulator for all RACMO and MAR forcings, as a function of annual surface melt and rainfall (y-axis) and snow accumulation (x-axis). The grid cells are grouped in melt and accumulation bins of 50 mm w.e. yr⁻¹. The dotted lines indicate the 0.7 and 1.7 Melt-over-Accumulation (MoA) thresholds.

4.4 Transient PFAs

305 As Figure 4b shows, PFAs can also develop, shrink and subsequently disappear, ~~henceforth~~ and henceforth will be referred to as transient PFAs. This shrinkage and disappearance occurs when the firn layer becomes too thin (i.e. the firn air too depleted) to host an aquifer. A transient PFA is defined as a PFA that existed for at least 5 years but has disappeared before 2100. Figure 11b shows the presence of transient PFAs simulated with the emulator, as a function of melt (plus rain) against accumulation rates for RACMO and MAR simulations. The white dotted lines indicate the Melt-over-Accumulation (MoA) ratios of 0.7 and 310 1.7, which are theoretical thresholds for indicating ice slabs and ablation zone, respectively (Brils et al., 2024). Transient PFAs mainly occur once the 0.7 and especially the 1.7 threshold have been exceeded, and mainly below accumulation rates of 1000 mm w.e. yr^{-1} .

Figure S4 shows the locations of these transient PFAs for the RACMO and MAR (SSP5-8.5) driven emulator simulations. Transient PFAs mainly occur along the boundaries of PFA regions. Most notably, transient PFAs occur on Wilkins ice shelf 315 for RACMO-CESM for the SSP5-8.5 scenarios. Transient PFAs mainly occur under SSP5-8.5, covering on average 28,000 km^2 by 2100, which is 13% compared to the 2100 PFA extent (Fig. 11). In the SSP1-2.6 and SSP2-4.5 scenarios the extent of transient PFAs is on average only 4,200 km^2 , 9% of the 2100 PFA extent.

5 Discussion

5.1 Firn model and observations

320 Since our emulator is trained on IMAU-FDM data, the uncertainties inherent to the firn model will also persist in the emulator. High surface melt and snowfall conditions are favorable for perennial firn aquifer (PFA) formation, however some processes, which are not included in IMAU-FDM can inhibit their formation despite these conditions. For instance, firn aquifers are known to drain in heavily crevassed areas (Poinar et al., 2017), a process that is not included in the firn model. In addition, ice lenses can also limit vertical meltwater percolation to deeper firn, thereby increasing refreezing (e.g., Culberg et al., 2021), although 325 they generally form under lower accumulation conditions (Brils et al., 2024). The bucket method does not include preferential flow, which can lead to an underestimation of the percolation depth of liquid water and therefore aquifer recharge as observed in Greenland (Miller et al., 2018; Vandecrux et al., 2020). The firn model also does not allow for lateral meltwater transport, which has been observed within aquifers, but with a limited measured specific discharge (60–140 m yr^{-1}) (Montgomery et al., 2020; Miller et al., 2018). On the other hand, not including lateral meltwater transport in streams can also result in underestimating 330 aquifer presence. For instance, on the firn-terminating Priestley Glacier, meltwater flows down the ice shelf across blue ice into the section of the ice shelf with thicker surface snow, where it likely saturates the firn and feeds firn aquifers (Bell et al., 2018). Similar firn-terminating streams have been observed on the Amery and Nivilsen ice shelves (e.g., Fricker et al., 2002; Kingslake et al., 2015; Leppäranta et al., 2013).

In addition, the use of irreducible water content hinders the ability to estimate the volume of meltwater stored within a PFA. 335 Nonetheless, the bucket method is used as it is a fast and stable meltwater scheme, in contrast to many Richards' equation-

based models that are computationally more expensive and prone to numerical instabilities. While it is impossible to estimate the volume of meltwater stored within a PFA with the bucket method, the presence of year-round liquid water in the firn can be used as an indication of PFA presence (Kuipers Munneke et al., 2014). sBrils et al. (2024) simulated aquifers in Greenland using IMAU-FDM, agreeing with 62% of the observed aquifers by airborne and ground-penetrating radar measurements. The mismatch can be explained by ~~underestimation of the~~ underestimation of observationally detected aquifers by e.g. widely spaced radar data, resolution and process limitations of the firn model, ~~drainage to e.g.~~ drainage of meltwater into crevasses that is not included in the model, ~~in addition to the above-described~~ the above-described firn model limitations, and/or biases in the climatic forcing.

In Antarctica, PFAs have thus far only been observed in situ on the Wilkins and Müller ice shelves (Montgomery et al., 2020; MacDonell et al., 2021). Using an integrated approach combining satellite and regional climate model (RCM) data, PFAs are also predicted along the north-western coast of the Antarctic Peninsula (AP) and on George VI ice shelf (Di Biase et al., 2024). PFAs along the north-western coast, indicated by the satellite as well as the RCM data, agree with our emulator results. However, historical PFA extent on Wilkins and George VI ice shelves appears underestimated in the emulator simulations. In contrast, on Larsen C ice shelf current PFAs are simulated by the emulator for all MAR forcing datasets, while PFA probability is low here (Di Biase et al., 2024). Müller ice shelf is located in a region with complex topography, which is not resolved in the RCMs. Overall, the modelled historical PFA extent shows moderate agreement with observations, especially for intermediate accumulation and surface melt rates. As future warming leads to increased melt and accumulation, the emulator is expected to produce more accurate PFA predictions compared to the contemporary estimates. This is because warming is likely to shift more locations into well-defined PFA climate regimes characterized by high melt and accumulation, where aquifer formation is highly certain, rather than into regimes with lower melt and accumulation, where aquifer formation is less consistent (Fig. 11).

5.2 XGBoost emulator

Our PFA emulator captures at least 89% of the simulated PFA variance. Despite its demonstrated generalisation capabilities, we emphasize that XGBoost models can be prone to errors when used beyond their range of training conditions (Chen and Guestrin, 2016). Given the large spatial variability in climatic conditions across the AIS, the range of the individual input features are likely covered in the training data. However, our findings suggest that some specific combinations of these features may be absent in the training data. This is illustrated by the lower melt and accumulation for a given temperature in HIRHAM compared to RACMO and MAR. As a result, the emulator predicts PFAs for HIRHAM forcings under unrealistically low melt and accumulation rates (Brils et al., 2024; van Wessem et al., 2021). RACMO and MAR, are in better agreement on the conditions of PFA formation, reinforcing our confidence in the emulator's performance. However, if a different RCM is used for either training or predicting, the consistency between the RCMs should be evaluated to ensure reliable results. Additionally, our work highlights the importance of using spatial and scenario blocking instead of randomly splitting training and testing data. In the latter approach, the training and testing data are very similar, while with spatial or scenario blocking, XGBoost is

trained and tested on datasets with slightly different characteristics. As a result, overfitting and overoptimistic perception of the
370 predictive power is prevented (Roberts et al., 2017).

5.3 PFA expansion uncertainties

The emulator predicts a large range of possible PFA distributions, caused by differences in the climatic forcing. This highlights the usefulness of the emulator as it allows incorporating a wide range of climate forcings to better capture the variability and uncertainty. Overall, PFA expansion responds primarily to the modelled warming, either by a higher emission scenario, or by
375 a warmer model. The Larsen C ice shelf stands out in this regard, as all emulator simulations for MAR clearly predict PFAs here, also in colder runs, such as MAR-norESM (Figs. 5, 7). Generally, PFAs occur on the windward side of Antarctic and Greenland mountain ranges, and receive moist and mild air masses originating from the sea, resulting in favorable conditions for PFA formation (Turner et al., 2019). The Larsen C ice shelf is considered too dry, as it is located on the leeward side of the AP mountain range. This ice shelf is frequently affected by föhn winds, resulting in dry and warm conditions (Luckman
380 et al., 2014). Nevertheless, substantial PFAs are predicted here by the emulator for MAR, as this is the only RCM that allows precipitation to be advected over the mountains before reaching the surface. Advection of precipitation is also included in the recently developed RACMO2.4 version (van Dalum et al., 2024). However, as described in Section 5.1, MAR overestimates current PFA extent on Larsen C ice shelf, suggesting that advection of precipitation might be overestimated. In addition, the coarse resolution of MAR (35 km), contributes to this effect, as blocking of humid air is underestimated (Datta et al., 2018).
385 This in turn, can cause overestimation of clouds, precipitation, longwave radiation and consequently melt on Larsen C ice shelf.

The horizontal resolutions of all forcing datasets and thus emulator results ranges from 12 to 35 km. However, Van Wessem et al. (2016) suggest that even at 5.5 km resolution, the underestimation of the height and slope of the orographic barrier may result in an underestimation of precipitation and föhn winds in the AP. Another uncertainty concerning the emulator results is that all SSP1-2.6 and SSP2-4.5 simulations are indirectly driven by CESM2, which exhibits a high climate sensitivity, along
390 with a cold bias and a high precipitation bias over the AIS (Dunmire et al., 2022). However, these biases do not necessarily persist in the downscaled RCM output (Veldhuijsen et al., 2024d).

5.4 Implication for future ice-shelf and ice-sheet stability

While there is a substantial spread in PFA predictions by the emulator between the climate models, some general patterns emerge. Under SSP1-2.6 and SSP2-4.5, the emulator predicts that PFAs remain restricted to but expand within the AP, the
395 region where aquifers are currently present. For SSP5-8.5, aquifer expansion is more pronounced within the AP. In addition, for SSP5-8.5, PFAs expand to Ellsworth Land in six out of the seven simulations and to Enderby Land in East Antarctica in five out of the seven simulations. This expansion within the AP can potentially contribute to ice-shelf collapse in this region, also for SSP1-2.6 and SSP2-4.5. Significant tensile stresses observed on the Wilkins and Larsen C ice shelves (Lai et al., 2020) suggest that these areas may be vulnerable to further instability caused by PFAs. Although the PFA expansion on Larsen C ice shelf by
400 MAR might be overestimated, as explained in Section 5.3. The stress regime on George VI suggests it is currently resilient to hydrofracturing. In Ellsworth Land, substantial tensile stresses on the Abbot ice shelf suggest that PFAs under SSP5-8.5 could

potentially induce hydrofracturing here. Similarly, ice shelves in Enderby Land are vulnerable to collapse, though they do not provide substantial buttressing (Lai et al., 2020). Furthermore, PFA expansion on the grounded ice, which mainly occurs in the northwestern AP, could cause meltwater drainage through crevasses to the bed (Poinar et al., 2017). This drainage may result in enhanced basal lubrication, increasing ice velocity and ice discharge into the ocean (Zwally et al., 2002).

As new PFAs are developing, the firm air content is also decreasing (Fig. 4). If warming continues, the firm air content eventually becomes fully depleted, after which the PFAs disappear, making them transient. Our results only show transient PFAs along the boundaries of PFA regions and on Wilkins ice shelf in SSP5-8.5 (Fig. S4), but this is expected to increase after 2100. In fact, explicit firn simulations for Greenland suggest that PFAs are a committed transition state from healthy firn to depleted firn in high accumulation regions ($>1,000 \text{ mm yr}^{-1}$) (Brils et al., 2024). The duration of such transient PFA presence will decrease with stronger warming rates or lower accumulation rates due to relatively quick firn air depletion. Thus, the difference in the timing of ice-shelf vulnerability because of PFAs and firn air depletion becomes smaller. This is illustrated by the stable PFA on Adelaide island (Fig. 4d), which has a high accumulation rate over the historical period (3100 mm yr^{-1}) and low warming rate over the 21st century ($+3.2 \text{ K}$). In contrast, on Wilkins ice shelf, the historical accumulation rate is relatively low (800 mm yr^{-1}), and the warming rate is high ($+7.2 \text{ K}$), resulting in the disappearance of the PFA after 40 years (Fig. 4b). However, the onset of the aquifer still occurs decades earlier than the firn air depletion, highlighting the importance of considering PFAs, rather than solely firn air depletion, when assessing vulnerability to hydrofracturing.

6 Conclusions

An XGBoost machine learning emulator was set up to predict future Antarctic perennial firn aquifers (PFAs) (2015-2100) for an ensemble of 12 ~~scenarios~~ simulations from three regional climate models (RCMs) (RACMO, MAR and HIRHAM) in combination with five global climate models (GCMs). The emulator was trained with simulations of three scenarios (SSP1-2.6, SSP2-4.5 and SSP5-8.5) from IMAU-FDM forced by RACMO-CESM. Using For training and testing, we used a scenario and spatial blocking evaluation approach ~~we found and not random selections of training, testing and evaluation data, as the latter approach is vulnerable for overfitting. We firstly demonstrate~~ that the emulator successfully estimates PFAs, it explains at least 89% of the variance in PFA presence and perennial liquid water content ~~variance. Under.~~ Secondly, we project, based on the available model simulations, that for SSP1-2.6 and SSP2-4.5, PFA presence remains restricted to the Antarctic Peninsula (AP). For SSP5-8.5, PFAs expand within the AP, and expand to Ellsworth Land in six out of the seven simulations, and to Enderby Land in East Antarctica in five out of the seven simulations and to Marie Byrd Land in two out of the seven simulations. ~~There is~~ Lastly, we observe a large spread among the RCMs and GCMs emulator predictions, related to differences in the climatic input, which highlights the necessity of analyzing multiple RCM and ESM combinations and usefulness of the emulator ~~The emulator results for that make such analysis feasible. Further analysis of the results shows that the emulator projects, while using RACMO and MAR agree on the simulations, PFAs for similar surface melt and accumulation conditions required for PFA formation. While these results for HIRHAM. When the emulator is fed with HIRHAM data, these conditions~~ are slightly different, caused by different ~~modelled relations~~ relation between temperature, accumulation and melt in HIRHAM compared

435 to RACMO. ~~Overall~~Concluding, our results show that, irrespective of the emission scenario, firm aquifers are likely to expand in a warmer Antarctica. This highlights the importance to understand the impacts PFAs have on ice sheet hydrology, instability and ice-shelf instability.

Code and data availability. The code of IMAU-FDM v1.2AD is available at: <https://doi.org/10.5281/zenodo.10723570> (Veldhuijsen et al., 2023). The code of the XGBoost model is published at: <https://zenodo.org/records/13750024> (Veldhuijsen et al., 2024a) The emulator results
440 (perennial liquid water content) are published at: <https://zenodo.org/records/13626884> (Veldhuijsen et al., 2024b). The firm model results (perennial liquid water content) are published at: <https://doi.org/10.5281/zenodo.10723570> (Veldhuijsen et al., 2024c).

Author contributions. SbmV and WJvdB defined the research goals and designed the study. SV developed the XGBoost model, performed the simulations and analyzed the results. NH and FB provided HIRHAM data, and CK and CA provided MAR data. All authors contributed to discussions on the manuscript.

445 *Competing interests.* MRvdB is a member of the editorial board of journal The Cryosphere.

Acknowledgements. This work was funded by the Netherlands Organization for Scientific Research (grant no. OCENW.GROOT.2019.091). FBO and NH is supported by the Danish State through the National Centre for Climate Research (NCKF), further NH is also supported by the Novo Nordisk Foundation project PRECISE (NNF23OC0081251). We thank Kamilla Hauknes Sjursen for her valuable feedback on the development procedure of the XGBoost model. MvdB acknowledges support from NESSC. SbmV used ChatGPT to enhance the readability
450 of the manuscript. Following this, the content was reviewed and edited as necessary and full responsibility is taken for the final publication.

References

- Anilkumar, R., Bharti, R., Chutia, D., and Aggarwal, S. P.: Modelling point mass balance for the glaciers of the Central European Alps using machine learning techniques, *The Cryosphere*, 17, 2811–2828, <https://doi.org/10.5194/tc-17-2811-2023>, 2023.
- Arthern, R. J., Vaughan, D. G., Rankin, A. M., Mulvaney, R., and Thomas, E. R.: In situ measurements of Antarctic snow compaction compared with predictions of models, *Journal of Geophysical Research: Earth Surface*, 115, <https://doi.org/10.1029/2009JF001306>, 2010.
- Banwell, A. F., MacAyeal, D. R., and Sergienko, O. V.: Breakup of the Larsen B Ice Shelf triggered by chain reaction drainage of supraglacial lakes, *Geophysical Research Letters*, 40, 5872–5876, <https://doi.org/10.1002/2013GL057694>, 2013.
- Bell, R. E., Banwell, A. F., Trusel, L. D., and Kingslake, J.: Antarctic surface hydrology and impacts on ice-sheet mass balance, *Nature Climate Change*, 8, 1044–1052, <https://doi.org/10.1038/s41558-018-0326-3>, 2018.
- 460 Boberg, F., Mottram, R., Hansen, N., Yang, S., and Langen, P. L.: Uncertainties in projected surface mass balance over the polar ice sheets from dynamically downscaled EC-Earth models, *The Cryosphere*, 16, 17–33, <https://doi.org/10.5194/tc-16-17-2022>, 2022.
- Bozkurt, D., Bromwich, D. H., Carrasco, J., and Rondanelli, R.: Temperature and precipitation projections for the Antarctic Peninsula over the next two decades: contrasting global and regional climate model simulations, *Climate Dynamics*, 56, 3853–3874, <https://doi.org/10.1007/s00382-021-05667-2>, 2021.
- 465 Braun, M., Humbert, A., and Moll, A.: Changes of Wilkins Ice Shelf over the past 15 years and inferences on its stability, *The Cryosphere*, 3, 41–56, <https://doi.org/10.5194/tc-3-41-2009>, 2009.
- Brils, M., Munneke, P. K., Jullien, N., Tedstone, A. J., Machguth, H., van de Berg, W., and van den Broeke, M.: Climatic drivers of ice slabs and firn aquifers in Greenland, *Geophysical Research Letters*, 51, e2023GL106613, <https://doi.org/10.1029/2023GL106613>, 2024.
- Calonne, N., Milliancourt, L., Burr, A., Philip, A., Martin, C. L., Flin, F., and Geindreau, C.: Thermal conductivity of snow, firn, and porous ice from 3-D image-based computations, *Geophysical Research Letters*, 46, 13 079–13 089, <https://doi.org/10.1029/2019GL085228>, 2019.
- 470 Carter, J., Leeson, A., Orr, A., Kittel, C., and van Wessem, J. M.: Variability in Antarctic surface climatology across regional climate models and reanalysis datasets, *The Cryosphere*, 16, 3815–3841, <https://doi.org/10.5194/tc-16-3815-2022>, 2022.
- Chen, T. and Guestrin, C.: Xgboost: A scalable tree boosting system, in: *Proceedings of the 22nd acm sigkdd international conference on knowledge discovery and data mining*, 13–17 August 2016, San Francisco, CA, USA, pp. 785–794, <https://doi.org/10.1145/2939672.2939785>, 2016.
- 475 Coléou, C. and Lesaffre, B.: Irreducible water saturation in snow: experimental results in a cold laboratory, *Annals of glaciology*, 26, 64–68, <https://doi.org/10.3189/1998AoG26-1-64-68>, 1998.
- Cook, A. J. and Vaughan, D. G.: Overview of areal changes of the ice shelves on the Antarctic Peninsula over the past 50 years, *The cryosphere*, 4, 77–98, <https://doi.org/10.5194/tc-4-77-2010>, 2010.
- 480 Culberg, R., Schroeder, D. M., and Chu, W.: Extreme melt season ice layers reduce firn permeability across Greenland, *Nature communications*, 12, 2336, <https://doi.org/10.1038/s41467-021-22656-5>, 2021.
- Danabasoglu, G., Lamarque, J.-F., Bacmeister, J., Bailey, D., DuVivier, A., Edwards, J., Emmons, L., Fasullo, J., Garcia, R., Gettelman, A., Hannay, C., Holland, M. M., Large, W. G., Lauritzen, P. H., Lawrence, D. M., Lenaerts, J. T. M., Lindsay, K., Lipscomb, W. H., Mills, M. J., Neale, R., Oleson, K. W., Otto-Bliesner, B., Phillips, A. S., Sacks, W., Tilmes, S., van Kampenhout, L., Vertenstein, M., Bertini, A., Dennis, J., Deser, C., Fischer, C., Fox-Kemper, B., Kay, J. E., Kinnison, D., Kushner, P. J., Larson, V. E., Long, M. C., Mickelson, S., Moore, J. K., Nienhouse, E., Polvani, L., Rasch, P. J., and Strand, W. G.: The community earth system model version 2 (CESM2), *Journal of Advances in Modeling Earth Systems*, 12, e2019MS001916, <https://doi.org/10.1029/2019MS001916>, 2020.

- Datta, R. T., Tedesco, M., Agosta, C., Fettweis, X., Kuipers Munneke, P., and van den Broeke, M. R.: Melting over the northeast Antarctic Peninsula (1999–2009): evaluation of a high-resolution regional climate model, *The Cryosphere*, 12, 2901–2922, 2018.
- 490 Di Biase, V., Kuipers Munneke, P., Veldhuijsen, S., de Roda Husman, S., van den Broeke, M., Noël, B., Buth, L., and Wouters, B.: Probability of firn aquifer presence in Antarctica by combining remote sensing and regional climate model data, *Geophysical Research Letters*, 51, e2024GL109367, <https://doi.org/10.1029/2024GL109367>, 2024.
- Dunmire, D., Lenaerts, J. T., Datta, R. T., and Gorte, T.: Antarctic surface climate and surface mass balance in the Community Earth System Model version 2 during the satellite era and into the future (1979–2100), *The Cryosphere*, 16, 4163–4184, <https://doi.org/10.5194/tc-16-4163-2022>, 2022.
- 495 Dunmire, D., Wever, N., Banwell, A. F., and Lenaerts, J. T.: Antarctic-wide ice-shelf firn emulation reveals robust future firn air depletion signal for the Antarctic Peninsula, *Communications Earth & Environment*, 5, 100, <https://doi.org/10.1038/s43247-024-01255-4>, 2024.
- Fatima, S., Hussain, A., Amir, S. B., Ahmed, S. H., and Aslam, S. M. H.: Xgboost and random forest algorithms: an in depth analysis, *Pakistan Journal of Scientific Research*, 3, 26–31, <https://doi.org/10.57041/pjosr.v3i1.946>, 2023.
- 500 Forster, R. R., Box, J. E., Van Den Broeke, M. R., Miège, C., Burgess, E. W., Van Angelen, J. H., Lenaerts, J. T., Koenig, L. S., Paden, J., Lewis, C., Gogineni, S. P., Lauschen, C., and McConnel, J. R.: Extensive liquid meltwater storage in firn within the Greenland ice sheet, *Nature Geoscience*, 7, 95–98, <https://doi.org/10.1038/ngeo2043>, 2014.
- Fricker, H. A., Allison, I., Craven, M., Hyland, G., Ruddell, A., Young, N., Coleman, R., King, M., Krebs, K., and Popov, S.: Redefinition of the Amery ice shelf, East Antarctica, grounding zone, *Journal of Geophysical Research: Solid Earth*, 107, ECV–1, <https://doi.org/10.1029/2001JB000383>, 2002.
- 505 González Herrero, S., Vasallo, F., Bech, J., Gorodetskaya, I., Elvira Montejó, B., and Justel, A.: Extreme precipitation records in Antarctica, *International Journal of Climatology*, pp. 1–14, <https://doi.org/10.1002/joc.8020>, 2023.
- Grinsztajn, L., Oyallon, E., and Varoquaux, G.: Why do tree-based models still outperform deep learning on typical tabular data?, *Advances in neural information processing systems*, 35, 507–520, https://proceedings.neurips.cc/paper_files/paper/2022/file/0378c7692da36807bdec87ab043cdadc-Paper-Datasets_and_Benchmarks.pdf, 2022.
- 510 Hansen, N., Langen, P. L., Boberg, F., Forsberg, R., Simonsen, S. B., Thejll, P., Vandecrux, B., and Mottram, R.: Downscaled surface mass balance in Antarctica: impacts of subsurface processes and large-scale atmospheric circulation, *The Cryosphere*, 15, 4315–4333, <https://doi.org/10.5194/tc-15-4315-2021>, 2021.
- Horlings, A. N., Christianson, K., and Miège, C.: Expansion of firn aquifers in southeast Greenland, *Journal of Geophysical Research: Earth Surface*, 127, e2022JF006753, <https://doi.org/10.1029/2022JF006753>, 2022.
- 515 Hubbard, B., Luckman, A., Ashmore, D. W., Bevan, S., Kulesa, B., Kuipers Munneke, P., Philippe, M., Jansen, D., Booth, A., Sevestre, H., Tison, J.-L., O’leary, M., and Rutt, I.: Massive subsurface ice formed by refreezing of ice-shelf melt ponds, *Nature communications*, 7, 11897, <https://doi.org/10.1038/ncomms11897>, 2016.
- Iban, M. C. and Bilgilioglu, S. S.: Snow avalanche susceptibility mapping using novel tree-based machine learning algorithms (XGBoost, NGBost, and LightGBM) with eXplainable Artificial Intelligence (XAI) approach, *Stochastic Environmental Research and Risk Assessment*, 37, 2243–2270, <https://doi.org/10.1007/s00477-023-02392-6>, 2023.
- 520 Jakobs, C. L., Reijmer, C. H., Kuipers Munneke, P., König-Langlo, G., and Van Den Broeke, M. R.: Quantifying the snowmelt–albedo feedback at Neumayer Station, East Antarctica, *The Cryosphere*, 13, 1473–1485, <https://doi.org/10.5194/tc-13-1473-2019>, 2019.
- Jourdain, N. C., Amory, C., Kittel, C., and Durand, G.: Changes in Antarctic surface conditions and potential for ice shelf hydrofracturing from 1850 to 2200, *EGU sphere*, 2024, 1–34, <https://doi.org/10.5194/egusphere-2024-58>, 2024.
- 525

- Kingslake, J., Ng, F., and Sole, A.: Modelling channelized surface drainage of supraglacial lakes, *Journal of Glaciology*, 61, 185–199, <https://doi.org/10.3189/2015JoG14J158>, 2015.
- Kittel, C., Amory, C., Agosta, C., Jourdain, N. C., Hofer, S., Delhasse, A., Doutreloup, S., Huot, P.-V., Lang, C., Fichet, T., and Fettweis, X.: Diverging future surface mass balance between the Antarctic ice shelves and grounded ice sheet., *The Cryosphere*, 15, 1215–1215, <https://doi.org/10.5194/tc-15-1215-2021>, 2021.
- Kuipers Munneke, P., M. Ligtenberg, S., Van Den Broeke, M., Van Angelen, J., and Forster, R.: Explaining the presence of perennial liquid water bodies in the firn of the Greenland Ice Sheet, *Geophysical Research Letters*, 41, 476–483, <https://doi.org/10.1002/2013GL058389>, 2014.
- Lai, C.-Y., Kingslake, J., Wearing, M. G., Chen, P.-H. C., Gentine, P., Li, H., Spergel, J. J., and van Wessem, J. M.: Vulnerability of Antarctica's ice shelves to meltwater-driven fracture, *Nature*, 584, 574–578, <https://doi.org/10.1038/s41586-020-2627-8>, 2020.
- Leppäranta, M., Järvinen, O., and Mattila, O.-P.: Structure and life cycle of supraglacial lakes in Dronning Maud Land, *Antarctic Science*, 25, 457–467, <https://doi.org/10.1017/S0954102012001009>, 2013.
- Luckman, A., Elvidge, A., Jansen, D., Kulesa, B., Munneke, P. K., King, J., and Barrand, N. E.: Surface melt and ponding on Larsen C Ice Shelf and the impact of föhn winds, *Antarctic Science*, 26, 625–635, <https://doi.org/doi:10.1017/S0954102014000339>, 2014.
- MacDonell, S., Fernandoy, F., Villar, P., and Hammann, A.: Stratigraphic analysis of firn cores from an Antarctic ice shelf firn aquifer, *Water*, 13, 731, <https://doi.org/10.3390/w13050731>, 2021.
- Miège, C., Forster, R. R., Brucker, L., Koenig, L. S., Solomon, D. K., Paden, J. D., Box, J. E., Burgess, E. W., Miller, J. Z., McNerney, L., Brautigam, N., Fausto, R. S., and Gogineni, S.: Spatial extent and temporal variability of Greenland firn aquifers detected by ground and airborne radars, *Journal of Geophysical Research: Earth Surface*, 121, 2381–2398, <https://doi.org/10.1002/2016JF003869>, 2016.
- Miller, O., Solomon, D. K., Miège, C., Koenig, L., Forster, R., Schmerr, N., Ligtenberg, S. R., and Montgomery, L.: Direct evidence of meltwater flow within a firn aquifer in southeast Greenland, *Geophysical Research Letters*, 45, 207–215, <https://doi.org/10.1002/2017GL075707>, 2018.
- Montgomery, L., Miège, C., Miller, J., Scambos, T. A., Wallin, B., Miller, O., Solomon, D. K., Forster, R., and Koenig, L.: Hydrologic properties of a highly permeable firn aquifer in the Wilkins Ice Shelf, Antarctica, *Geophysical Research Letters*, 47, e2020GL089552, <https://doi.org/10.1029/2020GL089552>, 2020.
- Mottram, R., Hansen, N., Kittel, C., van Wessem, J. M., Agosta, C., Amory, C., Boberg, F., van de Berg, W. J., Fettweis, X., Gossart, A., van Lipzig, N. P. M., van Meijgaard, E., Orr, A., Philips, T., Webster, S., Simonson, S. B., and Souverijns, N.: What is the surface mass balance of Antarctica? An intercomparison of regional climate model estimates, *The Cryosphere*, 15, 3751–3784, <https://doi.org/10.5194/tc-15-3751-2021>, 2021.
- Oppenheimer, M., Glavovic, B., Hinkel, J., van de Wal, R., Magnan, A. K., Abd-Elgawad, A., Cai, R., Cifuentes-Jara, M., Deconto, R. M., Ghosh, T., Hay, J., Ben, M., Meyssignac, B., Sebesvari, Z., A, S., Dangendorf, S., and Frederikse, T.: Sea level rise and implications for low lying islands, coasts and communities, *The Intergovernmental Panel on Climate Change*, <https://doi.org/10.1017/9781009157964.006>, 2019.
- Poinar, K., Joughin, I., Lilien, D., Brucker, L., Kehrl, L., and Nowicki, S.: Drainage of Southeast Greenland firn aquifer water through crevasses to the bed, *Frontiers in Earth Science*, 5, 5, <https://doi.org/10.3389/feart.2017.00005>, 2017.
- Roberts, D. R., Bahn, V., Ciuti, S., Boyce, M. S., Elith, J., Guillera-Arroita, G., Hauenstein, S., Lahoz-Monfort, J. J., Schröder, B., Thuiller, W., Warton, D. I., Wintle, B. A., Hartig, F., and Dormann, C. F.: Cross-validation strategies for data with temporal, spatial, hierarchical, or phylogenetic structure, *Ecography*, 40, 913–929, <https://doi.org/10.1111/ecog.02881>, 2017.

Sarofim, M. C., Smith, C. J., Malek, P., McDuffie, E. E., Hartin, C. A., Lay, C. R., and McGrath, S.: High radiative forcing climate scenario relevance analyzed with a ten-million-member ensemble, *Nature communications*, 15, 8185, <https://doi.org/10.1038/s41467-024-52437-9>, 2024.

Scambos, T., Fricker, H. A., Liu, C.-C., Bohlander, J., Fastook, J., Sargent, A., Massom, R., and Wu, A.-M.: Ice shelf disintegration by plate bending and hydro-fracture: Satellite observations and model results of the 2008 Wilkins ice shelf break-ups, *Earth and Planetary Science Letters*, 280, 51–60, <https://doi.org/10.1016/j.epsl.2008.12.027>, 2009.

Scambos, T. A., Hulbe, C., Fahnestock, M., and Bohlander, J.: The link between climate warming and break-up of ice shelves in the Antarctic Peninsula, *Journal of Glaciology*, 46, 516–530, <https://doi.org/doi:10.3189/172756500781833043>, 2000.

Shepherd, A., Gilbert, L., Muir, A. S., Konrad, H., McMillan, M., Slater, T., Briggs, K. H., Sundal, A. V., Hogg, A. E., and Engdahl, M. E.: Trends in Antarctic Ice Sheet elevation and mass, *Geophysical Research Letters*, 46, 8174–8183, <https://doi.org/10.1029/2019GL082182>, 2019.

Smith, B., Fricker, H. A., Gardner, A. S., Medley, B., Nilsson, J., Paolo, F. S., Holschuh, N., Adusumilli, S., Brunt, K., Csatho, B., Harbeck, K., Markus, T., Neumann, T., Siegfried, M. R., and Zwally, H. J.: Pervasive ice sheet mass loss reflects competing ocean and atmosphere processes, *Science*, 368, 1239–1242, <https://doi.org/10.5194/tc-5-809-2011>, 2020.

Sun, L., Zhang, X., Xiao, P., Wang, H., Wang, Y., and Zheng, Z.: Fusing daily snow water equivalent from 1980 to 2020 in China using a spatiotemporal XGBoost model, *Journal of Hydrology*, 632, 130 876, <https://doi.org/10.1016/j.jhydrol.2024.130876>, 2024.

Turner, J., Phillips, T., Thamban, M., Rahaman, W., Marshall, G. J., Wille, J. D., Favier, V., Winton, V. H. L., Thomas, E., Wang, Z., van den Broeke, M., Hosking, J. S., and Lachlan-Cope, T.: The dominant role of extreme precipitation events in Antarctic snowfall variability, *Geophysical Research Letters*, 46, 3502–3511, <https://doi.org/10.1029/2018GL081517>, 2019.

van Dalum, C. T., van de Berg, W. J., Gadde, S. N., van Tiggelen, M., van der Drift, T., van Meijgaard, E., van Ulf, L. H., and van den Broeke, M. R.: First results of the polar regional climate model RACMO2. 4, *The Cryosphere*, 18, 4065–4088, <https://doi.org/10.5194/tc-18-4065-2024>, 2024.

Van Katwyk, P., Fox-Kemper, B., Seroussi, H., Nowicki, S., and Bergen, K. J.: A variational LSTM emulator of sea level contribution from the Antarctic ice sheet, *Journal of Advances in Modeling Earth Systems*, 15, e2023MS003 899, <https://doi.org/10.1029/2023MS003899>, 2023.

Van Wessem, J., Ligtenberg, S., Reijmer, C., Van De Berg, W., Van Den Broeke, M., Barrand, N., Thomas, E., Turner, J., Wuite, J., Scambos, T., et al.: The modelled surface mass balance of the Antarctic Peninsula at 5.5 km horizontal resolution, *The Cryosphere*, 10, 271–285, <https://doi.org/10.5194/tc-10-271-2016>, 2016.

van Wessem, J. M., Steger, C. R., Wever, N., and van den Broeke, M. R.: An exploratory modelling study of perennial firn aquifers in the Antarctic Peninsula for the period 1979–2016, *The Cryosphere*, 15, 695–714, <https://doi.org/10.5194/tc-15-695-2021>, 2021.

van Wessem, J. M., van den Broeke, M. R., Wouters, B., and Lhermitte, S.: Variable temperature thresholds of melt pond formation on Antarctic ice shelves, *Nature Climate Change*, 13, 161–166, <https://doi.org/10.1038/s41558-022-01577-1>, 2023.

Vandecrux, B., Mottram, R., Langen, P. L., Fausto, R. S., Olesen, M., Stevens, C. M., Verjans, V., Leeson, A., Ligtenberg, S., Kuipers Munneke, P., Marchenko, S., van Pelt, W., Meyer, C. R., Simonsen, S. B., Heilig, A., Samimi, S., Marshall, S., Machguth, H., MacFerrin, M., Niwano, M., Miller, O., Voss, C. I., and Box, J. E.: The firn meltwater Retention Model Intercomparison Project (RetMIP): evaluation of nine firn models at four weather station sites on the Greenland ice sheet, *The Cryosphere*, 14, 3785–3810, <https://doi.org/10.5194/tc-14-3785-2020>, 2020.

- Veldhuijsen, S., van de Berg, W. J., Kuipers Munneke, P., and van den Broeke, M. R.: IMAU-FDM v12AD TC release, Zenodo [code], <https://doi.org/10.5281/zenodo.10723570>, 2023.
- Veldhuijsen, S., van de Berg, W. J., Kuipers Munneke, P., Hansen, N., Boberg, F., Kittel, C., Amory, C., and van den Broeke, M. R.: Code perennial firn aquifer emulator [code], <https://zenodo.org/records/13750024>, 2024a.
- 605 Veldhuijsen, S., van de Berg, W. J., Kuipers Munneke, P., Hansen, N., Boberg, F., Kittel, C., Amory, C., and van den Broeke, M. R.: Data perennial firn aquifer emulator, Zenodo [data set], <https://doi.org/10.5281/zenodo.13692364>, 2024b.
- Veldhuijsen, S., van de Berg, W. J., Kuipers Munneke, P., Hansen, N., Boberg, F., Kittel, C., Amory, C., and van den Broeke, M. R.: IMAU-FDMv1.2AD perennial LWC (1950-2100), Zenodo [data set], <https://doi.org/10.5281/zenodo.13692385>, 2024c.
- Veldhuijsen, S. B., Van De Berg, W. J., Kuipers Munneke, P., and Van Den Broeke, M. R.: Firn air content changes on Antarctic ice shelves
610 under three future warming scenarios, *The Cryosphere*, 18, 1983–1999, <https://doi.org/10.5194/tc-18-1983-2024>, 2024d.
- Verjans, V., Leeson, A., McMillan, M., Stevens, C., van Wessem, J. M., van de Berg, W. J., van den Broeke, M. R., Kittel, C., Amory, C., Fettweis, X., Hansen, N., Boberg, F., and Mottram, R.: Uncertainty in East Antarctic firn thickness constrained using a model ensemble approach, *Geophysical Research Letters*, 48, e2020GL092060, <https://doi.org/10.1029/2020GL092060>, 2021.
- Zwally, H. J., Abdalati, W., Herring, T., Larson, K., Saba, J., and Steffen, K.: Surface melt-induced acceleration of Greenland ice-sheet flow,
615 *Science*, 297, 218–222, <https://doi.org/10.1126/science.1072708>, 2002.