Vertical changes in volatile organic compounds (VOCs) and impacts on photochemical ozone formation

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Abstract

Volatile organic compounds (VOCs) play crucial roles in regulating the formation of tropospheric ozone. However, limited knowledge on the interactions between vertical VOC variations and photochemical ozone formation in the planetary boundary layer (PBL) has hindered effective ozone control strategies, especially in large cities. In this study, we investigated the vertical changes in concentrations, compositions, and key driving factors of a large suite of VOCs using online gradient measurements taken from a 325 m tall tower in urban Beijing. The impact of these vertical VOC variations on photochemical ozone formation were also analyzed using box model simulations. Our results indicate that VOCs exhibited distinct vertical variation patterns due to their differences in sources and chemical reactivities, along with the diurnal evolution of the PBL. During daytime, reactive VOCs (e.g., hydrocarbons) are rapidly oxidized as they mix upwards, accompanied by the formation and accumulation of oxygenated VOCs (OVOCs) in the middle and upper layers. In addition, the photochemical formation of ozone responds positively to changes in both NOx and VOCs. As a result, the production rate of ozone declines with height due to the simultaneous decreases in concentrations of reactive VOCs and NOx, but remains high in the middle and upper layers. The strong production of ozone aloft is primarily driven by the presence of high OVOCs concentrations. Therefore, careful consideration should be given to the vertical variations in both photochemical ozone production rates and formation regimes in the whole PBL when developing regional ozone control strategies.

1 Introduction

Volatile organic compounds (VOCs) are crucial constituents of atmospheric chemicals *(Li et al., 2022c)* and play important roles in regulating the atmospheric oxidation capacity and contributing to the photochemical formation of tropospheric ozone *(Zhao et al., 2022; Yang et al., 2024b)*. Ozone is a major air pollutant in urban environments, with increasing trends reported globally over recent decades *(Fleming et al., 2018; Cooper et al., 2020)*, despite stringent measures to control its precursor emissions *(Wang et al., 2020b; Yeo and Kim, 2021; Li et al., 2022b; Perdigones et al., 2022)*. As highlighted in previous studies, reducing emissions of reactive VOCs is key to controlling ozone pollution at present and in the foreseeable future *(Zhao et al., 2022; Wang et al., 2024)*.

The primary prerequisite for effective regional ozone pollution control is the determination of the photochemical ozone formation regime *(Souri et al., 2020; Zhao et al., 2022)*, which facilitates the development of reduction schemes for key precursor emissions *(Ou et al., 2016; Wang et al., 2019)*. The main challenges in controlling ozone pollution stem from the complex compositions of its precursors (e.g., VOCs and NOx) in ambient air *(Guo et al., 2017; Wu et al., 2020; Li et al., 2022c)*, as well as the complicated responses of photochemical ozone formation to changes in these precursors *(Shao et al., 2021; Perdigones et al., 2022)*. Furthermore, the interactions between vertical variations of ozone precursors and ozone formation remain unclear *(Tang et al., 2017; Sun et al., 2018; Li et al., 2024)*, adding to the complexity of ozone pollution control.

In most cases, the identification of key ozone precursors has been conducted using ground-level observations *(Qi et al., 2021; Lu et al., 2022)* or compiled source emission inventories *(Ou et al., 2015; An et al., 2021; Wang et al., 2022b)*. While these methods are undoubtedly helpful in determining key ozone precursors and corresponding reduction strategies, they often encounter unexpected uncertainties in urban regions. *(Mo et al., 2018; Mo et al., 2020)*. Consequently, ground-level measurements of ozone

precursors have been favored to constrain model calculations *(Lu et al., 2012; Wang et al., 2022a; Yang et al., 2022*) or provide empirical evidence for hypothesized theories *(Hofzumahaus et al., 2009; Wang et al., 2022c)*. However, these ground-level measurements cannot fully characterize atmospheric chemical processes in the entire planetary boundary layer (PBL) due to strong vertical variations in precursor concentrations *(Velasco et al., 2008; Li et al., 2018; Sun et al., 2018)*.

Ambient VOCs, as crucial ozone precursors, are composed of myriad species *(Wu et al., 2020; Gkatzelis et al., 2021; Ye et al., 2021; He et al., 2022)* and serve diverse functions in photochemical ozone formation *(Vo et al., 2018; Li et al., 2022a; Zhang et al., 2022)*. Owing to the impact of variations in emission sources, chemical removal, advection and convection transport, and secondary formation, the concentration and composition of VOCs typically display notable vertical variability within the PBL, especially in urban areas *(Li et al., 2022c)*. The ozone formation regime likely undergoes significant transitions from the ground to the upper boundary layer *(Li et al., 2024; Liu et al., 2024a)*. Ozone generated throughout the PBL can influence surface ozone levels due to enhanced atmospheric vertical mixing during the day. Consequently, it is imperative to comprehend the vertical variations and principal determinants of VOCs, as well as their effects on photochemical ozone formation within the PBL.

With the rapid development of cities in the recent two decades in China, a large number of pollution-emitting industries and factories have been relocated from city centers to alleviate air pollution. Concurrently, there's been a swift increase in the ownership of electric vehicles *(Guo et al., 2021)*. These shifts in energy consumption have driven the change in concentrations and compositions of VOCs in major cities like Beijing *(Liu et al., 2024b)*, subsequently affecting photochemical ozone formation *(Wang et al., 2024)*. However, the vertical variations and key drivers of VOCs and their impacts on photochemical ozone formation in the urban PBL remain elusive. A primary hurdle in studying these vertical changes in photochemical ozone formation is the scarcity of reliable vertical VOC data *(Dieu Hien et al., 2019; Li et al., 2022c)*. Engaging in vertical profiling of VOCs, ensuring all necessary species represented and

obtaining sufficient sample size, is especially challenging in the lower PBL where atmospheric chemical reactions are most intense *(Benish et al., 2020; Kim et al., 2021)*.

Previous studies on vertical distributions of photochemical ozone formation in the PBL have been conducted using measurements of a limited number of VOC species and samples *(Zhang et al., 2018; Benish et al., 2020; Geng et al., 2020)*. In this study, online gradient measurements of ozone, NOx, and a large suit of VOCs were made on a 325 m tall tower in urban Beijing during the summer of 2021. Additionally, box model simulations constrained by the gradient measurements were performed to analyze the vertical variations and key drivers of VOCs as well as their impacts on photochemical ozone formation.

2 Methods and materials

2.1 Description of the site, instrument, and field campaign

The data utilized in this study was derived from an intensive field campaign conducted at the Beijing Meteorological Tower (BMT: 39°58′ N, 116°23′ E) between July 6 and August 4, 2021. The BMT has a height of 325 m and is located in the northern part of downtown Beijing, positioned between the third and fourth ring roads (Fig. S1). A vertical observation system, established using long perfluoroalkoxy alkane (PFA) Teflon tubes (OD: 1/2 in.), was used to make online gradient measurements of ozone, NOx, and a set of VOCs on the BMT. Five specific heights, namely 15, 47, 102, 200, and 320 m above ground level, were selected to mount the tube inlets, as depicted in Fig. S3. An additional inlet, situated approximately 5 m above ground level, was mounted on the rooftop of the observation room that was adjacent to the tower. Consequently, the vertical observation system totally included a total of six sampling inlets. The sampling inlet at the 15 m height was not utilized during this field campaign.

Filters were installed downstream of the tubing inlets on the tower to remove fine particles. A rotary vane vacuum pump was used to simultaneously and continuously draw sample air from the five tubes, ensuring that all tubes were flushed by ambient air

to reduce tubing delays of sticky organic compounds *(Pagonis et al., 2017; Liu et al., 2019)*. Five critical orifices were employed to control the flow rate of the air stream in each tubing, resulting in flow rates ranging between 15 and 20 standard liter per minute (SLPM). Instruments drew sample air from the five tubes sequentially through a Teflon solenoid valve group at designated time intervals. The switching time intervals of the Teflon solenoid valve group were set as 4 minutes during this field campaign. The measurements of trace gases in the first and last 1 minute of a four-minute period were discarded to eliminate cross interferences between different inlet heights. Detailed information on the vertical observation system and the assessment of trace gas measurements through hundreds of meters long PFA tubes has been provided in our previous works *(Li et al., 2023; Song et al., 2024; Yang et al., 2024a)*.

Ozone was measured using the ultraviolet photometry method (49i, Thermo Fisher Scientific Inc., USA). NO, NO_2, and NOx were measured using the chemiluminescence method (42i, Thermo Fisher Scientific Inc., USA). Gradient measurements of ozone and NOx were conducted at a time resolution of 10 seconds. The photolysis frequencies of NO_2, represented by $j(NO_2)$, were measured by a spectrometer (PFS-100, Focused Photonics Inc., China) situated on the rooftop of the observation room and have a time resolution of 8 seconds. In situ measurements of meteorological parameters including wind speed, air temperature, and relative humidity were made at 15 heights between 8 m and 320 m on the BMT with a time resolution of 20 seconds. Planetary boundary layer height (PBLH) was obtained from the Air Resources Laboratory (*https://ready.arl.noaa.gov/READYamet.php, last access: 10 June 2024*) and was linearly interpolated to hourly values based on the initial time resolutions of three hours *(Li and Fan, 2022)*.

A high-resolution proton-transfer-reaction quadrupole interface time-of-flight mass spectrometer (PTR-ToF-MS, Ionicon Analytik, Austria) was employed to measure VOCs at a time resolution of 10 seconds. The PTR-ToF-MS used both hydronium ion (H_3O^+) *(Yuan et al., 2017; Wu et al., 2020; Li et al., 2022c)* and nitric oxide ion (NO^+) *(Wang et al., 2020a)* as reagent ions. These two reagent ions were automatically

switched every 60 min for H_3O^+ and every 22 min for NO^+ throughout the campaign. The PTR-ToF-MS operated at an E/N value of approximately 120 Td in H_3O^+ mode and an E/N value of around 60 Td in NO^+ mode. Instrument backgrounds were automatically measured during the last two minutes of each operation mode by passing ambient air through a platinum catalyst heated to 365 ℃. A gas standard containing 39 VOC species was used to calibrate the PTR-ToF-MS daily. Sensitivities for the remaining species were determined based on reaction kinetics of the PTR-ToF-MS *(Wu et al., 2020)*. Impacts of ambient humidity on the PTR-ToF-MS measurements were corrected by using humidity-dependence curves of VOCs obtained in our laboratory *(Wang et al., 2020a; Wu et al., 2020)*. Carbon dioxide (CO_2 in dry air) and humidity were measured using a CO_2 and H_2O gas analyzer (Li-840A, Licor Inc., USA) at a time resolution of 10 seconds.

Gradient measurements of the total OH reactivity (OHR) of atmospheric trace gases were made using the improved comparative reactivity method (ICRM) developed by our team *(Wang et al., 2021a)* from July 28 to 31. In addition, gradient measurements of carbon monoxide (CO), methane (CH_4), CO_2, and H_2O were simultaneously measured using the cavity ring-down spectroscopy (CRDS) method (G-2401, Picarro Inc., USA) at a time resolution of 10 seconds from May 15 to June 25. Sulfur dioxide (SO_2) was measured using the ultraviolet fluorescence method (43i, Thermo Fisher Scientific Inc., USA) at a time resolution of 10 seconds from June 25 to August 3. The total OHR of VOCs, denoted by OHR_{VOCs}, can be estimated by excluding those of the inorganic species (namely ozone, NOx, CO, SO_2, and CH_4). It should be noted that gradient measurements of CH_4 and CO were not made during July 28-31, and their average concentrations in daytime (11:00-16:00 LT) between May 15 and June 25 at 5 m were used for all altitudes to calculate OHR_{VOCs}. This method will bring minor uncertainties due to the minor vertical differences in concentrations of CH_4 and CO in daytime (Fig. S4). The OHR of VOCs can also be calculated by summing the products of their measured concentrations and their reaction rate coefficients with OH radicals, as formulated in Eq. (1):

$$\mathrm{OHR} = \sum k^{i}{}_{OH-R}[VOC_i] \qquad \text{Eq. (1)}$$

where $k^{i}{}_{OH-R}$ is the reaction rate coefficient of the i^{th} VOC species with OH radical and $[VOC_i]$ is the concentration of the i^{th} VOC species.

2.2 Estimation of NMHC concentrations at the BMT site

The PTR-ToF-MS is limited in its ability to measure VOC species with proton affinities higher than H_2O (691 kJ mol^{-1}) when operating in the H_3O^+ mode *(Yuan et al., 2017)*. This limitation results in the absence of certain nonmethane hydrocarbons (NMHCs), such as alkanes and many alkene species, which play important roles in photochemical ozone formation. To obtain a comprehensive understanding of vertical variations in concentrations, compositions, and environmental impacts of VOCs, this study estimated the vertical profiles of those unmeasured NMHC species based on the concentrations of measured VOCs using the PTR-ToF-MS. Detailed information on estimation of NMHC concentrations is provided in SI.

2.3 Box model setup

A zero-dimension box model (F0AM) coupled with the Master Chemical Mechanism (v3.3.1) *(Wolfe et al., 2016; Yang et al., 2022)* was used to compute the production rate of ozone, denoted by $P(O_3)$ as formulated in Eq. (2):

$$\mathrm{P}(O_3) = k_{HO_2+NO}[HO_2][NO] + \sum k^{i}{}_{RO_2+NO}\left[R^{i}O_2\right][NO] \qquad \text{Eq. (2)}$$

$$k_{RO_2+NO}[RO_2][NO]$$

where $[HO_2]$ and $[NO]$ is the concentrations of HO_2 and NO, $[R^{i}O_2]$ is the concentration of the i^{th} organic peroxyl radical. The relative incremental reactivity (RIR) of photochemical ozone production to changes in different precursors was determined using Eq. (3):

$$\mathrm{RIR(X)} = \frac{\left[P^{S}_{O_x}(\mathrm{X}) - P^{S}_{O_x}(\mathrm{X}-\Delta\mathrm{X})\right]/P^{S}_{O_x}(\mathrm{X})}{\Delta\mathrm{S(X)}/\mathrm{S(X)}} \qquad \text{Eq. (3)}$$

where X represents ozone precursors, $P_{O_x}^{S}(\mathrm{X})$ is the contribution of X to the production rate of Ox, $\Delta \mathrm{X}$ is the amount of change in ozone precursors, $\mathrm{S(X)}$ is the initial concentration of X. RIR values were used to discern sensitivities of photochemical ozone formation to changes in precursor gases. A positive RIR(X) value suggests that an increase in X enhances ozone formation, while a negative RIR value indicates that an increase in X inhibits ozone formation.

Model calculations were constrained by measurements of ozone, NOx, CO, a suit of VOCs, air temperature, and relative humidity. In addition to the measured or estimated concentrations of NMHCs, nine oxygenated VOC (OVOC) species (Table S1) measured by PTR-ToF-MS were used to constrain the model calculation. The model was run in a time-dependent mode with a time resolution of 5 minutes and a spin-up period of 2 days *(Lu et al., 2012; Wang et al., 2022c)*. The dry deposition velocity of ozone was set as 0.27 cm s^{-1} when calculating $P(O_3)$ 5 m and was zeroed out when calculating $P(O_3)$ at other heights.

3 Results and discussions

3.1 Temporal and vertical variations in concentrations of trace gases

As shown in Fig. 1, the meteorology in Beijing was characterized by high air temperature (27.3±2.9 ℃), high humidity (83.9%±16.2%), and gentle winds (1.1±0.4 m s^{-1}) throughout the campaign. The intense solar radiation, elevated air temperature, and mild winds favored the photochemical formation and accumulation of ozone, leading to frequent occurrences of ozone pollution episodes. Fig. 1 also presents time series of mixing ratios of ozone and its selected precursors (namely isoprene, toluene, monoterpenes, and NOx) along with $j(NO_2)$ measured at 5 m. The campaign mean ozone mixing ratio was 45.6±25.3 ppb, but the maximum hourly mean ozone mixing ratio reached 129.3 ppb, indicating strong photochemical reactions in urban Beijing during the campaign. Surface ozone concentrations exhibited a typical diurnal variation

pattern with the maximum occurring at 16:00 LT (Fig. S6), implying its predominant source from local photochemical production.

Isoprene is a typical tracer of biogenic emissions and is also a highly reactive VOC species *(Atkinson and Arey, 2003)*. Isoprene had a campaign mean mixing ratio of 0.7±0.6 ppb. The average diurnal profile of isoprene at 5 m has a unimodal pattern with the maximum occurring at 14:00 LT (Fig. S6), exhibiting strong dependence on solar radiation. Monoterpenes were also generally recognized as typical tracers of biogenic emissions *(Gómez et al., 2020)* and have a campaign mean mixing ratio of 0.3±0.3 ppb. The average diurnal profile of monoterpenes was characterized by low mixing ratios in daytime with two peaks occurring at 05:00 and 20:00 LT, respectively.

Toluene and NOx are recognized as typical tracers of anthropogenic emissions in urban regions *(Niu et al., 2017; Li et al., 2022c)*, with campaign mean mixing ratios of 0.7±0.7 and 8.1±4.8 ppb, respectively. The average diurnal profiles of toluene and NOx at 5 m exhibited similar variations with larger values at night than during the day. Based on the measured concentrations and diurnal variations of ozone and its key precursors at ground level, it can be inferred that urban Beijing is experiencing severe ozone pollution, which is predominantly contributed by local photochemical production. As key ozone precursors, ambient concentrations of VOCs are contributed by the mixture of anthropogenic and biogenic sources.

Fig. 2 shows the average diurnal and vertical variations in mixing ratios of ozone, NOx, Ox (O_3+NO_2), and six selected VOC species (three hydrocarbons and three OVOCs) within the measurement height range of 5-320 m. High mixing ratios of ozone were observed in the afternoon following the enhancement of solar radiation, which was consistent with the diurnal change pattern of ozone concentrations at the ground level. The vertical gradients of ozone mixing ratios were positive throughout the day but substantially enhanced at night (Fig. 3). The lower ozone mixing ratios near the surface than aloft were mainly caused by the enhancement of dry deposition and NO titration *(Brown et al., 2007; Ma et al., 2013; Li et al., 2022b)*.

NOx is a primary pollutant and mainly contributed by vehicular exhausts in urban regions. In contrast to ozone, NOx mixing ratios were low in daytime and exhibited negative vertical gradients throughout the day, as shown in Figs. 2B and 3A-B. In nighttime, large amounts of local NOx emissions were trapped and accumulated in a shallow boundary layer (<100 m). NOx concentrations rapidly decreased with height even in the overlying residual layer due to the suppression of turbulence vertical mixing. With the onset of sunlight, the PBL rapidly expanded due to the surface heating effect. The accumulated high concentrations of NOx in the shallow nocturnal boundary layer were thereupon diluted and removed by photochemical reactions.

Ox is frequently used as a conserved metric to investigate temporal and spatial variability of ozone by eliminating the NO titration effect. As shown in Fig. 2C, the mixing ratios of Ox had similar diurnal and vertical variation patterns to those of ozone, but the vertical gradients of Ox were weaker than those of ozone. This result suggests that the vertical distribution of NO concentrations played an important role in regulating the vertical change of ozone concentrations. The enhanced positive gradients of ozone mixing ratios at night were predominantly due to the strict suppression of turbulence vertical mixing *(Geyer and Stutz, 2004)*. The higher concentrations of ozone aloft are considered as the residual of the ozone produced in the daytime PBL and have been recognized as an important reservoir for the enhancement of surface ozone in morning periods *(Kaser et al., 2017; Li and Fan, 2022; He et al., 2023)*.

Benzene and toluene demonstrated similar diurnal and vertical variations to NOx, with low concentrations in daytime and high concentrations at night, as shown in Figs. 2D-F and 3A-B. The concentrations of both benzene and toluene decreased with height throughout the day, confirming their primary emissions from ground-level sources. However, unlike benzene, the diurnal and vertical variations of toluene were more pronounced. Isoprene emissions are highly dependent on solar radiation, resulting in its higher concentrations in the early afternoon compared to other times of the day. Isoprene mixing ratios also exhibited strong negative vertical gradients below 320 m throughout the day. In contrast to toluene, isoprene concentrations decreased more

rapidly with height in the daytime. For instance, the mixing ratios of isoprene decreased by approximately 70% from 5 to 320 m in the daytime, while it was only 30% for toluene.

Fig. 3A-B show the average vertical profiles of the NMHCs, normalized to their respective ground-level concentrations measured by the PTR-ToF-MS in daytime and nighttime. The normalized mixing ratios of the NMHCs exhibited significantly differentiated gradients in daytime. In contrast, apart from monoterpenes, the differences in vertical gradients of the normalized vertical profiles for other NMHCs were relatively small at night. The differentiated vertical gradients of the NMHCs in daytime were primarily caused by their intrinsic chemical reactivities, such as reactions with OH radicals. As shown in Fig. 4, concentration ratios of the NMHC species between 320 m and 5 m with k_{OH} values lower than 2.5×10^{-11} cm^{-3} $molecule^{-1}$ s^{-1} exhibited slight variability and rapidly declined with the further increases in k_{OH}. The lower NMHC concentrations at higher altitudes were predominantly caused by the combined effects of atmospheric diffusion and chemical removal *(Sangiorgi et al., 2011)*.

Considering the effects of atmospheric diffusion and chemical removal by reactions with OH radicals, concentration ratios of NMHC species between 320 m and 5 m in daytime can be estimated using Eq. (4):

$$y = A \times \exp(-k_{OH}[OH]\Delta t) \quad \text{Eq. (4)}$$

where *y* represents concentration ratios of the NMHC species between two altitudes, A represents the effect of atmospheric dilution, k_{OH} is the reaction rate coefficient of NMHCs with OH radicals, [OH] is the concentration of OH radical, Δt is the turbulence mixing time scale between the two altitudes. The term $[OH]\Delta t$ thus refers to the exposure of NHMCs to OH radicals between the two altitudes. As shown in Fig. 4, the average concentration ratios of NMHCs between 320 m and 5 m in daytime during the campaign can be well reproduced using Eq. (4) with the coefficients A of 0.88 and $[OH]\Delta t$ of 1.0×10^{10} molecules cm^{-3} s. Atmospheric diffusion processes have same impact on the vertical distributions of all trace gases. The differences in vertical

gradients of NMHCs were mainly determined by the differences in their chemical removal rates without considering influences from advection transport.

Methanol, as one of the most abundant OVOC species in the atmosphere, had its lowest concentrations during daytime and displayed negative vertical gradients throughout the day, as shown in Fig. 2G. The vertical and diurnal variations of methanol suggest that its ambient concentrations in urban Beijing were mainly contributed by local primary emissions. Conversely, formaldehyde and MVK+MACR (the first-generation oxidation products of isoprene), as the photochemical oxidation products of NMHCs, had higher concentrations during daytime than at night and exhibited relatively weak vertical concentration gradients (Fig. 2H-I). This is mainly because these OVOCs are produced from the oxidation of NMHCs during turbulence vertical mixing and will accumulate in high altitudes. These phenomena were also observed for other OVOC species, as shown in Fig. 3.

The vertical and diurnal variations in concentrations of ozone, NOx, and VOCs are intricately governed by their sources, chemical reactivities, and the evolution of the PBL (namely the vertical dilution conditions). A significant accumulation of VOCs in the shallow nocturnal PBL is subsequently vertically diluted and chemically removed during daytime, thereby impacting the photochemical formation of ozone within the daytime PBL. In addition, the observed vertical changes in concentrations of VOCs imply that they will play distinct roles in contributing to photochemical ozone formation.

3.2 Vertical variations in contributions of VOCs to OHR

During the daytime, VOCs are primarily oxidized by OH radicals and contribute to the photochemical formation of ozone. To provide an overview on the vertical variations in contributions of different VOCs to OHR, another 1204 ions measured by the PTR-ToF-MS and can be quantified were used for analysis. All the VOCs were classified into three large categories, namely C_xH_y (including alkanes, alkenes, aromatics, and other hydrocarbons; 121 species), OVOCs ($C_xH_yO_1$, 121 species;

$C_xH_yO_2$, 120 species; $C_xH_yO_{\geq 3}$, 256 species), and N/S-containing (653 species), as shown in Fig.5. Acetylene is included in alkenes.

Fig. 5A illustrates that the total mixing ratios of VOCs in daytime exhibited a slight downward trend from 5 m to 320 m, primarily due to the rapid decrease in mixing ratios of the C_xH_y category. The total mixing ratios of the C_xH_y category decreased from 16.8 to 10.6 ppb from 5 m to 320 m, with alkanes making the largest contribution, followed by alkenes, aromatics, and other C_xH_y. Alkanes constituted 58% of the total mixing ratios of C_xH_y at 5 m, but this proportion increased to 65% at 320 m. The fractional contributions of alkenes and aromatics in the total mixing ratios of C_xH_y slightly declined from 28% to 22% and from 12% to 10%, respectively, between these two altitudes. As for OVOCs, the $C_xH_yO_1$ category was the most abundant among the measurements, contributing to 52%-58% of the total mixing ratios at the five heights, followed by the $C_xH_yO_2$ (8%-10%), and $C_xH_yO_{\geq 3}$ (2%) categories. The mixing ratios of the N/S-containing category slightly varied around 2.8 ppb between 5-320 m, contributed to approximately 6% of the total VOC concentrations.

Similar to the vertical variations in concentrations, OHRs of the C_xH_y category, denoted by OHR_{CH}, also rapidly decreased from 6.9 s^{-1} to 2.5 s^{-1} between 5 and 320 m, accounting for 52%-31% in the total OHRs of VOCs (Fig. 5B). Fractional contributions of alkenes (40-18%), alkanes (5%), and aromatics (5%-4%) to the total OHRs of VOCs all exhibited decreasing tendencies from 5 m to 320 m. The total OHRs of alkenes decreased more quickly from 5 to 320 m than those of alkanes and aromatics. OHRs of the other C_xH_y category stabilized at approximately 0.3 s^{-1} below 320 m, exhibiting an increasing contribution (2%-4%) to the total OHRs of VOCs with the increase in height. The OHRs of other VOC categories only slightly varied without exhibiting a clear variation trend from 5 to 320 m during the day. As a result, fractional contributions of the $C_xH_yO_1$ (27%-42%), $C_xH_yO_2$ (12%-18%), and $C_xH_yO_{\geq 3}$ (5%-7%), and N/S-containing (2%-4%) categories in the total OHRs of VOCs all increased with height. The increased contributions of OVOCs and N/S-containing species to the total

concentrations and OHRs of VOCs implied that air masses became more aged with the increase in height.

As depicted in Fig. 6A-B, high OHR_{CH} values were mainly constrained in the PBL and is mainly contributed by biogenic hydrocarbons, specifically isoprene, during daytime due to their high OH reactivities and enhanced emissions. The fractional contributions of isoprene in OHR_{CH} decreased rapidly with increasing height (Fig. 7A). For instance, isoprene accounted for a campaign median fraction of 58% in OHR_{CH} at 5 m in daytime, making it a frequent contributor to photochemical ozone formation in urban regions. However, this faction decreased to 38% at 320 m. Therefore, it can be speculated that the total contributions of hydrocarbons to the total OHRs of VOCs will also rapidly decline from 320 m to the top of the PBL, which typically ranges between several hundreds of meters to approximately 2~3 km in daytime (Fig. S7).

The total concentrations and OHRs of OVOCs only slightly decreased with the increase in height below 320 m in daytime, as shown in Fig. 6C-D. This is consistent with the results of *(Wang et al., 2021b)*, which observed high concentrations of OVOCs in the upper PBL. Consequently, the ratio of OHR_{OVOC} to OHR_{CH}, denoted by OHR_{OVOC}/OHR_{CH}, rapidly increased from 0.87 at 5 m to 2.6 at 320 m (Fig. 7A). This suggests that OVOCs may play more important roles in regulating the photochemical ozone formation in the middle and upper layers. To assess their potential roles in contributing to the photochemical ozone formation throughout the PBL, we calculated the mean OHRs (MOHR) of different VOC categories in daytime using Eq. (5):

$$\mathrm{MOHR(X)} = (\sum([\mathrm{X}]_i + [\mathrm{X}]_{i-1})(h_i - h_{i-1})/2)/(320 - 5) \qquad \text{Eq. (5)}$$

where $\mathrm{MOHR(X)}$ is the MOHR of the VOC category X, $[\mathrm{X}]_i$ is the concentration of X at the i^{th} altitude (namely 5, 47, 102, 200, and 320 m for h_i) above ground level.

As shown in Fig. 7B, the campaign median MOHR for isoprene was 1.7 s^{-1} and accounted for 48% of the campaign median MOHR of the C_xH_y category. This fraction was significantly lower than that of isoprene (57%) in OHR_{CH} at 5 m. In addition, the campaign median MOHR of the C_xH_y category (3.5 s^{-1}) was also significantly lower

than the OHR_{CH} (6.0 s^{-1}) at 5 m. By contrast, the campaign median MOHR of OVOCs (4.8 s^{-1}) was comparable to that of OHR_{OVOC} (4.9 s^{-1}) at 5 m. As unsaturated hydrocarbons, most alkene species are more reactive than alkanes and aromatics *(Atkinson and Arey, 2003)*. As a result, alkenes had dominant contributions to the MOHR of the C_xH_y category and the OHR_{CH} at 5 m in daytime. As shown in Fig. 7C, the campaign mean OHRs of alkanes, alkenes, and aromatics at 5 m in daytime were 0.7, 5.2, and 0.7 s^{-1}, respectively, accounting for 10%, 75%, and 10% of the OHR_{CH}. However, the campaign mean MOHRs of alkanes, alkenes, and aromatics were 0.5, 2.7, and 0.5 s^{-1}, respectively, accounting for 12%, 68%, and 12% of the MOHR of NMHC. We can also expect that the total contributions of alkenes to the MOHR of the C_xH_y category in daytime will significantly decrease if their vertical distributions in the whole PBL are considered.

This study investigated and compared the vertical profiles of measured OHR_{VOCs} and calculated OHR_{CH} during daytime over the period of July 28-31, as shown in Fig. 7D. The campaign median of the measured OHR_{VOCs} exhibited a slow decrease from 38.4 s^{-1} at 5 m to 25.4 s^{-1} at 320 m. As anticipated, the OHR_{CH}/OHR_{VOCs} ratio declined rapidly from 16% to 7% from 5 to 320 m. It is important to note that the small OHR_{CH}/OHR_{VOCs} ratio and its declining trend with the increasing height do not imply the insignificant roles of hydrocarbons in regulating the secondary pollutant formation in higher altitudes. The measured concentrations of hydrocarbons are merely the remnants of chemical reactions. The oxidation products of NMHCs, such as OVOCs and organic nitrates, formed during vertical mixing in daytime, will continue to participate in atmospheric chemical reactions.

3.3 Vertical variations in photochemical ozone formation

The surface ozone budget is intimately linked to the vertical variations of photochemical ozone formation throughout the PBL. Previous studies have consistently reported that the photochemical formation of ozone, encompassing both $P(O_3)$ and ozone formation regimes (namely the NOx-limited, VOCs-limited, and transition

426 regimes), are highly dependent on the change in its precursors *(Shao et al., 2021; Yang*
427 *et al., 2022)*. Consequently, any changes in the concentrations and compositions of
428 VOCs and NOx within the PBL will inevitably lead to alternations in the vertical
429 distribution of $P(O_3)$ and ozone formation regimes *(Tang et al., 2017; Li et al., 2024)*.

430 Fig. 8A illustrates the average dependence of $P(O_3)$ on NOx concentrations along
431 with the normalized probability density (NPD) distribution of NOx concentrations at 5
432 m, 200 m, and 320 m in daytime during the field campaign. At different heights, $P(O_3)$
433 all rapidly increased with the rise in NOx until a critical NOx mixing ratio was reached,
434 after which $P(O_3)$ decreased slowly. The critical NOx mixing ratios decreased from
435 approximately 9.5 ppb at 5 m to 5.0 ppb at 320 m, primarily caused by the decreases in
436 both NOx concentrations and the OHRs of VOCs. As also shown in Fig. 8A, the
437 majority of the measured NOx mixing ratios fall into the transition zone of the $P(O_3)$-
438 NOx curves, suggesting that the photochemical ozone formation in Beijing belonged to
439 the transition regime below 320 m.

440 RIR values were also calculated using the box model results to further elucidate
441 the sensitivities of photochemical ozone formation to changes in multiple precursors at
442 different altitudes. As shown in Fig. 8B, positive RIR values were observed for both
443 NOx and various VOC groups at the five heights, further confirming that the
444 photochemical ozone formation belonged to the transition regime in the lower layer.
445 RIR values for NOx rapidly declined from 5 to 320 m, implying that the photochemical
446 ozone formation in higher altitudes was more prone to be controlled by the abundance
447 of VOCs. This is also manifested by the increasing RIR values for both AVOCs and
448 OVOCs from 5 m to 320 m. RIR values for BVOCs significantly decreased with height
449 due to their rapid removal by reactions with OH radicals when being vertically mixed.
450 These results are consistent with the results in section 3.3 that the less reactive AVOCs
451 and OVOCs are the dominant species in regulating the photochemical formation of
452 ozone in urban regions aloft.

453 According to the vertical distribution patterns of the photochemical ozone
454 formation regime, $P(O_3)$ decreases with increasing height alongside simultaneous

455 declines in concentrations of both NOx and VOCs. Fig. 8C presents the average diurnal
456 and vertical variations in $P(O_3)$ calculated by the box model during the campaign. The
457 $P(O_3)$ values were higher in daytime and correlated well with $j(NO_2)$. $P(O_3)$ decreased
458 from the ground to 320 m, where it still maintained a relatively high value of
459 approximately 10 ppb h^{-1} at noon. These results highlight that the photochemical
460 formation of ozone aloft also remained strong compared to those at ground level.
461 Consequently, the downward transport of ozone from high altitudes, driven by
462 turbulence mixing, can become significant sources of surface ozone during the day
463 *(Karl et al., 2023)*.

464 Due to the measurement height limitation, the vertical distributions of $P(O_3)$ in the
465 middle and upper parts of the PBL were not determined in this study. As reported by
466 the work in *(Benish et al., 2020)*, $P(O_3)$ typically exhibited weak and nearly linear
467 decline tendencies from 300 m to the top of the PBL during daytime. $P(O_3)$ at the PBL
468 top was approximately half of that at 300 m. Consequently, we can assume that $P(O_3)$
469 decreased linearly from 320 m to the top of the PBL. The integral of $P(O_3)$ at different
470 heights within the PBL can then be estimated using a similar method as described in
471 Eq. (5).

472 As shown in Fig. 8D, the total amount of ozone photochemically produced below
473 47 m constituted a mere 6% of the entire PBL. This fractional contribution increased to
474 approximately 35% at 320 m, further corroborating that the majority of the boundary-
475 layer ozone was produced in the middle and upper layers. Given the enhancement of
476 turbulence vertical mixing in daytime, ozone produced at high altitudes becomes a
477 significant source of surface ozone. This is substantiated by the widespread reports of
478 strong downward ozone fluxes in the bottom part of the PBL (tens of meters above
479 ground level) *(Fares et al., 2010; Liu et al., 2021; Karl et al., 2023)*. Consequently,
480 when devising ozone control strategies, particularly in urban regions with intricate
481 precursor emissions, careful considerations should be given to the vertical variations in
482 the formation regimes of ozone in the PBL.

4 Conclusions

In this study, we investigated the vertical variations, key drivers, and environmental impacts of VOCs in the PBL using tower-based online gradient measurements in urban Beijing during the summer of 2021. The diurnal and vertical variations of various VOC species were strictly regulated by the diurnal evolution of the PBL. In daytime, reactive NMHC species were rapidly oxidized when they were mixed upward along with the formation of OVOCs. As a result, OVOC species played more significant roles in regulating the photochemical ozone formation in urban regions aloft. The photochemical formation of ozone belongs to the transition regime in the lower part of the PBL and becomes more sensitive to changes in the concentrations of AVOCs and OVOCs with increasing height. $P(O_3)$ exhibited decreasing tendencies with height but remained very large in high altitudes, likely driven by the high concentrations of OVOCs and OH radicals. Therefore, careful consideration should be given to the vertical variations in both $P(O_3)$ and photochemical ozone formation regimes in the whole PBL when making regional ozone control strategies.

The vertical variations in concentrations and compositions of VOCs significantly influence the formation of secondary pollutants. Furthermore, vertical changes in chemical reaction environments (e.g., temperature, humidity, and solar radiation) and concentrations of other chemicals (e.g., particulate matters, NOx, ozone) can also impact the degradation pathways of VOCs. These factors also affect the formation pathways and production yields of secondary pollutants. This is particularly crucial for the highly reactive NMHCs in urban areas with complex anthropogenic emissions and is expected to be thoroughly elucidated in future studies.

Data availability

The observational data used in this study are available from corresponding authors upon request.

Author contributions

BY, XBL, and YH designed the research. XBL, BY, YH, XS, JQ, XH, SW, YC, QY, YS, YP, GT, JG, and MS contributed to the data collection and data analysis. XBL, SY, and BY designed and performed the box model simulations. XBL and BY wrote the paper with contributions from all coauthors. All the coauthors discussed the results and reviewed the paper.

Competing interests

The authors declare that they have no conflict of interest.

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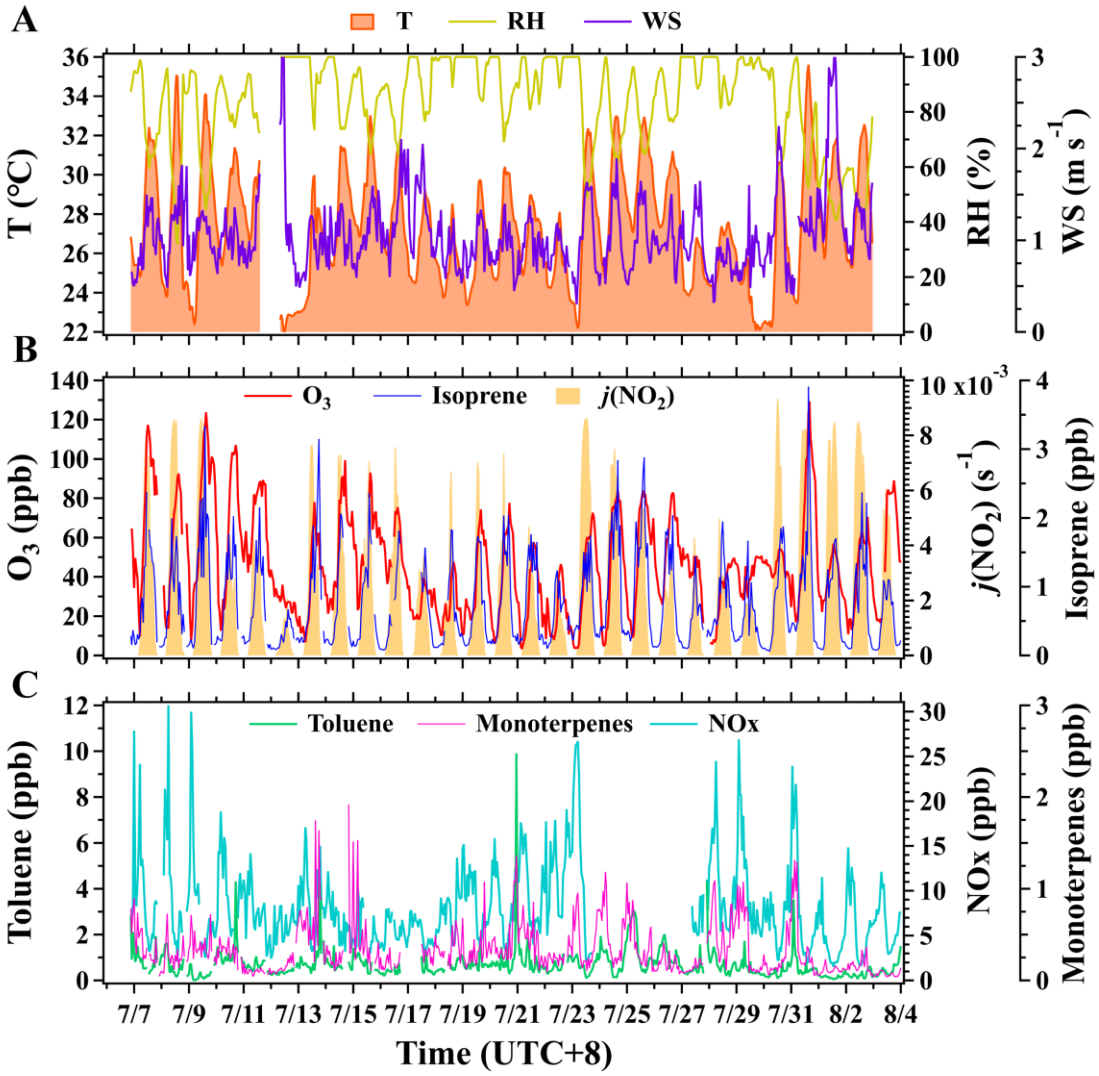
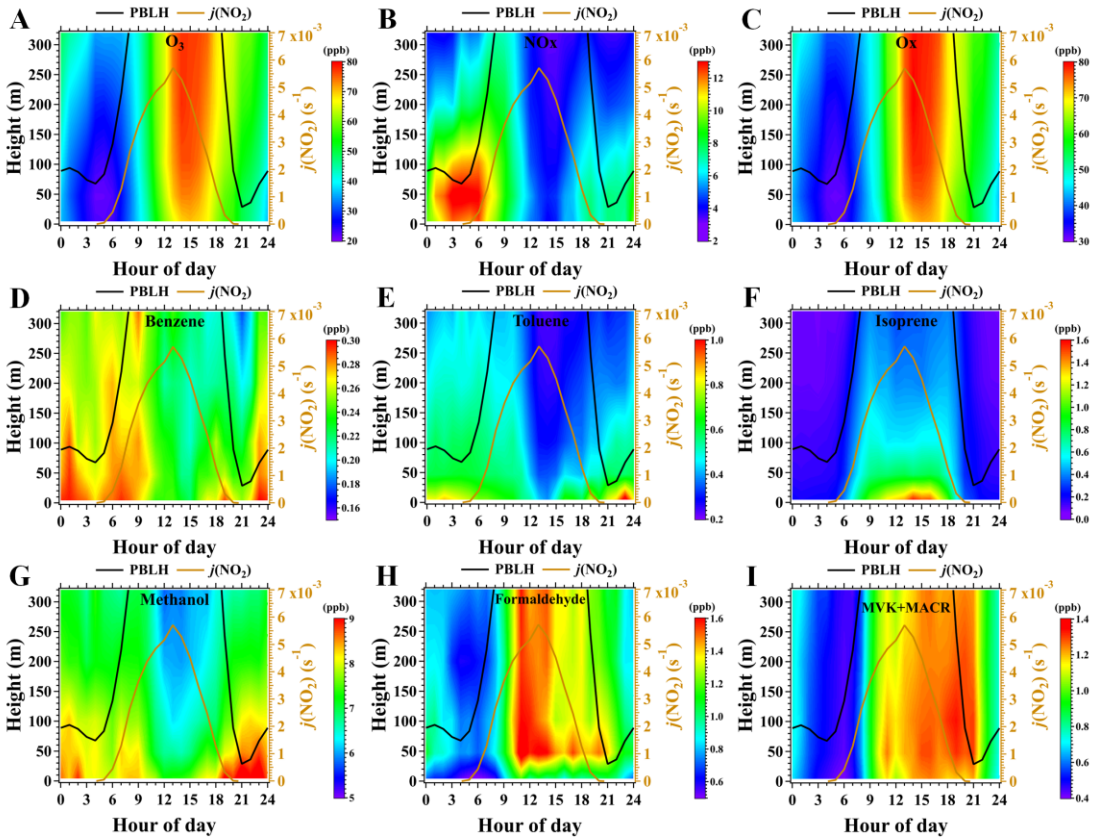


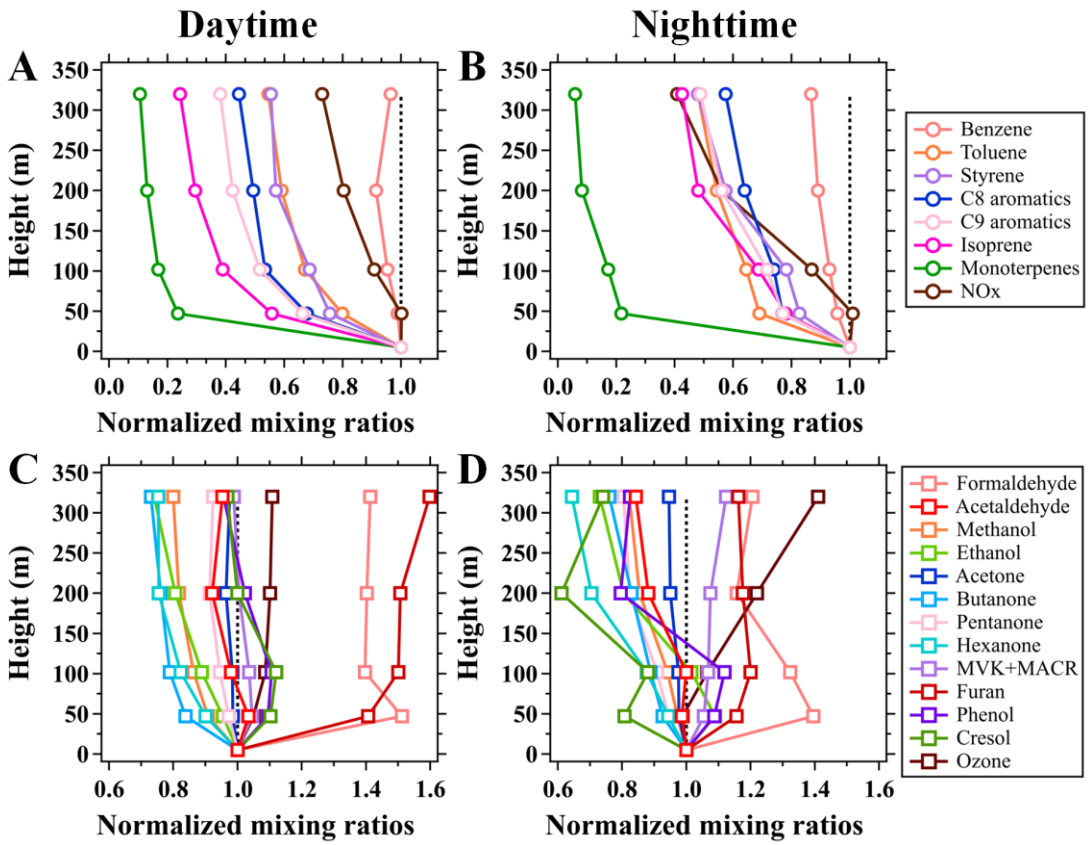
833

834 **Figure 1.** Time series of hourly mean air temperature (T), relative humidity (RH), wind
835 speed (WS), and mixing ratios of surface ozone, NOx, and VOC species along with
836 $j(NO_2)$ at the BMT site during the campaign. Meteorological parameters were measured
837 at 8 m above ground level and mixing ratios of ozone and its selected precursors were
838 measured at 5 m above ground level.



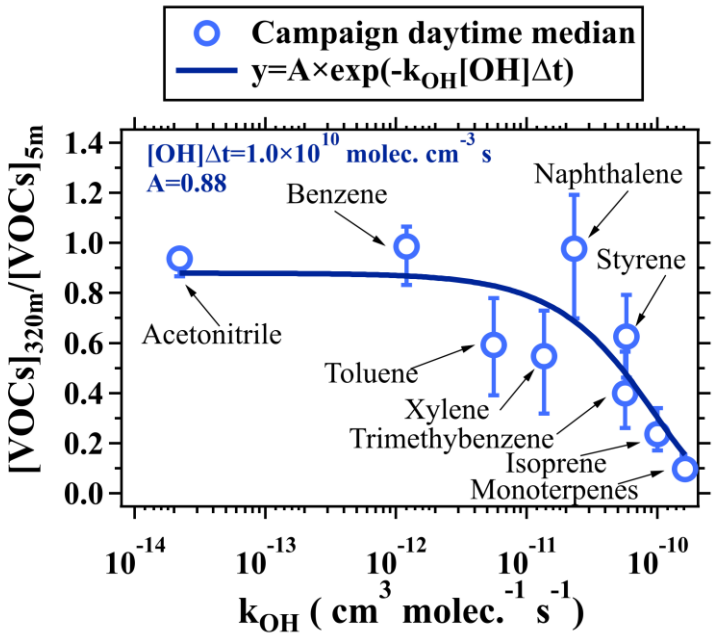
839

840 **Figure 2.** Average diurnal and vertical variations in mixing ratios ozone, NOx, Ox
841 (O_3+NO_2), and six selected VOC species along with the average diurnal profiles of
842 PBLH and $j(NO_2)$ during the campaign. The figures were obtained by linearly
843 interpolating the data at the five inlet heights on both altitude and temporal scales.



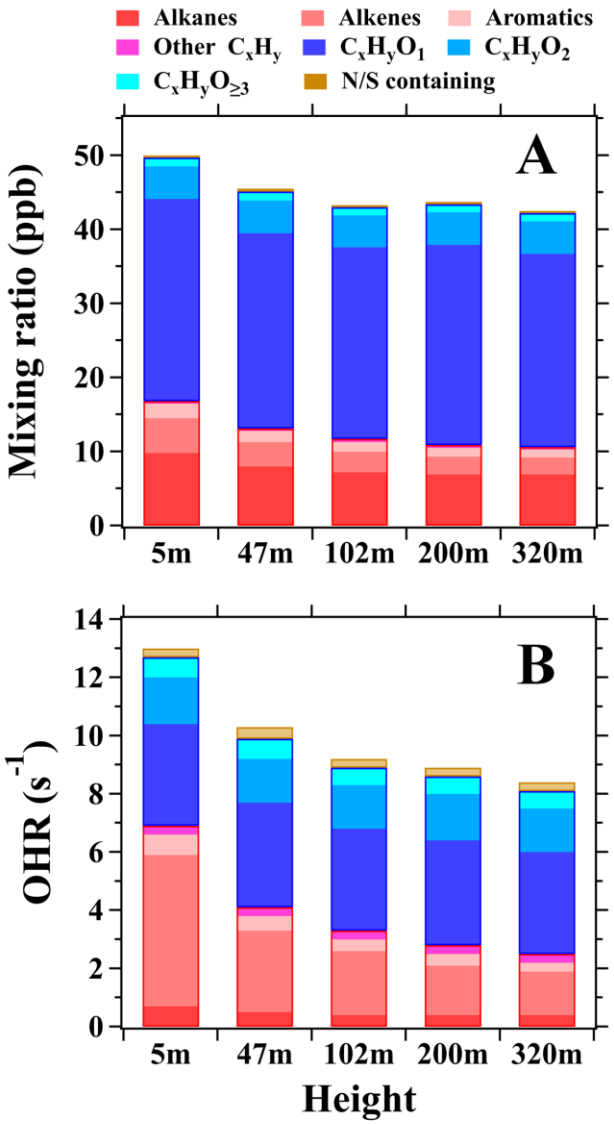
844

845 **Figure 3.** Average vertical profiles of (A-B) NMHCs and NOx, (C-D) OVOCs and O_3
846 during the daytime (11:00-16:00 LT) and nighttime (23:00-04:00 LT) of the campaign.
847 The mixing ratios of the chemical species measured above 5 m are normalized to those
848 at 5 m.



849

850 **Figure 4.** The change in ratios of NMHC concentrations (including acetonitrile)
851 between 320 m and 5 m as a function of k_{OH}. The vertically-resolved measurements of
852 VOCs made on the BMT in daytime during the campaign were used for analysis.
853 Hollow markers represent median values and error bars indicate the range between 25th
854 and 75th percentiles.



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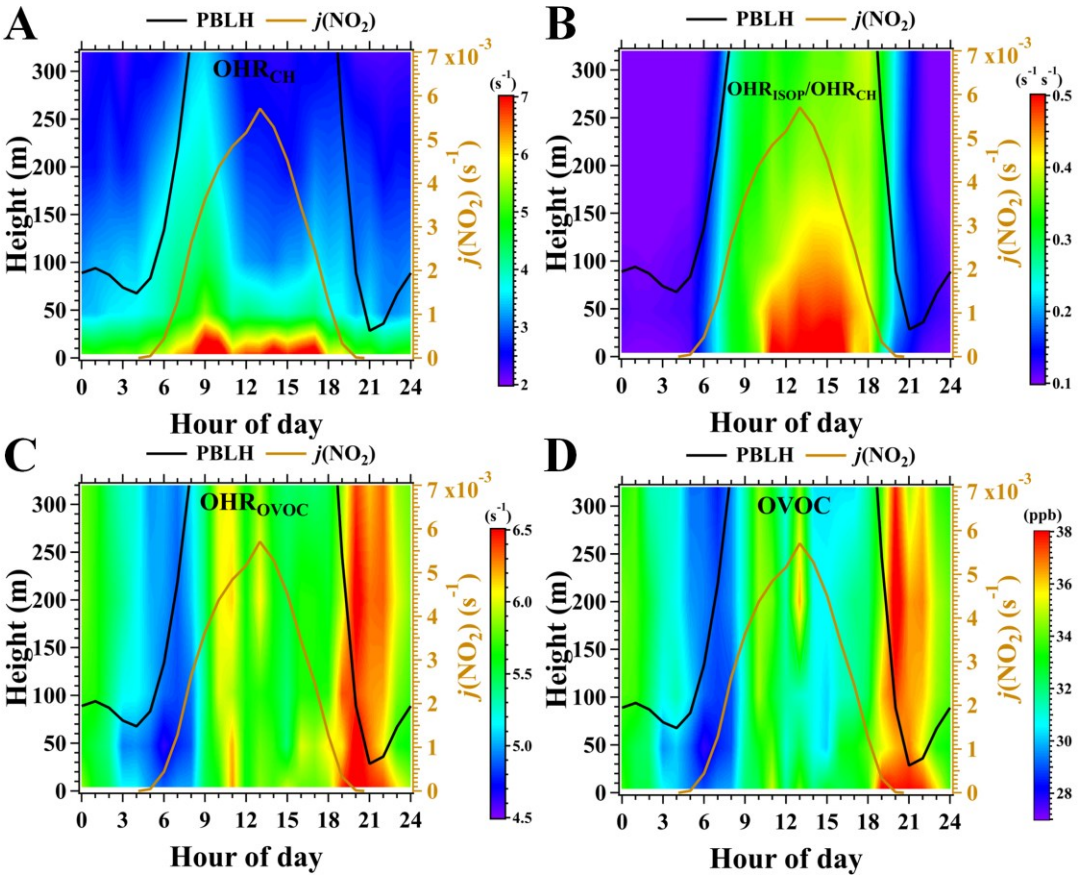
856 **Figure 5.** (A) Mean mixing ratios and (B) OHRs of different VOC categories at the five
857 inlet heights in daytime during the campaign.



858

859 **Figure 6.** (A-B) Average diurnal and vertical variations in OHRs of C_xH_y and the OHR
860 ratios of isoprene to C_xH_y (OHR_{ISOP}/OHR_{CH}) during the campaign. (C-D) Average
861 diurnal and vertical variations in mixing ratios and OHRs of OVOC. ISOP refers to
862 isoprene. The figures were obtained by linearly interpolating the data at the five
863 measurement heights on both altitude and temporal scales.

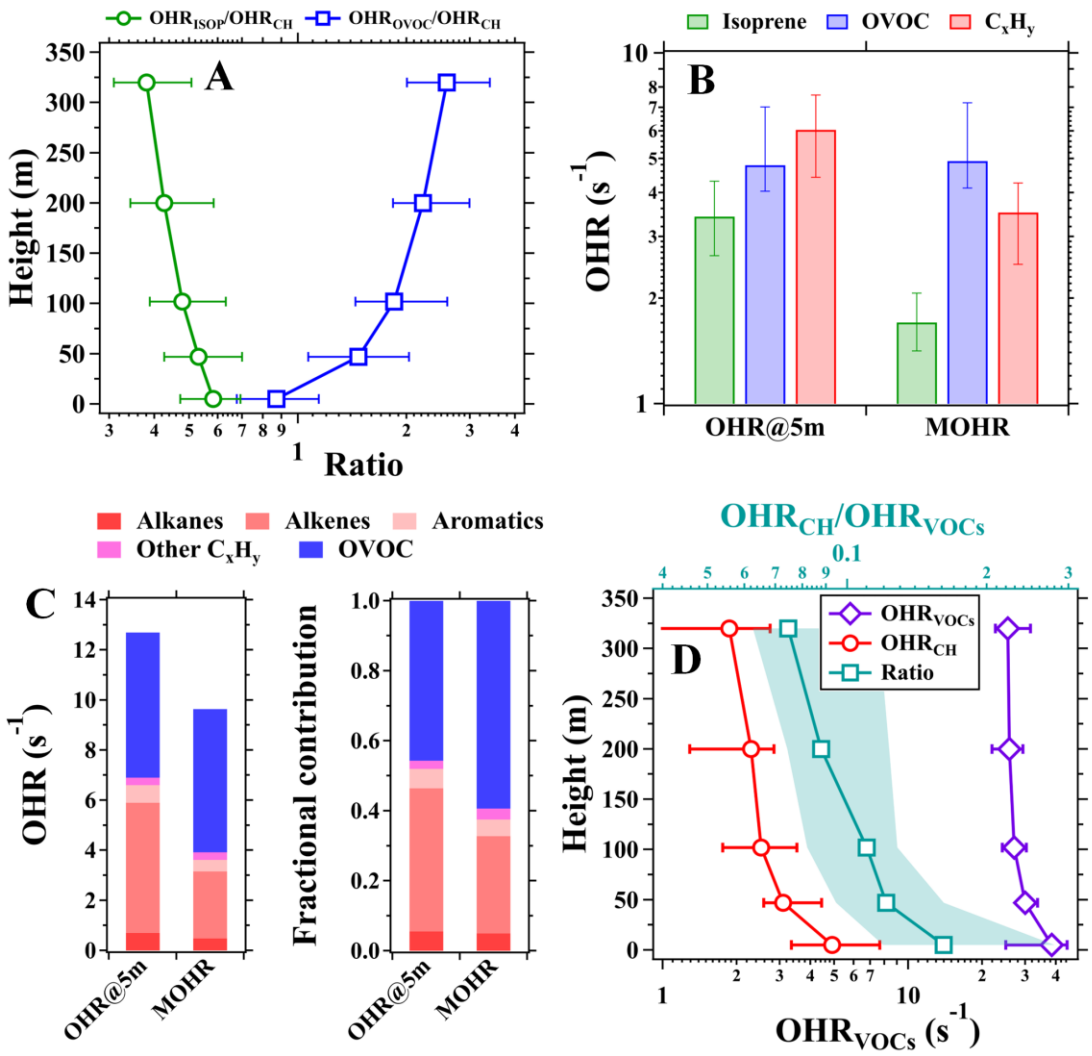
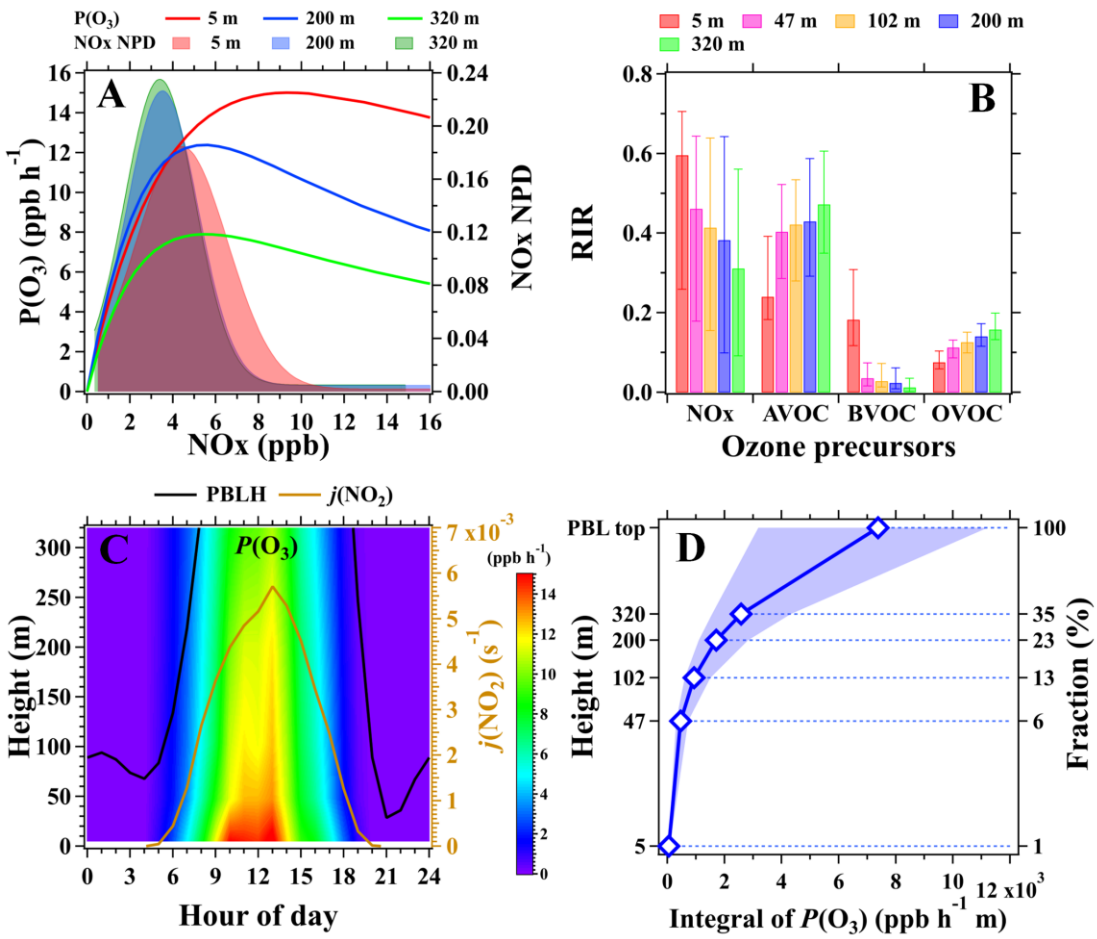


Figure 7. (A) Average vertical profiles of OHR ratios of isoprene to C_xH_y (OHR_{ISOP}/OHR_{CH}) and OVOC to NMHC (OHR_{OVOC}/OHR_{CH}). (B) Median values of the OHR at 5 m and the mean OHR (MOHR) between 5 m and 320 m for isoprene, OVOC, and C_xH_y. (C) Mean contributions of different VOC categories to the MOHR below 320 m and the OHR at 5 m. (D) Vertical profiles of the measured OHR_{VOCs} and the calculated OHR_{CH} (bottom axis) and the OHR_{CH}/OHR_{VOCs} ratios (top axis) during July 28-31, 2021. The data used for analysis in panels A-D was within the time window from 11:00 to 16:00 LT during the campaign. Markers in panels A and D represent median values. Shaded areas and error bars in panels A, B, and D indicate the range between 25th and 75th percentiles.



875

876 **Figure 8.** (A) Left axis: average dependence of $P(O_3)$ on NOx concentrations in
877 daytime during the campaign; Right axis: normalized probability density (NPD) of
878 NOx mixing ratios in daytime at the three inlet heights. (B) Median RIR values of
879 photochemical ozone formation to changes in NOx, AVOC (NMHCs excluding BVOC),
880 BVOC (isoprene), and OVOC (nine OVOC species in Table S1) at the five inlet heights;
881 Error bars indicate the range between 25th and 75th percentiles. (C) Average diurnal and
882 vertical variations in $P(O_3)$ during the campaign; The figure was obtained by linearly
883 interpolating the data at the five measurement heights on both altitude and temporal
884 scales. (D) The vertical profile of the integral of $P(O_3)$ in daytime during the campaign;
885 Markers indicate median values and Shaded areas indicate the range between 25th and
886 75th percentiles.