



1 **The Critical Role of Aqueous-Phase Processes in**
2 **Aromatic-Derived Nitrogen-Containing Organic**
3 **Aerosol Formation in Cities with Different**
4 **Energy Consumption Patterns**

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20 **Abstract.** Nitrogen-containing organic compounds (NOCs) impact air quality and
21 human health. Here, the abundance, potential precursors, and main formation
22 mechanisms of NOCs in PM_{2.5} during winter were compared for the first time among
23 Haerbin (coal-dependent for heating), Beijing (natural gas and coal as heating energy),
24 and Hangzhou (no centralized heating policy). The total signal intensity of CHON⁺,
25 CHN⁺, and CHON⁻ compounds was highest in Haerbin and lowest in Hangzhou.
26 Anthropogenic aromatics accounted for 73%–93% of all identified precursors of
27 CHON⁺, CHN⁺, and CHON⁻ compounds in Haerbin. Although the abundance of
28 aromatics-derived NOCs was lower in Beijing than in Haerbin, aromatics were also
29 the main contributors to NOC formation in Beijing. Hangzhou exhibited the lowest
30 levels of aromatic precursors. Furthermore, non-metric multidimensional scaling
31 analysis indicated an overall reduction in the impact of fossil fuel combustion on
32 NOC pollution along the route from Haerbin to Beijing to Hangzhou. We found that
33 aqueous-phase processes (mainly condensation, hydrolysis or dehydration processes
34 for reduced NOCs, and mainly oxidization or hydrolysis processes for oxidized NOCs)
35 can promote the transformation of precursors to produce NOCs, leading to the most
36 significant increase in aromatic NOC levels in Haerbin (particularly on haze days).
37 Reduced precursor emissions in Beijing and Hangzhou (the lowest) constrained the
38 aqueous-phase formation of NOCs. The overall results suggest that the aerosol NOC
39 pollution in coal-dependent cities is mainly controlled by anthropogenic aromatics



40 and aqueous-phase processes. Thus, without effective emission controls, the
41 formation of NOCs through aqueous-phase processes may still pose a large threat to
42 air quality.

43

44 **Keywords:** Aerosols, Nitrogen-containing organic compounds, Heating energy

45 consumption, Anthropogenic pollutants, Formation mechanism

46



47 **1. Introduction**

48 Nitrogen-containing organic compounds (NOCs) are abundant reactive nitrogen
49 species in aerosol particles, accounting for up to 40%–80% of total nitrogen
50 deposition (Li et al., 2023; Xi et al., 2023; Yu et al., 2020). Clearly, aerosol NOCs can
51 significantly contribute to the global nitrogen cycle (Li et al., 2023; Cape et al., 2011).
52 Moreover, the formation of secondary organic aerosols (SOA) and light-absorbing
53 organic aerosols (e.g., brown carbon) is also tightly associated with NOCs (Wang et
54 al., 2024; Liu et al., 2023b; Zeng et al., 2021), thus affecting the radiative balance and
55 air quality (Yuan et al., 2023; Jiang et al., 2023). In particular, certain NOCs, such as
56 nitroaromatics and peroxyacyl nitrates, are characterized as phytotoxins and potential
57 carcinogens, posing threats to ecosystems and human health (Shi et al., 2023; Singh
58 and Kumar, 2022; Huang et al., 2024). Therefore, understanding the characteristics,
59 origins, and atmospheric processes of NOCs is essential for comprehending their
60 climate and health effects.

61 Aerosol NOCs can be derived from primary emissions associated with
62 anthropogenic activities and natural sources (Cape et al., 2011; Lin et al., 2023; Xu et
63 al., 2020a; Wang et al., 2017; Song et al., 2018; Song et al., 2022; Ma et al., 2024).
64 Secondary formation processes may play a more crucial role in the formation of
65 NOCs in fine aerosol particles, which involve interactions among volatile organic
66 compounds (VOCs), atmospheric oxidants, and reactive inorganic nitrogen species



67 (Montoya-Aguilera et al., 2018; Perraud et al., 2012; Hallquist et al., 2009). For
68 instance, laboratory studies have observed the formation of organic nitrates from the
69 oxidation of isoprene and α -/ β -pinene by atmospheric oxidants and nitrogen oxide
70 (NO_x) (Surratt et al., 2010; Rollins et al., 2012; Nguyen et al., 2015). Additionally,
71 aqueous-phase reactions of NH_4^+ (or NH_3) with biogenic VOCs or carbonyl
72 compounds have been suggested to be important mechanisms of reduced NOC (Re-
73 NOCs) formation (Abudumutailifu et al., 2024; Laskin et al., 2014; Li et al., 2019b;
74 Liu et al., 2023b; Wang et al., 2024). However, understanding the origins, formation
75 mechanisms, and environmental impacts of NOCs is hindered by the elusive and
76 intractable molecular information regarding NOCs and their precursors.

77 Aerosol liquid water (ALW) can greatly increase the formation of aerosol NOCs
78 by facilitating the conversion of water-soluble organic gases into particles and
79 subsequently enabling aqueous-phase reactions (Li et al., 2019a; Lv et al., 2022; Liu
80 et al., 2023b). A positive correlation between aerosol NOC abundance and either ALW
81 or relative humidity (RH) has been observed in several observation studies (Jiang et
82 al., 2023; Liu et al., 2023b; Xu et al., 2020b). In particular, it has been suggested that
83 increased ALW levels can exacerbate winter haze in China (Wu et al., 2018; Hodas et
84 al., 2014; Lv et al., 2022; Wang et al., 2021d; Liu et al., 2023b; Wang et al., 2021a; Li
85 et al., 2019a). Presumably, precursors and ALW are the two key factors in the
86 formation of aerosol NOCs. Haze environments have potentially high RH levels and



87 large emissions of NOC precursors. Moreover, in Chinese cities with different energy
88 consumption (e.g., coal, biomass, and natural gas) for winter heating (Zhang et al.,
89 2021b; Zhang et al., 2023b), the types and emission intensities of pollutants released
90 from different heating sources are expected to vary considerably (Bond et al., 2006;
91 Stockwell et al., 2015; Křůmal et al., 2019). However, the potential effects of ALW in
92 the formation of NOCs in Chinese cities with different energy consumption during
93 winter, particularly in haze periods, are not well documented. Moreover, the roles of
94 ALW-related NOC formation processes in the formation of haze in cities with
95 different energy consumption types also remain largely unknown.

96 In this study, we present the measurements of the NOCs and other chemical
97 compositions in PM_{2.5} collected from three cities (Haerbin, Beijing, and Hangzhou)
98 with different energy consumption during winter. The specific objectives of this study
99 were: (1) to investigate the differences in the abundance, composition, and major
100 precursors of NOCs in different cities with different energy consumption, especially
101 on polluted days; and (2) to elucidate the potential effects of ALW on the formation of
102 oxidized NOCs (Ox-NOCs) and reduced NOCs (Re-NOCs) during winter
103 (particularly on polluted days) in cities with different energy consumption. The
104 research findings are expected to provide valuable implications for the mitigation of
105 aerosol NOCs pollution in urban environments.

106



107 **2. Materials and methods**

108 **2.1. Study site description and sample collection**

109 The study sites are located in three urban areas, including Haerbin (HEB, i.e.,
110 Harbin, 126.64°E, 45.77°N), Beijing (BJ, 116.41°E, 40.04°N), and Hangzhou (HZ,
111 120.16°E, 30.30°N) (**Fig. S1**). The city of HEB, with a population density of 9.95
112 million, is situated in the northeastern region of China. It relies heavily on coal for
113 centralized heating during winter. The rapid urbanization and increased coal
114 consumption have significantly deteriorated air quality in HEB in recent years (Ma et
115 al., 2020). In contrast, BJ has largely shifted towards the utilization of cleaner energy
116 sources (e.g., natural gas) for centralized heating in recent years, particularly
117 following the implementation of the "Beijing 2013–2017 Clean Air Action Plan" (Vu
118 et al., 2019; Yuan et al., 2023). HZ, situated within the Yangtze River Delta, is exempt
119 from the necessity of heating due to the relatively mild winter climate (average
120 temperature of 6.6 ± 2.4 °C during the sampling period, **Table S1**). Clearly, the
121 distinctive energy consumption patterns observed in these three cities during winter
122 provide a valuable opportunity to examine the impact of various precursors and
123 aqueous-phase processes on aerosol NOC formation.

124 Sample collection was carried out simultaneously in three cities from 16
125 December 2017 to 14 January 2018. PM_{2.5} samples were collected every two or three
126 days with a duration of 24 hours onto prebaked (450°C for ~10 hours) quartz fiber



127 filters (Pallflex, Pall Corporation, USA) using a high-volume air sampler (Series 2031,
128 Laoying, China). One blank sample was collected at each sampling site. A total of 39
129 samples were collected, all of which were stored at -30°C . Meteorological data (e.g.,
130 temperature, relative humidity (RH) and wind speed) together with concentrations of
131 various pollutants (e.g., SO_2 and NO_2) were obtained from nearby environmental
132 stations. The sampling periods were classified as either "clean" or "haze" based on
133 whether the daily average concentration of $\text{PM}_{2.5}$ was below or above $75 \mu\text{g m}^{-3}$
134 (Zhang and Cao, 2015; Xu et al., 2024).

135

136 **2.2. Chemical analysis**

137 The extraction and analysis methods for NOCs were consistent with those
138 described in our recent publication (Ma et al., 2024). Briefly, a portion of each filter
139 was extracted with methanol (LC-MS grade, CNW Technologies Ltd.) using
140 sonication in an ice bath ($\sim 4^{\circ}\text{C}$). The extracts were filtered through a $0.22 \mu\text{m}$
141 polytetrafluoroethylene syringe filter and then concentrated to $300 \mu\text{L}$ under a gentle
142 stream of nitrogen gas. The concentrated extracts underwent composition analysis via
143 an ultra-performance liquid chromatography coupled with quadrupole time-of-flight
144 mass spectrometry equipped with an electrospray ionization (ESI) source (UPLC-ESI-
145 QToFMS, Waters Acquity Xevo G2-XS) (Wang et al., 2021c; Ma et al., 2024). This
146 analysis was done in both ESI+ and ESI- modes. The organic compounds were



147 separated on a C₁₈ column (2.1 × 100 mm, 1.7 μm particle size, Waters) with an 18-
148 minute gradient elution. The mobile phases comprised ultrapure water with 0.1%
149 formic acid (A) and methanol with 0.1% formic acid (B). Gradient elution was
150 conducted according to the following protocol: 1% B was held for 1.5 minutes,
151 followed by an increase to 54% B over a period of 6.5 minutes. Thereafter, the B was
152 increased to 95% over a period of 3 minutes. After reaching 100% B in one minute,
153 this state was maintained for 3 minutes. Finally, the concentration was returned to 1%
154 B in 0.5 minutes and held for 2.5 minutes. Due to uncertainties in ionization
155 efficiencies for different compounds (Ditto et al., 2022; Yang et al., 2023), an
156 intercomparison (mainly compared among samples within this study) of compound
157 relative abundance was conducted without accounting for differences in ionization
158 efficiency in the present study. This consideration was consistent with previous
159 studies (Xu et al., 2023; Jiang et al., 2022; Ma et al., 2024).

160 Another filter portion was ultrasonically extracted using Milli-Q water (~4°C ice
161 bath) to analyze the concentrations of inorganic ions and organic acids. The inorganic
162 ions (e.g., NO₃⁻, SO₄²⁻, Cl⁻, Ca²⁺, Mg²⁺, Na⁺, and NH₄⁺) and organic acids (e.g.,
163 formic acid, acetic acid, oxalic acid, succinic acid, glutaric acid, and methanesulfonic
164 acid) were quantified using an ion chromatograph system (Dionex Aquion, Thermo
165 Scientific, USA) as described previously (Xu et al., 2022b; Yang et al., 2024b).

166



167 **2.3. Compound categorization and precursor identification**

168 The identified molecular formulas via UPLC-ESI-QToFMS were categorized
169 into different compound classes based on their elemental compositions, which
170 included CHO⁻, CHON⁻, CHONS⁻, and CHOS⁻ in ESI⁻ mode and CHO⁺, CHON⁺,
171 and CHN⁺ in ESI⁺ mode (Ma et al., 2024). Here, we mainly focus on NOCs (i.e.,
172 CHN⁺, CHON⁺, and CHON⁻ compounds) (Ma et al., 2024; Jiang et al., 2022; Wang
173 et al., 2017). The carbon oxidation state (OS_C) and double bond equivalent (DBE)
174 were calculated to indicate the oxidation level and unsaturation degree of the organics,
175 respectively (**Sect. S1**) (Kroll et al., 2011; Ma et al., 2024). Additionally, the modified
176 aromaticity index (AI_{mod}) and aromaticity equivalent (X_C) were used to evaluate
177 aromaticity of organics (Koch and Dittmar, 2006), as detailed in **Sect. S1**.

178 The potential precursors of NOCs were identified based on the methodology
179 outlined in previous studies (Nie et al., 2022; Guo et al., 2022; Jiang et al., 2023).
180 The classification of CHON⁺ and CHON⁻ compounds was refined into following
181 categories, including aliphatics-, heterocyclics-, and aromatics-derived Re-NOCs
182 and isoprene-, monoterpenes-, aliphatics-, and aromatics-derived Ox-NOCs.
183 Moreover, CHN⁺ compounds were classified into aliphatic, monoaromatic, and
184 polyaromatic CHN⁺ compound (Wang et al., 2021b; Yassine et al., 2014). A detailed
185 description of the revised workflow for classifying NOCs according to potential
186 precursors was provided in **Sect. S2** and **Fig. S2**.



187

188 **2.4. Classification of potential pathways for NOC formation**

189 To identify potential aqueous-phase processes for aerosol NOC formation, we
190 screened precursor-product pairs from the organic compounds that have been detected
191 (Su et al., 2021; Xu et al., 2023; Jiang et al., 2023). The reaction pathways of Re-
192 NOCs (mainly CHON⁺ compounds in this study) were refined into the following
193 categories, including condensation (cond_N), hydrolysis (hy_N), dehydration (de_N),
194 cond_hy_N (involving cond_N and hy_N), cond_de_N (involving cond_N and de_N),
195 hy_de_N (involving hy_N and de_N), cond_hy_de_N (involving cond_N, hy_N and
196 de_N) and unknown_N (unknown processes) formation pathways (**Fig. S3 and Table**
197 **S4**) (Sun et al., 2024; Abudumutailifu et al., 2024; Laskin et al., 2014; Liu et al.,
198 2023c). Another significant class of Re-NOCs is the CHN⁺ compounds. Their
199 potential formation mechanisms include cond_N, de_N, cond_de_N, and other
200 unidentified (unknown_N) pathways (**Fig. S4 and Table S4**) (Abudumutailifu et al.,
201 2024; Laskin et al., 2014; Liu et al., 2023c). In addition, the reaction pathways of Ox-
202 NOCs (mainly CHON⁻ compounds in this study) were refined into the following
203 categories, including ox_N, hy_N, ox_hy_N (involving ox_N and hy_N), and other
204 unidentified (unknown_N) pathways (Jiang et al., 2023; Su et al., 2021) (**Fig. S5 and**
205 **Table S4**). A detailed overview of the methodologies employed to discern potential
206 NOC formation pathways was shown in **Sect. S3, Table S4, and Figs. S3–S5**.



207 It is important to acknowledge the potential limitations in the categorization
208 methodology of NOC formation pathways described above. This is because the
209 approach applied here and in previous studies (Jiang et al., 2023; Su et al., 2021) may
210 classify NOCs from primary emissions as products of secondary aqueous-phase
211 reactions. Accordingly, our results should be regarded as indicating a maximal
212 potential (or an upper limit) for NOC generation from aqueous-phase reactions. In this
213 study, NOCs produced from the reaction pathways identified by the abovementioned
214 classification methodology can explain 76% of CHON⁺ compounds, 61% of CHN⁺
215 compounds, and 65% of CHON⁻ compounds. Thus, the classification of potential
216 pathways for NOC formation was representative, at least in this study.

217

218 **2.5. More parameter calculations and data analysis**

219 A thermodynamic model (ISORROPIA-II) was used to estimate the ALW
220 concentration and pH value, as described in previous studies (Xu et al., 2020b; Xu et
221 al., 2023; Xu et al., 2022c). Ambient hydroxyl radical ($\bullet\text{OH}$) concentrations were
222 predicted using empirical formulas proposed by Ehhalt and Rohrer (Ehhalt and Rohrer,
223 2000), which was reported in detail in our previous field observations (Liu et al.,
224 2023a; Lin et al., 2023). The ventilation coefficient (VC) is an indicator of the
225 potential for atmospheric dilution of pollutants, which was calculated by multiplying
226 the wind speed by the planetary boundary layer height (PBLH) (Gani et al., 2019).



227 Non-metric multidimensional scaling (NMDS) was employed to visualize the
228 distributions of NOCs (CHON⁺, CHN⁺, and CHON⁻ compounds) in two dimensions,
229 based on Bray-Curtis distances (Chao et al., 2006). The stress values ranged from
230 0.03 to 0.11 (less than 0.2, **Table S5**) in our analysis, indicating that the differences
231 among samples can be well represented in the two-dimensional pattern. To further
232 assess the influence of anthropogenic emissions and aqueous-phase processes on the
233 distribution of NOCs, the envfit function in the R package Vegan (Oksanen et al.,
234 2007) was utilized. Furthermore, the Spearman rank correlation, a non-parametric
235 measure with less sensitivity to outliers and independent of data distribution
236 assumptions, was employed to examine the association patterns between NOCs and
237 the parameters related to anthropogenic emissions and aqueous-phase processes
238 (Kellerman et al., 2014).

239

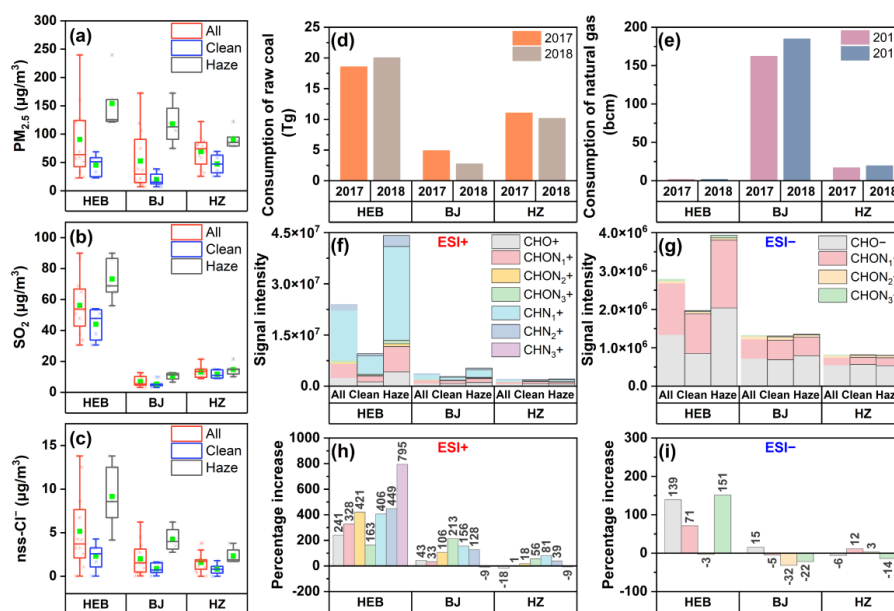
240 **3. Results and discussion**

241 **3.1. Overview of pollution and aerosol NOC characteristics in different cities**

242 **Figure 1a–c** and **Table S1** show the variations in major gaseous pollutants, PM_{2.5}
243 and its major compositions, as well as meteorological parameters in three Chinese
244 cities with different energy consumptions during winter. The average PM_{2.5}
245 concentration in HEB was $90.6 \pm 62.4 \mu\text{g m}^{-3}$, which was significantly higher than
246 that observed in BJ ($52.7 \pm 51.4 \mu\text{g m}^{-3}$) and HZ ($69.1 \pm 29.6 \mu\text{g m}^{-3}$). Similarly, the



247 concentrations of SO_2 and nss-Cl^- were higher in HEB than in BJ and HZ. In addition,
248 a lower $\text{NO}_3^-/\text{SO}_4^{2-}$ mass ratio (**Table S1**) was found in HEB. SO_2 and nss-Cl^- have
249 been suggested to be typical pollutants emitted from coal combustion during winter in
250 cities (Zhao and Sun, 1986; Streets and Waldhoff, 2000). The low $\text{NO}_3^-/\text{SO}_4^{2-}$ mass
251 ratio can indicate a predominance of stationary sources (e.g., coal combustion) (Wang
252 et al., 2006; Arimoto et al., 1996; Xiao and Liu, 2004). These results suggest that coal
253 combustion during the winter heating season in HEB may significantly contributed to
254 severe $\text{PM}_{2.5}$ pollution. This consideration can also be supported by the highest coal
255 consumption in HEB in 2017–2018 (**Fig. 1d**). Due to the large-scale use of clean
256 energy (i.e., natural gas) in BJ (**Fig. 1e**), the coal consumption in BJ was the lowest
257 (**Fig. 1d**). This results in the lowest pollutant levels in BJ. From clean to polluted days,
258 HEB and BJ showed larger increases in pollutant levels (e.g., $\text{PM}_{2.5}$, SO_2 , and CO),
259 followed by HZ. Thus, the release of pollutants caused by the use of fossil fuels for
260 centralized heating in winter (only occurred in HEB and BJ) is undoubtedly one of the
261 important factors contributing to the generation of haze in HEB and BJ.



262

263 **Figure 1.** Box and whisker plots showing variations in the concentration of (a) $PM_{2.5}$,

264 (b) SO_2 , and (c) $nss-Cl^-$ in all (gray), clean (blue), and haze (red) periods in different

265 cities. Each box encompasses the 25th–75th percentiles. Whiskers are the 5th and

266 95th percentiles. The green squares and solid lines inside boxes indicate the mean and

267 median value. The consumption of (d) raw coal and (e) natural gas in 2017 and 2018

268 in different cities was obtained from the local statistical yearbooks. Average

269 distributions in the signal intensity of species detected in $PM_{2.5}$ collected during

270 different winter periods in different cities in (f) ESI+ and (g) ESI- modes. Percentage

271 variations in the signal intensity of each subgroup from clean to haze periods in

272 different cities in (h) ESI+ and (i) ESI- modes.

273



274 **Figure 1f** and **g** show the average signal intensity distributions of organic
275 compounds detected in $PM_{2.5}$ across sampling periods in different cities. The detailed
276 mass spectra of organic compounds detected in ESI⁺ and ESI⁻ were shown in **Fig. S6**.
277 CHN_{1+} ($n = 437-448$) compounds were the main CHN molecules measured in ESI⁺
278 mode in all cities (**Fig. 1f** and **Table S6**), the signal intensity of which accounted for
279 over 77% of the total CHN_{1-3+} signal intensity. Similarly, $CHON_{1+}$ compounds ($n =$
280 $398-421$) dominated in $CHON_{1-3+}$ molecules, with a higher signal intensity than
281 $CHON_{2-3+}$ (**Fig. 1f** and **Table S6**). The high abundances of CHN_{1+} and $CHON_{1+}$
282 compounds in NOCs were similar to previous reports about the NOC characteristics
283 of urban aerosols (He et al., 2024; Abudumutailifu et al., 2024). The signal intensity
284 fractions (40%–77%) of CHN_{+} compounds in total NOCs in these three cities were
285 higher than those observed (8.20%–17.47%) during winter in Ürümqi where the same
286 NOC analysis method was conducted (Ma et al., 2024). However, the signal intensity
287 fractions of $CHON_{+}$ compounds in total NOCs were lower in these three cities
288 (23%–60%) than in Ürümqi (over 82.53%) (Ma et al., 2024). More frequent biomass
289 burning and relatively dry climate in Ürümqi (northwest China) (Ma et al., 2024) may
290 result in different sources and formation processes of NOCs compared to this study.
291 The signal intensity of these NOCs detected in ESI⁺ mode varied across cities, with
292 the highest CHN_{+} and $CHON_{+}$ signal intensities in HEB, followed by BJ and HZ.
293 Moreover, we found that the total signal intensities of CHN_{+} and $CHON_{+}$ compounds



294 increased by 382% in HEB from clean to haze periods, followed by increase of 102%
295 in BJ and increase of 31% in HZ (**Fig. 1h** and **Table S5**). This variation pattern of
296 CHN⁺ and CHON⁺ compounds from clean to haze periods was similar to that of the
297 pollutants mentioned previously (**Fig. 1a–c**). Given the high sensitivity of ESI⁺ mode
298 to protonatable species, reduced species (e.g., amine- and amide-like compounds)
299 were expected to predominate the NOCs (Han et al., 2023; Wang et al., 2018), the
300 formation of which was highly related to precursor emission level, aerosol acidity, and
301 ALW concentrations (Kuwata and Martin, 2012; Vione et al., 2005; Yang et al., 2024a;
302 Xu et al., 2020b). Thus, these results suggest that there may be significant differences
303 in the sources, precursor emission intensity, or main formation pathways of NOCs in
304 different energy consuming cities.

305 The number of NOCs identified in ESI[−] (296–301 molecules excluding sulfur-
306 containing compounds, **Table S7**) was found to be lower than that observed in ESI⁺
307 (1346–1361) (**Table S6**). This finding was similar to previous observations about the
308 NOCs of urban organic aerosols in Beijing, Mainz, Changchun, Guangzhou, and
309 Shanghai (Wang et al., 2021b; Wen et al., 2023; Wang et al., 2018). CHON₁[−]
310 compounds were the main NOC molecules in ESI[−] mode in all cities (**Fig. 1g** and
311 **Table S7**). The average signal intensity of CHON[−] compounds was highest in HEB,
312 followed by BJ and HZ. Moreover, the outbreak of CHON_{1–3} signal intensity during
313 polluted periods was found in HEB, whereas insignificant increases occurred in BJ



314 and HZ (**Fig. 1i**). Deprotonated NOCs with oxidized nitrogen-functional groups, such
315 as nitro ($-\text{NO}_2$) or nitrooxy ($-\text{ONO}_2$) groups, are more sensitive to the ESI⁻ mode
316 (Wang et al., 2017; Jiang et al., 2023; Yuan et al., 2023). Clearly, the formation of
317 aerosol CHON^- compounds was largely dependent on atmospheric oxidation capacity
318 and gas- and aqueous-phase reactions (Ng et al., 2017; Shi et al., 2020; Shi et al.,
319 2023). Thus, the differences in CHON^- compound abundance in different polluted
320 periods and cities together with the spatiotemporal changes in CHN^+ and CHON^+
321 abundances mentioned previously were likely attributed to variations in sources,
322 mechanisms, or key influencing factors of NOC formation in these three cities, which
323 will be further discussed in the following sections.

324

325 **3.2. Potential precursors of aerosol NOCs in different cities**

326 **Figure 2** presents the average signal intensity percentage and signal intensity
327 distributions of different NOCs from various precursors in different cities during
328 winter. Aromatics-, heterocyclics-, and aliphatics-derived Re-NOCs together
329 accounted for more than 74% (74%–79%) of the total signal intensity of CHON^+
330 compounds in the three study cities (**Fig. 2a–c** and **Table S8**). Specifically, the
331 proportion of the aromatics-derived CHON^+ signal intensity in the total CHON^+
332 signal intensity was much higher in HEB (73%) than in BJ (33%), with the lowest
333 proportion observed in HZ (23%) (**Fig. 2a–c**). Furthermore, we observed that

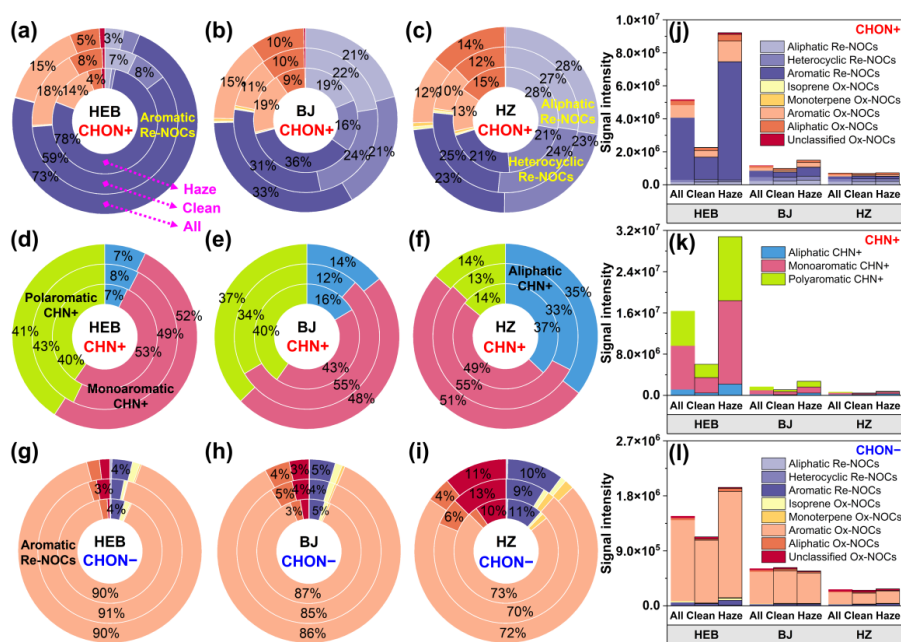


334 aromatic CHN⁺ compounds (mono- and poly-aromatics) dominated the total CHN⁺
335 compounds in both number and abundance in all investigated cities (**Table S9** and **Fig.**
336 **2d–f**). The average signal intensity percentage and signal intensity of aromatic CHN⁺
337 compounds were also highest in HEB (**Fig. 2d–f** and **k**). The calculated AI_{mod} values
338 for CHON⁺ and CHN⁺ compounds were higher in HEB than in BJ and HZ (**Table**
339 **S10**), which further indicated a higher aromaticity of these NOCs in HEB. It has been
340 suggested that coal combustion can release a large amount of aromatic compounds
341 (Zhang et al., 2023a), which potentially increased NOC aromaticity (Yuan et al.,
342 2023). Thus, the higher signal proportion of aromatics-derived Re-NOCs in HEB can
343 be explained by the higher coal combustion emissions during winter. In contrast, the
344 use of clean energy during the central heating season in BJ and the reduced emissions
345 in HZ without central heating weakened the formation of aerosol aromatic NOCs.

346 CHON[–] compounds were also primarily dominated by aromatics-derived Ox-
347 NOCs in all three cities, accounting for more than 73% (73%–90%) of the total signal
348 intensity of CHON[–] compounds, on average (**Fig. 3g–i**). This finding was consistent
349 with field observations conducted in other Chinese cities such as Shanghai,
350 Changchun, Guangzhou, and Wangdu during winter (Wang et al., 2021b; Jiang et al.,
351 2023). The abundance of aromatics-derived Ox-CHON[–] compounds and the AI_{mod}
352 value of CHON[–] were highest in HEB and decreased sequentially in BJ and HZ (**Fig.**
353 **2h** and **Table S10**), further indicating our previous consideration that coal combustion



354 heating in HEB can lead to higher NOC pollution. It is worth noting that the
 355 percentage of total signal intensity of Ox-NOCs with biogenic VOCs (BVOCs) as
 356 precursors was less than 3% (**Fig. 2g–i** and **Table S8**). This can be partly supported by
 357 the previous observations showing that anthropogenic VOCs (AVOCs) were the main
 358 contributors to the formation of Ox-NOCs (e.g., organic nitrates) in urban areas in
 359 China (Wang et al., 2021b; Jiang et al., 2023). The overall results suggest the
 360 significant role of AVOCs in the formation of NOCs in all investigated cities,
 361 particularly in HEB.



362
 363 **Figure 2.** Average percentage distributions of signal intensities for (a–c) CHON+,
 364 (d–f) CHN+, and (g–i) CHON– compounds from various sources in PM_{2.5} collected
 365 from different cities during winter. Average signal intensity distributions for (j)



366 CHON⁺, (k) CHN⁺, and (l) CHON⁻ compounds from various sources in PM_{2.5}
367 collected from different cities during winter.

368

369 From clean to haze periods, the signal intensities of all aromatics-derived CHON
370 compounds increased significantly in HEB (**Figs. 2a, j, g, l and S7**). In contrast, the
371 signal intensities of aromatics-derived CHON compounds in BJ and HZ showed an
372 insignificant increase during haze periods. In addition, the average values of O/C_w and
373 OS_{Cw} for CHON⁺ and CHON⁻ compounds were higher in HEB than in BJ (second
374 highest) and HZ, and their increases from clean to haze periods were also greater in
375 HEB (**Table S10**). This indicates that aerosol NOCs in HEB were more oxidized
376 aromatics, particularly during haze. The average N/C_w ratios of CHON⁺ and CHON⁻
377 compounds in HEB (0.13 and 0.15, respectively) (**Table S10**) were higher than those
378 of CHON⁺ (0.079) and CHON⁻ (0.07) compounds in aerosols directly emitted from
379 coal combustion (Song et al., 2022; Song et al., 2018). The N/C_w ratios were also
380 higher in HEB than in BJ and HZ and increased during hazy days (0.13 for CHON⁺
381 and 0.16 for CHON⁻ in hazy days in HEB). It has been suggested that the N/C_w ratio
382 of CHON⁻ compounds tended to increase (from 0.109 to 0.119) after aging of fuel
383 combustion-derived aerosols (Zhao et al., 2022a). Thus, these results, combined with
384 previous analysis of potential precursors for NOCs, suggest that anthropogenic
385 precursor emissions and their atmospheric transformation to form CHON compounds



386 were stronger in HEB than in BJ and HZ. Moreover, considering that the emission
387 intensity of precursors during clean and hazy days may not significantly change,
388 secondary processes may significantly promote the formation of NOCs in HEB during
389 hazy days (the most significant increase in NOC abundance). However, this
390 promoting effect during hazy days was insignificant in BJ and HZ (less increase in
391 NOC abundance).

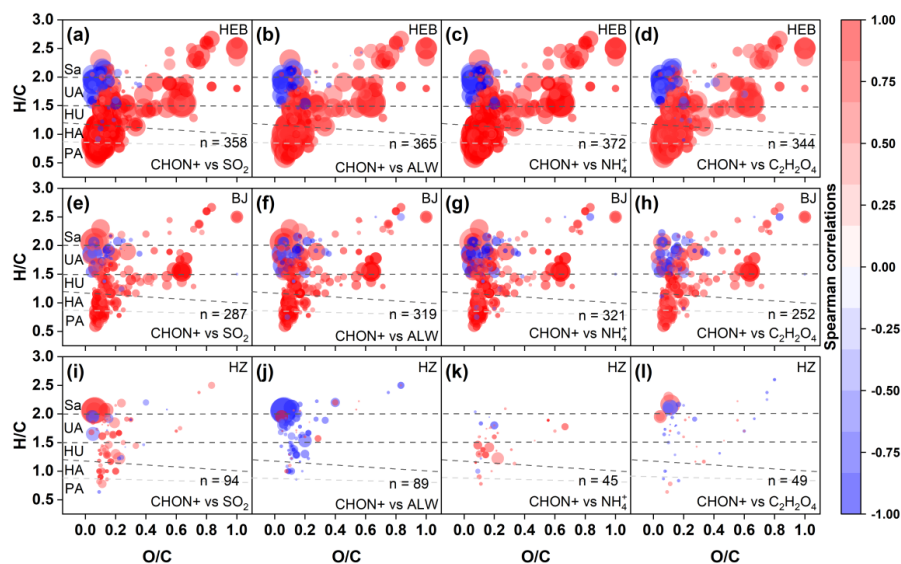
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393 **3.3. Main factors influencing aerosol NOC formation in different cities**

394 The Spearman correlation analysis between various parameters and NOCs was
395 conducted to investigate the key factors influencing NOC molecular distributions (**Fig.**
396 **3** and **Figs. S8–S12**). The peak intensity of most CHON⁺ compounds (mainly
397 aromatics, as mentioned previously) showed a strong correlation ($P < 0.01$) with the
398 concentrations of combustion source-related tracers (e.g., SO₂, nss-Cl⁻, nss-K⁺, CO,
399 and NO₂) (Zhao and Sun, 1986; Streets and Waldhoff, 2000; Shen et al., 2009; Zhang
400 et al., 2011; Mafusire et al., 2016; Liu et al., 2019; Zhang et al., 2021a; Wang et al.,
401 2020) in HEB (**Figs. 3a** and **S8a–d**). Although there was a significant correlation ($P <$
402 0.05) between most CHON⁺ compounds and those combustion source indicators in
403 BJ, the strength of this correlation was weaker in BJ than in HEB (**Figs. 3e** and **S8f–i**).
404 However, similar significant correlations between them were not observed in HZ
405 (**Figs. 3i** and **S8k–n**). Thus, the greatest contribution of anthropogenic activities to the



406 formation of CHON+ compounds in winter was in HEB (central heating with coal),
 407 followed by BJ (central heating with coal and natural gas) and HZ (without central
 408 heating). Most of CHN+ and CHON- compounds showed a similar spatial response
 409 pattern to those anthropogenic activities (**Figs. S9** and **S10**). These results are
 410 consistent with the previous analysis of NOC precursors (**Fig. 2**), which concluded
 411 that the intensity of anthropogenic pollutant emissions in different energy consuming
 412 cities was an important factor affecting the formation of NOC and causing spatial
 413 differences in NOC abundance.



414
 415 **Figure 3.** Spearman rank correlation coefficients (with $P < 0.01$ in HEB and $P < 0.05$
 416 in BJ and HZ) of individual CHON+ molecules with selected parameters in (a–d)
 417 HEB, (e–h) BJ and (i–l) HZ. The color scale indicates Spearman correlations between
 418 the intensity of individual CHON+ molecules and each parameter. The symbol n in



419 the bottom right corner of each panel indicates the number of molecular formulas
420 significantly correlated with the variables. The subgroups in the panels include
421 polycyclic aromatic-like (PA), highly aromatic-like (HA), highly unsaturated-like
422 (HU), unsaturated aliphatic-like (UA), and saturated-like (Sa) compounds.

423

424 Furthermore, we found that the peak intensities of most CHON⁺, CHN⁺, and
425 CHON⁻ compounds (mainly aromatics) were significantly correlated ($P < 0.01$) with
426 the concentrations of ALW, NH₄⁺, oxalic acid, and SO₄²⁻ (Figs. 3b–d, S8e, and
427 S11–S12) in HEB. The correlations between these NOCs and parameters weakened in
428 BJ and disappeared in HZ (Figs. 3, S8, and S11–12). It is generally accepted that
429 SO₄²⁻, NH₄⁺, and NO₃⁻ in fine aerosols are primarily formed through secondary
430 processes (Gao et al., 2021; Wang et al., 2021d). NH₄⁺ can serve as a key reactant in
431 the formation of aerosol NOCs (e.g., "carbonyl-to-imine" transformation) in the
432 aqueous-phase (Laskin et al., 2014; Lee et al., 2013; Li et al., 2019b). Previous studies
433 have identified oxalic acid (C₂H₂O₄) as a tracer for aqueous-phase SOA (Xu et al.,
434 2022a; Carlton and Turpin, 2013). Additionally, numerous laboratory and field
435 observational studies have shown that ALW can promote the formation of NOCs (Lv
436 et al., 2022; Liu et al., 2023b; Jimenez et al., 2022; Jiang et al., 2023). Thus, these
437 results indicate that aqueous-phase processes can significantly promote the formation
438 of NOCs in HEB, however, as the precursor emission intensity gradually decreased in

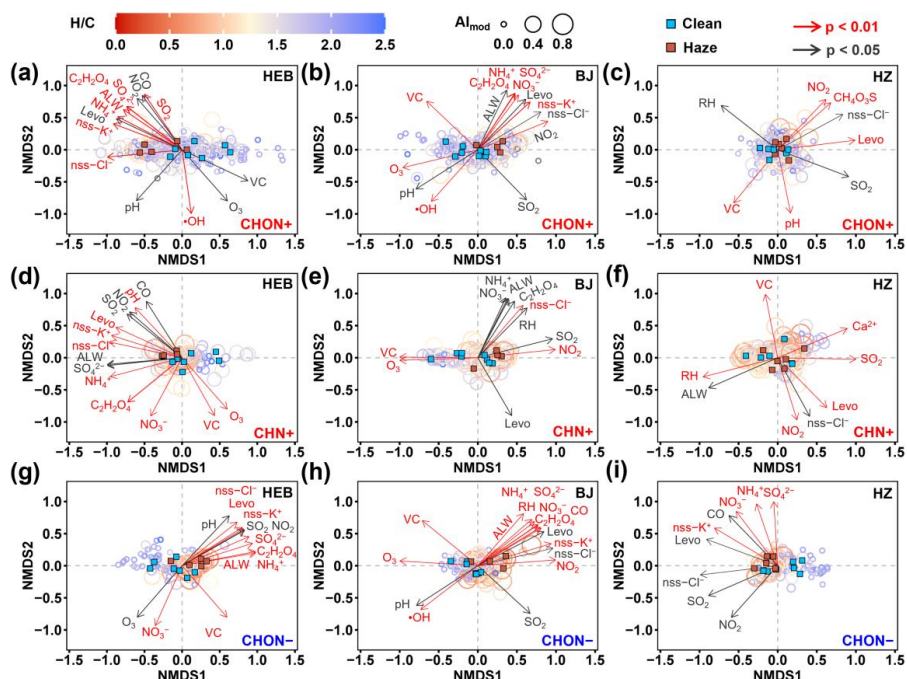


439 BJ and HZ, this aqueous-phase promoting effect also decreased.

440 The NMDS analysis between various parameters and NOCs was conducted to
441 further investigate the variations in key factors affecting the formation of NOCs from
442 clean to haze days (**Fig. 4**). The formation of CHON⁺, CHN⁺, and CHON⁻
443 compounds with higher AI_{mod} values (mainly aromatics, as mentioned previously)
444 during haze days in HEB and BJ were closely associated with the factors indicating
445 anthropogenic precursor emissions and aqueous-phase reaction processes. In contrast,
446 the level of oxidants (i.e., O₃ and •OH) played a more important role during clean
447 days in HEB and BJ, driving more highly saturated NOC formation during clean days
448 (**Fig. 4**). A reasonable explanation for this is that the solar radiation and •OH levels on
449 polluted days were lower than those on clean days (**Table S1**). The impacts of various
450 factors on the formation of aerosol NOCs showed a weak discrimination between
451 haze and clean days in HZ (**Fig. 4c, f and i**). Laboratory studies have shown that
452 reactive components (e.g., •OH and H₂O₂) in the aqueous phase can continuously
453 convert low-solubility organics to form aqueous phase SOA (Chen et al., 2008; Huang
454 et al., 2011). Field observations also suggested that precursors (most of them are
455 aromatic compounds) released from the combustion of fossil fuels significantly
456 contributed to the aqueous SOA formation (> 50% total molecules) (Xu et al., 2022a)
457 through the rapid aqueous-phase conversion of primary organic aerosol (POA) to
458 SOA at high RH (Wang et al., 2021a). This implies that higher precursor abundance



459 can drive more aerosol NOC formation via aqueous-phase processes. As mentioned
460 previously, the emission intensity of precursors decreased sequentially from HEB to
461 BJ and then to HZ. Moreover, the ALW concentrations were much higher on polluted
462 days than on clean days in three investigated cities. The rising ALW during the
463 pollution period and the quiescent steady state of the atmosphere favored the
464 formation of SOA from anthropogenic emission precursors (Guo et al., 2014; He et al.,
465 2018). Thus, the above discussion can suggest that the spatial differences in precursor
466 emission intensity (higher in HEB) and enhancement of aqueous-phase processes in
467 polluted days were the main factors leading to the differences in the proportion
468 (higher in HEB) of increase in NOC abundance from clean days to polluted days in
469 three different energy consuming cities. In addition, the increased VC value (**Table**
470 **S1**) in clean days (beneficial for the diffusion of pollutants) (Gani et al., 2019) was
471 also an important factor limiting the abundance of NOCs (**Fig. 4**), resulting in a lower
472 NOC abundance on clean days compared to polluted days (**Fig. 1**).



473

474 **Figure 4.** Nonmetric multidimensional scaling of (a–c) CHON+, (d–f) CHN+, and

475 (g–i) CHON– compounds from organic aerosol in different cities. The color and size

476 of the circle indicate the H/C ratio and Al_{mod} value of compounds, respectively.

477 Significant relationships between the variables and ordination (999 permutations) are

478 indicated by $p < 0.05$ (grey) and $p < 0.01$ (red). Insignificant correlations are not

479 shown. The scores of the samples collected during clean and haze periods were shown

480 as blue and brown squares, respectively.

481

482 As mentioned above, the aerosol NOCs of HZ were less affected by

483 anthropogenic pollutants emitted from coal and natural gas combustion compared to



484 HEB and BJ with centralized heating. Interestingly, we found that the molecular
485 distributions of most aromatic CHON⁺ compounds in HZ were not only influenced
486 by some anthropogenic pollutants (e.g., SO₂ and NO₂), but also by methanesulfonic
487 acid (CH₃SO₃H) (**Fig. 4c**). Moreover, neither CHN⁺ nor CHON⁺ exhibited significant
488 correlations with factors related to secondary processes in HZ (**Fig. 4c** and **f**).
489 Methanesulfonic acid has been suggested to be a tracer for ocean aerosols (Ayers and
490 Gras, 1991; Suess et al., 2019). These results suggest that aerosol CHON⁺
491 compounds in HZ may be influenced by long-distance transport air masses originating
492 from the ocean. This consideration can be also supported by the fact that only HZ was
493 affected by air masses originated from the ocean (**Fig. S13**). Thus, marine emissions
494 may be an important contributor to aerosol NOCs in HZ, which was significantly
495 different from the cases of HEB and BJ where aromatic pollutants from fossil fuel
496 combustion and aqueous-phase processes control the composition and abundance of
497 aerosol NOCs.

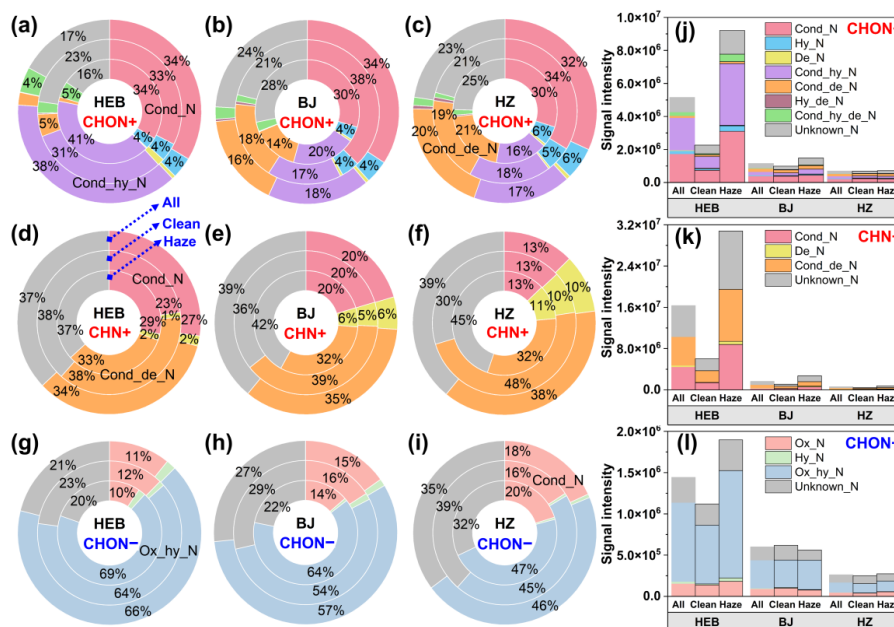
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499 **3.4. Potential formation mechanisms of aerosol NOCs in cities with different** 500 **energy consumption**

501 **Figure 5** shows the average signal intensity percentage and signal intensity
502 distributions of NOCs formed by different aqueous-phase processes (**Table S4** and
503 **Figs. S3–S5**) in different cities during winter. The identification of specific reaction



504 pathways was detailed in **Figs. S3–S5** and **Sect. S3**. During the entire study period,
505 the cond_N, cond_{hy}_N, and cond_{de}_N pathways together accounted for more than
506 68% (68%–74%) of the total signal intensity of CHON⁺ compounds in the three cities
507 (**Fig. 5a–c** and **Table S11**). Specifically, the formation of CHON⁺ compounds was
508 mainly dominated by the cond_N and cond_{hy}_N pathways in HEB, with less impact
509 from the cond_{de}_N pathway (**Fig. 5a**). However, CHON⁺ compounds derived from
510 the cond_{de}_N pathway showed a much higher proportion in BJ and HZ than in HEB
511 (**Fig. 5b** and **c**). The cond_{de}_N pathway involves both condensation and dehydration
512 processes (**Table S4** and **Fig. S3**). It has been suggested that higher temperatures can
513 facilitate the dehydration of amides into nitriles (Mekki-Berrada et al., 2013). The
514 temperatures in BJ and HZ were higher than those in HEB (**Table S1**), which may
515 partly explain the higher signal proportion of CHON⁺ compounds formed through the
516 cond_{hy}_N pathway in BJ and HZ than in HEB. Furthermore, the higher signal
517 proportions of CHN⁺ compounds formed through the de_N pathway in BJ (6%) and
518 HZ (11%) than in HEB (2%) may also be associated with this temperature-induced
519 dehydration mechanism (**Fig. 5d–f** and **Table S12**). For CHN⁺ compounds, the
520 cond_{de}_N process dominated their formation (**Fig. 5d–f**). In general, the cond_N,
521 cond_{hy}_N, and cond_{de}_N processes contributed most significantly to the formation
522 of Re-NOCs in HEB, followed BJ and HZ.



523

524 **Figure 5.** Average percentage distributions of signal intensities for aerosol (a–c)

525 CHON+, (d–f) CHN+, and (g–i) CHON– compounds from various reaction pathways

526 in different cities during winter. Average signal intensity distributions for aerosol (j)

527 CHON+, (k) CHN+, and (l) CHON– compounds from various reaction pathways in

528 different cities during winter.

529

530 A typical mechanism for Re-NOC formation is the aqueous-phase reactions

531 between carbonyl compounds and NH_4^+ (or NH_3) (Abudumutailifu et al., 2024;

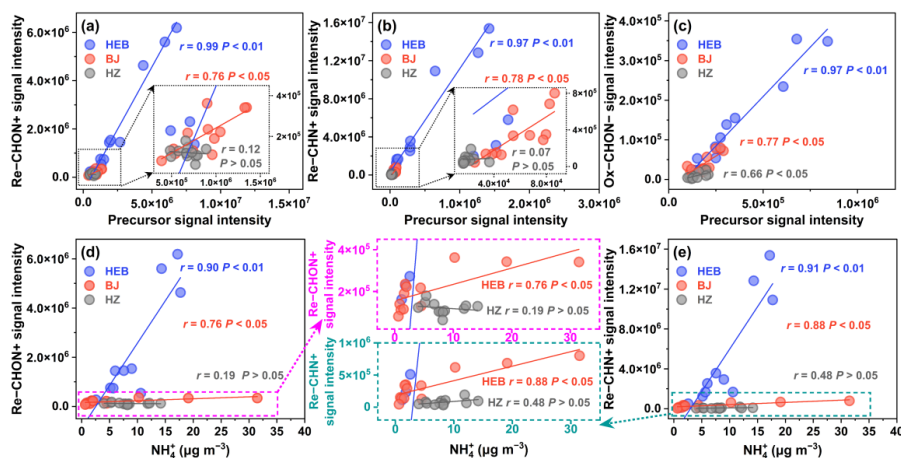
532 Laskin et al., 2014; Li et al., 2019b; Liu et al., 2023b; Wang et al., 2024). If this

533 mechanism is simplified as a second-order reaction (i.e., $[\text{Precursor}] + [\text{NH}_3 \text{ and } \text{NH}_4^+]$

534 \leftrightarrow [Re-NOCs]), the production of Re-NOCs is expected to be proportional to the



535 abundances of precursor and NH_4^+ (Yang et al., 2023; Lin et al., 2023). Indeed, the
 536 signal intensities of the Re-CHON+ and Re-CHN+ compounds were significantly
 537 positively correlated with the signal intensities of their CHO precursors (identified
 538 using the precursor-product pairs theory, **Figs. S3 and S4**) and NH_4^+ concentration in
 539 HEB (**Fig. 6a, b, d and e**). This correlation gradually weakened from BJ to HZ (**Fig.**
 540 **6a, b, d and e**). As previously discussed, differences in energy consumption patterns
 541 resulted in the highest levels of anthropogenic aromatic compound emissions in HEB
 542 during the winter, followed by BJ, with the lowest levels in HZ (**Figs. 2 and S14**).
 543 Thus, the signal intensities of CHON+ and CHN+ compounds from cond_N,
 544 cond_de_N, and cond_hy_N processes were higher in HEB than in BJ and lowest in
 545 HZ (**Fig. 5j and k**).



546 **Figure 6.** Signal intensity of (a) Re-CHON+, (b) Re-CHN+, and (c) Ox-CHON-
 547 compounds as functions of signal intensity of precursors (CHO compounds). Signal
 548 intensity of (d) Re-CHON+ and (e) Re-CHN+ compounds as functions of the
 549 intensity of NH_4^+ concentration.



550 concentrations of NH_4^+ .

551

552 Additionally, we noticed that the contribution of these aqueous-phase processes to
553 the formation of CHON^+ and CHN^+ compounds increased significantly from clean to
554 hazy days in HEB and BJ (**Fig. 5**). The increased ALW concentrations (**Table S1**) and
555 atmospheric stability during haze periods likely provided favorable conditions for the
556 precursors to undergo these aqueous-phase reactions, resulting in the formation of
557 NOCs. Clearly, high pollutant emission levels in HEB provided a greater potential to
558 convert precursors into more NOCs via the cond_N , $\text{cond}_\text{hy}_\text{N}$, and $\text{cond}_\text{de}_\text{N}$
559 processes during haze periods. Thus, the hazy days in the HEB showed the largest
560 increase in CHON^+ and CHN^+ compounds from the cond_N , $\text{cond}_\text{hy}_\text{N}$, and
561 $\text{cond}_\text{de}_\text{N}$ processes (**Fig. 5j** and **k**). In contrast, the lower precursor emissions in
562 HZ without centralized heating policy were not sufficient to support the production of
563 large amounts of NOCs in the aqueous phase. These results also indicate that emission
564 reduction is the key to controlling aerosol NOC pollution.

565 CHON^- compounds derived from the $\text{ox}_\text{hy}_\text{N}$ and ox_N processes together
566 accounted for more than 64% (64%–71%) of the total signal intensity of CHON^-
567 compounds in the three cities (**Fig. 5g–i, l** and **Table S13**). The signal intensity
568 proportions of CHON^- compounds formed by the $\text{ox}_\text{hy}_\text{N}$ process in these three
569 cities (ranging from 47% in HZ to 69% in HEB) were higher than that in Wangdu (<



570 20%) (Jiang et al., 2023). The observation study in Wangdu examined aerosol organic
571 components only in ESI⁻ mode (Jiang et al., 2023), which may underestimate the
572 importance of the CHO⁺ compounds that could serve as precursors of Ox-NOCs. In
573 general, CHON⁻ compounds formed through the ox_{hy}N and ox_N processes
574 showed the highest abundance in HEB, followed by BJ and HZ (**Fig. 5j-i**). According
575 to a simplified reaction mechanism for the formation of Ox-NOCs via aqueous-phase
576 processes (i.e., [Precursor] + [Oxidants] ↔ [Ox-NOCs]) (Shi et al., 2023; Kroflič et
577 al., 2015; Vione et al., 2005), we can infer that Ox-NOCs production is proportional
578 to precursor levels when oxidants (e.g., NO₂ radical or NO₂⁺) are in a steady state in
579 the atmosphere. Indeed, the signal intensities of the Ox-CHON⁻ compounds were
580 significantly positively correlated with the signal intensities of their CHO precursors
581 identified using the precursor-product pairs theory in HEB (**Fig. 6c**). Moreover, this
582 correlation gradually weakened from BJ to HZ (**Fig. 6c**). Thus, the spatial differences
583 in the contribution of the ox_{hy}N and ox_N processes to Ox-NOC production across
584 the three cities can also be explained by differences in precursor emission intensity, as
585 indicated by above mentioned Re-NOC formation.

586

587 **4. Conclusion**

588 The abundance, composition, potential precursors, and potential formation
589 mechanisms of NOCs in PM_{2.5} in three Chinese cities with different energy

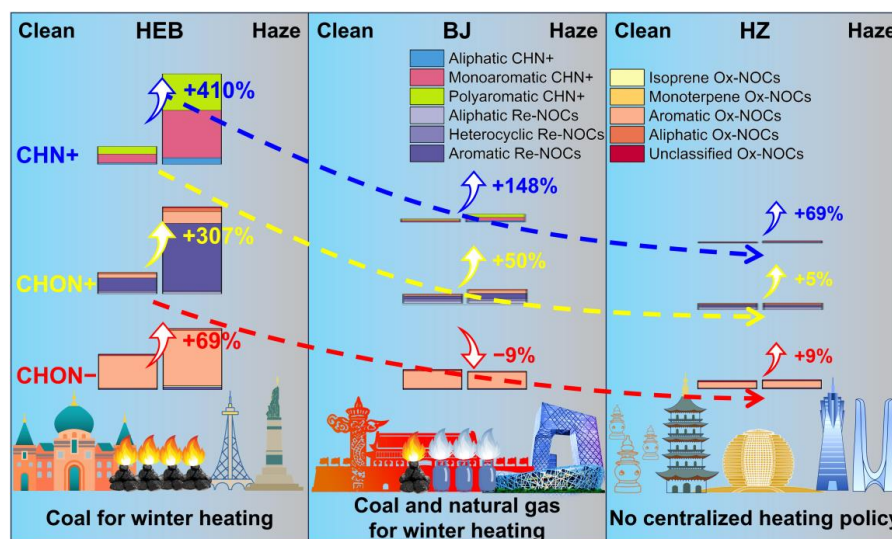


590 consumption types during the winter were systematically investigated. On average,
591 the total signal intensity of NOCs (i.e., CHN⁺, CHON⁺, and CHON⁻ compounds)
592 was highest in HEB, followed by BJ. The lowest total NOC signal intensity was found
593 in HZ. According to the identification of potential precursors of NOCs, we found that
594 anthropogenic aromatic compounds were the main precursors of NOCs during winter
595 in HEB where mainly relies on coal for winter heating, with less impact from BVOCs.
596 Anthropogenic aromatic precursors were also identified to be important contributors
597 to NOC formation in BJ which uses natural gas and coal for winter heating, although
598 the contribution ratio was lower in BJ than in HEB. In contrast, the lowest aromatic
599 precursor levels occurred in HZ without winter heating policy. Furthermore, the
600 NMDS analysis supported that the impact of anthropogenic fossil fuel combustion on
601 NOC pollution gradually decreased from HEB to BJ and then to HZ.

602 The formation of CHON⁺ compounds was mainly associated with the cond_N,
603 cond_{hy}_N, and cond_{de}_N processes. The cond_N and cond_{de}_N processes
604 dominated the formation of CHN⁺ compounds. The production of CHON⁺ and
605 CHN⁺ compounds from the cond_N, cond_{hy}_N, and cond_{de}_N processes was
606 highest in HEB, followed by BJ and HZ. The ox_{hy}_N and ox_N processes
607 contributed significantly to CHON⁻ compound formation, from which the highest
608 CHON⁻ production occurred in HEB and the lowest in HZ. The spatial differences in
609 the contribution of different aqueous-phase processes to NOC production in the three



610 different cities can be attributed to differences in precursor emission intensity. In
 611 particular, the contribution of these aqueous-phase processes to the formation of
 612 CHON+ and CHN+ compounds showed the most significant increase from clean to
 613 hazy days in HEB, followed by BJ. We concluded that high pollutant emission levels
 614 can provide a greater potential to convert precursors to produce more NOCs via
 615 aqueous-phase processes during haze periods. The above findings are summarized in
 616 a diagram (Fig. 7).



617
 618 **Figure 7.** Conceptual illustration showing the characteristics of different NOCs from
 619 the clean days to the haze days in different cities.

620 In general, the aerosol NOCs pollution during winter is closely linked to both the
 621 intensity of precursor emissions and the efficiency of aqueous-phase processes in
 622 converting these emissions into NOCs. The overall results highlight the importance of



623 emission reduction strategies in controlling aerosol NOCs pollution during winter.
624 Targeted reduction of precursor emissions, especially from coal combustion, could
625 significantly mitigate NOCs levels, thereby improving air quality and public health in
626 urban areas. Future research should focus on further elucidating the specific pathways
627 of aqueous-phase NOC formation and developing available models to predict NOC
628 dynamics under varying environmental conditions. Additionally, research into the
629 long-term effects of transitioning to cleaner energy sources on the reduction of NOC
630 pollution will be essential for guiding effective air quality management strategies.

631

632 **Data availability.** The data presented in this work are available upon request from the
633 corresponding authors.

634

635 **Competing interests.** The authors declare no conflicts of interest relevant to this
636 study.

637

638 **Supplement.** Details of parameter calculation, classification method for identifying
639 precursors of NOCs, classification of possible aqueous-phase processes NOCs based
640 on precursor-product pairs, thirteen tables (Tables S1–S13), and fourteen extensive
641 figures (Figures S1–S14) are provided in the Supplement.

642



643 **Author contributions.** YX designed the study. YJM, TY, LG, HX, and HWX
644 performed field measurements and sample collection; YJM performed chemical
645 analysis; YX and YJM performed data analysis; YJM and YX wrote the original
646 manuscript; and YX, YJM, and HX reviewed and edited the manuscript.

647

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652

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