

1 **The Critical Role of Aqueous-Phase Processes in**
2 **Aromatic-Derived Nitrogen-Containing Organic**
3 **Aerosol Formation in Cities with Different**
4 **Energy Consumption Patterns**

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20 **Abstract.** Nitrogen-containing organic compounds (NOCs) impact air quality and
21 human health. Here, the abundance, potential precursors, and main formation
22 mechanisms of NOCs in PM_{2.5} during winter were compared for the first time among
23 Haerbin (coal-dependent for heating), Beijing (natural gas and coal as heating energy),
24 and Hangzhou (no centralized heating policy). The total signal intensity of CHON⁺,
25 CHN⁺, and CHON⁻ compounds was highest in Haerbin and lowest in Hangzhou.
26 Anthropogenic aromatics accounted for 73%–93% of all identified precursors of
27 CHON⁺, CHN⁺, and CHON⁻ compounds in Haerbin. Although the abundance of
28 aromatics-derived NOCs was lower in Beijing than in Haerbin, aromatics were also
29 the main contributors to NOC formation in Beijing. Hangzhou exhibited the lowest
30 levels of aromatic precursors. Furthermore, non-metric multidimensional scaling
31 analysis indicated an overall reduction in the impact of fossil fuel combustion on
32 NOC pollution along the route from Haerbin to Beijing to Hangzhou. We found that
33 aqueous-phase processes (mainly condensation, hydrolysis or dehydration processes
34 for reduced NOCs, and mainly oxidization or hydrolysis processes for oxidized NOCs)
35 can promote the transformation of precursors to produce NOCs, leading to the most
36 significant increase in aromatic NOC levels in Haerbin (particularly on haze days).
37 Reduced precursor emissions in Beijing and Hangzhou (the lowest) constrained the
38 aqueous-phase formation of NOCs. The overall results suggest that the aerosol NOC
39 pollution in coal-dependent cities is mainly controlled by anthropogenic aromatics

40 and aqueous-phase processes. Thus, without effective emission controls, the
41 formation of NOCs through aqueous-phase processes may still pose a large threat to
42 air quality.

43

44 **Keywords:** Aerosols, Nitrogen-containing organic compounds, Energy structure,
45 Anthropogenic pollutants, Formation mechanisms

46

47 **1. Introduction**

48 Nitrogen-containing organic compounds (NOCs) are abundant reactive nitrogen
49 species in aerosol particles, accounting for up to 40%–80% of total nitrogen
50 deposition (Li et al., 2023; Xi et al., 2023; Yu et al., 2020). Clearly, aerosol NOCs can
51 significantly contribute to the global nitrogen cycle (Li et al., 2023; Cape et al., 2011).
52 Moreover, the formation of secondary organic aerosols (SOA) and light-absorbing
53 organic aerosols (e.g., brown carbon) is also tightly associated with NOCs (Wang et
54 al., 2024; Liu et al., 2023b; Zeng et al., 2021), thus affecting the radiative balance and
55 air quality (Yuan et al., 2023; Jiang et al., 2023). In particular, certain NOCs, such as
56 nitroaromatics and peroxyacyl nitrates, are characterized as phytotoxins and potential
57 carcinogens, posing threats to ecosystems and human health (Shi et al., 2023; Singh
58 and Kumar, 2022; Huang et al., 2024). Therefore, understanding the characteristics,
59 origins, and atmospheric processes of NOCs is essential for comprehending their
60 climate and health effects.

61 Aerosol NOCs can be derived from primary emissions associated with
62 anthropogenic activities and natural sources (Lin et al., 2023; Xu et al., 2020a; Wang
63 et al., 2017; Song et al., 2018; Song et al., 2022; Ma et al., 2024; Gui et al., 2024; Xu
64 et al., 2024a). Secondary formation processes may play a more crucial role in the
65 formation of NOCs in fine aerosol particles, which involve interactions among
66 volatile organic compounds (VOCs), atmospheric oxidants, and reactive inorganic

67 nitrogen species (Montoya-Aguilera et al., 2018; Perraud et al., 2012; Hallquist et al.,
68 2009). For instance, laboratory studies have observed the formation of organic nitrates
69 from the oxidation of isoprene and α -/ β -pinene by atmospheric oxidants and **nitrogen**
70 **oxides** (NO_x) (Surratt et al., 2010; Rollins et al., 2012; Nguyen et al., 2015).
71 Additionally, aqueous-phase reactions of NH_4^+ (or NH_3) with biogenic VOCs or
72 carbonyl compounds have been suggested to be important mechanisms of reduced
73 NOC (Re-NOCs) formation (Abudumutailifu et al., 2024; Laskin et al., 2014; Li et al.,
74 2019b; Liu et al., 2023b; Wang et al., 2024). However, understanding the origins,
75 formation mechanisms, and environmental impacts of NOCs is hindered by the
76 elusive and intractable molecular information regarding NOCs and their precursors.

77 Aerosol liquid water (ALW) can greatly increase the formation of aerosol NOCs
78 by facilitating the conversion of water-soluble organic gases into particles and
79 subsequently enabling aqueous-phase reactions (Li et al., 2019a; Lv et al., 2022; Liu
80 et al., 2023b). **Several observational studies have found a positive correlation between**
81 **aerosol NOC abundance and either ALW or relative humidity (RH)** (Jiang et al., 2023;
82 Liu et al., 2023b; Xu et al., 2020b). In particular, it has been suggested that increased
83 ALW levels can exacerbate winter haze in China (Wu et al., 2018; Hodas et al., 2014;
84 Lv et al., 2022; Wang et al., 2021d; Liu et al., 2023b; Wang et al., 2021a; Li et al.,
85 2019a). Presumably, precursors and ALW are the two key factors in the formation of
86 aerosol NOCs. Haze environments have potentially high RH levels and large

87 emissions of NOC precursors (Zheng et al., 2023; Nie et al., 2022; Liu et al., 2021;
88 Wang et al., 2021a). Moreover, in Chinese cities with different energy consumption
89 (e.g., coal, biomass, and natural gas) for winter heating (Zhang et al., 2021b; Zhang et
90 al., 2023b; Yang et al., 2024c), the types and emission intensities of pollutants
91 released from different heating sources are expected to vary considerably (Bond et al.,
92 2006; Stockwell et al., 2015; Krůmal et al., 2019). However, the potential effects of
93 ALW in the formation of NOCs in Chinese cities with different energy consumption
94 during winter, particularly in haze periods, are not well documented. Moreover, the
95 roles of ALW-related NOC formation processes in the formation of haze in cities with
96 different energy consumption types also remain largely unknown.

97 In this study, we present the measurements of the NOCs and other chemical
98 compositions in PM_{2.5} collected from three cities (Haerbin, Beijing, and Hangzhou)
99 with different energy consumption during winter. The specific objectives of this study
100 were: (1) to investigate the differences in the abundance, composition, and major
101 precursors of NOCs in different cities with different energy consumption, especially
102 on polluted days; and (2) to elucidate the potential effects of aqueous-phase processes
103 on the formation of oxidized NOCs (Ox-NOCs) and reduced NOCs (Re-NOCs)
104 during winter (particularly on polluted days) in cities with different energy
105 consumption. The research findings are expected to provide valuable implications for
106 the mitigation of aerosol NOCs pollution in urban environments.

107

108 **2. Materials and methods**

109 **2.1. Study site description and sample collection**

110 The study sites are located in three urban areas, including Haerbin (HEB, i.e.,
111 Harbin, 126.64°E, 45.77°N), Beijing (BJ, 116.41°E, 40.04°N), and Hangzhou (HZ,
112 120.16°E, 30.30°N) (**Fig. S1**). The city of HEB, with a population of 9.95 million, is
113 situated in the northeastern region of China. It relies heavily on coal for centralized
114 heating during winter. The rapid urbanization and increased coal consumption have
115 significantly deteriorated air quality in HEB in recent years (Ma et al., 2020). In
116 contrast, BJ has largely shifted towards the utilization of cleaner energy sources (e.g.,
117 natural gas) for centralized heating in recent years, particularly following the
118 implementation of the "Beijing 2013–2017 Clean Air Action Plan" (Vu et al., 2019;
119 Yuan et al., 2023). HZ, situated within the Yangtze River Delta, is exempt from the
120 necessity of heating due to the relatively mild winter climate (average temperature of
121 6.6 ± 2.4 °C during the sampling period, **Table S1**). Clearly, the distinctive energy
122 consumption patterns observed in these three cities during winter provide a valuable
123 opportunity to examine the impact of various precursors and aqueous-phase processes
124 on aerosol NOC formation.

125 Sample collection was carried out simultaneously in three cities from 16
126 December 2017 to 14 January 2018. PM_{2.5} samples were collected every two or three

127 days with a duration of 24 hours onto prebaked quartz fiber filters (Pallflex, Pall
128 Corporation, USA) using a high-volume air sampler (Series 2031, Laoying, China).
129 One blank sample was collected at each sampling site. A total of 39 samples were
130 collected, all of which were stored at -30°C . Meteorological data (e.g., temperature,
131 relative humidity (RH) and wind speed) together with concentrations of various
132 pollutants (e.g., SO_2 and NO_2) were obtained from nearby environmental stations. In
133 China, according to the Air Quality Index (MEEPRC, 2012), a pollution day is
134 defined as a day with a 24-hour average $\text{PM}_{2.5}$ concentration above $75 \mu\text{g m}^{-3}$. This
135 standard has also been used in other studies performed in China (Zhang and Cao,
136 2015; Xu et al., 2024b; Yan et al., 2024), showing that the sampling periods were
137 classified as either "clean" or "haze" based on whether the daily average concentration
138 of $\text{PM}_{2.5}$ was below or above $75 \mu\text{g m}^{-3}$.

139

140 **2.2. Chemical analysis**

141 The extraction and analysis methods for NOCs were consistent with those
142 described in our recent publication (Ma et al., 2024). Briefly, a portion of each filter
143 was extracted with methanol (LC-MS grade, CNW Technologies Ltd.) using
144 sonication in an ice bath ($\sim 4^{\circ}\text{C}$). The extracts were filtered through a $0.22 \mu\text{m}$
145 polytetrafluoroethylene syringe filter and then concentrated to $300 \mu\text{L}$ under a gentle
146 stream of nitrogen gas. The concentrated extracts underwent composition analysis via

147 an ultra-performance liquid chromatography coupled with quadrupole time-of-flight
148 mass spectrometry equipped with an electrospray ionization (ESI) source (UPLC-ESI-
149 QToFMS, Waters Acquity Xevo G2-XS) (Wang et al., 2021c; Ma et al., 2024). This
150 analysis was done in both ESI+ and ESI- modes. The organic compounds were
151 separated on an Acquity HSS T3 column (2.1 × 100 mm, 1.8 μm particle size, Waters)
152 with an 18-minute gradient elution. The mobile phases comprised ultrapure water
153 with 0.1% formic acid (A) and methanol with 0.1% formic acid (B). Gradient elution
154 was conducted according to the following protocol: 1% B was held for 1.5 minutes,
155 followed by an increase to 54% B over a period of 6.5 minutes. Thereafter, the B was
156 increased to 95% over a period of 3 minutes. After reaching 100% B in one minute,
157 this state was maintained for 3 minutes. Finally, the concentration was returned to 1%
158 B in 0.5 minutes and held for 2.5 minutes. **More detailed information about the**
159 **UPLC-ESI-QToFMS analysis can be found in Sect. S1.** Due to uncertainties in
160 ionization efficiencies for different compounds (Ditto et al., 2022; Yang et al., 2023),
161 an intercomparison (mainly compared among samples within this study) of compound
162 relative abundance was conducted without accounting for differences in ionization
163 efficiency in the present study. This consideration was consistent with previous
164 studies (Xu et al., 2023; Jiang et al., 2022; Ma et al., 2024).

165 Another filter portion was ultrasonically extracted using Milli-Q water (~4°C ice
166 bath) to analyze the concentrations of inorganic ions and organic acids. The inorganic

167 ions (e.g., NO_3^- , SO_4^{2-} , Cl^- , Ca^{2+} , Mg^{2+} , Na^+ , and NH_4^+) and organic acids (e.g.,
168 formic acid, acetic acid, oxalic acid, succinic acid, glutaric acid, and methanesulfonic
169 acid) were quantified using an ion chromatograph system (Dionex Aquion, Thermo
170 Scientific, USA), as described previously (Xu et al., 2022b; Yang et al., 2024b).

171

172 **2.3. Compound categorization and precursor identification**

173 The identified molecular formulas via UPLC-ESI-QToFMS were categorized
174 into different compound classes based on their elemental compositions, which
175 included CHO^- , CHON^- , CHONS^- , and CHOS^- in ESI⁻ mode and CHO^+ , CHON^+ ,
176 and CHN^+ in ESI⁺ mode (Ma et al., 2024). **Unless otherwise indicated, the molecular**
177 **formulas presented in the manuscript refer to neutral molecules. The "-" and "+"**
178 **symbols denote the detection ion modes, which correspond to ESI⁻ and ESI⁺ modes,**
179 **respectively.** Here, we mainly focus on NOCs (i.e., CHN^+ , CHON^+ , and CHON^-
180 compounds) (Ma et al., 2024; Jiang et al., 2022; Wang et al., 2017). The carbon
181 oxidation state (OS_C) and double bond equivalent (DBE) were calculated to indicate
182 the oxidation level and unsaturation degree of the organics, respectively (**Sect. S2**)
183 (Kroll et al., 2011; Ma et al., 2024). Additionally, the modified aromaticity index
184 (AI_{mod}) and aromaticity equivalent (X_C) were used to evaluate aromaticity of organics
185 (Koch and Dittmar, 2006), as detailed in **Sect. S2**.

186 The potential precursors of NOCs were identified based on the methodology

187 described in previous studies (Nie et al., 2022; Guo et al., 2022; Jiang et al., 2023).
188 The classification of CHON⁺ and CHON⁻ compounds was refined into following
189 categories, including aliphatics-, heterocyclics-, and aromatics-derived Re-NOCs
190 and isoprene-, monoterpenes-, aliphatics-, and aromatics-derived Ox-NOCs.
191 Moreover, CHN⁺ compounds were classified into aliphatic, monoaromatic, and
192 polyaromatic CHN⁺ compound (Wang et al., 2021b; Yassine et al., 2014). A detailed
193 description of the revised workflow for classifying NOCs according to potential
194 precursors was provided in **Sect. S3** and **Fig. S2**.

195

196 **2.4. Classification of potential pathways for NOC formation**

197 To identify potential aqueous-phase processes for aerosol NOC formation, we
198 screened precursor-product pairs from the organic compounds that have been detected
199 (Su et al., 2021; Xu et al., 2023; Jiang et al., 2023). The reaction pathways of Re-
200 NOCs (mainly CHON⁺ compounds in this study) were refined into the following
201 categories, including condensation (cond_N), hydrolysis (hy_N), dehydration (de_N),
202 cond_hy_N (involving cond_N and hy_N), cond_de_N (involving cond_N and de_N),
203 hy_de_N (involving hy_N and de_N), cond_hy_de_N (involving cond_N, hy_N and
204 de_N) and unknown_N (unknown processes) formation pathways (**Fig. S3 and Table**
205 **S4**) (Sun et al., 2024; Abudumutailifu et al., 2024; Laskin et al., 2014; Liu et al.,
206 2023c). Another significant class of Re-NOCs is the CHN⁺ compounds. Their

207 potential formation mechanisms include cond_N, de_N, cond_de_N, and other
208 unidentified (unknown_N) pathways (**Fig. S4** and **Table S4**) (Abudumutailifu et al.,
209 2024; Laskin et al., 2014; Liu et al., 2023c). In addition, the reaction pathways of Ox-
210 NOCs (mainly CHON⁻ compounds in this study) were refined into the following
211 categories, including ox_N, hy_N, ox_hy_N (involving ox_N and hy_N), and other
212 unidentified (unknown_N) pathways (Jiang et al., 2023; Su et al., 2021) (**Fig. S5** and
213 **Table S4**). A detailed overview of the methodologies employed to discern potential
214 NOC formation pathways was shown in **Sect. S4**, **Table S4**, and **Figs. S3–S5**.

215 It is important to acknowledge the potential limitations in the categorization
216 methodology of NOC formation pathways described above. This is because the
217 approach applied here and in previous studies (Jiang et al., 2023; Su et al., 2021) may
218 classify NOCs as products of aqueous-phase reactions from primary emissions.
219 Accordingly, our results can be regarded as a maximal potential (or an upper limit) for
220 NOC generation from aqueous-phase reactions. In particular, certain reaction
221 pathways (e.g., oligomerization) were not included in this study due to the complexity
222 of the atomic changes involved, which could not be effectively characterized using the
223 "precursor-product pairs" approach. In this study, NOCs produced from the reaction
224 pathways identified by the abovementioned classification methodology can explain 76%
225 of CHON⁺ compounds, 61% of CHN⁺ compounds, and 65% of CHON⁻ compounds.
226 Thus, the classification of potential pathways for NOC formation was representative,

227 at least in this study.

228

229 **2.5. More parameter calculations and data analysis**

230 A thermodynamic model (ISORROPIA-II) was used to estimate the ALW
231 concentration and pH value, as described in previous studies (Xu et al., 2020b; Xu et
232 al., 2023; Xu et al., 2022c). Ambient hydroxyl radical ($\bullet\text{OH}$) concentrations were
233 predicted using empirical formulas proposed by Ehhalt and Rohrer (2000), which was
234 reported in detail in our previous field observations (Liu et al., 2023a; Lin et al., 2023).
235 The ventilation coefficient (VC) is an indicator of the potential for atmospheric
236 dilution of pollutants, which was calculated by multiplying the wind speed by the
237 planetary boundary layer height (PBLH) (Gani et al., 2019) .

238 Non-metric multidimensional scaling (NMDS) was employed to visualize the
239 distributions of NOCs (CHON^+ , CHN^+ , and CHON^- compounds) in two dimensions,
240 based on Bray-Curtis distances (Chao et al., 2006). The stress values ranged from
241 0.03 to 0.11 (less than 0.2, **Table S5**) in our analysis, indicating that the differences
242 among samples can be well represented in the two-dimensional pattern. To further
243 assess the influence of anthropogenic emissions and aqueous-phase processes on the
244 distribution of NOCs, the envfit function in the R package Vegan (Oksanen et al.,
245 2007) was utilized. Furthermore, the Spearman rank correlation, a non-parametric
246 measure with less sensitivity to outliers and independent of data distribution

247 assumptions, was employed to examine the association patterns between NOCs and
248 the parameters related to anthropogenic emissions and aqueous-phase processes
249 (Kellerman et al., 2014).

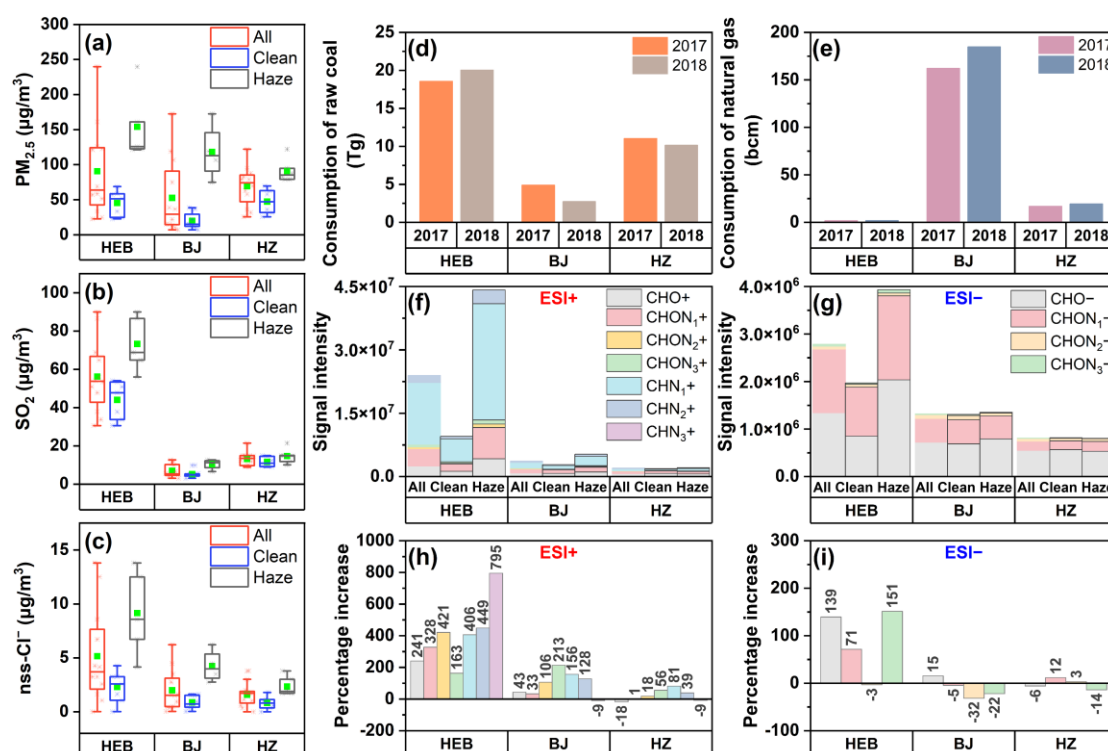
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251 **3. Results and discussion**

252 **3.1. Overview of pollution and aerosol NOC characteristics in different cities**

253 **Figure 1a–c** and **Table S1** show the variations in major gaseous pollutants, PM_{2.5}
254 and its major compositions, as well as meteorological parameters in three Chinese
255 cities with different energy consumptions during winter. The average PM_{2.5}
256 concentration in HEB was $90.6 \pm 62.4 \mu\text{g m}^{-3}$, which was significantly higher than
257 that observed in BJ ($52.7 \pm 51.4 \mu\text{g m}^{-3}$) and HZ ($69.1 \pm 29.6 \mu\text{g m}^{-3}$). Similarly, the
258 concentrations of SO₂ and nss-Cl⁻ were higher in HEB than in BJ and HZ. In addition,
259 a lower NO₃⁻/SO₄²⁻ mass ratio (**Table S1**) was found in HEB. SO₂ and nss-Cl⁻ have
260 been suggested to be typical pollutants emitted from coal combustion during winter in
261 cities (Zhao and Sun, 1986; Streets and Waldhoff, 2000). The low NO₃⁻/SO₄²⁻ mass
262 ratio can indicate a predominance of stationary sources (e.g., coal combustion) (Wang
263 et al., 2006; Arimoto et al., 1996; Xiao and Liu, 2004). These results suggest that coal
264 combustion during the winter heating season in HEB may significantly contributed to
265 severe PM_{2.5} pollution. This consideration can also be supported by the highest coal
266 consumption in HEB in 2017–2018 (**Fig. 1d**). Due to the large-scale use of clean

267 energy (i.e., natural gas) in BJ (**Fig. 1e**), the coal consumption in BJ was the lowest
 268 (**Fig. 1d**). This resulted in the lowest pollutant levels in BJ. From clean to polluted
 269 days, HEB and BJ showed larger increases in pollutant levels (e.g., PM_{2.5}, SO₂, and
 270 CO), followed by HZ. Thus, the release of pollutants caused by the use of fossil fuels
 271 for centralized heating in winter (only occurred in HEB and BJ) was undoubtedly one
 272 of the important factors contributing to the generation of haze in HEB and BJ.



273
 274 **Figure 1.** Box and whisker plots showing variations in the concentration of (a) PM_{2.5},
 275 (b) SO₂, and (c) nss-Cl⁻ in all (gray), clean (blue), and haze (red) periods in different
 276 cities. Each box encompasses the 25th–75th percentiles. Whiskers are the 5th and
 277 95th percentiles. The green squares and solid lines inside boxes indicate the mean and
 278 median value. The consumption of (d) raw coal and (e) natural gas in 2017 and 2018

279 in different cities was obtained from the local statistical yearbooks. Average
280 distributions in the signal intensity of species detected in PM_{2.5} collected during
281 different winter periods in different cities in (f) ESI+ and (g) ESI- modes. Percentage
282 variations in the signal intensity of each subgroup from clean to haze periods in
283 different cities in (h) ESI+ and (i) ESI- modes.

284

285 **Figure 1f** and **g** show the average signal intensity distributions of organic
286 compounds detected in PM_{2.5} across sampling periods in different cities. The detailed
287 mass spectra of organic compounds detected in ESI+ and ESI- were shown in **Fig. S6**.
288 CHN₁₊ ($n = 437-448$) compounds were the main CHN molecules measured in ESI+
289 mode in all cities (**Fig. 1f** and **Table S6**), the signal intensity of which accounted for
290 over 77% of the total CHN₁₋₃₊ signal intensity. Similarly, CHON₁₊ compounds ($n =$
291 $398-421$) dominated in CHON₁₋₃₊ molecules, with a higher signal intensity than
292 CHON₂₋₃₊ (**Fig. 1f** and **Table S6**). The high abundances of CHN₁₊ and CHON₁₊
293 compounds in NOCs were similar to previous reports about the NOC characteristics
294 of urban aerosols (He et al., 2024; Abudumutailifu et al., 2024). The signal intensity
295 fractions (40%–77%) of CHN+ compounds in total NOCs in these three cities were
296 higher than those observed (8.20%–17.47%) during winter in Ürümqi where the same
297 NOC analysis method was conducted (Ma et al., 2024). However, the signal intensity
298 fractions of CHON+ compounds in total NOCs were lower in these three cities

299 (23%–60%) than in Ürümqi (over 82.53%) (Ma et al., 2024). More frequent biomass
300 burning and relatively dry climate in Ürümqi (northwest China) (Ma et al., 2024) may
301 result in different sources and formation processes of NOCs compared to this study.
302 The signal intensity of these NOCs detected in ESI+ mode varied across cities, with
303 the highest CHN+ and CHON+ signal intensities in HEB, followed by BJ and HZ.
304 Moreover, we found that the total signal intensities of CHN+ and CHON+ compounds
305 increased by 382% in HEB from clean to haze periods, followed by increase of 102%
306 in BJ and increase of 31% in HZ (**Fig. 1h** and **Table S6**). This variation pattern of
307 CHN+ and CHON+ compounds from clean to haze periods was similar to that of the
308 pollutants mentioned previously (**Fig. 1a–c**). Given the high sensitivity of ESI+ mode
309 to protonatable species, reduced species (e.g., amine- and amide-like compounds)
310 were expected to predominate the NOCs (Han et al., 2023; Wang et al., 2018), the
311 formation of which was highly related to precursor emission level, aerosol acidity, and
312 ALW concentrations (Kuwata and Martin, 2012; Vione et al., 2005; Yang et al., 2024a;
313 Xu et al., 2020b). Thus, these results suggest that there may be significant differences
314 in the sources, precursor emission intensity, or main formation pathways of NOCs in
315 different energy consuming cities.

316 The number of NOCs identified in ESI– (296–301 molecules excluding sulfur-
317 containing compounds, **Table S7**) was found to be lower than that observed in ESI+
318 (1346–1361) (**Table S6**). This finding was similar to previous observations about the

319 NOCs of urban organic aerosols in Beijing, Mainz, Changchun, Guangzhou, and
320 Shanghai (Wang et al., 2021b; Wen et al., 2023; Wang et al., 2018). CHON₁-
321 compounds were the main NOC molecules in ESI⁻ mode in all cities (**Fig. 1g** and
322 **Table S7**). The average signal intensity of CHON⁻ compounds was highest in HEB,
323 followed by BJ and HZ. Moreover, the outbreak of CHON₁₋₃ signal intensity during
324 polluted periods was found in HEB, whereas insignificant increases occurred in BJ
325 and HZ (**Fig. 1i**). Deprotonated NOCs with oxidized nitrogen-functional groups, such
326 as nitro (-NO₂) or nitrooxy (-ONO₂) groups, are more sensitive to the ESI⁻ mode
327 (Wang et al., 2017; Jiang et al., 2023; Yuan et al., 2023). Clearly, the formation of
328 aerosol CHON⁻ compounds was largely dependent on atmospheric oxidation capacity
329 and gas- and aqueous-phase reactions (Ng et al., 2017; Shi et al., 2020; Shi et al.,
330 2023). Thus, the differences in CHON⁻ compound abundance in different polluted
331 periods and cities together with the spatiotemporal changes in CHN⁺ and CHON⁺
332 abundances mentioned previously were likely attributed to variations in sources,
333 mechanisms, or key influencing factors of NOC formation in these three cities, which
334 will be further discussed in the following sections.

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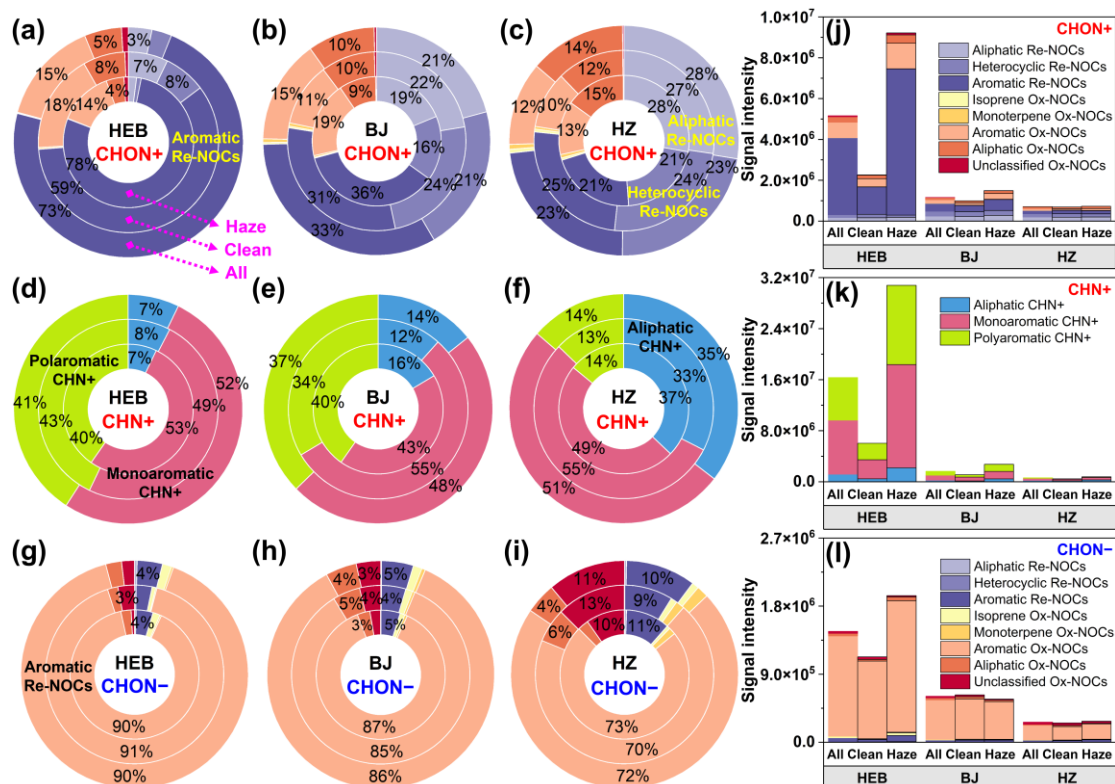
336 **3.2. Potential precursors of aerosol NOCs in different cities**

337 **Figure 2** presents the average signal intensity percentage and signal intensity
338 distributions of different NOCs from various precursors in different cities during

339 winter. Aromatics-, heterocyclics-, and aliphatics-derived Re-NOCs together
340 accounted for more than 74% (74%–79%) of the total signal intensity of CHON+
341 compounds in the three study cities (**Fig. 2a–c** and **Table S8**). Specifically, the
342 proportion of the aromatics-derived CHON+ signal intensity in the total CHON+
343 signal intensity was much higher in HEB (73%) than in BJ (33%), with the lowest
344 proportion observed in HZ (23%) (**Fig. 2a–c**). Furthermore, we observed that
345 aromatic CHN+ compounds (mono- and poly-aromatics) dominated the total CHN+
346 compounds in both number and abundance in all investigated cities (**Table S9** and **Fig.**
347 **2d–f**). The average signal intensity percentage and signal intensity of aromatic CHN+
348 compounds were also highest in HEB (**Fig. 2d–f** and **k**). The calculated AI_{mod} values
349 for CHON+ and CHN+ compounds were higher in HEB than in BJ and HZ (**Table**
350 **S10**), which further indicated a higher aromaticity of these NOCs in HEB. It has been
351 suggested that coal combustion can release a large amount of aromatic compounds
352 (Zhang et al., 2023a), which potentially increased NOC aromaticity (Yuan et al.,
353 2023). Thus, the higher signal proportion of aromatics-derived Re-NOCs in HEB can
354 be explained by the higher coal combustion emissions during winter. In contrast, the
355 use of clean energy during the central heating season in BJ and the reduced emissions
356 in HZ without central heating weakened the formation of aerosol aromatic NOCs.

357 CHON- compounds were also primarily dominated by aromatics-derived Ox-
358 NOCs in all three cities, accounting for more than 73% (73%–90%) of the total signal

359 intensity of CHON⁻ compounds, on average (**Fig. 2g-i**). This finding was consistent
360 with field observations conducted in other Chinese cities such as Shanghai,
361 Changchun, Guangzhou, and Wangdu during winter (Wang et al., 2021b; Jiang et al.,
362 2023). The abundance of aromatics-derived Ox-CHON⁻ compounds and the AI_{mod}
363 value of CHON⁻ were highest in HEB and decreased sequentially in BJ and HZ (**Fig.**
364 **2l** and **Table S10**), further indicating our previous consideration that coal combustion
365 heating in HEB can lead to higher NOC pollution. It is worth noting that the
366 percentage of total signal intensity of Ox-NOCs with biogenic VOCs (BVOCs) as
367 precursors was less than 3% (**Fig. 2g-i** and **Table S8**). This can be partly supported by
368 the previous observations showing that anthropogenic VOCs (AVOCs) were the main
369 contributors to the formation of Ox-NOCs (e.g., organic nitrates) in urban areas in
370 China (Wang et al., 2021b; Jiang et al., 2023). The overall results suggest the
371 significant role of AVOCs in the formation of NOCs in all investigated cities,
372 particularly in HEB.



373

374 **Figure 2.** Average percentage distributions of signal intensities for (a–c) CHON+,
 375 (d–f) CHN+, and (g–i) CHON– compounds from various sources in PM_{2.5} collected
 376 from different cities during winter. Average signal intensity distributions for (j)
 377 CHON+, (k) CHN+, and (l) CHON– compounds from various sources in PM_{2.5}
 378 collected from different cities during winter.

379

380 From clean to haze periods, the signal intensities of all aromatics-derived CHON
 381 compounds increased significantly in HEB (Figs. 2a, j, g, l and S7). In contrast, the
 382 signal intensities of aromatics-derived CHON compounds in BJ and HZ showed an
 383 insignificant increase during haze periods. In addition, the average values of O/C_w and
 384 OS_{Cw} for CHON+ and CHON– compounds were higher in HEB than in BJ (second

385 highest) and HZ, and their increases from clean to haze periods were also greater in
386 HEB (Table S10). Concurrently, the O/C_w ratio of aerosol NOCs in HEB was
387 observed to be higher than that of coal-derived aerosols (Song et al., 2018). Heald et
388 al. (2010) previously demonstrated that oxidation processes can lead to an increase in
389 the O:C ratio of organic aerosols. These results indicated that aerosol NOCs in HEB
390 were more oxidized aromatics (or aged aromatics), particularly during haze. The
391 average N/C_w ratios of $CHON^+$ and $CHON^-$ compounds in HEB (0.13 and 0.15,
392 respectively) (Table S10) were higher than those of $CHON^+$ (0.079) and $CHON^-$
393 (0.07) compounds in aerosols directly emitted from coal combustion (Song et al.,
394 2022; Song et al., 2018). The N/C_w ratios were also higher in HEB than in BJ and HZ
395 and increased during hazy days (0.13 for $CHON^+$ and 0.16 for $CHON^-$ in hazy days
396 in HEB). It has been suggested that the N/C_w ratio of $CHON^-$ compounds tended to
397 increase (from 0.109 to 0.119) after aging of fuel combustion-derived aerosols (Zhao
398 et al., 2022a). Thus, these results, combined with previous analysis of potential
399 precursors for NOCs, suggest that anthropogenic precursor emissions and their
400 atmospheric transformation to form CHON compounds were stronger in HEB than in
401 BJ and HZ. Moreover, considering that the emission intensity of precursors during
402 clean and hazy days may not significantly change, secondary processes may
403 significantly promote the formation of NOCs in HEB during hazy days (the most
404 significant increase in NOC abundance). However, this promoting effect during hazy

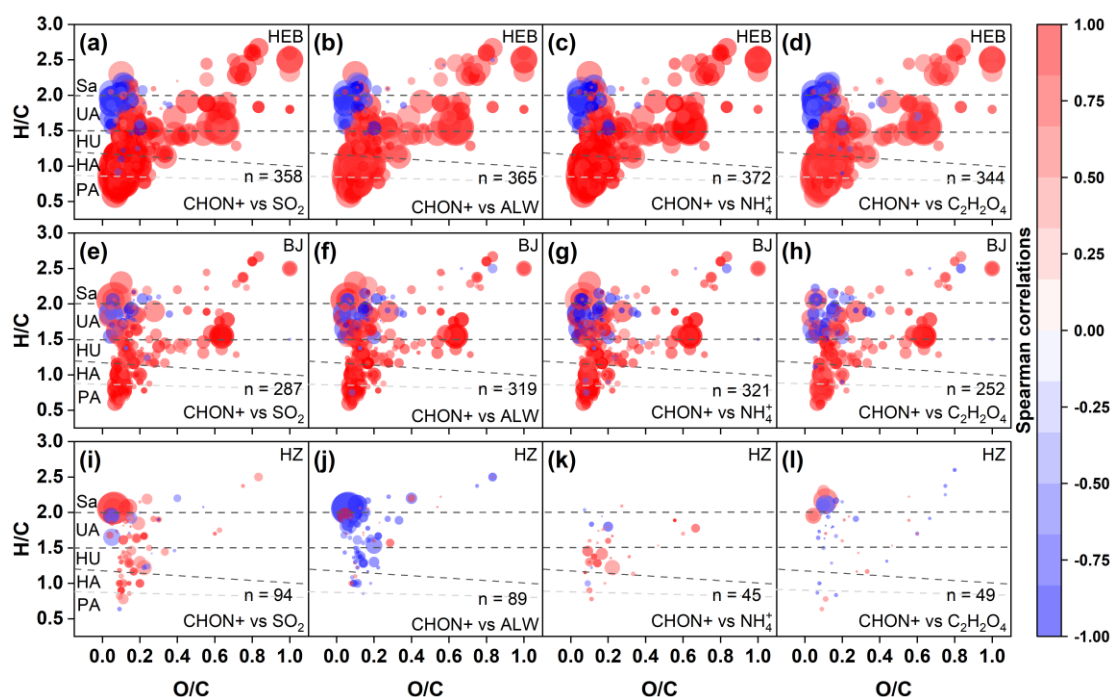
405 days was insignificant in BJ and HZ (less increase in NOC abundance).

406

407 **3.3. Main factors influencing aerosol NOC formation in different cities**

408 As discussed in the previous section, the results indicated that AVOCs play a
409 significant role in the formation of NOCs. Furthermore, secondary processes may
410 contribute to NOC formation to varying extents in different cities. This section
411 provides a detailed discussion of the key factors influencing the molecular distribution
412 of NOCs. First, a Spearman correlation analysis was performed to examine the
413 relationship between various parameters and NOCs (**Fig. 3** and **Figs. S8–S12**). The
414 peak intensity of most CHON⁺ compounds (mainly aromatics, as mentioned
415 previously) showed a strong correlation ($P < 0.01$) with the concentrations of
416 combustion source-related tracers (e.g., SO₂, nss-Cl⁻, nss-K⁺, CO, and NO₂) (Zhao
417 and Sun, 1986; Streets and Waldhoff, 2000; Shen et al., 2009; Zhang et al., 2011;
418 Mafusire et al., 2016; Liu et al., 2019; Zhang et al., 2021a; Wang et al., 2020) in HEB
419 (**Figs. 3a** and **S8a–d**). Although there was a significant correlation ($P < 0.05$) between
420 most CHON⁺ compounds and those combustion source indicators in BJ, the strength
421 of this correlation was weaker in BJ than in HEB (**Figs. 3e** and **S8f–i**). However,
422 similar significant correlations between them were not observed in HZ (**Figs. 3i** and
423 **S8k–n**). Thus, the greatest contribution of anthropogenic activities to the formation of
424 CHON⁺ compounds in winter was in HEB (central heating with coal), followed by BJ

425 (central heating with coal and natural gas) and HZ (without central heating). Most of
 426 CHN⁺ and CHON⁻ compounds showed a similar spatial response pattern to those
 427 anthropogenic activities (**Figs. S9** and **S10**). These results are consistent with the
 428 previous analysis of NOC precursors (**Fig. 2**), which concluded that the intensity of
 429 anthropogenic pollutant emissions in different energy consuming cities was an
 430 important factor affecting the formation of NOC and causing spatial differences in
 431 NOC abundance.



432
 433 **Figure 3.** Spearman rank correlation coefficients (with $P < 0.01$ in HEB and $P < 0.05$
 434 in BJ and HZ) of individual CHON⁺ molecules with selected parameters in (a–d)
 435 HEB, (e–h) BJ and (i–l) HZ. The color scale indicates Spearman correlations between
 436 the intensity of individual CHON⁺ molecules and each parameter. The symbol n in
 437 the bottom right corner of each panel indicates the number of molecular formulas

438 significantly correlated with the variables. The subgroups in the panels include
439 polycyclic aromatic-like (PA), highly aromatic-like (HA), highly unsaturated-like
440 (HU), unsaturated aliphatic-like (UA), and saturated-like (Sa) compounds.

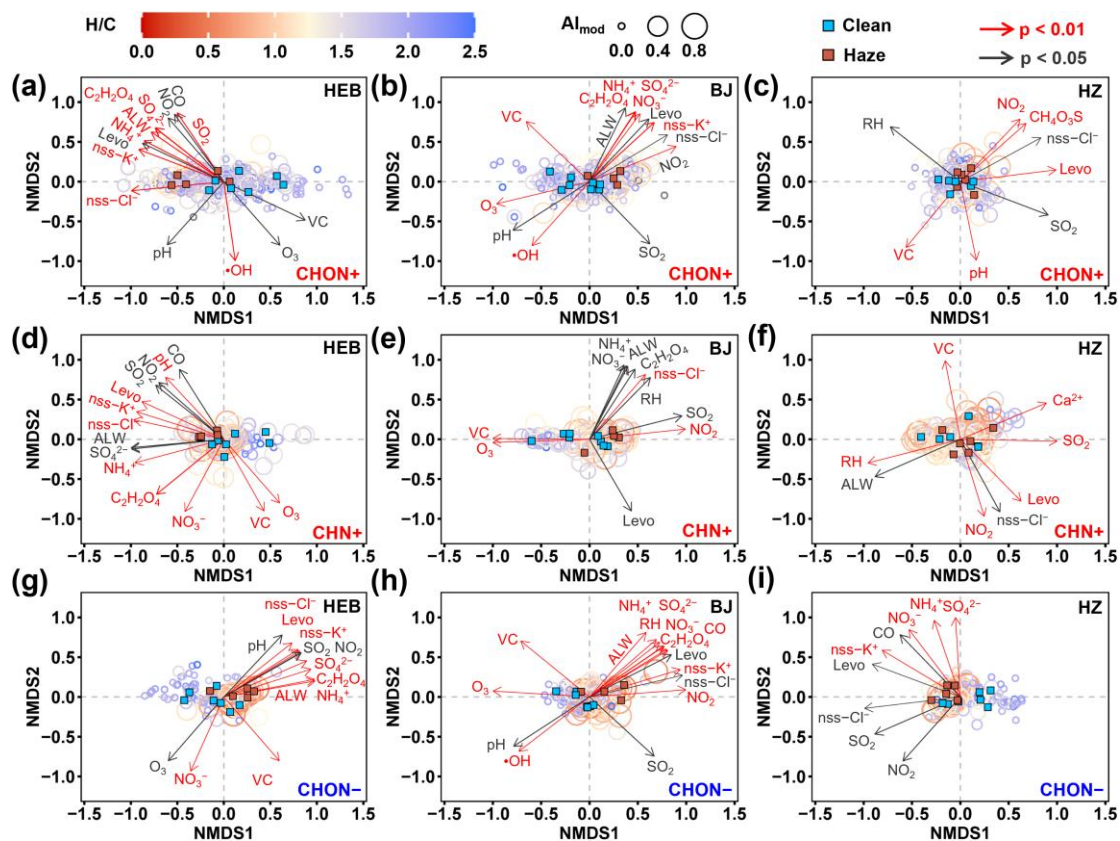
441

442 Furthermore, we found that the peak intensities of most CHON⁺, CHN⁺, and
443 CHON⁻ compounds (mainly aromatics) were significantly correlated ($P < 0.01$) with
444 the concentrations of ALW, NH₄⁺, oxalic acid, and SO₄²⁻ (Figs. 3b–d, S8e, and
445 S11–S12) in HEB. The correlations between these NOCs and parameters weakened in
446 BJ and disappeared in HZ (Figs. 3, S8, and S11–12). It is generally accepted that
447 SO₄²⁻, NH₄⁺, and NO₃⁻ in fine aerosols are primarily formed through secondary
448 processes (Gao et al., 2021; Wang et al., 2021d). NH₄⁺ can serve as a key reactant in
449 the formation of aerosol NOCs (e.g., "carbonyl-to-imine" transformation) in the
450 aqueous-phase (Laskin et al., 2014; Lee et al., 2013; Li et al., 2019b). Oxalic acid
451 (C₂H₂O₄) has been identified as a marker (defined by Nozière et al. (2015)) for
452 aqueous-phase SOA (Xu et al., 2022a; Chen et al., 2021). Additionally, numerous
453 laboratory and field observational studies have shown that ALW can promote the
454 formation of NOCs (Lv et al., 2022; Liu et al., 2023b; Jimenez et al., 2022; Jiang et al.,
455 2023). Thus, these results indicate that aqueous-phase processes can significantly
456 promote the formation of NOCs in HEB, however, as the precursor emission intensity
457 gradually decreased in BJ and HZ, this aqueous-phase promoting effect also

458 decreased.

459 The NMDS analysis between various parameters and NOCs was conducted to
460 further investigate the variations in key factors affecting the formation of NOCs from
461 clean to haze days (**Fig. 4**). The formation of CHON⁺, CHN⁺, and CHON⁻
462 compounds with higher AI_{mod} values (mainly aromatics, as mentioned previously)
463 during haze days in HEB and BJ were closely associated with the factors indicating
464 anthropogenic precursor emissions and aqueous-phase reaction processes. In contrast,
465 the level of oxidants (i.e., O₃ and •OH) played a more important role during clean
466 days in HEB and BJ, driving more highly saturated NOC formation during clean days
467 (**Fig. 4**). A reasonable explanation for this is that the solar radiation and •OH levels on
468 polluted days were lower than those on clean days (**Table S1**). The impacts of various
469 factors on the formation of aerosol NOCs showed a weak discrimination between
470 haze and clean days in HZ (**Fig. 4c, f and i**). Laboratory studies have shown that
471 reactive components (e.g., •OH and H₂O₂) in the aqueous phase can continuously
472 convert low-solubility organics to form aqueous phase SOA (Chen et al., 2008; Huang
473 et al., 2011; Dong et al., 2021). Field observations also suggested that precursors
474 (most of them are aromatic compounds) released from the combustion of fossil fuels
475 significantly contributed to the aqueous SOA formation (> 50% total molecules) (Xu
476 et al., 2022a) through the rapid aqueous-phase conversion of primary organic aerosol
477 (POA) to SOA at high RH (Wang et al., 2021a). This implies that higher precursor

478 abundance can drive more aerosol NOC formation via aqueous-phase processes. As
479 mentioned previously, the emission intensity of precursors decreased sequentially
480 from HEB to BJ and then to HZ. Moreover, the ALW concentrations were much
481 higher on polluted days than on clean days in three investigated cities. The rising
482 ALW during the pollution period and the quiescent steady state of the atmosphere
483 favored the formation of SOA from anthropogenic emission precursors (Guo et al.,
484 2014; He et al., 2018). Thus, the above discussion can suggest that the spatial
485 differences in precursor emission intensity (higher in HEB) and enhancement of
486 aqueous-phase processes in polluted days were the main factors leading to the
487 differences in the proportion (higher in HEB) of increase in NOC abundance from
488 clean days to polluted days in three different energy consuming cities. In addition, the
489 increased VC value (**Table S1**) in clean days (beneficial for the diffusion of pollutants)
490 (Gani et al., 2019) was also an important factor limiting the abundance of NOCs (**Fig.**
491 **4**), resulting in a lower NOC abundance on clean days compared to polluted days (**Fig.**
492 **1**).



493

494 **Figure 4.** Nonmetric multidimensional scaling of (a–c) CHON+, (d–f) CHN+, and

495 (g–i) CHON– compounds from organic aerosol in different cities. The color and size

496 of the circle indicate the H/C ratio and AI_{mod} value of compounds, respectively.

497 Significant relationships between the variables and ordination (999 permutations) are

498 indicated by $p < 0.05$ (grey) and $p < 0.01$ (red). Insignificant correlations are not

499 shown. The scores of the samples collected during clean and haze periods were shown

500 as blue and brown squares, HEB, respectively.

501

502 As mentioned above, the aerosol NOCs of HZ were less affected by

503 anthropogenic pollutants emitted from coal and natural gas combustion compared to

504 HEB and BJ with centralized heating. Interestingly, we found that the molecular
505 distributions of most aromatic CHON⁺ compounds in HZ were not only influenced
506 by some anthropogenic pollutants (e.g., SO₂ and NO₂), but also by methanesulfonic
507 acid (CH₄O₃S) (**Fig. 4c**). Moreover, neither CHN⁺ nor CHON⁺ exhibited significant
508 correlations with factors related to secondary processes in HZ (**Fig. 4c** and **f**).
509 Methanesulfonic acid has been suggested to be a tracer for ocean aerosols (Ayers and
510 Gras, 1991; Suess et al., 2019). These results suggest that aerosol CHON⁺
511 compounds in HZ may be influenced by long-distance transport air masses originating
512 from the ocean. This consideration can be also supported by the fact that only HZ was
513 affected by air masses originated from the ocean (**Fig. S13**). Thus, marine emissions
514 may be an important contributor to aerosol NOCs in HZ, which was significantly
515 different from the cases of HEB and BJ where aromatic pollutants from fossil fuel
516 combustion and aqueous-phase processes control the composition and abundance of
517 aerosol NOCs.

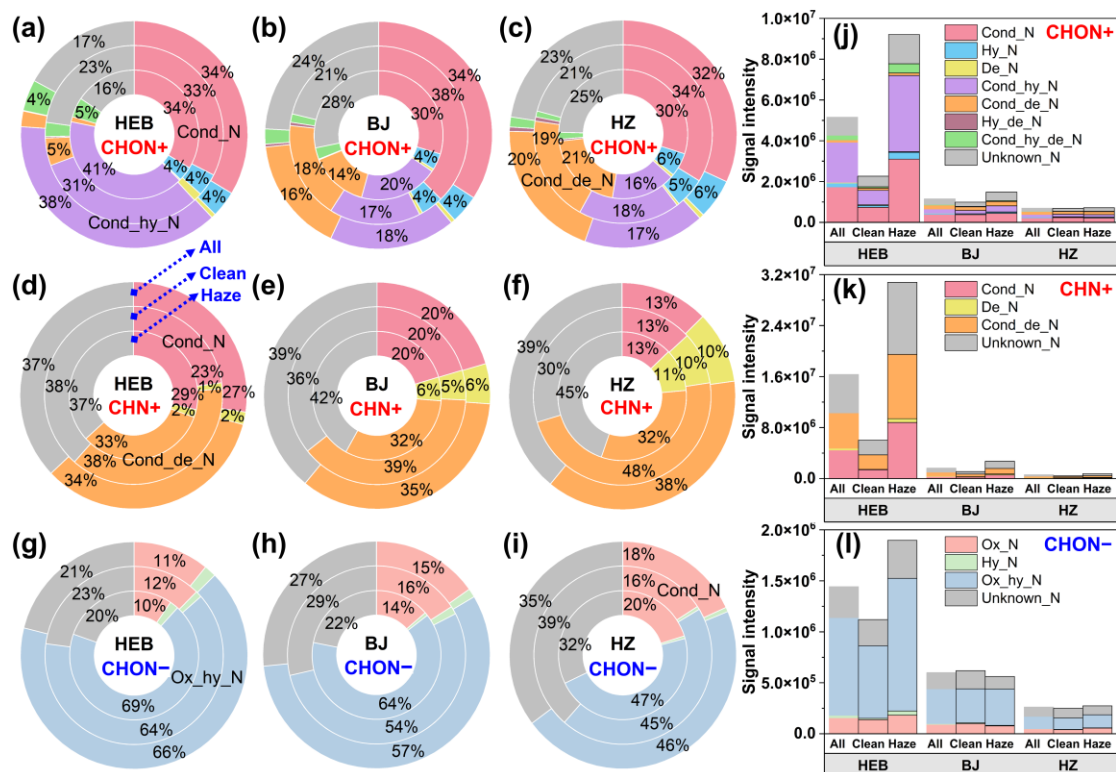
518

519 **3.4. Potential formation mechanisms of aerosol NOCs in cities with different** 520 **energy consumption**

521 **Figure 5** shows the average signal intensity percentage and signal intensity
522 distributions of NOCs formed by different aqueous-phase processes (**Table S4** and
523 **Figs. S3–S5**) in different cities during winter. The identification of specific reaction

524 pathways was detailed in **Figs. S3–S5** and **Sect. S4**. During the entire study period,
525 the cond_N, cond_hy_N, and cond_de_N pathways together accounted for more than
526 68% (68%–74%) of the total signal intensity of CHON⁺ compounds in the three cities
527 (**Fig. 5a–c** and **Table S11**). Specifically, the formation of CHON⁺ compounds was
528 mainly dominated by the cond_N and cond_hy_N pathways in HEB, with less impact
529 from the cond_de_N pathway (**Fig. 5a**). However, CHON⁺ compounds derived from
530 the cond_de_N pathway showed a much higher proportion in BJ and HZ than in HEB
531 (**Fig. 5b** and **c**). The cond_de_N pathway involves both condensation and dehydration
532 processes (**Table S4** and **Fig. S3**). Recent studies have identified that dehydration
533 reactions may occur in aerosols and fog water (Sun et al., 2024), as well as in
534 photochemical transformations of organic compounds in aqueous phase (Lian et al.,
535 2020). While the exact pathways of dehydration reactions in the particle phase remain
536 uncertain, stronger solar radiation in BJ and HZ than in HEB (**Table S1**) may partly
537 explain the higher signal proportion of CHON⁺ compounds formed through the
538 cond_hy_N pathway in BJ and HZ. Furthermore, the higher signal proportions of
539 CHN⁺ compounds formed through the de_N pathway in BJ (6%) and HZ (11%) than
540 in HEB (2%) may also be associated with this solar radiation-induced dehydration
541 mechanism (**Fig. 5d–f** and **Table S12**). For CHN⁺ compounds, the cond_de_N
542 process dominated their formation (**Fig. 5d–f**). In general, the cond_N, cond_hy_N,
543 and cond_de_N processes contributed most significantly to the formation of Re-NOCs

544 in HEB, followed by BJ and HZ.



545

546 **Figure 5.** Average percentage distributions of signal intensities for aerosol (a–c)

547 CHON+, (d–f) CHN+, and (g–i) CHON– compounds from various reaction pathways

548 in different cities during winter. Average signal intensity distributions for aerosol (j)

549 CHON+, (k) CHN+, and (l) CHON– compounds from various reaction pathways in

550 different cities during winter.

551

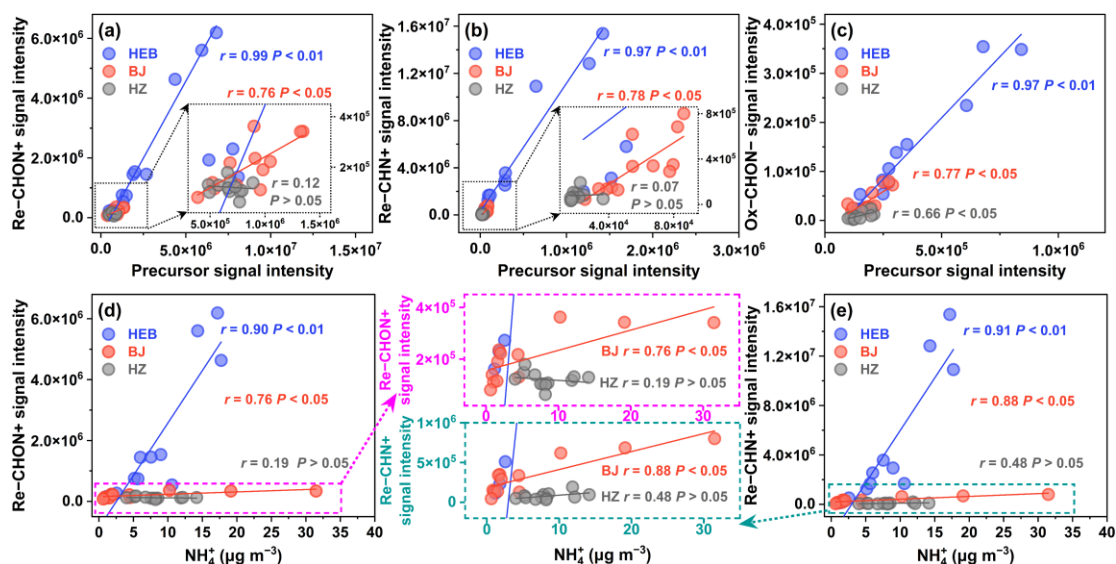
552 A typical mechanism for Re-NOC formation is the aqueous-phase reactions

553 between carbonyl compounds and NH_4^+ (or NH_3) (Abudumutailifu et al., 2024;

554 Laskin et al., 2014; Li et al., 2019b; Liu et al., 2023b; Wang et al., 2024). If this

555 mechanism is simplified as a second-order reaction (i.e., $[\text{Precursor}] + [\text{NH}_3 \text{ and } \text{NH}_4^+]$)

556 \leftrightarrow [Re-NOCs]), the production of Re-NOCs is expected to be proportional to the
 557 abundances of precursor and NH_4^+ (Yang et al., 2023; Lin et al., 2023). Indeed, the
 558 signal intensities of the Re-CHON+ and Re-CHN+ compounds were significantly
 559 positively correlated with the signal intensities of their CHO precursors (identified
 560 using the precursor-product pairs theory, **Figs. S3** and **S4**) and NH_4^+ concentration in
 561 HEB (**Fig. 6a, b, d** and **e**). This correlation gradually weakened from BJ to HZ (**Fig.**
 562 **6a, b, d** and **e**). As previously discussed, differences in energy consumption patterns
 563 resulted in the highest levels of anthropogenic aromatic compound emissions in HEB
 564 during the winter, followed by BJ, with the lowest levels in HZ (**Figs. 2** and **S14**).
 565 Thus, the signal intensities of CHON+ and CHN+ compounds from cond_N,
 566 cond_de_N, and cond_hy_N processes were higher in HEB than in BJ and lowest in
 567 HZ (**Fig. 5j** and **k**).



568
 569 **Figure 6.** Signal intensity of (a) Re-CHON+, (b) Re-CHN+, and (c) Ox-CHON-
 570 compounds as functions of signal intensity of precursors (CHO compounds). Signal

571 intensity of (d) Re-CHON⁺ and (e) Re-CHN⁺ compounds as functions of the
572 concentrations of NH₄⁺.

573

574 Additionally, we noticed that the contribution of these aqueous-phase processes to
575 the formation of CHON⁺ and CHN⁺ compounds increased significantly from clean to
576 hazy days in HEB and BJ (**Fig. 5**). The increased ALW concentrations (**Table S1**) and
577 atmospheric stability during haze periods likely provided favorable conditions for the
578 precursors to undergo these aqueous-phase reactions, resulting in the formation of
579 NOCs. Clearly, high pollutant emission levels in HEB provided a greater potential to
580 convert precursors into more NOCs via the cond_N, cond_{hy}_N, and cond_{de}_N
581 processes during haze periods. Thus, the hazy days in the HEB showed the largest
582 increase in CHON⁺ and CHN⁺ compounds from the cond_N, cond_{hy}_N, and
583 cond_{de}_N processes (**Fig. 5j** and **k**). In contrast, due to generally mild winters
584 leading to the absence of heating and the implementation of stricter pollution control
585 measures (more coal usage in HZ than in BJ, as shown in **Fig. 1d**), the precursor
586 emissions in HZ were lower. These emissions were insufficient to support the
587 production of large amounts of NOCs in the aqueous phase. These results also
588 indicate that emission reduction is the key to controlling aerosol NOC pollution.

589 CHON⁻ compounds derived from the ox_{hy}_N and ox_N processes together
590 accounted for more than 64% (64%–71%) of the total signal intensity of CHON⁻

591 compounds in the three cities (**Fig. 5g-i, l** and **Table S13**). The signal intensity
592 proportions of CHON⁻ compounds formed by the ox_hy_N process in these three
593 cities (ranging from 47% in HZ to 69% in HEB) were higher than that in Wangdu (<
594 20%) (Jiang et al., 2023). The observation study in Wangdu examined aerosol organic
595 components only in ESI⁻ mode (Jiang et al., 2023), which may underestimate the
596 importance of the CHO⁺ compounds that could serve as precursors of Ox-NOCs. In
597 general, CHON⁻ compounds formed through the ox_hy_N and ox_N processes
598 showed the highest abundance in HEB, followed by BJ and HZ (**Fig. 5g-i**). According
599 to a simplified reaction mechanism for the formation of Ox-NOCs via aqueous-phase
600 processes (i.e., [Precursor] + [Oxidants] ↔ [Ox-NOCs]) (Shi et al., 2023; Kroflič et
601 al., 2015; Vione et al., 2005), we can infer that Ox-NOCs production is proportional
602 to precursor levels when oxidants (e.g., NO₂ radical or NO₂⁺) are in a steady state in
603 the atmosphere. Indeed, the signal intensities of the Ox-CHON⁻ compounds were
604 significantly positively correlated with the signal intensities of their CHO precursors
605 identified using the precursor-product pairs theory in HEB (**Fig. 6c**). Moreover, this
606 correlation gradually weakened from BJ to HZ (**Fig. 6c**). Thus, the spatial differences
607 in the contribution of the ox_hy_N and ox_N processes to Ox-NOC production across
608 the three cities can also be explained by differences in precursor emission intensity, as
609 indicated by above mentioned Re-NOC formation.

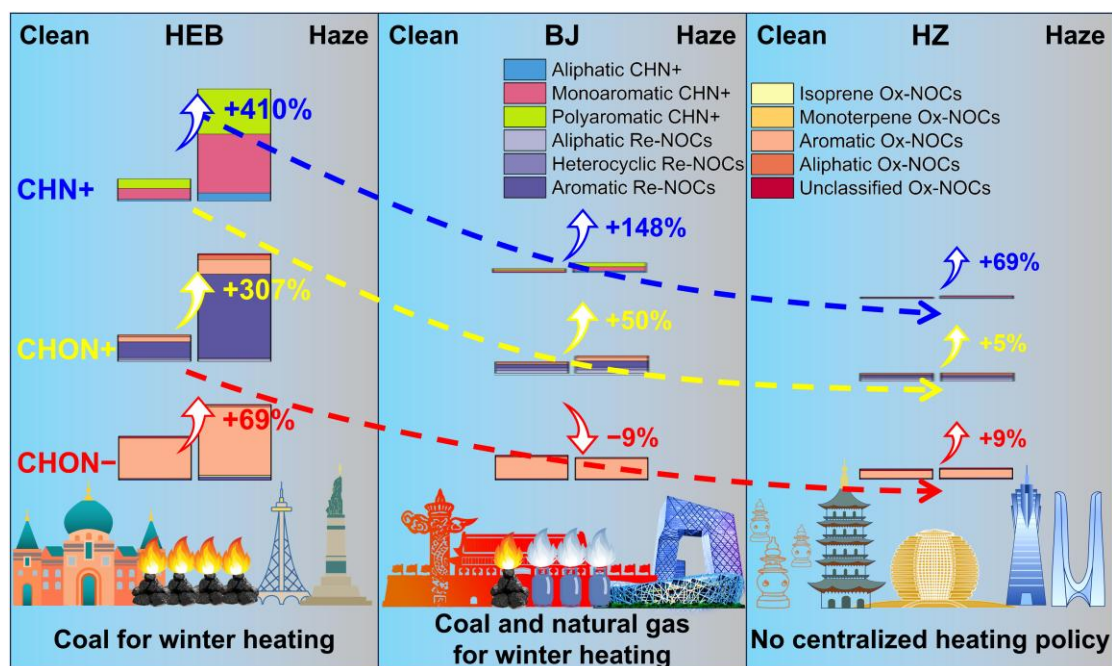
610

611 4. Conclusion

612 The abundance, composition, potential precursors, and potential formation
613 mechanisms of NOCs in PM_{2.5} in three Chinese cities with different energy
614 consumption types during the winter were systematically investigated. On average,
615 the total signal intensity of NOCs (i.e., CHN⁺, CHON⁺, and CHON⁻ compounds)
616 was highest in HEB, followed by BJ. The lowest total NOC signal intensity was found
617 in HZ. According to the identification of potential precursors of NOCs, we found that
618 anthropogenic aromatic compounds were the main precursors of NOCs during winter
619 in HEB where mainly relies on coal for winter heating, with less impact from BVOCs.
620 Anthropogenic aromatic precursors were also identified to be important contributors
621 to NOC formation in BJ which uses natural gas and coal for winter heating, although
622 the contribution ratio was lower in BJ than in HEB. **In contrast, due to generally mild
623 winters resulting in the absence of a winter heating policy and the implementation of
624 strict pollution control measures, as mentioned previously, aromatic precursor
625 emissions in HZ were expected to be the lowest.** Furthermore, the NMDS analysis
626 supported that the impact of anthropogenic fossil fuel combustion on NOC pollution
627 gradually decreased from HEB to BJ and then to HZ.

628 The formation of CHON⁺ compounds was mainly associated with the cond_N,
629 cond_{hy_N}, and cond_{de_N} processes. The cond_N and cond_{de_N} processes
630 dominated the formation of CHN⁺ compounds. The production of CHON⁺ and

631 CHN+ compounds from the cond_N, cond_hy_N, and cond_de_N processes was
 632 highest in HEB, followed by BJ and HZ. The ox_hy_N and ox_N processes
 633 contributed significantly to CHON- compound formation, from which the highest
 634 CHON- production occurred in HEB and the lowest in HZ. The spatial differences in
 635 the contribution of different aqueous-phase processes to NOC production in the three
 636 different cities can be attributed to differences in precursor emission intensity. In
 637 particular, the contribution of these aqueous-phase processes to the formation of
 638 CHON+ and CHN+ compounds showed the most significant increase from clean to
 639 hazy days in HEB, followed by BJ. We concluded that high pollutant emission levels
 640 can provide a greater potential to convert precursors to produce more NOCs via
 641 aqueous-phase processes during haze periods. The above findings are summarized in
 642 a diagram (Fig. 7).



643

644 **Figure 7.** Conceptual illustration showing the characteristics of different NOCs from
645 the clean days to the haze days in different cities.

646 In general, the aerosol NOCs pollution during winter is closely linked to both the
647 intensity of precursor emissions and the efficiency of aqueous-phase processes in
648 converting these emissions into NOCs. The overall results highlight the importance of
649 emission reduction strategies in controlling aerosol NOCs pollution during winter. **It**
650 **is imperative to manage precursor emissions during hazy episodes in order to restrict**
651 **the increased formation of secondary NOCs in conditions of high humidity. Moreover,**
652 targeted reduction of precursor emissions, especially from coal combustion, could
653 significantly mitigate NOCs levels, thereby improving air quality and public health in
654 urban areas. **The transition to cleaner energy sources, as evidenced by the decreased**
655 **gradient of NOC pollution from HEB to BJ to HZ, represents an effective pathway for**
656 **the mitigation of NOC pollution.** Future research should focus on further elucidating
657 the specific pathways of aqueous-phase NOC formation and developing available
658 models to predict NOC dynamics under varying environmental conditions.
659 Additionally, research into the long-term effects of transitioning to cleaner energy
660 sources on the reduction of NOC pollution will be essential for guiding effective air
661 quality management strategies.

662

663 **Data availability.** The data presented in this work are available upon request from the

664 corresponding authors.

665

666 **Competing interests.** The authors declare no conflicts of interest relevant to this
667 study.

668

669 **Supplement.** Details of parameter calculation, classification method for identifying
670 precursors of NOCs, classification of possible aqueous-phase processes NOCs based
671 on precursor-product pairs, thirteen tables (Tables S1–S13), and fourteen extensive
672 figures (Figures S1–S14) are provided in the Supplement.

673

674 **Author contributions.** YX designed the study. YJM, TY, LG, HX, and HWX
675 performed field measurements and sample collection; YJM performed chemical
676 analysis; YX and YJM performed data analysis; YX and YJM wrote the original
677 manuscript; and YX, YJM, and HYY reviewed and edited the manuscript.

678

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