This paper presents calculations of POD6 from current conditions to the year 2100. Although changes in crop yield due to ozone over this period are of interest, I find it difficult to know what to make of the results when I have not been given any impression about whether the models can actually predict ozone and especially POD6 to any satisfactory degree in the base run. The lack of comparison with measurements is even more surprising given that this manuscript was submitted to the TOAR-II Special issue! I am afraid I find this omission to be too significant to be ignored, and therefore cannot recommend this manuscript for publication. More detailed comments follow.

Here we clarify key features of our modeling choices and tools used and summarize key results of the new analyses performed to address the issues raised. Detailed answers can be found in the point-by-point responses and in the newly written Supplementary materials.

The skill of CMIP6 models (UKESM1-0-LL and GFDL-ESM4) in predicting ozone concentration during the baseline period has been evaluated against TOAR ozone data in Turnock et al. (2020) study, which we reference. That study underwent peer review and was published in Atmospheric Chemistry and Physics, so we believe it provides reliable results that we confidently reference.

Therefore, our work acknowledges and builds on those findings and assumes those bias estimates as valid, with the noted uncertainties. Namely, Turnock et al. (2020) reported that CMIP6 models generally overestimate O₃ concentrations, although their comparison is performed between the O₃ concentrations at the lowest model level (roughly between 15 and 20 m above the ground), and the measured O₃ concentrations in the TOAR database (typically between 2 and 3 meters). Since O₃ concentrations are lower the closer to the ground, and since our dry deposition model scales O₃ concentrations from the lowest model level to the canopy height, this overestimation would reasonably be mitigated. Furthermore, it should be noticed that, despite the acknowledged uncertainties, the ozone data used in this study represent the only currently available coupled-chemistry datasets offering both hourly resolution and the required time span.

Although we believe that further evaluation is unnecessary for the outlined reasons, we provide an additional assessment of the bias between the ozone concentrations predicted by CMIP6 models (scaled from lowest model level to the canopy height) and those measured at stations in the TOAR-II database (scaled to the canopy height as well), according to the reviewer's request. This comparison shows that, as expected, the O₃ concentrations modelled by the CMIP6 models have a smaller bias (global averages: +1.31 and +0.88 ppb in GFDL-ESM4 and UKESM1-0-LL, respectively; <u>see supplementary materials attached at the end of this document for details</u>), compared to the overestimation reported by Turnock et al. (2020). Furthermore, we estimate the effect of the O₃ bias on the POD₆ estimate, and show that it is relatively small on average, due to climate limiting stomatal conductance.

While regional differences in O₃ concentrations and associated bias against observations are present, East Asia shows the largest bias, but, in fact, the propagated bias leads to relatively

low overestimation of POD₆ (mean: +0.28 and +0.33 mmol m⁻², for GFDL-ESM4 and UKESM1-0-LL respectively). Therefore, the biases found in the modelled O₃ concentrations are acceptable for the purposes of of our study. The main findings regarding the bias will be mentioned in the manuscript, both in the results and in the discussion section, and the supplementary materials will be referenced therein.

Regarding the POD₆ metric used to compute the O₃ damage, we clarify that it is calculated based on a modified version of the DO3SE model and following the same formulation (as stated in the manuscript). As this model is widely accepted as reliable for predicting ozone deposition against field measurements (as affirmed by the reviewer and the literature), results in our study will be similarly valid as based on a fundamentally analogous deposition model.

Nevertheless, we provide additional evidence on the validity of our dry deposition model by assessing its performances against data from a measurement campaign over a wheat field (Gerosa et al., 2003) Specifically, we compare modelled and observed values of total O_3 flux, latent heat flux and friction velocity on a half-hour basis, as detailed in the Supplementary Material and responses below. These variables were chosen because they are directly measured, while on the contrary stomatal conductance and stomatal ozone fluxes are derived from water and O_3 fluxes under some assumptions using an indirect inferential methodology. Our results indicate high skills in reproducing the latent heat flux (R²=0.74, and MAE =0.02 W/m²; see supplementary materials attached at the end of this document for details), which is a key proxy for the stomatal conductance, and thus of stomatal fluxes. Therefore, this further supports the appropriateness of using our model. These results are comparable with the ones reported by Mills et al. (2018, with R²=0.63, also cited by the reviewer in the major comments). This similarity was expected, since Mills et al. (2018) employed the DO3SE model, of which our dry deposition model is a modified version.

Moreover, a companion TOAR-II paper will evaluate specifically the performances and the sensitivity to input variables of different stomatal models, including the DO3SE (Emmerichs et al., in prep). We remark that the main objectives of this work are to evaluate the POD for bread wheat across the 21st century, and to identify the regions vulnerable to future food security threats, and that the evaluation of the sensitivity to input variables, and an in-depth evaluation of the performances of the CMIP6 models are beyond the scope of this work.

Major comments

1. The major weaknesses of this paper are that the authors have chosen to model a very difficult ozone metric (POD6), and they present no evidence to show that the models used have any ability to model that metric (or indeed any other), even for present day conditions.

The POD6 metric is indeed more difficult to calculate, especially in comparison to exposurebased metrics. However, flux-based metrics are frequently referred to as the correct approach to study O_3 risk to vegetation (Emberson, 2020; Mills et al., 2011). As mentioned above, our model follows the same paradigm of the DO3SE model, which is widely employed in ozone risk assessments to vegetation, and constitutes the O_3 dry deposition scheme within of the EMEP chemical transport model (Simpson et al., 2012), which is cited as a positive example by the reviewer in the Major comment #4. An evaluation of the dry deposition model is now presented in the supplementary materials, which are also attached at the end of the response.

2. First, about the metric itself. Why was POD6 chosen? It is well known that ozone metrics such as PODY can be very difficult to estimate, especially when the Y threshold is very high (e.g. Sofiev and Tuovinen et al, 2001, Tuovinen 2000, Touvinen et al, 2007). POD is also a difficult metric to obtain from observations because its calculation requires a large number of parameters, assumptions and auxiliary measurements that are usually not available. Such problems explain why the otherwise comprehensive TOAR database of vegetation-relevant ozone metrics (Lefohn et al, 2018, Mills et al, 2018a) did not include estimates of POD.

The reviewer seems to refer to TOAR-I. In the current TOAR version (TOAR-II), POD is fully recognized as the only metric effectively bridging atmospheric chemistry with plant physiology (e.g., Emberson, 2020; Mills et al., 2011). The harmful effects of ozone on vegetation are not due to the mere exposure to high ozone concentrations, but to the ozone uptake through stomatal pores. The concept of dose, therefore, augment the one of exposure (reflected in metrics like M7, AOT40, etc.), as documented by the ICP Vegetation (LRTAP Convention, 2017) and extensively reported in related papers. This distinction is explicitly noted in the mapping manual, which is cited by the reviewer: "Scientific evidence suggests that observed effects of O_3 on vegetation are more strongly related to the uptake of O_3 through the stomatal leaf pores (stomatal flux) than to the concentration in the atmosphere around the plants (Mills et al., 2011b)." (see chapter II.3.1.2 of the Mapping manual, whose insert is reported below). This is the main reason we chose to use POD, specifically POD6, as recommended by the mapping manual for wheat.

III.3.1.2 METRICS FOR CRITICAL LEVELS OF O₃ FOR VEGETATION

A glossary for all terms used for O₃ critical levels is provided in Annex III.2.

For O_3 , two types of metrics are available for risk assessment, either based on the cumulative stomatal flux or the cumulative exposure. Scientific evidence suggests that observed effects of O_3 on vegetation are more strongly related to the uptake of O_3 through the stomatal leaf pores (stomatal flux) than to the concentration in the atmosphere around the plants (Mills et al., 2011b). Stomata are physiologically controlled

3. Further, the LRTAP mapping manual makes it clear that the so-called PODYSPEC metrics (including POD6SPEC) are intended for situations where ozone and meteorological variables can be accurately estimated at the flag leaf of a wheat plant.

Ozone and meteorological variables are indeed estimated and referenced at canopy height and flag leaves for wheat fields. This was accounted for by applying a resistive network and the big-leaf approach, as done in DO3SE and also described in detail in our previous paper (Guaita et al., 2023).

Global scale model simulations are not at all well suited to making accurate predictions of POD6. Indeed, the LRTAP manual suggests that large-scale simulations make use of the a lower Y threshold, and some simpler parameter settings, which they denoted POD3IAM.

We disagree, as the LRTAP manual explicitly advocates for using our approach. Specifically, text Box 9 on page 45 (reported below, LRTAP Convention, 2017) explicitly recommends using PODYSPEC for climate change contexts: "For applications in a climate change context, the PODYSPEC method is recommended as key factors such as phenology and soil moisture are not included in the parameterization of PODYIAM."

Text Box 9: Applications for vegetation-type flux models and critical levels, POD_YIAM

These flux models have simpler form than POD_YSPEC and have been developed specifically for use in large-scale integrated assessment modelling, including for scenario analysis and optimisation runs. Separate parametrisations are provided for Mediterranean and non-Mediterranean areas for application in risk assessments for crops, forest trees and (semi-)natural vegetation.

The flux-effect relationships can be used for:

- Crops: potential maximum yield loss calculation and indicative economic losses in worst case scenario;
- Forest trees and (semi-)natural vegetation: indicative of the potential maximum risk for estimating environmental cost, but not economic losses.

The critical levels can be used for calculating critical levels exceedances, both amount and area. For applications in a climate change context, the POD_YSPEC method is recommended as key factors such as phenology and soil moisture are not included in the parameterisation of POD_YIAM.

The PODYIAM model is indeed a simplification of more detailed flux models (such as those based on PODYSPEC), suitable only when the data required for PODYSPEC application is unavailable. Additionally, PODYIAM does not specify any particular plant species, unlike our study, that is specifically designed to target wheat.

Finally, we would like to highlight that PODYSPEC for wheat and PODYIAM for crops have precisely the same parameterization (see the table reported below, LRTAP Convention, 2017),

with the sole exception that PODYIAM does not account for soil moisture or phenology—factors that are indeed essential in the context of climate change.

OD3IAM definition						POD6SPEC definition						
							Parameter	Units	C	rop species p	arameterisa	
able III.15: Parameterisation of the DO ₃ SE model for POD _V IAM calculat. forests and (semi-) natural vegetation. Separate parameterisations							Region (may also be applicable in these regions)		Atlantic, Boreal, Continental (Pannonian, Stappic)	Mediterranean	Mediterranean	
Parameter	Units	Crop parameterisation POD₃IAM		Fore		Currier	Common name	(Bread) Wheat	(Bread) Wheat	(Durum) Wheat		
Biogeographic region		Atlantic, Boreal, Continental, Steppic, Pannonian		Mediterranean	Ai Cont		Species	Latin name	Triticum aestivum	Triticum aestivum	Triticum durum	
Based on species			Wheat		Wheat	Beech		g _{max}	mmol O ₃ m ⁻²	500	430	410
gmax	mmol O ₃ m ⁻² PLA s ⁻¹		500		430			f .	fraction	0.01	0.01	0.01
T _{min}	raction		0.0105		0.0105		1	light a	-	0.0105	0.0105	0.0105
Tmin	°C		12		13	<u> </u>	1	Tmin	°C	12	12	11
Topt	°C		26		28		L	Topt	°C	26	28	28
T _{max}	°c		40		39			T _{max}	°C	40	39	45
VPD _{max}	kPa		1.2		3.2		1	VPDmax	kPa	1.2	3.2	3.1
VPD _{min}	kPa		3.2		4.6		1	VPDmin	kPa	3.2	4.6	4.9
ΣVPD _{crit}	kPa		8		8		1	ΣVPD _{crit}	kPa	8	16	16
PAWt	%		f _{sw} = 1		f _{sw} = 1		1	PAWti	%	50	-	-
SWCmax	% volume	I	-		-	1	1	SWC _{max} i	% volume	-	18.6	18.0

4. For these reasons the global scale POD assessments of Mills et al. (2018b,c) made use of POD3IAM metric. And although neither of the Mills papers was able to evaluate even this POD3 metric globally, they did show that the EMEP chemical transport model that was used was able to satisfactorily reproduce some basic statistics, namely mean of daily maximum ozone and M7, at sites from around the globe (Mills et al., 2018b, SI), and that model had been extensively tested against field data relevant to ozone deposition and fluxes.

Mills et al. (2018a, b) assessed metrics relevant to TOAR-I. In the current TOAR-II framework, the Ozone Deposition Focus working group focuses on metrics such as PODY, as will be seen in other papers that will be submitted to the TOAR-II special issue.

Regarding the reviewer's claim that the models were extensively tested against field data relevant to ozone deposition and fluxes, this deserves further discussion. As we understand it, the EMEP model reproduces ozone fluxes by coupling with the DO3SE model (Simpson et al., 2012). If, as the reviewer suggests, this model reliably reproduces ozone deposition and fluxes, our deposition model—which is based on the DO3SE—should also be capable of doing so, and this was demonstrated even at fine temporal resolution as reported in the previously cited comparison exercise described in the supplementary materials.

To our knowledge, no published studies directly compare the ozone fluxes or POD metrics predicted by the DO3SE model against direct measurements of ozone fluxes in crop fields. If the reviewer is aware of such studies, we welcome references.

For instance, Mills et al. (2018b) compared satellite-estimates of evapotranspiration (i.e. latent heat flux, LE) in the US with the modelled POD3IAM, under the assumption that both POD3IAM and evapotranspiration are driven by stomatal conductance. However, their approach presents clear limitations: (i) there was no discussion on the temporal resolution of the satellite measurements; (ii) assuming 10% wheat cover in a grid cell does not ensure representativeness of the grid cell's ET for wheat's water exchange; (iii) using three-year averages excessively smooths and ease the comparison; (iv) the comparison was indirect, making claims of validation difficult. Therefore, we decided to directly validate our model against flux measurements obtained over a wheat field (the aim of this manuscript) with the eddy covariance technique and show that our model satisfactorily represents LE fluxes on a half-hourly basis. Please refer to the supplementary materials for detailed analyses and results.

5. The usual problems of accurately modeling O3 and its metrics are exacerbated when climate models are used. In this case the meteorology is not constrained by reanalysis, and hence diverges more from the real-world than usually seen in current day chemical transport models. So, how well can your models predict O3, M7, and AOT40 for example (ie the metrics which can be derived from global observations), and indeed the hourly frequency distribution of O3?

The reviewer suggests that meteorology might be distorted in climate models due to the lack of reanalysis constraints, potentially affecting the prediction of photochemical ozone production. While this could be theoretically valid, CMIP climate models undergo extensive testing and continuous refinement to address meteorological distortions, both in terms of biases and in terms of trends. At each stage of CMIP model development, evaluations of key meteorological variables (including biases, trends, and seasonal variability) is routinely conducted. These assessments are consistently reported in the literature associated with CMIP model documentation, to which we refer the reviewer (e.g., Dunne et al., 2020; Horowitz et al., 2020; Sellar et al., 2019). Furthermore, it is important to mention that in order to make future projections of ozone damage, it is not possible to use reanalysis products, but climate models are the only possibility, even with the obvious uncertainties attached.

In the supplementary material attached to this reply we present an evaluation of the O_3 bias only for the daylight hours and over the accumulation period for POD (please refer to it), and show that the O3 concentration bias is small on average, and that, even for the regions where the bias is larger (such as East Asia), the effect on the POD₆ is relatively small.

6. A related issue is also that this paper seems to use quite short slices of meteorology. The base simulation is for 15 years (2000-2014), and the climate runs seem to be for 10 years (though I am a little confused by the 10 year slices given on L134 and the 15-year slice mentioned on L116). With short time-slices there is an increased risk that changes seen are due to random variations rather than to a true climate signal. Even with 20 year time-slices Langner et al. (2012) showed that the changes seen in summertime ozone were not significant at the 95% level over large parts of Europe.

The 10-year time slices used for projecting variable fields in 2050 and 2100 are an established approach in climate change studies, especially for air quality studies, as supported by several papers (e.g., Griffiths et al., 2020; Ronan et al., 2020; Sellar et al., 2019; Turnock et al., 2020). Furthermore, a 10-years slice seems adequate to capture the rate of change of precursor emissions. A 30-years average, while being a conventional averaging period for climate, will not be appropriate to reflect the rapid changes in ozone precursor emissions. In any case, the interpretation of the regional POD changes, the shown spatial and temporal trends for POD₆ over the century (Figure 4), and the statistical significance of the results (Table 4), indicate that a signal exists and that the shown trends reflect real tendencies rather than random fluctuations.

In any case, following the comments in the review of Owen Cooper, we added specific pvalues and confidence intervals in the manuscript (section 3.3) where appropriate, in order to point out uncertain results. Furthermore, the supplementary materials will contain, in their final form, a table with p-values and 95% confidence intervals corresponding to table 4 in the manuscript, and a map of the p-value associated with the ANOVA (i.e. Figure 8).

7. In the manuscript here, there is no discussion of these key issues. Instead we are referred to Turnock et al. (2020) for information about model skill, but that paper states that "CMIP6 models consistently overestimate observed surface O3 concentrations across most regions and in most seasons by up to 16 ppb, with a large diversity in simulated values over Northern Hemisphere continental regions".

The "up to 16 ppb" value the reviewer mentions indicates the maximum mean bias, and thus does not apply universally across all stations. While CMIP6 models are found to generally overestimate O_3 (Turnock et al., 2020), this overestimation is neither uniform nor pervasive across all models, and moreover does not account for the deposition processes that we consider in our study (please see the new supplementary materials attached below). As mentioned above, an O_3 bias evaluation tailored to our case-study is provided now in the supplementary materials, where O_3 concentrations are found to be modelled adequately for the purposes of our study.

8. Given such issues with surface O3 and the modeling in general, I have no reason to believe that the POD6 values have satisfactory values, or that trends in this metric are any more reliable.

In conclusion, we emphasize that the limitations and uncertainties of our study were adequately reported and quantified in the manuscript, and that the evaluation requested by the reviewer does not affect the validity of our study nor its conclusions. Following the reviewer's request, we performed an additional evaluation of the O₃ concentrations simulated by the CMIP6 models, which represents an extension of was what performed in Turnock et al. (2020), and we also tested the performances of our deposition model in reproducing the half-hourly latent heat flux.

Regarding the choice of the POD₆ metric, we believe that the POD₆ metric is the most appropriate for climatic studies like ours, as supported by both the mapping manual (referenced by the reviewer, though) and the cited relevant literature.

In any case, we want to highlight that combining different approaches (i.e., exposure-based approaches in most earlier studies and dose-based approach in this study) in assessing O₃ damage on vegetation may better constrain the uncertainties in the modeled results, as has been demonstrated in ensemble modeling air-quality and climate studies in literature. Therefore, since existing dose-based studies of O₃ damage on vegetation are very rare, our study is an important contribution to the current knowledge, despite known model uncertainties mentioned above.

Other comments

p2, L48. Strange not to mention the Mills et al. papers here, or that of Van Dingenen et al., 2009; these papers both offer both global-scale assessments which involved a lot of work (including comparison with observations) and are widely cited.

Thanks for the suggestion. We read the papers and the citations were added.

p4, Table 1. This table should also include the thickness of the lowest model, as this is the important for deriving crop-height O3 concentrations.

Yes. The height of the lowest model level (cell centre) was added in the table

p5. on L116, we read that the baseline is calculated for 15 years, over 2000-2014. How many years are used for the 2100 simulations?

The POD6 was calculated every year from 2015 to 2100. Each individual year was compared to the baseline average. To avoid possible confusion, the beginning of that sentence has been rephrased as: "Yearly O_3 risk for wheat cultivation from 2015 to 2100 is quantified with respect to a POD baseline value,..."

p5, L139. What does "Contextually" mean here? It sounds odd.

Thanks. "Contextually" was substituted with "moreover".

P5. Sect.2.2 The text suggests that wilting point and field capacity are needed, and derived as volumetric soil moisture (VSM) values. One issue is that the ESMs will have their own systems for dealing with soil water, and their calculations of near-surface ozone and resistances in general will presumably reflect their interpretation of soil-water effects. Possibly more serious is the use of volumetric soil water (VSM). The same VSM can represent very wet conditions in some soils, but very dry in others, but as far as I can tell the methods don't distinguish between different soils at all.

Indeed, the ESMs do represent soil water and offer the corresponding output. However, the soil water values from the ESMs at a given node are influenced by various land covers and land uses within both the grid node itself and its surrounding areas, and for this reason they are not intended to be referred to a single wheat field, but rather to the average soil water content of a large area. On the contrary, we wanted to simulate the soil water content only in a wheat field, without taking into account other land covers. Therefore our deposition model uses an online water soil module, following the approach of Mintz and Walker (1993). This is described in detail in our previous paper (Guaita et al., 2023) that we refer to.

Further, we note that different soil types have different associated wilting point and field capacity values. Therefore, these parameters effectively distinguish between soil types (please, see the dataset referenced in the manuscript, Zhang et al., 2018).

p6. The text here omits any mention of the difference between leaf and canopy scale resistances, but this is a key part of the DO3SE methodology (e.g. Tuovinen et al., 2009)

Our model uses the same approach as the DO3SE methodology. This was thoroughly discussed in the paper Guaita et al., 2023, which is frequently cited across the manuscript (Please see eq. 46-52 in Guaita et al., 2023).

According to the reviewer's remark, in the revised paper we clarify the difference between the bulk resistances (upcase R) used to scale the ozone concentration from the lowest model level to the canopy height, and the leaf-level resistances (lowcase r), which are used to calculate the ozone stomatal flux for an upper canopy leaf as prescribed by the DO3SE methodology.

p7, L173. Again the word contextually is used. It fits better here than in the above example, but I think it is better to say "In the context of...".

Thanks, the sentence was modified as suggested.

p7, L174. The word parameterizations is a bit vague, and readers cannot be expected to know what this means. Please make a table with the parameter values.

Ok, we have now included a table in the supplementary materials which we refer to in the main manuscript.

p7, L184. In what way are the Feng et al. (2012) parameterizations incomplete? I would have thought that methods developed from China were more appropriate for global approaches than those from Spain.

The parameterization proposed by Feng et al. (2012) did not consider limitations to stomatal conductance from temperature and soil water content. Since these variables are important in a climate change context, we preferred to exclude this parameterization in our study. In fact, Feng et al. (2012) adopted the following formulation for the Jarvis model:

2.3. The multiplicative stomatal conductance model



Since temperature and soil moisture are important in a climate change context, in our study we preferred to exclude the parameterization of Feng et al. (2012). On the contrary, the parameterization from González-Fernández et al. (2013) includes the effect of soil water and temperature as well, and was based on field measurements datasets.

p7, L195 "A well-established dose-response relationship...". Is this so well established? The mapping manual states "the percentage effect due to O3 impact on crop yield estimated in large-scale modeling should be calculated as follows:

(PODYIAM – Ref10 PODYIAM) * (% reduction per mmol/m2 PODYIAM.POD3)

And indeed, Mills et al (2018c) used:

RYL = (POD3IAM-0.1)*0.64

but this manuscript uses POD6 rather than the recommended POD3IAM for unexplained reasons, and makes no mention of the "Ref10" correction.

We removed the adjective "well-established". However, as discussed in the major comments above, POD6 is indeed the right approach in our case (LRTAP Convention, 2017, pag. 45). Following the mapping manual, the relative yield loss in this case is given by:

RYL = (PODYSPEC - Ref10 PODYSPEC) * % reduction per mmol m-2 PODYSPEC

However, the Ref10 PODYSPEC is null in the case of wheat (see the Table III.10 of the Mapping Manual, LRTAP Convention, 2017, reported below), and therefore it was omitted.

Species	Effect parameter	Biogeo- graphi- cal region*	Potential effect at CL (% reduction)	Critical level (mmol m ⁻² PLA)**	Ref10 POD ₆ (mmol m ⁻² PLA)	Potential maximum rate of reduction (%) per mmor m ⁻² PLA of POD ₆ SPEC***
Wheat	Grain yield	A,B,C,M (S,P)****	5%	1.3	0.0	3.85
Wheat	1000-grain weight	A,B,C,M (S,P)****	5%	1.5	0.0	3.35
Wheat	Protein yield	A,B,C,M (S,P)****	5%	2.0	0.0	2.54
Potato	Tuber yield	A,B,C (M,S,P)	5%	3.8	0.0	1.34
Tomato	Fruit yield	M (A,B, C,S,P)	5%	2.0	0.0	2.53
Tomato	Fruit quality	M (A,B, C,S,P)	5%	3.8	0.0	1.30

Table III.10: POD₆SPEC critical levels (CL) for crops.

P7, L202—204. This sentence seems out of place compared to the preceding text.

Yes. The sentence was moved to the Appendix B, and it was referenced within the section 2.3.

p24, L456. Given all the uncertainties I mentioned in the major comments section, I wonder what the phrase "Our results may be associated with different degrees of confidence depending on the agreement between the two available CMIP6 models," means? The paper has barely mentioned the main sources of uncertainty I think.

Different physics-based models displaying similar features suggest that their results are associated with a greater degree of confidence, as opposed to the two models disagreeing. Of course, considering more models would improve the study, but the UKESM1-0-LL and GFDL-ESM4 remain to date the only models within CMIP6 that are suitable for our study. This is now clarified in the manuscript.

As shown in the newly written supplementary materials, the effect of the O3 bias on POD6 is found to be negligible for average climatic conditions in two of the main areas impacted by O3 risk at the present time (North America and Europe). As for the other region (East Asia), the O_3 bias could lead to an 0.28-0.33 mmol m⁻² overestimation of POD₆ on average, which however would not be enough to classify this region as not at risk, and therefore such possible overestimation does not affect our main conclusions. A discussion on this has been added to the discussion section.

p42, L890 UKESM - 20m. Is that cell depth, or cell-center?

It's the cell-center. Thanks. This has been added to the Appendix.

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Supplement of Global flux-based ozone risk assessment for wheat up to 2100 under different climate scenarios

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1. O3 mean bias and its effect on POD6

The bias of the O_3 concentrations (Table S1, Figure S1) taken as input data by our dry deposition model are evaluated against the surface O_3 observations taken from the TOAR-II dataset (Schröder et al., 2021) for the baseline years (2000-2014). For each measurement station, the mean bias (MB) is defined as:

$$MB(x) = \frac{\sum_{t=1}^{n} M(x,t) - O(x,t)}{n}$$
(S1)

where x indicates the location, and t = 1, ..., n are the considered timesteps (daylight hours, over the accumulation period for that location). M(x, t) is the O₃ output from the considered CMIP6 models (sfo3 in CMIP6 notation, indicating the O₃ at the lowest model level), scaled at the canopy height by means of the dry deposition scheme used in this study (Guaita et al., 2023), and O(x, t) is the observed surface O₃ concentration (also scaled to canopy height). The ground stations are selected to be representative of the agricultural context, and therefore only rural, background or crop locations are included. Since the dry deposition model of this study scales O₃ from the height of the lowest model level of the CMIP6 models to the crop canopy height, only ground stations with measurements height below 3 meters are considered. Furthermore, only stations with more than 10 years long timeseries are included. In this regard, an exception is made for China, as the stations that could represent wheat fields have timeseries in that region are shorter than 10 years.

To estimate the effect of the O₃ bias on the O₃ risk, the yearly POD₆ was regressed (with interaction) against O₃ concentration and f_{clim} averaged over the accumulation period:

$$POD_{6} = \beta_{0} + \beta_{1}[O_{3}] + \beta_{2}f_{clim} + \beta_{3}[O_{3}]f_{clim} + \epsilon$$
(S2)

where β_i (i = 0, ..., 3) are the regression coefficients, and $\epsilon \in N(0, \sigma^2)$ is the error. The regression is calibrated over the whole globe, using both the baseline and the three scenarios considered in this study (SSP1-2.6, SSP3-7.0, SSP5-8.5). In the eq. (S2), $\beta_1 + \beta_3 \cdot f_{clim}$ is the mean POD₆ increment per ppb given a certain value of f_{clim} . Therefore, the expression ($\beta_1 + \beta_3 \cdot f_{clim}$) \cdot *MB* represents the effect of the O₃ MB on the POD₆. In other words, when f_{clim} is equal to the regional (or global) average ($\mu(f_{clim})$), the resulting value is the average effect of the O₃ MB on the POD₆ over the specified region under mean f_{clim} conditions, while, when $f_{clim} = 1$, the resulting value is the average MB effect on POD₆ under climatic conditions optimal to the stomatal conductance (Table S1).

Table S1: MB of O₃ at canopy height, POD₆ increments per ppb of O₃ concentration conditional to f_{clim} equal to its regional mean $(\beta_1 + \beta_3 \cdot \mu(f_{clim}))$, and to $f_{clim} = 1$ $(\beta_1 + \beta_3 \cdot 1)$, and the corresponding projected effect on POD₆ for average climatic conditions and optimal conditions to the stomatal conductance.

Region	# station	MB ± SD [ppb]	$\beta_1 + \beta_3 \cdot \mu(f_{clim})$ (95%CI) [mmol m ⁻² /ppb]	$\beta_1 + \beta_3 \cdot 1$ (95%CI) [mmol m ⁻² /ppb]	$(\beta_1 + \beta_3 \cdot \mu(f_{clim})) \cdot MB \pm SD$ [mmol m ⁻²]	$(\beta_1 + \beta_3 \cdot 1) \cdot MB$ $\pm \text{SD}$ [mmol m ⁻²]
GFDL-ESM4	4					
Global	1330	1.31 ± 7.36	0.0154 (0.0153,0.0155)	0.0582 (0.0581,0.0584)	0.02±0.03	0.08 ± 0.37
N. America	579	$\textbf{-2.8} \pm \textbf{4.96}$	0.0188 (0.0186,0.0189)	0.0493 (0.049,0.0496)	-0.05 ± 0.02	-0.14±0.14
Europe	480	2.09 ± 5.59	0.0152 (0.015,0.0154)	0.0573 (0.0567,0.0579)	0.03 ± 0.02	$0.12{\pm}0.22$
East Asia	283	8.02 ± 8.53	0.0351 (0.0348,0.0354)	0.0913 (0.0907,0.0918)	0.28±0.26	0.73 ± 1.76
UKESM1-0-	LL					
Global	1350	0.88 ± 7.61	0.0316 (0.0313,0.0318)	0.1192 (0.1186,0.1198)	0.03±0.12	0.1±1.67
N. America	579	$\textbf{-3.27} \pm 5.91$	0.0441 (0.0436,0.0446)	0.1289 (0.1276,0.1302)	-0.14±0.16	-0.42±1.35
Europe	486	3.13 ± 5.58	0.0401 (0.0393,0.0409)	0.1614 (0.1579,0.165)	0.13±0.12	0.51±1.93
East Asia	276	5.59 ± 9.34	0.0588 (0.058,0.0597)	0.164 (0.162,0.1659)	0.33±0.72	0.92±5.61



O₃ MB (UKESM1-0-LL(hc) – TOAR)



Figure S1: MB of O₃ at canopy height for GFDL-ESM4 (a), and UKESM1-0-LL (b), obtained by scaling O₃ in output from the CMIP6 models (at the lowest model level height), to the canopy height, by means of the resistive network of the dry deposition model. O₃ is compared with ground measurements from the TOAR-II database.

2. Evaluation of the performance of the dry deposition model used for POD₆ calculations

The dry deposition model used in this study (Guaita et al., 2023) is compared to the ozone flux measurements made on a wheat field wheat in Comun Nuovo (Italy) (Gerosa et al., 2003).

The model is tested for its capability to reproduce the total O_3 flux (F_{O3}) over the wheat field which corresponds to testing the resistance network altogether, and the latent heat flux (LE, W/m²) which is a proxy for stomatal conductance (and therefore a proxy for stomatal resistance). Namely, the total stomatal flux is calculated as follows:

$$F_{03} = \frac{O_3(z_{m03})}{R_{aH}(d+z_{0m}) + R_{b03} + R_{surf,03}}$$
(S3)

where z_{m03} is the measurement height, d is the displacement height, z_{0m} is the roughness length for momentum, $O_3(z_{m03})$ is the ozone concentration at measured height, $R_{aH}(d + z_{0m}, z_{m03})$ is the aerodynamic resistance between $d + z_{0m}$ and z_{m03} , R_{b03} is the quasi-laminar resistance, and $R_{surf,03}$ is the bulk overall surface resistance to deposition. For details on the calculation for each of these variables, see the Appendix A in Guaita et al. (2023). Figure S2 shows the timeseries of F_{O3} and LE from the beginning of flowering to the harvest. Table S2 shows statistics for the model performance over the daylight hours (6 am-6 pm). The reported values are obtained by simply regressing the modelled values against the observed ones.



Figure S2: Timeseries for modelled and observed $F_{\rm O3}$ (a), and LE (b)

Table S2: Statistics for the comparison between observed and measured F ₀₃ , and LE. The regression uses the	ıe
modelled data as predictand and the observations as predictor.	

	$F_{O3} (nmol m^{-2} s^{-1})$	LE (W m ⁻²)
Regression intercept (p-	4.3377 (1.8448e-24)	-0.0209 (3.5875e-14)
value)		
Regression slope (p-value)	1.0481 (1.3058e-101)	1.1438 (1.0069e-197)
Mean Bias	3.3298	-0.0089
Mean Absolute Error	4.0317	0.0200
Root Mean Square Error	5.8807	0.0308
R-squared	0.4965	0.7403