1 **Isomer Molecular Structures and Formation Pathways of Oxygenated Organic Molecules**
2 **in Newly Formed Biogenic Particles**

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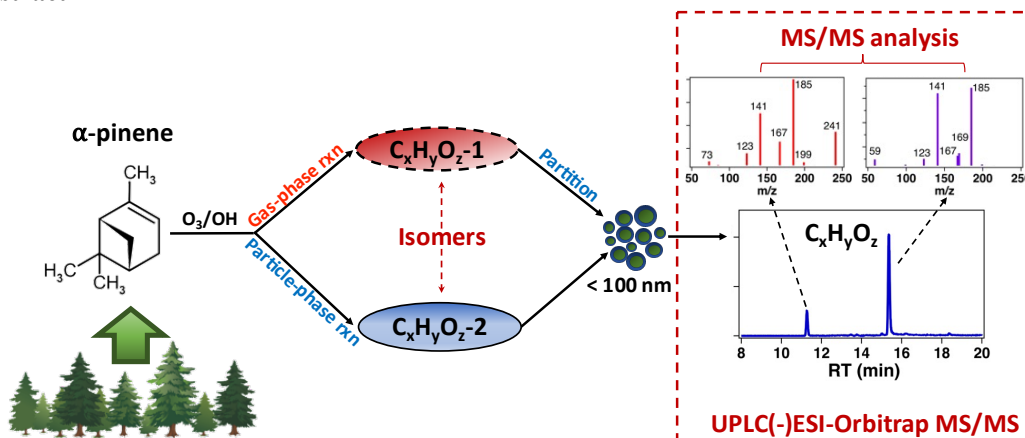15 States

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19

20 **Abstract**

23 Oxygenated organic molecules (OOMs) formed from oxidation of anthropogenic and biogenic
24 volatile organic compounds (VOCs) are essential ingredients for atmospheric new particle
25 formation (NPF) and secondary organic aerosol (SOA) formation, and thus impact air quality,
26 human health, and climate. There is a large variety of OOM compounds, but currently, for the vast
27 majority of OOMs, their molecular structures and formation pathways are still unknown. In this

28 study, we identified isomer-resolved molecular structures and reaction pathways for dimer OOMs
29 formed from α-pinene ozonolysis, using an ultrahigh-performance liquid chromatography-
30 electrospray ionization Orbitrap mass spectrometer (UPLC/(-)ESI-Orbitrap MS) tandem analysis
31 and a high-resolution time-of-flight chemical ionization mass spectrometer (HrTOF-CIMS)
32 attached to the filter inlet for gas and aerosol (FIGAERO), combined with explicit chemical
33 modeling simulations using the Generator of Explicit Chemistry and Kinetics of Organics in the
34 Atmosphere (GECKO-A). In general, each OOM identified in the newly formed biogenic particles
35 contains 2-8 isomers with distinctive MS/MS fragmentation ions. For $C_{19}H_{30}O_5$, which is one of
36 most abundant dimers identified from the boreal forests and laboratory biogenic NPF studies, one
37 isomer forms in the gas phase from a stabilized Crigee Intermediate (sCI) peroxy biradical and
38 aldehyde, followed by subsequent gas-to-particle conversion; and another isomer forms in the
39 particle phase via the Baeyer-Villiger reaction from a cyclic acylperoxyhemiacetal and α-
40 pinanediol. Two isomers of $C_{16}H_{26}O_6$ form in the particle phase via decarboxylation from two
41 different isomers of $C_{17}H_{26}O_8$ after the condensation from the gas phase. Thus, our results show
42 that biogenic OOMs can also form from particle-phase reactions and have different isomeric
43 structures than in the gas phase. Our study represents the first molecular-level chemical analysis
44 to identify particle-formation pathways for OOMs in the newly formed biogenic nanoparticles.
45 Currently, parameterizations of NPF (e.g., biogenic NPF) are based on the gas-to-particle
46 conversion of extremely low-volatility OOM dimers that form in the gas phase alone (e.g., via RO_2
47 + RO_2 reactions). Our study demonstrates that additional, independent particle-phase formation
48 pathways should also be considered for predictions of the formation and growth of new particles
49 in the atmosphere.

50 **Keywords**. OOMs, particle-phase reaction, gas-to-particle conversion, new particle formation,
51 isomer, MS/MS tandem analysis

52 **Synopsis**: This study highlights the importance of particle-phase reactions that are currently not
53 considered in the new particle formation processes.

54

1.   Introduction

55

56   Oxygenated organic molecules (OOMs) formed from oxidation reactions of biogenic volatile
57   organic compounds (BVOCs) can contribute to secondary organic aerosol (SOA) and new particle
58   formation (NPF) processes [*Barsanti et al.*, 2017; *Bianchi et al.*, 2019; *Hallquist et al.*, 2009; *S-H*
59   *Lee et al.*, 2019; *Ziemann and Atkinson*, 2012]. In the atmosphere, BVOCs are oxidized by ozone,
60   hydroxy (OH), and nitrate (NO$_3$) radicals. There are many OOMs in the atmosphere with different
61   chemical identities (e.g., chemical formulas and isomeric structures). For example, from
62   ozonolysis of α-pinene alone, thousands of OOMs were detected with different mass-to-charge
63   ratios (m/z) with a high-resolution time-of-flight chemical ionization mass spectrometer (HrTOF-
64   CIMS), hence, indicating different chemical formulas (e.g., [*Caudillo et al.*, 2021; *Lee Tiszenkel*
65   *and Lee*, 2023]). Yet, only an extremely limited number of OOMs have been identified with their
66   isomer molecular structures and formation pathways, even for the α-pinene ozonolysis system,
67   which is the most-studied biogenic oxidation system in the atmospheric chemistry field. Previous
68   studies have made isomer structural analysis of biogenic OOMs mostly for SOA particles and
69   still, detailed molecular structures of OOMs in newly formed pure biogenic particles are not
70   known.

71   In the gas phase, dimeric OOMs form from reactions involving organic peroxy radicals (RO$_2$)
72   [*Peräkylä et al.*, 2023; *Tomaz et al.*, 2021; *Wang et al.*, 2023] and stabilized Criegee intermediates
73   (sCI) [*Y Zhao et al.*, 2015; *Y Zhao et al.*, 2022]. Particle-phase formation of dimeric OOMs
74   involves esterification [*Müller et al.*, 2008; *Yasmeen et al.*, 2010], Baeyer-Villiger reactions
75   [*Claflin et al.*, 2018; *Kenseth et al.*, 2023; *Pospisilova et al.*, 2020], aldol condensation [*Tolocka*
76   *et al.*, 2004; *Witkowski and Gierczak*, 2014; *Ziemann and Atkinson*, 2012], diacyl decomposition
77   [*Zhang et al.*, 2015], and acyl trioxide decomposition [*Kahnt et al.*, 2018]. Isomer-resolved
78   characterization of these OOMs was made possible using liquid chromatographic techniques; and
79   high-resolution mass spectrometer coupled with tandem MS/MS analysis enabled the elucidation
80   of molecular structures of BVOC oxidation products [*Nozière et al.*, 2015]. For α-pinene oxidation
81   SOA, the dimer esters, C$_{17}$H$_{26}$O$_8$ (*m/z* = 358) and C$_{19}$H$_{28}$O$_7$ (*m/z* = 368), have been structurally
82   resolved using a linear ion trap mass spectrometer and it was shown that they form from acyl
83   trioxide decomposition of a gas-phase product dimer C$_{19}$H$_{28}$O$_{11}$ by *Kahnt et al.* [2018]. *Zhang et*
84   *al.* [2015] proposed the structure of the dimers in the particle phase using electrospray ionization-
85   quadrupole time of flight mass spectrometer (ESI-QTOF-MS), of which C$_{17}$H$_{26}$O$_6$ (m/z = 309)
86   forms from decarboxylation reaction. *Kristensen et al.* [2016] identified the structure of specific
87   dimer esters from α-pinene ozonolysis experiments, which were also identified from the boreal
88   forest ambient samples identified using ESI-QTOF-MS. *Kenseth et al.* [2023] illustrated the
89   particle-phase accretion reaction and subsequent formation of C$_{19}$H$_{30}$O$_5$ (*m/z* = 337) through
90   Baeyer-Villiger decomposition reaction. In addition to the α-pinene oxidation SOA, MS/MS
91   analysis has also been used to study other monoterpene oxidation products such as limonene-
92   derived products using online-nitrate chemical ionization Orbitrap MS/MS by *Tomaz et al.* [2021],
93   β-pinene-derived products using liquid chromatography (LC) coupled with ESI-QTOF-MS
94   [*Kenseth et al.*, 2018] and molecular analysis of SOA derived from monoterpenes such as α-pinene,
95   β-pinene, limonene, 3-carene and sabinene using LC coupled with ion trap mass spectrometer by
96   *Glasius et al.* [2000].

97 The current understanding of the NPF processes is based on the gas-to-particle conversion of low-
98 volatility chemical precursors [*S-H Lee et al.*, 2019]. It is assumed that OOMs form in the gas
99 phase, and the OOMs that have sufficiently low volatilities undergo gas-to-particle conversion to
100 contribute to the formation and growth of new particles. Therefore, particle phase formation of
101 OOMs is not considered in the current NPF theories and parameterizations. Typically, the volatility
102 (or saturation vapor concentration at a specific temperature) of an OOM moiety is estimated with
103 the Volatility Basis Set (VBS) parameterization [*Donahue et al.*, 2011], where the elemental
104 composition of C, O, and N dictates saturation vapor concentration [*Ye et al.*, 2019]. Volatilities
105 can be also derived from thermogram measurements using a filter inlet for gas and aerosols
106 (FIGAERO) attached to a high-resolution time-of-flight chemical ionization mass spectrometer
107 (HrTOF-CIMS) [*Lopez-Hilfiker et al.*, 2014]. Laboratory experiments have shown that the
108 volatilities derived from these two methods have systematic discrepancies for OOMs formed from
109 ozonolysis of α-pinene (e.g., VBS over-estimates volatilities for monomers and under-estimates
110 for dimers compared to FIGAERO) [*Lee Tiszenkel and Lee*, 2023].

111 In this work, we have analyzed the molecular composition of OOMs formed from the ozonolysis
112 of α-pinene using an ultra-performance liquid chromatography-electrospray ionization Orbitrap
113 mass spectrometry (UPLC/(-)ESI-Orbitrap MS) and HrTOF-CIMS coupled with a FIGAERO. The
114 chemical composition of OOMs in the gas and aerosol phase was measured simultaneously with
115 FIGAERO HrTOF-CIMS. The liquid chromatogram separates isomers of OOMs, and the most
116 likely molecular structures of an isomer were derived based on the fragmentation ions detected
117 with high-resolution Orbitrap MS/MS analysis. To help verify the molecular structures of OOMs
118 in the gas phase, we also performed simulations with the chemically explicit model, the Generator
119 of Explicit Chemistry and Kinetics of Organics in the Atmosphere (GECKO-A) ([*Aumont et al.*,
120 2005; *Camredon et al.*, 2007; *Valorso et al.*, 2011]; updated per [*Jenkin et al.*, 2020; *Jenkin et al.*,
121 2019; *Jenkin et al.*, 2018]). Based on these analyses, we propose the isomeric structures of two
122 dimer OOMs ($C_{19}H_{30}O_5$ and $C_{16}H_{26}O_6$) and their potential formation pathways in the gas and
123 aerosol phase.

2.   Materials and Methods

125 Experiments were carried out in the Tandem Aerosol Nucleation and Growth Environment Tube
126 (TANGENT) setup. A detailed description of this experimental setup has been published
127 previously [*Lee Tiszenkel and Lee*, 2023; *L. Tiszenkel et al.*, 2019]. Briefly, 240 ppb α-pinene in
128 nitrogen was mixed with 1200 ppb O_3 at room temperature (298 K) and dry conditions (RH <
129 10%). The total residence time in the flow tube was 150 s. The OH radicals present in the
130 experimental system were the byproduct of the monoterpene ozonolysis reaction. Experiments
131 were done without an OH scavenger and without seed aerosols. A chemical box model using the
132 Master Chemical Mechanism (MCM v3.3.1) [*Jenkin et al.*, 2015] was used to estimate the OH
133 radical concentration in the experimental setup at 1.6 ppt. Aerosol size distributions from 1-80 nm
134 were measured by combining data from a particle size magnifier (PSM; A10, Airmodus) and a
135 scanning mobility particle sizer (SMPS; Model 3080 electrostatic classifier, model 3776
136 condensation particle counter CPC, and model 3085 differential mobility analyzer DMA, TSI Inc.).

2.1.   Filter Inlet for Gases and Aerosols on High Resolution Time of Flight Chemical
138   Ionization Mass Spectrometer (FIGAERO Hr-TOF CIMS)

139   The chemical composition of OOMs in the gas and aerosol phase was measured online by an iodide
140   HrTOF-CIMS attached to FIGAERO [*B H Lee et al.*, 2014] (Aerodyne Inc.). The HrTOF has a
141   mass resolution (m/Δm) of 7,000. During gas-phase sampling, the FIGAERO drew 2 LPM from
142   the flow tube through a 0.64 cm OD stainless steel line onto a Teflon filter (Zerflour® 24 mm
143   diameter, 2 μm pore size, Pall Corp) for 20 minutes of particle collection. The FIGAERO then
144   directed the filter into a flow of dry ultra-high purity (UHP) nitrogen for thermal desorption of the
145   collected particles. The nitrogen was heated from room temperature at a rate of 35 °C per minute
146   for 5 minutes to a final temperature of 200 °C. The filter was then heat soaked with a flow of 200
147   °C nitrogen for an additional 15 minutes. The maximum vaporization temperature (T_{max}) during
148   thermal desorption was used to derive the saturation vapor concentration (C*) of the OOMs based
149   on calibration of the FIGAERO with polyethylene glycols with known C* [*Lee Tiszenkel and Lee*,
150   2023]. The FIGAERO T_{max} is subject to variability depending on the filter mass loading and
151   thermally driven particle-phase chemistry [*Huang et al.*, 2020]. Gas-phase concentrations of
152   species measured with the HrTOF-CIMS were calculated by assuming that OOM sensitivity was
153   equivalent to the formic acid sensitivity. This assumption is reasonable to within a factor of 1.5
154   [*Lopez-Hilfiker et al.*, 2015]. Calibration of the HrTOF-CIMS using a formic acid permeation
155   source (Kin-Tek) yielded a sensitivity of 6.7 Hz ppt^{-1} MHz^{-1}. The detection limit (defined as the
156   signal with a S/N ratio of 3) was 4.3 pptv. Uncertainties in this calibration factor, taking into
157   account errors in flow meters, permeation sources, and calibration, were calculated as 16%.

2.2.   Filter sample collection and extraction

159   Particles were collected on Whatman® glass microfiber filters (Grade GF/D, 2.7 μm pore size,
160   25mm diameter), which were pre-baked at 550 °C for 24 hours to remove any residual organics.
161   The particles were collected continuously for 28 hours with a 1 L/min flow through the filter from
162   the experimental tube to generate approximately 500 μg of aerosol samples. Filters were moved to
163   borosilicate glass vials for storage at -20 °C immediately after collection to prevent any significant
164   changes to the OOM chemical composition and their abundance [*Resch et al.*, 2023]. Three
165   replicate filters were collected from similar experimental conditions.

166   Filter samples were divided into quarters and each quarter was separately extracted by sonicating
167   for 30 minutes in borosilicate glass vials with 20 mL methanol (Optima™ LC/MS, Fisher
168   Scientific) placed in beakers filled with ice. After 30 minutes, the ice was replaced and the sample
169   was sonicated for additional 30 minutes. The sample was then dried down under a weak stream of
170   pure nitrogen gas in a room-temperature water bath. The dried sample was reconstituted in 2 mL
171   of methanol and filtered through a 0.2 um nylon syringe filter (ThermoFisher Scientific) into a
172   sampling vial for LC analysis. Previous studies have shown that organic peroxides decay relatively
173   slowly in 100% methanol [*R Zhao et al.*, 2018a; *Y Zhao et al.*, 2015; *Y Zhao et al.*, 2022].
174   Procedural blanks were prepared by subjecting filters that had undergone the pre-baking process
175   to the same extraction procedures.

2.3.   Ultrahigh-Performance   Liquid   Chromatography/Negative   Electrospray
Ionization Orbitrap Mass-Spectrometry (UPLC/(-)ESI-Orbitrap MS)

178   The filter samples were chromatographically separated and analyzed using ultrahigh-performance
179   liquid chromatography (UPLC) (Vanquish, ThermoFisher Scientific) coupled to an ultra-high
180   resolution Orbitrap mass spectrometer (Exploris 120 MS, ThermoFisher Scientific) with an

181 electrospray ionization (ESI) source. The orbitrap has a mass resolution (m/Δm) of 120,000 for
182 MS and 15,000 for MS/MS analysis. The separation of compounds including isomers was achieved
183 using a 100 mm x 2.1 mm reverse phase C-18 column with 1.8 um particle size (Waters, ACQUITY
184 Premier HSS T3). The column and autosampler temperatures were maintained at 40 °C and 4 °C ,
185 respectively. The polar (A) and non-polar (B) mobile phase solvents were 0.1% formic acid in
186 ultra-pure water (Optima™ LC/MS, Fisher Scientific) and 0.1% formic acid in MeOH (Optima™
187 LC/MS, Fisher Scientific) respectively. Gradient elution was performed starting with 1-minute
188 post-injection hold at 90% A and 10% B followed by a decrease to 10% A and 90% B over 26
189 minutes, then a return to 90% A and 10% B over 2 minutes, and ending with a 2 minute column
190 equilibration, as adopted from [*Shao et al.*, 2022]. The total flow rate was 0.3 mL min^{-1} with an
191 injection volume of 2 μL; these conditions were determined by optimizing the *cis*-pinic acid (*m/z*
192 185) signal for maximum peak area. The non-targeted mass spectrometric analysis was carried out
193 using optimized ESI parameters: 2.5 kV capillary voltage; 325 °C ion transfer tube temperature;
194 350 °C vaporizer temperature, and 50, 10, and 1 (arbitrary units) flow rates for sheath gas, auxiliary
195 gas and sweep gas, respectively. The parent molecules were deprotonated using (-)ESI and were
196 detected as [M-H]$^-$ ions, also protonated by (+)ESI, and were detected as [M+H]$^+$ and [M+Na]$^+$
197 ions. This study is mostly based on the negative ionization results. Unscheduled (RT not specified)
198 and non-targeted tandem mass spectrometric (MS/MS) analyses were performed using higher-
199 energy collisional dissociation (HCD) with a stepped normalized collision energy of 20%, 40%,
200 and 60% for those compounds detected with an ion intensity threshold of 5×10^3. Post-acquisition
201 data processing was carried out using a non-targeted compound identification method developed
202 in Compound Discoverer software (version 3.3 SP2, ThermoFisher Scientific). Additional details
203 are given in Supplementary Information.

204 Combining these two independent high-resolution mass spectrometer techniques, HrTOF-CIMS
205 and UPLC/(-)ESI-Orbitrap MS, provides a very powerful tool for analyzing the particle-phase
206 chemical composition of OOMs, for identifying molecular structures of different isomers. So far,
207 only a few studies have combined these two methods to make molecular-level chemical speciation
208 of OOMs (e.g., [*Du et al.*, 2022; *Huang et al.*, 2020; *Mehra et al.*, 2020]). It is noted that HrTOF-
209 CIMS (iodide ionization) and UPLC/(-)ESI-Orbitrap MS may have different ionization
210 efficiencies and detection efficiencies for different chemical compounds. Additionally, offline
211 filter sample analysis can introduce some artifacts during sampling collection, storage, and
212 extraction processes.

213 **2.4.   GECKO-A simulations**

214 To help verify the chemical reactants and products of a-pinene ozonolysis reactions in the gas
215 phase, as well as their molecular structures, we generated an explicit α-pinene degradation
216 mechanism using GECKO-A ([*Aumont et al.*, 2005; *Camredon et al.*, 2007; *Valorso et al.*, 2011];
217 updated per [*Jenkin et al.*, 2020; *Jenkin et al.*, 2019; *Jenkin et al.*, 2018]), and performed model
218 simulations using the box model. GECKO-A is a state-of-the-art automatic mechanism generator
219 designed to produce explicit oxidation degradation schemes for organic compounds. Detailed
220 mechanism description and simulation procedures used in this study are described in the
221 Supporting Information.

222

223 **3.   Results and discussions**

3.1. Overall molecular composition of oxygenated organic molecules in α-pinene ozonolysis system

Figure S1 shows the measured aerosol size distribution with a mean diameter of 70 nm during the typical experiment. We chose only the 50% most abundant OOMs in the present study. Figure 1 shows the total ion mass spectra and the mass defect plots the OOMs detected using LC-Orbitrap MS, FIGAERO-CIMS, and HrTOF-CIMS. In all mass spectra, clearly, monomers (C_5-C_{10}) and dimers (C_{15}-C_{20}) were resolved. In the gas phase, there were relatively high ion intensities for monomers than dimers, whereas relative ion intensities of dimers increased in the particle phase (as shown in both FIGAERO-CIMS and UPLC/(-)ESI-Orbitrap MS), indicating that dimers, in general, are more favorable for gas-to-particle conversion due to relatively lower volatilities, consistent with previous observations [*Lee Tiszenkel and Lee*, 2023]. Notably, the particle-phase OOM spectra taken by UPLC/(-)ESI-Orbitrap MS and FIGAERO-CIMS show strikingly similar patterns, despite the entirely different sampling (offline versus online) and ionization methods and potential artifacts introduced during the filter collection, storage, and extraction procedures. There were also many more small ions (C_{5-6}) detected with the FIGAERO-CIMS than in the LC-Orbitrap mass spectrometer, likely due to the thermal fragmentation of OOMs during the desorption cycle in the FIGAERO. As shown in Figure 1g, many OOMs detected by LC-Orbitrap MS showed 2-8 isomers (discussed in detail in Section 3.3), but these isomers were not resolved in FIGAERO-CIMS.

In total, 167 OOMs were identified with UPLC/(-)ESI-Orbitrap MS, 405 OOMs were identified in the particle phase with the FIGAERO -CIMS, and 437 OOMs were identified in the gas phase with the HrTOF-CIMS (Figure 1h). Of the 588 OOMs detected, 304 OOMs were found in both the gas and particle phases, suggesting that these are OOMs that likely underwent gas-to-particle conversion. However, there were 124 OOMs that were detected in the gas phase alone, and 151 OOMs that were detected only in the particle phase (by the FIGAERO-CIMS and UPLC/(-)ESI-Orbitrap MS). Further discussions of particle-phase OOMs will concern 77 compounds selected from the 108 detected in the particle-phase by the FIGAERO-CIMS and UPLC/(-)ESI-Orbitrap MS. Table S1 shows the chemical formula of OOMs detected in HrTOF-CIMS (gas and particle phases) and UPLC/(-)ESI-Orbitrap MS (particle phase). Table S2 lists these OOM compounds, along with their isomers defined by their distinct RT and tandem MS/MS fragmentation ions of each isomer. These compounds were selected by considering the reproducibility of their desorption thermograms.

3.2. FIGAERO measured versus VBS calculated saturation vapor concentration

Figure 2 shows the saturation vapor concentrations (C*) of the OOMs derived from FIGAERO thermogram measurements and VBS parameterization. FIGAERO measured saturation vapor concentrations (C*) classify the OOMs in the range between extremely-low-volatility organic compounds (ELVOC) and semi-volatile organic compounds (SVOC). For the same set of OOMs, VBS-derived saturation vapor concentrations have a wider range, from ultra-low-volatility organic compounds (ULVOC) to intermediate-volatility organic compounds (IVOC). Regardless, there was the same trend of decreasing saturation vapor concentration with increasing molecular weight for both methods (Figures 2c and S2). Some OOMs (e.g., $C_{10}H_{14}O_6$, $C_{10}H_{16}O_6$, $C_{11}H_{16}O_6$, $C_{12}H_{20}O_5$, $C_{14}H_{20}O_8$, $C_{15}H_{22}O_6$, $C_{15}H_{24}O_8$, $C_{15}H_{26}O_6$, $C_{17}H_{26}O_8$, $C_{17}H_{28}O_5$ and $C_{18}H_{30}O_7$) showed

266 a better agreement between FIGAERO measurements and VBS calculations. Remarkably, both
267 methods were able to replicate the "teeth-like" feature observed in measured saturation vapor
268 concentration with increasing molecular weight, oxygen, and carbon numbers, as predicted from
269 the VBS parameterizations.

270 However, there were general discrepancies in saturation vapor concentration from these two
271 methods. While volatilities derived with these two methods were more correlated to each in the
272 middle-number carbons (e.g., C_{10-15}), they deviated by several orders of magnitude at the smaller
273 and larger carbon numbers, consistent with our previous study [*Lee Tiszenkel and Lee*, 2023]. For
274 example, the VBS-calculated volatilities overestimate FIGAERO measurements by 4 orders of
275 magnitude for C5 and underestimate FIGAERO measurements by 4 orders of magnitude for C20
276 OOMs. FIGAERO measurements also show that the C_{16-17} compounds have the lowest measured
277 volatility, in contrast to the VBS calculations that estimate C_{20} OOMs are the least volatile. We
278 also find that the trend in measured volatilities for some series of OOMs deviated from the trend
279 expected with VBS estimation. For example, VBS estimated the decreasing volatility with
280 increasing oxygen number for $C_{10}H_{14}O_{4-6}$, while FIGAERO measurements showed an increase in
281 volatility from 4 to 6 oxygen atoms. Another discrepancy arises when comparing the FIGAERO
282 and VBS for compounds with a constant number of carbon or oxygen but differing hydrogen
283 numbers such as $C_8H_{12-14}O_6$ or $C_9H_{14-16}O_6$. Here, the saturation vapor pressure tended to be higher
284 by 1-3 orders of magnitude for the compounds containing fewer hydrogen. The opposite was true
285 with dimers such as $C_{19}H_{28-30}O_7$ and $C_{19}H_{28-30}O_6$. These observations are consistent with [*Hyttinen*
286 *et al.*, 2021; *Kurtén et al.*, 2016] that in the case of monomers, an increase in intermolecular
287 hydrogen bonding by hydrogen-donor functional groups likely decreases the volatility, while
288 dimers with more hydrogen atoms can favorably form intramolecular hydrogen bonds, which make
289 them more stable in the gas-phase. Hydrogen bonding, functional groups, and structural
290 differences between isomeric compounds affect volatilities [*Pankow and Asher*, 2008; *West et al.*,
291 2023], but these processes are not considered in the current VBS parameterizations. For some
292 OOMs (e.g., $C_8H_{14}O_5$ and $C_{20}H_{32}O_6$), we observed multimodal peaks in the FIGAERO
293 thermogram (Figure S3), which is likely due to the presence of isomeric compounds. Thus,
294 isomeric compounds can affect the shape (and peak location) of the desorption peak (T_{max}) for a
295 certain OOM moiety and hence, influence the estimation of the saturation vapor concentration.

3.3. Molecular structures and reaction pathways of dimer OOM isomers

297 Isomers were identified in the liquid chromatogram, as they have different interactions with the
298 column material due to their unique chemical structures and so can be effectively separated at
299 different retention times (RT). The UPLC/(-)ESI-Orbitrap mass spectrometer detected 77 OOM
300 compounds different m/z values. When including their isomers, which had distinctive MS/MS
301 fragmentation ions and hence different molecular structures, there were in total 277 compounds
302 (Figure 1g).

303 $C_{19}H_{30}O_5$ (m/z = 337) is one of the major dimers produced from α-pinene ozonolysis, as has been
304 previously detected in boreal forests [*Kristensen et al.*, 2016] and laboratory studies [*Kenseth et*
305 *al.*, 2023; *Y Zhao et al.*, 2018b]. Figure 3a shows the EIC of $C_{19}H_{30}O_5$, indicating six isomers that
306 eluted at different retention times. Figures 3b and 3c show MS/MS spectra of two main isomers
307 $C_{19}H_{30}O_5$-1 (RT = 15.3 min) and $C_{19}H_{30}O_5$-2 (RT = 20.2 min). The distinctive MS/MS

308   fragmentation ions from these two compounds indicate that these two isomeric compounds have
309   structural differences. Based on its MS/MS fragmentation ions, we have derived its chemical
310   structure containing secondary ozonide (SOZ) for $C_{19}H_{30}O_5$-1 (Scheme S1). Due to the unique
311   SOZ structure, there are three main fragmentations, leading to 3 pairs of daughter ions (m/z =
312   213/123, 167/169, and 183/153). These daughter ions further fragment by losing H_2O, CO, or
313   C_2H_3O to form smaller ions (m/z 195, 125, and 141). The parent SOZ also loses H_2O to form
314   another daughter ion (m/z = 319). All these fragmentation ions appeared in the MS/MS spectrum
315   (Figure 3b). There were additional two fragmentation ions (m/z = 293 and 275) which we cannot
316   identify their structures. Based on the SOZ structure, we propose that the $C_{19}H_{30}O_5$-1 isomer forms
317   from the gas phase dimerization of a stabilized Criegee intermediate (sCI) and norpinonaldehyde
318   ($C_9H_{14}O_2$), and further undergoes gas-to-particle conversion (Figure 3d). These reactants,
319   including their specific molecular structures, were reproduced by the GECKO-A modeling
320   simulations from a-pinene ozonolysis (Table S3). Previously, *Y Zhao et al.* [2022] also proposed a
321   SOZ molecular structure for $C_{20}H_{32}O_5$ from α-pinene ozonolysis, and indicated the same dimer
322   formation pathway from bimolecular reactions involving sCI. On the other hand, our proposed
323   reaction mechanism contrasts with the particle-phase aldol condensation route for $C_{19}H_{30}O_5$
324   proposed by *Witkowski and Gierczak* [2014], although they showed similar fragmentation patterns
325   (as shown in Scheme S1).

326   The MS/MS spectrum for $C_{19}H_{30}O_5$-2 (RT = 20.2 min) (Figure 3c) shows the same fragmentation
327   ions as shown previously by [*Kenseth et al.*, 2023; *Kristensen et al.*, 2016]. As illustrated in
328   Scheme S2, there are two main fragmentations near the ester function group, leading to 2 pairs of
329   daughter ions (m/z= 169/167 and 151/185). The m/z = 185 undergoes subsequent fragmentations
330   by losing CO_2, H_2O, then C_3H_6 to form m/z = 141, 123, then 81, respectively. Previously, [*Kenseth
331   et al.*, 2023] showed that this isomer forms in the particle phase from cyclic acylperoxyhemiacetal
332   (m/z = 185) and α-pinanediol (m/z 193), via nucleophilic addition followed by Baeyer-Villiger
333   decomposition reaction (Figure 3e). Cyclic acylperoxyhemiacetal is the isomerization product of
334   *cis*-3-peroxy pinalic acid. The GECKO-A model predicted the formation of *cis*-3-peroxy pinalic
335   acid and α-pinanediol in the gas phase (Table S3), with the same molecular structures as shown in
336   Figure 3e. These two compounds were also identified in the particle phase with UPLC/(-)ESI-
337   Orbitrap mass spectrometer (e.g., Figure S4). Thus, our analysis demonstrates that the accretion
338   reaction of $C_{19}H_{30}O_5$-2, which takes place within SOA particles [*Kenseth et al.*, 2023], also occurs
339   in newly formed biogenic particles.

340   Thus, while $C_{19}H_{30}O_5$-1 forms in the gas phase via bimolecular reactions involving sCI followed
341   by gas-to-particle conversion, $C_{19}H_{30}O_5$-2 forms within the particle phase via esterification. These
342   proposed mechanisms are qualitatively consistent with *Y Zhao et al.* [2022] that one of $C_{19}H_{30}O_5$
343   isomers was reduced when a sCI scavenger was introduced into the system (very likely $C_{19}H_{30}O_5$-
344   1), while another isomer of $C_{19}H_{30}O_5$ did not change in abundance when sCI scavenger was used
345   (possibly $C_{19}H_{30}O_5$-2).

346   To further understand the relative contribution of these two processes discussed above, gas-to-
347   particle conversion of $C_{19}H_{30}O_5$-1 versus particle-phase Baeyer-Villiger reaction for $C_{19}H_{30}O_5$-2,
348   we changed relative concentrations of oxidants (Figure S5). We observed a lower abundance of
349    $C_{19}H_{30}O_5$-2 compared to $C_{19}H_{30}O_5$-1 when the oxidation system was dominated by ozone rather

350   than OH (ozone = 1200 ppb, OH = 1.6 ppt, $k_{\alpha\text{-pinene* O3}}$ = 9 x 10^{-17} cm^3 molecule^{-1} s^{-1} and $k_{\alpha\text{-pinene*}}$
351    $_{OH}$ = 5.3 x 10^{-11} cm^3 molecule^{-1} s^{-1} [*Atkinson et al.*, 2006]). However, when ozone concentration
352   was varied from 250-1000 ppb, $C_{19}H_{30}O_5$-1 abundance was reduced compared to experiments at
353   higher ozone, but $C_{19}H_{30}O_5$-2 abundance remained constant. This is because $C_{19}H_{30}O_5$-1 formation
354   was dependent on gas-phase precursor concentration, while $C_{19}H_{30}O_5$-2 formed from particle-
355   phase reactions. Clearly, isomeric compounds can be formed independently from both gas- and
356   particle-phase reaction pathways to contribute to the particle mass, and which isomer becomes
357   dominant depends on the atmospheric conditions.

358   There were 12 compounds detected only in the particle phase, without their gas-phase counterparts
359   (Figure 1h and Table S1), implying that these compounds form via particle-phase reactions (as
360   opposed to gas-to-particle conversion). Among these compounds, $C_{16}H_{26}O_6$ (*m/z* = 313) has been
361   identified in the particle phase in previous studies [*Huang et al.*, 2020; *Kristensen et al.*, 2016;
362   *Zhang et al.*, 2015]. Here we propose that two isomers of $C_{16}H_{26}O_6$ form exclusively in the particle
363   phase via decarboxylation from two isomers of $C_{17}H_{26}O_8$ (as opposed to gas-to-particle
364   conversion).

365   The EIC of $C_{16}H_{26}O_6$ shows the presence of five isomers eluting out at different RTs (Figure 4a).
366   The fragmentation spectra of the two most abundant isomers, $C_{16}H_{26}O_6$-1 (RT = 16.6 min) and
367    $C_{16}H_{26}O_6$-2 (RT = 18.4 min), are shown in Figures 4b and c, respectively. The proposed structure
368   and fragmentation mechanism of $C_{16}H_{26}O_6$-1 is presented in Scheme S3. The fragmentation ions
369   of $C_{16}H_{26}O_6$-1 agree with those shown by [*Kristensen et al.*, 2016], although this cited study did
370   not show the molecular structure of this compound. Fragmentation of the parent ion of $C_{16}H_{26}O_6$-
371   1 yields one daughter ion (m/z = 185), which fragments into three smaller ions: m/z = 167 by
372   losing water, and two isomers with m/z = 141 by losing CO_2. The two isomers further fragment
373   into smaller ions with m/z = 123 and 71 by losing H_2O and C_5H_{10}, respectively. Another small ion
374   with *m/z* = 57 is fragmented from the parent molecule. These fragmentation ions are consistent
375   with those shown in the MS/MS spectrum (Figure 4b). Figure 4d illustrates the reaction pathway
376   for the $C_{16}H_{26}O_6$-1 isomer proposed in this study. First, the precursor molecule ($C_{17}H_{26}O_8$) forms
377   by the RO_2 + RO_2 gas-phase dimerization of $C_9H_{14}O_5$ and $C_8H_{14}O_5$. Upon partitioning to the
378   particle phase, $C_{17}H_{26}O_8$ undergoes the decarboxylation reaction to form $C_{16}H_{26}O_6$-1. The HrTOF-
379   CIMS measured both the reactants ($C_9H_{14}O_5$ and $C_8H_{14}O_5$) in the gas phase, while $C_{17}H_{26}O_8$ was
380   detected both in the gas and particle phase (Table S1). The GECKO-A model also simulated the
381   reactants with the same molecular structures (Table S3) as shown in Figure 4d.

382   For the second isomer, $C_{16}H_{26}O_6$-2, fragmentation of the parent molecule at the ester linkage
383   results in daughter ions with *m/z* = 171 and 141 (Scheme 4). Loss of CO_2 from *m/z* 171 produces
384   a fragmentation ion *m/z* = 123, which subsequently loses H_2O to form *m/z* = 109. These
385   fragmentation ions are consistent with the MS/MS spectrum shown in Figure 4c. We propose that
386    $C_{17}H_{26}O_8$ -2 forms in the gas phase by the RO_2· + RO_2· dimerization from $C_9H_{14}O_4$· and $C_8H_{14}O_6$·
387   (both were detected in the gas phase), followed by gas-to-particle conversion, and further
388   undergoes particle-phase decarboxylation to form the dimer $C_{16}H_{26}O_6$-2 in the particle phase
389   (Figure 4e). Thus, the two isomers of $C_{16}H_{26}O_6$ form from the particle-phase decomposition
390   reactions of two isomers of $C_{17}H_{26}O_8$, while $C_{16}H_{26}O_6$ was not detected in the gas phase.

391   The cumulative signal fractions of OOMs showed higher ion intensities of C_{15-19} species than C_{20}
392   in the UPLC/(-)ESI-Orbitrap MS (Figure S7). Interestingly, there were also more different
393   compounds (and their isomers) within C_{15-19} than C_{20}. The increased C_{15-19} dimer abundance
394   compared to C_{20} dimers was also seen in previous chamber studies, where the particle-phase
395   chemical evolution of dimers measured in real-time using extractive electrospray ionization-time
396   of flight (EESI-TOF) mass spectrometer showed a slowly increasing abundance of C_{16-18} dimers
397   with decreasing C_{20} dimers [*Pospisilova et al.*, 2020]. The particle-phase decomposition reactions
398   can explain these observations.

Atmospheric Implications

400   We have investigated the chemical composition of gas- and particles-phase OOMs generated from
401   ozonolysis of α-pinene, including their isomeric molecular structures and potential reaction
402   schemes, using two high-resolution mass spectrometers: HrTOF-CIMS equipped with the
403   FIGAERO inlet and UPLC/(-)ESI-Orbitrap MS. The key result of this study is particle-phase
404   formation of OOMs. For example, whereas one isomer $C_{19}H_{30}O_5$ (a key dimer formed from
405   ozonolysis of α-pinene) is from gas-to-particle formation upon the formation in the gas phase via
406   sCI and aldehyde dimerization, another isomer of $C_{19}H_{30}O_5$ forms in the particle phase via
407   accretion reaction of α-pinanediol and cyclic acylperoxyhemiacetal. Likewise, two isomers of
408    $C_{16}H_{26}O_6$ form from decarboxylation reactions exclusively in the particle phase. Our results
409   highlight the dynamic reaction medium of the particle phase, where OOMs readily undergo
410   particle-phase reactions such as accretion and decomposition to form OOMs with entirely different
411   molecular structures. Thus, biogenic new particle formation or growth cannot be predicted solely
412   based on the volatility of OOMs formed in the gas phase. Our study demonstrates that particle-
413   phase formation and transformations of OOMs should also be considered in the biogenic NPF
414   processes, as well as urban NPF (in the presence of monoterpenes).

415   Most OOMs produced from the ozonolysis of α-pinene have 2-8 isomers. These isomeric
416   compounds have not been considered in the NPF mechanisms. Rather, typically, an extremely
417   broader grouping of biogenic OOMs has been used, such as C_5, C_{10}, C_{15}, or C_{20} groups to
418   understand biogenic NPF chemical mechanisms [*Heinritzi et al.*, 2020; *McFiggans et al.*, 2019;
419   *Lee Tiszenkel and Lee*, 2023], regardless the different OOMs chemical identities and molecular
420   structures. Our results show that the VBS parameterization devised for determining the volatility
421   of organic compounds exhibits inconsistencies with the FIGAERO-measured saturation vapor
422   concentrations for some OOMs, depending on monomer vs. dimer, or whether OOMs containing
423   intermolecular vs. intramolecular hydrogen bonding. These two methods, VBS and FIGAERO
424   thermogram, do not consider isomers of OOMs in the volatility estimation.

425   The analysis of OOM formation pathways presented here highlights the effectiveness of combined
426   online HrTOF-CIMS and offline UPLC/(-)ESI-Orbitrap MS/MS analysis. HrTOF-CIMS is
427   essential for understanding the chemical evolution of an organic compound in real time. The liquid
428   chromatogram and high-resolution MS/MS analyses are necessary for identifying the molecular
429   structures of isomers. Our study using these two independent high-resolution mass spectrometers
430   demonstrates that, despite the various possible artifacts associated with sampling storage and
431   extraction procedures used in the off-line method, the mass spectra of OOMs in the particle show
432   strikingly similar features from these two methods. Thus, our results provide reestablished

433 confidence in using the UPLC/(-)ESI-Orbitrap MS/MS analysis for chemical composition and
434 molecular structural analysis.

435 This study investigates the OOMs chemical composition in biogenic new particles. However, it is
436 probable that similar particle-phase reactions also take place in SOA particles. Accounting for
437 structurally resolved OOMs formed in gas- or particle-phase reaction mechanisms may help close
438 the long-standing gap between model predictions and ambient measurements of SOA [*Hallquist*
439 *et al.*, 2009].

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Author Contributions

446 SHL designed the concept; VVG, LT, RR, and DJB performed experiments; VVG, LT, and SHL
447 made data analysis; ZW, JTL, and KCB made modeling simulations; VVG, LT, and SHL wrote the
448 manuscript; all co-authors commented on the manuscript.

Data and Code Availability

450 Data used in this study will be available upon request (shanhu.lee@uah.edu). The MS/MS spectra
451 of 139 select OOMs produced from α-pinene ozonolysis will be available at a public data
452 repository site.

Competing Interests

454 At least one of the (co-)authors is a member of the editorial board of Atmospheric Chemistry and
455 Physics. The authors declare no other competing interests.

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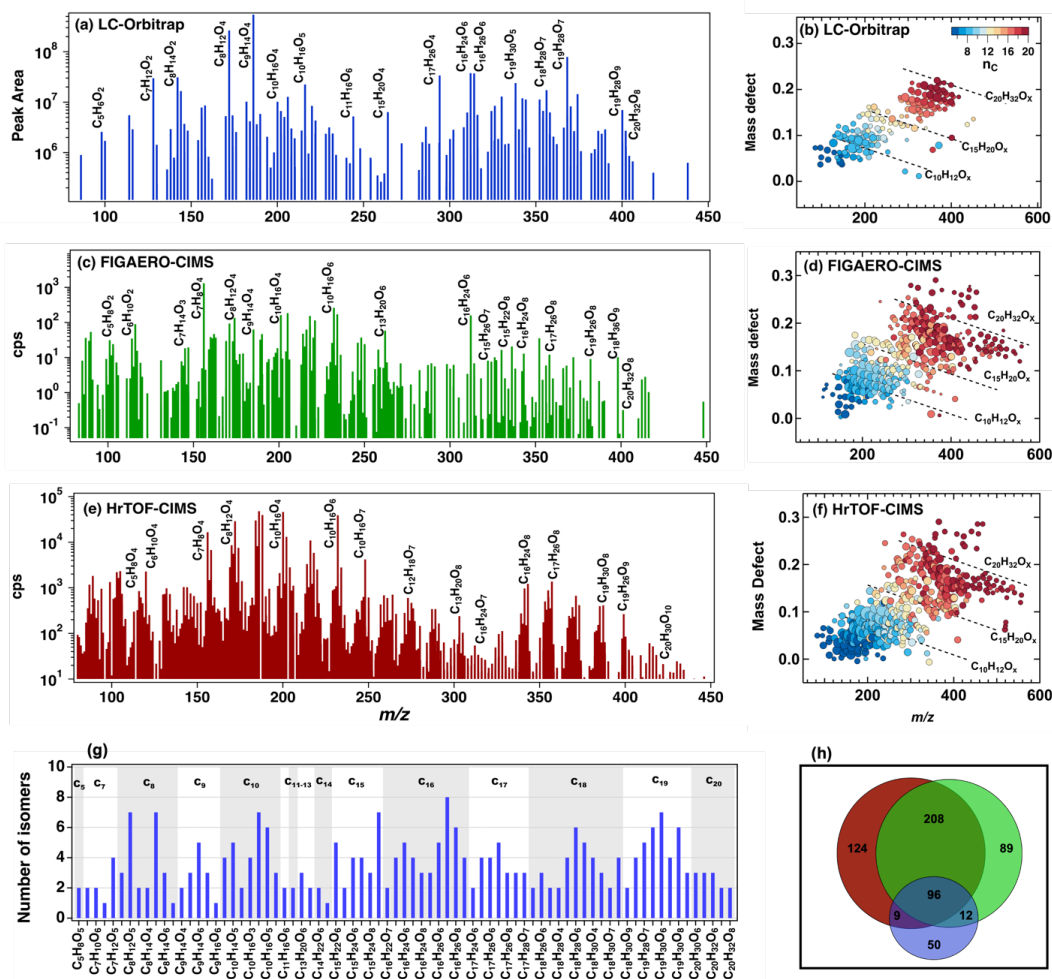
Figure 1. Mass spectra and mass defects of OOMs detected from α-pinene ozonolysis flow-tube experiment, with α-pinene of 238 ppb, ozone of 1.2 ppm, a temperature of 298 K, RH < 10 %, [OH] = 1.6 ppt and residence time = 150 s, in the particles phase using offline UPLC/(-)ESI-Orbitrap MS (a, b) and the particle (c, d) and gas phase (e, f) using FIGAERO-CIMS. (g) The number of isomers identified for each detected OOMs in the particle phase. For clarity, only every other OOMs are shown with their chemical formulas here (see Figure 2c for the entire formulas for all OOMs.) (h) Venn diagram showing the number of compounds detected from the gas phase (red) and particle phase (green) with FIGAERO-CIMS, and in the particle phase with UPLC/(-)ESI-Orbitrap MS (blue).

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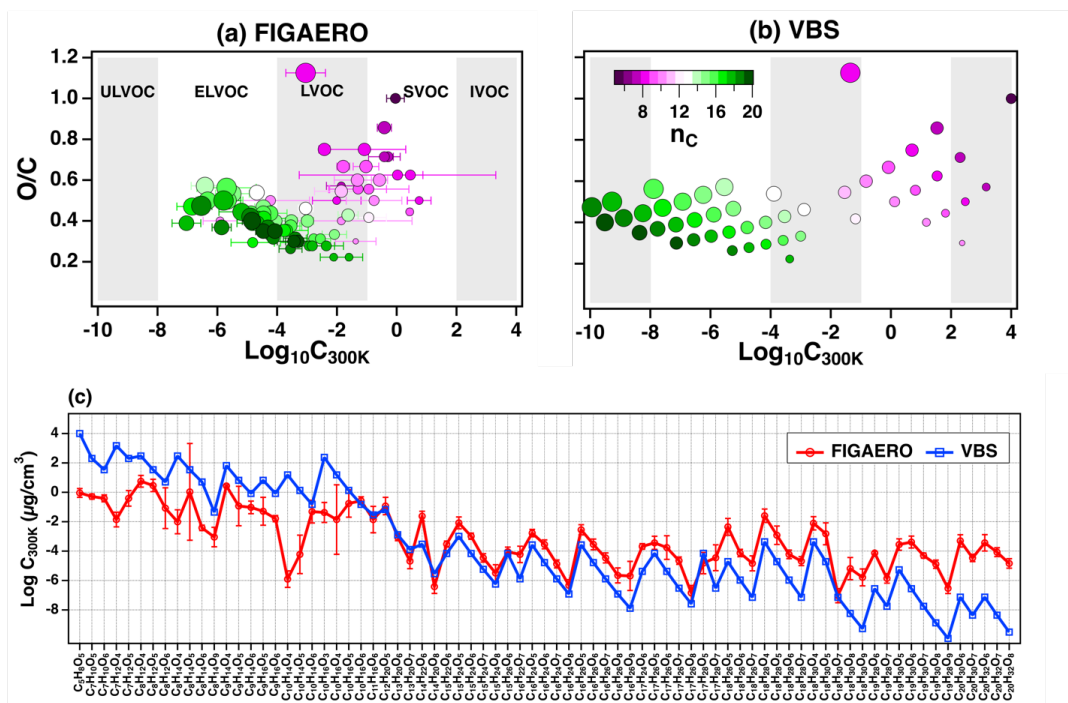
Figure 2. Comparison of saturation vapor concentrations of biogenic OOMs produced from α-pinene ozonolysis indicated by O:C ratio measured with the FIGAERO thermogram (a) and estimated from VBS (b). The color scale represents the number of carbons in the compound and bubble size represents the oxygen numbers. The O/C vs. $Log_{10}C_{300k}$ space is classified into different bins as ultra-low volatility organic compounds (ULVOC); extremely-low volatility organic compounds (ELVOC); low-volatility organic compounds (LVOC); semi-volatile organic compounds (SVOC); and intermediate volatile organic compounds (IVOC). (c) Comparison of the volatilities obtained from the VBS and FIGAERO for each chemical formula of OOMs.
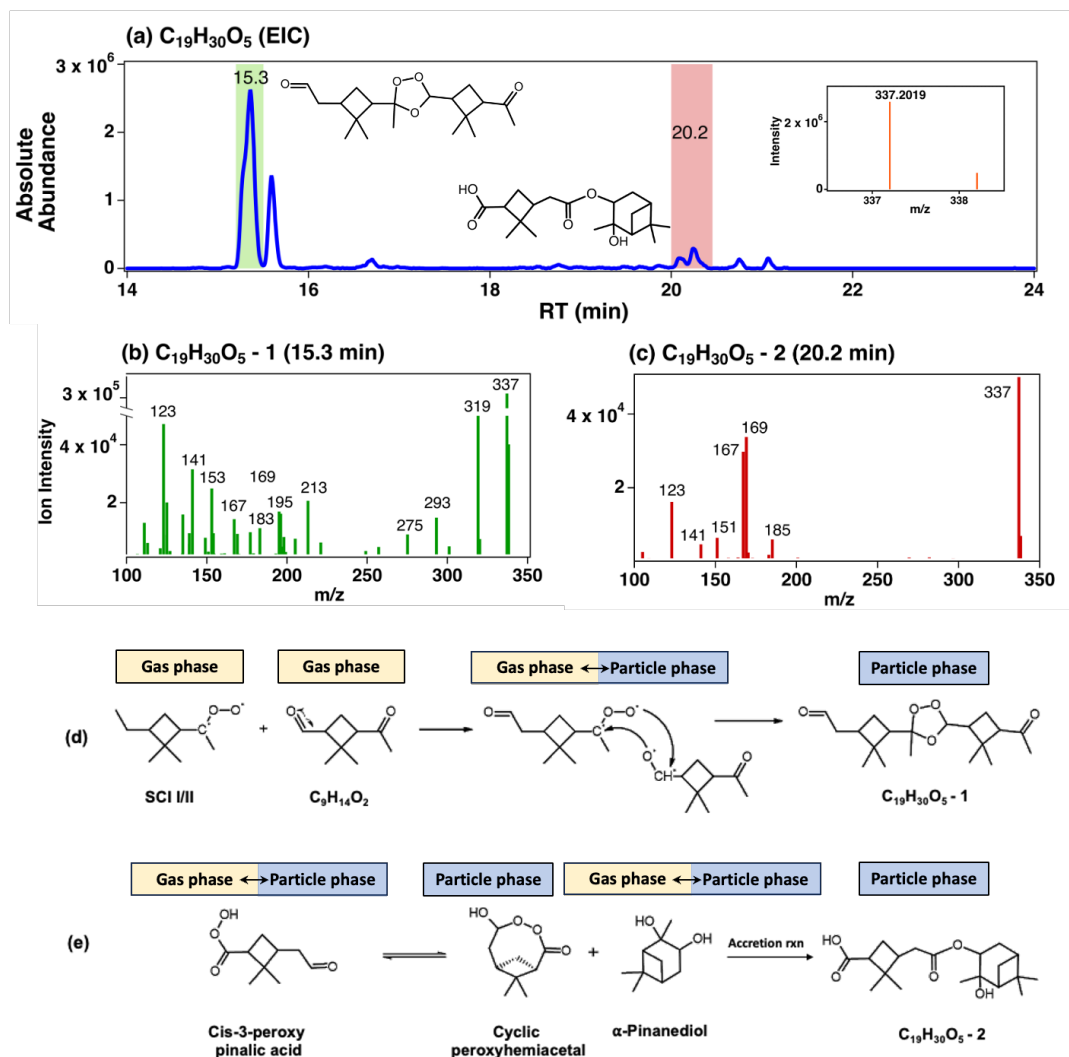
Figure 3. (a) EIC of $C_{19}H_{30}O_5$ (m/z = 337.2019). The inset figure shows the MS spectra showing the mass peak of the $C_{19}H_{30}O_5$ at the retention time (RT) of 15.3 min. MS/MS spectra of the isomers with the RT of 15.3 min (b) and 20.2 min (c). Reactions leading to the formation of two isomers, $C_{19}H_{30}O_5$ -1 (d) and (e) $C_{19}H_{30}O_5$ -2 (e). Reactions shown in Figure 3d were proposed in this study. Reactions shown in Figure 3e were adopted from [*Kenseth et al.*, 2023].

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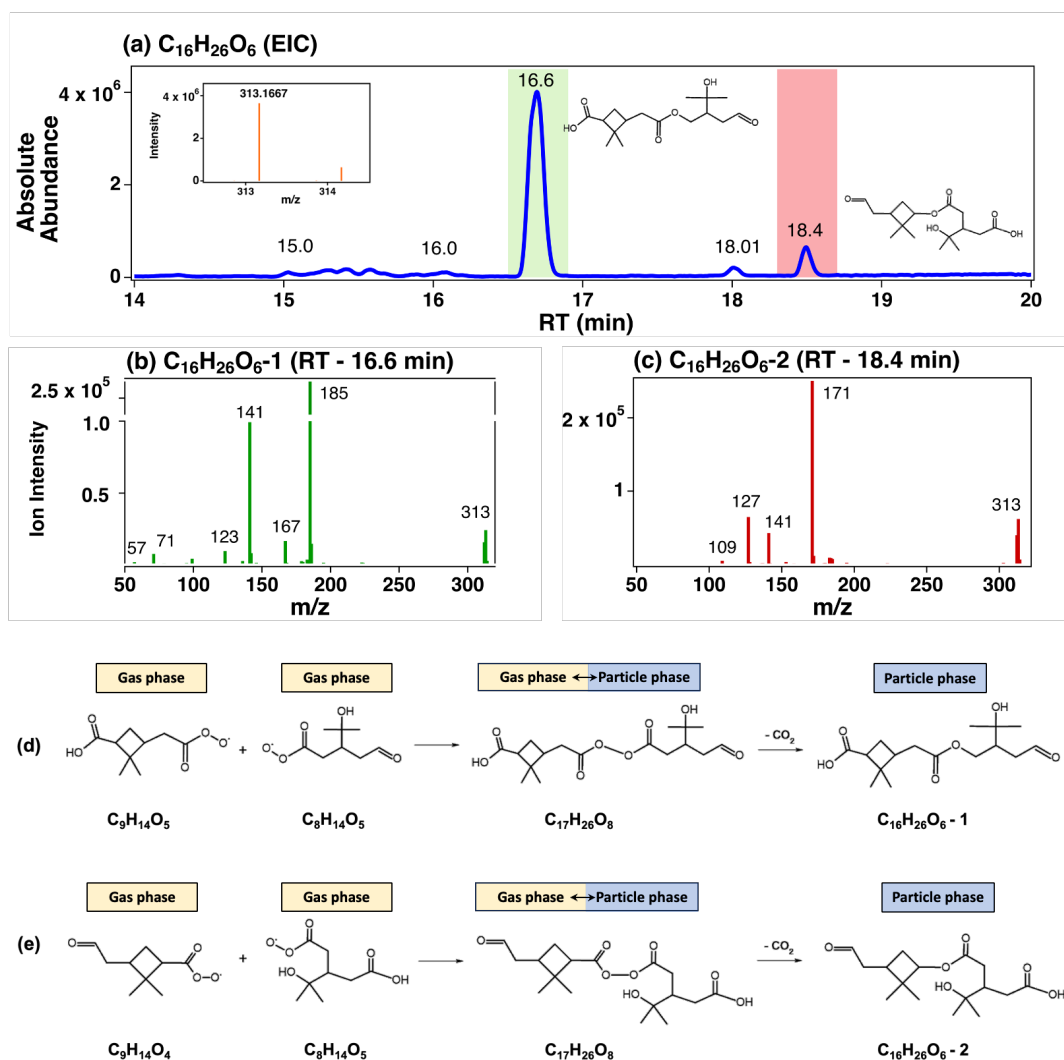
Figure 4. **(a)** EIC of $C_{16}H_{26}O_6$ with its five isomers labeled. The inset figure is the MS spectra showing the mass peak of $C_{16}H_{26}O_6$ at RT = 16.6 min. MS/MS spectrum of two isomers with RT = 16.6 min (b) and RT = 18.4 min (c). Reactions leading to the formation of two isomers, $C_{16}H_{26}O_6$-1 (d) and $C_{16}H_{26}O_6$-2 (e), both proposed in this study.

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