Total Air Content measurements from the RECAP ice core

Sindhu Vudayagiri[1], Bo Vinther[1], Johannes Freitag[2], Peter L. Langen[3], Thomas Blunier[1]*,

1 Physics of Ice, Climate and Earth, Niels Bohr Institute, University of Copenhagen, Tagensvej 16, 2200 Copenhagen, Denmark
5 2 Alfred Wegner Institute, Snow and Firn Section Glaciology, 27570 Bremerhaven, Germany
3 Department of Environmental Science, iClimate, Aarhus University, Frederiksborgvej 399, 4000 Roskilde, Denmark

*Correspondence to: Thomas Blunier (blunier@nbi.ku.dk), Tel: +45 35 32 05 84

Abstract. Total air content (TAC) of the REnland ice CAP project (RECAP) core, drilled in summer 2015, is presented. In principle, TAC is a proxy for the elevation at which the ice was originally formed as the TAC in ice cores is
10 predominantly influenced by surface air pressure and conditions like temperature and local summer insolation. This, however, presupposes dry sintering of the firn with no surface melting. The RECAP TAC data shows incoherently low values in the Holocene climatic optimum (6 to 9 kyr b2k) and in the last interglacial (119 to 121 kyr b2k) originating from melt layers which renders the TAC data unfit for paleo elevation interpolation. Melt instances can, however, be used to reconstruct summer temperatures and we find that Renland has been ~2° to 3°C warmer compared to today in the early
15 Holocene. Similarly, samples from the previous interglacial hint summer temperatures at least 5°C warmer than today. The glacial section (11.7 kyr to 119 kyr b2K) has consistent TAC values thus in principle facilitating the past elevation calculations. However, we observe TAC variations related to Dansgaard-Oeschger events (D-O) that cannot originate from elevation changes but must be linked to changes in the firn structure. We analyse the pattern of these structural changes in the RECAP and NGRIP cores and conclude that only samples from the stable portion of the Last Glacial
20 Maximum are suitable for elevation reconstructions. Within uncertainty, the elevation has been similar to today at the last glacial maximum.

Keywords: Greenland Ice sheet, Renland, Total air content, elevation change, Insolation, melt layers, firn structure.

1 Introduction

25 We present the first total air content (TAC) record from the Renland ice cap from the RECAP core, drilled in 2015.
The motivation for studying TAC at Renland was that Vinther et al (2009) used Renland as an anchor point in their reconstruction of the Holocene Greenland Ice Sheet (GIS) history. They argued that the ice cap has not experienced significant ice flow or elevation changes due to its isolation from the GIS owing to the surrounding topography. The TAC signal from the RECAP core could be used to infer the elevation changes thereby supporting or refuting their assumption.
30 However, in the course of the study we learned that the RECAP TAC is affected by melt during the Holocene making it impossible to use it for elevation reconstructions. On the other hand, the melt fractions in the samples construed by assuming a linear relationship between the TAC of a sample and the melt fraction of the sample allow for estimating summer temperatures. We carry that concept to samples from the previous interglacial (Eemian). The effect of local summer insolation on the RECAP TAC signal is analysed. The RECAP TAC signal from the glacial section, unaffected
35 by melt is, however, affected by rapid climate change events that hinder reconstruction of the past elevation of the Renland ice cap except for the climatically stable phase of the Last Glacial Maximum.

1.1 The RECAP ice core

The Renland ice cap is situated in Eastern Greenland on a high elevation plateau on the Renland peninsula in the Scoresbysund fjord (Fig. 1) with a present elevation of 2340 m a.s.l. at the summit. The RECAP core is ~584 m long and

40 was drilled to bedrock in 2015 at an elevation 2315 m a.s.l. (71° 18' 18" N; 26° 43' 24" W) near the summit, about 1.5

km away from the 1988 drill site (Johnsen et al., 1992). The present day annual mean temperature is -17.8°C measured

at 20 m depth in the firn, and the present accumulation rate is approximately 0.5 m of ice equivalent precipitation per

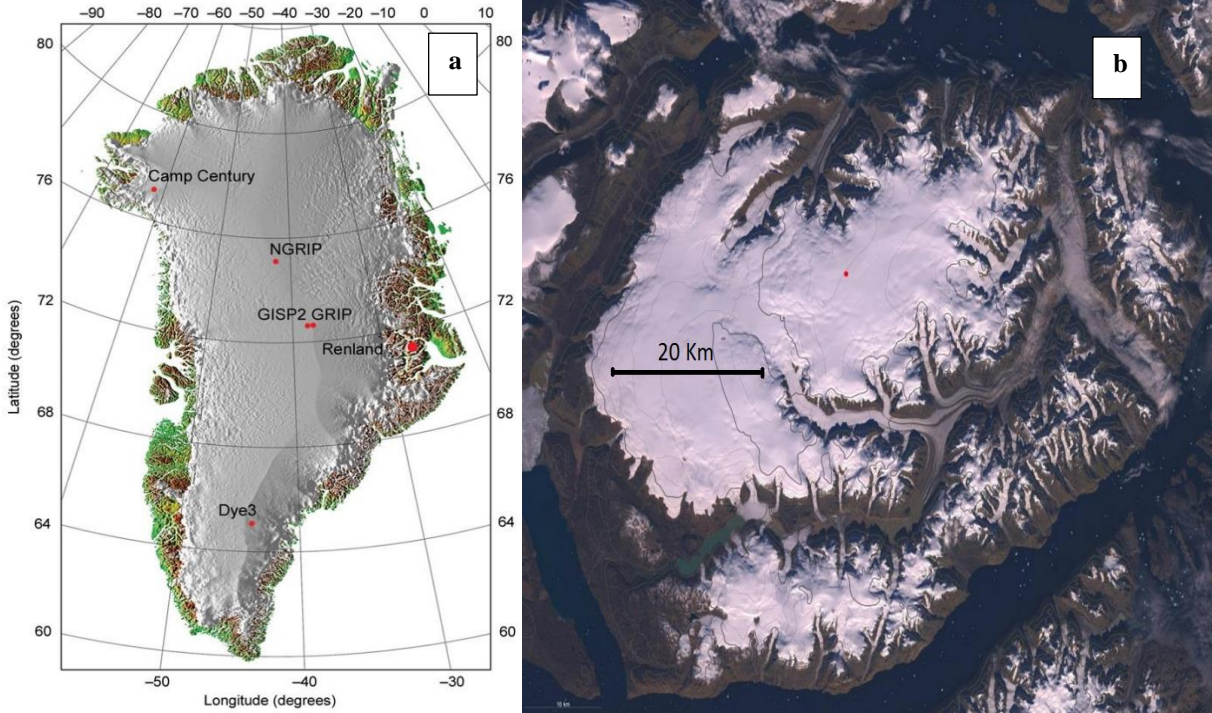
year. In the interior of the Greenland ice sheet the average size of the air bubbles is monotonically decreasing with depth

till they disappear forming clathrates (Shoji and Langway, 1982). The enclosed air in the RECAP core exists fully in the

45 form of air bubbles. At the given temperature, clathrate formation would start below the bedrock depth of ~584 m below

surface (Uchida and Hondoh, 2000). The bubble diameter, however, is increasing again from about 530 m below the

surface (supplemental Fig. S3). This may be due to fast thinning of the annual layers in the Renland ice cap causing the

small bubbles to coalesce to form bigger bubbles.

50 **Figure 1: (a) Map of Greenland, showing the location of the Renland ice cap and other cores (Danish Cadastre) (b) Satellite image of the Renland peninsula, which is almost entirely covered by the Renland ice cap (Landsat).**

1.2 Total air content in ice

1.2.1 Principle

55 The density of dry sintered snow at the surface of an ice sheet is typically 0.3–0.35 g cm^{-3}. This open porous firn is then

densified by compaction and dry sintering to a density of 0.81–0.84 g cm^{-3} where the open pores are isolated. The amount

of gas trapped at this time, the total air content (TAC), depends on the pore volume, the temperature, and the pressure

(e.g. Martinerie et al., 1992). It is usually expressed as cm^{3} of gas per kg of ice at standard temperature and pressure (STP)

(equation 1) where V_c, P_c, and T_c are pore volume per kg of ice, and pressure, temperature at close respectively. P and T

60 are standard temperature and standard pressure (1013 mbar and 273.15 K).

$$\text{TAC} = V_c \frac{P_c}{T_c} \frac{T}{P} \tag{1}$$

With known V_c and T_c elevation changes can be obtained applying the barometric formula (equation 2).

$$P_c = P_a \left[\frac{T_a}{T_c} \right]^{\frac{gM_{air}}{R(\frac{dT}{dz})}} \text{ with } T_c = T_a + \frac{dT}{dz}(h_c - h_a) \tag{2}$$

Where P_c is the pressure at altitude h_c, P_a, T_a, and h_a are pressure, temperature, and elevation at sea level, respectively. dT/dz is the lapse rate at the location, M_{air} the molecular mass of air, g the gravitational constant, and R the gas constant. A 1% change in TAC at the elevation of RECAP corresponds to a pressure change of about 7 mbar and 80 m change in elevation.

It has been observed that the pore volume (V_c) at the air isolation depth exhibits a slight dependence on temperature (Martinerie et al., 1994). This relationship can be described by the following linear equation, which has a correlation coefficient of 0.90:

$$V_c \, (\text{cm}^3 \, \text{g}^{-1}) = 6.95 \times 10^{-4} \, T_c(K) - 0.043 \tag{3}$$

We will apply this parametrization based on data from sites with a temperature range ~-15 to -60°C to calculate the pore volume for the RECAP core.

At any site short term sub-annual variability of V_c on the order of 20% is observed. It is explained by the variability of the density originating from summer to winter precipitation and successive metamorphosis throughout the firn column to the air insolation depth (Hörhold et al., 2011). Additionally high density wind crusts potentially add to the variability. (Martinerie et al., 1994).

1.2.2 TAC variations across millennia

First on O_2/N_2 ratios, later also on TAC, it has been discovered that they both anti-correlate with local summer insolation in Greenland and Antarctica (Suwa and Bender, 2008; Kawamura et al., 2007; Raynaud et al., 2007, and references therein). The reasoning brought forward is that summer time insolation influences the metamorphism of snow near the surface of polar ice as it causes evaporation and grain growth (Bender, 2002). It is explained that summer insolation causes rapid grain growth in the snow surface by creating an apparent summer temperature gradient. Thus, the increase in grain size below the surface affects the densification process. An increase in insolation thereby causes the grain size to increase, porosity at close-off to decrease and density at close off to increase. The proposed mechanism explains the anti-correlation between the integrated summer insolation and the TAC. As insolation increases, porosity at close off and pore volume decreases, causing an overall decrease of the TAC. The O_2/N_2 signal results from fractionation at the close off as a consequence of mentioned surface metamorphism processes (Suwa and Bender, 2008). Both TAC and O_2/N_2 have proven to be reliable proxies for local insolation and hence can be used for orbital dating of ice cores despite the remaining gaps in our understanding of the physical mechanisms (Lipenkov et al., 2011). After correcting for the effect of changing local solar insolation, TAC can be interpreted to give paleo surface elevations, with high TAC corresponding to lower elevations (Raynaud et al., 1997; Raynaud et al., 2007; NEEM Community Members, 2013).

1.2.3 Perennial TAC variations

TAC has also been found to be influenced by rapid climatic transitions in connection with Dansgaard-Oeschger (D-O) events during the last glacial in the Greenlandic NGRIP core(Eicher et al., 2016). Surprisingly this seems also to be the case for some Antarctic sites (Epifanio et al., 2023). Lacking understanding for those fast TAC variations it seems that only TAC measurements from climatically stable periods should be used for past elevation estimation.

2 Measurements

Measurements of the RECAP ice core were made at PICE (Physics of Ice, Climate and Earth) and PSU (Penn State University). While the system at PICE is dedicated to total air content measurements following the barometric method and giving absolute calibrated volumes (Lipenkov et al., 1995), the measurements at PSU are a by-product of measurements for $\delta^{15}N$ and CH_4 contents.

At PICE, air is extracted from cubical samples of 10 to 15 g of ice by two melt-refreeze cycles under vacuum. The extracted air is passed through a dry ice/ethanol water trap and quantitatively trapped on Haysep D at LN_2 temperature. The air is then expanded into a calibrated measuring volume by warming up the Haysep D trap. Experimental details are given in supplements S1. Data from measurements at PICE are published here: (Blunier, 2024).

Two sets of TAC measurements are obtained at PSU. The samples used for CH_4 measurements are cylinders of diameter 4.1 cm, height of 5.5±0.3 cm, weighing 65 ± 3 g each and the samples used in $\delta^{15}N$ measurements are rectangular cubes of ice (2 x 1.2 x 5 cm) weighing ~13 g each. In both these measurements, an automatic air extraction device (referred to as "The Spider"), which employs the vacuum volumetric principle is used. The volume of the extracted air is measured after which the air samples are used for CH_4 and $\delta^{15}N$ measurements (Fegyveresi, 2015).

The spider apparatus consists of 14 steel vessels used to hold ice samples, each with a total sampling volume of ~96 ± 2 cm^3 (Fegyveresi, 2015). During measurements, the system performs a single melt-refreeze cycle to free the trapped air from within the ice (Fegyveresi, 2015). Ice samples are placed in the respective vessels and isolated from the ambient atmosphere using copper gaskets. The entire system is then evacuated to 0.3 mbar to remove air in each vessel's head-space, and various leak-checks are performed to ensure the seals are intact with no contamination from ambient air. The ice samples are then melted allowing the air trapped in the ice samples to escape into the headspace of the enclosing vessels. The melt is then refrozen, leaving the liberated air separated above each of the refrozen samples. Once the temperature of the ice reaches -69°C, the air in each vessel is expanded into a vacuum manifold containing a 10 cm^3 sample loop, which is then connected to a gas chromatograph (Fegyveresi, 2015). The pressure in the vacuum manifold with the ice core air sample is noted (generally between 60 and 80 torr) before the loop is switched for CH_4 concentration or $\delta^{15}N$ measurements. Solubility correction in connection with the CH_4 measurements is quite large (~6%) in this method due to the high ratio of sample/vessel volume which yields a high headspace pressure that contributes to more gas getting stuck in the refrozen ice (Fegyveresi, 2015). The calibrations (volume and temperature) are briefed in the supplementary information (S2). Data from PSU-$\delta^{15}N$ measurements are published here: (Sowers, 2018).

3 Cut bubble correction

Air bubbles at the surface of the sample are cut during sample preparation resulting in air loss. Therefore, TAC measurements need to be corrected for the so called 'cut bubble effect' (CBE). The CBE correction approximates to 10% near the close off depth and decreases to around 1% in deeper strata (Martinerie et al., 1990). Martinerie et al. (1990) derived the formula for the cut bubble effect assuming spherical bubbles:

$$TAC = \left(1 - \frac{1}{2}\langle D\rangle \frac{A_S}{V_S}\right)^{-1} \cdot TAC_{raw} \tag{4}$$

Where, D is the average bubble diameter in the sample and A_S, V_S are sample surface area and volume, respectively. In the current study, only samples analysed at PICE had their bubble diameters measured. A photograph of each sample is taken (supplemental Fig. S4) from which an average of 20 bubble diameters is taken as the sample bubble diameter. The average bubble diameter of every sample and the corresponding CBE calculations are provided in the supplementary section S3. TAC data from PSU are a by-product of methane concentration and $\delta^{15}N$ measurements. Bubble diameters have not been measured for these samples. In this study we estimate the CBE for the PSU data from the PICE data. Through the Holocene section of the RECAP core bubble diameter is decreasing with depth. This is expected as the bubbles are compressed by the increasing pressure of the overlaying ice. Therefore, down to the YD-Preboreal transition at 532.6 m below surface we calculate the CBE for the PSU data from the linear regression in the PICE data (120-530 m). For samples below 532.6 m we use the corresponding average of bubble diameters in the PICE data.

4 Comparison of datasets

The sample sizes, extraction devices and measurement procedures are different at PICE and PSU. Correlation plots of the final data including all corrections from PSU are made to analyse their deviations from the TAC data obtained from the barometric method at PICE (supplementary section S4). The data sets show a good correlation with the vacuum volumetric data obtained at PICE. However, on individual data points differences can be significant resulting from up to one meter depth difference to the closest correspondent and rapid fluctuations of TAC. After applying all corrections, the pooled standard deviations of TAC for PICE, PSU-CH$_4$, and PSU-δ^{15}N are 6.67, 6.80, and 6.11 cm^3 kg^{-1}, respectively, excluding samples with obvious melt features. We observe no significant difference in the dispersion in the three datasets.

5 Results and discussion

The TAC data are presented on the RECAP GICC05 ice age time scale (Simonsen et al., 2019 and S5) in Fig. 2. The Holocene section of the record from 12 kyr (section a) shows a decrease of TAC to roughly 80 cm^3 kg^{-1} followed by an increase to present day values. The variations in that section are caused by melt layers as we will argue in the following. Similarly, section c which is part of the previous interglacial period (Eemian) is heavily affected by melt with TAC as low as 20 cm^3 kg^{-1}. In the first few meters of the record (section a) TAC is heavily affected by visible melt layers with TAC as low as 20 cm^3 kg^{-1}. Based on the melt fraction we will in the following reconstruct summer temperatures. The cold glacial section 115-12 kyr BP (second b) shows no signs of melt layers. However, fast variations occur which, as we will discuss below, seem related to rapid climate changes.

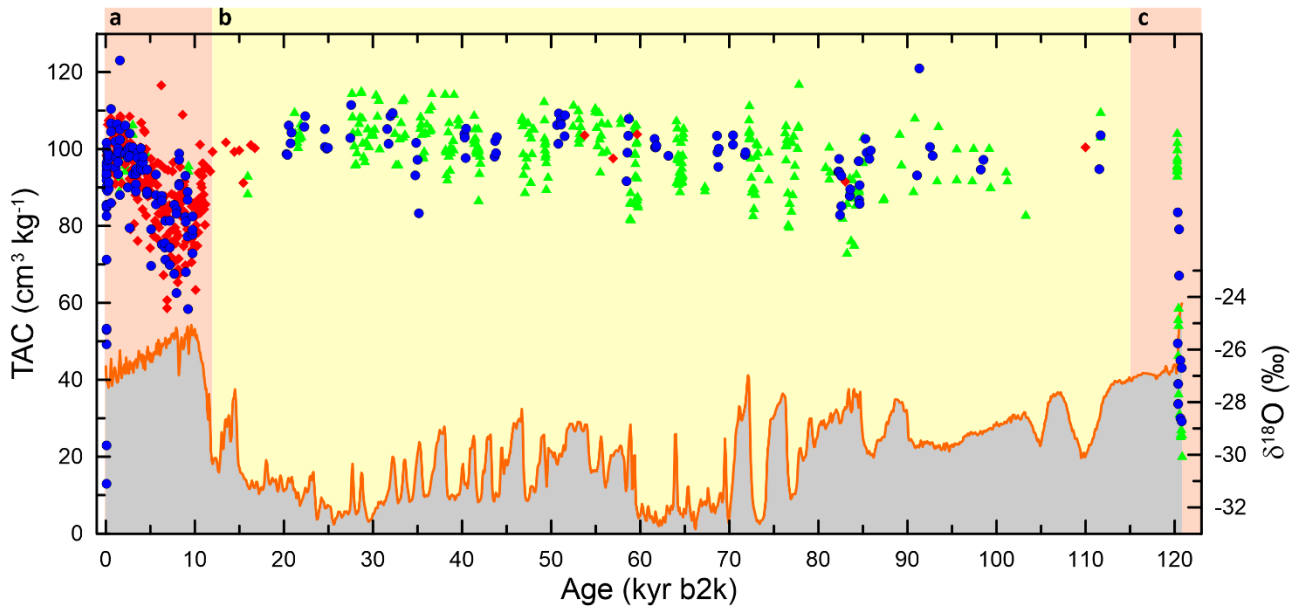



Figure 2: TAC and δ^{18}O of the RECAP ice core with sections a and c affected by melt (peach) and b unaffected by melt (light yellow). Top, TAC from PICE, PSU-CH$_4$, and PSU-δ^{15}N as blue dots, red diamonds, and green triangles, respectively. Bottom, RECAP δ^{18}O (red line) (Gkinis et al., 2024). The data are presented on the RECAP GICC05 ice time scale b2k (before A.D. 2000).

5.1 Holocene

As outlined in Vinther et al. (2009), it is expected that the altitude of the ice sheet was constant over the course of the Holocene. Consequently, we expect TAC to be constant, except for minor changes related to temperature and insolation. The expected TAC for present day Renland is 99 cm^3 kg^{-1} (see supplemental S6.1). From present day throughout the Holocene period (0 to 11.7 kyr b2k), we find TAC values are lower than expected, especially during the climatic optimum (6 ka to 9 kyr b2k) the values are as low as ~ 80 cm^3 kg^{-1} indicative of melt layers. Line scan can detect melt layers thicker than 2 mm and the RECAP line scan record indeed shows numerous melt layers (an example in Fig. 3). However,

observations of melt layers decrease with depth because they become quickly too thin to be detected (Taranczewski et al., 2019).

Any deviations from the expected near constant TAC are likely due to the presence of melt layers that form during periods of elevated summer temperatures. In the following we will determine the melt fraction in the RECAP TAC data and from 175 this estimate past summer temperatures.

First, we need to establish the expected TAC assuming no elevation change. Given by the ideal gas law TAC will change with temperature. This effect is below 1% for the Holocene period on Renland and we neglect it. The effect of insolation on the Greenland Holocene TAC has been estimated from data from NEEM, Camp Century, and GRIP (NEEM Community Members, 2013). The increased insolation at the beginning of the Holocene compared to today resulted in a 180 reduction in TAC of about 5 cm^3 kg^{-1}. We see the 5 cm^3 kg^{-1} change over the Holocene as a maximum for Renland. In fact, we suspect that Renland may experience very little insolation driven TAC change. The accumulation rate determines the exposure time of the surface layers to insolation which may result in more or less sensitivity of the O_2/N_2 ratio to insolation (Suwa and Bender, 2008) and also TAC. Given that Renland experiences more than double the accumulation rate than the central Greenland cores TAC may be significantly less affected (see S8 and subsequent discussion).

185

We derive the melt fraction by assuming a linear relationship between the TAC and the percentage melt in a sample (Herron and Langway Jr, 1987). The linear dependence is established by present day TAC of 99 cm^3 kg^{-1} for 0% melt and 21.5 cm^3 kg^{-1} for 100% melt. The latter is calculated with present day conditions and refrozen water equilibrated with the atmosphere based on Henry's solubility law (see S6.2 for details). We calculate the melt fraction from the insolation 190 corrected TAC (NEEM Community Members, 2013) as well as for the uncorrected TAC data from 100 year averaged TAC data (Fig. 4).

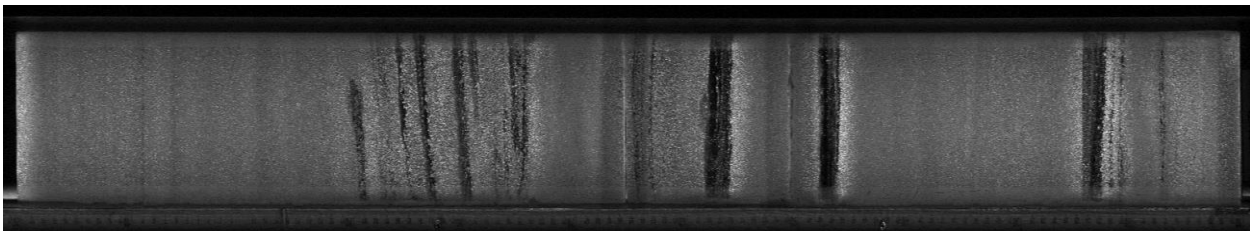



Figure 3: Line scan image of Bag 143 and 144 of the RECAP core showing melt layers.

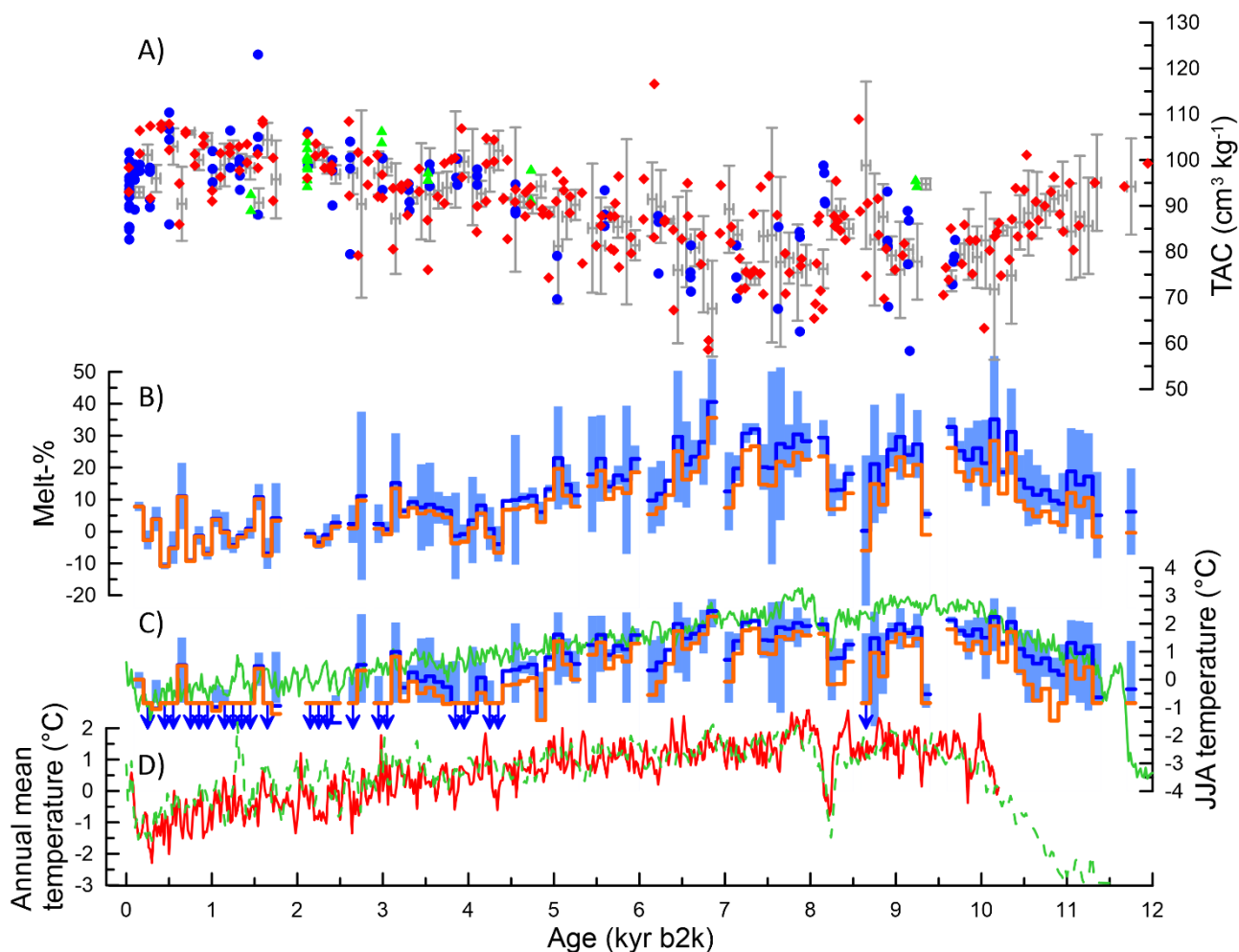



Figure 4: RECAP Holocene TAC with corresponding melt-%, and derived summer temperature all presented on the RECAP GICC05 ice time scale b2k (before A.D. 2000) together with other temperature estimates. A) TAC from PICE, PSU-CH₄, and PSU-δ¹⁵N as blue dots, red diamonds, and green triangles, respectively with 100-year averages and 1 sigma standard errors in gray. B) Melt fraction calculated from the 100-year averaged TAC as orange and dark blue step plots corrected and uncorrected for insolation effects, respectively. Light blue bar chart gives 1 sigma standard errors for the uncorrected melt fraction. C) Deviations from modern JJA temperatures calculated from the 100-year averaged melt fractions. 0 equals the average from 100-200 years b2k. Orange and dark blue step plots calculated from corrected and uncorrected melt fractions for insolation effects, respectively. Light blue bar chart gives 1 sigma standard errors for the uncorrected melt fraction. Melt-% below 2.5% indicate temperatures colder than -5.5°C according to the simulations. The arrow indicates that JJA temperatures are lower or equal to -5.5°C (corresponding to -0.84°C relative to present day as defined above). Solid green line JJA temperature calculations for Renland (Buizert et al., 2018). D) Red and green dotted line are Renland annual mean temperature reconstructions from Vinther et al. (2009) and Buizert et al. (2018), respectively.

5.1.1 Holocene summer temperatures inferred from melt fraction

We now use the estimated melt fraction to infer local summer temperature at Renland making use of an extension to the subsurface scheme of the HIRHAM5 regional climate model by Langen et al. (2017). The extrapolated temperatures from melt fractions suggest that Holocene summer temperatures in Renland were ~2 to 3°C warmer than the present day (Fig. 4). This is consistent with summer temperature reconstruction from Buizert et al. (2018). It is also consistent with the annual mean temperature reconstructions from the δ¹⁸O signals of Agassiz and Renland which reveal that, during the Holocene climatic optimum, Greenland temperatures were higher than the present day by ~2°C (Vinther et al., 2009). GRIP paleo temperatures interpreted from the δ¹⁸O profile and borehole temperature measurements also reveal that Greenland was warmer in the climatic optimum (8 kyr-10 kyr BP, boreal) by ~ 3 to 4°C (Johnsen et al., 1995; Dahl-Jensen et al., 1998). We note that melt layers are basically missing in the last 2kyr, increasing only in the last century. This is in line with the observations from Taranczewski et al. (2019) based on line scan images.

5.2 Previous interglacial (Eemian)

220 Greenland surface temperatures were warmer at the end of the previous interglacial period than in the Holocene. At NEEM, it is estimated that at 126 kyr b2k, the temperature peaked at 8±4 degrees Celsius above the mean of the past millennium (NEEM

225 Community Members, 2013). GISP 2 $\delta^{18}O$ records also indicate temperatures ~4 to 8°C warmer than the present around 126-128 kyr BP (Yau et al., 2016). The TAC signal of RECAP in the Eemian section (> 119 kyr b2k) has incongruously low values (as low

230 as ~ 20 cm^3 kg^{-1}, Fig. 5). It is likely that this is due to melt occurring due to increased temperatures. Applying the same metric as for the Holocene, the observed low TAC originates from temperatures at least 5°C warmer than today. This estimate

235 disregards insolation changes which are comparable to today 120 kyr ago. Higher temperature results in

Figure 5: RECAP section showing low gas content, starting around 555 m below surface. TAC from PICE, PSU-CH$_4$, and PSU-δ^{15}N as blue dots, red diamonds, and green triangles, respectively. RECAP δ^{18}O data (Gkinis et al., 2024) in red and on-line CH$_4$ data as black dots. The data are presented on depth with the RECAP GICC05 ice time scale b2k (before A.D. 2000) on top.

higher pore volume (Martinerie et al., 1994) resulting in higher TAC. For each degree increase, the pore volume becomes half a percent larger. Therefore, we see our estimate of 5°C warmer as a minimum. This assumes that none of the TAC changes are caused by elevation changes. If the higher temperature has led to a decrease of the Renland ice cap, TAC has

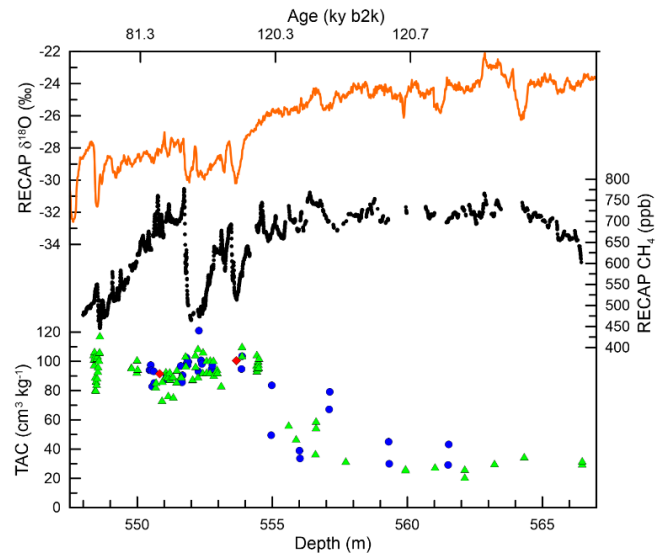
240 to increase, again making the estimated 5°C temperature change a minimum.

Generally, melt layers lead to spikes in the CH$_4$ record due to the higher solubility of methane compared to bulk air or in situ production (see e.g. NEEM Community Members, 2013). TAC from the NEEM ice core in the Eemian interglacial shows low TAC values, with spikes in the CH$_4$ and N$_2$O records which is a clear indication of the presence of surface melt (NEEM Community Members, 2013). Surprisingly, we do not see spikes in the on-line CH$_4$ record of RECAP in the

245 Eemian section (Fig. 5). The low gas content of the ice core in combination with the extremely low depth resolution (554-562 m corresponds to 119 to 120.8 kyr b2k) smoothed the CH$_4$ record. Potentially the melt spikes are just not visible any longer.

5.3 RECAP TAC during D-O events

250 The TAC of RECAP in the glacial section (11.7 to 119 kyr b2k) shows overall similar values as at present day (Fig. 2). However, like for NGRIP (Eicher et al., 2016), we find TAC variations associated with D-O events that are not related to elevation. Generally, in the vicinity of the D-O events, the RECAP TAC signal is dropping rapidly, recovering after a few hundred years (Fig. S9 a-d). The variations we see are on the order of 10-20%. If those changes were related to elevation changes, they would correspond to several hundred meter changes, which is unrealistic in only a couple of

255 hundred years. It is more likely that the changes are related to changes in pore volume. Similar effects have been observed in the NGRIP core (Eicher et al., 2016). An increase in temperature will with constant pore volume result in a small reduction of TAC. This effect is slow to take effect because changes in surface temperature must first reach the close-off depth through thermal diffusion. Once steady state is reached, the effect is counterbalanced by a slightly bigger pore volume (Martinerie et al., 1994). As for the NGRIP site (Eicher et al., 2016) we dismiss synoptic pressure changes as a

260   primary cause for the observed changes in TAC. To analyse the effect further and comparison to the NGRIP site, we took the following approach:

Dynamical effects in TAC can be expected from the moment of change till a new steady state is established. At a D-O event this is when the higher accumulation snow has reached close-off. To create a general picture of what is happening in the firn column, we decided to produce a stacked plot over D-O events for RECAP and also NGRIP. The time period

265   considered, corresponding to the time it takes for surface snow to arrive at close off is Δage. Methane and temperature changes have been found to occur in close temporal proximity during D-O events (e.g. Baumgartner et al., 2014). Changes in the methane concentration are recorded at the bottom of the firn column while other changes related to D-O events like $\delta^{18}O$ of H_2O or dust concentrations are recorded at the surface. Therefore, the depth interval to be considered for a dynamical firn change is between the depth when methane changes are observed and the depth where changes in

270   parameters recorded in the ice occur. For the RECAP ice core, we find that this depth interval corresponding to Δage is quite variable (see Fig. S9 a-d). We lack an understanding of why this is.

To create a general picture of what is happening in the firn column, we decided to produce a stacked plot over D-O events with a normalized time axis. For each event the time axis is normalized so that the methane transition (in some events defined by change in $\delta^{15}N$) is set to 1 and the decrease in dust (coincident with the change in $\delta^{18}O$) is set to 0. We treat

275   the Eicher et al. (2016) dataset for NGRIP in a similar way. The detailed results of this approach for RECAP and NGRIP can be found in the supplemental plots S10. The results for TAC are shown in Fig. 6 as a lowpass cubic spline fit with a 200-year cutoff period, according to Enting (1987) with 1 sigma uncertainties for the spline fit. The uncertainty is obtained by randomly varying the data points within their error before calculating 1000 Monte Carlo splines.

For both cores, on average, the TAC values start to decrease around the depth (time) when CH_4 starts to increase at the

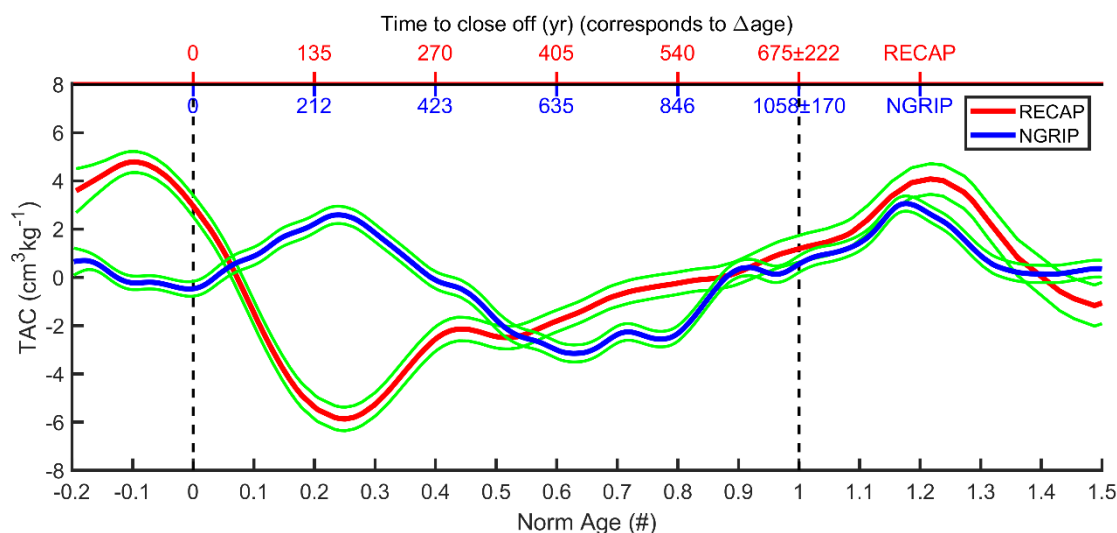
280   beginning of a D-O event. However, the minimum TAC is found before the depth (time) when D-O manifest as drop in dust or increase in $\delta^{18}O$. For NGRIP this minimum is reached some 600 years before the snow associated with the D-O event reaches close off while for RECAP it is about 150 years. Also, the drop in TAC is more significant for RECAP than for NGRIP. Overall, TAC variations associated with D-O events as recorded in RECAP are ~30% higher than in the NGRIP ice core. This may be a result of the higher accumulation and temperature at Renland.

285

Figure 6: Effect of D-O events on TAC signals for RECAP (red) and NGRIP (blue) on a normalized age/depth scale over the past firn column. At a D-O event transition, the surface at that time is at depth 0 and the corresponding close off is at depth 1. In other words, the zone between 0 and 1 is the past firn layer at a D-O event. On top is the duration it takes for the rapid change to reach close off. This corresponds to Δage. The uncertainty given ±222 and ±170 years for RECAP and NGRIP,

290   **respectively, is the standard deviation of the events considered for this stacked record (Fig. S10). Presented is the Enting spline (1987) with a 200 year cut off period with green lines giving the one sigma uncertainty of the spline.**

5.4 TAC and insolation

The influence of insolation on firn structure has been observed to be profound at Antarctic sites (e.g. Lipenkov et al., 2011; Bender, 2002). TAC records from Antarctic ice cores show a pronounced correlation with integrated summer insolation (ISI) as seen in EPICA DC (Raynaud et al., 2007) where the correlation is 0.86 (r^2), while the correlation of the Greenland NGRIP glacial TAC signal is only $r^2 = 0.3$ (Eicher et al., 2016).

For the RECAP core, the correlation (r^2) of spline of ISI (sum of annual insolation ≥ 380W/m^2) filtered with a cut off period (COP) of 3000 years (Enting, 1987) and the low pass spline of glacial TAC (11.7 to 119 kyr b2k) filtered with COP of 750 years is obtained as 0.004 (Fig. S11). As outlined earlier, we may see a pattern where higher accumulation rate, due to reduced exposure time of the firn structure, results in reduced influence of insolation on TAC. However, we do not rule out that the effect may be masked in the RECAP core by the high variability associated with D-O events. A higher sample resolution allowing to exclude data affected by rapid climate change from the analysis could clarify how accumulation rate and insolation interact.

5.5 Elevation calculation for the last glacial maximum

We finally make an attempt to use glacial TAC to reconstruct ice sheet elevation. To do so we need to avoid periods of rapid climate changes. Only the last glacial maximum section fulfils that criterion and has good data coverage (Fig. S12). For a meaningful interpretation of past elevation changes, TAC data generally need to be corrected for upstream flow, summer insolation influences, surface melting and effects of temporal variations. Since the RECAP ice core is drilled near the dome of the ice cap, upstream correction is not necessary. The melt affected TAC data in the Holocene and the last interglacial sections are not used for elevation calculations. No melting is expected in the glacial section.

As discussed in the previous section, since the RECAP TAC has negligible correlation with ISI, insolation correction may be unnecessary. We calculate elevation with and without accounting for insolation changes where we apply the TAC correction according to NEEM Community Members, 2013 (2013). From TAC the local ambient pressure (P_c) can be estimated where we need to estimate local temperature (T_c) and pore volume (V_c) applying Eq. 1. We estimate the past local temperature from NGRIP (Kindler et al., 2014) where the NGRIP record is increased by 13°C according to the present day difference between the NGRIP and RECAP sites. The average pressure of the 21 samples in the LGM period comes to 744 ± 5 mbar (1 standard error). The insolation correction for this period is -10 mbar. Uncertainty of T_c and V_c are significant. Each centigrade changes P_c by 4 mbar and 1% change in V_c results in a 7 mbar change in P_c.

The pressure P_c can now be interpreted in terms of elevation based on the barometric formula, Eq. (2). Unknowns are the past near surface lapse rate, and the pressure at sea level.

Along the Greenland ice sheet, the annual mean near surface lapse rate (dT/dz) has been calculated as -7.1 °C km^{-1} based on the data obtained from the 18 automatic weather stations for the period 1995-1999 (Steffen and Box, 2001). The lapse rate varies largely over the year from -4 °C km^{-1} in summer to -10 °C km^{-1} in winter (Steffen and Box, 2001). However, for present day Renland we calculate a near surface lapse rate of -4.5 °C km^{-1} where our point of reference is Illoqqortoormiut about 200km from RECAP with T_a of 265.5 °K, h_a=0 m, P_a =1012.2 mbar (Cappelen et al., 2001).

Our calculations are relative to the present sea level. Krinner et al. (2000) suggest that sea level pressure at current sea level was slightly higher than today during the LGM. From Fig. 2 in Krinner et al. (2000) this increase is between 0 and 5 mbar. In the following we disregard the uncertainty but increase the past sea level pressure to 1015 mbar. A model study on the LGM lapse rate concludes that the LGM was about 2°C km^{-1} lower than today (Erokhina et al., 2017). Based on the present day lapse rate we calculated a LGM lapse rate of -6.5°C km^{-1}. One could argue that the observed lapse rate for Renland of today is above the observation for Greenland since we measure the temperature in the RECAP firn. We do know that there is melting occurring today and we therefore may underestimate the annual mean temperature at Renland. Therefore, we also calculate with the lower lapse rate of -9.1 °C km^{-1} (again lowered by 2°C from modern).

Due to the topography of the ice sheet, the expectation is that the Renland ice sheet elevation is similar to today at 2340 m above present see level. Without insolation correction we calculate 2259 m and 2286 m for lapse rates of -7.1°C km^{-1} and -9.1°C km^{-1}, respectively. Including insolation correction, the numbers climb to 2354 m, and 2384 m for the two cases. The statistical uncertainty is ± 50 m (1 standard error) which does not include any uncertainty on V_c or T_c. E.g. including a 2°C uncertainty in T_c combined with a 2% uncertainty in V_c increases the uncertainty of the calculations to ± 220 m, enough that any of the four calculations covers the assumed ice sheet elevation of 2340 m above present see level for the LGM. Within errors, the RECAP site did not change significantly between the LGM and today. Hence, the results are coherent with the prime hypothesis in Vinther et al. (2009) that the Renland ice cap did not change elevation through time.

6 Conclusion

We measured TAC back to 121 kyr b2k from the Greenland RECAP ice core. The TAC signal has unexpectedly low values in the early Holocene (6 to 9 kyr b2k) and during the last interglacial (119 kyr to 121 kyr b2k). The low TAC values in the Holocene period point to melt events as corroborated by elevated CH_4 values in the RECAP core (Vladimirova, personal communication 2019). Melt fractions calculated from the RECAP TAC signal in the Holocene are in turn used to interpolate the summer surface temperatures (subsurface HIRHAM5 model). Summer temperatures in the early Holocene at Renland were ~2° to 3°C warmer compared to today. This finding is in agreement with similar findings from Greenland ice cores and model calculations. During the previous interglacial we see significant melting that let us conclude that temperatures at Renland were at least 5°C warmer than today.

The influence of local summer insolation ≥ 380 W/m^2 on the TAC signal of Renland is minimal as indicated by the correlation coefficient (r^2) of 0.004. Elevation of the Renland ice cap is calculated from the last glacial maximum TAC data. These elevation calculations encompass the uncertainties that arise from the assumption of the lapse rate, temperature and pressure gradients that existed in the past and sum up to ± 220 m. Within that large uncertainty the elevation has been similar to today. During D-O events, RECAP TAC shows significant variations that are larger than in other ice cores. How these variations come about is currently not understood. The stacked data analysis that we performed for RECAP and NGRIP show that changes in the firn structure must occur within the firn column during events of rapid climate change.

To predict how ice sheets will respond to warming, reconstructing their elevation history is crucial. However, recent findings cast doubt on the reliability of elevation reconstructions based on TAC - especially in Greenland over shorter millennial timescales. We need to explore the physical reasons behind short-term TAC changes and quantify insolation effects before we can confidently interpret TAC data for elevation shifts.

Appendix A: We use the estimated melt fraction to infer local summer temperature at Renland. Langen et al. (2017) extended the subsurface scheme of the HIRHAM5 regional climate model to include snow densification, varying hydraulic conductivity, irreducible water saturation and other effects on snow liquid water percolation and retention calculate melt water production. It allows us to derive an empirical relationship between melt fraction and temperature in the region. The model takes weather forcing at the surface from the regional climate model HIRHAM5 over the period 1980-2016 (forced in turn by ERA-Interim on the lateral boundaries) and calculates surface melting (of snow and bare ice), vertical percolation, retention, refreezing, densification, grain growth, runoff and surface mass balance. The subsurface calculations are performed on the 5.5x5.5 km grid of HIRHAM5. Over 15 grid cells centred on the RECAP drill site (Fig. A1), we gather total annual meltwater production and JJA average temperatures from each of the years 1980-2016 (giving 15 x 37=555 data points).

Meltwater production is converted into melt percentages using an observed approximate annual accumulation rate of 50 cm ice equivalent. The data points are then divided into 0.5 K bins with respect to JJA temperatures. For each bin, a mean melt percentage is calculated (Fig. A1). An exponential fit describes the resulting relation between JJA temperatures and melt percentages as: %-melt = b*exp(a·T), a= 0.6732, b = 179.02 where T is the mean JJA surface temperature in °C.

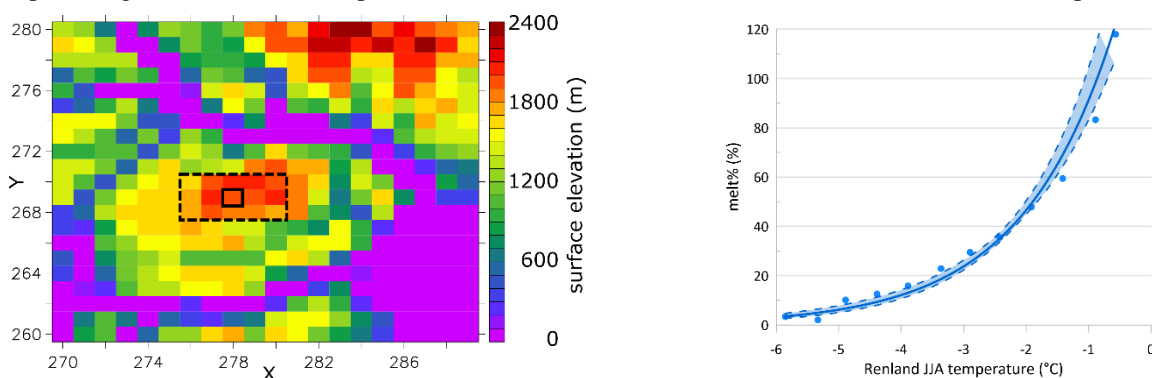



Figure A1: left) 15 grid cells (dashed square) centered on the RECAP drill site (solid square) on the subsurface calculations performed on the 5.5x5.5 km grid of HIRHAM5. right) Renland surface melt% versus JJA temperatures.

Acknowledgements: The RECAP ice coring effort was financed by the Danish Research Council through a Sapere Aude grant, the NSF through the Division of Polar Programs, the Alfred Wegener Institute, and the European Research Council under the European Community's Seventh Framework Program (FP7/2007-2013) / ERC grant agreement 610055 through the Ice2Ice project. The Centre for Ice and Climate is funded by the Danish National Research Foundation. PLL gratefully acknowledges the contributions of Aarhus University Interdisciplinary Centre for Climate Change (iClimate, Aarhus University). Thomas Blunier acknowledges support from the Carlsberg Foundation and Australian Antarctic Program Partnership. We thank Todd Sowers for measurements at PSU and for decades of fruitful collaborations.

7 Author contributions.

Sindhu Vudayagiri: TAC measurements at PICE, data collection, data analysis and drafted the manuscript.
Johannes Freitag: Line scan images.
Peter L. Langen: Simulated JJA surface temperatures based on the meltwater production (subsurface model calculations performed on the 5.5x5.5 km grid of HIRHAM5) in Renland.
Bo Vinther: Contributed to data analysis and manuscript preparation.
Thomas Blunier: Designed the experiments, made the final data analysis, and final manuscript preparation.

Data availability. Data not corrected for CBE from PSU-δ^{15}N measurements can be found here (Sowers, 2018). The full dataset is available at the Arctic Data Center (Blunier, 2024).

Competing interests. The contact author has declared that none of the authors has any competing interests.

8 References

Baumgartner, M., Kindler, P., Eicher, O., Floch, G., Schilt, A., Schwander, J., Spahni, R., Capron, E., Chappellaz, J., Leuenberger, M., Fischer, H., and Stocker, T. F.: NGRIP CH$_4$ concentration from 120 to 10 kyr before present and its relation to a δ^{15}N temperature reconstruction from the same ice core, Clim Past, 10, 903-920, 10.5194/cp-10-903-2014, 2014.

Bender, M. L.: Orbital tuning chronology for the Vostok climate record supported by trapped gas composition, Earth Planet Sc Lett, 204, 275-289, 2002.

Blunier, T.: Renland Ice Cap Project (ReCAP) total air content data, Arctic Data Center, 10.18739/A2ZK55P0N, 2024.

410 Buizert, C., Keisling, B. A., Box, J. E., He, F., Carlson, A. E., Sinclair, G., and DeConto, R. M.: Greenland-Wide Seasonal Temperatures During the Last Deglaciation, Geophys Res Lett, 45, 1905-1914, 10.1002/2017gl075601, 2018.

Cappelen, J., Jørgensen, B. V., Laursen, E. V., Stannius, L. S., and Thomsen, R. S.: The Observed Climate of Greenland , 1958-99 - with Climatological Standard Normals , 1961-90 Klimaobservationer i Grønland , 1958-99, Danish Meteorological Institute, CopenhagenISSN 1399-1388, 151, 2001.

415 Dahl-Jensen, D., Mosegaard, K., Gundestrup, N. S., Clow, G. D., Johnsen, S. J., Hansen, A. W., and Balling, N.: Past temperatures directly from the Greenland ice sheet, Science, 282, 268-271, 1998.

Eicher, O., Baumgartner, M., Schilt, A., Schmitt, J., Schwander, J., Stocker, T. F., and Fischer, H.: Climatic and insolation control on the high-resolution total air content in the NGRIP ice core, Clim. Past, 12, 1979-1993, 10.5194/cp-12-1979-2016, 2016.

420 Enting, I. G.: On the use of smoothing splines to filter CO_2 data, J Geophys Res-Atmos, 92, 10977-10984, 10.1029/JD092iD09p10977, 1987.

Epifanio, J. A., Brook, E. J., Buizert, C., Pettit, E. C., Edwards, J. S., Fegyveresi, J. M., Sowers, T. A., Severinghaus, J. P., and Kahle, E. C.: Millennial and orbital-scale variability in a 54,000-year record of total air content from the South Pole ice core, The Cryosphere, 17, 4837-4851, 10.5194/tc-17-4837-2023, 2023.

425 Erokhina, O., Rogozhina, I., Prange, M., Bakker, P., Bernales, J., Paul, A., and Schulz, M.: Dependence of slope lapse rate over the Greenland ice sheet on background climate, Journal of Glaciology, 63, 568-572, 10.1017/jog.2017.10, 2017.

Fegyveresi, J. M.: Physical properties of the West Antarctic Ice Sheet (WAIS) Divide deep core: Development, evolution, and interpretation, Ph.D. thesis, Department of Geosciences, The Pennsylvania State University, Ann Arbor, 304 pp., 430 2015.

Gkinis, V., Vinther, B. M., Cook, E., Freitag, J., Holme, C. T., Hughes, A. G., Kipfstuhl, S., Kjær, H. A., Maffrezzoli, N., Morris, V., Popp, T. J., Rasmussen, S. O., Simonsen, M. F., Svensson, A. M., Vallelonga, P. T., Vaughn, B. H., White, J. W. C., Winstrup, M., and Hansen, S. B.: An ultra-high resolution water isotope record (δ^{18}O, δD) from the Renland ice cap spanning 120,000 years of climate history, Pangaea, doi.pangaea.de/10.1594/PANGAEA.966693, 435 2024.

Herron, S. L., and Langway Jr, C. C.: Derivation of paleoelevations from total air content of two deep Greenland ice cores, The Physical Basis of Ice Sheet Modelling Proceedings of a symposium held during the XIX Assembly of the International Union of Geodesy and Geophysics, Vancouver, 1987, 283-295,

Hörhold, M. W., Kipfstuhl, S., Wilhelms, F., Freitag, J., and Frenzel, A.: The densification of layered polar firn, J. 440 Geophys. Res., 116, F01001, 10.1029/2009JF001630., 2011.

Johnsen, S., Dahl-Jensen, D., Dansgaard, W., and Gundestrup, N.: Greenland palaeotemperatures derived from GRIP bore hole temperature and ice core isotope profiles, Tellus, 47B, 624-629, 1995.

Johnsen, S. J., Clausen, H. B., Dansgaard, W., Gundestrup, N. S., Hanson, M., Jonson, P., Steffensen, J. P., and Sveinbjörnsdottir, A. E.: A "deep" ice core from East Greenland, Meddelelser om Grønland, Geoscience, 29, 22, 1992.

445 Kawamura, K., Parrenin, F., Lisiecki, L., Uemura, R., Vimeux, F., Severinghaus, J. P., Hutterli, M. A., Nakazawa, T., Aoki, S., Jouzel, J., Raymo, M. E., Matsumoto, K., Nakata, H., Motoyama, H., Fujita, S., Goto-Azuma, K., Fujii, Y., and Watanabe, O.: Northern Hemisphere forcing of climatic cycles in Antarctica over the past 360,000 years, Nature, 448, 912-U914, 2007.

Kindler, P., Guillevic, M., Baumgartner, M., Schwander, J., Landais, A., and Leuenberger, M.: Temperature 450 reconstruction from 10 to 120 kyr b2k from the NGRIP ice core, Clim Past, 10, 887-902, 10.5194/cp-10-887-2014, 2014.

Krinner, G., Raynaud, D., Doutriaux, C., and Dang, H.: Simulations of the Last Glacial Maximum ice sheet surface climate: Implications for the interpretation of ice core air content, J Geophys Res-Atmos, 105, 2059-2070, 10.1029/1999jd900399, 2000.

455 Langen, P. L., Fausto, R. S., Vandecrux, B., Mottram, R. H., and Box, J. E.: Liquid Water Flow and Retention on the Greenland Ice Sheet in the Regional Climate Model HIRHAM5: Local and Large-Scale Impacts, Frontiers in Earth Science, 4, 1-18, 10.3389/feart.2016.00110, 2017.

Lipenkov, V., Candaudap, F., Ravoire, J., Dulac, E., and Raynaud, D.: A New Device for the Measurement of Air Content in Polar Ice, Journal of Glaciology, 41, 423-429, 1995.

460 Lipenkov, V. Y., Raynaud, D., Loutre, M. F., and Duval, P.: On the potential of coupling air content and O_2/N_2 from trapped air for establishing an ice core chronology tuned on local insolation, Quaternary Sci Rev, 30, 3280-3289, 10.1016/j.quascirev.2011.07.013, 2011.

Martinerie, P., Lipenkov, V., and Raynaud, D.: Correction of air content measurements in polar ice for the effect of cut bubbles at the surface of the sample, Journal of Glaciology, 36, 299-303, Doi 10.3189/002214390793701282, 1990.

465 Martinerie, P., Raynaud, D., Etheridge, D. M., Barnola, J.-M., and Mazaudier, D.: Physical and climatic parameters which influence the air content in polar ice, Earth Planet Sc Lett, 112, 1-13, 1992.

Martinerie, P., Lipenkov, V. Y., Raynaud, D., Chappellaz, J., Barkov, N. I., and Lorius, C.: Air content paleo record in the Vostok ice core (Antarctica): A mixed record of climatic and glaciological parameters, J Geophys Res-Atmos, 99, 10565-10576, 10.1029/93jd03223, 1994.

470 NEEM Community Members: Eemian interglacial reconstructed from a Greenland folded ice core, Nature, 493, 489-494, 10.1038/nature11789, 2013.

Raynaud, D., Chappellaz, J., Ritz, C., and Martinerie, P.: Air content along the Greenland Ice Core Project core: A record of surface climatic parameters and elevation in central Greenland, J Geophys Res, 102, 26607-26614, 1997.

475 Raynaud, D., Lipenkov, V., Lemieux-Dudon, B., Duval, P., Loutre, M. F., and Lhomme, N.: The local insolation signature of air content in Antarctic ice. A new step toward an absolute dating of ice records, Earth Planet Sc Lett, 261, 337-349, 10.1016/j.epsl.2007.06.025, 2007.

Shoji, H., and Langway, C. C., Jr.: Air hydrate inclusions in fresh ice core, Nature, 298, 548-550, 1982.

Simonsen, M. F., Baccolo, G., Blunier, T., Borunda, A., Delmonte, B., Frei, R., Goldstein, S., Grinsted, A., Kjaer, H. A., Sowers, T., Svensson, A., Vinther, B., Vladimirova, D., Winckler, G., Winstrup, M., and Vallelonga, P.: East

480 Greenland ice core dust record reveals timing of Greenland ice sheet advance and retreat, Nature Communications, 10, 10.1038/s41467-019-12546-2, 2019.

Sowers, T.: Elemental and isotopic composition of oxygen (O_2) and nitrogen (N_2) in the Renland Ice Cap (RECAP) ice core, Arctic Data Center, 10.18739/A2C824F12, 2018.

Steffen, K., and Box, J.: Surface climatology of the Greenland Ice Sheet: Greenland Climate Network 1995–1999, J.

485 Geophys. Res. Atmos., 106, 33951-33964, 10.1029/2001jd900161, 2001.

Suwa, M., and Bender, M. L.: O_2/N_2 ratios of occluded air in the GISP2 ice core, J Geophys Res-Atmos, 113, 2008.

Taranczewski, T., Freitag, J., Eisen, O., Vinther, B., Wahl, S., and Kipfstuhl, S.: 10,000 years of melt history of the 2015 Renland ice core, EastGreenland, The Cryosphere Discuss., 2019, 1-16, 10.5194/tc-2018-280, 2019.

Uchida, T., and Hondoh, T.: Laboratory studies on air-hydrate crystals, in: Physics of ice core records, edited by: Hondoh,

490 T., Hokkaido University Press, Sapporo, 423-457, 2000.

Vinther, B. M., Buchardt, S. L., Clausen, H. B., Dahl-Jensen, D., Johnsen, S. J., Fisher, D. A., Koerner, R. M., Raynaud, D., Lipenkov, V., Andersen, K. K., Blunier, T., Rasmussen, S. O., Steffensen, J. P., and Svensson, A. M.: Holocene thinning of the Greenland ice sheet, Nature, 461, 385-388, 10.1038/nature08355, 2009.

Yau, A. M., Bender, M. L., Robinson, A., and Brook, E. J.: Reconstructing the last interglacial at Summit, Greenland:

495 Insights from GISP2, Proc Natl Acad Sci U S A, 113, 9710-9715, 10.1073/pnas.1524766113, 2016.