

A Comprehensive Global Modelling Assessment of Nitrate Heterogeneous Formation on Desert Dust

Rubén Soussé Villa¹, Oriol Jorba¹, María Gonçalves Ageitos¹, Dene Bowdalo¹, Marc Guevara¹, and Carlos Pérez García-Pando^{1,2}

¹Barcelona Supercomputing Center, Barcelona, Spain

²Catalan Institution for Research and Advanced Studies (ICREA), Barcelona, Spain

Correspondence: Rubén Soussé (ruben.sousse@bsc.es), Oriol Jorba (oriol.jorba@bsc)

Abstract.

Desert dust undergoes complex heterogeneous chemical reactions during atmospheric transport, forming nitrate coatings that influences hygroscopicity, gas partitioning, optical properties, and aerosol radiative forcing. Contemporary atmospheric chemistry models show significant disparities in aerosol nitrogen species due to varying parameterizations and inaccuracies in representing heterogeneous chemistry and dust alkalinity. This study investigates key processes in nitrate formation over dust and evaluates their representation in models. We incorporate varying levels of dust heterogeneous chemistry complexity into the MONARCH model, assessing sensitivity to critical processes. Our analyses address the condensation pathways of gas species onto dust (irreversible and reversible), the influence of nitrate representation on species burdens, lifetimes and size distribution, and the role of alkalinity. Using annual global simulations, we compare particulate and gas species surface concentrations against observations, and evaluate global budgets and spatial distributions. Findings show significant outcome dependence on methodology, particularly on reversible vs. irreversible gas condensation on dust, with wide ranges for particulate nitrate burdens (0.66 to 1.93 Tg) and correlations with observations (0.66 to 0.91). In contrast, particulate ammonium burdens show lower variability (0.19 to 0.31 Tg). Incorporating dust (together with sea-salt) alkalinity improves consistency with observations, with reversible condensation along with alkalinity representation yielding the best agreement, while showing consistent gas and particle partitioning. In contrast, irreversible uptake reactions overestimate coarse particulate nitrate formation. Our findings offer guidelines for integrating nitrate heterogeneous formation on dust in models, paving the road for improved estimates of aerosol radiative effects.

1 Introduction

Desert dust is produced by wind erosion of arid and semi-arid surfaces, contributing approximately 40% of the total dry aerosol mass globally, and between 70% and 80% if sea-salt aerosol (SS) is not considered (Boucher et al., 2013; Adebisi and Kok, 2020). Dust interacts with shortwave and longwave radiation, affects cloud formation and alters atmospheric composition, modifying the Earth's energy and water cycles (Pérez et al., 2011; Boucher et al., 2013; Randall et al., 2013). When deposited, dust also affects biogeochemical cycles both on ocean and continental areas (Mahowald et al., 2014; Li et al., 2016; Bergas-Massó et al., 2023). If inhaled, dust can be potentially harmful for animal and human health (Usher et al., 2003). Temporal

25 variations in dust emissions, from interannual to geological time scales, have been a key driver of the past climate of the Earth, as observed in ice cores and ocean sediment samples (Rea, 1994). All these considerations make desert dust particles a key component of the Earth system influencing climate (Semeniuk and Dastoor, 2020).

Climate perturbations by dust depend fundamentally upon dust particles' physical and chemical properties. These properties are mainly the particle size distribution (PSD), shape, surface characteristics, mineral composition, and mixing state (Usher et al., 2003; Riemer et al., 2019). These characteristics depend on the dust source region and on its chemical transformations while transported in the atmosphere (Claquin et al., 1999). These two factors determine the final dust optical properties, and consequently its radiative forcing.

A particularly significant factor driving the chemical evolution of dust in the atmosphere is *heterogeneous chemistry*: Chemical reactions involving more than one matter phase (e.g. gas, liquid and solid particles) that might occur in the surfaces of aerosol particles or within their liquid phases (Schwartz, 1986; Dentener et al., 1996; Usher et al., 2003; Bauer et al., 2004, 2007; Riemer et al., 2019).

For example, the condensation of atmospheric gas species on liquid or solid particles are key for particle growth and changes in optical properties during atmospheric transport (Vlasenko et al., 2009; Fairlie et al., 2010; Karydis et al., 2016). Heterogeneous reactions involving nitrogen, for example, can even cause dust to act as a transport medium of nitrates from nitrate-rich areas to regions downwind of dust sources (Ma et al., 2021).

Heterogeneous reactions mainly occur when dust mixes with anthropogenic pollutants emitted in urban and industrial areas. Nitric acid ($\text{HNO}_{3(g)}$), ammonia ($\text{NH}_{3(g)}$), sulfur dioxide ($\text{SO}_{2(g)}$) and sulfuric acid ($\text{H}_2\text{SO}_{4(g)}$) are the most important anthropogenic species that react with dust (Usher et al., 2003; Yue et al., 2022). These interactions lead to: 1) the formation of aqueous coatings around the particles (Usher et al., 2003; Krueger et al., 2003, 2004; Fountoukis and Nenes, 2007; Li and Shao, 2009) and 2) the reaction of gases with the non-volatile cations (NVC) present both at the particles' surfaces and dissolved in their liquid envelope (Dentener et al., 1996; Goodman, 2000; Usher et al., 2003; Li and Shao, 2009). These processes transfer mass from the gas to the aerosol phase, either through irreversible reactions of low volatile gas vapors with the bulk material of the particle or reversible condensation-evaporation processes between the gas phase and the liquid coating (Usher et al., 2002, 2003; Krueger et al., 2003; Crowley et al., 2010).

Each gas species' chemical transformation follows a distinct pathway depending on its solubility and reactivity with other atmospheric species dissolved in the particle's liquid coating and with the NVC in dust. Among these, $\text{HNO}_{3(g)}$ plays a major role in dust heterogeneous chemistry due to its relatively high solubility and reactivity with ammonium (NH_4^+) from dissolved $\text{NH}_{3(g)}$, leading to the formation of ammonium nitrate (NH_4NO_3) (Metzger et al., 2002; Usher et al., 2003). Aqueous $\text{H}_2\text{SO}_{4(g)}$ is also paramount, competing with dissolved $\text{HNO}_{3(g)}$ to neutralize NH_4^+ . However, the neutralization by $\text{H}_2\text{SO}_{4(g)}$ predominates due to its very low vapor pressure, preventing its evaporation back to the gas phase and resulting in the formation of ammonium sulfate ($(\text{NH}_4)_2\text{SO}_4$) (Usher et al., 2003; Uno et al., 2020). These are the main formation pathways of particulate nitrate (NO_3^-), particulate ammonium (NH_4^+) and particulate sulfate (SO_4^{2-}) from the HNO_3 - NH_3 - H_2SO_4 neutralization system in the particles' liquid coating. Additionally, $\text{SO}_{2(g)}$ is relevant in the aqueous medium as it converts to $\text{H}_2\text{SO}_{4(g)}$ through oxidation with O_3 and H_2O_2 (Seinfeld and Pandis, 1998; Usher et al., 2002).

60 The dust minerals also provide additional reactive surfaces to neutralize $\text{HNO}_{3(\text{g})}$, with their reactivity depending on their solubility and environmental relative humidity (RH). Under low RH, solid minerals may serve as active sites for reactions with gas species on the particle surfaces (Usher et al., 2003). Conversely, at high RH, minerals can dissociate in the aqueous medium, releasing NVC such as Ca^{2+} , K^+ , Mg^{2+} and Na^+ . The active sites in dust minerals and NVC play a key role in neutralizing dissolved $\text{HNO}_{3(\text{g})}$ and forming NO_3^- compounds (i.e. $\text{Ca}(\text{NO}_3)_{2(\text{a})}$, $\text{Mg}(\text{NO}_3)_{2(\text{a})}$, $\text{KNO}_{3(\text{a})}$ and $\text{NaNO}_{3(\text{a})}$)

65 (Fenter et al., 1995; Krueger et al., 2004). Consequently, these nitrate salts may form on the surface of the dust under low RH or dissociate in the particle's liquid coating (Usher et al., 2003; Jones et al., 2021). These reactions represent the primary pathways for NO_3^- formation in the presence of dust, and are highly sensitive to NO_x and sulfate concentrations (Fenter et al., 1995; Riemer et al., 2003). For instance, the combination of reduced sulfate emissions along with unchanged or even enhanced $\text{NH}_{3(\text{g})}$ emissions, as predicted by 21st century emission scenarios, implies a reduction in particle acidity (Bauer et al., 2007;

70 Bellouin et al., 2011; Boucher et al., 2013; Hauglustaine et al., 2014; Bian et al., 2017; Karydis et al., 2021). This scenario, alongside a potential rise in dust (Usher et al., 2002; Adebisi et al., 2023), would lead to an increase in NO_3^- formation, especially in the fine mode, if NO_x emissions are not concurrently reduced (Bauer et al., 2016; Bian et al., 2017; Zaveri et al., 2021). Therefore, accurate modelling of dust heterogeneous chemistry in atmospheric models is important for present and future air quality control (Myhre et al., 2006).

75 In the last decades, several approximations have been introduced in atmospheric chemistry models to address nitrate heterogeneous reactions on both dust and SS, with a particular focus on $\text{HNO}_{3(\text{g})}$ condensation. These approaches range from dynamic mass transfer (DMT) calculations between gas and aerosol phases (Meng and Seinfeld, 1996; Lurmann et al., 1997; Song and Carmichael, 2001; Feng and Penner, 2007; Zaveri et al., 2008; Trump et al., 2015), to the assumption that the bulk gas-aerosol phases instantly reach thermodynamic equilibrium (TEQ) and calculating their correspondent concentrations

80 (Lurmann et al., 1997). While DMT can accurately capture processes far from TEQ (e.g. the condensation of gas species at low temperatures, under extreme RH conditions or onto coarse particles), the inherent stiffness of inorganic heterogeneous chemistry renders DMT a rigorous but computationally-expensive methodology (Feng and Penner, 2007; Zaveri et al., 2008; Trump et al., 2015; Benduhn et al., 2016). On the other hand, assuming instantaneous TEQ is more efficient and has gained popularity despite its tendency to overestimate coarse nitrate formation (Nenes et al., 1998; Feng and Penner, 2007; Bauer et al., 2007; Hauglustaine et al., 2014; Paulot et al., 2016; Bian et al., 2017). To balance accuracy and computational cost, several intermediate strategies have been developed, including 1) simplifying the DMT equations to a first-order irreversible uptake reaction (UPTK), which ignores the evaporation of taken up species back to the gas phase (Jacob, 2000; Bauer et al., 2004; Feng and Penner, 2007; Fairlie et al., 2010), 2) the double call of the TEQ concentrations calculation for both fine (diameter up to $2.5\mu\text{m}$) and coarse (diameter above $2.5\mu\text{m}$) modes of dust and SS (double call method, DBCLL), either after

90 kinetically limiting the gas condensing on each bin or mode (Pringle et al., 2010) or redistributing the condensed mass from the bulk TEQ using kinetic coefficients (Karydis et al., 2016), and 3) employing an hybrid approach (HYB) that applies TEQ to the fine bins or modes and UPTK to the coarse ones (Capaldo et al., 2000; Hodzic et al., 2006; Hauglustaine et al., 2014; Trump et al., 2015). Overall, methods involving DMT or TEQ calculations allow to simulate the reversible heterogeneous

reactions (condensation-evaporation dynamics), and the UPTK calculates the irreversible uptake of gas species, accounting for gas specifications, particle's alkalinity and environmental RH (Fairlie et al., 2010; Paulot et al., 2016).

Despite these efforts to incorporate nitrate heterogeneous reactions on coarse particles, atmospheric models still significantly diverge in their predictions of the tropospheric burden of oxidized ($\text{HNO}_{3(g)} + \text{NO}_3^-$) and reduced nitrogen ($\text{NH}_{3(g)} + \text{NH}_4^+$), often struggling to reproduce observational data of these species (Fairlie et al., 2010; Hauglustaine et al., 2014; Paulot et al., 2016; Zakoura and Pandis, 2018; Luo et al., 2019; Jones et al., 2021; Rémy et al., 2022). For instance, the particulate nitrate AeroCom phase III experiment (Bian et al., 2017), an extensive intercomparison study of atmospheric models incorporating NO_3^- formation processes on dust and SS, highlights substantial disagreements. The average NO_3^- atmospheric burden among models is 0.63 Tg, with a standard deviation of 0.56 Tg, nearly 90% of the mean value. Similar variability is observed for NH_4^+ (0.32 ± 0.20 Tg). The study also highlights the general inaccuracy of current models in reproducing observations of NO_3^- concentrations after long-range transport of precursor species, indicating that nitrogen heterogeneous chemistry processes on dust and SS are often misrepresented in models.

The scope of the present work is to understand the role of dust in NO_3^- formation through a systematic investigation of the underlying processes governing dust heterogeneous chemistry. To achieve this goal, we incorporate a variety of mechanisms of different complexities into a global model. This enables a comprehensive analysis of the partitioning between gas and aerosol phases, the suitability of irreversible and reversible parameterizations for the condensation of gas species on dust, and the role of explicit representation of alkalinity. While our primary emphasis is on the heterogeneous chemistry on dust surfaces, we also account for nitrate formation on SS and its alkalinity.

This paper is structured as follows. Section 2 introduces the MONARCH atmospheric chemistry model (2.1), detailing the specific developments implemented for this study (2.2), the setup of the simulations conducted (2.5) and the datasets used for evaluation (2.6). Section 3 presents an analysis of the global simulations and their evaluation against observational data. This section includes a comparison of the spatial distributions and an examination of the total nitrogen burden and gas/particle partitioning. Additionally, we discuss the budgets of reduced and oxidized nitrogen species, depositions, production/loss rates and lifetimes. Our results are contextualized with findings from previous studies, providing a comprehensive understanding of the results. Section 4 provides a summary of our key findings.

2 Methods

2.1 The MONARCH model

MONARCH is an atmospheric chemistry model developed at the Earth Sciences department of the Barcelona Supercomputing Center (Pérez et al., 2011; Jorba et al., 2012; Badia et al., 2017; Klose et al., 2021; Gonçalves Ageitos et al., 2023; Navarro-Barboza et al., 2024). It simulates the atmospheric life cycle of aerosol and gas-phase species through an online coupling with the Nonhydrostatic Multiscale Model on the B-grid (NMMB) (Janjic and Gall, 2012). NMMB allows running both global and regional atmospheric simulations with embedded telescoping nests. The Arakawa B grid is used in the horizontal direction and the Lorenz hybrid pressure-sigma coordinate in the vertical direction. MONARCH global simulations use a

regular latitude-longitude grid with polar filtering and a rotated longitude-latitude grid is adopted for regional applications. The NMMB numerical schemes are based on principles described in Janjic and Gall (2012). The physical parameterizations used in the model include (1) a surface layer scheme based on the Monin-Obukhov similarity theory (Monin and Obukhov, 1954) combined with a viscous sublayer on continental and water surfaces (Zilitinkevich, 1965; Janjic, 1984, 1996), (2) the Mellor–Yamada–Janjic (MYJ) planetary boundary layer and the free troposphere turbulence scheme (Janjić, 2001), (3) the Unified NCEP-NCAR-AFWA Noah land surface model (Ek et al., 2003) to compute the surface heat and moisture fluxes, (4) the 1D Rapid Radiative Transfer Model for Global circulation Models (RRTMG) (Iacono et al., 2008) for the calculation shortwave and longwave radiative fluxes, (5) the Ferrier microphysics scheme (Ferrier et al., 2002) for grid-scale clouds, and (6) the Betts–Miller–Janjic convective clouds scheme (Betts and Miller, 1986; Emanuel and Živković Rothman, 1999; Janjić, 2000). The same advection and vertical mixing schemes formulated in NMMB are used for both meteorological and chemistry species for consistency.

MONARCH includes a gas-phase module combined with a hybrid sectional–bulk multi-component mass-based aerosol module. The gas-phase chemistry is based on the Carbon Bond 2005 (CB05) chemical mechanism extended with chlorine chemistry (Yarwood et al., 2005; Whitten et al., 2010) designed to describe urban to remote tropospheric conditions. The photolysis rates are computed using the Fast-J scheme (Wild et al., 2000) accounting for aerosols, clouds, and absorbers such as ozone. A resistance approach is adopted for dry deposition (Wesely, 1989), and in-cloud and below-cloud scavenging and wet deposition follow Byun (1999) and Foley et al. (2010).

The aerosol representation in MONARCH considers 8 main components, namely dust, SS, black carbon, organic matter (both primary and secondary), SO_4^{2-} , NH_4^+ , NO_3^- and unspeciated aerosol mass. Mineral dust and SS are described with a sectional size distribution of 8 bins spanning from $0.2\mu\text{m}$ to $20\mu\text{m}$ and $30\mu\text{m}$ in diameter, respectively. All the other aerosol components are represented by a fine mode, except NO_3^- , represented with both fine and coarse modes to consider the condensation of $\text{HNO}_{3(\text{g})}$ on coarse particles. Table S2 in the Supplementary material reports the bin volumetric and effective radii, density and their fractional contribution to $\text{PM}_{2.5}$ and PM_{10} .

Black carbon is represented in two primary modes: hydrophobic and hydrophilic, with 80% of its emitted mass initially classified as hydrophobic. During atmospheric transport, an aging process with an e-folding time of 1.2 days facilitates the transition of mass from the hydrophobic to the hydrophilic mode (Chin et al., 2002). Organic aerosols are represented using the simplified scheme of Pai et al. (2020), which assumes fixed secondary organic aerosol (SOA) yields calibrated to align with more complex volatility-based approaches. For primary organic aerosols, a dual-mode representation—hydrophobic and hydrophilic—is adopted, similar to black carbon. In this scheme, 50% of the emitted mass is classified as hydrophobic, with an aging e-folding time of 1.15 days transitioning it to the hydrophilic mode. Originally designed for global models, this approach has also demonstrated strong performance at regional scales (Navarro-Barboza et al., 2024).

A simplified gas–aqueous–aerosol mechanism accounts for sulfur chemistry through the oxidation of $\text{SO}_{2(\text{g})}$ and dimethyl sulfide (DMS) to $\text{H}_2\text{SO}_{4(\text{g})}$. At the end of each chemistry integration time step, the remaining $\text{H}_2\text{SO}_{4(\text{g})}$ that has not formed aqueous sulfate is assumed to fully nucleate into fine particulate SO_4^{2-} (Spada, 2015). The heterogeneous hydrolysis of $\text{N}_2\text{O}_{5(\text{g})}$ on aqueous sulfate particles is included to account for additional $\text{HNO}_{3(\text{g})}$ formation, following the formulation of Riemer et al.

(2003). Prior to this study, secondary nitrate–ammonium aerosols were modeled using the TEQ module EQuilibrium Simplified Aerosol Model version v03b (EQSAM v03b, Metzger et al. (2002)) for fine particles. It is important to note that EQSAM v03b exclusively considers sulfate-nitrate-ammonium partitioning and does not account for the presence of other species (i.e. Dust or SS alkalinity). In Section 2.2.2 we detail the adoption of ISORROPIA-II v1 (Fountoukis and Nenes, 2007) for this study. To account for secondary nitrate aerosol formation on coarse dust and SS particles, an hybrid (HYB) approach was employed through an uptake reaction (UPTK) of $\text{HNO}_{3(g)}$. The reaction uses the uptake rate (K) defined by Jacob (2000) as a first-order function (Schwartz, 1986):

$$K = \left(\frac{r}{D_g} + \frac{4}{v\gamma} \right)^{-1} \cdot S \quad (1)$$

where r is the aerosol bin radius, D_g the gas-phase diffusion coefficient, v the mean molecular speed, S the aerosol specific surface area, and γ the uptake coefficient, defined as the ratio of the number of gas molecules reacting with the particle’s surface over the fraction of molecules being absorbed by the given surface (i.e. the accommodation coefficient) (Phadnis and Carmichael, 2000; Guimbaud et al., 2002). A γ value of 0.1 was assumed for dust (Hanisch and Crowley, 2001; Vlasenko et al., 2006) and 0.01 for SS (Tolocka et al., 2004). The production of fine and coarse NO_3^- was traced in separated bins.

Finally, MONARCH includes meteorology-driven emission modules for key species. Emissions of biogenic non-methane volatile organic compounds (NMVOC) and NO are calculated from the Model of Emissions of Gases and Aerosols from Nature (MEGAN) v2.04 (Guenther et al., 2006). Several SS source functions are available in the model (Spada et al., 2013); here we use the Jaeglé et al. (2011) formulation. Similarly, different parameterizations for dust emissions are available ranging from more simplified to more physics-based descriptions (Klose et al., 2021). Following Gonçalves Ageitos et al. (2023), the G01-UST scheme based on Ginoux et al. (2001) described in Klose et al. (2021) is used in this work. Dust emission is limited to areas presenting a frequency of occurrence of dust optical depth above 0.2, identified using maps created from Moderate Resolution Imaging Spectroradiometer (MODIS) Deep Blue retrievals (Hsu et al., 2004; Ginoux et al., 2012). Surface roughness influence on dust emission is parameterized based on Raupach et al. (1993), whose vegetation cover is determined using surface reflectance from Landsat and MODIS monthly data (Raupach et al., 1993; Guerschman et al., 2015).

For the evaluation of the model’s dust cycle, the reader is referred to Klose et al. (2021) and Gonçalves Ageitos et al. (2023), and for SS to Spada et al. (2013).

2.2 Model updates

In this study, we investigate the primary chemical pathways responsible for NO_3^- formation on preexisting particles, with a particular focus on coarse dust particles, by integrating mechanisms of varying complexity within the global model MONARCH (Fig. 1). Table 1 lists the irreversible and reversible heterogeneous reactions considered in our analysis. Here, we detail the enhancements implemented in MONARCH to address partially or fully the array of reactions of interest, with a primary focus on maintaining a balance between complexity, accuracy, and computational efficiency in the resulting solution.

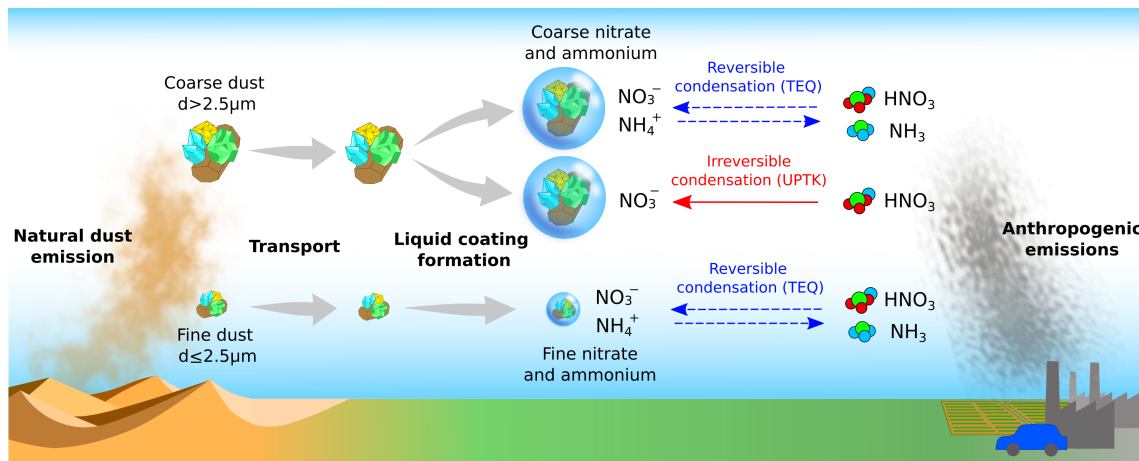


Figure 1. Illustration of the respective mineral dust heterogeneous chemistry mechanisms implemented in this work for fine and coarse particles. Fine dust aqueous coating is assumed to reach thermodynamic equilibrium (TEQ) with anthropogenic gas species to form fine particulate nitrate (NO_3^-). Conversely, coarse particulate nitrate is formed either through also the reversible condensation (TEQ) of $\text{HNO}_{3(g)}$ and $\text{NH}_{3(g)}$ after kinetic limitation, or through the irreversible uptake reaction (UPTK) of $\text{HNO}_{3(g)}$ on coarse dust particles. Particulate ammonium (NH_4^+) is formed in both modes in those mechanisms assuming TEQ.

To trace the formation of fine and coarse NO_3^- , NH_4^+ , and SO_4^{2-} under moist conditions — the primary regime where these formation pathways occur (Usher et al., 2003; Jordan et al., 2003) — an additional hydrophilic bin for the coarse mode of these species is added to the default MONARCH size parametrization, as detailed in Supplementary Table S2. The new bins account for the total mass of NO_3^- , NH_4^+ , and SO_4^{2-} formed on both dust and SS particles indiscriminately. Sensitivity tests, with and without dust and SS in the UPTK and TEQ processes, assess their relative contributions (see Section 2.4).

2.2.1 Irreversible heterogeneous chemistry of nitrate and sulfate

A widely adopted method to simulate NO_3^- and SO_4^{2-} formation on coarse particles involves incorporating irreversible heterogeneous reactions of gas species on dust and SS through a first-order uptake parameterization. Specifically, the uptake of $\text{HNO}_{3(g)}$ on coarse particles is commonly assumed to drive coarse NO_3^- formation (Jacob, 2000; Hodzic et al., 2006; Bauer et al., 2004; Feng and Penner, 2007; Fairlie et al., 2010; Hauglustaine et al., 2014; Paulot et al., 2016; Jones et al., 2021), while the uptake of $\text{SO}_{2(g)}$ is known to lead to the formation of coarse SO_4^{2-} (Phadnis and Carmichael, 2000; Song and Carmichael, 2001; Usher et al., 2002; Prince et al., 2007; Fairlie et al., 2010; Li et al., 2012; Liu and Abbatt, 2021; Yue et al., 2022). Most models assume constant uptake coefficients (γ) for these reactions, for example $\gamma = 0.1$ for $\text{HNO}_{3(g)}$ uptake on dust (Dentener et al., 1996; Hanisch and Crowley, 2001, 2003; Bauer et al., 2004; Hodzic et al., 2006; Wei, 2010)). However, recent studies have shown that using this value tends to overestimate NO_3^- formation (Vlasenko et al., 2006; Mashburn et al., 2006; Fairlie et al., 2010). This suggests that the uptake coefficient for $\text{HNO}_{3(g)}$ should be lower and that it is highly influenced by RH (Goodman, 2000; Krueger et al., 2003; Vlasenko et al., 2006, 2009; Fairlie et al., 2010; Wei, 2010) and dust alkalinity (Good-

Table 1. Heterogeneous reactions implemented in MONARCH. Dust and sea-salt particles are referred as *DU* and *SS*, respectively.

Irreversible reactions			
	Reaction	Process	Notes
(R1)	$\text{DU}_{(\text{aer})} + \text{H}_2\text{SO}_{4(\text{g})} \rightarrow \text{SO}_{4(\text{aq})}^{2-}$	condensation	1
(R2)	$\text{SS}_{(\text{aer})} + \text{H}_2\text{SO}_{4(\text{g})} \rightarrow \text{SO}_{4(\text{aq})}^{2-}$	condensation	1
(R3)	$\text{DU}_{(\text{aer})} + \text{SO}_{2(\text{g})} \rightarrow \text{SO}_{4(\text{aer})}$	UPTK (γ_1)	2
(R4)	$\text{CaCO}_{3(\text{aer})} + 2\text{HNO}_{3(\text{g})} \rightarrow \text{Ca}(\text{NO}_3)_{2(\text{aer})} + \text{H}_2\text{O} + \text{CO}_2$	UPTK (γ_2)	3, 4
(R5)	$\text{SS}_{(\text{aer})} + \text{HNO}_{3(\text{g})} \rightarrow \text{NaNO}_{3(\text{aer})} + \text{HCl}_{(\text{g})}$	UPTK (γ_3)	4, 5
Gas-particle equilibrium reactions			
	Reaction	Process	Notes
(R6)	$\text{HNO}_{3(\text{g})} + \text{NH}_{3(\text{g})} \rightleftharpoons \text{NH}_{4(\text{aq})}^+ + \text{NO}_{3(\text{aq})}^-$	TEQ	6
(R7)	$\text{H}_2\text{SO}_{4(\text{g})} + \text{NH}_{3(\text{g})} \rightleftharpoons \text{NH}_{4(\text{aq})}^+ + \text{SO}_{4(\text{aq})}^{2-}$	TEQ	6, 7
(R8)	$\text{CaCO}_{3(\text{aq})} + 2\text{HNO}_{3(\text{g})} \rightleftharpoons \text{Ca}_{(\text{aq})}^{2+} + 2\text{NO}_{3(\text{aq})}^- + \text{H}_2\text{O} + \text{CO}_2$	TEQ	4, 8
(R9)	$\text{MgCO}_{3(\text{aq})} + 2\text{HNO}_{3(\text{g})} \rightleftharpoons \text{Mg}_{(\text{aq})}^{2+} + 2\text{NO}_{3(\text{aq})}^- + \text{H}_2\text{O} + \text{CO}_2$	TEQ	4, 8
(R10)	$\text{K}_2\text{CO}_{3(\text{aq})} + 2\text{HNO}_{3(\text{g})} \rightleftharpoons 2\text{K}_{(\text{aq})}^+ + 2\text{NO}_{3(\text{aq})}^- + \text{H}_2\text{O} + \text{CO}_2$	TEQ	4, 8
(R11)	$\text{Na}_2\text{CO}_{3(\text{aq})} + 2\text{HNO}_{3(\text{g})} \rightleftharpoons 2\text{Na}_{(\text{aq})}^+ + 2\text{NO}_{3(\text{aq})}^- + \text{H}_2\text{O} + \text{CO}_2$	TEQ	4, 8
(R12)	$\text{NaCl}_{(\text{aq})} + \text{HNO}_{3(\text{g})} \rightleftharpoons \text{Na}_{(\text{aq})}^+ + \text{NO}_{3(\text{aq})}^- + \text{HCl}_{(\text{g})}$	TEQ	4, 9

1. Sulfuric acid ($\text{H}_2\text{SO}_{4(\text{g})}$) is assumed to completely condense on fine and coarse dust and SS, transferring mass to the respective size mode. The condensation reaction does not depend on dust and SS alkalinity, but solely on their specific surface area based on Pringle et al. (2010).

2. Sulfur dioxide ($\text{SO}_{2(\text{g})}$) uptake coefficient on dust (γ_1) is a function of RH as defined in Fairlie et al. (2010) and based on experimental studies performed on calcite particles by Prince et al. (2007). However, to account for $\text{SO}_{2(\text{g})}$ oxidation by deliquesced O_3 and NO_2 , the UPTK reaction of $\text{SO}_{2(\text{g})}$ is performed even in the absence of alkalinity. That is why this reaction is considered to happen over $\text{DU}_{(\text{aer})}$ and not only over $\text{CaCO}_{3(\text{aq})}$, although it assumes same dust alkalinity as in nitric acid uptake (R4).

3. Nitric acid ($\text{HNO}_{3(\text{g})}$) uptake coefficient on $\text{CaCO}_{3(\text{aq})}$ (γ_2) is a function of RH as defined by Fairlie et al. (2010) based on experimental studies performed on calcite particles by Liu et al. (2008). $\text{CaCO}_{3(\text{aq})}$ concentration is used instead of $\text{DU}_{(\text{aer})}$ concentration because the uptake coefficient is scaled for alkalinity as shown in Eq. 2.

4. $\text{CaCO}_{3(\text{aq})}$, $\text{MgCO}_{3(\text{aq})}$, $\text{K}_2\text{CO}_{3(\text{aq})}$, $\text{Na}_2\text{CO}_{3(\text{aq})}$ and $\text{NaCl}_{(\text{aq})}$ refers to the NVC content derived from the $\text{DU}_{(\text{aer})}$ and $\text{SS}_{(\text{aer})}$ concentrations using fractions from Gonçalves Ageitos et al. (2023) for dust and from Seinfeld and Pandis (2006) for SS (Section 2.2.3).

5. $\text{HNO}_{3(\text{g})}$ uptake coefficient on SS (γ_3) is based on the experimental study Liu et al. (2007), that reports $\text{HNO}_{3(\text{g})}$ uptake kinetics for different RH and sea-salt particles sizes. The average value $\gamma_3 = 0.05$ for particles from 2.5 to 10 μm (coarse mode) is assumed, considering ambient RH=80%.

6. Neutralization of $\text{HNO}_{3(\text{g})}$ and $\text{H}_2\text{SO}_{4(\text{g})}$ by ammonia is calculated through TEQ with ISORROPIA-II. It is assumed to happen in the fine mode for all mechanisms, and additionally in the coarse mode over coarse particles in those mechanisms considering coarse NO_3^- and NH_4^+ formation (Myhre et al., 2006; Usher et al., 2003; Uno et al., 2020). All reactants and products are assumed to remain in the aqueous phase (metastable assumption).

7. The result of the neutralization of $\text{H}_2\text{SO}_{4(\text{g})}$ can be $(\text{NH}_4)_2\text{SO}_{4(\text{aer})}$, $\text{NH}_4\text{HSO}_{4(\text{aer})}$ or $(\text{NH}_4)_3\text{H}(\text{SO}_4)_2(\text{aer})$ if solid results were assumed (Liu et al., 2022), but under metastable assumption only aqueous ions of SO_4^{2-} are considered.

8. Calcium, magnesium, potassium and sodium (NVC) deliquesced from carbonates present in the bulk of dust particles neutralize $\text{HNO}_{3(\text{g})}$ in the liquid coating of dust aerosols (Usher et al., 2003; Krueger et al., 2004; Fountoukis and Nenes, 2007; Hauglustaine et al., 2014). Dust NVC content (i.e. alkalinity) is dependent on particle size and globally averaged from Journet et al. (2014) mineral data.

9. Sodium chloride from sea-salt particles dissolves and reacts with $\text{HNO}_{3(\text{g})}$ in the liquid coating of sea-salt particles (Myhre et al., 2006). Sea-salt also presents other NVC that are included in reactions R8-11, which are assumed globally homogeneous from Seinfeld and Pandis (2006) and Karydis et al. (2016).

man, 2000; Hanisch and Crowley, 2001; Krueger et al., 2004; Liu et al., 2007; Prince et al., 2007; Wei, 2010; Crowley et al., 2010). Recent studies increasingly implement γ as a function of RH and employ different parameterizations of this function to account for dust alkalinity.

For our study, we extended the chemical mechanism of MONARCH to incorporate pathways for the formation of coarse NO_3^- and sulfate aerosols (Fig. 2). This extension involved refining irreversible heterogeneous parameterizations within the model, specifically the uptake of $\text{HNO}_{3(\text{g})}$ on dust and SS particles (R4-5 in Table 1), as well as the uptake of $\text{SO}_{2(\text{g})}$ on dust particles (R3 in Table 1). No additional heterogeneous chemistry, such as the transformation of $\text{NO}_{2(\text{g})}$ to $\text{HNO}_{3(\text{g})}$ on the surface of dust particles, was considered due to its relatively low significance (Jacob, 2000; Jordan et al., 2003; Liao et al., 2003; Li et al., 2024).

Our implementation incorporates dependencies of γ on RH and the alkalinity of dust (Vlasenko et al., 2006, 2009; Liu et al., 2007, 2008; Crowley et al., 2010; Fairlie et al., 2010; Wei, 2010). The γ dependency on RH is modeled akin to a Brunauer–Emmett–Teller (BET) isotherm, which characterizes water adsorption on dust particles (Vlasenko et al., 2006). We employed a modified BET function to formulate γ , extending it to account for dust alkalinity. This formulation is represented by the following equation:

$$\gamma = Sc \cdot \frac{|c_1 \cdot RH|}{(1 - RH)|1 + c_2 \cdot RH|} \quad (2)$$

where RH is the relative humidity (ranging from 0 to 1), c_1 and c_2 denote the water adsorption scaling factors (Vlasenko et al., 2006), and Sc is a factor dependent on dust alkalinity. For the uptake of $\text{HNO}_{3(\text{g})}$ on dust (R4 in Table 1), typical values assumed for c_1 and c_2 are 8.0 and 7.0, respectively (Li et al., 2012; Paulot et al., 2016; Wang et al., 2017). For Sc , however, literature reports varying values based on dust alkalinity assumptions, ranging from $Sc = \frac{1}{30}$ for the industrially-standardized Arizona Test Dust (Möhler et al., 2006; Herich et al., 2009; Suman et al., 2024) to $Sc = 0.018$ for samples from the China Loess with 39% CaCO_3 content (Krueger et al., 2004; Wei, 2010).

We adopt the uptake RH functions for both $\text{HNO}_{3(\text{g})}$ and $\text{SO}_{2(\text{g})}$ on dust from Fairlie et al. (2010). To fit experimental data from Song et al. (2007) and the RH function reported by Fairlie et al. (2010), for $\gamma(\text{HNO}_3)$ (Reaction R4 in Table 1) we determine $c_1 = 3.84 \cdot 10^{-4}$ and $c_2 = 0.56$. Additionally, Fairlie et al. (2010) assumes a NVC content of 3.0% of Ca and 0.6% of Mg, which differs from the NVC values used in the present study. Therefore, we use the alkalinity scaling factor Sc to normalize the Fairlie et al. (2010) function accordingly. Specifically, values for Sc are defined as the ratio of Ca and Mg percentage used in our study relative to those assumed by Fairlie et al. (2010), resulting in $Sc = 1.80$ and $Sc = 1.52$ respectively for the two average alkalinity values derived from Journet et al. (2014) and Claquin et al. (1999) used in our experiments (see Section 2.2.3 below). Note that Sc is zero if no alkalinity is considered and that a constant value of $\gamma(\text{HNO}_3) = Sc \cdot 1.05 \cdot 10^{-3}$ is used for RH higher than 80%.

Similarly, we determined the $\gamma(\text{SO}_2)$ on dust (R3 in Table 1) fitting Eq. 2 to experimental data from Prince et al. (2007) and the RH function from Fairlie et al. (2010), yielding values of $c_1 = 2.7 \cdot 10^{-6}$ and $c_2 = -1.06$. The same Sc values as used for $\gamma(\text{HNO}_3)$ are applied for $\gamma(\text{SO}_2)$: $Sc = 1.80$ and $Sc = 1.52$. However, if alkalinity is not considered, Sc is set to 1.0 (and not

zero, as the case for $\gamma(\text{HNO}_3)$) to account for the oxidation of $\text{SO}_{2(\text{g})}$ by deliquesced O_3 and NO_2 (Usher et al., 2002; Prince et al., 2007; Yu et al., 2017; Li et al., 2024). For RH above 90%, $\gamma(\text{SO}_2)$ remains constant at $\gamma(\text{SO}_2) = Sc \cdot 5.0 \cdot 10^{-4}$.

245 For the $\text{HNO}_{3(\text{g})}$ uptake on SS (R5 in Table 1), we adopted the $\gamma(\text{HNO}_3)$ values of Liu et al. (2007), which provide experimental estimates of this factor for different particle sizes and RH. However, a clear uptake function on these parameters has not been found in the literature. Therefore, for the sake of simplicity, we defer an implementation of an uptake coefficient dependent on these metrics to future research. In this study, we did not account for the RH dependency of $\gamma(\text{HNO}_3)$, and instead we used average values at 80% RH resulting in $\gamma(\text{HNO}_3) = 0.15$ for SS particles in the range of 0.1 to 2.5 μm and
250 $\gamma(\text{HNO}_3) = 0.05$ for particles larger than 2.5 μm . While larger values have been reported by Guimbaud et al. (2002), we opted for these values as they align with more widely accepted ranges found in the literature (Saul et al., 2006; Pratte and Rossi, 2006; Liu et al., 2007; Fagerli et al., 2015). Using higher values could potentially overestimate the uptake on SS particles.

The condensation of $\text{H}_2\text{SO}_{4(\text{g})}$ (R1-2 in Table 1) on dust and SS is another relevant source of SO_4^{2-} introduced in our model. Due to the extremely low volatility of $\text{H}_2\text{SO}_{4(\text{g})}$ at atmospheric temperatures, its condensation onto existing particles
255 is assumed as irreversible and complete. This process signifies a direct mass transfer from the gas to the aerosol phase (Zaveri et al., 2008; Hauglustaine et al., 2014). The amount of $\text{H}_2\text{SO}_{4(\text{g})}$ that condenses in the fine and coarse modes is determined using kinetic diffusive coefficients calculated as described in Section 2.3.

2.2.2 Reversible heterogeneous chemistry of nitrate and ammonium

The gas-aerosol partitioning of semivolatile inorganic aerosols in previous studies with MONARCH was based on the EQSAM
260 v03b TEQ model. EQSAM provides a computationally efficient approach that bypasses the expensive iterative activity coefficient calculation employed in other thermodynamic models. EQSAM was originally designed to handle the partitioning of ammonium-sulfate-nitrate-water system excluding solid components and was extended to include solids, HCl and Cl-/Na+ in the version v03b, the one used in MONARCH. One of the limitations of such version is the lack of information on NVC and/or mineral species in traced species.

265 For this study, we implemented the ISORROPIA-II v1 (Fountoukis and Nenes, 2007) TEQ model as an additional option in MONARCH to investigate the sensitivity of the partitioning of semivolatile inorganic compounds to NVCs. While a more recent version of ISORROPIA-II (v2.3) exists, which improves aerosol pH estimations at near pH-neutral conditions (Song et al., 2018), global scale simulations have shown only minor differences when compared to ISORROPIA-II v1 (Milousis et al., 2024).

270 ISORROPIA-II v1 determines TEQ concentrations of gas, liquid and solid phases. It can assume either stable conditions, where compounds precipitate into solids, or metastable conditions, where compounds remain as supersaturated liquid solutions. To enhance computational efficiency, ISORROPIA-II employs a segmented approach for calculating TEQ concentrations. This approach defines five different regimes based on the ratios of precursor species (i.e. sulfate, sodium and crustal species), RH and temperature. Each regime addresses a specific subset of relevant species and equilibrium equations. Efficiency is
275 further improved by retrieving species' activity coefficients from look-up tables (Fountoukis and Nenes, 2007; Milousis et al., 2024). The medium acidity is determined by the concentrations of acidic/basic gaseous species ($\text{HNO}_{3(\text{g})}$, $\text{NH}_{3(\text{g})}$, $\text{H}_2\text{SO}_{4(\text{g})}$),

particle (NH_4^+ , SO_4^{2-} , NO_3^-), and crustal ions (K^+ , Ca^{2+} , Mg^{2+} , Na^{2+} , Cl^-), which are inputs to ISORROPIA-II. After TEQ is calculated with these species, the resulting pH is provided by the thermodynamic model.

In this work, we use the metastable solution of ISORROPIA-II, assuming all the resulting particulate compounds from the TEQ computation to remain in the liquid phase. Previous studies comparing stable and metastable methodologies with ISORROPIA-II have reported only marginal differences in global nitrate budgets between both modes. At global scales, these differences noted slightly higher pH values (0.5) and nitrate formation (2%) when using the metastable assumption (Karydis et al., 2016, 2021; Milousis et al., 2024), although these differences are reported to be more important (<2 pH units and <60% nitrate concentrations) close to regions with low RH and high concentration of crustal species, or their downwind areas. However, given the global scale scope of the present study, we used the metastable assumption since it allows for full traceability of total aerosol nitrate, ammonium and sulfate formation (reactions R6-12 in Table 1).

We also adopt the temperature and pressure applicability range for ISORROPIA-II proposed by Sulprizio (2022), which highlights potential instabilities in reactions occurring below 250°K and 200hPa. Consequently, ISORROPIA-II computations are limited to cells with temperature and pressure values above these thresholds.

2.2.3 Dust and sea-salt alkalinity

Alkalinity refers to the ability of a substance to neutralize acids and maintain a stable pH level. Both dust and SS particles contain NVCs that contribute to the overall alkalinity of the aerosol, thereby neutralizing gas acidic species such as $\text{HNO}_{3(g)}$ and sulfates.

To investigate the importance of representing dust alkalinity, we derive a global average size-dependent NVCs content from 5-year long MONARCH simulations that explicitly track dust mineral species (Gonçalves Ageitos et al., 2023). To assess the uncertainty arising from our limited knowledge of the soil mineralogy of dust sources, we relied on two different MONARCH experiments detailed in Gonçalves Ageitos et al. (2023), which utilized the Claquin et al. (1999) and Journet et al. (2014) soil mineralogical datasets, respectively. The simulation based on Claquin et al. (1999) accounts for eight distinct minerals, whereas the simulations based on Journet et al. (2014) for twelve minerals (Table S3). The Claquin et al. (1999) dataset includes quartz, feldspar, illite, smectite, kaolinite, calcite, gypsum, and hematite. The Journet et al. (2014) dataset includes these minerals as well as chlorite, vermiculite, mica, and goethite (Table S5). In this study, we adopt an upper bound for the minerals' solubility and reactivity with gas species based on Hanisch and Crowley (2001). Moreover, we assume size-dependent but globally homogeneous values for dust mineralogy, and consequently dust alkalinity and NVC, to focus on understanding heterogeneous reaction parameterizations. We defer the analysis of the potential importance of geographical dust mineralogical variations upon dust heterogeneous chemistry to a forthcoming study.

Based on the global average mineral mass fraction for each dust size bin derived from the mineralogy simulations and the elemental composition associated to each mineral (see the Supplementary Tables S4 and S6) we estimate the average NVC content per dust size bin at each time step following Eq. 3:

	t0: Fine SO4	t1: Fine DU	t2: Coarse DU	t3: Gas HNO3, NH3 and H2SO4	t4: Fine DU & SS	t5: Coarse DU	t6: Coarse SS	t7: Fine/Coarse SO4
TEQ	Cond. (x1.0 H2SO4)	UPTK (SO2)	UPTK (SO2)	-	TEQ (NH3 & HNO3)	-	-	-
HYB				-		UPTK (HNO3)	UPTK (HNO3)	-
DBCLL	Cond. (x0.5 H2SO4)			DIFFLIM		TEQ (NH3 & HNO3)		Cond. H2SO4

Table 2. Sequence of reactions and calculations performed for each scheme (rows). In columns, the sequence order (tN) and the gas/aerosol mode of each process is indicated. Dust and sea-salt particles are referred as *DU* and *SS*, respectively, and condensation is abbreviated as *Cond.*.

$$NVC_{i,j} = \left(\sum_k fNVC_{k,j} \cdot M_{i,k} \right) \cdot DU_i \quad (3)$$

where the NVC concentration for each element j (i.e., Ca, Mg, K, Na) and size bin i ($NVC_{i,j}$) at a given location is derived by considering the molar fraction of each element in each mineral k ($fNVC_{k,j}$), the global average mass fraction of each mineral and size bin ($M_{i,k}$) and the bin's dust concentration at the specified location (DU_i). $NVC_{i,j}$ serves as input to the TEQ calculation. We consider only minerals soluble in water or acids that may at least partly dissolve in the liquid coating of the particles (Usher et al., 2003), and only NVC reacting with the gas species in ISORROPIA-II are used (calcite, magnesium, potassium, sulfate, chlorite or sodium) for the calculation.

The dust NVC global average content results in: 5.17% Ca^{2+} , 0.79% Na^+ , 2.37% K^+ , 1.32% Mg^{2+} for the Journet et al. (2014) dataset, and 3.68% Ca^{2+} , 0.87% Na^+ , 3.15% K^+ , 1.75% Mg^{2+} for Claquin et al. (1999). The size-resolved NVC percentages for each dust bin are reported in the Supplementary Tables S3, S4, S5 and S6.

In most of the sensitivity runs, the Journet et al. (2014) global average is employed, if not stated otherwise (see Table 3). Values for Claquin et al. (1999) are used solely in one sensitivity test, as explained in Section 2.4. These values are within the range reported by Karydis et al. (2016) ($5.36 \pm 3.69\%$ Ca^{2+} , $2.46 \pm 1.90\%$ Na^+ , $2.08 \pm 1.34\%$ K^+ , $1.96 \pm 2.20\%$ Mg^{2+}). The Journet et al. (2014) dataset results in a higher proportion of Ca^{2+} compared with Claquin et al. (1999), while similar fractions for Na^+ and K^+ are reported.

Additionally, as discussed in Section 2.2.1, the derived dust NVC implies the application of the scaling factors $Sc = 1.80$ and $Sc = 1.52$ in Eq. 2 for the irreversible uptake experiments, assuming the average alkalinity derived from the Journet et al. (2014) and Claquin et al. (1999) simulations, respectively.

Regarding SS, we use a global average composition from Seinfeld and Pandis (2006) with 55% Cl^- , 30.6% Na^+ , 7.7% SO_4^{2-} , 3.7% Mg^{2+} , 1.2% Ca^{2+} , and 1.1% K^+ .

2.3 Nitrate mechanisms under study

In atmospheric conditions, $\text{HNO}_{3(\text{g})}$ and $\text{NH}_{3(\text{g})}$ exhibit higher volatility compared to $\text{H}_2\text{SO}_{4(\text{g})}$. As a result, their condensation onto liquid coatings around particles is a reversible process and should not be assumed as irreversible uptake reactions (Usher et al., 2003). Different mechanisms have been proposed to model the partitioning of nitrate and ammonium across the

entire aerosol size range, aiming to mitigate the computationally expensive cost of solving the dynamic mass transfer equations (Capaldo et al., 2000; Feng and Penner, 2007; Hauglustaine et al., 2014).

335 The assumption of TEQ between gas and aerosol phases provides a practical approximation to account for the potential evaporation of already dissolved molecules in the liquid coating of fine particles. Equilibrium timescales for fine ammonium nitrate (diameter less than $1\mu m$) are typically on the order of minutes under typical atmospheric conditions (Wexler and Seinfeld, 1990; Dassios and Pandis, 1999). However, in TEQ models, it is assumed that TEQ is reached within each model time-step (on the order of few minutes). This assumption is reasonable for fine particles, but less so for coarse particles, where
340 achieving equilibrium can take minutes to hours (Feng and Penner, 2007).

To overcome this limitation, different approaches have been proposed in the literature to incorporate the condensation-evaporation of $\text{HNO}_{3(g)}$ and $\text{NH}_{3(g)}$ on coarse particles while minimizing computational costs. In this study, we explore two such methods. (1) The hybrid method (HYB) (Hodzic et al., 2006; Feng and Penner, 2007) solves the partitioning over fine particles using a TEQ model and employs a first-order irreversible uptake (UPTK) reaction for condensation over coarse
345 particles (section 2.2.1). (2) A more refined approach treats the formation of coarse NO_3^- as a reversible process through the combination of a double call of the TEQ calculation, one for the fine and one for the coarse mode (DBCLL), together with a kinetic limitation of the gas species involved in the partitioning (Pringle et al., 2010). These two mechanisms are evaluated in this study to assess their impact on the formation of coarse NO_3^- and NH_4^+ , as illustrated in Fig. 2. Additionally, for the purpose of comparison, a scheme neglecting coarse NO_3^- formation (fTEQ) is also employed. We briefly describe each approach below.

350 The fTEQ approach solves the partitioning of semivolatile inorganic species with ISORROPIA-II exclusively within the fine mode to assess the effect of neglecting coarse NO_3^- and NH_4^+ formation on atmospheric composition. This mechanism solves the nitric-ammonia-sulfate neutralization (R6-7 in Table 1), accounting for alkalinity effects (R8-12 in Table 1) if fine dust and SS are considered in the mixture. Since $\text{H}_2\text{SO}_{4(g)}$ and SO_4^{2-} influence the ambient pH, they are always involved in any TEQ calculation. As described in Section 2.1, aqueous sulfate formation is solved through the oxidation of $\text{SO}_{2(g)}$ and
355 DMS. Here, however, only 50% of the remaining $\text{H}_2\text{SO}_{4(g)}$ is assumed to directly nucleate as fine SO_4^{2-} through the aqueous phase chemistry, while the rest condenses into fine SO_4^{2-} through reactions R1 and R2 in Table 1 (Fig. 2 (a)). While in our study fTEQ serves as a sensitivity test to assess the impact of neglecting coarse NO_3^- and NH_4^+ formation, the fTEQ approach may be appropriate in environments where coarse particles are sparse or in applications focusing primarily on fine particle formation (Bian et al., 2017).

360 Conversely, the HYB mechanism (Fig. 2 (a)) is a commonly employed strategy in atmospheric chemistry models to represent nitrate formation on coarse particles. It solves NO_3^- formation through a sequential implementation of 1) TEQ reaction between $\text{HNO}_{3(g)}$ and $\text{NH}_{3(g)}$ considering internal mixing with fine dust and SS modes (R6-12 in Table 1), and 2) an irreversible first-order UPTK reaction of the remaining $\text{HNO}_{3(g)}$ on the coarse modes of dust and SS, excluding its evaporation back to the gas phase (R4 and R5 in Table 1). The UPTK reaction of $\text{HNO}_{3(g)}$ follows the implementation detailed in Section 2.2.1. Sulfate is
365 treated in an analogous manner to the fTEQ mechanism.

Finally, the DBCLL mechanism (Fig. 2 (b)) (Pringle et al., 2010) treats coarse NO_3^- and NH_4^+ formation as a reversible condensation-evaporation process. Firstly, the DBCLL methodology involves the calculation of kinetic diffusion limitation

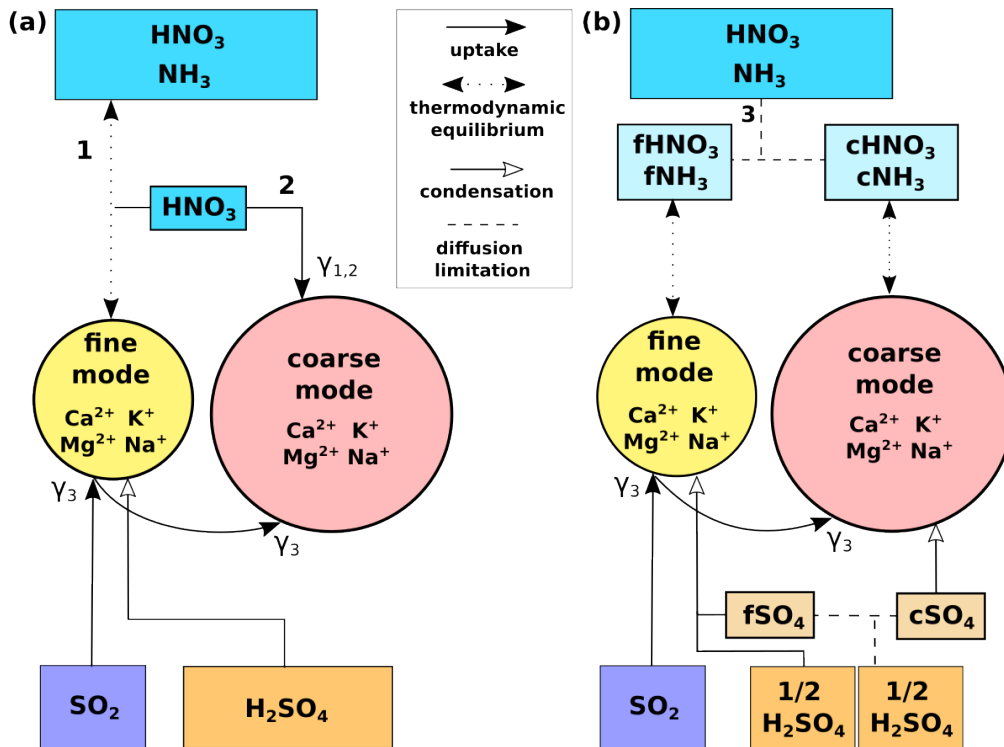


Figure 2. Scheme of the heterogeneous chemistry reactions on dust and SS developed in this work. (a) Schemes for the fTEQ and HYB mechanisms. (b) Scheme for DBCLL mechanism, that includes the kinetic diffusion limitation of gas species. The legend for line patterns is included on the center box. Numbers represent (1) fTEQ, (2) HYB and (3) DBCLL. In the bottom, the pathways for processing $\text{SO}_{2(g)}$ and $\text{H}_2\text{SO}_{4(g)}$ gases for each scheme are represented; $\text{SO}_{2(g)}$ is uptaken by fine particles and the remaining gas is uptaken by coarse particles. For $\text{H}_2\text{SO}_{4(g)}$, (a) the fTEQ and HYB mechanisms assume that it nucleates directly as fine SO_4^{2-} , while (b) DBCLL mechanism assumes that 50% of its concentration nucleates, and the rest is divided by diffusion limitation between fine and coarse modes, completely condensing on each of them. The uptake coefficients used for each uptake process (solid arrows) are indicated as $\gamma_{1,2}$ for $\text{HNO}_{3(g)}$ uptake on dust and SS, and γ_3 for $\text{SO}_{2(g)}$ uptake on dust.

coefficients (DIFFLIM) for both the fine and coarse size modes of each of the condensing gas species, which restricts the amount of gas available to condense on each mode (Table 2). DIFFLIM has been implemented based on the formulation by Vignati et al. (2004) for $\text{H}_2\text{SO}_{4(g)}$, and its extension to other gases by Pringle et al. (2010). Following DIFFLIM, sequential TEQ calculations are conducted over fine and coarse modes (double call of the TEQ model, reactions R8-12 in Table 1), using the DIFFLIM coefficients to limit the availability of gas condensing on each mode.

Regarding the sulfate treatment in DBCLL, a similar approach to fTEQ is employed, but the DIFFLIM coefficients calculated for $\text{H}_2\text{SO}_{4(g)}$ are used to condense the available $\text{H}_2\text{SO}_{4(g)}$ into either fine or coarse modes of SO_4^{2-} (reactions R1 and R2 in Table 1 and Fig. 2 (b)).

Experiment	Fine TEQ	Coarse TEQ	Coarse DU UPTK	Coarse SS UPTK	DIFFLIM	DU Alk.	SS Alk.
noHC							
fTEQ_noAlk							
fTEQ_du-ssAlk						J	
HYB_duUPTK						J	
HYB_du-ssUPTK						J	
HYB_DL						J	
HYB_g0p1						J	
DBCLL_noAlk							
DBCLL_duAlk						J	
DBCLL_du-ssAlk						J	
DBCLL_ClaqAlk						C	

Table 3. The sensitivity experiments conducted in this study and the processes performed in each one of them (green boxes). *Fine TEQ* and *Coarse TEQ* refer to the respective calculation of the TEQ on the fine and coarse modes of dust and SS, referred as *DU* and *SS*, respectively. *Coarse DU UPTK* and *Coarse SS UPTK* indicate the irreversible UPTK of $\text{HNO}_{3(g)}$ on coarse dust and SS, respectively. Alkalinity is denoted as *Alk.* and for the dust alkalinity column, alkalinity derived from the Journet et al. (2014) and Claquin et al. (1999) simulations is indicated by *J* and *C*, respectively.

2.4 Sensitivity runs

We performed global simulations based on the mechanisms described in Section 2.3. Overall, eleven different runs (Table 3) are analyzed to test different degrees of complexity and sensitivity to parameterizations when simulating heterogeneous chemistry of dust, such as the hypothesis on reversibility on nitrate formation, and the role that dust (and SS) alkalinity play in the partitioning of gas and aerosol species. Unless stated to the contrary, all the experiments employ dust alkalinity derived from the average of the Journet et al. (2014) simulation, as explained in Section 2.2.3.

Three initial run sets are conducted neglecting coarse NO_3^- formation, serving as a comparison reference for following experiments including coarse NO_3^- . The noHC run assumes that there is no formation of NO_3^- , NH_4^+ or SO_4^{2-} aerosol through heterogeneous chemistry on any aerosol particle, serving as a baseline to estimate the burden of gas condensing in particles in other configurations and the influence of particle formation on nitrogen deposition rates. Additionally, two sensitivity experiments with fTEQ are included to discuss the impact of ignoring the partitioning of semivolatile inorganic species on coarse particles: fTEQ_noAlk neglects the presence of dust or SS in the aerosol mixture (R6-7 in Table 1), while fTEQ_du-ssAlk considers TEQ between gas and NVC in the fine modes of dust and SS particles (R6-12 in Table 1).

Next, we addressed the condensation of nitrate across the entire particle size range with runs employing the HYB and DBCLL mechanisms. We conducted two sensitivity runs with the HYB mechanism to explore the impact of implementing the $\text{HNO}_{3(g)}$ UPTK reaction on coarse dust only (HYB_duUPTK, reaction R4 from Table 1)) and on both coarse dust and SS (HYB_du-ssUPTK, reactions R4 and R5), comparing their results to assess the relative contributions of dust and SS in heterogeneous chemistry under the assumption of $\text{HNO}_{3(g)}$ irreversible UPTK on the coarse mode.

In the HYB approach, all $\text{HNO}_{3(g)}$ and $\text{NH}_{3(g)}$ concentrations are initially available to condense in the fine mode through
395 TEQ reactions. Only the remaining $\text{HNO}_{3(g)}$ after the TEQ calculation is considered for UPTK reactions in the coarse mode.
This assumption may potentially lead to a misrepresentation of fine and coarse NO_3^- formation, such as the underproduction
of fine and overproduction of coarse NO_3^- (Feng and Penner, 2007; Hauglustaine et al., 2014; Bian et al., 2017; Jones et al.,
2021). To address this potential limitation, an additional simulation (HYB_DL) was conducted using the DIFFLIM calculation
to distribute the $\text{HNO}_{3(g)}$ and $\text{NH}_{3(g)}$ gas that is kinetically available for condensation in the fine mode through TEQ and the
400 $\text{HNO}_{3(g)}$ that can form coarse NO_3^- through UPTK reactions.

Furthermore, to assess the influence of the UPTK coefficient of $\text{HNO}_{3(g)}$ on the results, instead of using a RH function (see
Section 2.2.1) we also conducted a simulation setting $\gamma(\text{HNO}_3) = 0.1$ for dust (HYB_g0p1), following experimental findings
(Fenter et al., 1995; Hanisch and Crowley, 2001, 2003) and various modelling studies (Dentener et al., 1996; Liao et al., 2003;
Bauer et al., 2004; Hodzic et al., 2006; Bauer et al., 2007; Feng and Penner, 2007).

405 Finally, we conducted four sensitivity simulations using the DBCLL mechanism, which accounts for reversible heteroge-
neous nitrate chemistry both on fine and coarse modes. These simulations evaluate the influence of alkalinity, including the
DBCLL_noAlk run that excludes both dust and SS NVC content (reactions R6-7, Table 1), the DBCLL_duAlk run accounting
for dust alkalinity only (reactions R8-11), and the DBCLL_du-ssAlk run accounting for both dust and SS alkalinity (reactions
R8-12). Additionally, since all three cases use dust alkalinity from the average of the Journet et al. (2014) simulation, an ad-
410 ditional simulation was performed using the average dust alkalinity from Claquin et al. (1999) simulations (DBCLL_Claq),
instead of Journet et al. (2014), to assess the effect of the specific dust alkalinity content used.

2.5 Experimental setup

The model simulations were conducted on a global domain at a spatial resolution of 1.4 degrees longitude by 1.0 degree
latitude, utilizing 48 hybrid pressure-sigma vertical layers up to 5hPa. The dynamics timestep was set to 180s and results were
415 stored every 6 hours. The analysis period is year 2018 after a spin-up period of half a year to initialize the concentration fields.
Meteorological variables were initialized from the ERA5 reanalysis (Hersbach et al., 2023) every 24h to keep the modeled
circulation close to observations. A meteorological spin-up of 12h was used in each daily cycle before solving the chemistry.
The initial state of the chemistry fields are those prognostically calculated by MONARCH the day before.

In addition to the meteorology-driven online emissions described in Sect. 2.1, the High-Selective Resolution Modelling Emis-
420 sion System version 3 (HERMESv3; Guevara et al., 2019) was employed to process both anthropogenic and biomass burning
primary emissions. The global inventory CAMS-GLOB-ANT_v4.2 (Soulie et al., 2024) for 2016 was used for anthropogenic
sources with updated temporal profiles that provide gridded monthly, day-of-the-year, day-of-the-week and hourly weight fac-
tors for the temporal disaggregation of emitted fluxes (Guevara et al., 2021). The biomass burning emissions are provided
by the GFASv1.2 dataset (Kaiser et al., 2012), which accounts for forest, grassland, and agricultural waste fires derived from
425 satellite products. Oceanic natural emissions of DMS are provided by CAMS-GLOB-OCE_v3.1 (Lana et al., 2011; Denier van
der Gon et al., 2023).

Table S1 summarizes the total emissions (anthropogenic, biogenic and biomass burning) used in this work. The emitted mass of the main anthropogenic aerosol precursors are 104.4 Tg for $\text{SO}_{2(g)}$, 93.8 Tg for $\text{NO}_{x(g)}$, and 61.8 Tg for $\text{NH}_{3(g)}$.

2.6 Model evaluation

430 The model results are evaluated against surface observational datasets from several networks sourced from the Globally Harmonised Observational Surface Treatment (GHOST) project, an initiative of the BSC's Earth Department dedicated to the harmonisation of publicly available global surface observations (Bowdalo et al., 2024). Figure S13 in the Supplement shows the stations used for each of the species analyzed. For gas and aerosol nitrate, ammonia and sulfate species, GHOST include datasets from the Clean Air Status and Trends Network (US-EPA-CASTNET), the US EPA Air Quality System (US-EPA-
435 AQS) and the Canada National Air Pollution Surveillance Program (NAPS) for North and Central America, the East Asia Acid Deposition Monitoring Network (EANET) for Asia, and the EBAS and the European Environmental Agency Air Quality (EEA AQ eReporting) for Europe.

These observations are filtered for rural and background sites only, in order to exclude stations near emission sources not representative of the background conditions depicted by the model resolution, which describes the long-range transport of
440 NO_3^- , NH_4^+ and SO_4^{2-} formation. However, information on the station type was not available for the EANET and US-EPA-CASTNET data. For these networks, all stations were used, a factor that has to be accounted for when evaluating the results.

Regarding $\text{PM}_{2.5}$ and PM_{10} , GHOST includes data from the CHILE-SINCA network for Chile, the Beijing Municipal Ecological and Environmental Monitoring Center (BJMEMC), the China National Environmental Monitoring Centre (CNEMC), WMO World Data Center for Aerosols (EBAS-WMO-WDCA) for Europe, the Japan National Institute for Environmental
445 Studies (NIES), the Ministerio de Transición Ecológica (MITECO) for Spain, the UK AIR network for United Kingdom, and the global AirNow DOS network (US-EPA-AirNow-DOS). For these networks, the station type was not included as a criteria to include it in the evaluation. The statistical metrics used in the evaluation are outlined in the Supplementary Section S3, and the quality flags for the selected stations in Table S8.

Additionally, we compare our model results with the budgets reported in the literature (Hauglustaine et al., 2014; Bian
450 et al., 2017; Rémy et al., 2022). Namely, the AeroCom phase III nitrate experiment (Bian et al., 2017) intercompares global NO_3^- budgets from nine global models for the year 2008. Particular discussion is devoted to results shown there for the GMI Bian et al. (2009) and EMAC (Karydis et al., 2016) models, that introduce relevant approaches of interest to our work. Complementarily, results from Hauglustaine et al. (2014) and Rémy et al. (2022) are also used. The comparison provides a qualitative view of current estimates of particulate NO_3^- formation in the atmosphere and the role of representing key processes
455 in models. Since results from the literature are provided for different years, some of the differences may be attributed to changes in emissions and environmental conditions.

3 Results and discussion

3.1 Spatial distributions

The spatial distribution of nitrate species exhibits significant variability depending on the assumed formation mechanisms. Here, we analyze the results of the main sensitivity runs, arranged in order of increasing complexity as detailed in Section 2.4. These simulations range from a basic scenario that considers only fine NO_3^- formation while excluding dust and sea-salt alkalinity (i.e. fTEQ), to a comprehensive approach that incorporates reversible NO_3^- formation in both fine and coarse modes (i.e. HYB and DBCLL).

Spatial distributions of surface concentration, column load and zonal average concentration for all sensitivity simulations (see Table 3) are averaged for 2018 and presented for $\text{HNO}_{3(\text{g})}$ and particulate NO_3^- (Fig. 3, Fig. S1, and Fig. S2 in the Supplement), $\text{NH}_{3(\text{g})}$ and NH_4^+ (Fig. S4, Fig. S5, and Fig. S6 in the Supplement), as well as $\text{SO}_{2(\text{g})}$ and SO_4^{2-} (Fig. S7, Fig. S8, and Fig. S9 in the Supplement). To facilitate readability, the analysis of the HYB_g0p1, DBCLL_duAlk and DBCLL_ClaqAlk runs is provided in Supplementary Section S5.

3.1.1 Effects of omitting coarse nitrate formation

In the fTEQ_noAlk sensitivity run, fine particulate NO_3^- primarily forms in regions with significant anthropogenic pollution—areas where emissions of $\text{NO}_{\text{x}(\text{g})}$, $\text{NH}_{3(\text{g})}$, and $\text{SO}_{2(\text{g})}$ are dominant—such as Northern China, India, Europe and eastern North America. In these regions, NO_3^- reaches average concentrations of 2 to 10 $\mu\text{g m}^{-3}$ (Fig. 3e). At the surface, NH_4^+ is mostly associated with fine NO_3^- , presenting concentrations between 1 and 3 $\mu\text{g m}^{-3}$ (Fig. 3e and Fig. S4e in the Supplement). At higher altitudes, fine NH_4^+ also forms in the presence of dust and SS, resulting in column burdens ranging from 1 to 3 mg m^{-2} (Fig. S5 and Fig. S8 in the Supplement). This occurs because NH_4^+ neutralizes particulate SO_4^{2-} , which forms when $\text{H}_2\text{SO}_{4(\text{g})}$ and $\text{SO}_{2(\text{g})}$ condense on fine dust and SS particles.

The omission of dust and SS alkalinity leads to globally low pH values. Over oceanic and dusty regions, aerosol pH typically ranges from 1 to 2, while in industrialized regions, such as Europe and Asia, pH values can reach up to 5 (Fig. 4a). Aerosol pH increases significantly when dust and SS NVC are included in the fTEQ_du-ssAlk run (Fig. 4b). This effect is especially pronounced over the open ocean, where pH increases from 1 to 5, and over dust source regions such as the Sahara desert, where pH increases from 1 to 9.

Column burdens of $\text{HNO}_{3(\text{g})}$ are notably affected by the elevated pH, decreasing by approximately 5 mg m^{-2} across the Northern Hemisphere compared to the fTEQ_noAlk scenario (Fig. S1d, g in the Supplement). This decline predominantly happens in the upper troposphere (above 500hPa, Fig. S2d, g) where concentrations decrease by about 0.5 $\mu\text{g m}^{-3}$ (Fig. S2 in the Supplement). Conversely, column burdens of fine NO_3^- increase substantially, by more than 5 mg m^{-2} over dusty regions such as the Saharan desert and Middle East (Fig. S1e, h). In remote oceanic regions, fine NO_3^- column burdens increase by approximately 1 mg m^{-2} , with transoceanic transport at low latitudes showing column loads ranging from 4 to 7 mg m^{-2} (Fig. S1h in the Supplement).

490 At the surface, the inclusion of dust and SS alkalinity has a minimal impact on continental $\text{HNO}_{3(\text{g})}$ concentrations (Fig. 3d, g). Surface fine NO_3^- concentrations of about $0.5 \mu\text{g m}^{-3}$ are present across the equatorial belt, with higher values, up to $2 \mu\text{g m}^{-3}$, over remote dusty regions. However, no significant increase in fine NO_3^- concentrations is simulated in polluted areas such as Europe and Asia (Fig. 3e, h and Fig. S2e, h).

These results suggest that the condensation of $\text{HNO}_{3(\text{g})}$ on dust and SS happens mostly during long-range transport, representing a key driver of nitrate formation in this scenario. They also indicate that incorporating reactions R8-12 increases fine NO_3^- concentrations, with the majority forming on dust rather than on SS. This is attributed to the relatively higher alkalinity of dust compared to SS, as derived from the Journet et al. (2014) dataset (see Sect. 2.2.3).

3.1.2 Effects of assuming irreversible formation of coarse nitrate

The HYB methods exclude the possibility of coarse NO_3^- evaporating back to the gas phase, which may lead to positive biases in global NO_3^- burdens.

In the HYB_duUPTK sensitivity run, coarse NO_3^- forms exclusively on dust, redistributing the available $\text{HNO}_{3(\text{g})}$ towards coarse NO_3^- production in regions affected by dust, such as the Middle East and East Asia. Coarse NO_3^- column loads average between $0.5\text{-}4.0 \text{ mg m}^{-2}$ across the Northern Hemisphere, with peaks reaching $10\text{-}20 \text{ mg m}^{-2}$ (Fig. S1 l in the Supplement). Notably, this impact is less pronounced in the Saharan dust belt, where limited $\text{HNO}_{3(\text{g})}$ availability constrains coarse NO_3^- formation. The production of coarse NO_3^- is accompanied by a modest reduction in fine NO_3^- column loads over dusty regions (around 1 mg m^{-2}) compared to the fTEQ_du-ssAlk simulation, driven by availability of $\text{HNO}_{3(\text{g})}$ for fine NO_3^- formation. The uptake of $\text{HNO}_{3(\text{g})}$ significantly influences the long-range transport of NO_3^- , showcasing notable transatlantic and transpacific transport (Fig. S1 in the Supplement). This transport predominantly occurs below 800hPa, with concentrations around $1 \mu\text{g m}^{-3}$, although lower concentrations of coarse NO_3^- are also discernible at altitudes up to 400hPa (Fig. S2 in the Supplement). At the surface, coarse NO_3^- concentrations also increase, ranging from 2 to $10 \mu\text{g m}^{-3}$ over regions such as China, India, and the Middle East (Fig. 3 l).

The HYB_du-ssUPTK run investigates the role of SS by enabling both R4 and R5 (Table 1). The uptake of $\text{HNO}_{3(\text{g})}$ on SS enhances the formation of coarse NO_3^- over open ocean and those coastal areas affected by SS outbreaks (Fig. S11 in the Supplement). Surface concentrations in Europe and North-America reach $1\text{-}2 \mu\text{g m}^{-3}$, with regions exceeding $5 \mu\text{g m}^{-3}$ significantly expanding (Fig. 3o). Coarse NO_3^- becomes also notable over West Siberia, East South-America and South Africa ($0.5\text{-}1 \mu\text{g m}^{-3}$) and over remote oceanic regions ($0.2 \mu\text{g m}^{-3}$). Its presence extends to high altitudes in the Southern Hemisphere (Fig. S2o in the Supplement). Long-range transport of coarse NO_3^- is enhanced over oceans, with column burdens increasing by approximately 3.5 mg m^{-2} (Fig. S1o in the Supplement), compared to the HYB_duUPTK case. This increase is attributed to greater depletion of $\text{HNO}_{3(\text{g})}$ through irreversible uptake over dust and SS (Fig. S1m in the Supplement), resulting in lower $\text{HNO}_{3(\text{g})}$ concentrations in the atmosphere compared to the other simulations.

Results obtained by increasing the value of $\gamma(\text{HNO}_3)$ in HYB_g0p1 scenario (Supplementary Section S5 and Fig. S3) provide insights into the nature of the coarse NO_3^- transoceanic transport. While the inclusion of UPTK on SS enhances long-range transport, the HYB_g0p1 run shows a slight increase in coarse NO_3^- compared to HYB_du-ssUPTK, particularly

over the Sahara and across the Atlantic and Pacific oceans. Although these differences are not major, they suggest that the
525 availability of $\text{HNO}_{3(g)}$ may act as a limiting factor for coarse NO_3^- formation.

To further explore this phenomenon, we used the HYB_DL run, which is equivalent to HYB_du-ssUPTK run but constrains the available gas using the DIFFLIM calculation (see Methods). This approach results in coarse NO_3^- column burdens just below those obtained in HYB_g0p1. The similarity between HYB_DL and HYB_du-ssUPTK can be attributed to two possible factors: 1) HYB_DL conserves more $\text{HNO}_{3(g)}$ for reactions on coarse dust during long-range transport via reactions R4 and R5
530 (Table 1) or 2) an intrinsic overestimation of the uptake coefficients, leading to similar coarse NO_3^- formation rates regardless of the $\text{HNO}_{3(g)}$ available. This is further discussed in Section 3.2.

Overall, results show that the HYB mechanisms preferentially condense $\text{HNO}_{3(g)}$ via TEQ in the fine mode and over SS through R5 in the coarse mode, rather than over coarse dust through R4, as noted here and also assessed in subsequent sections.

535 A comparison of the spatial distributions obtained by the EMEP, INCA and GMI models in Bian et al. (2017), which employ HYB approaches similar to our HYB_du-ssUPTK run, reveals comparable trends. All models show significant formation of total NO_3^- (fine + coarse) over polluted regions, although our results, ranging from 10 to 25 mg m^{-2} , generally exceed the column loads reported by AeroCom models, that range from 7 to 16 mg m^{-2} (Fig. S1n, o in the Supplement). AeroCom models do not show such the pronounced transport of coarse NO_3^- across the North Atlantic observed in our study, with
540 column burdens of 0.2-0.5 mg m^{-2} , compared to 4 mg m^{-2} in our results. Section 3.4 further investigates the excessive formation of particulate NO_3^- in the HYB mechanism in terms of the total nitrogen budget.

A closer comparison with Hauglustaine et al. (2014), using the LMDz-INCA model with UPTK reactions on dust and SS, shows strong agreement with our HYB_duUPTK run but not with HYB_du-ssUPTK. For instance, concentrations of total NO_3^- over polluted areas in Hauglustaine et al. (2014) align with HYB_duUPTK in both geographical distribution and magnitude
545 (14-20 mg m^{-2} in HYB_duUPTK vs. 10-20 mg m^{-2} in Hauglustaine et al. (2014), Fig. S1 l). Similarly, Hauglustaine et al. (2014) simulates fine NO_3^- transport downwind of the Sahara and coarse NO_3^- across the North Atlantic, with column burdens around 1-2 mg m^{-2} , slightly below our HYB_duUPTK results of 1-5 mg m^{-2} . This preferential similarity of Hauglustaine et al. (2014) to HYB_duUPTK rather than HYB_du-ssUPTK, may be attributed to the $\text{HNO}_{3(g)}$ UPTK coefficient for SS employed in Hauglustaine et al. (2014). For this reaction, instead of a constant value as used in HYB_du-ssUPTK, Hauglustaine
550 et al. (2014) uses a function of RH ranging from $1 \cdot 10^{-3}$ to 0.1, leading to a lower NO_3^- formation over SS in their model. This comparison suggests a potential overestimation of the UPTK coefficient for $\text{HNO}_{3(g)}$ on SS in the HYB_du-ssUPTK simulation. This conclusion is further discussed in the observational evaluation (Section 3.2) and the budgets analysis (Section 3.4). As a final remark, it is important to note that, similar to the findings of Hauglustaine et al. (2014), the HYB_duUPTK simulations overestimates particulate NO_3^- concentrations compared to observational data, as discussed in Section 3.2.

555 To further investigate the source of the coarse NO_3^- overestimation in the HYB_du-ssUPTK simulation, we compare our spatial distributions with those from Jones et al. (2021), which employs a similar HYB mechanism using the Met Office Unified Model (UM). In Jones et al. (2021), fine NO_3^- formation is computed using adaptations from Hauglustaine et al. (2014), specifically by testing different accommodation coefficients for $\text{HNO}_{3(g)}$ onto preexisting ammonium-nitrate aerosols—though

this adjustment has minimal impact on coarse NO_3^- concentrations. Coarse NO_3^- is formed through the uptake of $\text{HNO}_{3(\text{g})}$,
 560 with coefficients sourced from Fairlie et al. (2010) for dust and Burkholder et al. (2020) for SS. Despite using different uptake
 coefficients for $\text{HNO}_{3(\text{g})}$ on SS, the resulting coarse NO_3^- distribution is broadly comparable to that of the HYB_du-ssUPTK
 simulation. However, key differences emerge between the two models. Compared to Jones et al. (2021), HYB_du-ssUPTK
 exhibits a stronger tendency for coarse NO_3^- formation over polluted areas such as Europe, eastern North-America and Asia
 ($8\text{--}20\text{ mg m}^{-2}$ in HYB_du-ssUPTK vs. 4 mg m^{-2} in Jones et al. (2021)), as well as over oceanic regions ($1.5\text{--}3.0$ vs. 0.4--
 565 2.0 mg m^{-2}), rather than over sub-Saharan areas. Remarkably, significant North-Atlantic transport of coarse NO_3^- ($2.0\text{--}4.0$
 mg m^{-2}) is only present in our simulation and is absent in Jones et al. (2021). The global budget comparison in Section 3.4
 and Table 5 further highlights that Jones et al. (2021) reports significantly lower global NO_3^- burdens compared to our HYB
 runs, including HYB_duUPTK. This suggests that the HYB mechanisms in our simulations may excessively produce coarse
 NO_3^- , as both the burdens and distributions of coarse NO_3^- in Jones et al. (2021) are consistently lower, despite also accounting
 570 for SS in nitrate formation.

Nevertheless, several alternative explanations could account for the differences between the models. These include differ-
 ences in the fine NO_3^- formation mechanism, variations in the study period (2018 in our study vs. a 20-year average in Jones
 et al. (2021)), and differing parameterizations of $\text{HNO}_{3(\text{g})}$ uptake on SS.

Overall, the comparison of spatial distributions between our HYB mechanisms and those reported in the literature support
 575 the initial hypothesis that the assumption of irreversible $\text{HNO}_{3(\text{g})}$ uptake, and potentially the uptake coefficients used, may
 contribute to the overestimation of NO_3^- formation. This issue is examined in greater detail in Sections 3.2 and 3.4.

3.1.3 Effects of accounting for reversible formation of coarse nitrate

The DBCLL simulations (described in Section 2.4 and Tables 2 and 3) assess the impact of reversible chemistry and the role
 of dust and SS NVC in nitrate partitioning.

580 The DBCLL_noAlk run produces spatial distributions of $\text{HNO}_{3(\text{g})}$ and fine NO_3^- similar to those in the fTEQ_noAlk run.
 $\text{HNO}_{3(\text{g})}$ remains predominantly in the gas phase across the Northern Hemisphere (Fig. S1p in the Supplement), while fine
 NO_3^- primarily associates with fine NH_4^+ over regions lacking dust or SS influence, such as parts of China, India, Europe
 and eastern North-America (Fig. 3q). Minimal concentrations of fine NO_3^- are simulated at higher altitudes (Fig. S2q in the
 Supplement), though fine NH_4^+ coexists with sulfate particles at elevated levels, as observed in the fTEQ_du-ssAlk case. In
 585 contrast, coarse NO_3^- is nearly absent in DBCLL_noAlk, with only minimal surface concentrations over northern China (0.4
 $\mu\text{g m}^{-3}$) and no clear correlation with dust and SS presence (Fig. 3r). Coarse NH_4^+ formation is similarly limited, with slight
 presence over the Persian Gulf and the Sahara ($0.02\text{--}0.1\text{ mg m}^{-2}$), primarily in anthropologically polluted regions where it
 forms alongside coarse SO_4^{2-} (0.2 mg m^{-2}) (Fig. S5r and Fig. S8r). The simultaneous formation of coarse NH_4^+ and SO_4^{2-}
 particles is attributed to diffusion limitation, where $\text{NH}_{3(\text{g})}$ preferentially neutralizes coarse sulfate particles.

590 The differences between fTEQ_noAlk and DBCLL_noAlk illustrate the impact of diffusion limitation before the partitioning
 of semi-volatile species. While fine NO_3^- formation remains consistent between both runs, the assumption of allowing all acid
 species to partition into the fine mode in fTEQ_noAlk leads to enhanced fine particle production and long-range transport—an

effect not observed in DBCLL_noAlk. Consistently, the pH in the fine mode closely mirrors that of fTEQ_noAlk, while the pH of particles in the coarse mode indicates even greater acidity (Fig. 4f1, f2)).

595 Including dust and SS alkalinity in DBCLL_du-ssAlk produces an important increase in the pH over arid regions and oceanic regions—from 1 to 5-6 in the fine mode and above 7 in the coarse mode—compared to DBCLL_noAlk (Fig. 4h1, h2). Relative to DBCLL_du-Alk (dust alkalinity only, see Supplementary Section S5), DBCLL_du-ssAlk redistributes the NO_3^- partitioning toward the coarse mode: fine NO_3^- concentrations approximately halve across regions, while coarse NO_3^- surface concentrations and column loads roughly double over dusty regions but not over the ocean (Fig. 3w, x and Fig. S1w, x). This shift leads to enhanced transatlantic and transpacific transport, with column loads of $2\text{--}4\text{ mg m}^{-2}$, surface concentrations of $0.4\text{--}1.0\text{ }\mu\text{g m}^{-3}$ and zonal averages of $0.5\text{--}1.0\text{ }\mu\text{g m}^{-3}$. Over continental regions with dust concentration, coarse NO_3^- reaches column loads of $4\text{--}5\text{ mg m}^{-2}$ and surface concentrations of $0.5\text{--}1.0\text{ }\mu\text{g m}^{-3}$. Additionally, $\text{HNO}_{3(\text{g})}$ surface concentrations increase from 0.1 to $1.0\text{ }\mu\text{g m}^{-3}$ in the heavily polluted areas (Fig. 3s, v). The enhancement of coarse NO_3^- formation is attributed to SS NVC (see Section 2.2.3), with Cl^- suppressing fine NO_3^- formation and leaving more $\text{HNO}_{3(\text{g})}$ available for transport and subsequent coarse NO_3^- formation over dusty regions.

Including SS alkalinity reduces fine NH_4^+ loads over Europe and Eastern North America from 3 to 1 mg m^{-2} and slightly lowers values over Asia ($2\text{--}3\text{ mg m}^{-2}$). Conversely, coarse NH_4^+ increases over Europe and Eastern North America ($0.1\text{--}0.2\text{ mg m}^{-2}$), Asia (up to 1.0 mg m^{-2}), and transoceanic regions (0.1 mg m^{-2}) (Fig. S5x in the Supplement). This redistribution mirrors the particulate NO_3^- patterns, highlighting the strong coupling between these species when TEQ is applied to both fine and coarse partitioning. These results emphasize the sensitivity of nitrate formation to SS alkalinity, particularly due to basic NVCs like Na^+ .

Compared to the HYB_du-ssUPTK run, DBCLL_du-ssAlk shows lower coarse NO_3^- concentrations and shifts coarse NO_3^- formation toward dusty rather than anthropologically polluted areas. For example, over Europe, surface concentrations in DBCLL_du-ssAlk do not exceed $0.5\text{ }\mu\text{g m}^{-3}$, while HYB_du-ssUPTK reports $1\text{--}2\text{ }\mu\text{g m}^{-3}$ (Fig. 3x).

615 Overall, regardless of the mechanism used, the results demonstrate the high sensitivity of NO_3^- formation to dust and SS alkalinity, significantly altering its spatial distribution (reactions R8 to R12, Table 1). However, implementing the DIFFLIM calculation in the DBCLL mechanism has a limited impact, as shown by the comparison between fTEQ_noAlk and DBCLL_noAlk.

We compare DBCLL_du-ssAlk and the results from the EMAC model (Karydis et al., 2016), which similarly employs TEQ for the bulk gas and aerosol mass, followed by DIFFLIM. Additionally, EMAC makes use of globally heterogeneous dust alkalinity based on 12 mineralogy source data points. Results from the EMAC model are reported in Bian et al. (2017) and Karydis et al. (2016) (Table 5 and Supplementary Section S4). Surface distributions of aerosol NO_3^- are closely aligned in both models over Europe and North America, with surface concentrations ranging from 1 to $3\text{ }\mu\text{g m}^{-3}$. However, DBCLL_du-ssAlk shows higher concentrations over Asia, India and Middle East compared to EMAC ($5\text{--}10\text{ }\mu\text{g m}^{-3}$ vs. $6\text{--}15\text{ }\mu\text{g m}^{-3}$, respectively). Additionally, our results show some variations over secondary areas compared to Karydis et al. (2016). EMAC reports biases, such as underestimating coarse NO_3^- over southern Europe and central-east Asia, attributed to excessive sulfate condensation on dust—a limitation mitigated in our study due to lower sulfate levels (Fig. S7 in the Supplement). Additionally, while EMAC

overestimates fine NO_3^- over the Arctic, this is not observed in our results due to the RH and temperature restrictions applied in ISORROPIA-II (see Section 2.2.2). Over central Africa, EMAC overpredicts coarse NO_3^- due to excessive $\text{HNO}_{3(g)}$ from biomass burning coupled with low $\text{H}_2\text{SO}_{4(g)}$ concentrations, from where Karydis et al. (2016) suggested that a HYB approach could reduce the bias. Our DBCLL_du-ssAlk and HYB runs report lower concentrations, though still potentially overestimated due to too low $\text{HNO}_{3(g)}$ levels. Total particulate SO_4^{2-} and NH_4^+ are in close agreement in both EMAC and the DBCLL_du-ssAlk run (Fig. S4w, x and S7w in the Supplement, x), with only slightly higher formation of coarse NH_4^+ over the India peninsula in MONARCH compared to EMAC.

635 3.2 Evaluation with observations

To assess the performance of each sensitivity run, we compare simulated surface concentrations of key species involved in nitrate formation with observational data, as described in Section 2.6. This evaluation includes analyses of both statistics and time series of monthly mean values, focusing on how effectively each modeled mechanism captures observed nitrate variability on a global scale. Globally averaged results are shown in Fig. 5, while regio-specific evaluations for Europe, Asia, and Central-
640 North America are presented in Fig. S12 in the Supplement. It is important to note that the number of stations and spatial coverage vary significantly by species, as illustrated in Fig. S13. Consequently, certain evaluation metrics may not fully reflect global model performance, and could instead represent accuracy in data-rich regions.

Correlation coefficients, bias and RMSE metrics for the different sensitivity simulations are reported in Table 4, with methodological details outlined in Supplementary Section S3. A detailed evaluation of the HYB_g0p1, DBCLL_duAlk and DB-
645 CLL_ClaqAlk simulations is available in Supplementary Section S5. Additionally, a similar evaluation for total reduced and oxidized nitrogen species is presented in Supplementary Section S6 and further discussed in Section 3.3.2.

3.2.1 Nitrate species

Fine particulate nitrate ($\text{PM}_{2.5}\text{NO}_3$) is consistently formed across all mechanisms (through TEQ), with the exception of DB-CLL scheme, which incorporates a DIFFLIM coefficient to limit its formation. Consequently, the fTEQ and HYB mechanisms
650 produce higher $\text{PM}_{2.5}\text{NO}_3$ from October to April ($0.5\text{--}1.0 \mu\text{g m}^{-3}$) than DBCLL, which reports approximately half of these values. This discrepancy is more pronounced over North and Central America than in Europe, although both regions exhibit similar seasonal patterns. From April to October, differences among mechanisms diminish, except for the HYB_du-ssUPTK, which significantly overestimates $\text{PM}_{2.5}\text{NO}_3$ during these months. Notably, the runs considering the effect of dust and SS alkalinity in the partitioning process reduce the negative bias of $\text{PM}_{2.5}\text{NO}_3$ significantly, especially during February.

Greater variability among experiments emerges when evaluating $\text{HNO}_{3(g)}$ and Total NO_3^- , with more observational data available over North-Central America than Europe or Asia. This variability is highly influenced by how each mechanism models coarse NO_3^- formation, especially over Asia (Fig. S12 in the Supplement). For example, the HYB runs, which rely on irreversible uptake reactions, consistently overestimate total NO_3^- across all regions. In contrast, the DBCLL runs, which assume reversible partitioning, better align with observations.

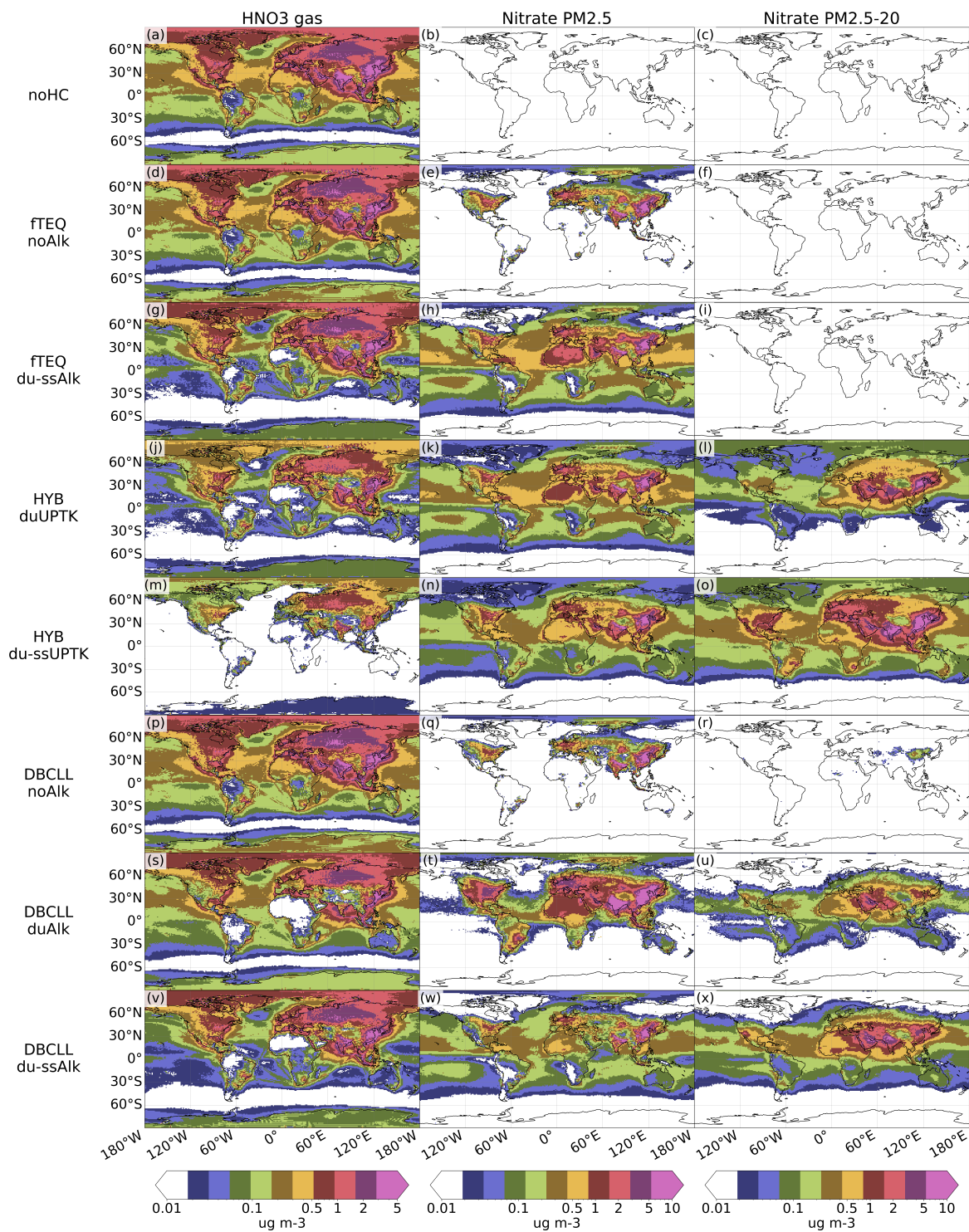


Figure 3. Surface concentrations ($\mu\text{g m}^{-3}$) of $\text{HNO}_3(\text{g})$, fine and coarse NO_3^- simulated by the different mechanisms, averaged for 2018.

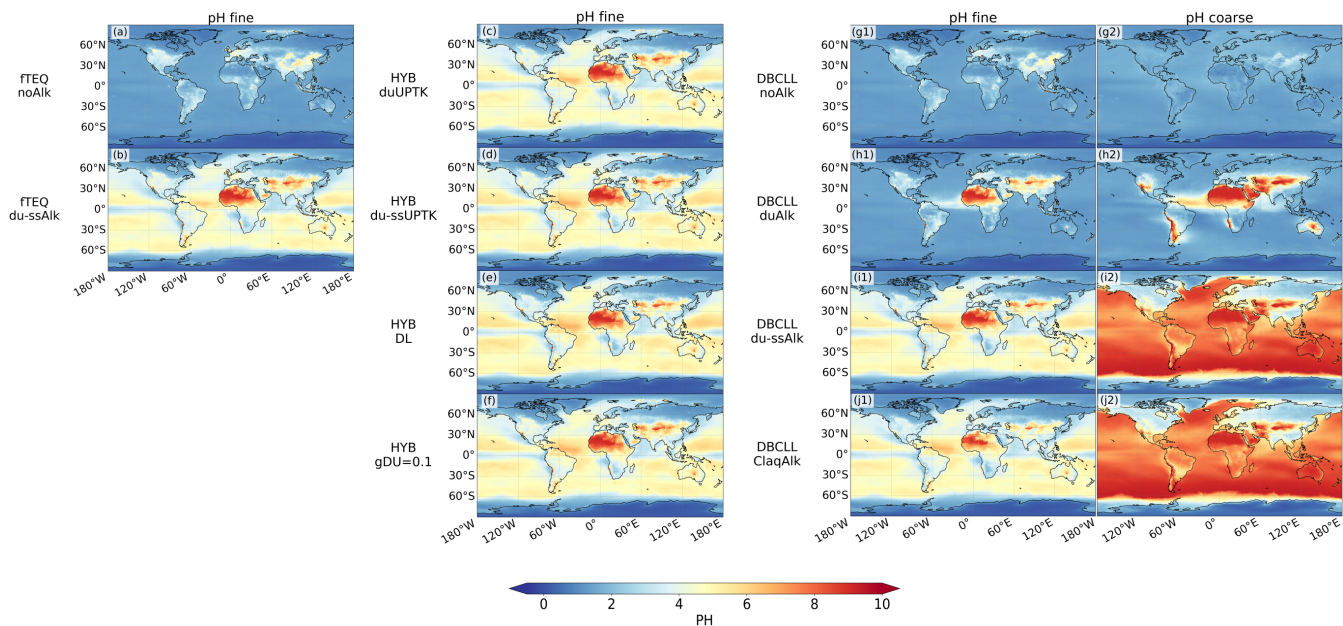


Figure 4. Surface pH of the studied mechanisms for fine and coarse aerosol size modes, averaged for 2018.

From the HYB simulations, the HYB_duUPTK run demonstrates strong correlation coefficients for both fine and total NO_3^- concentrations (0.8 and 0.9, respectively). The seasonal variability in this run closely resembles that of the fTEQ_noAlk and fTEQ_du-ssAlk experiments, which fully neglect coarse NO_3^- formation. This similarity suggests a limited amount of sites affected significantly by coarse NO_3^- formation, while highlighting the relevant role of the TEQ calculation in the fine mode. Despite showing a systematic positive bias (of $0.03/0.04/0.29 \mu\text{g m}^{-3}$ for $\text{HNO}_{3(\text{g})}$, fine NO_3^- , and Total NO_3^- , respectively), HYB_duUPTK maintains consistency across continents. However, it notably overestimates Total NO_3^- over Asia in November and over Europe from May to September, likely due to dust events affecting the monitoring sites.

Including the UPTK reactions on SS in the HYB_du-ssUPTK run exacerbates this overestimation due to excessive $\text{HNO}_{3(\text{g})}$ condensation into coarse NO_3^- . Specifically, the correlation for Total NO_3^- drops from 0.9 to 0.66, $\text{HNO}_{3(\text{g})}$ concentrations are underestimated (bias of $-0.10 \mu\text{g m}^{-3}$), and Total NO_3^- surface concentrations are overestimated by $1.50 \mu\text{g m}^{-3}$ —the highest bias among all sensitivity runs (Table 4). This overestimation is most pronounced over Asia ($3\text{--}5 \mu\text{g m}^{-3}$ vs. $1 \mu\text{g m}^{-3}$) and North-Central America ($1.5\text{--}2.5 \mu\text{g m}^{-3}$ vs. $1.0 \mu\text{g m}^{-3}$). Discrepancies in European sites are more pronounced during summer and fall (Fig. S12 in the Supplement).

To evaluate whether this overestimation results from unlimited $\text{HNO}_{3(\text{g})}$, we examined the HYB_g0p1 and HYB_DL configurations (Supplementary Section S5). Increasing the uptake coefficient $\gamma(\text{HNO}_3)$ in HYB_g0p1 further enhances the overestimation of total NO_3^- by an additional 30%, suggesting that $\text{HNO}_{3(\text{g})}$ availability is not the limiting factor. Similarly, HYB_DL shows only a marginal reduction in total NO_3^- bias (from 1.5 to $1.3 \mu\text{g m}^{-3}$), reinforcing that the overestimation stems from

the overly high $\gamma(\text{HNO}_3)$ values for SS particles, as previously discussed in Section 3.1.2. This highlights the need for refining HYB schemes.

Conversely, DBCLL mechanisms markedly reduce surface concentration biases. Without alkalinity, the total NO_3^- bias shifts from 0.29 to $-0.43 \mu\text{g m}^{-3}$ between the HYB and DBCLL runs. Incorporating dust and SS NVC further improves the bias, reducing it from 1.5 to $-0.1 \mu\text{g m}^{-3}$ between the HYB_du-ssUPTK and DBCLL_du-ssAlk runs (Table 4). All consistently achieve high correlation coefficients (aprox. 0.8).

However, the DBCLL_noAlk run tends to underestimate total NO_3^- (bias of $-0.43 \mu\text{g m}^{-3}$) and overestimates $\text{HNO}_{3(\text{g})}$ (bias of $+0.13 \mu\text{g m}^{-3}$), failing to capture its seasonal cycle (correlations of 0.06). Surprisingly, the fTEQ runs outperform DBCLL_noAlk despite neglecting coarse NO_3^- formation.

Including dust and SS alkalinity in DBCLL_du-ssAlk substantially improves these biases. Enhanced condensation of $\text{HNO}_{3(\text{g})}$ into both fine and coarse NO_3^- (Fig. 5) yields a reduced bias in total NO_3^- ($-0.10 \mu\text{g m}^{-3}$) and a strong seasonal agreement, with correlation coefficients of 0.82 and 0.78 for fine and total NO_3^- , respectively. These results also outperform the DBCLL_duAlk run, which excludes SS alkalinity (Supplementary Section S5 and Table S7).

The evaluation of the DBCLL experiments highlights the paramount importance of accounting for NO_3^- formation on both dust and SS to accurately represent $\text{HNO}_{3(\text{g})}$ and particulate NO_3^- concentrations.

The sensitivity of model performance to the dust NVC representation is further assessed using the DBCLL_Claq run, which applies an average dust alkalinity based on Claquin et al. (1999) instead of Journet et al. (2014) (see Sect. 2.2.3). Results show limited differences compared to DBCLL_du-ssAlk, with correlation and bias for total NO_3^- just slightly improved in DBCLL_Claq with 0.81 and $-0.09 \mu\text{g m}^{-3}$, respectively (Supplementary Section S5 and Table S7). Timeseries from Fig. S14, compared to Fig. 5, also shows comparably similar results when using DBCLL with Claquin et al. (1999) or Journet et al. (2014) mineral datasets. Overall, based on the limited number of observation sites used in our evaluation, the impact of dust NVC representation on surface NO_3^- concentrations does not appear to be significant in the observational points.

3.2.2 Ammonia and particulate ammonium

Results for $\text{NH}_{3(\text{g})}$ and particulate NH_4^+ are generally more consistent across all mechanisms compared to NO_3^- outcomes. This consistency is expected, as all sensitivity runs use the same condensation pathway for the conversion of $\text{NH}_{3(\text{g})}$ to fine NH_4^+ through reactions R6 and R7 (Table 1). All mechanisms effectively capture the seasonal cycle of both $\text{NH}_{3(\text{g})}$ and fine NH_4^+ . Correlation coefficients for $\text{NH}_{3(\text{g})}$ range from 0.85 to 0.88, while those for fine NH_4^+ range from 0.59 to 0.77. However, model performance is slightly weaker for total NH_4^+ , with correlations between 0.41 and 0.54.

The DBCLL_du-ssAlk and DBCLL_Claq runs yields the lowest errors for NH_4^+ . RMSEs are $0.17 \mu\text{g m}^{-3}$ for fine and $0.12 \mu\text{g m}^{-3}$ for total concentrations, with corresponding biases of $-0.02 \mu\text{g m}^{-3}$ and $0.08 \mu\text{g m}^{-3}$ (Tables 4 and S7). These results highlight the significant sensitivity of NH_4^+ predictions to the treatment of dust and SS non-volatile cations (NVCs). Including or omitting NVC representation notably affects the biases across different schemes. Unlike NO_3^- , increased particle alkalinity reduces the condensation of $\text{NH}_{3(\text{g})}$, thereby limiting NH_4^+ formation. This effect helps mitigate biases when compared to observational data.

Overall, biases for both fine and total NH_4^+ remain relatively low in most experiments, ranging from -0.02 to $0.29 \mu\text{g m}^{-3}$.

3.2.3 Sulfur species

Results for $\text{SO}_{2(\text{g})}$ and particulate SO_4^{2-} are consistent across all sensitivity simulations. This consistency is expected since sulfate formation is independent of the pathways used for coarse NO_3^- formation. In all cases, particulate SO_4^{2-} is produced from $\text{SO}_{2(\text{g})}$ and $\text{H}_2\text{SO}_{4(\text{g})}$ in a similar manner (see Section 2.3). Differences solely arise from variations in the treatment of the $\gamma(\text{SO}_2)$ coefficient when accounting for dust alkalinity, which introduces slight differences in SO_4^{2-} formation. Overall, the model slightly underestimates both $\text{SO}_{2(\text{g})}$ and total SO_4^{2-} , particularly at European monitoring sites. In contrast, fine SO_4^{2-} is marginally overestimated, mainly driven by observations from Central-North American sites. Temporal correlation coefficients average around 0.6, with slightly lower values for total SO_4^{2-} . For $\text{SO}_{2(\text{g})}$, the slight mismatch with observational data primarily stems from the European evaluation, which presents a consistent negative bias of $-0.5 \mu\text{g m}^{-3}$ throughout the studied period (Fig. S12 in the Supplement). Fine and total SO_4^{2-} concentrations also show generally good agreement with observations. However, there is a modest overestimation in fine SO_4^{2-} concentrations, with an average bias of $0.30 \mu\text{g m}^{-3}$, particularly over Central and North America. Conversely, total particulate SO_4^{2-} surface concentrations are underestimated over Europe, with an average bias of -0.43 (Table 4 and Fig. S12 in the Supplement). The negative bias in SO_4^{2-} concentrations may be attributed to the sulfate formation scheme and sulfur emission inventories employed in our runs. Reevaluating the $\text{SO}_{2(\text{g})}$ uptake coefficient could provide valuable insights. The current implementation uses the uptake coefficient function from Fairlie et al. (2010). Alternative formulations have been suggested in the literature, such as the uptake coefficient values proposed by Phadnis and Carmichael (2000) for $\text{SO}_{2(\text{g})}$ uptake on dust, and by Song and Carmichael (2001) for uptake on SS, as reported in Li et al. (2012). Exploring these alternatives may help improve the model's representation of sulfur species.

3.2.4 Fine and total particulate matter

Results for $\text{PM}_{2.5}$ and PM_{10} are consistent across the different sensitivity runs (Fig. 5), indicating that the total mass is more influenced by components other than secondary inorganic species.

Globally, the seasonal cycles of $\text{PM}_{2.5}$ are well reproduced across all sensitivity runs, with correlation coefficients consistently around 0.75. However, annual $\text{PM}_{2.5}$ concentrations are overestimated by an average of $20 \mu\text{g m}^{-3}$ (Table 4). In contrast, PM_{10} concentrations over Asia show good agreement with observations, indicating that the model underestimates coarse particulate matter (PM), especially at monitoring sites in mainland China (Supplementary Figures S12 and S13). It is important to note that regions with significant $\text{PM}_{2.5}$ overestimation do not overlap with the limited sites where chemical composition measurements are available—locations where the model accurately represents total particulate NO_3^- and NH_4^+ (Supplementary Fig. S12). Our results suggest that the absence of anthropogenic coarse dust emissions in the CAMS emission inventory may hinder the formation of coarse inorganic species, including NO_3^- . The formation of coarse NO_3^- is critical for scavenging $\text{HNO}_{3(\text{g})}$ and suppressing the formation of fine particulate NO_3^- , by providing surfaces for heterogeneous reactions. The significant influence of typically neglected anthropogenic coarse PM emissions on nitrate formation has been

emphasized in previous studies (e.g. Zhai et al. (2023)), highlighting the need for improved emission inventories to better capture the interactions between coarse particles and nitrogen species.

745 In contrast, PM_{2.5} model results over Europe and North and Central America are well aligned with observations (Fig. S12). The evaluation of PM₁₀ shows just slightly worsened RMSE (ranging from 22.49 to 30.42 $\mu\text{g m}^{-3}$) and correlation coefficients (from 0.36 to 0.46) compared to PM_{2.5} (RMSE from 24.74 to 26.16 $\mu\text{g m}^{-3}$ and correlations from 0.74 to 0.76, Table 4). This is attributed to the PM₁₀ surface concentrations underestimation over North and Central America, specially from April to September, which does not seem to be caused by the underestimation of any secondary inorganic aerosol species studied in
750 this work. Over Europe, our results are in agreement with those reported by Jones et al. (2021).

3.3 Nitrogen partitioning

In this section, we explore the impact of coarse particulate NO₃⁻ formation on the overall partitioning of atmospheric nitrogen between the gas and particle phases (Section 3.3.1). Additionally, we analyze how this formation influences the distribution of
755 nitrogen between its oxidized and reduced forms (Section 3.3.2). The observational evaluation of surface concentrations for total reduced and oxidized nitrogen is presented in Supplementary Section S6.

3.3.1 Partitioning between gas and particle phases

The partitioning of atmospheric nitrogen species in gas and aerosol phases varies according to the underlying chemical assumptions. We analyze the average atmospheric burden and deposition of nitrogen species in the MONARCH model across
760 sensitivity runs (Fig. 6 and Table S13 in the Supplement). The simulation labeled as noHC, excluding heterogeneous chemistry involving nitrate, serves as the baseline for understanding the distribution of nitrogen species in the gas phase. Under this configuration, the dominant nitrogen species are HNO_{3(g)}, peroxyacetyl nitrate (PAN), NH_{3(g)}, and NO_x, with burdens of 1.06, 0.8, 0.71, and 0.38 TgN, respectively. The longer atmospheric lifetimes of HNO_{3(g)} and PAN result from their chemical stability compared to more reactive NO_x species. Notably, N₂O_{5(g)} plays a key role in nighttime HNO_{3(g)} production via
765 heterogeneous hydrolysis (Riemer et al., 2003).

Total nitrogen burden is highly sensitive to the nitrate formation mechanism. Schemes promoting efficient coarse-mode NO₃⁻ formation significantly reduce nitrogen burdens. For example, assuming irreversible condensation of HNO_{3(g)} onto coarse particles reduces nitrogen load by 32% and 40% when considering dust only (HYB_duUPTK) and both dust and SS (HYB_du-ssUPTK), respectively. In HYB_duUPTK, 0.13 TgN partitions into fine and coarse NO₃⁻ on dust. Including SS
770 (HYB_du-ssUPTK) further increases consumption of HNO_{3(g)} and coarse NO₃⁻ formation leading to a total NO₃⁻ burden 0.27 TgN, while the fine NO₃⁻ burden remains unchanged at 0.12 TgN (Supplementary Table S13).

Comparing HYB runs to fTEQ_du-ssAlk reveals that dust and SS uptake reduces HNO_{3(g)} burdens by 0.21 TgN (dust) and 0.50 TgN (dust + SS) relative to TEQ-only partitioning. This is driven by irreversible uptake in the coarse mode. Consequently, the deposition rate of NO₃⁻ increases from 80 to 140 Tg/y, and its atmospheric lifetime decreases from 2.6 to 2.3 days (Table 5).

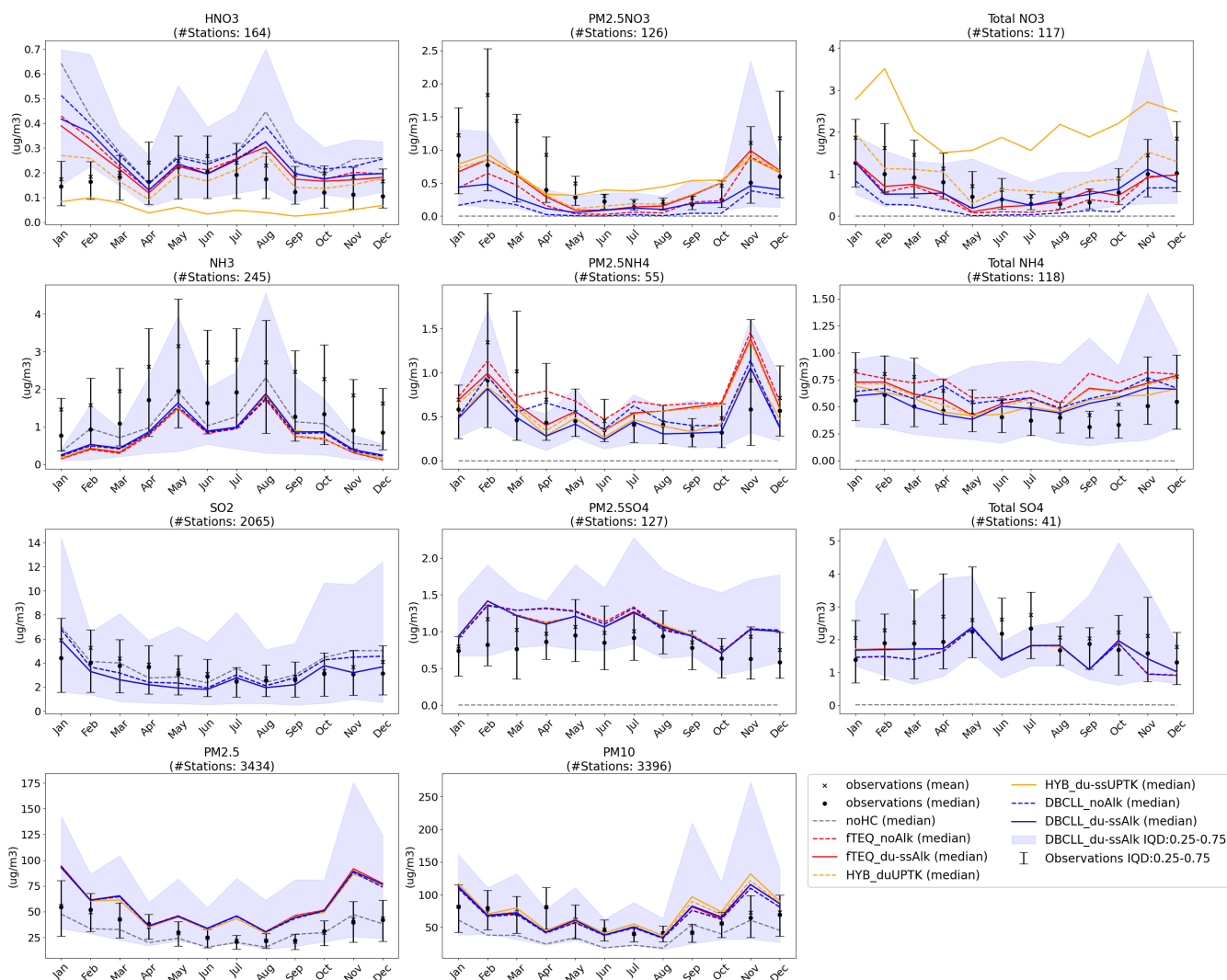


Figure 5. Observational evaluation of gas and particulate species surface concentrations. Black solid dots and crosses represent observations' monthly median and mean, respectively. Colored lines represent each configuration's monthly median over the observational points. Error bars are the observational interquartile 0.25 to 0.75 distance. Blue shading is the interquartile 0.25 - 0.75 distance for the DBCLL_du-ssAlk simulation. For the Total NO_3^- modes, the available data from Europe, Asia and North-Central America stations have been averaged together, despite data from Asia and North-Central America refer to total particle concentration, while data from Europe is limited strictly to particles with diameter up to $10 \mu\text{m}$.

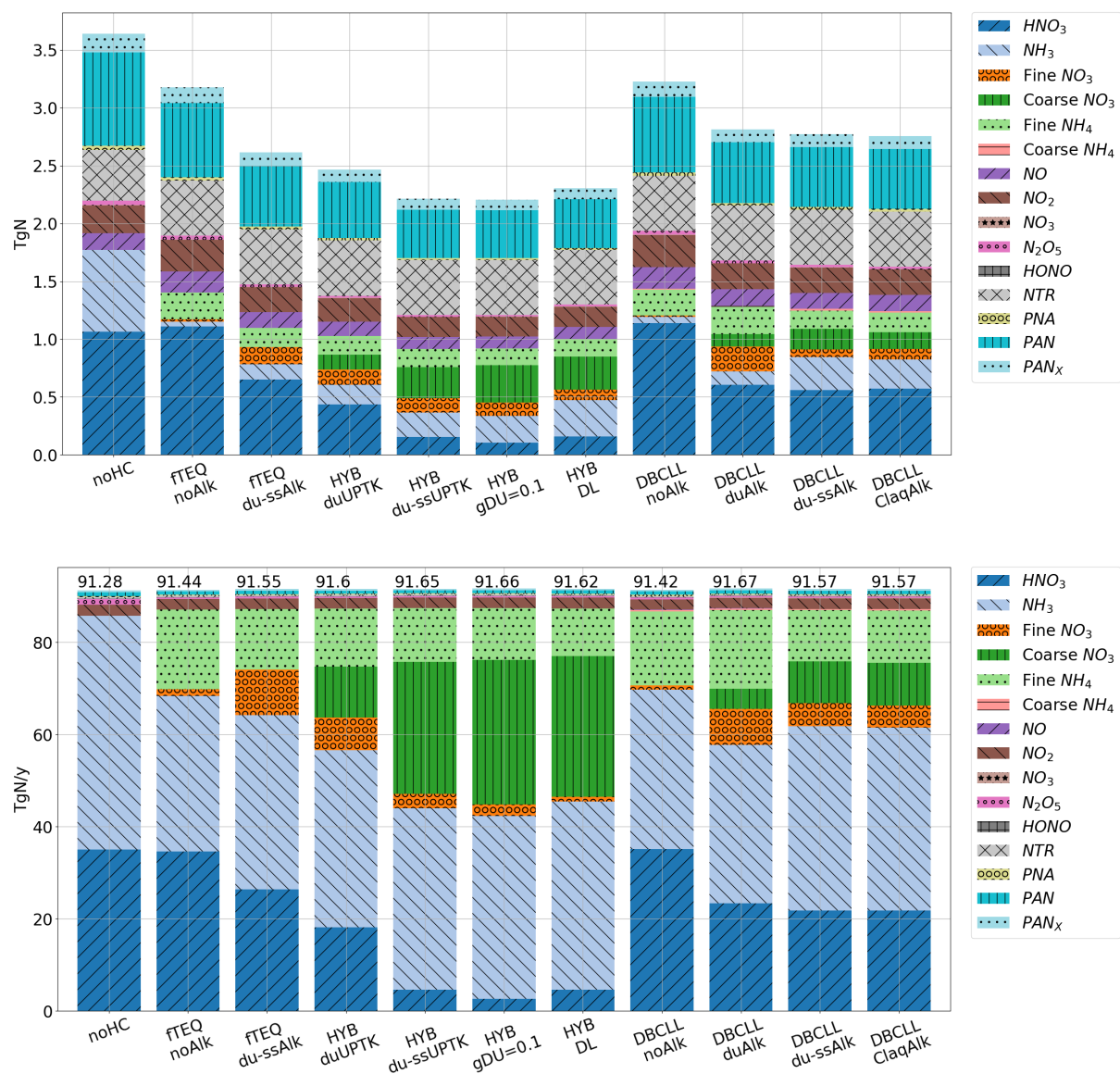


Figure 6. Atmospheric nitrogen average burdens (top) and accumulated depositions (bottom) in Tg of nitrogen (TgN) for the different dust heterogeneous chemistry mechanisms. The corresponding fractions of the nitrogen-containing species simulated in MONARCH is shown in each bar plot. Total nitrogen deposition is reported on the top of each bar of the depositions plot.

Table 4: Correlation coefficients, bias ($\mu g m^{-3}$) and root mean square error (rmse, $\mu g m^{-3}$) of each configuration's median with respect to the median of observations, corresponding to the timeseries shown in Fig. 5.

	HNO ₃			PM _{2.5} NO ₃			Total NO ₃		
	corr	bias	rmse	corr	bias	rmse	corr	bias	rmse
noHC	0.03	0.15	0.20	0.00	-0.43	0.49	0.00	-0.70	0.77
fTEQ_noAlk	0.15	0.08	0.12	0.70	-0.10	0.23	0.91	-0.19	0.25
fTEQ_du-ssAlk	0.23	0.07	0.10	0.76	0.02	0.20	0.88	-0.09	0.19
HYB_duUPTK	0.23	0.03	0.06	0.80	0.04	0.18	0.90	0.29	0.36
HYB_du-ssUPTK	0.06	-0.10	0.11	0.77	0.14	0.21	0.66	1.50	1.56
DBCLL_noAlk	0.06	0.13	0.16	0.68	-0.31	0.36	0.86	-0.43	0.46
DBCLL_du-ssAlk	0.16	0.08	0.12	0.82	-0.18	0.24	0.78	-0.10	0.23

	NH ₃			PM _{2.5} NH ₄			Total NH ₄		
	corr	bias	rmse	corr	bias	rmse	corr	bias	rmse
noHC	0.80	-0.29	0.44	0.00	-0.48	0.51	0.00	-0.45	0.46
fTEQ_noAlk	0.88	-0.65	0.69	0.64	0.29	0.35	0.41	0.26	0.28
fTEQ_du-ssAlk	0.87	-0.62	0.67	0.59	0.18	0.28	0.51	0.17	0.20
HYB_duUPTK	0.87	-0.61	0.66	0.56	0.16	0.27	0.48	0.15	0.18
HYB_du-ssUPTK	0.86	-0.55	0.60	0.70	0.03	0.16	0.54	0.10	0.13
DBCLL_noAlk	0.86	-0.58	0.62	0.66	0.10	0.20	0.51	0.16	0.18
DBCLL_du-ssAlk	0.85	-0.53	0.59	0.70	-0.02	0.17	0.47	0.08	0.12

	SO ₂			PM _{2.5} SO ₄			Total SO ₄		
	corr	bias	rmse	corr	bias	rmse	corr	bias	rmse
noHC	0.61	0.65	1.24	0.59	-0.79	0.80	0.62	-1.91	1.93
fTEQ_noAlk	0.59	0.23	1.12	0.61	0.33	0.36	0.51	-0.47	0.58
fTEQ_du-ssAlk	0.63	-0.30	0.91	0.57	0.30	0.33	0.60	-0.38	0.49
HYB_duUPTK	0.63	-0.29	0.91	0.57	0.29	0.33	0.59	-0.38	0.50
HYB_du-ssUPTK	0.63	-0.29	0.91	0.57	0.29	0.33	0.60	-0.38	0.49
DBCLL_noAlk	0.59	0.23	1.12	0.58	0.32	0.36	0.49	-0.47	0.59
DBCLL_du-ssAlk	0.63	-0.29	0.91	0.56	0.29	0.32	0.69	-0.52	0.67

Table 4 continued from previous page

	PM _{2.5}			PM ₁₀		
	corr	bias	rmse	corr	bias	rmse
noHC	0.75	-5.62	9.55	0.40	-23.16	28.42
fTEQ_noAlk	0.76	20.91	24.98	0.46	5.14	22.49
fTEQ_du-ssAlk	0.74	21.75	26.16	0.44	7.16	24.38
HYB_duUPTK	0.74	21.05	25.32	0.40	10.16	27.41
HYB_du-ssUPTK	0.76	19.59	23.91	0.36	13.92	30.45
DBCLL_noAlk	0.76	20.59	24.74	0.46	4.98	22.50
DBCLL_du-ssAlk	0.75	21.02	25.22	0.43	8.09	24.80

775 The combined UPTK on dust and SS nearly depletes $\text{HNO}_{3(g)}$, leaving only 0.15 TgN, which contributes to the overestimation of surface NO_3^- (Sections 3.1 and 3.2).

Irreversible condensation also reduces $\text{NO}_{(g)}$, $\text{NO}_{2(g)}$, and $\text{N}_2\text{O}_{5(g)}$ burdens to 0.11, 0.17, and 0.01 TgN, respectively, due to enhanced $\text{N}_2\text{O}_{5(g)}$ hydrolysis on increased particle surfaces. This reduction can be attributed to increased NO_3^- particle formation, and consequently, available surface where hydrolysis of $\text{N}_2\text{O}_{5(g)}$ can take place. Ultimately, this produces more
780 $\text{HNO}_{3(g)}$ available for partitioning into the particle phase.

Conversely, nitrogen loads increase when reversible condensation of $\text{HNO}_{3(g)}$ and $\text{NH}_{3(g)}$ is considered in the fTEQ and DBCLL sensitivity simulations. A larger proportion of nitrogen remains in the gas phase compared to HYB schemes. Notably, neglecting dust and SS alkalinity (fTEQ_noAlk, DBCLL_noAlk) slightly increases $\text{HNO}_{3(g)}$ burdens (from 1.11 to 1.14 TgN) and elevates $\text{NO}_{(g)}$ and $\text{NO}_{2(g)}$ by 0.29% and 0.12%, respectively. When dust and SS alkalinity are considered (fTEQ_du-ssAlk, DBCLL_duAlk, DBCLL_du-ssAlk), $\text{HNO}_{3(g)}$ burdens decrease by 40–50%, enhancing particulate NO_3^- formation.
785 This shift towards the aerosol phase results from the neutralization of NVC from dust and SS by $\text{HNO}_{3(g)}$. In these cases, $\text{NO}_{(g)}$, $\text{NO}_{2(g)}$ and $\text{N}_2\text{O}_{5(g)}$ budgets decrease by approximately 20% compared to analogous mechanisms without alkalinity, driven by the more efficient $\text{N}_2\text{O}_{5(g)}$ aqueous dissociation due to the increased particle presence. SS alkalinity alters the partitioning of aerosol NO_3^- from fine (-72%) towards coarse mode (+63%) when comparing DBCLL_duAlk and DBCLL_du-ssAlk, resulting in 0.06 (fine) and 0.18 TgN (coarse) burdens (Fig. 6). This surge in coarse NO_3^- formation with dust and SS
790 alkalinity causes a 26% decrease in total NO_3^- burden and a 14% increase deposition rates compared to the DBCLL_duAlk case, mainly driven by enhanced wet deposition of coarse NO_3^- from 11.0 to 27.5 Tg/y (Table 5). This can be attributed to the higher deposition efficiency of SS particles due their larger size ranges and abundance near the surface. Consequently, the presence of SS decreases the atmospheric lifetime of particulate NO_3^- from 4.9 to 3.1 days, contrasting with scenarios that
795 consider only dust alkalinity.

Additionally, SS NVCs (e.g., Na^+) inhibit $\text{NH}_{3(g)}$ condensation into NH_4^+ , reducing NH_4^+ burdens from 0.31 to 0.22 Tg (Table S11).

3.3.2 Reduced and oxidized nitrogen

Figure 7 presents the annual averages burdens and depositions of oxidized ($\text{HNO}_{3(g)} + \text{NO}_3^-$) and reduced ($\text{NH}_{3(g)} + \text{NH}_4^+$) nitrogen for each sensitivity run. These results are compared with global estimates from the AeroCom phase III nitrate experiment (Bian et al., 2017), which offers a comprehensive range of budgets derived from different global aerosol models employing mechanisms similar to those analyzed in our study. As additional references, the results from the EMAC model (Karydis et al., 2016) and the LMDz-INCA model (Hauglustaine et al., 2014) are shown separately from the average of AeroCom in Fig. 7 due to their similarity with the DBCLL and HYB mechanisms, respectively.

The HYB mechanisms yield an annual average burden of oxidized nitrogen of 0.55 TgN, slightly below the AeroCom multi-model mean but in close agreement with Hauglustaine et al. (2014). Compared to the fTEQ and DBCLL mechanisms that account for dust and sea salt (SS) alkalinity, which show higher oxidized nitrogen burdens (0.6–0.8 TgN), the HYB mechanisms fall on the lower end of the distribution. This reduction is primarily due to the efficient deposition of particulate NO_3^- and the limited atmospheric presence of $\text{HNO}_{3(g)}$ under the assumption of irreversible uptake, as discussed in Sections 3.2 and 3.3.1.

For reduced nitrogen species, the HYB mechanisms produce an average burden of approximately 0.3 TgN, aligning well with the AeroCom mean. However, HYB configurations yield lower particulate NH_4^+ burdens (0.15 TgN, Table S11 in the Supplement) and higher $\text{NH}_{3(g)}$ levels in the gas phase compared to Hauglustaine et al. (2014) and other mechanisms. Among HYB cases, $\text{NH}_{3(g)}$ budgets show an inverse relationship with $\text{HNO}_{3(g)}$, indicating that $\text{HNO}_{3(g)}$ acts as a limiting factor for the $\text{NH}_{3(g)}$ neutralization in the fine mode. Given that most $\text{HNO}_{3(g)}$ is consumed by the UPTK reactions in the HYB mechanism, these results suggest that the NH_4^+ mass formed with this configuration is mainly a product of $\text{NH}_{3(g)}$ neutralization with sulfate through the fTEQ calculation.

The fTEQ and DBCLL mechanisms exhibit a high sensitivity in the partitioning of oxidized and reduced nitrogen to alkalinity, regardless of whether coarse NO_3^- formation is considered. When dust and SS alkalinity are neglected (fTEQ_noAlk and DBCLL_noAlk), oxidized nitrogen burdens rise significantly to 1.1 TgN, primarily due to the accumulation of unreacted $\text{HNO}_{3(g)}$, which slightly exceeds the burden in the noHC run (1.05 TgN). This is attributed to the sensitivity of $\text{N}_2\text{O}_{5(g)}$ hydrolysis to aerosol loading. Conversely, reduced nitrogen burdens in these schemes remain relatively low compared to other mechanisms due to the high consumption of $\text{NH}_{3(g)}$ by neutralization of $\text{HNO}_{3(g)}$, which acts as an important sink of reduced nitrogen as particulate NH_4^+ . The introduction of dust and SS NVC in the reactions decreases oxidized nitrogen by an enhanced consumption of $\text{HNO}_{3(g)}$ to form NO_3^- . Furthermore, reduced nitrogen concentrations increase as a consequence of the lower availability of $\text{HNO}_{3(g)}$ to neutralize $\text{NH}_{3(g)}$. This phenomenon is similarly observed in simulations including TEQ calculations over fine (fTEQ_du-ssAlk) and over both fine and coarse modes (DBCLL_duAlk and DBCLL_du-ssAlk). Comparatively, the presence of dust and SS alkalinity decreases the oxidized nitrogen burdens by approximately 27% compared to analogous simulations neglecting NVC. However, contrasting the results between DBCLL_duAlk and DBCLL_du-ssAlk shows that while the presence of SS NVC further enhances the consumption of $\text{HNO}_{3(g)}$ by 20%, it concurrently reduces NO_3^- burdens by 5%. This is likely due to SS contributing with acidic anions such as Cl^- and SO_4^{2-} (see Section 2.2.3), which

hinders the formation of particulate NO_3^- in the fine mode and leaves more $\text{HNO}_{3(\text{g})}$ available to form coarse NO_3^- over dust. This is consistent with the spatial distributions discussed in Section 3.1.3 and the nitrogen budgets presented in Section 3.3.1 and Fig. 6.

835 Alkalinity leads to higher reduced nitrogen budgets due to enhanced basic conditions (high pH, as shown in Fig. 4) facilitated by dust and SS NVC. This environment inhibits $\text{NH}_{3(\text{g})}$ condensation over dust and SS particles, thereby increasing the $\text{NH}_{3(\text{g})}$ atmospheric lifetime (Table S10) and allowing it to mix with sulfate from polluted areas, as can be seen in the spatial correlation between particulate SO_4^{2-} and NH_4^+ (see discussion in Section 3.1 and Figures S4 and S7). Consequently, alkalinity facilitates the formation of NH_4^+ through the neutralization of $\text{NH}_{3(\text{g})}$ with sulfate, contrasting with the neutralization process involving
840 $\text{HNO}_{3(\text{g})}$ observed in the DBCLL_noAlk simulation.

Overall, the DBCLL_du-ssAlk run aligns with the average of AeroCom in both oxidized and reduced partitions, although its reduced phase falls below EMAC levels. This discrepancy can be attributed to EMAC showing very low deposition rates of $\text{NH}_{3(\text{g})}$ (see Table S10), which enhances its atmospheric burden.

The observational evaluation of both reduced and oxidized species, along with the monitoring stations used, can be found
845 in the Supplementary Section S6. Overall, a general good agreement with observations is obtained. Oxidized nitrogen species exhibit low biases in all mechanisms except HYB_du-ssUPTK, which overestimates observations throughout the studied period, supporting the conclusion that UPTK coefficients used are excessively efficient. Reduced nitrogen species, while slightly underestimated, remain well within the observational variability range.

850 3.4 Nitrate budgets

The NO_3^- budgets from the sensitivity runs are summarized in Table 5, including burdens, wet and dry deposition, productivities, and lifetime of fine, coarse and total NO_3^- . Comparative data from previous studies (Bian et al., 2017; Karydis et al., 2016; Rémy et al., 2022; Hauglustaine et al., 2014; Jones et al., 2021) are also included. Analogous information on $\text{HNO}_{3(\text{g})}$, $\text{NH}_{3(\text{g})}$, NH_4^+ , and SO_4^{2-} can be found in Supplementary Section S4 (Tables S9 to S12).

855 The sensitivity runs conducted in this study yield a wide range of total NO_3^- burdens, ranging from 0.09 to 1.93 Tg. Experiments neglecting coarse NO_3^- and the role of non-volatile components (NVC) fall on the lower end of this spectrum, while the HYB schemes consistently report higher burdens. The coarse to fine ratio decreases in the mechanisms that neglect part or completely NVC in the nitrate partitioning, and HYB runs consistently simulate the highest coarse burdens (0.57 to 1.43 Tg). Compared with the reported values in the literature, both HYB (1.15 to 1.93 Tg) and DBCLL (1.04 to 1.44 Tg) schemes
860 considering NVC fall within the upper range of the AeroCom total NO_3^- burden (Bian et al., 2017) and above specific global models such as IFS (Rémy et al., 2022) or MetOffice UM (Jones et al., 2021). Similar differences are found in total annual depositions with ranges that span a factor of two among schemes. For wet deposition, DBCLL mechanisms show values ranging from 37 to 45 Tg/y, within the lower half of the range found in the literature (45 ± 30 Tg/y for AeroCom to 75 Tg/y for IFS), while HYB consistently simulates higher estimates even beyond reported values (63 to 120 Tg). This feature is also seen in
865 the dry deposition. Consequently, the lifetimes of total NO_3^- shown in Table 5 range from 2.3 to 2.6 days for HYB schemes,



Figure 7. Average 2018 nitrogen budgets (top) and depositions (bottom) for the different dust heterogeneous chemistry mechanisms. Each plot shows oxidized (top rows) and reduced (bottom rows) mass for each mechanism (in columns). They are compared with literature references (three last right columns).

Table 5. Results for fine, coarse and total particulate NO_3^- obtained with the studied heterogeneous chemistry mechanisms. Results from the literature references are reported at the end of the table: The average of all the participating models in the intercomparison AeroCom phase III nitrate experiment for 2008 (*AeroCom*), specifying their standard deviation (*STD AeroCom*), results from the GMI model (*GMI*, using a similar HYB approach with UPTK reactions on dust and SS), and results from the EMAC model (*EMAC 2008*, using a similar approach to DBCLL_du-ssAlk). Also using the EMAC model, results obtained by the Karydis et al. (2016) study are reported as *EMAC 2005-2008*. Results from models using a similar approach to HYB_du-ssUPTK are reported as *IFS* for Rémy et al. (2022), *LMDz-INCA* for Hauglustaine et al. (2014) and *MetOffice UM* for Jones et al. (2021).

NO_3^-	Burden (Tg)			Wet Dep. (Tg y-1)			Dry Dep. (Tg y-1)			Total Dep. (Tg y-1)			Production (Tg y-1)			Lifetime (days)		
Experiment	Fine	Coarse	Total	Fine	Coarse	Total	Fine	Coarse	Total	Fine	Coarse	Total	Fine	Coarse	Total	Fine	Coarse	Total
fTEQ_noAlk	0.09	-	0.09	4.6	-	4.6	2.0	-	2.0	6.6	-	6.6	6.6	-	6.6	2.6	0.0	2.6
fTEQ_du-ssAlk	0.66	-	0.66	36.1	-	36.1	7.4	-	7.4	43.5	-	43.5	43.6	-	43.6	2.7	0.0	2.7
HYB_duUPTK	0.58	0.57	1.15	25.8	37.9	63.7	5.4	11.1	16.5	31.2	49.0	80.2	31.3	48.9	80.2	3.4	2.1	2.6
HYB_du-ssUPTK	0.55	1.20	1.75	10.7	103.4	114.1	2.9	23.6	26.5	13.6	127.0	140.6	13.6	126.9	140.5	7.4	1.7	2.3
HYB_gdust=0.1	0.50	1.43	1.93	8.4	111.9	120.4	2.3	27.0	29.3	10.7	138.9	149.7	10.8	138.9	149.6	8.5	1.9	2.4
HYB_DL	0.41	1.26	1.68	2.8	109.7	112.6	1.4	25.7	27.1	4.2	135.4	139.6	4.2	135.3	139.5	17.8	1.7	2.2
DBCLL_noAlk	0.06	0.00	0.06	2.5	0.0	2.6	1.6	0.1	1.7	4.1	0.1	4.2	4.2	0.1	4.3	2.6	2.8	2.6
DBCLL_duAlk	0.96	0.47	1.44	26.5	11.0	37.5	8.3	8.2	16.5	34.8	19.3	54.0	34.7	19.2	53.9	5.1	4.5	4.9
DBCLL_du-ssAlk	0.28	0.78	1.07	18.2	27.5	45.7	4.0	12.6	16.5	22.2	40.1	62.3	22.3	40.0	62.2	2.3	3.6	3.1
DBCLL_ClaqAlk	0.41	0.63	1.04	17.4	27.8	45.2	3.8	12.9	16.7	21.2	40.7	61.9	21.3	40.6	61.9	3.5	2.8	3.1
AeroCom	-	-	0.63	-	-	45.9	-	-	20.7	-	-	66.6	-	-	60.6	-	-	5.0
STD AeroCom	-	-	±0.56	-	-	±30.7	-	-	±19.5	-	-	±50.2	-	-	±45.9	-	-	±2.3
GMI	-	-	0.97	-	-	43.3	-	-	14.8	-	-	58.1	-	-	59.3	-	-	6.0
EMAC 2008 ¹	-	-	0.67	-	-	0.0	-	-	46.3	-	-	46.3	-	-	-	-	-	-
EMAC 2005-2008 ²	-	-	0.44	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-
IFS ³	0.47	0.35	0.82	35.9	39.6	75.5	4.6	23.8	28.4	40.5	63.4	103.9	40.5	63.4	103.6	4.2	2.1	3.0
LMDz-INCA	0.22	0.58	0.80	-	-	56.1	-	-	7.4	-	-	63.5	14.1	49.5	63.6	-	-	4.6
MetOffice UM ⁴	0.49	0.40	0.89	-	-	63.3	-	-	39.4	-	-	102.8	27.9	73.5	101.4	6.2	2.0	3.2

1. From Bian et al. (2017).

2. From Karydis et al. (2016).

3. Fine NO_3^- reported from neutralization of nitric acid, ammonia and sulfate. Coarse NO_3^- from heterogeneous chemistry. (Rémy et al., 2022).

4. Jones et al. (2021) performs two sensitivity tests to the accommodation coefficient used for NO_3^- formation in the fine mode: *FAST* with 0.193 and *SLOW* with 0.001. Here, results from the *FAST* test are reported on the basis that they present similar fine NO_3^- formation rates to our average fine NO_3^- results.

systematically estimating the lowest lifetimes. Conversely, models introducing reversible partitioning provide longer lifetimes (2.6 to 4.9 days). To better understand the possible reasons for such a wide range of estimates, in the next lines we analyze some of our sensitivity runs compared to similar systems in the literature.

870 The assumption of irreversibility of $\text{HNO}_{3(g)}$ condensation onto dust and SS particles (HYB_du-ssUPTK run) produces the largest increase in coarse NO_3^- formation, exceeding values reported in the literature. For instance, the HYB_du-ssUPTK model reports a total NO_3^- burden of 1.75 Tg, which is higher than the AeroCom range (0.63 ± 0.56 Tg) and values from models such as EMAC in AeroCom (Bian et al., 2017), as well as findings from Bian et al. (2017), Jones et al. (2021), Rémy et al. (2022) and Hauglustaine et al. (2014) (0.67, 0.89, 0.82, and 0.80 Tg, respectively). Interestingly, the results surpass
875 those from three AeroCom models (EMEP, INCA, and GMI) using a similar HYB approach with UPTK reactions on dust and SS, which range from 0.26 to 0.95 Tg of total NO_3^- (see Bian et al. (2017)). Multiple reasons could explain the excessive NO_3^- formation found in this comparison. For instance, Hauglustaine et al. (2014) employs a similar HYB implementation and UPTK coefficients for dust as the HYB_du-ssUPTK run, although a different parametrization for the $\text{HNO}_{3(g)}$ UPTK on SS. Despite Hauglustaine et al. (2014) also overestimating particulate NO_3^- with a global normalized mean bias of +68%, the
880 HYB_du-ssUPTK simulation still reports higher burdens for NO_3^- (0.80 vs. 1.75 Tg) and lower burdens for $\text{HNO}_{3(g)}$ (1.35 vs. 0.69 Tg, Table S9). Our study also presents lower NO_x emissions than Hauglustaine et al. (2014), which would imply a reduced availability of precursor gases that would typically result in lower NO_3^- formation (40.8 vs. 46.TgN, Table S1 in the Supplement). Additionally, the run neglecting the uptake on SS (HYB_duUPTK) reports values closer to observations (Section 3.2) and to the budgets in the literature.

885 This suggests that the high coefficients employed for $\text{HNO}_{3(g)}$ UPTK on SS may be excessively efficient. This conclusion is consistent with analyses discussed in previous Sections 3.1.2, 3.2 and 3.3. By comparison, other systems (Hauglustaine et al., 2014; Jones et al., 2021; Rémy et al., 2022) typically employ much lower UPTK coefficients for SS compared to our implementation.

Therefore, a function for the uptake coefficient on SS dependent on relative humidity and on SS particle size should be
890 considered in future works, following the same approach as for dust (Eq. 2). However, such a function has not clearly been determined in the literature. While Hauglustaine et al. (2014) assumes a BET isotherm for the $\text{HNO}_{3(g)}$ UPTK on SS similar to the UPTK on dust (Fairlie et al., 2010), it remains unclear if this approach is suitable for SS as it is for dust. To our best knowledge, the most appropriate alternative to the assumption of averaged UPTK coefficients for SS would be to fit a function to the experimental values reported by Liu et al. (2007), which provides uptake coefficient dependencies with relative humidity and
895 particle size. However, several discrepancies between Liu et al. (2007) and previous studies still remain unresolved (Tolocka et al., 2004; Saul et al., 2006), presenting difficulties in reaching a common agreement on a $\text{HNO}_{3(g)}$ UPTK coefficient function on SS.

Beyond the excessive NO_3^- formation driven by the UPTK on SS, the sole implementation of $\text{HNO}_{3(g)}$ UPTK on dust
900 (HYB_duUPTK run) still slightly exceeds particulate NO_3^- burdens and deposition rates reported by the references (Table 5).

This is also observed when compared to observational surface concentrations (see Section 3.2), suggesting that the sole UPTK on dust results in excessive NO_3^- formation. Given that previous studies have consistently shown non negligible $\text{HNO}_{3(\text{g})}$ uptake on SS (Myhre et al., 2006; Athanasopoulou et al., 2008), these biases could be explained by 1) an excessive UPTK coefficient employed for $\text{HNO}_{3(\text{g})}$ UPTK on dust, 2) the inherent inappropriateness of the irreversible condensation assumption of gas species on particles to simulate particulate NO_3^- formation, or 3) the different atmospheric lifetime of coarse particles among different systems. Regarding possible issues in the UPTK parameterization on dust adopted in our study, the inclusion of the scaling factor for alkalinity added to the dust uptake coefficients (Sc in Eq. 2) could contribute to excessively high UPTK rates. This alkalinity scaling factor is based on the NVC fractions provided by Fairlie et al. (2010), namely calcium and magnesium (see Section 2.2.1). Since our study includes additional NVC (i.e. potassium and sodium), this estimation might not be representative and could result in excessive condensation. However, additional information on the NVC content in Fairlie et al. (2010) is missing to refine our approach. Concerning the possible inappropriateness of the irreversible condensation assumption, indeed models implementing such an approach with similar UPTK coefficients tend to overestimate particulate NO_3^- formation (Hauglustaine et al., 2014; Jones et al., 2021; Rémy et al., 2022). This might point at the assumption of reversible evaporation-condensation of gas species or at the inclusion of alkalinity consumption as necessary implementations to improve the results.

Lastly, a longer lifetime of coarse particles such as dust or SS could contribute to enhance NO_3^- production regardless of the adopted mechanism. Spada et al. (2013) intercompared different SS emission schemes in the MONARCH model and reported lifetimes ranging from 4 to 12 days including results from literature. These findings highlight significant differences in emission, transport, and sedimentation schemes among global models, which are not negligible. Such differences could explain part of the variations identified in our analysis. For instance, if a model assumes a longer lifetime for coarse particles, it would result in a prolonged period during which $\text{HNO}_{3(\text{g})}$ can condense onto these particles, thus increasing the overall burden of particulate NO_3^- . Understanding and standardizing these lifetimes could be crucial for achieving more consistent and accurate predictions of NO_3^- aerosol formation across different modelling frameworks.

Linking with the limitations of irreversible approaches, we finally analyze the DBCLL runs to illustrate the paramount role of alkalinity in NO_3^- formation. Neglecting NVC in the partitioning (DBCLL_noAlk) results in negligible NO_3^- formation, with a burden of just 0.06 Tg in the fine mode (Table 5). In contrast, only considering dust alkalinity (DBCLL_duAlk) significantly increases the burden of both fine and coarse modes to 0.96 and 0.47 Tg, respectively, which exceeds estimates reported in other works (Hauglustaine et al., 2014; Jones et al., 2021; Rémy et al., 2022).

Additionally, accounting for SS alkalinity (reaction R12 from Table 1), the DBCLL_du-ssAlk run shifts particulate NO_3^- formation from fine towards the coarse mode, increasing coarse burdens from 0.47 in DBCLL_duAlk to 0.78 Tg. Consequently, this leads to higher total NO_3^- deposition rates, rising from 54 to 62 Tg/y. Including both dust and SS alkalinity results in a notable agreement for total NO_3^- (1.07 Tg) and $\text{HNO}_{3(\text{g})}$ (2.53 Tg) burdens with values reported in the literature. This improvement highlights the importance of considering both dust and SS alkalinity in accurately modelling NO_3^- formation and deposition rates in atmospheric chemistry simulations.

Some differences emerge when comparing the NO_3^- size distribution against references that report fine and coarse NO_3^- budgets. Specifically, the fine NO_3^- burden tends to be on the lower end of the references range (0.28 Tg vs. 0.39 Tg on average), while the coarse NO_3^- burden is on the higher end (0.78 Tg vs. 0.54 Tg on average), although the relatively high variability between the references in both modes should be considered. Despite these variations, accounting for alkalinity significantly improves the agreement of NO_3^- lifetime (3.1 days) with AeroCom (5.0 ± 2.3 days), Rémy et al. (2022) (3.0 days) and Jones et al. (2021) (3.2 days).

To delve into the differences observed between simulations employing constant dust alkalinity derived from Journet et al. (2014) (DBCLL_du-ssAlk) and from Claquin et al. (1999) (DBCLL_ClaqAlk, discussed in Supplementary Section S5) as shown in Table 5, it becomes evident that DBCLL_ClaqAlk results in a noticeable shift towards fine NO_3^- formation (from 0.28 Tg to 0.40 Tg), with fine NO_3^- increasing its lifetime (2.3 to 3.5 days) and coarse NO_3^- decreasing it (3.6 to 2.8 days). Also, we observe a slightly higher burden of $\text{HNO}_{3(g)}$ (from 2.53 Tg to 2.58 Tg) and a lower burden of $\text{NH}_{3(g)}$ (from 0.35 Tg to 0.31 Tg) (see Supplementary Tables S9 and S10). This can be attributed to the higher fraction of Ca^{2+} NVC in the fine mode derived from Journet et al. (2014) (5.73%) compared to Claquin et al. (1999) (3.68%). Both datasets present more similar fractions of Ca^{2+} for the coarse mode (4.61% vs. 3.74%) (see Supplementary Tables S4 and S6). This underscores the important role that Ca^{2+} NVC plays in heterogeneous chemistry. As a matter of fact, Ca^{2+} fractions surpasses the fractions of other NVC in the Claquin et al. (1999) dataset (0.07% Na, 0.78% K and 0.43% Mg, Tables S3 to S6). The comparison between DBCLL_du-ssAlk and DBCLL_ClaqAlk provides a quantification of the impact that mineral distributions among dust sources (especially calcite) has on gas-aerosol partitioning, specially in the formation of particulate NO_3^- . It also highlights the importance of advancing on the representation of mineralogical dust composition at the global scale and its regional variability.

Finally, we compare the outcomes from the DBCLL_du-ssAlk simulation with those reported by the EMAC model, which employs a similar approach to the DBCLL mechanism and also accounts for dust and SS alkalinity (see Section 3.1.3). The analysis is performed against two studies where EMAC budgets were reported: the AeroCom intercomparison nitrate experiment (Bian et al., 2017) and Karydis et al. (2016) (Table 5). It's notable that the total NO_3^- burden (1.07 Tg) in DBCLL_du-ssAlk exceeds both the EMAC results from AeroCom (0.67 Tg) and Karydis et al. (2016) (0.44 Tg), although our findings show reasonable agreement with observations, especially over Asia (see Section 3.2). Regarding the total budget of $\text{HNO}_{3(g)}$ (2.53 Tg), despite our result closely matching the AeroCom experiment average (2.50 ± 1.83 Tg, Table S9), it is below the values reported by EMAC in this experiment (3.10 Tg) while significantly exceeding results from Karydis et al. (2016) (1.65 Tg). These outcomes reveal intrinsic differences in the rates of $\text{HNO}_{3(g)}$ condensation on dust and SS between both models. Conversely, the $\text{NH}_{3(g)}$ burden (0.35 Tg) in our simulation falls below those reported by both studies (0.85 and 0.82 Tg, respectively), potentially due to the high formation of particulate NO_3^- over dust and SS NVC in our study. Nevertheless, the resulting burdens for NH_4^+ align with both references (0.22 vs. 0.19 and 0.17 Tg for EMAC in AeroCom and Karydis et al. (2016), respectively). Various factors could explain the differences observed between EMAC and the DBCLL_du-ssAlk simulation: the different simulated periods between our study, Karydis et al. (2016) and AeroCom (2018 vs. 2005-2008 average and 2008, respectively), the varying NVC content assumed for dust and SS between DBCLL_du-ssAlk and EMAC (Section

3.1.3), and the intrinsic differences in the heterogeneous chemistry mechanisms employed in both models (DIFFLIM and DB-CLL in MONARCH and bulk TEQ followed by DIFFLIM used in EMAC). Despite their conceptual equivalence, our results demonstrate that these mechanisms may yield different outcomes.

975 4 Conclusions

In this study, we conducted a comprehensive exploration of the processes driving nitrate formation on fine and coarse particles at a global scale using the MONARCH atmospheric model. Our sensitivity simulations incorporated state-of-the-art dust heterogeneous chemistry mechanisms, including reversible condensation-evaporation through thermodynamic equilibrium (TEQ) and irreversible uptake reactions (UPTK) between gas and aerosol phases. Three mechanisms for particulate nitrate formation
980 were implemented: (1) fTEQ, which considers fine nitrate formation through reversible TEQ reactions and neglects coarse nitrate; (2) HYB, where fine nitrate forms via TEQ with a subsequent irreversible uptake of $\text{HNO}_{3(\text{g})}$ on coarse particles; and (3) DBCLL, which allows for both fine and coarse nitrate formation through reversible TEQ processes with kinetic gas limitation.

Key assumptions such as uptake coefficients, reversible partitioning, and the inclusion of dust and sea-salt alkalinity were thoroughly assessed. Global average dust alkalinity was sourced from previous dust mineralogy simulations (Gonçalves Ageitos
985 et al., 2023) using Journet et al. (2014) and Claquin et al. (1999) mineral atlases, while globally homogeneous sea-salt alkalinity values are derived from Seinfeld and Pandis (2006). We evaluated annual cycle surface concentrations against observations for various atmospheric species and compared global spatial distributions and budgets with existing literature. Additionally, we investigated the partitioning of nitrogen species and assessed the impact of dust and sea-salt heterogeneous chemistry on total nitrogen budgets.

Neglecting coarse nitrate formation (fTEQ_du-ssAlk run) results in budgets and distributions of fine nitrate that closely
990 match observational surface concentrations of total nitrate ($-0.09 \mu\text{g m}^{-3}$ bias, 0.88 correlation). This good agreement suggests that observational stations are dominated by fine nitrate concentrations or that fine nitrate formation is overestimated in this run. Since this mechanism aligns with references forming both fine and coarse nitrate (Bian et al., 2017; Rémy et al., 2019; Hauglustaine et al., 2014), it likely compensates for the lack of coarse nitrate formation through the instantaneous TEQ of gas
995 species in the fine mode, thus not accurately reflecting the size distribution of particles.

The formation of coarse nitrate through the irreversible uptake of $\text{HNO}_{3(\text{g})}$ on coarse particles (HYB methodologies) is highly sensitive to whether the $\text{HNO}_{3(\text{g})}$ uptake is assumed to occur solely on dust or on both dust and sea-salt particles. The uptake solely on dust (HYB_duUPTK run) closely reproduces fine and total particulate nitrate seasonality (correlations of 0.8 and 0.9, respectively), although it slightly overestimates observations ($+0.29 \mu\text{g m}^{-3}$ bias), particularly over Asia, and nitrate
1000 budgets compared to the literature (1.15 vs. 0.74 Tg on average). An excessively efficient $\text{HNO}_{3(\text{g})}$ uptake on dust particles could explain such a performance, which could be improved through more accurate scaling factors adopted for alkalinity. However, the lack of information in the literature poses a challenge in implementing such adjustments.

The introduction of the $\text{HNO}_{3(g)}$ uptake on sea-salt (HYB_du-ssUPTK run) further increases the formation of coarse nitrate, exceeding the reported burdens in the literature (1.75 vs. 0.74 Tg from the references' average) and overestimating observational surface concentrations ($+1.50 \mu\text{g m}^{-3}$ bias). Moreover, some deviations with respect to references are observed in the latitude and magnitude of the transatlantic transport of coarse nitrate formed during long-range transport of dust and sea-salt. These findings indicate the need for a revision of the $\text{HNO}_{3(g)}$ uptake coefficients on sea-salt. A potential alternative could involve its implementation as a function of relative humidity and particle size, aligning it with experimental data from Liu et al. (2007), although discrepancies with earlier experimental studies present difficulties in determining consistent uptake coefficients for sea-salt.

The reversible condensation-evaporation of gas species on both fine and coarse modes of dust and sea-salt (DBCLL mechanism) highlights the paramount importance of accounting for alkalinity to derive consistent results both globally and across continents. Remarkably, the DBCLL_du-ssAlk run, which accounts for dust and sea-salt alkalinity, effectively captures monthly global concentrations of fine and total nitrate, with correlations of 0.82 and 0.78 and concentrations slightly underestimated with -0.18 and $-0.10 \mu\text{g m}^{-3}$ bias, respectively. These good results extend to particulate ammonium for both fine and total fractions (-0.02 and $0.08 \mu\text{g m}^{-3}$ bias, respectively). The consistent excessive formation of total $\text{PM}_{2.5}$ ($+20 \mu\text{g m}^{-3}$ bias) over Asia across the experiments may be explained by the absence of anthropogenic coarse dust emissions in the employed CAMS inventory. This omission leads to the overestimation of fine particulate matter formation, compensating for the misrepresentation of coarse particulate matter. Including both dust and sea-salt alkalinity in the DBCLL mechanism significantly raises global pH levels, which enhances particulate nitrate formation by 94% in comparison to the same mechanism with no alkalinity. This incorporation also reduces NO_x budgets and lowers the reduced nitrogen burden due to an increase of ammonium formation. Moreover, the sole inclusion of sea-salt alkalinity is identified as responsible of substantially shifting the size partitioning of particulate nitrate from fine (-72%) to coarse ($+63\%$), along with the rise in wet deposition of particulate nitrate due to high scavenging rate of sea-salt particles, overall reducing the total aerosol nitrate atmospheric lifetime from 4.9 to 3.1 days.

Comparison with references reveals that while surface concentrations in the DBCLL_du-ssAlk simulation align well with observations, its global burden (1.07 Tg) sits at the upper limit of AeroCom's reported range (0.63 ± 0.56 Tg) and slightly exceeds the average reported by other studies (0.74 Tg). Fundamental differences between the models, like the alkalinity factors, intrinsic differences in the heterogeneous chemistry, aerosol representation and transport processes can partly explain the wide range of results reported in the literature. Overall, the DBCLL_du-ssAlk scheme demonstrates the best accuracy compared to the other tested configurations, as evidenced by its closer alignment with observations.

A comparison between runs adopting the two different dust alkalinity fractions derived from averaging simulations using Journet et al. (2014) and Claquin et al. (1999) soil mineralogy datasets as described in Gonçalves Ageitos et al. (2023) reveals that assumptions on the dust composition are crucial, particularly influencing particulate nitrate size distribution. A decrease from 5.17% to 3.68% in ionic Ca^{2+} fractions, along with slight increases in Na, K and Mg between Journet et al. (2014) and Claquin et al. (1999) leads to a +35% increase in fine and -21% decrease in coarse nitrate formation. This high sensitivity is

especially attributed to the lower calcite content and its different size distribution in the Claquin et al. (1999) database compared to Journet et al. (2014).

1040 It is important to note that our computational cost analysis reveals highly similar processing times across the sensitivity runs, with a standard variation of only 5%. This is within the estimated variability of the supercomputing resources utilized for the present work. Consequently, the computational cost does not indicate a clear advantage in efficiency for any of the methodologies assessed.

1045 This study establishes a crucial benchmark for future investigations into the impact of incorporating regional variations in dust alkalinity, as derived from Claquin et al. (1999), Journet et al. (2014), and the upcoming spectroscopically-based EMIT surface mineralogical dataset (Green et al., 2018; Thompson et al., 2024; Brodrick et al., 2023), on the formation of particulate nitrate and atmospheric composition. The findings also evaluate the significance of dust and sea-salt alkalinity in inorganic aerosol heterogeneous chemistry, providing insights into the optimal representation of dust alkalinity in atmospheric models. Ultimately, these results aim to enhance the capability of atmospheric and climate models to simulate the formation of aerosol nitrate, ammonium, and sulfate, potentially improving our ability to estimate the radiative effects of these species in climate
1050 projections.

Code availability. The MONARCH code is available at <https://earth.bsc.es/gitlab/es/monarch> (last access: 21 July 2024) and the HERMESv3_GR code is accessible at https://earth.bsc.es/gitlab/es/hermesv3_gr (last access: 21 July 2024).

Data availability. The GHOST dataset is made freely available via the following repository: <https://doi.org/10.5281/zenodo.10637449> (Bowdalo et al., 2024). The model output used in this work is available in the Zenodo data repository at <https://doi.org/10.5281/zenodo.12789730> (Sousse, 2024).
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Author contributions. RS, OJ and CPG-P developed the model, designed the methodology and the conceptualization, and performed the investigation and analysis of the results. RS designed and conducted the sensitivity simulations, and the results' postprocessing, validation and visualization, assisted by OJ and CP. MGA provided the simulations with the average mineralogy from the utilized mineral dust datasets. DB provided the observational evaluation data under request, developed the GHOST dataset and the evaluation software. MGv provided the
1060 CAMS-ANTv4.2 emission dataset under request and developed the HERMESv3_GR emission model. RS wrote the manuscript, which was re-edited by OJ and CPG-P, with contributions from all other co-authors.

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