

# Mount Pinatubo's effect on the moisture-based drivers of plant productivity

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## Abstract

Large volcanic eruptions can significantly affect the state of the climate, including stratospheric sulfate concentrations, surface and top-of-atmosphere radiative fluxes, stratospheric and surface temperature, and regional hydroclimate. The prevalence of higher natural variability in the regional rainfall response to volcanic-induced climate perturbations creates a gap in our understanding of how eruptions affect ecohydrological conditions and plant productivity. Here, we will explore the understudied store (soil moisture) and flux (evapotranspiration) of water as the short-term ecohydrological control over plant productivity in response to the 1991 eruption of Mt. Pinatubo. We used the NASA's Earth system model for modeling of the Mt. Pinatubo eruption and detect the ensuing hydroclimate responses. The model simulates a mean surface cooling of  $\sim 0.5$  °C following the Mt. Pinatubo eruption. The rainfall response is spatially heterogeneous with large temporal variability, yet still shows suppressed rainfall in the northern hemisphere after the eruption. We find that up to 10-15% of land regions show a statistically significant hydroclimate response (wet and dry) as calculated by the Soil Moisture Deficit Index (SMDI) and Evapotranspiration Deficit Index (ETDI). Results confirm that these impact metrics successfully present a more robust understanding of plant productivity. Our results also explain the

geographical dependence of various contributing factors to the compound response and their implications for exploring the climate impacts of such episodic forcings.

## **Introduction**

Volcanic eruptions are the most prominent source of sulfate aerosols in the stratosphere and are natural drivers of climate variability. Volcanically injected sulfate aerosols in the stratosphere alter the Earth's radiative balance by simultaneously reflecting incoming solar radiation and absorbing outgoing longwave radiation emitted from the Earth's surface (Robock, 2000). The presence of sulfate aerosol for months to years after an eruption and its microphysical transformation in the stratosphere affect the climate system through numerous direct and indirect effects (Barnes and Hofmann, 1997; Briffa et al., 1998; Deshler et al., 2003; Kremser et al., 2016; LeGrande et al., 2016; Marshall et al., 2022; Tejedor et al., 2021; Timmreck, 2012; Toohey et al., 2019). The Mt. Pinatubo eruption (June 1991) remains the largest eruption in the satellite era, and it has been explicitly documented and analyzed for its radiative and climate impacts. Numerous studies based on satellite observations and supported through different modeling efforts estimated that Mt. Pinatubo injected 10-20 Tg of SO<sub>2</sub> at a range of 18-25 km of plume height (Aquila et al., 2012; Bluth et al., 1992; Dhomse et al., 2014; Gao et al., 2023; Sheng et al., 2015a, b; Stenchikov et al., 1998). The radiative impacts of Mt. Pinatubo's eruption estimate an aerosol optical depth of 0.15 for 550 nm wavelength, with an effective radius in the range of 0.16 to 1 micrometer (μm) and net radiative forcing on the order of 5-6 Wm<sup>-2</sup> (Lacis, 2015; Lacis et al., 1992; Sato et al., 1993; Stenchikov et al., 1998). Estimates of the induced surface cooling range between 0.3-0.7 °C; lower stratosphere warming estimates are in the range of 2-3 °C (Dutton and Christy, 1992; Hansen et al., 1992; Labitzke and McCormick, 1992; Lacis et al., 1992; Minnis et al., 1993; Ramachandran et al., 2000; Stenchikov et al., 1998).

In this study, we aim to explore the mechanisms by which the Mt. Pinatubo eruption affected the hydroclimatic conditions and water-based drivers of plant productivity. Agricultural productivity is sensitive to changes in temperature and precipitation (Lobell and Field, 2007; Olesen and Bindi, 2002; Rosenzweig and Parry, 1994). Although much work has been devoted to understanding Mt. Pinatubo's impacts on plant productivity, the literature has been dominated by studies focusing on impacts from changes to the quantity and quality of incoming solar radiation (Farquhar and Roderick, 2003; Gu et al., 2002, 2003; Jones and Cox, 2001; Robock, 2005). Proctor et al. (2018) have estimated a decrease in C4 (maize) and C3 (soy, rice, and wheat) agricultural crop production in response to Mt. Pinatubo driven mainly by changes in incoming radiation. Krakauer and Randerson (2003) evaluated the role of surface cooling in reduced net primary productivity (NPP) using tree ring growth patterns following multiple Mt. Pinatubo-sized eruptions in the last millennium record. Reduced NPP was found in northern mid to high latitudes, while the signal in the lower latitudes and tropics was either insignificant or constrained by other factors. Other studies have further expanded into societal impact research focusing on volcanically induced changes to harvest and agricultural productivity over different regions (van Dijk et al., 2023; Hao et al., 2020; Huhtamaa and Helama, 2017; Manning et al., 2017; Singh et al., 2023; Toohey et al., 2016).

Similarly, a plenty of work has been devoted to understanding the hydroclimate response to Mt. Pinatubo through changes in atmospheric precipitation (Barnes et al., 2016; Paik et al., 2020; Trenberth and Dai, 2007), as the monsoon seasonal rainfall decreases in the season following the eruption (Colose et al., 2016; Iles et al., 2013; Liu et al., 2016; Singh et al., 2023; Tejedor et al., 2021). It is also shown that volcanic eruptions can alter regional rainfall and hydroclimate in general, which could prominently affect regional plant productivity (Zuo et al.,

2019a, b). However, rainfall alone provides an incomplete understanding of the drought conditions relevant to plant productivity; a rainfall deficit could, in principle, be overcome by moisture stored in the soil. Hence, meteorological drought indices (e.g. SPI (McKee et al., 1993) or PDSI (Palmer 1965)) ignore a full water balance approach. Furthermore, meteorological drought indices tend to be designed to evaluate prolonged periods of abnormally dry weather conditions. For instance, PDSI is an indicator of drought with a 9-month horizon (Mullapudi et al., 2023). Yet, agricultural crops are heterogeneously sensitive to the timing and degree of moisture deficits during particular portions of the crop growth cycle (Hane and Pumphrey, 1984). Thus, consideration of indices with high temporal frequency can be critical when focusing on agriculture.

Soil moisture is the stock of water stored underground and is a primary source of water flux to the atmosphere and plants through evapotranspiration. Evaporation of water from bare surface soil or transpiration of water during photosynthesis in plants from the root zone soils uses a large portion of absorbed solar energy (Trenberth et al., 2009). Plant transpiration is the largest contributor to land evapotranspiration (Dirmeyer et al., 2006; Lawrence et al., 2007; Nilson and Assmann, 2007; Seneviratne et al., 2010). Soil moisture decrease in the root zone establishes an essential control over plant productivity as transpiration is an integral component of photosynthesis (Chen and Coughenour, 2004; Denissen et al., 2022). Multiple studies have established that water supply is the limiting factor for climatic evapotranspiration over tropical and subtropical land areas, while temperature is an important controlling factor in northern mid- and high latitudes (Dong and Dai, 2017; Mintz and Walker, 1993; Nemani et al., 2003; Pan et al., 2015). However, soil moisture changes in response to the Mt. Pinatubo eruption (1991) are

underreported in the literature, and it is unclear how soil moisture would respond given volcanically forced changes in primary drivers (temperature and precipitation).

Studying a large (10xPinatubo) volcanic eruption, (Frölicher et al., 2011) have shown that the terrestrial carbon pool is sensitive to the regional (in the tropics and sub-tropics) soil-moisture content through the net-ecosystem productivity. Using the geoengineering large ensemble simulations with the CESM model, (Cheng et al., 2019) have analysed the changes in terrestrial hydrological cycle and discussed the future soil-moisture response and its drivers under a geoengineering scenario. To our knowledge, no study has yet investigated multiple indicators of water use with agricultural productivity after a short-duration event like Mt. Pinatubo eruption. Hence, this study looks to explicitly investigate changes in agricultural drought indices from Mt. Pinatubo by considering the store (soil moisture) and flux (evapotranspiration) of water as potential short-term controls over productivity in particular regions. We use NASA's state-of-the-art Earth system model, which incorporates prognostically evolving aerosols, to conduct simulation experiments following the counterfactual inference of causation approach for the Mt. Pinatubo eruption. We assess the impact of the Mt. Pinatubo eruption on the model-simulated climate via multiple pathways, ranging from primary dependent variables to higher-order responses that influence plant productivity. The use of prognostic aerosols enhances the simulations by capturing dynamically consistent feedbacks between the climate response and volcanic aerosols, including aerosol-radiation interactions and stratosphere-troposphere energy flux exchanges (McGraw and Polvani, 2024). Considering the complexity of modeling the terrestrial system, vegetation demographics, and physiological characteristics, we use the soil moisture and evapotranspiration-based agricultural drought indices SMDI (soil moisture deficit index) and ETDI (evapotranspiration deficit index) for short-term and long-term

scale to account for agricultural productivity. By focusing on soil moisture and evapotranspiration metrics, the major water-based drivers of plant productivity are explored to deepen our understanding of the Mt. Pinatubo impacts on plant productivity.

## **2.0 Method, Experiment, and Data**

**2.1 NASA GISS ModelE2.1 (MATRIX):** We use the state-of-the-art Earth system model from the NASA (National Aeronautics and Space Administration) Goddard Institute for Space Studies, NASA GISS ModelE2.1 (Bauer et al., 2020; Kelley et al., 2020). NASA GISS ModelE2.1 has an atmospheric horizontal latitude-longitude grid spacing of 2.0x2.5 degrees (at the equator) with 40 vertical levels and a model top at 0.1 hPa. We used the interactive chemistry version MATRIX (Multiconfiguration Aerosol TRacker of mIXing state) aerosol microphysics module (Bauer et al., 2008, 2020), which is based on the Quadrature Method of Moment (QMOM) to predict aerosol particle number, mass, and size distribution for 16 different mixed modes of the aerosol population. New particle formation is represented by Vehkamäki et al. (2002), along with aerosol-phase chemistry, condensational growth, coagulation, and mixing states (Bauer et al., 2013). 16 mixing states with 51 aerosol tracers for sulfate, nitrate, ammonium, aerosol water, black carbon, organic carbon, sea salt, and mineral dust are resolved in this microphysical module (Bauer et al., 2008, 2020). The first indirect effect of aerosols in terms of changes in cloud properties through nucleation is also computed within MATRIX.

The model's ocean component (GISS Ocean v1) has a horizontal resolution of 1x1.25 degrees, with 40 vertical layers. The land component is the Ent Terrestrial Biosphere Model (TBM) (Kim et al., 2015; Kiang 2012) which includes an interactive carbon cycle (Ito et al., 2020), satellite-derived (MODIS-Moderate Resolution Imaging Spectroradiometer) plant functional type, and monthly variation of leaf area index (Gao et al., 2008; Myneni et al., 2002),

and tree height (Simard et al., 2011). Interannual variations in the vegetation properties are controlled by rescaling the vegetation fractions (Figure S8) using historical crops and pasture at the grid scale to account for land cover and land use changes (Ito et al., 2020; Miller et al., 2021). The land model has two defined tiles for the soil layer: bare and vegetated, and each has six vertical levels to a depth of 3.5 m (11.5 feet) (Rosenzweig and Abramopoulos, 1997). Rooting depths for different plant functional types are also given by Rosenzweig and Abramopoulos (1997), and more than 60% of roots for crop plant functional types are located within 0.6 m (1.96 feet) of soil depth. In this version of the model, the agricultural grid cells, crop plant functional type, and crop calendar are prescribed according to McDermid et al., (2019). Irrigation in the GISS ModelE is implemented using the water irrigation demand data (IWD) (Wisser et al., 2010) and irrigation potential calculations based on (Wada et al., 2013) as discussed in (Cook et al., 2020).

**2.2 Experiment Design:** The MATRIX version of GISS ModelE2.1 with active tracers is three times more computationally expensive than the non-interactive (prescribed pre-calculated aerosol concentration and extinction) version. We extended an equilibrated 1400-yearlong PI control run with non-interactive tracers with an additional 500 years using the MATRIX version with prognostic tracers before starting the ‘historical’ run. MATRIX includes the tropospheric chemistry scheme that contains the inorganic (Ox NOx Hox and CO), organic chemistry of CH<sub>4</sub>, and higher hydrocarbons (Gery et al., 1989; Shindell, 2001; Shindell et al., 2003). The stratospheric chemistry includes bromine, chlorine, and polar stratospheric clouds (Shindell et al., 2006). Dust emission in the model is controlled by climate variables such as winds and soil moisture at the spatial and temporal scales (Miller et al., 2006). However, anthropogenic dust is not included in GISS ModelE2.1. Other anthropogenic emissions, including biomass burning

(pre-1997 from (van Marle et al., 2017) and 1997 onwards from the GFED4s inventory (van der Werf et al., 2017)), are taken from the Community Emission Data System (CEDS) inventory (Hoesly et al., 2018). Most importantly, the volcanic SO<sub>2</sub> forcing for the ‘historical’ run (1850-1977) is the daily emission rate from VolcanEESM (Neely and Schmidt 2016 : <https://catalogue.ceda.ac.uk/uuid/a8a7e52b299a46c9b09d8e56b283d385> ), and satellite measurement based SO<sub>2</sub> inventory (Carn et al., 2016) for 1978 to 2014 (extended up to 2021). The cumulative Mt. Pinatubo emission is 15194 kt (~15.2 Tg) of SO<sub>2</sub> injected from 13<sup>th</sup> to 16<sup>th</sup> of June 1991 above the Mt. Pinatubo vent, with a maximum of 15000 kt (15 Tg) emitted on June 15<sup>th</sup> at a plume height of 25 km Carn et al., (2016). The MATRIX version of the GISS ModelE2.1 used for all of our simulations predicts the nucleation, evolution, and removal of sulfate aerosols prognostically.

The model simulations we performed (Table 1) are described here. We started from the 1400-year-long preindustrial control run from CMIP6 (GISS-CMIP6-PI) with the prescribed average AOD historical period, which is further extended for 500 years using the GISS ModelE2.1 – MATRIX with prognostic tracers (GISS-PI). Then, the CMIP6 historical run (GISS-HIST-SO<sub>2</sub>; 1850-2014) started with all forcings as specified by CMIP6 except the daily emission rate of injection of SO<sub>2</sub> (VolcanEESM) (Carn et al., 2016; Neely III and Schmidt, 2016). We branched out the experiment ensemble with Mt. Pinatubo eruption (GISS-PIN-SO<sub>2</sub>) and the counterfactual ensemble without Mt. Pinatubo (GISS-NOPIN-SO<sub>2</sub>) from the historical (GISS-HIST-SO<sub>2</sub>) using perturbed initial conditions (1st Jan 1986) from the year 1986 to 1999. The perturbation to the initial conditions is generated by altering the radiation-related random number generator that deals with fractional cloudiness in the column.

**Table 1: Simulation experiment details.**

| EXP Name                   | Description   | Time period /run length      | # of ensembles |
|----------------------------|---------------|------------------------------|----------------|
| GISS-CMIP6-PI              | Preindustrial | 1850 climatology /1400 years | 1              |
| GISS-PI                    | Preindustrial | 1850 climatology /500 years  | 1              |
| GISS-HIST-SO <sub>2</sub>  | historical    | 1850-2014/165 years          | 1              |
| GISS-PIN-SO <sub>2</sub>   | historical    | 1986-1999/ 15 years          | 11*            |
| GISS-NOPIN-SO <sub>2</sub> | historical    | 1986-1999/ 15 years          | 11*            |

\*These ensemble members are branched out from the GISS-HIST-SO<sub>2</sub> by perturbing the initial conditions.

**2.3 Methods:** This study investigates the impacts of the Mt. Pinatubo eruption on the major drivers of primary productivity, focusing on soil moisture-related metrics and evapotranspiration. Hereafter, we use the terminology ‘PCH’ (Mt. Pinatubo and Cerro Hudson) to refer to the ‘GISS-PIN-SO<sub>2</sub>’ and ‘NP’ for the counter-factual ensemble ‘GISS-NOPIN-SO<sub>2</sub>’. We have included the Cerro Hudson eruption in both ensembles since we are focusing on the Mt. Pinatubo-driven climate response.

### 2.3.1 Statistical analysis for detecting Mt. Pinatubo-significant regions and anomalies calculations.

We treat the no-Pinatubo ensemble (NP) as a counter-factual climate simulation and utilize it to perform the paired Student’s t-test. The null hypothesis is that the ensemble means of a quantity of interest (QoI) in a region over a time period are the same between ensembles (i.e.  $H_0: \bar{\mu}_{PCH} = \bar{\mu}_{NP}$ ). In the subsequent figures in this document, gray regions indicate non-rejection of the null hypothesis at the 95% confidence level, while the colored areas highlight its rejection and show anomalies relative to 1950-2014 simulated climatology (see Supplementary information section S1.0). However, directly comparing the difference between the two ensembles (PCH and NP) is

an alternative approach to presenting the Pinatubo effect (see Supplement Figure S2). Using either approach leads to the same general conclusions, with only small quantitative differences. Nevertheless, we chose to remain consistent with the baseline requirements for other metrics as well and used the historical climatology for period 1950-2014 as the baseline for the core of our analysis. The coloring emphasizes the significant regions of anomalies. But we also emphasize the difference in calculations: the grey areas show no significant change between the PCH and NP ensembles, while the anomalies are PCH ensemble mean minus climatology.

## **2.4 Impact metrics**

The distribution of incoming and outgoing radiation influences the hydrological cycle (Kiehl and Trenberth, 1997; Trenberth and Dai, 2007). A reduction of solar radiation at the surface has the potential to reduce rainfall and change the latent heat-dominated atmospheric heating pattern (Trenberth and Stepaniak, 2004). The perturbed atmospheric conditions and surface energy budget could affect soil moisture. Along with the surface air temperature and precipitation, we use soil moisture and surface energy budget-oriented drought indices (the soil moisture deficit index (SMDI) and evapotranspiration deficit index (ETDI)) to evaluate the land-atmosphere interaction and account for the potential drivers of the crop plant productivity in the model simulated post- Mt. Pinatubo environmental conditions (Narasimhan and Srinivasan, 2005). SMDI represents the land-based soil moisture state in selected depth horizons (i.e., SMDI\_2 means Soil Moisture Deficit Index for 2 feet (0.6 m) depth). ETDI represents the atmospheric conditions governing the land-atmosphere interaction and is also an indicator of plant health. Lastly, plant transpiration is analyzed to explore the simulated physiological response to the volcanically induced hydroclimatic conditions. The Palmer drought severity index (PDSI) and other indices are commonly used to represent climatological drought conditions. However, we

focus on SMDI and ETDI because these can capture short-term developing agricultural drought conditions that impact plant productivity and are free from the limitations of other metrics like PDSI. For example, SMDI and ETDI are seasonally independent measures and are comparable across space, even for different climatic zones.

SMDI and ETDI were calculated as described in Narasimhan and Srinivasan (2005) using model output at monthly and daily scales. Daily model output is resampled to a weekly scale to compute the indices. The weekly frequency is used because it is suitable for agricultural applications, and the daily frequency is comparatively higher, which makes these indices computationally expensive. Below, we reproduce the weekly calculation of SMDI and ETDI as presented in Narasimhan and Srinivasan (2005).

**2.4.1 Soil Moisture Deficit Index (SMDI) :** The soil moisture deficit index measures the wetness/dryness of soil moisture conditions in comparison to long-term records spanning 1950-2014.

$$SD_{i,j} = \frac{SW_{i,j} - MSW_j}{MSW_j - \min SW_j} \times 100 \quad \text{if } SW_{i,j} \leq MSW_j \quad \dots (2.4.1a)$$

And

$$SD_{i,j} = \frac{SW_{i,j} - MSW_j}{\max SW_j - MSW_j} \times 100 \quad \text{if } SW_{i,j} > MSW_j \quad \dots (2.4.1b)$$

$SD_{i,j}$  is the soil water deficit (%) for week  $j$  of the year  $i$ .  $SW_{i,j}$  is the mean weekly soil water available in the soil profile (mm) for week  $j$  of the year  $i$ ,  $MSW_j$  is the long-term (calibration period) median available water in the soil profile (mm) for week  $j$ , and  $\min SW_j$  and  $\max SW_j$  are the  $j^{\text{th}}$  weekly minimum and maximum of soil water available in the soil profile across the calibration period (1950-2014). The soil moisture deficit index for any given week can be calculated as

$$\text{SMDI}_j = 0.5 * \text{SMDI}_{j-1} + \text{SD}_j/50 \quad \dots (2.4.1c)$$

SMDI can be calculated for different soil depths; we used the 2, 4, and 6-feet depths for SMDI estimation, approximately 0.6, 1.2, and 1.8 meters, respectively. For SMDI, it is typical to use feet instead of meters in the literature, which is why we use the same units along with the conventional SI units in brackets here. SMDI-4 means we considered the soil moisture content between 2 (0.6 m) to 4 feet (1.2 m) in depth. Similarly, SMDI-6 indicates the soil moisture content between 4 (1.2 m) to 6 feet (1.8 m) in depth.

**2.4.2 Evapotranspiration Deficit Index (ETDI):** The limitations of the Palmer Drought Severity Index (PDSI) (Palmer, 1965) and Crop Moisture Index (CMI) (Palmer, 1968) in the formulation used for PET (Thornthwaite, 1948) calculation and lack of accountability to the land cover type on water balance encouraged the exploration of ETDI for agricultural productivity. Also, in the climate models, surface energy fluxes are parameterized in terms of the thermodynamic gradient of atmosphere and land models and thus represent the land-atmosphere interactions, which are not accounted for by these atmosphere-only indices. We utilized model simulated surface energy fluxes (Latent and Sensible heat) to calculate the potential (PET) and actual evapotranspiration (AET) to estimate the water stress ratio. However, the applicability of the Penman-Monteith equation for reference crops (Allen et al., 1998) provides a substitute method for PET calculation, which broadly produced similar results (not shown). In Equations 2.4.2a and 2.4.2b, we used the model simulated energy fluxes to calculate AET and PET as suggested in (Milly and Dunne, 2016; Scheff and Frierson, 2015).

The energy budget equation at the surface is given by  $R_n = G + LH + SH$ , where  $R_n$  is net solar radiation,  $G$  is net ground heating (heat flux),  $LH$  and  $SH$  represent the Latent and Sensible

279 heat fluxes, respectively. We then use these to calculate PET and AET (unit as mm per day; 1  
 280  $Wm^{-2} = 0.0353$  mm/day):

$$281 \quad PET = 0.8(R_n - G) = (0.8 * 0.0353) * (LH + SH) \quad \dots\dots (2.4.2a)$$

282 And

$$283 \quad AET = LH * (0.0353). \quad \dots\dots (2.4.2b)$$

284 The evapotranspiration deficit index is estimated using the water stress condition using the actual  
 285 evapotranspiration (AET) and potential evapotranspiration (PET) per grid cell, as given below.

$$286 \quad WS = \frac{PET - AET}{PET} \quad \dots\dots (2.4.2c)$$

287 WS range, from 0 to 1, where 0 signifies that evapotranspiration is happening at the potential rate  
 288 and 1 stand for no actual evapotranspiration. WS represents the water stress ratio on a monthly or  
 289 weekly basis ( $WS_j$ ), which is further utilized to calculate water stress anomaly ( $WSA_{i,j}$ ) for week  
 290  $j$  of year  $i$  as given below.

$$291 \quad WSA_{i,j} = \frac{MWS_j - WS_{i,j}}{MWS_j - \min WS_j} \times 100 \quad \text{if } WS_{i,j} \leq MWS_j \quad \dots\dots (2.4.2d)$$

292 And

$$293 \quad WSA_{i,j} = \frac{MWS_j - WS_{i,j}}{\max WS_j - MWS_j} \times 100 \quad \text{if } WS_{i,j} > MWS_j \quad \dots\dots (2.4.2e)$$

294 Here,  $MWS_j$ ,  $\min WS_j$ , and  $\max WS_j$  represent the long-term median, minimum, and maximum of  
 295 the water stress ratio over the calibration period. Water stress anomaly ranges between -100% to  
 296 100%, indicating very dry to wet conditions over the region.

297 Finally, drought severity is calculated as ETDI, similar to SMDI (equation 2.4.1c) at a  
 298 monthly/weekly time scale.

$$299 \quad ETDI_j = 0.5 * ETDI_{j-1} + WSA_j / 50. \quad \dots\dots (2.4.2f)$$

The indices SMDI and ETDI range from -4 to +4, representing the excessive wet and dry conditions. The bounding values -4 or +4 represent extremely dry/wet conditions as the deficit/excess of soil-moisture deficit (SM) or water stress anomaly (WSA) reached relative to the maximum over the reference calibration period.

We also highlight the justification for selecting 1950-2014 as the base period for analyzing the response in climate variables and the long-term calibration period for drought indices calculations. (Supplementary information section S1).

### **3.0 Results**

The result section of this study first presents the NASA GISS model's simulated properties of the 1991 Mt. Pinatubo eruption, and then further the evaluation of the primary (aerosol optical depth, radiation, and temperature) and secondary (precipitation, soil moisture, evapotranspiration, and transpiration) impacts on plant productivity.

#### **3.1 Radiative forcing response**

In the Mt. Pinatubo eruption simulation, the volcanically injected SO<sub>2</sub> in the stratosphere oxidizes in the presence of prognostically evolving OH radicals to form the stratospheric sulfate aerosols (Supplementary section S2.0). GISS ModelE (MATRIX) PCH also simulated a peak global mean aerosol optical depth (AOD; for 550 nm wavelength) of 0.22 (Supplementary Fig S4 bottom panel) a few months after the eruption, which then decreases with time due to the deposition of volcanic aerosols (English et al., 2013; Sato et al., 1993). In this study, the model-simulated aerosol optical depth (AOD) due to volcanic aerosol and radiative forcing is larger than the previously reported AOD of 0.15 and forcing of -4.0 to -5.0 Wm<sup>-2</sup> due to the Mt. Pinatubo eruption (Hansen et al., 1992; Lacis et al., 1992). The mass and size of volcanic sulfate aerosol firmly control the scattering of the incoming shortwave radiation and the absorption of

longwave (Kinne et al., 1992; Lacis, 2015; Lacis et al., 1992; Lacis and Hansen, 1974). The first-order climate response to the volcanically-injected sulfate aerosol in the stratosphere is the perturbation of the radiative balance of the Earth system (Brown et al., 2024; Hansen et al., 1980; Lacis et al., 1992; Stenchikov et al., 1998). Supplementary figure S3 shows that the GISS ModelE PCH has simulated a peak longwave, shortwave, and net radiative response of  $+3.0 \text{ Wm}^{-2}$ ,  $-8.0 \text{ Wm}^{-2}$ , and  $-5.0 \text{ Wm}^{-2}$  respectively, a few months after the eruption, which recovers slowly in next 24 months and is consistent with previous studies (Stenchikov et al., 1998; Hansen et al., 1992; Minnis et al., 1993; Brown et al., 2024). The GISS model also simulated a smaller peak ranging within  $-1 \text{ Wm}^{-2}$  in the counterfactual (without Mt. Pinatubo) runs, likely due to the Cerro Hudson eruption in August 1991.

### **3.2 Aerosol dispersion and temperature response**

Figure 1 shows the zonal mean anomaly for the aerosol optical depth (AOD) and surface air temperature. The zonal AOD shows the dispersion and transport of aerosol poleward after the eruption. Horizontal dispersion and transport of the aerosols are strictly influenced by the stratospheric meteorology and atmospheric circulation, which is independent in each ensemble member, and depends on the plume height and season. GISS ModelE has simulated AOD consistently with previous studies (Aquila et al., 2012; Rogers et al., 1998; Timmreck et al., 1999; Trepte et al., 1993). Cross-equatorial dispersion to the southern hemisphere might be due to the more robust Brewer-Dobson circulation in the austral winter (Aquila et al., 2012). Meanwhile, the phases of QBO and local heating also play a crucial role in the poleward and vertical dispersion of stratospheric aerosols (Hitchman et al., 1994). A smaller peak in the southern hemisphere ( $45^\circ \text{ S}$ ) in the late 1991, likely due to the Cerro Hudson eruption, which injected  $\sim 1.5 \text{ Tg}$  of  $\text{SO}_2$  at a height of 15 km.

The lower panels in Figure 1 show the surface air temperature response due to the net radiative perturbation, which is dominated by the scattering of incoming solar radiation. The zonal structure of surface temperature shows that the surface cooling follows the aerosol optical depth pattern, and the most significant cooling is simulated in the northern hemisphere high latitudes. Temporal characteristics of lower stratosphere warming, and surface cooling also show the seasonal variations in incoming solar radiation in northern polar latitudes. Model simulated the lower stratospheric warming due to longwave radiation absorption in lower stratosphere consistent with previous studies (Supplementary text section S3.0).

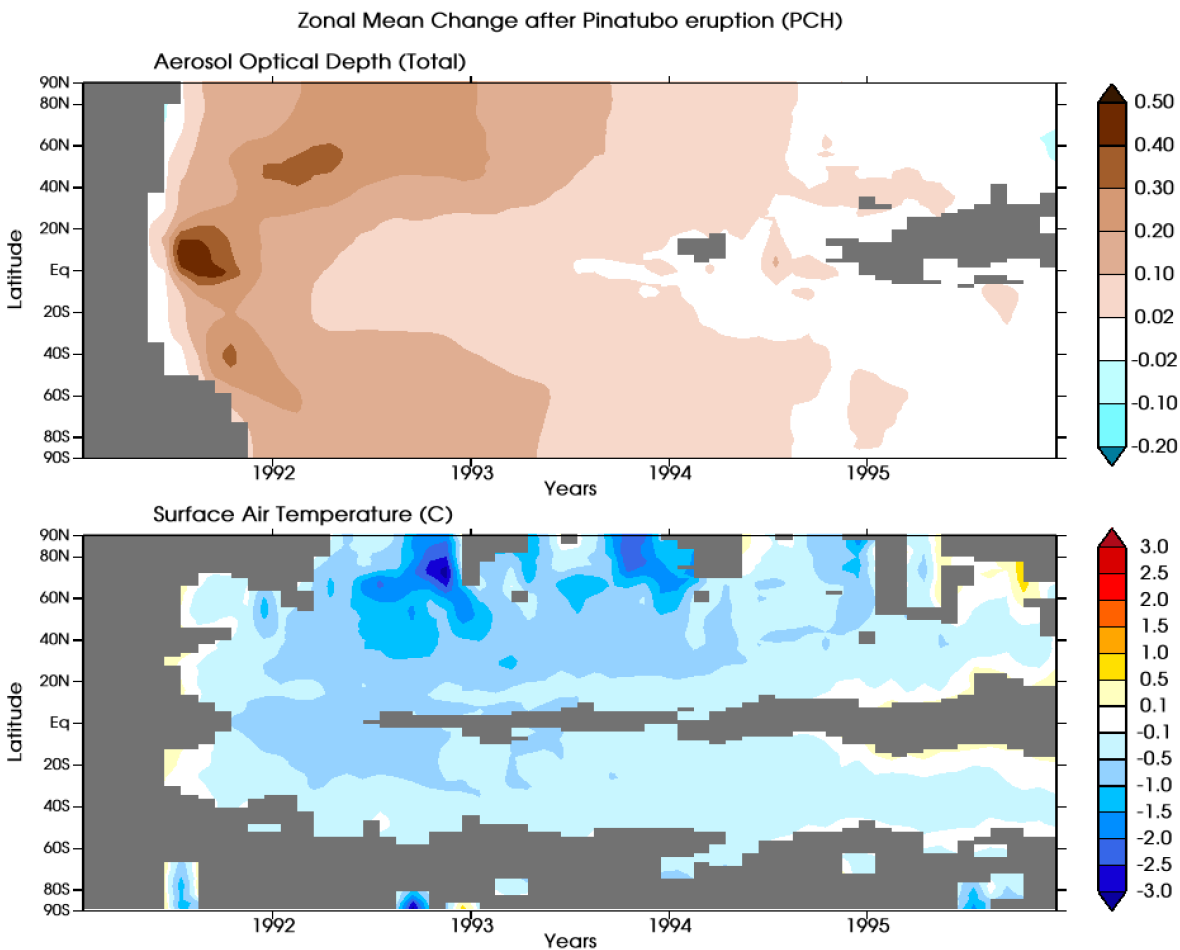


Figure 1. Zonal mean of monthly anomalies for multi-ensemble means for aerosol optical depth at 550 nm (upper panel) and surface air temperature (lower panel) with respect to the 1950-2014 climatology. Gray regions show no statistically significant difference between the PCH and NP responses. The colored areas show anomalies of PCH with respect to the climatology from 1950-2014.

The spatial pattern of surface air temperature response is evaluated at the seasonal scale for each year from 1991 to 1995, as shown in Figure 2. We conclude that the volcanic forcing from the Mt. Pinatubo eruption results in a detectable seasonal mean surface air temperature response. Figure 2 shows that a spatial pattern of surface cooling starts appearing after a few months of the eruption (during the SON season of 1991) when the gaseous  $\text{SO}_2$  is oxidized into sulfate aerosols. The surface cooling signature due to the volcanic aerosols was significant in 1992 and 1993 before recovering in 1994 towards pre-eruption temperature conditions. The highest surface cooling is noticed over the sub-tropics and higher latitude land regions in the northern hemisphere and reaches up to  $2.5^\circ\text{C}$  at a regional scale. To summarize: the PCH GISS ModelE simulated global mean peak cooling response is  $\sim 0.5^\circ\text{C}$  after the eruption, as shown in Supplement Figure S4 (Middle panel), with a range between  $0.25 - 1.0^\circ\text{C}$  for individual ensemble members, and this is consistent with the various observation and modeling studies (Dutton and Christy, 1992; Hansen et al., 1996; Kirchner et al., 1999; Minnis et al., 1993; Parker et al., 1996; Ramachandran et al., 2000; Stenchikov et al., 1998).

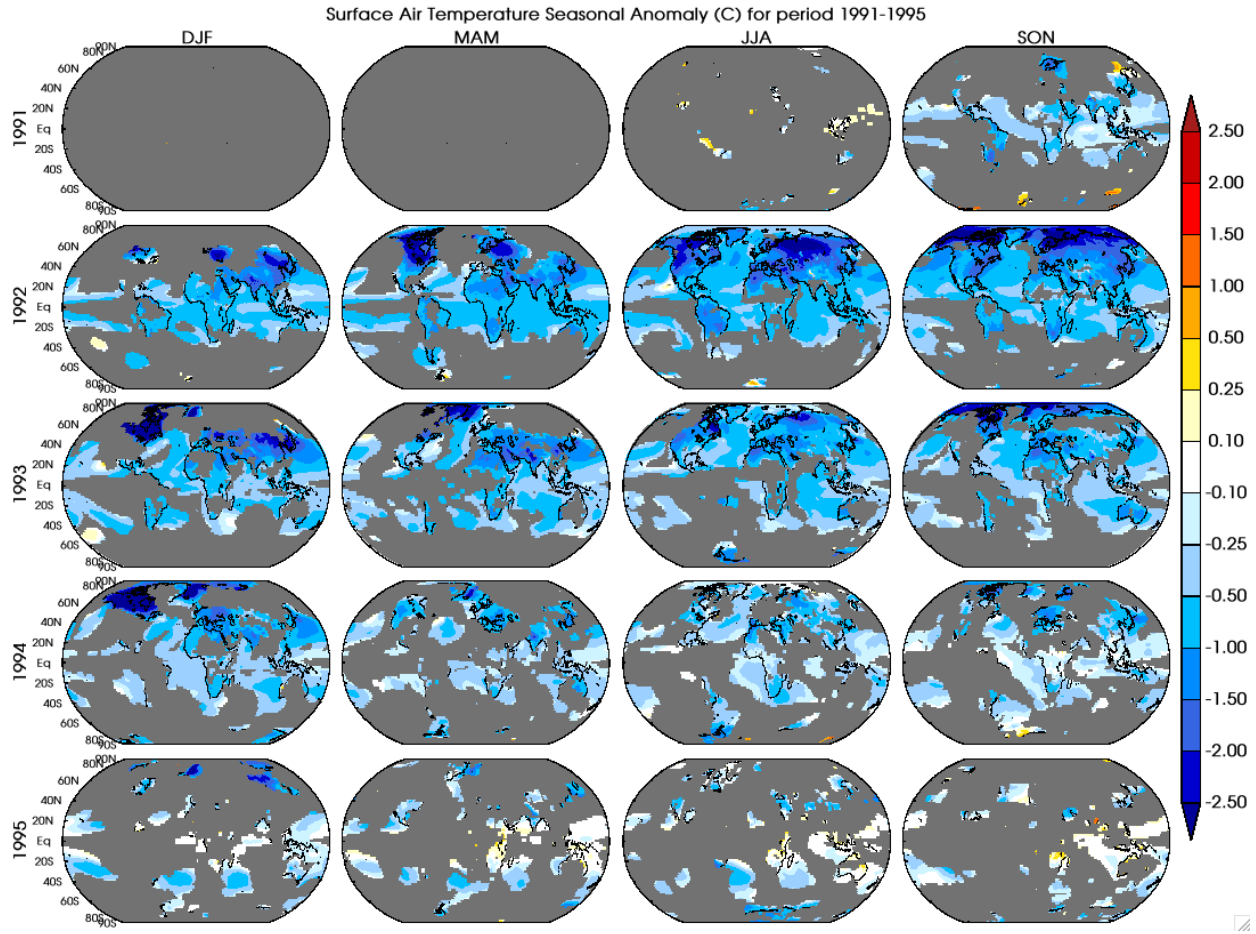


Figure 2: Seasonal mean surface temperature anomalies ( $^{\circ}\text{C}$ ) from the year 1991 to 1995 with respect to the reference period of 1950-2014. A grey color is painted over the grid cells where the surface temperature anomalies are not statistically significant in comparison to the counterfactual ensemble. The colored areas show anomalies of PCH with respect to the climatology from 1950-2014.

### 3.3 Rainfall Response

Precipitation, presented seasonally for the year of eruption (1991) and the following year (1992) in Figure 3, shows a highly complex and variable response to the volcanically induced tropospheric cooling and radiative balance perturbation because of its sensitivity to the other climate system components. Studies have shown that global mean precipitation decreases after large volcanic eruptions (Gu et al., 2007; Gu and Adler, 2012; Iles et al., 2013; Robock and Liu,

1994; Singh et al., 2023; Trenberth and Dai, 2007). Colose et al., (2016) have postulated that the asymmetrical surface cooling and radiative balance perturbation create an energetic deficit in the hemisphere of eruption that constrains the poleward propagation of tropical rainfall belt (ITCZ) in that hemisphere. In the case of the Mt. Pinatubo eruption, the PCH simulations show that regional patches of significant decrease of up to 1 mm per day are noticed over tropical and northern hemispheres (Africa, eastern and northern Asia) after the eruption (Figure 3). Also, increasing rainfall patterns are simulated over the Mediterranean and European regions. Broadly, the confidence level of precipitation response due to volcanic aerosols is strongly influenced by the uncertainty due to many possible factors and prominent modes of atmospheric variability, such as the strength of El Nino (Paik et al., 2020).

The zonal mean of the rainfall response (Figure S5) shows a clear decreasing trend in the northern hemisphere tropical and higher latitudes with a positive rainfall response band around 20° N. The PCH modelled rainfall response due to the Mt. Pinatubo eruption is broadly consistent with the previous studies (Joseph and Zeng, 2011; Liu et al., 2016; Trenberth and Dai, 2007), but the uncertainty in rainfall response is still high. Although we use statistical significance at 95% confidence level as our metric for determining significant anomalies and only a few regions exhibit a spatially coherent and detectable forced response (Figure 3). We acknowledge the signals due to the model's internal variability when averaging the impacts across multiple ensembles, but 11 ensembles are a good compromise between few vs. many ensemble members which was shown to be sufficient to represent significance in climate response at the regional scale (Polvani et al., 2019; Singh and AchutaRao, 2019). The inconsistency and complexity in the precipitation response drive us to explore the compound hydroclimatic pathways that lead to drought risk.

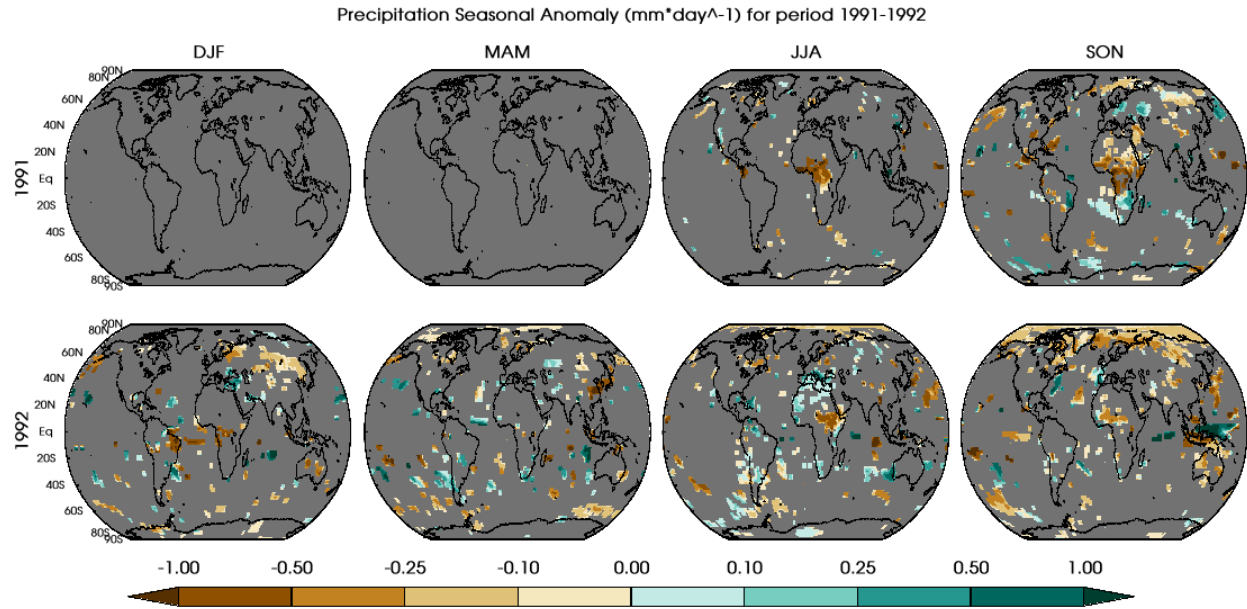


Figure 3. Seasonal mean precipitation anomalies (mm per day) from the year 1991 and 1992 with respect to the reference period of 1950-2014. A grey color is painted over the grid cells where the precipitation anomalies are not statistically significant in comparison to the counterfactual ensemble. The colored areas show anomalies of PCH with respect to the climatology from 1950-2014.

### 3.4 Drought Conditions

Land-atmosphere interactions under a radiatively perturbed environment are crucial in regulating the climate response at regional and sub-regional scales. On short timescales, changes in these interactions can strongly affect plant productivity. Even short-lived adverse conditions in the growth cycle have the potential for outsized impacts, especially if they happen at a particular time in the growing cycle. Hence, we explore the weekly aspects of these drought conditions in Section 3.6 to explore the temporal characteristics of variability in the conditions.

#### 3.4.1 Seasonal Soil Moisture Drought Index (SMDI)

The root zone is commonly defined as the top 3 – 6 feet (0.9-1.8 m) of the soil column (Keshavarz et al., 2014) but most agricultural crops have shallower root systems confined to the

top 2 feet (0.6 m) (Narasimhan and Srinivasan 2005). Hence, we focus on the soil moisture deficit index (SMDI) (Narasimhan and Srinivasan 2005) for the top 2 feet (0.6 m) of ground depth (SMDI\_2) as shown in Figure 4. As anticipated, more land area is covered by statistically different SMDI\_2 than in Figure 3. This enhances our analysis of water-driven impacts on plant productivity, offering greater insights than precipitation.

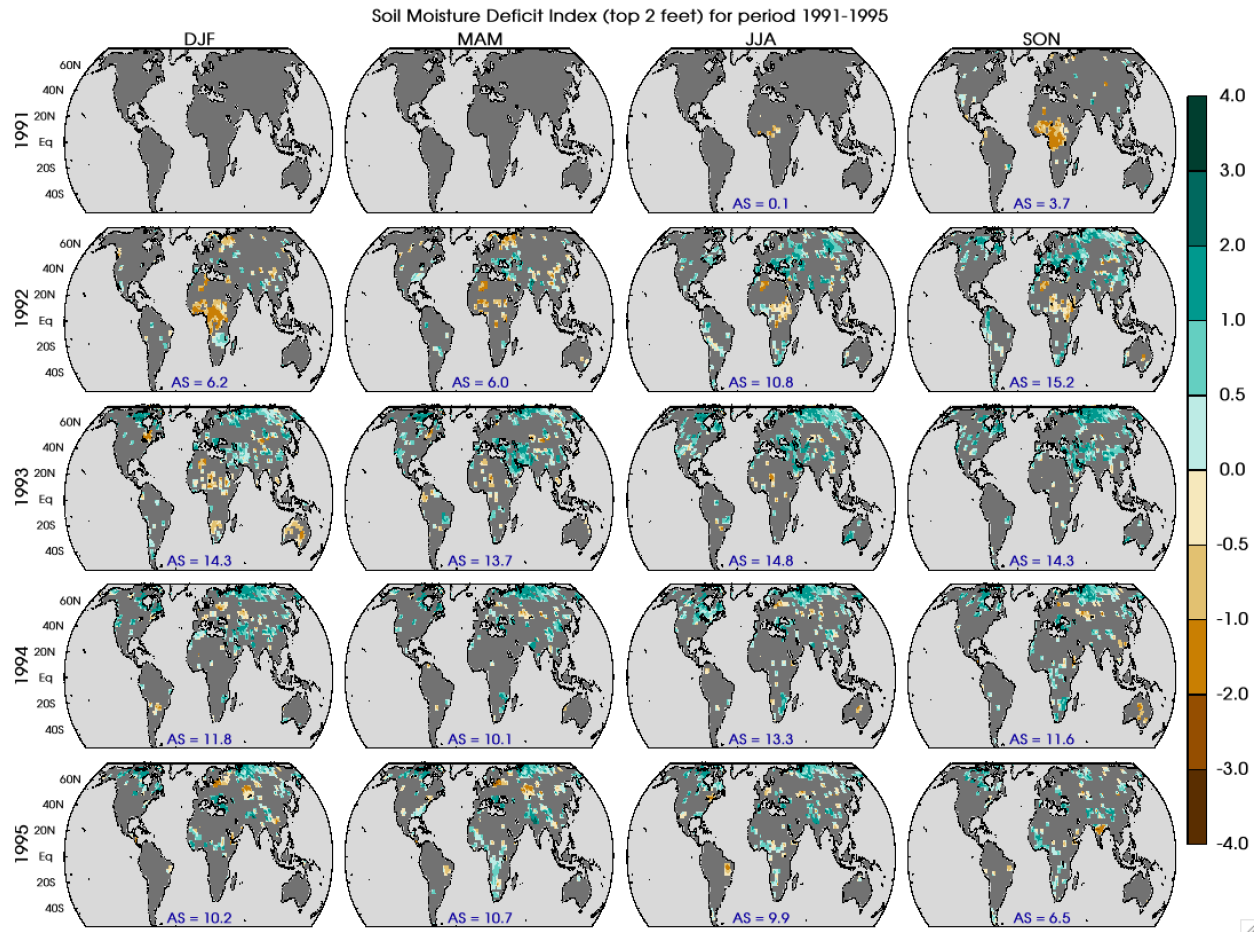


Figure 4. Soil moisture deficit index (SMDI\_2) for the top 2 (0.6 m) of ground depth evaluated seasonally from 1991 to 1995. Grey color is painted over the grid cell where the SMDI\_2 is not statistically significant in contrast to counter-factual ensemble. The parameter AS on each panel marks the percentage of land area that has shown statistically significant dry or wet response after Mt. Pinatubo eruption.

Figure 4 clearly shows that the equatorial region, especially over Africa, has a significant drying response due to Mt. Pinatubo compared to long-term historical data from the SON season

of 1991 through the following DJF season. Although less robust, the dryness in this region lasted through MAM of 1993. The severity of drying response reaches up to -2.0 on a scale of extreme wet/dry at 4.0/-4.0, where the severity of -4.0 reflects the maximum dryness (rarest case) over the entire 1950-2014 calibration period. Figure 3 shows a similar pattern in equatorial African rainfall decrease. A decrease in rainfall was present in the first season post-Mt. Pinatubo eruption, indicating an expected lagged response in SMDI\_2. Spatial coherence between these signals is again re-established in the JJA and SON 1992 seasons, albeit with more variation in the strength of the signal.

Meanwhile, in the high latitudes of the northern hemisphere, we see an increase in the store of soil moisture despite a decrease in rainfall in higher latitudes. An exception is the Mediterranean (extending towards the east Mediterranean and western Asia) region, where soil moisture and rainfall both show an increase after the Mt. Pinatubo eruption. This increase in soil moisture in the northern hemisphere is comparatively more pronounced in the summer months than in the winter season. Thus, despite less water supply through rainfall, there has been a persistent increase in the soil moisture in the root zone layer since the JJA season of 1992. This is likely due to less water extracted from this layer through evaporation and transpiration as well as due to the implemented irrigation in GISS modelE (details in further sections). Overall, Figure 4 shows equatorial drying signals mostly dominated through the DJF season of 1993, but the wet conditions over higher latitudes lasted till 1995. Broadly, 6-13% of the land region has shown a statistically significance response in terms of dry/wet conditions by the end of year 1995 because of the Mt. Pinatubo forcing.

Deeper soil layers better approximate longer-term meteorologically defined drought indices (Narasimhan and Srinivasan 2005). This makes intuitive sense: precipitation provides

the recharge for the store of soil moisture, and if there is a longer-term decline in precipitation, all available moisture will be used hydraulically for plant transpiration (both the deeper stores of water and the soil-penetrating precipitation available), not allowing for deeper depth recharge. Here, we evaluate discrete layer depths instead of cumulative depths for two reasons. First, the soil permeability changes with the depth, and the inclusion of top layers erroneously reflects the SMDI\_2 signal in potentially impermeable regions; second, the SMDI\_2 signal gets superposed over the deeper layer response and misleads the actual soil moisture response for the deeper layers.

As expected, when we evaluate the soil moisture deficit response between 2-4 feet (0.6-1.2 m) soil depth (SMDI\_4) in Figure 5 and 4-6-feet (1.2-1.8 m) soil depth (SMDI\_6) in Figure S6, we see similar spatial and temporal distributions as shown in Figure 4 with a corresponding decrease in the percentage of area response. Spatially, we see high latitudes across North America and Northeastern and western Asia, equatorial Africa, Europe, and Mediterranean regions maintain their SMDI-2 trend in Figure 4. However, the total response area decreases from peak coverages of 12-13% in SMDI-2 to less than 10-12% in SMDI-4 and 7-10% in SMDI\_6 (as shown in Figure S6). Additional decreases in the degree of impact are also seen between the three soil layers. Note that the light grey colored regions in Figures 5 and S6 represent regions of impermeability, which does affect the total area of response.

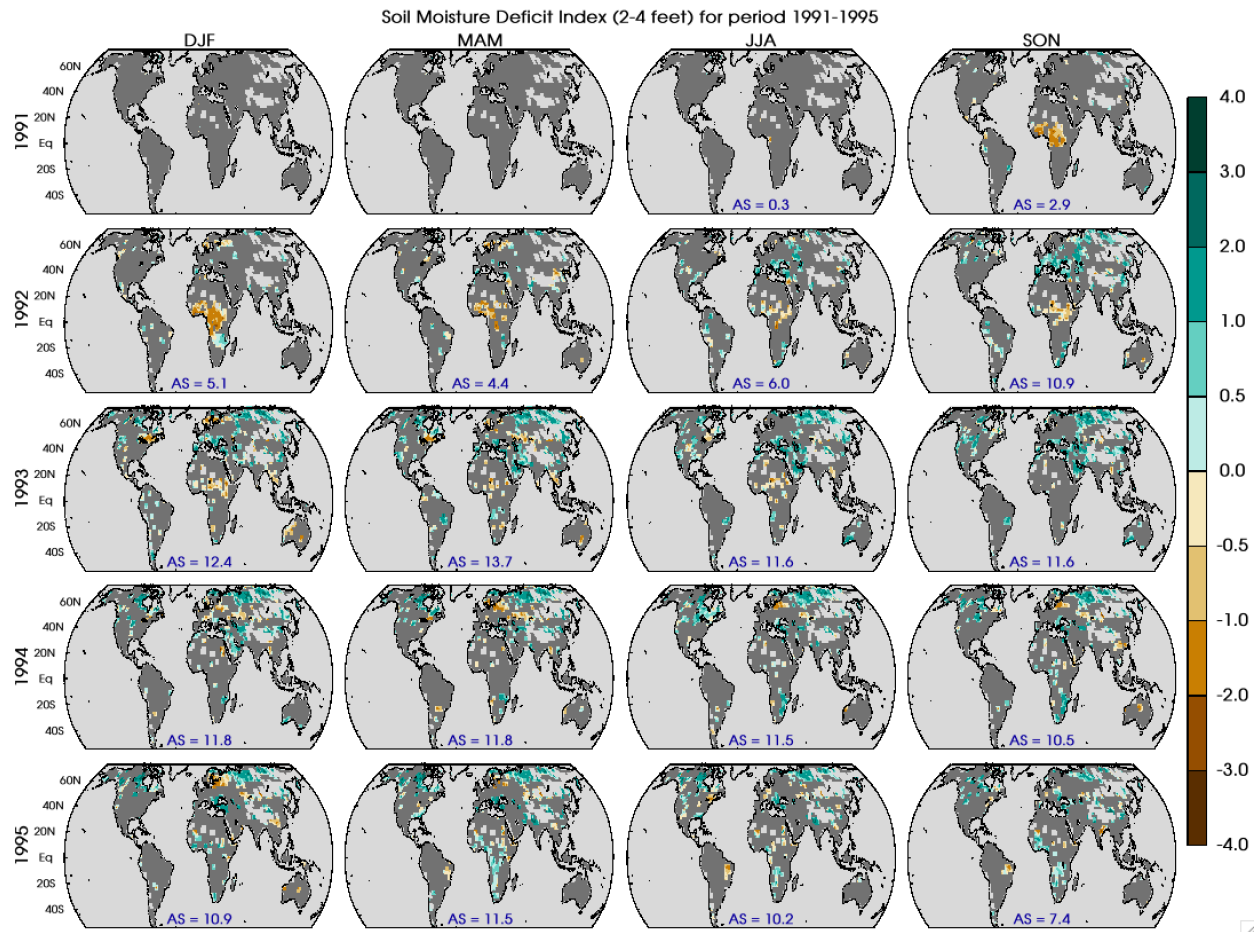


Figure 5. Soil moisture deficit index (SMDI<sub>4</sub>) for soil depths between 2-4 feet (0.6-1.2 meter) evaluated seasonally from 1991 to 1995. The grey color is painted over the grid cell where the SMDI<sub>4</sub> is not statistically significant in contrast to the counterfactual ensemble. The light grey colored regions represent regions of impermeability. The parameter AS on each panel marks the percentage of land area showing statistically significant dry or wet response after the Mt. Pinatubo eruption.

### 3.4.2 Seasonal Evapotranspiration Deficit Index (ETDI)

As indicated in the methodology section, ETDI calculation is similar to SMDI but is based on the water stress anomaly, which accounts for the difference between actual and potential evapotranspiration. ETDI is a measure of the flux of water between land and atmosphere, and like SMDI<sub>2</sub> in Figure 5, it shows robust statistical differences over land.

Figure 6 illustrates the reduction in ETDI response in the equatorial region started developing during the DJF season of the year 1992, and these conditions persisted over the year. Like SMDI\_2, ETDI increases over the regions encompassing the Mediterranean and western Asia. However, ETDI differs from the SMDI\_2 over some northern hemisphere regions, especially Northeastern Asia. A drying response in terms of ETDI in the northern hemisphere regions persisted during 1993 and 1994, whereas SMDI\_2 shows an opposite response. This contrasting response through the ETDI and SMDI\_2 points to the complexity of land-atmosphere interactions over these regions. We utilized model simulated surface energy fluxes (Latent and Sensible heat) to calculate the potential (PET) and actual evapotranspiration (AET) to estimate the water stress ratio. In these regions, soil moisture is available during the summer and early winter months. Still, an evapotranspiration deficit reflects the decrease in plant transpiration (latent heat flux), which may be due to the unavailability of plants. Also, the surface temperature (sensible heat flux) response supports the non-water-stressed atmospheric conditions, and thus, overall, it shows an evapotranspiration deficit. Areas of significant response in terms of ETDI vary from 7 to 14.5% on a seasonal basis during the years following the eruption. The most significant areas of ETDI coverage occur during the same periods as SMDI\_2 (between JJA 1992 – JJA 1993).

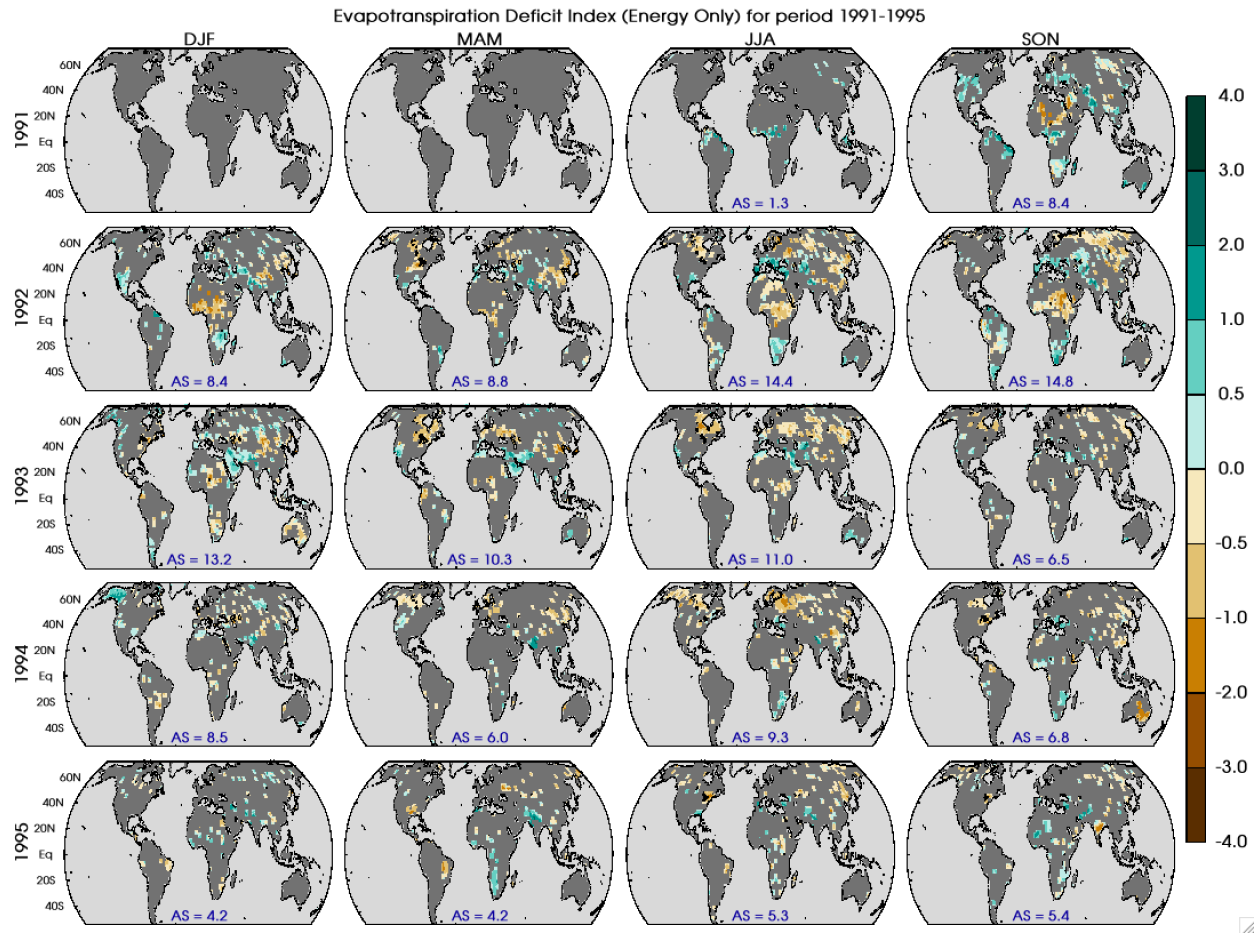


Figure 6. Evapotranspiration deficit index (ETDI) at seasonal scale from 1991 to 1995. The grey color is painted over the grid cell where the ETDI is not statistically significant compared to the counterfactual ensemble. The parameter AS on each panel marks the percentage of land area that has shown statistically significant dry or wet response after Mt. Pinatubo eruption.

### 3.5 Seasonal plant productivity inferences

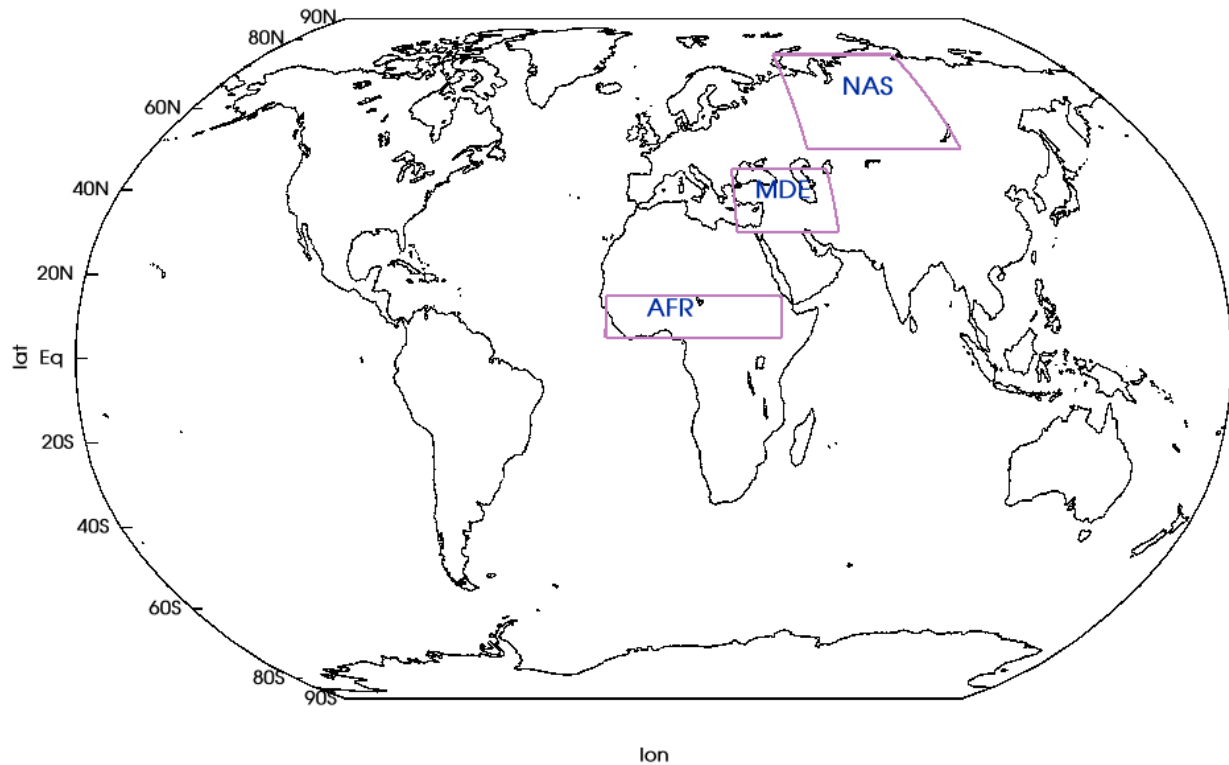
SMDI (at depths of 0-2 (0.6m), 2-4 (0.6m-1.2m), and 4-6 feet (1.2m -1.6m)) and ETDI have proven helpful in analyzing the climatic impact of the Mt. Pinatubo eruption on a seasonal scale. Additionally, SMDI\_2 (top 2 feet or 0.6 m) and ETDI have demonstrated a slow development of drought conditions, beginning by the end of the year 1991 (SON season), reflecting a time lag between seasonal precipitation patterns (Narasimhan and Srinivasan 2005). Crucially, the seasonal depiction of drying/wet conditions via SMDI and ETDI provides a comprehensive

overview of prolonged or recurrent dry/wet conditions in susceptible regions. Moreover, understanding these typical agricultural drought indices indicates potential effects on plant productivity at the seasonal scale.

In equatorial Africa, decreases in both SMDI and ETDI indicated that there was likely a negative impact on plant productivity. The Mediterranean region (encompassing the eastern Mediterranean and western Asian region) showed increased SMDI and ETDI, indicating a positive effect on plant productivity. Northern Asia, on the other hand, exhibited an increase in SMDI with a decrease in ETDI, suggesting that plant productivity likely decreased, but not because of water-based drivers.

### **3.6 High frequency impact pathways evaluation**

Here, we use the daily model output to calculate weekly drought indices in each grid cell. These weekly scale drought indices and changes in other atmospheric variables are explored at the regional scale to understand the associated land-atmosphere interactions pertaining to higher-order impacts. The higher temporal resolution of these parameters is crucial for analyzing different stages of the crop cycle in a region. Considering the complexity of the representation of spatial features, we selected three distinct regions (shown in Figure 7 and detailed in caption) in the northern hemisphere based on the climate response to Mt. Pinatubo in the seasonal analyses presented in Section 3.0. We followed the same strategy described in Section 2.3.1 to mask out the statistically insignificant grid cells using the counterfactual ensemble after creating the weekly time series for different drought indices and atmospheric parameters.



553  
 554 Figure 7: Demarcation of the regions selected over tropics, Equatorial Africa Region (AFR; Lat:  
 555 5°N – 15°N, Lon: 15°W– 40°E), mid-latitude, Middle East Region (MDE; Lat: 30°N – 45°N,  
 556 Lon: 27°E – 60°E), and in the high latitudes, Northern Asia Region (NAS; Lat: 50°N –75°N,  
 557 Lon: 55°W –110°E ).



Figure 8. Spatially averaged drought indices (SMDI\_2 & ETDI) and anomalies for other drivers (Actual and Potential Evapotranspiration, Transpiration, Moisture Source (Precipitation plus irrigation) and Surface Temperature) at weekly scale for the equatorial Africa region (Latitude = 5°-15° N, Longitude = 15° W- 40° E).

### 3.6.1 Equatorial Africa

Figure 8 shows the weekly response to volcanic forcing for the years 1991-1995 in terms of agricultural drought indices (SMDI\_2 & ETDI), AET, PET, transpiration, total soil moisture source, and surface temperature for an equatorial region in northern Africa. This region exhibits consistent statistical differences across the drivers on a weekly scale; thus, for most of the time period, this is unmasked, revealing the degree of anomalous conditions.

This region lies between the latitude 5-15° N, where the precipitation during the monsoon season shows a decrease in response to a southern migration of the inter-tropical convergence zone (Iles et al., 2013; Colose et al., 2016; Singh et al., 2023). Weekly precipitation change in the equatorial Africa shows a significant deficit of more than 1.5 mm per day consistently for several weeks, especially during the JJAS monsoon season. This region also shows that a deficit in precipitation during the major precipitation season (JJAS) can result in a soil moisture deficit in the root zone in the following seasons (DJF and MAM in SMDI\_2) and consequently affect the entire crop cycle. The root zone soil moisture, SMDI\_2, also shows a persistent drying through 1993, and combined with the lack of precipitation, the potential for recharge is limited. Also, this region has no contribution from irrigation as a source of additional soil moisture, as shown in Figure S7 (bottom panel) (Cook et al., 2020). Cumulative annual rainfall change over this region shows a deficit of 33.2, 9.5, and 3.2 mm per day for the years 1991, 1992, and 1993 and an increase of 10.5 and 13.6 mm per day for the years 1994 and 1995, respectively, where soil-moisture response shows a recovery from the dry conditions.

Hence, it is unsurprising that a corresponding decrease in ETDI through 1993 is consistent with this lack of moisture. However, the evaporative demand, as shown by surface temperature change, does not consistently decrease until September of 1991, and hence, ETDI is slow to show a response in the deficit index. After that point, evaporative demand decreases with lower temperatures, but the evapotranspiration is dominated by transpiration (Seneviratne et al 2010; Nilson and Assmann 2007), and so the majority of the decrease in ETDI is explained by the shown decrease in plant transpiration. As expected, this decrease in plant transpiration is well correlated with decreases in AET. Conclusively, precipitation response in this region shows dominance in regulating the ecohydrological conditions. A substantial decrease in the weekly

rainfall over the region perpetuates a root-zone water deficit, resulting in decreased plant transpiration. Decreases in both SMDI\_2 and ETDI thus indicate developing agricultural drought conditions, which are confirmed by a decrease in the direct measure of plant transpiration.

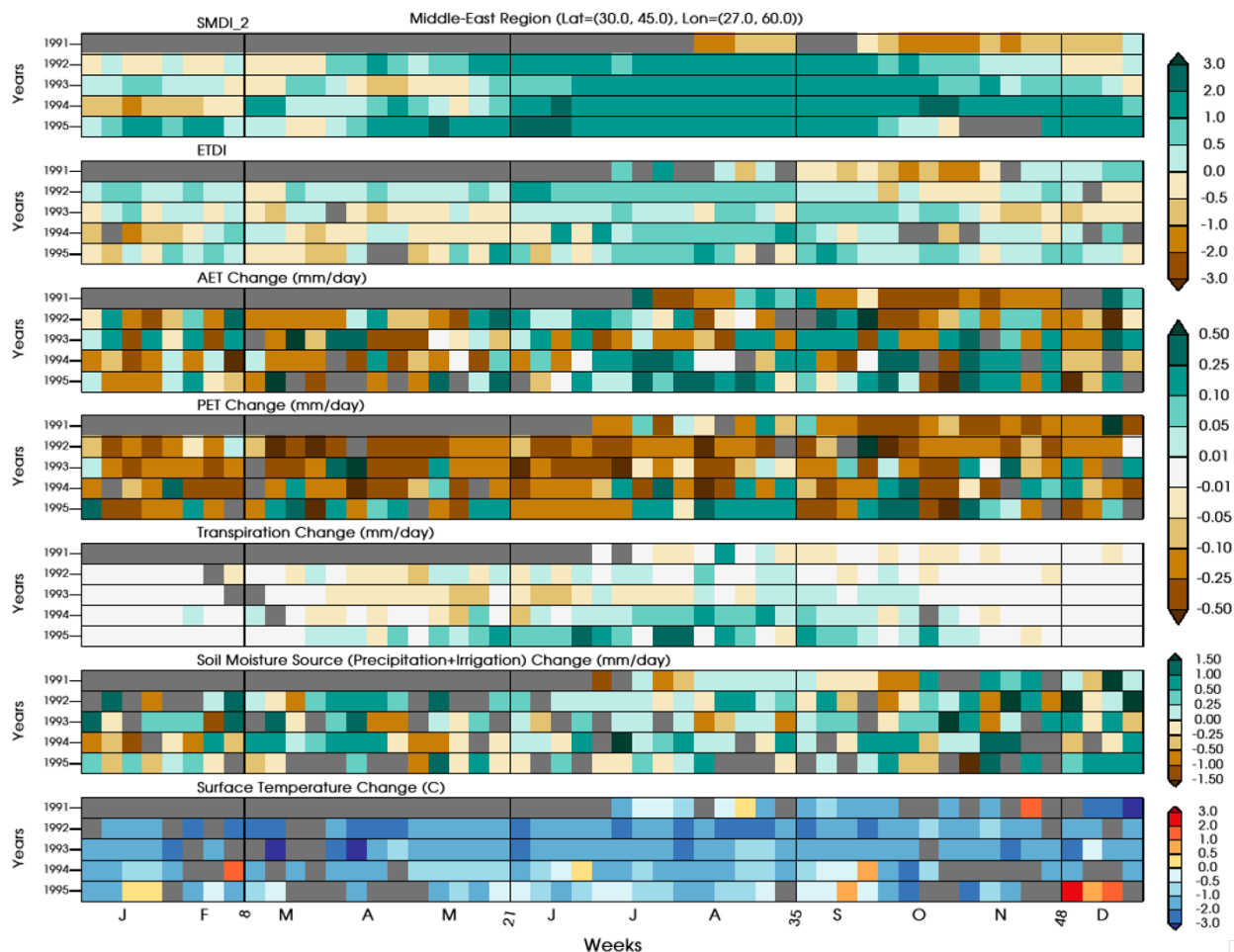


Figure 9. Spatially averaged drought indices (SMDI\_2 & ETDI) and anomalies for other drivers (Actual and Potential Evapotranspiration, Transpiration, Moisture Source (Precipitation plus irrigation) and Surface Temperature) at weekly scale for Middle East (MDE) Region (Latitude =30° N - 45° N, Longitude=27°-60° E; Eastern Mediterranean /Western Asian).

### 3.6.2 Middle East Region

Figure 9 shows the region covering the eastern Mediterranean and the western Asian regions where rainfall slightly increased after the Mt. Pinatubo eruption. Additionally, this region exhibits a significantly positive trend in irrigation practices post-1950, with a substantial peak

over the eastern Mediterranean region following the Mt. Pinatubo eruption (Cook et al., 2020; Figure 1 and 2).

In the eastern Mediterranean, wet and cold autumns and winters persist for several years after the Mt. Pinatubo eruption, offering significant root zone recharge potential. The summer months, in general, reflect a slightly uncertain model response in the regional rainfall, with some weeks of deficit and some excess, but an additional water supply through irrigation contributes to the overall moisture content in the region (Figure S7). Root zone soil moisture (SMDI\_2) indicates sufficient water availability throughout the growing seasons during the entire analysis period. Overall, the findings demonstrate that this region is not moisture-limited, with adequate precipitation and irrigation ensuring the replenishment of root zone moisture as plants grow. Cumulative weekly anomalies show that precipitation change in 1991 is slightly negative (-0.5 mm per day), but an increase in annual rainfall of 13.8, 8.0, 10.9, and 4.5 mm per day is simulated for the years 1992, 1993, 1994, and 1995, respectively. Irrigation implemented across this Middle East region exhibits a strong positive trend from 1950 to 2005 (Cook et al., 2020). Notably, a significant cumulative increase in irrigation of 0.5, 1.3, 1.3, 0.8, and 0.9 mm per day during the years 1991 to 1995 provides an additional source of moisture supply for the region (Cook et al., 2020). This irrigation is particularly crucial during the summer months, compensating for rainfall deficits lasting several weeks and contributing 10–20% of soil moisture source changes (Figure S7).

The corresponding increases in ETDI and AET indicate the abundant water availability for transpiration in the region. While transpiration remains temporally correlated with AET, the increases are less pronounced. Simultaneously, there is a decline in PET response, which is associated with a more significant temperature decrease in this region compared to equatorial

629 Africa. The reduction in PET, combined with the increase or maintenance of AET through  
630 transpiration, leads to an increase in ETDI. Consequently, agricultural productivity in this region  
631 is generally positively impacted, as ample moisture provides essential support. Nevertheless, the  
632 data reveals heterogeneous patterns, indicating that 1993 may have influenced plant productivity  
633 with a positive but lower-magnitude ETDI, inconsistent AET, and reduced transpiration.

634       Regardless of the presence of a volcanically induced response, the weekly scale analysis  
635 demonstrates its importance by virtue of an example from 1993, where rainfall deficit is  
636 produced during the 15th and 16th weeks (April) of the year. Combined with lower SMDI\_2, this  
637 may lead to insufficient moisture availability during a critical stage of the crop cycle. The  
638 duration of such moisture deficits could significantly impact overall seasonal crop production.  
639 Consequently, even if favorable conditions prevail for most of the crop cycle, adverse effects  
640 during essential phases can critically influence yields in ways that seasonal averages fail to  
641 capture.

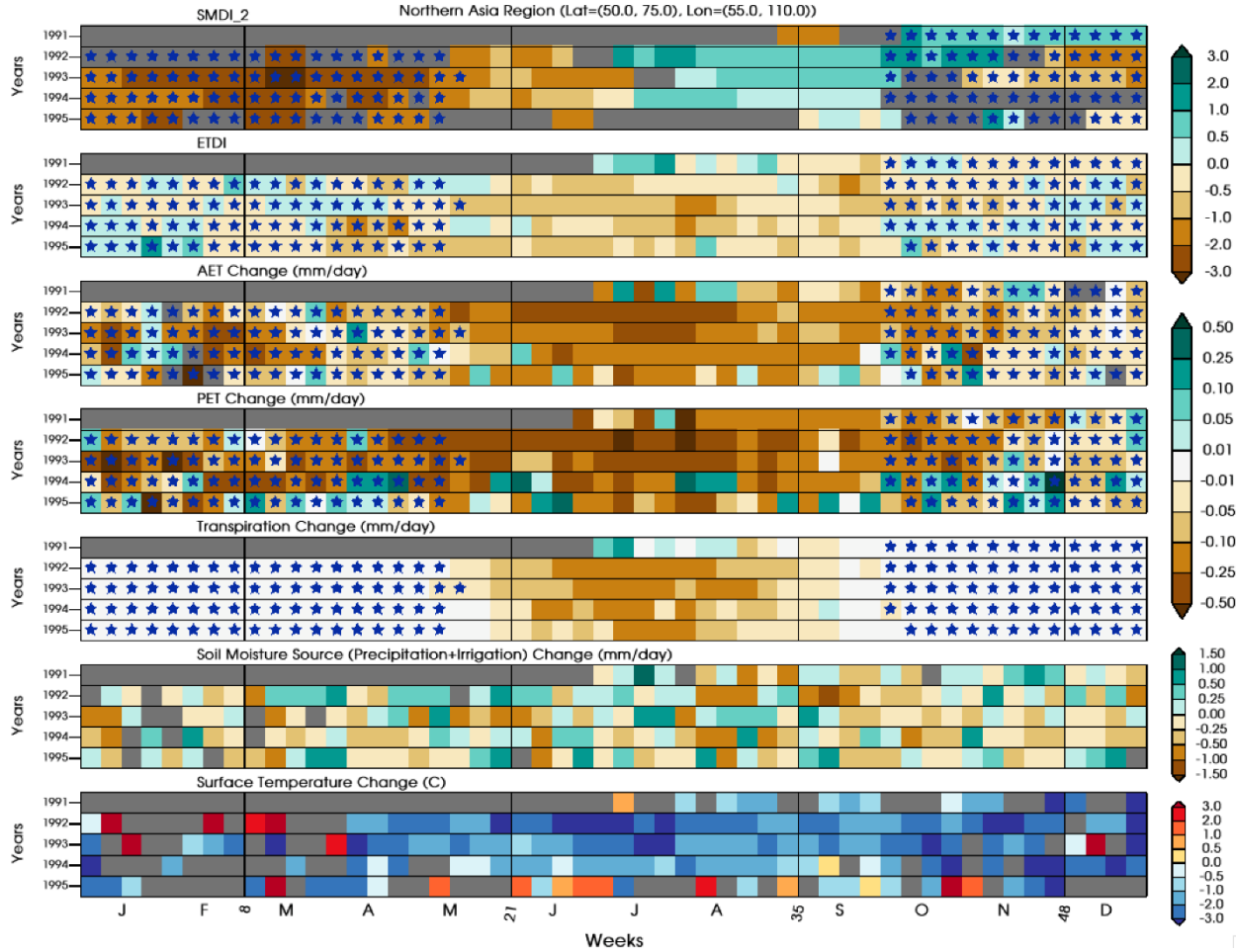


Figure 10. Spatially averaged drought indices (SMDI\_2 & ETDI) and anomalies for other drivers (Actual and Potential Evapotranspiration, Transpiration, Moisture Source (Precipitation plus irrigation) at weekly scale Northern Asia Region (Latitude = 50° N - 75° N, Longitude = 55° - 110° E). Blue stars represent the weeks with average surface temperature below freezing point.

### 3.6.3 Northern Asia

Finally, we selected a region (NAS) in higher latitudes to explore the interplay between the various drivers governing the conditions for plant productivity. This region consistently exhibits statistically significant differences across drivers on a weekly scale. However, higher latitudes experience strong seasonal controls on plant productivity, with below-freezing temperatures (indicated by blue stars) halting growth. Therefore, our analysis focuses on months when plant growth is possible (~MJJAS).

Precipitation changes over the entire analysis period are highly uncertain. However, there is generally a slight trend toward increased precipitation from November 1991 to June 1992, followed by a decline through 1994. As shown in Figure S5, the contribution of irrigation to soil moisture in this region is negligible. Cumulative weekly precipitation anomalies indicate an annual increase of 0.05 mm per day in 1991, followed by decreases of -2.0, -4.1, -6.8, and -1.0 mm per day from 1992 to 1995, respectively. Alternatively, root zone moisture shows that sufficient water available to plants during the JAS (July-August-September) growing months after a substantial deficit in the early MJ (May-June) months. Certainly, during the summer months, the melting of frozen surfaces and snow supplies moisture to the upper layers, resulting in wet conditions, which accounts for this strong dichotomy. However, there were no corresponding increases in ETDI and AET following the 1991 season. This indicates that despite ample water, plants are still not growing; this is conclusively confirmed by the decrease in transpiration starting in 1992. Meanwhile, the simultaneous reduction of PET response is correlated with the most substantial decline in temperature, on the order of 2-3° C, for this region.

Unlike the other two regions for which SMDI\_2 and ETDI exhibited similar wet/dry patterns, this region shows diverging patterns. Overall, this indicates that although moisture is available to support plant productivity, it is not being effectively utilized. Therefore, other factors must be responsible for the decrease in plant transpiration and ETDI. The more significant decline in PET compared to AET suggests that temperature is playing a role. Since temperature is directly related to reduced incident radiation, the combined effects of temperature and radiation are likely the primary factors controlling the reduced plant productivity in this region, rather than moisture conditions.

## 4.0 Conclusions

This study utilized the NASA GISS ModelE2.1 (MATRIX) Earth system modeling framework to investigate the mechanisms by which the 1991 Mt. Pinatubo eruption influenced hydroclimatic conditions and water-based drivers of plant productivity. The simulation successfully reproduced key microphysical properties ( $R_{eff} \approx 0.5 \mu\text{m}$ ,  $\text{AOD} \approx 0.22$ ), radiative forcing ( $\sim -5 \text{ Wm}^{-2}$ ), surface cooling ( $\sim 0.5^\circ\text{C}$ ), and regional rainfall changes consistent with previous studies. The study further examined secondary impacts of the eruption, beyond changes in radiation and temperature, by analyzing agricultural drought indices to better assess its effects on plant productivity. Two metrics, SMDI and ETDI, which account for land-atmosphere interactions, were utilized to capture the progression of short- and long-term conditions affecting plant productivity, particularly in agricultural contexts. These drought indices confirm the moisture-driven dry and wet patterns observed in early 1992 and the following years over the tropical regions and mid-latitudes of the Northern Hemisphere, respectively, as a response to the radiative perturbation caused by the Mt. Pinatubo eruption. Based on both drought indices, we conclude that approximately 10-15% of land areas exhibit statistically significant dry or wet patterns under the volcanically altered climate conditions of 1992 and 1993. The fraction of land region showing a significant dry or wet response range between 5-10% for the next two (1994 and 1995) years. In equatorial Africa, seasonal decreases in both SMDI and ETDI suggested a likely negative impact on plant productivity. In contrast, the Middle East showed increases in both SMDI and ETDI, indicating a positive effect on plant productivity. Northern Asia, by comparison, exhibited an increase in SMDI alongside a decrease in ETDI, implying a reduction in plant productivity, though not driven by water-related factors.

Motivated by these key pattern differences, we extended our analysis of the drought indices by incorporating higher temporal (weekly) frequencies along with AET, PET, and transpiration. These regional analyses generally exhibit much stronger statistical significance on a weekly scale and further confirm the season-based inferences described above. Further, weekly drought indices show the temporal variability characteristics in the signal, which also demonstrates the utility of explaining the effectiveness of short-term dry/wet conditions corresponding to a regional crop cycle. In locations with insufficient/excess soil moisture, there is a corresponding decrease/increase in evapotranspiration (AFR/MDE) and hence decreased/increased plant productivity.

Kandlbauer et al., (2013) examined crop responses (using C3 and C4 grasses as proxies) to the 1815 Tambora eruption using the HadGEM-ES model in three regions very similar to those in our study. Their findings suggest that plant productivity decreases with positive changes in soil moisture in the higher-latitude Asian region. In the mid-latitudes over the Southern Europe/Middle East region (adjacent to our MDE region), volcanic eruptions may enhance plant productivity by providing additional soil moisture through increased rainfall. However, in the MDE region in our study, we found that the applied irrigation also benefits soil moisture supply along with the increased rainfall. Furthermore, both studies report a decrease in productivity in the tropical region. In general, these results complement the findings of this study, which suggest that if sufficient water is available in the Southern Europe/Middle East region, volcanic eruptions may enhance plant productivity. In contrast, in the far northern latitudes, water is not the primary driver of plant responses, and productivity is likely to decline. Seasonal-scale changes in gross primary productivity (GPP) confirm the regional trends in plant productivity following the eruption. The simulations show a more pronounced decrease in GPP in the northern high-latitude

region and a significant increase in GPP over the European and Mediterranean regions. Additionally, distinct patterns of decrease and increase in GPP are simulated in the tropical northern and southern regions, respectively (Figure S9).

This study is the first to conclusively demonstrate that there is an excess of root-zone soil moisture in high-latitude regions (NAS) that plants are not utilizing for growth, indicating that temperature and radiation are likely the primary controlling factors, thus confirming previous findings (Krakauer and Randerson, 2003) and (Dong and Dai, 2017). The intricate nature of the compounded response, particularly in relation to soil moisture-driven impact pathways in tropical regions and high-latitude areas of the Northern Hemisphere, highlights the need to expand the investigation beyond soil moisture and land-atmosphere interactions. The current configuration of the NASA GISS model operates with prescribed vegetation with static plant functional types and leaf area index. Incorporating dynamic vegetation could be essential for capturing interactive land surface responses. Additionally, evaluating the influence of regional and local biomes on photosynthesis rates could offer deeper insights into how these processes respond to the climatic impacts of volcanic forcings. McDermid et al. (2022) have demonstrated the sensitivity of regional hydroclimate to the local changes in soil organic carbon changes using the soil moisture content. The results presented in this study regarding soil-moisture-based drivers to plant productivity and surface temperature response in the northern hemisphere's high latitudes also hinted towards the dominance of temperature effects on enhanced carbon sink in terms of soil and plant respiration and reduced NPP (Krakauer and Randerson, 2003; Lucht et al., 2002). Meanwhile, water-based drivers predominantly influence productivity responses in many tropical and subtropical regions. Our findings demonstrate that soil moisture conditions across different regions can provide valuable insights into the full impacts on agricultural yields and

regional carbon sink responses, particularly under scenarios involving changes in dynamic vegetation and crop cover. A recently developed fully demographic dynamic vegetation model (ModelE-BiomE v.1.0 (Weng et al., 2022)) incorporates interactive biophysical and biogeochemical feedbacks between climate and land systems within the NASA GISS ModelE framework. This model could be instrumental in evaluating carbon cycle responses under such forcings.

### **Code/Data availability.**

Details to support the results in the manuscript is available as supplementary information is provided with the manuscript. GISS Model code snapshots are available at <https://simplex.giss.nasa.gov/snapshots/> (National Aeronautics and Space Administration, 2024) and calculated diagnostics are available at zenodo repository (<https://zenodo.org/records/12734905>) (Singh et al., 2024). However, raw model output and data at high temporal (daily) resolution and codes are available on request from author due to large data volume.

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### **Author's contributions**

RS, KT, DB, LS and KM identified the study period in consultation with the other authors and RS, KT, DB, LS and BW designed the underlying simulations strategies. RS and KT implemented it and performed the simulations using NASA GISS ModelE. RS and KT have performed the analysis. RS created the figures in close collaboration with all authors. RS wrote the first draft of the manuscript, and all other authors has contributed the writing of subsequent drafts. All authors contributed to the interpretation of results.

#### **Competing interests**

One of the co-authors is member of the editorial board of Atmospheric Chemistry and Physics.

#### **Short Summary**

Analysis of post-eruption climate conditions using the impact metrics is crucial for understanding the hydroclimatic responses. We used NASA's Earth system model to perform the experiments and utilized the moisture-based impact metrics and hydrological variables to investigate the effect of volcanically induced conditions that govern plant productivity. This study highlights the impact of the Mt. Pinatubo eruption on the drivers of plant productivity, as well as the regional and seasonal dependence of these drivers.

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