

Mount Pinatubo's effect on the moisture-based drivers of plant productivity

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Abstract

Large volcanic eruptions can significantly affect the state of the climate, including stratospheric sulfate concentrations, surface and top-of-atmosphere radiative fluxes, stratospheric and surface temperature, and regional hydroclimate. The prevalence of higher natural variability in the regional rainfall response to volcanic-induced climate perturbations creates a gap in our understanding of how eruptions affect ecohydrological conditions and plant productivity. Here, we will explore the understudied store (soil moisture) and flux (evapotranspiration) of water as the short-term ecohydrological control over plant productivity in response to the 1991 eruption of Mt. Pinatubo. We used the NASA's Earth system model for modeling of the Mt. Pinatubo eruption and detect the ensuing hydroclimate responses. The model simulates a mean surface cooling of ~0.5 °C following the Mt. Pinatubo eruption. The rainfall response is spatially heterogeneous with large temporal variability, yet still shows suppressed rainfall in the northern hemisphere after the eruption. We find that up to 10-15% of land regions show a statistically significant hydroclimate response (wet and dry) as calculated by the Soil Moisture Deficit Index (SMDI) and Evapotranspiration Deficit Index (ETDI). Results confirm that these impact metrics successfully present a more robust understanding of plant productivity. Our results also explain the

geographical dependence of various contributing factors to the compound response and their implications for exploring the climate impacts of such episodic forcings.

Introduction

Volcanic eruptions are the most prominent source of sulfate aerosols in the stratosphere and are natural drivers of climate variability. Volcanically injected sulfate aerosols in the stratosphere alter the Earth's radiative balance by simultaneously reflecting incoming solar radiation and absorbing outgoing longwave radiation emitted from the Earth's surface (Robock, 2000). The presence of sulfate aerosol for months to years after an eruption and its microphysical transformation in the stratosphere affect the climate system through numerous direct and indirect effects (Barnes and Hofmann, 1997; Briffa et al., 1998; Deshler et al., 2003; Kremser et al., 2016; LeGrande et al., 2016; Marshall et al., 2022; Tejedor et al., 2021; Timmreck, 2012; Toohey et al., 2019). The Mt. Pinatubo eruption (June 1991) remains the largest eruption in the satellite era, and it has been explicitly documented and analyzed for its radiative and climate impacts. Numerous studies based on satellite observations and supported through different modeling efforts estimated that Mt. Pinatubo injected 10-20 Tg of SO₂ at a range of 18-25 km of plume height (Aquila et al., 2012; Bluth et al., 1992; Dhomse et al., 2014; Gao et al., 2023; Sheng et al., 2015a, b; Stenchikov et al., 1998). The radiative impacts of Mt. Pinatubo's eruption estimate an aerosol optical depth of 0.15 for 550 nm wavelength, with an effective radius in the range of 0.16 to 1 micrometer (μm) and net radiative forcing on the order of 5-6 Wm⁻² (Lacis, 2015; Lacis et al., 1992; Sato et al., 1993; Stenchikov et al., 1998). Estimates of the induced surface cooling range between 0.3-0.7 °C; lower stratosphere warming estimates are in the range of 2-3 °C (Dutton and Christy, 1992; Hansen et al., 1992; Labitzke and McCormick, 1992; Lacis et al., 1992; Minnis et al., 1993; Ramachandran et al., 2000; Stenchikov et al., 1998).

In this study, we aim to explore the mechanisms by which the Mt. Pinatubo eruption affected the hydroclimatic conditions and water-based drivers of plant productivity. Agricultural productivity is sensitive to changes in temperature and precipitation (Lobell and Field, 2007; Olesen and Bindi, 2002; Rosenzweig and Parry, 1994). Although much work has been devoted to understanding Mt. Pinatubo's impacts on plant productivity, the literature has been dominated by studies focusing on impacts from changes to the quantity and quality of incoming solar radiation (Farquhar and Roderick, 2003; Gu et al., 2002, 2003; Jones and Cox, 2001; Robock, 2005). Proctor et al. (2018) have estimated a decrease in C4 (maize) and C3 (soy, rice, and wheat) agricultural crop production in response to Mt. Pinatubo driven mainly by changes in incoming radiation. Krakauer and Randerson (2003) evaluated the role of surface cooling in reduced net primary productivity (NPP) using tree ring growth patterns following multiple Mt. Pinatubo-sized eruptions in the last millennium record. Reduced NPP was found in northern mid to high latitudes, while the signal in the lower latitudes and tropics was either insignificant or constrained by other factors. Other studies have further expanded into societal impact research focusing on volcanically induced changes to harvest and agricultural productivity over different regions (van Dijk et al., 2023; Hao et al., 2020; Huhtamaa and Helama, 2017; Manning et al., 2017; Singh et al., 2023; Toohey et al., 2016).

Similarly, a plenty of work has been devoted to understanding the hydroclimate response to Mt. Pinatubo through changes in atmospheric precipitation (Barnes et al., 2016; Paik et al., 2020; Trenberth and Dai, 2007), as the monsoon seasonal rainfall decreases in the season following the eruption (Colose et al., 2016; Iles et al., 2013; Liu et al., 2016; Singh et al., 2023; Tejedor et al., 2021). It is also shown that volcanic eruptions can alter regional rainfall and hydroclimate in general, which could prominently affect regional plant productivity (Zuo et al.,

2019a, b). However, rainfall alone provides an incomplete understanding of the drought conditions relevant to plant productivity; a rainfall deficit could, in principle, be overcome by moisture stored in the soil. Hence, meteorological drought indices (e.g. SPI (McKee et al., 1993) or PDSI (Palmer 1965)) ignore a full water balance approach. Furthermore, meteorological drought indices tend to be designed to evaluate prolonged periods of abnormally dry weather conditions. For instance, PDSI is an indicator of drought with a 9-month horizon (Mullapudi et al., 2023). Yet, agricultural crops are heterogeneously sensitive to the timing and degree of moisture deficits during particular portions of the crop growth cycle (Hane and Pumphrey, 1984). Thus, consideration of indices with high temporal frequency can be critical when focusing on agriculture.

Soil moisture is the stock of water stored underground and is a primary source of water flux to the atmosphere and plants through evapotranspiration. Evaporation of water from bare surface soil or transpiration of water during photosynthesis in plants from the root zone soils uses a large portion of absorbed solar energy (Trenberth et al., 2009). Plant transpiration is the largest contributor to land evapotranspiration (Dirmeyer et al., 2006; Lawrence et al., 2007; Nilson and Assmann, 2007; Seneviratne et al., 2010). Soil moisture decrease in the root zone establishes an essential control over plant productivity as transpiration is an integral component of photosynthesis (Chen and Coughenour, 2004; Denissen et al., 2022). Multiple studies have established that water supply is the limiting factor for climatic evapotranspiration over tropical and subtropical land areas, while temperature is an important controlling factor in northern mid- and high latitudes (Dong and Dai, 2017; Mintz and Walker, 1993; Nemani et al., 2003; Pan et al., 2015). However, soil moisture changes in response to the Mt. Pinatubo eruption (1991) are

underreported in the literature, and it is unclear how soil moisture would respond given volcanically forced changes in primary drivers (temperature and precipitation).

Studying a large (10xPinatubo) volcanic eruption, (Frölicher et al., 2011) have shown that the terrestrial carbon pool is sensitive to the regional (in the tropics and sub-tropics) soil-moisture content through the net-ecosystem productivity. Using the geoengineering large ensemble simulations with the CESM model, (Cheng et al., 2019) have analysed the changes in terrestrial hydrological cycle and discussed the future soil-moisture response and its drivers under a geoengineering scenario. To our knowledge, no study has yet investigated multiple indicators of water use with agricultural productivity after a short-duration event like Mt. Pinatubo eruption. Hence, this study looks to explicitly investigate changes in agricultural drought indices from Mt. Pinatubo by considering the store (soil moisture) and flux (evapotranspiration) of water as potential short-term controls over productivity in particular regions. We use NASA's state-of-the-art Earth system model, which incorporates prognostically evolving aerosols, to conduct simulation experiments following the counterfactual inference of causation approach for the Mt. Pinatubo eruption. We assess the impact of the Mt. Pinatubo eruption on the model-simulated climate via multiple pathways, ranging from primary dependent variables to higher-order responses that influence plant productivity. The use of prognostic aerosols enhances the simulations by capturing dynamically consistent feedbacks between the climate response and volcanic aerosols, including aerosol-radiation interactions and stratosphere-troposphere energy flux exchanges (McGraw and Polvani, 2024). Considering the complexity of modeling the terrestrial system, vegetation demographics, and physiological characteristics, we use the soil moisture and evapotranspiration-based agricultural drought indices SMDI (soil moisture deficit index) and ETDI (evapotranspiration deficit index) for short-term and long-term

scale to account for agricultural productivity. By focusing on soil moisture and evapotranspiration metrics, the major water-based drivers of plant productivity are explored to deepen our understanding of the Mt. Pinatubo impacts on plant productivity.

2.0 Method, Experiment, and Data

2.1 NASA GISS ModelE2.1 (MATRIX): We use the state-of-the-art Earth system model from the NASA (National Aeronautics and Space Administration) Goddard Institute for Space Studies, NASA GISS ModelE2.1 (Bauer et al., 2020; Kelley et al., 2020). NASA GISS ModelE2.1 has an atmospheric horizontal latitude-longitude grid spacing of 2.0x2.5 degrees (at the equator) with 40 vertical levels and a model top at 0.1 hPa. We used the interactive chemistry version MATRIX (Multiconfiguration Aerosol TRacker of mIXing state) aerosol microphysics module (Bauer et al., 2008, 2020), which is based on the Quadrature Method of Moment (QMOM) to predict aerosol particle number, mass, and size distribution for 16 different mixed modes of the aerosol population. New particle formation is represented by Vehkamäki et al. (2002), along with aerosol-phase chemistry, condensational growth, coagulation, and mixing states (Bauer et al., 2013). 16 mixing states with 51 aerosol tracers for sulfate, nitrate, ammonium, aerosol water, black carbon, organic carbon, sea salt, and mineral dust are resolved in this microphysical module (Bauer et al., 2008, 2020). The first indirect effect of aerosols in terms of changes in cloud properties through nucleation is also computed within MATRIX.

The model's ocean component (GISS Ocean v1) has a horizontal resolution of 1x1.25 degrees, with 40 vertical layers. The land component is the Ent Terrestrial Biosphere Model (TBM) (Kim et al., 2015; Kiang 2012) which includes an interactive carbon cycle (Ito et al., 2020), satellite-derived (MODIS-Moderate Resolution Imaging Spectroradiometer) plant functional type, and monthly variation of leaf area index (Gao et al., 2008; Myneni et al., 2002),

147 and tree height (Simard et al., 2011). Interannual variations in the vegetation properties are
148 controlled by rescaling the vegetation fractions (Figure S9) using historical crops and pasture at
149 the grid scale to account for land cover and land use changes (Ito et al., 2020; Miller et al., 2021).
150 The land model has two defined tiles for the soil layer: bare and vegetated, and each has six
151 vertical levels to a depth of 3.5 m (11.5 feet) (Rosenzweig and Abramopoulos, 1997). Rooting
152 depths for different plant functional types are also given by Rosenzweig and Abramopoulos
153 (1997), and more than 60% of roots for crop plant functional types are located within 0.6 m (1.96
154 feet) of soil depth. In this version of the model, the agricultural grid cells, crop plant functional
155 type, and crop calendar are prescribed according to McDermid et al., (2019). Irrigation in the
156 GISS ModelE is implemented using the water irrigation demand data (IWD) (Wisser et al., 2010)
157 and irrigation potential calculations based on (Wada et al., 2013) as discussed in (Cook et al.,
158 2020).

159 **2.2 Experiment Design:** The MATRIX version of GISS ModelE2.1 with active tracers is three
160 times more computationally expensive than the non-interactive (prescribed pre-calculated
161 aerosol concentration and extinction) version. We extended an equilibrated 1400-yearlong PI
162 control run with non-interactive tracers with an additional 500 years using the MATRIX version
163 with prognostic tracers before starting the ‘historical’ run. MATRIX includes the tropospheric
164 chemistry scheme that contains the inorganic (Ox NOx Hox and CO), organic chemistry of CH₄,
165 and higher hydrocarbons (Gery et al., 1989; Shindell, 2001; Shindell et al., 2003). The
166 stratospheric chemistry includes bromine, chlorine, and polar stratospheric clouds (Shindell et
167 al., 2006). Dust emission in the model is controlled by climate variables such as winds and soil
168 moisture at the spatial and temporal scales (Miller et al., 2006). However, anthropogenic dust is
169 not included in GISS ModelE2.1. Other anthropogenic emissions, including biomass burning

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(pre-1997 from (van Marle et al., 2017) and 1997 onwards from the GFED4s inventory (van der Werf et al., 2017)), are taken from the Community Emission Data System (CEDS) inventory (Hoesly et al., 2018). Most importantly, the volcanic SO₂ forcing for the ‘historical’ run (1850-1977) is the daily emission rate from VolcanEESM (Neely and Schmidt 2016 : <https://catalogue.ceda.ac.uk/uuid/a8a7e52b299a46c9b09d8e56b283d385>), and satellite measurement based SO₂ inventory (Carn et al., 2016) for 1978 to 2014 (extended up to 2021). The cumulative Mt. Pinatubo emission is 15194 kt (~15.2 Tg) of SO₂ injected from 13th to 16th of June 1991 above the Mt. Pinatubo vent, with a maximum of 15000 kt (15 Tg) emitted on June 15th at a plume height of 25 km Carn et al., (2016). The MATRIX version of the GISS ModelE2.1 used for all of our simulations predicts the nucleation, evolution, and removal of sulfate aerosols prognostically.

The model simulations we performed (Table 1) are described here. We started from the 1400-year-long preindustrial control run from CMIP6 (GISS-CMIP6-PI) with the prescribed average AOD historical period, which is further extended for 500 years using the GISS ModelE2.1 – MATRIX with prognostic tracers (GISS-PI). Then, the CMIP6 historical run (GISS-HIST-SO₂; 1850-2014) started with all forcings as specified by CMIP6 except the daily emission rate of injection of SO₂ (VolcanEESM) (Carn et al., 2016; Neely III and Schmidt, 2016). We branched out the experiment ensemble with Mt. Pinatubo eruption (GISS-PIN-SO₂) and the ensemble without Mt. Pinatubo (GISS-NOPIN-SO₂) from the historical (GISS-HIST-SO₂) using perturbed initial conditions (1st Jan 1986) from the year 1986 to 1999. The perturbation to the initial conditions is generated by altering the radiation-related random number generator that deals with fractional cloudiness in the column.

Table 1: Simulation experiment details.

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EXP Name	Description	Time period /run length	# of ensembles
GISS-CMIP6-PI	Preindustrial	1850 climatology /1400 years	1
GISS-PI	Preindustrial	1850 climatology /500 years	1
GISS-HIST-SO ₂	historical	1850-2014/165 years	1
GISS-PIN-SO ₂	historical	1986-1999/ 15 years	11*
GISS-NOPIN-SO ₂	historical	1986-1999/ 15 years	11*

*These ensemble members are branched out from the GISS-HIST-SO₂ by perturbing the initial conditions.

2.3 Methods: This study investigates the impacts of the Mt. Pinatubo eruption on the major

drivers of primary productivity, focusing on soil moisture and evapotranspiration related metrics.

We followed the counterfactual inference approach to draw causal inference of the Pinatubo

eruption. Hereafter, we use ‘PCH’ (Mt. Pinatubo and Cerro Hudson) to refer to the ‘GISS-PIN-

SO₂’ and ‘NP’ for the ‘GISS-NOPIN-SO₂’ ensembles. We have included the Cerro Hudson

eruption in both ensembles since we are focusing on the Mt. Pinatubo-driven climate response.

2.3.1 Statistical analysis for detecting Mt. Pinatubo-significant regions and anomalies

calculations.

We treat the no-Pinatubo ensemble (NP) as a counterfactual climate simulation and utilize it to perform the paired Student’s t-test for causal inference. The null hypothesis is that the ensemble

means of a quantity of interest (QoI) in a region over a time period are the same between

ensembles (i.e. $H_0: \bar{\mu}_{PCH} = \bar{\mu}_{NP}$). In the subsequent figures in this document, gray regions

indicate acceptance of the null hypothesis at the 95% confidence level, while the coloring

emphasizes the rejection of null hypothesis and significant regions of anomalies relative to 1950-

2014 climatology (see Supplementary information section S1.0). However, we also explored the

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223 ~~alternate approach of~~ directly comparing the difference between the two ensembles (PCH and
 224 NP) ~~for~~ presenting the Pinatubo effect (see Supplement Figure S2). ~~It is concluded that both of~~
 225 ~~the approaches~~ ~~led~~ to the same general conclusions, with only small quantitative differences.
 226 Nevertheless, we chose to remain consistent with the baseline requirements for other metrics,
 227 ~~and~~ used the historical climatology for ~~the same~~ period 1950-2014 as the baseline for the core of
 228 our analysis. ~~Thus,~~ the grey areas ~~indicate~~ no significant ~~differences~~ between the PCH and NP
 229 ensembles, while the ~~colored regions represents the statistically significant~~ anomalies, ~~calculated~~
 230 ~~as~~ PCH ensemble mean minus climatology.

231 2.4 Impact metrics

232 The distribution of incoming and outgoing radiation influences the hydrological cycle (Kiehl and
 233 Trenberth, 1997; Trenberth and Dai, 2007). A reduction of solar radiation at the surface has the
 234 potential to reduce rainfall and change the latent heat-dominated atmospheric heating pattern
 235 (Trenberth and Stepaniak, 2004). The perturbed atmospheric conditions and surface energy
 236 budget could affect soil moisture. Along with the surface air temperature and precipitation, we
 237 use soil moisture and surface energy budget-oriented drought indices (the soil moisture deficit
 238 index (SMDI) and evapotranspiration deficit index (ETDI)) to evaluate the land-atmosphere
 239 interaction and account for the potential drivers of the crop plant productivity in the model
 240 simulated post- Mt. Pinatubo environmental conditions (Narasimhan and Srinivasan, 2005).
 241 SMDI represents the land-based soil moisture state in selected depth horizons (i.e., SMDI_2
 242 means Soil Moisture Deficit Index for 2 feet (0.6 m) depth). ETDI represents the atmospheric
 243 conditions governing the land-atmosphere interaction and is also an indicator of plant health.
 244 Lastly, plant transpiration is analyzed to the explore the simulated physiological response to the
 245 volcanically induced hydroclimatic conditions. The Palmer drought severity index (PDSI) and

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other indices are commonly used to represent climatological drought conditions. However, we focus on SMDI and ETDI because these can capture short-term developing agricultural drought conditions that impact plant productivity and are free from the limitations of other metrics like PDSI. For example, SMDI and ETDI are seasonally independent measures and are comparable across space, even for different climatic zones.

SMDI and ETDI were calculated as described in Narasimhan and Srinivasan (2005) using model output at monthly and daily scales. Daily model output is resampled to a weekly scale to compute the indices. The weekly frequency is used because it is suitable for agricultural applications, and the daily frequency is comparatively higher, which makes these indices computationally expensive. Below, we reproduce the weekly calculation of SMDI and ETDI as presented in Narasimhan and Srinivasan (2005).

2.4.1 Soil Moisture Deficit Index (SMDI) : The soil moisture deficit index measures the wetness/dryness of soil moisture conditions in comparison to long-term records spanning 1950-2014.

$$SD_{i,j} = \frac{SW_{i,j} - MSW_j}{MSW_j - \min SW_j} \times 100 \quad \text{if } SW_{i,j} \leq MSW_j \quad \dots (2.4.1a)$$

And

$$SD_{i,j} = \frac{SW_{i,j} - MSW_j}{\max SW_j - MSW_j} \times 100 \quad \text{if } SW_{i,j} > MSW_j \quad \dots (2.4.1b)$$

$SD_{i,j}$ is the soil water deficit (%) for week j of the year i . $SW_{i,j}$ is the mean weekly soil water available in the soil profile (mm) for week j of the year i , MSW_j is the long-term (calibration period) median available water in the soil profile (mm) for week j , and $\min SW_j$ and $\max SW_j$ are the j^{th} weekly minimum and maximum of soil water available in the soil profile across the

279 calibration period (1950-2014). The soil moisture deficit index for any given week can be
280 calculated as

281
$$SMDI_j = 0.5 * SMDI_{j-1} + SD_j/50 \quad \dots (2.4.1c)$$

282 SMDI can be calculated for different soil depths; we used the 2, 4, and 6-feet depths for SMDI
283 estimation, approximately 0.6, 1.2, and 1.8 meters, respectively. For SMDI, it is typical to use
284 feet instead of meters in the literature, which is why we use the same units along with the
285 conventional SI units in brackets here. SMDI-4 means we considered the soil moisture content
286 between 2 (0.6 m) to 4 feet (1.2 m) in depth. Similarly, SMDI-6 indicates the soil moisture
287 content between 4 (1.2 m) to 6 feet (1.8 m) in depth.

288 **2.4.2 Evapotranspiration Deficit Index (ETDI):** The limitations of PDSI (Palmer, 1965) and
289 CMI (Palmer, 1968) in the formulation used for PET calculation (Thorntwaite, 1948), together
290 with the lack of consideration of the land cover type on water balance, have encouraged the
291 exploration of ETDI for agricultural productivity. Also, in the climate models, surface energy
292 fluxes are parameterized in terms of the thermodynamic gradient of atmosphere and land models
293 and thus represent the land-atmosphere interactions, which are not accounted for by these
294 atmosphere-only indices. We utilized model simulated surface energy fluxes (Latent and Sensible
295 heat) to calculate the potential (PET) and actual evapotranspiration (AET) to estimate the water
296 stress ratio. However, the applicability of the Penman-Monteith equation for reference crops
297 (Allen et al., 1998) provides a substitute method for PET calculation, which broadly produced
298 similar results (not shown).
299 In Equations 2.4.2a and 2.4.2b, we used the model simulated energy fluxes to calculate AET and
300 PET as suggested in (Milly and Dunne, 2016; Scheff and Frierson, 2015).

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308 The energy budget equation at the surface is given by $R_n = G + LH + SH$, where R_n is net
309 solar radiation, G is net ground heating (heat flux), LH and SH represent the Latent and Sensible
310 heat fluxes, respectively. We then use these to calculate PET and AET (unit as mm per day; 1
311 $Wm^{-2} = 0.0353$ mm/day):

$$312 \text{ PET} = 0.8(R_n - G) = (0.8 * 0.0353) * (LH + SH) \dots\dots (2.4.2a)$$

313 And

$$314 \text{ AET} = LH * (0.0353). \dots\dots (2.4.2b)$$

315 The evapotranspiration deficit index is estimated using the water stress condition using the actual
316 evapotranspiration (AET) and potential evapotranspiration (PET) per grid cell, as given below.

$$317 \text{ WS} = \frac{\text{PET} - \text{AET}}{\text{PET}} \dots\dots (2.4.2c)$$

318 WS range, from 0 to 1, where 0 signifies that evapotranspiration is happening at the potential rate
319 and 1 stand for no actual evapotranspiration. WS represents the water stress ratio on a monthly or
320 weekly basis (WS_j), which is further utilized to calculate water stress anomaly ($WSA_{i,j}$) for week
321 j of year i as given below.

$$322 \text{ WSA}_{i,j} = \frac{\text{MWS}_j - \text{WS}_{i,j}}{\text{MWS}_j - \text{minWS}_j} \times 100 \text{ if } \text{WS}_{i,j} \leq \text{MWS}_j \dots\dots (2.4.2d)$$

323 And

$$324 \text{ WSA}_{i,j} = \frac{\text{MWS}_j - \text{WS}_{i,j}}{\text{maxWS}_j - \text{MWS}_j} \times 100 \text{ if } \text{WS}_{i,j} > \text{MWS}_j \dots\dots (2.4.2e)$$

325 Here, MWS_j , minWS_j , and maxWS_j represent the long-term median, minimum, and maximum of
326 the water stress ratio over the calibration period. Water stress anomaly ranges between -100% to
327 100%, indicating very dry to wet conditions over the region.

328 Finally, drought severity is calculated as ETDI, similar to SMDI (equation 2.4.1c) at a
329 monthly/weekly time scale.

330 $ETDI_j = 0.5 * ETDI_{j-1} + WSA_j / 50$ (2.4.2f)

331 The indices SMDI and ETDI range from -4 to +4, representing the excessive wet and dry
332 conditions. The bounding values -4 or +4 represent extremely dry/wet conditions as the
333 deficit/excess of soil-moisture deficit (SM) or water stress anomaly (WSA) reached relative to
334 the maximum over the reference calibration period.

335 We also highlight the justification for selecting 1950-2014 as the base period for
336 analyzing the response in climate variables and the long-term calibration period for drought
337 indices calculations. (Supplementary information section S1).

338 **3.0 Results**

339 The result section of this study first presents the NASA GISS model's simulated
340 properties of the 1991 Mt. Pinatubo eruption, and then further the evaluation of the primary
341 (aerosol optical depth, radiation, and temperature) and secondary (precipitation, soil moisture,
342 evapotranspiration, and transpiration) impacts on plant productivity.

343 **3.1 Radiative forcing response**

344 In the Mt. Pinatubo eruption simulation, the volcanically injected SO_2 in the stratosphere
345 oxidizes in the presence of prognostically evolving OH radicals to form the stratospheric sulfate
346 aerosols (Supplementary section S2.0). GISS ModelE (MATRIX) PCH also simulated a peak
347 global mean aerosol optical depth (AOD; for 550 nm wavelength) of 0.22 (Supplementary Fig
348 S4 bottom panel) a few months after the eruption, which then decreases with time due to the
349 deposition of volcanic aerosols (English et al., 2013; Sato et al., 1993). In this study, the model-
350 simulated aerosol optical depth (AOD) due to volcanic aerosol and radiative forcing is larger
351 than the previously reported AOD of 0.15 and forcing of -4.0 to -5.0 Wm^{-2} due to the Mt.
352 Pinatubo eruption (Hansen et al., 1992; Lacis et al., 1992). The mass and size of volcanic sulfate

aerosol firmly control the scattering of the incoming shortwave radiation and the absorption of longwave (Kinne et al., 1992; Lacis, 2015; Lacis et al., 1992; Lacis and Hansen, 1974). The first-order climate response to the volcanically-injected sulfate aerosol in the stratosphere is the perturbation of the radiative balance of the Earth system (Brown et al., 2024; Hansen et al., 1980; Lacis et al., 1992; Stenchikov et al., 1998). Supplementary figure S3 shows that the GISS ModelE PCH has simulated a peak longwave, shortwave, and net radiative response of +3.0 Wm⁻², -8.0 Wm⁻², and -5.0 Wm⁻² respectively, a few months after the eruption, which recovers slowly in next 24 months and is consistent with previous studies (Stenchikov et al., 1998; Hansen et al., 1992; Minnis et al., 1993; Brown et al., 2024). The GISS model also simulated a smaller peak ranging within -1 Wm⁻² in the runs without Mt. Pinatubo eruption (NP), likely due to the Cerro Hudson eruption in August 1991.

3.2 Aerosol dispersion and temperature response

Figure 1 shows the zonal mean anomaly for the aerosol optical depth (AOD) and surface air temperature. The zonal AOD shows the dispersion and transport of aerosol poleward after the eruption. Horizontal dispersion and transport of the aerosols are strictly influenced by the stratospheric meteorology and atmospheric circulation, which is independent in each ensemble member, and depends on the plume height and season. GISS ModelE has simulated AOD consistently with previous studies (Aquila et al., 2012; Rogers et al., 1998; Timmreck et al., 1999; Trepte et al., 1993). Cross-equatorial dispersion to the southern hemisphere might be due to the more robust Brewer-Dobson circulation in the austral winter (Aquila et al., 2012). Meanwhile, the phases of QBO (Quasi-Biennial Oscillation) and local heating also play a crucial role in the poleward and vertical dispersion of stratospheric aerosols (Hitchman et al., 1994). A

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378 smaller peak in the southern hemisphere (45° S) in the late 1991, likely due to the Cerro Hudson
379 eruption, which injected ~1.5 Tg of SO₂ at a height of 15 km.

380 The lower panels in Figure 1 show the surface air temperature response due to the net
381 radiative perturbation, which is dominated by the scattering of incoming solar radiation. The
382 zonal structure of surface temperature shows that the surface cooling follows the aerosol optical
383 depth pattern, and the most significant cooling is simulated in the northern hemisphere high
384 latitudes. Temporal characteristics of lower stratosphere warming, and surface cooling also show
385 the seasonal variations in incoming solar radiation in northern polar latitudes. Model simulated
386 the lower stratospheric warming due to longwave radiation absorption in lower stratosphere is,
387 consistent with previous studies (Supplementary text section S3.0).

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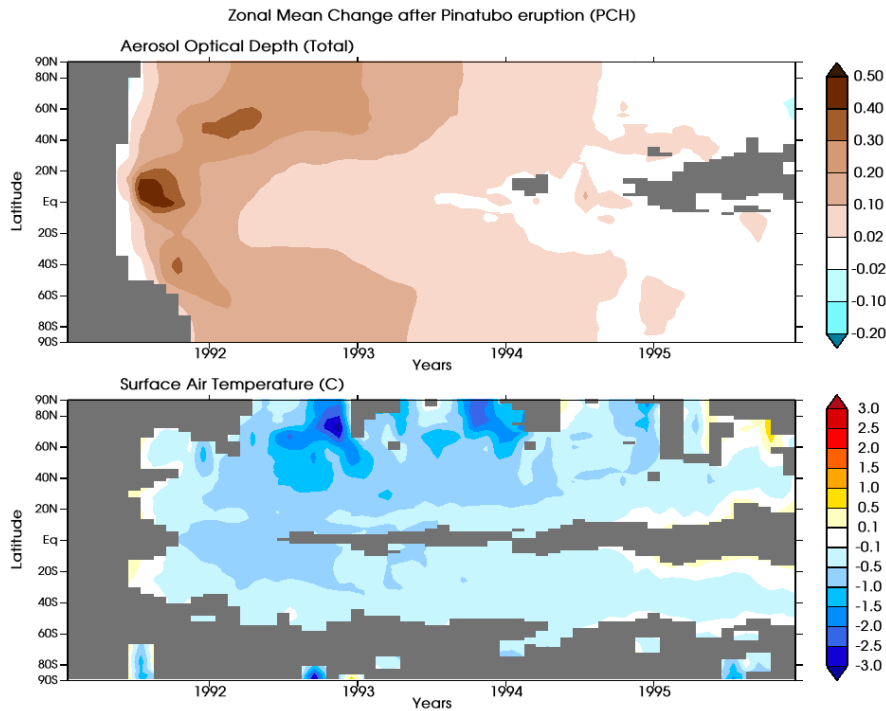


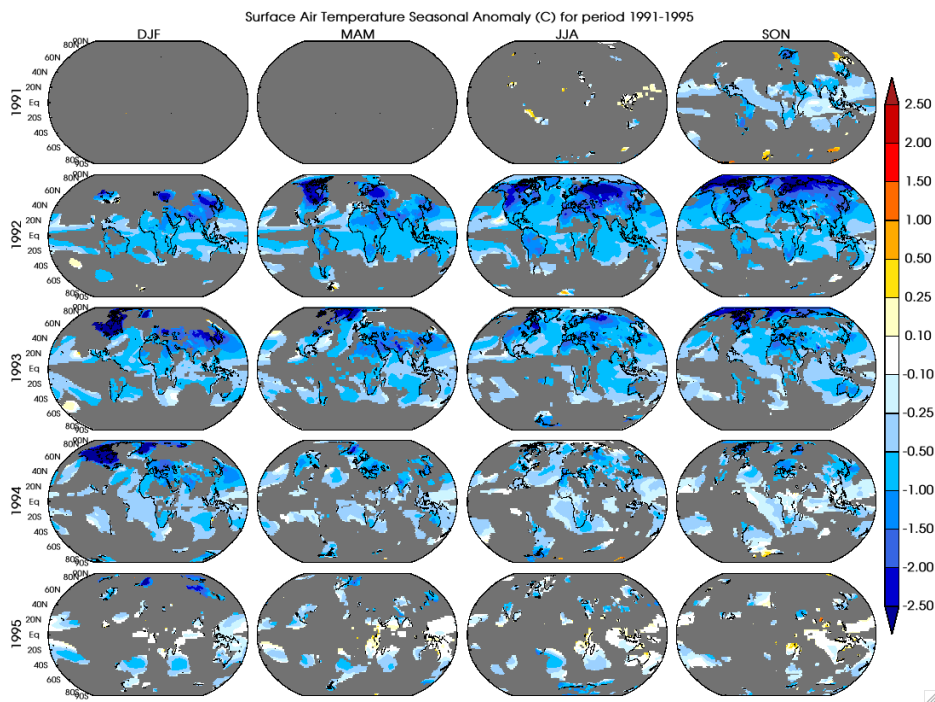
Figure 1. Zonal mean of monthly anomalies for multi-ensemble means for aerosol optical depth at 550 nm (upper panel) and surface air temperature (lower panel) with respect to the 1950-2014 climatology. Gray regions show no statistically significant difference between the PCH and NP responses. The colored areas show anomalies of PCH with respect to the climatology from 1950-2014.

The spatial pattern of surface air temperature response is evaluated at the seasonal scale for each year from 1991 to 1995, as shown in Figure 2. We conclude that the volcanic forcing from the Mt. Pinatubo eruption results in a detectable seasonal mean surface air temperature response. Figure 2 shows that a spatial pattern of surface cooling starts appearing after a few months of the eruption (during the SON season of 1991) when the gaseous SO_2 is oxidized into

402 sulfate aerosols. The surface cooling signature due to the volcanic aerosols was significant in
403 1992 and 1993 before recovering in 1994 towards pre-eruption temperature conditions. The
404 highest surface cooling is noticed over the sub-tropics and higher latitude land regions in the
405 northern hemisphere and reaches up to 2.5 °C at a regional scale. To summarize: the PCH GISS
406 ModelE simulated global mean peak cooling response is ~0.5 °C after the eruption, as shown in
407 Supplement Figure S4 (Middle panel), with a range between 0.25 – 1.0 °C for individual
408 ensemble members, and this is consistent with the various observation and modeling studies
409 (Dutton and Christy, 1992; Hansen et al., 1996; Kirchner et al., 1999; Minnis et al., 1993; Parker
410 et al., 1996; Ramachandran et al., 2000; Stenchikov et al., 1998).

411

412
413



414 Figure 2: Seasonal mean surface temperature anomalies (°C) from the year 1991 to 1995 with
415 respect to the reference period of 1950-2014. A grey color is painted over the grid cells where
416 the surface temperature anomalies are not statistically significant in comparison to the ~~NP~~
417 ensemble. The colored areas show anomalies of PCH with respect to the climatology from 1950-
418 2014.

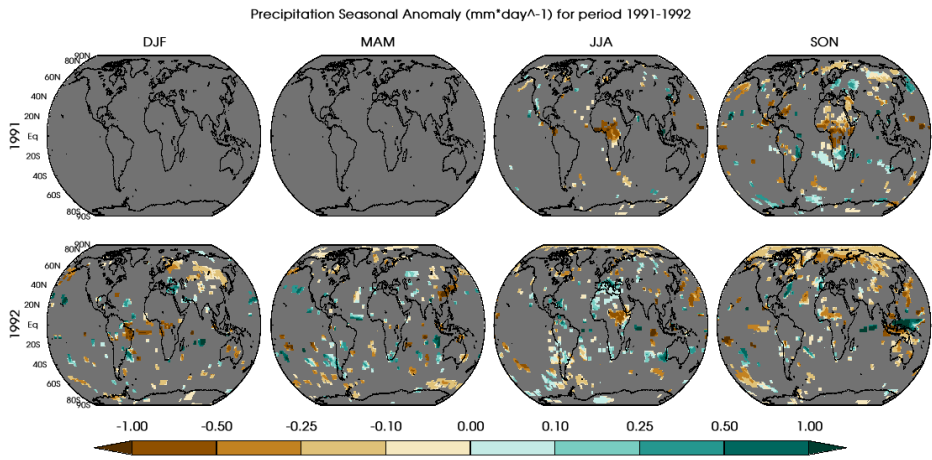
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419 420 3.3 Rainfall Response

421
422 Precipitation, presented seasonally for the year of eruption (1991) and the following year
423 (1992) in Figure 3, shows a highly complex and variable response to the volcanically induced
424 tropospheric cooling and radiative balance perturbation because of its sensitivity to the other
425 climate system components. Studies have shown that global mean precipitation decreases after
426 large volcanic eruptions (Gu et al., 2007; Gu and Adler, 2012; Iles et al., 2013; Robock and Liu,
427 1994; Singh et al., 2023; Trenberth and Dai, 2007). Colose et al., (2016) have postulated that the
428 asymmetrical surface cooling and radiative balance perturbation create an energetic deficit in the
429 hemisphere of eruption that constrains the poleward propagation of tropical rainfall belt (ITCZ)
430 in that hemisphere. In the case of the Mt. Pinatubo eruption, the PCH simulations show that
431 regional patches of significant decrease of up to 1 mm per day are noticed over tropical and
432 northern hemispheres (Africa, eastern and northern Asia) after the eruption (Figure 3). Also,
433 increasing rainfall patterns are simulated over the Mediterranean and European regions. Broadly,
434 the confidence level of precipitation response due to volcanic aerosols is strongly influenced by
435 the uncertainty due to many possible factors and prominent modes of atmospheric variability,
436 such as the strength of El Nino (Paik et al., 2020).

437 The zonal mean of the rainfall response (Figure S5) shows a clear decreasing trend in the
438 northern hemisphere tropical and higher latitudes with a positive rainfall response band around
439 20° N. The PCH modelled rainfall response due to the Mt. Pinatubo eruption is broadly

441 consistent with the previous studies (Joseph and Zeng, 2011; Liu et al., 2016; Trenberth and Dai,
 442 2007), but the uncertainty in rainfall response is still high. Although we use statistical
 443 significance at 95% confidence level as our metric for determining significant anomalies and
 444 only a few regions exhibit a spatially coherent and detectable forced response (Figure 3). We
 445 acknowledge the signals due to the model's internal variability when averaging the impacts
 446 across multiple ensembles, but 11 ensembles are a good compromise between few vs. many
 447 ensemble members which was shown to be sufficient to represent significance in climate
 448 response at the regional scale (Polvani et al., 2019; Singh and AchutaRao, 2019). The
 449 inconsistency and complexity in the precipitation response drive us to explore the compound
 450 hydroclimatic pathways that lead to drought risk.



451
 452
 453 Figure 3. Seasonal mean precipitation anomalies (mm per day) from the year 1991 and 1992 with
 454 respect to the reference period of 1950-2014. A grey color is painted over the grid cells where the
 455 precipitation anomalies are not statistically significant in comparison to the **NP** ensemble. The
 456 colored areas show anomalies of PCH with respect to the climatology from 1950-2014.

457
 458 **3.4 Drought Conditions**

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Land-atmosphere interactions under a radiatively perturbed environment are crucial in regulating the climate response at regional and sub-regional scales. On short timescales, changes in these interactions can strongly affect plant productivity. Even short-lived adverse conditions in the growth cycle have the potential for outsized impacts, especially if they happen at a particular time in the growing cycle. Hence, we explore the weekly aspects of these drought conditions in Section 3.6 to explore the temporal characteristics of variability in the conditions.

3.4.1 Seasonal Soil Moisture Drought Index (SMDI)

The root zone is commonly defined as the top 3 – 6 feet (0.9-1.8 m) of the soil column (Keshavarz et al., 2014) but most agricultural crops have shallower root systems confined to the top 2 feet (0.6 m) (Narasimhan and Srinivasan 2005). Hence, we focus on the soil moisture deficit index (SMDI) (Narasimhan and Srinivasan 2005) for the top 2 feet (0.6 m) of ground depth (SMDI_2) as shown in Figure 4. As anticipated, more land area is covered by statistically different SMDI_2 than in Figure 3. This enhances our analysis of water-driven impacts on plant productivity, offering greater insights than precipitation.

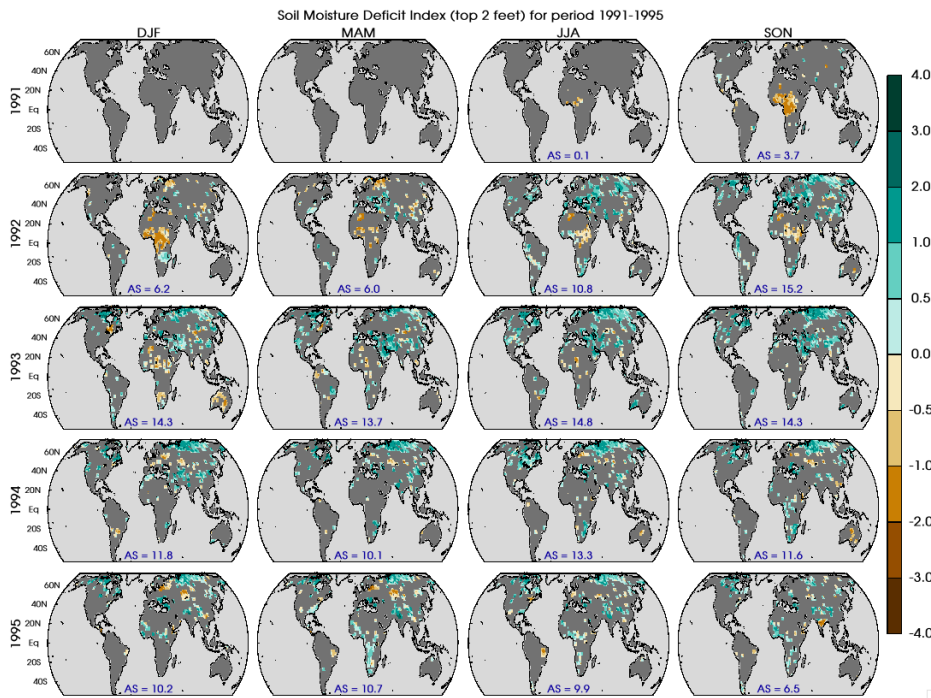


Figure 4. Soil moisture deficit index (SMDI_2) for the top 2 (0.6 m) of ground depth evaluated seasonally from 1991 to 1995. Grey color is painted over the grid cell where the SMDI_2 is not statistically significant in contrast to **NP** ensemble. The parameter AS on each panel marks the percentage of land area that has shown statistically significant dry or wet response after Mt. Pinatubo eruption.

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Figure 4 clearly shows that the equatorial region, especially over Africa, has a significant drying response due to Mt. Pinatubo compared to long-term historical data from the SON season of 1991 through the following DJF season. Although less robust, the dryness in this region lasted from 1991 through the following DJF season. The severity of drying response reaches up to -2.0 on a scale of extreme wet/dry at 4.0/-4.0, where the severity of -4.0 reflects the maximum dryness (rarest case) over the entire 1950-2014 calibration period. Figure 3 shows a similar pattern in equatorial African

489 rainfall decrease. A decrease in rainfall was present in the first season post-Mt. Pinatubo
490 eruption, indicating an expected lagged response in SMDI_2. Spatial coherence between these
491 signals is again re-established in the JJA and SON 1992 seasons, albeit with more variation in the
492 strength of the signal.

493 Meanwhile, in the high latitudes of the northern hemisphere, we see an increase in the
494 store of soil moisture despite a decrease in rainfall in higher latitudes. An exception is the
495 Mediterranean (extending towards the east Mediterranean and western Asia) region, where soil
496 moisture and rainfall both show an increase after the Mt. Pinatubo eruption. This increase in soil
497 moisture in the northern hemisphere is comparatively more pronounced in the summer months
498 than in the winter season. Thus, despite less water supply through rainfall, there has been a
499 persistent increase in the soil moisture in the root zone layer since the JJA season of 1992. This is
500 likely due to less water extracted from this layer through evaporation and transpiration as well as
501 due to the implemented irrigation in GISS modelE (details in further sections). Overall, Figure 4
502 shows equatorial drying signals mostly dominated through the DJF season of 1993, but the wet
503 conditions over higher latitudes lasted till 1995. Broadly, 6-13% of the land region has shown a
504 statistically significance response in terms of dry/wet conditions by the end of year 1995 because
505 of the Mt. Pinatubo forcing.

506 Deeper soil layers better approximate longer-term meteorologically defined drought
507 indices (Narasimhan and Srinivasan 2005). This makes intuitive sense: precipitation provides
508 the recharge for the store of soil moisture, and if there is a longer-term decline in precipitation,
509 all available moisture will be used hydraulically for plant transpiration (both the deeper stores of
510 water and the soil-penetrating precipitation available), not allowing for deeper depth
511 recharge. Here, we evaluate discrete layer depths instead of cumulative depths for two

512 reasons. First, the soil permeability changes with the depth, and the inclusion of top layers
513 erroneously reflects the SMDI_2 signal in potentially impermeable regions; second, the SMDI_2
514 signal gets superposed over the deeper layer response and misleads the actual soil moisture
515 response for the deeper layers.

516 As expected, when we evaluate the soil moisture deficit response between 2-4 feet (0.6-
517 1.2 m) soil depth (SMDI_4) in Figure 5 and 4-6-feet (1.2-1.8 m) soil depth (SMDI_6) in Figure
518 S6, we see similar spatial and temporal distributions as shown in Figure 4 with a corresponding
519 decrease in the percentage of area response. Spatially, we see high latitudes across North
520 America and Northeastern and western Asia, equatorial Africa, Europe, and Mediterranean
521 regions maintain their SMDI-2 trend in Figure 4. However, the total response area decreases
522 from peak coverages of 12-13% in SMDI-2 to less than 10-12% in SMDI-4 and 7-10% in
523 SMDI_6 (as shown in Figure S6). Additional decreases in the degree of impact are also seen
524 between the three soil layers. Note that the light grey colored regions in Figures 5 and S6
525 represent regions of impermeability, which does affect the total area of response.

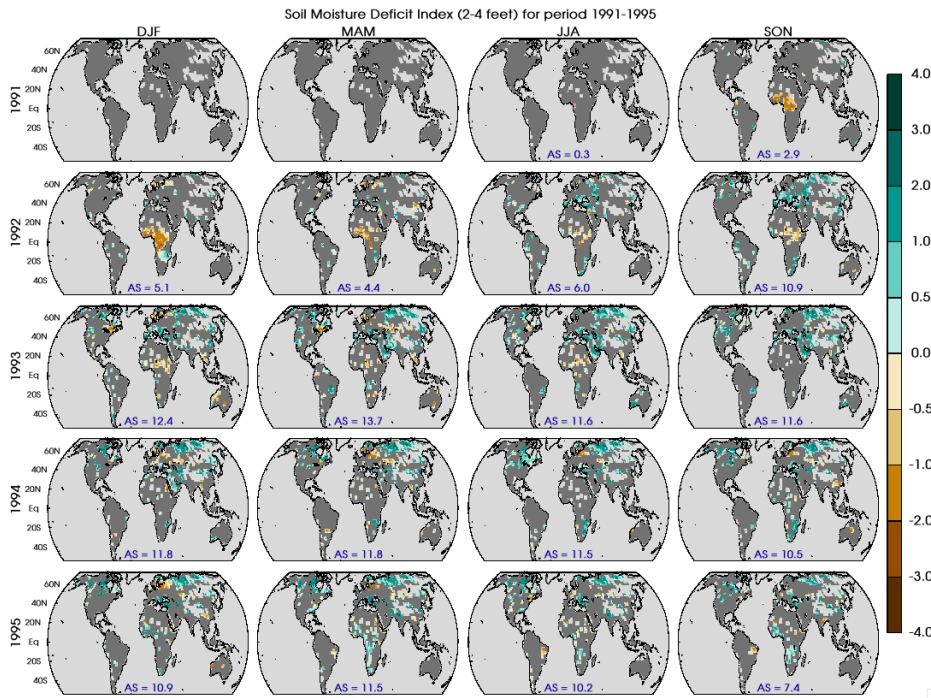


Figure 5. Soil moisture deficit index (SMDI_4) for soil depths between 2-4 feet (0.6-1.2 meter) evaluated seasonally from 1991 to 1995. The grey color is painted over the grid cell where the SMDI_4 is not statistically significant in contrast to the **NP** ensemble. The light grey colored regions represent regions of impermeability. The parameter AS on each panel marks the percentage of land area showing statistically significant dry or wet response after the Mt. Pinatubo eruption.

3.4.2 Seasonal Evapotranspiration Deficit Index (ETDI)

As indicated in the methodology section, ETDI calculation is similar to SMDI but is based on the water stress anomaly, which accounts for the difference between actual and potential evapotranspiration. ETDI is a measure of the flux of water between land and atmosphere, and like SMDI_2 in Figure 5, it shows robust statistical differences over land.

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Figure 6 illustrates the reduction in ETDI response in the equatorial region started developing during the DJF season of the year 1992, and these conditions persisted over the year. Like SMDI_2, ETDI increases over the regions encompassing the Mediterranean and western Asia. However, ETDI differs from the SMDI_2 over some northern hemisphere regions, especially Northeastern Asia. A drying response in terms of ETDI in the northern hemisphere regions persisted during 1993 and 1994, whereas SMDI_2 shows an opposite response. This contrasting response through the ETDI and SMDI_2 points to the complexity of land-atmosphere interactions over these regions. We utilized model simulated surface energy fluxes (Latent and Sensible heat) to calculate the potential (PET) and actual evapotranspiration (AET) to estimate the water stress ratio. In these regions, soil moisture is available during the summer and early winter months. Still, an evapotranspiration deficit reflects the decrease in plant transpiration (latent heat flux), which may be due to the unavailability of plants. Also, the surface temperature (sensible heat flux) response supports the non-water-stressed atmospheric conditions, and thus, overall, it shows an evapotranspiration deficit. Areas of significant response in terms of ETDI vary from 7 to 14.5% on a seasonal basis during the years following the eruption. The most significant areas of ETDI coverage occur during the same periods as SMDI_2 (between JJA 1992 – JJA 1993).

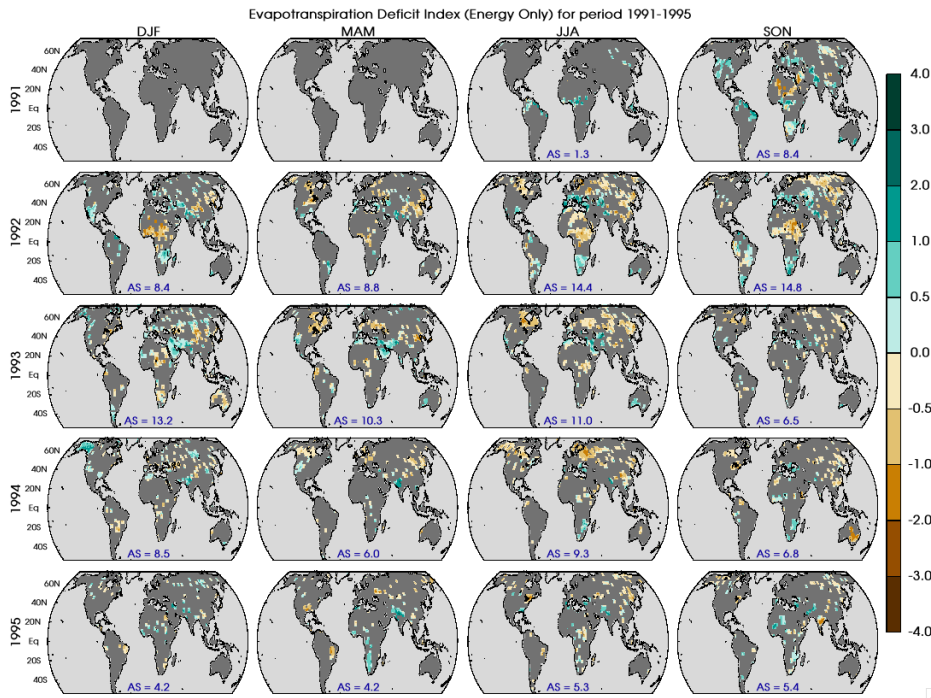


Figure 6. Evapotranspiration deficit index (ETDI) at seasonal scale from 1991 to 1995. The grey color is painted over the grid cell where the ETDI is not statistically significant compared to the ~~NP~~ ensemble. The parameter AS on each panel marks the percentage of land area that has shown statistically significant dry or wet response after Mt. Pinatubo eruption.

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3.5 Seasonal plant productivity inferences

SMDI (at depths of 0-2 (0.6m), 2-4 (0.6m-1.2m), and 4-6 feet (1.2m -1.6m)) and ETDI have proven helpful in analyzing the climatic impact of the Mt. Pinatubo eruption on a seasonal scale. Additionally, SMDI_2 (top 2 feet or 0.6 m) and ETDI have demonstrated a slow development of drought conditions, beginning by the end of the year 1991 (SON season), reflecting a time lag between seasonal precipitation patterns (Narasimhan and Srinivasan 2005). Crucially, the seasonal depiction of drying/wet conditions via SMDI and ETDI provides a comprehensive

571 overview of prolonged or recurrent dry/wet conditions in susceptible regions. Moreover,
572 understanding these typical agricultural drought indices indicates potential effects on plant
573 productivity at the seasonal scale.

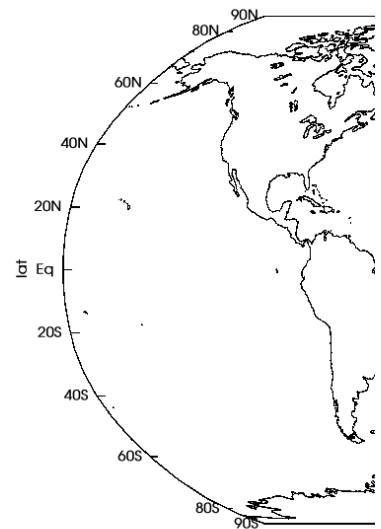
574 In equatorial Africa, decreases in both SMDI and ETDI indicated that there was likely a
575 negative impact on plant productivity. The Mediterranean region (encompassing the eastern
576 Mediterranean and western Asian region) showed increased SMDI and ETDI, indicating a
577 positive effect on plant productivity. Northern Asia, on the other hand, exhibited an increase in
578 SMDI with a decrease in ETDI, suggesting that plant productivity likely decreased, but not
579 because of water-based drivers.

580 **3.6 High frequency impact pathways evaluation**

581 Here, we use the daily model output to calculate weekly drought indices in each grid cell. These
582 weekly scale drought indices and changes in other atmospheric variables are explored at the
583 regional scale to understand the associated land-atmosphere interactions pertaining to higher-
584 order impacts. The higher temporal resolution of these parameters is crucial for analyzing
585 different stages of the crop cycle in a region. Considering the complexity of the representation of
586 spatial features, we selected three distinct regions (shown in Figure S7 and detailed in caption) in
587 the northern hemisphere based on the climate response to Mt. Pinatubo in the seasonal analyses
588 presented in Section 3.0. We followed the same strategy described in Section 2.3.1 to mask out
589 the statistically insignificant grid cells using the **NP** ensemble after creating the weekly time
590 series for different drought indices and atmospheric parameters.

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Figure 7. Spatially averaged drought indices (SMDI_2 & ETDI) and anomalies for other drivers (Actual and Potential Evapotranspiration, Transpiration, Moisture Source (Precipitation plus irrigation) and Surface Temperature) at weekly scale for the equatorial Africa region (Latitude = 5°-15° N, Longitude = 15° W- 40° E).

3.6.1 Equatorial Africa

Figure 7 shows the weekly response to volcanic forcing for the years 1991-1995 in terms of agricultural drought indices (SMDI_2 & ETDI), AET, PET, transpiration, total soil moisture source, and surface temperature for an equatorial region in northern Africa. This region exhibits consistent statistical differences across the drivers on a weekly scale; thus, for most of the time period, this is unmasked, revealing the degree of anomalous conditions.

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616 This region lies between the latitude 5-15° N, where the precipitation during the monsoon
617 season shows a decrease in response to a southern migration of the inter-tropical convergence
618 zone (Iles et al., 2013; Colose et al., 2016; Singh et al., 2023). Weekly precipitation change in the
619 equatorial Africa shows a significant deficit of more than 1.5 mm per day consistently for several
620 weeks, especially during the JJAS monsoon season. This region also shows that a deficit in
621 precipitation during the major precipitation season (JJAS) can result in a soil moisture deficit in
622 the root zone in the following seasons (DJF and MAM in SMDI_2) and consequently affect the
623 entire crop cycle. The root zone soil moisture, SMDI_2, also shows a persistent drying through
624 1993, and combined with the lack of precipitation, the potential for recharge is limited. Also, this
625 region has no contribution from irrigation as a source of additional soil moisture, as shown in
626 Figure S8 (bottom panel) (Cook et al., 2020). Cumulative annual rainfall change over this region
627 shows a deficit of 33.2, 9.5, and 3.2 mm per day for the years 1991, 1992, and 1993 and an
628 increase of 10.5 and 13.6 mm per day for the years 1994 and 1995, respectively, where soil-
629 moisture response shows a recovery from the dry conditions.

630 Hence, it is unsurprising that a corresponding decrease in ETDI through 1993 is
631 consistent with this lack of moisture. However, the evaporative demand, as shown by surface
632 temperature change, does not consistently decrease until September of 1991, and hence, ETDI is
633 slow to show a response in the deficit index. After that point, evaporative demand decreases with
634 lower temperatures, but the evapotranspiration is dominated by transpiration (Seneviratne et al
635 2010; Nilson and Assmann 2007), and so the majority of the decrease in ETDI is explained by
636 the shown decrease in plant transpiration. As expected, this decrease in plant transpiration is well
637 correlated with decreases in AET. Conclusively, precipitation response in this region shows
638 dominance in regulating the ecohydrological conditions. A substantial decrease in the weekly

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rainfall over the region perpetuates a root-zone water deficit, resulting in decreased plant transpiration. Decreases in both SMDI_2 and ETDI thus indicate developing agricultural drought conditions, which are confirmed by a decrease in the direct measure of plant transpiration.

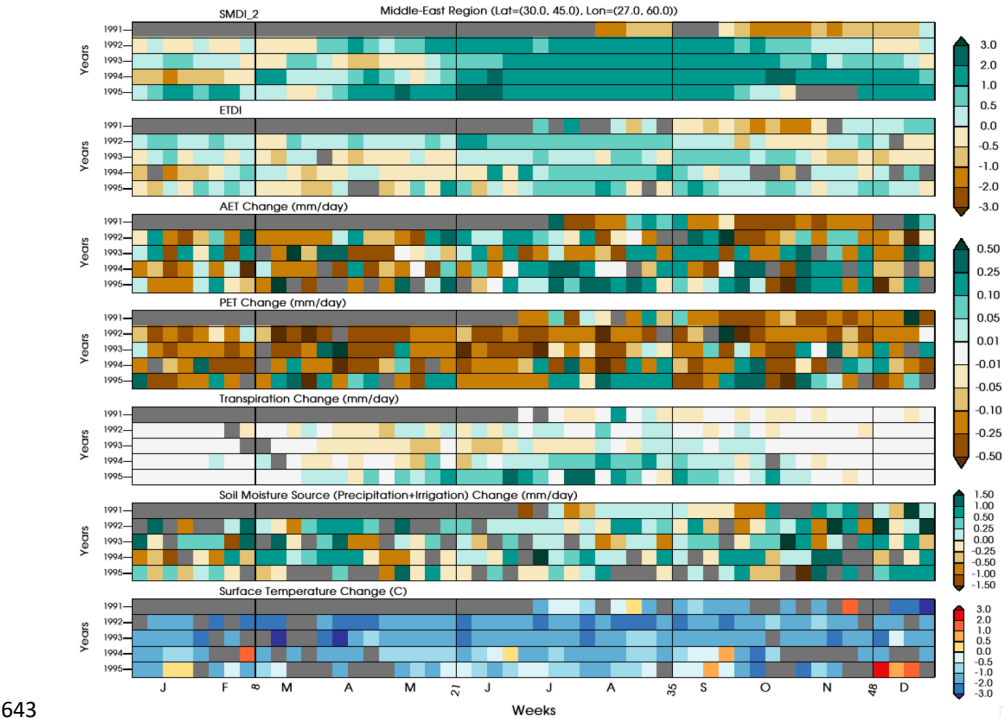


Figure 8. Spatially averaged drought indices (SMDI_2 & ETDI) and anomalies for other drivers (Actual and Potential Evapotranspiration, Transpiration, Moisture Source (Precipitation plus irrigation) and Surface Temperature) at weekly scale for Middle East (MDE) Region (Latitude =30° N - 45° N, Longitude=27°-60° E; Eastern Mediterranean /Western Asian).

3.6.2 Middle East Region

Figure 8 shows the region covering the eastern Mediterranean and the western Asian regions where rainfall slightly increased after the Mt. Pinatubo eruption. Additionally, this region exhibits a significantly positive trend in irrigation practices post-1950, with a substantial peak

655 over the eastern Mediterranean region following the Mt. Pinatubo eruption (Cook et al., 2020;
656 Figure 1 and 2).

657 In the eastern Mediterranean, wet and cold autumns and winters persist for several years
658 after the Mt. Pinatubo eruption, offering significant root zone recharge potential. The summer
659 months, in general, reflect a slightly uncertain model response in the regional rainfall, with some
660 weeks of deficit and some excess, but an additional water supply through irrigation contributes to
661 the overall moisture content in the region (Figure S8). Root zone soil moisture (SMDI_2)
662 indicates sufficient water availability throughout the growing seasons during the entire analysis
663 period. Overall, the findings demonstrate that this region is not moisture-limited, with adequate
664 precipitation and irrigation ensuring the replenishment of root zone moisture as plants grow.
665 Cumulative weekly anomalies show that precipitation change in 1991 is slightly negative (-0.5
666 mm per day), but an increase in annual rainfall of 13.8, 8.0, 10.9, and 4.5 mm per day is
667 simulated for the years 1992, 1993, 1994, and 1995, respectively. Irrigation implemented across
668 this Middle East region exhibits a strong positive trend from 1950 to 2005 (Cook et al., 2020).
669 Notably, a significant cumulative increase in irrigation of 0.5, 1.3, 1.3, 0.8, and 0.9 mm per day
670 during the years 1991 to 1995 provides an additional source of moisture supply for the region
671 (Cook et al., 2020). This irrigation is particularly crucial during the summer months,
672 compensating for rainfall deficits lasting several weeks and contributing 10–20% of soil moisture
673 source changes (Figure S8).

674 The corresponding increases in ETDI and AET indicate the abundant water availability
675 for transpiration in the region. While transpiration remains temporally correlated with AET, the
676 increases are less pronounced. Simultaneously, there is a decline in PET response, which is
677 associated with a more significant temperature decrease in this region compared to equatorial

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680 Africa. The reduction in PET, combined with the increase or maintenance of AET through
681 transpiration, leads to an increase in ETDI. Consequently, agricultural productivity in this region
682 is generally positively impacted, as ample moisture provides essential support. Nevertheless, the
683 data reveals heterogeneous patterns, indicating that 1993 may have influenced plant productivity
684 with a positive but lower-magnitude ETDI, inconsistent AET, and reduced transpiration.

685 Regardless of the presence of a volcanically induced response, the weekly scale analysis
686 demonstrates its importance by virtue of an example from 1993, where rainfall deficit is
687 produced during the 15th and 16th weeks (April) of the year. Combined with lower SMDI_2, this
688 may lead to insufficient moisture availability during a critical stage of the crop cycle. The
689 duration of such moisture deficits could significantly impact overall seasonal crop production.
690 Consequently, even if favorable conditions prevail for most of the crop cycle, adverse effects
691 during essential phases can critically influence yields in ways that seasonal averages fail to
692 capture.

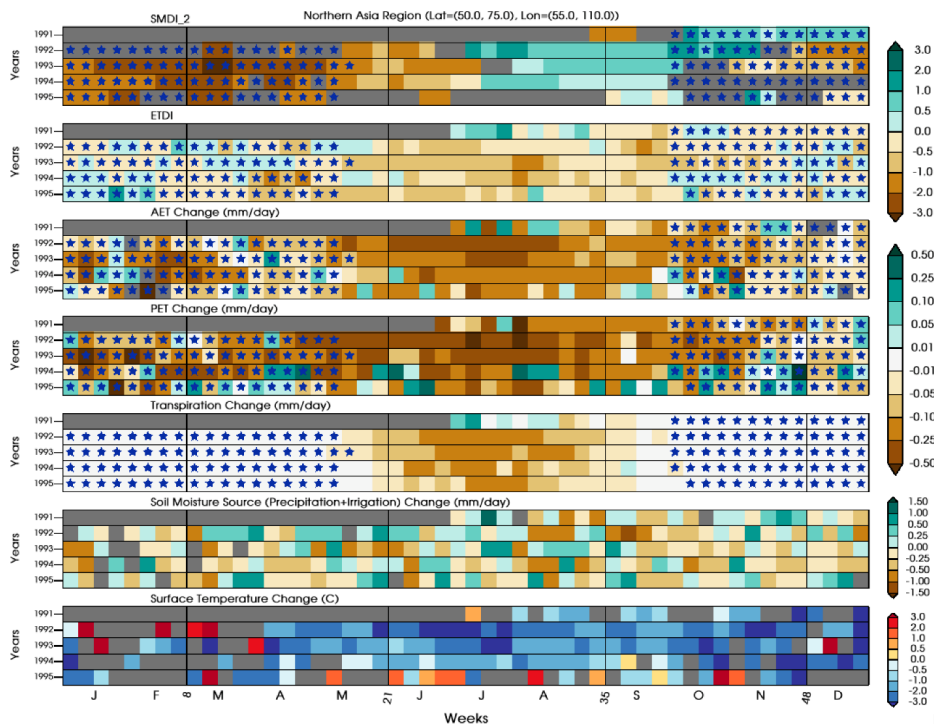


Figure 9. Spatially averaged drought indices (SMDI 2 & ETDI) and anomalies for other drivers (Actual and Potential Evapotranspiration, Transpiration, Moisture Source (Precipitation plus irrigation) at weekly scale Northern Asia Region (Latitude =50° N - 75° N, Longitude=55°- 110° E). Blue stars represent the weeks with average surface temperature below freezing point.

3.6.3 Northern Asia

Finally, we selected a region (NAS; Figure 9) in higher latitudes to explore the interplay between the various drivers governing the conditions for plant productivity. This region consistently exhibits statistically significant differences across drivers on a weekly scale. However, higher latitudes experience strong seasonal controls on plant productivity, with below-freezing temperatures (indicated by blue stars) halting growth. Therefore, our analysis focuses on months when plant growth is possible (~MJJAS).

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Precipitation changes over the entire analysis period are highly uncertain. However, there is generally a slight trend toward increased precipitation from November 1991 to June 1992, followed by a decline through 1994. As shown in Figure S8, the contribution of irrigation to soil moisture in this region is negligible. Cumulative weekly precipitation anomalies indicate an annual increase of 0.05 mm per day in 1991, followed by decreases of -2.0, -4.1, -6.8, and -1.0 mm per day from 1992 to 1995, respectively. Alternatively, root zone moisture shows that sufficient water available to plants during the JAS (July-August-September) growing months after a substantial deficit in the early MJ (May-June) months. Certainly, during the summer months, the melting of frozen surfaces and snow supplies moisture to the upper layers, resulting in wet conditions, which accounts for this strong dichotomy. However, there were no corresponding increases in ETDI and AET following the 1991 season. This indicates that despite ample water, plants are still not growing; this is conclusively confirmed by the decrease in transpiration starting in 1992. Meanwhile, the simultaneous reduction of PET response is correlated with the most substantial decline in temperature, on the order of 2-3° C, for this region.

Unlike the other two regions for which SMDI_2 and ETDI exhibited similar wet/dry patterns, this region shows diverging patterns. Overall, this indicates that although moisture is available to support plant productivity, it is not being effectively utilized. Therefore, other factors must be responsible for the decrease in plant transpiration and ETDI. The more significant decline in PET compared to AET suggests that temperature is playing a role. Since temperature is directly related to reduced incident radiation, the combined effects of temperature and radiation are likely the primary factors controlling the reduced plant productivity in this region, rather than moisture conditions.

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732 4.0 Conclusions

733 This study utilized the NASA GISS ModelE2.1 (MATRIX) Earth system modeling
734 framework to investigate the mechanisms by which the 1991 Mt. Pinatubo eruption influenced
735 hydroclimatic conditions and water-based drivers of plant productivity. The simulation
736 successfully reproduced key microphysical properties ($R_{eff} \approx 0.5 \mu\text{m}$, $\text{AOD} \approx 0.22$), radiative
737 forcing ($\sim -5 \text{ Wm}^{-2}$), surface cooling ($\sim 0.5^\circ\text{C}$), and regional rainfall changes consistent with
738 previous studies. The study further examined secondary impacts of the eruption, beyond changes
739 in radiation and temperature, by analyzing agricultural drought indices to better assess its effects
740 on plant productivity. Two metrics, SMDI and ETDI, which account for land-atmosphere
741 interactions, were utilized to capture the progression of short- and long-term conditions affecting
742 plant productivity, particularly in agricultural contexts. These drought indices confirm the
743 moisture-driven dry and wet patterns observed in early 1992 and the following years over the
744 tropical regions and mid-latitudes of the Northern Hemisphere, respectively, as a response to the
745 radiative perturbation caused by the Mt. Pinatubo eruption. Based on both drought indices, we
746 conclude that approximately 10-15% of land areas exhibit statistically significant dry or wet
747 patterns under the volcanically altered climate conditions of 1992 and 1993. The fraction of land
748 region showing a significant dry or wet response range between 5-10% for the next two (1994
749 and 1995) years. In equatorial Africa, seasonal decreases in both SMDI and ETDI suggested a
750 likely negative impact on plant productivity. In contrast, the Middle East showed increases in
751 both SMDI and ETDI, indicating a positive effect on plant productivity. Northern Asia, by
752 comparison, exhibited an increase in SMDI alongside a decrease in ETDI, implying a reduction
753 in plant productivity, though not driven by water-related factors.

754 Motivated by these key pattern differences, we extended our analysis of the drought
755 indices by incorporating higher temporal (weekly) frequencies along with AET, PET, and
756 transpiration. These regional analyses generally exhibit much stronger statistical significance on
757 a weekly scale and further confirm the season-based inferences described above. Further,
758 weekly drought indices show the temporal variability characteristics in the signal, which also
759 demonstrates the utility of explaining the effectiveness of short-term dry/wet conditions
760 corresponding to a regional crop cycle. In locations with insufficient/excess soil moisture, there
761 is a corresponding decrease/increase in evapotranspiration (AFR/MDE) and hence
762 decreased/increased plant productivity.

763 Kandlbauer et al., (2013) examined crop responses (using C3 and C4 grasses as proxies)
764 to the 1815 Tambora eruption using the HadGEM-ES model in three regions very similar to
765 those in our study. Their findings suggest that plant productivity decreases with positive changes
766 in soil moisture in the higher-latitude Asian region. In the mid-latitudes over the Southern
767 Europe/Middle East region (adjacent to our MDE region), volcanic eruptions may enhance plant
768 productivity by providing additional soil moisture through increased rainfall. However, in the
769 MDE region in our study, we found that the applied irrigation also benefits soil moisture supply
770 along with the increased rainfall. Furthermore, both studies report a decrease in productivity in
771 the tropical region. In general, these results complement the findings of this study, which suggest
772 that if sufficient water is available in the Southern Europe/Middle East region, volcanic eruptions
773 may enhance plant productivity. In contrast, in the far northern latitudes, water is not the primary
774 driver of plant responses, and productivity is likely to decline. Seasonal-scale changes in gross
775 primary productivity (GPP) confirm the regional trends in plant productivity following the
776 eruption. The simulations show a more pronounced decrease in GPP in the northern high-latitude

777 region and a significant increase in GPP over the European and Mediterranean regions.
778 Additionally, distinct patterns of decrease and increase in GPP are simulated in the tropical
779 northern and southern regions, respectively (Figure S10).

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780 This study is the first to conclusively demonstrate that there is an excess of root-zone soil
781 moisture in high-latitude regions (NAS) that plants are not utilizing for growth, indicating that
782 temperature and radiation are likely the primary controlling factors, thus confirming previous
783 findings (Krakauer and Randerson, 2003) and (Dong and Dai, 2017). The intricate nature of the
784 compounded response, particularly in relation to soil moisture-driven impact pathways in
785 tropical regions and high-latitude areas of the Northern Hemisphere, highlights the need to
786 expand the investigation beyond soil moisture and land-atmosphere interactions. The current
787 configuration of the NASA GISS model operates with prescribed vegetation with static plant
788 functional types and leaf area index. Incorporating dynamic vegetation could be essential for
789 capturing interactive land surface responses. Additionally, evaluating the influence of regional
790 and local biomes on photosynthesis rates could offer deeper insights into how these processes
791 respond to the climatic impacts of volcanic forcings. McDermid et al. (2022) have demonstrated
792 the sensitivity of regional hydroclimate to the local changes in soil organic carbon changes using
793 the soil moisture content. The results presented in this study regarding soil-moisture-based
794 drivers to plant productivity and surface temperature response in the northern hemisphere's high
795 latitudes also hinted towards the dominance of temperature effects on enhanced carbon sink in
796 terms of soil and plant respiration and reduced NPP (Krakauer and Randerson, 2003; Lucht et al.,
797 2002). Meanwhile, water-based drivers predominantly influence productivity responses in many
798 tropical and subtropical regions. Our findings demonstrate that soil moisture conditions across
799 different regions can provide valuable insights into the full impacts on agricultural yields and

801 regional carbon sink responses, particularly under scenarios involving changes in dynamic
802 vegetation and crop cover. A recently developed fully demographic dynamic vegetation model
803 (ModelE-BiomE v.1.0 (Weng et al., 2022)) incorporates interactive biophysical and
804 biogeochemical feedbacks between climate and land systems within the NASA GISS ModelE
805 framework. This model could be instrumental in evaluating carbon cycle responses under such
806 forcings.

807 **Code/Data availability.**

808 Details to support the results in the manuscript is available as supplementary information is
809 provided with the manuscript. GISS Model code snapshots are available at
810 <https://simplex.giss.nasa.gov/snapshots/> (National Aeronautics and Space Administration, 2024)
811 and calculated diagnostics are available at zenodo repository
812 (<https://zenodo.org/records/12734905>) (Singh et al., 2024). However, raw model output and data
813 at high temporal (daily) resolution and codes are available on request from author due to large data
814 volume.

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827 **Author's contributions**

828 RS, KT, DB, LS and KM identified the study period in consultation with the other authors and RS,
829 KT, DB, LS and BW designed the underlying simulations strategies. RS and KT implemented it
830 and performed the simulations using NASA GISS ModelE. RS and KT have performed the
831 analysis. RS created the figures in close collaboration with all authors. RS wrote the first draft of
832 the manuscript, and all other authors has contributed the writing of subsequent drafts. All authors
833 contributed to the interpretation of results.

834 **Competing interests**

835 One of the co-authors is member of the editorial board of Atmospheric Chemistry and Physics.

836 **Short Summary**

837 Analysis of post-eruption climate conditions using the impact metrics is crucial for understanding
838 the hydroclimatic responses. We used NASA's Earth system model to perform the experiments and
839 utilized the moisture-based impact metrics and hydrological variables to investigate the effect of
840 volcanically induced conditions that govern plant productivity. This study highlights the impact of
841 the Mt. Pinatubo eruption on the drivers of plant productivity, as well as the regional and seasonal
842 dependence of these drivers.

843

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