1	Supplementary Information for "The effectiveness of solar radiation management for
2	marine cloud brightening geoengineering by fine sea spray in worldwide different
3	climatic regions"
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28 We use the method of Martin et al. (1994) to calculate the cloud effective radius:

$$r_e = \left(\frac{3q_L}{4\pi\rho_w kN_{tot}}\right)^{\frac{1}{3}}$$

where q_L is the liquid water content, ρ_w is the density of water, N_{tot} is the cloud droplet number concentration (CDNC), and k is the ratio between the cube of the mean volume radius and the cube of the effective radius. Martin et al. (1994) estimated k for unpolluted (marine) stratocumulus clouds to be equal to 0.80 and for polluted (continental) stratocumulus clouds to be equal to 0.67. Here, we refer to the method of Goddard et al. (2022) and similarly set k to 0.80 for cloud condensation nuclei (CCN) concentrations of 0–50 cm⁻³ at 0.1% saturation, 0.74 for CCN concentrations of 50–150 cm⁻³ and equal to 0.67 where CCN concentrations are greater than 150 cm⁻³.

We use the method of Wood (2007) and Stephens (1978) to approximate the column cloud optical thickness (COT):

39
$$\tau \approx \frac{3}{2} \int_{z=0}^{h} \frac{q_L(z)}{\rho_w r_e(z)} dz$$

40 when integrated from the surface to a specified height, h. The height is determined by the highest grid cell 41 containing a liquid cloud (Goddard et al., 2022).

42 We use the method of Schwartz et al. (2002) to approximate the column mean cloud albedo:

43
$$\alpha_c \approx \frac{\tau(1-g) + 0.097}{\tau(1-g) + 1.43}$$

44 where *g* is the asymmetry parameter we assume *g* to be 0.834 for $r_e \le 6 \mu m$, 0.873 for $r_e \ge 19 \mu m$, and to 45 increase linearly between these re boundaries (Goddard et al., 2022).



Variations (Exp - Base) of total upward shortwave radiative flux (SW_TOT) at the TOA

Figure S1. Spatial distribution of the differences (Exp - Base) in the SW_TOT at the TOA due to the injection
of sea-salt aerosols in different ways in five oceanic regions.

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Figure S2. Spatial distribution of the differences (Exp - Base) in the sea-salt emissions due to the injection of
 sea-salt aerosols in different ways in five oceanic regions.



Figure S3. Comparison of MCB efficiency (E_{MCB}) resulting from uniform injection of 10⁻⁹ kg m⁻² s⁻¹ sea-salt aerosols within sensitive areas and over the full domains in five regions. The blue columns represent the E_{MCB} resulting from the uniform injection of sea-salt aerosols over the entire domain, while the yellow columns represent the E_{MCB} resulting from injection only within the sensitive areas to the entire region.

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Figure S7. Spatial distribution of dry diameter of accumulation mode aerosols for five ocean regions. The first column is for Base, the second is for sensitivity experiments with uniform injections of 10⁻⁹ kg m⁻² s⁻¹, and the third is for sensitivity experiments after uniform injections in sensitive areas. The black rectangles are sensitive areas.



Figure S8. Same caption as Fig. S7, but for the wet diameter of accumulation mode aerosols.



Figure S9. Spatial distribution of aerosol single scattering albedo ($\lambda = 0.533 \mu m$) for Base and sensitivity experiments in five regions. The first row for each region shows the results for Base, the sensitivity experiment with a uniform injection of 10^{-9} kg m⁻² s⁻¹, and the sensitivity experiment with a uniform injection in the sensitive area. The second row shows the difference between the sensitivity experiment and Base (Exp - Base), respectively. The black rectangles are sensitive areas.



- **Figure S10**. Same caption as Fig. S9, but for the aerosol asymmetry factor.





Figure S11. Relationship between changes in AOD and SW_AER responses due to uniform injection of 10⁻⁹
 kg m⁻² s⁻¹ sea-salt aerosols over the entire region (first column) and injection only within sensitive areas
 (second column). The Pearson's correlation coefficient (R) is given for each relationship.

Mean Liquid Cloud Fraction (from surface to 3000 m)



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Figure S12. Column mean liquid cloud fraction from the surface to 3000 m altitude for five regions. The first to fourth columns are Base, the sensitivity experiment with a uniform injection of 10⁻⁹ kg m⁻² s⁻¹ sea-salt aerosols over the entire region, Exp - Base, and the percent change of Exp - Base, respectively.



Figure S13. Same caption as Fig. S12, but for the sensitivity experiment with a uniform injection of 10^{-9} kg m⁻² s⁻¹ sea-salt aerosols only in the sensitive area. The black rectangles are sensitive areas.



Figure S14. Vertical cross sections of the mean liquid cloud fraction from the surface to 3000 m altitude for five regions, with cross sections longitudinally averaged. The first to third columns are Base, the sensitivity experiment with a uniform injection of 10^{-9} kg m⁻² s⁻¹ of sea-salt aerosols over the entire region, and Exp -Base, respectively.

For Asia (10⁻⁹ kg m⁻² s⁻¹ injection)



- **Figure S15**. Same caption as Fig. 10, but for the Asia region.



Figure S16. Same caption as Fig. 10, but for the NP region.



Figure S17. Same caption as Fig. 10, but for the Equa region.



- **Figure S18.** Same caption as Fig. 10, but for the SA region.



122 Figure S19. Same caption as Fig. 10, but showing the spatial distribution of the liquid cloud property response

- to a uniform injection of sea-salt aerosols within the sensitive area in the SP. The black rectangles are the
- 124 sensitive areas.



For Asia (10^{-9} kg m⁻² s⁻¹ injection in sensitive area)

Figure S20. Same caption as Fig. S19, but for the Asia region.



Figure S21. Same caption as Fig. S19, but for the NP region.



Figure S22. Same caption as Fig. S19, but for the Equa region.



Figure S23. Same caption as Fig. S19, but for the SA region.

Mean Cloud Radiative Forcing Parameter



Figure S24. The cloud radiative forcing (CRF) parameters after injection of sea-salt aerosols in the five regions. The first to fourth columns are Base, the sensitivity experiment with a uniform injection of 10^{-9} kg m⁻² s⁻¹ sea-salt aerosols over the entire region, Exp - Base, and the CRF'_{param} approximated by the perturbation method, respectively.

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Mean Cloud Radiative Forcing Parameter



Figure S25. Same caption as Fig. S24, but for the sensitivity experiment with a uniform injection of 10^{-9} kg m⁻² s⁻¹ sea-salt aerosols only in the sensitive area. The black rectangles are sensitive areas.



Figure S26. The three additive perturbation terms of the CRF'_{*param*} after uniform injection of sea-salt aerosols within the five regions (First column: driven by the perturbation of cloud albedo. Second column: driven by the change in cloud fraction. Third column: jointly driven by the interaction of the two.), as well as the CRF'_{*param*} approximated using the perturbation method (fourth column, see Equation 6 and 7). The percentage contribution of each item to the total CRF'_{*param*} is labeled in the lower right corner for the entire region.



Figure S27. Same caption as Fig. S26, but for the sensitivity experiment with a uniform injection of 10^{-9} kg m⁻² s⁻¹ sea-salt aerosols only in the sensitive area. The black rectangles are sensitive areas.



Figure S28. Same caption as Fig. 11, but for the sensitivity experiment with a uniform injection of 10^{-9} kg m⁻ 2 s⁻¹ sea-salt aerosols only in the sensitive area. The black rectangles are sensitive areas.



Figure S29. Box plots of total column water vapor (TCWV) for the five ocean regions from ERA5. The left 164 column shows daily mean data from ERA5 for the years 1990-2020 (1990-2023 for SP and SA), listed by 165 month. For Asia, NP and Equa, each year is wetter in August and drier in February. For SP and SA, March is 166 wetter and September is drier. The right column shows the daily average of the ERA5 data for the simulated 167 (wetter) and dry months of the year. For Asia, NP, and Equa, the initial simulation time period was August 168 2018, which was wetter, so we chose the dry time period of February of the same year to simulate again. SP 169 and SA were initially simulated in March 2023, which was wetter, so we chose the dry time period of 170 September of the same year to simulate again. 171

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Models	Strategies	Increased emissions	Locations	References
	set N = 400 cm ⁻³	10 ⁹ in mass (NaCl) 10 ²⁶ in number of droplets	global	(Latham, 2002)
A simplified version of the model of marine stratocumulus clouds	$\bigtriangleup N$ = 10, 30, 100, 300 and 1000 cm $^{-3}$			(Bower et al., 2006)
HadGAM	set Nd = 375 cm^{-3}		alabal	(Latham et
CAM	set Nd = 375 and 1000 cm ⁻³		giodai	al., 2008)
HadGEM2	set CDNC = 375 cm^{-3}		North Pacific (NP), South Pacific (SP) and South Atlantic (SA)	(Jones et al., 2009)
CCSM	set CDNC = 1000 cm^{-3}		20%, 30%, 40% and 70% of ocean area	(Rasch et al., 2009)
GLOMAP	set geoengineering particle number flux (GEO and 5GEO) according to U10		off the coast of California (North Pacific), Chile (South Pacific), Namibia (South Atlantic) and Western Australia (Indian Ocean)	(Korhonen et al., 2010)
HadGEM2-AO	set CDNC = 375 cm^{-3}		North Pacific (NP), South Pacific (SP) and South Atlantic (SA)	(Jones et al., 2011)
WRF	inject CCN	1.45×10 ⁶ m ⁻² s ⁻¹ (375 cm ⁻³ hour ⁻¹)		(Wang et al., 2011)
CAM3.5-CLM3.5	r_d over the ocean is reduced from 14 to 11.5 μ m		global	(Bala et al., 2011)
NorESM	increase sea-salt emissions	10 ⁻⁹ kg m ⁻² s ⁻¹ (350 tons s ⁻¹)	global	(Alterskjær et al., 2012)
GFDL-CM2G	set CCN = 500 and 1000 cm ⁻³		North Pacific and the Southern Ocean	(Baughman et al., 2012)
HadGEM1	set CDNC = 375 cm^{-3}		global	(Gadian, 2012)
AGCM	increase sea salt aerosols	fivefold	tropical North Pacific (NP), South Pacific (SP), and South Atlantic (SA)	(Hill and Ming, 2012)
HadGEM2-ES	follow Korhonen et al., (2010)	GEO and 5GEO	global and 10% of optimal sea-spray emission areas	(Jones and Haywood,

Table S1. Summar	y of modeling studi	es on marine cloud l	brightening (MCB), mari	ine sky brighteni	ıg (MSI	3) and in	jection of	f sea salt	aerosols
	5 0		0 0 (1 0	\mathcal{O}	/	J		

				2012)
HadGEM1	set CDNC = 375 cm^{-3}		off the western coasts of California, Peru and Namibia	(Latham et
HadGEM1	set CDNC = 375 cm^{-3}		global and off the western coasts of California, Peru, Namibia	(Latham et al. 2012b)
ECHAM5.5-HAM2	set geoengineering particle number flux (GEO) according to U10	20.6–443.9 Tg yr ⁻¹	global and North Pacific, South Pacific and South Atlantic (3.3% of the Earth's surface)	(Partanen et al., 2012)
0-D model GLOMAP-MODE EMAC ECHAM-HAM	follow Korhonen et al., (2010)		global	(Pringle et al., 2012)
HadGEM1	set CDNC = 375 cm^{-3}		follow Jones et al. (2009)	(Parkes et al., 2012)
NorESM IPSL-CM5A MPI-ESM	increase sea salt aerosols use the output of the NorESM	266–560 Tg yr ⁻¹	between 30°S and 30°N	(Alterskjær et al., 2013)
NorESM	increase sea salt emissions	10 ⁻¹¹ –10 ⁻⁸ kg m ⁻² s ⁻¹	between 30°S and 30°N	(Alterskjær and Kristjánsson, 2013)
WRF-Chem	inject aerosols	3–15 kg s ⁻¹	point source injection	(Jenkins et al., 2013)
MPI-ESM and NorESM	increase sea salt aerosols	$10^{-9} \text{ kg m}^{-2} \text{ s}^{-1}$	between 30°S and 30°N	(Niemeier et al., 2013)
Gaussian plume model LES ECHAM5.5-HAM2	inject aerosols	20.6 Tg yr ⁻¹	North Pacific, South Pacific and South Atlantic	(Stuart et al., 2013)
HadGEM2-ES	50% increase in CDNC increase sea salt aerosols	100–400 Tg yr ⁻¹	global between 30°S and 30°N	(Kravitz et al., 2013)
Lagrangian Cloud Model	inject aerosols	100, 200, 400, 800 cm ⁻³		(Andrejczuk et al., 2014)
ACPIM	inject aerosols	NaCl mixing ratios = 10^{-14}		(Connolly et

		10 ⁻⁴ kg kg ⁻¹		al., 2014)
cloud-resolving model (WRF)	inject CCN	1.45×10 ⁶ m ⁻² s ⁻¹	single moving point source injection (Arctic, 71.32°N, 156.61°W)	(Kravitz et al., 2014)
UCLALES	Inject particles	15 kg s ⁻¹		(Maalick et al., 2014)
HadGEM1	set CDNC = 375 cm^{-3}		Antarctic, off the West coasts of North and South America, and Africa	(Latham et al., 2014)
HadGEM2	increase sea salt aerosols	1.8×10 ⁸ m ⁻³	between 30°S and 30°N	(Crook et al., 2015)
HadGEM1 GLAM	set CDNC = 375 cm^{-3}		off the western coasts of California, Peru and Namibia	(Parkes et al., 2015)
NorESM1-M IPSL-CM5A-LR MPI-ESM-LR	increase sea salt aerosols use the output of the NorESM		between 30°S and 30°N	(Muri et al., 2015)
MPI-ESM NorESM IPSL-CM5	follow Alterskjær et al. (2013) and Muri et al. (2015)		between 30°S and 30°N	(Aswathy et al., 2015)
UVic ESCM ECHAM5.5-HAM2	use the radiative forcing from Partanen et al. (2012)		off the west coasts of North America, South America, and Southern Africa	(Partanen et al., 2016)
LMDZ5B	prescribe an additional concentration sea salt		between 30°S and 30°N	(Boucher et al., 2017)
NorESM1-M GISS-E2-R HadGEM2-ES	inject sea salt particles	250 Tg yr ⁻¹ 590 Tg yr ⁻¹ 200 Tg yr ⁻¹	between 30°S and 30°N	(Ahlm et al., 2017)
CESM	r_d over the ocean is reduced from 14 to 11 μ m		global	(Duan et al., 2018)
NorESM1-ME	increase sea salt emissions	460 Tg yr ⁻¹	between 45°S and 45°N	(Muri et al., 2018)
BNU-ESM CanESM2 CSIRO-Mk3L-1-2 GISS-E2-R	set CDNC = 375 cm^{-3}		global	(Stjern et al., 2018)

HadGEM2-ES				
IPSL-CM5A-LR				
MIROC-ESM				
MPI-ESM1-LR				
NorESM1-M				
BNU-ESM				
CanESM2				(Kim et al
CSIRO-Mk31-1-2	set CDNC = 375 cm^{-3}		global	(Killi et al.,
HadGem2-ES				2020)
MIROC-ESM				
CEOS Cham	inicat and calt norticles	212–569 Tg yr ⁻¹	hotwood 2005 and 200N	(Horowitz et
GEOS-Chem	inject sea sait particles	$(3.0-8.0\times10^{-12} \text{ kg m}^{-2} \text{ s}^{-1})$	between 50°S and 50°N	al., 2020)
LCM				(Hoffmann
LES	inicat annov duanlata	1.2 18274.0 mc ⁻¹		and
	inject spray dropiets	1.2–18374.9 mg		Feingold,
parcel model				2021)
	::	50 70 To	5 4 0/ - 6 4 - 5 6	(Wood,
	inject sea sait particles	50-70 Ig yr	54 % of the Earth's surface	2021)
HadGEM2 ES	50% increase in CDNC		the Schere Schel Archien Depingula zone	(Zhu et al.,
HauOEWI2-ES	50% increase in CDIVE		the Sanara-Sanci-Arabian remissing zone	2021)
CESM	r_d over the ocean is reduced from 14		alahal	(Zhao et al.,
CESIM	to 11 µm		giotai	2021)
BNU-ESM				
CanESM2				
HadGEM2-ES	50% increase in CDNC		alahal	(Xie et al.,
ISPL-CM5A-LR	50% increase in CDNC		global	2022)
MIROC-ESM				
NorESM1-M				
WDE Chart	inicat and solt norticles	10.9 To rec	Culf of Marian	(Goddard et
w Kr-Chem	inject sea sait particles	10.8 Ig yr	Gull of Mexico	al., 2022)
	14 increase sea salt emissions	7.66×10 ⁻¹¹ kg m ⁻² s ⁻¹	haturaan 2008 and 200N	(Mahfouz et
GFDL-AM4		(456 Tg yr ⁻¹)	between 50°S and 50°N	al., 2023)
LCM	inject aerosols	1 μg kg ⁻¹ of air	25°N, 120°W	(Prabhakaran

LES				et al., 2023)
			NP (north Pacific: 30°–50°N,170°–240°E), NEP (north-	
UKESM1	modify see selt emissions	413 Tg yr ⁻¹	east Pacific: 0°–30°N, 210°–250°E), SEP (south-east	(Haywood et
OKLSWI	moury sea sait emissions		Pacific: 0°–30°S, 250°–290°E) and SP (south Pacific: 30°–	al., 2023)
			50°S, 190°–270°E)	

Note: Some studies included multiple sensitivity experiments with aerosol injection, and only representative experiments may be listed in the table.

Table S2. The total upward shortwave radiation flux (SW_TOT) at the TOA and the corresponding sea-salt aerosol injections resulting from different strategies of injecting sea-salt aerosols in five areas, and the MCB efficiency (E_{MCB}).

Strategies	Areas	SW_TOT (W m ⁻²)	Add Sea-salt aerosols (×10 ⁻⁹ kg m ⁻² s ⁻¹)	E _{MCB} (GW kg ⁻¹ s)
	Asia	0.60	0.09	6.97
	NP	2.08	0.06	37.67
Natural×5	Equa	0.06	0.05	1.11
	SP	1.55	0.04	43.46
	SA	1.43	0.03	47.07
	Asia	4.02	0.19	21.41
	NP	8.47	0.20	41.72
Wind-adjusted	Equa	1.35	0.19	7.13
	SP	7.75	0.18	42.74
	SA	7.91	0.21	37.02
	Asia	17.34	1.00	17.45
$\Gamma_{1} = 1 + 10^{-9}$	NP	23.11	1.00	23.22
Fixed at 10^{-7} kg m	Equa	10.96	1.00	10.96
- 8 -	SP	24.50	1.00	24.58
	SA	22.36	1.00	22.43
	Asia	0.65	0.05	13.36
10-9121	NP	2.69	0.05	55.43
10 ⁻ kg m ⁻ s ⁻ in the	Equa	0.74	0.05	14.67
sensitive area	SP	3.27	0.05	65.53
	SA	1.81	0.05	36.48
	Asia	7.21	1.00	7.22
First stird	NP	16.07	1.00	16.22
F1Xed-W1nd-	Equa	5.00	1.00	5.02
adjusted	SP	16.40	1.00	16.46
	SA	19.78	1.00	19.82

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