



Size-resolved hygroscopicity and volatility properties of ambient urban aerosol particles measured by the VH-TDMA system in the autumn of 2023

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Abstract.

Hygroscopicity, volatility and the hygroscopicity of the non-volatile core of submicron aerosols with diameters of 50 nm, 80 nm, 110 nm and 150 nm were measured using a Volatility Hygroscopicity Tandem Differential Mobility Analyzer (VH-TDMA) system with a relative humidity of 90 % and a thermal denuder temperature of 270 °C from 15 11 October to 6 November 2023 in Beijing. The mean hygroscopic growth factor (HGF) for particles ranging from 50 nm to 150 nm was 1.15 ± 0.07 , 1.24 ± 0.08 , 1.30 ± 0.09 , and 1.36 ± 0.10 , respectively, while the mean volatile shrink factor (VSF) was 0.51 ± 0.05 , 0.55 ± 0.04 , 0.56 ± 0.05 , and 0.56 ± 0.07 , respectively. Both the HGF probability density function (HGF-PDF) and the VSF probability density function (VSF-PDF) for all particle sizes exhibited a pronounced bimodal distribution, indicating that the particles were primarily in an external mixing state. 20 Hygroscopicity was observed to increase with particle size in both clean and pollution periods, while volatility decreased slightly with particle size during the clean period, without apparent trend during the pollution period. Additionally, a positive correlation was identified between hygroscopicity and volatility, as well as between the number fraction of nearly hydrophobic (NH) and non-volatile (NV) particles, which was higher correlation during the pollution period. Furthermore, this study measured the HGF of the non-volatile core (HGF_{core}) of submicron 25 aerosols heated at 270 °C and derived the HGF of the volatile coating (HGF_{coating}). The mean HGF_{coating} for particles ranging from 50 nm to 150 nm particles was 1.17 ± 0.08 , 1.27 ± 0.10 , 1.35 ± 0.10 and 1.41 ± 0.10 , respectively, which is 2 % to 7 % higher than the mean HGF for the same particle sizes. The mean HGF_{core} for particles ranging from 50 nm to 150 nm was 1.08 ± 0.03 , 1.07 ± 0.03 , 1.07 ± 0.03 and 1.09 ± 0.04 , respectively. Notably, the HGF_{core} values were influenced by marine aerosols during air mass passages through and originating from the Bohai Sea.



30 **1 Introduction**

Atmospheric aerosols significantly influence the radiation balance and climate change (IPCC, 2013), play a crucial role in reducing visibility (Zhang et al., 2012), and affect public health (Ou et al., 2020; Lippmann and Albert, 1969). The impact of atmospheric aerosols on global climate, human health, and air quality depends largely on their physicochemical properties (Pöschl, 2005; Buseck and Posfai, 1999). A deep understanding of aerosol particles' physicochemical properties, including hygroscopicity and volatility, is crucial for determining their potential impacts. Hygroscopicity, the ability of aerosols to uptake water vapor (Su et al., 2010), determines the size and optical properties of aerosol particles, thus affecting visibility and direct radiative forcing (Zhao et al., 2019; Xia et al., 2019; Freney et al., 2010). Moreover, hygroscopicity significantly influences CCN activation and cloud droplet formation (Zhang et al., 2023; Tang et al., 2016; Reutter et al., 2009), and the degree of aerosol particles mixing also can be determined by measuring their hygroscopicity distribution (Wang et al., 2019; Ye et al., 2013). The Humidity Tandem Differential Mobility Analyzer (H-TDMA) is extensively used to measure aerosol hygroscopicity in various field studies (Tang et al., 2019; Swietlicki et al., 2008).

Volatility affects the gas-particle distribution of aerosol particles and influences their dry and wet deposition rates (Hakala et al., 2016; Bidleman, 1988), as well as atmospheric lifetime (Huffman et al., 2009). Therefore, studying particle volatility across different regions and environments is essential. The volatility properties of atmospheric aerosols have been investigated using several instruments and techniques (Xu et al., 2019; Jiang et al., 2018; Karnezzi et al., 2014). Volatility Tandem Differential Mobility Analyzer (V-TDMA) is one of the commonly used instruments for measuring the volatility of particles. By comparing changes in the diameters of particles before and after heated at a certain temperature, the volatility of particles can be inferred. Volatile compounds such as sulfates, nitrates, and most of the organic species evaporate at 300 °C (Hong et al., 2014). However, some compounds like black carbon (BC), soot, sea salt, and crustal materials do not volatilize even at higher temperatures. The V-TDMA is usually used to assess the mixing state of typically non-volatile BC particles (Chen et al., 2022; Wehner et al., 2009) and to determine the number, and volume fractions of non-volatile material at a selected temperature up to 300 °C for particles (Zhang et al., 2016; Jiang et al., 2018; Ghadikolaei et al., 2020). Studies indicate that BC's internal mixing is higher in winter than in summer (Chen et al., 2022) and increases significantly during the pollution period compared to the cleaner period (Wehner et al., 2009; Wang et al., 2017b).

Although previous studies have focused on the hygroscopicity or volatility of aerosol particles at different sampling locations under varying atmospheric conditions using H-TDMA or V-TDMA (Zhang et al., 2023; Wang et al., 2019; Jiang et al., 2018; Wu et al., 2016; Ye et al., 2013), simultaneous measurements of both properties,



60 particularly in China, have been scarce. The Volatility Hygroscopicity Tandem Differential Mobility Analyzer (VH-TDMA) combines the measurement characteristics of the V-TDMA and H-TDMA systems, providing an opportunity to simultaneously study the hygroscopicity and volatility of particles, as well as the hygroscopicity of the particle core after the volatilization (Häkkinen et al., 2012; Villani et al., 2008; Johnson et al., 2005). This instrument has been applied in different environments such as oceans (Modini et al., 2009; Johnson et al., 2005),
65 forests (Ristovski et al., 2010; Hong et al., 2014) and urban areas (Enroth et al., 2018; Hakala et al., 2016; Villani et al., 2013) to investigate aerosol mixing states and reveal correlations between hydrophobic and non-volatile particles (Modini et al., 2009; Ristovski et al., 2010; Enroth et al., 2018). For instance, a study in Budapest demonstrated that while hygroscopic particles are generally volatile, hydrophobic and low-volatile particles are predominantly combustion particles emitted from vehicles (Enroth et al., 2018). Zhang et al. (2016) observed strong
70 correlations between hydrophobic and non-volatile particles at a suburban site in Beijing. However, a weak linear relationship was reported between hydrophobic and non-volatile particles in urban Beijing (Wang et al., 2017a). The four cases in previous study by Villani et al. (2008) showed that gentle thermodesorption potentially has a significant impact on the hygroscopicity of particles at a given RH. It suggests that the loss of volatile material from the particle surface may significantly increase or decrease the hygroscopic growth, depending on the nature of the
75 condensed material. Research on forest aerosol particles shows that hygroscopic materials remain in the particles even after heated at 280 °C (Hong et al., 2014).

Simultaneous measurement of the hygroscopicity and volatility of aerosol particles remain scarce. To our knowledge, no study in China has yet focused on the hygroscopicity of particles after heating. Investigating the hygroscopicity of the particle core after the volatilization is crucial for understanding their composition and sources.
80 In recent years, China has gained significant progress in air pollution control. However, atmospheric aerosol pollution events still frequently occur in major cities (Zhang et al., 2018; Wang et al., 2019; Fan et al., 2020a; Miao et al., 2021; Lu et al., 2024). Beijing, the capital of China, regularly experiences haze events due to its large population, substantial industrial emissions, and an increasing number of vehicles (Zhong et al., 2018; Huang et al., 2021; Yang et al., 2023). In this study, we utilized the VH-TDMA system for the first time to investigate the
85 hygroscopicity and volatility, and the hygroscopicity of the non-volatile core of submicron aerosols in urban Beijing during the autumn of 2023. We explored the variations in hygroscopicity, volatility, and particle mixing states under different atmospheric conditions and further analyzed the relationship between the hygroscopicity and volatility of aerosol particles. We also studied the hygroscopicity of non-volatility particles after heated at 270 °C.



2 Experimental set up

90 2.1 Sampling site

The measurement of aerosol hygroscopicity and volatility were conducted from 11 October to 6 November 2023, in a laboratory on the second floor (~ 5 m above the ground) of Chinese Academy of Meteorological Sciences (CAMS) building, located on the campus of the China Meteorological Administration in Beijing, China. The site is located between the 2nd and 3rd ring roads. West of the site is a major road with heavy traffic within a distance of 95 less than 200 m and this site is predominantly affected by residential and traffic emissions (Wang et al., 2018).

2.2 VH-TDMA system description

The schematic diagram of the Volatility Hygroscopicity Tandem Differential Mobility Analyzer (VH-TDMA) (TROPOS, Germany) is shown in Fig. 1. The VH-TDMA mainly consists a DMA1, SMPS1 (DMA2+CPC1), SMPS2 (DMA3+CPC2), thermal denuders (TD) and humidifiers. DMA1 is utilized for the size selection of 100 monodisperse particles, while the two SMPS system are used to measure ambient, thermally treated or humidified aerosol. The temperature of the thermal denuders (TD) can be controlled from ambient temperature up to 350 °C, whereas the RH of aerosol and sheath air could be controlled up to 90 %. The system also includes a nebulizer setup for ammonium sulfate calibrations, as well as ball valves, which control the system operation modes through the home-made program.

105 In this study, VH-TDMA was alternately operated in combinations of V-TDMA mode and H-TDMA mode or V-TDMA mode and VH-TDMA mode. Ambient air was dried to an RH of less than 20% using a silica and a Nafion dryer, and then neutralized with an X-ray aerosol neutralizer (TSI 3088). In the first combination, the monodisperse particles from DMA1, at a flow rate of 2 L min⁻¹, were equally split to two flows. One flow went to the thermal denuder set at 270 °C, then to the SMPS1 for measuring the size distribution of heated particles. The other flow 110 went to the aerosol humidifier set at 90 % RH, then to the SMPS2 for measuring the size distribution of humidified particles. Therefore, the volatility and hygroscopicity of monodisperse particles were simultaneously measured in this combination. Similarly, in the latter combination, the monodisperse particles from DMA1 with a flow rate of 2 L min⁻¹ went through the thermal denuder first, then equally split to two flows. One flow directly went to the SMPS1 for measuring the size distribution of heated particles, and another flow went to the aerosol humidifier at 90 % RH, 115 then to the SMPS2 for measuring hygroscopicity of these monodisperse particles after heating. Particles were selected at diameters of 50 nm, 80 nm, 110 nm, and 150 nm using DMA1. Each particle size was sampled for approximately 15 minutes, including about 8 minutes for measurement and 6 minutes for system flushing between combination change or size change, with each cycle lasting about 1 h. The thermal denuder is well-insulated



120 cylindrical metal tube, maintained a consistent heating temperature at 270 ± 0.5 °C. The RH of the aerosol flow and sheath flow in the humidification system was set to 90 %, allowing for was maintained at 90 ± 1 %. The hygroscopicity of 100 nm ammonium sulphate ($(\text{NH}_4)_2\text{SO}_4$) particles generated by nebulizer were measured regularly to check the accuracy of RH measurement. Size calibration was also conducted for DMAs using the Polystyrene Latex Spheres (PSL) particles with a nominal size of 203 nm. Data inversion was accomplished using the TDMAinv program package (Gysel et al., 2009) in IGOR Pro software.

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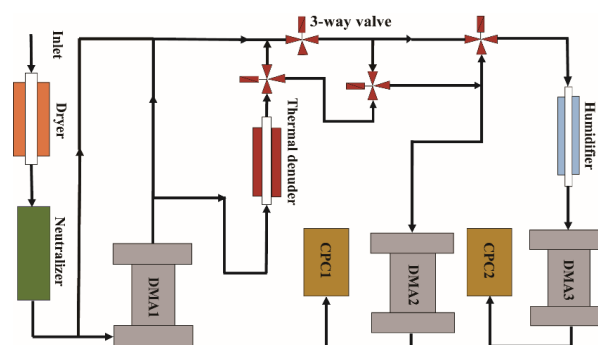


Figure 1. The schematic diagram of the VH-TDMA.

2.3. Particle number size distribution

130 The particle number size distribution (PNSD) in the range of 3-850 nm is measured by a twin scanning mobility particle sizer (TSMPS, TROPOS, Germany). This system is comprised by two differential mobility analyzers (DMA and UDMA, TROPOS, Germany) and two CPCs (CPC 3772 and ultrafine CPC 3776, TSI Inc., St Paul, USA). The combination of CPCs and DMAs can measure the PNSD in the size ranges of 3-80 nm and 40-850 nm, respectively. The combined PNSD covers the size range of 3-850 nm with 40 size bins in a 10 minutes time resolution. More information about the instrument setup and data evaluation can be referred to Shen et al. (2018).

135 2.4. Chemical composition

A high-resolution time-of-flight aerosol mass spectrometer (HR-ToF-AMS, Aerodyne Research, Inc., USA) was used to measure the mass concentrations of the main chemical composition of non-refractory PM_{10} , including organic matter, sulfate, nitrate, ammonium and chloride from 24 October to 6 November 2023, with a time resolution of 5 min. For more detailed information on the working principles of the instrument and data processing, refer to
140 Zhang et al. (2012).



2.5. Supporting data

The mass concentrations of PM_{2.5} and PM₁₀ were from GuanYuan Air Quality Monitoring Site, China National Environmental Monitoring Center (CNEMC, <http://www.cnemc.cn>), located ~3 km from the CAMS site. The meteorological data used in this study were from the Haidian National Basic Meteorological Station (station No.54399) of Beijing Meteorological Bureau, located ~ 5 km northwest of the CAMS site. The GDAS data from the National Center for Environmental Prediction (NCEP) were used to derive 48 h backward trajectory at an altitude of 500 m above ground level of the CAMS station, using the trajectory calculation software Trajstat developed by Wang et al. (2011).

3 Data analysis

150 The hygroscopic growth factor (HGF) is defined as the ratio of the particle mobility diameter under a given relative humidity $D_p(T_{\text{room}}, \text{RH})$ to the dry diameter $D_p(T_{\text{room}}, \text{RH}_{\text{dry}})$ at room temperature:

$$\text{HGF} = \frac{D_p(T_{\text{room}}, \text{RH})}{D_p(T_{\text{room}}, \text{RH}_{\text{dry}})} \quad (1),$$

In this study, the relative humidity was set at 90 %.

The volatile shrink factor (VSF) is defined as the ratio of the particle mobility diameter under a given temperature $D_p(T, \text{RH}_{\text{dry}})$ to the dry diameter $D_p(T_{\text{room}}, \text{RH}_{\text{dry}})$ at room temperature:

$$\text{VSF} = \frac{D_p(T, \text{RH}_{\text{dry}})}{D_p(T_{\text{room}}, \text{RH}_{\text{dry}})} \quad (2),$$

In this study, the heating temperature was set at 270 °C.

The hygroscopic growth factor of the non-volatile core of aerosol after heating, denoted as HGF_{core} are derived as:

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$$\text{HGF}_{\text{core}} = \frac{D_p(270^\circ\text{C}, 90\% \text{RH})}{D_p(270^\circ\text{C}, \text{RH}_{\text{dry}})} \quad (3).$$

HGF-PDF was retrieved from the measured number size distributions of the humidified aerosols. $c(\text{HGF}, D_p)$ is the normalized probability density distribution of the particle HGF, such that $\int_0^\infty c(\text{HGF}, D_p) d\text{HGF} = 1$. The number-weighted mean hygroscopic growth factor was calculated as follows:

$$\text{HGF}_{\text{mean}} = \int_0^\infty \text{HGF} c(\text{HGF}, D_p) d\text{HGF} \quad (4),$$

165 The Mean VSF (VSF_{mean}) was calculated same as the HGF_{mean}.

According to the κ -Köhler theory (Petters and Kreidenweis, 2007), the single hygroscopic parameter κ can be calculated using the HGF as follow:

$$\kappa = (\text{HGF}^3 - 1) \left[\frac{1}{\text{RH}} \exp\left(\frac{4\sigma M_w}{RT\rho_w D_p \text{HGF}}\right) - 1 \right] \quad (5),$$



where RH is the relative humidity, HGF is the hygroscopic growth factor from Eq. (1), D_p is the diameter of
170 dry particles, σ is droplet surface tension (assumed to be pure water, $\sigma = 0.0728 \text{ N m}^{-2}$), M_w is the molecular mass
of water, R is the universal gas constant, T and ρ_w is the temperature (in K) and density of water, respectively.

Based on the measured HGF-PDFs and VSF-PDFs, aerosol particles were classified into two hygroscopic
groups and two volatile groups in this study, i.e. nearly hydrophobic (NH) mode particles ($\text{HGF} \leq 1.2$) and more
hygroscopic (MH) mode particles ($\text{HGF} > 1.2$), as well as non-volatile (NV) mode particles ($\text{VSF} \geq 0.8$) and very
175 volatile (VV) mode particles ($\text{VSF} < 0.8$).

The number fraction (NF) is defined as the proportion of the number of particles with HGF in the boundary of
(a, b) and computed as follow:

$$\text{NF} = \int_a^b c(\text{HGF}, D_p) d\text{HGF} \quad (6),$$

The number of particles with VSF in the boundary was calculated same as the number of particles with HGF.

180 The standard deviation (σ) of the HGF-PDF is used as a measure for the spread of growth factors (Liu et al.,
2011; Sjogren et al., 2008), is defined as follow:

$$\sigma_{\text{HGF-PDF}} = \left[\int_0^\infty (\text{HGF} - \text{HGF}_{\text{mean}})^2 c(\text{HGF}, D_p) \right]^{1/2} \quad (7).$$

$\sigma_{\text{HGF-PDF}} \leq 0.10$ indicates that particles are in a state of internal mixing with limited spread of growth factors,
while a $\sigma_{\text{HGF-PDF}} \geq 0.15$ indicates substantial external mixing (Sjogren et al., 2008).

185 **4 Results and discussion**

4.1. Meteorological conditions during the sampling period

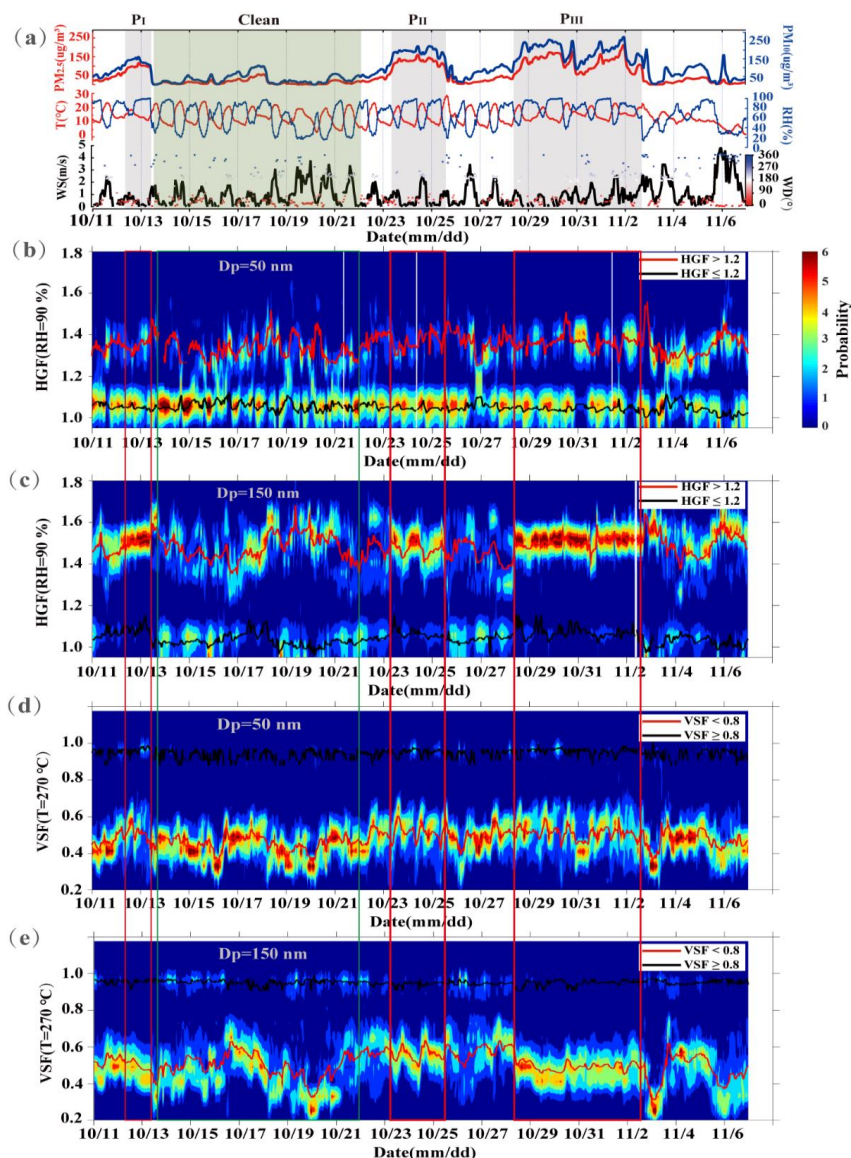
Time series of hourly meteorological parameters and $\text{PM}_{2.5}$, PM_{10} mass concentrations from 11 October to 6
November 2023 were shown in Figure 2a. The average mass concentrations of $\text{PM}_{2.5}$ and PM_{10} were $50.7 \pm 51.7 \mu\text{g}$
 m^{-3} and $95.0 \pm 72.2 \mu\text{g m}^{-3}$, respectively. Temperature and relative humidity ranged from 1.8 to 28.4 °C and 14 to
190 100 %, with average of $13.9 \pm 5.1 \text{ }^\circ\text{C}$ and $66.3 \pm 25.8 \%$, respectively. The average wind speed was $0.9 \pm 0.9 \text{ m s}^{-1}$
during the sampling period, predominantly from the northeast. Pollution periods were defined as time when hourly
 $\text{PM}_{2.5}$ concentrations exceeded $75 \mu\text{g m}^{-3}$, based on Grade II standard values of Chinese Ambient Air Quality
standards (GB3095-2012). Three pollution periods are identified as: P_I (12 October, 9:00–13 October, 9:00 LT), P_{II}
(23 October, 9:00–25 October, 13:00 LT), and P_{III} (28 October, 10:00–2 November, 17:00 LT). It is worth to mention
195 that more than 90 % time period of an orange alert for severe air pollution (from 30 October, 12:00 to 2 November,
24:00 LT) was included in P_{III}. We classified the days with new particle formation (NPF) days as clean period (13
October, 13:00–21 October, 23:00 LT), during which seven NPF events were observed.



During the pollution period, average concentrations of PM_{10} and $PM_{2.5}$ were $185.2 \pm 41.6 \mu\text{g m}^{-3}$ and $118.7 \pm 33.1 \mu\text{g m}^{-3}$, respectively, compared to $27.4 \pm 17.5 \mu\text{g m}^{-3}$ for PM_{10} and $9.5 \pm 6.9 \mu\text{g m}^{-3}$ for $PM_{2.5}$ during the clean period. The average wind speed during the pollution period ($0.7 \pm 0.6 \text{ m s}^{-1}$) was lower than that during the clean period ($1.0 \pm 0.9 \text{ m s}^{-1}$), which facilitates the accumulation of pollutants (Wehner et al., 2008). Analysis of air mass back trajectories during the pollution period (Fig. S1b) reveals that 76 % of the air masses passed through densely populated southern region of Beijing (such as Hebei, northwestern Shandong and northern Shanxi) and 98.5 % of the air masses traveled below 1.0 km before reaching the sampling site. In contrast, over 80 % of air masses originated from the northwest during the clean period (Fig. S1a). Air masses from the northwest moved more quickly and accumulated fewer pollutants during the transport before reaching the sampling site.

4.2 Time series of HGF- PDF、 VSF- PDF and averages of hygroscopicity and volatility

Figure 2 and S2 show the time series of HGF-PDFs and VSF-PDFs for different size particles throughout the sampling period. Both HGF-PDFs and VSF-PDFs exhibited clearly bimodal distributions for 50 nm and 150 nm particles, similar for 80 nm and 110 nm particles shown in Fig. S2. As mentioned above, the lower (HGF ≤ 1.2) and higher modes (HGF > 1.2) of HGF-PDFs are referred to as NH and MH modes, respectively. The lower (VSF < 0.8) and higher mode (VSF ≥ 0.8) of VSF-PDFs correspond to the VV and NV mode, respectively. As can be seen from Fig. 2b-e, MH mode for HGF-PDFs and VV mode for VSF-PDFs showed larger variations than NH mode and NV mode during the course of sampling period. However, HGF-PDFs and VSF-PDFs for 150 nm particles exhibited smaller variations and higher proportions of more hygroscopic (MH) and very volatile (VV) modes during the pollution period, especially during P_{III} (Fig. 2c and 2e). The variations in the HGF-PDFs and VSF-PDFs over time could be influenced by different emission sources, atmospheric processes, meteorological factors and air mass origins (Enroth et al., 2018). Although the unimodal or bimodal distribution of the HGF-PDFs and VSF-PDFs could give some information on the mixing state of particles, the standard deviation of the HGF-PDF is used to quantify the degree of external mixing of particles (Sjogren et al., 2008; Liu et al., 2011). A value of $\sigma_{\text{HGF-PDF}} \geq 0.15$ indicating an externally mixed state, as defined by Sjogren et al. (2008). In this study, more than 85 % of HGF-PDF for 80 nm, 110 nm and 150 nm aerosol particles was exhibited a $\sigma_{\text{HGF-PDF}} \geq 0.15$, while less than 50 % of the 50 nm aerosol particles. The mean $\sigma_{\text{HGF-PDF}}$ for 50 nm, 80 nm, 110 nm, and 150 nm particles was 0.15 ± 0.03 , 0.19 ± 0.04 , 0.20 ± 0.04 , and 0.20 ± 0.04 , respectively. These results indicate a stronger degree of external mixing for larger particles.



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Figure 2. Time series of PM_{2.5} and PM₁₀ mass concentrations and meteorological parameters (a), hygroscopic growth factor probability density function (HGF-PDF) for 50 nm (b) and 150 nm (c) particles at RH=90 %, and volatile shrink factor probability density function (VSF-PDF) for 50 nm (d) and 150 nm (e) particles at $T=270$ °C. In the HGF-PDFs, the blackline represents the mean HGF of the nearly hydrophobic mode (HGF_{NH}), and the red line represents the mean HGF of the more hygroscopic mode (HGF_{MH}). In the VSF-PDFs, the blackline represents the mean VSF of the non-volatility mode (VSF_{NV}), and the red line represents the mean VSF of the very volatility mode (VSF_{VV}).

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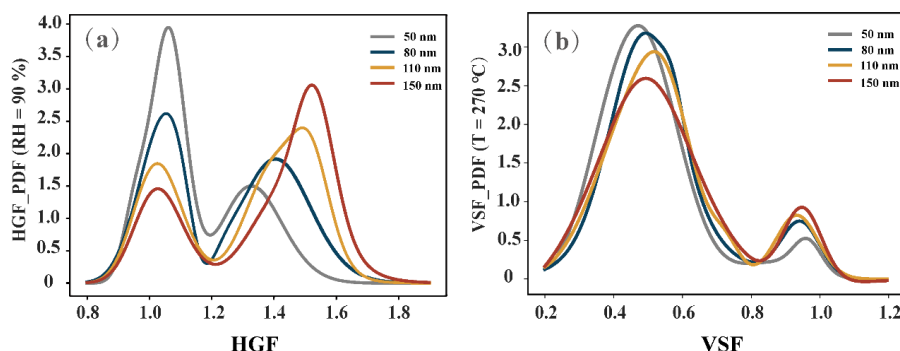


Figure 3 displays the size-resolved mean HGF-PDFs and VSF-PDFs for the sampling periods. As mentioned above, both the HGF-PDFs and VSF-PDFs exhibited a clearly bimodal distribution. The averaged HGF-PDFs at 90 % RH typically show a minimum at HGF around 1.2 (Fig. 3a), which is why we use HGF=1.2 as the separation between NH and MH mode in this study. The contribution of the NH mode in the HGF-PDF decreased with increasing particle diameter, with the average number fraction of NH (NF_{NH}) decreasing monotonically from 66 % to 28 % as particle diameter increased from 50 nm to 150 nm (Fig. 3a and Table 1). The peak values of the NH mode remained around 1.06, independent of particle diameter. The peak values of the MH mode increase from 1.33 to 1.52 as particle size increased from 50 nm to 150 nm. The MH modes for 150 nm particle sizes were narrower than other particles (Fig. 3a), indicating the uniformity in particle compositions of this mode. Previous study has found that the proportion of the NH mode is higher in Beijing than that in Shanghai and Guangzhou, likely due to higher number of vehicles, resulting from higher emissions of hydrophobic primary aerosols in Beijing (Ye et al., 2013). The strength of the MH mode increased with particle size, indicating that the proportion of strongly hygroscopic particles gradually increases with particle size. MH-mode particles are mainly from secondary formation or the aging of primary particles (Wang et al., 2019).

The averaged VSF-PDFs typically showed a minimum at VSF around 0.8 (Fig. 3b), thus we use VSF = 0.8 as the separation between NV and VV mode. The contribution of the NV mode in the VSF-PDF increased with particle diameter, and the average number fraction of NV (NF_{NV}) increased from 8 % to 14 % (Fig. 3b). This trend aligns with observations in Beijing and Shanghai (Jiang et al., 2018; Wehner et al., 2009). Particles in the NV mode contained non-volatile species, likely BC, highly oxidized organics or polymers (Hong et al., 2014). Larger particles showed stronger NV modes, which are significantly influenced by highly oxidized organics aerosols, and volatility is generally inversely correlated with the oxidation state (Wang et al., 2017a). The size-resolved VSF-PDFs exhibited an NV mode peak near 0.95 and a broad, approximately symmetric VV mode peak ranging between 0.47 and 0.52 for all measured particle diameters. Unlike the MH mode, the peak values in the VV mode showed no clear dependency on diameter. Additionally, NF_{NV} was observed to be lower than NF_{NH} in this study, indicating a significant presence of hydrophobic but volatile particles. Previous researches suggest that NF_{NH} and NF_{NV} are predominantly influenced by the same component, associated with newly emitted combustion particles (Liu et al., 2013; Cheung et al., 2016; Zhang et al., 2016). The lower proportion of NF_{NV} observed in this study could be attributed to Beijing's recent restrictions on coal burning, which have led to a reduction in non-volatile and non-hygroscopic carbon emissions (Hu et al., 2024). Currently, traffic emissions are the primary pollution source in



Beijing, characterized by freshly emitted organics from vehicles with a low oxidation state, high volatility, and low hygroscopicity (Tiitta et al., 2010; Feng et al., 2023; Liu et al., 2023).



265 **Figure 3.** Size-resolved mean hygroscopicity growth factor probability density functions (HGF-PDFs) (a) and volatility shrink factor probability density function (VSF-PDFs) (b) for particles ranging from 50 nm to 150 nm throughout the sampling period.

Table 1. Summary of the hygroscopicity and volatility parameters throughout the sampling period, as well as during clean and pollution periods.

	Particles	50 nm	80 nm	110 nm	150 nm
	property	Mean ± SD	Mean ± SD	Mean ± SD	Mean ± SD
Total	N _{FNV}	0.08 ± 0.06	0.11 ± 0.06	0.13 ± 0.06	0.14 ± 0.07
	VSF	0.51 ± 0.05	0.55 ± 0.04	0.56 ± 0.05	0.56 ± 0.07
	N _{FNH}	0.66 ± 0.20	0.47 ± 0.20	0.35 ± 0.18	0.28 ± 0.17
	HGF	1.15 ± 0.07	1.24 ± 0.08	1.30 ± 0.09	1.36 ± 0.10
	κ	0.062 ± 0.032	0.103 ± 0.043	0.138 ± 0.050	0.174 ± 0.058
	σ _{HGF-PDF}	0.15 ± 0.03	0.19 ± 0.04	0.20 ± 0.04	0.20 ± 0.04
Clean	N _{FNV}	0.06 ± 0.05	0.12 ± 0.06	0.15 ± 0.06	0.17 ± 0.07
	VSF	0.49 ± 0.05	0.55 ± 0.05	0.57 ± 0.06	0.57 ± 0.07
	N _{FNH}	0.68 ± 0.19	0.57 ± 0.20	0.47 ± 0.19	0.40 ± 0.18
	HGF	1.15 ± 0.06	1.20 ± 0.09	1.25 ± 0.10	1.31 ± 0.11
	κ	0.058 ± 0.029	0.084 ± 0.044	0.112 ± 0.054	0.142 ± 0.061
	σ _{HGF-PDF}	0.14 ± 0.04	0.19 ± 0.04	0.21 ± 0.04	0.23 ± 0.04
Pollution	N _{FNV}	0.10 ± 0.07	0.12 ± 0.05	0.11 ± 0.04	0.09 ± 0.03
	VSF	0.56 ± 0.05	0.58 ± 0.04	0.57 ± 0.05	0.55 ± 0.05
	N _{FNH}	0.67 ± 0.21	0.40 ± 0.16	0.23 ± 0.10	0.14 ± 0.06
	HGF	1.15 ± 0.07	1.28 ± 0.07	1.36 ± 0.05	1.44 ± 0.04
	κ	0.060 ± 0.033	0.121 ± 0.037	0.172 ± 0.032	0.218 ± 0.029
	σ _{HGF-PDF}	0.15 ± 0.03	0.19 ± 0.02	0.18 ± 0.03	0.16 ± 0.03

270 **4.3 Diurnal variations of single hygroscopicity parameter and VSF_{mean}**

Figure 4a and 4b display the diurnal variation of the mean single hygroscopicity parameter (κ_{mean}) for particles of different sizes during the clean and the pollution periods, respectively. The κ_{mean} shows similar diurnal patterns



during both periods, increasing in the afternoon and early morning, and decreasing at noon and during the evening rush hours. However, peak values of κ_{mean} occurred at different times: in the late afternoon (15:00-16:00 LT) during the clean period (Fig. 4a) and in the early morning (05:00-06:00 LT) during the pollution period (Fig. 4b). This discrepancy is likely due to an increase in aged particles from daytime photochemical reactions during the clean period, consistent with findings from the same sampling site in 2018 (Zhang et al., 2023). During the pollution period, the early morning peak in κ_{mean} can be attributed to more active heterogeneous hydrolysis reactions of N_2O_5 , which rapidly convert to hygroscopic nitrates at lower temperatures (Zhou et al., 2018; Zhao et al., 2017), coupled with lower wind speeds and higher humidity that enhance condensation on aerosol particles. This results in the transformation of low-hygroscopicity primary particles into more hygroscopic and internally mixed aerosol particles (Fan et al., 2020b; Zhang et al., 2016). The two low κ_{mean} values occurred at noon (12:00-13:00 LT) and evening rush time (19:00-20:00 LT) (Fig. 4). During these periods, the average organic mass fraction of PM_{10} increased (Fig. S3). The decrease in hygroscopicity around noon was more pronounced for Aitken mode particles compared to accumulation mode particles, which could be attributed to the cooking emissions. HGF values of 1.13 and 1.17 were observed for 50 nm and 80 nm particles respectively at noon during the clean period, while HGF values of 1.10 and 1.25 were observed during the pollution period. These findings are consistent with the previous results that reported HGF value of 1.17 ± 0.11 for 40 nm and 80 nm particles during cooking activities in Beijing (Ren et al., 2023). The lowest particle hygroscopicity was observed at evening rush time, likely due to the reduction in the boundary layer at night, which traps local pollutants from cooking and vehicle emissions, leading to decreases in aerosol hygroscopicity (Zhang et al., 2016).

Figures 4c and 4d show notable differences in the diurnal variations of κ -PDF for 50 nm aerosol particles during clean and pollution periods. During the clean period, coinciding with the NPF event at approximately 10:00 LT (Fig. S4), the number fraction of the MH mode for 50 nm particles increased (Fig. 4c). This suggests that intense photochemical reactions generate condensable species such as sulfuric, nitrate, and organics, which condense onto pre-existing aerosol particles, transforming them from externally mixed to internally mixed states (Wang et al., 2019). The MH mode was predominantly observed in the afternoon (14:00-17:00 LT). After 17:00 LT, the number fraction of NH mode particles increased, peaking at 20:00 LT (Fig. 4c and Fig. S5), primarily due to substantial traffic emissions during evening rush hour. During the pollution period, the κ -PDF of 50 nm aerosol particles exhibited a bimodal distribution throughout the day. A higher number fraction of MH mode particles was observed in the early morning (Fig. 4d), likely attributable to the intensified heterogeneous reactions of NO_x during this time.



For 150 nm particles, the κ -PDF showed a similar diurnal variation pattern during both clean and pollution periods, displaying a bimodal distribution throughout the day (Fig. 4e and 4f). During the clean period, the number fraction of MH for 150 nm particles was higher during the daytime than that at nighttime (Fig. 4e). As the sun rises, the elevation of the boundary layer allows more solar radiation to reach the ground, resulting in enhanced photochemical reactions and consequently an increase in aged particles, thereby enhancing hygroscopicity of particles and the number fraction of MH. Influenced by local primary emissions, the number fraction of NH increased after 18:00 LT (Fig. 4e). During the pollution period, the MH mode for 150 nm particles was stronger and more stable throughout the day compared to the clean period (Fig. 4f). Additionally, the MH mode for 150 nm particles was narrower and more concentrated than that during the clean period, as the mean HGF-PDF shown in Fig. S6. The proportion of number fraction of NH consistently remained below 25 % throughout the day (Fig. S5). This could be attributed to the heterogeneous NO_x reactions on the surface of aerosol particles, which significantly contribute to hygroscopicity by increasing the proportion of inorganic materials (Sun et al., 2016; Wang et al., 2016). AMS results show that nitrate is the main inorganic component of PM_1 (Fig. 8a), further supporting this viewpoint. The average nitrate mass fraction in PM_1 increased from $13 \% \pm 8 \%$ before P_{III} to $33 \% \pm 4 \%$ during the P_{III} , indicating that the increase of ammonium nitrate is the main cause of heavy pollution in Beijing in autumn. In recent years, with the implementation of strict air pollution control strategies, there has been a significant reduction in the concentrations of gaseous pollutants, especially SO_2 (Wang et al., 2022; Chu et al., 2020). During pollution events in Beijing, nitrates have replaced sulfates as the predominant inorganic component (Chen et al., 2019; Wang et al., 2019; Wang et al., 2022).

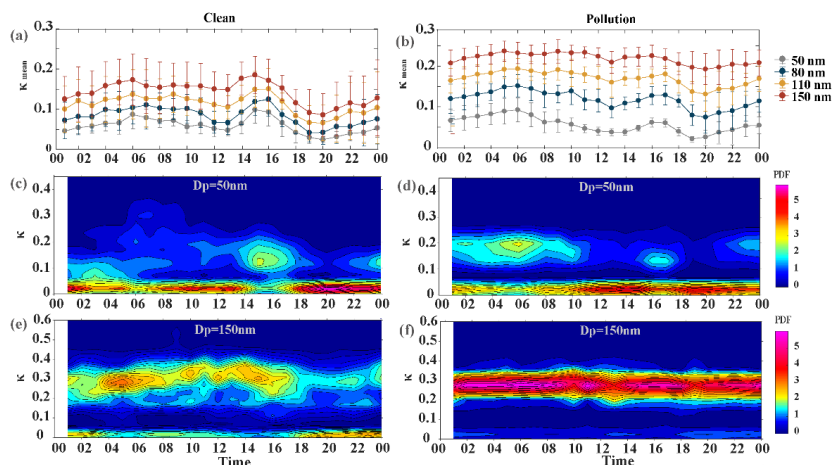




Figure 4. Diurnal variation of mean κ for different diameters and the error bars represent the standard deviations (a, b); κ -PDF for particles with $D_p = 50$ nm (c, d) and κ -PDF for particles with $D_p = 150$ nm (e, f) during the clean and pollution periods, respectively.

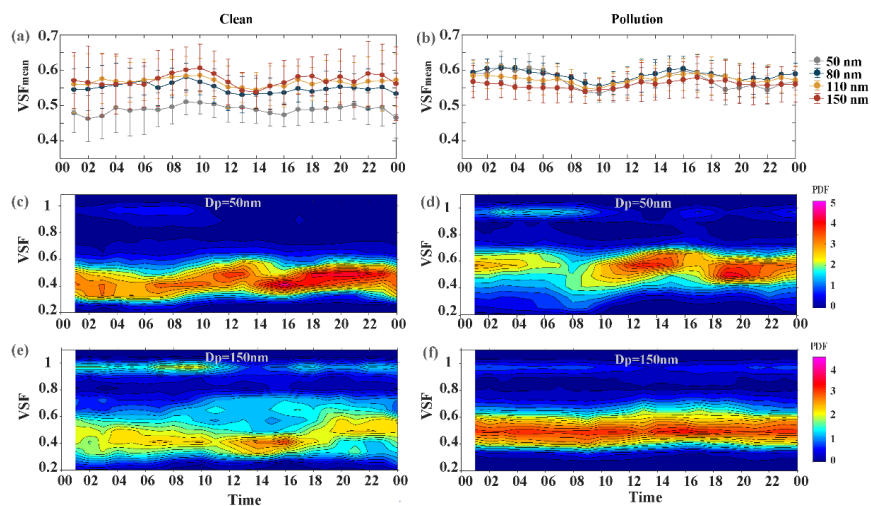
325 Figure 5a and 5b show the significant differences in the diurnal variation of VSF_{mean} for particles with diameters ranging from 50 to 150 nm during two different periods. During the clean period, the highest VSF_{mean} was observed in the morning (09:00-10:00 LT) and the lowest in the afternoon. Conversely, during the pollution period, the highest VSF_{mean} was observed in the afternoon (15:00-16:00 LT) and the lowest observed in the morning for all particle sizes. During the clean period, particle volatility increased dramatically around 10:00 LT (VSF_{mean} decreased) along
330 with the occurrence of NPF events, indicating that the newly formed matter was more volatile. The VSF_{mean} for 50 nm particles was noticeably lower than that for larger particles throughout the day. The volatility of 80 nm, 100 nm, and 150 nm particles reached peaked around 14:00 LT (lowest VSF) and two hours later for 50 nm particles. During the pollution period, the volatility of different diameters exhibited similar diurnal variation patterns (Fig. 5b). In addition, the variability in volatility for Aitken mode particles was greater than that for accumulation mode particles,
335 consistent with findings from previous studies (Wang et al., 2017a; Zhang et al., 2016). The highest volatility for 50 nm and 80 nm particles occurred during the morning and evening rush hours, while it was lower in the afternoon and at night. This result was likely influenced by traffic emissions and meteorological conditions. The CAMS station, located approximately 200 m from a major road to the west, indicated significant influence from traffic emissions. The median diameter range of exhaust particles from gasoline vehicles is between 55-73 nm, with average diameter
340 of 65 nm (Momenimovahed and Olfert, 2015). Thus, the volatility of 50 nm and 80 nm aerosol particles was particularly impacted by traffic emissions in this study. Additionally, the lower boundary layer height and wind speed during the pollution period hindered the dispersion of primary emissions. Several studies have demonstrated that the volatility of organic compounds from traffic emissions decreases as the oxidation level increases (Huffman et al., 2009; Cao et al., 2018; Zhu et al., 2021; Yang et al., 2023). Therefore, the reduced volatility of particles in the
345 afternoon and at night could be due to an increased oxidation degree of organic component, resulting in the production of more refractory organics (Cheung et al., 2016; Wang et al., 2017a; Wang et al., 2022).

Figure 5c-5f show the diurnal variation in VSF-PDFs for 50 nm and 150 nm particles during the clean and pollution periods, indicating that all VSF-PDFs exhibited both NV mode and VV mode. Notably, external mixing was more apparent during the night and early morning, especially for 150 nm particles during the clean period and
350 50 nm particles during the pollution period (Fig. 5d and 5e). This phenomenon could be attributed to the reduced boundary layer height, which leads to the accumulation of nonvolatile particulate matter emissions (e.g. BC, soot



aggregates) from cooking or vehicles emissions. Diesel trucks are permitted to enter Beijing's 5th Ring Road from 00:00 to 06:00 LT (Beijing Municipal Bureau of Transportation, 2017), resulting in elevated nonvolatile particulate matter emissions during the night and early morning (Hakkaraian et al., 2023). As the boundary layer develops during the day, promoting vertical mixing, the number fraction of NV mode particles decreases. During the clean period, the number fraction of NV mode for 50 nm particles decreases around 10:00 LT (Fig. 5c), as these newly formed particles are primarily composed of volatile substances such as sulfates, ammonium, and organic matter during NPF days (Yue et al., 2010; Wu et al., 2016; Yang et al., 2021). During the pollution period, external mixing of 50 nm particles with volatile (e.g., organic compounds) and non-volatile matter primarily due to local primary emissions (Wehner et al., 2009; Wang et al., 2019).

The VSF-PDF of 150 nm aerosol particles during the pollution period exhibits a stronger and more stable VV mode, and a lower NV mode compared to the clean period (Fig. 5e and 5f). During the pollution period, primary emitted particles can be rapidly coated and aged through heterogeneous reactions (Hakala et al., 2016; Wang et al., 2016), causing the non-volatile particles to transition from an externally mixed mode with volatile particles to an internally mixed mode (Wehner et al., 2009; Yue et al., 2009; Hong et al., 2018). During the clean period, the active aging process in the afternoon promotes the mixing of primary particles with secondary particles, increasing the number fraction of VV mode for 150 nm particles. Influenced by primary emissions, the volatility of VV mode began to weaken after 18:00 LT.

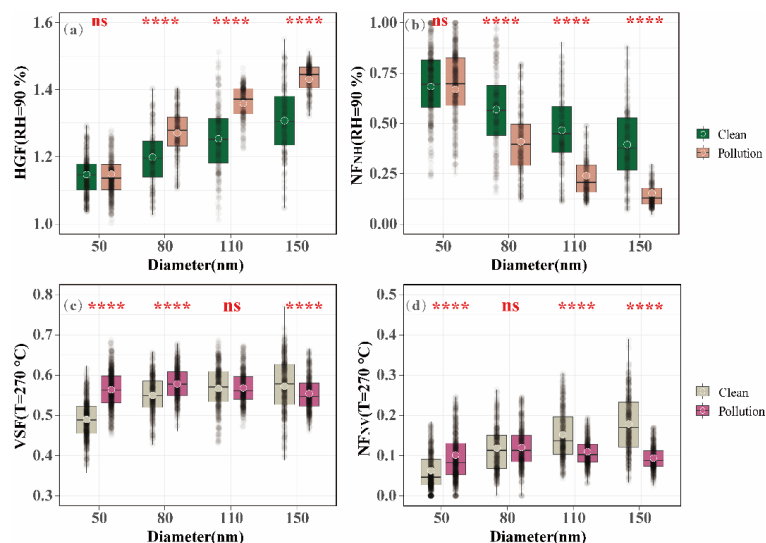


370 **Fig. 5.** Diurnal variation of mean VSF for different diameters and the error bars represent the standard deviations (a, b); VSF-PDF for particles with $D_p = 50$ nm (c, d) and VSF-PDF for particles with $D_p = 150$ nm (e, f) during the clean and pollution periods, respectively.



4.4 Size-resolved hygroscopicity and volatility during the clean and pollution periods

Figure 6 shows the statistics of size-resolved hygroscopicity and volatility, as well as the number fraction of
375 nearly hydrophobic mode particles ($N_{F_{NH}}$) and non-volatile mode particles ($N_{F_{NV}}$) during the clean and pollution
periods. As particle diameter increases, both the mean and median values of the HGF showed noticeable increases
during both periods. Furthermore, for particles ranging from 80 nm to 150 nm, the hygroscopicity exhibited
significant increases ($p < 0.0001$) during the pollution period compared to the clean period (Fig. 6a). The mean HGF
increased from 1.15 to 1.31 for particle sizes ranging from 50 nm to 150 nm during the clean period and from 1.15
380 to 1.44 during the pollution period (Table 1). In contrast to HGF, both the mean and median values of $N_{F_{NH}}$
decreased with increasing particle size (Fig. 6b). The $N_{F_{NH}}$ for particles ranging from 80 nm to 150 nm exhibited
significant decreases ($p < 0.0001$) during the pollution period compared to the clean period. The mean $N_{F_{NH}}$
decreased from 0.68 to 0.40 and from 0.67 to 0.14 for particle sizes ranging from 50 nm to 150 nm during the clean
period and the pollution period, respectively (Table 1). Larger particles under the pollution period typically
385 comprised more inorganic salts, enhancing their hygroscopicity. The variation in parameters for different particle
sizes was much larger during the clean period than during the pollution period, as shown in Fig. 6. Additionally, we
found more external mixed larger particles during the clean period, as indicated by larger $\sigma_{HGF-PDF}$ values (Table 1).
Thus, the chemical composition may deviate significantly during the clean period. During the clean period, NPF
events occur frequently, and larger particles are likely formed by the condensation of different substances (such as
390 sulfate, ammonium, and organics) on pre-existing particles, leading to a more complex chemical composition
(Cheung et al., 2016; Wang et al., 2017a). However, during the pollution period, the larger particles are more aged
and contain a higher inorganic fraction, resulting the chemical composition more homogeneous (Wehner et al., 2009;
Sun et al., 2016; Wu et al., 2016).



395 **Figure 6.** Boxplot of (a) size-resolved hygroscopic growth factor (HGF) and (b) the number fraction of nearly hydrophobic mode particles (NF_{NH}) during the clean and pollution periods. (c) (d) same as (a) (b) but for volatile shrink factor (VSF) and number fraction of non-volatile mode particles (NF_{NV}), respectively. The two borders of box indicate the 25th and 75th percentiles, the line in box represents the median, whiskers indicate the 5th and 95th percentiles, and the circle in each box indicates the mean. Ns means the observed difference is not statistically
400 significant, **** means P value ≤ 0.0001 .

In terms of aerosol volatility, the mean and median values of VSF slightly increased with particle size during the clean period, indicating reduced volatility (higher VSF corresponds to lower volatility); However, no distinct trend was observed during the pollution period (Fig. 6c). The NF_{NV} followed a similar trend to VSF (Fig. 6d). During the clean period, the mean VSF increased from 0.49 to 0.57, and NF_{NV} increased from 0.06 to 0.17 for particles
405 ranging from 50 nm to 150 nm. However, during the pollution period, the mean VSF was around 0.56 for all selected sizes, with less than 5% variability among different particle sizes (Table 1). Notably, Aitken mode particles exhibited higher volatility and a lower NF_{NV} during the clean period compared to the pollution period (Fig. 6c and 6d). This difference was more pronounced for 50 nm particles compared to 80 nm particles. For accumulation mode particles, the situation was opposite, and the difference between two periods for 150nm particles was greater than for 110 nm
410 particles. Previous research has shown that particles smaller than 100 nm exhibit lower NF_{NV} on nucleation days than non-nucleation days, while results are reversed for particles larger than 100 nm (Wehner et al., 2009). During the clean period, nucleation events primarily contribute to the volatility of Aitken mode particles, while the growth of larger particles is mainly due to the coating effect of condensed vapors (primarily organic vapors) on pre-existing particles during NPF events (Wang et al., 2017a). Wu et al. (2017) suggest that during NPF events, the increase in



415 non-volatile materials with particle size is caused by photochemically induced nucleation of condensing vapors or
rapid chemical transformations occurring within the particle phase. The volatility observed in this study is higher
than previous results from Beijing in 2015, where Wang et al. (2017) reported that the VSF of particles ranging from
40 nm to 150 nm was between 0.6 to 0.72 during the clean period and 0.65 during the pollution period. This increase
in volatility could be attributed to the continuous reduction in BC concentration, which has decreased significantly
420 from $6.25 \pm 5.73 \mu\text{g m}^{-3}$ in 2012 to $1.80 \pm 1.54 \mu\text{g m}^{-3}$ in 2020, and further to $0.79 \pm 0.63 \mu\text{g m}^{-3}$ in 2022, due to the
implementation of the Clean Air Action Plan in 2013 (Sun et al., 2022; Hu et al., 2024).

4.5 Relationship of hygroscopic and volatile properties

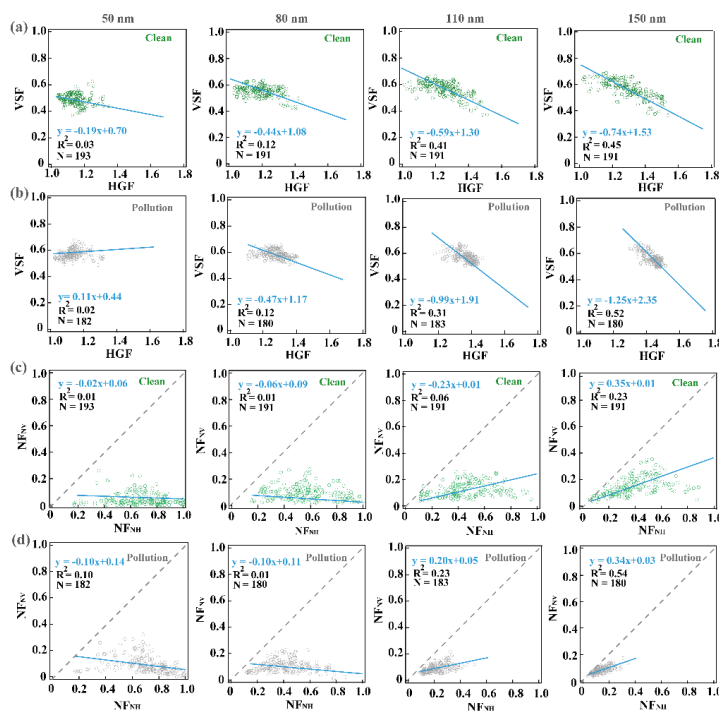
In order to understand the relationship between hygroscopicity and volatility of particles, we investigated the
correlation between the HGF and VSF, as well as NF_{NH} and NF_{LV} across different aerosol size during the clean and
425 pollution periods. As shown in Fig. 7a and 7b, except for 50 nm particles, VSF exhibited a decreasing trend with
increasing HGF for all selected particle sizes during both periods, indicating that volatility increases with an increase
in hygroscopicity. This trend suggests that higher hygroscopic and volatile components within aerosol particles are
predominantly inorganics. Furthermore, the correlations between HGF and VSF were size-dependent, exhibiting
stronger correlation and higher absolute slope values for larger particles, likely due to their higher proportion of
430 inorganic components. Previous study has shown that as particle size increases, the inorganic mass fraction increases
from approximately 35% to 70% (Wu et al., 2016). Another notable observation from Figures 7a and 7b is that the
hygroscopicity and volatility of particles during the pollution period was less scattered compared with those during
the clean period. This could be attributed to higher degree of aging and more uniform chemical composition of
particles during pollution period. However, 50 nm particles, primarily consisting of freshly emitted particulates,
435 exhibit significant chemical composition variability across different emission sources (Wang et al., 2019).

To further understand the relationship between hygroscopicity and volatility, we also investigated the
correlation between NF_{NH} and NF_{NV} during the clean and pollution periods (Fig. 7c and 7d). For submicron particles,
those remaining non-volatile particles after heated at 300 °C are generally considered as nearly hydrophobic
particles, primarily consisting of soot and carbonaceous particles (Massling et al., 2009; Wehner et al., 2009; Zhang
440 et al., 2016). Our result shows a weak correlation between NF_{NH} and NF_{NV} , especially for Aitken mode particles.
The correlation coefficients (R^2) between NF_{NH} and NF_{NV} for particles with diameters of 50 nm, 80 nm, 110 nm and
150 nm were 0.01, 0.01, 0.06, 0.23, and 0.10, 0.01, 0.23, 0.54 during the clean and pollution periods, respectively.
These findings are consistent with previous research in Beijing (Wang et al., 2017a). However, the R^2 values from
our study are lower than those for same diameter particles in the suburban Xiang He station. This discrepancy is



445 influenced by presence of more soot and local light-absorbing carbonaceous aerosols in Xiang He, and these materials are neither hygroscopic nor volatile (Zhang et al., 2016). As shown in Fig.7c and 7d, all the points are below the 1:1 line, indicating that a certain proportion of nearly hydrophobic particles are volatile. NF_{NV} for all particles were lower 0.4, NF_{NH} for Aitken mode particles mainly ranged from 0.2 to 1.0, and for accumulation mode particles, it was focused below 0.8 and 0.4 during the clean and pollution periods, respectively. This indicates that

450 NF_{NH} and NH_{NV} for accumulation mode particles were composed of similar components during the pollution period, likely BC and soot. A larger proportion of hydrophobic and volatile particles was found in the Aitken mode, which can be attributed to particles from NPF events and traffic emissions. These particles contain a variety of organic materials that are typically highly volatile but exhibit low hygroscopicity (Cai et al., 2022; Feng et al., 2023).



455 **Figure 7.** Relationship between the hygroscopic growth factors (HGFs) and the volatility shrink factors (VSFs) during the clean and pollution periods for different dry particle diameters (a, b); relationship between the NF_{NH} and the NF_{NV} during the clean and pollution periods (c, d) for different dry particle diameters.

4.6 Hygroscopicity of non-volatile core and volatile coating

For ambient aerosols, most compounds can volatilize at temperatures of 300 °C (Hong et al., 2014). However,

460 certain compounds such as BC, sea salt, and crustal materials do not volatilize even at higher temperatures. Previous researches have shown that submicron atmospheric aerosols contain non-volatile components other than BC, possibly including organic polymers (Backman et al., 2010; Häkkinen et al., 2012; Hong et al., 2014). Even newly

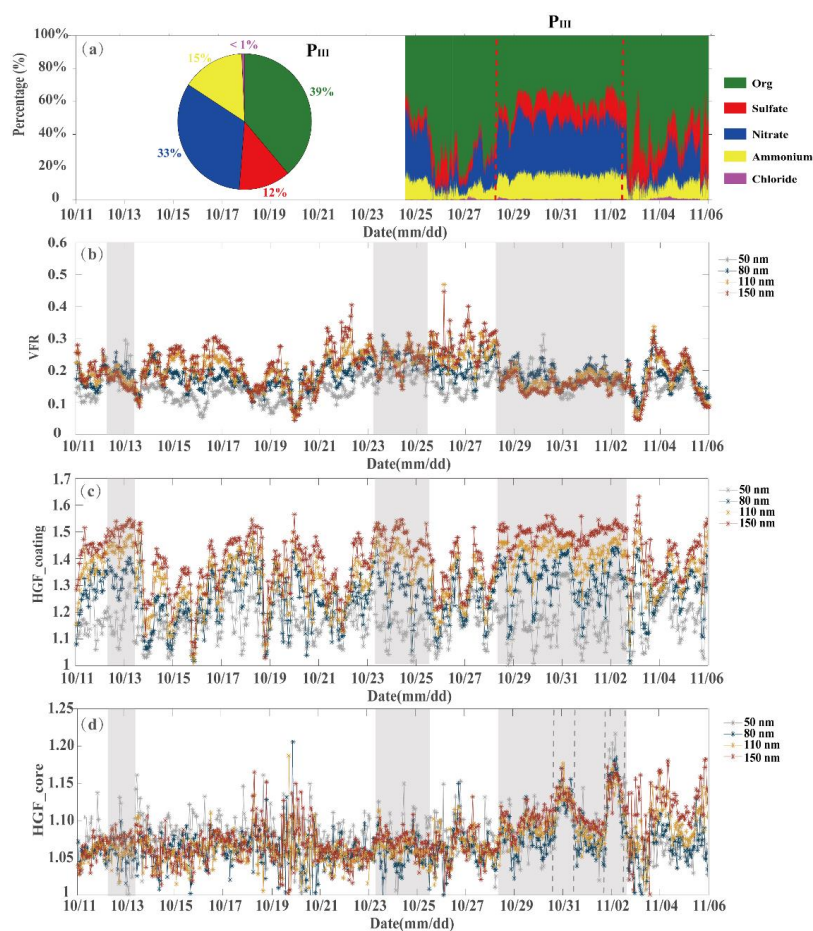


nucleated particles also retain a core that does not volatilize at 300 °C (Wehner et al., 2005; Ehn et al., 2007a; Wu et al., 2017). This section presents the measured hygroscopic growth of the non-volatile core after heated at 270 °C during the sampling period. Based on the simultaneously measured hygroscopic growth factor of particles and their non-volatile core, we also calculated the hygroscopic growth factor of volatile coatings under the assumption that the particles are a two-component system composed of a non-volatile core and a volatile coating. According to the Zdanovskii - Stokes - Robinson (ZSR) relation (Stokes and Robinson, 1966), the HGF of a mixed particle can be estimated from the HGFs of its components according to their volume fractions (see Gysel et al., 2009; Swietlicki et al., 2008): $HGF_{\text{mix}} = \sum_i \varepsilon_i HGF_i$. Therefore, the observed hygroscopic growth factor of particles is calculated as follows: $HGF_{\text{measured}}^3 = \text{VFR} * HGF_{\text{core}}^3 + (1 - \text{VFR}) * HGF_{\text{coating}}^3$, where HGF_{measured} , HGF_{core} , and HGF_{coating} are the hygroscopic growth factors for dry particles, the non-volatile core and the volatile coating, respectively. The VFR (=VSF³) is the volume fraction ratio of the non-volatile core, and 1-VFR is the volume fraction ratio of the volatile coating. The VFR intuitively shows the variations of volume fraction of non-volatile core for different size particles (Fig. 8b). The mean VFR of particles ranging from 50 nm to 150 nm was 0.15 ± 0.05 , 0.19 ± 0.05 , 0.20 ± 0.06 and 0.20 ± 0.07 , respectively. During the pollution period, especially during P_{III}, the VFR exhibits minimal variation. The mean HGF_{coating} for particles ranging from 50 nm to 150 nm was 1.17 ± 0.08 , 1.27 ± 0.10 , 1.35 ± 0.10 and 1.41 ± 0.10 , respectively, which was 2% -7% higher than mean HGF for corresponding particles size. As shown in Fig. 8c, variations in HGF_{coating} basically similar to those hygroscopicity of unheated particles (HGF_{mean}) (Fig.S7) and exhibit significant size dependency. It is clear that the particle coating exhibits higher hygroscopicity during the pollution period. Figure 9 shows that HGF_{coating} was influenced by the volume fraction of the coating in particles during the pollution period. The correlations between the HGF_{coating} and the volume fraction of the coating in particles were stronger for the largest particles, suggesting that these particles were typically highly aged and had a higher proportion of inorganic components with a greater volume proportion of the coating. AMS data also show that the relative mass fraction of inorganic components in the particulate phase significantly increases during the pollution period (Fig. 8a). The nitrate content dominated the inorganic components of aerosol particles during the pollution period, reaching 33 % during the P_{III} and likely being the major contributor to the increase in HGF_{coating} .

The mean HGF_{core} for particles ranging from 50 nm to 150 nm in this study was 1.08 ± 0.03 , 1.07 ± 0.03 , 1.07 ± 0.03 , and 1.09 ± 0.04 , respectively. Unlike HGF_{coating} , HGF_{core} showed no size-dependent variation, likely due to the similar proportions of residual hygroscopic components in particles across different sizes (Fig. 8d). Given that BC and mineral dust are not hygroscopic, the non-volatility fraction of particles still contains other hygroscopic components, likely highly oxidized organic polymers (Liu et al., 2021; Hong et al., 2014). Studies on forest aerosols



have shown that the mass fraction of non-BC material accounts for 13 %–45 % of the total remaining mass after heated at 280 °C, correlating significantly with mass fraction of organic nitrates (Häkkinen et al., 2012). In Beijing, 495 the organic nitrates were effectively quantified to contribute a notable fraction to organic aerosol (OA), namely 9 %–20 % in autumn (Yu et al., 2019). Additionally, such non-volatile organic compounds may also be related to humic-like substances (Wu et al., 2009), which are significant components of fine particulate matter in Beijing (Ma et al., 2018). These oxygenated organic aerosols do not volatilize even after heated to 270 °C, remaining as part of the non-volatile components contributing to the core's hygroscopicity.



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Figure 8. Time series of the main mass fraction of chemical components in PM₁ during the sampling period based on the available AMS data (a), The volume fraction of non-volatile core (VFR) (b), calculated HGF_{coating} (c) and measured HGF_{core} (d) for particles from 50 nm to 150 nm during the sampling period. The shadows represent the pollution period and the dashed lines represent the period during which the hygroscopicity of non-volatile core was 505 relatively high.

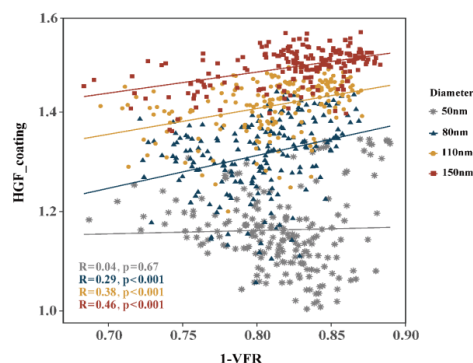


Figure 9. Correlations between the hygroscopic growth factor of volatile coating and its volume fraction for different particle sizes during the pollution period.

510 Figure 8d shows two significant increases in the hygroscopicity of the non-volatile core during periods of severe pollution. The first increase, referred to as event I, occurred from 30 October, 15:00 to 31 October, 15:00 LT. The second, event II, occurred from 1 November, 17:00 to 2 November, 17:00 LT. To further investigate these events, we analyzed the 48 h backward trajectories reaching the sampling site at a height of 500 m above ground level (Fig. 10). During event I, the air mass originated from Shanxi (Fig. 10a). Four hours later, the air mass passed through Shandong, Bohai Sea, and Tianjin before reaching the sampling site, eventually shifting to northwest Mongolian at the end of event I. The early stage of event II was initiated by a short-distance air mass from Hebei (Fig. 10b). Subsequently, the air mass shifted to the Bohai Sea, Shanxi, and then transitioned into a long-range, clean air mass from the northwest at the end of event II. During these two events, part of air masses reached the sampling site either passing through or originating from the Bohai Sea. Moreover, these air masses traveled below 1 km at lower speeds, enhancing the possibility of marine aerosol entering the atmosphere (Fig. 10c and 10d). Thus, the observed increases in the hygroscopicity of non-volatile core particles were probably influenced by marine aerosol.

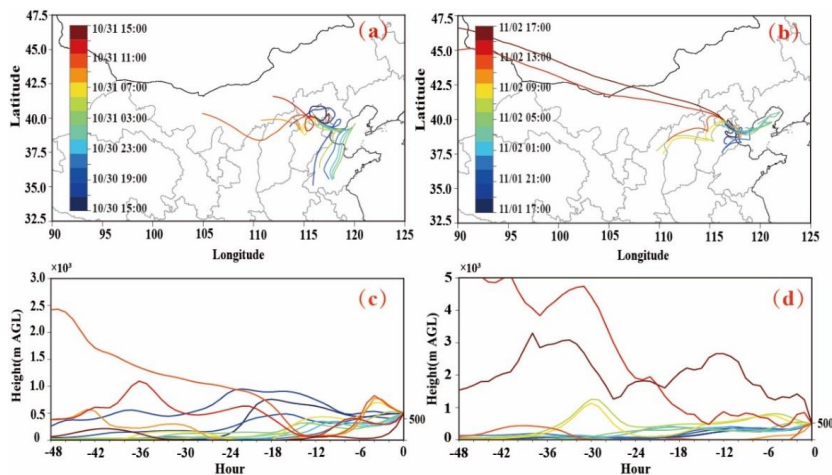
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Figure 11 displays the VSF-PDF and VHGF-PDF for 150 nm particles during the periods when the air mass passed through (30 October, 21:00 to 31 October, 03:00 LT in event I) and originated from the Bohai Sea (2 November, 01:00 to 07:00 LT in event II), respectively. The top axis of the plot denotes the particle diameter. The shift of VHGF-PDF relative to VSF-PDF was attributed to the hygroscopic growth of non-volatile core. During the air mass passed through the Bohai Sea in event I, the diameter of 150 nm particles was reduced to 75 nm (D_p (270 °C, RH_{dry})) after heating (Fig. 11a). Subsequently, 75 nm particles exhibited a trimodal distribution in the VHGF-PDF. This distribution comprised a nearly hydrophobic mode growing to 77 nm (D_{p1} (270 °C, 90 %RH)), a less hygroscopic mode growing to 96 nm (D_{p2} (270 °C, 90 %RH)), and a more hygroscopic mode growing to 135 nm

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530 (D_{p3} (270 °C, 90 %RH)) after rehydration, corresponding to the HGFs of 1.03, 1.28 and 1.80, respectively. When the air mass originated from the Bohai Sea during event II, the 150 nm particles were reduced to 77 nm (D_p (270 °C, RH_{dry})) after heating (Fig. 11b), the VHGF-PDF also exhibited a trimodal distribution after rehydration. The HGFs corresponding to the three peaks were 1.21, 1.41 and 1.81, respectively.



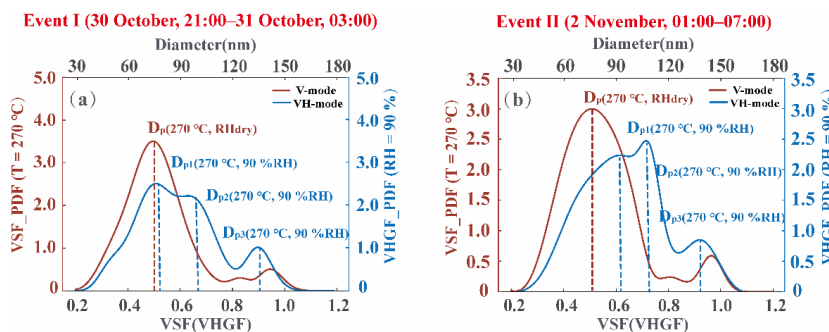
535 **Figure 10.** The 48 h back trajectories arriving at the height of 500 m AGL at the CAMS site (a and b) and change of air mass height of the backward trajectories (c and d) during two events when the hygroscopicity of non-volatile core was relatively high. Event I was from 30 October, 15:00 to 31 October, 15:00 LT and event II was from 1 November, 17:00 to 2 November, 17:00 LT. The trajectory was calculated every 2 h.

Sea-spray aerosol (SSA) consists of particles directly produced at the sea surface and suspended in the air, with sizes ranging from less than 10 nm to several millimeters (Bates et al., 2012; De Leeuw et al., 2011), and SSA can even stay in the atmosphere for days (Grythe et al., 2014). Sub-100 nm SSA non-volatile components at temperatures 250-550 °C include sea salt and non-sea salt components, which could be organic fraction, biological materials, BC, and minerals (Xu et al., 2020; Mallet et al., 2016; Cravigan et al., 2015). Even when heated at 550 °C, 29 % to 48 % of nonvolatile organic fraction still remains in 60 nm SSA (Mallet et al., 2016). The HGFs of SSA ranging from 1.0 to 1.1, typically composed of biological organic components or anthropogenic emissions such as BC (Kim et al., 2015), while HGFs between approximately 1.1 and 1.4 are associated with organics aerosols (Gantt and Meskhidze, 2013; Ovadnevaite et al., 2011; Asmi et al., 2010). In the Aitken mode, the HGF of 1.8 in non-volatile marine particles is likely due to the presence of sea salt (Sellegrri et al., 2008). In summary, the non-volatile SSA from these two events likely comprise a mixture of sea salt, non-volatile organic compounds, and BC. Before reaching the sampling site, SSA particles in the air masses were further aged and grown during transport, forming a volatile coating on their surface that reduced their hygroscopic properties. Therefore, no significant increase in

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particle hygroscopicity was found before heating (Fig. S7). However, this study cannot definitively identify the specific components of marine aerosols transported by the air masses due to available measurements.



555 **Figure 11.** Representative case studies of VH-TDMA measurements. The mean VSF-PDF and mean VHGF-PDF of 150 nm particles from 30 October, 21:00 to 31 October, 03:00 LT in event I (11a), and from 2 November, 01:00 to 07:00 LT in event II (11b).

5 Conclusions

Volatility Hygroscopicity Tandem Differential Mobility Analyzer (VH-TDMA) was used to measure the
 560 hygroscopicity and volatility of the aerosol particles (50 nm, 80 nm, 110 nm, and 150 nm) from 11 October to 6
 November 2023, in Beijing. The HGF-PDF and VSF-PDF for different particle sizes displayed a bimodal
 distribution, indicating a clear state of external mixing throughout the study. The mean HGF for particles from 50
 to 150 nm during the sampling period was 1.15 ± 0.07 , 1.24 ± 0.08 , 1.30 ± 0.09 , and 1.36 ± 0.10 , respectively, while
 the mean VSF was 0.51 ± 0.05 , 0.55 ± 0.04 , 0.56 ± 0.05 , and 0.56 ± 0.07 , respectively. We divided the sampling
 565 period into clean and pollution periods to better compare the hygroscopic and volatile properties of particles under
 different weather conditions. The air mass arriving at the sampling site was mainly from the northwest and the south
 during the clean period and pollution period, respectively. Hygroscopicity increased with particle size in both clean
 and pollution periods, with a more pronounced increase during the pollution period, where mean HGF rose from
 1.15 to 1.44, compared to 1.15 to 1.31 during the clean period. In contrast, the mean NF_{NH} decreased with an increase
 570 in particle size, from 0.68 to 0.40 during the clean period and from 0.67 to 0.14 during the pollution period. The
 average nitrate mass fraction in PM_{10} increased from $13 \% \pm 8 \%$ before P_{III} to $33 \% \pm 4 \%$ during P_{III} , reflecting the
 increased hygroscopicity of larger particles could be attributed to their greater age and higher proportion of inorganic
 component. Volatility decreased and NF_{NV} increased slightly with increase of particle size during the clean periods,
 no apparent trend was observed for the pollution period.

575 The κ_{mean} exhibited similar diurnal variations during both clean and pollution periods. However, it peaked in
 the late afternoon during the clean period and in the early morning during the pollution period. This result could be



due to the dominance of photochemical reactions during the daytime in the clean period and stronger heterogeneous reactions at night during the polluted period. The mean VSF displayed the different diurnal variations during the clean and pollution periods. The highest VSF was observed in the morning and the lowest in the afternoon for particles during the clean period. Conversely, the highest VSF in the afternoon with the lowest observed in the morning during the pollution period. During the clean period, increased volatility was observed during daytime due to NPF events, driven by the formation of volatile matter and the coating effect of condensable vapors. During the pollution period, the reduced volatility of particles in the afternoon could be attributed to an enhanced oxidation degree of organic components. Accumulation mode particles exhibited a stronger and less variable VV and MH modes, due to the particles get aged and mixed with newly emitted particles during pollution period. The results of the correlation between the hygroscopicity and volatility of aerosol particles indicated that the correlation between HGF and VSF, $N_{F_{NH}}$ and $N_{F_{LV}}$ was stronger during pollution period. This highlights the significant impact of environmental conditions on aerosol particle properties.

Furthermore, the study calculated the HGF_{core} and $HGF_{coating}$ of particles after heating. The mean $HGF_{coating}$ for particles ranging from 50 nm to 150 nm was 1.17 ± 0.08 , 1.27 ± 0.10 , 1.35 ± 0.10 and 1.41 ± 0.10 , respectively, which is 2%–7% higher than the mean HGF for similarly sized particles. The mean HGF_{core} ranging from 50 nm to 150 nm is 1.08 ± 0.03 , 1.07 ± 0.03 , 1.07 ± 0.03 and 1.09 ± 0.04 , respectively. The HGF_{core} values were significantly influenced by marine aerosols during air mass passages and originating from the Bohai Sea. These findings provide valuable scientific data for comprehensively understanding the hygroscopic and volatile properties of urban aerosols in Beijing. In the future, obtaining information on the hygroscopic properties of low-volatile organics will be important and will aid in interpreting the VH-TDMA results, thus improving our understanding of the chemical properties of submicron aerosol. Additionally, investigating the hygroscopicity of aerosol particles heated at different temperatures and locations will further elucidate the structure and formation mechanisms of aerosol particles.

600 **Data availability.**

The data in the study are available from the authors upon request (jysun@cma.gov.cn).

Competing interests.

The authors declare that they have no conflict of interest.



Author contributions.

605 AYY analyzed the observational data and prepared the figures of the manuscript. JYS designed the study and outlined the manuscript. JYS, AYY and JYL performed the instrument deployment and operation. All co-authors discussed the results and commented on the manuscript. AYY prepared the manuscript with contributions from all authors.

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