

1 Title

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3 **Eddy Covariance Evaluation of Ecosystem Fluxes at a Temperate Saltmarsh in**  
4 **Victoria, Australia Shows Large CO<sub>2</sub> Uptake**

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## 36 Key Points

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38 This is the first study using eddy covariance to measure CO<sub>2</sub> fluxes at an Australian  
39 temperate saltmarsh, revealing temperature and light limitations to CO<sub>2</sub> uptake.

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41 CO<sub>2</sub> fluxes varied seasonally; growing season net ecosystem productivity was 10.54 g CO<sub>2</sub>  
42 m<sup>-2</sup> day<sup>-1</sup>, dropping to 1.64 g CO<sub>2</sub> m<sup>-2</sup> day<sup>-1</sup> in winter.

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44 Productivity at the French Island saltmarsh is high relative to global saltmarsh estimates but  
45 below global mangrove averages.

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## 49 Abstract

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51 Recent studies highlight the important role of vegetated coastal ecosystems in atmospheric  
52 carbon sequestration. Saltmarshes constitute 30% of these ecosystems globally and are the  
53 primary intertidal coastal wetland habitat outside the tropics. Eddy covariance (EC) is the  
54 main method for measuring biosphere-atmosphere fluxes, but its use in coastal environments  
55 is rare. At an Australian temperate saltmarsh site on French Island, Victoria, we measured  
56 CO<sub>2</sub> and water gas concentration gradients, temperature, wind speed and radiation. The  
57 marsh was dominated by a dense cover of *Sarcocornia quinqueflora*. Fluxes were seasonal,  
58 with minima in winter when vegetation is dormant. Net ecosystem productivity (NEP) during  
59 the growing season averaged 10.54 g CO<sub>2</sub> m<sup>-2</sup> day<sup>-1</sup> decreasing to 1.64 g CO<sub>2</sub> m<sup>-2</sup> day<sup>-1</sup> in  
60 the dormant period, yet the marsh remained a CO<sub>2</sub> sink due to some sempervirent species.

61 Ecosystem respiration rates were lower during the dormant period compared with the  
62 growing season (1.00 vs 1.77 μmol CO<sub>2</sub> m<sup>-2</sup> s<sup>-1</sup>) with a slight positive relationship with

63 temperature. During the growing season, fluxes were significantly influenced by light levels,

64 ambient temperatures and humidity with cool temperatures and cloud cover limiting NEP.

65 Ecosystem water use efficiency of 0.86 g C kg<sup>-1</sup> H<sub>2</sub>O was similar to other C3 intertidal

66 marshes and evapotranspiration averaged 2.48 mm day<sup>-1</sup> during the growing season.

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## 68 EGUsphere Topics

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Deleted: During the growing season, fluxes were significantly influenced by light levels, ambient temperatures and humidity. Evapotranspiration peaked at 0.27 mm h<sup>-1</sup>. We cautiously estimate the annual NEP budget at this marsh to be 753 (±112.7) g C m<sup>-2</sup> y<sup>-1</sup> which is similar to carbon uptake by temperate saltmarshes in Europe and within the range measured at some US saltmarshes. This value is higher than the value hypothesised for global saltmarshes of 382 g C m<sup>-2</sup> y<sup>-1</sup> but is only half the mean value estimated for global mangroves.

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84 Emissions, Marine and Freshwater Biogeosciences, Earth System Biogeosciences

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86 Short Summary

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88 Studies show that saltmarshes excel at capturing carbon from the atmosphere. In this study,  
89 we measured CO<sub>2</sub> flux in an Australian temperate saltmarsh on French Island. The temperate  
90 saltmarsh exhibited strong seasonality. During the warmer growing season, the saltmarsh  
91 absorbed on average 10.5 grams of CO<sub>2</sub> from the atmosphere per m<sup>2</sup> daily. Even in winter,  
92 when plants were dormant, it continued to be a CO<sub>2</sub> sink, albeit smaller. Cool temperatures  
93 and high cloud cover inhibit carbon sequestration.

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98 1. Introduction

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100 Despite their relatively small global footprint of 54,650 km<sup>2</sup> (Mcowen et al., 2017), salt  
101 marshes provide a range of ecosystem services, including shoreline protection (Shepard et al.,  
102 2011), nutrient uptake, nursery grounds for fish populations (Whitfield, 2017) as well as  
103 functioning as significant carbon sinks through CO<sub>2</sub> uptake and storage in their organic rich  
104 sediments (McLeod et al., 2011). These ‘blue carbon’ habitats are recognised for their  
105 significant contribution to the global carbon cycle, as coastal wetlands more broadly are  
106 estimated to have accumulated more than a quarter of global organic soil carbon (Duarte,  
107 2017).

108

109 Saltmarshes are a widely distributed intertidal habitat but are floristically divergent globally  
110 (Adam, 2002), such that commonalities in function and form do not extend across  
111 biogeographic realms. US saltmarshes, for example, are extensively dominated by a single  
112 grassy species, *Spartina alterniflora*, as opposed to the dominance of C<sub>3</sub> Chenopodioideae  
113 species in the southern hemisphere (Adam, 2002). Temperate saltmarshes occupy a  
114 latitudinal range spanning from approximately 30° to 60° (Mcowen et al., 2017) and are most  
115 commonly found along protected coastlines such as bays, estuaries, and lagoons, where they  
116 are sheltered from the full force of wave action (Mitsch and Gosselink, 2000). In the  
117 Southern Hemisphere, temperate saltmarshes have a strong Gondwanan element with high  
118 floristic similarity among the marshes of New Zealand, the southernmost coasts of South  
119 America and South Africa and the southern coastlines of Australia (Adam, 1990). These  
120 marshes are often associated with extensive seagrass meadows and mudflats, and in parts of  
121 their range, mangroves, forming complex coastal mosaics (Huxham et al., 2018).

122 Saltmarshes have been heavily degraded across their range, and it is estimated that perhaps  
123 up to 50% of the global saltmarsh area has been lost since 1900 (Gedan et al., 2009),  
124 primarily due to land use change.

125

126 In most areas where they occur, seasonality plays a major role in the functioning of temperate  
127 saltmarshes (Ghosh and Mishra, 2017). These ecosystems experience distinct growing and  
128 dormant seasons, primarily driven by temperature, light availability, and precipitation  
129 patterns (Adam, 2000). During the growing season (typically spring and summer), increased  
130 temperatures and longer daylight hours stimulate plant growth, photosynthetic activity, and

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132 decomposition processes. Photosynthesis typically outpaces decomposition during this  
133 period, resulting in the temperate saltmarsh acting as a net CO<sub>2</sub> sink (Chmura et al., 2003).  
134 Conversely, the dormant season (usually fall and winter) is characterized by cooler  
135 temperatures and shorter days (Adam, 2000; Howe et al., 2010). These factors lead to  
136 reduced plant growth and photosynthetic activity (Adam, 2000) and while decomposition  
137 processes also slow down due to cooler temperatures, CO<sub>2</sub> release through decomposition  
138 often exceeds CO<sub>2</sub> uptake during this period (Artigas et al., 2015). In Australia, saltmarshes  
139 have been assumed to not exhibit seasonality (Owers et al., 2018) despite there being a  
140 scarcity of data on saltmarsh phenology and the implication this untested assumption could  
141 have on carbon budget estimations.

142  
143 Gross primary production (GPP) of saltmarshes is the total amount of CO<sub>2</sub> uptake by plants  
144 through photosynthesis. Respiration (R<sub>e</sub>) leads to a CO<sub>2</sub> flux directed back to the atmosphere  
145 due to all respiration processes occurring within the saltmarsh, involving both autotrophs and  
146 heterotrophs. The difference between these two fluxes is the net ecosystem exchange (NEE).  
147 Saltmarsh ecosystems can act as both sources and sinks of carbon dioxide (CO<sub>2</sub>), influencing  
148 atmospheric CO<sub>2</sub> concentrations (Chmura et al., 2003). However, quantifying their net  
149 exchange remains challenging (Lu et al., 2017) hindering their effective inclusion in Earth  
150 System Models (Ward et al., 2020) and confounding the incorporation of saltmarsh  
151 restoration in emission reduction targets. Eddy covariance (EC) provides a powerful method  
152 for near-continuous, high-frequency monitoring of gas exchange between a vegetated surface  
153 and the atmosphere (Baldocchi, 2003), enabling the determination of net ecosystem exchange  
154 (NEE) of CO<sub>2</sub>, and identifying the forcings that determine how CO<sub>2</sub> fluxes will respond to  
155 global climate change (Borges et al., 2006; Cai, 2011).

156  
157 Previous EC studies in coastal saltmarshes have been focused on the Northern Hemisphere, in  
158 sites in the USA (e.g. Hill and Vargas, 2022; Kathilankal et al., 2008; Moffett et al., 2010;  
159 Nahrawi et al., 2020; Schäfer et al., 2019), France (Mayen et al., 2024), Japan (Otani and  
160 Endo, 2019) and China (Wei et al., 2020) but interest in the southern hemisphere is growing  
161 (Bautista et al., 2023). The NEE values from these studies indicate that there is high inter-site  
162 (as well as interannual, Erickson et al., (2013)) variability in carbon dynamics of saltmarshes,  
163 with a link to species types, salinity, hydrology (Moffett et al., 2010; Nahrawi et al., 2020),  
164 site specific biochemical conditions (Seyfferth et al., 2020) and latitude (Feagin et al., 2020).  
165 While generally considered important carbon sinks (e.g. ranging between 130 to 775 g C m<sup>-2</sup>

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171 yr<sup>-1</sup> in the USA, according to Kathilankal et al. (2008) and Wang et al.(2016) respectively  
172 and globally hypothesised to average 382 g C m<sup>-2</sup> y<sup>-1</sup> (Alongi, 2020), some EC studies  
173 revealed saltmarshes to be net sources of CO<sub>2</sub> to the atmosphere (Vázquez-Lule and Vargas,  
174 2021) especially in temperate saltmarshes that experience long dormant periods.

175  
176 The aim of this study is to estimate CO<sub>2</sub> and water fluxes in a temperate saltmarsh in  
177 Victoria, southern Australia, to better characterise the effect of seasonality and environmental  
178 variables on the saltmarsh CO<sub>2</sub> budgets. This is the first study in an Australian coastal  
179 saltmarsh where CO<sub>2</sub> fluxes are estimated using the EC method.

## 180 2. Methods

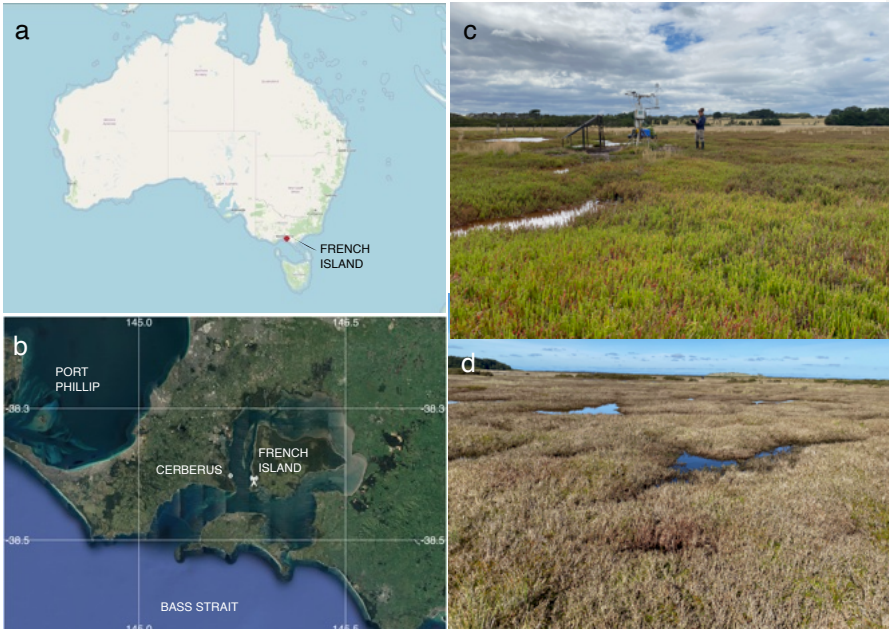
### 181 2.82 Site Description

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183 Ecosystem flux measurements were collected at the Tortoise Head Ramsar coastal wetland on  
184 French Island, Victoria (38.388°S, 145.278°E, Fig. 1) within the Western Port embayment.  
185 French Island is within the Cfb climate zone (temperate oceanic climate) and experiences  
186 distinct seasonal variations in temperature and precipitation. Long term (30 year) climate data  
187 averaged from the nearby Cerberus Station (Australian Bureau of Meteorology, site 86361)  
188 indicated that summers, spanning from December through February, are generally mild to  
189 warm, with maximum temperatures typically ranging from 17°C to 25°C although occasional  
190 heatwaves lead to temporary spikes in temperature that can exceed 30°C. Winters, from June  
191 to September, are cooler, with maximum temperatures ranging between 7°C and 14°C and a  
192 mean minimum temperature of 6°C. Frost is infrequent due to maritime influence, though  
193 crisp mornings below 0°C occur 10% of the time in winter. Rainfall, evenly distributed  
194 throughout the year, averages ca. 715 mm y<sup>-1</sup>, although in 2020 the site ~~led~~ higher than  
195 average rainfall (860 mm y<sup>-1</sup>). The island is exposed to weather patterns influenced by the  
196 Southern Ocean and Bass Strait, leading to occasional storm systems, particularly in winter,  
197 bringing gusty winds and increased precipitation. Western Port has semi-diurnal tides with a  
198 range of nearly 3 m, resulting in wide intertidal flats occupied by mangroves of the species  
199 *Avicennia marina* and saltmarshes.  
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207 Figure 1: a) The location of French Island along the Bass Strait coast of Australia, and b) The  
208 location of the flux tower on French Island as well as the nearby Cerberus meteorological  
209 station (Bureau of Meteorology, Australia), © Google Earth. c) An image of the saltmarsh  
210 within the flux tower footprint during the growing season (with the tower and the author in  
211 the background), taken in February 2020 by Prudence Perry. d) an image of the saltmarsh  
212 during the dormant period, taken at the same location in September 2020 by Ruth Reef.

213

214 The site at French Island is dominated by an extensive temperate coastal saltmarsh  
215 community that is a particularly good natural representation of a broader biogeographic  
216 saltmarsh grouping which covers an area of ca. 7000 ha along Victoria's central coast  
217 embayments (Navarro et al., 2021). While the wetland at the site is a saltmarsh-mangrove-  
218 seagrass wetland system, the footprint of the flux tower was limited to the saltmarsh alone,  
219 which extends more than a kilometre from the shoreline in places. This geography provided  
220 the critical horizontally homogenous area with flat terrain required for ecosystem flux  
221 measurements. Floristically this saltmarsh is species poor, dominated by *Sarcocornia*  
222 *quinqueflora*. Stands of *Tecticornia arbuscula* are common in this saltmarsh, while *Atriplex*  
223 *cinerea* <sup>approx. 7%</sup> *australia* and *Distichis distichophylla* can be prevalent depending on

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225 elevation and soil drainage conditions. *Sarcocornia quinqueflora* is a perennial succulent and  
226 at the temperate ranges of its distribution it has a distinct growing season from October to  
227 May (Fig. 1c) when the stems turn red, followed by a woody and fibrous dormant period  
228 during the colder months of June through September (Fig. 1d). The height of the dominant  
229 vegetation ranged between 0.3 m.

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## 231 2.83 2.2 Data Collection and Analysis

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233 Eddy covariance measurements were made between November 2019 and August 2021  
234 capturing both the saltmarsh growing season (October-~~May~~) as well as a dormant period  
235 (~~June~~-September). An array of standard micro-meteorological instruments included a 3-  
236 dimensional sonic anemometer (CSAT3, Campbell Scientific, USA), an open-path infra-red  
237 carbon dioxide (CO<sub>2</sub>) gas and water vapour (H<sub>2</sub>O) analyser (Li-7500, Li-Cor, USA) and 2  
238 data-loggers. The tower was powered by a solar array with two accompanying 12V DC  
239 storage batteries. The sonic anemometer was mounted 2.3 m above ground. The CO<sub>2</sub>/H<sub>2</sub>O  
240 gas analyser was mounted 0.11 m longitudinally displaced from the anemometer. A CR3000  
241 datalogger (Campbell Scientific, USA), recorded the Li-7500, anemometer, short- and long-  
242 wave radiation (CNR4, Klip & Zonen, the Netherlands), air temperature and humidity (083E,  
243 Met One, USA) readings at 10 Hz frequency. Due to the location of the site in the Bass Strait  
244 (a region that experiences regular winter storms, high wind speeds and higher than national  
245 average cloud cover) the tower sustained damage due to winter storms several times during  
246 the deployment, as well as suffered periods of poor power supply due to short day lengths  
247 and high cloud cover; this was exacerbated by poor accessibility to the remote location during  
248 COVID-19 travel restrictions. The analysis, thus focused on extended periods of continuous  
249 daily records and periods with large gaps in the dataset were removed.

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251 Ecosystem fluxes were calculated for 30 min intervals using Eddy Pro software v.7 (LI-COR  
252 Inc., USA) Express Mode protocols. This processing step includes coordinate axis rotation  
253 correction, trend correction, data synchronisation, statistical tests for quality, density  
254 corrections and spectrum corrections. As part of this step, flux quality flags were assigned to  
255 the calculated CO<sub>2</sub> fluxes using the 0–2 flag policy ‘Mauder and Foken 2004’, based on the  
256 steady state test and the developed turbulent conditions test. The steady state test checks if  
257 fluxes remain consistent over the 30-minute averaging period by comparing the mean and



262 standard deviation (SD) of fluxes in the first and second halves of the period. The developed  
 263 turbulent conditions test ensures turbulence is well-developed and its energy spectra fits the  
 264 Kolmogorov spectrum. Both tests assign partial flags that are combined into a single flag (0–  
 265 2) in Eddy Pro, indicating the overall data quality. Only data that met the criteria of being in  
 266 quality class 0 ('best quality fluxes') for CO<sub>2</sub> flux were chosen for further analysis. We  
 267 further removed anomalous data points defined as values that exceed four SDs from the mean  
 268 CO<sub>2</sub> flux; this resulted in the additional loss of ca. 1% of the dataset. Gap filling was not  
 269 applied. Additional filtering was applied to nighttime data due to known weak convection at  
 270 night, thus CO<sub>2</sub> flux data during periods of atmospheric stability, i.e. when night friction wind  
 271 velocities (u\*) were below 0.2 m s<sup>-1</sup>, were excluded following inspection of the nightly NEE  
 272 vs. u\* curve to detect the threshold where NEE fall-off occurs. 0.2 m s<sup>-1</sup> is the typical  
 273 threshold value used in eddy-covariance studies (Davis et al., 2003). This resulted in a dataset  
 274 of 674 day-time and 606 nighttime flux measurements during the dormant period and 4124  
 275 day-time and 3020 nighttime flux measurements for the growing period (Table 1). The  
 276 growing season dataset included 90 days with 85% or more flux data coverage, while the  
 277 dormant season dataset included 18 days, and these days were used for 24-hour flux  
 278 integrations.

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280 Table 1: Mean (±SD) net ecosystem exchange (μmol CO<sub>2</sub> m<sup>-2</sup> s<sup>-1</sup>) during day- and nighttime  
 281 respectively, as well as the corresponding number of half hourly measurements from each  
 282 month, following filter applications (n). Pink shading indicates the dormant season at the  
 283 French Island saltmarsh.

Month	Daytime Mean NEE (SD); n	Nighttime Mean NEE (SD); n
October 2019	-2.29 (3.08); 121	2.04 (1.28); 70
November 2019	-1.84 (3.89); 151	2.85 (1.75); 110
December 2019	-3.33 (4.59); 96	1.14 (1.70); 15
January 2020	-1.31 (3.31); 63	2.10 (0.79); 27
February 2020	-3.83 (4.11); 540	1.89 (1.10); 280
March 2020	-3.86 (3.90); 494	1.63 (0.78); 351
August 2020	0.05 (2.05); 150	1.76 (1.22); 39
September 2020	-0.98 (2.04); 147	1.27 (0.96); 101
January 2021	-4.81 (5.04); 602	2.15 (1.55); 373
February 2021	-3.62 (4.27); 615	2.00 (1.19); 423
March 2021	-3.07 (3.95); 660	1.76 (1.20); 556
April 2021	-2.08 (3.02); 409	1.15 (0.87); 403

May 2021	-0.98 (2.57); 377	1.14 (1.04); 423
June 2021	0.58 (1.67); 271	0.93 (1.30); 328
July 2021	1.07 (1.38); 102	0.82 (0.62); 127

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292

293 Half-hourly average CO<sub>2</sub> flux was measured in  $\mu\text{mol m}^{-2} \text{s}^{-1}$ , with positive fluxes indicating a  
 294 flux direction from the Earth's surface to the atmosphere. Net ecosystem exchange (NEE)  
 295 was defined as the net flux of CO<sub>2</sub> from the atmosphere to the marsh and was often negative  
 296 during daytime, indicating that Gross Primary Productivity (GPP) was larger than ecosystem  
 297 respiration (R<sub>e</sub>). Evapotranspiration (ET) was calculated by Eddy Pro as the ratio between the  
 298 latent heat flux (LE) and latent heat of vaporisation ( $\lambda$ ). Ecosystem water use efficiency  
 299 (WUE<sub>e</sub>) was then expressed as the ratio between daytime net ecosystem productivity in g  
 300 CO<sub>2</sub> m<sup>-2</sup> h<sup>-1</sup> and evapotranspiration in mm h<sup>-1</sup>.

301

302 A two-dimensional footprint estimation was provided according to the simple footprint  
 303 parameterisation described in Kljun et al. (2015) calculating the ground position of the  
 304 cumulative fraction of flux source contribution by distance for each 30-minute interval. We  
 305 assessed the short-term effects of environmental factors on CO<sub>2</sub> fluxes at a half-hourly time  
 306 scale (e.g. the effects of light, air temperature and vapour pressure deficit) using a series of  
 307 non-linear or linear models. These analyses were limited to the growing season, when the  
 308 plants were actively photosynthesising. ~~To calculate the daily-integrated CO<sub>2</sub> and H<sub>2</sub>O fluxes,~~  
 309 ~~the daily sum of these fluxes was determined for days with at least 85% data coverage. This~~  
 310 ~~involved using the trapezoid rule to estimate the area under the curve for each of these 24-~~  
 311 ~~hour periods. The trapezoid rule approximates the total flux by dividing the day into smaller~~  
 312 ~~intervals, each lasting 1,800 seconds. For each interval, the area is calculated by averaging~~  
 313 ~~the flux values at the beginning and end of the interval, then multiplying by the interval~~  
 314 ~~duration. These areas are then summed to obtain the total daily flux. This method ensures that~~  
 315 ~~even with some missing data points, a reliable estimate of the daily flux can be obtained.~~ All  
 316 post-processing and statistical analyses were performed in R 4.3.2 (R Core Team, 2024)  
 317 including the packages *ggplot2*, *clifro*, *MASS*, *dismo*, *amerifluxr*, *rmarkdown*, *geosphere*,  
 318 *ggmap* and *gbm*.

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320 Because of the large data gaps, it was not possible to model the partition of the NEE in GEP  
 321 and R<sub>e</sub> using common partitioning methods (Lasslop et al., 2010). For simplicity, it was

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- Deleted: over time (i.e.,
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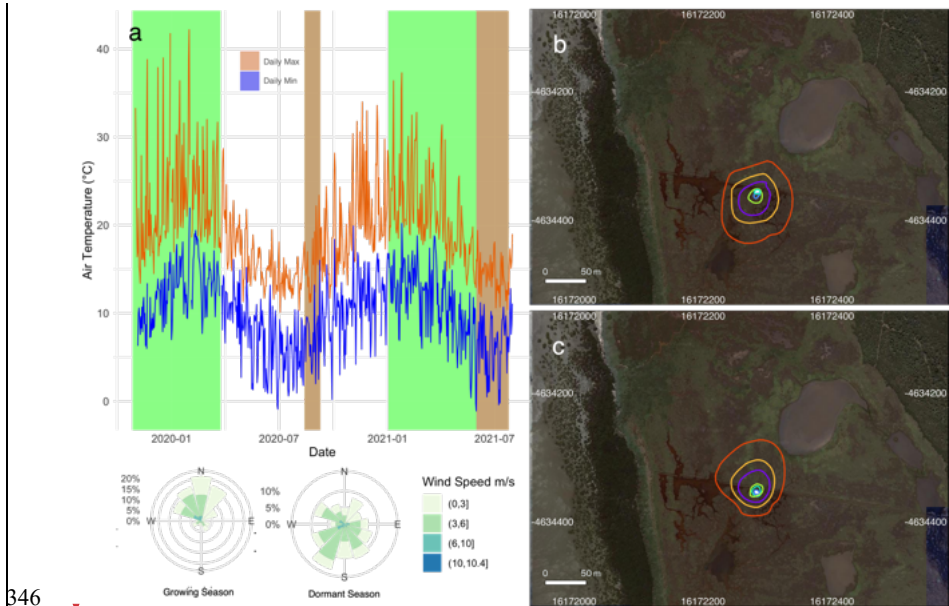
328 assumed that NEE at night coincided with  $R_e$ .  $R_e$  was corrected for temperature effects on  
 329 respiration using a linear slope of the relationship between nighttime NEE and temperature.  
 330 For the CO<sub>2</sub> budget, Net Ecosystem Production (NEP), was defined as  $NEP = -NEE$ .

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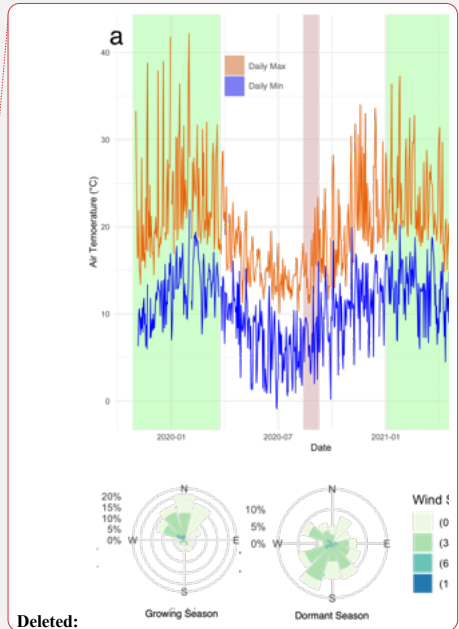
332 3. Results

333  
 334 The observations were divided into a growing season and a dormant season to reflect the  
 335 seasonal phenology of the dominant vegetation type within the flux tower footprint, which  
 336 has a relatively short growing season during the summer. During the growing season, mean  
 337 temperature averaged 22.3°C. Several heatwaves occurred during this period, with  
 338 temperatures exceeding 40°C on a few occasions in 2019. The dormant season was  
 339 significantly colder and windier, with frequent southerly winds (Fig. 2a). Footprint models  
 340 showed a slight variation in flux source between the two seasons, although in both cases the  
 341 size of the footprint and the vegetation composition within the footprint was similar (Figs. 2b  
 342 and 2c), but the shape was skewed to the north during winter due to the prevalent southerly  
 343 winds in that season (Fig. 2a). 70% of the flux measurement source was from within 50 m of  
 344 the tower, while the maximum length of the source location was 73 m.

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352

353 Figure 2: a) The minimum and maximum daily temperature recorded at the Cerberus  
354 meteorological station (Bureau of Meteorology, Fig. 1b) during 2019-2021. The marsh  
355 growing (October-May) and dormant (June-September) periods observed during this study  
356 are shaded in green and pink respectively. A corresponding wind rose diagram summarises  
357 the wind speeds and directions measured at the tower site during the observation periods. The  
358 flux source footprint surrounding the tower during the dormant season (b) and the growing  
359 season (c) shows the cumulative flux source contribution to the flux measurements, with the  
360 outer red line representing the distance by which 90% of the calculated flux is sourced and  
361 the other isolines from the tower outwards correspond to 10%, 20%, 40%, 60% and 80% of  
362 the flux.

363

364 The growing season dataset included 90 days with 85% or more flux data coverage, while the  
365 dormant season dataset included 18 days. There was a strong temporal variability in net  
366 ecosystem exchange (NEE) across both short (daily) and long (seasonal) temporal scales  
367 (Fig. 3). Daytime fluxes were defined as flux points where the global radiation values in the  
368 flux averaging half-hour interval were  $>12 \text{ W m}^{-2}$ , (as per EddyPro methodology). At the  
369 diurnal scale, saltmarsh NEE were negative mostly during the day and positive mostly during  
370 the night and ranged between  $-19.1$  and  $10.86 \mu\text{mol m}^{-2} \text{ s}^{-1}$  across the measurement periods.

371 Monthly averages and data coverage are shown in Table 1.

372

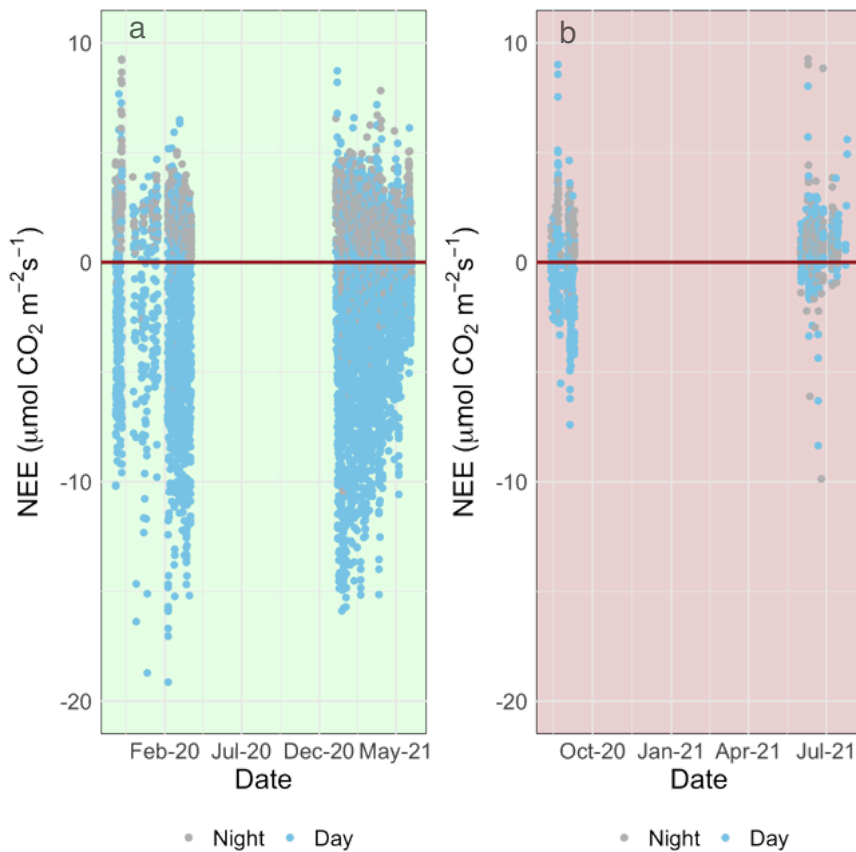
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380 Figure 3: A time series of half-hourly measurements of CO<sub>2</sub> flux between a temperate  
381 saltmarsh and the atmosphere measured by eddy covariance during the marsh growing season  
382 (a) and the dormant season (b). Blue and grey points indicate measurements taken during  
383 daytime and nighttime respectively. Positive fluxes indicate a direction of flux from the Earth  
384 surface to the atmosphere.

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386 Flux rates varied across the day, with CO<sub>2</sub> uptake peaking at 11:00 during the growing  
387 season, and later in the day (14:00) during the dormant period (Fig. 4). Ecosystem respiration  
388 rates ( $R_e$ , defined as nighttime CO<sub>2</sub> flux) were on average ( $\pm$ SD) 1.77 ( $\pm$ 1.12)  $\mu\text{mol m}^{-2} \text{s}^{-1}$   
389 during the growing season and 1.0 ( $\pm$  0.93)  $\mu\text{mol m}^{-2} \text{s}^{-1}$  during the dormant period. The

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393 difference in ecosystem respiration between the growing and dormant seasons is highly  
394 significant (t-test,  $p < 0.01$ ). Daytime  $\text{CO}_2$  flux was on average ( $\pm$ SD)  $-3.53 (\pm 4.15) \mu\text{mol m}^{-2}$   
395  $\text{s}^{-1}$  during the growing season and  $-0.25 (\pm 2.18) \mu\text{mol m}^{-2} \text{s}^{-1}$  during the dormant season.

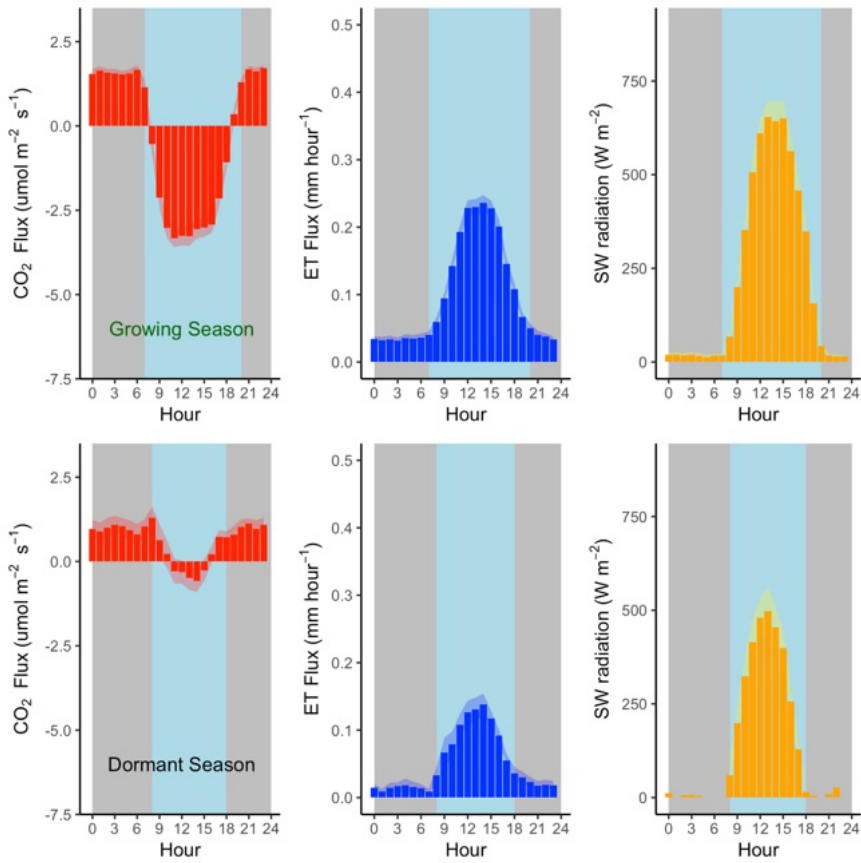
396 Thus, we derive that the maximum Gross Primary Productivity (GPP) of this ecosystem from  
397 NEE and temperature-corrected  $R_e$  (Fig. 5), measured during the growing season, is ca.  $-5.34$   
398  $\pm 4.3 \mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$  ( $-5.53 \pm 4.45 \text{ g C m}^{-2} \text{ day}^{-1}$ ). Average  $R_e$  is thus estimated to comprise  
399 33% of GPP.

400

401 Mean ( $\pm$ SD) daily evapotranspiration was 2.48 mm ( $\pm 2.79$  mm) during the growing season  
402 and 0.97 mm ( $\pm 1.35$  mm) during the dormant season (Fig. 4). Evapotranspiration peaked at  
403 noon AEST during the growing season ( $0.26 \text{ mm h}^{-1}$ ), and later in the day (14:00 AEST)  
404 during the dormant season ( $0.14 \text{ mm h}^{-1}$ ).

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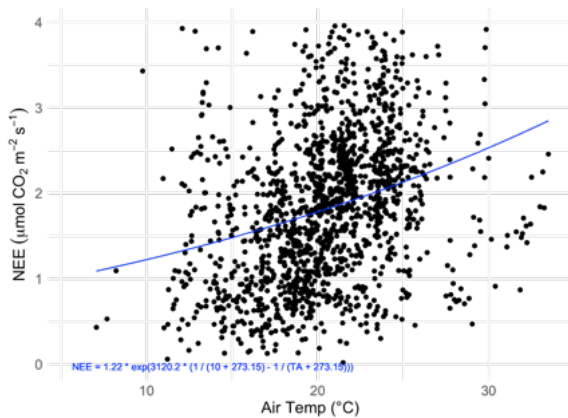
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Figure 4: Mean hourly CO<sub>2</sub> and H<sub>2</sub>O flux (evapotranspiration) rates during the growing season (top) and the dormant season (bottom) alongside mean short wave incoming radiation. Shading corresponds to 1 standard deviation (SD) around the mean. Grey plot background approximates nighttime periods, while light blue approximates daytime (actual day length varies within each season).

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416

417 Figure 5: The relationship between nighttime half-hourly flux measurements (NEE) taken  
 418 between the hours of 22:00 and 02:00 and air temperature (TA). The fitted curve (blue line) is  
 419 the fitted Lloyd & Taylor Arrhenius non-linear model:  $NEE = 1.22 * \exp(3120.2 * (1/283.2 -$   
 420  $1/(TA+273.2)))$ ,  $R^2 = 0.09$ .

421

422 The effect of some environmental forcings on daytime NEE during the saltmarsh growing  
 423 season were explored (Fig. 6). To distinguish this daytime-only value from the 24-hour  
 424 carbon balance integration, and to better highlight CO<sub>2</sub> uptake, NEP values are shown.

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426 Short wave radiation (visible light) was a limiting factor to NEP below approximately 300 W  
 427 m<sup>-2</sup>, but radiation did not reach damaging levels that would lead to a drop in NEP throughout  
 428 the measurement range, which reached a maximum level of ca. 800 W m<sup>-2</sup>. Unlike light, the  
 429 NEP-air temperature relationship followed a Gaussian response, with the highest NEP

430 achieved at the optimal temperature of 25.3°C with a SD of 3.8°C followed by a decline in  
 431 CO<sub>2</sub> uptake by the marsh at higher temperatures. The minimum and maximum air

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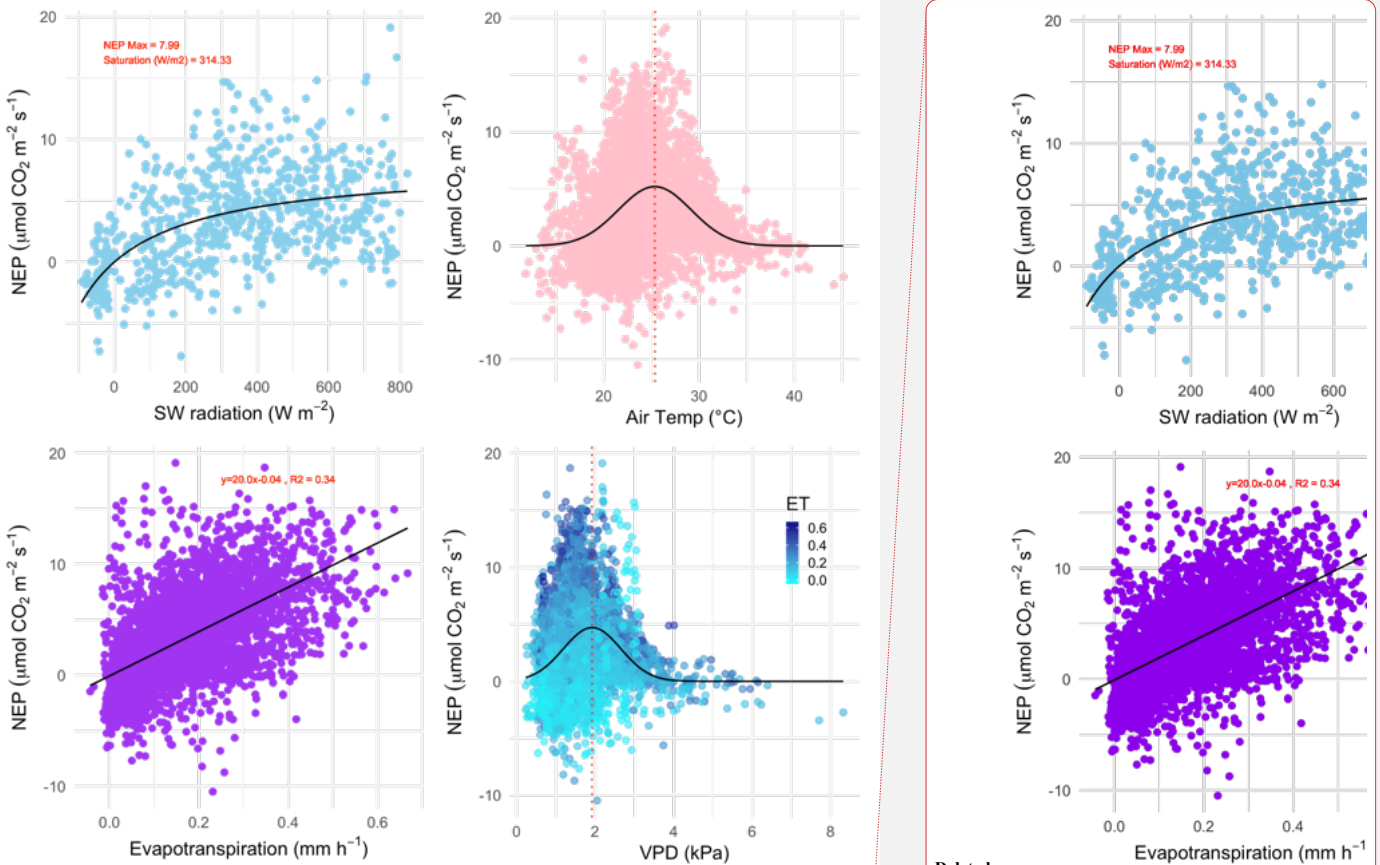
432 temperatures for which modelled NEP nears zero (defined here as 3 SDs from the mean) are  
 433 13.9°C and 36.7°C respectively. Temperature also had a slight but significant positive linear  
 434 relationship with ecosystem respiration (slope=0.07 µmol CO<sub>2</sub> m<sup>-2</sup> s<sup>-1</sup> °C<sup>-1</sup>, p<0.01, data not  
 435 shown).

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436



440 NEP was positively correlated with evapotranspiration during the growing season (Pearson  $r$   
 441 = 0.59, Fig. 6 C). The slope of the NEP/ET relationship was 20.0, indicating an ecosystem  
 442 water use efficiency (WUE<sub>e</sub>) of 0.86 g C kg<sup>-1</sup> H<sub>2</sub>O ( $R^2 = 0.34$ ,  $p < 0.001$ ). The response of  
 443 NEP to atmospheric vapour pressure deficit (VPD) fit a Gaussian relationship (the commonly  
 444 observed inverse U-shaped curve relationship in response to VPD in plants), with NEP  
 445 declining rapidly when VPD exceeded 2.39 kPa. The optimal range of VPD within which  
 446 NEP was maximised in this ecosystem was 1.92 kPa ( $\pm 0.73$  kPa).  
 447



448  
 449 Figure 6: The relationship between growing season daytime half-hourly values of net  
 450 ecosystem productivity (NEP,  $\mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$ ) and corresponding environmental variables.  
 451 a) Net shortwave (SW) radiation (visible light); black line is the Michaelis-Menten model of

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457 best fit. The coefficient of saturation is at  $314 \text{ W m}^{-2}$  and maximum net productivity is  $8.0$   
458  $\mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$ . b) Air temperature; black line is a Gaussian model of best fit with a  
459 temperature optimum at  $25.3^\circ\text{C}$ . c) Evapotranspiration; linear model ( $R^2 = 0.34$ ) has a slope  
460 of  $20.0$ . d) Vapour Pressure Deficit; black line is a Gaussian model of best fit with a VPD  
461 optimum at  $1.92 \text{ kPa}$ . points are coloured by the level of evapotranspiration during the half  
462 hourly NEP measurement.

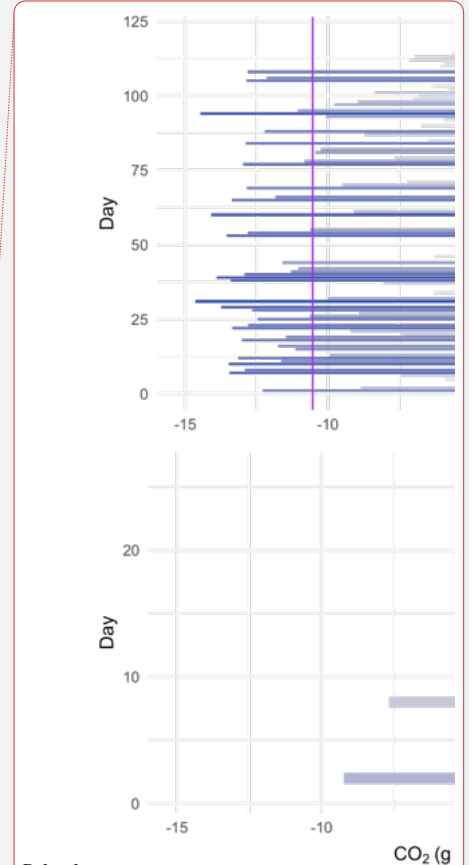
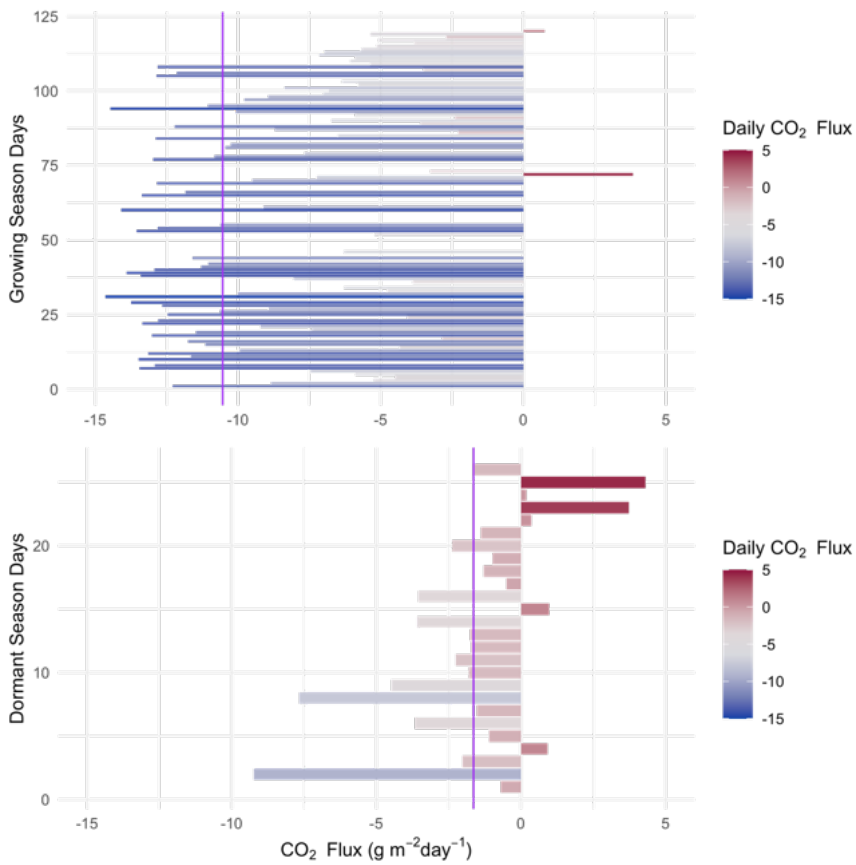
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464 When integrated over a 24-hour period, the saltmarsh is on average a daily  $\text{CO}_2$  sink during  
465 all canopy phenological phases (Fig. 7), although during the dormant season the sink is  
466 weaker, with an average uptake of  $-2.42 \text{ g CO}_2 \text{ m}^{-2} \text{ day}^{-1}$  ( $\pm 2.54$ ). During the growing season  
467 (defined as the non-dormant period and thus reflecting several phenological stages), the  
468 marsh is a substantial sink with a mean ( $\pm\text{SD}$ ) daily NEP of  $10.95 \text{ g CO}_2 \text{ m}^{-2} \text{ day}^{-1}$  ( $\pm 4.98$ )  
469 over a 24-hour period (ranging between  $-22.8$  and  $4.3 \text{ g}$  of  $\text{CO}_2$  emission to the atmosphere  
470  $\text{m}^{-2} \text{ day}^{-1}$ ). The daily  $\text{CO}_2$  budget during the growing season showed some variability among  
471 days ( $\text{CV}=0.46$ , Fig. 7) and days with lower average light levels (i.e. cloudy days) had a  
472 significant negative impact on the  $\text{CO}_2$  budget (multiple linear regression,  $p < 0.02$ ,  $R^2 =$   
473  $0.27$ ). Daily maximum air temperatures did not have a significant impact on the daily  $\text{CO}_2$   
474 budget ( $p = 0.77$ ) at this location, although NEE was significantly affected by temperature at  
475 finer temporal scales (Figure 6).

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Deleted: 5). Assuming the dormant period spans a third of the year, we cautiously estimate an annual NEP value of  $753$  ( $\pm 112.7$ )  $\text{g C m}^{-2} \text{ yr}^{-1}$ .



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484  
 485 Figure 7: Daily (24 h) integrated NEE in g CO<sub>2</sub> m<sup>-2</sup> day<sup>-1</sup> during the saltmarsh growing  
 486 season (top) and the dormant season (bottom) for days with data density > 85%. Purple lines  
 487 indicate the mean daily integrated flux for each season (-10.54 and -1.64 g CO<sub>2</sub> m<sup>-2</sup> day<sup>-1</sup> with  
 488 an SD of 4.98 and 2.54 for growing and dormant respectively). A positive balance indicates  
 489 an integrated net flux of CO<sub>2</sub> from the Earth's surface to the atmosphere over the 24-hour  
 490 period.

493 4. Discussion

501 ~~The study provided high-frequency measurements of an abundant greenhouse gas (CO<sub>2</sub>)~~  
502 ~~using a precise technique (eddy covariance flux) in an ecosystem with limited historical~~  
503 ~~measurements. Time series analysis was performed on CO<sub>2</sub> flux measurements across various~~  
504 ~~scales (daily, nightly, diel, half-hourly, hourly, seasonally) to assess the impacts of ET, SW~~  
505 ~~Rad, VPD, and Tair on CO<sub>2</sub> flux and how these relationships change throughout the year.~~  
506 ~~Seasonality was observed for the first time in an Australian saltmarsh and~~ had a significant  
507 effect on carbon and water flux. Growing season net ecosystem productivity was five times  
508 greater than during the dormant period. Seasonality in Australian marshes has not been  
509 previously reported in the scientific literature, and ~~contradicts previous~~ assumptions that  
510 Australian saltmarshes do not exhibit the growing and dormant phenology observed on other  
511 continents (Clarke and Jacoby, 1994). ~~Seasonality had a significant impact on the carbon~~  
512 ~~budget in this marsh and is an~~ important characteristic of this habitat ~~that has been~~  
513 ~~overlooked (Owers et al., 2018). Seasonality can also have other broader implications yet to~~  
514 ~~be considered in Australian marshes.~~ For example, in the USA, the saltmarsh greening up  
515 period was shown to be an important range-wide timing event for migratory birds (Smith et  
516 al., 2020) with plant-growth metrics predicting the timing of nest initiation for shorebirds.  
517 Saltmarshes in Australia are important roosting and feeding sites along the East Asian  
518 Australasian Flyway, particularly ~~for~~ waders, thus potentially a similar relationship between  
519 migration timing and saltmarsh phenology could be occurring. Seasonality also affects other  
520 significant ecosystem functions such as the bio-geomorphological feedback between  
521 saltmarshes, coastal hydrodynamics and landscape evolution (Reents et al., 2022).

522  
523 We derived the light-response and associated coefficients of light regulation of saltmarsh  
524 NEE using the Michaelis Menten model (Chen et al., 2002). Quantum (or production)  
525 efficiency is the predominant input in remote sensing techniques to model productivity, and is  
526 specific to the biome (Hilker et al., 2010). While not directly comparable to leaf level  
527 quantum efficiency measurements, the quantum efficiency ( $\alpha$ ) of the NEP light response  
528 curve was estimated from the slope of the Michaelis-Menten model to be 0.025  $\mu\text{mol CO}_2 \text{ J}^{-1}$ .  
529 The ecosystem reached light saturation at an insolation of 314  $\text{W m}^{-2}$ , but daytime insolation  
530 was below this value more than 50% of the time suggesting that light might be a significant  
531 limiting factor to NEP at this marsh, especially during winter. The level of light limitation we  
532 observed is an underestimation, due to the loss of high-quality EC data during periods of rain.  
533 The solar geometry at this latitude and the length of day result in an annual average top of

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540 atmosphere SW radiation of 250 W m<sup>-2</sup>, but clouds can strongly modulate the SW radiation  
541 balance (SWCRE), and apart from the months of January and February when cloudy days are  
542 less frequent (10-12 days per month), cloudy days are frequent at this site, averaging 15-17  
543 days per month (Bureau of Meteorology) and could significantly impact on NEP.

544

545 Temperature is another forcing that significantly impacts NEE at this marsh, with an optimal  
546 range for maximum NEP at 25.3°C (21.5°C-29.1°C). Data for Australian saltmarshes is not  
547 available, but this optimal temperature response range is similar to that measured  
548 experimentally in a saltmarsh species in an equivalent climate zone (e.g. Georgia,  
549 (Giurgevich and Dunn, 1981)) and to the values hypothesised for the habitat from data  
550 collected along the US Atlantic Coast, (Feher et al., 2017). The long-term average maximum  
551 daytime temperature at this site is 19.2°C, which is cooler than the optimal range for NEE  
552 suggesting temperature can be a significant limiting factor to productivity, especially during  
553 the dormancy period where average monthly maximum temperatures are only 13.7°C to  
554 16.6°C (Bureau of Meteorology). During the growing season the average maximum  
555 temperatures are within the range of optimal NEE (20.6°C to 23.1°C), although hot days  
556 (>30°C) significantly depress NEE and depending on the year, can be common during  
557 summer months (averaging 2-6 days per month). Within the diversity of saltmarsh species  
558 found globally, some species have C4 photosynthetic pathways (Drake, 1989). C4  
559 photosynthesis plants often exhibit higher optimum temperature ranges (30-35°C, Berry and  
560 Björkman, 1980) than C3 photosynthesis plants (20-25), and the cooler conditions at this site  
561 could explain the absence of C4 plants from this bioregion. The parabolic relationship  
562 between NEP and air temperature and NEP and VPD suggest that higher air temperatures and  
563 VPD (which are expected with climate change) could negatively impact CO<sub>2</sub> uptake by these  
564 coastal ecosystems. High VPD was related to lower NEP, and to a lesser extent, lower ET  
565 (Fig. 6d). However, VPD increases atmospheric demand for water, increasing the evaporation  
566 from the saturated marsh surfaces in the footprint, and this atmospheric demand could be  
567 forcing ET at high VPD rather than plant moderation via reduced transpiration, even if  
568 transpiration is reduced. Thus, despite maintained ET during VPD periods we cannot  
569 conclude a non-closure of stomata. NEP also reduced below a VPD of 1.92 KPa, but at our  
570 field site low VPD correlated with low temperatures (r = 0.88), and low temperatures were  
571 shown to limit NEP.

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575 In saltmarshes, evapotranspiration occurs from plant mediated transpiration but also from soil  
576 pores (which tend to be saturated), wetted leaves and open water. We observed average  
577 evaporation rates of 2.48 mm day<sup>-1</sup> during the growing season and 0.97 mm day<sup>-1</sup> during the  
578 dormant season. Actual evapotranspiration in this region modelled using the CMRSET  
579 algorithm is estimated to range between 0.6 and 3.2 mm day<sup>-1</sup> during winter and summer  
580 respectively (McVicar et al., 2022); our field measurements support the model. Overall,  
581 rainfall is in excess of the requirements for maintaining ET at this site, although deficits can  
582 develop for short periods during the growing season, when ET is higher, perhaps explaining  
583 the drier saltmarsh surface during this period. Conversely, long term rainfall excess could be  
584 contributing to the complicated hydrology at this location, where inundation is not strictly  
585 associated with tidal stage (data not shown) and our observation of long (5-day) periods of  
586 inundation during winter.

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587  
588 Growing season ET rates are significantly higher than those of the dormant season, partly due  
589 to the solar configuration in winter as opposed to summer, but also due to phenological  
590 changes. A big leaf model estimation of evapotranspiration from saltmarshes in New South  
591 Wales estimates ET to be highly sensitive to vegetation height, increasing by more than 1 mm  
592 day<sup>-1</sup> as vegetation height increases from 0.1 to 0.4 m (Hughes et al., 2001) and transpiration  
593 in saltmarsh plants in the cold season has been shown to account for only 20% of the annual  
594 transpiration budget (Giurgevich and Dunn, 1981) following the same pattern as the seasonal  
595 distribution of productivity.

596  
597 The rate of carbon uptake per unit of water loss (WUE) is a key ecosystem characteristic,  
598 which is a result of a suite of physical and canopy physiological forcings, and has direct  
599 implications for ecosystem function and global water and carbon cycling. Mean water use  
600 efficiency (WUEe) of this saltmarsh was estimated at 0.86 g C kg<sup>-1</sup> H<sub>2</sub>O, which is markedly  
601 lower than for grass dominated saltmarshes in China (2.9 g C kg<sup>-1</sup> H<sub>2</sub>O, Xiao et al. (2013))  
602 but similar to the value for WUEe based on NEP and ET in mangroves (0.77 g C kg<sup>-1</sup> H<sub>2</sub>O,  
603 Krauss et al. (2022)), which are also C<sub>3</sub> plants. The Chinese saltmarshes studied in Xiao et al.  
604 (2013) are dominated by *Spartina alterniflora*, a C<sub>4</sub> perennial grass. C<sub>4</sub> plants have higher  
605 (often double) water use efficiencies than C<sub>3</sub> plants due to CO<sub>2</sub> concentrating mechanisms  
606 (Osborne and Freckleton, 2009). The saltmarsh at French Island includes only C<sub>3</sub> plants, and  
607 the dominant chenopod *Sarcocornia quinqueflora* has been suspected to have higher  
608 evapotranspiration rates than saltmarsh by approx. 15% (Hughes et al., 2001), but while

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615 *Sarcocornia quinqueflora* dominates at this site, the footprint is a mix of species, and the  
616 lower WUEe cannot be directly linked to the presence of *Sarcocornia quinqueflora*.  
617 Furthermore, like most wetlands, the wetland surface is a mixed composition of emergent  
618 vegetation, unsaturated soil and water bodies thus the spatial scale at which WUEe is  
619 determined encompasses both the canopy (Ec) as well as any open water present in the  
620 footprint. Transpiration is predicted to account for only 55% of ET in these systems (Hughes  
621 et al., 2001), which is an Ec to ET ratio similar to that of mangroves (Krauss et al., 2022) but  
622 significantly lower than terrestrial forests where more than 90% of ET can be attributed to  
623 transpiration. Thus, regional variations in WUEe can be attributed to multiple forcings that  
624 form complex spatiotemporal patterns.

625  
626 Saltmarshes are considered among the most productive ecosystems on Earth with an  
627 estimated global NEP of 634 Tg C y<sup>-1</sup> (Fagherazzi et al., 2013) and 601 634 Tg C y<sup>-1</sup>  
628 (Rosentreter et al., 2023). Productivity of southern Australian marshes was previously  
629 estimated at 0.8 kg m<sup>-2</sup> y<sup>-1</sup> by repeated measurements of above ground standing crops (Clarke  
630 and Jacoby, 1994), which if not accounting for season, equates to 2.2 g C m<sup>-2</sup> d<sup>-1</sup>. Similar  
631 studies on saltmarshes in France report lower productivity (483 g C m<sup>-2</sup> y<sup>-1</sup>, (Mayen et al.,  
632 2024)) and daily growing season rates of 1.53 g C m<sup>-2</sup> d<sup>-1</sup>, but mid-latitude saltmarsh sites in  
633 the USA and China show productivity rates of 775 g C m<sup>-2</sup> y<sup>-1</sup>, (Wang et al., 2016) and 668 g  
634 C m<sup>-2</sup> y<sup>-1</sup>, (Xiao et al., 2013) respectively. It is clear that productivity across climate zones  
635 and biogeographic regions varies widely with some studies even reporting net emissions over  
636 an annual period from some marshes and a global average estimated between 382 (Alongi,  
637 2020) and 1,585 g C m<sup>-2</sup> y<sup>-1</sup> (Chmura et al., 2003), albeit based on a small subset of studies.  
638 An analysis of GPP across latitudes in the USA show that warmer sites (including mangrove  
639 wetlands in southern USA) had significantly higher GPP than mid-latitude saltmarshes such  
640 as the one on French Island (Feagin et al., 2020). Mangroves have higher NEE than  
641 saltmarshes, estimated by Krauss et al. (2022) to average 1200 g C m<sup>-2</sup> y<sup>-1</sup>. While our data  
642 does not provide enough coverage for a long-term annual estimate of carbon flux, our daily  
643 values of an average of 2.88 g C m<sup>-2</sup> d<sup>-1</sup> during the growing season, combined with the  
644 relatively short dormant season relative to other temperate locations, suggest a high carbon  
645 sequestration rate for this ecosystem type. In another southern hemisphere study, growing  
646 season rates at an EC tower site in Argentina, are extrapolated by us to average 1.6 g C m<sup>-2</sup> d<sup>-1</sup>  
647 (Bautista et al., 2023) but in that saltmarsh, flooding reduced vegetation biomass and  
648 productivity.

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662

663 The data presented here is the exchange of carbon between the land surface and the  
664 atmosphere, but saltmarshes, like other marine connected communities, exchange carbon also  
665 through dissolved carbon pathways, which can be significant (Cai, 2011). Thus, the fluxes  
666 presented here do not constitute the entire carbon budget of this ecosystem.

667

## 668 5. Conclusions

669

670 The response of the French Island saltmarsh to environmental drivers is indicative of the  
671 complex interactions determining saltmarsh productivity. ~~The unique long-term, high-~~  
672 resolution record enabled us to derive temperature, VPD and light response functions, thus  
673 formulating equations that describe how climate-change sensitive parameters such as  
674 temperature, relative humidity, and cloud cover, affect CO<sub>2</sub> uptake, respiration and  
675 evapotranspiration. The marsh operated as a CO<sub>2</sub> sink throughout the various canopy  
676 phenological phases, but during the dormant period, CO<sub>2</sub> uptake was less than 25% that of  
677 the growing season. Seasonality ~~of greenhouse gas fluxes~~ in Australian saltmarshes ~~is an~~  
678 ~~understudied but important aspect of global~~ carbon ~~budgeting~~.

679

## 680 Competing interests

681

682 ~~The contact author has declared that none of the authors has any competing interests.~~

683

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685

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691 Networks of Excellence award to RR.

692

## 693 Data Availability

**Deleted:** While the overall carbon sequestration rate we measured was in the range of other temperate saltmarsh estimates (ca. 750 g C m<sup>-2</sup> y<sup>-1</sup>),

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702 Data used for this analysis is available at <https://figshare.com/s/ba62aafd1a4049248a08> (note  
703 that this is a temporary private link to an embargoed dataset which will be replaced with a  
704 publicly available DOI upon publication).

705

706 Author contribution

707 RR conceptualised the study, acquired funding, prepared the manuscript, designed and  
708 carried out the field campaign, and performed the analysis. ED acquired funding, developed  
709 methodology and prepared the manuscript. AG developed methodology and prepared the  
710 manuscript. TA, EJ VH, HR and MP were involved in the field investigation and  
711 administration of the project and provided edits on the manuscript.

712

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