

Assessing BC and CO Emissions from China Using EMeRGe Aircraft Observations and WRF/CMAQ Modelling

Phuc Thi Minh Ha¹, Yugo Kanaya¹, Kazuyo Yamaji², Syuichi Itahashi³, Satoru Chatani⁴, Takashi Sekiya¹, Maria Dolores Andrés Hernández⁵, John Philip Burrows⁵, Hans Schlager⁶, Michael Lichtenstern⁶, Mira Poehlker⁷, Bruna Holanda⁷

¹Research Institute for Global Change, JAMSTEC, Yokohama 236-0001, Japan,

²Graduate School of Maritime Sciences, Kobe University, Kobe 658-0022, Japan,

³Research Institute for Applied Mechanics, Kyushu University, Fukuoka 816-8580, Japan,

⁴National Institute for Environmental Studies, Tsukuba, Ibaraki 305-8506, Japan,

⁵Institut für Umweltp Physik, Universität Bremen, Otto-Hahn-Allee 1, Bremen 28359, Germany,

⁶DLR Oberpfaffenhofen, Institut für Physik der Atmosphäre, 82234 Wessling, Germany,

⁷Max Planck Institute for Chemistry, Multiphase Chemistry Department, Hahn-Meitner-Weg 1, 55128 Mainz, Germany

Correspondence to: Phuc T. M. Ha (phucha@jamstec.go.jp)

Abstract. Accurate estimates of short-lived climate forcers emissions are essential for effective climate change mitigation, yet large uncertainties remain across Asia. This study integrates the Weather Research and Forecasting / Community Multiscale Air Quality version 5.0.2 (WRF/CMAQv5.0.2) simulations with Effect of Megacities on the Transport and Transformation of Pollutants at Regional and Global Scales (EMeRGe) airborne observations from spring 2018 to refine combustion-related emissions of black carbon (BC) and CO over East Asia. Aircraft observations, which are less affected by near-surface processes, confirmed the accuracy of the Global Fire Emissions Database (GFEDv4.1s) inventory near Thailand. In contrast, anthropogenic BC and CO from the Regional Emission Inventory for Asia (REASv2.1) in the Philippines showed negative biases. For Chinese air masses, the Hemispheric Transport of Air Pollution with corrections (HTAPv2.2z) resulted in overestimation for BC and underestimation for CO, agreeing with our prior ground-based findings on Fukue Island. Scaling Chinese BC emissions using an observation/model ratio (0.48 ± 0.13) gave our best estimate of 0.65 ± 0.25 Tg BC yr⁻¹. Chinese emissions were further estimated at 166 ± 65 Tg CO yr⁻¹ and 12.4 ± 4.8 Pg CO₂ yr⁻¹ using the BC/CO and CO/CO₂ ratios from observations. These results suggest HTAPv2.2z requires revisions of 50% downward for BC and 20% upward for CO, particularly over Central East China during spring. Other inventories including Community Emissions Data System (CEDSV2021_02_05) remain notably biased, while Evaluating the Climate and Air Quality Impacts of Short-Lived Pollutants (ECLIPSEv6b) and Tropospheric Chemistry Reanalysis (TCR2) estimates fall within our identified uncertainty ranges.

1 Introduction

Short-lived climate forcers (SLCFs) are atmospheric gases or particles that persist for days to months, with some lasting up to two decades. Similar to long-lived greenhouse gases, SLCFs alter the energy balance of the Earth, making their mitigation crucial for addressing human-induced climate change (Myhre et al., 2013; Szopa et al., 2021; IPCC AR6 WG1 Ch6). Carbon

monoxide (CO) remains in the atmosphere for one to four months and indirectly contributes to warming by removing hydroxyl radicals (OH), which is the dominant oxidising agent for methane (CH₄) oxidation. CO also facilitates ozone (O₃) formation in the presence of sufficient nitrogen oxides (NO_x = NO + NO₂).

Black carbon (BC) is primarily released through the incomplete combustion of carbon-based fuels in boilers, cooking stoves, and heating stoves that use raw coal. BC drives radiative heating by absorbing sunlight and altering snow and ice albedo, thereby intensifying climate forcing. As nanometer-sized particles, BC poses severe health risks in megacities (Shindell et al., 2012; WHO 2021 new AQ guideline).

A deeper understanding of SLCFs is essential for devising strategies to reduce air pollution, limit man-made climate change, and support sustainable development (Szopa et al., 2021). However, the complexity of SLCFs and their interactions with physical and chemical atmospheric processes poses significant challenges.

Incomplete combustion from human activities results in significant BC and CO emissions. Since Asia accounts for more than 50 % of global BC and CO emissions (Hoesly et al., 2018; Szopa et al., 2021), getting precise emission data in this part of the world is essential. However, both emission estimates and our understanding of BC and CO sources need improvement (Kurokawa et al., 2013). The changes in national SLCF emissions due to the potentially uneven adoption of emission reduction techniques across fast-growing Asian economies over the past two decades are not well understood (Chen and Chen, 2019; Kanaya et al., 2020; Ikeda et al., 2023; Zhang et al., 2022). Moreover, biomass-burning in Southeast Asia has become the primary source of carbon emissions from forest fires in spring (Reid et al., 2013; Heald et al., 2003; Palmer et al., 2006; Johnston et al., 2012), driven by the regional reliance on biomass for residential and agricultural purposes. Fires result from both lightning strikes and human activities.

Assessments of combustion-related BC and CO emissions benefit from airborne measurements. Compared to ground-based measurements, aircraft observations are less affected by local emissions and near-surface processes. The EMeRGe airborne campaign, conducted over East Asia on the High Altitude and Long range (HALO) aircraft in March and April 2018 (Andrés Hernández et al., 2022), captured pollution outflows that offer valuable input for such emission assessments. The EMeRGe aircraft campaign contributed to the evaluation of the Chemical Atmospheric General Circulation Model for the Study of Atmospheric Environment and Radiative Forcing (CHASER) v4.0, revealing large negative biases in daytime HONO concentrations over East Asian coastal regions (Ha et al., 2023). Comparisons of Asian and European outflows revealed underestimation of BC and CO outflows from the Manila plume in the Philippines, as well as overestimation of the BC/CO ratio for overall Asia (4.1 ng m⁻³ ppb⁻¹) in multi-model simulations (Deroubaix et al., 2024a, b). However, detailed region-specific emissions have not been explored. The captured pollution plumes from EMeRGe campaign with a typical travel time of less than 5 d from the source enable the careful selection of air masses unaffected by precipitation, and provide information about emissions of the airborne observed species including CO₂. Estimations of the emissions from urban areas in Asia can then be compared with previous studies (Suntharalingam et al., 2004; Tohjima et al., 2010; Wada et al., 2011). These studies of outflows from Asia have highlighted the roles of deep convection (Folkens et al., 1997), middle latitude cyclone cold fronts (Liu et al., 2003; Sawa et al., 2007), and orographic effects (Liu et al., 2003) in pollutant transport.

Bottom-up emission inventories for Asia, particularly China, have improved markedly in recent years, e.g., reducing the uncertainty range for BC emissions to $\pm 208\%$ (Kurokawa et al., 2013; Lei et al., 2011; Li et al., 2017a; Lu et al., 2011; Zhao et al., 2011, 2013; Zhang et al., 2009; Zheng et al., 2018). Despite this progress, independent observational verification is still
70 required to refine the uncertainty ranges further. Chinese CO₂ emission uncertainties range from -10% to $+9\%$ across all sectors in 2010 (Zhao et al., 2013), 31% for REAS v2.1 in 2010 (Kurokawa et al., 2013; Li et al., 2017b), and 19% for REAS v3.2 in 2015 (Kurokawa and Ohara., 2020). Meanwhile, CO emissions in China exhibit an uncertainty range of $\pm 86\%$ (Kurokawa et al., 2013; Li et al., 2017a, b; Zhang et al., 2009; Zhao et al., 2013; Zheng et al., 2018). Uncertainties on primary particle emissions, such as BC, are subject to more uncertain emission factors in residential fuel combustion than gaseous
75 species (Bond et al., 2002; Zhao et al., 2013; Zhi et al., 2008). Bottom-up estimates often diverge from observation-based (top-down) estimates when local contexts are not adequately incorporated in a timely manner (Choi et al., 2020; Kanaya et al., 2020; Suntharalingam et al., 2004; Zhao et al., 2013). Therefore, emission inventories must be tested using independent observations to improve accuracy.

This study aims to improve the understanding of BC, CO, and CO₂ emissions in Asia to refine regional emission inventories.
80 The key objectives of this study are: (1) to compare aircraft observations with model simulations for four key regions: fire emissions (including of biomass burning) near the Gulf of Thailand (THL), and urban pollution in the Philippines (PHL), Japan (JPN), and China (CHN); (2) to evaluate residual ratios against emission ratios in inventories across distinct air masses using a multi-species analysis; and (3) to estimate CHN emissions of BC, CO, and CO₂ using a combined model–observation framework integrating WRF/CMAQ and EMeRGe aircraft data. Previous studies (Choi et al., 2020; Kanaya et al., 2020; Zhao
85 et al., 2013) evaluated the WRF/CMAQ performance at ground level, where transport/deposition uncertainties for near-ground levels may arise. Incorporating EMeRGe airborne BC and CO data improves confidence in prior ground-level assessments of the CHN case. Additionally, CO₂ observation data, largely unaffected by vegetation interactions (i.e., CO₂ uptake during photosynthesis and losses through respiration), enabled precise BC and CO emission factor calculations for combustion sources. The text is structured as follows: Sect. 2 details the observational data, model configuration settings, and applied methods;
90 Sect. 3.1 examines three polluted cases: THL, JPN, and JPN, focusing on comparisons of concentrations and emission ratios; Sect. 3.2 focuses on CHN case and covers concentration comparisons, emission ratios, emission estimates, and an assessment of analysis uncertainties; Sect. 4 summarizes the key findings and conclusions.

2 Material and methodology

In the following, aircraft observation (Sect. 2.1), modelling system (Sect. 2.2), HYSPLIT backward trajectory model and
95 VIIRS active fire data (Sect. 2.3) are described. In Sect. 2.4 and 2.5, the methodology using residual ratios of co-emitted pollutants (BC/CO, CO/CO₂, and BC/CO₂) to identify emission sources and the introduction of a combined model–observation method are detailed. Additionally, the long-term, ground-based observations from Fukue Island (32.75°N , 128.68°E ; Kanaya et al., 2020) is also used to support the analysis of CHN air masses.

2.1 Airborne observations in Asia during the EMERGE campaign

- 100 The EMERGE campaign (<https://www.iup.uni-bremen.de/emerge/home/>; Andrés Hernández et al., 2022) conducted two field experiments: one over Europe in summer 2017 and another in East Asia during the inter-monsoon period in spring 2018. This study focuses on the data obtained during the Asia experiment, when peak biomass burning in Southeast Asia and continental outflow from China transported pollution to the observable regions. Measurements on-board the German HALO aircraft, which is operated by the German Aerospace Center (Deutsches Zentrum für Luft- und Raumfahrt - DLR), covered a wide region
- 105 from Taiwan north and southward. Pollution plumes from major population centres near Gulf of Thailand (THL case), Manila (Philippines) (PHL case), Osaka (Japan) (JPN case), and Central East China (including Beijing, Hebei, Henan, Shanghai) (CHN case) were detected during flights E-AS-03 (12 March), E-AS-06 (20 March), E-AS-08 (24 March), E-AS-09 (26 March), E-AS-10 (28 March), and E-AS-13 (4 April). Flight altitudes during the measurements ranged from ground level to ~12 km, with ranges around ~0.3–2 km specifically evaluated for designated polluted areas.
- 110 This study analyses combustion-related SLCFs: BC, CO, and CO₂. BC was measured using a Single-Particle Soot Photometer (SP2, Droplet Measurement Techniques, USA) (HALO-CNN by the Max Planck Institute for Chemistry in Mainz) (Holanda et al., 2020), with an uncertainty of ~10 % according to Ohata et al. (2021). SP2 detects BC cores with mass-equivalent diameters of 70–500 nm by measuring time-dependent scattering and incandescence signals as particles cross a Gaussian-shaped laser beam ($\lambda=1064$ nm) (Schwarz et al., 2006). Avalanche photo-diode detectors capture scattering and incandescence
- 115 at high / low gain across two wavelength ranges (350–800 nm and 630–880 nm). BC-containing particles absorb laser energy, heating to ~4000°C and emitting incandescence proportional to their mass, which is determined after applying a correction factor (Laborde et al., 2013). BC concentrations were normalized to standard temperature and pressure (STP, $T_0=273.15$ K, $p_0=1013.25$ hPa). The SP2 incandescence signal was calibrated using size-selected fullerene soot particles at the beginning, during, and end of the campaign (Holanda et al., 2020).
- 120 CO was measured by UV photometry (AMTEX, Institute of Atmospheric Physics, DLR), with a detection limit (3σ) of 5 ppb at 30 s resolution, precision 2 ppb or 1 %, and errors of 4 % (Gerbig et al., 1996). CO₂ was measured using cavity ring-down spectroscopy (CATS by the Institute of Atmospheric Physics, DLR), with a total uncertainty of 0.14 ppm and accuracy better than 0.05 ppm (Chen et al., 2010).
- Other atmospheric constituents, including ozone, its precursors, volatile organic compounds (VOCs, including oxygenated
- 125 VOCs), and NO_x, additionally measured on HALO (Andrés Hernández et al., 2022) were beyond the scope of this study. A custom merged dataset of 15 s averages (until flight E-AS-07) or 30 s averages (from flight E-AS-08 onward) was compiled.

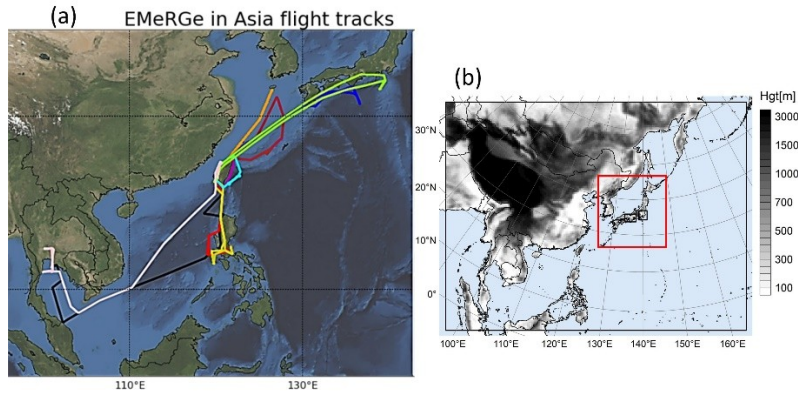


Figure 1. (a) EMeRGe flight tracks across Asia, color-coded by mission: E-AS-03 (black), E-AS-04 (purple), E-AS-05 (brown), E-AS-06 (red), E-AS-08 (orange) (overlapped with E-AS-09), E-AS-10 (yellow), E-AS-11 (blue), E-AS-12 (cyan), E-AS-13 (light green), E-AS-14 (white). (b) WRF/CMAQ model target domains: d01 (entire figure) and d02 (red box).

2.2 WRF/CMAQ modelling system

The CMAQ modelling system version 5.0.2 (Byun and Schere, 2006) is suitable for accurately simulating regional concentrations of pollutant gases and particulates in Asia, including JPN. The model configuration is shown in Table 1 (Ikeda et al., 2014; Yamaji et al. 2020; Chatani et al., 2018, 2020). CMAQ was coupled with the Weather Research and Forecasting, Advanced Research Weather Research and Forecasting (WRF-ARW) model version 3.7.1 (Skamarock et al., 2008), nudged by the US National Centers for Environmental Prediction Final Operational Model Global Tropospheric Analyses (ds083.2) at a $1^\circ \times 1^\circ$ resolution (Gemmill et al., 2007). It was also supplemented with a real-time global sea surface temperature high-resolution analysis at a $1/12^\circ \times 1/12^\circ$ resolution.

Two target nested domains in this campaign were considered: the following outer domain covering East Asia (d01; $45 \text{ km} \times 45 \text{ km}$) and the inner domain covering JPN (d02; $15 \text{ km} \times 15 \text{ km}$) (Fig. 1b). The model extends vertically through 27 layers to 50 hPa. Initial concentrations followed the default conditions of the Statewide Air Pollution Research Center model version 07 (SAPRC07) chemical category, for which spin-up calculations were performed for several months. Boundary concentrations were derived from CHASER simulation (Sudo et al., 2002) using the HTAPv2.2 emission inventory (Huang et al., 2017; Janssens-Maenhout et al., 2015). BC simulations were performed using the sixth-generation CMAQ aerosol module AERO6. In this module, BC concentration is unaffected by chemical or heterogeneous reactions, while coagulation alters particle size without affecting BC mass. Dry deposition loss has been shown to be negligible over a $\sim 3 \text{ d}$ transport period (see Fig. 6, Kanaya et al., 2016).

For domain d01 (East Asia, excluding JPN), monthly anthropogenic emissions were sourced from HTAPv2.2 with a reference year of 2010 unless otherwise stated. The HTAPv2.2 inventory integrated the mosaic Asian anthropogenic emission inventory MIX version 1.1, which couples the Multi-resolution Emission Inventory for China (MEICv1.0) developed by Tsinghua University (<http://www.meicmodel.org>) (Li et al., 2017b), local emission inventories from Korea and India, and the REASv2.1

for the rest of Asia (Kurokawa et al., 2013). MEICv1.0 provides gridded month-to-month emissions at a 0.25° resolution for five sectors (<http://meicmodel.org.cn/>, last accessed 15 May 2023). Sectoral annual emissions in the MEICv1.0 were multiplied by selected factors to represent recent changes in precursor emissions, as described by Zheng et al. (2018) to form the HTAPv2.2z inventory. Open biomass burning emissions were obtained from the GFEDv4.1s as three-hourly data (Giglio et al., 2013; Van der Werf et al., 2017).

For domain d02 (JPN), anthropogenic emissions were taken from the updated Japan Clean Air Program and Japan Auto-Oil Program emission inventory database module (updated JEI-DB; 2013), except for vehicle emissions (2010). Ship emissions (~1 km × 1 km resolution) were provided by the Sasakawa Peace Foundation and were estimated from activities determined using an automatic identification system, with averaging performed on an hourly and daily basis. The JEI-DB and ship emissions data have advanced air quality modelling research in JPN (Chatani et al., 2018, 2023; Li et al., 2017b; Kurokawa et al., 2013; Shibata and Morikawa, 2021). Biogenic VOC emissions were derived from the Model of Emissions of Gases and Aerosols from Nature (MEGANv2.1) for both domains as hourly data (Guenther et al., 2012). The d01 simulation period covered 3 y (2016–2018).

Over the less than 5 d transport, CO, CO₂, and BC are, to first approximation, inert. CO undergoes slow oxidation with OH, given its several-week atmospheric lifetime. As a stable gas, CO₂ is highly inert, particularly when transported at high altitudes without vegetation contact, and deposition is negligible. Although BC can be influenced by wet deposition, precipitation was minimal during the study periods (Sect. 2.3). While the model wet deposition scheme may require testing and refinement, its impact on this analysis is minimal. VOCs, including biogenic, undergo NO_x-driven photochemical oxidation forming O₃ and secondary organic aerosols (SOA), potentially accelerating BC aging, but this effect was minor due to strong emission sources and negligible wet deposition. Although biogenic VOC oxidation produces CO, this contribution was minor given the plumes analysed originated from metropolitan areas. In this study, model uncertainties in the aircraft observation range will be demonstrated in Sect. 3.2.4.

Table 1. Weather Research and Forecasting / Community Multiscale Air Quality (WRF/CMAQ) model configurations

Chemistry Transport Model	CMAQ5.0.2 (Byun and Schere, 2006)	Regional Meteorology Model	WRF(ARW) v3.7.1 (WRF-CMAQ Model User Guide, GitHub)
Boundary condition (d01)	CHASER (MIROC-ESM v4.0) (Sudo et al., 2002)	Nudging for WRF	NCEP-FNL
Chemical Mechanism	SAPRC–07 (CMAQ version 5.0.2 Technical Documentation, 2014)		
Aerosol Module	AERO6 (CMAQ version 5.0.2 Technical Documentation, 2014)		
Domains and resolution	d01 (out of Japan): 45×45 km	d02 (within Japan): 15×15 km	

Emissions	•Anthropogenic HTAPv2.2z (month) REASv2.1 for the rest of Asia (Kurokawa et al., 2013) MEICv1.0 (Li et al., 2017b) + Zheng et al. (2018) for China •Biomass burning GFEDv4.1s (day, 3-hour) (Giglio et al., 2013; Van der Werf et al., 2017) •Biogenic VOCs MEGANv2.1 (day, hour) (Guenther et al., 2012)	• Vehicles JEI-DB (month, weekday/end, hour) (Chatani et al., 2018, 2020) • Other anthropogenic JEI-DB (Updated) (month, hour) (Chatani et al., 2018, 2020) • Ship SPF (month, day of week) (Chatani et al., 2018) •Biomass burning GFEDv4.1s (day, 3-hour) (Giglio et al., 2013; Van der Werf et al., 2017) • Biogenic VOCs MEGANv2.1 (day, hour) (Guenther et al., 2012)
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CHASER(MIROC-ESM) = Circulation Model for the Study of Atmospheric Environment and Radiative Forcing coupled with the Model for Interdisciplinary Research on Climate–Earth System Model, NCEP-FNL = US National Centers for Environmental Prediction Final, SAPRC-07 = Statewide Air Pollution Research Center model version 07, AERO6 = sixth-generation CMAQ aerosol module, HTAP = Hemispheric Transport of Air Pollution emission inventory, JEI-DB = Japan Clean Air Program and Japan Auto-Oil Program emission inventory database module, REAS = Regional Emission Inventory for Asia, MEIC = Multi-resolution Emission Inventory for China, GFED = Global Fire Emissions Database, SPF = Sasakawa Peace Foundation, VOCs = volatile organic compounds, MEGAN = Model of Emissions of Gases and Aerosols from Nature, AeroCom = Aerosol Comparison between Observations and Models, JMA = Japan Meteorological Agency.

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2.3 Backward trajectories and active fire data

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We performed backward trajectory analyses using the US National Oceanic and Atmospheric Association Hybrid Single Particle Lagrangian Integrated Trajectory (HYSPLIT) 4 model (Draxler et al., 2018; Stein et al., 2015) to identify pollutant emission sources. HYSPLIT is an extensive system for computing simple air parcel trajectories (<https://www.arl.noaa.gov/hysplit/>, last accessed 30 May 2023). Back trajectories for the investigated flight segments were traced for 120 h (5 d) from individual three-dimensional locations with a resolution of 15–30 s, which matched the averaging steps of the aircraft data. A 120 h range is generally employed for back-trajectory analyses in long-range transport studies,

190 capturing key atmospheric transport dynamics and ensures source coverage for studied events (traveling time to the aircraft was estimated to be 40–120 h from CHN, 0–10 h (E-AS-03 S2) and 0–24 h (E-AS-06 and E-AS-10) from Manila (PHL), 10–120 h from the broad fire region near THL, 0–36 h from JPN) with extra margin to identify additional influences. This approach agrees with prior studies (Choi et al., 2020; Kanaya et al., 2016; Kanaya et al., 2020; Miyakawa et al., 2017; Zhu et al., 2019). The meteorological field used in HYSPLIT was the Global Data Assimilation System at $1^\circ \times 1^\circ$ resolution.

195 Active fire data were incorporated to validate emission types and back-trajectory results. Near-real-time hotspot data were obtained from the US NASA Fire Information for Resource Management System (NASA FIRMS; <https://firms.modaps.eosdis.nasa.gov/map/>, last accessed 30 May 2023) via the Visible Infrared Imaging Radiometer Suite (VIIRS) sensors aboard the Suomi National Polar-orbiting Partnership and US National Oceanic and Atmospheric Association 20 satellites (bright_ti5).

200 Since the flights were conducted away from rain events, the accumulated precipitation along the trajectories (APT), derived from HYSPLIT rainfall data, helps exclude data significantly affected by wet removal. A detailed analysis of APT was conducted, applying APT criterion during their travel time to the aircraft for the air masses from CHN, THL (APT 72 h < 1 mm), PHL E-AS-06 and E-AS-10 (APT 24 h = 0 mm), PHL E-AS-03 S2 (APT 10 h = 0 mm), JPN (APT 36 h = 0 mm). The APT 72 h < 1 mm criterion showed minimal influence of wet deposition on the emission-concentration relationship (Kanaya et al., 2016, 2020), while APT = 0 mm criterion excludes the effects of rain entirely. Compared to the full dataset, the APT-based analysis reduced data availability in all cases except PHL E-AS-03 S2 and JPN (Table S6). In CHN and PHL E-AS-06, statistical metrics changed only marginally (Table S7). For PHL E-AS-10, great data loss affected the statistical metrics, though residual ratios remained largely stable. In THL case, a decline in the observed BC/CO ratio was noted (Sect. 3.1.1). The full dataset was retained for the main analysis.

205 Moreover, the measured relative humidity remained below 100 % and observed and simulated water concentrations showed strong correlations (Fig. S1 in the supplement). In addition, we carefully discuss deposition, dilution, convection effects, and the representativeness of the air masses in Sect. 2.5 and for each case in Sect. 3.

2.4 Residual ratios for co-emitted pollutants

The EMeRGe-Asia campaign successfully collected pollution samples across CHN, JPN, PHL, and near THL, assessing spatiotemporal pollutant transport within the planetary boundary layer and determining residual ratios. These ratios (X/Y) between co-emitted pollutants, target species X and tracer species Y, are derived from enhancements (Δ) above baseline concentrations. In this study, the residual ratios were obtained from regression slopes fitted to concentration scatter plots (BC/CO, CO/CO₂, and BC/CO₂) for each polluted case, similar to previous aircraft-based studies (e.g., Takegawa et al., 2004; Tang et al., 2018). These residual ratios from rain-free transport cases can be used to identify emission sources and changes in combustion technology (Choi et al., 2020; Kanaya et al., 2020).

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Because CO₂ is a major product of combustion processes, it serves as an ideal reference gas for tracking fresh and aged anthropogenic plumes, although biogenic interactions complicate source attribution (Takegawa et al., 2004). CO and CO₂

ratios have been used to assess incomplete combustion in developing and developed countries (Chandra et al., 2016; Palmer et al., 2006; Suntharalingam et al., 2004; Takegawa et al., 2004; Tang et al., 2018; Wang et al., 2010). BC and CO are byproducts of incomplete combustion of carbon-based fuels, serve as valuable indicators for characterizing emission sources and validating BC estimates in bottom-up inventories (Choi et al., 2020; Guo et al., 2017; Zhu et al., 2019; Kondo et al., 2006, 2011). BC/CO, CO/CO₂, and BC/CO₂ emission ratios differ among combustion sources, techniques, and environmental influences (related to biogenic CO₂ sinks). Depending on fire type, BC/CO ratios vary widely (1.7–29.8 ng m⁻³ ppb⁻¹), while CO/CO₂ ranges from 1.5–22.2 % (Akagi et al., 2011; Chi et al., 2013; Cristofanelli et al., 2013; Kondo et al., 2011). BC/CO ratios from CHN anthropogenic sources spans 1.8–7.9 ng m⁻³ ppb⁻¹ (ground/ship-based: Choi et al., 2020; Guo et al., 2017; Kanaya et al., 2020), while CO/CO₂ ranges 28–60 ppb ppm⁻¹ (airborne: Takegawa et al., 2004; Tang et al., 2018; Suntharalingam et al., 2004). Emission inventories have reported CO/CO₂ from 15.5 ppb ppm⁻¹ (Emissions Database for Global Atmospheric Research inventory - EDGARv6.1) to 23.2 ppb ppm⁻¹ (CEDS v_2021_02_05).

Residual ratios in the polluted cases were slightly influenced by wet deposition, as shown by the APT analysis (Sect. 2.3). The exception was the THL case, though this resulted from air mass selection rather than APT criterion (Sect. 3.1.1). In the CHN case, residual ratios remained consistent even under stricter APT criteria at 72 h (Sect. 3.2.2). In this study, observed and simulated residual ratios are compared to assess model performance against aircraft measurements and to evaluate emission ratios in frequently used inventories for potential improvement (Sect. 3.1 and 3.2.2). Observed residual ratios enable the CO and CO₂ retrieval from BC emission estimates (Sect. 3.2.3). Residual ratios in plumes must be interpreted with caution when assessing long-range transport, convective transport to the lower stratosphere, or interactions with different ecosystems, as noted by Yokelson et al. (2013).

2.5 Combined model–observation approach for estimating Chinese emissions

A combined model–observation approach was used to estimate emission correction factors for pollutants in the HTAPv2.2z inventory used in the CMAQ model ($E(x)_{\text{HTAPv2.2z}}$ or $E(x)$ for short, where x is BC or CO). This was achieved by calculating observed-to-modelled concentration ratios by dividing the mean observed value by the mean simulated value for plumes.

$$E(\text{BC})_{\text{HTAPv2.2z}} = \text{mean observed [BC]} / \text{mean modelled [BC]} \quad (\text{Eq. 1})$$

$$E(\text{CO})_{\text{HTAPv2.2z}} = \text{mean observed } [\Delta\text{CO}] / \text{mean modelled } [\Delta\text{CO}] \quad (\text{Eq. 2})$$

As wet and dry deposition during the transport from CHN have been analysed and found negligible (Kanaya et al., 2016, 2020), linear responses between emissions and concentrations were assumed for both observed and simulated BC and enhanced CO (from the baseline levels; ΔCO), which has been verified by Ikeda et al. (2022). The observation-to-model ratios derived from Eqs. 1 and 2 are discussed for each case (Sect. 3.1 and 3.2.1), and most deeply in the emission estimate for CHN (Sect. 3.2.3). Similarly, emission estimates for BC, CO, and CO₂ were achieved based on linear emission-to-enhanced-concentration relationships using correction factors (i.e. observation-to-model ratios derived from Eqs. 1, 2), assuming the geographic distribution in HTAPv2.2z was accurate.

255 For enhanced concentrations, BC baseline concentrations were assumed to be zero, while non-zero CO baselines were considered (hence $[\Delta\text{CO}]$ was used in Eq. 2) (Kondo et al., 2011; Kanaya et al., 2016; 2020). Several methods are available for estimating CO baseline concentration (Matsui et al., 2011; Miyakawa et al., 2017; Oshima et al., 2012; Takegawa et al., 2004; Verma et al., 2011), including the 5th percentile of a 14 d moving dataset (Choi et al., 2020; Kanaya et al., 2020). However, this method is challenging for aircraft data due to rapid altitude and location variations. In this study, the baseline
 260 CO concentration was determined from the slope and intercept of the linear regression fitted to the BC/CO scatter plots. Estimated BC and CO emissions ($\text{Est}(x)$) were derived by multiplying $\text{E}(\text{BC})$ and $\text{E}(\text{CO})$, respectively, with the total CHN emission value from the HTAPv2.2z inventory (Eq. 3). Given the uniformity of emission structures across China, as evidenced by long-term ground-based observations capturing various air masses from CHN (eg., Kanaya et al., 2020), applying the correction factor to the entire region of CHN is a reasonable approach. Estimated emissions of other pollutants ($\text{Est}^*(x)$) were
 265 translated from BC and CO estimates using observed residual ratios of BC/CO and CO/CO₂ (Sect. 2.4) (Eqs. 4–6).

$$\text{Est}(x) = \text{E}(x)_{\text{HTAPv2.2z}} \times \text{HTAPv2.2z}(x), \text{ where } x \text{ is BC or CO} \quad (\text{Eq. 3})$$

$$\text{Est}^*(\text{CO}) = \text{Est}(\text{BC}) / (\text{observed BC/CO ratio}) \quad (\text{Eq. 4})$$

$$\text{Est}^*(\text{BC}) = \text{Est}(\text{CO}) \times (\text{observed BC/CO ratio}) \quad (\text{Eq. 5})$$

$$\text{Est}^*(\text{CO}_2) = \text{Est}(\text{CO}) / (\text{observed CO/CO}_2 \text{ ratio}) \quad (\text{Eq. 6})$$

270 In CHN case, the air mass footprint spans two major emission regions – North- and South-Central East China (NS-CEC; Fig. S7) - covering complex anthropogenic sources. Chinese emissions were well-mixed before reaching the aircraft, as sector-specific sources exhibited largely spatial and temporal overlap in emission inventories. To ensure a nationally representative estimate by mitigating local transport and source biases, we used the averaged observed residual ratios and correction factors applied to inventory-prescribed national emissions. To minimise discrepancies in BC/CO measured by aircraft, we applied the
 275 spring-averaged 2018 BC/CO ratio obtained from Fukue Island to Eqs. 4 and 5. This long-term surface observation, ongoing since 2009, frequently captures air masses from CHN, including those during the EMERGe campaign (Kanaya et al., 2020).

3 Results and discussion

The data analysis comprises (1) comparison of observational and model datasets, (2) assessment of emission ratios, and (3) combination of model and observations to estimate emissions. The following subsections will discuss the results of four
 280 pollution events by regions (THL, PHL JPN and CHN). The corresponding figures and tables, however, are organized by analysis objectives: (1) Fig. 2–4 and Table 2, (2) Fig. 5–6, and (3) Fig. 7–9 and Table S2.

3.1 Fire emission near THL and anthropogenic emissions from PHL and JPN

3.1.1 THL case study: fire emissions from fresh flaming fires in GFEDv4.1s inventory

The pollutant peak for the THL case during flight E-AS-03 S1 (12 March 2018, 0517–0546 UTC) was caused by emissions
 285 from fires (Fig. 2a, b, Fig. 3a). The HYSPLIT model indicated that the sources were in Cambodia, Vietnam, and Myanmar,

with 9116 fire hotspots on 11 March were identified using the VIIRS. The highest concentrations at an altitude of 1.5 km were $3.26 \mu\text{g m}^{-3}$ for BC (mean $1.63 \mu\text{g m}^{-3}$, $\sigma 0.80 \mu\text{g m}^{-3}$), 580 ppbv for CO (mean 374 ppbv, $\sigma 106$ ppbv), and 420 ppm for CO₂ (Fig. 2a and b). The model predicted similar maximum BC of $3.11 \mu\text{g m}^{-3}$ (mean $1.40 \mu\text{g m}^{-3}$, $\sigma 0.76 \mu\text{g m}^{-3}$) and CO of 722 ppbv (mean 398 ppbv, $\sigma 147$ ppbv) during the flight, though peaks occurred 10 min earlier than observed, likely due to spatial and temporal limitations in model inputs and outputs. While maximum and mean concentrations agreed, correlation coefficients were affected by timing discrepancies (Table 2, Fig. 4a, b). The agreement for THL suggests the GFEDv4.1s inventory is suitable for fire emissions in Southeast Asia, as it incorporates small fires with increased carbon and CO emissions which is not included in GFEDv4.0 (Van der Werf et al., 2017).

In the THL case, fire emission resulted in maxima of $7.1 \pm 0.3 \text{ ng m}^{-3} \text{ ppb}^{-1}$ for BC/CO, $39.5 \pm 4.1 \text{ ppb ppm}^{-1}$ ($\sim 4\%$) for CO/CO₂, and $281 \pm 33 \text{ ng m}^{-3} \text{ ppm}^{-1}$ for BC/CO₂, with strong correlations ($R_{\text{BC/CO}} = 0.95$, $R_{\text{CO/CO}_2} = 0.78$, and $R_{\text{BC/CO}_2} = 0.74$, p -values < 0.0001) (Fig. 5a–c), indicating a common source for these species. The CMAQ model showed a comparable BC/CO ratio ($5.20 \pm 0.04 \text{ ng m}^{-3} \text{ ppb}^{-1}$) with $R_{\text{BC/CO}} > 0.99$. The simulated BC/CO ratio was slightly lower than the observed one, corresponding to higher CO concentrations modelled (Fig. 2b). However, spatial mean BC/CO emission ratio in the model for the THL domain (13° – 18° N, 100° – 113° E, sparsely dashed box in Fig. S5 in the supplement) was $8.2 \text{ ng m}^{-3} \text{ ppb}^{-1}$ across CMAQ sectors and $7.3 \text{ ng m}^{-3} \text{ ppb}^{-1}$ using GFEDv4.1s, which agreed more with observations. Moreover, the APT analysis (Sect. 2.3) excluded long-range transport from Myanmar and southern China via the East Ocean while retaining more regional air masses from Cambodia, Thailand, and southern Laos (Fig. S5e and f). This lowered the observed BC/CO ratio to $5.21 \pm 0.68 \text{ ng m}^{-3} \text{ ppb}^{-1}$ (Table S7), agreeing more closely with modelled values. Although observed BC, CO, and CO₂ concentrations decreased under this criterion, simulated values remained unchanged, suggesting that air mass selection rather than APT threshold accounts for the improved model – observation agreement.

The BC/CO emission ratio from the aircraft-based measurements ($7.1 \pm 0.3 \text{ ng m}^{-3} \text{ ppb}^{-1}$) closely matches previous studies, including emission factor database (Akagi et al., 2011), WRF-Chem (FINNV1.5) simulations (Lee et al., 2018), airborne observations (Kondo et al., 2011; Warneke et al., 2009), and ground-based biomass-burning study (Zhu et al., 2019) (Table S5 in the supplement). High and low BC/CO ratios have also been found in earlier studies (Chi et al., 2013, Cristofanelli et al., 2013; Kondo et al., 2011; Paris et al., 2009; Vasileva et al. 2017; Warneke et al., 2009), with Asia showing higher values than North America due to increased modified combustion efficiencies, inversely proportional to $\Delta\text{CO}/\Delta\text{CO}_2$ ratios during the flame phase (Kondo et al., 2011). BC/CO ratios tend to be lower and CO/CO₂ ratios tend to be higher in the smouldering phase (higher CO emissions), whereas BC and CO₂ emissions rise in flame phase (Laursen et al., 1992; Ward et al., 1992; Nance et al., 1993; Le Canut et al., 1996; Yokelson et al., 1999; Goode et al., 2000). However, CO/CO₂ ranges vary by fire type may not be explicit from observations (Table S5 in the supplement). The BC/CO ratios in aircraft-based measurements and the low CO/CO₂ ratio for fires ($\sim 4\%$), even with changes in the APT analysis (Table S7), indicate that the THL fires were dominated by very fresh Asian spring savanna flaming fires, consistent with the GFEDv4.1s classification (Andela et al., 2019; Van der Werf et al., 2017).

3.1.2 PHL case study: anthropogenic emissions in REASv2.1 inventory

320 In the PHL case, maximum BC and CO concentrations of $1.51 \mu\text{g m}^{-3}$ and 241 ppbv, respectively, were found for flight E-AS-10 (0111–0419 UTC), while CO peaked at 211 ppbv for flight E-AS-06 (0058–0347 UTC), and maximum BC and CO concentrations of $0.98 \mu\text{g m}^{-3}$ and 197 ppbv, respectively, were found for flight E-AS-03 S2 (0905–0922 UTC) over PHL at ~1–2 km altitude (S2 boxes in Fig. 2a, b, and red boxes in c–e). The primary BC and CO source was Manila (Fig. 3b–d). The BC and CO concentrations were lower than those in THL and CHN (Sect. 3.1.1 and 3.2.1). CMAQ simulations underestimated
325 the observed maxima, with biases of –23 % to –30 % for BC (maxima of $0.25\text{--}0.29 \mu\text{g m}^{-3}$) and –5 % to –23 % for CO (maxima of 135–153 ppbv). The CMAQ BC underestimations agreed with simulations by CAMchem–CESM2.1 and WRFchem v4.3.3 (Deroubaix et al., 2024a), while IFS–CAMS simulation better predicted the CO maxima (Deroubaix et al., 2024a) possibly due to its use of the CAMS–GLOB–ANTv4.2 inventory (Granier et al., 2019). In our study, correlation coefficients between observations and simulations were $R_{\text{BC}} = 0.49\text{--}0.54$ and $R_{\text{CO}} = 0.51\text{--}0.59$ (p -values ≤ 0.001), except for
330 $R_{\text{CO_E-AS-03}}$ (0.47; p -values > 0.001) (Table 2). Observation/model ratios (Eqs. 1 and 2) were 2.00 ± 0.77 and 3.24 ± 1.92 for BC and 1.07 ± 0.55 , 1.20 ± 0.16 , and 1.37 ± 0.74 for CO (Fig. 4c–g), indicating REASv2.1 underestimated BC more than CO emissions for PHL, likely due to underestimated residential sector emissions (contributing 76 % BC, 63 % CO). The lower concentration variability of the model (Fig. 4c–g) most probably resulted from the coarse resolution ($\sim 0.5^\circ$ grid vs. $\sim 0.01^\circ$ per 15 s), limited temporal resolution in REASv2.1, and the model biases in dispersion and mixing, which might smooth out
335 concentration contrasts near emission hotspots in PHL.

Kurokawa et al. (2013) identified large uncertainties in REASv2.1 BC (± 257 %) and CO (± 131 %) emissions for Southeast Asian. These uncertainties were primarily related to gaps in the residential sector (± 351 % for BC and ± 208 % for CO) due to insufficient data of emission factors, removal efficiency, and road emissions. BC emissions in REASv2.1 were half of those in the EDGARv4.2, while CO amounts were similar (Lee et al., 2018). Lee et al. (2018) also found that the underestimated
340 fine particulate ($\text{PM}_{2.5}$) concentrations in Manila in the REASv2.1 may stem from missing aerosol sources (e.g., roads, construction, and industrial dust). In our study, BC and CO discrepancies likely related to combustion sources. Our CMAQ simulations assumed the un-changed 2010 emissions, despite a 65 % increase in PHL’s national carbon budget from 2010–2018 (Friedlingstein et al., 2020). BC, CO, and CO_2 emissions in 2015 were 13 %, 14 %, and 32 % higher, respectively, in the newer version REASv3.2 (Kurokawa and Ohara, 2020) and other bottom-up inventories, probably more appropriate for PHL.
345 During flights E-AS-03 S2 and E-AS-10, CO_2 concentrations from Manila area were not positively correlated with BC and CO, probably due to variations in biogenic fluxes (Fig. 2a, b (segment S2), c–e). However, other combinations are of interest. The observed BC/CO ratios converged around $11.4\pm 1.3 \text{ ng m}^{-3} \text{ ppb}^{-1}$ for flight E-AS-03 ($R = 0.83$) and $11\pm 0.1 \text{ ng m}^{-3} \text{ ppb}^{-1}$ for flight E-AS-10 ($R = 0.97$), suggesting that BC emissions are higher relative to CO emissions in PHL than in other investigated urban areas. CMAQ simulations for E-AS-03 S2 ($10.2\pm 0.4 \text{ ng m}^{-3} \text{ ppb}^{-1}$, $R = 0.97$) agreed well with observations,
350 validating the BC/CO emission ratio in REASv2.1 for Manila. In contrast, the simulation for E-AS-10 ($4.3\pm 0.1 \text{ ng m}^{-3} \text{ ppb}^{-1}$,

$R = 0.91$) underestimated BC/CO ratio because of differing biases in the BC and CO concentrations. The CMAQ (REASv2.1) model estimated a BC/CO ratio of $8.0 \text{ ng m}^{-3} \text{ ppb}^{-1}$ for PHL, still underestimating observations by $\sim 30 \%$ (Fig. 6a). The low simulated BC/CO ratio in E-AS-10 may have resulted from mixing of multiple sources, inferred from an unusually low CO intercept ($\sim 75 \text{ ppb}$ on x-axis, Fig. 5e). The observed CO/CO₂ ratio for E-AS-06 was $21.2 \pm 1.1 \text{ ppb ppm}^{-1}$, which was 34 % lower than the lowest ratio among bottom-up inventories (28 ppb ppm^{-1} in ECLIPSEv6b interpolated for 2018).

In conclusion, CMAQ (REASv2.1) simulations gave insufficient BC and CO emissions for PHL, with BC emissions more inadequate than CO emissions, though the BC/CO ratio remained appropriate. Adjusting the REASv2.1 activity data upward by 1.2 for both BC and CO, based on the observation/model ratios for E-AS-03 S2 (Fig. 4c, d) and the constraint of a good relationship between the modeled and observed BC/CO ratios for E-AS-03 S2 (Fig. 5d), could improve the agreement of the model with the observations. Unlike the emission gaps, concentration ratio gaps primarily stem from emission factors or pollutant activity levels, where specific sectoral activity levels require more detailed characterization. Other bottom-up inventories, including REASv3.2, require BC/CO and CO/CO₂ ratios that are 10–30 % and 25–50 % lower, respectively, to match observations (Fig. 6a, b). The underestimation of REASv2.1 CO emission and the overestimation of CO/CO₂ ratios suggest that the CO₂ emission rates in the PHL inventory may need to be higher.

3.1.3 JPN case study: evaluating emissions in JEI-DB inventory

For the JPN case, higher BC and CO concentrations were observed during flight E-AS-13 (4 April 2018) between approximately 0300 and 0600 UTC. The CMAQ model predicted this period well (the dashed box in Fig. 2f, g). However, the observed BC concentration fluctuated, resulting in weaker correlations with the simulation. A sub-segment (0428–0455 UTC, the black box in Fig. 2f, g) at an altitude of $\sim 2.5 \text{ km}$ over Sagami Bay and Suruga Bay (112 data points) was selected (Fig. 3k).

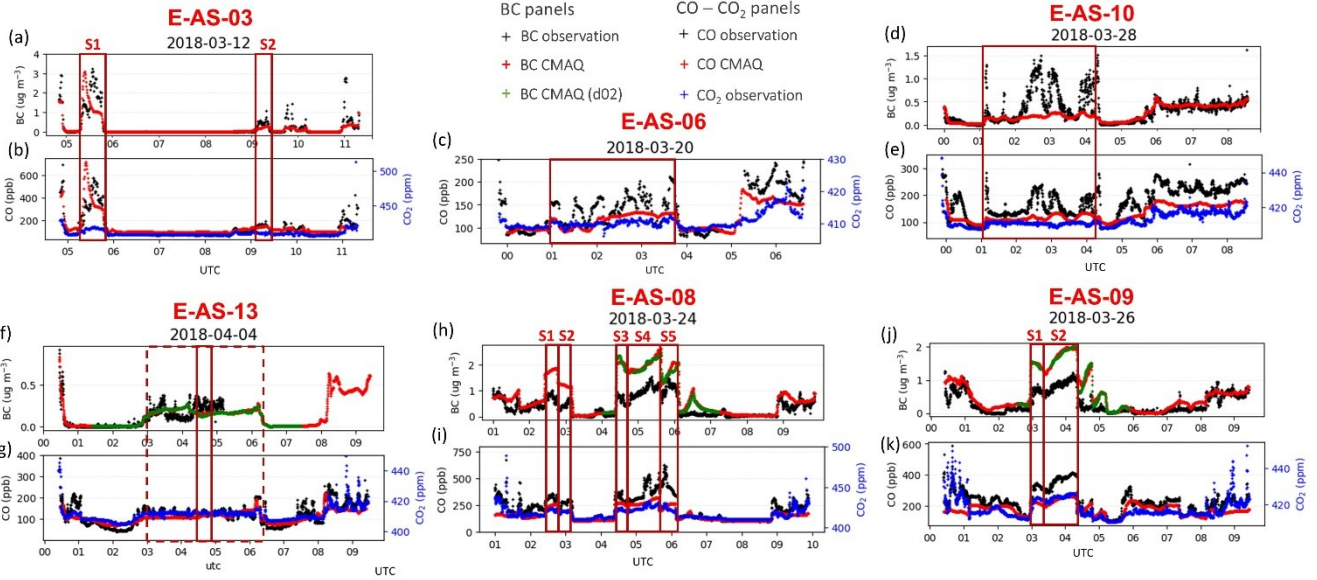
In this sub-segment, the observed CO and CO₂ concentrations correlated best ($R = 0.90$). The maximum BC concentration was $0.37 \text{ } \mu\text{g m}^{-3}$ (mean $0.26 \pm 0.04 \text{ } \mu\text{g m}^{-3}$), and back trajectory indicated transport across Kagoshima, Osaka, Nagoya, and Tokyo (Fig. 3e). CMAQ model underestimated the maximum BC concentration by maintaining $0.13\text{--}0.17 \text{ } \mu\text{g m}^{-3}$ (mean $0.16 \pm 0.01 \text{ } \mu\text{g m}^{-3}$). The observations did not contain a clear maximum CO concentration (mean $130 \pm 7 \text{ ppbv}$) and was slightly underestimated by CMAQ (mean $119 \pm 5 \text{ ppbv}$), resulting in low correlation coefficients. Similar to THL, the poor correlations were attributed to the spatial and temporal limitations of the CMAQ model inputs and outputs. CMAQ overestimated BC concentrations below $\sim 1.5 \text{ km}$ but underestimated BC at $\sim 2\text{--}4 \text{ km}$ over Sagami Bay and Suruga Bay ($33^\circ\text{--}35^\circ \text{ N}$, $138^\circ\text{--}140^\circ \text{ E}$), as shown in Fig. S2b in the supplement. A previous study found that CMAQ (version 4.7) underestimated BC over Fukue Island when air masses passed over Japan's main islands (Kanaya et al., 2020). While it is unclear whether the JEI-DB inventory contributed to the underestimation of BC in this study, JEI-DB has been found to underestimate BC emissions in Tokyo (Shibata and Morikawa, 2021).

The observed BC/CO, CO/CO₂, and BC/CO₂ ratios were $4.5 \pm 0.4 \text{ ng m}^{-3} \text{ ppb}^{-1}$ ($R = 0.71$), $13.0 \pm 0.6 \text{ ppb ppm}^{-1}$ ($R = 0.90$), and $60 \pm 6 \text{ ng m}^{-3} \text{ ppm}^{-1}$ ($R = 0.67$), respectively. However, the CMAQ model failed to capture the observed BC maximum at an

altitude of ~2.5 km, underestimating the BC/CO ratio to $1.8 \pm 0.06 \text{ ng m}^{-3} \text{ ppb}^{-1}$ ($R = 0.95$). An unacceptably low x-axis intercept (34 ppb, which corresponded to CO baseline concentration) suggested a mixing of two sources rather than a valid emission ratio in the simulation. The measured BC/CO ratio was slightly lower than the values reported in other ground-based ($6.8 \pm 0.2 \text{ ng m}^{-3} \text{ ppb}^{-1}$ for 2010–2016 (Choi et al., 2020), $5.9 \pm 3.4 \text{ ng m}^{-3} \text{ ppb}^{-1}$ for 2009–2015 (Kanaya et al., 2016; yellow diamonds in Fig. 6d), $5.7 \pm 0.9 \text{ ng m}^{-3} \text{ ppb}^{-1}$ for 2003–2005 (Kondo et al., 2006)) and aircraft-based ($6.3 \pm 0.5 \text{ ng m}^{-3} \text{ ppb}^{-1}$ for 2003 (Kondo et al., 2006)) studies. The observed BC/CO ratio closely matched the values from HTAPv2.2z (JEI-DB) ($5.15 \text{ ng m}^{-3} \text{ ppb}^{-1}$, open black square in Fig. 6d), ECLIPSEv6b inventory (dark grey circles in Fig. 6d), and EDGARv6.1 (dark blue line with circles in Fig. 6d), in which updated transport activity and shipping emission data were used (Klimont et al., 2017; https://edgar.jrc.ec.europa.eu/index.php/dataset_ap61, last accessed 22 February 2024). However, the observed BC/CO ratios were clearly overestimated in CEDS (CMIP6) (Hoesly et al., 2018) (blue line with circles in Fig. 6d) and CEDS v_2021_02_05 (O'Rourke et al., 2021) ($\sim 15 \text{ ng m}^{-3} \text{ ppb}^{-1}$, orange line with circles in Fig. 6d).

The observed high CO/CO₂ ratio ($13.0 \pm 0.6 \text{ ppb ppm}^{-1}$), which exceeded Yokosuka (JPN) 2022 value ($\sim 5.8 \text{ ppb ppm}^{-1}$), suggests unusual sources such as engineering vessels (Zhang et al., 2016; Wu et al., 2022). Historical observations of 14–18 ppb ppm⁻¹ (Takegawa et al., 2004), $11.2 \pm 0.4 \text{ ppb ppm}^{-1}$, and $12.6 \pm 0.7 \text{ ppb ppm}^{-1}$ (Kondo et al., 2006) were two–three times higher than the ratio for Yokosuka. The narrow range of CO₂ concentrations (411–413 ppm) and the shallow mixing depth (determined from the HYSPLIT results, data not shown) suggest that the air masses were from mixing of multiple background sources via convective transport (Yokelson et al., 2013) or long-range transport (over 5 d) from North China. Due to these uncertainties (high CO/CO₂ ratios, narrow CO₂ variability, and shallow mixing depths during transport), this air mass lacked sufficient emission information to be considered a typical JPN air mass, despite the agreement for BC/CO ratio with other observations and emission inventories. Accurate analysis of national emission ratios depends on selecting representative air masses, which are assessed through concentrations of co-emitted species and relevant meteorological parameters.

THL case E-AS-03 S1 **PHL case** E-AS-03 S2, E-AS-06, E-AS-10
JPN case E-AS-13 **CHN case** E-AS-08 S1,2,3,4,5, E-AS-09 S1,2



405 **Figure 2.** Black carbon (BC), CO, and CO₂ concentrations during selected EMerge flights. Polluted cases highlighted by red boxes include: THL (E-AS-03 S1; (a), (b)), PHL (E-AS-03 S2 in (a), (b); E-AS-06 in (c); E-AS-10 in (d), (e)), JPN (E-AS-13; (f), (g)), and CHN (E-AS-08 S1–S5 in (h), (i); E-AS-09 S1, S2 in (j), (k)). Observed BC and CO (black) are compared with CMAQ simulations from domain d01 (45 × 45 km, red) and domain d02 (15 × 15 km, green). Observed CO₂ is shown in blue.

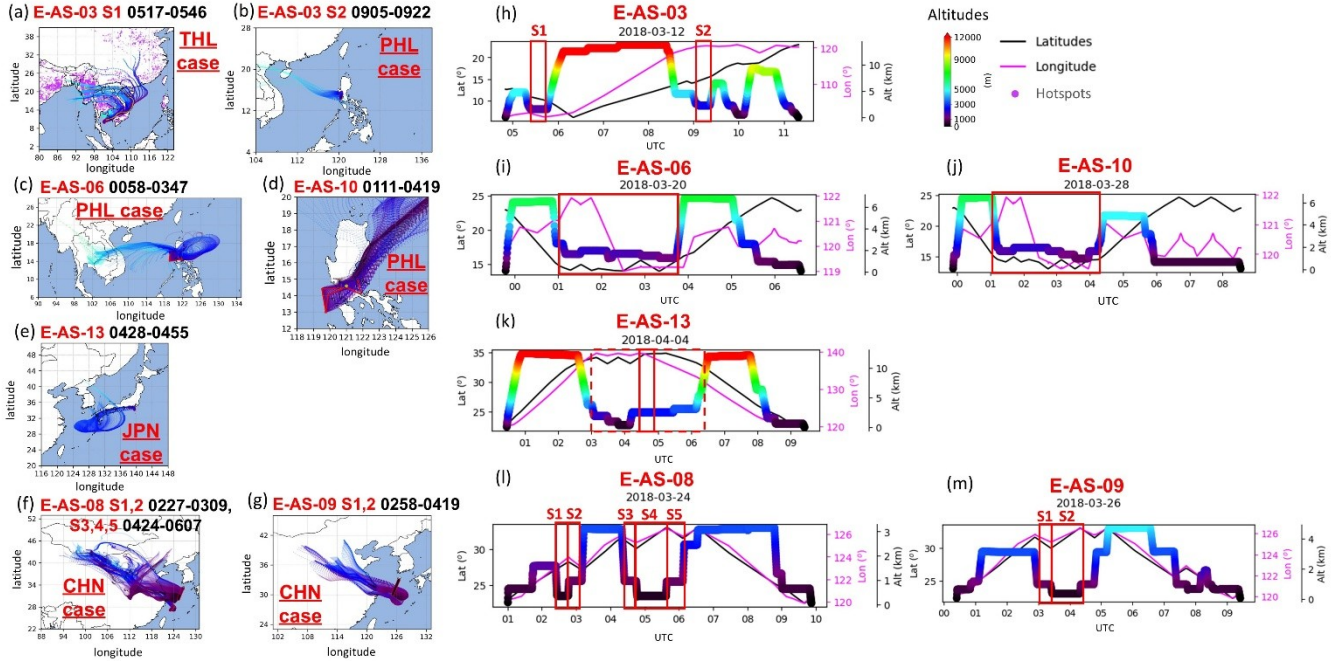


Figure 3. 120 h back trajectories (a–g) and flight coordinates (h–m) highlighting selected polluted cases. Trajectory altitudes in (a–g) are color-coded using the top-right bar. In (a), magenta dots mark VIIRS hotspots (760 m² fires; dot size for visibility). Panels (h–m) show latitudes (black, left axis), longitudes (magenta, right axis), and altitudes (color-coded, second right axis; color-scaled at top right). Red boxes highlight selected cases.

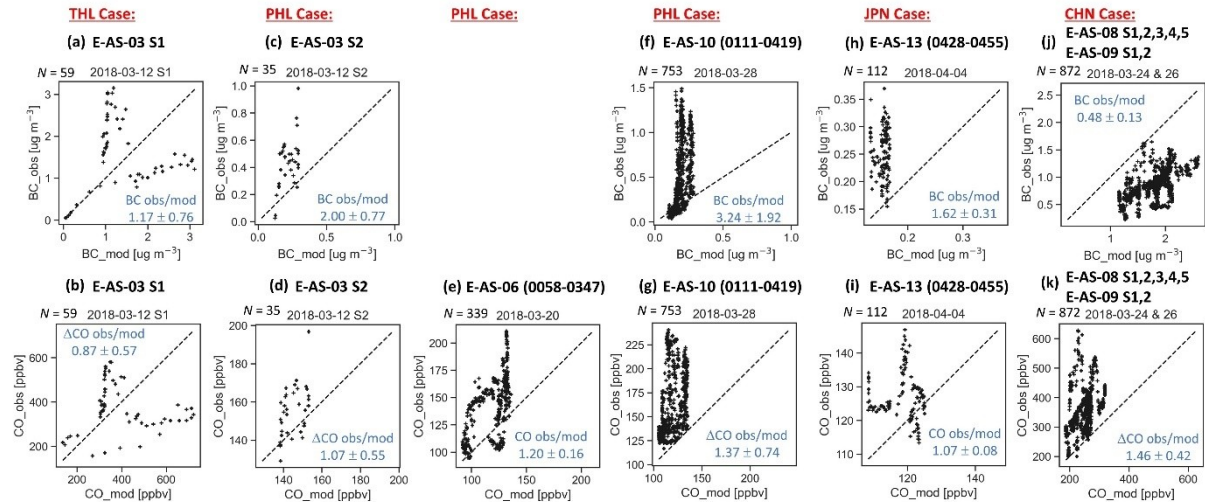
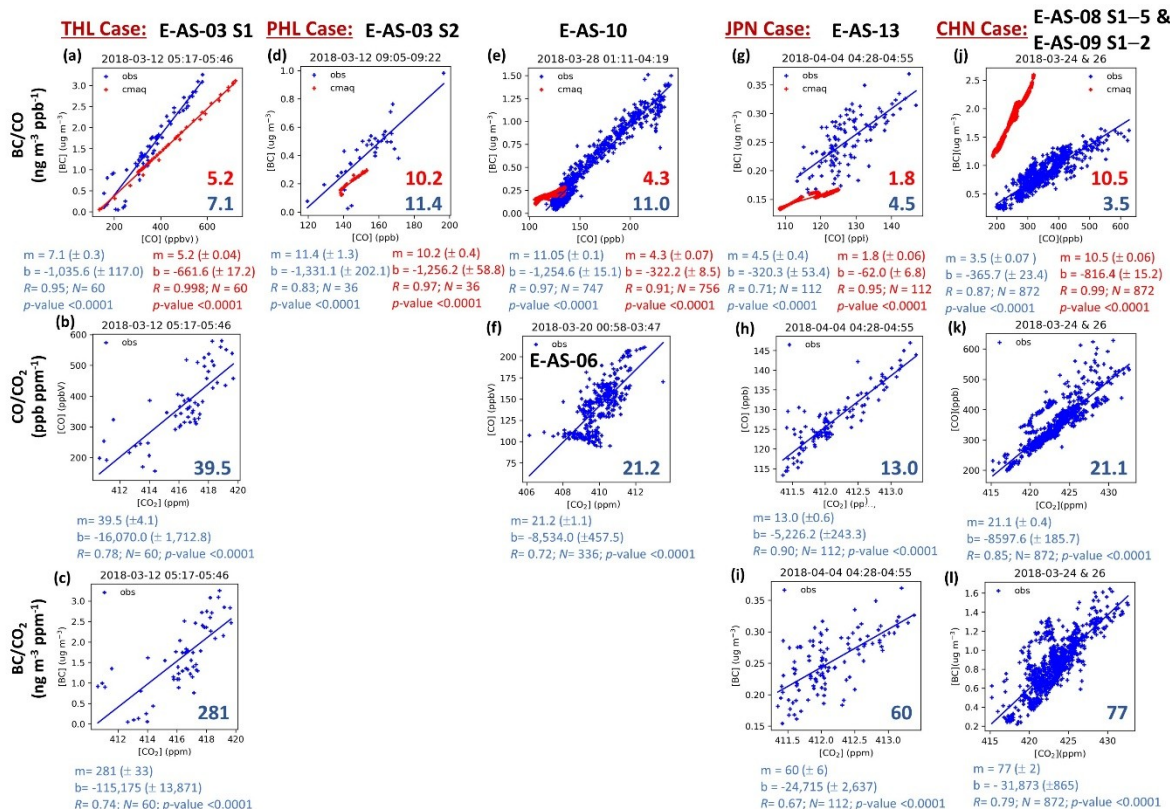


Figure 4. Observed vs. simulated black carbon (BC; upper panels) and CO (lower panels) concentrations. Dashed black lines indicate 1:1 agreement. Observation/model ratios for BC and Δ CO (Eqs. 1, 2) and standard errors are shown in blue text. N denotes data counts. For (e) and (i), CO ratios are used in place of Δ CO due to unavailable baseline CO. Standard errors are based on full data unless invalid, including: (a–b) E-AS-03 S1 (2-min averages), (e) E-AS-06 (5-min averages), and (f–g) E-AS-10 (filtered with APT 72 h $\leq$ 1 mm). Obs. = observed; Mod. = simulated. Reduced model variability relative to observations in PHL, JPN, and CHN likely stemmed from coarse model resolution ($\sim$ 0.5° vs. $\sim$ 0.01° per 15 s), limited temporal resolution in HTAPv2.2z (CHN) and REASv2.1 (PHL), atmospheric mixing limitations which might reduce spatial contrasts near emission hotspots (PHL), and meteorological mismatches affecting air masses detection (JPN).

Table 2. Statistical summary of observed (Obs.) and simulated (Mod.) BC and CO concentrations for THL (Thailand), PHL (the Philippines), JPN (Japan), and CHN (China), including correlation coefficient (R), data count (N), p -values, maximum (Max.), mean, and standard deviation (σ).

Case studies (Flight legs)		THL case (E-AS-03 S1)	PHL case (E-AS-03 S2)	PHL case (E-AS-06)	PHL case (E-AS-10)	JPN case (E-AS-13)	CHN case (E-AS-08 S1–5 & E-AS-09 S1,2)
Time (UTC)		2018/3/12 0517-0546 (30-s)	2018/3/12 0905-0922 (30-s)	2018/3/20 0058-0347 (30-s)	2018/3/28 0111-0419 (15-s)	2018/4/4 0428-0455 (15-s)	2018/3/24,26 Table S1 (15-s)
BC	R	-0.01	0.49	/	0.54	-0.01	0.63
	N	60	35	/	753	112	872
	P	> 0.05	< 0.001	/	< 0.001	> 0.05	< 0.001
Max. (Mean $\pm\sigma$) ($\mu\text{g m}^{-3}$)	Obs.	3.11 (1.63 $\pm$ 0.80)	0.98 (0.44 $\pm$ 0.18)	/	1.49 (0.48 $\pm$ 0.38)	0.37 (0.26 $\pm$ 0.04)	1.66 (0.84 $\pm$ 0.27)
	Mod.	3.26 (1.40 $\pm$ 0.76)	0.29 (0.22 $\pm$ 0.06)	/	0.28 (0.18 $\pm$ 0.04)	0.17 (0.16 $\pm$ 0.01)	2.60 (1.77 $\pm$ 0.33)
CO	R	-0.08	0.47	0.59	0.51	-0.07	0.4
	N	60	35	339	747	112	872
	P	> 0.05	< 0.05	< 0.001	< 0.001	> 0.05	< 0.001

Max. (Mean±σ) (ppbv)	Obs.	580 (374±106)	197 (155±13)	211 (142±26)	241 (157±34)	149 (130±7)	628 (351±69)
	Mod.	722 (398±147)	153 (145±6)	136 (119±14)	135 (117±9)	125 (119±5)	320 (246±31)



430 **Figure 5. Residual BC/CO, CO/CO₂, and BC/CO₂ ratios, expressed as regression slopes, for (a–c) Thailand (THL), (d–f) Philippines (PHL), (g–i) Japan (JPN), and (j–l) China (CHN). Blue (observed) and red (simulated) lines show linear fits; regression slopes (m), intercepts (b), correlation coefficients (*R*), data counts (*N*), and *p*-values. CMAQ simulations shown only for BC/CO.**

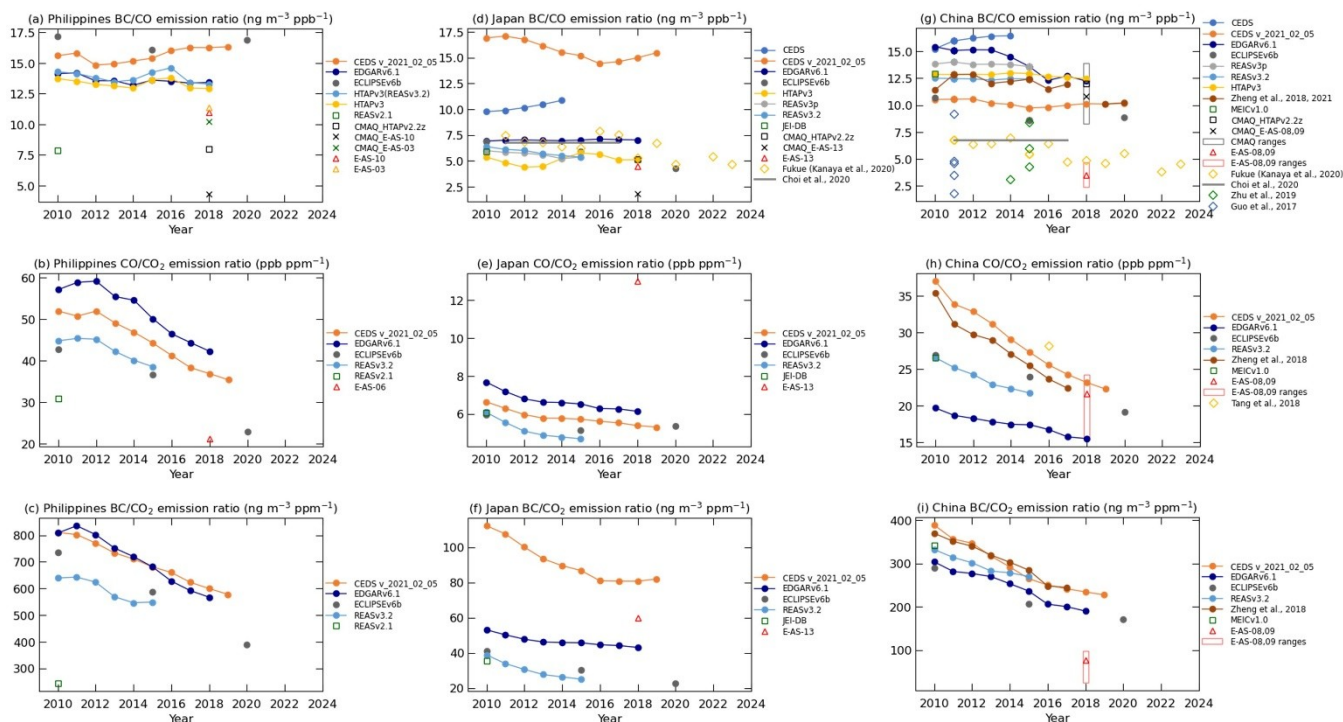


Figure 6. Residual BC/CO, CO/CO₂, and BC/CO₂ ratios for (a–c) Philippines (PHL), (d–f) Japan (JPN), and (g–i) China (CHN), based on aircraft-based observations (red triangles) and CMAQ model simulations (crosses). In panels (g–i), boxes represent the observed (red) and simulated (black) ratio ranges in sub-air masses. Decadal trends of these ratios based on previous observations (diamonds, grey lines) and bottom-up inventories including MIX (green squares), HTAPv2.2z (black squares), and other inventories (solid markers, with or without connecting lines) are presented. These panels showed that CO/CO₂ and BC/CO₂ ratios have declined in these regions while BC/CO ratios have remained largely unchanged except in the EDGARv6.1 inventory. The BC/CO ratio agreement in (a) for E-AS-03 likely reflects accurate emission factors but underestimated activity levels in REASv2.1. Reasonable variations in deposition parameters cannot account for this match, as modelled BC and CO distributions showed no evidence of enhanced emissions with rapid removal below the flight altitudes (not shown).

3.2 Comprehensive reassessment of CHN emissions from HTAPv2.2z inventory

Taking advantage of the long-range transport from China captured by aircraft and supported by long-term, ground-based observations at Fukue Island (Kanaya et al., 2020), the observed CHN air masses are explored in greater detail, including an observation–model comparison, assessment of emission ratios (Sect. 2.4), and an estimate of emissions using a combined model-observation approach (Sect. 2.5).

3.2.1 Observed and simulated BC, CO, CO₂ concentrations in CHN anthropogenic emissions

During flights E-AS-08 (24 March 2018, 0227–0309 UTC and 0424–0607 UTC) and E-AS-09 (26 March 2018, 0258–0419 UTC), air masses with peak BC and CO concentrations were observed near Taiwan, Jeju (South Korea), and Okinawa (JPN), which originated in Central East China (CEC) (Fig. 3f, g), and totalled 872 data points. These flights captured pollution plumes characteristic of dominant Chinese anthropogenic sources and seasonal patterns. Fire activity, typically peaking in March and

April, contributed only ~10 % of CHN anthropogenic emissions (7 % BC, 13 % CO; according to GFED and HTAPv2.2z used in CMAQ). Trajectory analysis showed plumes originating from Chinese source regions around 22 March 2018 (Thursday), with transport time of 40–120 h. The air mass was likely influenced by typical weekday emissions from the transportation
455 sector. This plume represented a significant pollution transport event with a typical transport route and negligible deposition loss (via APT analysis; see Sect. 2.3 and Fig. S6). Under the high-pressure system, low specific humidity (mean 3.6 ± 1.1 g/kg) and relative humidity (mean 57 ± 13 %) over 72 h transport from HYSPLIT model suggested negligible cloud formation and high moisture influence.

The maximum observed BC and CO concentrations were $1.66 \mu\text{g m}^{-3}$ and 628 ppbv, respectively, with CO_2 increased from
460 419 to 433 ppm. The correlation coefficients between the modelled and observed concentrations were moderate ($R_{\text{BC}} = 0.63$ and $R_{\text{CO}} = 0.40$, p -values < 0.001 ; see Table 2). The model overestimated BC by maxima of $+1.62 \mu\text{g m}^{-3}$ and underestimated CO by maxima of -400 ppbv, predicting maximum concentrations of $2.60 \mu\text{g m}^{-3}$ for BC and 320 ppbv for CO (Fig. 2h–k). Kanaya et al. (2020) reported similar tendencies for CMAQv4.7.1 (REASv2.1) model at Fukue Island (JPN), including for E-AS-08 and E-AS-09 periods. The observation/simulation ratios (Eqs. 1 and 2) were 0.48 ± 0.13 for BC and 1.46 ± 0.42 for ΔCO
465 (Fig. 4j and 6k). Although the HTAPv2.2z inventory considers recent reductions in Chinese emissions (Chatani et al., 2020; Zheng et al., 2018), these updates may still be insufficient for accurately predicting BC concentrations.

CHN air masses were investigated separately for two groups of origin (Table S1). S-CEC (southern CEC) group included flights E-AS-08 S1–S3 and E-AS-09 S2 (mean $\text{BC} = 0.73 \mu\text{g m}^{-3}$ and mean $\Delta\text{CO} = 181$ ppbv). NS-CEC (northern and southern CEC) group included flights E-AS-08 S4–S5 and E-AS-09 S1 (mean $\text{BC} = 0.85 \mu\text{g m}^{-3}$ and mean $\Delta\text{CO} = 338$ ppbv). Kanaya et al. (2020) recorded similar peaks for N-CEC (northern CEC) and S-CEC on Fukue Island on March 24–26 and 26–28. Our
470 mean S-CEC results closely matched those of Kanaya et al. (2020) ($0.67 \mu\text{g m}^{-3}$ BC and 190 ppbv ΔCO). Our NS-CEC mean values were similar to the N-CEC values on Fukue Island ($0.92 \mu\text{g m}^{-3}$ BC and 254 ppbv ΔCO) (Table S4 in the supplement). The largest model biases occurred for flight E-AS-08 S3 and S5 ($+1.18 \mu\text{g m}^{-3}$ BC and -364 ppbv CO, respectively), both at an altitude of ~900 m.

475 3.2.2 CHN emission ratios of BC/CO, CO/CO₂, and BC/CO₂

The observed residual ratios for CHN air masses were $3.5 \pm 0.1 \text{ ng m}^{-3} \text{ ppb}^{-1}$ ($R = 0.87$) for BC/CO, $21.1 \pm 0.4 \text{ ppb ppm}^{-1}$ for CO/CO₂ ($R = 0.85$), and $77 \pm 2 \text{ ng m}^{-3} \text{ ppb}^{-1}$ for BC/CO₂ ($R = 0.79$). These ratios remained largely unchanged despite slight rainfall, shown by the APT analysis (Sect. 2.3; Table S7) and minimal variation (1.5–2.3 %) between the full dataset and APT-filtered subsets (APT 72 h = 0 mm; Fig. S6), supporting their robustness for diagnosing CHN anthropogenic emissions (p -
480 value < 0.0001). In contrast, the CMAQ model overestimated the BC/CO ratio ($10.5 \pm 0.1 \text{ ng m}^{-3} \text{ ppb}^{-1}$; $R = 0.99$). The observed ratio agreed with the lower end of ground- and ship-based measurements, ranging from $(6.2\text{--}7.9) \pm 0.7 \text{ ng m}^{-3} \text{ ppb}^{-1}$ (Choi et al., 2020), $3.8\text{--}7.0 \text{ ng m}^{-3} \text{ ppb}^{-1}$ (Kanaya et al., 2020), $1.8\text{--}3.5 \text{ ng m}^{-3} \text{ ppb}^{-1}$ (Guo et al., 2017), and $3\text{--}6 \text{ ng m}^{-3} \text{ ppb}^{-1}$ (Zhu et al., 2019). Note that the aethalometer-based value (Zhu et al., 2019) required correction for non-BC/BC mass ratio. Fig. 6g–i

compares observed ratios (red triangles) and CHN sub-air mass ranges (red boxes) to other observations and bottom-up
485 inventories. The observed BC/CO ratios were less than one-third of the HTAPv2.2z value ($12 \text{ ng m}^{-3} \text{ ppb}^{-1}$; black square) and
half of other inventories ($8.4\text{--}16.4 \text{ ng m}^{-3} \text{ ppb}^{-1}$). Even the lowest inventory estimate, CEDS v_2021_02_05 (10.1 ng m^{-3}
 ppb^{-1} ; O'Rourke et al., 2021), exceeds observations, despite reductions from its predecessor (CEDS CMIP6), indicating a
positive bias in CHN BC/CO ratios in bottom-up inventories, consistent with REASv2.1 analyses (Choi et al., 2020).

The observed CO/CO₂ ratio ($21.6 \text{ ppb ppm}^{-1}$) closely matched CEDS v_2021_02_05 value ($23.2 \text{ ppb ppm}^{-1}$), though lower
490 than the 2016 KORUS-AQ measurement ($28.2 \text{ ppb ppm}^{-1}$) (Tang et al., 2018). CHN sub-air masses CO/CO₂ ratios ($15.5\text{--}23.2$
 ppb ppm^{-1}) agreed well with the 2018 REASv3.2 and ECLIPSEv6b (Fig. 6h), confirming the reliability of CO/CO₂ ratios in
bottom-up inventories for CHN. Given the lower uncertainties in CO₂ emissions in these inventories than CO and BC (Zhao
et al., 2013), this agreement supported the reasonable CO emissions for 2018 in inventories. The decline in CO/CO₂ ratio
compared to past airborne measurements ($28\text{--}60 \text{ ppb ppm}^{-1}$) taken ~20 years ago (Takegawa et al., 2004; Suntharalingam et
495 al., 2004) reflected dramatically improved combustion efficiencies in CHN (Wang et al., 2010; Tang et al., 2018). Conversely,
the observed BC/CO₂ ratio ($77 \text{ ng m}^{-3} \text{ ppm}^{-1}$) was less than half the inventory range ($190\text{--}234 \text{ ng m}^{-3} \text{ ppm}^{-1}$) (Fig. 6i),
indicating BC emissions are overestimated in bottom-up inventories.

HTAPv2.2z identifies the residential sector as the dominant source of BC (59 %) and CO (54 %) emissions in CHN. Other
inventories, such as ECLIPSEv6b, also overestimate BC emissions. Detailed review of ECLIPSEv6b revealed the likely cause:
500 high assumed emission factors for raw coal in traditional stoves ($0.06\text{--}5.76 \text{ g kg}^{-1}$), which represent the largest activity share
(51 %) and contributed 80 % of BC emissions (Klimont, personal communications, 2021; Table S3). Emission factors are
highly sensitive to fuel quality and combustion conditions (Li et al., 2017a; Zhao et al., 2013). For example, medium-volatile
bituminous raw coal has higher emission factors ($>10.1 \text{ g kg}^{-1}$) than other fuels (Chen et al., 2006, 2009; Sun et al., 2017;
Zhang et al., 2008; Zhi et al., 2008). However, such emissions were expected to decline following the 2013 National Air
505 Quality Action Plan and 2015 Clean Coal Technology Policies (Chen et al., 2019; Zhang et al., 2022), with cleaner technologies
like honeycomb-coal-briquette. While BC emissions from this source were projected at 16.7 Gg in 2020 (Chen et al., 2009),
ECLIPSEv6b assigns 401 Gg, indicating substantial overestimation.

BC emissions showed greater spatial and temporal heterogeneity than CO and CO₂. NS-CEC air masses exhibited BC/CO
ratios $< 2.9 \text{ ng m}^{-3} \text{ ppb}^{-1}$ ($2.4\text{--}2.7 \text{ ng m}^{-3} \text{ ppb}^{-1}$), while S-CEC air masses had higher ratios ($3.5 \text{ ng m}^{-3} \text{ ppb}^{-1}$, ranging $3.6\text{--}4.7$
510 $\text{ng m}^{-3} \text{ ppb}^{-1}$) (Figs. S3 and S4). This likely reflected different regional emission changes between 2010 and 2018, with greater
reductions in BC/CO ratios in N-CEC than in S-CEC (Kanaya et al., 2020). The airborne BC/CO ratios for S-CEC agreed with
Fukue Island values during the same episode ($3.5 \text{ ng m}^{-3} \text{ ppb}^{-1}$; Table S4), confirming for the first time the stability of BC/CO
ratio during transport, proving that the emissions are reliable. However, these ratios were markedly lower than Fukue's spring
means for N-CEC ($4.4 \text{ ng m}^{-3} \text{ ppb}^{-1}$) and S-CEC ($5.2 \text{ ng m}^{-3} \text{ ppb}^{-1}$). While airborne BC/CO ratios may not fully capture
515 national-scale variability, the data demonstrated vertical consistency from ground to flight altitude. Future airborne
measurements with larger footprints will further improve national emission estimates.

While emission ratios reflect source characteristics and support national-scale diagnostics, this observation alone cannot attribute emissions to specific sectors. However, the BC/CO ratio implies that model BC and CO emissions should be respectively reduced and increased. Although indirectly, typical dominant sources, e.g. coal-fuel stoves, can be potential contributors. A thorough review is essential to identify outdated input and incorporate recent policy shifts or experimental data.

3.2.3 Estimates of BC, CO, and CO₂ emissions from CHN using the combined model-observation approach

The combined model–observation approach (Sect. 2.5) was applied to CHN air masses using either E(BC) or E(CO) values as a basis. The HTAPv2.2z inventory prescribed Chinese emissions of 1.36 Tg yr⁻¹ for BC and 134 Tg yr⁻¹ for CO (Fig. 7) to use in Eq. 3. The other CHN emission rates (Eqs. 4–6) were translated using the mean BC/CO ratio from Fukue Island observations (Kanaya et al., 2020) in spring 2018 ($4.9 \pm 0.1 \text{ ng m}^{-3} \text{ ppb}^{-1}$) and airborne CO/CO₂ ratio ($21.1 \pm 0.4 \text{ ppb ppm}^{-1}$) (Fig. 7). The method used to estimate uncertainties is discussed later (Sect. 3.2.4).

Using E(BC)_{HTAPv2.2z} as a basis to estimate emissions

Eq. 1 yielded an emission correction factor $E(\text{BC})_{\text{HTAPv2.2z}} = 0.48 \pm 0.13$ (Fig. 7a). The estimation results and uncertainty ranges (Sect. 3.2.4) are shown in Fig. 8 (red triangles and boxes) for comparison with previous studies. The calculated BC emission rate was $0.65 \pm 0.25 \text{ Tg yr}^{-1}$, which was ~52 % lower than the HTAPv2.2z inventory value of 1.36 (Tg BC) yr⁻¹ (this study, marked with a black cross in Fig. 8a). Similar to recent estimates made using observations on Fukue Island and in the Noto region (Kanaya et al., 2020; Ikeda et al., 2022, 2023; Miyakawa et al., 2023), our estimated BC emission rate was lower than bottom-up inventory values. Our BC emission rate was 38 % lower than emission rates estimated using the long-term surface observations on Fukue Island during 2009–2018 ($1.06 \pm 0.29 \text{ Tg yr}^{-1}$; Kanaya et al., 2020; green line and shade in Fig. 8a). This difference may in part arise from different selections of the CHN air mass. The ECLIPSEv6b BC emission rate (0.96 Tg yr^{-1} ; dark grey circles in Fig. 8a) incorporated CHN air quality improvements under the 12th Five–Year plan (Klimont et al. 2017) closely matched our estimated uncertainty range ($0.4\text{--}0.9 \text{ Tg yr}^{-1}$; red box in Fig. 8a). The CEDS (CMIP6) BC rate (2.54 Tg yr^{-1} for 2014; blue line with circles in Fig. 8a) significantly decreased in CEDS v_2021_02_05 (1.58 Tg yr^{-1} for 2014, 1.22 Tg yr^{-1} for 2018; orange line with circles in Fig. 8a) reflecting updated International Energy Agency energy data with less biomass consumed after 2000 and improved BC and OC industrial emissions from CHN (O'Rourke et al., 2021). Despite the reductions, CEDS v_2021_02_05 BC rate remained higher than our estimate.

The estimated CO emission rate was $166 \pm 65 \text{ Tg yr}^{-1}$, which was ~24 % higher than the value in the HTAPv2.2z inventory (134 Tg yr^{-1} ; this study, marked with a black cross in Fig. 8b). Our estimate was 10–32 % higher than those of other inventories, including the highest (CEDS v_2021_02_05: 150 Tg yr^{-1} ; orange line with circles, Fig. 8b) and lowest (EDGARv6.1: 114 Tg yr^{-1} ; dark blue line with circles, Fig. 8b). These values in CEDS v_2021_02_05 and EDGARv6.1 relied on activity data and emission factors without observation-based corrections (Crippa et al., 2018). The top-down TCR2 CO rate of 153 Tg yr^{-1} for 2019–2020 (red line with stars in Fig. 8b) was close to our estimate. Our estimated CO₂ emission rate was $(12.4 \pm 4.8) \times 10^3 \text{ Tg yr}^{-1}$, which was similar to the rates estimated by EDGARv6.1 ($11.5 \times 10^3 \text{ Tg yr}^{-1}$; dark blue line with circles in Fig. 8c) and

550 CEDS v_2021_02_05 ($10.2 \times 10^3 \text{ Tg yr}^{-1}$; orange line with circles in Fig. 8c), to take our uncertainty into account. All inventories indicated an increase in CO₂ emissions from 2010 to 2020 that exceeded the reported 2018 carbon budget for CHN of $9.9 \times 10^3 \text{ Tg yr}^{-1}$ (black dashes in Fig. 8c) (Friedlingstein et al., 2020).

Using E(CO)_{HTAPv2.2z} as a basis to estimate emissions

Eq. 2 resulted in a correction factor of $E(\text{CO})_{\text{HTAPv2.2z}} = 1.46 \pm 0.42$ (Fig. 7b). The results and uncertainty ranges are shown as green triangles and boxes in Fig. 8. The estimated CO emission rate was $195 \pm 59 \text{ Tg yr}^{-1}$; the HTAPv2.2z inventory value was 46 % lower, yet it was still close to the edge of our uncertainty range (black cross and green box in Fig. 8b). Only the CEDS v_2021_02_05 CO rate (orange line with circles in Fig. 8b) among other inventories fell within this range. This method yielded a higher BC emission rate ($0.77 \pm 0.23 \text{ Tg yr}^{-1}$) and lower uncertainty than the E(BC)-based method due to reduced instrument and model-specific transport errors (Sect. 3.2.4). However, the BC rate remained 44 % lower than the HTAPv2.2z inventory value (black cross in Fig. 8a). The estimated CO₂ emission rate was $(14.5 \pm 4.4) \times 10^3 \text{ Tg yr}^{-1}$, which covered the EDGARv6.1 ($11.5 \times 10^3 \text{ Tg yr}^{-1}$; dark blue line with circles in Fig. 8c) and CEDS v_2021_02_05 ($10.2 \times 10^3 \text{ Tg yr}^{-1}$; orange line with circles in Fig. 8c). Our CO₂ estimates remained higher than those of previous studies. An extra 7 % of CO₂ emissions from humans and livestock respiration in CHN (Cai et al., 2022), which is typically excluded from bottom-up inventories, could partially explain this high CO₂ emission rate. When compared to the CO and CO₂ emission reference ranges, the E(CO)-based method produced worse estimates of CO and CO₂ than the E(BC)-based method.

565 Overall, this multi-species analysis using either E(BC) and E(CO) method both underscores the necessity of adjusting CHN emission inventories. The HTAPv2.2z (this study), EDGARv6.1, and CEDS_2021_02_05 inventories require downward corrections for BC emissions and upward revisions for CO emissions to better agree with observed data. BC emissions in HTAPv2.2z require a 45–54 % reduction, while BC emissions in the EDGARv6.1, HTAPv3, and CEDS v_2021_02_05 inventories require a 31–49 % decrease to agree with our estimates. Lower BC emission rates would more accurately reflect bottom-up CO₂ emissions, which have been confirmed for the first time. CO emissions in HTAPv2.2z require a 24–46 % increase and CO emissions in the EDGARv6.1, HTAPv3, and CEDS v_2021_02_05 inventories require a 10–42 % increase to agree with our estimates (see Table S2 in the supplement).

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3.2.4 Uncertainty quantification in CHN emission estimates

The uncertainties in emission estimates were propagated from four sources. (1) Error in observation-to-model ratios and multi-species emission ratios was ~26 % for the E(BC)-based method, including standard errors. (2) Systematic instrument error was 10 % for the SP2 (Ohata et al., 2021). We estimated (3) representation errors as ± 4 % due to limited opportunities for aircraft observations to be made in terms of seasonal and spatial variabilities, estimated by comparing spring and annual data across CEC region and large footprint for CHN recorded at Fukue Island, assuming CEC emissions dominate BC sources (69.7 %) (Fig. S5 in the supplement; Kanaya et al., 2020). (4) A model-specific transport error (26 %) was determined by comparing CMAQ-simulated BC concentrations with the Comprehensive Air Quality Model with eXtensions (CAMx) system simulations

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using identical emission data (Chatani et al., 2018) for the model comparison framework (Itahashi et al., 2018; Yamaji et al., 2020). Despite differences in meteorology and aerosol components, CAMx produced a correction factor of 0.54, closely matching CMAQ's E(BC) of 0.48. This result indicated small transport and meteorology errors, suggesting that the discrepancies between model and observations primarily stemmed from the underlying emission inventory. Overall, the estimated uncertainties in the E(BC)-based emission rates were ~39 % (0.25 Tg yr^{-1} for BC, 65 Tg yr^{-1} for CO, and $4.8 \times 10^3 \text{ Tg yr}^{-1}$ for CO_2). The BC emission uncertainty ($\pm 0.25 \text{ Tg yr}^{-1}$) closely matched the surface-based estimates at Fukue Island ($\pm 0.29 \text{ Tg yr}^{-1}$), although they studied for longer period (2009–2018) (Kanaya et al., 2020).

For emissions estimated using E(CO), the uncertainties include: (1) observation-to-model errors (~28 %), (2) instrument errors (4 % for the UV photometer; Gerbig et al., 1996), (3) representation errors ($\pm 4 \%$), and (4) transport errors (7 %). These contributed to an overall uncertainty of ~30 % (0.23 Tg yr^{-1} for BC, 59 Tg yr^{-1} for CO, and $4.4 \times 10^3 \text{ Tg yr}^{-1}$ for CO_2).

Fig. 9 illustrates the consistency of BC, CO, and CO_2 emission rates using the E(BC) (orange triangles) and E(CO) (green triangles) method. Estimate uncertainties (orange and green lines with circles) were compared to reference values for CO ($113\text{--}150 \text{ (Tg CO) yr}^{-1}$) and CO_2 ($(9.9\text{--}11.5) \times 10^3 \text{ (Tg CO}_2\text{) yr}^{-1}$; red lines). The E(BC) method yielded CO and CO_2 estimates more agreed with references, while its BC estimate deviated further from the reference range ($1.11\text{--}1.29 \text{ Tg yr}^{-1}$; red line on BC axis) than E(CO). As uncertainties in CO_2 emissions across emission inventories are lower than those in BC and CO, the strong CO/ CO_2 ratio agreement (Sect. 3.2.2) suggested that bottom-up inventories provide more accurate CO than BC emissions. Thus, the E(BC)-based was the superior estimation method, producing more reliable CO and CO_2 estimates with smaller uncertainties (standard errors of 26 % vs 31 %; Fig. 7). Large variations were also found in CO baselines calculated from BC/CO ratio of the sub-air masses, as well as E(CO) between NS-CEC and S-CEC sub-air masses (1.93 vs. 1.29).

This study focuses on total national emissions rather than sector-specific contributions. Although the event footprint was limited, it covered key emission regions, i.e. CEC region which accounts for ~ 69.7 % of CHN BC sources. To support the extrapolation to annual, nation-wide emission estimates, a long-term ground-based analysis with broader spatial footprint (Kanaya et al., 2020) was utilized. Temporal, spatial, and model-specific uncertainties were effectively incorporated, ensuring the estimates are representative for China in 2018. Given the consistent regional emission patterns across inventories exhibiting near-linear concentration-emission response (Ikeda et al., 2022), the choice of emission inventory does not affect the finding of this study.

Refining emission inventories for BC can be achieved by targeting major sources based on emission factors and activity levels. For CO and CO_2 , where source attribution is more challenging, improving accuracy may involve updating emission factors in BC-associated sectors and identifying additional potential missing sources.

In conclusion, multi-species analysis suggested that bottom-up emission estimates for CHN in the HTAPv2.2z emission inventory need to be decreased by ~50 % for BC and increased by ~20 % for CO to minimize the differences from measured data based on the E(BC) method. This implies broader trends showing a decline in BC emissions since 2010 (Kanaya et al., 2020), with levels remaining low in 2019–2020 ($\sim 0.6 \text{ Tg BC yr}^{-1}$) according to the translation from TCR2 CO data. However, inventories such as CEDS v2021-02-05 and MEICv1.0 (Zhang et al., 2021) still report BC emissions over 1 Tg yr^{-1} for

2018–2020, reinforcing the need for downward adjustment. The decreasing trend and overestimations of BC emissions in CHN suggests a smaller warming effect than previously estimated in the CMIP6 model (Ikeda et al., 2023). Post-2018 changes in China such as promoted biomass energy use (IEA Bioenergy Country Report 2021; China Energy Transformation Outlook 2023) and major pollution from vehicles despite advances in electrification and emissions control, however, may shift BC, CO, and CO₂ trends unpredictably, warranting future investigations. Nevertheless, the inventory revision for 2018 suggested in this study remain worth considering.

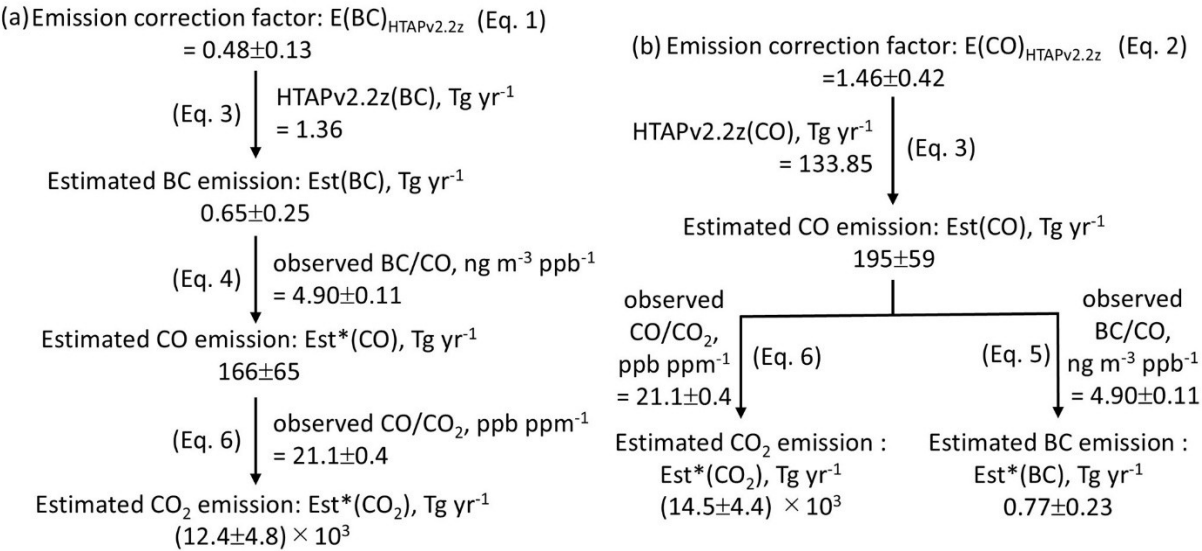
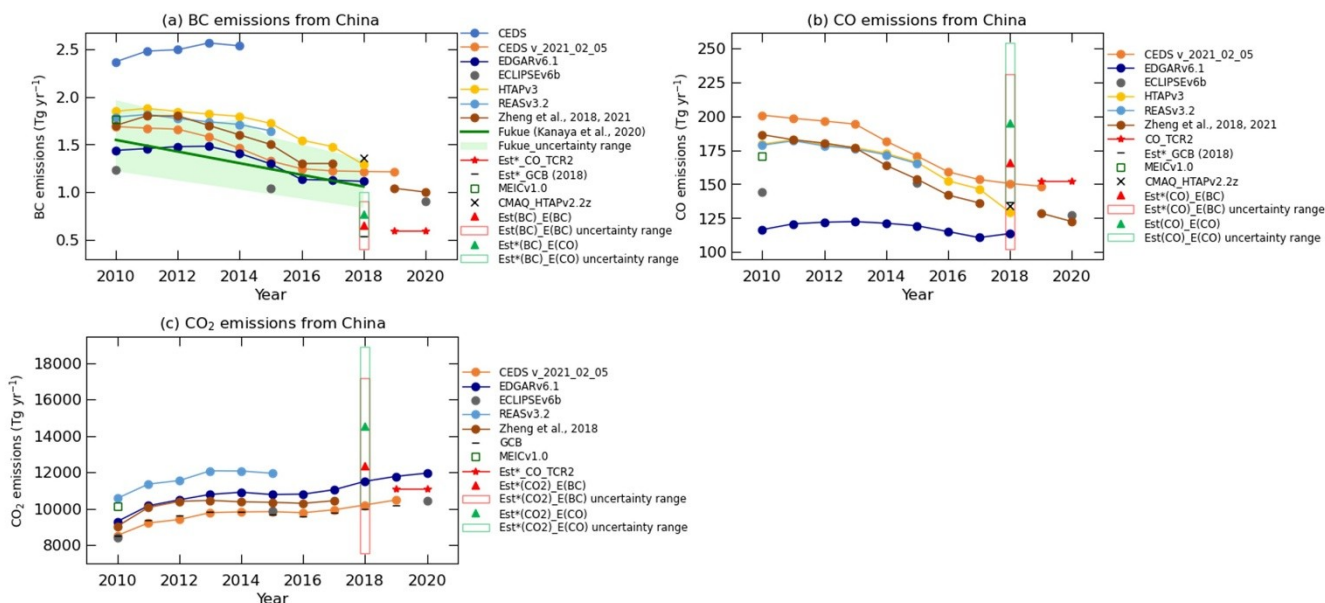


Figure 7. Combined model–observation approach for estimating black carbon (BC), CO, and CO₂ emissions using emission correction factors (a) E(BC) and (b) E(CO) applied to aircraft-based observations relative to the HTAPv2.2z inventory



625 **Figure 8.** Estimated Chinese BC, CO, and CO₂ emissions based on E(BC)_{HTAPv2.2z} (solid red triangles) and E(CO)_{HTAPv2.2z} (solid green triangles), with uncertainty ranges (Sect. 3.2.4) shown as red and green boxes. Comparative data are included from **bottom-up inventories and databases**, MEIC v1.0 (open green squares), HTAPv2.2z (black crosses), the global carbon budget (Friedlingstein et al., 2020; black dashes), TCR2 (Miyazaki et al., 2020; red lines with stars), and BC estimates from Fukue Island (Kanaya et al., 2020; green) along with its **uncertainty** range (light green **shade**). Full references are provided in Table S2 in the supplement.

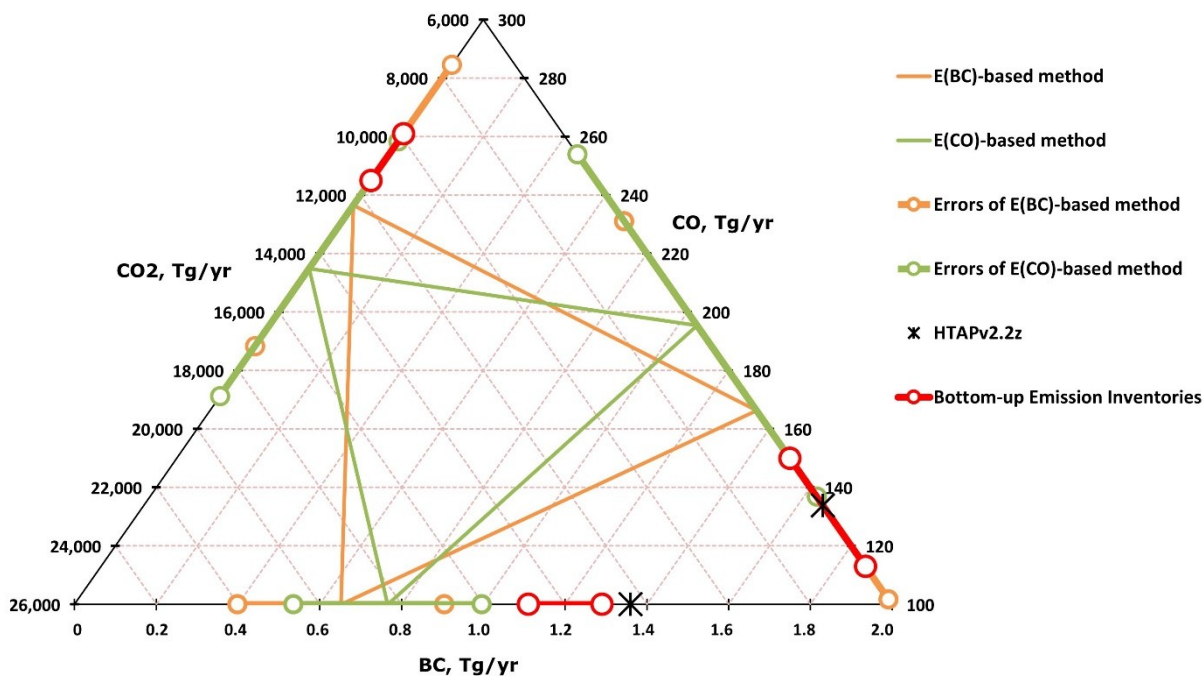


Figure 9. Cross-species consistency of estimated BC, CO, and CO₂ emissions. Orange and green triangles denote E(BC)_{HTAPv2.2z}- and E(CO)_{HTAPv2.2z}-based estimates, respectively, with associated uncertainty ranges shown as orange and green lines with circles. Red lines with circles indicate reference values from bottom-up inventories, while black stars represent HTAPv2.2z emissions.

4 Conclusions

635 This study integrated EMeRGe aircraft measurements with WRF/CMAQ model simulations to assess combustion-related SLCF emissions (BC, CO, and CO₂) in East Asia. Four pollution events from March–April 2018 were investigated: biomass burning near the Gulf of Thailand (THL) and urban pollution in the Philippines (PHL), Japan (JPN), and China (CHN). Distinctive emission inventories were used for each pollution case. Our analyses included: (1) model–observation comparisons; (2) assessment of residual and emission ratios across air masses; and (3) estimation of CHN emissions using a
640 combined model–observation approach. The analysis results led to following main conclusions:

a) The GFEDv4.1s inventory of fire emissions in the THL case provided accurate BC and CO emissions. Residual ratios of co-emitted pollutants were used to assess emission characteristics. The BC/CO concentration ratio was found to be $7.1 \pm 0.3 \text{ ng m}^{-3} \text{ ppb}^{-1}$ for flaming-dominant combustion over the Indochina Peninsula.

b) Simulations of PHL air masses showed that REASv2.1 (2010) underestimated the concentrations of BC and CO in
645 Manila. The BC/CO concentration ratio for urban emissions was $11.4 \pm 1.3 \text{ ng m}^{-3} \text{ ppb}^{-1}$, whereas REASv2.1 also gave a good range ($10.2 \pm 0.4 \text{ ng m}^{-3} \text{ ppb}^{-1}$). This suggests that the underestimation in REASv2.1 is likely linked to the activity data.

c) The JPN air masses exhibited a high CO/CO₂ ratio, which may not be representative of typical JPN emissions, despite the fact that BC/CO ratios agree with bottom-up inventories and previous observations.

d) Aircraft observations of CHN air masses confirmed that the CMAQ model using HTAPv2.2z inventory overestimated
650 BC concentrations by $+1.62 \text{ } \mu\text{g m}^{-3}$ and underestimated CO concentrations by -400 ppbv , supporting ground-based findings (Kanaya et al., 2020). For the first time, the BC/CO ratio for CHN air masses found in Kanaya et al. (2020) was validated using aircraft-based measurements ($3.5 \pm 0.1 \text{ ng m}^{-3} \text{ ppb}^{-1}$). This ratio agreed with other ground- and ship-based observations, but it was markedly lower than CMAQ (HTAPv2.2z) simulations ($10.5 \pm 0.1 \text{ ng m}^{-3} \text{ ppb}^{-1}$) and other inventories. This implied that BC emissions were overestimated and CO emissions were underestimated in HTAPv2.2z and other inventories. This
655 discrepancy may be due to the rapid decrease in BC emissions from coal combustions and the use of old-fashioned heating stoves in CHN during the decade leading up to 2018.

e) A combined model–observation approach was applied to estimate BC, CO, and CO₂ emission for CHN. The best estimates were based on the E(BC)_{HTAPv2.2z} correction factor using spring mean residual BC/CO ratio ($4.9 \pm 0.1 \text{ ng m}^{-3} \text{ ppb}^{-1}$) and the aircraft-measured CO/CO₂ ratio ($21.1 \pm 0.4 \text{ ppb ppm}^{-1}$). The estimated emission rates were $0.65 \pm 0.25 \text{ Tg BC yr}^{-1}$,
660 $166 \pm 65 \text{ Tg CO yr}^{-1}$, and $(12.4 \pm 4.8) \times 10^3 \text{ Tg CO}_2 \text{ yr}^{-1}$. The BC and CO emission estimates were tied to the CO₂ estimates to reduce uncertainty. The low BC emission rate implies that BC emission in the HTAPv2.2z inventory for CHN needs to be decreased by $\sim 50 \%$, while CO emission requires a $\sim 20 \%$ increase to improve agreement with airborne measurements.

The findings highlight the need of verifying bottom-up emission inventories and identifying gaps to improve air quality simulations and climate impact mitigation strategies. Updated observational and experimental data might support inventory developers to refine and update BC and CO emission factors and adjust the activity levels of dominant BC sources. The overestimations of BC emissions in CHN suggests a smaller warming effect than previously estimated in the CMIP6 model (Ikeda et al., 2023).

Data availability

The primary data collected by HALO aircraft during the EMeRGe campaign are publicly available and can be requested from the HALO database (<https://halo-db.pa.op.dlr.de/>). The data simulated using the CMAQ model can be acquired upon request to Kazuyo Yamaji (kazuyo@maritime.kobe-u.ac.jp).

Author contributions

P Ha conducted all the analysis and composed the text. Y Kanaya supervised the findings of this study and contributed significantly to manuscript revisions. K Yamaji owns the CMAQ model code and provides the model's inputs and results. S Itahashi provided Comprehensive Air quality Model with eXtensions data of BC and CO concentrations, which was used to calculate the error estimate for the transport. S Chatani provided the model-ready emission inputs. T Sekiya provided the CO_TCR2 data. Y Kanaya, K Yamaji, and T Sekiya thoroughly discussed the analysis results before and during the manuscript writing. MD Andrés Hernández, JP Burrows, H Schlager, M Lichtenstern, M Poehlker, B Holanda provided the EMeRGe-Asia campaign's data used in the study. All authors have contributed to the manuscript, corrections, post-writing formatting, and revisions in reviewing and interpreting the results presented.

Competing interests

At least one of the (co-)authors is a member of the editorial board of Atmospheric Chemistry and Physics.

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