Quantifying the Impacts of Marine Aerosols over the Southeast Atlantic Ocean using a chemical transport model: Implications for aerosol-cloud interactions

Mashiat Hossain[1], Rebecca M. Garland[2], Hannah M. Horowitz[1]

[1] Civil and Environmental Engineering, University of Illinois Urbana-Champaign, IL, USA
[2]Department of Geography, Geoinformatics and Meteorology, University of Pretoria, Pretoria, South Africa

Correspondence to: Mashiat Hossain (mashiat3@illinois.edu), Hannah M. Horowitz (hmhorow@illinois.edu)

Abstract. The southeast Atlantic region, characterized by persistent stratocumulus clouds, has one of the highest uncertainties in aerosol radiative forcing and significant variability across climate models. In this study, we analyze the seasonally varying role of marine aerosol sources and identify key uncertainties in aerosol composition at cloud-relevant altitudes over the southeast Atlantic using the GEOS-Chem chemical transport model. We evaluate simulated aerosol optical depth (AOD) and speciated aerosol concentrations against those collected from ground observations and aircraft campaigns such as LASIC, ORACLES, and CLARIFY, conducted during 2017. The model consistently underestimates AOD relative to AERONET, particularly at remote locations like Ascension Island. However, when compared with aerosol mass concentrations from aircraft campaigns during the biomass burning period, it performs adequately at cloud-relevant altitudes, with a normalized mean bias (NMB) between -3.5% (CLARIFY) and -7.5% (ORACLES). At these altitudes, in the model organic aerosols (63%) dominate during the biomass burning period, while sulfate (41%) prevails during austral summer, when dimethylsulfide (DMS) emissions peak in the model. Our findings indicate that marine sulfate can account for up to 69% of total sulfate during high DMS period. Sensitivity analyses indicate that refining DMS emissions and oxidation chemistry may increase sulfate aerosol produced from marine sources, highlighting that there remains large uncertainty in the role of DMS emissions in the marine boundary layer. Additionally, we find marine primary organic aerosol emissions may substantially increase total organic aerosol concentrations, particularly during austral summer. This study underscores the imperative need to refine marine emissions and their chemical transformations, as aerosols from marine sources are a major component of total aerosols at cloud-relevant altitudes and may impact uncertainties in aerosol radiative forcing over the southeast Atlantic.

1 Introduction

Marine aerosols are a primary contributor to natural atmospheric aerosols, and consequently influence the Earth's radiative balance (Spracklen et al., 2008; Vignati et al., 2001). Aerosols in the marine boundary layer have significant impact on the properties of low-altitude marine clouds, particularly their ability to reflect solar radiation and cool the climate (Seinfeld and Pandis, 2016; Wood, 2012; Chen et al., 2014; Quinn et al., 2017). The southeast Atlantic (SEA) is marked by a persistent deck of low-level stratocumulus (Sc) clouds. However, this region exhibits highest uncertainty in aerosol radiative forcings in the AeroCom intercomparison across CMIP5 general circulation models (GCMs) and chemical transport models (Stier et al., 2013). This uncertainty is primarily driven by challenges in

accurately representing cloud fraction, aerosol-cloud properties, and vertical structure, both in the presence and absence of smoke (Stier et al., 2013; Doherty et al., 2022). In this study, we investigate the role of marine aerosols and sources of uncertainty affecting aerosol composition within the boundary layer, particularly in this critical region of aerosol-cloud interactions over the SEA.

The SEA region encompasses the Benguela upwelling system (BUS), renowned for its high primary production of marine phytoplankton and fish populations (Shannon and Nelson, 1996; Jarre et al., 2015). This elevated phytoplankton activity serves as the main natural source of the volatile organic compound dimethylsulfide (DMS), thereby influencing the global tropospheric sulfur budget (Andreae, 1990; Bates et al., 1992). Once released into the atmosphere through air-sea exchange, DMS undergoes complex chemical transformations. In the gas phase, it is oxidized to form H_2SO_4 and methanesulfonic acid (MSA), which has implications for new particle formation (Chen et al., 2015); while in the aqueous phase, it leads to the production of MSA and sulfate aerosols, impacting cloud microphysical properties (Kaufman and Tanré, 1994). Although DMS is a critical source of natural aerosols, contributing over 50% of natural gas-phase sulfur emissions (Chin et al., 1996; Kilgour et al., 2021), the exact mechanisms of DMS oxidation and subsequent formation of sulfate and MSA aerosol remain inadequately understood (Ravishankara et al., 1997; Barnes et al., 2006; Hoffmann et al., 2016). This gap in understanding contributes to substantial uncertainty in aerosol radiative forcing, which is highly sensitive to uncertainties in natural aerosols (Carslaw et al., 2013; Fung et al., 2022). Additionally, marine aerosols comprise primary aerosols such as sea spray aerosols, which consist of salts, sulfate, and organic matter, released into the atmosphere primarily by the bubble-bursting process (O'Dowd and De Leeuw, 2007; Russell et al., 2010; Prather et al., 2013; Brooks and Thornton, 2018; Russell et al., 2023). Investigating the uncertainties related to marine emissions and chemistry are crucial to refine our understanding of the impacts of marine aerosols on climate.

The SEA lies at the confluence of not only marine aerosols, but other natural and anthropogenic aerosols from local and distant origin (Andreae et al., 1995; Swap et al., 1996; Formenti et al., 1999; Swap et al., 2003; Tournadre, 2014). During the austral spring (August to October), seasonal fires in the neighboring southern African region contribute nearly one-third of global total biomass burning emissions (van der Werf et al., 2010). This seasonal influx of biomass burning aerosols aloft interacts with the underlying Sc deck, introducing considerable variability into aerosol forcing assessments in the SEA region (Lindesay et al., 1996; Swap et al., 2003). To address these uncertainties, several international field campaigns were conducted between 1992 and 2018 during the peak biomass burning season (Swap et al., 2003; Formenti et al., 2019; Haywood et al., 2021; Redemann et al., 2021). Despite the region being a prolific source of marine aerosols throughout the year, the potential impact of aerosols on regional climate dynamics through interactions with the persistent low-level marine clouds outside of the biomass burning season has been largely overlooked.

Here, we use the GEOS-Chem global chemical transport model to analyze high-resolution, seasonally varying aerosol composition at the altitudes of persistent stratocumulus clouds over the SEA. We specifically focus on the role of marine aerosols, analyzing their contributions to sulfate and organic aerosol concentrations. We evaluate simulated

aerosol optical depth (AOD) and speciated aerosol concentrations against observational data from the Aerosol Robotic Network (AERONET) and the Layered Atlantic Smoke Interactions with Clouds (LASIC; Zuidema et al., 2018), ObseRvations of Aerosols above CLouds and their intEractionS (ORACLES; Redemann et al., 2021), and CLoud–Aerosol–Radiation Interaction and Forcing (CLARIFY; Haywood et al., 2021) field campaigns during the year 2017. We assess the sensitivity of our results to uncertainty in DMS oxidation mechanisms and emissions of DMS, SO_2, and marine primary organics. Our findings aim to enhance understanding of the seasonally varying role of marine aerosols in aerosol-cloud interactions in the SEA by a comprehensive evaluation of aerosol composition at cloud altitudes.

2 Methodology

2.1 Model Description

Here, we use the GEOS-Chem 3D atmospheric chemical transport model version 13.3.3 with detailed gas- and aerosol-phase tropospheric chemistry (https://zenodo.org/records/5748260). The model is driven by meteorology from the Modern-Era Retrospective Analysis for Research and Applications, Version-2 (MERRA2) reanalysis, from the NASA Global Modeling Assimilation Office (GMAO) (Gelaro et al., 2017). We perform nested grid simulations over the southwestern coast of Africa (40°W-20°E, 0-40°S) with a horizontal resolution of 0.5° by 0.625° and extending over 47 vertical layers from the surface to 0.01hPa. A chemical time step of 20 minutes and transport time step of 10 minutes is applied, as recommended by Philip et al. (2016). Prior to the target year, 2017, we conduct a 6-month spin-up simulation. Boundary conditions are obtained from global simulations performed at 4° latitude × 5° longitude horizontal resolution for the same year after a 6-month initialization.

In GEOS-Chem, carbonaceous aerosol includes organic aerosols (OA) and black carbon (BC). Organic aerosol is simulated using the "simple" scheme which treats primary organic aerosol (POA) as non-volatile and includes irreversible direct yield of SOA from precursors (Pai et al., 2020). The BC simulation follows the methodologies of Park et al. (2003) and Wang et al. (2014). Sulfate (Alexander et al., 2009), nitrate (Jaeglé et al., 2018), and ammonium (Fountoukis and Nenes, 2007) thermodynamic partitioning is estimated using the ISORROPIA II thermodynamic model (Fountoukis and Nenes, 2007). Monthly anthropogenic emissions follow the Community Emissions Data System (CEDSv2) inventory (Hoesly et al., 2018). Biomass burning emissions are calculated using the Global Fire Emissions Database (GFED4.1s) at 0.25°×0.25° spatial resolution, with fractional daily and 3-hourly scaling factors applied to the cumulative monthly data (van der Werf et al., 2017). DMS emissions in the standard model use the Lana et al. (2011) climatology, which compiles DMS concentrations using data from the Global Surface Seawater DMS Database (http://saga.pmel.noaa.gov/dms/) collected from 1972 to 2009, incorporated with additional observations from the South Pacific (Lee et al., 2010). The standard DMS oxidation mechanism in the model includes only three gas-phase DMS reactions, which directly yield SO_2 and MSA according to the reaction mechanism outlined by Chin et al. (1996), and incorporates updated reaction rate coefficients from Burkholder et al. (2015). Sea-salt aerosol (SSA) emissions from the open ocean are both windspeed- (Gong et al., 2003) and sea surface temperature-dependent (Jaeglé et al., 2011). Dust emissions include natural dust (Fairlie et al., 2007) and anthropogenic dust from the AFCID inventory (Philip et al., 2017).

In this study, we carry out multiple simulations to explore the sensitivity of marine aerosols to various emission sources. To quantify the impact of marine sources on sulfate aerosols within the stratocumulus cloud layer, we perform a high-resolution (0.5° x 0.625°) marine emissions only sensitivity simulation where SO_2 and SO_4 emissions from anthropogenic sources, biomass burning, volcanic activity, ships and aviation were turned off. Additionally, to investigate the sensitivity of DMS emission fluxes to surface ocean DMS concentrations, we perform an additional simulation with DMS concentrations from Galí et al. (2018). In this dataset, DMS concentrations are estimated through a remote-sensing algorithm that integrates satellite-derived estimates of chlorophyll and light penetration, along with climatological mixed layer depth (Galí et al., 2018). Furthermore, we assess the impact of adding marine POA, co-emitted with sea-salt aerosols (Gantt et al., 2015), on the overall organic aerosol burden, which is not included in the standard model configuration. Finally, to evaluate how uncertainty in biomass burning SO_2 emissions affects the relative importance of marine emissions to sulfate aerosol, we conduct two sets of sensitivity simulations using the Quick Fire Emissions Dataset (QFED) (Darmenov & da Silva, 2013; Das et al., 2017), and the Global Fire Assimilation System (GFAS) (Kaiser et al., 2012; Su et al., 2023). Each of these inventories differ in data sources, methodology, temporal resolution and plume injection height. These sensitivity analyses were conducted for the year 2017, following a six-month spin-up period. Details regarding the spatial resolution used in each sensitivity analysis are provided in Table A1.

2.2 Ground-based measurements

We evaluate simulated aerosol optical depth (AOD) against AOD retrieved from the ground-based Aerosol Robotic Network (AERONET) of sun photometers with direct sun measurements every 15 min (Holben et al., 1998). We use Level 2.0 Version 3 data that have improved cloud screening algorithms (Giles et al., 2019). We strategically select nine sites in the study domain along coastal and oceanic regions, as shown in Fig. 1. Site information, including the coordinates, number of months with available data and the monthly average AOD for three distinct time periods, is summarized in Table A2. The AERONET monthly average AOD is computed from daily averages for sites with at least 3 months of observations during the model simulation period (year 2017) and months with at least 15 days of measurements. These are then compared with the monthly mean AOD from the GEOS-Chem model.

The modeled AOD is sampled at each AERONET site location and computed at 550 nm wavelength by vertically integrating scattering and absorption coefficients based on the properties of various aerosol components, such as size distributions, hygroscopicity, refractive indices, and densities (Latimer and Martin, 2019). For comparison with modeled monthly AOD, daily measurements at each site at 440 nm are first interpolated to the standard wavelength of 550 nm using the local Ångström exponent between 440 and 870 nm channels, following the Ångström power law (Eq. (1); Martínez-Lozano et al., 1998). These interpolated values are then averaged to calculate the observed mean monthly AOD. The interpolation formula used is:

$$AOD_{(550nm)} = AOD_{(440nm)} * \left(\frac{550}{440}\right)^{-\alpha ext\left(\frac{440}{870}\right)} \quad (1)$$

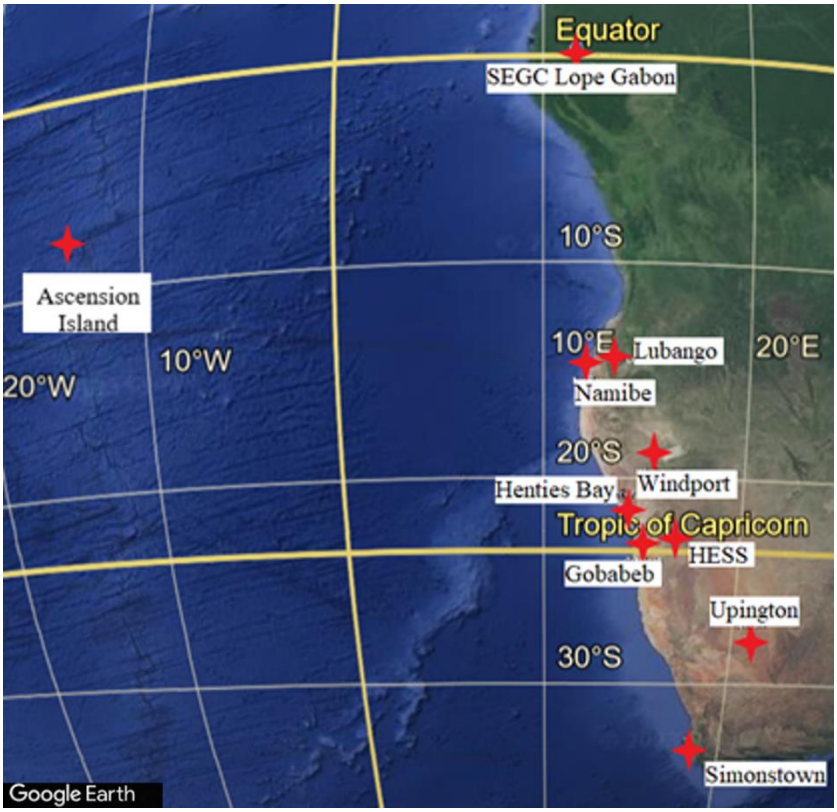

Figure 1: Map of AERONET sites used for model evaluation (© Google Earth).

In addition, we evaluate the model's relative aerosol composition against measurements from the Atmospheric Radiation Measurement (ARM) facility on Ascension Island during the LASIC campaign, conducted from January to November 2017. LASIC employed an Aerodyne aerosol chemical speciation monitor (ACSM) to provide quantitative measurement of the chemical composition of non-refractory aerosol components including sulfate, nitrate, ammonium, and organics. For comparative analysis, we use aerosol concentrations corrected for composition-dependent collection efficiency (CDCE) obtained from the ARM Data Archive. Barrett et al. (2022) reported that aerosol mass concentrations of individual components observed by the LASIC ACSM were 2 to 4.5 times lower than those measured by the aerosol mass spectrometer (AMS) aboard the CLARIFY campaign aircraft. Hence, we evaluate the relative rather than absolute aerosol speciation in GEOS-Chem against the LASIC ACSM.

2.3 Aircraft measurements

We evaluate simulated aerosol composition against airborne measurements from two campaigns, NASA ORACLES (Redemann et al., 2021; Ryoo et al., 2021) and UK CLARIFY (Haywood et al., 2021). The ORACLES field campaign used the NASA P-3 aircraft to make measurements based out of São Tomé and Príncipe while CLARIFY used the FAAM BAe-146 aircraft around Ascension Island for data collection. The ORACLES aircraft primarily conducted morning sampling, between 8:00-13:00 UTC, while the CLARIFY aircraft often sampled extended hours, typically from 7:00-18:00 UTC. Both campaigns occurred during the austral winter/spring (August-September), corresponding

with peak biomass burning events in southern Africa (Adebiyi et al., 2015). Figure 2 shows the flight tracks for these campaigns. The primary instruments and references for each campaign are listed in Table 1.

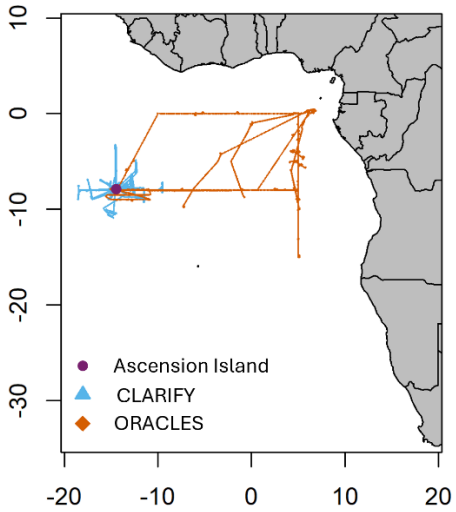

Figure 2: Flight tracks from the two aircraft campaigns used to evaluate the model, CLARIFY (in blue) and ORACLES (in orange), conducted during August-September 2017 over the southeast Atlantic region. Ascension Island is marked by the purple dot.

To facilitate comparison between airborne measurements and the GEOS-Chem model, we sampled the model to the nearest grid box, both temporally and spatially, along the flight tracks. Observations from both campaigns are reported at 1-minute averaging intervals, while the model operates at a 10-minute temporal resolution (see Sect. 2.1). Aerosol concentrations from the campaigns are reported as mass concentrations at standard temperature and pressure (STP: 273 K, 1 atm). The modeled concentrations are thus also standardized to STP conditions.

Table 1: Aircraft campaigns in the southeast Atlantic used for model evaluation during the biomass burning season

Campaign	Date range (Duration)	Instruments*	Aerodynamic Diameter (μm)	Altitude from surface (km)	Primary Reference
CLARIFY	7th August–4th September 2017 (99h)	C-ToF-AMS	0.05 to 0.60	0 to 8	Haywood et al., 2021

ORACLES	16th August–6th September 2017 (112h)	HR-ToF-AMS	0.07 and 0.70	0 to 7	Redemann et al., 2021

*Compact Time-of-Flight (C-ToF), High Resolution Time-of-Flight (HR-ToF), Aerosol Mass Spectrometer (AMS)

3.1 Model Evaluation

3.1.1 Seasonal variation of AOD

The spatial distribution of seasonal mean AOD from GEOS-Chem for the year 2017 is presented in Fig. 3. Three distinct seasonal periods reflect dominant atmospheric and oceanic processes. These include the high DMS emission period in the SEA, during the months of January, February, November, and December (JFND); the peak biomass burning season in southern Africa, spanning from July to October (JASO); and the transitional season, encompassing March, April, May, and June (MAMJ).

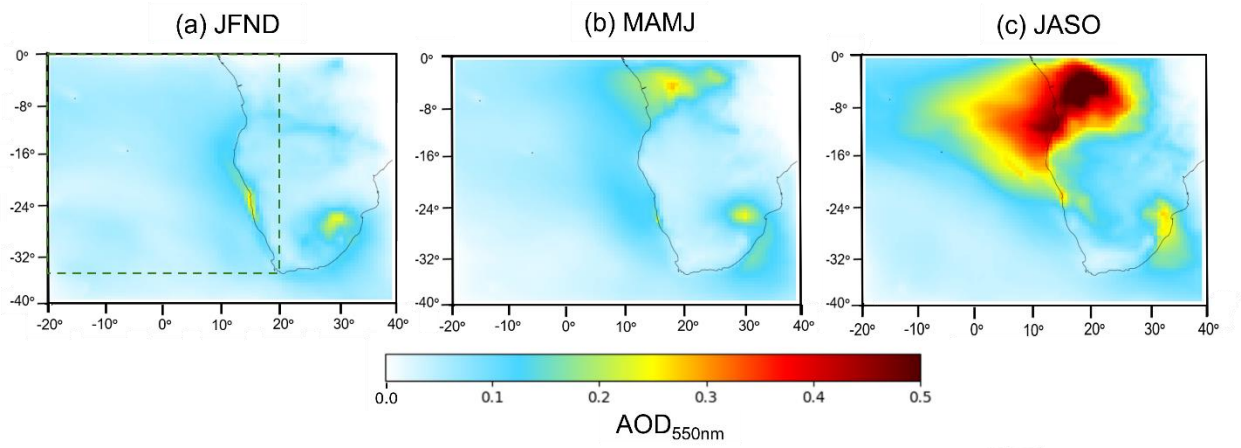

Figure 3: Spatial distribution of seasonal mean modeled AOD at 550 nm for 2017. Seasons are as follows: (a) the peak DMS emission period (JFND), (b) the transitional period (MAMJ), and (c) the peak biomass burning period (JASO). The sub-domain (0–35° S, 20° E–20° W) is highlighted with a green dotted rectangle in panel (a) for reference.

The simulated DMS emissions, based on Lana climatology (2011), indicates that emissions in the BUS region peak in January, leading to elevated concentrations of sulfate aerosols. This increased sulfate (~20%), combined with dust (59%) emissions from the Namib desert, contributes to an AOD hotspot as depicted in Fig. 3a on the southwestern coast. In the JASO period (Fig. 3c), modeled AOD increases due to biomass burning aerosols, originating from savannah fires in Central and southern Africa and transported westward towards the SEA region by the southern African easterly jet (Adebiyi and Zuidema, 2016). The spatial distribution of mean transitional period AOD (Fig. 3b) features hotspots in Congo and Angola which coincide with the onset of biomass burning in Central Africa. Additionally, a year-round AOD hotspot is observed in northeastern South Africa (Gauteng province; Fig. 3), which is associated with elevated aerosol concentrations due to industrial and mining activities, as well as domestic fuel burning (Arowosegbe et al. 2021; Zhang et al. 2021).

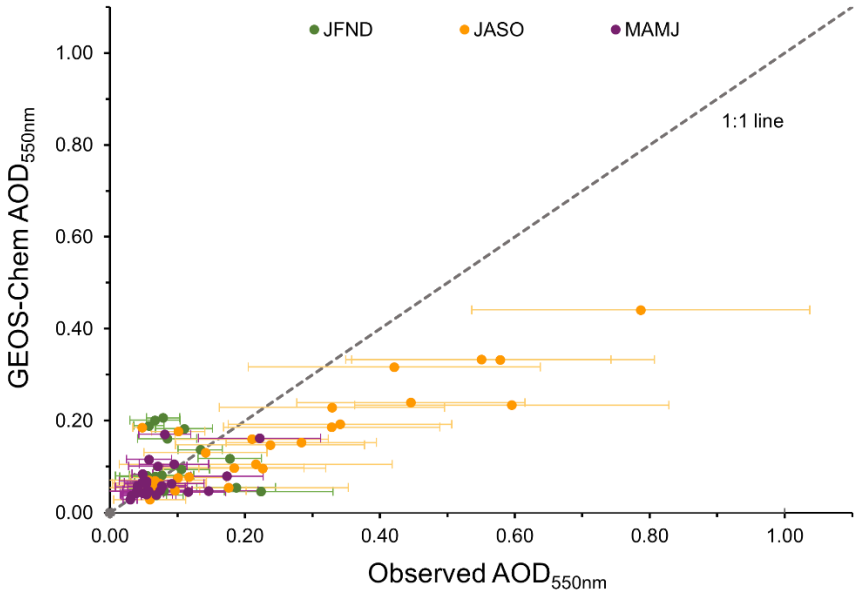

Figure 4: Modeled AOD_{550nm} (Y-axis) with respect to AERONET AOD_{550nm} (X-axis). Each data point represents the monthly mean values for each station color-coded by season (green- DMS period, yellow - biomass burning period, purple-transitional period). Error bars indicate the standard deviation of the AERONET AOD_{550nm} values, and the dotted line depicts the 1:1 relationship.

Figure 4 shows the correlation of monthly average AERONET and GEOS-Chem AOD across the nine selected sites (see Sect. 2.1 and Fig. 1), with the three seasonal periods distinguished by color: green for peak DMS emission season (JFND), yellow for biomass burning season (JASO), and purple for the transition period (MAMJ). Each data point corresponds to the monthly mean AOD values at distinct AERONET sites. The error bars in Fig. 4 represent the ±1 standard deviation in monthly AOD measurements at these sites, with higher deviations observed during the biomass burning months (up to ±0.25 at Namibe site). The comparison of monthly mean AOD across individual sites (see Table A2 in the Appendix) shows that, with the exception of Ascension Island, Gobabeb, and Upington, the mean AOD at the remaining sites during the biomass burning season (JASO) is atleast one standard deviation higher than the mean AOD in other seasons (JFND & MAMJ).

Table 2 compiles the performance of monthly mean GEOS-Chem AOD with respect to AERONET AOD by season. JASO exhibits the strongest correlation (R = 0.901), which is statistically significant ($p < 0.05$). The transitional period (MAMJ) shows a moderate correlation (R = 0.48) with a normalized mean bias (NMB) of 4.5%. Negligible negative correlation coefficient (R = -0.058) with a positive bias (29.8%) is seen during the summer period (JFND), predominantly due to anomalies at two sites. This period witnesses a considerable underestimation of AOD at Ascension Island, alongside an overestimation of dust aerosol at Gobabeb. Excluding these two sites, improves the

model's correlation to 0.67 (p = 0.55) and reduces the NMB to 4.7%. This underestimate of AOD at Ascension Island (Fig. A1 in the Appendix) during summer (JFND) suggests potential model limitations in accurately simulating natural aerosol emissions such as sea salt and marine biogenic emissions. Meanwhile, the AOD discrepancy at Ascension Island in the biomass burning season, may be due to the underestimate of transatlantic transport of light-absorbing carbon aerosols (Das et al., 2017) and deviations in its spatial distribution from typical zonal patterns over the Atlantic (Adebiyi et al., 2023). The sources of these model biases are discussed in further detail in Section 3.1.2.

Table 2: Statistical parameters of monthly mean modeled AOD with respect to observed AOD at the AERONET sites by season

Time period	Number of observations	Correlation coefficient (R)	Normalized mean bias (NMB) (%)	Root-mean square error (RMSE)
JFND	20	-0.058 (p = 0.62)	29.8	0.079
MAMJ	26	0.48 (p = 0.59)	4.5	0.043
JASO	28	0.901 (p = 0.044)	-18.6	0.15

We evaluate the relative aerosol speciation simulated at Ascension Island against monthly mean ACSM observations during the LASIC campaign (see Sect. 2.2) available for January–November 2017 (Fig. A2 in the Appendix). The seasonality of the relative contributions of organic aerosols and sulfate are consistent between the model and observations. However, the model underestimates the relative contribution of sulfate during most months, while generally overestimating the proportion of organics. An increase in the transport of biomass burning organic aerosols would further worsen the model underestimate of sulfate. A slight overestimate in the modeled relative contribution of sulfate is observed in February, and November, when simulated DMS emissions in the region are high (Lana et al., 2011), largely due to enhanced underestimations of organics and nitrates.

3.1.2 Vertical profiles of aerosol composition

Figure 5 depicts the mean vertical profiles of speciated aerosol mass concentrations observed during ORACLES and CLARIFY aircraft campaigns in August–September 2017 (the biomass burning season), compared to GEOS-Chem (see Sect. 2.2 and Table 1). The cloud top height in the SEA region generally falls between 0 to 2 km (Redemann et al., 2021). Findings from Diamond et al. (2018) indicate that aerosols below clouds in this lower atmospheric layer can also substantially impact cloud microphysics. At these altitudes (0–2 km), GEOS-Chem performs well against AMS measurements of total aerosol mass, which includes sulfate, nitrate, ammonium and organics from these campaigns, with an NMB between -3.5% (CLARIFY) to -7.5% (ORACLES). At mid-altitudes (2–4 km), the model

is biased low, with NMB values spanning -19% (ORACLES) to -57% (CLARIFY). However, the model demonstrates a pronounced bias at higher altitudes (4–7 km), where NMB values drop to -92% (ORACLES) to -93.5% (CLARIFY), underscoring challenges in accurately modeling aerosol concentrations at these elevations. These significant low biases in aerosol concentrations at higher altitude likely contribute to the model's underestimation of AOD during the biomass burning period (see Section 3.1.1). This underestimation may also be affected by the model's bulk aerosol scheme, which inadequately captures the optical properties of aerosols and is compounded by a low relative humidity bias (Zhai et al., 2021). The bulk scheme also assumes all aerosols are externally mixed, which contrasts with the variable degree of particle mixing states in the atmosphere (Yu et al., 2012; Dang et al., 2022). Moreover, studies like Hodzic et al. (2020) using NASA ATom aircraft data indicate that GEOS-Chem substantially underestimates oxidation levels of organic aerosols in remote areas, which could affect estimates of their burden and optical properties. Pai et al. (2020) further suggests that the model underestimation of organic aerosol loading at mid-tropospheric heights is linked to the surface injection treatment of fire emissions in GFED4.1s. Recent studies by Wizenberg et al. (2023) and Marvin et al. (2024) concur that fire injection scheme is a critical source of model uncertainty, emphasizing the potential importance of accurate fire injection modeling in the free troposphere. Nonetheless, our study focuses on aerosol composition within cloud-relevant altitudes to improve our understanding of aerosol-cloud interactions and their climate implications. The observed vertical distribution of aerosol mass concentrations (left panels of Fig. 5), indicates that 18% and 36% of the aerosol mass for the ORACLES and CLARIFY campaigns, respectively, is located below 2 km, within columns extending up to flight altitudes of 7 km and 8 km. However, the model simulates elevated aerosol mass at these lower altitudes, 24% and 50% of the column for ORACLES and CLARIFY, respectively.

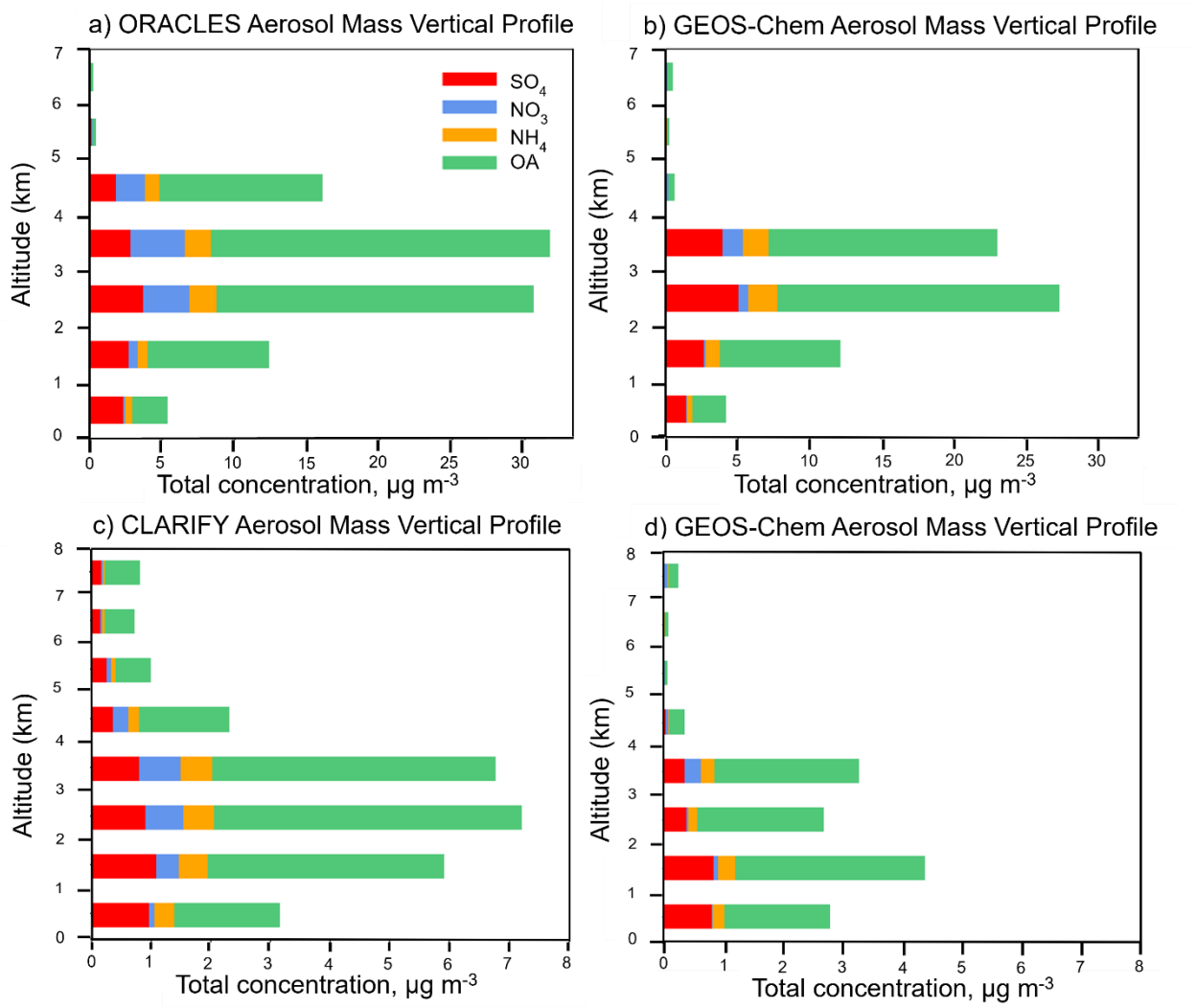

Figure 5: Average vertical profiles of simulated and observed aerosol mass during August–September 2017 (peak biomass burning season) from aircraft campaigns. The left column presents the vertical distribution of aerosols observed during the ORACLES flight campaign (panel a) and the CLARIFY flight campaign (panel c) at STP (see Sect. 2.3). The right column displays the GEOS-Chem model simulations along the respective flight tracks of each campaign (panels b and d). All data are averaged over 1 km vertical bins.

At altitudes where clouds persist in the domain (0 to 2 km), sulfate and organic aerosols are the dominant aerosol types. Here, the model effectively captures the mass concentration of organic aerosols, with an NMB ranging from -0.40% for ORACLES to -14% for CLARIFY. However, it underestimates sulfate aerosol concentrations by 19% at cloud altitudes for both campaigns. For other aerosol types and altitudes, the model consistently underestimates concentrations, except for sulfate and ammonium aerosols between 2 to 4 km during the ORACLES campaign, which the model overestimates by 40% and 4.6%, respectively. The model captures the total aerosol mass from 0 through 7 km for sulfate and ammonium aerosols during the ORACLES campaign, with only minimal underestimations of 1.5% and 0.7%, respectively. This indicates a potential discrepancy in the vertical distribution of these aerosols rather than in total mass.

3.2 Seasonal variation in aerosol composition and sources at cloud altitudes

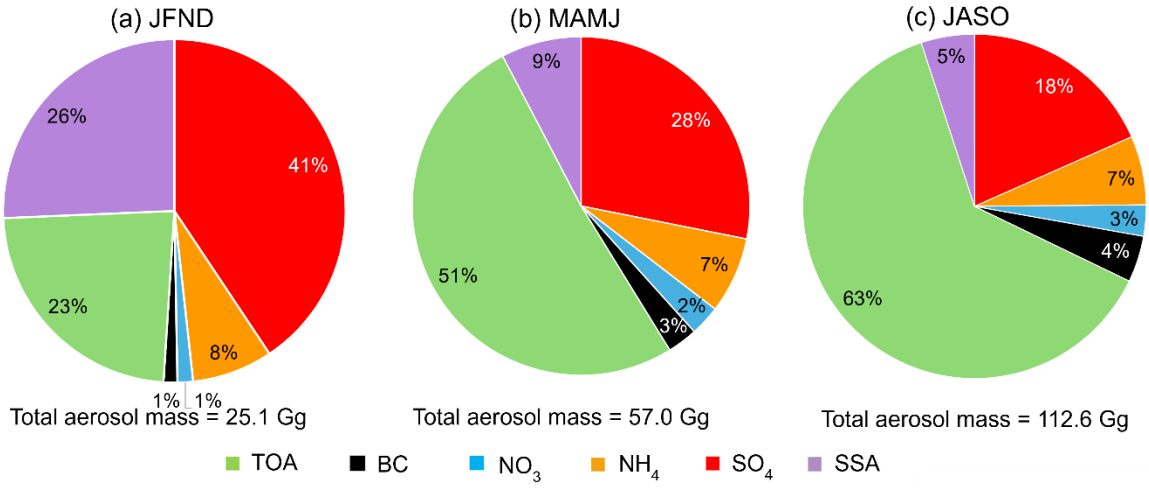

Figure 6: Simulated mean fractional aerosol composition at cloud heights (0 –2 km) over the ocean in the stratocumulus sub-domain (0–35° S, 20° E–20° W) by season: (a) JFND, (b) MAMJ, and (c) JASO. Here SO_4, NH_4, NO_3, BC, TOA, SSA represents sulfate, ammonium, nitrate, black carbon, total organic aerosol and accumulation-mode sea salt aerosols, respectively.

Figure 6 presents the simulated seasonal mean aerosol fractional composition within cloud-relevant altitudes (0-2 km), averaged over the ocean only across the subdomain (0–35° S, 20° E – 20° W) (see the map shown in Fig. 7). This area is strategically selected to coincide with the persistent Sc cloud deck and enhance our analysis of aerosol-cloud interactions. Organic aerosols, an indicator of biomass burning, predominate during both the biomass burning (JASO) and transitional (MAMJ) periods. In contrast, sulfate aerosols dominate during austral summer, likely influenced by the high primary production from coastal upwelling that leads to DMS emissions. We investigate the model representation of sulfate and these processes further in subsequent sections. An increase in the accumulation-mode sea-salt aerosols (radius 0.01–0.5 μm) contribution (total mass of 6.7 Gg) is observed in summer (Fig. 6a) as well, compared to other seasons (5.2 Gg during MAMJ and 5.8 Gg during JASO), owing to the peak wind speeds in the southern Benguela region in this season (Hutchings et al., 2009). Black carbon, ammonium, and nitrate aerosols make minor contributions to simulated aerosol mass at cloud height throughout the year.

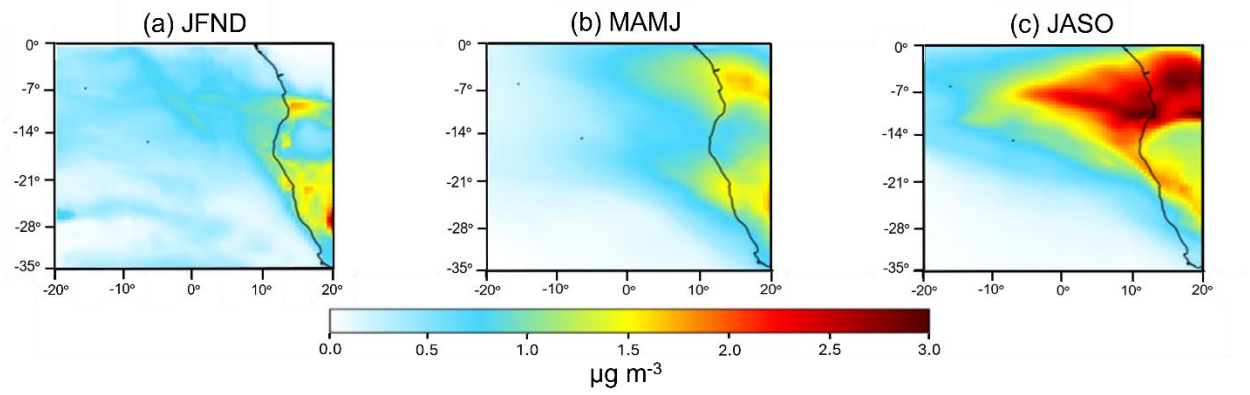

Figure 7: Spatial distribution of simulated mean sulfate aerosol concentrations averaged over cloud altitudes (0–2 km) in the sub-domain (0–35° S, 20° E–20° W) by season in 2017: (a) peak DMS emission season (JFND), (b) transitional phase (MAMJ), and (c) biomass burning season (JASO).

3.2.1 Drivers of sulfate aerosol and importance of marine precursor emissions

Sulfate aerosols are the most or 2nd most important aerosol component in cloud heights over the SEA (Fig. 6). We examine the sources of sulfur emissions within the model in Figure 8. Within the broader domain (0–40° S, 40° E–20° W), anthropogenic activities are the largest source of sulfur emissions throughout the year (Fig. 8). However, the model default CEDS inventory (Hoesly et al., 2018) fails to capture the seasonality of these emissions due to absence of regional inventories and reliance on the global datasets such as the International Energy Agency (IEA) energy statistics. The anthropogenic emissions are followed by DMS emissions from the ocean, which become more pronounced during the austral summer, peaking in January. Additionally, biomass burning contributes to SO_2 emissions seasonally, becoming the 3rd most important source of total sulfur emissions during July - September (Fig. 8). In contrast, sulfur contributions from volcanic, shipping, and aircraft emissions remain minimal and constant year-round, reflecting assumptions of static fuel burned and emission levels across inventories.

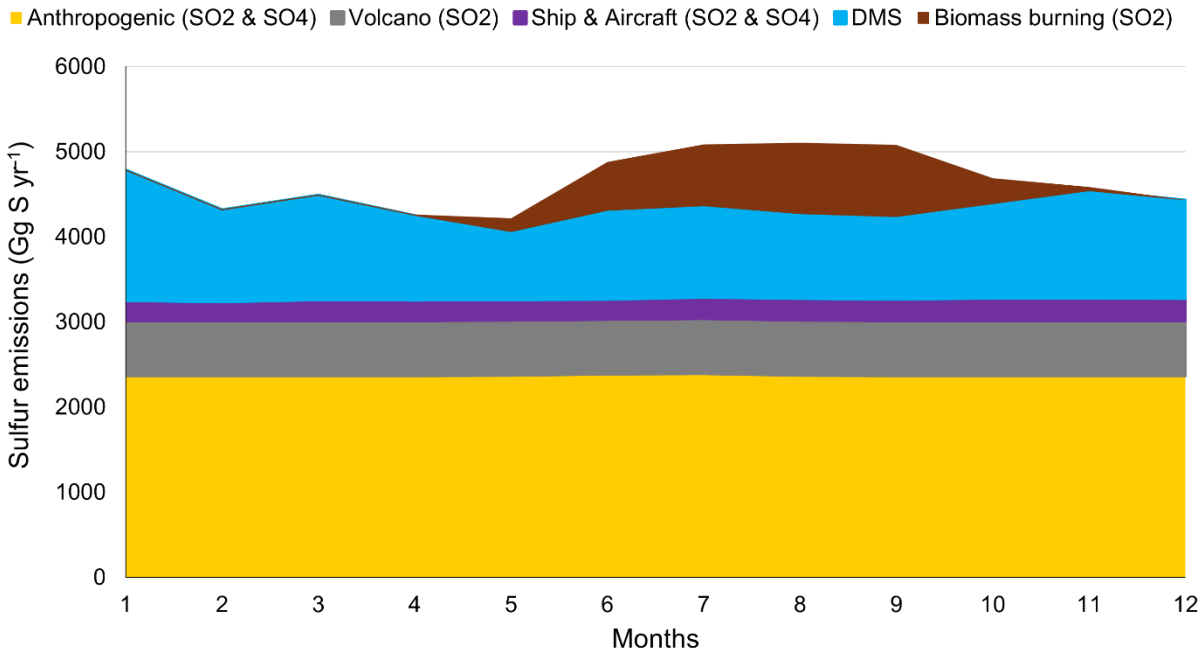

Figure 8: Stacked area chart of monthly total sulfur emissions by source for 2017 across the study domain (0–40° S, 40° E–20° W) in gigagrams of sulfur per year (Gg S yr^{-1}). Sources are indicated by color and encompass anthropogenic activities, volcanic activity, ship and aircraft emissions, biomass burning and natural emissions of dimethyl sulfide (DMS).

To improve understanding of the processes driving sulfate aerosol concentrations in the region, we examine its simulated spatial distribution averaged by season over the cloud height (0–2 km) in Fig. 7. Elevated concentrations of DMS, resulting from higher rates of primary production (Lana et al., 2011; Galí et al., 2018), lead to an increase in

sulfate concentrations along the coastline of the Benguela region and the inner shelf of Namibia during JNFD (Fig. 7a), aligning with the AOD hotspot observed in Fig. 3a. This is consistent with the simulated dominance of sulfate aerosols at cloud-relevant altitudes during JFND (Fig. 6a). During the biomass burning months (JASO), while their relative contribution decreases (Fig. 6c), sulfate aerosols display a pronounced increase in absolute concentration (Fig. 7c) as a consequence of savanna fire emissions from southwestern Africa (van der Werf et al., 2010; Das et al., 2017). As outlined in the AOD evaluation (Sect. 3.1.1), the model underestimates the transport of emissions to remote sites (Fig A1), resulting in a steep gradient in sulfate concentrations from the eastern landmass towards the western open ocean.

To quantitatively estimate the contribution of marine precursor emissions to sulfate aerosols, we compare the sulfate mass between the standard and marine emissions only sensitivity simulations (Sect. 2.1). Figure 9 shows seasonally averaged vertical profiles over the ocean region of the Sc sub-domain (0°–35° S, 20° E–20° W). The figure presents the marine-only sulfate mass and the total sulfate mass from the standard simulation (left panels), and the ratio of marine sulfate to total sulfate (right panels). Vertical profiles were computed by summing the sulfate mass within each grid box, scaled by the grid box ocean fraction, across latitude and longitude within each vertical layer of the model, and then averaged temporally across each season.

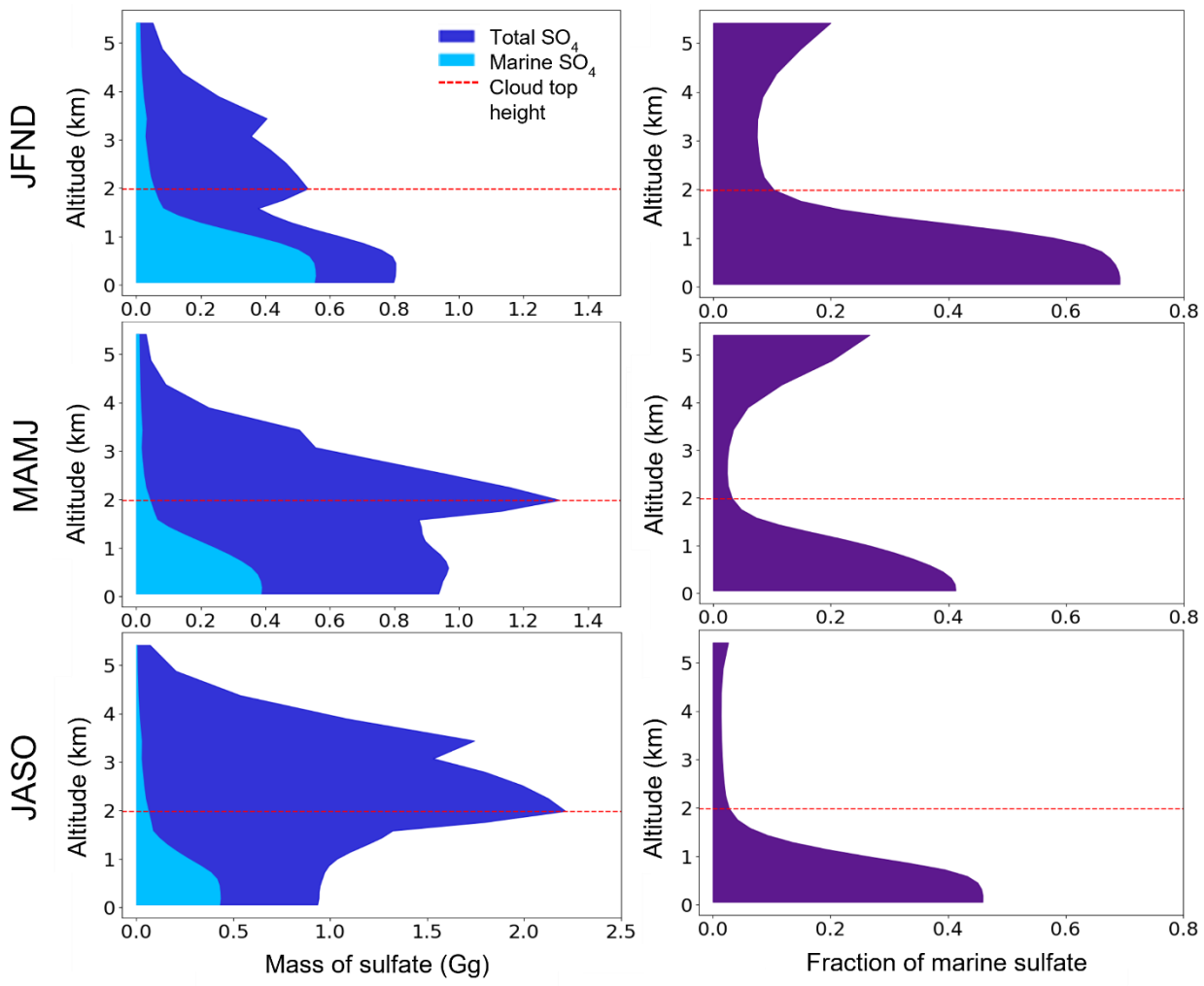

Figure 9: Simulated vertical profiles of sulfate aerosol mass over oceanic regions within the sub-domain (0°–35° S, 20° E–20° W) by season. The left panel shows the mass of total and marine sulfate aerosols, and the right panel indicates the sulfate fraction from marine sources. The top row corresponds to the peak dimethyl sulfide (DMS) emission period (JFND); the middle row to the transitional period (MAMJ); and the bottom row to the peak biomass burning period (JASO) (note: the bottom left panel displays a higher x-axis scale). The upper red dashed line denotes the typical maximum cloud top height (Redemann et al., 2021).

Our analysis highlights the substantial influence of marine sulfur sources on sulfate during JFND, as evidenced in the top left panel of Fig. 9. During this period the proportion of marine sulfate reaches up to 69.1% within-cloud (from surface to 2 km). The contribution of marine sulfate within the cloud in the subsequent periods is reduced (ranging between 2.7–45.9%; Fig. 9). We find that marine-sourced sulfate mass remains fairly consistent throughout the year (Fig. 9, left panels), with variations in the marine sulfate fraction (Fig. 9, right panels) mainly due to changes in land-based sulfate sources. Total sulfate mass during seasons influenced by biomass burning (MAMJ and JASO) peaks at 2 km, with greater mass above 2 km during peak biomass burning (JASO) in contrast to JFND where mass peaks within clouds (0–2 km).

Table 3 summarizes the monthly mean percent contribution of marine sulfate averaged across cloud altitudes (0–2 km). The annual average total sulfate mass and marine sulfate mass is 16.2 Gg and 3.5 Gg, respectively. The within-cloud marine sulfate contribution peaks in January (57.7%) and is smallest in September (10.3%). Thus, our analysis suggests that DMS emissions influence sulfate aerosol formation and their interactions with clouds in the region throughout most of the year, excepting only the peak biomass burning season. This emphasizes that constraining marine sulfur sources and chemistry both in chemical transport and climate models may improve representation of aerosol-climate dynamics in the SEA region. Limited available observations suggest the model is biased low in AOD throughout most of the year (Sect. 3.1.1), and underestimates sulfate aerosol concentrations in August and September at cloud altitudes (Sect. 3.1.2, Fig. 5). We explore related uncertainties and their implications in the following sections.

Table 3: Seasonal variation of percentage of monthly mean percent contribution of marine sulfate within cloud height

Month	Percentage of marine sulfate
January	57.7
February	54.8
March	25.3
April	26.6
May	15.3
June	15.0
July	14.8
August	14.7
September	10.3
October	22.4
November	39.1
December	44.3

3.3 Uncertainties

3.3.1 Assessing variations in DMS emission rates and oxidation mechanism on sulfate aerosol formation

The Benguela region has substantial uncertainties in DMS concentrations in surface seawater (Asher et al., 2011; Tortell et al., 2011) and the corresponding emission fluxes owing to the limited availability of biogenic sulfur measurements. To investigate the sensitivity of DMS emission fluxes to changes in surface seawater DMS concentrations, we conducted two simulations with DMS concentrations from Lana et al. (2011) and Galí et al. (2018) (see Sect. 2.1). The standard results presented thus far were conducted using the Lana dataset.

In the southern Benguela, south of approximately 27° S, marked upwelling during the austral summer (Shannon and Nelson, 1996; Hutchings et al., 2009) promotes phytoplankton growth and elevates DMS emissions. Figure 10 indicates that the Lana dataset aligns with this phenomenon, displaying peak DMS emission fluxes over the Sc sub-domain in January. However, it lacks clear seasonality for the remaining months. In contrast, satellite-based DMS estimates from Galí show pronounced emissions throughout the austral summer (JFND), as shown in Fig. 10. Both datasets concur in magnitude for January and February, a period with better data coverage in the Lana et al. (2011) climatological data set over the domain. However, the Lana dataset DMS emissions are up to 38% less during December, while 51% higher in July relative to the Galí dataset. This suggests the marine contribution to sulfate in our standard simulation using the Lana dataset may be underestimated from October through December (encompassing two months of the peak DMS season) and overestimated from March through August (Fig. 10).

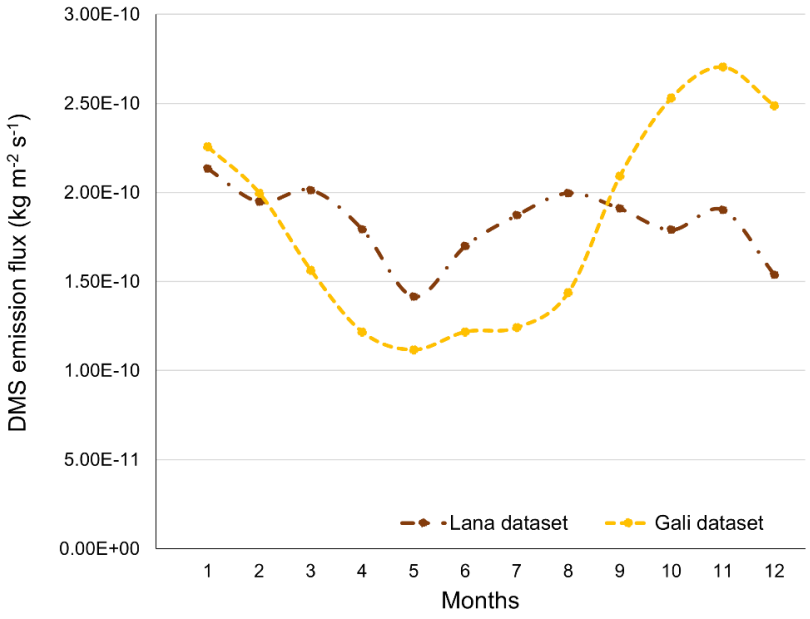

Figure 10: Monthly DMS emissions over the stratocumulus sub-domain (0–35° S, 20° E–20° W) using two distinct datasets for surface seawater DMS concentrations. The brown dashed line presents emissions calculated using Lana

et al. (2011) climatology, which compiles data across 1972-2009 from multiple sources. In contrast, the yellow dashed line depicts emissions based on satellite-derived estimates of surface seawater DMS concentrations (Galí et al., 2018).

The ongoing discovery of complexities within DMS oxidation mechanisms, along with the incomplete incorporation of these mechanisms into atmospheric chemistry models, further contributes to uncertainties in predicting the impact of DMS emissions on aerosols and climate (Faloona, 2009; Quinn and Bates, 2011; Carslaw et al., 2013). Chen et al. (2018) highlighted the impacts of changes to DMS chemistry in the GEOS-Chem model, integrating a series of multiphase sulfur oxidation mechanisms and two DMS intermediates, which led to a decrease in the global DMS burden, thereby decreasing SO_2 and sulfate levels. On the other hand, Novak et al. (2021) found that the cloud uptake of hydroperoxymethyl thioformate (HPMTF), a newly identified oxidation product of DMS (Wu et al., 2015; Veres et al., 2020), lowers near-surface SO_2 concentration while elevating sulfate concentration in the model. Most recently, Tashmim et al. (2024) implemented an advanced DMS oxidation mechanism in GEOS-Chem that incorporates the latest developments in DMS chemistry, including those previously mentioned, which led to a lower SO_2 mixing ratio (~70%) and a higher SO_4 mixing ratio (~35%) over the SEA during austral summer. Thus, an improved representation of DMS emissions and oxidation chemistry in the model could enhance the sulfate aerosol estimations during the peak DMS season. This refinement may address model underestimates of aerosol concentrations during this period (Sect. 3.1.1).

3.3.2 Exploring the impact of marine organic aerosol emissions on organic aerosol concentrations

Beyond marine sulfate and sea-salt aerosols, organic matter also makes a significant contribution to marine aerosol mass (Middlebrook et al., 1998; Oppo et al., 1999; Russell et al., 2010). Notably, substantial concentrations of organic carbon aerosols have been observed in marine regions, particularly during periods of intense biological activity (O'Dowd et al., 2004). These aerosols can also increase CCN, affecting cloud properties and radiative balance (Arnold et al., 2009; Gantt and Meskhidze, 2013). However, the standard GEOS-Chem model does not account for these marine organic aerosol emissions. Here, we analyzed the impact of marine POA on cloud-altitude aerosols over the SEA by incorporating marine POA emissions based on satellite-derived chlorophyll-a concentrations (Gantt et al., 2015; See Sect. 2.1) in the model.

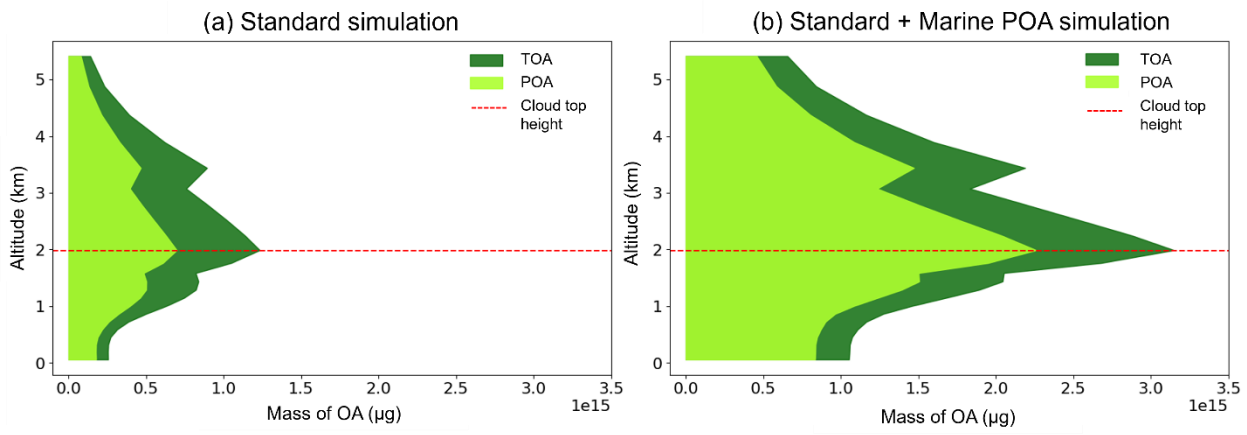

Figure 11: Vertical distribution of organic aerosol mass during November 2017, the month of maximum discrepancy between the standard and MPOA simulations, over the Sc sub-domain (0–35° S, 20° E–20° W). Left: mass profile for total organic aerosols (TOA) and primary organic aerosols (POA) under standard simulation conditions (Std); right: when marine primary organic aerosol (MPOA) emissions are included (Std + MPOA). The red dashed line indicates the typical maximum cloud top height.

We find that the inclusion of MPOA emissions consistently resulted in higher organic aerosol mass, with the greatest increase in November. Figure 11 shows the vertical distribution of total organic aerosols (TOA) mass and POA mass (including MPOA and other POA sources) with and without MPOA emissions during this month. Similar to our earlier vertical profile analysis (refer to Sect. 3.2.2), we find that the maximum organic aerosol mass occurred at the highest cloud top height (2 km). The Standard + MPOA simulated peak total organic aerosol mass was approximately three times higher than that in the Standard simulation, highlighting the potential contribution of marine sources to total organic aerosol mass concentrations. However, during the biomass burning season, the sensitivity simulation showed only a minimal increase, indicating that it does not adequately address the model's underestimation (refer to Fig. 5). Gantt et al. (2015) demonstrated that including MPOA emissions in GEOS-Chem reduced the normalized mean bias (NMB) of surface organic aerosol concentrations at coastal sites by 67%. Additionally, Pai et al. (2020) noted that without a marine POA source, the model fails to accurately reproduce lower-tropospheric concentrations over oceans, although the marine POA scheme might be biased high. Despite the limitations of a chlorophyll-based parameterization like the one used here in providing mechanistic understanding of the seasonal and geographical variability of organic matter emissions from sea spray (Burrows et al., 2022), our findings suggest that MPOA may play a role in aerosol-cloud interactions outside of the biomass burning season, in addition to marine-derived sulfate from DMS (Sect. 3.2).

3.3.3 Impacts of uncertainties in biomass burning emissions of SO_2

To assess the impact of uncertainty in biomass burning emissions of SO_2 on the relative contribution of marine vs. land sources to aerosol, we performed a sensitivity analysis using two alternative inventories, QFED and GFAS (see Sect. 2.1 and Fig. A3). The standard simulations, as detailed in Sect. 2.3, use the default biomass burning inventory in GEOS-Chem, GFED. The GFAS inventory SO_2 and CO emissions over the domain are constant in time, aligning with QFED during the non-biomass burning months (Fig. A3). We find that CO emissions from GFED and QFED align closely; however, there is a notable difference in SO_2 emissions between the two inventories (Fig. A3). These discrepancies likely originate from variations in SO_2 emission factors employed by each inventory. In July, which exhibits the largest difference between the two inventories, peak SO_2 emissions in QFED are almost five-fold higher than those in GFED. This discrepancy leads to a 25% increase in sulfate aerosol concentrations at cloud altitudes relative to the standard results using GFED (not shown). Consequently, the contribution of marine sulfate to total sulfate (see Sect. 3.2.1) may further decrease during the peak biomass burning season if QFED is used, highlighting the sensitivity of aerosol source attributions to the selected biomass burning inventory.

4 Implications

In this study, monthly marine sulfate constitutes between 10.3% and 57.7% of total sulfate within the cloud height, peaking during the high DMS emission period. However, the default Lana et al. (2011) climatology largely underestimates DMS emissions during the austral summer (November and December) by up to 38%, compared to the satellite-derived estimates from Galí et al. (2018). Moreover, improvement of DMS chemistry in the model by incorporating new oxidation mechanisms and intermediate products could shift the balance towards increased sulfate aerosol production (with Tashmim et al., 2024 suggesting an increase of up to 35% over the SEA). Marine primary organic aerosol emissions may also contribute substantially to the organic aerosol mass during the peak primary production period (JNFD), highlighting the importance of marine contributions to overall aerosol concentrations. Meanwhile, discrepancies in SO_2 emissions from biomass burning can increase sulfate aerosol from biomass burning by up to 25%. These changes would improve the model underestimate of AOD relative to AERONET observations; however, observations of aerosol composition outside of August-September are very limited and this is a large gap. Our results suggest marine-sourced sulfate and organics significantly influence aerosol loading and composition in the SEA, particularly during the non-biomass burning period. Accurately characterizing the seasonal dynamics of aerosols within cloud heights is imperative for quantifying aerosol-cloud interactions and understanding the dynamics of marine aerosols in the SEA region, where uncertainties in aerosol radiative forcing are most pronounced. This understanding is essential for improving the reliability of climate models in areas critical to both regional and global climate dynamics.

5 Conclusion

Aerosols over the southeast Atlantic strongly influence global climate dynamics due to the presence of persistent stratocumulus clouds and large uncertainties in aerosol-cloud interactions. However, precisely representing these interactions in global climate models remains challenging, in part due to sparse available observations, especially outside of the biomass burning season. In this study, we employed the GEOS-Chem chemical transport model to assess the aerosol composition at cloud-relevant altitudes (0–2 km) and identify the sensitivities to marine emissions and chemistry in the southeast Atlantic. This analysis aims to enhance our understanding of the role of marine aerosols and the associated uncertainties affecting aerosol-cloud interactions within this climate-sensitive region.

We performed nested grid simulations with a 0.5° x 0.625° horizontal resolution and evaluated the model against ground-based and aircraft campaign observations throughout 2017. We analyzed results for three seasonal periods with distinct dominant processes including (a) the high DMS emission season (JFND), (b) the peak biomass burning season (JASO), and (c) the transitional season (MAMJ). Our analysis showed that simulated monthly average aerosol optical depth (AOD) exhibits the strongest correlation (R = 0.901) with the AERONET AOD observations during the JASO season. However, the model generally underestimates AOD throughout the year, except in the JFND period, where an overestimate at Gobabeb site offset underestimations at other sites. Moreover, a comparison of aerosol speciation measured at Ascension Island during the LASIC campaign indicates that the model consistently underestimates sulfate aerosols. We further evaluated the simulated vertical profile of aerosol mass concentrations and composition against measurements from the ORACLES and CLARIFY campaigns. These comparisons showed that sulfate aerosols were underestimated by 19% at cloud-relevant altitudes of 0–2 km by both campaigns. The

underestimate of sulfate aerosols at lower altitudes (0-2 km), coupled with an underestimate of other aerosols at higher altitudes (4-8 km), likely contributes to the overall low bias in modeled AOD. The misrepresentation of natural aerosol emissions and transatlantic aerosol transport may be responsible for these underestimates. Nevertheless, discrepancies increase with altitude, reflecting challenges in accurately modeling high-altitude aerosol concentrations.

Analysis of seasonal mean aerosol composition at cloud height showed that organic aerosols predominate during JASO (63%) and MAMJ (51%), while sulfate aerosols are most prevalent (41%) during the austral summer (JFND). Given the prominence of sulfate as a marine sourced aerosol in remote oceanic environments, we investigated the processes influencing the sulfate aerosol concentrations in our domain. Throughout the year, anthropogenic sources and oceanic DMS emissions are the primary atmospheric sulfur contributors. Spatial mapping across the sub-domain (0–35° S, 20° E–20° W) showed high sulfate concentrations (up to 3μg m^{-3}) at cloud height during the peak biomass burning season (JASO), primarily from savannah fires in southern Africa. Despite this, sulfate aerosols only account for 18% of the total aerosol mass in JASO.

Sulfate, primarily from marine sources, is the dominant aerosol at cloud-relevant altitudes during JFND in the model (up to 69% marine contribution); however, significant uncertainties regarding the treatment of DMS persist that may affect this finding. To assess the impact of these uncertainties on sulfate aerosols, we compared DMS emission fluxes from Lana et al. (2011) climatological data, which has limited spatial and temporal coverage in our domain, with those from Galí et al. (2018), which are based on satellite-based estimates of surface seawater DMS concentrations. We find that, within our domain, the Lana dataset emissions estimates are 51% higher in July and 38% lower in December compared to the Galí dataset. Moreover, improvement of DMS chemistry in the model by incorporating new oxidation mechanisms and intermediate products could shift the balance towards increased sulfate aerosol production (with Tashmim et al., 2024 suggesting an increase of up to 35% over the SEA). Additionally, emissions of marine primary organic aerosols during the peak primary production period (JNFD) may substantially contribute to the mass of organic aerosols which can also act as CCN. This emphasizes the critical role of marine sources in influencing aerosol concentrations, even in oceanic regions impacted by large seasonal biomass burning. Variations in SO_2 emissions from biomass burning could potentially increase sulfate aerosol concentrations at cloud altitudes by up to 25%. Addressing these discrepancies is essential for improving the model's underestimation of AOD and aerosol concentrations compared to observations.

This study highlights the importance of constraining marine emissions and their chemical transformations by incorporating satellite-retrieved datasets and extending field campaign efforts during non-biomass burning periods. Such initiatives are essential to accurately characterize seasonal aerosol dynamics at cloud heights and to improve our understanding of aerosol-cloud interactions in regions with persistent low-altitude clouds. These advancements could substantially minimize uncertainties in model estimates of radiative forcing and enhance the reliability of climate model projections in the southeast Atlantic region.

Appendix A

Table A1: Configuration of sensitivity analysis simulations

Simulations	Resolution	Emission inventories (References)
Marine sulfur emissions only	0.5° x 0.625°	Standard model inventories (see Section 2.1 in the main text)
DMS emissions	4° x 5°	Climatological product for seawater DMS (Lana et al., 2011); Satellite-derived DMS estimates (Galí et al., 2018)
Biomass burning inventories	4° x 5°	Global Fire Emissions Database (van der Werf et al., 2017); Quick Fire Emissions Dataset (Darmenov & da Silva, 2013); Global Fire Assimilation System (Kaiser et al., 2012).
Marine primary organics	0.5° x 0.625°	Marine primary organic aerosol emission estimates from satellite-derived chlorophyll-a concentrations (Gantt et al., 2015)

Table A2: AERONET site information and the monthly average AOD_{550} value (±1 standard deviation) for three distinct time periods per site are shown. An asterisk (first column) indicates that the mean AOD during the biomass burning season (JASO) at these sites is more than 1 standard deviation higher than the mean AOD in other season(s) (JFND and MAMJ).

Site	Latitude (°)	Longitude (°)	Months of data availability for 2017	Monthly Average AOD_{550nm} ± 1 SD (JFND)	Monthly Average AOD_{550nm} ± 1 SD (MAMJ)	Monthly Average AOD_{550nm} ± 1 SD (JASO)
Ascension Island	-7.976	-14.415	7	0.320 ± 0.061	0.256 ± 0.046	0.396 ± 0.078
Gobabeb	-23.562	15.041	12	0.268 ± 0.033	0.254 ± 0.032	0.399 ± 0.115
HESS*	-23.273	16.503	10	0.162 ± 0.024	0.211 ± 0.033	0.368 ± 0.108
Henties_Bay*	-22.095	14.26	3	0.166 ± 0.021	0.138 ± 0.021	0.380 ± 0.114
Lubango*	-14.958	13.445	9	0.120 ± 0.011	0.251 ± 0.011	0.534 ± 0.139
Namibe*	-15.159	12.178	8	0.279 ± 0.028	0.314 ± 0.029	0.689 ± 0.192
Simonstown_IMT*	-34.193	18.446	7	0.246 ± 0.049	0.179 ± 0.018	N/A
Upington	-28.379	21.156	8	0.131 ± 0.019	0.193 ± 0.035	0.321 ± 0.115
Windport*	-19.366	15.483	10	0.249 ± 0.041	0.242 ± 0.037	0.514 ± 0.159

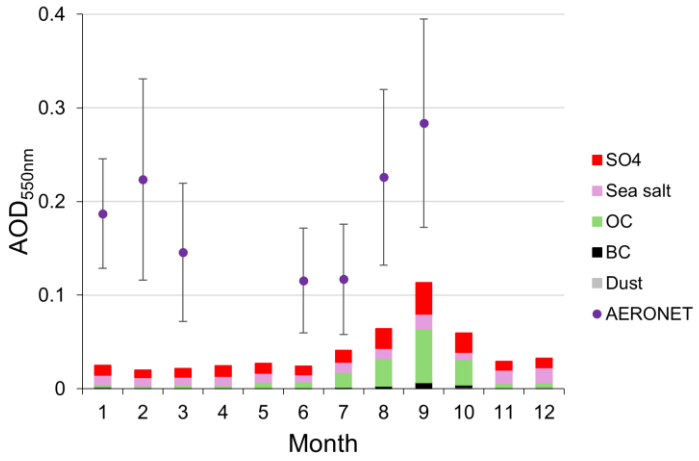

Figure A1: Comparative analysis of aerosol optical depth at 550 nm (AOD_{550nm}) for Ascension Island in 2017. The purple dots present the measured mean monthly AOD values, with vertical error bars illustrating the range of AOD_{550nm} measurements captured by the AERONET ground station. The stacked bars represent the GEOS-Chem model's

simulated AOD values, with each layer corresponding to the major aerosol components, such as sulfate (SO_4), sea salt, organic carbon (OC), black carbon (BC), and dust, providing insight into the model's aerosol composition representation.

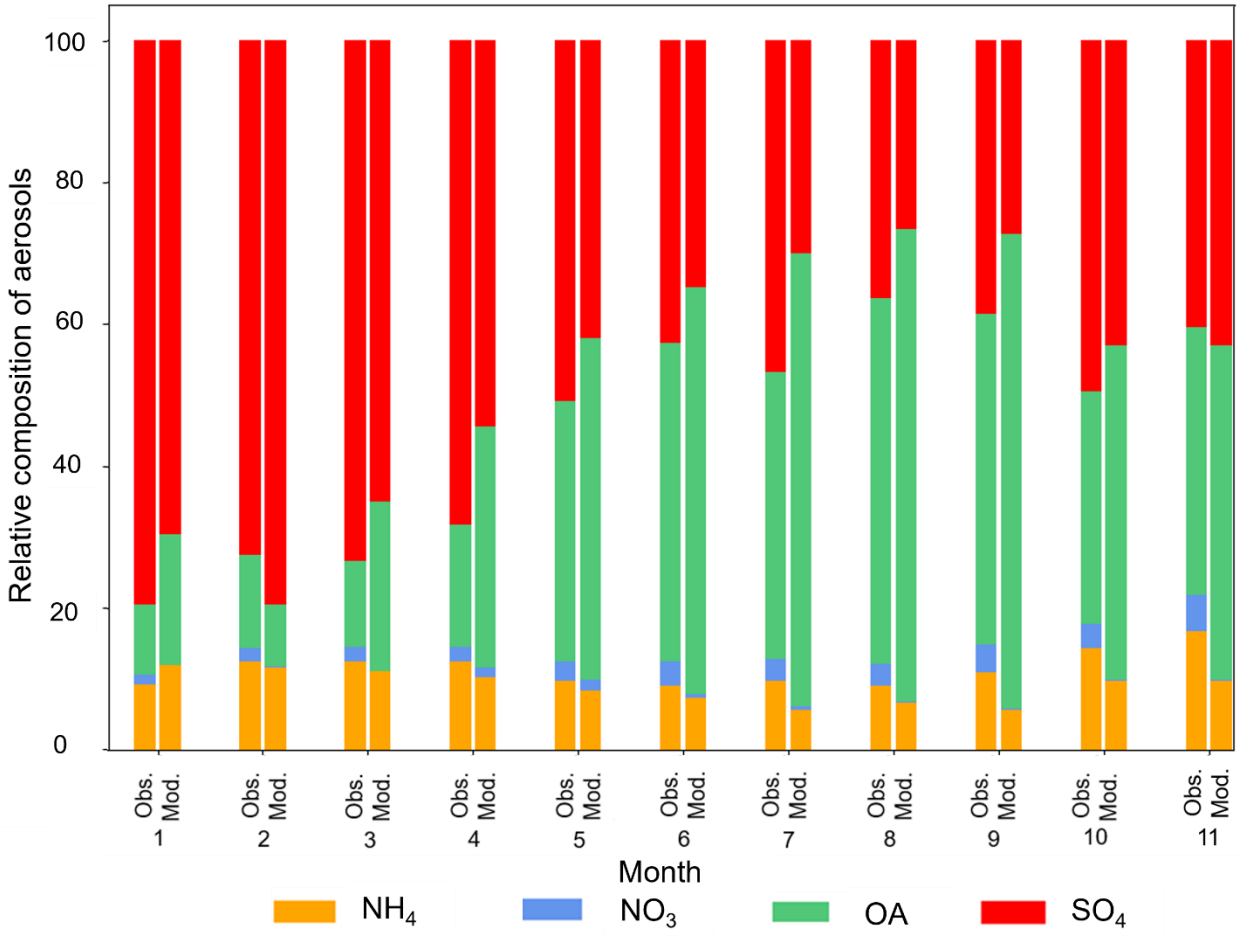

Figure A2: Comparative analysis of the relative aerosol composition at Ascension Island in 2017. The stacked bars on the left depict observations of chemical composition taken during the LASIC campaign at the ARM facility on Ascension Island, utilizing an aerosol chemical speciation monitor (ACSM) at 341 meters. The bars on the right illustrate the GEOS-Chem simulated aerosol composition at the same coordinates. Each segment of the stack represents different aerosol components: ammonium (NH_4), nitrates (NO_3), organic aerosols (OA), and sulfate (SO_4).

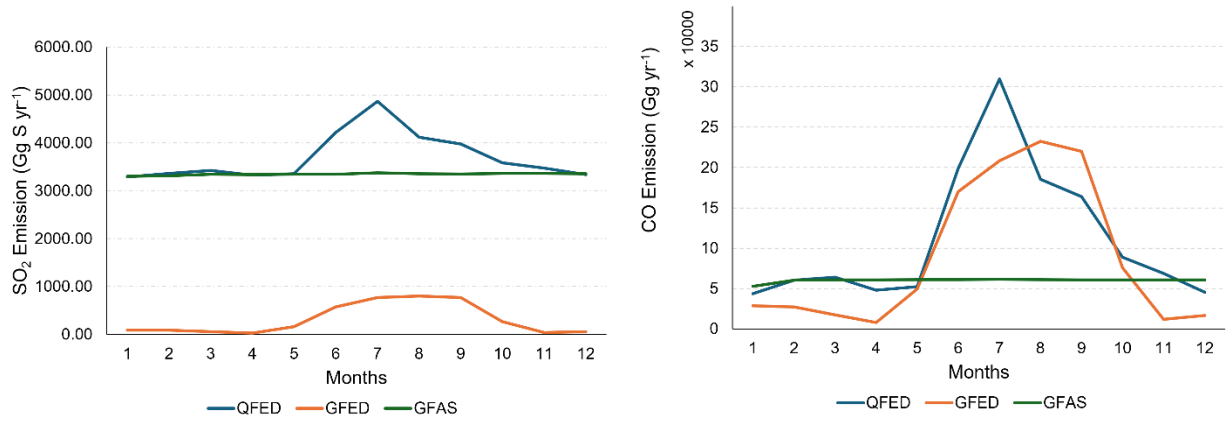

Figure A3: Comparison of biomass burning emissions across various inventories, namely GFED, QFED, and GFAS across the domain (0–40° S, 40° E–20° W). The panels depict the interannual variability of biomass burning emissions, with the left panel illustrating sulfur dioxide (SO_2) emissions, and the right panel displaying carbon monoxide (CO) emissions. Both GFED and QFED indicate similar emission trends; however, GFED exhibits lower SO_2 emission magnitudes compared to QFED. GFAS presents emission magnitudes similar to QFED during non-biomass burning period.

Code and data availability

The GEOS-Chem model used here is publicly available at https://zenodo.org/records/5748260. The ORACLES campaign data from 2017 are available at: https://doi.org/10.5067/Suborbital/ORACLES/P3/2017_V2 (last access January 26, 2024). The CLARIFY campaign data are available at https://catalogue.ceda.ac.uk/uuid/38ab7089781a4560b067dd6c20af3769 (last access February 12, 2024). The ACSM dataset from LASIC campaign is available at https://doi.org/10.5439/1763029 (last access September 10, 2024).

Author contributions

MH, HMH, and RMG designed the research. MH conducted the model simulations, analysis, and visualization, with expert advice from HMH. MH drafted the manuscript, which was then revised by all co-authors.

Competing interests

At least one of the (co-)authors is a member of the editorial board of Atmospheric Chemistry and Physics.

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