

1 **Brown carbon aerosol in rural Germany: sources, chemistry, and** 2 **diurnal variations**

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11
12 **Abstract.** Brown carbon aerosol (BrC) is one major contributor to atmospheric air pollution in Europe, especially in
13 winter. Therefore, we studied the chemical composition, diurnal variation, and sources of BrC from 17th February to
14 16th March at a rural location in southwest Germany. In total, 178 potential BrC molecules (including 7 nitro aromatic
15 compounds, NACs) were identified in the particle phase comprising on average 83 ± 44 ng m⁻³, and 31 potential BrC
16 (including 4 NACs) molecules were identified in the gas phase contributing on average 8.5 ± 6.7 ng m⁻³ during the
17 whole campaign. The 178 potential BrC molecules only accounted for $2.6 \pm 1.5\%$ of the total organic mass, but can
18 explain $14 \pm 13\%$ of the total BrC absorption at 370 nm, assuming an average mass absorption coefficient at 370 nm
19 (MAC_{370}) of 9.5 m² g⁻¹. A few BrC molecules dominated the total BrC absorption. In addition, diurnal variations show
20 that gas phase BrC was higher at daytime and lower at night. It was mainly controlled by secondary formation (e.g.
21 photooxidation) and particle-to-gas partitioning. Correspondingly, the particle phase BrC was lower at daytime and
22 higher at nighttime. Secondary formation dominates the particle-phase BrC with $61 \pm 21\%$, while $39 \pm 21\%$ originated
23 from biomass burning. Furthermore, the particle-phase BrC showed decreasing light absorption due to photochemical
24 aging. This study extends the current understanding of real-time behaviors of brown carbon aerosol in the gas and
25 particle phase at a location characteristic for the central Europe.

26 **1. Introduction**

27 The Brown Carbon (BrC) aerosol has significant impact on air quality and climate, since it absorbs the solar radiation
28 in the near-ultraviolet and visible region (Laskin et al., 2015; Moise et al., 2015). Global simulation showed that the
29 mean radiative forcing of BrC aerosol was -0.43 W m^{-2} and 0.05 W m^{-2} at the surface and at the top of the atmosphere,
30 accounting for 15% of total radiative forcing by the absorbing aerosol (Park et al., 2010). In addition, global
31 measurements of BrC found that the average direct radiative effect of BrC absorption accounted between 7% to 48%
32 at the top of the atmosphere (Zeng et al., 2020).

33 Some typical molecules of BrC have been identified, such as nitro-aromatic compounds (NACs), imidazoles, and
34 polycyclic aromatic hydrocarbons (PAH), etc., (Jiang et al., 2022; Wu et al., 2018; Huang et al., 2018; Liu et al., 2023).

35 In western Europe, the concentration levels of NACs range between $1\text{--}20 \text{ ng m}^{-3}$, accounting for 0.3%–4% of total
36 absorption of BrC at UV wavelengths (Jiang et al., 2022; Mohr et al., 2013; Teich et al., 2017). In addition, imidazoles
37 were detected with concentrations ranging between $0.2\text{--}14 \text{ ng m}^{-3}$ in ambient aerosol samples from different
38 environments in Europe and China (Teich et al., 2016). Furthermore, parent-PAHs and carbonyl-OPAHs accounted
39 for on average $\sim 1.7\%$ of the overall absorption of methanol-soluble BrC in Urban Xi'an, Northwest China (Huang et
40 al., 2018). Even though many studies have investigated the chemical composition of brown carbon and calculated the
41 absorption contribution from BrC molecules, there are still many unknown brown carbon molecules to allow a
42 quantitative assessment of their sources and atmospheric impact.

43 Sources of BrC can be separated as primary emissions and secondary formation. The primary sources of BrC are
44 biomass burning and fossil fuel combustion (Andreae and Gelencser, 2006). On a global scale, a majority of BrC
45 aerosol mass is associated with biomass burning dominating BrC absorption (Zeng et al., 2020). The major secondary
46 sources of brown carbon are from oxidation of aromatic volatile organic compounds, such as toluene (Lin et al., 2015),
47 naphthalene (Siemens et al., 2022), ethylbenzene (Yang et al., 2022), and indole (Montoya-Aguilera et al., 2017; Jiang
48 et al., 2023), especially in the presence of NO_2 .

49 BrC in the atmosphere can be suspended in the gas phase or particle phase. However, only a few studies have
50 investigated the sources and chemical composition of BrC in the gas phase. For example, NACs in the gas phase were
51 highest during the daytime at a rural site in China (Salvador et al., 2021). The major sources of NACs were from
52 secondary formation on days without extensive biomass burning emissions, but mainly from primary emissions in
53 biomass burning events (Salvador et al., 2021). The source of nitrophenol, a typical BrC molecule, was mainly from
54 secondary formation outweighing losses by photolysis in polluted urban environments, Beijing (Cheng et al., 2021).

55 The major chromophores of BrC in the gas phase were rich in phenol- and protein-like substances in Xi'an, China,
56 during the summer (Chen et al., 2021). Therefore, the previous studies mainly focus on sources and chromophores of
57 BrC, especially NACs. However, the real-time diurnal variation and sources of BrC in the gas phase in the atmosphere
58 have rarely been investigated in central Europe.

59 Previous field studies have investigated the sources of BrC in the particle phase which are mainly from secondary
60 formation and primary emissions (Wang et al., 2019a; Moschos et al., 2018; Satish et al., 2017). In the central Europe,
61 the secondary biogenic organic aerosol (OA) contributes less BrC in summer. However, the primary and secondary
62 wood burning emissions dominated the BrC (Moschos et al., 2018). The primary emissions of BrC contributed more
63 to organic aerosol light absorption than those from secondary processes in the North China Plain, China (Wang et al.,
64 2019a). However, secondary sources for BrC were more important for absorption than primary ones in the Southeastern
65 Margin of Tibetan Plateau (Wang et al., 2019b). Loss pathways of BrC in the particle phase mainly comprise
66 photooxidation and photobleaching, but also dilution of BrC e.g. by rising boundary layer height influences its
67 concentration levels (Satish et al., 2017; Laskin et al., 2015; Moise et al., 2015). The absorption of BrC was high in
68 the early morning and later decreased due to the bleaching of chromophores (Wang et al., 2019a; Satish et al., 2017).
69 A diurnal cycle showed that secondary chromophores can be formed from photochemical oxidation after sunrise
70 followed by photobleaching of the chromophores under the oxidizing conditions as the day progressed (Wang et al.,
71 2019b). Lower BrC concentrations during noon were explained by the fact that planetary boundary layer heights were
72 highest during the middle of the day (Liu et al., 2023). However, also nighttime aqueous-phase chemistry can promote
73 the formation of secondary light absorbing compounds and the production of strongly absorbing particles (Wang et al.,
74 2019a). In addition, higher emissions of biomass burning BrC were observed at nighttime. Actually, the BrC in the
75 particle phase undergoes complex photochemical processing during the whole day. The time dependent sources and
76 diurnal variations of BrC in aerosol particles are still reported rarely and not well understood.

77 To better understand the chemical characterization, diurnal variation, and sources of BrC in central Europe, we
78 performed online measurements of BrC during February-March 2021 at a rural location in southwest Germany. In the
79 following, we will describe the experimental methods used in this study. Subsequently, the mass concentrations of BrC
80 in gas and particle phase will be determined. Furthermore, the contribution of BrC to light absorption in the particle
81 phase will be estimated. Then, the diurnal variations and sources of BrC in the gas and particle phase will be analyzed.
82 Finally, the atmospheric implications of our findings will be discussed.

83 **2. Experimental methods**

84 **2.1. Measurement site**

85 We performed particle and trace gas measurements from February 17th–March 16th 2021 at KIT Campus Nord, a rather
86 rural area in Germany (49°05'43.1"N 8°25'45.6"E). The sampling site is located at the building number 322 of the
87 IMK-AAF on KIT Campus Nord, as shown in Figure S1. The campus is mostly surrounded by the Hardwald forest
88 dominated by pine trees. The sampling site is also near some villages e.g. 3–4 km east of the village “Eggenstein-
89 Leopoldshafen”, 6–7 km northeast of the village “Neureut”, 3–4 km west of the village “Friedrichstal”, 4–5 km
90 northwest of the village “Stutensee”, and 5–6 km southeast of the village “Linkenheim”. Therefore, influences by
91 biomass burning emissions from wood stove combustion in these residential areas during winter time can be expected
92 (Thieringer et al., 2022). Furthermore, the city of Karlsruhe with 3000000 inhabitants is 10 km south of the
93 measurement site. The city includes industrial areas with a coal-fired power plant “Rheinhafen” and a refinery “MIRO”.
94 Therefore, the measurement site is potentially affected by different aerosol sources.

95 **2.2. Meteorological, aerosol particle, and traces gas instruments**

96 All instruments were set up in a temperature-controlled measurement building. The samples were collected above the
97 roof top about 8 m above ground level via stainless steel tubes and a PM_{2.5} and a TSP inlet as well as FEP tubes for the
98 VOC measurements. An overview of the instruments used and the parameters measured is given in Table S1 of the
99 Supplement.

100 Temperature, relative humidity (RH), pressure, wind speed, wind direction, precipitation, and global radiation were
101 measured by a meteorological sensor (WS700, Lufft GmbH; see Table S1) about 8 m above the ground level. The main
102 wind directions during the campaign were southwest, northeast, and southeast, since winds were channeled by the
103 Rhine River valley. O₃ and NO₂ were measured with standard gas monitors (Table S1). The particle number
104 concentrations (>2.5 nm) were measured by a water-based condensation particle counter (CPC3789, TSI Inc.). PM_{2.5}
105 was measured by an optical particle counter (OPC-FIDAS 200, Palas Inc.). The particle number size distributions were
106 measured by a nanoparticle sizer (NanoScan, TSI Inc.) ranging from 10-410 nm at a time resolution of 1 min. Black
107 carbon (BC) concentrations were measured with aethalometers (AE33, Aerosol Magee Scientific).

108 2.3. Online FIGAERO-CIMS measurement and identifications of potential BrC molecules

109 The individual organic compounds in both the gas and particle phase were measured with a filter inlet for gases and
110 aerosols coupled to a high-resolution time-of-flight chemical ionization mass spectrometer (FIGAERO-HR-ToF-CIMS,
111 Aerodyne Research Inc. hereafter CIMS) employing iodide (I⁻) for chemical ionization (Lopez-Hilfiker et al., 2014;
112 Jiang et al., 2022). During the gas-phase measurement, the ambient air was sampled via a fluorinated ethylene
113 propylene (FEP) tube of 4.5 m length (flow rate 8 L min⁻¹, residence time 0.9 s). At the same time, the particles were
114 collected on a Teflon (Polytetrafluoroethylene, PTFE) filter via a separate sampling port connected to a PM_{2.5} inlet
115 (total flow rate 16.7 L min⁻¹) and an 8 m long stainless-steel tube. The loading time and sampling flow of Teflon filters
116 were 30 minutes and 4 L min⁻¹, respectively. At regular intervals (46 min), the gas-phase measurement was switched
117 off and particles on the filter were desorbed by a flow of ultra-high-purity nitrogen (99.9999 %) heated from room
118 temperature to 200 °C over the course of 35 min (Lopez-Hilfiker et al., 2014; Huang et al., 2019a). The resulting mass
119 spectral signal evolutions as a function of desorption temperature are termed thermograms (Lopez-Hilfiker et al., 2014).
120 Integration of thermograms of individual compounds yielded their signal in counts per second, which were converted
121 to mass concentrations using an average sensitivity of 22 count s⁻¹ ppt⁻¹ (Lopez-Hilfiker et al., 2014). After the field
122 campaign, the calibration of 4-nitrophenol, 4-nitrocatechol, 2-methyl-4-nitrophenol, and 4-methyl-5-nitrocatechol was
123 utilized to characterize the sensitivity factor of nitro aromatic compounds (NACs), as shown in the Supplement. The
124 sensitivity factors of our iodide CIMS for 4-nitrophenol, 4-nitrocatechol, 2-methyl-4-nitrophenol, and 4-methyl-5-
125 nitrocatechol were 0.80 ± 0.44, 0.50 ± 0.32, 0.96 ± 0.52, 0.97 ± 0.63, respectively (Figure S9). The average sensitivity
126 factor of 4 NACs was 0.81 ± 0.53. We used this average sensitivity factor to calibrate other potential brown carbon
127 molecules in this study. The sensitivity factor of levoglucosan was 0.40 ± 0.14 in this study (Figure S10). We used the
128 sensitivity factor of 0.40 ± 0.14 to estimate the concentrations of molecules, which are not identified as potential BrC
129 molecules. Please note that the sensitivity of CIMS for different organic compounds varies by a few orders of
130 magnitude. Sensitivity uncertainties were taken into account in the calculation of the overall uncertainties of CIMS
131 concentrations (±60%) following the approach by Thompson et al. (2017).

132 During the measurements, the mass resolution of FIGAERO-CIMS was relatively stable with about 4000 m/Δ. The
133 interference from isomers with different vapor pressures or thermal fragmentation of larger oligomeric molecules can
134 lead to more complex, multimodal and broader thermograms (Lopez-Hilfiker et al., 2014). The signal integration can
135 include the different isomers or thermal fragmentation of larger oligomers. Therefore, the isomers or thermal

136 decomposition can lead to increase errors of estimating the organic mass concentrations. In this study, BrC molecules
137 were identified and partially quantified in atmospheric aerosol by FIGAERO-CIMS. Please note that the iodide CIMS
138 has sensitivities varying over several orders magnitude for different compounds e.g. of different oxidation states
139 (Lopez-Hilfiker et al., 2016). Therefore, the quantitative interpretation is limited to the small amount of compounds
140 for which we could do calibration with authentic standards. Keeping this in mind, it can still be meaning to a relative
141 comparison of the large number of high oxidized compounds assuming the same sensitivity. The raw data were
142 analysed by using the toolkit Tofware (v3.1.2, Tofwerk, Thun, Switzerland, and Aerodyne, Billerica) with the Igor Pro
143 software (v7.08, Wavemetrics, Portland, OR). Gas phase background was determined by sampling zero air (high purity
144 synthetic air). Particle phase backgrounds were assessed by putting an additional Teflon filter upstream of the particle
145 phase sampling port during the deposition (Huang et al., 2019a; Lee et al., 2018).
146 We observed typically about 1500 mass peaks from particles and 120 mass peaks in gases corresponding to different
147 oxygenated organic compounds by using FIGAERO-CIMS. Individual compounds were assigned to the mass peaks
148 by fitting, $C_cH_hO_oN_n$, different numbers of atoms: c carbon, h hydrogen, o oxygen, n nitrogen (Lopez-Hilfiker et al.,
149 2014). A double bond equivalent (DBE) can be calculated as follows (Daumit et al., 2013):

$$150 \quad DBE = \frac{n-h}{2} + c + 1 \quad (1)$$

151 Lin et al., (2016, 2018) employed high-resolution mass spectrometry to analyze biomass burning organic aerosol. They
152 assigned potential brown carbon compounds according to the correlation of double bond equivalents (DBE) with the
153 number of carbon atoms per molecule (Figure S12). We used this method to assign 178 potential BrC molecules
154 (including 7 NACs) in the particle phase and 31 potential BrC molecules (including 4 NACs) in the gas phase, as
155 shown in Figure 1 in the corresponding mass spectra. The gas to particle phase partitioning coefficients of those semi
156 volatile potential brown carbon molecules which could be measured in both phases with sufficient sensitivity are listed
157 in table S6. A few other studies used this method also to assign more brown carbon molecules. For example, good
158 correlations ($r = 0.9$) between mass absorption efficiency at 365 nm and potential brown carbon molecules of larger
159 molecular weight were found by Tang et al., (2020). Xu et al., (2020) used this method to assign 149 nitrogen-
160 containing potential BrC chromophores at the Tibetan Plateau and we used this method to assign potential BrC
161 molecules in downtown Karlsruhe (Jiang et al., 2022). The potential BrC molecules we assigned according to this
162 method for the particle and the gas phase are listed in Tables S2 and S3.

163 2.4. Particle light absorption from aethalometer measurements

164 In the aethalometer AE33 (Magee Scientific), aerosol particles are continually sampled on a quartz filter and the optical
165 attenuation is measured with time resolutions 1 minute at seven wavelengths (370, 470, 520, 590, 660, 880, and 950
166 nm) during this campaign. The light absorption at seven wavelengths was calculated from the measured attenuation.
167 Attenuation is measured on two spots with different sample flows and on the reference spot without sample flow. The
168 two loading spots with different flow are used to allow for loading effect corrections (Drinovec et al., 2015). Since our
169 aethalometer has been used two loading spots, the loading effect was corrected by a Dual-spot loading compensation
170 algorithm (Drinovec et al., 2015). To further address the scattering effect (Yus-Díez et al., 2021), we did comparison
171 experiments in the Aerosol Preparation and Characterization (APC) chamber (Huang et al., 2018). Black carbon was
172 injected into the APC chamber by using the PALAS soot generator (GfG 1000, Palas) (Saathoff et al., 2003). The APC
173 chamber was connected to a photoacoustic spectrometer (PAS) operating at three wavelengths (405, 520, and 658 nm)
174 (Linke et al., 2016) and an aethalometer AE33. As shown in Figure S11, for three wavelengths (370, 520, and 660 nm),
175 the correlation slopes were 1.88, 1.94, and 1.98, respectively. The average multiple-scattering correction factor was
176 1.90 ± 0.06 in this study.

177 The BC mass concentration is calculated from the change in optical attenuation at 880 nm in the selected time interval
178 using the mass absorption cross section $7.77 \text{ m}^2 \text{ g}^{-1}$ (Gundel et al., 1984), since other aerosol particles (organic aerosol
179 or mineral) have less absorption at this wavelength and major absorption is contributed from BC alone. The attenuation
180 mass absorption coefficients of AE33 from 370 – 880nm were 18.47, 14.54, 13.14, 11.58, 10.35, and $7.77 \text{ m}^2 \text{ g}^{-1}$,
181 respectively. The absorption measurements by aethalometer have the filter-based lensing effect (Moschos et al. 2021).
182 According to previous studies, the uncertainty from lensing effect for BC and BrC measurement were 8%-27% and
183 6%-20%, respectively (Moschos et al. 2021). We assumed an AAE_{BC} value of 1.0 in this study. However, this
184 assumption introduces an uncertainty in the estimations of BC and BrC light absorptions. According to previous studies,
185 the AAE_{BC} ranges between 0.8-1.4 (Lack and Langridge 2013). This range although maybe not fully applicable to our
186 measurement location, potentially causes relatively large uncertainties of up 81% (at 370nm) in splitting between BrC
187 and BC absorption (Figure S13) (Duan et al. 2024). Despite these potentially large uncertainties on absolute absorption
188 values, we consider this method still useful. Our assumption of $\text{AAE}_{\text{BC}} = 1.0$ is reasonable for our location as based on
189 previous measurements and it should still allow to discuss the relative evolution of BC and BrC absorption.

190 We assumed that the absorption from dust and other aerosol was negligible. Hence, the absorption was only
 191 contributed from BC and BrC. Therefore, $Abs(\lambda)$ can be divided in BC and BrC absorption:

$$192 \quad Abs = Abs_{BrC}(\lambda) + Abs_{BC}(\lambda) \quad (2)$$

193 where $Abs_{BrC}(\lambda)$ is the absorption caused by BrC at the following aethalometer wavelengths, $\lambda = 370, 470, 520, 590,$
 194 or 660 nm while $Abs_{BC}(\lambda)$ is the absorption contributed by BC at the same wavelength (Wang et al., 2019). To
 195 determine $Abs_{BC}(\lambda)$ at each wavelength, we assumed that BC was the only absorber at $\lambda = 880$ nm, and thus the $Abs_{BC}(\lambda)$
 196 ($\lambda = 370, 470, 520, 590,$ and 660) can be extrapolated from the following equation:

$$197 \quad Abs_{BC}(\lambda) = Abs_{880} \times \left(\frac{\lambda}{880}\right)^{-AAE_{BC}} \quad (3)$$

198 where AAE_{BC} represents the spectral dependence of $Abs_{BC}(\lambda)$, and a value of 1.0 was chosen for AAE_{BC} based on
 199 previous studies in Germany (Teich et al., 2017). Finally, one can obtain the $Abs_{BrC}(\lambda)$ as follows:

$$200 \quad Abs_{BrC}(\lambda) = Abs(\lambda) - Abs(880) \times \left(\frac{\lambda}{880}\right)^{-AAE_{BC}} \quad (4)$$

201 The fraction of wood burning black carbon (BC_{wb}) was calculated by using the Aethalometer model (Sandradewi et
 202 al., 2008a; Sandradewi et al., 2008b):

$$204 \quad BC_{wb} = \left[\frac{b_{abs(470nm)} - b_{abs(950nm)} \times \left(\frac{470}{950}\right)^{-\alpha_{ff}}}{\left(\frac{470}{950}\right)^{-\alpha_{wb}} - \left(\frac{470}{950}\right)^{-\alpha_{ff}}} \right] / b_{abs(950nm)} * BC \quad (5)$$

205 Where two pairs of Ångström exponents values were utilized to obtain BC associated with fossil fuel (BC_{ff}) and wood
 206 burning (BC_{wb}). One of the largest sources of uncertainty in the Aethalometer model is related to the section of α_{ff} and
 207 α_{wb} values (Healy et al., 2017; Zotter et al., 2017). In addition, the α_{ff} was typically in the range of $\sim 0.8 - 1.2$ in
 208 ambient air whereas α_{wb} can vary from 1.6 to 2.2 (Saarikoski et al., 2021). However, we used the α_{ff} and α_{wb} values
 209 as 0.95 and 1.68 to calculate the BC source (Helin et al., 2018), since our measurement site is in a rural area and nearby
 210 a suburban area.

211 **3. RESULTS AND DISCUSSION**

212 **3.1. Overview of the field observations**

213 Figures S1 and S2 give an overview of the measurement location and the meteorological parameters, traces gases,
 214 particle concentrations, and their optical properties during the campaign. The major wind directions at KIT Campus
 215 Nord, 3 km east of the village of Eggenstein-Leopoldshafen, were northeast and southwest (Figure S1) caused by
 216 channeling of the wind in the Rhine valley. The average wind speeds were 1.1 ± 0.8 (average \pm standard deviation) m

217 s⁻¹. Depending on meteorological conditions, local sources and regional transport had a major impact on air quality in
218 Leopoldshafen in summer (Shen et al., 2019). As shown in Figure S5, O₃ had diurnal variations with peaks at daytime
219 and an average of 41.3 ± 26.2 μg m⁻³ during the campaign. In contrast, the relative humidity (RH) showed diurnal
220 variations with peaks at nighttime and an average of 68 ± 16% during the campaign (Figure S5). The average
221 temperature during the winter campaign was 6.5 ± 5.6 °C and slowly increased from beginning to the end of the
222 campaign. NO₂ had high concentrations at some periods e.g. from 20th to 23th February with 22 ± 8.6 μg m⁻³ and from
223 2nd to 4th March with 35 ± 14 μg m⁻³. The average SO₂ concentration was 0.8±1.0 μg m⁻³, significantly lower than the
224 NO₂ concentrations. During some Saharan dust events, the PM_{2.5} and PM₁₀ mass concentrations were 21 ± 6 and 45 ±
225 20 μg m⁻³, respectively, from 18th to 26th February and 19 ± 6 and 24 ± 7 μg m⁻³, respectively, from 1st to 4th March as
226 indicated by red boxes in the lowest panel of Figure S2. In addition, BC showed many spikes and a good correlation
227 (r = 0.8) with NO₂ (Figure S3). This indicates that there were many combustion events during the campaign (Figure
228 S3). The absorption Ångström exponents of particles between 370 and 520 nm (AAE₃₇₀₋₅₂₀) and AAE₆₆₀₋₉₅₀ had diurnal
229 variations with peaks at nighttime. We calculated the fraction of wood burning BC and fossil fuel BC as shown in
230 Figure S3 using the Aethalometer model (Sandradewi et al., 2008a). During the winter campaign, the biomass burning
231 BC was on average 0.61 ± 0.049 μg m⁻³, mostly higher than 0.025 ± 0.027 μg m⁻³ for fossil fuel BC. The AAE₃₇₀₋₅₂₀,
232 AAE₆₆₀₋₉₅₀, biomass burning BC, and NO₂ values were enhanced from 20th to 23th February and 2nd to 4th March. This
233 indicates that strong biomass burning (BB) events were on these days. During this winter campaign, the BrC absorption
234 accounted for ~40% of total absorption caused by BC and BrC at 370 nm. This points to the at least regional or seasonal
235 importance of BrC absorption which has an important effect on air quality and climate.

236 3.2. Mass concentrations and volatility of potential brown carbon molecules

237 Figure 2 shows an overview of levoglucosan concentrations, BC concentrations, absorption of brown carbon at 370
238 nm (b_{brC370}), AAE₃₇₀₋₅₂₀, volatility and mass concentrations of 178 potential brown carbon molecules identified in the
239 particle phase and 31 potential brown carbon molecules in the gas phase during the whole winter campaign. We
240 identified 178 potential BrC molecules according to the method developed by Lin et al., (2018) (cf. section 2.3.). The
241 mass of these molecules shows a good correlation (r = 0.7 ± 0.1) with the absorption at 370 nm (b_{BrC370}) of BrC (cf.
242 Figure S6). This indicates that it is meaningful to extract these 178 potential BrC molecules from more than one
243 thousand and five hundred molecules detected by FIGAERO-CIMS based on the double bond equivalent/carbon
244 number ratio (DBE/C) of each molecule being higher than 0.5 and less than 0.9. The levoglucosan showed a good

245 correlation ($r = 0.7$) with BC. This is in line with the large fraction of biomass burning contributing to BC during the
246 winter campaign. Biomass burning BC accounted for $(71 \pm 40)\%$ of total BC as we discussed above. The 178 potential
247 BrC molecules detected in the particle phase correspond to an average mass concentrations of $83 \pm 44 \text{ ng m}^{-3}$. In
248 addition, the nitro aromatic compounds (NACs) were also detected during the winter campaign. The mass
249 concentration of $\sum\text{NACs}$ in the gas phase and particle phase were $1.9 \pm 1.5 \text{ ng m}^{-3}$ and $17.5 \pm 18.4 \text{ ng m}^{-3}$, respectively
250 (Table S4 and S5). Mohr et al. (2013) found that five BrC molecules (nitro aromatic compounds) were 20 ng m^{-3}
251 detected by CIMS during winter in Detling, United Kingdom. Jiang et al. (2022) measured an average concentration
252 of five BrC molecules (nitro aromatic compounds) of $1.6 \pm 0.9 \text{ ng m}^{-3}$ during the winter at a kerbside in downtown
253 Karlsruhe, a city in southwest Germany and close to our measurement site. Therefore, the detection of the 178 potential
254 BrC molecules allows more complete assessment of the BrC concentrations during this winter campaign. Their
255 concentrations were significantly higher for biomass burning (BB) events e.g. $144 \pm 41 \text{ ng m}^{-3}$ at BB event 1 and 124
256 $\pm 39 \text{ ng m}^{-3}$ at BB event 2, respectively. In addition, the absorption of brown carbon at 370 nm ($b_{\text{brc}370}$) had high peaks
257 with $\sim 100 \text{ Mm}^{-1}$ and the $\text{AAE}_{370-520}$ of particles increased from ~ 1.5 to ~ 2 during the BB events. The average
258 concentration of potential BrC in the gas phase was $8.5 \pm 6.7 \text{ ng m}^{-3}$ during the winter campaign. At BB events, their
259 concentration can reach up to 38 ng m^{-3} . Therefore, biomass burning had a significant impact on optical properties of
260 aerosol and brown carbon concentrations. The lowermost panel of Figure 2 shows the temporal variation of the average
261 volatility of brown carbon molecules in the gas and particle phase. The average volatility or saturation concentration
262 ($\log_{10}C_{\text{sat}}$) of potential BrC in the particle phase was with $-1.1 \pm 0.5 \text{ } \mu\text{g m}^{-3}$ lower than $0.9 \pm 0.6 \text{ } \mu\text{g m}^{-3}$ of potential BrC
263 in the gas phase during the winter campaign. Organic compounds with $\log_{10}C_{\text{sat}}$ lower than $-4.5 \text{ } \mu\text{g m}^{-3}$, between -4.5
264 and $-0.5 \text{ } \mu\text{g m}^{-3}$, between -0.5 and $2.5 \text{ } \mu\text{g m}^{-3}$, and between 2.5 and $6.5 \text{ } \mu\text{g m}^{-3}$ are termed extremely low-volatility
265 organic compounds (ELVOCs), low-volatility organic compounds (LVOCs), semi-volatile organic compounds
266 (SVOCs), and intermediate-volatility organic compounds (IVOCs), respectively (Donahue et al., 2009). Therefore,
267 BrC in the particle phase can be classified on average to the LVOCs and BrC in the gas phase to the SVOCs.

268 **3.3 Absorption contribution of nitroaromatic compounds and potential brown carbon molecules**

269 Black carbon dominated light absorption of aerosol particles with a contribution of 100% at 880 nm and decreasing to
270 73% at 370 nm. With shorter wavelengths, the brown carbon absorption contribution significantly increased
271 contributing 27% of total aerosol absorption at 370 nm (Figure 3a). We have no independent quantification of the total
272 organic aerosol mass loadings. However, we estimated the total organic mass as a fraction of $50 \pm 20 \%$ of $\text{PM}_{2.5}$ which

273 is a typical fraction for at the location (Song et al., 2022; Song et al., 2024). According to this assumption, the average
274 organic mass concentration was $4.5 \pm 3.1 \mu\text{g m}^{-3}$. The organic mass detected by FIGAERO-CIMS based on calibrated
275 sensitivity factors was $37 \pm 20\%$ of the estimated organic mass. This is in a similar range as observed in previous
276 studies (Ye et al., 2021). We calculated the light absorption of NACs by using molecular MAC_{365} (Xie et al., 2017),
277 as shown in Table S5. Based on this, the mean light absorption of the sum of the seven NACs was calculated to be 0.2
278 $\pm 0.2 \text{ Mm}^{-1}$, contributing to $2.2 \pm 2.1\%$ of total BrC absorption at 370 nm, but they only contributed $0.45 \pm 0.32\%$ of
279 the total organic mass.

280 In order to calculate the light absorption from the other 171 potential brown carbon molecules identified, we assumed
281 an average MAC value of $9.5 \text{ m}^2 \text{ g}^{-1}$ at 370 nm for all BrC molecules to estimate their absorption (Jiang et al., 2022).
282 So far, the MAC_{370} of most potential brown carbon molecules are still unknown. In addition, since the potential BrC
283 molecules detected by FIGAERO-CIMS could have isomers effect, we did not calibrate mass absorption coefficients
284 of 171 potential BrC. Despite these uncertainties, we think it is reasonable to estimate the order magnitude of the total
285 BrC absorption based on this assumption. Based on this assumption, we calculated the light absorption of the 171
286 potential brown carbon molecules identified to $0.6 \pm 0.3 \text{ Mm}^{-1}$ at 370 nm as average for the whole winter campaign.
287 This is half the values Jiang et al. (2022) found as mean light absorption of 316 potential BrC molecules of 1.2 ± 0.2
288 Mm^{-1} at 365 nm for downtown Karlsruhe in winter. Relative to this total organic aerosol particle mass and the measured
289 brown carbon absorption, the 171 potential identified brown carbon molecules and 7 NACs only accounted for $2.6 \pm$
290 1.5% of the total organic mass, but explain $14 \pm 13\%$ of total brown carbon absorption at 370 nm (Figure 3b and 3c).
291 Palm et al. (2020) found that particulate nitroaromatic compounds (BrC molecules) can explain $29 \pm 15\%$ of average
292 BrC light absorption at 405 nm, despite accounting for just $4 \pm 2\%$ of average OA mass in fresh wildfire plumes. Mohr
293 et al. (2013) found that five nitroaromatic compounds (BrC molecules) are potentially important contributors to
294 absorption at 370 nm measured by an aethalometer and account for $4 \pm 2\%$ of UV light absorption by brown carbon in
295 Detling, United Kingdom during winter. Jiang et al. (2022) determined a mean light absorption of the 316 potential
296 BrC molecules accounting for $32 \pm 15\%$ of methanol-soluble BrC absorption at 365nm, but only accounted for $2.5 \pm$
297 0.6% of the organic aerosol mass. Therefore, even small mass fractions of strongly absorbing brown carbon molecules
298 can dominate the brown carbon absorption.

299 3.4 Diurnal variations and sources of potential BrC in the gas phase

300 As shown in Figure 4a, the 31 gas-phase potential BrC (GBrC) molecules showed higher concentrations at daytime
301 (09:00-17:00) and lower concentration between evening and early morning (18:00-08:00). Salvador et al. (2021) also
302 found that 16 gas-phase nitro-aromatic compounds (BrC molecules) measured by FIGAERO-CIMS were higher during
303 daytime and lower at nighttime during winter in rural China. As discussed above, strong biomass burning emission
304 were mostly observed at evening and early morning hours. However, gas-phase BrC had no peaks during those time
305 periods. Therefore, the primary emission from biomass burning was not a major source for GBrC at KIT Campus Nord.
306 It seems to be mainly controlled by secondary formation (e.g. photochemical smog) or/and particle-to-gas partitioning
307 (Salvador et al., 2021).

308 To demonstrate how secondary formation and partitioning control the gas-phase BrC in rural Germany, we plotted
309 diurnal profiles of the average volatility and volatility fractions of IVOC, SVOC, and LVOC of the gas-phase BrC
310 (Figure 4b). The LVOC of BrC increased at evenings and decreased at daytime. In contrast, the IVOC of BrC increased
311 at daytime and reached ~17% of total $\log_{10}C^*$ (volatility) in gas-phase BrC while SVOC remained with a relative
312 constant fraction (~60%). Furthermore, the IVOC fraction of BrC in the particle-phase was only 1.5% with a flat
313 diurnal profile (Figure S7). The O/C ratio of gas-phase BrC also increased during daytime (Figure 4d). Therefore, the
314 higher fraction of IVOC in the gas phase at daytime is most likely caused by secondary formation e.g. photochemical
315 conversion/aging because of higher oxidant levels as indicated e.g. by higher concentration of ozone at same time
316 (Figure 4c) (Saarikoski et al., 2021). Figure S8 shows that BrC in the gas phase had a moderate positive correlation (r
317 = 0.4) with temperature. This explains why the temperature shows a similar diurnal profile as the gas-phase BrC.
318 Therefore, particle-to-gas partitioning was also an important source for gas-phase BrC. However, our results are not
319 consistent with previous studies where 16 BrC molecules in gas phase were mainly from primary emission during the
320 biomass burning evenings and secondary formation during the clear days in rural China (Salvador et al., 2021). Our
321 measurement site was several km away from biomass burning sites with ~7-10 km. And the 31 potential BrC in the
322 gas-phase sum up to $8.5 \pm 6.7 \text{ ng m}^{-3}$, significantly lower than 1720 ng m^{-3} of 16 BrC (Salvador et al., 2021). Cheng et
323 al. (2021) found that secondary formation was a strong source for five BrC molecules in the gas-phase. Therefore, BrC
324 in the gas-phase are less influenced from primary emissions from biomass burning but are mainly controlled by
325 secondary formation and partitioning in rural Germany.

326 3.5 Diurnal variations and sources of potential BrC in the particle phase

327 The 178 potential BrC molecules in the particle phase (PBrC) exhibited two peaks in the diurnal profile (Figure 4a)
328 averaged over the whole winter campaign. They increased from 19:00 to 01:00 with a peak at $82 \pm 35 \text{ ng m}^{-3}$ around
329 midnight. Then the PBrC slowly decreased after midnight. However, they increased again from 6:00 to 08:00 and
330 forming a second peak with $102 \pm 49 \text{ ng m}^{-3}$ in the morning. During daytime, they decreased reaching lowest values
331 with $61 \pm 31 \text{ ng m}^{-3}$ at 14:00-15:00. During the nighttime and morning hours, the higher mass concentrations of PBrC
332 were caused by residential wood burning emissions. Consistently, higher $\text{PM}_{2.5}$ concentration levels at nighttime at a
333 rural site near Karlsruhe, Germany, could be assigned to wood burning emissions from wood stove operation during
334 winter (Thieringer et al., 2022). The low mass concentrations of PBrC at daytime could be explained by photobleaching
335 and evaporation of BrC, and/or dilution by the increasing planetary boundary layer heights (Satish et al., 2017). Satish
336 et al. (2017) found that BrC over the Indo-Gangetic Plain had two peaks of BrC at evening and morning hours, and
337 lowest values during daytime.

338 To determine the sources of brown carbon, we used the edge approach (Day et al., 2015). It allows to estimate the
339 contribution of primary biomass burning (BB) to the measured BrC concentrations using levoglucosan as a primary
340 source tracer. This approach is analogous to the widely used elemental carbon (EC) tracer approach, in which EC is
341 used to distinguish the primary organic carbon (POC) and secondary organic carbon (SOC) in total organic carbon
342 (OC) measurements (Day et al., 2015; Cabada et al., 2004). Levoglucosan (lev) and potential BrC were measured
343 online by the same instruments and under the same conditions. As discussed above, we observed a good correlation (r
344 $= 0.8$) between levoglucosan and BC during the winter campaign. Therefore, levoglucosan is a suitable tracer for
345 primary BB. Please note that we did not calibrate the sensitivities of levoglucosan detected by FIGAERO-CIMS.
346 Therefore, it could cause some uncertainties to estimate brown carbon from biomass burning and secondary formation.
347 Figure 5a shows that the blue points can be used as edge points to determine the ratio of BrC/levoglucosan at the
348 primary emissions from biomass burning. The relative contributions of primary emissions (BB) and secondary (sec)
349 formation for total BrC molecules were estimated using the following expression:

$$350 \quad BrC_{BB} = \left(\frac{[BrC]}{[lev]} \right)_{BB} * [lev.]$$

$$351 \quad [BrC_{sec}] = [BrC_{Tot}] - [BrC_{BB}]$$

352 Where $([\text{BrC}]/[\text{lev}])_{\text{BB}}$ is the ratio of the concentration of the BrC to that of levoglucosan in the primary emissions from
353 biomass burning and this value is 1.1 ± 0.1 (Figure 5a), BrC_{BB} and BrC_{sec} are the fractions of BrC generated through
354 biomass burning and secondary production, respectively, BrC_{Tot} and lev. are the measured concentrations of BrC and
355 levoglucosan during the winter campaign. Using this approach, we calculated the diurnal profiles of BrC from primary
356 emissions (BrC_{BB}) and secondary formation (BrC_{sec}) shown in Figure 5b. The uncertainty of the splitting between BrC
357 from biomass burning and of secondary origin is mainly based on the levoglucosan concentration for which we have
358 included the calibration. Based on this we estimated the uncertainty of the BrC source splitting to $\pm 35\%$. The mass
359 fraction of BrC_{sec} increased at daytime and decreased at evening. This indicates that the secondary formation for BrC
360 in the particle phase was enhanced during daytime, facilitated by the higher levels of oxidants e.g. O_3 (Figure 4c). The
361 mass fraction of BrC_{BB} had two peaks at early morning and in the evening hours, respectively. This may be caused by
362 residential wood burning emissions. BrC_{BB} accounts for $39 \pm 21\%$ of the total BrC as averaged for the whole
363 measurement period. During biomass burning events, the BrC_{BB} is a major mass fraction for total BrC that accounts
364 for $61 \pm 13\%$ during BB-event1 and $65 \pm 12\%$ during BB-event-2, respectively. Therefore, the primary emissions of
365 BrC have a significant impact on BrC, especially, at biomass burning events. However, on average over the whole
366 campaign, BrC_{sec} dominates the mass fraction of BrC with $61 \pm 21\%$. Therefore, the secondary formation can be
367 considered as an important source for BrC in rural Germany. Consistently, secondary formation from biomass burning
368 emission is important for the brown carbon absorption in the Switzerland, the central Europe. (Moschos et al., 2018).
369 Secondary sources for BrC were more important for absorption than primary ones in the Southeastern Margin of the
370 Tibetan Plateau (Wang et al., 2019b).

371 To further investigate the oxidation of BrC in the particle phase we plotted, the diurnal profiles of O/C ratios of BrC
372 during the whole campaign was measured, as shown in Figure 6. The O/C ratio of the potential BrC molecules increased
373 during daytime and decreased at nighttime. This is an indication for an impact of photo-oxidation on BrC either during
374 formation or aging leading to an increase of its O/C ratio. Consequently, the O/C ratio of the potential BrC molecules
375 shows a positive correlation ($r = 0.8$) with ozone, another product of photo chemistry. In contrast, the light absorption
376 of BrC at 370 nm ($b_{\text{brc}370}$) and the double bond equivalent (DBE) decreased at daytime and increased at nighttime.
377 During daytime, the absorption of brown carbon at 370 nm decreased due to lower DBE and higher O/C values of
378 brown carbon caused by photooxidation. This is in accordance with previous studies where atmospheric photooxidation

379 diminishes light absorption of primary brown carbon aerosol from biomass burning (Sumlin et al., 2017). Oxidative
380 whitening can reduce light absorption of brown carbon during the day (Hems et al., 2021).

381 **Conclusions**

382 The chemical composition, diurnal variation, and sources of brown carbon aerosol were investigated during February-
383 March 2021 in a rural area, at KIT Campus Nord, a location characteristic for central Europe. The 178 potential brown
384 carbon molecules (including 7 nitro aromatic compounds, NACs) identified in the particle phase contributed on average
385 $83 \pm 44 \text{ ng m}^{-3}$ and 31 potential brown carbon molecules (including 4 NACs) identified in the gas phase contributed
386 on average $8.5 \pm 6.7 \text{ ng m}^{-3}$ during the whole campaign. During dedicated biomass burning events, potential BrC
387 concentrations in the particle phase were significantly higher with up to $\sim 254 \text{ ng m}^{-3}$. The 178 identified potential
388 brown carbon molecules only accounted for $2.6 \pm 1.5\%$ of the total organic mass, but explained $14 \pm 13\%$ of the total
389 brown carbon absorption at 370 nm, assuming a MAC_{370} as $9.5 \text{ m}^2 \text{ g}^{-1}$. This shows that a small fraction of the brown
390 carbon molecules dominates the overall absorption. This indicates the great importance of identifying these molecules,
391 the strong absorbers, to predict aerosol absorption.

392 Diurnal variations show that the particle-phase potential BrC had two peaks at early morning and evening hours,
393 respectively. These were mainly caused by residential wood burning emissions. In contrast, the gas-phase potential
394 BrC showed higher concentrations at daytime and lower concentrations at nighttime. The gas-phase BrC molecules
395 were mainly controlled by secondary formation (e.g. by photochemical processes) and particle-to-gas partitioning. The
396 two main sources contributed to particle-phase BrC were primary emission from biomass burning and secondary
397 formation. Secondary formation, e.g. by photooxidation, is an important source of particle-phase BrC corresponding
398 to increasing O/C ratios of BrC during daytime and a positive correlation ($r = 0.8$) with ozone concentrations. In
399 addition, the DBE of the particle-phase decreased during daytime. This indicates that the absorption of brown carbon
400 at 370 nm decreased due to lower DBE and higher O/C ratio due to the photooxidation of brown carbon. Compared
401 with previous measurements in central Europe (Lukács et al., 2007; Zhang et al., 2020), our study found that secondary
402 formation, e.g., photochemical processes, was an important source for BrC in gas and particle phases. To improve air
403 quality in winter, we need to reduce biomass burning emissions (e.g., regulate wood stoves) but also reduce the
404 precursors to form secondary aerosol. Overall, this study provides good insight into the light absorption, sources, and

405 diurnal variation from real-time observations of brown carbon molecules in central Europe by using mass spectrometry
406 and aethalometer.

407 ***Data availability***

408 Data are available upon request to the corresponding author.

409 **Competing interests**

410 At least one of the (co-)authors is a member of the editorial board of Atmospheric Chemistry and Physics

411 **Author contributions**

412 FJ and HS designed the measurement campaign. FJ, LG, JS, and HS performed the experimental work. FJ did
413 FIGAERO-CIMS and AE33 data analysis. HS and HZ processed the trace gas and meteorological data, respectively.
414 TL gave general comments for this paper. FJ wrote the paper with contributions from all co-authors.

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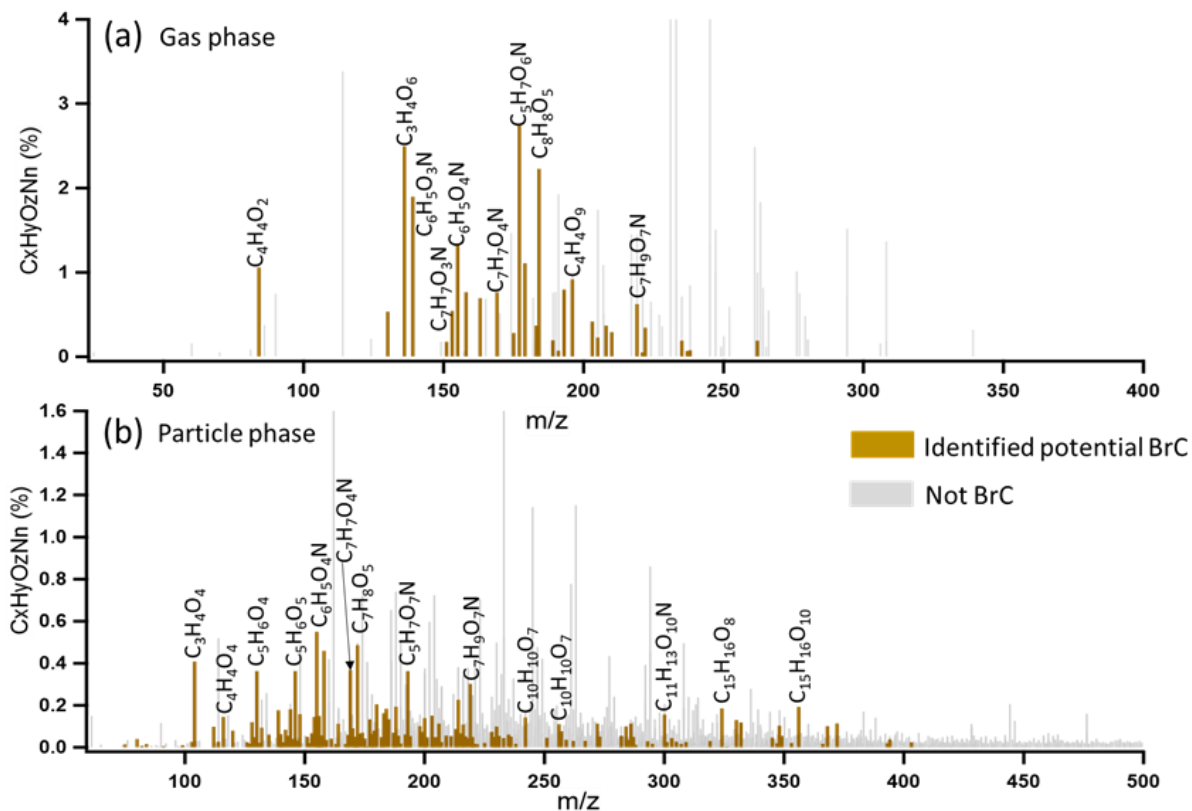
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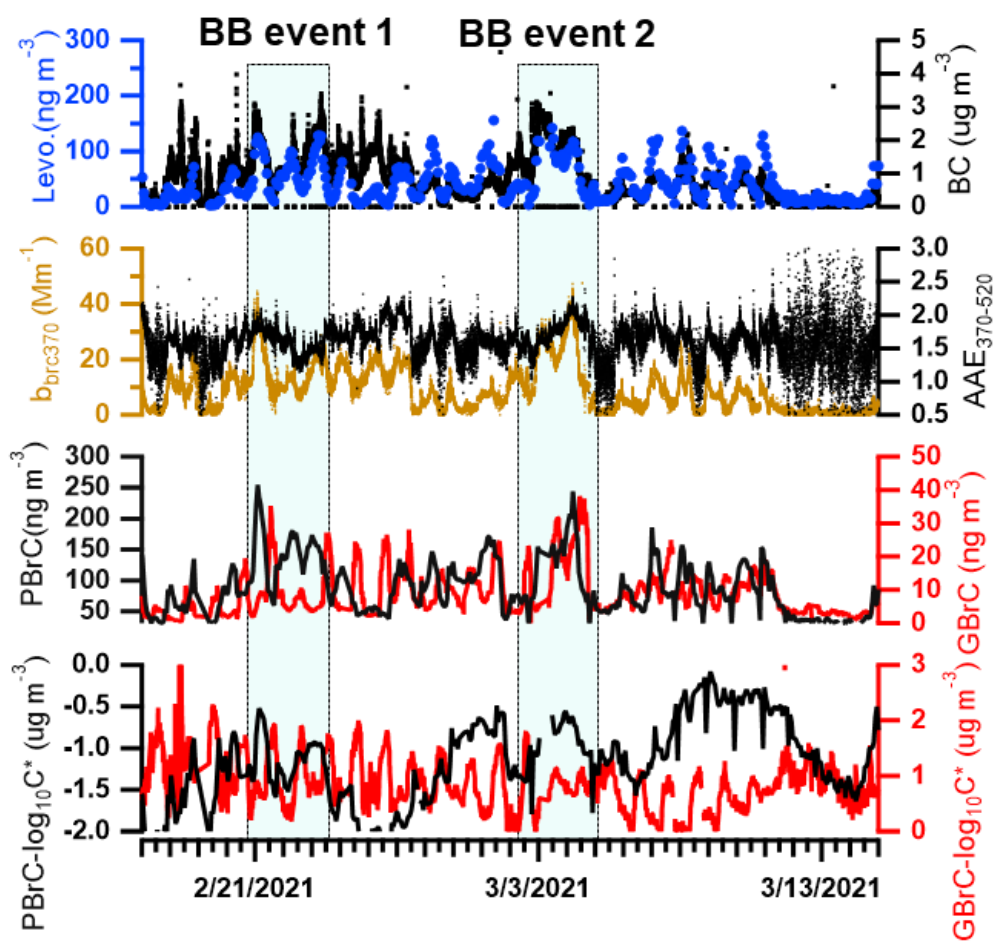
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634 **Figure 1. CIMS mass spectra of organic aerosol measured by FIGAERO-CIMS for a biomass burning event on**
 635 **March 1st, 2021, a: gas phase, b: particle phase. The CI source employs reactions of I^- ions, which convert**
 636 **analyte molecules into $[M+I]^-$ ions. Legends above MS features correspond to neutral molecules. The brown**
 637 **peaks in mass spectra were assigned as potential BrC molecules, while the gray peaks refer to the other organic**
 638 **molecules.**

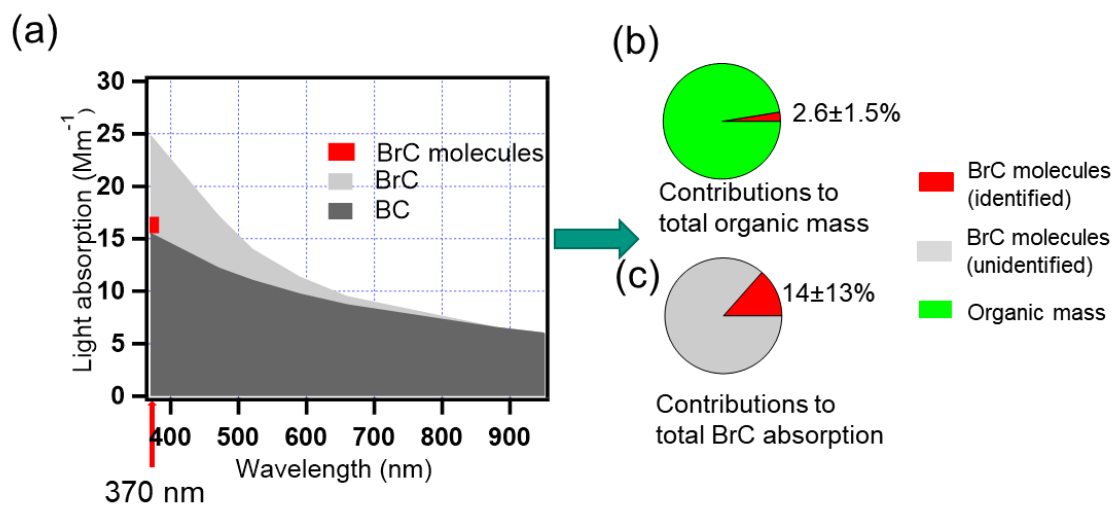
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 641 **Figure 2.** Time series of levoglucosan (Levo.) concentrations in particle phase from FIGAERO-CIMS, BC
 642 concentrations from aethalometer (AE33), absorption of brown carbon at 370 nm (b_{bc370}), absorption Ångström
 643 exponents between 370 nm and 520 nm ($AAE_{370-520}$), brown carbon concentrations in particle phase (PBrC) and
 644 gas phase (GBrC) and volatility ($\log_{10}C^*$) of brown carbon in particle phase (PBrC_ $\log_{10}C^*$) and gas phase
 645 (GBrC_ $\log_{10}C^*$) during the winter campaign.

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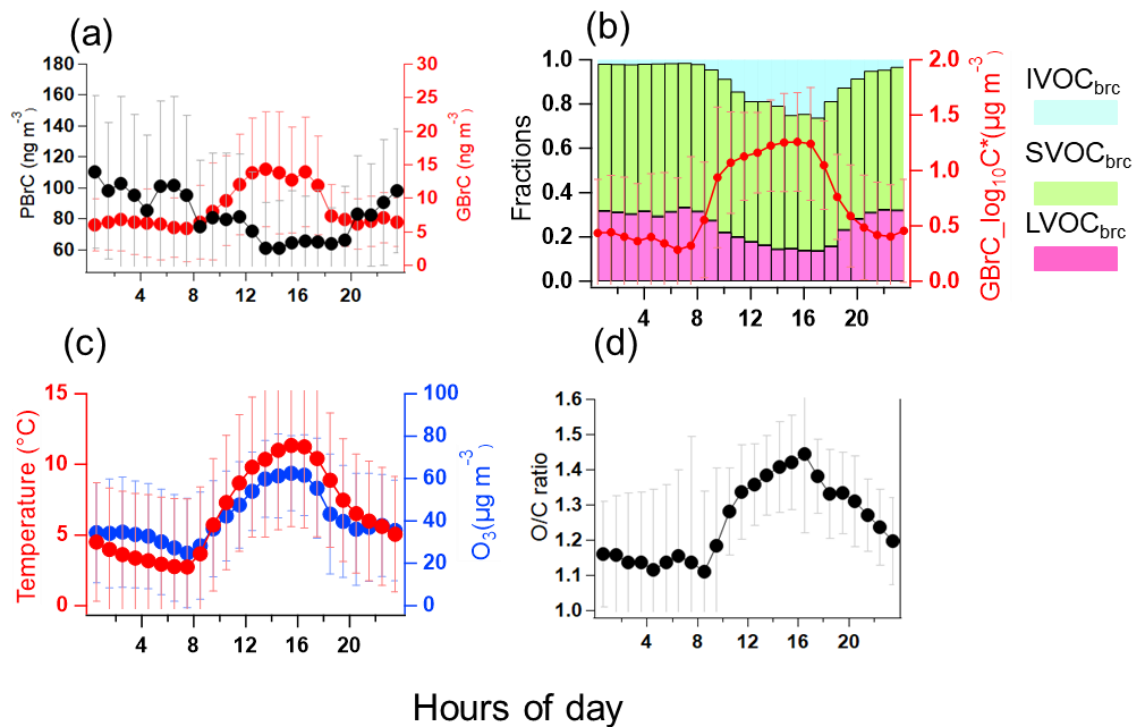


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649 **Figure 3. (a) A stacked plot showing the main contributions to aerosol absorption from brown carbon and black**
 650 **carbon based on the seven wavelengths measured by the aethalometer AE33. The contribution of the identified**
 651 **brown carbon molecules to the total aerosol absorption is indicated in red at 370 nm. (b) Average mass**
 652 **contribution of the potential BrC molecules to estimated total organic mass and (c) absorption contribution of**
 653 **the potential BrC molecules identified to total absorption by BrC. The red pie: identified BrC molecules; the**
 654 **gray pie: unidentified-BrC molecules; the green pie: all organic mass.**

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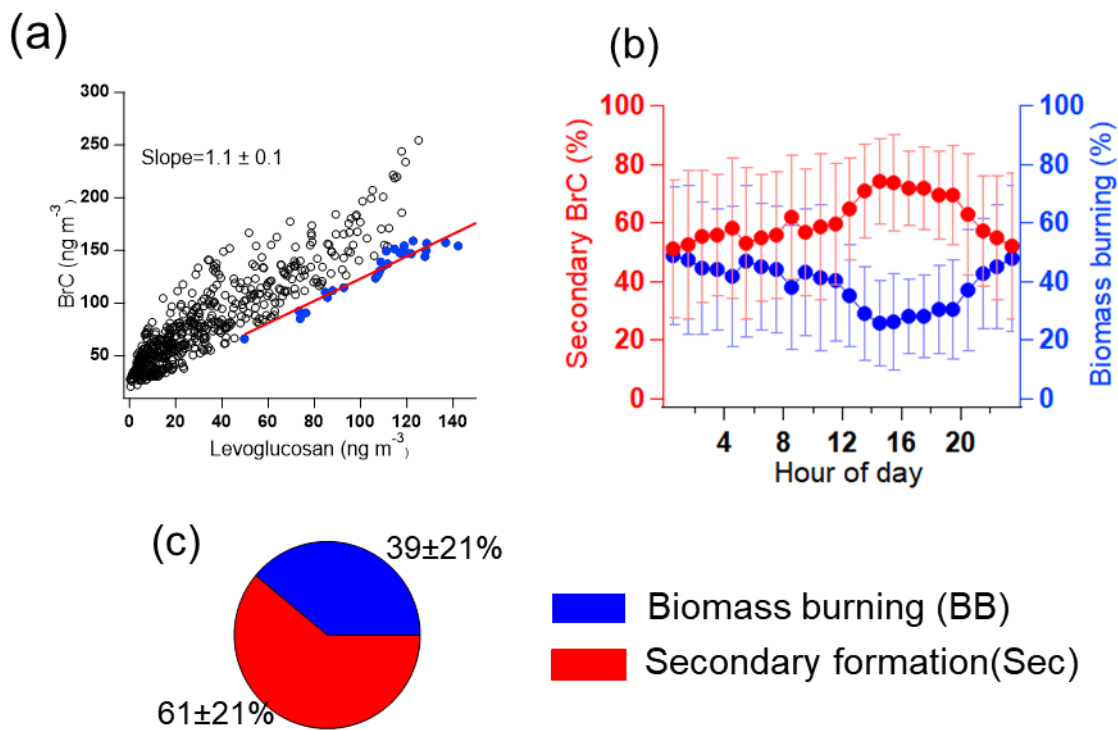
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 658 **Figure 4. Diurnal profiles averaged over the whole winter campaign of (a) BrC in the particle (PBrC) and gas**
 659 **phase (GBrC), (b) BrC volatility fractions in LVOC_{brc}, SVOC_{brc}, IVOC_{brc}, and mean BrC volatility in the gas**
 660 **phase (red line), (c) temperature and ozone concentration. (d) O/C ratio of the oxidized organic components in**
 661 **the gas phase.**

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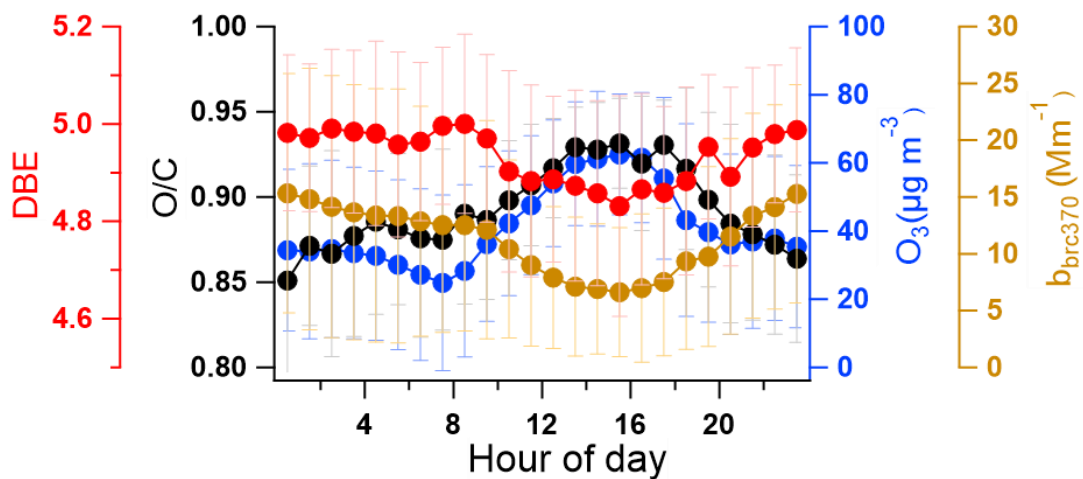


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 665 **Figure 5. (a) Correlation analysis of BrC and levoglucosan in the particle phase for the analysis of the**
 666 **contribution of biomass burning using the edge method (Day et al., 2015). Blue points are the data used to**
 667 **determine [BrC/lev.]_{BB}. (b) diurnal profile of secondary-formation BrC and biomass-burning BrC for the whole**
 668 **measurement campaign. (c) Average mass fractions of secondary formed BrC and biomass-burning primary**
 669 **BrC for the whole campaign.**

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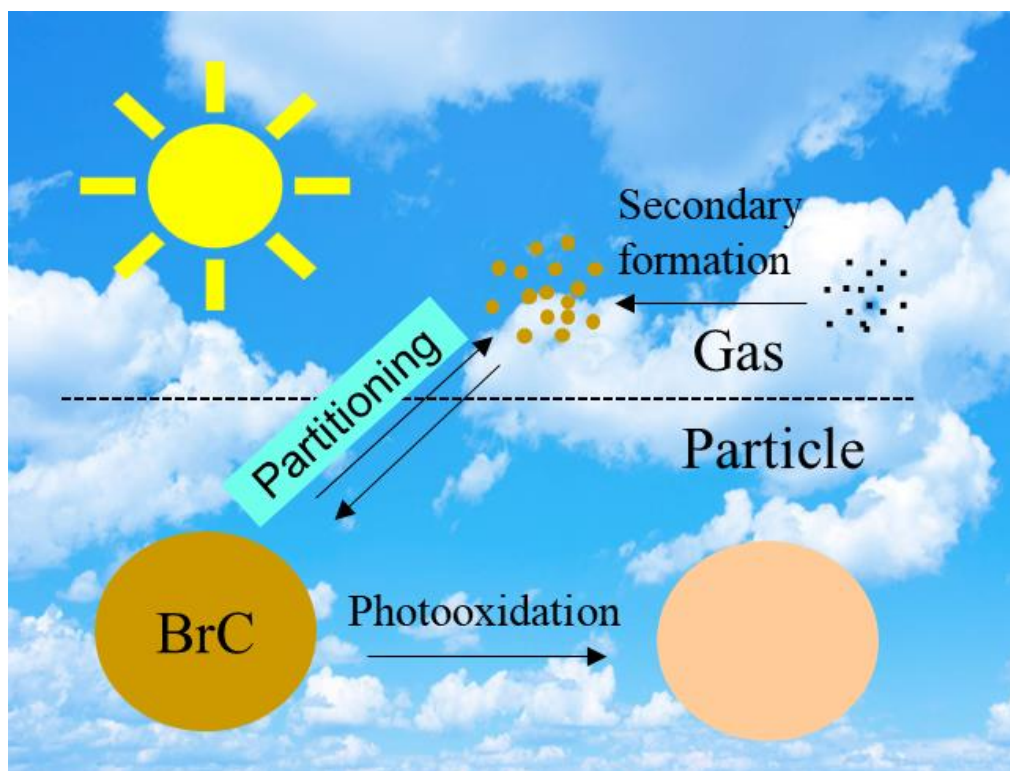
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674 **Figure 6. The diurnal profile of DBE (double bond equivalent), O/C ratio of BrC, O₃, and b_{brc370} (absorption of**
675 **BrC at 370 nm) during the whole measured period.**

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680 **Figure 7. A conceptual picture of the abstract**

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