

review of egosphere-2024-1796

In this study, the authors investigate the effect of changing the (complex) refractive index of black carbon (BC) aerosol on the aerosol absorption optical depth and on the effective radiative forcing from BC-radiation interactions. They isolate the effect of the BC refractive index by running ensembles of simulations in the atmospheric model CanAM5.1-PAM and varying only the BC refractive index among three values commonly used in Earth system models. The authors state that no previous study has isolated the impact of the refractive index of BC in this way, using a single model with other aerosol treatments held constant and the simulations being otherwise identical. If so, then this is an important assessment to have conducted. The manuscript is well written, and the results are important. As such, in my opinion, this manuscript is appropriate for publication in Atmospheric Chemistry and Physics pending some clarifications and small corrections, as I list below.

We thank the reviewer for their very helpful comments on our manuscript. Our responses to their feedback are provided in green, below. We also wish to highlight the addition of the Bescond et al. (2016) refractive index to our main analysis.

1. It is mentioned in Appendix A, but the authors should also mention in the main part of the text whether the radiative properties of the BC aerosol are affected by humidity.

In recognition of its importance for interpreting the manuscript, we have moved the model description from the appendix to the main text (lines 150-192). We have also expanded the description of the calculation of aerosol optical properties in the model (lines 181-191):

Aerosol optical properties in PAM are determined from pre-computed lookup tables. The tables are generated using Mie theory to determine optical properties as a function of relative humidity, wavelength, and particle size. Although the assumption of spherical particles may be inappropriate for freshly-emitted BC, the majority of the BC simulated by an Earth system model is hours to days old and will be relatively compact and/or internally mixed. Mie theory is thus a reasonable approximation. For internally mixed aerosol an effective refractive index is computed using the Maxwell-Garnett approximation (Wu et al., 2018), which has been demonstrated to describe the optical properties of coated BC better than volume-weighted or core-shell approximations (Adachi et al., 2010; Stevens and Dastoor, 2019).

Other treatments of aerosol morphology and mixing state would likely yield different absolute values of simulated AAOD and ERFari. However, the focus of this work is on the difference in these quantities between simulations conducted with different BCRI schemes, with other parameterizations held constant. As it is, CanAM5.1-PAM's simulated AAOD is in good agreement with other Earth system models and with observations...

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Stevens, R. and Dastoor, A.: A review of the representation of aerosol mixing state in atmospheric models, *Atmosphere*, 10, <https://doi.org/10.3390/atmos10040168>, 2019.

Wu, K., Li, J., von Salzen, K., and Zhang, F.: Explicit solutions to the mixing rules with three-component inclusions, *Journal of Quantitative Spectroscopy and Radiative Transfer*, 207, 78–82, <https://doi.org/10.1016/j.jqsrt.2017.12.020>, 2018.

2. lines 51-31: “For instance, measurements of scattering and absorption by flame-generated particle may be fit to Mie theory (assuming spherical particles) or Rayleigh-Debye-Gans theory (assuming aggregates).” – Rayleigh-Debye-Gans theory is only approximate, and thus the authors should mention other theoretical frameworks that have been used to fit/model scattering and absorption by BC aggregates, e.g., the multiple sphere T-matrix method (MSTM), the discrete dipole approximation (DDA), and the generalized multi-particle Mie method (GMM). (See for example, Liu and Mishchenko, (2005, 2007), Liu et al. (2008), Sorensen et al. (2018), Kahnert and Kanngießer (2020), and Haspel et al. (2023).) Also, “flame-generated particle” should be “flame-generated particles”.

We thank the reviewer for this suggestion, and for catching our typo. We have expanded our description of available optical models (lines 59-74):

The choice of optical model can introduce substantial uncertainty into the derived refractive index. Historically, many estimates of the refractive index of black carbon have used Mie theory (Mie, 1908) which provides an exact solution to Maxwell's equations for scattering from spherical particles. However, freshly emitted BC particles are not spherical, but instead consist of fractal-like aggregates of individual monomers. Assuming Mie theory can result in substantial underprediction of these particles' absorption and scattering; furthermore, the inferred BCRI describes a combination of pure BC and the air contained within the aggregates' voids, whereas the objective is to measure the refractive index of pure BC (Bond and Bergstrom, 2006). Improving on Mie theory, a number of optical models for aggregate particles exist. Rayleigh-Debye-Gans (RDG) theory remains the most frequently used due to its simplicity (Liu et al., 2020), but it provides an approximate solution only and does not account for multiple scattering between the monomers, which may lead to underestimation of the mass absorption cross section (Mackowski, 1995; Bond and Bergstrom, 2006). Numerically exact solutions for aggregate particles include the multi-sphere T-matrix (MSTM; Mackowski and Mishchenko, 1996) and generalized multi-particle Mie (GMM; Xu, 1995) methods for aggregates composed of non-overlapping spheres, or the discrete dipole approximation (DDA; Purcell and Pennypacker, 1973; Yurkin and Hoekstra, 2007) for more complex morphologies. For more on these methods, the interested reader is referred to Kahnert and Kanngießer (2020). In practice, however, the RDG approximation is often sufficient for BC since its scattering is so low and other uncertainties are so high (Kahnert and Kanngießer, 2020; Mackowski, 1995). None of the BCRI schemes assessed in this work were derived using these more complex optical models.

We have also included reference to the inversion techniques used to determine the full spectrum of the BCRI from a set of measurements at discrete wavelengths (lines 54-58):

The measurements are then inverted to yield the full spectrally-varying refractive index, for example using the Kramers-Kronig relations (Chang and Charalampopoulos, 1990) or the Drude-Lorentz dispersion relation (Dalzell and Sarofim, 1969; Lee and Tien, 1981). The former method is exact, but requires measurements over a greater range of wavelengths; the latter requires fewer measurements but yields poor results at visible wavelengths (Chang and Charalampopoulos, 1990; Bond and Bergstrom, 2006; Menna and D'Alessio, 1982).

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3. lines 99-100: Why is BC set to 1850 levels in the perturbed ensemble? That sounds more like what a control ensemble would be.

We have slightly rephrased our description of the control and perturbed experiments (lines 201-203)

In the control ensemble, all emissions of aerosols and greenhouse gases are transient; in the perturbed ensemble, BC emissions are fixed at 1850 levels while other emissions evolve as in the control scenario.

and added the following sentence to clarify our methodology (lines 204-206):

This method of calculating ERF is chosen over the alternative approach in which the control scenario uses preindustrial emissions and the perturbed scenario adds transient emissions of the forcer of interest because the latter method does not account for interactions between species.

4. line 227, lines 266-267: Can the authors explain why the impact on total BC effective radiative forcing is not statistically significant even though impact on the effective radiative forcing from BC-radiation interactions is large?

We have expanded the paragraph to read (lines 340-347):

Varying the BCRI, and thus the absorption, of atmospheric BC is found not to have a statistically significant impact on the aerosol-cloud forcing in this experiment. We did not vary the BCRI within cloud droplets, so the only impact on clouds would be via the impact of changes in atmospheric temperature profiles, discussed below. These changes are found to be small relative to the variability of simulated cloud fields. Similarly, the radiative forcing from albedo changes was not found to vary with BCRI scheme because we did not vary the refractive index of BC deposited on snow and ice. The only change in the total BC ERF was thus from the aerosol-radiation component. For the BB2006low and BB2006high schemes, this change was too small to result in a statistically significant increase in total ERF. However, the Besc2016 scheme led to a statistically significant increase in global-mean BC ERF relative to the dA1991 scheme, from -0.02 W/m² to +0.24 W/m².

5. lines 282-291: Even higher values for the complex refractive index of BC have been measured and used in previous studies (e.g., Janzen, 1979; soot G of Fuller et al., 1999; Liu and Mishchenko, 2005, 2007; Liu et al., 2008; Moteki et al., 2010). The authors should mention these and also assess the impact of one of these even higher values.

We thank the reviewer for highlighting these additional refractive indices.

Our original manuscript restricted its main analysis to BCRI commonly used in Earth system models, and only touched on the impacts of using more strongly-absorbing schemes in the discussion section. In the revised manuscript, we have moved the Bescond et al. (2016) refractive index, which was recommended by Liu et al. (2020), from the discussion to our main analysis. In the process we have renamed our schemes from “low”, “medium”, and “high” to “dA1991”, “BB2006low”, “BB2006high”, and “Besc2016” for improved clarity. The Besc2016 results are included in all of our figures. In brief, increasing BC absorption from the dA1991 to BB2006low (BB2006high, Besc2016) BCRI scheme increases global-mean 2015-2019 AAOD by 27% (42%, 59%) and ERF_{air} by 32% (47%, 100%).

We then address a number of other BCRI in the discussion (lines 402-419):

We have assessed four BCRI schemes here, but many others exist. As well as the Bescond et al. (2016) estimate assessed here, the review by Liu et al. (2020) highlighted the Williams et al. (2007) values of $m_{635nm} = 1.75-1.03i$ and $E(m_{635nm}) = 0.365$ as being consistent with current estimates of the absorption function. Williams et al. (2007) reported measurements at 635nm and 1310nm and did not extrapolate to other wavelengths, but assuming a relatively flat $E(m)$ through the visible range, this would indicate a degree of absorption somewhere between our BB2006high and Besc2016 schemes. Other estimates which have been widely used in the combustion science literature, such as the Janzen (1979) value of $m=2.0-1.0i$ for all visible wavelengths or the more recent Moteki et al. (2010) $m_{1064nm} = (2.26\pm 0.13) - (1.26\pm 0.13)i$, also yield $E(m)$ between BB2006high and Besc2016, but with values low enough that they are not recommended by Liu et al. (2020).

Refractive indices determined from laboratory measurements may not be representative of atmospheric black carbon. For instance, BC generated by a simple, clean laboratory flame will likely have a different temperature history – and thus, different optical and structural properties – than that generated by the more complex sources responsible for most atmospheric BC (Bond and Bergstrom, 2006). Combustion experiments also measure freshly emitted particles, which may have substantial morphological differences from hours-to-days old atmospheric BC. The recent work by Moteki et al. (2023), who measured the refractive indices of atmospheric BC particles sampled during a scientific cruise in the northwest Pacific may be more suitable for use in climate models. By combining their optical measurements with the constraints imposed by the accepted mass absorption cross section of BC, they obtained a range of plausible refractive indices suitable for describing atmospheric BC. Their recommended value, $m_{633nm} = 1.95 - 0.96i$ with $E(m_{633nm}) = 0.297$, also falls between the BB2006high and Besc2016 schemes.

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