

Multi-hazard susceptibility mapping in ~~the~~a karst context using a machine-learning method (MaxEnt)

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Abstract. In this study, we extend the application of the Maximum Entropy model (MaxEnt), traditionally applied to ecological research and less explored in natural hazard studies, to a novel context by characterising a multi-hazard scenario (i.e., flood-triggered sinkholes) in the Orléans karst region (Val d'Orléans) of France. Many regions of the world exhibit complex hazard landscapes where networks of multi-hazard interrelationships (cascades) pose challenges due to the potential interactions between hazards and the different temporal and spatial scales of hazard events. While mountainous, coastal and volcanic regions have been recognised as multi-hazard forming zones, karst terrains have received little attention despite being prone to multi-hazard events due to their distinct geology, geomorphology, hydrogeology and other environmental characteristics. Incorporating karst-specific multi-hazard scenarios [into resilience planning](#) processes supports disaster risk reduction efforts by raising the awareness of citizens, protecting elements at risk and facilitating decisions on disaster prevention. To support this aim, we developed a multi-hazard susceptibility map for the karst region of Val d'Orléans that characterises flood-triggered sinkholes. We applied MaxEnt, a machine learning method, to forecast the spatial probability distribution of flood-triggered sinkholes. Model inputs included the location of past sinkhole occurrences and geo-environmental factors contributing to sinkhole formation (e.g., topography, local geology, hydrology and flood hazard). We validated the performance of the model by initially using 70% of the sinkhole inventory data and keeping the remaining 30% for testing. This validation process assessed the model's performance using the Area Under the Curve of the Receiver Operating Characteristic (AUC-ROC). The resulting map reveals areas located up to 1 km south of the Loire River and areas with lowest elevation with highest susceptibility to flood-triggered sinkholes. We conclude that our approach to producing this type of multi-hazard scenario and map is useful for identifying flood-triggered sinkholes in Val d'Orléans and other karst areas around the globe, supporting effective land use planning.

40 1 Introduction

The identification and characterisation of multi-hazard regions (i.e., contexts affected by multiple, interrelating hazards) and the delineation of potential multi-hazards scenarios can support forward-looking approaches to risk reduction (Liu et al., 2016; Ward et al., 2022). The effectiveness of hazard mitigation measures partially depends on the capacity to collectively consider, visualize, and evaluate the complexity of hazards (Pourghasemi et al., 2020). As emphasized in the Sendai Framework for Disaster Risk Reduction (UNDRR, 2015), multi-hazard approaches are critical for a range of stakeholders including local developers, urban planners, and local and national authorities.

50 Across the globe, there are several types of geographical regions affected by multiple natural hazards, which can interrelate, and therefore, have the intrinsic potential to form a network of hazard interrelationships. Mountainous, coastal and volcanic regions are commonly recognised multi-hazard-forming zones, where hazards interact with each other in different forms such as simultaneously, cascadingly or cumulatively (Kappes et al., 2012; Gill and Malamud, 2016; de Ruiter et al., 2020; 55 Gallina et al., 2020). Several studies have mapped out the typical multi-hazard scenarios in mountainous regions (see Yousefi et al., 2020; Tsereteli et al., 2018; Pourghasemi et al., 2019; Terzi et al., 2019). In mountain environments, an earthquake or heavy rainfall can trigger flooding, landslides or debris flows, which then by blocking a river or breaching a river's dam can generate further hazards, such as flooding (Kappes et al., 2012; van den Bout et al., 2022). For example, the 2015 Gorkha earthquake event in the 60 Himalayan nation of Nepal triggered more than 24,000 landslides adding to a death toll of nearly 9000 people (Collins et al., 2015).

In coastal environments, there have been several attempts to characterise and map coastal multi-hazards (see Marignani et al., 2017; Gallina et al., 2020). In such settings, a hurricane can cause strong wind and storm surges, which trigger flooding, mudslides and landslides (Appelquist and Balstrøm, 2015; 65 López-Saavedra and Martí, 2023). For example, on 14th September 2022, in Puerto Rico, hurricane Fiona (Category 1) unleashed more than 30 inches/76 cm of rain resulting in widespread flooding. Consequently, this flooding triggered numerous mudslides and landslides (Richard et al., 2023). There have been several attempts to characterize and map coastal multi hazards (see Marignani et al., 2017; 70 Gallina et al., 2020).

Similarly, in volcanic regions, several multi-hazard scenario identification and mapping tools for volcanic settings can be found in the literature (Neri et al., 2013; López-Saavedra and Martí, 2023). a Volcanic eruptions can cause a range of hazards (e.g., tephra fall, pyroclastic density currents, lava 75 flows, and gas emissions) that can combine with other hazards to trigger new hazards (e.g., rainfall mobilising tephra deposits as lahars). Volcanic unrest may trigger landslides or earthquakes, and in some cases tsunamis (Gill et al., 2021). The eruption of Mount Pinatubo in the Philippines in 1991 can be a notable example, which was accompanied by an earthquake triggered earthquake-triggered lahars.

80 The combination of these events led to over 300 casualties, displacement of more than 2 million people,
and damaged 800 houses (de Ruiter et al., 2020). ~~Several multi-hazard scenario identification and
mapping tools for volcanic settings can be found in the literature (Neri et al., 2013; López-Saavedra and
Martí, 2023).~~

85 Another multi-hazard landscape context, often overlooked in the multi-hazard literature, is karst terrain.
Karst terrain is created from the dissolution of soluble rocks (carbonates and/or evaporites) and is characteri
sized by distinct landforms such as springs, caves and sinkholes (Waltham et al., 2005). These
landscapes, with their distinct geology, geomorphology, hydrogeology, and other environmental
characteristics, are prone to multi-hazard events (Gill and Malamud, 2014; Parise et al., 2018). While
90 previous studies recogni
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ze the potential for heavy rainfall events and flooding to trigger sinkholes in
karst environments and emphas
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ze the need to include karst landscapes in multi-hazard discussions
(Lei et al., 2013; Lei et al., 2016; Xiao et al., 2016; Xiao et al., 2018; Kwak et al., 2020), sinkhole
mapping in karst areas has predominantly focused on single hazards (Subedi et al., 2019; Kim et al.,
2022).

95 Methods used for mapping sinkhole hazard can be broadly categori
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zed into qualitative, quantitative,
and hybrid/semi-quantitative approaches (Galve et al., 2008; Galve et al., 2009). Qualitative methods
typically encompass descriptive representations that rely on the expertise of local specialists such as the
Analytical Hierarchy Process (AHP), which may introduce subjective biases (TaHERI et al., 2015; Subedi
et al., 2019; Qiu et al., 2020). Hybrid/semi-quantitative approaches often combine expert judgment with
100 spatial pattern analyses such as the Nearest Neighbour Index and kernel density estimation, and field
observations. Both of these methods are data-limited and rely on interpretive perspectives of the experts
involved (Clark and Evans, 1954; Hyatt and Jacobs, 1996; Galve et al., 2009).

Quantitative approaches, however, rely on numerical data and statistical analysis to explain phenomena.
105 The most applied ones are the frequency ratio (Ozdemir, 2015), multivariate statistical methods such as
logistic regression, artificial neural network, and support vector machine (TaHERI et al., 2015; Kim ~~and
Namet al.~~, 2018; Kim et al., 2022) and weight-of evidence (Perrin et al., 2015). Recently Machine
Learning (ML) approaches have advanced prediction techniques, resulting in enhanced and more
effective solutions (TaHERI et al., 2019; Nachappa et al., 2020; Tehrani et al., 2022). Different ML
110 algorithms have been applied to sinkhole mapping including decision trees (Gao and Alexander, 2008),
Random Forest which is an extension to decision trees promising less overfitting models compared to
the traditional decision tree models (Breiman, 2001; Elmahdy et al., 2022), and Bayes-based machine
learning algorithms (e.g., Naïve Bayes, Bayes Net), Logistic Regression, and Bayesian Logistic
Regression (TaHERI et al., 2019; Tehrani et al., 2022). In this study, we use Maximum Entropy Model
115 (MaxEnt), an ML algorithm, originally introduced by ecologists (Phillips et al., 2006) with wide
applications in natural, social and computational sciences. Bianchin et al. (2022) carried out one of the
first applications of MaxEnt to sinkhole mapping, demonstrating its potential in karst contexts. Here,

we build on the work of Bianchin et al. (2022) to explore the use of MaxEnt in mapping flood-triggered sinkholes multi-hazard.

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We applied the MaxEnt method to Val d'Orléans in France with an area of 260 km² and about 70,000 inhabitants belonging to 31 municipalities (Serrhini et al., 2023). In 2016, Val d'Orléans experienced a multi-hazard event where a flood triggered by rainfall resulted in the collapse of more than 100 sinkholes (Noury et al., 2018). This impacted the main highway between Orléans and Paris, four small towns, and an industrial park, and resulted in financial loss of over €20 million (Luu et al., 2019). It remains unclear if the flood directly triggered sinkhole formations or triggered the collapse of pre-existing sinkholes. Therefore, here we use the term "flood-triggered sinkholes" for both scenarios.

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Previous research in the region can be divided to before and after the 2016 event. Cerema (2014) mapped the karst hazard, assessing five criteria contributing to sinkhole formation (proximity to the Loire River, density of sinkholes, thickness of alluvium, presence of paleochannels, and the presence of drainage axes in the Beauce Aquifer). Perrin et al. (2015) produced a karst susceptibility map applying the statistical "weight-of-evidence theory" method using four factors (thickness of low permeability and saturated layers, distance to streams, and alluvium thickness). Their resulting map indicated that nearly the entire region falls into high hazard zones, challenging its usability in land use planning, management and decision-making. Following the 2016 flood and sinkhole collapses event, Noury et al. (2018) and Luu et al. (2019) focused on simulating the internal soil erosion during floods (including the 2016 flood in Val d'Orléans). Noury et al. (2018) found that intense rainfall and floods significantly increased sinkhole formation in this region with rates up to 24,000 times higher than normal. In a subsequent study aiming to understand how different erosion regimes accelerated internal erosion during the 2016 flood, which led to various sinkhole shapes (hourglass and inverted bowl sinkholes), Luu et al. (2019) applied different hydraulic pressures, underground conduit sizes, and the particle cohesion of granular media. They identified two distinct erosion processes: the upward growth of cavities (resulting in dropout sinkholes) and the downward flow of granules (resulting in subsidence sinkholes) in cohesive soil.

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In the context of increasing momentum and demand for a shift from single layer multi-hazard approaches to more comprehensive multi-hazard assessments that consider hazard interrelationships (e.g., see Kappes et al., 2012; Gill and Malamud, 2014; Ward et al., 2022), this study has three main objectives: (1) to enhance the recognition of karst terrains as multi-hazard forming area, (2) to advocate for a shift in multi-hazard assessment methodologies, moving from a multi-layer single hazards to a multi-hazard approach (see Gill and Malamud, 2014; Gill and Malamud, 2016), considering the interconnected effects between floods and sinkholes, and (3) to explore the applicability of the MaxEnt model for representing a multi-hazard susceptibility scenario in a karst environment, specifically flood-triggered sinkholes to evaluate the potential for the MaxEnt model to be used in multi-hazard mapping in the context of flood-triggered sinkholes.

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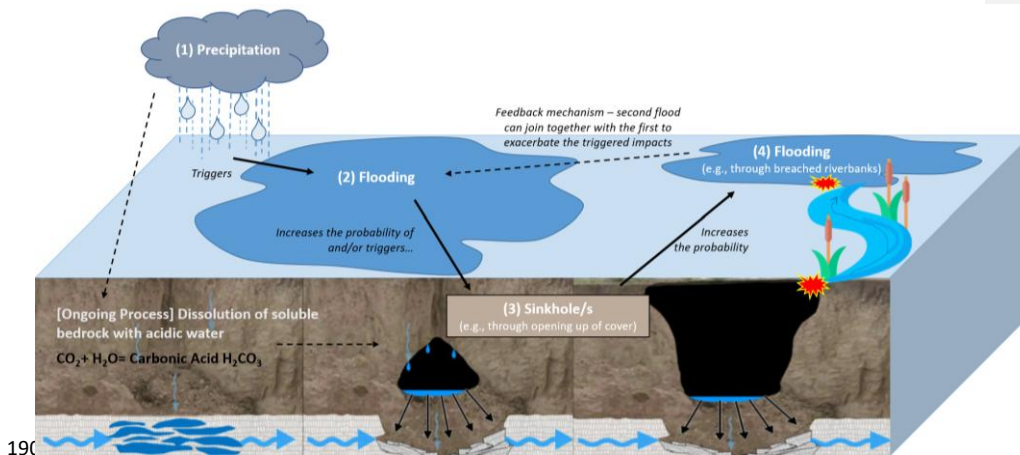
155

160 **2 The multi-hazard environment of karst terrains**~~Karst terrains as a multi-hazard-forming environment~~

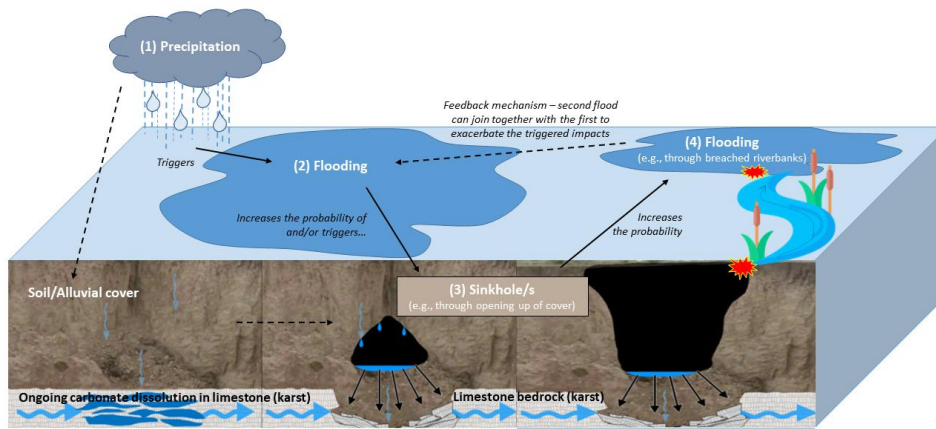
2.1 Exploring a potential multi-hazard scenario

165 Given the definition of the multi-hazard term as “the specific contexts where hazardous events may occur simultaneously, cascadingly or cumulatively over time” (UNDRR, 2017), multi-hazard forming environments are where different geophysical processes have the potential to overlap in various ways (adopted from UNDRR, 2017). It can, therefore, be argued that karst landscapes, covering about 20 to 25% of the land surface (Waltham et al., 2005), exhibit such properties. In karst terrains, meteorological events could act as a trigger to activate a cascade of hazards. The existing body of literature extensively explores the impact of severe rainfall and flood events on sinkhole collapses (Hyatt and Jacobs, 1996; 170 Brinkmann and Parise, 2008; Martinotti et al., 2017; Parise et al., 2018; Noury et al., 2018; Xiao et al., 2018; Luu et al., 2019; Kwak et al., 2020). For instance, the flooding of the Flint River in 1994 triggered 312 sinkholes in the karst Dougherty Plain at Albany, Georgia, with 88% of these collapses occurring within the spatial extent affected by flooding (Hyatt and Jacobs, 1996). In Florida, the tropical storm Debby in June 2012 caused a prolonged rainfall event followed by a period of drought. The flood caused 175 by the storm triggered hundreds of sinkholes collapsing, damaging infrastructure, buildings and forcing the inhabitants to evacuate their homes (Brinkmann, 2013). Similarly, during a 2012 heavy rainfall in Liuzhou city of Guangxi province, south China, a cascade of sinkhole collapse events occurred with an initial sinkhole collapse resulting in 37 further sinkholes within a span of 5 days (Lei et al., 2016).

180 In a karst terrain (**Fig. 1**), the slightly acidic rainwater stored in subterranean soluble rocks percolate down through the existing cracks. The exposure of groundwater ~~and to~~ easily eroded rocks (carbonate and evaporate rocks) results in the rock dissolution creating small cavities or voids deep underground (Lewin and Woodward, 2009; Xiao et al., 2016). Over time, those cavities grow larger and the overlying surficial soils move downwards to fill in the cavities, resulting in upward/downward erosion of soil 185 particles beginning from ~~bottom~~base/above of the overlying surficial soils (to know more about this process please see Noury et al., 2018; Luu et al., 2019). Eventually, when surface soils fall into the subterranean voids due to a loss of compaction, a sinkhole, the most common hazard in karst areas, occurs (Waltham et al., 2005).



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195 **Figure 1.** A simplified sequential diagram of the potential multi-hazard interactions in karst covered terrains. The diagram illustrates the sequence of events starting with (1) Precipitation, which triggers (2) Flooding (solid arrow). Flooding increases the likelihood of or trigger (3) Sinkhole formation due to the dissolution of soluble bedrock by acidic water (dashed arrow from precipitation to dissolution, an ongoing

process over time). Sinkhole collapses can then lead to (4) Secondary flooding (e.g., through breached
200 riverbanks, indicated by the solid arrow from sinkholes to flooding). The dashed bidirectional arrow
between stages (4) and (2)—shows the feedback mechanism where initial flooding can exacerbate
subsequent flooding events. The arrows underneath the sinkholes depict the flow of water into
underground cavities, contributing to the dissolution process and potential further collapse.

205 As illustrated in **Fig. 1**, intense rainfall and flooding can apply load to underground cavities that are
close to the surface through saturating and liquefying overburden (Hyatt and Jacobs, 1996) and can
accelerate the internal erosion processes by a sudden massive infiltration of water (Noury et al., 2018;
Luu et al., 2019). The failure happens when the applied stress by flood or hydraulic load is sufficient to
210 overcome the resisting forces of the material covering up the cavities structure. Additionally, in a
hypothetical scenario, where multiple cavities exist close to each other, the collapse of one could cause
a cascade effect for others, creating bigger sinkhole collapses. Depending on the location of sinkholes,
the collapses can increase the probability for another hazard. For example, if the collapse of a sinkhole
breaks a river's dykes, another flood could be triggered, adding to the hazard chain. In multi-hazard
literature, this relationship is called a bi-directional relationship or feedback mechanism where a
215 primary hazard can produce itself (Gill and Malamud, 2016; Ciurean et al., 2018). Furthermore, this
chain of hazards can potentially continue if debris produced by a flood blocks the sinkholes' conduits,
resulting in water not being able to infiltrate into the ground at sufficient speed, and therefore, causing
flooding (Zhou, 2007; Xiao et al., 2016).

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2.2 Study area: the region of Orléans (Val d'Orléans)

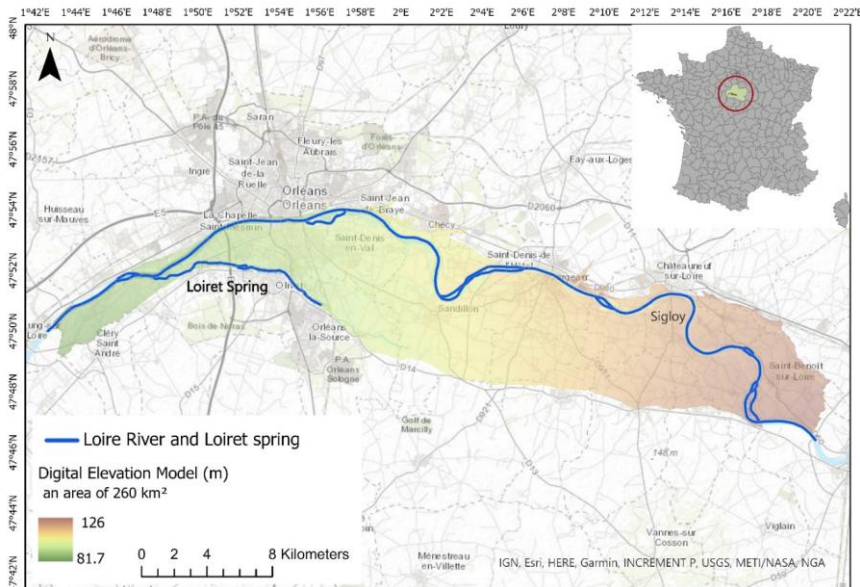
Val d'Orléans is located in the Loiret department, within the Centre-Val de Loire region of France. The
geography of Val d'Orléans is characterized by an extended basin that stretches for approximately 36
km from east to west and spans 7 km from north to south (Blanchard, 1903) (**Fig 2**). This basin is
225 bordered by the Forest of Orléans to the north and the plateau of Sologne to the south. The Loire River,
the longest river in France with a length of 1012 km, (Auterives et al., 2014), meanders with sinuous
curves through the region. It extends over a distance of 33 km, from Sigloy in the east to the confluence
of the Loire-Loiret in the west within this larger basin. The Val d'Orléans is a large depression and an
alluvial plain within the major bed of the Loire River (Desprez, 1967).

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The bedrock in this region is composed of highly fissured karstic limestone called Beauce limestone
("Calcaire de Beauce"), a carbonate lacustrine deposit, with a thickness ranging from 50 to 90 meters
and situated at a depth of up to 10 meters below the surface (Jozja et al., 2010). This geological structure
is overlaid with alluvial sediments originating from the Loire River. The intensive karstification of the
235 Beauce limestone in this region is attributed to the subterranean water flows of the Loire. Consequently,
the formation of voids frequently results in the appearance of sinkholes and subsidence on the valley's
surface (Cerema, 2014). The process of karstification in the region likely started during the Würm

glacial period (approximately 115,000 to 11,700 years ago) when the recession of the sea triggered increased erosion activities; leading to the gradual erosion of the less permeable sediments covering the limestone in Val d'Orléans. This erosion event caused a significant influx of water from the Loire River and its tributaries into the limestone formations (Perrin et al., 2015). The karst aquifer overlain by the Quaternary Alluvium alluvial of the Loire River (Auterives et al., 2014) is characterised by a complex geological setting, which consists of a multi-layered system of clay interbedded with limestone (Lepiller, 2006).

The region of Orléans experienced a multi-hazard event in 2016. Five consecutive days of heavy rainfall led to flooding across a 30-km area including the overflow of an artificial canal near the Loire River spreading the flood over a 1 km² area (Noury et al., 2018; Luu et al., 2019). The floodwater as a primary hazard triggered more than 100 sinkhole collapses, which were mainly attributed to karst collapses ~~which are attributed to be mainly karst collapses~~ as opposed to anthropogenic, with 12 sinkholes observed within a 1 km² area during a 10-day flood period (Noury et al., 2018; Luu et al., 2019). The flood-induced sinkholes occurred ~~happened~~ near the Loire River, breaching the river dyke, and causing another flooding.



260 **Figure 2.** Geographical location of the study area situated within the Loiret department, France. The inset map shows the location of the study area (red circle) in France.

3 Methods and datasets

265 This section outlines the application of the MaxEnt model to multi-hazard susceptibility assessment, discussing its advantages, limitations, and the steps taken to address common issues such as model variability and collinearity among variables. Additionally, we provide an overview of the data sources and variables used in our analysis.

3.1 Multi-hazard susceptibility assessment (MaxEnt application)

270 We applied the MaxEnt method to identify the spatial distribution of flood-triggered sinkholes and represent these in the form of a multi-hazard susceptibility map. Susceptibility assessment denotes the relative probability of a hazard event without reference to any specific time interval (Gutiérrez et al., 2008). As post-event sinkhole surveys often cannot provide the necessary information to precisely support hazard modelling, for example, the exact occurrence time (Parise et al., 2018), this study adopts a susceptibility assessment approach as opposed to hazard assessment.

280 Over the past decade, many studies have assessed the MaxEnt model in addressing various single natural hazards including landslides (Park, 2015; [Mokhrari-Mokhtari](#) and Abedian, 2019; Liu et al., 2022), floods (Siahkamari et al., 2018; Mobley et al., 2019), gully erosion (Pournader et al., 2018; Azareh et al., 2019) and more recently, sinkholes (Bianchini et al., 2022). MaxEnt model has also been used to explore multi-hazards, including flood, landslides, wildfire and gully erosion (Javidan et al., 2021; Rusk et al., 2022). Previous work suggests that the MaxEnt model can generate outcomes useful for identifying areas that are susceptible to given natural hazards. Some key advantages of MaxEnt include: 1) a solid mathematical foundation avoiding the black box nature of many machine learning models, 2) an open-source software for public users, and 3) robust multivariate analyses enabling data iteration for enhancing the model's predictive accuracy (Phillips et al., 2006; Sillero et al., 2021). Note that in the context of this research, the term "prediction" refers to the process of forecasting the potential future trends or outcomes based on current data and models.

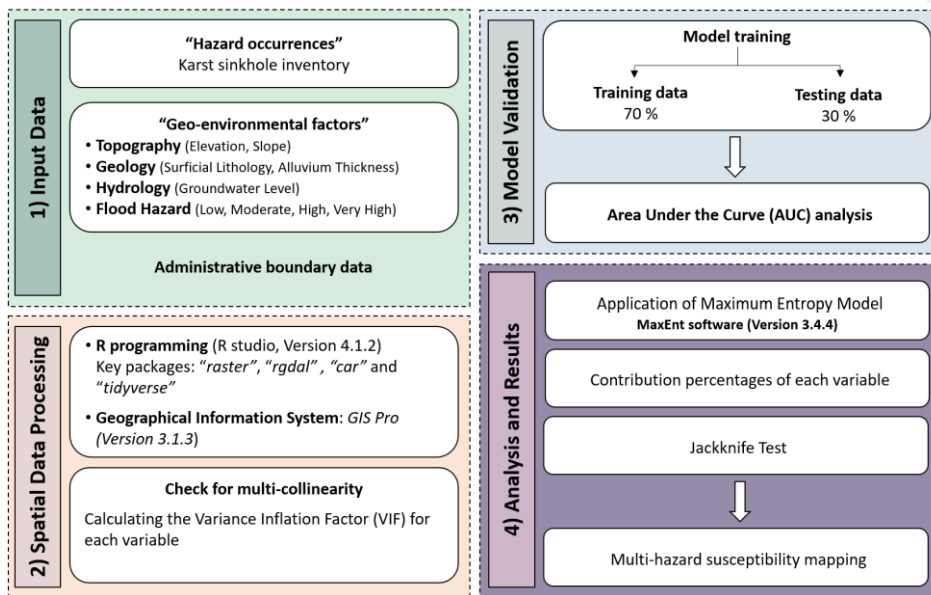
290 A notable limitation in using MaxEnt and other machine learning methods arises from the inherent variability of model outputs when the same dataset is applied repeatedly (Phillips et al., 2006; 2017; Sillero and Barbosa, 2021) and the issue of collinearity among variables used in the model (Sillero and Barbosa, 2021; Sillero et al., 2021). The former can be addressed by running the model multiple times, ideally a minimum of 10 to higher numbers like 50 or 100 times, depending on the available computer time and storage (Phillips et al., 2006; Sillero and Barbosa, 2021), on the same dataset and calculating

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the average and standard deviation of the results to evaluate if outcomes are consistent across datasets. The latter (variables collinearity) can occur when a wide range of variables is used in the model without assessing the potential collinearity among them. High collinearity among variables can lead to unstable estimates of the model parameters, making it difficult to distinguish the individual effect of each predictor on the response variable. Yet, these crucial steps are frequently overlooked in the literature. In the current research, we have addressed these common issues by both calculating the collinearity among variables and replicating the model 50 times.

A methodological flowchart (Fig. 3) summarises the workflow used to generate and analyse the MaxEnt outputs. The following sub-sections describe the MaxEnt framework, the selected contributing factors, and the application of the model for multi-hazard susceptibility assessment. A methodological flowchart, as illustrated in Fig. 3, provides a comprehensive overview of the process used to generate and explore outputs from MaxEnt. In the following sub-sections, we will explain the MaxEnt theory, discuss each selected contributing factor, and explore the application of MaxEnt in generating a multi-hazard susceptibility assessment.

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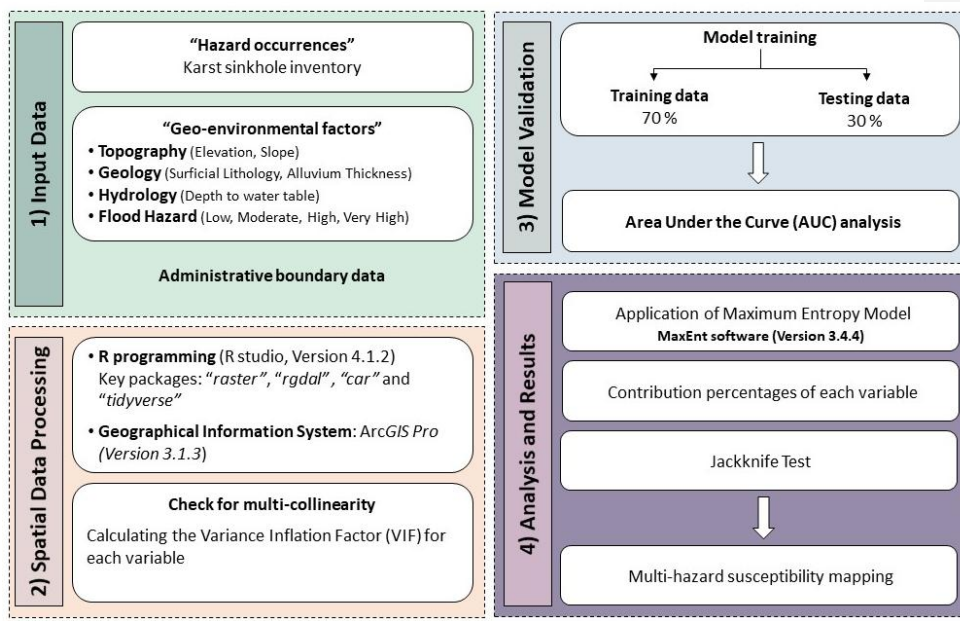


Figure 3. The methodological flowchart. The modelling process is divided into four stages: **1)** input data (data collection); **2)** spatial data processing, **3)** model validation; and **4)** model application (analysis and results).

3.2 Maximum Entropy Model (MaxEnt)

When there is incomplete and unknown information about a system's probability distribution, Jaynes (1957) suggests that the best estimation is to choose the distribution that maximizes entropy, given the constraints (i.e., environmental variables) (Jaynes, 1957). This concept is known as the maximum-entropy principle (Phillips et al., 2006). Maximum Entropy Model (MaxEnt) is a statistical-probabilistic machine learning algorithm which quantitatively estimates the probability distribution (P) of target occurrences over the set locations/ pixels (X) within the study area based on known contributing factors (Elith et al., 2011) using the Bayesian rule (Rahmati et al., 2016; Shi, 2022). MaxEnt achieves this by dividing the study area into locations or pixels X, each representing a

computing unit with a probability value ($\pi(x)$) indicating the relative chance of a target event occurrence. The distribution π allocates a positive probability $\pi(x)$ to every point x , and the sum of these probabilities equals 1.

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The response variable (y) represents whether a computing unit experiences a target event (e.g., flood-triggered sinkholes) ($y = 1$) or not ($y = 0$). To calculate the probability of a given event occurring at a particular point ($P(y=1|x)$), Rahmati et al (2016) summarises the main aspect of the model extracted from Phillips et al. (2006) and Elith et al. (2011) as follows:

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$$P(y = 1|x) = \frac{P(y=1)P(x|y = 1)}{P(x)} = \frac{P(y=1)\Phi(x)}{\frac{1}{|x|}} \quad (1)$$

Where $P(y = 1|x)$ is the probability of a hazard at a specific point, $P(y = 1)$ is the probability of event y being equal to 1, without considering any other variables. It is the marginal probability of variable y taking the specific value 1. $P(x | y = 1)$ is the conditional probability of event x occurring, given that event y is equal to 1. $P(x)$ is the probability of event x occurring without considering any other variables. It is the marginal probability of variable x . The MaxEnt algorithm estimates $\Phi(x)$ as equivalent to a Gibbs probability distribution. The Gibbs probability distribution is denoted as Eq. (2):

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$$q\lambda(\mathbf{x}) = \frac{1}{Z\lambda(x)} \exp\left(\sum_{i=1}^n \lambda_i f_i(\mathbf{x})\right) \quad (2)$$

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$Z\lambda(x)$ and λ_i represent a normalization constant (ensuring $q\lambda(x)$ sums to one across the study area) and the vector of weights assigned to the features, respectively. During the estimation phase of $q\lambda(x)$, MaxEnt modeling seeks to pinpoint the distribution that follows closely to the constraints. To prevent overfitting, it utilizes l_1 regularization. Therefore, the MaxEnt model aims to discover the Gibbs probability distribution that maximizes penalized log-likelihood values.

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Additionally, if there are m occurrences in the study area, the difference between regularization and log-likelihood, which should be maximized, is denoted as:

$$\Psi(\lambda) = 1/m \sum_{i=1}^m \ln(q\lambda(x_i)) - \sum_{j=1}^n \beta_j |\lambda_j| \quad (3)$$

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Where β_j represents the regularization parameter for the j^{th} feature (f_j). All input conditioning factors are treated as random variables in the model, following the MaxEnt algorithm.

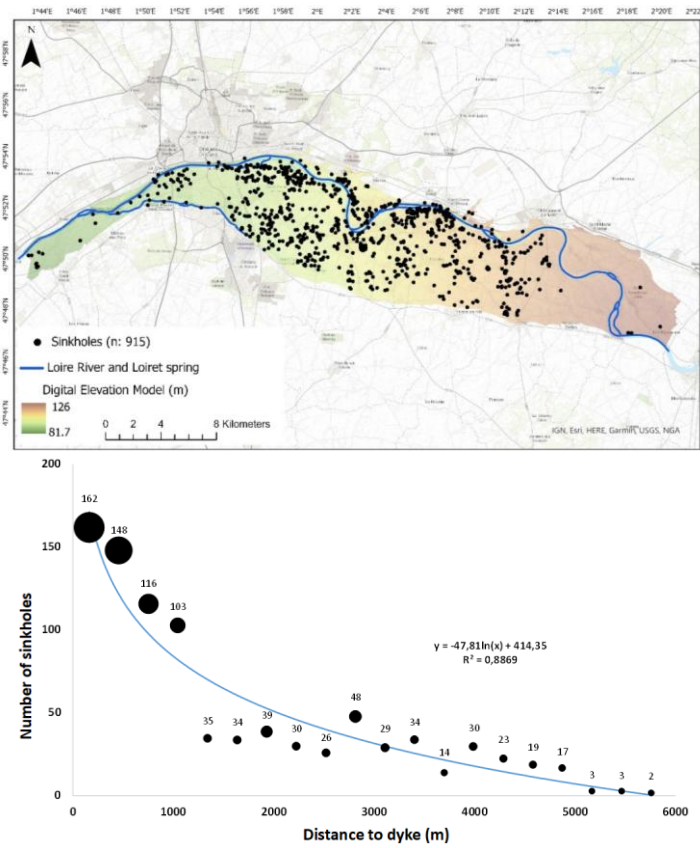
3.3 Natural sinkhole inventory

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The first step in sinkhole-related hazard mapping is to create a sinkhole inventory map showing the locations of the previously identified sinkholes (Kim and Namet al., 2018). In France, sinkhole events under the title of ground movements (BD Cavit , BDMVT database) are recorded by the French Geological Survey (BRGM) and are publicly available on the national cavity database website

(www.georisques.gouv.fr/, last access: 5 June 2024). In this study, 915 sinkhole locations with natural origins are considered. Even though it is unclear how many are triggered by floods, the location of sinkholes suggests that about 58% of sinkholes have occurred within approximately one kilometre of the Loire River (Fig. 4). This observation raises concern about the occurrence of more collapses in case of a flood event.

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380 **Figure 4.** (Top) The location of past sinkhole events shown in black circles (915 events). (Bottom)
Number of sinkholes and their distance from Loire River's dyke.

3.4 Geo-environmental factors for flood-triggered sinkholes

385 The process of selection of geo- environmental factors was driven by the principle of parsimony, aiming
to strike a balance between model simplicity and explanatory power. The concept of parsimonious
models, rooted in Occam's razor, emphasizes the importance of using the fewest necessary elements
to explain a phenomenon effectively (Baker, 2003; Shatz, 2019). We selected six primary factors (**Table
2**), with the ability to provide substantial explanatory and predictive capabilities while minimizing
unnecessary complexity. These factors are topography (elevation and slope), geology (surficial geology
and alluvium thickness) and hydrology (~~groundwater level~~[depth to water table](#) and flood hazard zones).
390 The selection process was informed by previous studies including geotechnical work carried out in the
Val d'Orléans area (Cerema, 2014; Noury et al., 2018; Luu et al., 2019). For each factor, a map was
generated using Arc GIS Pro 3.1.3 software (**Fig. 5**). Below, we explain each factor and the datasets
used in map generation.

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Table 2. Geo- environmental factors included in the MaxEnt model and data sources.

Variables	Data Source
Digital Elevation Model	BD ALTI® (géoservices.ign.fr)
Slope (m)	Extracted from DEM layer using GIS Pro
Surficial geology	BD – Charm-50, the geological map (1/50,000), (BRGM)
Depth to water table Groundwater (m)	Groundwater national portal (ades.eaufrance.fr)
Flood hazard zones (floodwater height for a 100-year return period)	Plan de Prévention des Risques d'Inondation (PPRi, 2015)
Alluvium thickness (m)	Extracted from Perrin et al. (2015)

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3.4.1 Topography (Elevation and Slope)

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~~Both floods and sinkholes typically occur in low-elevation areas. Low-elevation karst landscapes facilitate water accumulation during intense rainfall events. This may result in a reduced depth to water table (i.e., a rise in groundwater level), which can increase hydraulic pressures within subsurface voids and potentially enhance instability. Both flood and sinkholes typically occur in low-elevation areas.~~
410 ~~Low elevation karst landscapes facilitate the overflow of water during intense rainfall events. This overflow can lead to higher groundwater levels, which may accelerate karst maturation processes and increase pressure on subsurface voids~~ (Kovačič and Ravbar, 2008). These combined effects can increase the risk of sinkhole formation and collapse during flooding events. The Elevation data used in this study were obtained from the Digital Elevation Model (DEM) available on the national website BD ALTI@
415 (géoservices.ign.fr, last access: 5 June 2024). The region of Orléans is a relatively low-lying region with the elevation ranging from nearly 82 to 126 m (Fig. 5a). The elevation decreases from east to west in the area where more sinkholes are concentrated. The terrain is characterised by a low to moderate slope gradient (0-71 degrees), with noticeable variation observed alongside the river, where the slope becomes relatively steeper (above 20 degrees) (Fig. 5b).

420

3.4.2 Surficial geology

In most karst areas, sinkholes do not result from the collapse of the soluble bedrock. Instead, they predominantly occur due to ground failure where soil rapidly washes into bedrock cavities that have
425 developed over geological timescales (Waltham et al., 2005). For this reason, we used surficial geological data instead of the underlying bedrock (limestone) in our analysis. The spatial data were extracted from the geological map of the area, published by BRGM at 1/50,000 scale, using the BD-charm-50 database available on the BRGM official website (infoterre.brgm.fr/, last access date: 5 June 2024) (Fig. 5c). The area is mainly covered by ~~mostly~~ permeable Holocene and Pleistocene alluvium.

430

3.4.3 Alluvium thickness

The accumulation of river sediments over time contributes to the increased thickness of alluvium in the central part of the region (2°2'-2°10'E), which also has a low elevation (81 m to 95 m). In these areas, the Beauce limestone is covered by Quaternary alluvial deposits with a thickness of up to 20 m. Noury
435 et al. (2018) note that sinkholes in this region mainly occur due to suffosion (i.e., down washing of the cover material). Therefore, it is reasonable to expect that the thinner the alluvium is, the more likely that cavities are exposed at the surface. Originally generated by BRGM, the alluvium thickness data were extracted from Perrin et al. (2015) by applying the Inverse Distance Weighting (IDW) method to create a raster map (Fig. 5d). Sinkholes in the area mainly occur in the alluvial deposits ranging from 1
440 to 5 m in diameter (median 2.8 m) and between 1 and 4 m (median 2 m) in depth (Perrin et al., 2015).

3.4.4 ~~Depth to water table~~Groundwater level

~~Subsurface cavities in karst systems may enlarge due to groundwater circulation, which enhances dissolution processes in soluble bedrock (Pazzi et al., 2018; Intrieri et al., 2018; Xiao et al., 2018). Variations in hydraulic head can influence internal erosion processes and the mechanical stability of overlying materials (Kim et al., 2018). Subsurface cavities, when exposed to groundwater flow, can enlarge and migrate towards the surface, driven by groundwater circulation, which facilitates the dissolution of specific rock formations (Pazzi et al., 2018). In karst regions underlain by soluble bedrock, continuous exposure to both subterranean and surface water can enhance the process of dissolution (Intrieri et al., 2018; Xiao et al., 2018). In this study, we used depth to water table (m below ground surface, m b.g.s.) as a proxy for groundwater conditions. Depth to water represents the vertical distance between the ground surface and the water table. Smaller values indicate shallower groundwater conditions.~~

~~In Val d'Orléans, seasonal hydraulic head differences are relatively limited (approximately 1 m; Desprez, 1967). Therefore, we used the annual average depth to water over an 11-year period (2012–2023) to represent typical groundwater conditions in the area.~~

~~Groundwater data were obtained from the national groundwater portal (ades.eaufrance.fr, last access: 23 April 2023). Five active piezometers are located within the study area. To improve spatial representation, five additional nearby stations were included. Depth-to-water values were interpolated using the Inverse Distance Weighting (IDW) method in ArcGIS Pro (Fig. 5e).~~

~~When the level of groundwater increases, it can lead to soil erosion by weakening the soil structure and facilitating the process of soil structure ravelling (Kim and Nam, 2018). Therefore, the level and the variation in piezometric levels are among the main factors to assess when evaluating the flood triggered sinkhole susceptibility. However, in the case of Val d'Orléans the hydraulic head difference measured in warm and cold seasons is about one meter (see Desprez, 1967). Therefore, in this study, we have considered the annual average of available depth to water data over an 11-year period (2012–2023).~~

~~The groundwater data were obtained from the groundwater national portal (ades.eaufrance.fr, last access: 23 April 2023). According to this portal, there are five active piezometer water wells in the Val d'Orléans. To enhance the accuracy of the interpolation for measuring groundwater levels, the five closest stations near the north, west, and northwest boundaries of the study area were selected. To interpolate and rasterize the data, we used Inverse Distance Weighting (IDW) in Arc GIS Pro (Fig. 5e).~~

3.4.5 Flood hazard map

We used the official flood hazard map for a 100-return period produced by PPRi (*Plan de Prévention des Risques d'Inondation*, 2015) (Fig. 5f). The map has been built based on topographical data from the

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airborne laser digital terrain model (i.e., represents only the bare ground surface) and the analysis of past floods for frequency and intensity calculations. Three major floods from the 19th century (1846, 1856, and 1866) were chosen as the reference floods (PPRi, 2015). The hazard zones represent the floodwater height. There are four hazard zones, each separately incorporated into the MaxEnt model:
485 Zone 1: Low (0 to 0.5 m), Zone 2: Moderate (0.5 to 1 m), Zone 3: High (1 to 2.5 m), and Zone 4: Very high (greater than 2.5 m).

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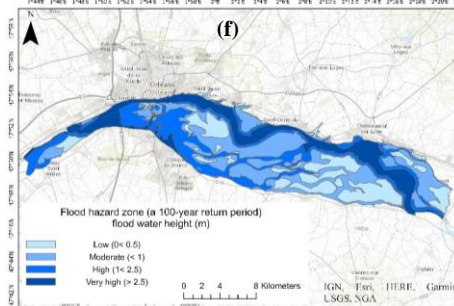
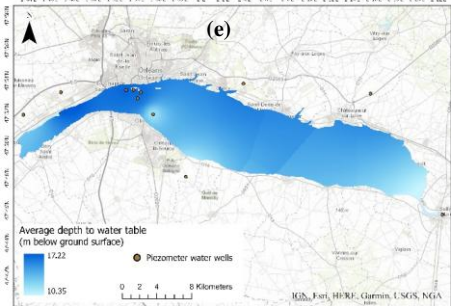
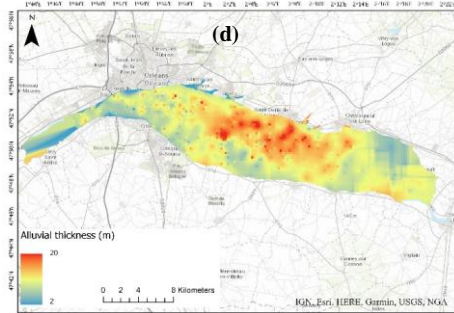
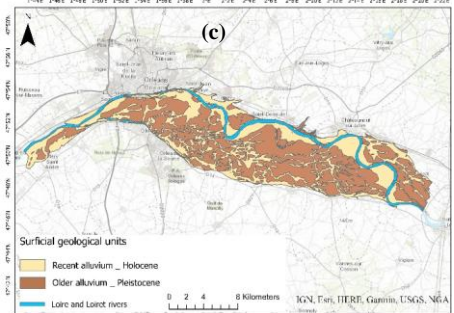
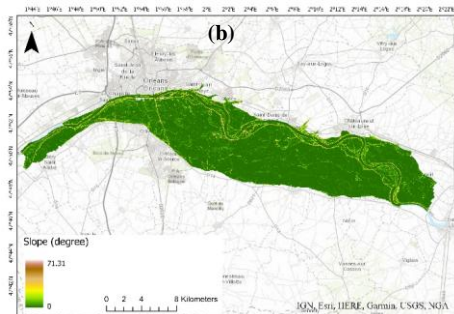
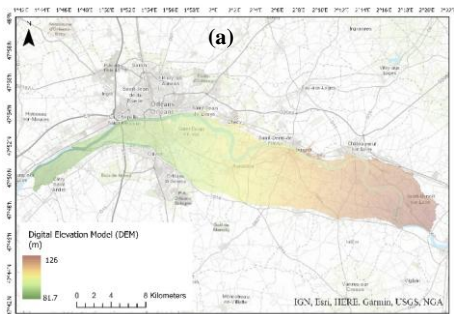
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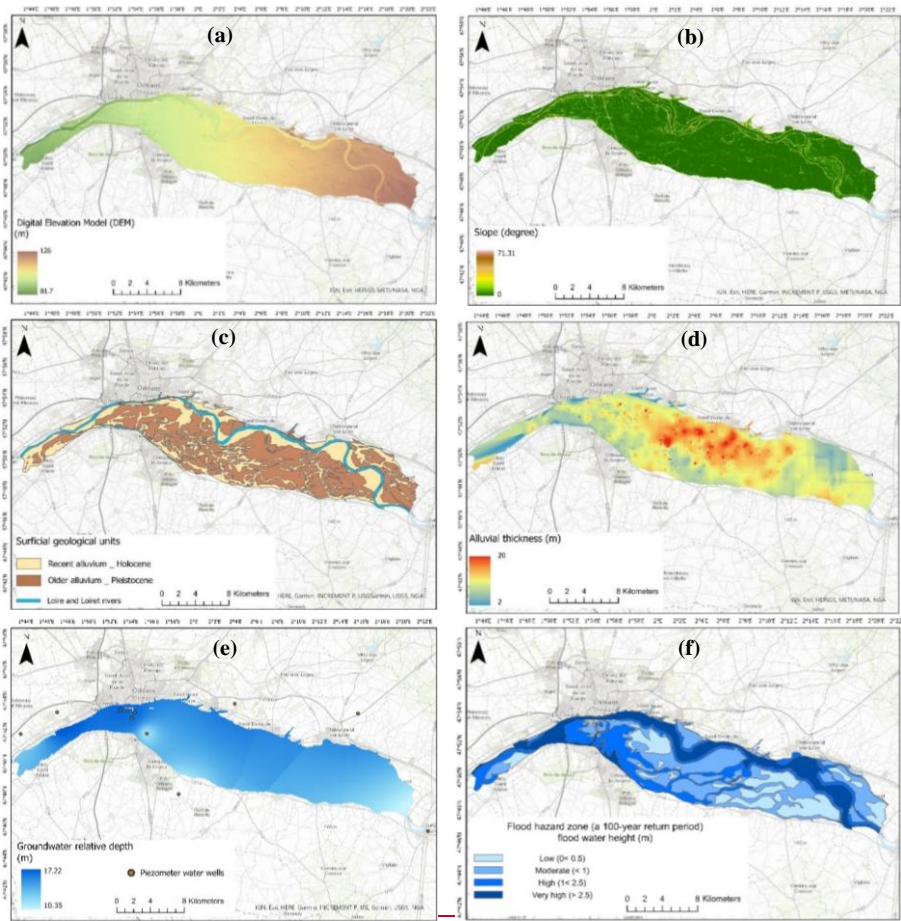


Figure 5. The six thematic maps of geo-environmental factors for assessing the spatial distribution of flood-triggered sinkholes over the study are; **a)** Elevation: showing a low-lying area with elevations ranging from 81.7 m to 126 m; **b)** Slope: depicting a degree of slopes from 0 to 71; **c)** Surficial geology: consisting of new and old alluvium layers; **d)** Alluvium thickness: Quaternary alluvial deposits with a thickness up to 20 m; **e)** Average depth to water table (m below ground surface): interpolated annual mean values based on 11 years of piezometric data (2012–2023), ranging from approximately 10 to 17 m; ~~undwater level: the annual average level of groundwater based on 11 years of data; ranging from about 10 to 17 m;~~ **f)** Official flood hazard map: classified into four levels of hazard (low, moderate, high and very high) (PPRi, 2015).

3.5 Data processing and parameter setting

To prepare the input data, we used MS Excel®, Arc GIS Pro 3.1.3, and R Studio program version 4.1.2. MaxEnt software version 3.4.4 was then used to generate the model and the susceptibility map. Using R Studio, we rasterized all the geo-environmental layers so that they have the same geographic boundary, cell size (10-meter pixel), and map projection (French coordinate system RGF 1993 Lambert-93). The collinearity test was then applied to all the rasterised layers by calculating Variance Inflation Factor (VIF) using the ‘car’ package in R Studio. As discussed in **Sect. 3.1**, many modelling algorithms are affected by high correlation among predictor variables, with the following consequences: i) over-fitted results, and ii) dependent response curves (Sillero and Barbosa, 2021; Sillero et al., 2021). The VIF measures the extent to which each variable is correlated with a combination of all other variables within the model (Sillero et al., 2021). According to Chatterjee and Hadi (2013), VIF values are interpreted as shown below:

- **VIF = 1:** There is no correlation between the variable in question and the other variables.
- **1 < VIF < 5:** This range suggests a moderate correlation among predictors, which is generally not problematic.
- **VIF >= 5:** A VIF value in this range signals a high level of multi-collinearity, potentially problematic, and might require further analysis or corrective measures.

The VIF for each variable is less than 1.50 (**Table 3**), suggesting there is no significant multi-collinearity among the variables.

Table 3. The VIF matrix for the six variables calculated by the ‘car’ package in R.

Variable	Elevation	Slope	Surficial Geology	Alluvium Thickness	Groundwater	Flood Zones
----------	-----------	-------	-------------------	--------------------	-------------	-------------

					<u>Level</u> <u>Depth to</u> <u>water table</u>	
Elevation	1.00	1.056	1.145	1.010	1.075	1.245
Slope	1.432	1.00	1.143	1.073	1.342	1.243
Geology	1.435	1.056	1.00	1.074	1.353	1.158
Thickness	1.349	1.058	1.146	1.00	1.306	1.295
<u>Groundwater</u> <u>Depth to</u> <u>water table</u>	1.140	1.050	1.146	1.037	1.00	1.287
Flood	1.379	1.016	1.025	1.075	1.345	1.00

570 From the sinkhole occurrence dataset, we randomly allocated 70% as training data to calculate the
575 model, while the remaining 30% was set aside for independent testing of model performance, validating
the resultant models. This ratio is commonly used in machine learning methods, striking a balance
between avoiding overfitting and model accuracy (Rahmati et al., 2016; Yousefi et al., 2020; Rusk et
al., 2022), Because sinkholes are spatially clustered along the Loire River, the random split may not
fully ensure spatial independence between training and testing datasets. Therefore, the reported AUC
primarily reflects within-area discrimination rather than fully spatially independent predictive
performance. -To ensure the convergence of our model towards an optimal solution, we set the
maximum iterations parameter to 1000. This iteration number allowed the MaxEnt algorithm sufficient
opportunity to adjust its parameters iteratively, seeking to minimise prediction error and accurately
reflect the observed distribution patterns of the data. To assess the reliability, stability, and uncertainty
of predictions, we employed 50 replicates in our MaxEnt analysis considering the available
computational resources.

585 MaxEnt has four output formats (raw, cumulative, logistic and cloglog). Each has different theoretical
justifications, with the choice of selecting one being highly dependent on the specific objectives of the
analysis and data characteristics (Phillips et al., 2017). For this research, the logistic output was chosen
which provides interpretable probability values (ranging from 0 to 1). This choice allows for the setting
of thresholds that are consistent with the binary nature of presence/absence data, making the threshold
settings clear and directly related to the probability values. Additionally, MaxEnt controls model
complexity and prevents overfitting through regularization (Radosavljevic and Anderson, et al., 2014).
590 For an in-depth understanding of the different tuning processes, refer to Radosavljevic and Anderson,
et al.(2014).

4 Results

595 In this section, we present the outcomes of the MaxEnt modeling analysis, which include assessing
model performance using the Receiver Operating Characteristic (ROC) test, generating a forecasted

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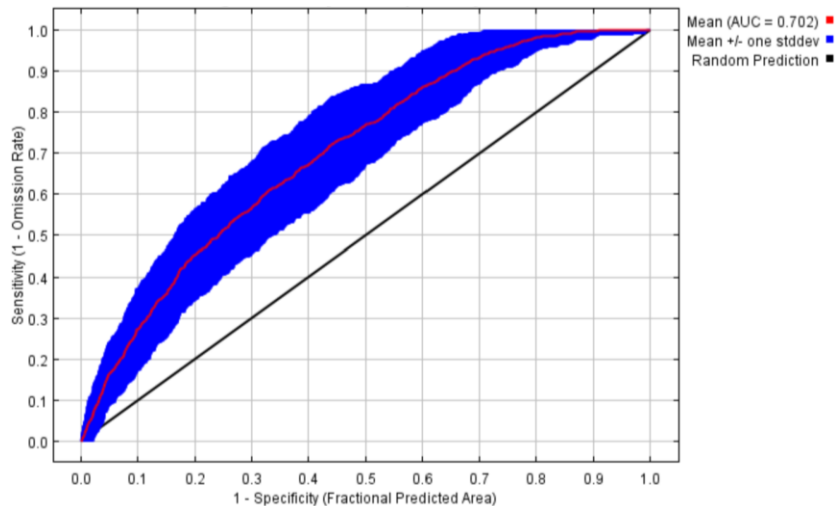
multi-hazard susceptibility map, and producing response curves for six flood-triggered sinkhole susceptibility factors. Additionally, we assess the importance of each factor using a jackknife test. Finally, we compare the existing flood hazard data with the resultant multi-hazard susceptibility map.

4.1 ~~Model performance assessment~~ **Model validation**

The Receiver Operating Characteristic (ROC) graph is a commonly used method for quantitatively assessing the performance (discriminatory power) of a diagnostic test. To assess the model's discriminatory power, we examine the relative trade-offs between the true positive rate (Sensitivity) and the false positive rate (1-Specificity) across all potential classification thresholds (Fawcett, 2006). This relationship is captured by the Area Under the Curve (AUC) metric, which ranges from 0 to 1. An AUC value of 0.5 signifies a model's performance that is no better than random discrimination, serving as a critical benchmark. Thus, any AUC value above 0.5 indicates a model that discriminates between presence and absence more effectively than random guessing, demonstrating its predictive accuracy (Fawcett, 2006; Phillips et al., 2006; Sillero et al., 2021). The reported AUC (Fig. 6) corresponds to the mean test AUC (30% withheld presence data) averaged over 50 replicate runs. In MaxEnt, AUC measures discrimination between presence locations and randomly sampled background (pseudo-absence) points rather than true presence-absence classification.

~~Figure 6 shows the ROC curve for flood-triggered sinkholes, plotted using the validation data, which includes both training and testing datasets.~~ The average test AUC ~~over 50 MaxEnt replicate runs~~ is 0.702, ~~with a~~ (standard deviation ~~of~~ 0.055), indicating a satisfactory predictive performance of the model (to better understand the threshold performance, see Fawcett, 2006 and Phillips et al., 2006). This AUC value reflects moderate discrimination ability; therefore, the results should be interpreted in terms of relative susceptibility patterns rather than deterministic prediction.

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625 **Figure 6.** Receiver Operating Characteristic (ROC) curve and Area Under the Curve (AUC) results based on the test dataset (30% withheld presence records) averaged over 50 MaxEnt replicate runs. The black diagonal line represents random discrimination (AUC = 0.5). The red curve shows the mean ROC curve (mean test AUC = 0.702). The shaded blue area represents \pm one standard deviation across the 50 replicate runs. The Area Under the Curve (AUC) results of the Receiver Operating Characteristic (ROC) analysis applied to both training and testing data. The black diagonal line represents a random chance. 630 The red curve represents the mean response of the 50 replicate MaxEnt runs which is 0.702. The shaded blue area around the mean ROC curve represents the variability of the AUC (\pm one standard deviation).

635 4.2 Multi-hazard susceptibility mapping

Based on MaxEnt logistic output, the calculated value of the model ranges from 0 to 1, with higher values indicating a higher probability of multi-hazard susceptibility. The map generated by MaxEnt model was reclassified into four susceptible zones: “Low”, “Moderate”, “High”, and “Very high” using the Natural Breaks (Jenks optimization) method, in ArcGIS Pro 3.1.3. Considering the actual distribution in the hazard data, this method, which is commonly used for hazard data classification, proves useful for datasets that are not evenly distributed. It is particularly effective in identifying natural groupings of similar values, highlighting outliers (Herries, 2018).

645 The classification of the multi-hazard susceptibility map into four distinct classes aligns strategically with the map classification guidelines outlined in the Plan de Prévention des Risques d'Inondation (PPRi, 2015), the official French document for flood risk prevention planning. Figure 7 illustrates the potential spatial distribution of the flood-triggered sinkholes across Val d'Orléans, presented as a raster surface with a resolution of 10 m per cell. The red-coloured areas denote a 'very high' level of susceptibility, covering 21.7% of the region with the relative probability distribution quantified from (0.53 - 0.90). Additionally, the distribution of susceptibility levels includes 'high' in orange (33.5%, probability 0.39 - 0.52), 'moderate' in light green (21.6%, probability 0.20 - 0.38), and 'low' in dark green (23.2%, probability 0.00 - 0.19). Most of the area is classified as High to Very high (~ 55 %, about 146 km²) where the majority of sinkholes are located. The western and most of the eastern parts are classified as Low to Moderate.

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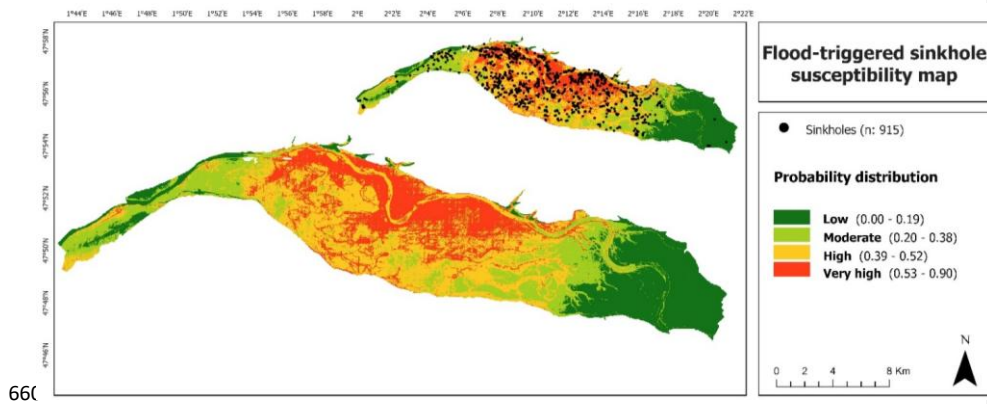


Figure 7. Classified multi-hazard susceptibility map generated by MaxEnt model using the natural breaks method. The map on top is the same map overlaid with the sinkhole locations. Dark green indicates comparatively low hazard susceptibility, with a relative index of probability distribution ranging from 0.00 to 0.19; Light green corresponds to a moderate hazard level ranging from 0.29 to 0.38. The orange and red colors represent the high and very high susceptible areas with probability distribution of 0.39 to 0.52 and 0.53 to 0.90, respectively.

4.1 Model-specific response curves and variable contributions

670 Model-specific response curves (Fig. 8) are generated from MaxEnt models using one variable at a time to capture the influence of that variable in the presence of its correlations with other variables. The probability distribution of the flood-triggered sinkhole occurrences across six contributing factors can be summarised as follows.

- 675 • The **elevation** response curve (Fig. 8a) shows low predicted susceptibility at the lowest elevations, with a clear peak between 90 m and 105 m. Within the model, this range corresponds to the highest relative probability of flood-triggered sinkhole occurrence. Above 105 m, the predicted probability declines markedly and remains low beyond approximately 110 m, indicating lower modelled susceptibility at higher elevations. The elevation curve (Fig. 8a) shows that starting at the lower elevation range, there is a relatively low probability of susceptibility. There is a distinct peak in the curve at around 90 m to 105 m. This indicates the relative optimal elevation range, where the model forecasts the highest relative probability of finding flood-triggered sinkholes. After this peak, the curve drops sharply (from 105 m to 126 m), showing that the probability of the susceptibility decreases as the elevation continues to increase beyond the optimal range. As the elevation reaches the higher end of 110 m and above, the probability of presence levels off to a very low value, suggesting that flood-triggered sinkholes are very unlikely to be present at these elevations (which are the highest elevation in the area).
- 680
- 685
- 690 • The **slope** curve (Fig. 8b) exhibits a typical logistic (S-shaped) pattern, indicating an increasing probability of sinkhole occurrences with increasing slope steepness. Slopes with steepness of up to 10 degrees show a very low probability of flood-triggered sinkhole occurrences. Between approximately 10 to 50 degrees, there is a sharp, consistent increase in the probability. Between 50 and 71 degrees, the probability of sinkhole occurrence remains relatively high.
- 695 • The **geological response** histogram (Fig. 8c), presented as a categorical layer, indicates that the highest probability value corresponds to the recent alluvium (Holocene in age). The predominant surficial geological layers in the area consist of both recent and older alluvium (Holocene and Pleistocene, respectively).
- 700 • The **alluvial thickness** curve (Fig. 8d) shows a very low probability of flood-triggered sinkhole occurrences where the alluvial thickness is minimal (close to 0). Probability increases sharply up to approximately 10 m and continues to increase more gradually toward 20 m. As the alluvial thickness increases, there is a sharp and steady increase in the logistic output up to about 10 m, indicating a higher probability of susceptibility with increased thickness of alluvial deposits. Beyond this point, the curve shows a more gradual yet consistent increase in the probability of sinkhole occurrence, which continues as the alluvial thickness reaches 20 m.
- 705
- 705 • The response curve for **depth to water of groundwater level** (Fig. 8e) indicates relatively low predicted susceptibility at very shallow groundwater conditions.

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- ~~a low probability of flood-triggered sinkholes at minimal groundwater levels in Val d'Orléans. The model suggests increasing susceptibility at intermediate depth-to-water values (approximately 14-16 m below ground surface), followed by a decline at greater depths. This pattern may reflect spatial associations between groundwater conditions and other geo-environmental factors rather than a direct mechanistic threshold. Therefore, the observed peak around ~15 m should be interpreted as a statistical tendency within the model rather than evidence of a specific hydrogeological optimum for sinkhole formation. This may reflect a buoyant force that stabilizes the ground, mirroring the isostatic principles seen in geological structures. As groundwater levels rise to around 15 m, sinkhole probability sharply increases. This is likely due to erosion of underground materials. At 15.5 m, conditions are optimal for sinkhole occurrence, but beyond this level, the probability swiftly declines.~~

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- The curve of **flood hazard** zones (**Fig. 8f**) shows that as the flood hazard level increases, ~~there is a general trend of increasing probability~~, suggesting a higher probability of sinkholes occurring in areas with more severe flood hazards where the level of floodwater during a 100-year flood event could reach up to 2.5 m. This observation aligns with the empirical data, as about 58 % of the sinkholes were recorded within one kilometre from the river (**Fig.4**).

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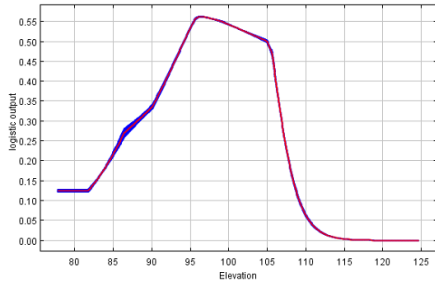
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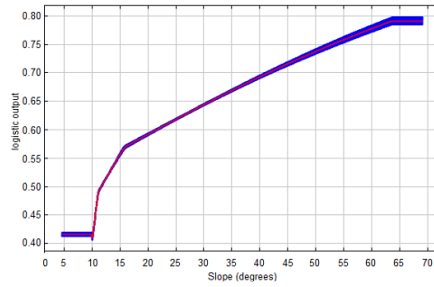
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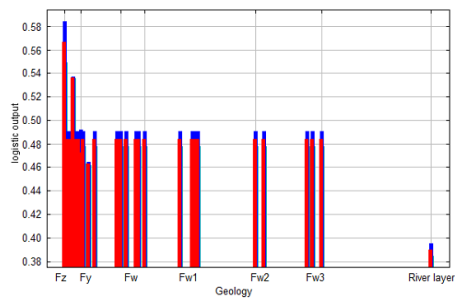
(a) Elevation (m)



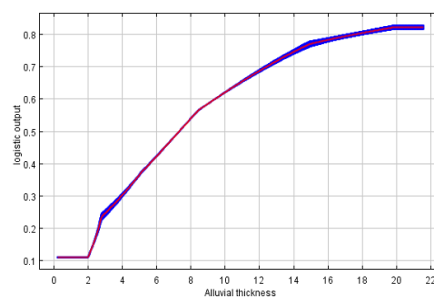
(b) Slope (degrees)



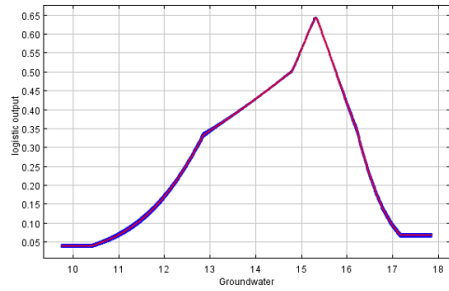
(c) Surficial geology



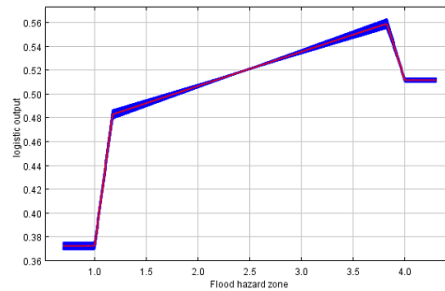
(d) Alluvial thickness (m)



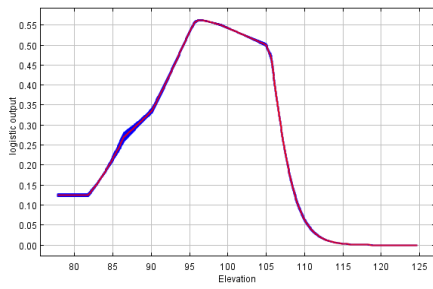
(e) Groundwater level (m)



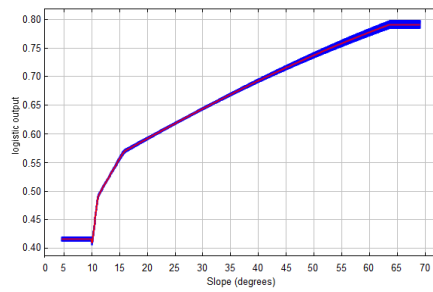
(f) Flood hazard zones



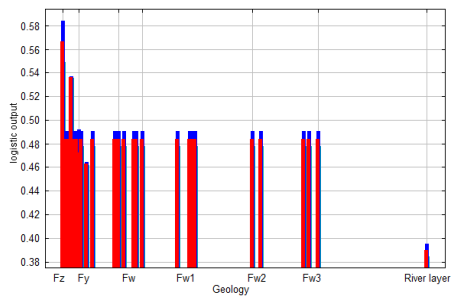
(a) Elevation (m)



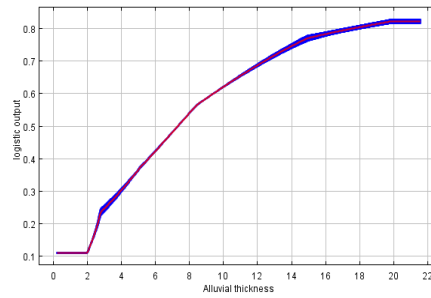
(b) Slope (degrees)



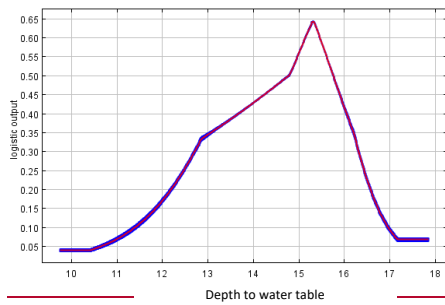
(c) Surficial geology



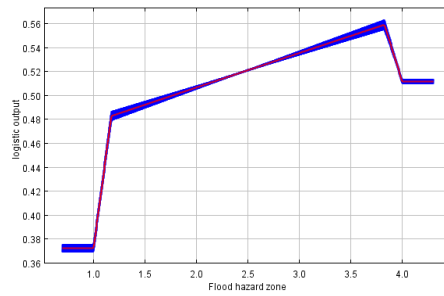
(d) Alluvial thickness (m)



(e) Depth to water table (m)



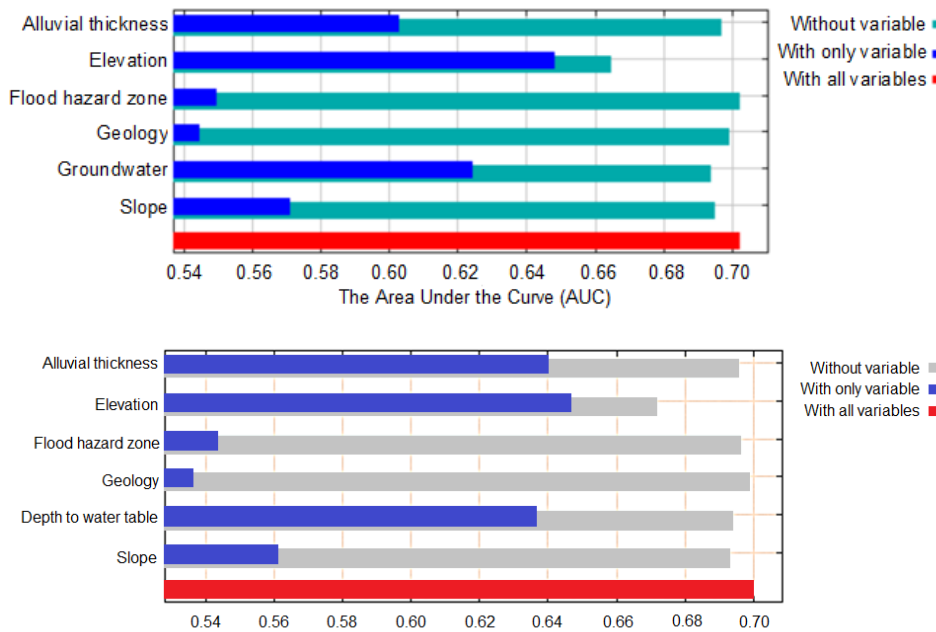
(f) Flood hazard zones



755 **Figure 8.** MaxEnt model response curves and histogram for flood-triggered sinkhole susceptibility factors. The curves and the histogram show the mean response of the 50 replicate MaxEnt runs (red) and the mean +/- one standard deviation (blue); each graph provides information of the impact of each factor and their correlations with other variables. **a) Elevation Response Curve:** Indicates a peak sinkhole occurrence probability between elevations of 90 m to 105 m. **b) Slope Response Curve:** ~~Indicates that, within the model, higher slopes correspond to higher predicted probability. Demonstrates an increasing probability of sinkholes with increasing slope values, suggesting higher susceptibility in relatively sloped areas.~~ **c) Geology Histogram:** Displays the distribution of sinkhole probability across different geological units, highlighting that the highest value corresponds to the recent alluvium deposits (Holocene) signified the letter “Fz and Fy”; the letters “Fw, Fw_{1,2,3}” are associated with layers of older alluvium deposits (Pleistocene). **d) Alluvial Thickness Response Curve:** Shows a steady increase in the suitability of the occurrences as the thickness increases. **e) Average depth to water table**~~**Groundwater Response Curve:** Shows a peak in predicted susceptibility at intermediate depth-to-water values (around 15 m below ground surface). Reveals a distinct peak at about 15 m indicating the optimal groundwater level associated with higher sinkhole probability.~~ **f) Flood Hazard Response**
770 **Curve:** Shows a positive correlation between flood hazard and sinkhole formation potential.

The jackknife test graph (**Fig. 9**) illustrates the influence of each contributing variable on the model’s AUC in two scenarios: with the variable (blue bar) and without it (green bar). The jackknife test, a resampling technique recognized for its efficacy in assessing variable importance within a model
775 (Phillips et al., 2006), systematically omits one variable at a time from the dataset to re-evaluate the model’s predictive performance (Phillips et al., 2006). ~~Results from the jackknife test reveal that elevation yields the highest AUC value when used as the sole predictor, highlighting its strong contribution to overall model performance. Results from the jackknife test reveals “Elevation” as a predominant variable, demonstrating the highest AUC value when it is the sole predictor, thus underseoring its criticality to the model’s forecasting ability.~~ The second significant AUC value belongs to ~~groundwater level~~**depth to water** followed by alluvial thickness and slope. Meanwhile, “Flood hazard zones” and “Geology”, although contributory to the overall AUC, shows a reduced impact when considered in isolation.

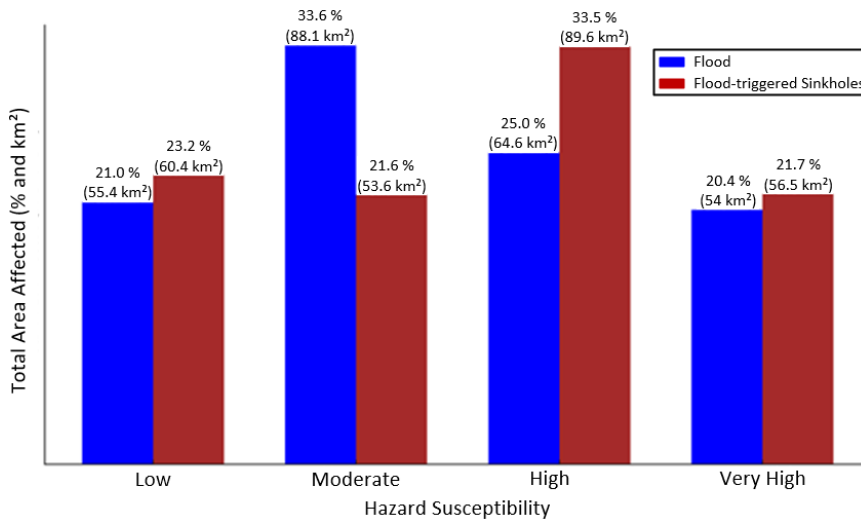
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790 **Figure 9.** The plot of the jackknife test results of the averages over 50 replicate runs contributing factor
 importance. The green-grey bar represents the AUC of the model without the corresponding variable,
 while the blue bar shows the AUC processed with only the corresponding variable. The longer the blue
 bar, the more important is the variable for model probability distribution. The red bar represents the
 AUC of the final model considering all the variables.

795 For the purpose of land use planning and natural hazard mitigation measures in Val d'Orléans, we
 compared and visualised the spatial extents affected by flooding as a single hazard and those susceptible
 to flood-induced sinkholes, representing a complex multi-hazard scenario. It is important to
 acknowledge the inherent differences in the nature and methodologies employed in hazard mapping
 800 (considering magnitude and frequency) and susceptibility assessments (spatial probability distribution).
 However, we have focused on comparing the spatial overlaps by means of a classification approach
 representing an effort towards standardisation to partially overcome the existing challenges in
 comparison of hazards within multi-hazard studies (Kappes et al., 2012). **Figure 10** shows the extent to

805 which different areas are affected by flood as a single hazard (the official map produced by PPRi, (2015), see Fig. 5f) and susceptible to sinkholes triggered by flooding as a multi-hazard (**Fig. 7**). The areas classified as having a “Low” flood hazard and flood-triggered sinkholes account for 21% and 23.2%, respectively. The “Moderate” flood hazard areas cover 33.6% of the region, whereas the areas susceptible to flood-triggered sinkholes constitute significantly less, at 21.6%. This 12% discrepancy suggests that the factors contributing to flooding do not uniformly increase the risk for sinkholes, 810 indicating that the other factors also are involved in flood-triggered sinkholes. The “High” susceptibility class shows a considerable difference between the two hazards, with a higher percentage of areas susceptible to flood-triggered sinkholes (33.5%) than to flooding (25%). For areas of “Very High” susceptibility, there is a notable spatial overlap between the flood hazard, which cover 20.4% of the area, and flood-triggered sinkholes, which cover 21.7%. This ~~proves-suggests~~ the spatial overlap of 815 these two hazards.



820 **Figure 10.** Spatially comparing areas covered by different susceptibility classes for two related yet distinct hazards in the study area: flood and flood-triggered sinkholes. The ~~black-blue~~ bars show the total areas affected by flood in both percentage and kilometre square. The ~~grey-red~~ bars illustrate the total areas affected by flood-triggered sinkholes in both percentage and kilometre square.

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5 Discussion and limitations

830 The MaxEnt method was applied to model susceptibility to flood-triggered sinkholes, achieving satisfactory discriminatory performance (mean test AUC = 0.702), achieving a satisfactory result with a predictive accuracy of 0.702 for AUC. The model's spatial distribution of relative probabilities indicates that approximately 55% of the study area (about 146 km²) falls within the High and Very high susceptibility categories. Elevation shows the highest relative contribution to overall model performance, with the 90 m to 105 m range corresponding to the highest predicted susceptibility values. Elevation appears to be the main variable in detecting areas of higher susceptibility, particularly at elevations ranging from 90 m to 105 m. Additionally, areas with groundwater levels depth to water table of about 15 m appear to be relatively more prone to sinkholes. This underlines that in the Val d'Orléans, characterized by its flood plains overlying karstic bedrock, factors such as elevation and groundwater—depth to water table levels are strongly associated with higher modelled susceptibility significantly contribute to and a greater the likelihood of sinkholes occurring during flood events. Furthermore, areas within the region that have relatively steeper slopes (above 20 degrees), greater alluvial thickness (about 10 m to 20 m), covered by recent alluvial (Holocene) units, and higher flood hazard scores, which correspond to higher water levels are more susceptible to flood-triggered sinkholes. The low predicted probability at minimal alluvial thickness is consistent with the dominant cover-collapse and cover-suffosion mechanisms documented in the Val d'Orléans, where sinkholes typically develop within alluvial deposits overlying karstified limestone (Noury et al., 2018; Luu et al., 2019). The region is largely characterised by mantled karst conditions, and the sinkhole inventory is dominated by collapse features within alluvial cover. While isolated solution sinkholes in areas of minimal cover cannot be excluded, they are not the predominant type represented in the inventory used

840 for model calibration. Therefore, the susceptibility map primarily reflects the spatial controls on cover-collapse processes in this specific geomorphological context.

855 -The jackknife results should be interpreted cautiously, as they provide relative importance values within a correlative modelling framework. The comparatively low standalone contribution of flood hazard zones should not be interpreted as evidence that flooding is unimportant in sinkhole formation. Rather, it reflects the distinction between triggering mechanisms and spatial susceptibility conditions. In the Val d'Orléans, flood-prone areas spatially overlap with low elevations, depth to water table around 15 m, and specific geological units. Consequently, much of the spatial signal associated with flood exposure is already represented by these conditioning variables. Within a static susceptibility model, triggering processes may therefore exhibit lower standalone discriminatory power, even though they remain critical in the temporal activation of collapse events.

The resultant jackknife test in MaxEnt outputs should be interpreted with caution as this test provides relative importance values that may not fully capture the multidimensional interactions among variables

865 ~~or the model's transferability to different areas and conditions. Therefore, in our modelling, flood hazard~~
~~zones and surficial geology datasets emerged as the least impactful factors compared to others. This~~
~~cannot be directly interpreted as less important for flood-triggered sinkhole susceptibility. Instead, it~~
~~may indicate that within the context of our dataset and model structure, these variables are~~
870 ~~overshadowed by more dominant variables. Moreover, this observation could be partly attributed to the~~
~~spatial overlap of flood hazard zones and surficial geology factors as indicated by their slightly moderate~~
~~correlation in the Variance Inflation Factor (VIF) test. However, as the model response histogram (Fig.~~
~~8e) and curve (Fig. 8f) show, these two factors contribute the entire model performance.~~

▲
Our findings reveal that areas with the highest susceptibility are primarily located along the Loire River
875 and close to the Loiret spring, pinpointing these zones as critical regions of concern for potential multi-
hazard occurrences. This pattern of susceptibility is underpinned by historical hydrological events;
notably, Martin et al. (2003) document a significant event in 1907 where flooding led to the Loire
River's backflow into the Loiret. Consequently, the spatial analysis delineates the historically affected
areas by this backflow between the Loire and Loiret as areas with High to Very high susceptibility.
880 Meanwhile, the western sections of the study area (approximately ranging from longitude 1°43'E to
1°54'E) are classified as having Low to Moderate susceptibility, although sporadic areas of High and
Very high susceptibility exist near the Loire River, suggesting localized areas of susceptibility. In
contrast, the eastern parts (approximately ranging from longitude 2°14'E to 2°20'E) are characteriszed
by relatively Low susceptibility (areas in dark green), largely attributed to higher elevations (about 110
885 to 126 m), ~~which are associated with lower modelled susceptibility which likely mitigate the impact of~~
~~flood and sinkhole formations.~~ The highest susceptibility probability in the most susceptible areas is a
relative probability value of 9 and the lowest susceptibility probability is a relative value of 0.0001
suggesting that the model forecast a very low, but not necessarily zero of susceptibility. This
information has relevance for policymakers and disaster risk management officials when prioritizing
890 resources and efforts for areas with the highest risk, while not neglecting the lower risk areas.

Our comparison of the existing flood hazard map and the new multi-hazard susceptibility map (this
study) reveals that areas exhibiting both Very high and High overlaps in flood hazard and flood-
triggered sinkhole susceptibility, such as those observed along the Loire River (from 1°54'E to 2°12'E),
895 require close evaluation for their potential sinkhole susceptibility. This underscores the necessity for
integrated hazard assessments and mitigation strategies that consider both surface (i.e., floods) and
subsurface (i.e., sinkholes) conditions. Furthermore, it becomes evident that areas prone to flood hazard
do not always directly correlate with those susceptible to flood-triggered sinkholes, suggesting a
spatially non-linear relationship, as seen in the eastern part (2°14'E to 2°20'E) and western part (1°43'E
900 to 1°54'E). ~~This indicates that although the two hazards are related, their spatial distributions are not~~
~~fully coincident. This may indicate that even though the two hazards are related, they do not always~~
~~happen in the same locations.~~ This information can be used to enhance emergency planning and public

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awareness campaigns, ensuring they address specific risks presented by both hazards and support efforts to advance multi-hazard understanding and frameworks in risk assessment and disaster preparedness.

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There are some limitations that introduce uncertainties into the modelling of susceptibility in this study. The quality of predictive models depends heavily on the completeness and representativeness of the sinkhole inventory (epistemic uncertainty) (Galve et al., 2009; Taheri et al., 2019). Specifically, for mapping the susceptibility of flood-triggered sinkholes, having an inventory that documents only sinkhole collapses triggered by flood events might yield more accurate results. Nevertheless, obtaining this information typically presents challenges, such as data scarcity and potential limitations in coverage. This also underscores the importance of recording sinkhole triggers along with their respective locations, itself requiring sufficient monitoring to support identification of newly triggered sinkholes. In addition, the limited elevation variability within the study area and its systematic east-west gradient mean that elevation may partly function as a spatial proxy variable in the modelling framework. Therefore, its high contribution to model performance should not be interpreted as evidence of a strictly causal elevation-sinkhole mechanism but rather as reflecting broader spatial and geomorphological patterns.

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Another limitation is that in an area characterised by a high flood hazard, particularly in the context of flood-triggered sinkholes, it becomes evident that there exist localized geological features that pose significant challenges for inclusion in modelling efforts. These localized factors are primarily related to the complex geological characteristics of the region. For instance, certain areas may exhibit distinct lines of geological weakness and fractures that facilitate the movement of water. Consequently, water is more likely to erode materials in these areas, increasing the risk of sinkhole formation. Integrating such fine-scale geological information into a model can be challenging, as it requires highly detailed geological maps with fine resolution. These features represent specific susceptibilities that may not be adequately accounted for in standard modelling approaches like MaxEnt. As a result, a more in-depth investigation, such as detailed mapping and site-specific assessments, becomes essential for gaining a comprehensive understanding of the likelihood and precise locations of flood-triggered sinkhole occurrences, especially in areas with complex geological characteristics. We also acknowledge that despite our careful and informed approach in choosing parameters for the MaxEnt model, there remains an inherent uncertainty associated with tuning the model parameter settings.

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The multi-hazard susceptibility map generated through this study may serve as a communication tool to support decision makers involved in resource allocation and risk mitigation efforts. Additionally, it can be used for raising awareness among the public about the diverse risks they may encounter, fostering a culture of preparedness and resilience within the community. While the probability of simultaneous occurrences of sinkholes triggered by flood events may be low, the consequences could be disproportionately high, requiring a proactive and preventive approach to community safety and disaster risk reduction. Since multi-hazard maps are a relatively recent addition to disaster risk management

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practices, further research is essential to seek stakeholders' perspectives on the utility of multi-hazard maps, ensuring that generated maps support decision-making processes.

945 6 Conclusions

In this study, we focus on a typical multi-hazard scenario prevalent in karst environments, specifically flood-induced sinkholes, using the Orléans region (Val d'Orléans), France, as a case study. We develop a multi-hazard susceptibility map that recognises the spatial interrelationships of flooding and sinkhole occurrences, which to our best knowledge is the first attempt in producing multi-hazard susceptibility map in karst area of Val d'Orléans and in the multi-hazard literature. This map incorporates the official flood hazard map as a factor influencing sinkhole formation/collapses in the area. Using the MaxEnt model, we find elevation (90 m to 105 m) to be the variable with the highest relative contribution the key driving factor for flood-triggered sinkholes, followed by groundwater leveldepth to water table (at or near 15 m) and alluvial thickness (10 m to 20 m). Our findings indicate that the middle parts of the study area and areas along the Loire River are highly susceptible to the combined flood/sinkhole hazard. Additionally, our comparison of existing single flood hazard map and the multi-hazard susceptibility map reveals a spatially non-linear relationship between the locations prone to flood hazards and sinkhole development. This finding indicates that while flood management is essential, it must be coupled with an understanding and anticipation of secondary hazards, such as sinkholes in karst terrains that may be triggered or exacerbated by flooding. This is particularly crucial in areas of High and Very high flood hazard zone, where the susceptibility to sinkholes is increased. This research demonstrates the applicability of MaxEnt for representing flood-triggered sinkhole susceptibility within a multi-hazard framework in a karst environment.~~This research contributes to our understanding of the application of the MaxEnt model to support characterisation of multi hazard scenarios in karst environments.~~

The application of the MaxEnt model and its 50 times replications in our study yielded results with the satisfactory forecasting outcomes meaning that the model could fitted well into the data. However, it is essential to approach MaxEnt findings with caution due to several limitations of the model at its current stage. Specifically, the model is sensitive to the extent of the study area, which can lead to inflation in the AUC score, and it is highly dependent on the choice of geo-environmental variables. MaxEnt is relatively new in the field of natural hazard and there is a need to explore this model further. Future research may benefit from comparative benchmarking against alternative modelling frameworks, implemented under consistent predictor selection and spatial resolution. Our study also provides a framework that can serve as guidance for multi-hazard scenarios in other karst environments, and motivate further work aimed at improving our forecasting of flood-triggered sinkholes.

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985 *Author contributions.* All authors contributed to the preparation of this paper. HS: conceptualization, formal analysis, visualisation, methodology, and writing (original draft preparation). KS: supervision, methodology and reviewing. JCG: supervision, visualisation, formal analysis, writing and editing. SF and SM: reviewing and editing.

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Code and data availability. The flood-triggered sinkhole susceptibility map was developed using publicly available datasets: Sinkhole inventory is available on (georisques.gouv.fr/, last access: 5 June 2024) in BD Cavit , BDMVT database. The Digital Elevation Model (DEM) was sourced from BD ALTI  (geoservices.ign.fr, last access: 5 June 2024), and the slope data were extracted from the DEM layer using GIS Pro. Surficial geology information was obtained from BD – Charm-50, the geological map (1/50,000) provided by BRGM available on (infoterre.brgm.fr/, last access date: 5 June 2024). Groundwater data were sourced from the Groundwater National Portal (ades.eaufrance.fr, last access: 23 April 2023). Flood hazard zones were provided by the Plan de Pr vention des Risques d'Inondation (PPRi, 2015). Finally, data on alluvium thickness were extracted from Perrin et al. (2015, <https://doi.org/10.1016/j.enggeo.2015.09.001>).

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