



A new approach to continuous monitoring of carbon use efficiency and biosynthesis in soil microbes from measurement of CO2 and O2

Kyle E. Smart¹, Daniel O. Breecker², Christopher B. Blackwood^{3,4,5}, and Timothy M. Gallagher¹

- ¹Department of Earth Sciences, Kent State University, Kent, OH, 44242, USA.
 - ²Department of Geological Sciences, University of Texas at Austin, Austin, TX, 78712, USA.
 - ³Department of Plant, Soil and Microbial Sciences, Michigan State University, East Lansing, MI, 48824, USA.
 - ⁴Department of Plant Biology, Michigan State University, East Lansing, MI, 48824, USA.
 - ⁵Ecology, Evolution, and Behaviour Program, Michigan State University, East Lansing, MI, 48824, USA.
- 10 Correspondence to: Kyle E. Smart (ksmart4@kent.edu)

Abstract

Soils comprise the largest terrestrial carbon pool. Therefore, understanding processes that control soil carbon stabilization and release is vital to improving our understanding of the global carbon cycle. Heterotrophic respiration is the main pathway by which soil organic carbon is returned to the atmosphere, however not all carbon utilized by heterotrophs shares this fate, as some portion is retained in the soil as biomass and biosynthesized extracellular compounds. The fraction of carbon consumed by microbes that is used for biomass growth (the carbon use efficiency or CUE) is an important variable controlling soil carbon stocks but is difficult to measure. Here we show that CUE can be continuously monitored in laboratory glucose-amended soil incubations by measuring CO2 and O2 gas concentrations, allowing instantaneous estimates of microbial biomass growth. We derive a theoretical relationship between the respiratory quotient (RQ), the ratio of carbon dioxide produced to oxygen consumed during respiration, and CUE that recognizes the influence of both substrate and biosynthesized product oxidation states on RQ. Assuming the biosynthesized product has the stoichiometry of an average microbe, and that the substrate is primarily the glucose used for amendment, we measure RQ and use our theoretical relationship to calculate CUE, and from that, biomass production. Extractions of microbial biomass carbon at the end of the experiments reveal minimal net increases in standing biomass across all amended treatments, suggesting that much of this newly produced biomass is likely converted to necromass as substrate availability declines and this results in a net storage of new soil organic matter. Carbon budgets compiled from measurements of relevant pools account for the amended carbon and suggest that with larger carbon amendments, increases in C:N ratios lead to increases in the relative portion of the amendment acutely lost from the soil. These findings demonstrate that soil RQ values may be used to monitor changes in CUE and that studies which monitor soil RQ values should consider CUE as a key factor when changes in RQ are observed, for instance, with changing environmental conditions or changes in production of plant derived compounds. This new approach may be leveraged to provide information on the storage of soil organic matter. These findings demonstrate how measurements of soil RQ may be leveraged to understand soil carbon transformations, specifically the fate of fresh carbon inputs.





1 Introduction

35

40

55

60

Soils represent one of the largest pools of carbon on the Earth's surface, with around 1477 Gt of carbon stored as soil organic matter (Scharlemann et al., 2014). The makeup of this pool can change dynamically as organic carbon is added through litter and root inputs, transformed by soil biogeochemical processes, and ultimately released back to the atmosphere via respiration (Dynarski et al., 2020; Kögel-Knabner, 2002; McDaniel et al., 2014; E. Paul, 2016). These exchanges of carbon are of particular importance, because as climate conditions continue to change and natural ecosystems exist in a state of increasing disequilibrium from antecedent conditions, it is difficult to predict the rates at which soils will accumulate or lose carbon. The processes that control soil carbon cycling are crucial to understand, not only in the context of global climate (Scharlemann et al., 2014), but also because soil organic carbon impacts soil fertility directly by providing essential nutrients and compounds for plants and microbes and indirectly by affecting soil physicochemical properties like wettability and drainage (Gaiser & Stahr, 2013). Therefore, improving our understanding of these processes may also better our efforts of conserving soil organic carbon in the context of global food security.

To understand if soils are experiencing a net gain or loss of carbon, it is necessary to first examine the interplay of biosynthesis and respiration (Adingo et al., 2021; Blagodatskaya et al., 2014; KM Geyer, 2016; Manzoni et al., 2018; Sinsabaugh et al., 2013). Accurately quantifying heterotrophic respiration is critical because it is the main mechanism by which carbon is released from soils (Landsberg & Gower, 1997; Mukul et al., 2020; Walker et al., 2018). Microbes consume soil organic matter not only as a source of energy via respiration, but also as a source of reduced carbon compounds for biosynthesis (Adingo et al., 2021; J. P. Schimel & Weintraub, 2003; J. Schimel & Schaeffer, 2012; Sinsabaugh et al., 2013). The proportion of carbon consumed by microbes that is retained in biomass, rather than respired, is known as the Carbon-Use Efficiency (CUE). Biosynthesis of microbial biomass and extracellular compounds is important to constrain because it is thought to be an important pathway for long-term stabilization of organic carbon within soils (A Miltner, 2012; Cotrufo et al., 2015a; MF Cotrufo, 2013; W Wieder, 2014). As soil microbes take up new organic carbon from fresh plant litter or other soil organic matter, CUE is the first crucial step in determining the fate of the consumed carbon (Kästner et al., 2021; Kindler et al., 2009; Liang et al., 2019; Miltner et al., 2011; E. A. Paul, 2016; Wang et al., 2021). After the stimulation of growth, newly produced microbial biomass is converted to necromass, as cell death occurs on the order of hours to days (Buckeridge et al., 2020). This necromass contains an abundance of molecules which may be further metabolized or recycled for molecular maintenance. However, not all this necromass is likely to be immediately accessible, due to factors including physical occlusion, chemical lability vs. recalcitrance, stabilization onto mineral surfaces, or continued supply of more desirable compounds (Buckeridge et al., 2020, 2022; Cotrufo et al., 2015b; Hu et al., 2023; Kästner et al., 2021; Liang et al., 2019; Lützow et al., 2006; E. A.



80

95



Paul, 2016; Wang et al., 2021). Regardless of the exact mechanism, many studies have shown that microbial necromass residues should be considered an important pool through which organic matter cycles and stabilizes in soils.

The concept of CUE can be applied at different spatial and temporal scales, depending on the question of interest (Adingo et al., 2021; Geyer et al., 2019; KM Geyer, 2016). For example, it may be useful to consider the CUE of individual microbial community members when studying ecological processes like competition or response to changes in environmental conditions. The CUE of the community as a whole may also be estimated when studying factors like ecosystem oxidation state (Geyer et al., 2019; Sinsabaugh et al., 2013). There is also debate as to whether CUE is an inherent species-specific value, and constant, or if CUE is a variable that can change over time given the needs of the microbes and the environmental conditions (Adingo et al., 2021; KM Geyer, 2016; Manzoni et al., 2012, 2018; Sinsabaugh et al., 2013). Regardless, CUE is crucial for understanding soil organic carbon stability because at low values, soil carbon is 'burned off' where at high values it is efficiently recycled. Unfortunately, CUE has been difficult to measure and nearly impossible to monitor continuously.

An emerging approach that can be used to study soil metabolisms and other soil processes is known as Respiratory Quotient (RQ), which is the ratio of CO₂ produced to the O₂ consumed during respiration (Dilly, 2001, 2003). The study of RQ can potentially provide insight into the substrate being metabolized because the stoichiometry of the compound should determine the reaction stoichiometry during aerobic respiration (Masiello et al., 2008). For example, respiration of compounds like sugars and other carbohydrates are predicted to produce an RQ of 1.0, lipids are predicted to have RQ values around 0.7, and most organic acids around 1.3 (Hicks Pries et al., 2020; Hilman et al., 2022; Masiello et al., 2008). While some studies report RQ values that resemble substrate-based predictions, other studies observed systematic deviations that were linked to non-metabolic processes which can affect soil CO₂ and O₂ concentrations, such as different diffusion constants of CO₂ and O₂ which can be accounted for and presented as Apparent Respiratory Quotient (ARQ), calcite dissolution/precipitation which can cause a transient decoupling of these two gases, and oxidation of reduced metal species which can cause additional draw down of oxygen (Angert et al., 2015; Bergel et al., 2017; Gallagher & Breecker, 2020; Hicks Pries et al., 2020; Hodges et al., 2019; Sánchez-Cañete et al., 2018).

The potential effect of microbial CUE on soil RQ values has received less attention to date, although Dilly (2001) suggested that incorporation of available substrates into microbial biomass could explain initial RQ values >1 observed during the initial stimulation period in soils amended with glucose. If microbial biosynthesis causes divergence of observed RQ values from expectations derived from substrate stoichiometry alone, then by examining the effects of CUE on RQ we may enable indirect monitoring of biosynthesis through the measurement of RQ. In order to examine the utility of RQ as a CUE-indicator, we designed incubation experiments in which glucose was added as a substrate to induce respiration, CO₂ and O₂ in the incubation vessel headspace were measured at high temporal resolution (every 2 h) and biomass was measured by the chloroform-fumigation extraction method once respiration rates declined to baseline values. We further explore the implications of the biosynthesis we infer from the measurements in the context of the fate of soil organic carbon transformations.



105

110

115



2 Connecting Carbon Use Efficiency and Respiratory Quotient

When substrate is converted entirely to CO₂ and yields no net biomass production, carbon use efficiency is zero and does not influence RQ. When CUE is non-zero, RQ values are driven by the difference between the oxidation states of carbon in substrate and reaction product (i.e., between the molecule consumed and the molecule produced through anabolism). To understand how biosynthetic processes influence RQ, we must describe how changing CUE will influence this stoichiometry by considering the production of microbial biomass as a key reaction product. Using a mass balance approach, we can explore the relationship between RQ and CUE in the reaction:

Eq 1.

$$A C_6H_{12}O_6 + B O_2 + F NO_3^- = C CO_2 + D H_2O + E C_1H_{1.8}O_{0.5}N_{0.2}$$

where $C_1H_{1.8}O_{0.5}N_{0.2}$ is a representative microbial biomass stoichiometry (Roels, 1980) normalized per mole of carbon, and the letters A - F serve as coefficients. Due to its relative importance in microbial makeup, nitrogen was included in the calculations. We chose to use nitrate as the nitrogen bearing substrate due to its impact on RQ values by its redox state and widespread occurrence in soils. A derived theoretical relationship between RQ and CUE, following Eq. 1 is shown below (Fig.1), and is further applied to experimental data to address our research questions . Further information on the derivation can be found in the appendix.

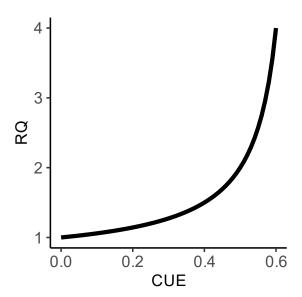


Figure 1: The calculated relationship between carbon use efficiency and respiratory quotient for Eq.1. This modeled relationship shows that as CUE increases, RQ values will also increase, which may seem counter-intuitive at first, given that an increase in CUE would cause a net decrease in CO_2 production, all else being equal. However, the concurrent O_2 consumption decreases more substantially as uptake of NO_3^- increases, which in turn results in RQ values increasing. The slope of the modelled relationship shows that RQ increases rapidly as CUE values approach towards 60%. We limited our



125

130

135



calculations to a maximum of 60% CUE, as this is referenced as a theoretical thermodynamic limit for microbial metabolisms (Sinsabaugh et al., 2013).

120 3 Materials and Methods

This study consists of two soil incubations designed to investigate the effects of labile substrate (glucose) amendment on RQ values at high temporal resolution, and to evaluate the effects of CUE on RQ. Control samples (addition of Type 1 deionized water, Millipore Milli-Q, to the soil) were incubated and measured for comparison. Treatment samples involved amendment with various masses of glucose (100 mg, 200 mg, 500 mg, or 1.0 g). Each of the two incubations consisted of two control samples, and six treatment samples. All incubated samples contained 20 g of soil. RQ was determined by monitoring the composition of headspace gas in the incubation vessels every 2 h for the duration of the incubations (262 h at longest).

The soils used in incubations were collected from a temperate deciduous forest in Portage County, Northeast Ohio. Soils in this location are designated as Chili Loam by the USDA Soil Survey. Soil collection was performed with a shovel, excavation included approximately the top four inches of the profile to include the Oi-horizon, and top 5 cm of the A horizon. Soil was then returned to the lab and homogenized. For purposes of incubation, the field moist soils were passed through a 2 mm sieve to remove large detritus and leaf litter and to break up large aggregates. Soils were then allowed to dry down, open to lab air, for 2 weeks to encourage the depletion of any preexisting labile carbon and reduction of standing microbial biomass. Soil aliquots of 20 g (approximately 30 mL) were added to each incubation bottle. Glucose amendments were weighed and added to the soils as a fine solid powder, homogenized through physical mixing, and placed in 500 mL bottles. Once in the bottles, 10 mL of Type 1 deionized water (Millipore Milli-Q) was dripped evenly over the soils to encourage glucose dissolution before the bottles were capped and connected to the gas sampling apparatus. The addition of water led to an average soil moisture of 36% by mass at the start of incubation. Incubations were carried out in an incubator held to constant temperature of 20 °C.

3.2 Automated Gas Sampling Apparatus

An automated gas sampling apparatus was constructed that allowed gas samples to be continuously collected and measured from soil incubations every two hours. Soils were incubated in 500 mL glass bottles (PYREXTM) with 3 gas-tight tube ports in the lid (Duran® GL45). One port on each bottle was connected to a Calibrated Instruments (McHenry, USA) Cali-5-BondTM gas sampling bag, filled with an additional 300mL of CO₂-Free Air to give the incubation vessel a variable volume, which enabled gas samples to be collected and new gas to be added while maintaining atmospheric pressure. Bev-A-Line IV tubing connected the bottles through a second port in the lid to a central manifold block with solenoid valves. The third port was closed off and was not used in this study. A Sable Systems (Las Vegas, USA) FOXBOX was used to measure high precision CO₂ and O₂ gas concentrations. All sampled gas was dried using PermaPure (Lakewood, NJ, USA) NafionTM Tubing, passing through a separate 500 mL bottle containing magnesium perchlorate, and held at partial vacuum, prior to measurement.



155

160

165



The configuration of the sampling apparatus is depicted below in Figure . From the central manifold system gas flow could be (1) closed, (2) directed from the bottles into the FOXBOX, or (3) directed from compressed gas cylinders into the bottles. The manifold system could also direct flow of the compressed cylinders directly to the FOXBOX.

The entire system was controlled by a programmable logic controller (PLC), which automatically opened and closed solenoid valves, directed the flow of gas through the system, and logged data from the FOXBOX. Every two hours a measurement sequence would begin whereby bottles were sequentially measured for 3.5 min at a flow rate of 50 mL min⁻¹ for a total sample of 175 mL of gas. To maintain high temporal resolution measurements (2 h), a maximum of eight individual samples could be incubated simultaneously. To account for any short-term drift in measured O₂ values, ambient air was automatically measured directly from the laboratory HVAC inlet vent, between sample measurements. Sampling from HVAC inlet vent was preferred over lab air because HVAC air is a mixture of air sources from throughout the building and would provide a more stable measurement of O₂, whereas lab air O₂ concentration may fluctuate more dramatically with changes in room occupancy or sampling exhaust. Additionally, gas cylinders were measured containing zero (CO2-Free Air) and calibration (5000 ppm CO2) gasses to account for long-term measurement reproducibility. Lastly, after sampling from each bottle, the 175 mL of gas removed for analysis was replaced with CO2-Free Air by directing cylinder flow through a needle valve and a mass-flow meter into the incubation vial-gas bag system. The resulting dilution of CO₂ and addition of O₂ within each bottle was accounted for when calculating moles of CO₂ produced and O₂ consumed between measurements.

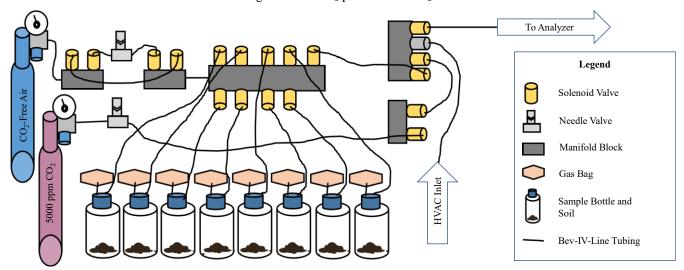


Figure 2: Experimental apparatus. This diagram displays the configuration of components used to construct the automated gas sampling apparatus. Blue and pink gas cylinders on the left are labelled by type. All other components are identified in the legend. Arrows indicate "HVAC Inlet" used between sample measurements to separate measurement periods, and "To Analyzer" as the final outflow to the sample drier and FoxBox.



175

180

185

195

200



170 3.3 Microbial Biomass

Microbial biomass carbon was measured via the Chloroform Fumigation Extraction method following the methods of (McDaniel et al., 2014; Vance et al., 1987) on initial material at the start of the incubation and on the control and treatment samples at the end of the incubations. In short, duplicate subsamples (~5g) were weighed out and one set were immediately extracted with 0.5M K₂SO₄, on a rotator table for 1 hour; these samples served as unfumigated water (K₂SO₄) extractable carbon. Next, the remaining samples were furnigated using ethanol-free chloroform (1mL) and capped for 24 h in a furne hood, then extracted with K₂SO₄; this set would serve as fumigated extractable carbon All extracts were filtered through a Whatman #1 filter with a vacuum filtration apparatus immediately following extraction. Soil moisture measurements were carried out with the use of a drying oven and were determined gravimetrically on a third subsample of ~ 5 g of soil. Non particulate organic carbon was measured using a Shimadzu TOC-L Analyzer (Shimadzu Scientific Instruments Inc.) and reported in dissolved organic carbon (DOC) in mg L⁻¹. Dissolved organic carbon for both fumigated and unfumigated subsamples were used to calculate biomass carbon as Fumigated DOC - Unfumigated DOC = Biomass Associated DOC in mg C g⁻¹ dry soil. Final values are reported on a per bottle basis. A correction factor (K_{ec} = 0.45) was applied to account for the extraction efficiency of biomass carbon by chloroform, to convert Biomass Associated DOC to Biomass C (Vance et al., 1987). Salt Extractable carbon is presented as the unfumigated DOC and reported in mg C per bottle. Microbial biomass extraction was conducted on initial soil, on incubated control soil, and incubated amended soil. Incubated samples were harvested for biomass extraction immediately following decline in respiration stimulated by the amendment, and when measured RO values drop below 1.0 for all replicates in each treatment group.

3.4 C:N Measurement

C:N values were determined for soil samples using a Costech Elemental Analyzer (EA) ECS 4010 configured with a CNH combustion column. In short, the dried subsamples used to collect soil moisture information as part of the CFE method were ground to a fine powder and weighed out into tin capsules. Reported values are given as the ratio of carbon to nitrogen in percent weight.

3.5 Data Analysis

Following the incubation, raw gas concentration data were processed in RStudio to quantify sample CO_2 and O_2 concentrations, and apply a baseline correction. The baseline correction is done with a linear fit to HVAC air measurements made immediately preceding sample measurements. These HVAC measurements were corrected to 20.95% O_2 . This correction is necessary to account for short-term drift on the fuel cell O_2 sensor, mostly caused through changes in temperature either by ambient temperature or through heat dissipation within the instrument. Once the HVAC measurement corrections are established, the same correction is applied to sample measurement windows. Reported values of each sample are taken as the average value during the last 20 s (measurements are recorded every 2 s, 10 consecutive measurements are used) of the sampling window



205

210

215

220

225

230



and an uncertainty is reported as the standard deviation. These drift-corrected data are then exported from RStudio into Excel for further processing. In Excel, measured CO₂ concentrations were corrected using a 2-point linear calibration curve produced from measurements of CO₂-Free Air and 5000 ppm CO₂ gases. A mass balance approach was then used to calculate the moles of CO₂ produced (Fig. 3a) and O₂ consumed (Fig. 3b) during each 2 h incubation window, accounting for the dilution effect of replacing the sampled gas volume with 175 mL of CO2-free air after each analysis. With these data, RQ values for each 2 h interval were calculated (Figure c). The variables of interest are saved in .csv files and imported to RStudio equipped with R version 4.2.2. Variables of interest include: time, CO₂ production rate, O₂ consumption rate, RQ, treatment, and replicate. Periods of substrate induced respiration are defined here as being represented by an RQ \geq 1.0 and occurring during periods of elevated CO₂ production. CUE values were then calculated at each measurement of RQ using the relationship in Fig. 1, during the previously defined periods of substrate induced respiration. Following this, CUE and CO2 production rates were used to calculate moles of biomass carbon produced for each 2 h measurement interval. All variables, both measured and calculated, were then plotted. Packages employed in R include tidyverse, gridExtra, cowplot, and svglite. After analysis of the data, it was determined that one of the eight bottles had an unnoticed leak during both incubation runs, so for these experiments only duplicate results are presented in Figure 3, and these data from the problematic bottle were removed from presented averages in Fig. 4 and Fig. 5 (200 mg and 1000 mg incubations). Data from the first 2 hours was not plotted in Fig. 3 and Fig. 4, as initial measurements produced a transient signal showing incredibly large O₂ consumption, which was likely the result of the bottles equilibrating with the new system. Presented values begin at 4 h.

4 Results and Discussion

4.1 High Temporal Resolution RQ

Respiration of glucose, and other simple carbohydrates, should produce a RQ value of 1.0, if CUE = 0 (Masiello et al., 2008). We observe RQs systematically greater than 1.0 post amendment, suggesting CUE > 0. Using mass balance calculations, we determined RQ values with a 2 h resolution (Figure c), over the duration of 262 hours (10 days and 22 hours). Initial rates of CO₂ production over the first 24 hours (Figure a) show a similar overall trend regardless of amendment quantity, with all four amended treatments resulting in almost identical values. The rate of increase in CO₂ production initially appears to be inversely related to the amendment quantity, as the smaller amendment treatments begin to grow slightly faster than the larger amendments. Around 80 h of incubation, the CO₂ production rate of the 100 mg treatments peaked and declined over the remainder of the incubation. Peak CO₂ production for the 200, 500, and 1000 mg treatments occurred at 46-60, 78, and 84-92 hours, respectively. Notably, the 500 mg treatments reached comparable maximum CO₂ production rates with the 1000 mg treatments, suggesting that substrate availability alone is not a reliable predictor of yield in peak microbial respiration. One possible explanation for this trend is slower dissolution of the glucose amendment in the 1000 mg treatment, which could be described by the data, because of the extended period of enhanced CO₂ production, resulting in greater cumulative CO₂ production, in the 1000 mg treatment (Figure a). Oxygen consumption rates displayed in Fig. 3b show a similar behavior to



235

240

245

250



CO₂ production rates in Figure a, apart from variability between timepoints and maximum values reached. Oxygen consumption rates occurred in a smoother, less erratic trend. Also, important to note is that the control bottles oxygen consumption and production of carbon dioxide did not respond during the incubation period in any meaningful way.

Initially, at 4 h of incubation, RQ values across all treatments were noisy and ranged between ~0.3- 1.5, probably related to error associated with determining RQ when respiration rates are small. From 4 h onward, RQ values in amended treatments start an overall ascent. After ~24 h of incubation, coinciding with an increase in CO₂ production and O₂ consumption, RQ values across most treatments are > 1.0. While the rates of gas exchange continue to climb, RQ values also increase. RQ values observed during peak respiration (~1.3-1.6) are similar across treatments. As the rates of CO₂ production begin to decline, RQ values also decline. Although treatment replicates are variable with respect to time, the overall trends are in good agreement. RQ is not shown for control samples because we observed no overall trend (i.e. no increase or decrease). We see that RQ values are dynamic at this temporal resolution, even during the period which should be dominated by substrate induced respiration, meaning that RQ values are not simply a direct result of the substrate being oxidized to produce CO₂.

Peak RQ values were observed during peak respiration and are similar to those observed in (Dilly, 2001), with RQ values \sim 1.5. Notably, all treatments measure \sim 1.5 despite an order of magnitude increase in glucose amendment. This suggests that biosynthetic processes are limited by the rate of synthesis of biomolecules perhaps by temperature or by the availability of other nutrients (eg. N or P). Also importantly, we see that the overall range in RQ values is quite large (0.3 - 1.9). These higher values could be explained through partially anabolic metabolism; however values below 1.0 likely indicate the use of some other substrate in which the carbon is more reduced. This other carbon substrate could be a form of less labile organic matter contained within the initial soil samples, or metabolites that were produced during the respiration of glucose.





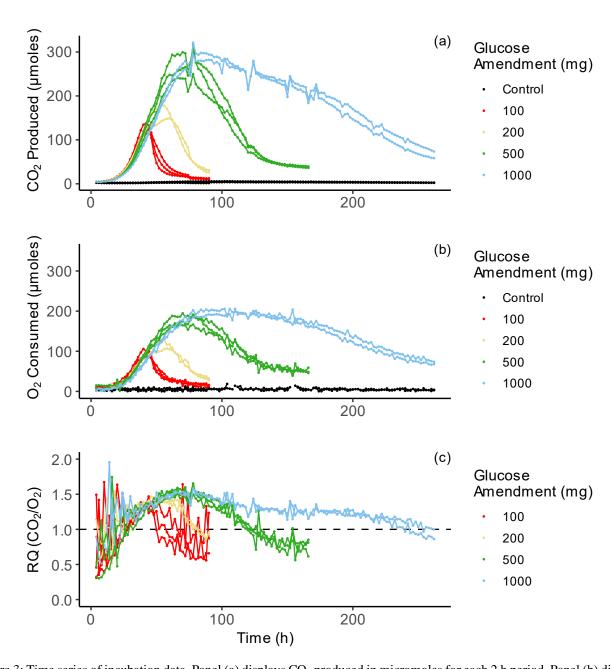


Figure 3: Time series of incubation data. Panel (a) displays CO₂ produced in micromoles for each 2 h period. Panel (b) displays O₂ consumption in micromoles for each 2 h period. Panel (c) displays Respiratory Quotient (RQ) for each 2 h timepoint, calculated as [CO₂ Produced / O₂ Consumed].





4.2 High Temporal Resolution Carbon Use Efficiency Estimates

Applying the RQ—CUE relationship (Fig. 1) to the incubation data (Fig. 3) allows CUE values to be estimated for each 2 h interval of the experiments. Then, using the CO2 production rate in moles and CUE, biomass production rate in moles-C per 2 hr period can be estimated during each time step (Figure 4b), using the following equation:

Eq. 2

Biomass Produced = CO_2 Produced / (1 - CUE) * CUE

(Fig. 4c) and cumulatively throughout the experiment (Figure 4d). It is important to note that once RQ values drop below a value 1.0, the modeled RQ—CUE relationship for glucose as the sole substrate no longer applies. Further, when RQ values drop below 1.0, this coincides with the point that respiration rates are returning to new basal respiration rates that are elevated over the basal respiration observed in control bottles (Fig. 3). We infer that most if not all available glucose provided in the amendment has been utilized by this point of the incubation. Any further activity is likely driven by metabolism of an alternative substrate, or biomass turnover. Biomass production rates closely resemble respiration rate trends for the incubation. Curves of cumulative biomass produced (Fig. 4d) show all treatments display a sigmoidal shape, which is to be expected as production rates begin low, increase, and then decline back to zero.

Maximum estimated CUE was ~0.56, and the highest values were seen near the beginning of the incubation when RQ values were around 1.9, which may indicate highly efficient growth of small microbial populations, although the small signals produced at the beginning of the incubation may have also been dominated by measurement noise because respiration rates were still low. As respiration rates begin to increase, CUE estimates stabilize at ~0.3, and then continue to increase with respiration rates to ~0.4. After respiratory decline, when CO₂ production and O₂ consumption rates return to new basal conditions, RQ values decline to <1.0 initially and CUE estimates fall to zero. This period may represent a transitionary phase, when the high-lability glucose amendment has been depleted and the microbes begin to turn over and/or target alternate sources of organic carbon. Masiello et al. (2008) provides RQ values for other common organic compounds in soils which may serve as these alternate sources. From the list of compounds and their associated RQ's several candidate compounds could satisfy the requirements of our observations; for example, proteins produce RQ's ranging from 0.67-1.01, lignin ranges from 0.88-0.94, and lipids range from 0.68-0.80. Oxidation of any or all of these classes of compounds could explain our observations given that they are basic constituents of plant and microbial biomass and are ubiquitous in soil organic matter.





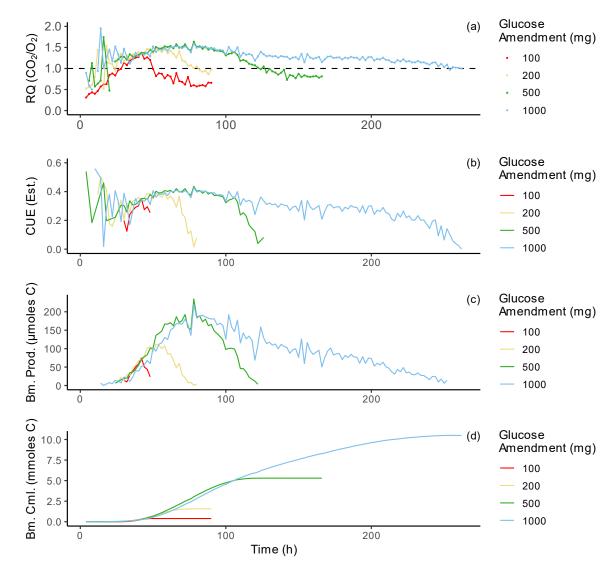


Figure 4: Time series of values calculated from incubation data. All panels present one replicate from each amended treatment for a visual example, additional replicates were hidden from this plot for purposes of clarity. All treatments were carried out through these same calculations. Panel (a) displays RQ over the incubation period, the same data presented in Figure 3(c) are shown here. Panel (b) displays Carbon Use Efficiency estimated for this incubation using the RQ—CUE relationship presented in Fig. 1. Panel (c) displays micromoles of biomass (carbon) produced at each 2 h time point for the incubation period. Panel (d) displays the cumulative sum of biomass produced during the incubation in millimoles carbon.



300

305

315



4.3 Understanding the Fate of Amended Carbon

Through the use of the CFE method for microbial biomass carbon measurements along with our gas-based measurements and determinations of respiration and biomass production, we can construct a carbon budget for each treatment (Fig. 5). Measurements of net biomass produced through CFE (Net Biomass produced = Biomass Amended - Biomass Control) on a per bottle basis, show minimal, but direct increases with amendment. Necromass values, calculated as Necromass = cumulative RQ Biomass Produced - Net Biomass Produced, show an increase with amendment size. Salt extractable carbon is presented as the increase in final unfumigated DOC in amended treatments relative to the unfumigated DOC measured from initial unamended soil, which was measured as part of the CFE biomass calculation, and shows minimal increase with amendment size. This increase in salt extractable carbon could be the result of leftover amendment, or from enzymes and other intercellular or extracellular compounds produced from the stimulated microbial activity, however, it is possible that some overlap between necromass and salt extractable carbon is possible. Respired carbon, calculated as the cumulative carbon lost through respiration, shows direct and relatively proportional increase with amendment size. Overall, the sum of these carbon pools nearly equals the amount of carbon amended to the soil (calculated as 0.4 mg C/mg glucose), as expected for a closed system (Fig. 5), which provides strong support for our predicted CUE and RQ relationship (Fig. 1).

Treatment	Replication	Amendment (mg C)	Respired (mg C)	Salt Extractable (mg C)	Biomass (mg C)	Necromass (mg C)
A100	n=3	41.7 ± 1.2	20.1 ± 0.2	4.2 ± 2.3	3.7 ± 5.0	2.5 ± 5.0
A200	n=2	82.4 ± 1.1	39.4 ± 0.3	8.5 ± 1.5	3.5 ± 1.6	14.5 ± 1.6
A500	n=3	200.5 ± 0.5	117.2 ± 1.1	5.0 ± 1.8	4.3 ± 3.3	52.3 ± 3.5
A1000	n=2	400.4 ± 2.3	263.7 ± 1.9	14.5 ± 3.4	5.6 ± 1.7	118.0 ± 2.6

Table 1. Reported mean ± uncertainty of each respective carbon pool in mg C per bottle. Amendment uncertainty is reported as standard deviation of the replicates, all other uncertainties reported are uncertainty propagated through calculation using standard deviation of replicates.

Taken together the Biomass, Necromass, and Salt Extractable carbon pools represent carbon that is remaining within the soil from the amendment after incubation, whereas respired carbon can be considered lost from the soil. With these results, we see that across the treatments, as amendment size increases a larger portion of the amendment is lost through respiration (~50% for the 100 mg amendment to ~66% for the 1000 mg amendment), and a smaller fraction of carbon initially amended as glucose remains in the soil after incubation. However, there are many aspects to this trend that must be considered, such as duration of incubation, long term stability of this necromass, and stoichiometric limitations.



330



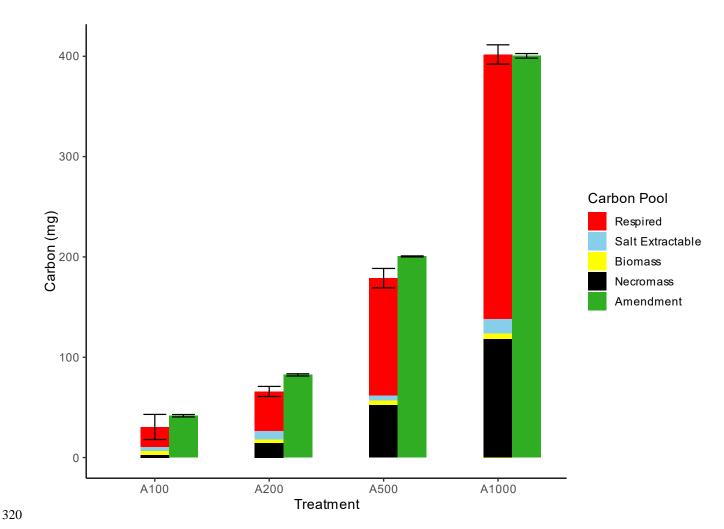


Figure 5. Fate of Amended Carbon. This bar chart shows the respective pools of carbon ascertained through direct measurement, or calculation presented as average with error bars representing uncertainty. Replication varies by treatment (A100: n=3,A200: n=2, A500:n=3,A1000:n=2).

Stoichiometric limitation may be driving the observed increase in the fraction of amended carbon lost via respiration with increasing glucose amendment, considering the carbon amendment was applied without the addition of any other nutrients like N or P. This treatment would drive C:N ratios up, placing the bulk SOM pool in a more carbon enriched state (Table 2), thus would likely drive more waste respiration as other critical nutrients would then be placed in relative limitation (Brown et al., 2022). Further support for this interpretation can be drawn from the slope of the declining RQ values following peak respiratory activity. During the 100 mg incubation, RQ values declined sharply once respiration slowed, whereas the decline became more gradual with increasing amendment size. Following our model, these decreasing RQ values correspond to decreasing CUE. With larger amendment sizes, there was a longer time interval during which RQ values remained above 1.0 but below the ~1.5



345

350

355

360



values observed during peak respiration. This period where RQ values are closer but remain above 1.0 could be explained through a mixture of ongoing glucose fueled metabolism and the onset of microbial necromass turnover, with the latter expected to produce an RQ of ~0.7-0.8. An alternative explanation could be a slower rate of microbial biosynthesis than during peak activity, as increasing nutrient limitation imposes thermodynamic/stoichiometric limitations on biosynthesis and this could be directly reflected in lower measured RQ values as a result of smaller CUE's during the later stages of glucose fueled metabolism.

Treatment	Replication	C:N	Carbon Weight %	Nitrogen Weight %
Initial Soil	n=3	20.2 ± 1.5	5.12 ± 0.02	0.25 ± 0.02
A100	n=3	22.2 ± 2.0	5.73 ± 0.24	0.26 ± 0.01
A200	n=2	24.9 ± 1.3	5.76 ± 0.04	0.23 ± 0.01
A500	n=3	25.7 ± 4.2	6.17 ± 0.24	0.25 ± 0.05
A1000	n=2	28.8 ± 1.7	6.65 ± 0.09	0.23 ± 0.02

Table 2. Measured C:N ratios, carbon weight percent, and nitrogen weight percent of initial soil and incubated treatment soils, reported as mean ± standard deviation.

Recent research shows that after a long (weeks to months) period of incubation, around half of biomass derived carbon may persist within soil as small fragments of cellular envelopes within soil organic matter (SOM) (Kindler et al., 2009; Liang et al., 2019; Miltner et al., 2011). Kästner et al. (2021) highlights a large discrepancy between small quantities of standing live biomass and massive quantities of necromass residue which make up a meaningful portion of SOM. Further, Liang et al. (2019) examined this contribution across ecotypes and found that in temperate forest systems it can account for ~30% of soil organic carbon (SOC), though they claim that this lower contribution in temperate forests may be the result of dilution from large continuous inputs of plant material and the lack of tillage. These findings warrant further investigation on the quantification of microbially derived accumulation of SOM, especially through understanding short term microbial metabolism and propagation. The short-term stability of freshly produced necromass in soils remains uncertain. Kästner et al. (2021) describes microbial turnover as a multi-step process where initial cell lysis results in a rapid release of compounds which can quickly stimulate continued biosynthesis, and this cell lysis can be driven through a slower process of starving as substrate availability declines or through more rapid process such as viral activity and microbial grazing (Santos-Medellín et al., 2023). Reflecting on the findings in Fig. 5, we can assess if our data are better explained by substrate depletion and starvation or viral activity and grazing. If substrate depletion and starvation is the dominant driving force behind the formation of necromass, then we might expect greater production of necromass later in the incubation, only once the substrate availability has declined significantly and RQ values drop below 1.0 and approach ~0.8. In contrast, if viral or grazing activity is the dominant mechanism by which necromass formation occurs, then we would expect a continued formation of necromass relatively in line with the rate of formation of new biomass (Jansson, 2023; Williamson, 2011; Williamson et al., 2005, 2017; Wu et al., 2021).





The CFE measurement of biomass carbon occurred immediately after gas measurements ceased, allowing minimal time for further biomass decline. These CFE measurements show that minimal increases in standing biomass production occurred with increasing amendment size, even though very little time passed between the end of the period explained by glucose metabolism (RQ values ≥ 1.0) and the harvesting for CFE.

This minor increase in standing biomass contrasts strongly with large quantities of total biomass production estimated from the observed RQ values. Taken together, these observations suggest that the rate of new biomass formation during the experiment was similar to the rate of necromass production. Otherwise, we would expect more substantial increases in living biomass once the incubations were stopped. Therefore, viral activity and microbial grazing are considered more suitable explanation, especially considering the treatment of samples as soil was allowed to dry down and were then re-wet (Santos-Medellín et al., 2023; Wu et al., 2021), as recent literature has shown rewetting of dry soil leads to elevated viral activity in a "culling of the victor" strategy. Additionally, considering the shifts in C:N ratios within these samples caused through the amendment of increasing quantities of carbon with no corresponding amendment of nitrogen, we likely drove stoichiometric limitation on the production of new biomass and could have created conditions in this soil which require elevated nutrient mining through strategies such as microbial grazing.

375 **5 Conclusions**

380

385

390

A new automated gas sampling apparatus design enabled measurement of high-precision RQ values at a high temporal resolution (~2 h). The non-destructive sampling method allows samples to continuously incubate for a wide range of experimental durations without needing to disturb the incubation chamber. Our results demonstrate that RO values observed throughout glucose-stimulated incubations display systematic deviations from the value predicted (1.0) for pure respiration of simple carbohydrates. During peak respiration, RQ values were >1.5, which cannot be explained by a shift to other substrates. Instead, these elevated RQ values during peak activity are best explained by some fraction of the substrate consumed being used to biosynthesize other compounds. Derivation of a stoichiometric relationship between RQ and CUE values enabled measurements of RQ to provide contextual information regarding microbial respiration and biosynthesis. Not only can this approach provide estimates of CUE at a temporal resolution matching that of RQ measurements, but simultaneous estimates of biomass production can also be calculated by combining this information with CO₂ production rate. Importantly, our derived CUE—RQ relation, may be one way forward in real time monitoring of CUE which has proven difficult to measure. Other relationships between CUE and RQ of alternative compounds could be derived for further applications. Although our study demonstrates that microbial CUE can impact measured RQ values, the RQ values can still act as a rough index of shifts in dominant metabolism, as evidenced by the observed shift to RQ values of ~0.8 after respiration declined significantly. Importantly, these shifts in dominant metabolism after the amendment of labile substrate are likely driven by the turnover and metabolism of newly produced necromass. However, RQ derived estimates of biomass production were much greater than CFE estimates of standing biomass, suggesting that much of the necromass was not rapidly consumed, although the longer-



395



term stability of this necromass is uncertain. Further consideration of these measurements with a carbon budget reveal that stoichiometric limitation of C:N ratio, could be driving enhanced microbial turnover. Implications of these findings must be considered in the context of environmental conditions, where heterogeneity of resource availability, and the synergistic mechanisms of a broad microbial community could act to support enhanced carbon stabilization over the long term.

Appendix A

Equations Used in Stoichiometric Modelling of RQ as a function of CUE.

400 REACTION: $A(C_6H_{12}O_6) + B(O_2) + F(NO_3^-) = C(CO_2) + D(H_2O_1) + E(C_1H_{1.8}O_{0.5}N_{0.2})$

First step: Elemental mass balance expressions

Carbon: 6A=C+E or E=6A-C

Hydrogen: 12A = 2D + 1.8E

Oxygen: 6A + 2B + 3F = 2C + D + 0.5E

405 Nitrogen: F= 0.2E

Second step: Define CUE and RQ

CUE = E / (E + C)

RQ = C/B

Third Step: solve for RQ as a function of CUE, and CUE as a function of RQ

410 RQ = (2 - 2CUE) / (2 - 3CUE) or CUE= (2RQ - 2) / (3RQ - 2)

Data Availability

The data used in the production of this manuscript is hosted in a Zenodo repository (Smart, K.E., 2024).

Author Contributions

415 KES: Carried out experiments and laboratory analyses, analyzed data, produced figures, and produced manuscript.

DOB: Developed RQ—CUE relationship, assisted in planning incubation, assistance in interpreting the results, and provided helpful guidance in the production of the manuscript.

CBB: Assisted in planning incubations, interpreting the results, and provided helpful feedback and suggestions in the production of the manuscript.

420 TMG: Directly advised KES throughout the entire project from conceptualization through the production of the manuscript.





Competing Interests

The authors declare that they have no conflict of interest.

Acknowledgements

The authors would like to thank the Kent State University Department of Earth Sciences for funding. This material is also based upon work supported by a GSA graduate student research grant awarded to K.E.Smart.

References

- A Miltner, P. B. B. S.-B. M. K. (2012). SOM genesis: microbial biomass as a significant source. *Biogeochemistry*, 111, 41–55.
- Adingo, S., Yu, J. R., Xuelu, L., Li, X., Jing, S., & Xiaong, Z. (2021). Variation of soil microbial carbon use efficiency (CUE)

 and its Influence mechanism in the context of global environmental change: a review. *PeerJ*, 9.

 https://doi.org/10.7717/PEERJ.12131
 - Angert, A., Yakir, D., Rodeghiero, M., Preisler, Y., Davidson, E. A., & Weiner, T. (2015). Using O2 to study the relationships between soil CO2 efflux and soil respiration. *Biogeosciences*, 12(7), 2089–2099. https://doi.org/10.5194/bg-12-2089-2015
- Bergel, S. J., Carlson, P. E., Larson, T. E., Wood, C. T., Johnson, K. R., Banner, J. L., & Breecker, D. O. (2017). Constraining the subsoil carbon source to cave-air CO2 and speleothem calcite in central Texas. *Geochimica et Cosmochimica Acta*, 217, 112–127. https://doi.org/10.1016/J.GCA.2017.08.017
 - Blagodatskaya, E., Blagodatsky, S., Anderson, T. H., & Kuzyakov, Y. (2014). Microbial Growth and Carbon Use Efficiency in the Rhizosphere and Root-Free Soil. *PLOS ONE*, *9*(4), e93282. https://doi.org/10.1371/JOURNAL.PONE.0093282
- 440 Brown, R. W., Chadwick, D. R., Bending, G. D., Collins, C. D., Whelton, H. L., Daulton, E., Covington, J. A., Bull, I. D., & Jones, D. L. (2022). *Nutrient (C, N and P) enrichment induces significant changes in the soil metabolite profile and microbial carbon partitioning*. https://doi.org/10.1016/j.soilbio.2022.108779
 - Buckeridge, K. M., Creamer, C., & Whitaker, J. (2022). Deconstructing the microbial necromass continuum to inform soil carbon sequestration. *Functional Ecology*, *36*(6), 1396–1410. https://doi.org/10.1111/1365-2435.14014
- Buckeridge, K. M., La Rosa, A. F., Mason, K. E., Whitaker, J., McNamara, N. P., Grant, H. K., & Ostle, N. J. (2020). Sticky dead microbes: Rapid abiotic retention of microbial necromass in soil. *Soil Biology and Biochemistry*, 149. https://doi.org/10.1016/J.SOILBIO.2020.107929
 - Cotrufo, M. F., Soong, J. L., Horton, A. J., Campbell, E. E., Haddix, M. L., Wall, D. H., & Parton, W. J. (2015a). Formation of soil organic matter via biochemical and physical pathways of litter mass loss. *Nature Geoscience 2015 8:10*, 8(10), 776–779. https://doi.org/10.1038/ngeo2520





- Cotrufo, M. F., Soong, J. L., Horton, A. J., Campbell, E. E., Haddix, M. L., Wall, D. H., & Parton, W. J. (2015b). Formation of soil organic matter via biochemical and physical pathways of litter mass loss. *Nature Geoscience*, 8(10), 776–779. https://doi.org/10.1038/NGEO2520
- Dilly, O. (2001). Microbial respiratory quotient during basal metabolism and after glucose amendment in soils and litter. *Soil Biology and Biochemistry*, *33*(1), 117–127. https://doi.org/10.1016/S0038-0717(00)00123-1
 - Dilly, O. (2003). Regulation of the respiratory quotient of soil microbiota by availability of nutrients. *FEMS Microbiology Ecology*, 43(3), 375–381. https://doi.org/10.1111/J.1574-6941.2003.TB01078.X
 - Dynarski, K. A., Bossio, D. A., & Scow, K. M. (2020). Dynamic Stability of Soil Carbon: Reassessing the "Permanence" of Soil Carbon Sequestration. *Frontiers in Environmental Science*, 8, 218. https://doi.org/10.3389/FENVS.2020.514701/BIBTEX
 - Gaiser, T., & Stahr, K. (2013). Soil organic carbon, soil formation and soil fertility. *Ecosystem Services and Carbon Sequestration in the Biosphere*, 407–418. https://doi.org/10.1007/978-94-007-6455-2_17/FIGURES/00177
 - Gallagher, T. M., & Breecker, D. O. (2020). The Obscuring Effects of Calcite Dissolution and Formation on Quantifying Soil Respiration. *Global Biogeochemical Cycles*, *34*(12), e2020GB006584. https://doi.org/10.1029/2020GB006584
- Geyer, K. M., Dijkstra, P., Sinsabaugh, R., & Frey, S. D. (2019). Clarifying the interpretation of carbon use efficiency in soil through methods comparison. *Soil Biology and Biochemistry*, 128, 79–88. https://doi.org/10.1016/J.SOILBIO.2018.09.036
- Hicks Pries, C., Angert, A., Castanha, C., Hilman, B., & Torn, M. S. (2020). Using respiration quotients to track changing sources of soil respiration seasonally and with experimental warming. *Biogeosciences*, 17(12), 3045–3055.
 https://doi.org/10.5194/BG-17-3045-2020
 - Hilman, B., Weiner, T., Haran, T., Masiello, C. A., Gao, X., & Angert, A. (2022). The Apparent Respiratory Quotient of Soils and Tree Stems and the Processes That Control It. *Journal of Geophysical Research: Biogeosciences*, 127(3), e2021JG006676. https://doi.org/10.1029/2021JG006676
- Hodges, C., Kim, H., Brantley, S. L., & Kaye, J. (2019). Soil CO 2 and O 2 Concentrations Illuminate the Relative Importance
 of Weathering and Respiration to Seasonal Soil Gas Fluctuations . Soil Science Society of America Journal, 83(4), 1167–1180. https://doi.org/10.2136/sssaj2019.02.0049
 - Hu, J., Du, M., Chen, J., Tie, L., Zhou, S., Buckeridge, K. M., Cornelissen, J. H. C., Huang, C., & Kuzyakov, Y. (2023). Microbial necromass under global change and implications for soil organic matter. *Global Change Biology*, 29(12), 3503–3515. https://doi.org/10.1111/GCB.16676
- 480 Jansson, J. K. (2023). Soil viruses: Understudied agents of soil ecology. *Environmental Microbiology*, 25(1), 143. https://doi.org/10.1111/1462-2920.16258
 - Kästner, M., Miltner, A., Thiele-Bruhn, S., & Liang, C. (2021). Microbial Necromass in Soils—Linking Microbes to Soil Processes and Carbon Turnover. *Frontiers in Environmental Science*, *9*, 756378. https://doi.org/10.3389/FENVS.2021.756378/BIBTEX





- Kindler, R., Miltner, A., Thullner, M., Richnow, H. H., & Kästner, M. (2009). Fate of bacterial biomass derived fatty acids in soil and their contribution to soil organic matter. *Organic Geochemistry*, 40(1), 29–37. https://doi.org/10.1016/J.ORGGEOCHEM.2008.09.005
 - KM Geyer, E. K.-S. A. G. S. F. (2016). Microbial carbon use efficiency: accounting for population, community, and ecosystem-scale controls over the fate of metabolized organic matter. *Biogeochemistry*, 127, 173–188.
- 490 Kögel-Knabner, I. (2002). The macromolecular organic composition of plant and microbial residues as inputs to soil organic matter. *Soil Biol. Biochem.*, *34*, 139–162.
 - Landsberg, J. J., & Gower, S. T. (1997). Carbon Balance of Forests. *Applications of Physiological Ecology to Forest Management*, 125–160. https://doi.org/10.1016/B978-012435955-0/50005-4
- Liang, C., Amelung, W., Lehmann, J., & Kästner, M. (2019). Quantitative assessment of microbial necromass contribution to soil organic matter. *Global Change Biology*, 25(11), 3578–3590. https://doi.org/10.1111/GCB.14781
 - Lützow, M. V., Kögel-Knabner, I., Ekschmitt, K., Matzner, E., Guggenberger, G., Marschner, B., & Flessa, H. (2006). Stabilization of organic matter in temperate soils: Mechanisms and their relevance under different soil conditions A review. *European Journal of Soil Science*, *57*(4), 426–445. https://doi.org/10.1111/J.1365-2389.2006.00809.X
- Manzoni, S., Čapek, P., Porada, P., Thurner, M., Winterdahl, M., Beer, C., Brüchert, V., Frouz, J., Herrmann, A. M., Lyon, S.
 W., SantrÅčkova, H., Vico, G., & Way, D. (2018). Reviews and syntheses: Carbon use efficiency from organisms to ecosystems Definitions, theories, and empirical evidence. *Biogeosciences*, 15(19), 5929–5949. https://doi.org/10.5194/BG-15-5929-2018
 - Manzoni, S., Taylor, P., Richter, A., Porporato, A., & Ågren, G. I. (2012). Environmental and stoichiometric controls on microbial carbon-use efficiency in soils. *New Phytologist*, *196*(1), 79–91. https://doi.org/10.1111/j.1469-8137.2012.04225.x
 - Masiello, C. A., Gallagher, M. E., Randerson, J. T., Deco, R. M., & Chadwick, O. A. (2008). Evaluating two experimental approaches for measuring ecosystem carbon oxidation state and oxidative ratio. *Journal of Geophysical Research: Biogeosciences*, 113(G3), 3010. https://doi.org/10.1029/2007JG000534
- McDaniel, M. D., Grandy, A. S., Tiemann, L. K., & Weintraub, M. N. (2014). Crop rotation complexity regulates the decomposition of high and low quality residues. *Soil Biology and Biochemistry*, 78, 243–254. https://doi.org/10.1016/J.SOILBIO.2014.07.027
 - MF Cotrufo, M. W. C. B. K. D. E. P. (2013). The Microbial Efficiency-Matrix Stabilization (MEMS) framework integrates plant litter decomposition with soil organic matter stabilization: do labile plant inputs form stable soil organic matter? *Glob. Change Biol.*, 19, 988–995.
- Miltner, A., Bombach, P., Schmidt-Brü Cken, B., & Kästner, M. (2011). SYNTHESIS AND EMERGING IDEAS SOM genesis: microbial biomass as a significant source. https://doi.org/10.1007/s10533-011-9658-z
 - Mukul, S. A., Halim, Md. A., & Herbohn, J. (2020). Forest Carbon Stock and Fluxes: Distribution, Biogeochemical Cycles, and Measurement Techniques. 1–16. https://doi.org/10.1007/978-3-319-71065-5_23-1



530



- Paul, E. (2016). The nature and dynamics of soil organic matter: plant inputs, microbial transformations, and organic matter stabilization. *Soil Biol. Biochem.*, *98*, 109–126.
 - Paul, E. A. (2016). The nature and dynamics of soil organic matter: Plant inputs, microbial transformations, and organic matter stabilization. *Soil Biology and Biochemistry*, 98, 109–126. https://doi.org/10.1016/J.SOILBIO.2016.04.001
 - Roels, J. A. (1980). Application of Macroscopic Principles To Microbial Metabolism. *Biotechnology and Bioengineering*, 22(12), 2457–2514. https://doi.org/10.1111/j.1749-6632.1981.tb14182.x
- Sánchez-Cañete, E. P., Barron-Gafford, G. A., & Chorover, J. (2018). A considerable fraction of soil-respired CO2 is not emitted directly to the atmosphere. *Scientific Reports 2018 8:1*, 8(1), 1–10. https://doi.org/10.1038/s41598-018-29803-x
 - Santos-Medellín, C., Blazewicz, S. J., Pett-Ridge, J., Firestone, M. K., & Emerson, J. B. (2023). Viral but not bacterial community successional patterns reflect extreme turnover shortly after rewetting dry soils. *Nature Ecology & Evolution* 2023 7:11, 7(11), 1809–1822. https://doi.org/10.1038/s41559-023-02207-5
 - Scharlemann, J. P. W., Tanner, E. V. J., Hiederer, R., & Kapos, V. (2014). Global soil carbon: Understanding and managing the largest terrestrial carbon pool. *Carbon Management*, 5(1), 81–91. https://doi.org/10.4155/CMT.13.77/SUPPL_FILE/TCMT_A_10816421_SM0001.DOC
- Schimel, J. P., & Weintraub, M. N. (2003). The implications of exoenzyme activity on microbial carbon and nitrogen limitation in soil: a theoretical model. *Soil Biology and Biochemistry*, *35*(4), 549–563. https://doi.org/10.1016/S0038-0717(03)00015-4
 - Schimel, J., & Schaeffer, S. (2012). Microbial control over carbon cycling in soil. Frontiers in Microbiology, 3.
 - Sinsabaugh, R. L., Manzoni, S., Moorhead, D. L., & Richter, A. (2013). Carbon use efficiency of microbial communities: Stoichiometry, methodology and modelling. *Ecology Letters*, 16(7), 930–939. https://doi.org/10.1111/ELE.12113
- 540 Smart, K. E. (2024). RQ-CUE_EGU_Biogeosciences (Repository). Zenodo. https://doi.org/10.5281/zenodo.11507912
 - Vance, E. D., Brookes, P. C., & Jenkinson, D. S. (1987). An extraction method for measuring soil microbial biomass C. *Soil Biology and Biochemistry*, 19(6), 703–707. https://doi.org/10.1016/0038-0717(87)90052-6
 - W Wieder, A. G. C. K. G. B. (2014). Integrating microbial physiology and physiochemical principles in soils with the MIcrobial-MIneral Carbon Stabilization (MIMICS) model. *Biogeoscience*, 11, 3899–3917.
- Walker, T. W. N., Kaiser, C., Strasser, F., Herbold, C. W., Leblans, N. I. W., Woebken, D., Janssens, I. A., Sigurdsson, B. D., & Richter, A. (2018). Microbial temperature sensitivity and biomass change explain soil carbon loss with warming.

 Nature Climate Change 2018 8:10, 8(10), 885–889. https://doi.org/10.1038/s41558-018-0259-x
 - Wang, C., Lingrui Qu, |, Yang, L., Liu, D., Morrissey, E., Miao, R., Liu, Z., Wang, Q., Fang, Y., & Bai, | Edith. (2021). Large-scale importance of microbial carbon use efficiency and necromass to soil organic carbon. *Glob Change Biol*, 27, 2039–2048. https://doi.org/10.1111/gcb.15550
 - Williamson, K. E. (2011). Soil Phage Ecology: Abundance, Distribution, and Interactions with Bacterial Hosts. 113–136. https://doi.org/10.1007/978-3-642-14512-4_4





- Williamson, K. E., Fuhrmann, J. J., Wommack, K. E., & Radosevich, M. (2017). Viruses in Soil Ecosystems: An Unknown Quantity Within an Unexplored Territory. *Annual Review of Virology*, 4(Volume 4, 2017), 201–219. https://doi.org/10.1146/ANNUREV-VIROLOGY-101416-041639/CITE/REFWORKS
- Williamson, K. E., Radosevich, M., & Wommack, K. E. (2005). Abundance and diversity of viruses in six Delaware soils. *Applied and Environmental Microbiology*, 71(6), 3119–3125. https://doi.org/10.1128/AEM.71.6.3119-3125.2005/ASSET/0241A010-5F52-4D29-9384-D6D4D5EB8861/ASSETS/GRAPHIC/ZAM0060555220003.JPEG
- Wu, R., Davison, M. R., Gao, Y., Nicora, C. D., Mcdermott, J. E., Burnum-Johnson, K. E., Hofmockel, K. S., & Jansson, J.
 K. (2021). Moisture modulates soil reservoirs of active DNA and RNA viruses. *Communications Biology*, 4(1). https://doi.org/10.1038/S42003-021-02514-2