



1	Changes in Aerosol/Gas-Phase Distribution Ratio of Semi-Volatile Products
2	Affect Secondary Organic Aerosol Formation with NOx from $\alpha$ -Pinene
3	Photooxidation
4	
5 6	Shijie Liu <sup>1,2</sup> , Xinbei Xu <sup>1</sup> , Si Zhang <sup>1</sup> , Rongjie Li <sup>1</sup> , Zheng Li <sup>1</sup> , Can Wu <sup>1</sup> , Rui Li <sup>1</sup> , Guiqin Zhang <sup>2</sup> , and Gehui Wang <sup>1,3*</sup>
7	
8 9 10	<sup>1</sup> Key Laboratory of Geographic Information Science of the Ministry of Education, School of Geographic Sciences, East China Normal University, Shanghai 200241, China
11 12	<sup>2</sup> Resources and Environment Innovation Institute, Shandong Jianzhu University, Jinan 250101, China
13 14 15 16 17 18 19 20 21 22	<sup>3</sup> Institute of Eco-Chongming, Cuiniao Road, Chenjia Zhen, Chongming, Shanghai 202150, China
23 24	*Correspondence to: Prof. Gehui Wang (ghwang@geo.ecnu.edu.cn)





# 25 Abstract

26	Atmospheric $\alpha$ -pinene is one of the most important precursors of secondary
27	organic aerosols (SOA). The formation of $\alpha$ -pinene derived SOA is strongly affected
28	by NOx. However, we still do not comprehensively understand the effects of $NO_{\boldsymbol{x}}$ on
29	$\alpha\text{-pinene}$ derived SOA formation. Therefore, we conducted $\alpha\text{-pinene}$ photooxidation
30	experiments in an atmospheric chamber at different NO <sub>x</sub> concentrations. The yields of
31	$\alpha\mbox{-pinene}$ SOA increased with $NO_x$ concentration under low-NO_x conditions, but were
32	suppressed under high-NOx conditions. The maximum SOA yields were $8.0\%$ and $26.2\%$
33	in the low- and high-volatility organic compound (VOC) experiments, respectively. We
34	found the increased SOA yields under low-NOx conditions were related to increased
35	consumption of $\alpha$ -pinene. The products of $\alpha$ -pinene photooxidation were mainly semi-
36	volatile, and the change in the aerosol/gas-phase distribution ratio as the formation of
37	$\alpha$ -pinene photooxidation products increased was identified as the main reason for the
38	enhanced SOA yields with increasing NOx. The sensitivity of the SOA yield to changes
39	in NOx and VOCs under different experimental conditions was also analyzed. This
40	study also quantified the nitrogen-containing organic compound (NOC) concentrations.
41	The mass fraction of NOCs in SOA increased monotonically with NOx in the $\alpha\mbox{-pinene}$
42	photooxidation process, and the maximum NOC mass fraction made up as much as
43	two-fifths of the $\alpha$ -pinene SOA.

44

45





### 46 **1. Introduction**

47 Fine particulate matter (PM<sub>2.5</sub>) is an important atmospheric pollutant. Recently,  $PM_{2.5}$  has attracted attention because of its negative impacts on human health, air 48 quality, atmospheric chemical reactions, and even climate radiation balances (Bellouin 49 et al., 2020; Wang et al., 2016). Secondary organic aerosols (SOA), which form via the 50 atmospheric photooxidation of volatile organic compounds (VOCs), are some of the 51 most significant sources of PM2.5 (Lv et al., 2022; Mcfiggans et al., 2019; Zhang et al., 52 2018). However, due to the diversity of SOA production processes and influencing 53 54 factors, much researches are still needed for the fully understand the formation mechanisms of SOA. 55

NO<sub>x</sub>, which is mainly emitted from human activities, is a key substance that 56 facilitates photooxidation. Differences in concentrations of NO<sub>x</sub> can lead to significant 57 changes in photooxidation processes and further affect SOA formation (Sarrafzadeh et 58 al., 2016; Nussbaumer et al., 2022). Organic peroxyl radicals (RO<sub>2</sub>) are the main 59 intermediate oxidation product when VOCs are oxidized by OH. The interaction 60 between NO<sub>x</sub> and RO<sub>2</sub> plays a key role in the SOA formation process, and nonlinear 61 relationships between NO<sub>x</sub> and SOA yields have been observed in a large number of 62 experimental investigations (Xu et al., 2021; Zhao et al., 2018; Sarrafzadeh et al., 2016; 63 Xu et al., 2014). Hydroperoxides (ROOH), which are formed from the reactions of RO<sub>2</sub> 64 with HO2 or RO2, have low volatilities and are responsible for SOA formation. However, 65 as the concentration of NOx increases, RO2 reacts with NO and is rapidly converted to 66 67 alkoxy radicals (RO). RO act as an intermediate and goes on to produce more high volatility compounds, which suppresses SOA formation (Sarrafzadeh et al., 2016; 68 Atkinson, 2000). Highly oxygenated organic molecules (HOMs) formed by the 69 70 autoxidation of RO<sub>2</sub> are also key compounds involved in SOA formation (Rissanen, 71 2021). HOM formation is suppressed when the  $NO_2$  concentration increases and  $RO_2$ 72 + NO becomes the dominant sink of RO2 in the photooxidation process, which also 73 contributes to the inhibition of SOA formation as NOx concentrations increase.





74 Clearly, the effects of NO<sub>x</sub> on photooxidation and SOA formation are quite complicated, but they have been widely studied in controlled chamber experiments. 75 Based on the semi-volatile partitioning theory in SOA formation, it has been established 76 77 that SOA yield is a function of SOA mass concentration when other experimental conditions are held constant (Odum et al., 1996; Takeuchi et al., 2022). Furthermore, 78 the mass concentration of formed SOA is directly related to the mass concentration of 79 available VOC precursors. However, the SOA yield is often discussed as a constant, 80 and the nonlinear relationships between SOA yield and initial NOx concentration 81 reported in chamber studies do not account for the consumption of VOCs (Chen et al., 82 2022b; Aruffo et al., 2022; Qi et al., 2020). To control for the impact of the reacted 83 VOC concentration on SOA yields under different NOx conditions, Sarrafzadeh et al. 84 (2016) conducted experiments that maintained similar levels of VOC consumption. 85 However, due to differences in OH concentrations during photooxidation, there were 86 87 clear differences in the reaction times in each experiment. Under normal circumstances, when exploring the impact of NOx on SOA yield, SOA yields are normalized based on 88 the reacted VOCs concentration, but this does not account for differences in the 89 90 consumption of VOCs. This approach inevitably overlooks the impact of semi-volatile partitioning theory in SOA formation under different NOx conditions. The roles of 91 92 chemical processes are often considered due to the impacts of NOx on SOA yields, but physical processes in SOA formation are equally significant and should be given more 93 attention. 94

α-Pinene is one of the most abundant monoterpene VOCs in the atmosphere 95 (Sindelarova et al., 2014; Guenther et al., 2012). Due to the high concentrations and 96 97 high SOA formation potential,  $\alpha$ -pinene is one of the most important sources of SOA in the atmosphere (Xu et al., 2015a; Zhang et al., 2018). In fact,  $\alpha$ -pinene derived SOA 98 99 is commonly utilized in atmospheric models to represent biogenic SOA (Henry et al., 100 2012). However, further in-depth research on the mechanisms of  $\alpha$ -pinene SOA formation is required to improve the fundamental basis and accuracy of atmospheric 101 models. In order to fully examine the effects of NOx on SOA formation, we used an 102





indoor chamber to investigate the yields of α-pinene derived SOA with different initial
NOx concentrations. The relationships among NO<sub>x</sub> concentration, VOCs consumption,
distribution of photooxidation products, and SOA mass yield were considered. Based
on the semi-volatile partitioning theory, this study aimed to better characterize the
mechanisms by which NO<sub>x</sub> affect SOA yield. **2. Experimental methods**

### 109 **2.1. Chamber studies**

A series of a-pinene photooxidation experiments initiated by NOx were performed 110 111 in a temperature controlled photooxidation chamber. The chamber and its characterization capabilities have been described in detail in our previous studies (Liu 112 et al., 2021a; Liu et al., 2021b). In brief, the photooxidation chamber was constructed 113 114 of Teflon-FEP film (0.06 mm) and surrounded with black light lamps (GE F40BLB) as the light source for the photooxidation reaction. The black light lamps provided an 115 effective light intensity (J<sub>NO2</sub>) of 0.165 min<sup>-1</sup> at full illumination. Before each 116 117 experiment, the chamber was cleaned by first evacuating all air and then filling it with purified air. This filling-purging cycle was repeated 5 times between experiments to 118 ensure the residual particulate and  $\alpha$ -pinene concentrations were less than 5 cm<sup>-3</sup> and 119 0.5 ppb, respectively. 5 m<sup>3</sup> zero air, which was supplied by a zero-air generator (111-120 D3N, Thermo Scientific<sup>TM</sup>, USA), was introduced into the chamber for the 121 photooxidation experiment. Analytically-pure, liquid  $\alpha$ -pinene (Sigma-Aldrich) was 122 injected into a Teflon tube using a micro syringe, and then the  $\alpha$ -pinene was evaporated 123 into gas and flushed into the chamber with zero air. NOx (Air Liquid Shanghai, 510 ppm 124 NO<sub>2</sub> in N<sub>2</sub>) were introduced directly into the chamber. After all the reactants were well 125 mixed, photooxidation of α-pinene was initiated by turning the black light lamps on. 126 All the experiments were conducted in dry conditions with a relative humidity of 15  $\pm$ 127 5 %. Both OH and O3 were formed under black light irradiation in the presence of NOx. 128 The detailed conditions for the α-pinene photooxidation experiments are listed in Table 129 130 1. The concentrations of  $\alpha$ -pinene were kept as constant as possible across different





- 131 experiments to ensure the effects of NO<sub>x</sub> were not obscured. No seed particles were
- 132 used in the chamber experiments.

#### 133

1.54 <b>Table 1.</b> Details of the <i>u</i> -difference $1.00x$ systems used in the difference $1.00x$ systems	Details of the $\alpha$ -pinene/NO <sub>x</sub> systems used in the photooxidation ex	periments.
---	---	------------

No.	α-pinene (ppb)	Δα-pinene (ppb)	NOx (ppb)	[VOCs] <sub>0</sub> /[NOx] <sub>0</sub>	SOA mass conc. <sup>a</sup> (µg m <sup>-3</sup> )	SOA yield
Exp. 1	116.2	67.3	12	9.7	26.0	6.5%
Exp. 2	117.2	96.2	25	4.7	43.8	7.7%
Exp. 3	117.8	115.3	68	1.7	54.3	8.0%
Exp. 4	119.3	117.6	150	0.8	34.6	4.9%
Exp. 5	116.9	114.9	337	0.3	15.3	2.2%
Exp. 6	115.6	112.8	600	0.2	11.4	1.7%
Exp. 7	258.3	70.7	8	32.3	20.9	5.0%
Exp. 8	251.0	168.7	26	9.7	144.0	14.5%
Exp. 9	263.4	234.0	52	4.1	370.3	26.2%
Exp. 10	250.8	245.3	113	2.2	347.6	23.7%
Exp. 11	247.8	244.2	237	1.5	253.5	17.1%
Exp. 12	260.0	256.0	369	0.7	226.4	14.6%
Exp. 14	254.2	249.9	669	0.4	175.0	11.8%

135 <sup>a</sup> All mass concentrations were wall-loss corrected.

### 136 **2.2. On-line instruments**

137 The concentration of NOx was constantly monitored with a NO/NO2/NOx analyzer (Model 42C, Thermo Corporation, USA). The  $\alpha$ -pinene concentration was measured 138 throughout the photooxidation process with a proton-transfer-reaction time-of-flight 139 mass spectrometer (PTR-tof-MS, Ionicon Analytik, Austria) using hydronium (H<sub>3</sub>O<sup>+</sup>) 140 ions. The drift tube of the PTR-tof-MS was operated at 60.0 °C (Tdrift), 2.30 mbar (Pdrift), 141 and 600V (Udrift). Three ions which are commonly observed in ambient air, i.e., H<sub>3</sub>O<sup>+</sup> 142 (21.0226 m/z), NO<sup>+</sup> (29.9980 m/z), and C<sub>3</sub>H<sub>6</sub>O<sup>+</sup> (59.0497 m/z), were used for PTR-tof-143 MS mass calibration. 144

The formation and chemical characteristics of bulk SOA were measured with a high-resolution time-of-flight aerosol mass spectrometer (HR-tof-AMS, Aerodyne Research Inc., USA). During which, the AMS was only operated in the V-mode in order to avoid W-mode data too noisy to analyze. The ionization efficiency (IE) of the AMS was calibrated via monodisperse with dried ammonium nitrate (AN) aerosols (300 nm).





- 150 The SOA mass concentrations obtained from AMS measurements were compared and corrected through a scanning mobility particle sizer. For SOA mass concentration 151 calculations, the relative ionization efficiencies (RIE) of 1.4 and 1.1 were applied for 152 153 organic compounds and nitrate, respectively (Middlebrook et al., 2012). The AMS data were analyzed based on the standard applications of SQUIRREL 1.51H and PIKA 154 1.10H in Igor Pro (WaveMetrics, Inc., USA).  $NO_2^+(m/z = 46)/NO^+(m/z = 30)$  was used 155 to differentiate inorganic and organic nitrate in the AMS measurements (Xu et al., 156 2015b; Kiendler-Scharr et al., 2016; Ng et al., 2017). 157
- 158 **3. Result and discussion**

### 159 **3.1 Effect of NOx on α-pinene SOA formation**

The relationship between SOA yields and NOx is shown in Fig. 1. Here, SOA yield 160 161 was calculated as the SOA mass concentration divided by the reacted VOCs. Clearly, the SOA yield increased rapidly at first, reached a maximum and then decreased 162 163 gradually with increasing initial NOx concentration under both high- and low-VOC conditions. In this study, we defend the positive correlation between SOA yield and 164 NOx as low-NOx conditions, and a negative correlation of them as high-NOx 165 conditions. In the low- and high-VOC experiments, the maximum SOA yields were 166 8.0% and 26.2% respectively. Since the change trend in SOA yield with NOx was not 167 affected by initial VOC concentration, the effect of NOx on SOA formation is discussed 168 in this section using the data from the low-VOCs experiments. The concentration of 169 VOCs in the chamber is generally higher than that in the real atmospheric environment. 170 Lower VOC concentrations would be closer to their levels in the actual atmosphere, 171 which would make experimental results more representative of the real atmospheric 172 environment. 173







176 177

174

175

Like our study, similar relationships between SOA yields and initial NOx 178 concentrations have been widely observed in previous studies (Aruffo et al., 2022; Liu 179 et al., 2019b; Zhao et al., 2018; Lane et al., 2008; Ng et al., 2007; Kroll et al., 2006). 180 181 RO<sub>2</sub> was the main intermediate in the photooxidation of VOCs and the concentration 182 and fate of RO<sub>2</sub> depended on the NO<sub>x</sub> concentrations. RO<sub>2</sub> mainly reacted with NO and was rapidly converted to alkoxy radicals (RO) under high-NO<sub>x</sub> conditions. RO 183 intermediates produce more high-volatility compounds than hydroperoxides (ROOH), 184 which form through the reaction of RO2 with RO2/HO2, and decrease SOA yields 185 (Atkinson, 2000; Sarrafzadeh et al., 2016). Furthermore, the autooxidation of RO2 can 186 be inhibited through the RO<sub>2</sub> + NO / NO<sub>2</sub> reaction, and the reduction in HOMs further 187 contributes to the decreased SOA yields under high-NO<sub>x</sub> conditions (Yu et al., 2022; 188 189 Laskin et al., 2018).

However, the mechanisms by which SOA yields increase with increasing NO<sub>x</sub> concentrations under low-NO<sub>x</sub> conditions remain poorly understood (Camredon et al., 2007; Kroll et al., 2006). The atmospheric oxidizing capacity (AOC), which indicates the oxidizing ability of the atmosphere, is significantly influenced by NOx (Wang et al., 2023). In this study, we assessed AOC based on the decay ratio of VOCs (AOC<sub>t</sub> = -d





195  $[VOCs]_t / [VOCs]_t$ ). The time-dependent consumption curves of VOC and the decay ratios of VOCs under different NOx concentrations are shown in Fig. 2. All experiments 196 were allowed the same photooxidation time. For the low-NOx experiments, the 197 consumption of  $\alpha$ -pinene increased with NO<sub>x</sub> concentration. Only when the initial NO<sub>x</sub> 198 concentration was higher than 68 ppb was  $\alpha$ -pinene completely consumed by the end 199 of the photooxidation period. The average decay ratio of VOCs increased from 4.75×10<sup>-</sup> 200  $^{3}$  to  $4.53 \times 10^{-2}$  as the initial NOx increased from 12 ppb to 150 ppb, and gradually 201 decreased to 2.64×10<sup>-2</sup> with an initial NOx of 600 ppb. Hence, the AOC in the chamber 202 also showed a trend of first increasing and then decreasing with increasing NOx 203 concentration. Sarrafzadeh et al. (2016) noted that the OH recycling reaction NO + HO<sub>2</sub> 204  $\rightarrow$  NO<sub>2</sub> + OH was responsible for the NO<sub>x</sub>-induced increase in OH concentration. Chen 205 et al. (2022b) revealed that the increasing NOx level obviously enhanced the 206 atmospheric oxidation ability in a NOx-sensitive (low-NOx) regime. The higher 207 208 consumption of VOCs and the faster VOC oxidation rate both suggested there was a higher atmospheric oxidation ability in the reaction system (Ng et al., 2007). In addition, 209 NOx is an important sink for OH, so the competition between NOx and VOCs for OH 210 211 might have been responsible for the decreases in VOC consumption and AOC under high-NOx conditions. 212

Although there were similar trends in the yield of SOA and AOC as the NOx 213 concentration increased, no correlation between SOA yield and AOC was observed in 214 215 our study (Fig. S1). It has been shown that increases in AOC are essential drivers of increases in SOA mass concentration in the troposphere (Feng et al., 2019; Li et al., 216 217 2023). The higher SOA mass concentration under high AOC conditions was due to 218 more VOCs being oxidized, rather than an increase in SOA yield. Therefore, the increased AOC was not the direct mechanism by which increasing NOx concentration 219 220 influenced SOA yield.

> 221 222

> 223

224











The function of SOA yield with SOA mass concentration (M<sub>0</sub>) was recalculated based on Odum's SOA yield model (Fig. 3). In each experiment, an increase in M<sub>0</sub> directly resulted in an increase in SOA yield. Notably, if the photooxidation products can be classified as low-volatility oxidation products, the SOA yield should remain constant with increasing M<sub>0</sub> (Krechmer et al., 2015; Ehn et al., 2014; Odum et al., 1996). However, the gas-particle distribution coefficients of semi-volatile substances are





directly related to their concentrations, and the distribution coefficients of semi-volatile
substances into the aerosol phase are larger at higher concentrations (Akherati et al.,
2019; Odum et al., 1996). Hence, the increasing distribution ratios of semi-volatile
organic products between aerosol and gas phases at high M<sub>0</sub> were responsible for the
increasing SOA yields with increasing photooxidation time (Kolesar et al., 2015;
Valorso et al., 2011; Takeuchi et al., 2022).

237 A two-product model is an effective method for fitting the relationship between 238 SOA yield and  $M_0$  (Liu et al., 2019a; Odum et al., 1996). The model was calculated via 239 equation (1), shown below:

$$Y = M_0 \times \left(\frac{\alpha_1 K_{om,1}}{1 + K_{om,1} M_0} + \frac{\alpha_2 K_{om,2}}{1 + K_{om,2} M_0}\right)$$
(1)

where, M<sub>0</sub> is SOA mass concentration (mg m<sup>-3</sup>);  $\alpha_1$  and  $\alpha_2$  are the mass-based stoichiometric coefficients of species with low-volatility and semi-volatility products, respectively; and K<sub>om,1</sub> and K<sub>om,2</sub> (m<sup>3</sup> µg<sup>-1</sup>) are the gas-particle partitioning equilibrium constants for low-volatility and semi-volatility products, respectively. The  $\alpha_1$ ,  $\alpha_2$ , K<sub>om,1</sub>, and K<sub>om,2</sub> of  $\alpha$ -pinene SOA formed under different initial NOx concentrations are shown in Table 2, and the simulated SOA yields with changing M<sub>0</sub> are shown in Fig. 3.

- 246
- 247 248

Table 2 Parameters of the two product model for α-derived SOA under different initial NOx concentration.

Initial NOx conc. (ppb)	$\alpha_1$	$K_{om,1} (m^3 \mu g^{-1})$	$\alpha_2$	$K_{om,2} (m^3 \mu g^{-1})$
12	0.048	0.19	0.28	0.0040
25	0.038	0.19	0.30	0.0039
68	0.028	0.19	0.32	0.0037
150	0.019	0.19	0.33	0.0031
337	0.017	0.19	0.35	0.0019
600	0.014	0.19	0.38	0.0016

249

It was clear that the curves of SOA yield to  $M_0$  moved lower on the SOA yield axis in Fig. 3 with increasing NO<sub>x</sub> concentration. According to Table 2,  $\alpha_1$  decreased while  $\alpha_2$  increased as the initial NOx concentration increased. For the experiments conducted with 12 ppb NOx, the  $\alpha_1/\alpha_2$  ratio of  $\alpha$ -pinene derived SOA was 0.17. This  $\alpha_1/\alpha_2$  ratio decreased to 0.13, 0.086, 0.058, 0.049, and 0.037 as the initial NOx concentration





255 increased to 25, 68, 150, 337, and 600 ppb, respectively. Lower  $\alpha_1$  values indicate lower proportions of low-volatility products in the SOA, while higher  $\alpha_2$  values indicate 256 higher proportions of semi-volatility products in SOA. In addition, the Kom,2 value also 257 decreased, dropping continuously from 0.0040 to 0.0016 m<sup>3</sup>  $\mu$ g<sup>-1</sup> as the initial NOx 258 concentration increased from 12 to 600 ppb. This meant that the volatility of semi-259 volatility products produced through  $\alpha$ -pinene photooxidation increased with initial 260 NOx concentration. Overall, with increasing initial NOx concentration, the low-261 volatility products made up decreasing proportions (lower  $\alpha_1$ ) of the total products 262 while semi-volatility products made up increasing proportions (higher  $\alpha_2$ ) and higher 263 volatilities (lower Kom,2). 264

To determine why the volatility of SOA increase with increasing NOx, we 265 analyzed the consumption of VOCs under different NOx conditions (Fig. 2). Due to the 266 lower consumption rate of VOCs and low AOC, a-pinene was not completely 267 consumed at the end of the photooxidation period under low-NOx conditions. Increased 268 consumption of VOCs will lead to higher concentrations of photooxidation products 269 generated in the chamber, so when the initial NOx concentration increased from 12 ppb 270 to 25 ppb to 68 ppb, the SOA mass concentration increased from 26.0 µg m<sup>-3</sup> to 43.8 271  $\mu$ g m<sup>-3</sup> to 54.3  $\mu$ g m<sup>-3</sup>. As mentioned above, SOA yield increased with the mass 272 concentration of SOA. Even if the fitting curve of the two-product model gradually 273 moved lower on the graph with increases in NOx, the SOA yield still increased from 274 275 6.5% to 8.0% when the initial NOx concentration increased from 12 ppb to 68 ppb. Hence, under low-NOx conditions, more SOA was generated due to the increased AOC 276 277 and VOC consumption with increasing NOx. The enhancement of the SOA yield with 278 increasing NOx concentrations can be attributed to the increased ratio of the aerosol/gas phase distribution resulting from higher concentrations of semi-volatile photooxidation 279 280 products. Chen et al. (2022c) compared the relative content of intermediate products 281 with different volatilities and found that the proportion of semi-volatile oxidized products in gas-phase intermediate products was lower when experiments had higher 282 VOC consumption and SOA yields. Assuming that the proportions of different volatile 283





oxidation products remain constant, the smaller proportion of semi-volatile organic products in the gas phase suggested that a larger proportion of semi-volatile organic products was condensed into the particulate phase when more VOCs were consumed. This result also supported our conclusion that the increased aerosol/gas phase distribution ratio of semi-volatile products was the dominant driver underlying the enhanced SOA yields with increased VOC consumption under low-NOx conditions.

It should be noted that if the SOA yield is simulated based only on the two-product 290 291 model, the SOA yield should increase by 34% and 51% when NOx increases from 12 ppb to 25 ppb and 68 ppb, respectively. However, in reality, the SOA yield only 292 increased 18% and 23%, respectively (Fig. S2). The real SOA yields were not only 293 lower than the simulated yields, but the discrepancy between the real and simulated 294 SOA yields increased with increasing NOx concentration. As mentioned earlier, 295 reactions between RO2 and NOx form highly volatile oxidized compounds. Therefore, 296 the increasing discrepancy between real and simulated SOA yields with increasing NOx 297 concentration indicated that inhibition of SOA formation through the RO2 + NOx 298 299 reaction pathway was always occurring. This indicated that, under low-NOx conditions, the increase in SOA yield resulted from a combination of effects, the positive effect of 300 the increased aerosol/gas phase distribution ratios of semi-volatile products and the 301 negative effect of the  $RO_2 + NOx$  reaction. 302

303

304 305

306

307







Figure 3. SOA yields as a function of organic aerosol mass concentration  $M_0$  of  $\alpha$ -pinene at different initial NO<sub>x</sub> concentrations. The simulated SOA yields based on the two-product model are shown by the solid lines.

The effects of NOx on the yields and chemical properties of SOA have been well 308 309 documented in previous studies (Chen et al., 2020; Eddingsaas et al., 2012; Sarrafzadeh et al., 2016; Zhao et al., 2018). However, few studies have focused on how NO<sub>x</sub> 310 influences VOC consumption during photooxidation, so the relationship between the 311 concentrations of NOx and oxidation products has not been fully elucidated. For SOA, 312 which is mainly composed of semi-volatile oxidation products, the gas-particle 313 314 distribution coefficients of semi-volatile substances are directly related to their concentrations. Therefore, SOA yield can be affected by the concentrations of semi-315 316 volatile oxidation products, and this should be taken into consideration when studying 317 SOA formation process. Hence, future studies of photooxidation and SOA formation should pay more attention to the formation processes of semi-volatile substances. 318

### 319 **3.2 SOA chemical composition at different NO<sub>x</sub> concentrations**

320 SOA chemistry must be studied at the molecular level to better understand the 321 characteristics of SOA formation. The bulk chemical properties of SOA generated







322 under low- and high-NO<sub>x</sub> concentrations are shown in Fig. 4.



326 327

323

324 325





328 Organic fragments of  $C_2H_3O^+(m/z = 43)$ , which originated from the fragmentation of aldehydes and ketones, were the dominant peaks in the AMS. The strong organic 329 AMS signal at m/z = 43 (f<sub>43</sub>) indicated that the SOA oxidation level was relatively low. 330 331 The intensity of the  $CO_2^+$  (m/z = 44), representing the thermal decarboxylation of organic acid groups, was rather low, which indicated there was low carboxylic acid 332 content in the  $\alpha$ -pinene derived SOA. Carboxylic acid has a relatively low volatility and 333 enters the particulate phase relatively easily, which means that the carboxylic acid 334 produced through the photooxidation of  $\alpha$ -pinene was likely less than that detected in 335 the AMS. In our study, the  $f_{43}$  and  $f_{44}$  of  $\alpha$ -pinene SOA ranged from 0.160 to 0.175 and 336 from 0.069 to 0.074, respectively. According to the "triangle plot" of the AMS, the SOA 337 derived from a-pinene photooxidation mainly fell in the lower area designated semi-338 volatile oxygenated organic aerosols (SV-OOA) (Ng et al., 2010). The AMS results 339 provided direct evidence that semi-volatile products were the main components of α-340 341 pinene SOA formed through NO<sub>x</sub> photooxidation. This result further supported our 342 observations in Section 3.1, wherein the aerosol/gas-phase distribution ratio of photooxidation products increased with increasing SOA mass concentration and VOC 343 344 consumption, resulting in the enhancement of SOA yield with increasing NO<sub>x</sub> at low-NOx conditions. 345

The differences in the AMS spectra between low- and high-NOx conditions are 346 compared in Fig. 4 (c). Although the inhibitory effect of the reaction between NOx and 347 348 RO<sub>2</sub> on SOA formation was confirmed, the chemical compositions of the SOA formed under different NOx conditions showed only minor differences. Photooxidation 349 350 products formed through the reaction of NO with RO2 radicals usually have relatively 351 high volatilities. However, the  $\alpha$ -pinene derived SOA was mainly composed of semivolatile photooxidation products, and these photooxidation products formed under 352 353 high-NOx conditions with higher volatilities may not easily be condensed into the particulate phase, so they were mainly present in the gas phase. This means that the 354 high volatility products formed via the RO<sub>2</sub> + NO reaction path under high-NOx 355 356 conditions did not affect the chemical composition of SOA. Therefore, the AMS spectra





of the α-pinene SOA were not significantly different between low- and high-NOxconditions.

In addition to the generation of RO, the reaction of RO2 with NO can also form N-359 containing compounds (NOCs). The NOC contents in SOA under different NOx 360 conditions were calculated based on the AMS fragments of NO<sup>+</sup> and NO<sub>2</sub><sup>+</sup>. NO<sup>+</sup> and 361 NO2<sup>+</sup> are characteristic fragments of NOCs, but the same NO<sup>+</sup> and NO2<sup>+</sup> fragments can 362 also be detected by AMS in the forms of inorganic nitrates. In order to distinguish 363 364 organic nitrates from total nitrates, we used the same method as described in our previous study based on the differences in the ratios of NO<sup>+</sup>/NO<sub>2</sub><sup>+</sup> for organic and 365 inorganic nitrates in the AMS mass spectra (Xu et al., 2015b; Ng et al., 2017; Day et 366 al., 2022). The concentrations of NO<sup>+</sup> and NO<sub>2</sub><sup>+</sup> from organic nitrates were calculated 367 using the following equations: 368

$$NO_{2,org} = \frac{NO_{2,meas} \times (R_{meas} - R_{AN})}{R_{ON} - R_{AN}}$$
(4)

$$NO_{org} = R_{ON} \times NO_{2,org}$$
(5)

369 where  $R_{\text{meas}}$  is the NO<sup>+</sup>/NO<sub>2</sub><sup>+</sup> ratio from the AMS results;  $R_{\text{ON}}$  and  $R_{\text{AN}}$  are the 370 NO<sup>+</sup>/NO<sub>2</sub><sup>+</sup> ratios for organic nitrate and ammonium nitrate, respectively; and NO<sub>2,meas</sub> 371 represents the total NO2<sup>+</sup> fragments obtained from the AMS data. In this study, R<sub>AN</sub> was about 1.1 based on the measurements and RON was assumed to be about 10 referring to 372 the study by Takeuchi and Ng (2019). Previous studies have shown monoterpene 373 hydroxyl nitrate (m/z = 215), pinene keto nitrate (m/z = 229), and monoterpene 374 dicarbonyl nitrate (m/z = 247) to be the main NOCs from  $\alpha$ -pinene photooxidation 375 (Chen et al., 2022a; Li et al., 2018). Here, the formation of NOCs was calculated based 376 on organic nitrate contents and the molecular weights of NOCs. 377







#### 378 379

380

Figure 5 Content and mass concentration of NOCs under different initial NOx conditions.

381 The contents and mass concentrations of NOCs under different NOx levels are shown in Fig. 5. The increase in NOC contents with NOx could be divided into two 382 383 stages. The first stage occurred when the initial NOx concentration increased from 12 384 ppb to 150 ppb. During this stage the content of NOCs in the SOA increased linearly, indicating that the NOx was the limiting factor for NOC formation in low-NOx 385 386 conditions. The second stage occurred as NOx concentrations continued to rise to 600 ppb. During this stage, the growth in NOC content gradually slowed while approaching 387 the maximum value. Based on the nonlinear fit between NOx concentration and NOC 388 content, the maximum value of NOCs content in SOA was predicted to be about 39  $\pm$ 389 390 3.8%. This indicated that, under extremely high-NOx conditions, NOCs could account for up to two-fifths of the α-pinene SOA. These results not only showed that NOCs play 391 a more important role in biogenic SOA formation under high-NOx conditions, but can 392 also serve as a basis for estimating the generation of NOCs in SOA. 393

## **394 3.3 Effect of initial VOC concentration on SOA formation**

395 The SOA yields with different  $\alpha$ -pinene concentrations were also compared in this 396 study. As shown in Fig. 1, the change trends in SOA yields with increasing NO<sub>x</sub>

18





397 concentrations were similar across all experiments, including low- and high-VOC conditions. However, higher initial VOC concentrations led to increases in both 398 maximum SOA yield and the initial NOx concentration required for the SOA yield to 399 reach its maximum value. Under low-VOCs conditions, based on the no-linear fitting 400 of SOA yield with initial NOx, the peak SOA yield was 8.8% when the initial NOx 401 concentration was 41 ppb. The maximum SOA yield increased to 26.7% under high-402 VOC conditions, and the concentration of NOx corresponding to the maximum SOA 403 yield increased to 66 ppb. Under high-VOC conditions, the maximum SOA mass 404 concentration was 370.3 µg m<sup>-3</sup>. However, the maximum mass concentration of SOA 405 formed under high-VOC conditions was only 54.3 µg m<sup>-3</sup>, which was 6.8 times lower 406 than that under high-VOCs. 407

Previous studies have suggested that the formation of SOA from  $\alpha$ -pinene 408 photooxidation is primarily limited by NOx concentration under low-NOx conditions, 409 and variation in VOC concentrations have little effect. However, as shown in Fig. 1, the 410 SOA yield under high-VOC conditions was consistently higher than that under low-411 VOC conditions. Moreover, the ratio and gap in SOA yields between high- and low-412 VOC conditions significantly increased with increasing NOx concentrations when NOx 413 concentrations were lower than 60 ppb (Fig. S3). Therefore, it can be concluded that 414 under low-NOx conditions, the formation of SOA from  $\alpha$ -pinene photooxidation is 415 limited by both VOCs and NOx concentration. Based on differences in VOC 416 417 consumption between high- and low-VOCs conditions (Fig. S4), it appeared that, at the same NOx concentration, the consumption of VOCs was higher under high-VOC 418 conditions and the difference in VOC consumption between high- and low-VOC 419 420 conditions gradually increased with increasing NOx. According to the aerosol/gasphase distribution ratio of semi-volatile products described in Section 3.1, higher VOC 421 422 consumption can generate more semi-volatile oxidation products, which can enhance the SOA yield. The continuously increasing concentration gradient in VOC 423 consumption between high- and low-VOC conditions indicated that the SOA yield 424 under low-NOx conditions was also affected by the concentration of VOCs. 425





426 When the initial NOx concentration was higher than 100 ppb, the growth rate of the SOA yield ratio between high- and low-VOC conditions was 67.7% lower than that 427 under low-NOx conditions. The ratio of SOA yield from high-VOC experiments was 428 about 3–8 times higher than that from the low-VOC experiments, which surpassed the 429 VOC ratio between different VOC conditions. This indicated that, under high-NOx 430 conditions, SOA yield has a high sensitivity to changes in VOC concentrations, and the 431 inhibition effect of reducing VOCs on SOA yields would be more pronounced under 432 high-NOx conditions. It should be noted that, while the difference in VOC consumption 433 did not change further under high-NOx conditions, the ratio of the SOA yields between 434 high- and low-VOCs conditions continued to increase with increasing NOx 435 concentrations. At lower VOC conditions, the RO2 + NO pathway was more 436 competitive, which enhanced the production of high-volatile products compared to in 437 the higher VOC experiments. This sustained increase in the ratio of SOA yields may 438 439 have been mainly due to the differences in volatilities of the oxidation products produced under high- and low-VOC conditions. Therefore, as illustrated in Fig. S3, it 440 can be inferred that changes in VOC concentrations have a more pronounced effect on 441 442 SOA yield under higher NOx conditions.

Although the trend in SOA yields with increasing NOx was similar under different 443 VOC conditions, the change trend in AOC with increasing NOx was different (Fig. S5). 444 First, under low-NOx conditions, the AOC under high-VOC conditions was lower than 445 446 when under low-VOCs. The consumption of  $\alpha$ -pinene was higher in the high-VOC experiments than in the low-VOC experiments and more oxidants were consumed by 447  $\alpha$ -pinene under high-VOC conditions, which may have led to the inhibition of AOC. 448 449 Secondly, AOC only increased with increasing NOx concentrations under high-VOC conditions. According to the fitting results in Fig. S5, under low-VOC conditions, AOC 450 451 initially increased with NOx and then decreased, with an inflection point at VOCs/NOx  $\approx$  0.47. However, under high-VOC conditions, AOC increased with initial NOx 452 concentration across the entire NOx gradient. Although the rate of change in AOC 453 gradually slowed as NOx increased, no inflection point occurred in the AOC change 454

> 460 461

> 462

463





455 trend even when NOx reached 669 ppb, at which point VOCs/NOx decreased to 0.38.
456 These results suggested that AOC is jointly affected by NOx and VOCs, and that
457 VOCs/NOx is not sufficient as a direct basis for the evaluation of AOC. This
458 demonstrated the complicated nature of accurately simulating AOC, which is an
459 important research topic that needs to be studied further.



Figure 6 Relationship between consumed VOCs and SOA yields under different NOx and VOC conditions

The SOA yield trends at different AOCs are shown in Fig. S1. Based on the 464 photooxidation experiments with different VOC concentrations, there was no clear 465 correlation between SOA yield and AOC. This result further explains why increases in 466 AOC are not the direct cause of increased SOA yields from a-pinene photooxidation. 467 468 We further compared the consumption of VOCs with SOA yield (Fig. 6) and found a significant positive correlation between the two. This result was consistent with the 469 study by Jiang et al. (2022) which reported that SOA yield increased with increasing 470 471 consumption of VOCs, even when the OH concentration was held constant across experimental photooxidation groups. The continuous growth trend in SOA yield with 472 473 increasing consumption of VOCs, which overlapped among different VOC conditions, 474 further supported the processes discussed in Section 3.1; that is to say, changes in the aerosol/gas-phase distribution ratio for semi-volatile products with changing VOC 475





- 476 concentrations lead to increased SOA yields with increasing NOx concentration under
- 477 low-NOx conditions.

However, when the NOx concentrations exceeded 68 ppb and 113 ppb under low-478 and high-VOC conditions, respectively, a-pinene was almost entirely consumed during 479 the photooxidation process. This meant that, although AOC continued to increase with 480 increasing NOx, further increases in AOC only accelerated VOC consumption time 481 without impacting total VOC consumption or photooxidation products formation. As a 482 483 result, the SOA yield did not continue to increase. In contrast, more NOx would promote the  $RO_2 + NO$  reaction, which would result in a decreasing trend in the SOA 484 yield. 485

### 486 **4.** Atmospheric implications

NO<sub>x</sub> is a key substrate for photooxidation in the atmosphere. Changes in NO<sub>x</sub> 487 concentrations can lead to significant changes in the photooxidation of VOCs and affect 488 subsequent SOA formation. In this study, SOA yields were observed to increase sharply 489 at first, and then decrease gradually with increasing NOx concentration. The maximum 490 SOA yields in the low- and high-VOC conditions were 8.0% and 26.2%, respectively. 491 Based on the relationship between SOA yield and M<sub>0</sub>, as well as the SOA chemical 492 composition, semi-volatile oxidation products were the main components of a-pinene 493 494 derived SOA. Under low-NOx conditions, increasing NOx concentration increased the consumption of VOCs during photooxidation and lead to an increase in the SOA yield. 495 We concluded that the combination of increased distribution ratios of the semi-volatile 496 oxidation products with increasing amounts formed was key in the promotion of SOA 497 yield with increasing NOx. Conversely, the effect of NOx on AOC did not directly cause 498 499 the increase in SOA yield. NOx can significantly increase the content of NOCs in SOA. At their highest, under extremely high-NOx concentrations, NOC contents made up as 500 much as two-fifths of  $\alpha$ -pinene derived SOA. 501

With the progression of urbanization and the changes to the natural environments
 surrounding urban areas, atmospheric environments of have become quite different (Xu





504	et al., 2015b; Domínguez-López et al., 2014; Shon et al., 2007; Agbo et al., 2022). Most
505	rural areas exhibit $\mathrm{NO}_{x}$ limitation, while urban areas more often exhibit VOC limitation
506	(Tan et al., 2018; Santos et al., 2021; Hui et al., 2018). In areas with low-NOx
507	concentrations, the consumption of biogenic VOCs might not be significant despite
508	there being high biogenic VOC emissions, which would limit the formation of SOA.
509	However, if unreacted VOCs are transported to areas with high-NOx concentrations (i.e.,
510	from rural to urban areas), or vice versa, there could be explosive increases in SOA
511	concentrations and noticeable changes of SOA chemical compositions. Therefore, the
512	regional transport of pollutants between urban and forested areas will significantly
513	impact the formation of atmospheric aerosols and environmental pollution. This was
514	supported by our study, which demonstrated how human activities can enhance SOA
515	formation from biogenic emissions.

Finally, as a typical goal, research often works toward creating or improving 516 models to simulate atmospheric phenomena. However, past simulations have generally 517 underestimated SOA concentrations when compared to field observations, especially 518 during pollution periods (Ling et al., 2022; Kelly et al., 2018). Our study noted that, 519 under high-NOx and high-VOC conditions, SOA yield exhibited significant changes 520 with changes in reaction conditions (Fig. S3). Under polluted conditions, the faster 521 consumption of VOCs might lead to underestimates of VOC concentrations in the 522 atmosphere and, subsequently, the generation of oxidation products, ultimately 523 524 resulting in underestimates of SOA yields. Hence, our research provides crucial information for improving the accuracy of air quality models simulating SOA formation. 525

526

### 527 Data availability

The data used to support the conclusions in this study are available at a public data
repository of Figshare via https://figshare.com/articles/dataset/NOx\_photooxidation\_pinene/25200929 (S. Liu, 2024)

531

### 532 Author contributions





- 533 SL and GW designed the whole work and wrote the paper. SL did the experiment, 534 collected the samples, conducted the sample analysis and performed the data 535 interpretation. All authors contributed to the paper with useful scientific discussions
- 536 and comments.
- 537

# 538 Competing interests

- 539 The authors declared that they have no conflict of interest.
- 540

#### 541 Acknowledgements

This work was funded by the National Natural Science Foundation of China (No.
42130704, U23A2030), the China Postdoctoral Science Foundation (No.
2022T150215), and ECNU Happiness Flower Project.

545

#### 546 **Reference**

- Agbo, K. E., Walgraeve, C., Vandermeersch, L., Eze, J. I., Ukoha, P. O., and Van Langenhove, H.:
  Residential VOCs concentration levels in Nsukka, Nigeria, Atmos. Environ., 289,
  10.1016/j.atmosenv.2022.119307, 2022.
- Akherati, A., Cappa, C. D., Kleeman, M. J., Docherty, K. S., Jimenez, J. L., Griffith, S. M., Dusanter, S.,
  Stevens, P. S., and Jathar, S. H.: Simulating secondary organic aerosol in a regional air quality model
  using the statistical oxidation model Part 3: Assessing the influence of semi-volatile and
  intermediate-volatility organic compounds and NOx, Atmos. Chem. Phys., 19, 4561-4594,
  10.5194/acp-19-4561-2019, 2019.
- Aruffo, E., Wang, J., Ye, J., Ohno, P., Qin, Y., Stewart, M., McKinney, K., Di Carlo, P., and Martin, S. T.:
   Partitioning of organonitrates in the production of secondary organic aerosols from α-pinene photooxidation, Environ. Sci. Technol., 56, 5421-5429, 10.1021/acs.est.1c08380, 2022.
- Atkinson, R.: Atmospheric chemistry of VOCs and NOx, Atmos. Environ., 34, 2063-2101,
   10.1016/S1352-2310(99)00460-4, 2000.
- Bellouin, N., Quaas, J., Gryspeerdt, E., Kinne, S., Stier, P., Watson-Parris, D., Boucher, O., Carslaw, K.
  S., Christensen, M., Daniau, A. L., Dufresne, J. L., Feingold, G., Fiedler, S., Forster, P., Gettelman,
  A., Haywood, J. M., Lohmann, U., Malavelle, F., Mauritsen, T., McCoy, D. T., Myhre, G.,
  Mulmenstadt, J., Neubauer, D., Possner, A., Rugenstein, M., Sato, Y., Schulz, M., Schwartz, S. E.,
  Sourdeval, O., Storelvmo, T., Toll, V., Winker, D., and Stevens, B.: Bounding global aerosol
  radiative forcing of climate change, Rev. Geophys., 58, e2019RG000660, 10.1029/2019RG000660,
  2020.
- Camredon, M., Aumont, B., Lee-Taylor, J., and Madronich, S.: The SOA/VOC/NOx system: an explicit
   model of secondary organic aerosol formation, Atmos. Chem. Phys., 7, 5599-5610, 10.5194/acp-7 5599-2007, 2007.
- 570 Chen, J., Wang, X., Zhang, J., Li, M., Li, H., Liu, Z., Bi, Y., Wu, D., Yin, X., Gu, R., Jiang, Y., Shan, Y.,
  571 Zhao, Y., Xue, L., and Wang, W.: Particulate organic nitrates at Mount Tai in winter and spring:
  572 Variation characteristics and effects of mountain-valley breezes and elevated emission sources,
  573 Environ. Res., 212, 113182, 10.1016/j.envres.2022.113182, 2022a.
- 574 Chen, L. H., Bao, Z. E., Wu, X. C., Li, K. W., Han, L. X., Zhao, X. Y., Zhang, X., Wang, Z. H., Azzi, M.,
  575 and Cen, K. F.: The effects of humidity and ammonia on the chemical composition of secondary
  576 aerosols from toluene/NOx photo-oxidation, Sci Total Environ, 728, ARTN 138671,
  577 10.1016/j.scitotenv.2020.138671, 2020.





578	Chen, T., Zhang, P., Ma, Q., Chu, B., Liu, J., Ge, Y., and He, H.: Smog chamber study on the role of NOx
579	in SOA and O <sub>3</sub> formation from aromatic hydrocarbons, Environ. Sci. Technol., 56, 13654-13663,
580	10.1021/acs.est.2c04022, 2022b.
581 582 583	Chen, T. Z., Zhang, P., Chu, B. W., Ma, Q. X., Ge, Y. L., Liu, J., and He, H.: Secondary organic aerosol formation from mixed volatile organic compounds: Effect of RO2 chemistry and precursor concentration, NPJ Clim. Atmos. Sci., 5, 95, 10.1038/s41612-022-00321-y, 2022c.
584 585 586 587	Day, D. A., Campuzano-Jost, P., Nault, B. A., Palm, B. B., Hu, W., Guo, H., Wooldridge, P. J., Cohen, R. C., Docherty, K. S., Huffman, J. A., de Sá, S. S., Martin, S. T., and Jimenez, J. L.: A systematic re-evaluation of methods for quantification of bulk particle-phase organic nitrates using real-time aerosol mass spectrometry, Atmos. Meas. Tech., 15, 459-483, 10.5194/amt-15-459-2022, 2022.
588	Domínguez-López, D., Adame, J. A., Hernández-Ceballos, M. A., Vaca, F., De la Morena, B. A., and
589	Bolívar, J. P.: Spatial and temporal variation of surface ozone, NO and NO <sub>2</sub> at urban, suburban, rural
590	and industrial sites in the southwest of the Iberian Peninsula, Environ. Monit. Assess., 186, 5337-
591	5351, 10.1007/s10661-014-3783-9, 2014.
592 593 594 595	Eddingsaas, N. C., Loza, C. L., Yee, L. D., Chan, M., Schilling, K. A., Chhabra, P. S., Seinfeld, J. H., and Wennberg, P. O.: α-pinene photooxidation under controlled chemical conditions - Part 2: SOA yield and composition in low- and high-NOx environments, Atmos. Chem. Phys., 12, 7413-7427, 10.5194/acp-12-7413-2012, 2012.
596	Ehn, M., Thornton, J. A., Kleist, E., Sipila, M., Junninen, H., Pullinen, I., Springer, M., Rubach, F.,
597	Tillmann, R., Lee, B., Lopez-Hilfiker, F., Andres, S., Acir, I. H., Rissanen, M., Jokinen, T.,
598	Schobesberger, S., Kangasluoma, J., Kontkanen, J., Nieminen, T., Kurten, T., Nielsen, L. B.,
599	Jorgensen, S., Kjaergaard, H. G., Canagaratna, M., Maso, M. D., Berndt, T., Petaja, T., Wahner, A.,
600	Kerminen, V. M., Kulmala, M., Worsnop, D. R., Wildt, J., and Mentel, T. F.: A large source of low-
601	volatility secondary organic aerosol, Nature, 506, 476-479, 10.1038/nature13032, 2014.
602	Feng, T., Zhao, S. Y., Bei, N. F., Wu, J. R., Liu, S. X., Li, X., Liu, L., Qian, Y., Yang, Q. C., Wang, Y. C.,
603	Zhou, W. J., Cao, J. J., and Li, G. H.: Secondary organic aerosol enhanced by increasing atmospheric
604	oxidizing capacity in Beijing-Tianjin-Hebei (BTH), China, Atmos. Chem. Phys., 19, 7429-7443,
605	10.5194/acp-19-7429-2019, 2019.
606	Guenther, A. B., Jiang, X., Heald, C. L., Sakulyanontvittaya, T., Duhl, T., Emmons, L. K., and Wang, X.:
607	The Model of Emissions of Gases and Aerosols from Nature version 2.1 (MEGAN2.1): an extended
608	and updated framework for modeling biogenic emissions, Geosci. Model Dev., 5, 1471-1492,
609	10.5194/gmd-5-1471-2012, 2012.
610	Henry, K. M., Lohaus, T., and Donahue, N. M.: Organic aerosol yields from α-pinene oxidation: bridging
611	the gap between first-generation yields and aging chemistry, Environ. Sci. Technol., 46, 12347-
612	12354, 10.1021/es302060y, 2012.
613	Hui, L. R., Liu, X. G., Tan, Q. W., Feng, M., An, J. L., Qu, Y., Zhang, Y. H., and Jiang, M. Q.:
614	Characteristics, source apportionment and contribution of VOCs to ozone formation in Wuhan,
615	Central China, Atmos. Environ., 192, 55-71, 10.1016/j.atmosenv.2018.08.042, 2018.
616	Jiang, X. T., Liu, D. T., Xu, L., Tsona, N. T., and Du, L.: Assessing the influence of environmental
617	conditions on secondary organic aerosol formation from a typical biomass burning compound, J.
618	Environ. Sci., 114, 136-148, 10.1016/j.jes.2021.08.016, 2022.
619	Kelly, J. M., Doherty, R. M., O'Connor, F. M., and Mann, G. W.: The impact of biogenic, anthropogenic,
620	and biomass burning volatile organic compound emissions on regional and seasonal variations in
621	secondary organic aerosol, Atmos. Chem. Phys., 18, 7393-7422, 10.5194/acp-18-7393-2018, 2018.
622	Kiendler-Scharr, A., Mensah, A. A., Friese, E., Topping, D., Nemitz, E., Prevot, A. S. H., Äijälä, M.,
623	Allan, J., Canonaco, F., Canagaratna, M., Carbone, S., Crippa, M., Dall Osto, M., Day, D. A., De
624	Carlo, P., Di Marco, C. F., Elbern, H., Eriksson, A., Freney, E., Hao, L., Herrmann, H., Hildebrandt,
625	L., Hillamo, R., Jimenez, J. L., Laaksonen, A., McFiggans, G., Mohr, C., O'Dowd, C., Otjes, R.,
626	Ovadnevaite, J., Pandis, S. N., Poulain, L., Schlag, P., Sellegri, K., Swietlicki, E., Tiitta, P.,
627	Vermeulen, A., Wahner, A., Worsnop, D., and Wu, HC.: Ubiquity of organic nitrates from
628	nighttime chemistry in the European submicron aerosol, Geophys. Res. Lett., 43, 7735-7744,
629	10.1002/2016GL069239, 2016.





630	Kolesar, K. R., Chen, C., Johnson, D., and Cappa, C. D.: The influences of mass loading and rapid
631	dilution of secondary organic aerosol on particle volatility, Atmos. Chem. Phys., 15, 9327-9343,
632	10.5194/acp-15-9327-2015, 2015.
633 634 635 636 637 638 639	Krechmer, J. E., Coggon, M. M., Massoli, P., Nguyen, T. B., Crounse, J. D., Hu, W., Day, D. A., Tyndall, G. S., Henze, D. K., Rivera-Rios, J. C., Nowak, J. B., Kimmel, J. R., Mauldin, R. L., 3rd, Stark, H., Jayne, J. T., Sipila, M., Junninen, H., Clair, J. M., Zhang, X., Feiner, P. A., Zhang, L., Miller, D. O., Brune, W. H., Keutsch, F. N., Wennberg, P. O., Seinfeld, J. H., Worsnop, D. R., Jimenez, J. L., and Canagaratna, M. R.: Formation of low volatility organic compounds and secondary organic aerosol from isoprene hydroxyhydroperoxide low-NO oxidation, Environ. Sci. Technol., 49, 10330-10339, 10.1021/acs.est.5b02031, 2015.
640	Kroll, J. H., Ng, N. L., Murphy, S. M., Flagan, R. C., and Seinfeld, J. H.: Secondary organic aerosol
641	formation from isoprene photooxidation, Environ. Sci. Technol., 40, 1869-1877,
642	10.1021/es0524301, 2006.
643	Lane, T. E., Donahue, N. M., and Pandis, S. N.: Effect of NOx on secondary organic aerosol
644	concentrations, Environ. Sci. Technol., 42, 6022-6027, 10.1021/es703225a, 2008.
645	Laskin, J., Laskin, A., and Nizkorodov, S. A.: Mass spectrometry analysis in atmospheric chemistry, Anal.
646	Chem., 90, 166-189, 10.1021/acs.analchem.7b04249, 2018.
647	Li, R., Wang, X. F., Gu, R. R., Lu, C. Y., Zhu, F. P., Xue, L. K., Xie, H. J., Du, L., Chen, J. M., and Wang,
648	W. X.: Identification and semi-quantification of biogenic organic nitrates in ambient particulate
649	matters by UHPLC/ESI-MS, Atmos. Environ., 176, 140-147, 10.1016/j.atmosenv.2017.12.038,
650	2018.
651	Li, Y., Han, Z. W., Song, Y., Li, J. W., Sun, Y. L., and Wang, T. T.: Impacts of the COVID-19 lockdown
652	on atmospheric oxidizing capacity and secondary aerosol formation over the Beijing-Tianjin-Hebei
653	region in Winter-Spring 2020, Atmos. Environ., 295, 10.1016/j.atmosenv.2022.119540, 2023.
654	Ling, Z., Wu, L., Wang, Y., Shao, M., Wang, X., and Huang, W.: Roles of semivolatile and intermediate-
655	volatility organic compounds in secondary organic aerosol formation and its implication: A review,
656	J. Environ. Sci., 114, 259-285, 10.1016/j.jes.2021.08.055, 2022.
657	Liu, C. G., Liu, J., Liu, Y. C., Chen, T. Z., and He, H.: Secondary organic aerosol formation from the
658	OH-initiated oxidation of guaiacol under different experimental conditions, Atmos. Environ., 207,
659	30-37, 10.1016/j.atmosenv.2019.03.021, 2019a.
660	Liu, S. J., Jiang, X. T., Tsona, N. T., Lv, C., and Du, L.: Effects of NOx, SO <sub>2</sub> and RH on the SOA formation
661	from cyclohexene photooxidation, Chemosphere, 216, 794-804,
662	10.1016/j.chemosphere.2018.10.180, 2019b.
663 664 665 666	Liu, S. J., Huang, D. D., Wang, Y. Q., Zhang, S., Liu, X. D., Wu, C., Du, W., and Wang, G. H.: Synergetic effects of NH <sub>3</sub> and NOx on the production and optical absorption of secondary organic aerosol formation from toluene photooxidation, Atmos. Chem. Phys., 21, 17759-17773, 10.5194/acp-21-17759-2021, 2021a.
667 668 669	Liu, S. J., Wang, Y. Q., Wang, G. H., Zhang, S., Li, D. P., Du, L., Wu, C., Du, W., and Ge, S. S.: Enhancing effect of NO <sub>2</sub> on the formation of light-absorbing secondary organic aerosols from toluene photooxidation, Sci Total Environ, 794, 148714, 10.1016/j.scitotenv.2021.148714, 2021b.
670	Lv, S., Wang, F., Wu, C., Chen, Y., Liu, S., Zhang, S., Li, D., Du, W., Zhang, F., Wang, H., Huang, C.,
671	Fu, Q., Duan, Y., and Wang, G.: Gas-to-aerosol phase partitioning of atmospheric water-soluble
672	organic compounds at a rural site in china: An enhancing effect of NH <sub>3</sub> on SOA formation, Environ.
673	Sci. Technol., 56, 3915-3924, 10.1021/acs.est.1c06855, 2022.
674 675 676 677 678	<ul> <li>McFiggans, G., Mentel, T. F., Wildt, J., Pullinen, I., Kang, S., Kleist, E., Schmitt, S., Springer, M., Tillmann, R., Wu, C., Zhao, D., Hallquist, M., Faxon, C., Le Breton, M., Hallquist, A. M., Simpson, D., Bergstrom, R., Jenkin, M. E., Ehn, M., Thornton, J. A., Alfarra, M. R., Bannan, T. J., Percival, C. J., Priestley, M., Topping, D., and Kiendler-Scharr, A.: Secondary organic aerosol reduced by mixture of atmospheric vapours, Nature, 565, 587-593, 10.1038/s41586-018-0871-y, 2019.</li> </ul>
679	Middlebrook, A. M., Bahreini, R., Jimenez, J. L., and Canagaratna, M. R.: Evaluation of composition-
680	dependent collection efficiencies for the aerodyne aerosol mass spectrometer using field data,

26





681	Aerosol Sci. Tech., 46, 258-271, 10.1080/02786826.2011.620041, 2012.
682 683 684	Ng, N. L., Kroll, J. H., Chan, A. W. H., Chhabra, P. S., Flagan, R. C., and Seinfeld, J. H.: Secondary organic aerosol formation from <i>m</i> -xylene, toluene, and benzene, Atmos. Chem. Phys., 7, 3909-3922, 10.5194/acp-7-3909-2007, 2007.
685	Ng, N. L., Canagaratna, M. R., Zhang, Q., Jimenez, J. L., Tian, J., Ulbrich, I. M., Kroll, J. H., Docherty,
686	K. S., Chhabra, P. S., Bahreini, R., Murphy, S. M., Seinfeld, J. H., Hildebrandt, L., Donahue, N. M.,
687	DeCarlo, P. F., Lanz, V. A., Prévôt, A. S. H., Dinar, E., Rudich, Y., and Worsnop, D. R.: Organic
688	aerosol components observed in Northern Hemispheric datasets from Aerosol Mass Spectrometry,
689	Atmos. Chem. Phys., 10, 4625-4641, 10.5194/acp-10-4625-2010, 2010.
690	Ng, N. L., Brown, S. S., Archibald, A. T., Atlas, E., Cohen, R. C., Crowley, J. N., Day, D. A., Donahue,
691	N. M., Fry, J. L., Fuchs, H., Griffin, R. J., Guzman, M. I., Herrmann, H., Hodzic, A., Iinuma, Y.,
692	Jimenez, J. L., Kiendler-Scharr, A., Lee, B. H., Luecken, D. J., Mao, J., McLaren, R., Mutzel, A.,
693	Osthoff, H. D., Ouyang, B., Picquet-Varrault, B., Platt, U., Pye, H. O. T., Rudich, Y., Schwantes, R.
694	H., Shiraiwa, M., Stutz, J., Thornton, J. A., Tilgner, A., Williams, B. J., and Zaveri, R. A.: Nitrate
695	radicals and biogenic volatile organic compounds: oxidation, mechanisms, and organic aerosol,
696	Atmos. Chem. Phys., 17, 2103-2162, 10.5194/acp-17-2103-2017, 2017.
697	Nussbaumer, C. M., Pozzer, A., Tadic, I., Röder, L., Obersteiner, F., Harder, H., Lelieveld, J., and Fischer,
698	H.: Tropospheric ozone production and chemical regime analysis during the COVID-19 lockdown
699	over Europe, Atmos. Chem. Phys., 22, 6151-6165, 10.5194/acp-22-6151-2022, 2022.
700 701 702	Odum, J. R., Hoffmann, T., Bowman, F., Collins, D., Flagan, R. C., and Seinfeld, J. H.: Gas/particle partitioning and secondary organic aerosol yields, Environ. Sci. Technol., 30, 2580-2585, 10.1021/es950943+, 1996.
703	Qi, X., Zhu, S. P., Zhu, C. Z., Hu, J., Lou, S. R., Xu, L., Dong, J. G., and Cheng, P.: Smog chamber study
704	of the effects of NOx and NH <sub>3</sub> on the formation of secondary organic aerosols and optical properties
705	from photo-oxidation of toluene, Sci Total Environ, 727, ARTN 138632,
706	10.1016/j.scitotenv.2020.138632, 2020.
707	Rissanen, M.: Anthropogenic volatile organic compound (AVOC) autoxidation as a source of highly
708	oxygenated organic molecules (HOM), J. Phys. Chem. A, 125, 9027-9039,
709	10.1021/acs.jpca.1c06465, 2021.
710	Santos, F. M., Gómez-Losada, A., and Pires, J. C. M.: Empirical ozone isopleths at urban and suburban
711	sites through evolutionary procedure-based models, J. Hazard. Mater., 419, ARTN 126386,
712	10.1016/j.jhazmat.2021.126386, 2021.
713	Sarrafzadeh, M., Wildt, J., Pullinen, I., Springer, M., Kleist, E., Tillmann, R., Schmitt, S. H., Wu, C.,
714	Mentel, T. F., Zhao, D., Hastie, D. R., and Kiendler-Scharr, A.: Impact of NOx and OH on secondary
715	organic aerosol formation from β-pinene photooxidation, Atmos. Chem. Phys., 16, 11237-11248,
716	10.5194/acp-16-11237-2016, 2016.
717	Shon, Z. H., Lee, G., Song, S. K., Lee, M., Han, J., and Lee, D.: Characteristics of reactive nitrogen
718	compounds and other relevant trace gases in the atmosphere at urban and rural areas of Korea during
719	May-June, 2004, J Atmos. Chem., 58, 203-218, 10.1007/s10874-007-9088-4, 2007.
720 721 722	Sindelarova, K., Granier, C., Bouarar, I., Guenther, A., Tilmes, S., Stavrakou, T., Müller, J. F., Kuhn, U., Stefani, P., and Knorr, W.: Global data set of biogenic VOC emissions calculated by the MEGAN model over the last 30 years, Atmos. Chem. Phys., 14, 9317-9341, 10.5194/acp-14-9317-2014, 2014.
723 724 725	Takeuchi, M. and Ng, N. L.: Chemical composition and hydrolysis of organic nitrate aerosol formed from hydroxyl and nitrate radical oxidation of α-pinene and β-pinene, Atmos. Chem. Phys., 19, 12749-12766, 10.5194/acp-19-12749-2019, 2019.
726 727 728	Takeuchi, M., Berkemeier, T., Eris, G., and Ng, N. L.: Non-linear effects of secondary organic aerosol formation and properties in multi-precursor systems, Nat. Commun., 13, 10.1038/s41467-022-35546-1, 2022.
729	Tan, Z. F., Lu, K. D., Dong, H. B., Hu, M., Li, X., Liu, Y. H., Lu, S. H., Shao, M., Su, R., Wang, H. C.,

731 Wu, Y. S., Wahner, A., and Zhang, Y. H.: Explicit diagnosis of the local ozone production rate and the ozone-NOx-VOC sensitivities, Sci. Bull., 63, 1067-1076, 10.1016/j.scib.2018.07.001, 2018. 27





732	Valorso, R., Aumont, B., Camredon, M., Raventos-Duran, T., Mouchel-Vallon, C., Ng, N. L., Seinfeld,
733	J. H., Lee-Taylor, J., and Madronich, S.: Explicit modelling of SOA formation from α-pinene
734	photooxidation: sensitivity to vapour pressure estimation, Atmos. Chem. Phys., 11, 6895-6910,
735	10.5194/acp-11-6895-2011, 2011.
736	Wang, G., Zhang, R., Gomez, M. E., Yang, L., Levy Zamora, M., Hu, M., Lin, Y., Peng, J., Guo, S.,
737	Meng, J., Li, J., Cheng, C., Hu, T., Ren, Y., Wang, Y., Gao, J., Cao, J., An, Z., Zhou, W., Li, G.,

- 738 Wang, J., Tian, P., Marrero-Ortiz, W., Secrest, J., Du, Z., Zheng, J., Shang, D., Zeng, L., Shao, M., 739 Wang, W., Huang, Y., Wang, Y., Zhu, Y., Li, Y., Hu, J., Pan, B., Cai, L., Cheng, Y., Ji, Y., Zhang, F., 740 Rosenfeld, D., Liss, P. S., Duce, R. A., Kolb, C. E., and Molina, M. J.: Persistent sulfate formation 741 from London Fog to Chinese haze, Proc. Natl. Acad. Sci. U. S. A., 113, 13630-13635, 742 10.1073/pnas.1616540113, 2016.
- Wang, Y., Jin, X., Liu, Z., Wang, G., Tang, G., Lu, K., Hu, B., Wang, S., Li, G., An, X., Wang, C., Hu, 743 744 Q., He, L., Zhang, F., and Zhang, Y.: Progress in quantitative research on the relationship between 745 atmospheric oxidation and air quality, J. Environ. Sci., 123, 350-366, 10.1016/j.jes.2022.06.029, 746 2023
- 747 Xu, J. L., Griffin, R. J., Liu, Y., Nakao, S., and Cocker, D. R.: Simulated impact of NOx on SOA 748 formation from oxidation of toluene and m-xylene, Atmos. Environ., 101, 217-225, 749 10.1016/j.atmosenv.2014.11.008, 2015a.
- 750 Xu, L., Kollman, M. S., Song, C., Shilling, J. E., and Ng, N. L.: Effects of NOx on the volatility of 751 secondary organic aerosol from isoprene photooxidation, Environ. Sci. Technol., 48, 2253-2262, 752 10.1021/es404842g, 2014.
- 753 Xu, L., Guo, H. Y., Boyd, C. M., Klein, M., Bougiatioti, A., Cerully, K. M., Hite, J. R., Isaacman-754 VanWertz, G., Kreisberg, N. M., Knote, C., Olson, K., Koss, A., Goldstein, A. H., Hering, S. V., de 755 Gouw, J., Baumann, K., Lee, S. H., Nenes, A., Weber, R. J., and Ng, N. L.: Effects of anthropogenic 756 emissions on aerosol formation from isoprene and monoterpenes in the southeastern United States, Proc. Natl. Acad. Sci. U. S. A., 112, 37-42, 10.1073/pnas.1417609112, 2015b. 757
- 758 Xu, Z. N., Nie, W., Liu, Y. L., Sun, P., Huang, D. D., Yan, C., Krechmer, J., Ye, P. L., Xu, Z., Qi, X. M., 759 Zhu, C. J., Li, Y. Y., Wang, T. Y., Wang, L., Huang, X., Tang, R. Z., Guo, S., Xiu, G. L., Fu, Q. Y., 760 Worsnop, D., Chi, X. G., and Ding, A. J.: Multifunctional products of isoprene oxidation in polluted 761 atmosphere and their contribution to SOA, Geophys. Res. Lett., 48, ARTN e2020GL089276, 10.1029/2020GL089276, 2021. 762
- 763 Yu, S. S., Jia, L., Xu, Y. F., and Pan, Y. P.: Molecular composition of secondary organic aerosol from 764 styrene under different NOx and humidity conditions, Atmos Res, 266, ARTN 105950, 765 10.1016/j.atmosres.2021.105950, 2022.
- 766 Zhang, H., Yee, L. D., Lee, B. H., Curtis, M. P., Worton, D. R., Isaacman-VanWertz, G., Offenberg, J. H., 767 Lewandowski, M., Kleindienst, T. E., Beaver, M. R., Holder, A. L., Lonneman, W. A., Docherty, K. 768 S., Jaoui, M., Pye, H. O. T., Hu, W., Day, D. A., Campuzano-Jost, P., Jimenez, J. L., Guo, H., Weber, 769 R. J., de Gouw, J., Koss, A. R., Edgerton, E. S., Brune, W., Mohr, C., Lopez-Hilfiker, F. D., Lutz, 770 A., Kreisberg, N. M., Spielman, S. R., Hering, S. V., Wilson, K. R., Thornton, J. A., and Goldstein, A. H.: Monoterpenes are the largest source of summertime organic aerosol in the southeastern 771 772 United States, Proc. Natl. Acad. Sci. U. S. A., 115, 2038-2043, 10.1073/pnas.1717513115, 2018.
- 773 Zhao, D., Schmitt, S. H., Wang, M., Acir, I. H., Tillmann, R., Tan, Z., Novelli, A., Fuchs, H., Pullinen, I., 774 Wegener, R., Rohrer, F., Wildt, J., Kiendler-Scharr, A., Wahner, A., and Mentel, T. F.: Effects of 775 NOx and SO<sub>2</sub> on the secondary organic aerosol formation from photooxidation of  $\alpha$ -pinene and 776 limonene, Atmos. Chem. Phys., 18, 1611-1628, 10.5194/acp-18-1611-2018, 2018.

777